1 Size distribution and mixing state of black carbon particles during a heavy air

2 pollution episode in Shanghai

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Abstract: A Single Particle Aerosol Mass Spectrometer (SPAMS), a Single Particle Soot 12 13 Photometer (SP2) and various meteorological instruments were employed to investigate the chemical and physical properties of black carbon (BC) aerosols during a regional air 14 pollution episode in urban Shanghai over a five-day period in December 2013. The 15 refractory black carbon (rBC) mass concentrations measured by SP2 averaged 3.2 μ g m⁻³, 16 with the peak value of 12.1 μ g m⁻³ at 04:26 LT on 7 December. The number of 17 BC-containing particles captured by SPAMS in the size range 200-1200 nm agreed very 18 well with that detected by SP2 ($R^2 = 0.87$). A cluster analysis of the single particle mass 19 spectra allowed for the separation of BC-containing particles into five major classes: (1) 20 21 Pure BC; (2) BC attributed to biomass burning (BBBC); (3) K-rich BC-containing (KBC); (4) BC internally-mixed with OC and ammonium sulfate (BCOC-SOx); (5) BC 22 23 internally-mixed with OC and ammonium nitrate (BCOC-NOx). The size distribution of 24 internally-mixed BC particles was bimodal. Detected by SP2, the condensation mode 25 peaked around ~230 nm and droplet mode peaked around ~380 nm, with a clear valley in 26 the size distribution around ~320 nm. The condensation mode mainly consisted of traffic 27 emissions, with particles featuring a small rBC core (~60-80 nm) and a relatively thin 28 absolute coating thickness (ACT, ~50-130 nm). The droplet mode included highly aged 29 traffic emission particles and biomass burning particles. The biomass burning particles

had a larger rBC core (~80-130 nm) and a thick ACT (~110-300 nm). The highly aged traffic emissions had a smaller core (~60-80 nm) and a very thick ACT (~130-300 nm), which is larger than reported in any previous literature. A fast growth rate (~20 nm h⁻¹) of rBC with small core sizes was observed during the experiment. High concentrations pollutants like NO₂ likely accelerated the aging process and resulted in a continuous size growth of rBC-containing particles from traffic emission.

36 **1 Introduction**

Aerosols represent the largest uncertainty in estimating radiative forcing of atmospheric species, by strongly affecting the energy balance of the Earth by scattering and/or absorbing solar radiation (Pöschl, 2005), and influencing cloud formation (Jacobson, 2006). Emitted from incomplete combustion of fossil fuel and biomass (Bond et al., 2013), black carbon (BC) is a strongly light-absorbing carbonaceous material in aerosols, second to carbon dioxide as a contributor to positive radiative forcing (Ramanathan and Carmichael, 2008;Jacobson, 2001).

The physical (e.g., size distribution and morphology) and chemical (e.g., mixing state and 44 composition) properties of ambient BC are very complex and are constantly changing in 45 46 the atmosphere. For example, BC particles exposed to sub-saturated sulfuric acid vapor 47 exhibit a marked change in morphology, characterized by a decreased mobility-based diameter but an increased fractal dimension and effective density (Zhang et al., 2008). By 48 with a transmission electron microscope 49 using electron tomography and 50 three-dimensional (3-D) imaging, Adachi et al. (2010) found that many BC particles have 51 open, chainlike morphology even after being surrounded by organic matter, and are located in off-center positions within their host materials. China et al. (2013) analyzed the 52 53 morphology of single BC particles using electron microscopy and classified them into 54 four categories: ~50% were embedded (heavily coated), ~34% were partly coated, ~12% had inclusions and $\sim 4\%$ were bare. The organic coating is known to strongly affect the 55 optical properties of the soot aggregates by acting as a lens that amplifies the absorption 56 57 cross section of the BC core (Lack and Cappa, 2010; Shiraiwa et al., 2010). Schnaiter et al. 58 (2005) observed amplification factors of the internally-mixed BC of 1.8 to 2.1 relative to

59 the specific absorption cross section of externally-mixed BC. Zhang et al. (2008) 60 observed that the internally-mixed particles can increase their absorption efficiency by nearly 2-fold and scattering efficiency by approximately 10-fold at 80% relative humidity 61 62 relative to fresh particles. On the other hand, Cappa et al. (2012) and Lan et al. (2013) 63 observed a limited enhancement due to the mixing state of ambient BC, suggesting that 64 other factors may affect their absorption properties. Through coagulation and 65 condensation, BC can form an internal mixture, which increases its cloud nucleation activity (Khalizov et al., 2009; Moffet and Prather, 2009). Most BC is removed from the 66 troposphere via wet deposition with a short lifetime of 5 to 10 days (Kanakidou et al., 67 2005; Chung and Seinfeld, 2002). 68

Many measurement methods for refractory BC (rBC) particles have been developed and 69 used in recent years (Petzold et al., 2013). Among them, the Single Particle Soot 70 71 Photometer (SP2) has become increasingly recognized as a valuable tool for 72 characterizing rBC-containing particles (Stephens et al., 2003;Schwarz et al., 2006). SP2 73 can quantitatively measure the mass and determine the mixing state of an individual 74 rBC-containing particle (Schwarz et al., 2010). Taylor et al. (2014) evaluated the 75 capability of the SP2 to determine the particle mixing state with help of the concentric 76 core/shell model. Liu et al. (2014) analyzed the size distribution and mixing state of rBC 77 aerosols in London during winter time using the same technique. Furthermore, Moteki et 78 al. (2014) identified two morphological types of the mixed rBC-containing particles as 79 attached and coated, an important finding for understanding the climate impact of rBC 80 particles. Recently, a soot particle aerosol mass spectrometer (SP-AMS) was developed 81 to characterize rBC and non-refractory particulate matter simultaneously (Cross et al., 82 2010;Onasch et al., 2012;Corbin et al., 2014). SP-AMS was previously used to quantify rBC mass concentration, mixing state and chemical composition in urban environment 83 and biomass burning influenced air (Lee et al., 2015a;Lee et al., 2015b;Willis et al., 84 85 2015).

As a highly complementary instrument, single particle aerosol mass spectrometer (SPAMS, not to be confused with the SP-AMS instrument mentioned above) can detect

88 the chemical properties of BC particles. Moffet and Prather (2009) observed a rapid 89 coating process of organic carbon and sulfate on the BC core and assessed the related 90 absorption enhancement during an air pollution episode of the Mexico City. Healy et al. 91 (2012b) found that the mass size distribution for BC-containing particles was bimodal at 92 an urban background site in Paris. The smaller mode was attributed to local emission, 93 mostly externally-mixed BC particles, while the larger mode was dominated by aged 94 particles associated with continental transport events. Zhang et al. (2014) found that an 95 active photochemical formation of secondary organic aerosol (SOA) led to a distinct 96 diurnal pattern of mixing state of BC with SOA in the condensation mode, while the 97 photochemical aging had limited or negligible influence on the mixing state and growth 98 of BC in the droplet mode. The size ranges of condensation mode (Vacuum Aerodynamic Diameter (D_{va}) = ~100-300 nm) and droplet mode (D_{va} = ~300- 1000 nm) were defined by 99 100 John et al. (1990) and Seinfeld and Pandis (2012).

Depending on the experimental method, different terms are used in the literature for the most refractory and light-absorbing components of carbonaceous aerosols: black carbon (BC), refractory black carbon (rBC) and elemental carbon (EC). The definitions of BC, rBC and EC have been discussed in details elsewhere (Bond and Bergstrom, 2006;Almeida et al., 2013;Petzold et al., 2013). In this paper, we use rBC and BC to illustrate the SP2 and SPAMS data, respectively.

107 All of the studies mentioned above relied on either an SP2 instrument or a single particle 108 aerosol mass spectrometer to characterize BC particles, but not both. Combining these 109 two methods would provide the chemical and physical prosperities of individual BC 110 particles simultaneously and greatly enhance our understanding of their sources and 111 evolution processes. Furthermore, most previous SP2 studies focused on the rBC particles during relatively clean days. Quantitative analysis on the mixing state of BC 112 particles during heavy pollution episodes is still lacking. In this study, we deployed two 113 114 complementary techniques, with single particle resolution and high time resolution, to 115 detect the evolution of the urban BC aerosols in Shanghai during an extreme pollution period. We used an SP2 instrument to measure the mass and size distribution, and the 116

mixing state of individual rBC particles. A SPAMS instrument was used in parallel to
record chemical characteristics and mixing state of individual BC particles.

119 **2 Experimental**

120 **2.1 Single Particle Soot Photometer**

121 **2.1.1 Description**

122 The number and mass size distribution, as well as the mixing state of individual rBC particles were characterized using a single particle soot photometer (SP2, Droplet 123 124 Measurement Technologies, Inc., Boulder, CO) (Stephens et al., 2003; Baumgardner et al., 2004). In brief, SP2 detects incandescence and scattering signals of rBC-containing 125 particles induced by a 1064 nm Nd: YAG intra-cavity laser. The mass of rBC is 126 127 proportional to the intensity of the laser induced incandescence signal. Any measured 128 particle with a detectable incandescence signal is treated as an rBC particle; whereas a particle that only exhibits scattering signal is considered as a non-rBC particle. The total 129 rBC mass loading is reported as the sum of all detected single rBC masses. The SP2 130 instrument samples at a low flow rate (30 cm³ min⁻¹) in order to avoid multiple particles 131 crossing the laser at the same time. We only saved data for every 50th particle in order to 132 extend the sampling time without generating excessively large data sets. 133

134 **2.1.2** Calibration and detection efficiency

The SP2 incandescence signal was calibrated using Aquadag[®] black carbon particles 135 136 (Aqueous Deflocculated Acheson Graphite, manufactured by Acheson Inc., USA). The 137 Aquadag[®] black carbon particles were selected by mobility diameter using a differential mobility analyzer (DMA) and the corresponding particle masses were calculated using 138 139 the effective density data provided in Gysel et al. (2011) (Fig. S1). The scattering signal was calibrated using mono-disperse polystyrene latex spheres (Nanosphere Size 140 Standards, Duke Scientific Corp., Palo Alto, CA, USA) with known diameters (80-350 141 142 nm). More details about the SP2 calibration can be found in Gysel et al. (2011), Baumgardner et al. (2012) and Laborde et al. (2012). A diagram of the calibration systemis shown in the supplement (Fig. S2).

The detection efficiency was measured using Aquadag® black carbon particles, and the 145 146 results are shown in Fig. S3. The details of the measurement method were described in 147 Schwarz et al. (2010). SP2 detection efficiency was nearly unity for larger rBC particles. The minimum rBC mass that could be observed with near-unity detection efficiency was 148 149 ~ 0.7 fg rBC, corresponding to 90 nm mass-equivalent diameter; the detection efficiency 150 declined rapidly at lower sizes (Fig. S3). The total ambient mass concentrations of rBC 151 were underestimated because of the low detection efficiency of the smaller rBC particles, 152 likely by ~20 % (Schwarz et al., 2006;McMeeking et al., 2010). During the calibration and sampling time, the SP2 was operated at a stable temperature of 20 °C and pressure 153 154 of ~1013 hPa. The SP2 laser current was around 1750 mA through the whole experiment.

155 **2.1.3 Data analysis**

156 The rBC mass in each individual particle was determined from the peak intensity of the incandescence signal according to the Aquadag[®] black carbon calibration (Sect. 2.1.2). 157 158 The conversion from the mass to the effective rBC core diameter requires making 159 assumptions about the morphology and effective density of the rBC cores in the particles. 160 Zhang et al. (2015) recently found the ambient rBC cores had an average shape factor of 1.2 and an average density of 1.2 g cm⁻³, suggesting a near- spherical shape with an 161 internal void of 30 % by volume. With the aging process, Zhang et al. (2015) observed 162 163 that the effective density of BC core increased and the cores transformed to a more compact shape. In this study, we focused on the aged BC-containing particles. Therefore, 164 a density of 1.8 g cm⁻³ was used to convert the ambient rBC mass to the mass equivalent 165 166 diameter. This value was also recommended in many previous studies (Bond and 167 Bergstrom, 2006; Moteki and Kondo, 2010; Moteki et al., 2010; McMeeking et al., 2011).

In addition to the rBC mass, the measurement of the scattering signal of an
rBC-containing particle allows for the determination of its scattering cross section.
However, the scattering properties of externally- and internally-mixed rBC particles, as

171 detected by the SP2, may be distorted, because the mass of each particle is reduced by the laser heating. Thus, scattered light from a sampled rBC particle does not yield a full 172 173 Gaussian waveform. The Gaussian scattering function was reconstructed from the leading 174 edge of the scattering signal (before the particle is perturbed by the laser), which was 175 measured with a two-element avalanche photodiode (APD). This method allows SP2 to 176 determine the scattering properties of individual rBC particles as well as the rBC mass 177 and to distinguish the mixing state of a single rBC particle (so called, LEO-fit method 178 (Gao et al., 2007)). The optical diameter of a rBC particle or the coated rBC size (D_p) was 179 derived by inputting the LEO fitted scattering signal and rBC core size (D_c) into Mie 180 calculations with a core refractive index m = 2.26-1.26i (Moteki et al., 2010;Liu et al., 181 2014;Laborde et al., 2013) and a coating refractive index m = 1.5 + 0i (Laborde et al., 182 2013). The absolute coating thickness (ACT) of an rBC particle was calculated as $(D_p - D_p)$ 183 D_c)/2, based on the assumption of a concentric core-shell morphology. However, rBC aging processes in the real atmosphere may result in aerosols with particles that deviate 184 185 from the core-shell morphology (Matsui et al., 2013). For example, when a small rBC particle coagulates with a relatively large rBC-free particle, the small rBC particle may 186 187 stay at the surface and lead to an effective negative coating thickness if determined by the 188 methods used here. In this study, the negative coating thickness was observed for less than 2% of all rBC-containing particles; we did not take those particles into account when 189 190 we calculated the average ACT. More details of data analysis and uncertainties are 191 discussed in supplement, as well as in Liu et al. (2014) and Laborde et al. (2013).

192 2.2 Single Particle Aerosol Mass Spectrometer

A SPAMS instrument (Hexin Analytical Instrument Co., Ltd., Guangdong, China) was deployed simultaneously with SP2 to detect chemical composition of BC-containing particles. The technical details of SPAMS have been described elsewhere (Li et al., 2011). Briefly, aerosols in the size range of 0.2–2.0 μ m are introduced into the focus lens through a 0.1 mm critical orifice at a flow of 80 mL min⁻¹ due to the pressure drop from ~760 to ~2.2 Torr. Then particles are accelerated to a terminal size-dependent aerodynamic velocity, which is measured by two orthogonally-oriented continuous lasers

200 (532 nm) separated by a fixed 6.0 cm distance. A pulsed desorption/ionization laser (Qswitched Nd: YAG laser, 266 nm) is triggered when a particle arrives at the ion source 201 202 region. Both positive and negative ions are detected simultaneously by the time-of-flight 203 mass spectrometer. In this work, the power of the desorption/ionization laser was kept at 204 ~ 0.6 mJ per pulse. The particle size was calculated from the measured speed using a 205 calibration curve generated for mono-disperse polystyrene latex spheres (Nanosphere 206 Size Standards, Duke Scientific Corp., Palo Alto, CA, USA) with known diameters 207 $(0.22-2.00 \ \mu m).$

208 All single particle mass spectra were converted into a list of peaks at each m/z using TSI 209 MS-Analyze software with a minimum signal threshold of 30 arbitrary units above the baseline. The resulting peak lists were then imported into YAADA (www.yaada.org), a 210 211 software toolkit in Matlab (version R2012b) for further analysis of particle sizes and 212 chemical components. A total of 385 683 particles were chemically analyzed with both positive and negative ion spectra, accounting for about 56 % of all sized particles. Based 213 214 on the similarities of the mass-to-charge ratio and peak intensity, particles were clustered 215 by using the ART-2a (adaptive resonance theory) method (Song et al., 1999) with a 216 vigilance factor of 0.85, a learning rate of 0.05 and 20 iterations. Then BC-containing particles, with D_{va} in the size range of 200-1200 nm, were chosen from the clusters, since 217 this size range was consistent with the dominant fraction of BC-containing mass in the 218 atmosphere (Zhang et al., 2014; Healy et al., 2012b). Finally, a total of 86 057 219 220 BC-containing particles were grouped into six general particle types according to mass 221 spectral patterns.

222 **2.3 Monitor for AeRosols and GAses (MARGA)**

A MARGA instrument (ADI 2080, Applikon Analytical B. B. Corp., Netherlands) was used to measure water-soluble inorganic ions in particles. The details of MARGA have been described previously (Jongejan et al., 1995;Du et al., 2011). Briefly, air to be analyzed enters into sample boxes via a PM_1 cyclone. The air flow is maintained at 1 m³ h⁻¹ by a mass flow controlled air pump. In the sample box, water-soluble gases (HCl, HONO, SO₂, HNO₃, NH₃) are completely absorbed in a dilute solution of hydrogen peroxide by using a wet rotating denuder (WRD). Aerosols pass through the WRD and are subsequently collected in a steam-jet aerosol collector (SJAC). The two liquid samples with absorbed gases and particles are accumulated in syringes in the analytical box. After filling the syringes for one hour, the samples are then injected into an ion chromatograph (IC). The IC is continuously controlled by an internal calibration method using a standard LiBr solution. In this study, the water-soluble inorganic ions (i.e., K⁺, SO₄²⁻, NO₃⁻) in bulk particles were analyzed.

236 2.4 Sampling period and site

The sampling lasted for almost 5 days, from 5 to 10 December, 2013. The instruments 237 238 were operated in the building of the Department of Environmental Science and Engineering, Fudan University (FDU, 31° 14' N, 121° 29' E) in urban Shanghai, close 239 to both residential and traffic emissions sources. Aerosols were sampled with a PM_{25} 240 241 cyclone positioned 2 m above the roof of the building and transferred to the instruments 242 through a 6 m long stainless steel pipe (45 mm inner diameter). A pump was used to pull air through the sampling system at 30 L min⁻¹, minimizing the particle residence time in 243 the sampling line. Aerosols were dried by diffusion drying tubes before they reached the 244 245 SP2 and SPAMS inlets, which were connected in parallel. The measurement system is 246 presented in Fig. S2. Because of the extremely high particle mass loading, the inlets of 247 SP2 and SPAMS were clogged two times during the sampling period.

248 **3 Result and discussion**

3.1 Overview of the meteorology and air quality

250 Temporal variations of measured relative humidity, temperature, CO, O₃, NO, NO₂, SO₂, PM_{2.5} and PM₁₀ in Shanghai from 12:00 LT on 5 December to 14:00 LT on 10 December 251 252 are shown in Fig. 1. The meteorology and air quality information were provided by the 253 Shanghai Environmental Monitoring Center, Hongkou Station 254 (http://www.semc.com.cn/aqi/home/Index.aspx). The station is 3.3 km north from the 255 sampling site. The temperature and relative humidity varied between 2-19 $^{\circ}$ C and 256 30-100%, with an average of 19°C and 73%, respectively, during the study. The O_3 concentration was relatively low from 18:00 LT on 5 December to 8:00 LT on 7 257

December. The CO concentration showed two peaks during this period, and its peak 258 value reached 4.1 mg m⁻³ at 14:00 LT on 6 December. The NO₂ concentration increased 259 quickly at the beginning, reached 202.5 µg m⁻³ at 21:00 LT on 5 December, and then 260 decreased slowly until 12:00 LT on 7 December. After 12:00 LT on 7 December, the 261 concentrations of O₃, CO, NO₂ and SO₂ fluctuated without an obvious pattern. The 262 concentrations of O₃ and NO₂ showed the expected anti-correlation, because NO was 263 oxidized to NO₂ by O₃. The CO concentration was found to correlate reasonably well 264 with rBC mass concentration ($R^2 = 0.59$, slope=0.33), as shown in Fig. S4. 265

The mass loading of PM_{2.5} was extremely high during this period. Its maximum value reached 636 μ g m⁻³ at 12:00 LT on 6 December, which was a record-breaking hourly concentration for Shanghai. The daily average concentration was 221 μ g m⁻³. Meanwhile, PM₁₀ varied from 47 to 691 μ g m⁻³, with an average of 252 μ g m⁻³. Concentrations of CO, O₃, NO, NO₂, SO₂, PM₁₀ and PM_{2.5} during 5-10 December all exceeded the Chinese national ambient air quality standards.

272 **3.2 BC size distributions and concentration measurement by SP2**

273 We fitted a log-normal distribution to the rBC core number and mass size measurements 274 during the entire sampling period, as shown in Fig. 2. The number size distribution spanned the range from ~ 60 to ~ 400 nm and the peak was around ~ 60 nm. The measured 275 276 number concentrations dropped below 60 nm because the SP2 detection efficiency 277 greatly decreases (Sect. 2.1.2) below this particle size. Using the same method, Schwarz 278 et al. (2008) also found that the peak concentration was around 60 nm in boundary layer. 279 The rBC core mass size distribution had a peak around 200 nm, and the majority of the 280 rBC mass was distributed between 70-500 nm.

As shown in Fig. 1, the rBC mass concentration varied from 0.6 μ g m⁻³ at 00:02 LT on 10 December to 12.1 μ g m⁻³ at 04:26 LT on 7 December, with an average of 3.2 μ g m⁻³. The rBC mass concentration observed in Shanghai was similar to other cities in China, e.g., ~4.1 μ g m⁻³ in Shenzhen (Huang et al., 2012) and ~3.3 μ g m⁻³ in Kaiping (Huang et al., 2011). However, it was much higher than in other mega-cities around the world, e.g., $\sim 0.9 \ \mu g \ m^{-3}$ in Paris (Laborde et al., 2013) and $\sim 1.3 \ \mu g \ m^{-3}$ in London (Liu et al., 2014). All of the values quoted above were based on SP2 measurement so a direct comparison is possible. The rBC mass accounted for 1.45% of PM_{2.5} mass on average in our measurements.

3.3 BC particles classification by SPAMS

Classification of particles analyzed by the SPAMS can help elucidate the sources, degree of aging, and mixing state of BC particles. We classified BC-containing particles into six groups according to their mass spectral characteristics. The names of these groups and their number fractions are shown in Table 1. The average mass spectral patterns of each group are shown in Fig. S5.

Pure BC particles only presented strong signals for black carbon fragment ions (C_n^- and C_n^+) in both positive and negative ion mass spectra without any signal of secondary species like sulfate or nitrate, suggesting they were fresh BC particles that had not undergone any aging process.

300 Biomass burning BC-containing (BBBC) particles were characterized by an intense K⁺ signal for +39 (the charge and m/z of the observed ion), +113 (K₂Cl⁺) and +213 (K₃SO₄⁺) 301 in the positive ion mass spectra and a strong signal for -26 (CN⁻) and -42 (CNO⁻) in the 302 negative ion mass spectra. A significant fragment of levoglucosan, -71 (C₃H₃O₂), was 303 also observed. Typical black carbon fragments (C_n) appeared in the negative ion mass 304 spectra. A high signal at -46 (NO₂), -62 (NO₃) and a relatively low signal at -97 (HSO₄) 305 306 were also observed, suggesting a significant accumulation of nitrate ions on BC particles 307 throughout the air pollution period. The criteria for the identification of BBBC particles is discussed in supplementary section. Potassium-containing soot is a well-established 308 tracer for biomass combustion (Andreae, 1983;Soto-Garcia et al., 2011). Water-soluble 309 310 K⁺ in ambient particles measured by an online MARGA method correlated reasonably well with the BBBC particles number ($R^2=0.64$), as shown in Fig. S6. Particles with 311 312 similar mass spectral patterns were previously observed in several urban field studies and assigned to biomass burning sources (Moffet et al., 2008;Healy et al., 2012b;Bi et al., 313

314 2011).

315 BC internally-mixed with organic carbon and ammonium sulfate (BCOC-SOx) particles 316 exhibited signals for ammonium +17 (NH₃⁺), +18 (NH₄⁺), organic carbon +37 (C₃H⁺), +43 (CH₃CO⁺), +50 (C₄H₂⁺), +51 (C₄H₃⁺), +61 (CH₃C(OH)=OH⁺), +62 ((CH₃)₂NHOH⁺), 317 318 and a small signal for sodium +23 (Na⁺) in the positive ion mass spectra, along with black carbon fragment ions (C_n^+). There was a high signal for sulfate -97 (HSO₄⁻) and a 319 320 relatively low signal for nitrate -46 (NO₃), -62 (NO₃) in the negative ion spectra. BC 321 internally-mixed with organic carbon and ammonium nitrate (BCOC-NOx) particles are 322 characterized by very similar positive ion mass spectra to BCOC-SOx, but exhibit lower 323 signals for sulfate and higher signals for nitrate in the negative ion spectra, i.e., -46 324 (NO_3) , -62 (NO_3) . BC particles with various intensities for organic carbon, nitrate and 325 sulfate were commonly detected in urban ATOFMS field studies (Moffet et al., 326 2008; Ault et al., 2009; Dall'Osto and Harrison, 2006) and were assigned to traffic 327 emissions (Healy et al., 2012b).

328 K-rich BC-containing (KBC) particles exhibited strong signals for black carbon fraction 329 in both positive and negative ion mass spectra. This class also had signals for potassium 330 +39 (K⁺), sodium +23 (Na⁺) and ammonium +17 (NH₃⁺), +18 (NH₄⁺) in positive ion mass spectra and nitrate -46 (NO₂), -62 (NO₃), and sulfate -97 (HSO₄) in the negative ion 331 332 mass spectra. This class was detected from diesel vehicle emissions in a previous study 333 by (Li et al., 2013). The KBC exhibited pronounced diurnal variation, with two major peaks during early morning (4:00-7:00 LT) and night hours (20:00-22:00 LT) (Fig. S7). 334 335 Shanghai municipal government regulates that the heavily loaded diesel trucks cannot go into downtown area from 7:00- 20:00 LT. The diurnal variation of KBC is consistent 336 337 with the traffic flow of diesel trucks based on our results.

NOx can be used as a tracer of local traffic emissions in urban areas. In this study, the NOx concentrations agreed well with the sum of KBC, BCOC-NOx and BCOC-SOx particles numbers ($R^2=0.65$) (Fig. S8). Based on the above analysis, we believed that the BBBC came from biomass burning, KBC, and BCOC-NOx and BCOC-SOx came from traffic emissions. We should note that SPAMS preferentially detected internally-mixed BC particles, and had reduced detection efficiency for pure BC particles. The particles detected and chemically analyzed by SPAMS range from 200 to 2000 nm in size, and the detection efficiency decreases rapidly below 400 nm and above 1200 nm (Li et al., 2011). The majority of the pure BC particles diameter are smaller than 200 nm in diameter (Kondo et al., 2006), and therefore, they are missed by SPAMS.

349 **3.4 Mixing state and size distribution of internally-mixed BC particles**

350 **3.4.1 Temporal variations of internally-mixed BC particles**

A comparison of the internally-mixed BC particles number concentration between SP2 351 and SPAMS is given in Fig. 3. The agreement observed is reasonably good ($R^2 = 0.87$) 352 considering the combined experimental uncertainties of the methods and the different 353 cut-off diameters of SP2 ($D_p > 170$ nm) and SPAMS (200 nm $< D_{va} < 1200$ nm). Detected 354 by SP2, the internally-mixed rBC particles accounted for approximately 70% number 355 fraction of BC-containing particles during the whole period. Moteki et al. (2007) also 356 357 found the internally-mixed rBC particles accounted for 63% number fraction of BC-containing particles in the aged urban plume. The high correlation coefficient 358 359 indicates that we can use the two complementary techniques to analyze the mixing state 360 and chemical composition of internally-mixed BC particles with single particle resolution at the same time (although not for the same particle since both methods are destructive). 361

The temporal variation of number size distribution and particle types changed rapidly and 362 intricately, as shown in Fig. 4. From 12:00 LT on 5 December to 00:00 LT on 7 363 December, the PM_{2.5} and rBC mass increased slowly to an extremely polluted state. The 364 number fraction of BBBC particles also increased during this period (Fig. 4(b)) and D_p of 365 rBC showed two distinct modes (Fig. 4(a)). Then, the BC-containing particles number 366 367 increased rapidly at 02:00 LT on 7 December. Presumably, boundary layer compression 368 during the night led the fast change of BC-containing particles. After that, the number concentration of BC-containing particles exhibited diurnal variation, with two major 369 370 peaks at the rush hours, i.e., from 8:00-12:00 LT or from 16:00-20:00 LT.

371 3.4.2 Size distribution and source apportionments of internally-mixed BC 372 particles

Fig. 5(a) shows the entire diameter (D_p) number size distribution histogram of 373 374 internally-mixed rBC particles detected by SP2 during the entire sampling period. The 375 BC-containing particles were detected in both the condensation and droplet modes in this 376 study. Since the two modes are overlapped, it is possible that some of the particles in the 377 droplet mode are from the tail of condensation mode. Here, we used the minimum value between the two peaks in Fig. 5(a) (black line) and its corresponding D_p (320 nm) as the 378 separation of the condensation mode and droplet mode. We added this separation line in 379 380 Fig. 5(c) to separate the condensation mode particles (left side) and the droplet mode particles (right side). The condensation mode peak was centered around ~230 nm and 381 382 droplet mode peak was centered around ~380 nm. The presence of condensation mode $(D_{va} = \sim 200-500 \text{ nm})$ and droplet mode $(D_{va} = \sim 550-1200 \text{ nm})$ was confirmed by the 383 SPAMS data (Fig. 5(b)). Here the SPAMS size distribution was based on the number 384 385 fraction of BC-containing particles in all detected particles. Similar particle size 386 distributions were also found in other studies in China (Huang and Yu, 2008; Zhang et al., 2014). 387

The specific composition in condensation and droplet modes were quite different (Fig.5 388 (b)). BBBC particles exhibited a higher number fraction in the droplet mode than in the 389 390 condensation mode. Ammonium nitrate can condense on particle surfaces during 391 atmospheric transport if sulfate is fully neutralized and excess ammonia is available 392 (Riemer et al., 2004). The sulfate condensation on BC surfaces occurs soon after the BC 393 emission, while ammonium nitrate condensation occurs over longer timescales during 394 transport (Healy et al., 2012a). In this work, most KBC and BBBC particles and all the BCOC-NOx particles showed stronger NO_3^- signals than SO_4^- signals (as shown in Fig. 395 396 S5), suggesting that most BC-containing were deeply aged. Based on the particle 397 classification and source apportionment analysis, the internally-mixed BC particles from 398 traffic emissions accounted for almost all of the particles observed in the condensation mode. However, the particle sources in the droplet mode were more diverse, includingtraffic emissions and biomass burning.

401 Previous studies revealed that different sources emit different core diameters for rBC-containing particles (Liu et al., 2014;Takahama et al., 2014;Reddington et al., 402 403 2013;Schwarz et al., 2008) and the aging processes affect the coating thickness (Laborde et al., 2013;Liu et al., 2014). We identified the sources and estimated aging process of 404 rBC-containing particles by using 2-D image plot "fingerprint" of D_c and absolute 405 coating thickness (ACT) information. Fig. 5(c) shows the dependence of ACT on D_c , 406 407 weighed by the number concentration. In the condensation mode, the particles were characterized by small D_c values (~60-80 nm) with thin ACT (~50-130 nm). In 408 combination with the SPAMS information, these particles with small D_c and thin ACT 409 should be mainly from the traffic sources (Fig. 5(b)). 410

411 However, the droplet mode was very different from the condensation mode and showed a diversity of sources. In the droplet mode, the "fingerprint" showed two peaks in the size 412 distribution. The first peak had small D_c values (~60-80 nm) and thick ACT (~130-300 413 414 nm). We assume that the rBC-containing particles in the first peak were from traffic 415 emissions. In previous studies, the particles associated with traffic emissions had small 416 core sizes and thin coating thickness (Laborde et al., 2013;Liu et al., 2014). However, in 417 this study, we found that the rBC-containing particles from traffic could be highly-aged, 418 resulting in a much thicker coating than previously observed. This could be because 419 polluted air masses promote faster rBC aging processes (Matsui et al., 2013). The second 420 peak showed larger D_c (~80-130 nm) and thick ACT (~110-300 nm). These particles 421 were presumably from biomass burning. It has been reported using SP2 measurements 422 that fresh biomass burning rBC particles are thickly coated (Schwarz et al., 2008;Sahu et al., 2012;Liu et al., 2014). 423

Since there was no clear-cut separation between traffic emissions and biomass burning rBC-containing particles in the droplet mode, it was hard to distinguish them when we just used the core and shell information from SP2 (Liu et al., 2014). We selected SP2-detected particles with larger core sizes (80-130 nm) and thicker coating (120-300

nm) and compared with the biomass burning particles number concentration from 428 SPAMS, as shown in Fig. S9. The good correlation ($R^2 = 0.71$) verified the conclusion 429 that the rBC-containing particles with larger cores and thicker coating were from biomass 430 431 burning. Even though these larger rBC-containing particles only accounted for less than 432 20% number fraction, they are likely to be more hygroscopic (Liu et al., 2013; Wang et al., 433 2014) and be scavenged by wet deposition (Moteki et al., 2012). Such particles will have 434 greater potential to enhance the semi-direct effect (Koch and Del Genio, 2010) through 435 interaction with cloud processes.

The diversity of sources of the droplet mode BC-containing particles was also detected in 436 437 SPAMS, as we discussed before. SPAMS data showed that the internally-mixed BC particles from traffic emissions were more abundant in the droplet mode than those form 438 439 biomass burning (Fig. 5(b)). However, the SP2 data showed that particles with a small core and thick ACT (major traffic emission) were less abundant than particles with a 440 larger core with thick ACT (major biomass burning) (Fig. 5(c)). As we discussed in part 441 442 2.1.2, rBC-containing particle with smaller cores are not efficiently detected by SP2, 443 which may result in an underestimation of the fraction of traffic emission of 444 rBC-containing particles in the droplet mode.

The aging of traffic-emitted rBC-containing particles during the heavy air pollution 445 episode (12:00 LT, 5 December 2013 - 12:00 LT, 7 December 2013) was elucidated 446 using the temporal variation of relative coating thickness (RCT, entire particle 447 diameter/rBC core diameter, D_p/D_c) of rBC-containing particles ($D_c = 60-80$ nm), as 448 shown in Fig. 6a. Note that we could only obtain optical sizing information from 449 450 sufficiently coated particles because of the SP2 minimal detectable optical diameter of 451 ~170 nm. From 16:00 LT to 22:00 LT on 5 December, the RCT of rBC-containing particles and PM_{2.5} concentration grew rapidly. Even though the SP2's inlet was blocked 452 from 23:00 LT on 5 December to 10:00 LT on 6 December due to the extremely high PM 453 454 mass loading, the data collected around that time suggest the rBC-containing particles 455 growth was continuous until 13:00 LT on 6 December. The absolute coating growth rate

456 was around 20 nm/hour during this period (16:00 LT, 5 December – 13:00 LT, 6
457 December).

458 Variations of the major chemical species in the vehicle-emitted BC-containing particles 459 (selected by SPAMS) were also analyzed. The relative peak areas of nitrate -63 (NO₃⁻) and organic carbon (i.e., +27 (C₂H₃⁺), +43 (CH₃CO⁺)) showed a relatively high level 460 during 16:00 LT on 5 December- 13:00 LT on 6 December (Fig. 6(b)). Guo et al. (2014) 461 462 observed that gaseous emissions of volatile organic compounds, nitrogen oxides from urban transportation and sulfur dioxide from region industry were responsible for large 463 secondary particle matter formation in Beijing. Fig. S10 shows the mass concentrations 464 of SO₂, NO₂ the mass ratio of NO₂/SO₂, MARGA-measured mass concentrations of 465 particulate sulfate and nitrate, and the mass ratio of NO_3^{-1}/SO_4^{-2} in PM₁ during the whole 466 sampling period. The average mass ratios of NO_2/SO_2 in gas phase and NO_3^2/SO_4^{2-} in 467 particles phase were 2.8 and 1.4 respectively. During the heavy air pollution episode 468 (12:00 LT, 5 December 2013 – 13:00 LT, 6 December 2013), both NO₂ and particulate 469 nitrate increased dramatically along with the traffic emitted BC particle growth (as shown 470 471 Fig. 6), while the SO_2 and particulate sulfate had a slight increase and remained at a relatively lower level. Apparently, the gas to particle conversion of NO₂ to nitrate played 472 a more important role than the condensation of SO₂ in the particle growth during this 473 pollution episode. In the previous field studies (Huebert et al., 1988; Yao et al., 2002), the 474 high mass ratio of NO_3^{-1}/SO_4^{-2} (>1.0) was regarded as a sign of dominant traffic emission. 475 Wang et al. (2015) found that the high mass ratio of NO₂/SO₂ resulting from traffic 476 477 emissions was a major reason in triggering the heavy haze in Shanghai. In this work, the evaluation of BC-containing particles also suggested that high concentrations of NO₂ and 478 possibly volatile organics and their transformations play a vital role for particle growth 479 480 and the increase of PM loading in urban area especially during a heavy pollution episode. 481 Reductions in the emissions of gaseous precursors are critical for remediation of the 482 severe urban haze pollution in China.

483 **4 Conclusions**

In this study, we characterized BC-containing particles during a heavy air pollution episode in Shanghai. The rBC mass loading in Shanghai was similar to other cities in China but much higher than in other mega-cities around the world, with an average of 3.2 μ g m⁻³ and the peak value of 12.1 μ g m⁻³ at 04:26 LT on 7 December 2013. The rBC mass accounted for 1.45% of PM_{2.5} mass on average. The number- and mass-weighted BC core size distributions were around ~60-400 and 70-500 nm, with peaks around ~60 and ~200 nm, respectively.

Using SPAMS, we classified the BC-containing particles into 6 groups, according to their mass spectral patterns. The pure BC particles accounted for 0.62% number fraction of BC-containing particles (although this number could be underestimated because of the low detection efficiency for pure BC in SPAMS). The BBBC particles from biomass burning accounted for 25.57%. The KBC, BCOC-NOx and BCOC-SOx from traffic emissions accounted for 70.18%. The remaining unidentified particles accounted for 3.63%.

498 The size distribution of internally-mixed rBC particles was bimodal. The condensation 499 mode mainly consisted of traffic emissions, which had a small core (~60-80 nm) with 500 thin ACT (~50-130 nm). The droplet mode included biomass burning and deeply aged traffic-emitted rBC-containing particles. The biomass burning particles had larger core 501 502 sizes (~80-130 nm) with thick ACT (~110-300 nm) and the highly aged traffic emissions 503 had small core sizes (~60-80 nm) with thick ACT (~130-300 nm). It is rare to see the traffic-emitted rBC growing so quickly to the droplet mode. The high concentration of 504 NO₂ and its rapid conversion to particulate nitrate accelerated the growth of 505 506 BC-containing particles and contributed to the high particle mass concentration during 507 this heavy air pollution episode.

The quantitative number and mass information provided by SP2 supplemented the SPAMS chemical analysis in the entire experiment. The two complementary techniques can detect the physical and chemical properties of BC aerosol with single particle resolution. The combined use of SP2 and SPAMS have great promise for wider applications in future atmospheric measurements.

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520 **References**

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522 Adachi, K., Chung, S. H., and Buseck, P. R.: Shapes of soot aerosol particles and implications

- for their effects on climate, Journal of Geophysical Research: Atmospheres, 115, D15206,
 10.1029/2009JD012868, 2010.
- 525 Almeida, J., Schobesberger, S., Kurten, A., Ortega, I. K., Kupiainen-Maatta, O., Praplan, A. P.,
- Adamov, A., Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Dommen, J., Donahue, N. M.,
- 527 Downard, A., Dunne, E., Duplissy, J., Ehrhart, S., Flagan, R. C., Franchin, A., Guida, R., Hakala, J.,
- 528 Hansel, A., Heinritzi, M., Henschel, H., Jokinen, T., Junninen, H., Kajos, M., Kangasluoma, J.,
- 529 Keskinen, H., Kupc, A., Kurten, T., Kvashin, A. N., Laaksonen, A., Lehtipalo, K., Leiminger, M.,
- 530 Leppa, J., Loukonen, V., Makhmutov, V., Mathot, S., McGrath, M. J., Nieminen, T., Olenius, T.,
- 531 Onnela, A., Petaja, T., Riccobono, F., Riipinen, I., Rissanen, M., Rondo, L., Ruuskanen, T.,
- 532 Santos, F. D., Sarnela, N., Schallhart, S., Schnitzhofer, R., Seinfeld, J. H., Simon, M., Sipila, M.,
- 533 Stozhkov, Y., Stratmann, F., Tome, A., Trostl, J., Tsagkogeorgas, G., Vaattovaara, P., Viisanen, Y.,
- 534 Virtanen, A., Vrtala, A., Wagner, P. E., Weingartner, E., Wex, H., Williamson, C., Wimmer, D., Ye,
- P. L., Yli-Juuti, T., Carslaw, K. S., Kulmala, M., Curtius, J., Baltensperger, U., Worsnop, D. R.,
 Vehkamaki, H., and Kirkby, J.: Molecular understanding of sulphuric acid-amine particle
 nucleation in the atmosphere, Nature, 502, 359-+, 10.1038/nature12663, 2013.
- Andreae, M. O.: Soot Carbon and Excess Fine Potassium: Long-Range Transport of
 Combustion-Derived Aerosols, Science, 220, 1148-1151, 10.1126/science.220.4602.1148,
 1983.
- Ault, A. P., Moore, M. J., Furutani, H., and Prather, K. A.: Impact of Emissions from the Los
 Angeles Port Region on San Diego Air Quality during Regional Transport Events,
 Environmental Science & Technology, 43, 3500-3506, 10.1021/es8018918, 2009.
- Baumgardner, D., Kok, G., and Raga, G.: Warming of the Arctic lower stratosphere by light
 absorbing particles, Geophysical Research Letters, 31, 10.1029/2003gl018883, 2004.
- 546 Baumgardner, D., Popovicheva, O., Allan, J., Bernardoni, V., Cao, J., Cavalli, F., Cozic, J.,
- 547 Diapouli, E., Eleftheriadis, K., Genberg, P. J., Gonzalez, C., Gysel, M., John, A., Kirchstetter, T. W.,
- 548 Kuhlbusch, T. A. J., Laborde, M., Lack, D., Muller, T., Niessner, R., Petzold, A., Piazzalunga, A.,
- 549 Putaud, J. P., Schwarz, J., Sheridan, P., Subramanian, R., Swietlicki, E., Valli, G., Vecchi, R., and

- Viana, M.: Soot reference materials for instrument calibration and intercomparisons: a
 workshop summary with recommendations, Atmos. Meas. Tech., 5, 1869-1887,
 10.5194/amt-5-1869-2012, 2012.
- 553 Bi, X. H., Zhang, G. H., Li, L., Wang, X. M., Li, M., Sheng, G. Y., Fu, J. M., and Zhou, Z.: Mixing 554 state of biomass burning particles by single particle aerosol mass spectrometer in the 555 of PRD, 45, urban China, Atmos. Environ., 3447-3453, area 556 10.1016/j.atmosenv.2011.03.034, 2011.
- 557 Bond, T. C., and Bergstrom, R. W.: Light absorption by carbonaceous particles: An 558 investigative review, Aerosol Sci. Technol., 40, 27-67, 10.1080/02786820500421521, 2006.
- Bond, T. C., Doherty, S. J., Fahey, D. W., Forster, P. M., Berntsen, T., DeAngelo, B. J., Flanner, M.
- 560 G., Ghan, S., Karcher, B., Koch, D., Kinne, S., Kondo, Y., Quinn, P. K., Sarofim, M. C., Schultz, M.
- 561 G., Schulz, M., Venkataraman, C., Zhang, H., Zhang, S., Bellouin, N., Guttikunda, S. K., Hopke, P.
- 562 K., Jacobson, M. Z., Kaiser, J. W., Klimont, Z., Lohmann, U., Schwarz, J. P., Shindell, D.,
- 563 Storelvmo, T., Warren, S. G., and Zender, C. S.: Bounding the role of black carbon in the
- 564 climate system: A scientific assessment, Journal of Geophysical Research-Atmospheres, 118,
 565 5380-5552, 10.1002/jgrd.50171, 2013.
- 566 Cappa, C. D., Onasch, T. B., Massoli, P., Worsnop, D. R., Bates, T. S., Cross, E. S., Davidovits, P.,
- 567 Hakala, J., Hayden, K. L., Jobson, B. T., Kolesar, K. R., Lack, D. A., Lerner, B. M., Li, S. M., Mellon,
- 568 D., Nuaaman, I., Olfert, J. S., Petaja, T., Quinn, P. K., Song, C., Subramanian, R., Williams, E. J.,
- and Zaveri, R. A.: Radiative Absorption Enhancements Due to the Mixing State of
 Atmospheric Black Carbon, Science, 337, 1078-1081, 10.1126/science.1223447, 2012.
- 571 China, S., Mazzoleni, C., Gorkowski, K., Aiken, A. C., and Dubey, M. K.: Morphology and mixing
 572 state of individual freshly emitted wildfire carbonaceous particles, Nature communications,
 573 4, 2013.
- 574 Chung, S. H., and Seinfeld, J. H.: Global distribution and climate forcing of carbonaceous
 575 aerosols, Journal of Geophysical Research: Atmospheres, 107, 4407,
 576 10.1029/2001JD001397, 2002.
- 577 Corbin, J. C., Sierau, B., Gysel, M., Laborde, M., Keller, A., Kim, J., Petzold, A., Onasch, T. B.,
 578 Lohmann, U., and Mensah, A. A.: Mass spectrometry of refractory black carbon particles
- from six sources: carbon-cluster and oxygenated ions, Atmos. Chem. Phys., 14, 2591-2603,

- 580 10.5194/acp-14-2591-2014, 2014.
- 581 Cross, E. S., Onasch, T. B., Ahern, A., Wrobel, W., Slowik, J. G., Olfert, J., Lack, D. A., Massoli, P.,
- 582 Cappa, C. D., Schwarz, J. P., Spackman, J. R., Fahey, D. W., Sedlacek, A., Trimborn, A., Jayne, J. T., 583 Freedman, A., Williams, L. R., Ng, N. L., Mazzoleni, C., Dubey, M., Brem, B., Kok, G., 584 Subramanian, R., Freitag, S., Clarke, A., Thornhill, D., Marr, L. C., Kolb, C. E., Worsnop, D. R., 585 and Davidovits, P.: Soot Particle Studies Instrument Inter-ComparisonProject Overview, 586 Aerosol Sci. Technol., 44, 592-611, 10.1080/02786826.2010.482113, 2010.
- 587 Dall'Osto, M., and Harrison, R. M.: Chemical characterisation of single airborne particles in 588 Athens (Greece) by ATOFMS, Atmos. Environ., 40, 7614-7631, 589 http://dx.doi.org/10.1016/j.atmosenv.2006.06.053, 2006.
- 590 Du, H., Kong, L., Cheng, T., Chen, J., Du, J., Li, L., Xia, X., Leng, C., and Huang, G.: Insights into 591 summertime haze pollution events over Shanghai based on online water-soluble ionic 592 of 45, composition aerosols, Atmos. Environ., 5131-5137, 593 http://dx.doi.org/10.1016/j.atmosenv.2011.06.027, 2011.
- 594 Gao, R., Schwarz, J., Kelly, K., Fahey, D., Watts, L., Thompson, T., Spackman, J., Slowik, J., Cross,
- 595 E., and Han, J.-H.: A novel method for estimating light-scattering properties of soot aerosols 596 using a modified single-particle soot photometer, Aerosol Sci. Technol., 41, 125-135, 2007.
- 597 Guo, S., Hu, M., Zamora, M. L., Peng, J., Shang, D., Zheng, J., Du, Z., Wu, Z., Shao, M., Zeng, L.,
- 598 Molina, M. J., and Zhang, R.: Elucidating severe urban haze formation in China, Proc. Nat. 599
- Acad. Sci., 111, 17373-17378, 2014.
- 600 Gysel, M., Laborde, M., Olfert, J. S., Subramanian, R., and Grohn, A. J.: Effective density of 601 Aquadag and fullerene soot black carbon reference materials used for SP2 calibration, 602 Atmos. Meas. Tech., 4, 2851-2858, 10.5194/amt-4-2851-2011, 2011.
- 603 Healy, R. M., Chen, Y., Kourtchev, I., Kalberer, M., O'Shea, D., and Wenger, J. C.: Rapid 604 Formation of Secondary Organic Aerosol from the Photolysis of 1-Nitronaphthalene: Role of
- 605 Naphthoxy Radical Self-reaction, Environmental Science & Technology, 46, 11813-11820,
- 606 10.1021/es302841j, 2012a.
- 607 Healy, R. M., Sciare, J., Poulain, L., Kamili, K., Merkel, M., Muller, T., Wiedensohler, A.,
- 608 Eckhardt, S., Stohl, A., Sarda-Esteve, R., McGillicuddy, E., O'Connor, I. P., Sodeau, J. R., and
- 609 Wenger, J. C.: Sources and mixing state of size-resolved elemental carbon particles in a

- European megacity: Paris, Atmospheric Chemistry and Physics, 12, 1681-1700,
 10.5194/acp-12-1681-2012, 2012b.
- Huang, X. F., and Yu, J. Z.: Size distributions of elemental carbon in the atmosphere of a
- 613 coastal urban area in South China: characteristics, evolution processes, and implications for
- 614 the mixing state, Atmospheric Chemistry and Physics, 8, 5843-5853, 2008.
- Huang, X. F., Gao, R. S., Schwarz, J. P., He, L. Y., Fahey, D. W., Watts, L. A., McComiskey, A.,
- 616 Cooper, O. R., Sun, T. L., Zeng, L. W., Hu, M., and Zhang, Y. H.: Black carbon measurements in
- 617 the Pearl River Delta region of China, Journal of Geophysical Research-Atmospheres, 116,
- 618 10.1029/2010jd014933, 2011.
- Huang, X. F., Sun, T. L., Zeng, L. W., Yu, G. H., and Luan, S. J.: Black carbon aerosol
- 620 characterization in a coastal city in South China using a single particle soot photometer,
- 621 Atmos. Environ., 51, 21-28, 10.1016/j.atmosenv.2012.01.056, 2012.
- Huebert, B., Mingxing, W., and Weixiu, L.: Atmospheric nitrate, sulfate, ammonium andcalcium concentrations in China, Tellus B, 40, 1988.
- 624Jacobson, M. Z.: Strong radiative heating due to the mixing state of black carbon in625atmosphericaerosols,Nature,409,695-697,
- 626 http://www.nature.com/nature/journal/v409/n6821/suppinfo/409695a0_S1.html, 2001.
- 627 Jacobson, M. Z.: Effects of Externally-Through-Internally-Mixed Soot Inclusions within
- 628 Clouds and Precipitation on Global Climate[†], The Journal of Physical Chemistry A, 110,
- 629 6860-6873, 10.1021/jp056391r, 2006.
- 630 John, W., Wall, S. M., Ondo, J. L., and Winklmayr, W.: MODES IN THE SIZE DISTRIBUTIONS OF
- 631 ATMOSPHERIC INORGANIC AEROSOL, Atmospheric Environment Part a-General Topics, 24,
- 632 2349-2359, 10.1016/0960-1686(90)90327-j, 1990.
- 633 Jongejan, P., Bai, Y., Veltkamp, A., Wye, G., and Slaninaa, J.: An Automated Field Instrument
- 634 for The Determination of Acidic Gases in Air, Int. J. Environ. Anal. Chem., 66, 241-251, 1995.
- 635 Kanakidou, M., Seinfeld, J. H., Pandis, S. N., Barnes, I., Dentener, F. J., Facchini, M. C., Van
- 636 Dingenen, R., Ervens, B., Nenes, A., Nielsen, C. J., Swietlicki, E., Putaud, J. P., Balkanski, Y.,
- 637 Fuzzi, S., Horth, J., Moortgat, G. K., Winterhalter, R., Myhre, C. E. L., Tsigaridis, K., Vignati, E.,
- 638 Stephanou, E. G., and Wilson, J.: Organic aerosol and global climate modelling: a review,
- 639 Atmos. Chem. Phys., 5, 1053-1123, 10.5194/acp-5-1053-2005, 2005.

- Khalizov, A. F., Zhang, R., Zhang, D., Xue, H., Pagels, J., and McMurry, P. H.: Formation of highly
 hygroscopic soot aerosols upon internal mixing with sulfuric acid vapor, Journal of
 Geophysical Research: Atmospheres, 114, D05208, 10.1029/2008JD010595, 2009.
- 643 Koch, D., and Del Genio, A. D.: Black carbon semi-direct effects on cloud cover: review and
- 644 synthesis, Atmospheric Chemistry and Physics, 10, 7685-7696, 10.5194/acp-10-7685-2010,
- 645 2010.
- Kondo, Y., Komazaki, Y., Miyazaki, Y., Moteki, N., Takegawa, N., Kodama, D., Deguchi, S.,
 Nogami, M., Fukuda, M., Miyakawa, T., Morino, Y., Koike, M., Sakurai, H., and Ehara, K.:
 Temporal variations of elemental carbon in Tokyo, Journal of Geophysical Research:
 Atmospheres, 111, D12205, 10.1029/2005JD006257, 2006.
- Laborde, M., Schnaiter, M., Linke, C., Saathoff, H., Naumann, K. H., Mohler, O., Berlenz, S.,
- 651 Wagner, U., Taylor, J. W., Liu, D., Flynn, M., Allan, J. D., Coe, H., Heimerl, K., Dahlkotter, F.,
- 652 Weinzierl, B., Wollny, A. G., Zanatta, M., Cozic, J., Laj, P., Hitzenberger, R., Schwarz, J. P., and
- 653 Gysel, M.: Single Particle Soot Photometer intercomparison at the AIDA chamber, Atmos.
- 654 Meas. Tech., 5, 3077-3097, 10.5194/amt-5-3077-2012, 2012.
- Laborde, M., Crippa, M., Tritscher, T., Jurányi, Z., Decarlo, P. F., Temime-Roussel, B., Marchand,
- 656 N., Eckhardt, S., Stohl, A., Baltensperger, U., Prévôt, A. S. H., Weingartner, E., and Gysel, M.:
- 657 Black carbon physical properties and mixing state in the European megacity Paris, Atmos.
- 658 Chem. Phys., 13, 5831-5856, 10.5194/acp-13-5831-2013, 2013.
- Lack, D. A., and Cappa, C. D.: Impact of brown and clear carbon on light absorption
- 660 enhancement, single scatter albedo and absorption wavelength dependence of black carbon,
- 661 Atm. Chem. Phys., 10, 4207-4220, 10.5194/acp-10-4207-2010, 2010.
- Lan, Z. J., Huang, X. F., Yu, K. Y., Sun, T. L., Zeng, L. W., and Hu, M.: Light absorption of black
- 663 carbon aerosol and its enhancement by mixing state in an urban atmosphere in South China,
- 664 Atmos. Environ., 69, 118-123, 10.1016/j.atmosenv.2012.12.009, 2013.
- Lee, A. K. Y., Willis, M. D., Healy, R. M., Onasch, T. B., and Abbatt, J. P. D.: Mixing state of
- 666 carbonaceous aerosol in an urban environment: single particle characterization using the
- 667 soot particle aerosol mass spectrometer (SP-AMS), Atmos. Chem. Phys., 15, 1823-1841,
- 668 10.5194/acp-15-1823-2015, 2015a.
- Lee, A. K. Y., Willis, M. D., Healy, R. M., Wang, J. M., Jeong, C. H., Wenger, J. C., Evans, G. J., and

- 670 Abbatt, J. P. D.: Single particle characterization of biomass burning organic aerosol (BBOA):
- evidence for non-uniform mixing of high molecular weight organics and potassium, Atmos.
- 672 Chem. Phys. Discuss., 15, 32157-32183, 10.5194/acpd-15-32157-2015, 2015b.
- Li, L., Huang, Z., Dong, J., Li, M., Gao, W., Nian, H., Fu, Z., Zhang, G., Bi, X., Cheng, P., and Zhou,
- 674 Z.: Real time bipolar time-of-flight mass spectrometer for analyzing single aerosol particles,
- 675 International Journal of Mass Spectrometry, 303, 118-124,
 676 http://dx.doi.org/10.1016/j.ijms.2011.01.017, 2011.
- Li, L., Tan, G. B., Zhang, L., Fu, Z., Nian, H. Q., Huang, Z. X., Zhou, Z., and Li, M.: Analysis of
 Diesel Exhaust Particles Using Single Particle Aerosol Mass Spectrometry, Chinese Journal

679 of Analytical Chemistry, 41, 1831-1836, 10.3724/sp.j.1096.2013.30545, 2013.

- Liu, D., Allan, J., Whitehead, J., Young, D., Flynn, M., Coe, H., McFiggans, G., Fleming, Z. L., and
- Bandy, B.: Ambient black carbon particle hygroscopic properties controlled by mixing state
 and composition, Atmospheric Chemistry and Physics, 13, 2015-2029,
 10.5194/acp-13-2015-2013, 2013.
- Liu, D., Allan, J. D., Young, D. E., Coe, H., Beddows, D., Fleming, Z. L., Flynn, M. J., Gallagher, M.
- W., Harrison, R. M., Lee, J., Prevot, A. S. H., Taylor, J. W., Yin, J., Williams, P. I., and Zotter, P.:
 Size distribution, mixing state and source apportionments of black carbon aerosols in
 London during winter time, Atmos. Chem. Phys. Discuss., 14, 16291-16349,
 10.5194/acpd-14-16291-2014, 2014.
- Matsui, H., Koike, M., Kondo, Y., Moteki, N., Fast, J. D., and Zaveri, R. A.: Development and
 validation of a black carbon mixing state resolved three-dimensional model: Aging
 processes and radiative impact, Journal of Geophysical Research: Atmospheres, 118,
 2304-2326, 10.1029/2012JD018446, 2013.
- 693 McMeeking, G. R., Hamburger, T., Liu, D., Flynn, M., Morgan, W. T., Northway, M., Highwood, E.
- 694 J., Krejci, R., Allan, J. D., Minikin, A., and Coe, H.: Black carbon measurements in the
- 695 boundary layer over western and northern Europe, Atmospheric Chemistry and Physics, 10,
- 696 9393-9414, 10.5194/acp-10-9393-2010, 2010.
- 697 McMeeking, G. R., Morgan, W. T., Flynn, M., Highwood, E. J., Turnbull, K., Haywood, J., and Coe,
- H.: Black carbon aerosol mixing state, organic aerosols and aerosol optical properties over
- 699 the United Kingdom, Atmospheric Chemistry and Physics, 11, 9037-9052,

- 700 10.5194/acp-11-9037-2011, 2011.
- Moffet, R. C., de Foy, B., Molina, L. T., Molina, M. J., and Prather, K. A.: Measurement of
 ambient aerosols in northern Mexico City by single particle mass spectrometry,
 Atmospheric Chemistry and Physics, 8, 4499-4516, 2008.
- Moffet, R. C., and Prather, K. A.: In-situ measurements of the mixing state and optical properties of soot with implications for radiative forcing estimates, Proceedings of the National Academy of Sciences of the United States of America, 106, 11872-11877, 10.1073/pnas.0900040106, 2009.
- Moteki, N., Kondo, Y., Miyazaki, Y., Takegawa, N., Komazaki, Y., Kurata, G., Shirai, T., Blake, D.,
 Miyakawa, T., and Koike, M.: Evolution of mixing state of black carbon particles: Aircraft
 measurements over the western Pacific in March 2004, Geophysical research letters, 34,
 2007.
- Moteki, N., and Kondo, Y.: Dependence of laser-induced incandescence on physical
 properties of black carbon aerosols: Measurements and theoretical interpretation, Aerosol
 Sci. Technol., 44, 663-675, 2010.
- 715 Moteki, N., Kondo, Y., and Nakamura, S.: Method to measure refractive indices of small
- nonspherical particles: Application to black carbon particles, Journal of Aerosol Science, 41,
 - 717 513-521, 10.1016/j.jaerosci.2010.02.013, 2010.
 - Moteki, N., Kondo, Y., Oshima, N., Takegawa, N., Koike, M., Kita, K., Matsui, H., and Kajino, M.:
 Size dependence of wet removal of black carbon aerosols during transport from the
 boundary layer to the free troposphere, Geophysical Research Letters, 39, L13802,
 10.1029/2012GL052034, 2012.
 - Moteki, N., Kondo, Y., and Adachi, K.: Identification by single particle soot photometer of
 black carbon particles attached to other particles: Laboratory experiments and ground
 observations in Tokyo, Journal of Geophysical Research: Atmospheres, 2014.
 - 725 Onasch, T. B., Trimborn, A., Fortner, E. C., Jayne, J. T., Kok, G. L., Williams, L. R., Davidovits, P.,
 - and Worsnop, D. R.: Soot Particle Aerosol Mass Spectrometer: Development, Validation, and
 - 727 Initial Application, Aerosol Sci. Technol., 46, 804-817, 10.1080/02786826.2012.663948,
 - 728 2012.
 - 729 Petzold, A., Ogren, J., Fiebig, M., Laj, P., Li, S.-M., Baltensperger, U., Holzer-Popp, T., Kinne, S.,

- Pappalardo, G., and Sugimoto, N.: Recommendations for reporting" black carbon"
 measurements, Atmospheric Chemistry and Physics, 13, 8365-8379, 2013.
- 732 Pöschl, U.: Atmospheric aerosols: Composition, transformation, climate and health effects,
- Angewandte Chemie International Edition, 44, 7520-7540, 2005.
- Ramanathan, V., and Carmichael, G.: Global and regional climate changes due to black
 carbon, Nature geoscience, 1, 221-227, 2008.
- 736 Reddington, C. L., McMeeking, G., Mann, G. W., Coe, H., Frontoso, M. G., Liu, D., Flynn, M.,
- 737 Spracklen, D. V., and Carslaw, K. S.: The mass and number size distributions of black carbon
- 738 aerosol over Europe, Atmospheric Chemistry and Physics, 13, 4917-4939,
 739 10.5194/acp-13-4917-2013, 2013.
- 740 Riemer, N., Vogel, H., and Vogel, B.: Soot aging time scales in polluted regions during day and
- night, Atmos. Chem. Phys., 4, 1885-1893, 10.5194/acp-4-1885-2004, 2004.
- 742 Sahu, L. K., Kondo, Y., Moteki, N., Takegawa, N., Zhao, Y., Cubison, M. J., Jimenez, J. L., Vay, S.,
- 743 Diskin, G. S., Wisthaler, A., Mikoviny, T., Huey, L. G., Weinheimer, A. J., and Knapp, D. J.:
- 744 Emission characteristics of black carbon in anthropogenic and biomass burning plumes
- 745 over California during ARCTAS-CARB 2008, Journal of Geophysical Research-Atmospheres,
- 746 117, 10.1029/2011jd017401, 2012.
- 747 Schnaiter, M., Linke, C., Mohler, O., Naumann, K. H., Saathoff, H., Wagner, R., Schurath, U., and
- Wehner, B.: Absorption amplification of black carbon internally mixed with secondary
 organic aerosol, Journal of Geophysical Research-Atmospheres, 110,
 10.1029/2005jd006046, 2005.
- 751 Schwarz, J. P., Gao, R. S., Fahey, D. W., Thomson, D. S., Watts, L. A., Wilson, J. C., Reeves, J. M.,
- 752 Darbeheshti, M., Baumgardner, D. G., Kok, G. L., Chung, S. H., Schulz, M., Hendricks, J., Lauer,
- A., Karcher, B., Slowik, J. G., Rosenlof, K. H., Thompson, T. L., Langford, A. O., Loewenstein, M.,
- 754 and Aikin, K. C.: Single-particle measurements of midlatitude black carbon and
- light-scattering aerosols from the boundary layer to the lower stratosphere, Journal ofGeophysical Research-Atmospheres, 111, 10.1029/2006jd007076, 2006.
- 757 Schwarz, J. P., Gao, R. S., Spackman, J. R., Watts, L. A., Thomson, D. S., Fahey, D. W., Ryerson, T.
- 758 B., Peischl, J., Holloway, J. S., Trainer, M., Frost, G. J., Baynard, T., Lack, D. A., de Gouw, J. A.,
- 759 Warneke, C., and Del Negro, L. A.: Measurement of the mixing state, mass, and optical size of

- individual black carbon particles in urban and biomass burning emissions, Geophysical
 Research Letters, 35, 10.1029/2008gl033968, 2008.
- 762 Schwarz, J. P., Spackman, J. R., Gao, R. S., Perring, A. E., Cross, E., Onasch, T. B., Ahern, A.,
- 763 Wrobel, W., Davidovits, P., Olfert, J., Dubey, M. K., Mazzoleni, C., and Fahey, D. W.: The
- 764 Detection Efficiency of the Single Particle Soot Photometer, Aerosol Sci. Technol., 44,
- 765 612-628, 10.1080/02786826.2010.481298, 2010.
- Seinfeld, J. H., and Pandis, S. N.: Atmospheric chemistry and physics: from air pollution toclimate change, John Wiley & Sons, 2012.
- 768 Shiraiwa, M., Kondo, Y., Iwamoto, T., and Kita, K.: Amplification of light absorption of black
- 769 carbon by organic coating, Aerosol Sci. Technol., 44, 46-54, 10.1080/02786820903357686,
- 770 2010.
- 771 Song, X. H., Hopke, P. K., Fergenson, D. P., and Prather, K. A.: Classification of single particles
- analyzed by ATOFMS using an artificial neural network, ART-2A, Analytical Chemistry, 71,
- 773 860-865, 10.1021/ac9809682, 1999.
- 774 Soto-Garcia, L. L., Andreae, M. O., Andreae, T. W., Artaxo, P., Maenhaut, W., Kirchstetter, T.,
- 775 Novakov, T., Chow, J. C., and Mayol-Bracero, O. L.: Evaluation of the carbon content of
- aerosols from the burning of biomass in the Brazilian Amazon using thermal, optical and
- thermal-optical analysis methods, Atmospheric Chemistry and Physics, 11, 4425-4444,
- 778 2011.
- Stephens, M., Turner, N., and Sandberg, J.: Particle identification by laser-induced
 incandescence in a solid-state laser cavity, Appl. Optics, 42, 3726-3736,
 10.1364/ao.42.003726, 2003.
- 782 Takahama, S., Russell, L. M., Shores, C. A., Marr, L. C., Zheng, J., Levy, M., Zhang, R., Castillo, E.,
- 783 Rodriguez-Ventura, J. G., Quintana, P. J. E., Subramanian, R., Zavala, M., and Molina, L. T.:
- 784 Diesel vehicle and urban burning contributions to black carbon concentrations and size
- distributions in Tijuana, Mexico, during the Cal-Mex 2010 campaign, Atmos. Environ., 88,
- 786 341-352, 10.1016/j.atmosenv.2013.09.057, 2014.
- 787 Taylor, J. W., Allan, J. D., Liu, D., Flynn, M., Weber, R., Zhang, X., Lefer, B. L., Grossberg, N.,
- Flynn, J., and Coe, H.: Assessment of the sensitivity of core/shell parameters derived using
- the single-particle soot photometer to density and refractive index, Atmos. Meas. Tech.

- 790 Discuss., 7, 5491-5532, 10.5194/amtd-7-5491-2014, 2014.
- Wang, Q., Zhuang, G., Huang, K., Liu, T., Deng, C., Xu, J., Lin, Y., Guo, Z., Chen, Y., Fu, Q., Fu, J. S.,
- and Chen, J.: Probing the severe haze pollution in three typical regions of China:
- 793 Characteristics, sources and regional impacts, Atmos. Environ., 120, 76-88,
- 794 <u>http://dx.doi.org/10.1016/j.atmosenv.2015.08.076</u>, 2015.
- 795 Wang, X., Ye, X., Chen, H., Chen, J., Yang, X., and Gross, D. S.: Online hygroscopicity and
- chemical measurement of urban aerosol in Shanghai, China, Atmos. Environ., 95, 318-326,
- 797 <u>http://dx.doi.org/10.1016/j.atmosenv.2014.06.051</u>, 2014.
- Willis, M. D., Healy, R. M., Riemer, N., West, M., Wang, J. M., Jeong, C. H., Wenger, J. C., Evans, G.
- J., Abbatt, J. P. D., and Lee, A. K. Y.: Quantification of black carbon mixing state from traffic:
- 800 implications for aerosol optical properties, Atmos. Chem. Phys. Discuss., 15, 33555-33582,
- 801 10.5194/acpd-15-33555-2015, 2015.
- Yao, X., Chan, C. K., Fang, M., Cadle, S., Chan, T., Mulawa, P., He, K., and Ye, B.: The
 water-soluble ionic composition of PM2.5 in Shanghai and Beijing, China, Atmos. Environ.,
 36, 4223-4234, http://dx.doi.org/10.1016/S1352-2310(02)00342-4, 2002.
- 805 Zhang, G., Bi, X., He, J., Chen, D., Chan, L. Y., Xie, G., Wang, X., Sheng, G., Fu, J., and Zhou, Z.:
- 806 Variation of secondary coatings associated with elemental carbon by single particle analysis,
- 807 Atmos. Environ., 92, 162-170, <u>http://dx.doi.org/10.1016/j.atmosenv.2014.04.018</u>, 2014.
- 808 Zhang, R., Khalizov, A. F., Pagels, J., Zhang, D., Xue, H., and McMurry, P. H.: Variability in
- 809 morphology, hygroscopicity, and optical properties of soot aerosols during atmospheric
- 810 processing, Proceedings of the National Academy of Sciences, 105, 10291-10296,
- 811 10.1073/pnas.0804860105, 2008.
- 812 Zhang, Y. X., Zhang, Q., Cheng, Y. F., Su, H., Kecorius, S., Wang, Z. B., Wu, Z. J., Hu, M., Zhu, T., 813 Wiedensohler, A., and He, K. B.: Measuring morphology and density of internally mixed 814 black carbon with SP2 and VTDMA: new insight to absorption enhancement of black carbon 815 Meas. Tech. Discuss.. 2015, 12025-12050, in the atmosphere, Atmos. 816 10.5194/amtd-8-12025-2015, 2015.
- 817
- 818

Table 1. Names, numbers and fractions of the six types of BC-containing particlesdetected by the SPAMS instrument.

Group	Number of particles	Fraction of particles
Pure BC	535	0.62%
Biomass Burning BC-containing (BBBC)	22 007	25.57%
K-rich BC-containing (KBC)	11 343	13.18%
BC internally-mixed with OC and ammonium nitrate (BCOC-NOx)	33 760	39.23%
BC internally-mixed with OC and ammonium sulfate (BCOC-SOx)	15 291	17.77%
Unidentified	3121	3.63%
Total BC-containing	86 057	100%



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Figure 1. Temporal profiles of temperature and relative humidity with 30 min resolution, gaseous pollutants (CO, O_3 , SO₂, NO and NO₂) with 60 min resolution, and PM_{2.5} and PM₁₀ mass concentrations with 60 min resolution. The concentration of rBC mass (black trace in the bottom panel, 10 min resolution) was continuously measured by SP2.



Figure 2. The measured rBC core mass size distribution and number size distribution are shown in open red circles and black triangles, respectively. The log-normal fits to the observed distributions are shown by the dashed lines.



Figure 3. Comparison of the SPMAS-detected internally-mixed BC-containing particles and
SP2-detected internally-mixed rBC-containing particles.





Figure 4. (a) Temporal variations of number and size distributions for rBC-containing
particles detected by SP2 with 30 min resolution. (b) Temporal variation of number
fractions of different BC-containing particle types with 10 min time resolution (detected
by SPAMS).



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Figure 5. (a) D_p number size distribution histogram for the SP2-detected rBC-containing particles. (b) D_{va} number fraction distribution of SPAMS-detected BC-containing particles color-coded by the particle type. (c) D_c and ACT with number size distribution in the condensation and droplet modes.





Figure 6. (a) Temporal variation of the relative coating thickness distribution of traffic-emitted rBC-containing particles (SP2) with resolutions of 0.2 RCT and 30 min and the temporal profile of NO₂ and PM_{2.5} concentrations with 60 min resolution. (b) Temporal variation of relative peak areas of +27 ($C_2H_3^+$), +43 (CH_3CO^+) and -62 (NO_3^-) of traffic-emitted BC-containing particles (SPAMS) and NO₂ concentration with 30 min resolution.