1 2	Vertical variation of optical properties of mixed Asian dust/pollution plumes according to pathway of airmass transport over East Asia
3 4 5	SK. Shin ^{1,a} , D. Müller ^{2,b} , Chulkyu Lee ³ , K. H. Lee ⁴ , D. Shin ^{5,*} , Y. J.Kim ¹ , Y. M. Noh ^{6,*}
6 7 8	¹ School of Environmental Science and Engineering, Gwangju Institute of Science & Technology Gwangju Republic of Korea
9 10	² School of Physics, Astronomy and Mathematics, University of Hertfordshire, Hertfordshire, United Kingdom
11 12	³ Korea Meteorological Administration, Seoul, Republic of Korea ⁴ Department of Atmospheric and Environmental Science, Gangneung-Wonju
13 14 15	National University, Gangneung, Republic of Korea ⁵ Air Quality Forecasting Centre, Climate and Air Quality Research Department, National Institute of Environmental Research Incheon Republic of Korea
16 17	⁶ International Environmental Research Centre (IERC), Gwangju Institute of Science & Technology, Gwangju, Republic of Korea
18 19	^a now at Department of Atmospheric and Environmental Science, Gangneung-Wonju National University, Gangneung, Republic of Korea
20 21 22	formerly at School of Environmental Science and Engineering, Gwangju Institute of Science & Technology, Gwangju, Republic of Korea
23 24	
25	
20 27	
28	
29	
30	
31	
32 33	
34	
35 36 37 38 39 40	* Corresponding Author Y. M. Noh Tel) +82-62-715-3392 Fax) +82-62-715-3402 Email) <u>nym@gist.ac.kr</u>

41 Abstract:

We use five years (2009 – 2013) of multiwavelength Raman lidar measurements at Gwangju, 42 Korea (35.10° N, 126.53° E) for the identification of changes of optical properties of East 43 Asian dust in dependence of its transport path over China. Profiles of backscatter and 44extinction coefficients, lidar ratios, and backscatter-related Ångström exponents (wavelength 45 pair 355/532nm) were measured at Gwangju. Linear particle depolarization ratios were used 46 to identify East Asian dust layers. We used backward trajectory modelling to identify the 47pathway and the vertical position of dust-laden air masses over China during long-range 48 transport. Most cases of Asian dust events can be described by the emission of dust in desert 49 areas and subsequent transport over highly polluted regions of China. The Asian dust plumes 50 51 could be categorized into two classes according to the height above ground at which these 52 plumes were transported: (cases I) the dust layers passed over China at high altitude levels (> 3 km) until arrival over Gwangju, and (case II) the Asian dust layers were transported near 53 54 the surface and within the lower troposphere (< 3 km) over industrialized areas before they arrived over Gwangju. We find that the optical characteristics of these mixed Asian dust 55 layers over Gwangju differ in dependence of their vertical position above ground over China 56 57 and the change of height above ground during transport. The mean linear particle depolarization ratio was 0.21±0.06 (at 532 nm), the mean lidar ratios were 52±7 sr at 355 nm 58 59 and 53±8 sr at 532 nm, and the mean Ångström exponent was 0.74±0.31 for case I. In contrast, plumes transported at lower altitudes (case II) showed low depolarization ratios 60 61 $(0.13\pm0.04 \text{ at } 532 \text{ nm})$, and higher lidar ratio $(63\pm9 \text{ sr at } 355 \text{ nm} \text{ and } 62\pm8 \text{ sr at } 532 \text{ nm})$ and 62 Ångström exponents (0.98±0.51). These numbers show that the optical characteristics of 63 mixed Asian plumes are more similar to optical characteristics of urban pollution. We find a decrease of the linear depolarization ratio of the mixed dust/pollution plume in dependence of 64 transport time if the pollution layer travelled over China at low heights, i.e., below 65 approximately 3 km above ground. In contrast we do not find such a trend if the dust plumes 66 travelled at heights above 3 km over China. We need a longer time series of lidar 67 measurements in order to determine in a quantitative way the change of optical properties of 68 69 dust with transport time.

- 70
- 71

72 **Key words:** lidar, Asian dust, optical properties, particle depolarization ratio, mixing, vertical

73 distribution, long-range transport

74 **1. Introduction**

Desert dust is the most abundant natural source of atmospheric particles over land. Its 75 light-absorption capacity is strong in the ultraviolet regions of the solar spectrum (Jacobson 76 2012). The light-absorption capacity depends on the proportions of Fe₂O₃, Al₂O₃, SiO₂, 77 CaCO₃, MgCO₃(s), clays, and other substances. The transport patterns of dust over North 7879 Africa and East Asia as well as the vertical distribution of dust change intra- and interannually. Thus the influence of dust on the atmosphere's radiation balance is complex (Griggs 80 81 and Noguer, 2002; Mahowald et al., 2006; Durant et al., 2009). Central East Asia has large 82 desert regions. Asian dust particles that originate from the Taklamakan desert in west China and the Gobi desert in Mongolia and northwest China (Figure. 1) influence the regional 83 84 climate over East Asia and can be found as far as the west coast of North America (Husar et al., 2001; McKendry., 2001; Huang et al., 2008). East Asian dust is particularly complicated 85 as it usually travels over densely populated and highly industrialized areas of China before it 86 moves out over Pacific Ocean. During transport over East Asia dust mixes with pollutants 87 88 such as industrial soot, toxic material, and acidic gases (Sun et al., 2005).

Field campaigns, such as ACE-Asia (Huebert et al., 2003) and ADEC (Mikami et al., 2006) significantly added to our knowledge of the radiative effects of Asian dust. Carrico et al. (2003) and Yu et al. (2006) found differences of dust optical properties as the result of the mixing of dust with anthropogenic pollution between source regions of dust and observation sites downwind of its source regions. The mixing between Asian dust and industrial pollutant particles has significant influence on the size distribution and the chemical composition of aerosol plumes (Wang et al., 2007; Sun et al., 2010).

There exist few studies on the degree of mixing that occurs between dust and pollution during transport, the effect of the direction of dust transport across China, and the vertical distribution of Asian dust layers during long-range transport over China. There still is a lack of understanding of how much of the mixing of dust with pollutants depends on the
vertical distribution of dust when it passes over source regions of anthropogenic pollution in
East Asia. One reason of our limited knowledge is that there are only few vertically-resolved,
long-term observations of pollution over East Asia.

LIDAR (LIght Detection And Ranging) is a powerful technique for measuring the vertical distribution of atmospheric aerosols with high temporal and spatial resolution. In this study we use Raman lidar data taken at Gwangju, South Korea, between 2009 and 2013. In our study we focus specifically on lidar observations of Asian dust layers as they passed over China. We use backward trajectory analysis with HYSPLIT (HYbrid Single Particle Lagrangian Integrated Trajectory) (Draxler and Rolph, 2003) model to identify the transport pathway and the vertical distribution of the Asian dust layers during long-range transport.

110 The main objective of this study is to investigate the variation of optical properties of 111 mixtures of Asian dust with anthropogenic pollution in dependence of the pathways and 112 vertical distributions of these mixed dust layers during long-range transport. In this 113 contribution we used the linear particle depolarization ratio to identify these dust layers. We 114 present vertically-resolved optical properties such as lidar ratio and Ångström exponent.

Section 2 presents the methods used in this study. Section 3 presents our results. We
 discuss our results and summarize our findings in section 4.

117

118 **2. Methodology**

119 2.1 GIST Multi-wavelength Raman lidar

120 The lidar station, dubbed MRS.LEA (Multi-wavelength Raman Spectrometer Lidar

121 in East Asia) of the Gwangju Institute of Science and Technology (GIST) is located at 35.10°

122 N, 126.53° E in the west-south-western part of the Korean peninsula (Figure. 1).

123 A description of the lidar system is given by Noh et al. (2007, 2008). The light source

124 of the lidar is a pulsed Nd: YAG laser that emits light pulses at 355 nm, 532 nm, and 1064 nm. The laser output power is 140, 154, and 640 mJ at the three emission wavelengths, 125126 respectively. The pulse repetition rate is 10 Hz. We use a beam expander at 532 nm and 1064 nm in order to reduce the divergence of the emitted light. The receiver consists of a 14-inch 127 Schmidt-Cassegrain telescope. The signals collected by the receiver telescope are separated 128 according to wavelength with beam splitters and then transmitted to photomultiplier tubes 129 (PMT). Transient recorders with 12-bit analog-to-digital converters and 250-MHz photon 130 131 counters are used for processing the output signals of the PMTs. The system allows us to 132 retrieve vertical profiles of the particle backscatter coefficients at 355, 532, and 1064 nm, the particle extinction coefficients at 355 and 532 nm, the linear particle depolarization ratio at 133 134 532 nm, the water-vapor mixing-ratio, and profiles of silicon-dioxide (Müller et al., 2010; Tatarov et al., 2011). Profiles of silicon-dioxide (quartz) can be used as a proxy of the 135 concentration of mineral dust. In this contribution we use the signals needed for measuring 136 particle backscatter and extinction coefficients at 355 and 532 nm and the linear particle 137 depolarization ratio at 532 nm. The measurements were carried out at night time under cloud-138 139 free conditions.

The profiles of particle backscatter coefficients (β_n) at 355 and 532 nm were 140 calculated with the Raman method (Ansmann et al., 1992b). The overlap effect which 141 142 describes the incomplete overlap between outgoing laser beam and field of view of the receiver telescope is cancelled out for the case of profiles of the backscatter coefficient 143 because the ratios of two signals (elastic signals from particles and molecules and the 144145 nitrogen Raman signals) are computed (Wandinger and Ansmann, 2002). In that way we can retrieve vertical profiles of the backscatter coefficient to 400 m above ground. The vertical 146 profiles of the aerosol extinction coefficients (a_p) at 355 and 532 nm were derived with the 147use of the nitrogen vibration Raman signals at 387 and 607 nm (Ansmann et al., 1990), 148

respectively. The aerosol extinction coefficients can be retrieved above 780 m and 540 m 149 above ground at the measurement wavelengths of 355 nm and 532 nm, respectively. We 150 derive particle extinction-to-backscatter ratios (lidar ratios, denoted as S in this contribution) 151152at 355 and 532 nm from the profiles of β_p and α_p . The lidar ratios can be used for aerosol typing (Müller et al., 2007). Murayama et al., (2004) find values of S = 48.6 sr at 355 nm and 153S = 43.1 sr at 532 in a well-isolated Gobi dust-laden layer observed above 4 km over Tokyo. 154De Tomasi et al. (2003) report an S value less than 50 sr at 351 nm for a Saharan dust layer. 155 Values of S at 355 nm ranged between 50 sr and 80 sr for dust observed over Leipzig, 156 Germany (Mattis et al., 2002). In contrast, Ferrare et al. (2002) report a high value of 68±12 157 sr of the lidar ratio at 355 nm. This high lidar ratio was associated with air masses advected 158from urban/industrial areas. Omar et al. (2009) finds values of 65-70 sr for the lidar ratio at 159 532 nm. The numbers describe continental-polluted aerosols and polluted dust. 160

The backscatter-related Ångström exponent for the wavelength pair of 355/532 nm 161 (denoted as \mathring{A}_{β}) is computed, too. The backscatter-related Ångström exponent is a good 162 indicator of the size of particles. High values (>1) are typically observed for accumulation 163 164mode particles such as fresh biomass-burning particles. Low values (~0) are observed for coarse mode particles such as Saharan dust or Asian dust (Eck et al., 1999; Sakai et al., 2002, 165 Chen et al., 2007). The values of 0.2-0.3 are reported as the values of \mathring{A}_{β} for Saharan dust 166 (Murayama et al., 2002; Tesche et al., 2009). Chen et al. (2007) and Müller et al. (2010) find 167 values of 0.7-1.5 for \mathring{A}_{β} for a mixture of mineral dust with urban haze. Values of 0.8-1.4 for 168 \mathring{A}_{β} were found for heavily polluted continental aerosol layers (Franke et al., 2003). 169

170 The depolarization ratio is used as indicator of particle shape (Bohren and Huffman, 171 1983). High values of the depolarization ratio of 0.3 to 0.35 at 532 nm indicate nearly pure 172 dust (Sugimoto and Lee, 2006; Freudenthaler et al., 2009). For example, Freudenthaler et al. 173 (2009) report a value of $\delta_{\mathbf{p}} = 0.31$ at 532 nm for Saharan dust observed during SAMUM 2006. 174Lidar observations were carried out close to the Taklamakan desert (Iwasaka et al., 2003) and the Gobi desert (Yi et al., 2014). We assume that these dust layers exhibit nearly pure dust 175conditions as anthropogenic pollution sources in these isolated areas are sparse. Values of $\delta_{\rm p}$ 176are in the range of 0.3 to 0.35 at 532 nm (Iwasaka et al., 2003; Yi et al., 2014). Small values, 177178e.g., values from 0.08 to 0.1 usually are an indicator that dust is mixed with spherical particles (Murayama et al., 2004; Chen et al., 2009; Tesche et al., 2009; Burton et al., 2013). 179 Anthropogenic aerosols normally are spherical with a small depolarization ratio (Murayama 180 et al., 2003). The degree of depolarization decreases as the sphericity of particles increases. 181 The depolarization ratio is dependent on the mixing ratio of dust with spherical particles. For 182 instance, Burton et al. (2013) report values of $\delta_{p} = 0.13-0.20$ and 0.03-0.07 at 532 nm for 183 polluted dust and urban aerosol particles, respectively. 184

Parallel polarized and perpendicular polarized signals are measured at 532 nm. The linear volume depolarization ratio (aerosols + molecules) $\boldsymbol{\delta}$ is defined as

187

 $\delta = \frac{\mathsf{P}_{\perp}}{\mathsf{P}_{\parallel} + \mathsf{P}_{\perp}}.\tag{1}$

188 P_{\perp} and P_{\parallel} denote the backscatter signal intensities that are polarized perpendicular 189 and parallel with respect to the plane of polarization of the emitted laser beam, respectively. 190 The δ can be also defined as P_{\perp}/P_{\parallel} (Cairo et al., 1999). We calculated the δ by using both 191 definitions and compared the difference between the derived values. The results from each 192 individual definition agree within the uncertainty of our depolarization ratio measurements 193 (Tesche et al., 2009; Shin et al., 2013).

194 The linear particle depolarization ratio $\delta_{\mathbf{p}}$ differs from δ as it depends on the 195 concentration of particles in relation to the concentration of air molecules. In this contribution 196 we use the linear particle depolarization ratio ($\delta_{\mathbf{p}}$) according to the definition by Shimizu et al. 197 (2004):

$$\delta_{\rm p} = \frac{\delta(z) R_{\rm B}(z) - \delta_{\rm m}}{R_{\rm B}(z) - 1}.$$
(2)

199 The term $\delta_{\mathbf{m}}$ is the linear depolarization ratio of air molecules at the wavelength of 200 the emitted laser wavelength. We used the value $\delta_{\mathbf{m}}=0.0044$ (Behrendt and Nakamura, 2002). 201 This value takes account of our interference filters which have a full width at half maximum 202 of 1.0 nm. $R_{\mathbf{B}}(\mathbf{z})$ is the backscatter ratio, expressed as $(\beta_{\mathbf{p}}+\beta_{\mathbf{m}})/\beta_{\mathbf{m}}$ at altitude *z*. $\beta_{\mathbf{m}}$ denotes the 203 backscatter coefficient of atmospheric molecules. The calibration of the polarization channels 204 was carried out by using rotating polarizers following the methodology explained by 205 Freudenthaler et al. (2009).

206

207 2.2 Dust Layer Identification

We use the profiles of the linear particle depolarization ratio for the identification of the Asian dust layers. An example of how the Asian dust layer was determined is shown in Figure 2. The Asian dust plume reached Korea on 22 April 2012. Figure 2 shows the timeheight cross section of the range-corrected backscatter signals and the linear volume depolarization ratio at 532 nm. Figure 2 also shows the mean profiles of δ and δ_p , *S* at 355 and 532 nm, and \mathring{A}_{B} for the measurement from 13:15 to 14:05 UTC.

Values of δ_p for individual aerosol types are reported in literature, e.g. δ_p for Asian 214 dust particles varies from 0.08-0.35 (Murayama et al., 2004; Shimizu et al., 2004; Chen et al., 215 2009; Shin et al., 2013) at 532 nm. Asian dust generally mixes with pollution during long-216 range transport which leads to variable $\delta_{\rm p}$. Thus, this range of 0.08-0.35 likely describes 217 mixtures of dust with anthropogenic pollution. For instance, Chen et al. (2009) uses 0.08 as 218 threshold value to identify dust in pollution. Furthermore, optical properties may also change 219 220 during long-range transport. Shimizu et al. (2004) define 0.1 as threshold value for the determination of polluted dust. In this study we used 0.08 as threshold value of δ_p to identify 221

222 dust.

In figure 2, the layer between 2.7 km 4.6 km (layer II) contains Asian dust particles as suggested from the values of δ_p , which are higher than 0.16. The mean value of δ_p in the layer between 1.2 km and 2.5 km (layer I) is 0.11 and thus also points to the presence of dust particles though the concentration of dust particles compared to the concentration of particles of anthropogenic pollution may be lower in layer I compared to layer II.

Other aerosol optical properties in layer I and layer II differ, too. The values of the *S* in layer I are 64 ± 4 sr and 66 ± 4 sr at 355 and 532 nm, respectively. The values of the *S* in layer II are as low 55 ± 4 sr and 55 ± 3 sr at 355 and 532 nm, respectively, see Fig. 2d. The standard deviations were computed for the lidar ratios in each of the layers we could identify.

The values of $Å_{B}$ in layer I are ~0.93 and thus considerably higher than in layer II 232 where we find a value of ~ 0.42 . These numbers suggest that the concentration of small 233 particles is higher in layer I than in layer II, respectively that the mean size of particles in 234 layer I is smaller than the mean size of particles in layer II. Regarding the interpretation of the 235 numbers of \mathring{A}_{β} we need to keep in mind that the backscatter-related Ångström exponent not 236 237 only depends on particle size but also on the complex refractive index and particle shape. The same holds true for the values of S. The different numbers thus could also result from 238 differences in particle shape and their absorption properties in these mixed Asian dust layers. 239

240

241 **2.3 Analysis of Backward Trajectories and Model Simulations of Pollution Emissions**

We used the HYSPLIT model (Draxler and Rolph, 2003) to generate 120 hours backward trajectories for air parcels arriving above our lidar site. The trajectories describe the different altitude levels in which dust was transported prior to the lidar observations. They also allow us to trace back the origin of the dust layers and the transport path.

246 The Monitoring Atmospheric Composition and Climate (MACC) global air quality

247 service of the European Centre for Medium-Range Weather Forecasts (ECMWF) provides a re-analysis of global atmospheric composition. The re-analysis assimilates satellite data, e.g. 248total aerosol optical depth (AOD) which is provided by the Moderate Resolution Imaging 249 Spectroradiometer (MODIS), into a global model and data assimilation system to correct for 250 model departures from observational data (Bellouin et al., 2013; Inness et al., 2013). This re-251 analysis provides fields of aerosols, namely mineral dust, black carbon, organic matter, and 252 sulphate, as well as chemically reactive gases, and greenhouse gases. We used the aerosol 253 AOD from the MACC re-analysis to determine the intensity of pollution (AOD) in densely 254 populated and industrialized regions along the transport path of the dust layers and to 255investigate the influence of anthropogenic pollution particles on the variation of the optical 256257 properties of Asian dust.

258

259 **3. Results and discussion**

We present data that cover the time from 2009-2013. During this time we observed 38 Asian dust layers on 32 days. These Asian dust layers were identified on the basis of the linear particle depolarization ratio measurements as described in section 2,2. The vertical profiles of the linear particle depolarization ratio allow us to determine the vertical distribution of the Asian dust layers.

Figure 3 shows the frequency distribution of $\delta_{\mathbf{p}}$, *S*, and \mathring{A}_{β} of Asian dust plumes observed during the observation period. The average value of $\delta_{\mathbf{p}}$ for all observed Asian dust layers is 0.17±0.02. The average values of *S* are 57±6 sr at 355 nm and 57±7 sr at 532 nm. The mean value of \mathring{A}_{β} is 0.84±0.37. The optical properties of each individual Asian dust layer vary over a wide range of values. We find values of 0.08-0.33 for $\delta_{\mathbf{p}}$, 38-83 sr for *S* at 355 nm, 41-73 sr for *S* at 532 nm, and 0.38-1.71 for \mathring{A}_{β} . The maximum value of $\delta_{\mathbf{p}}$ is 0.33 at 532 nm. The minimum values of *S* at 355 nm and 532 nm are 38 sr and 41 sr, respectively. The minimum value of \mathring{A}_{β} is 0.38. This maximum value of \eth_{p} and the minimum values of S at 355 nm and 532 nm and \mathring{A}_{β} are similar to the values of optical properties for pure dust particles. In contrast with these values, low values of \eth_{p} and high values of S and \mathring{A}_{β} are also measured. We find that the minimum value of \eth_{p} is 0.08 at 532 nm. The maximum values of of S at 355 nm and 532 nm are 83 sr and 73 sr, respectively. The maximum value of \mathring{A}_{β} is 1.71. These values are remarkably different from the values of optical properties of pure dust.

We speculate that these differences of the values of the optical properties of dust particles are caused by the effect of long-range transport during which dust mixes with anthropogenic pollution or biomass burning smoke when passing over industrialized/densely populated regions in China.

282

3.1 Qualitative Analysis of the Variation of Optical Properties of Mixed-Dust in Dependence of Pollution Levels

We divided the dust layers into two episodes. The two episodes differ according to 285 the level of pollution emissions along the transport pathway of the dust plumes. The 286 287 separation of our measurements into these two episodes was done on the basis of the distribution of aerosol optical depth (AOD) of anthropogenic pollution over China. The 288 Asian dust layers were classified as "more polluted", i.e., "MP" Asian dust when the 289 290 modelled AOD of anthropogenic pollution on that day was higher than the average AOD (modelled) of all 32 observation days considered in this study. In contrast, Asian dust layers 291 that passed over China during episodes of lower AOD, i.e., AOD was below the mean value 292 293 of modelled AOD of all 32 observation days, are denoted as "less polluted", i.e., "LP" Asian dust. 294

We used model results by MACC and backward trajectory analysis (see section 2.3) for the interpretation of our lidar results as we do not have direct observations of pollution, e.g. particle optical depth, lidar ratios, the linear particle depolarization ratios, and Ångström
exponents along the transport path of the pollution plumes. The reliability of inferring AOD
of pollution from MACC re-analysis is validated by comparing it to results from AERONET
sunphotometer measurements. MACC model is widely used to estimate AOD of pollution
(Bellouin et al., 2013; Cesnulyte et al., 2014).

Figure 4 shows the distribution of aerosol optical depth (AOD) at 550 nm for dust 302 and anthropogenic pollution on 10 April 2010 and 8 March 2013. These pollutants include 303 304 organic matter, black carbon, and sulphate aerosol. The pollution AOD was computed with the MACC model using re-analysis data of ECMWF. The re-analysis data from the MACC 305 model be downloaded the web of ECMWF 306 can at page 307 (http://apps.ecmwf.int/datasets/data/macc-reanalysis/). Fig. 4 shows that Asian dust particles emitted from the Taklamakan and the Gobi desert were transported across China. The model 308 results of AOD of anthropogenic pollutants over China for 10 April 2010 are significantly 309 higher than the model results of AOD on 8 March 2013. 310

Figure 5 shows the scatter diagram of \mathring{A}_{β} and S at 355 nm and 532 nm versus δ_{p} in 311 dependence of the transport events denoted as MP and LP. The mean value of $\delta_{\rm p}$ of the Asian 312 dust layers denoted as "LP" cases ranges between 0.08 (threshold value that we use to 313 identify dust) and 0.33. The corresponding values of \mathring{A}_{β} vary between 0.38 and 1.71. The 314 315 lidar ratios range between 38 sr and 83 sr at 355 nm and between 41 sr and 73 sr at 532 nm. The negative correlation of δ_p with \mathring{A}_β indicates that the impact of the non-spherical particles 316 (Asian dust with high δ_p) on the backscattered light decreases with increasing \mathring{A}_{β} . Higher 317 values of \mathring{A}_{β} indicate a considerable concentration of anthropogenic pollution particles which 318 in turn results in lower values of δ_{p} , of the mixed dust/pollution plumes. 319

Lower values of δ_p are dominantly found in the domain where lidar ratios are above 60-70 sr, except for a few cases. Comparably high lidar ratios are associated with air masses from urban/industrial areas (Noh et al., 2007; Müller et al., 2007; Burton et al., 2012). We find high values of δ_p for lidar ratios of 57±7 sr at 355 nm and 55±7 sr at 532 nm.

With regard to the MP cases the mean δ_p varies from 0.08 to 0.30. The corresponding values of \mathring{A}_{β} vary between 0.42 and 1.56. The lidar ratios vary between 44 sr and 74 sr at 325 mm and between 48 sr and 72 sr at 532 nm, respectively.

Figure 5(d-f) shows a negative correlation of δ_{p} with \mathring{A}_{β} and S at 355 nm and 532 nm. The mean values of the LP and MP cases are summarized in table 1. The transport pathway of dust over eastern China should influence the degree to which anthropogenic aerosols in the industrial areas contribute to the change of optical properties of dust. However, we do not find significant differences between the LP cases and MP cases. We assume that there is another factor that influences the change of the optical properties of the dust layers we observed.

334

335 3.2 Influence of pathway and vertical distribution of anthropogenic pollution on optical properties of Asian dust

337 We classified the Asian dust plumes into 2 categories with regard to height above ground when they passed over regions of anthropogenic emissions. We used 3 km height 338 above ground for the classification. The height of 3 km is reported as the planetary boundary 339 340 layer. Pollutants emitted at the surface predominantly stay in the planetary boundary layer (Noh et al., 2007; Xie et al., 2015). We assume that height above ground influences how 341 much anthropogenic pollution may mix with the dust layers and thus changes the optical 342 343 properties of the dust layers. The vertical positions of the dust plumes above ground during transport over China were inferred from the model results. We assume that the height of the 344 dust plumes above ground can be distinguished by HYSPLIT model results although the 345 results may have a certain error because of the spatial and temporal complexity of the 346

347 meteorological fields involved in the computations.

Figure 6 shows the transport pathway and the change of the vertical position of the dust plumes during transport to our lidar site. It is clear that backward trajectories cannot provide us with information on the concentration of dust and anthropogenic pollution in the air masses prior to observation over Korea. Still, backward trajectories show if the air masses originated from or nearby the desert regions, and whether the air masses passed over densely populated/industrialized regions.

Case I includes those Asian dust plumes that passed over industrialized areas in China at high altitude level (> 3km height above ground) as shown in figure 6a. The Asian dust plumes were classified as Case II when they were transported through the near surface/lower troposphere (< 3km height above ground) over industrialized areas in China, i.e., longitude range between 110° E and 125° E; the locations of industrialized and densely populated regions in China are shown in figure 1.

The mean values of the linear particle depolarization ratios of the Asian dust plumes we observed are lower compared to the linear particle depolarization ratios of pure dust particles. For example, Freudenthaler et al. (2009) report a value of $\delta_p = 0.31$ at 532 nm for pure Saharan dust observed during SAMUM 2006.

The values of δ_p and the corresponding values of \mathring{A}_{β} and S at 355 nm and 532 nm for the cases I and II are also shown in figure 6. The corresponding mean values of the parameters of these two cases are also summarized in table 1.

We find different clusters of the optical properties of the dust layers when we take into consideration their vertical position during transport. The cases I show larger values of $\delta_{\rm p}$ compared to the depolarization ratios in cases II. On average, \mathring{A}_{β} of case I is smaller than \mathring{A}_{β} of case II. The average values of $\delta_{\rm p}$ and \mathring{A}_{β} are 0.21±0.06 and 0.74±0.31, respectively, for case I. In contrast, $\delta_{\rm p}$ and \mathring{A}_{β} are 0.13±0.04 and 0.98±0.35, respectively, for case II. The lowest values of *S* at 355 nm and 532 nm are also measured for high values of $\delta_{\mathbf{p}}$ (0.21±0.06). We find values of 52±7 sr at 355 nm and 53±8 sr at 532 nm, respectively, for case I. Comparably high values of *S* were found for case II, i.e. 63±9 sr at 355 nm and 62±8 sr at 532 nm. In that case the value of $\delta_{\mathbf{p}}$ is 0.13±0.04.

There are several previous studies that report on linear particle depolarization ratios of polluted dust after long-range transport. According to these studies the observed dust particles were partly/completely mixed with anthropogenic pollution (Sakai et al., 2002; Müller et al., 2003; Shimizu et al., 2004; Chen et al., 2007). As a result of the mixing of dust with anthropogenic pollution, the values of δ_p were lower than the values of pure dust, which is estimated to be 0.3-0.35 (Murayama et al., 2004; Freudenthaler et al., 2009). Likewise, the values of \mathring{A}_{β} and S also differ compared to the values of \mathring{A}_{β} and S of pure dust.

We assume that the dust particles carried more anthropogenic pollution in cases where the air masses travelled near the surface. Consequently, the optical characteristics of the dust/pollution layers of case II are dominated by the optical properties of anthropogenic pollutants. In contrast, the optical properties of dust layers that travelled at high altitudes (case I) are less influenced by urban/industrial pollutants. Thus, the optical properties of these dust layers are more likely to be those of pure dust.

The Asian dust plumes were classified into 4 categories. We considered not only the 389 level of pollution emissions along the transport pathway, i.e., "MP" Asian dust and "LP" 390 Asian dust, but also the vertical position of the layers when they passed over polluted regions 391 of China ("below 3km" and "above 3km"). Figure 7 shows scatter diagrams of $Å_{\beta}$ 392 393 (wavelength range 355/532 nm), and S at 355 nm and 532 nm versus $\delta_{\rm p}$ at 532 nm in dependence of the level of pollution emission and the vertical position. The corresponding 394 mean values of the optical parameters of those clusters are given in Table 1. We expect that 395 the optical properties of Asian dust change most if pollution levels (in terms of AOD) are 396

397 high, (MP Asian case) and when the corresponding air masses passed over industrialized area of China at low altitude (below 3km height above ground). The mean values of δ_p at 532 nm 398 and \mathring{A}_{β} are 0.13±0.04 and 1.09±0.30, respectively, for this case which is denoted as 399 "MP_below 3km". The mean values of S are 61 ± 10 sr at 355 nm and 64 ± 7 sr at 532 nm. 400 However, these values of optical properties of dust for MP_below 3km are not significantly 401 different from the case of "LP_below 3km". In that case the mean values of δ_p at 532 nm and 402 $Å_{B}$ are 0.13±0.03 and 1.00±0.38, respectively. The mean values of S are 64±9 sr at 355 nm and 403 62±8 sr at 532 nm. 404

The values of optical properties between "MP" and "LP" at high altitude also do not differ significantly. The mean values of $\delta_{\mathbf{p}}$ at 532 nm, \mathring{A}_{β} , and S are 0.24±0.05, 0.58±0.14, and 53±5 sr at 355 nm and 53±2 sr at 532 nm, respectively, for the case "MP_above 3km". The highest values of $\delta_{\mathbf{p}}$ and lowest values of \mathring{A}_{β} are found for this case.

In the case of "LP_above 3km" the mean values of δ_p at 532 nm and \mathring{A}_{β} are 0.21±0.05 and 0.65±0.20, respectively. The mean values of *S* are 51±8 sr at 355 nm and 49±9 sr, respectively. We believe that the changes in the optical properties of Asian dust depend on the vertical position of the dust plume rather than the level of pollution emission during transport.

The clusters denoted as Case I and Case II were classified according to the altitude 413 (above ground) at which the dust-laden air masses passed over industrialized/populated 414regions of China. The differences of the optical properties of the dust layers are shown in 415figure 8. The corresponding values of the optical characteristics of the Asian dust layers at 416 417 each individual height are summarized in table 2. The difference of the optical characteristics 418of East Asian dust layers that travelled in surface-near heights and at high altitudes is obvious. The values of $\delta_{\mathbf{p}}$, \mathring{A}_{β} , and S are 0.12±0.01, 1.00±0.43, and 63±7 sr at 355 nm and 64±6 sr at 419 532 nm, respectively, when Asian dust passed over China below 1 km height above ground. 420 These values reflect the fact that the optical properties of the dust/pollution plumes are 421

dominated by the anthropogenic part of the particles in these plumes. Lower values of $\delta_{\mathbf{p}}$ 422 represent the dominance of spherical particles, i.e. the presence of urban pollution. High 423 values of \mathring{A}_{β} indicate that small particles dominate in the lower altitude level. The high lidar 424 ratio also indicates the presence of urban pollution which tends to be more light-absorbing 425(Müller et al., 2007). In contrast, values for $\delta_{\mathbf{p}}$, $\dot{A}_{\mathbf{\beta}}$, and S are 0.23±0.04, 0.60±0.17, and 50±6 426 427 sr at 355 nm and 49±5 sr at 532 nm, respectively, after the dust layers had passed over China at high altitudes, i.e., above 3 km. These values more likely reflect the optical characteristic 428 429 of Asian dust particles that are less affected by the contribution of anthropogenic pollution. The optical properties of Asian dust layer observed in our study reflect mixtures between 430 431 different aerosol types.

432 We notice that these variations of the optical properties of Asian dust layers may not only result from external mixing. Hygroscopic growth, aging and deposition during transport, 433 and internal mixing might be also affect dust properties (Burton et al., 2014). The 434 interpretation of the mixing state of Asian dust is a challenging task. The mixing state 435depends on many variables which are poorly known. Sugimoto et al. (2015) tried to identify 436 437 the mixing state of Asian dust (internal mixing or external mixing) by using analytical relationships inferred from lidar observation. However, we will not go into details here. We 438 assume that most of the Asian dust observed in this study was externally mixed. 439

The altitude in which the Asian dust layers passed over China have significant influence on their optical characteristics. In our study, we took 3 km above ground as threshold value as we observed a notable change of optical properties of the dust/pollution plumes if they travelled above or below 3 km height above ground. Pollution particles below 3 km could mix and interact with Asian dust particles (more influence). In contrast, we assume that optical properties of dust particles above 3 km are not that much influenced by anthropogenic pollution as the mixing of pollution into these heights is less intense. We emphasize that this threshold value of 3 km is merely a best estimate which is governed by the set of data we have at hand. We lack in additional information that would allow us to refine our data analysis. For example a longer time series of lidar measurements, (vertically resolved) observations of pollution transported over China, measurements under much more variable meteorological conditions, additional modelling results, just to name a few reasons, might change this threshold value.

We also investigated the optical properties of Asian dust with respect to transport 453 454time at different height level. Figure 9 shows scatter diagrams of optical properties of Asian dust versus the transport time. The correlation study is based on HYSPLIT model results, our 455 profiles of $\delta_{\rm p}$, ${\rm \AA}_{\rm \beta}$, and S, and the time (in hours) the Asian dust spent in polluted regions over 456 457 China during the transport. We can only use HYSPLIT results as an estimate of the total transport time and the time the plumes spent over pollution regions of China. The total 458 transport time may have considerable uncertainty. We need to decide from the trajectories the 459 start point of dust emission and this means we take the time when the air parcel (defined by 460 its trajectory) left one of the desert regions in Central Asia. The height above ground during 461 462 transport and the time the plumes spent over pollution regions also contains uncertainty as we neither have direct measurements of the height distribution of the plumes over China during 463 transport nor do we have information on the pollution levels over China while the desert 464465 plumes travelled over China in the various height layers. We can merely assume that the likelihood of mixing with dust and pollution increases the lower the dust travels above 466 467 ground and the longer it travels at low heights.

We again used our classification of Case I and Case II. However, we refined the vertical resolution to 5 height layers, i.e. transport occurred below 1 km, from 1-2 km, from 2-3 km, from 3-4 km, and above 4 km. We wanted to test if a more refined height separation would give us more insight on the change of optical properties with transport time and transport height.

The absolute time the dust layers spent in these different height levels is presented in figure 9. We also tested the effect of relative time in relation to total transport time but could not find a clear pattern. We find a maximum value of 0.3 for δ_p at 532 nm. On average, the depolarization decreases with increasing residence time over China. However, this dependence differs with respect to the height above ground of dust layers. The change of the depolarization ratio of dust layers travelling above 3 km above ground seems less dependent on the residence time over a given area.

We believe that short residence times (fast transport to Korea, 20 hours or less) reduces the chances that pollution may mix with dust, particularly if dust travelled below 3 km above ground. In contrast, longer residence times (slow transport to Korea, >50 hours) of the dust plumes may have increased the chances that pollution mixed with dust if dust travelled below 3km height above ground.

Regarding \mathring{A}_{β} we find a maximum value of 1.75 which decreases to 0.5 for slow transport of the plumes. The decrease of \mathring{A}_{β} with transport time seems to be correlated for plumes that mainly stay below 3 km height above ground. In contrast, if plumes were above 3 km, \mathring{A}_{β} does not seem to change with transport time. In that case, the mixing of pollution with dust may have been less likely because of the fast transport.

With regard to S at 355 nm and 532 nm we find a maximum value of approximately 75 sr which drops to approximately 40 sr for slow transport. Again, we see that for plumes below 3 km height above ground transport time seems to matter. S drops with increasing transport time. For the case of plumes above 3 km, i.e. dust that likely is not too much affected by mixing with anthropogenic pollution, the lidar ratios do not seem to depend on transport time. This result may however again be caused by the fact that transport times to Korea are comparably short.

We further investigated these results. We initially assumed that \mathring{A}_{β} either should 497 increase with transport time or does not drop significantly for pollution that travels near the 498 ground as there should be a higher share of small anthropogenic pollution particles in the dust 499plume (large particles). This opposite behaviour may be caused by the state of mixing, i.e., 500 pollution particles attach to the dust particles, thus increasing their mean size. Hygroscopic 501 growth of particles attached to dust may further contribute to the increase of mean size. One 502 point that complicates this interpretation is that \mathring{A}_{B} does not only depend on particle size but 503 504 also on particle shape and the real and imaginary part (scattering and absorption) of the particles. 505

With regard to S we also expected that S would increase with increasing transport times. If the particles travel at low height above ground more anthropogenic pollution should mix with dust. The decrease of S however suggests an increase of particle size and a decrease of the light-absorption capacity. Hygroscopic particle growth, i.e. increase of mean particle size and decrease of light-absorption by uptake of water might be responsible for this behavior.

We stress that other reasons may be responsible for these results. We have a comparably small set of observations. We have insufficient information whether the plumes consisted of internal and/or external mixtures. The shape and size of particles of mixed Asian dust might be influenced in a much more significant way by transport time. The kind of mixing between the dust particles and pollution particles could influence the light-absorption properties.

518

519 **4. Summary and Conclusion**

In this study we presented the differences of optical properties of mixed Asian dust
 layers in dependence of their vertical position over China during transport from the Chinese

dust source regions to Korea, downwind region of the source regions. The data cover the time frame from 2009-2013. The dust layers are divided into several categories which can be characterized by different heights above ground during transport. The change of height above ground during transport of the dust layers was identified by backward trajectory analysis.

The optical properties of Asian dust significantly change in dependence of the dust 526 527 plumes, the vertical position, and the change of vertical position above ground level during transport over China. We find lower values of the lidar ratios at 355 and 532 nm, lower 528 backscatter-related Ångström exponents (wavelength pair 355/532 nm), and higher linear 529 particle depolarization ratios at 532 nm for Asian dust that was transported at high altitudes 530 531 (> 3km height above ground) compared to the situation in which the dust plumes moved at 532 low altitudes across China. The mean linear particle depolarization ratio is 0.21±0.06 for 533 transport at high altitudes. The mean lidar ratios in that case are 52 ± 7 sr and 53 ± 8 sr at 355 nm and 532 nm, respectively. The mean Ångström exponent is 0.74±0.31. These values likely 534 reflect properties of dust little affected by anthropogenic pollution. However, we cannot 535 quantify the amount of anthropogenic pollution that may still be present in these dust layers. 536 In contrast, higher values of the lidar ratios and the backscatter-related Ångström exponents, 537 and lower values of the linear particle depolarization were found for dust layers that crossed 538 highly polluted regions in China at low altitudes. The value of the mean linear particle 539 540 depolarization ratio is 0.13 ± 0.04 . The mean lidar ratios are 63 ± 9 sr and 62 ± 8 sr at 355 nm and 532 nm, respectively. The mean backscatter-related Ångström exponent is 0.98±0.35. 541 These values more likely describe strong influence by anthropogenic pollution, i.e. the uptake 542 543 of urban pollution by dust may have been significant.

544 Our results suggest that the transport pathway as well as the vertical position of Asian 545 dust during long-range transport may have significant impact on the optical properties of 546 mixed Asian dust layers.

547	Acknowledgements
-----	------------------

549	This work was supported by a National Research Foundation of Korea (NRF) grant funded
550	by the Korean government (MEST) (No. 2012R1A1A2002983). This work was also
551	supported by the Korea Meteorological Administration Research and Development Program
552	under grant KMIPA2015-2012.
553	
554	
555	
556	
557	
558	
559	
560 561	
562	
563	
564	
565	
566	
567	
568	
569	
570 571	
572	
573	
574	
575	
576	
577	
578	
579	
580 501	
582	
583	
584	
585	
586	
587	
588	
589	
590 501	
091	

592 **References**

593 594

595

596 597

598

Michaelis, W.: Combined Raman elastic-backscatter lidar for vertical profiling of moisture, 599600 aerosol extinction, backscatter, and lidar ratio, Applied Physics B, 55, 18-28, 1992a. 601 Ansmann, A., Riebesell, M., and Weitkamp, C.: Measurement of atmospheric aerosol 602 603 extinction profiles with a Raman lidar, Optics Letters, 15, 746-748, 1990. 604 Ansmann, A., Wandinger, U., Riebesell, M., Weitkamp, C., and Michaelis, W.: Independent 605 606 measurement of extinction and backscatter profiles in cirrus clouds by using a combined Raman elastic-backscatter lidar, Applied Optics, 31, 7113-7131, 1992b. 607 608 609 Behrendt, A., and Nakamura, T.: Calculation of the calibration constant of polarization lidar and its dependency on atmospheric temperature, Optics express, 10, 805-817, 2002. 610 611 Bellouin, N., Quaas, J., Morcrette, J.-J., and Boucher, O.: Estimates of aerosol radiative 612 613 forcing from the MACC re-analysis, Atmospheric Chemistry and Physics, 13, 2045-2062, 2013. 614 615 Bohren, C. F., and Huffman, D. R.: Absorption and scattering by a sphere, Absorption and 616 Scattering of Light by Small Particles, 82-129, 1983. 617 618 Burton, S., Ferrare, R., Hostetler, C., Hair, J., Rogers, R., Obland, M., Butler, C., Cook, A., 619 Harper, D., and Froyd, K.: Aerosol classification using airborne High Spectral Resolution 620 Lidar measurements-methodology and examples, Atmospheric Measurement Techniques, 5, 621 622 73-98, 2012. 623 Burton, S., Ferrare, R., Vaughan, M., Omar, A., Rogers, R., Hostetler, C., and Hair, J.: 624 625 Aerosol classification from airborne HSRL and comparisons with the CALIPSO vertical feature mask, Atmospheric Measurement Techniques, 6, 1397-1412, 2013. 626 627 628 Burton, S., Vaughan, M., Ferrare, R., Hostetler, C.: Separating mixtures of aerosol types in airborne High Spectral Resolution Lidar data, Atmospheric Measurement Techniques, 7, 419-629 630 436, 2014. 631 Cairo, F., Di Donfrancesco, G., Adriani, A., Pulvirenti, L., and Fierli, F.: Comparison of 632 various linear depolarization parameters measured by lidar, Applied Optics, 38, 4425-4432, 633 1999. 634 635 Carrico, C. M., Kus, P., Rood, M. J., Quinn, P. K., and Bates, T. S.: Mixtures of pollution, 636

Anderson, T. L., Masonis, S. J., Covert, D. S., Charlson, R. J., and Rood, M. J.: In situ

measurement of the aerosol extinction-to-backscatter ratio at a polluted continental site,

Ansmann, A., Riebesell, M., Wandinger, U., Weitkamp, C., Voss, E., Lahmann, W., and

Journal of Geophysical Research: Atmospheres (1984–2012), 105, 26907-26915, 2000.

- dust, sea salt, and volcanic aerosol during ACE-Asia: radiative properties as a function of
 relative humidity, J. Geophys. Res.-Atmos., 108(D23), 8650, doi: 10.1029/2003JD003405,
 2003.
- 640

- 641 Cattrall, C., Reagan, J., Thome, K., and Dubovik, O.: Variability of aerosol and spectral lidar
- and backscatter and extinction ratios of key aerosol types derived from selected Aerosol
- 643 Robotic Network locations, Journal of Geophysical Research: Atmospheres (1984–2012),
- 644 110, doi: 10.1029/2004JD005124, 2005.
- 645

657

- Cesnulyte, V., Lindfors, A., Pitkänen, M., Lehtinen, K., Morcrette, J.-J., and Arola, A.:
 Comparing ECMWF AOD with AERONET observations at visible and UV wavelengths,
 Atmospheric Chemistry and Physics, 14, 593-608, 2014.
- 649
 650 Chen, W.-N., Chen, Y.-W., Chou, C. C., Chang, S.-Y., Lin, P.-H., and Chen, J.-P.: Columnar
 651 optical properties of tropospheric aerosol by combined lidar and sunphotometer
 652 measurements at Taipei, Taiwan, Atmospheric Environment, 43, 2700-2708, 2009.
- Chen, W.-N., Tsai, F.-J., Chou, C. C.-K., Chang, S.-Y., Chen, Y.-W., and Chen, J.-P.: Optical
 properties of Asian dusts in the free atmosphere measured by Raman lidar at Taipei, Taiwan,
 Atmospheric Environment, 41, 7698-7714, 2007.
- De Tomasi, F., Blanco, A., and Perrone, M. R.: Raman lidar monitoring of extinction and
 backscattering of African dust layers and dust characterization, Applied Optics, 42, 16991709, 2003.
- 661

670

675

- Draxler, R. R. and Rolph, G.,: HYSPLIT (Hybrid Single-Particle Lagrangian Integrated
 Trajectory) model access via NOAA ARL READY website, NOAA Air Resources
 Laboratory, Silver Spring, MD, available at: <u>http://www.arl.noaa.gov/ready/hysplit4.html</u>
 (last access: October 2014), 2003.
- Durant, A. J., Harrison, S. P., Watson, I. M., and Balkanski, Y.: Sensitivity of direct radiative
 forcing by mineral dust to particle characteristics, Progress in Physical Geography, 33, 80102, 2009.
- Eck, T., Holben, B., Reid, J., Dubovik, O., Smirnov, A., O'neill, N., Slutsker, I., and Kinne,
 S.: Wavelength dependence of the optical depth of biomass burning, urban, and desert dust
 aerosols, Journal of Geophysical Research: Atmospheres (1984–2012), 104, 31333-31349,
 1999.
- Ferrare, R. A., Turner, D. D., Brasseur, L. H., Feltz, W. F., Dubovik, O., and Tooman, T. P.:
 Raman lidar measurements of the aerosol extinction-to-backscatter ratio over the Southern
 Great Plains, Journal of Geophysical Research: Atmospheres (1984–2012), 106, 2033320347, doi:10.1029/2000JD000144, 2001.
- 680
- Franke, K., Ansmann, A., Müller, D., Althausen, D., Venkataraman, C., Reddy, M. S.,
 Wagner, F., and Scheele, R.: Optical properties of the Indo-Asian haze layer over the tropical
 Indian Ocean, Journal of Geophysical Research: Atmospheres (1984–2012), doi:
 10.1029/2002JD002473, 108, 2003.
- 686 Freudenthaler, V., Esselborn, M., Wiegner, M., Heese, B., Tesche, M., Ansmann, A., Müller,
- D., Althausen, D., Wirth, M., and Fix, A.: Depolarization ratio profiling at several
- wavelengths in pure Saharan dust during SAMUM 2006, Tellus B, 61, 165-179, 2009.
- 689

- 690 Griggs, D. J., and Noguer, M.: Climate change 2001: the scientific basis. Contribution of 691 working group I to the third assessment report of the intergovernmental panel on climate 692 change, Weather, 57, 267-269, 2002.
- Huang, J., Minnis, P., Chen, B., Huang, Z., Liu, Z., Zhao, Q., Yi, Y., and Ayers, J. K.: Longrange transport and vertical structure of Asian dust from CALIPSO and surface
 measurements during PACDEX, J. Geophys. Res.-Atmos., 113(D23), D23212, doi:
 10.1029/2008JD010620, 2008.
- Huebert, B. J., Bates, T., Russell, P. B., Shi, G., Kim, Y. J., Kawamura, K., Carmichael, G.,
 and Nakajima, T.: An overview of ACE-Asia: strategies for quantifying the relationships
 between Asian aerosols and their climatic impacts, J. Geophys. Res.-Atmos., 108(D23), 8633,
 doi: 10.1029/2003JD003550, 2003.
- 702

- Husar, R. B., Tratt, D., Schichtel, B. A., Falke, S., Li, F., Jaffe, D., Gasso, S., Gill, T.,
 Laulainen, N. S., and Lu, F.: Asian dust events of April 1998, Journal of Geophysical
 Research: Atmospheres (1984–2012), 106, 18317-18330, 2001.
- 706
- Inness, A., Baier, F., Benedetti, A., Bouarar, I., Chabrillat, S., Clark, H., Clerbaux, C.,
 Coheur, P., Engelen, R., and Errera, Q.: The MACC reanalysis: an 8 yr data set of
 atmospheric composition, Atmos Chem Phys, 13, 4073-4109, 2013.
- 710

716

720

- Iwasaka, Y., Shibata, T., Nagatani, T., Shi, G. Y., Kim, Y., Matsuki, A., Trochkine, D.,
 Zhang, D., Yamada, M., and Nagatani, M.: Large depolarization ratio of free tropospheric
 aerosols over the Taklamakan Desert revealed by lidar measurements: Possible diffusion and
 transport of dust particles, J. Geophys. Res.-Atmos., 108(D23), 8652, doi:
 10.1029/2002JD003267, 2003.
- Jacobson, M. Z.: Investigating cloud absorption effects: Global absorption properties of black
 carbon, tar balls, and soil dust in clouds and aerosols, Journal of Geophysical Research:
 Atmospheres (1984–2012), 117, D06205, doi: 10.1029/2011JD017218, 2012
- Mahowald, N. M., Muhs, D. R., Levis, S., Rasch, P. J., Yoshioka, M., Zender, C. S., and Luo,
 C.: Change in atmospheric mineral aerosols in response to climate: last glacial period,
 preindustrial, modern, and doubled carbon dioxide climates, J. Geophys. Res.-Atmos.,
 111(D10), D10202, doi: 10.1029/2005JD006653, 2006.
- Mattis, I., Ansmann, A., Müller, D., Wandinger, U., and Althausen, D.: Dual-wavelength
 Raman lidar observations of the extinction-to-backscatter ratio of Saharan dust, Geophysical
 Research Letters, 29, 20-21-20-24, 2002.
- 729
- McKendry, I., Hacker, J., Stull, R., Sakiyama, S., Mignacca, D., and Reid, K.: Long-range
 transport of Asian dust to the Lower Fraser Valley, British Columbia, Canada, Journal of
 Geophysical Research: Atmospheres (1984–2012), 106, 18361-18370, 2001.
- 733
- Mikami, M., Shi, G., Uno, I., Yabuki, S., Iwasaka, Y., Yasui, M., Aoki, T., Tanaka, T.,
 Kurosaki, Y., and Masuda, K.: Aeolian dust experiment on climate impact: An overview of
 Japan–China joint project ADEC, Global and Planetary Change, 52, 142-172, 2006.
- 737
- Müller, D., Ansmann, A., Mattis, I., Tesche, M., Wandinger, U., Althausen, D., and Pisani,

- G.: Aerosol-type-dependent lidar ratios observed with Raman lidar, J. Geophys. Res.Atmos.,112(D16), D16202, doi: 10.1029/2006JD008292, 2007.
- Müller, D., Franke, K., Ansmann, A., Althausen, D., and Wagner, F.: Indo-Asian pollution
 during INDOEX: microphysical particle properties and single-scattering albedo inferred from
 multiwavelength lidar observations, J. Geophys. Res.-Atmos., 108(D19), 4600, doi:
 10.1029/2003JD003538, 2003.
- Müller, D., Mattis, I., Tatarov, B., Noh, Y., Shin, D., Shin, S., Lee, K., Kim, Y., and
 Sugimoto, N.: Mineral quartz concentration measurements of mixed mineral dust/urban haze
 pollution plumes over Korea with multiwavelength aerosol Raman-quartz lidar, Geophys.
 Res. Lett., 37(20), L20810, doi: 10.1029/2010GL044633, 2010.
- Murayama, T., Masonis, S. J., Redemann, J., Anderson, T. L., Schmid, B., Livingston, J. M.,
 Russell, P. B., Huebert, B., Howell, S. G., and McNaughton, C. S.: An intercomparison of
 lidar-derived aerosol optical properties with airborne measurements near Tokyo during
 ACE-Asia, Journal of Geophysical Research: Atmospheres (1984–2012), 108, doi:
 10.1029/2002JD003259, 2003.
- Murayama, T., Müller, D., Wada, K., Shimizu, A., Sekiguchi, M., and Tsukamoto, T.:
 Characterization of Asian dust and Siberian smoke with multi-wavelength Raman lidar over
 Tokyo, Japan in spring 2003, Geophys. Res. Lett., 31(23), L23103, doi:
 10.1029/2004GL021105, 2004.
- 762

741

746

751

- Murayama, T.: Optical properties of Asian dust aerosol lofted over Tokyo observed by
 Raman lidar, Lidar Remote Sensing in Atmospheric and Earth Sciences, edited by
 Bissonnette, LR, Roy, G., and Vallée, G., Defence R&D Canada, Val-Bélair, 1, 331-334,
 2002.
- Noh, Y. M., Kim, Y. J., Choi, B. C., and Murayama, T.: Aerosol lidar ratio characteristics
 measured by a multi-wavelength Raman lidar system at Anmyeon Island, Korea,
 Atmospheric Research, 86, 76-87, 2007.
- Noh, Y. M., Kim, Y. J., and Müller, D.: Seasonal characteristics of lidar ratios measured with
 a Raman lidar at Gwangju, Korea in spring and autumn, Atmospheric Environment, 42,
 2208-2224, 2008.
- Omar, A. H., Winker, D. M., Vaughan, M. A., Hu, Y., Trepte, C. R., Ferrare, R. A., Lee, K.P., Hostetler, C. A., Kittaka, C., and Rogers, R. R.: The CALIPSO automated aerosol
 classification and lidar ratio selection algorithm, Journal of Atmospheric and Oceanic
 Technology, 26, 1994-2014, 2009.
- 780

771

- Sakai, T., Shibata, T., Iwasaka, Y., Nagai, T., Nakazato, M., Matsumura, T., Ichiki, A., Kim,
 Y.-S., Tamura, K., and Troshkin, D.: Case study of Raman lidar measurements of Asian dust
 events in 2000 and 2001 at Nagoya and Tsukuba, Japan, Atmospheric Environment, 36,
 5479-5489, 2002.
- 785
- Shimizu, A., Sugimoto, N., Matsui, I., Arao, K., Uno, I., Murayama, T., Kagawa, N., Aoki,
 K., Uchiyama, A., and Yamazaki, A.: Continuous observations of Asian dust and other

aerosols by polarization lidars in China and Japan during ACE-Asia, J. Geophys. Res.Atmos., 109(D19), D19S17, doi: 10.1029/2002JD003253, 2004.

Shin, S., Müller, D., Kim, Y., Tatarov, B., Shin, D., Seifert, P., and Noh, Y. M.: The retrieval
of the Asian dust depolarization ratio in Korea with the correction of the polarizationdependent transmission, Asia-Pacific Journal of Atmospheric Sciences, 49, 19-25, 2013.

- Sugimoto, N., and Lee, C. H.: Characteristics of dust aerosols inferred from lidar
 depolarization measurements at two wavelengths, Applied Optics, 45, 7468-7474, 2006.
- Sugimoto, N., Nishizawa, T., Shimizu, A., Matsui, I., and Kobayashi, H.: Detection of
 internally mixed Asian dust with air pollution aerosols using a polarization optical particle
 counter and a polarization-sensitive two-wavelength lidar, Journal of Quantitative
 Spectroscopy and Radiative Transfer, 150, 107-113, 2015.
- Sun, Y., Zhuang, G., Huang, K., Li, J., Wang, Q., Wang, Y., Lin, Y., Fu, J. S., Zhang, W., 803 and Tang, A.: Asian dust over northern China and its impact on the downstream aerosol 804 805 chemistry in 2004, J. Geophys. Res.-Atmos., 115(D17), D00K09, doi: 10.1029/2009JD012757, 2010. 806
- Sun, Y., Zhuang, G., Wang, Y., Zhao, X., Li, J., Wang, Z., and An, Z.: Chemical composition
 of dust storms in Beijing and implications for the mixing of mineral aerosol with pollution
 aerosol on the pathway, J. Geophys. Res.-Atmos., 110(D24), D24209, doi:
 10.1029/2005JD006054, 2005.
- 812

790

794

797

802

807

- Tatarov, B., Müller, D., Shin, D. H., Shin, S. K., Mattis, I., Seifert, P., Noh, Y. M., Kim, Y.,
 and Sugimoto, N.: Lidar measurements of Raman scattering at ultraviolet wavelength from
 mineral dust over East Asia, Optics express, 19, 1569-1581, 2011.
- 816

- Tesche, M., Ansmann, A., Müller, D., Althausen, D., Engelmann, R., Freudenthaler, V., and 817 818 Groß, S.: Vertically resolved separation of dust and smoke over Cape Verde using multiwavelength Raman and polarization lidars during Saharan Mineral Dust Experiment 819 820 2008. Journal of Geophysical Research: Atmospheres (1984 - 2012),114. 821 doi:10.1029/2009JD011862, 2009.
- Tesche, M., Ansmann, A., Müller, D., Althausen, D., Mattis, I., Heese, B., Freudenthaler, V.,
 Wiegner, M., Esselborn, M., and Pisani, G.: Vertical profiling of Saharan dust with Raman
 lidars and airborne HSRL in southern Morocco during SAMUM, Tellus B, 61, 144-164, 2009.
- Wandinger, U., and Ansmann, A.: Experimental determination of the lidar overlap profile
 with Raman lidar, Applied Optics, 41, 511-514, 2002.
- 829
- Wang, Y., Zhuang, G., Tang, A., Zhang, W., Sun, Y., Wang, Z., and An, Z.: The evolution of
 chemical components of aerosols at five monitoring sites of China during dust storms,
 Atmospheric Environment, 41, 1091-1106, 2007.
- Xie, C., Zhao, M., Wang, B., Zhong, Z., Wang, L., Liu, D., and Wang, Y.: Study of the
 scanning lidar on the atmospheric detection, Journal of Quantitative Spectroscopy and
 Radiative Transfer, 150, 114-120, 2015.

837 838 839 840	Yi, B., Yang, P., and Baum, B. A.: Impact of pollution on the optical properties of trans- Pacific East Asian dust from satellite and ground-based measurements, J. Geophys. Res.Atmos., 119(9), 5397-5409, doi: 10.1002/2014JD021721, 2014.
841 842 843 844	Yu, X., Cheng, T., Chen, J., and Liu, Y.: A comparison of dust properties between China continent and Korea, Japan in East Asia, Atmospheric Environment, 40, 5787-5797, 2006.
845	
846	
847	
848	
849	
850	
851	
852	
853	
854	
855	
856	
857	
858	
859	
860	
861	
862	
863	
864	

865 Table captions

Table 1. Summary of the linear particle depolarization ratio at 532 nm, lidar ratios, and backscatter-related Ångström exponents of Asian dust layers for each classification. Asian dust layers were classified according to (a) levels of anthropogenic pollution emission; LP denotes that Asian dust layers which are considered as less polluted and MP denotes that Asian dust layers which are considered as more polluted, (b) their vertical position at polluted region; Case I indicates Asian dust layers passed over China at high altitude (> 3 km) before they arrived over Gwangju, and Case II indicates Asian dust layers were transported at low altitude (< 3 km) over industrialized areas before they arrived over Gwangju, and (c) their vertical position (below 3 km or above 3 km) and level of pollution (LP or MP) when they passed over China

Table 2. Linear particle depolarization ratio at 532 nm, lidar ratios, and backscatter-related Ångström exponents of East Asian dust layers according to altitude range in which these plumes passed over polluted regions of China. Case I describes the layer from 3-4 km and above 4 km. Case II describes the layers from 0-1 km, from 1-2 km, and from 2-3 km height above ground.

898 Figure captions

899

Figure 1. Map of the desert regions (Taklimakan desert, Gobi desert, Badain Jaran desert, Ordos
Desert, Inner Mongolia plateau, and Manchuria) and loess regions (Loess Plateau and Manchuria).
The location of some major cities (Beijing and Shanghai) and industrialized areas of China (Hebei,
Shandong, Henan, and Zhejiang province) is also shown. MRS.LEA is located in Gwangju, Korea.

Figure 2. Measurement on 22 April 2012, 13:15-14:05 UTC. Shown are (a) the time-height cross
section of the range-corrected signal and (b) the volume depolarization ratio at 532 nm. Also shown
are the profiles of (c) the volume depolarization ratio and the linear particle depolarization ratio at 532
nm, and (d) the lidar ratio at 355 and 532 nm and the backscatter-related Ångström exponents.

- Figure 3. Frequency distributions of optical properties of Asian dust observed between 2009 and 2013. Shown are (a,b) lidar ratios at 355 and 532 nm, (c) linear particle depolarization ratios at 532 nm, and (d) Ångström exponents for the wavelength pair 355/532 nm. The numbers in each plot indicate the mean value and its standard deviation, the median (shown in brackets), and the minimum and maximum value of each distribution.
- 915

Figure 4. Distribution of AOD at 550 nm over East Asia retrieved from ECMWF for (a) and (e) dust,
(b) and (f) organic matter, (c) and (g) black carbon, and (d) and (h) sulphate aerosol. (a) - (d) refers to
8 March 2013. That day is classified as a relatively "low polluted" day over East China. (e) - (h) refers
to 10 April 2011 which is classified as a comparably "highly polluted" day over East China.

920

Figure 5. Scatter diagram of the linear particle depolarization at 532 nm versus (a), (d) the backscatter-related Ångström exponent (355/532 nm wavelength pair), (b), (e) the lidar ratio at 355 nm and (c), (f) the lidar ratio at 532 nm. The left column (a-c) shows the optical properties of Asian dust layers considered as less polluted (LP), the right column (d-f) shows the more polluted cases.

926 Figure 6. (top panel) Transport pattern of the dust plumes that originated in the desert regions and 927 passed over industrialized/populated regions of China before arrival over the Korean peninsula. (middle panel) Vertical position of the dust layers during transport: (a) Dust layers passed over China 928 929 at high altitude (Case I) (b) dust layers were transported over China through the near surface/lower 930 troposphere (Case II). (bottom panel) Scatter diagram of the linear particle depolarization at 532 nm 931 versus (c) the backscatter-related Ångström exponent (355/532 nm wavelength pair), and the (d), (e) 932 lidar ratio (at 355 nm and at 532 nm) with respect to Case I and Case II. The two categories I, II are 933 denoted by different colors. Case I is indicated by red circles. Case II is indicated by black circles. 934

934 935

Figure 7. Scatter diagram of the linear particle depolarization at 532 nm versus (a), (d) the backscatter-related Ångström exponent (355/532 nm wavelength pair), (b), (e) the lidar ratio at 355 nm and (c), (f) the lidar ratio at 532 nm. The left column (a-c) shows the optical properties of Asian dust layers considered as less polluted (LP), the right column (d-f) shows the more polluted cases. The Asian dust layers that passed over polluted regions in China at low altitude are denoted by black circles. The Asian dust layers transported at high altitude are denoted by red squares.

942

Figure 8. (top panel) (a) transport path and classification of East Asian dust layers with respect to (b) their altitude above ground when they passed over industrial regions of China. (bottom panel) transport path and corresponding altitude of Asian dust layers are distinguished by color. (black: 0 km -1 km; green: 1 km -2 km; purple: 2 km -3 km; blue: 3 km -4 km; red: above 4 km). Scatter plots of the linear particle depolarization at 532 nm (dark yellow), the backscatter-related Ångström exponent (355/532 nm wavelength pair, red), the lidar ratio at 355 nm (blue), the lidar ratio at 532 nm (green) in dependence of the 5 altitude categories (c). The height of the Asian dust layers above ground is separated by vertical lines. Case I included the layers from 3 – 4 km and above 4 km. Case II includes the layers from 0 - 1 km, from 1-2 km, and from 2-3 km height above ground.

Figure 9. Scatter diagram of optical dust properties versus the time the Asian dust layers travelled over polluted regions in China. Shown are (a) the particle depolarization ratio, (b) the backscatter-related Ångström exponent (355/532 nm wavelength pair), (c) the lidar ratio at 355 nm, and (d) the lidar ratio at 532 nm with respect to their altitude above ground when they passed over industrial regions of China. The corresponding altitude of Asian dust layer are distinguished by color. (black: below 0 km; green: 1 km – 2km; purple: 2km – 3 km; blue: 3 km – 4 km; red: above 4 km)

978 Tables

979 [1]

Closeifi	Classification		2	S	Å	
Classing			0 _p	355 nm	532 nm	- Α _β
Dollution laval ^(a)	Less Polluted	25	0.17 ± 0.02	57±7	55±7	0.82±0.37
Pollution level	More Polluted	13	0.17 ± 0.02	58±6	59±8	0.89±0.38
Vartical positor ^(b)	Case I	16	0.21±0.06	52±7	53±8	0.74±0.31
vertical positon	Case II	22	0.13 ± 0.04	63±9	62 <u>±</u> 8	0.98±0.35
	LP_below 3 km	12	0.13±0.03	64±9	62 <u>±</u> 8	1.00 ± 0.38
Pollution leve &	LP_above 3 km	13	0.21±0.05	51±8	49±9	0.65±0.20
Vertical position ^(c)	MP_below 3 km	8	0.13±0.04	61±10	64±7	1.09±0.30
	MP_above 3 km	5	0.24±0.05	53±5	53±2	0.58±0.14

980

981 [2]

	Height of dust layer at		Number of	2	S	[sr]	Å	
	pollut	ion regions	observed layers	o _p	355nm	532nm	A _β	
	Case I	Above 4 km	14	0.23±0.02	50±7	49±8	0.60±0.27	
		3 km-4 km	1	0.20 ± 0.04	44±2	47±7	0.67±0.29	
	Case II	2 km-3 km	7	0.13±0.02	61±7	66±5	1.11±0.47	
		1 km-2 km	6	0.15±0.03	65±7	59±9	0.94 ± 0.42	
		Below 1 km	10	0.12±0.01	63±7	64±6	1.00±0.43	
982								
983								
984								
985								
986								
987								
988								
989								
990								
991								
992								
993								
994								

- 995 Figures
- 996 [1]





[3]





l	mars-atls03-20140925105914-63601-4506.grib										
l	03/08/2013 03:00 GMT										
l	Sulphate_Aerosol_Optical_Depth_at_550nm_surface ()										
l											
l	0.00	0.20	0.40	0.60	0.80	1.00					



03/00/2013 03/00 GMT									
Black_Carbon_Aerosol_Optical_Depth_at_550nm_surface ()									
0.00 0.02 0.04 0.06 0.08 0.	10								



mars-atls03-20140925105914-63601-4506.grib 03/08/2013 03:00 GMT									
Organic_	Matter_Aer	osol_Optic	al_Depth_a	at_550nm_	surface (
0.00	0.02	0.04	0.06	0.08	0.10				



mars-a Dust_Ae	atls03-20 03, rosol_0	1409251 /08/2013 ptical_De	.05914-6 : 03:00 G :pth_at_5	3601-45 MT 50nm_s	06.grib urface
0.00	0.14	0.28	0.42	0.56	0.70





mars-atls03-20140925102747-63659-4243.grib 04/10/2011 03:00 GMT Black_Carbon_Aerosol_Optical_Depth_at_550nm_surface ()					mars Sulphate	-atls03-2 04 _Aerosol	01409251 l/10/2011 _Optical_l	102747-63 03:00 GM Depth_at_	3659-4243 MT _550nm_s	}.grib ;urface ()	
0.00	0.02	0.04	0.06	0.08	0.10	0.00	0.20	0.40	0.60	0.80	1.00



mars-atls03-20140925102747-63659-4243.grib 04/10/2011 03:00 GMT								
Organic_Matter_Aerosol_Optical_Depth_at_550nm_surface()								
0.00	0.02	0.04	0.06	0.08	0.10			



mars-atls03-20140925102747-63659-4243.grib 04/10/2011 03:00 GMT Dust_Aerosol_Optical_Depth_at_550nm_surface (
0.00	0.14	0.28	0.42	0.56	0.70















[9]