

Effects of Aerosols on Clear Sky Solar Radiation in the ALADIN-HIRLAM NWP System

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Abstract

The direct shortwave radiative effect of aerosols under clear sky conditions in the ALADIN-HIRLAM numerical weather prediction system was investigated using three shortwave radiation schemes in diagnostic single-column experiments: the IFS scheme, the acraneb2 scheme and the hlradia radiation scheme. The aim was to evaluate the strengths and weaknesses of the NWP system in this regard and to prepare it for use of real-time aerosol information. The experiments were run with particular focus on the August 2010 Russian wildfire case. Each of the three radiation schemes accurately (within $\pm 4\%$ at midday) simulates the direct shortwave aerosol effect when observed aerosol optical properties are used. When the aerosols were excluded from the simulations, errors of more than +15 % in global shortwave irradiance were found at midday, with the error reduced to +10 % when standard climatological aerosols were used. An error of -11 % was seen at midday if only observed aerosol optical depths at 550 nm, and not observation-based spectral dependence

of aerosol optical depth, single scattering albedos and asymmetry factors, were included in the simulations. This demonstrates the importance of using the correct aerosol optical properties. The dependency of the direct radiative effect of aerosols on relative humidity was tested and shown to be within $\pm 6\%$ in this case. By modifying the assumptions about the shape of the IFS climatological vertical aerosol profile, the inherent uncertainties associated with assuming fixed vertical profiles were investigated. The shortwave heating rates in the boundary layer changed by up to a factor of 2 in response to the aerosol vertical distribution. Finally, we tested the radiative transfer approximations used in the three radiation schemes for typical aerosol optical properties compared to the accurate DISORT model. These approximations are found to be accurate to within $\pm 13\%$ even for large aerosol loads.

1 Introduction

The direct radiative effect of aerosols resulting from scattering and absorption of electromagnetic radiation at shortwave (SW) and longwave (LW) wavelengths has an impact on the Earth's radiation budget (e.g. Haywood and Boucher, 2000; Bellouin et al., 2005; Jacobson, 2001; Myhre et al., 2013; Yu et al., 2006; Loeb and Manalo-Smith, 2005) and on meteorology (e.g. Cook and Highwood, 2004; Takemura et al., 2005; Wang, 2004; Mulcahy et al. 2014, Bangert et al., 2012) which needs to be accounted for in numerical weather prediction (NWP) models. Climatological distributions of aerosols are commonly used in present-day operational NWP models for calculating the direct radiative effect of aerosols.

Using unrealistic aerosol distributions can lead to considerable errors in meteorological forecasts. Milton et al. (2008) showed that excluding the direct radiative effect of mineral dust and biomass burning aerosols in forecasts using the UK Met Office Unified Model during the dry season in West Africa resulted in an inaccurate representation of the surface energy budget and a warm bias in screen level temperature. Carmona et al. (2008) presented significant correlations between errors in the aerosol optical depth (AOD) assumed in an NWP model and temperature forecast errors. Accurate simulation of the direct radiative effect of aerosols on SW radiation is important to the growing solar energy industry because under clear-sky conditions aerosols are the main modulator of SW fluxes (Breitkreuz et al.,

2009).

The monthly aerosol climatology described in Tegen et al. (1997) is used in ECMWF's (*the European Centre for Medium Range Weather Forecasts*) global Integrated Forecast System, IFS, and in the Aire Limitee Adaptation dynamique Developpement INternational – High Resolution Limited Area Model (ALADIN-HIRLAM) limited area modelling system used in this study. Tompkins et al. (2005) showed that replacing the Tanre et al. (1994) fixed average aerosol distribution in ECMWF's IFS model by the Tegen climatology improved forecasts of the African Easterly Jet. This change in the aerosol climatology also improved the forecast skill and seasonal mean errors (Rodwell and Jung, 2008).

Including a more complete representation of the effects of aerosols in NWP models can improve the meteorological forecasts and is an active area of research (e.g. Mulcahy et al., 2014, Bangert et al., 2012). Using real-time aerosol distributions, rather than climatological datasets, to account for the direct radiative effect of aerosols further improves the quality of the forecasts. Toll et al. (2015b) showed that the accuracy of the forecasts of near-surface conditions by the ALADIN-HIRLAM system during severe wildfires in summer 2010 in Eastern Europe were improved when the direct radiative effect of the realistic aerosol distribution was included in the model hindcasts. Palamarchuk et al. (2016) also found a noticeable sensitivity of the ALADIN-HIRLAM forecasts to the treatment of aerosols. On the other hand, Toll et al. (2016) showed that when observed aerosol distributions are close to average, improvements in the SW radiation, temperature and humidity forecasts in the lower troposphere are only slightly greater when time-varying realistic aerosol data from the Monitoring Atmospheric Composition and Climate (MACC) reanalysis (Inness et al., 2013) is used in place of the Tegen climatology. Similar conclusions were drawn by Zamora et al. (2005) who showed that for small AODs accounting for the climatological average direct radiative effect of aerosols gives very good estimates of SW fluxes, but large biases occur when the AOD is large.

Baklanov et al. (2014), Grell and Baklanov (2011) and Zhang (2008) have suggested using coupled air quality and NWP models to improve forecasts of both air quality and weather. However, for operational NWP such coupled models are still too demanding computationally, and this added cost has to be evaluated versus improvements in the meteorological forecasts. Mulcahy et al. (2014), Morcrette et al. (2011) and Reale et al. (2011) describe improved forecasts of the radiation budget and near surface conditions in

1 global NWP models when prognostic aerosols are included; however the impact of aerosols
2 on large scale atmospheric dynamics is generally weak.

3 The AOD at 550 nm (AOD550 hereafter in this paper) and aerosol inherent optical
4 properties (IOPs: spectral dependence of AOD, single scattering albedo (SSA) and
5 asymmetry factor (g)) depend on the size, shape and the complex refractive indices of the
6 aerosols and have a significant effect on global downwelling SW (SWD) fluxes.
7 Hygroscopic induced changes in the IOPs of different aerosols types also alter the radiative
8 effect of aerosols (e.g. Cheng et al., 2008; Bian et al., 2009; Markowicz et al., 2003; Zieger
9 et al., 2013). For example, Magi and Hobbs (2003) present measurements of enhanced
10 backscatter by biomass burning aerosols when the relative humidity (RH) is high. Pilinis et
11 al. (1995) estimated that the global forcing due to aerosols doubles for a relative humidity
12 increase from 40% to 80%.

13 The vertical profile of aerosols is also very important when estimating their direct radiative
14 effect. There are considerable variations in the vertical distributions of aerosols over Europe
15 (Guibert et al., 2005; Matthias et al., 2004). Therefore, inaccuracies result when constant
16 climatological profiles per aerosol species are used (as is the case in ALADIN-HIRLAM
17 which uses the profiles of Tanre et al., 1984). For example, Guibert et al. (2005) analysed
18 the vertical profiles of aerosol extinction over Europe and found that aerosols over southern
19 Europe are concentrated higher in the atmosphere due to the occurrence of dust storm
20 episodes. Meloni et al. (2005) showed that under clear-sky conditions the direct radiative
21 effect of aerosols on surface radiation has a low dependence on the aerosol vertical profile,
22 but that the profile has an impact on the top of the atmosphere forcing, especially for
23 absorbing aerosols. Toll et al. (2015) evaluated the profile of the aerosol attenuation
24 coefficient for land aerosols in ALADIN-HIRLAM against observations for the summer
25 2010 Russian wildfires. They found good agreement between the distribution assumed in
26 the model and CALIOP measurements. However, a more general evaluation of the vertical
27 profile of aerosols in the system has not been performed.

28 The main goal of the present study is to focus on the impact of AOD550, aerosol IOPs, the
29 vertical distribution of aerosols, relative humidity and radiative transfer algorithms on SW
30 fluxes in diagnostic single column clear-sky experiments using the ALADIN-HIRLAM
31 system. Such experiments are extremely useful for developing and testing parametrizations
32 and for running idealised experiments that focus on atmospheric physics in a simplified

framework. With these experiments we can evaluate the strengths and weakness of the NWP model regarding the treatment of the direct radiative effect of aerosols.

The paper is structured as follows: the model set-up and radiation schemes are described in Section 2; the aerosol datasets and atmospheric and surface input used in the experiments are detailed in Section 3; descriptions of each of the experiments and sensitivity tests are provided in Section 4; the results and discussion are presented in Section 5 and conclusions and future work are summarised in Section 6.

2 Model Set-up

2.1 ALADIN-HIRLAM

The ALADIN-HIRLAM numerical weather prediction system is used for operational weather forecasting by 26 national meteorological services in Europe and North Africa which form the HIRLAM (*High Resolution Limited Area Model*) and ALADIN (*Aire Limitée Adaptation dynamique Développement International*) consortia. Pottier (2016) summarises 42 limited area configurations of the system used by the consortia members. This system can also be used for regional climate simulations (Lindstedt et al., 2015), where the direct radiative effect of aerosols can be of greater importance than in short range NWP applications.

The HARMONIE-AROME configuration based on Seity et al., 2011 was used in this study. HARMONIE (*HIRLAM ALADIN Regional Mesoscale Operational NWP in Europe*) denotes the specific configuration of the ALADIN-HIRLAM system maintained by the HIRLAM consortium; AROME is a limited area model developed at Météo-France. Its default set-up for operational NWP uses a 2.5 km horizontal grid and 65 hybrid model levels with deep convection treated explicitly. This configuration uses ALADIN non-hydrostatic dynamics (Bénard et al., 2010), Meso-NH physics (Mascart and Bougeault, 2011) and the SURFEX externalised surface scheme (Masson et al., 2013). Surface physiographies are prescribed using the 1 km resolution ECOCLIMAP II database (Faroux et al., 2012) and surface elevation is based on GTOPO30 (USGS, 1998).

We used the single column version of HARMONIE-AROME based on Malardel et al.

(2006) for the experiments detailed in this paper. As in Malardel et al. we will refer to this model configuration as MUSC (*Modèle Unifié Simple Colonne*). It includes all of the atmospheric and surface parametrizations of HARMONIE-AROME but lacks the large-scale dynamics, horizontal advection, pressure gradient force and large-scale vertical motion. Because of the simplifying assumptions, MUSC is not suitable for operational weather forecasting. However, its value lies in the fact that it provides a useful means of studying the sensitivity of the model output to realistic atmospheric conditions and different physical parametrizations. The input to MUSC is derived from the output of a 3D HARMONIE-AROME experiment. This includes the initial conditions in the atmosphere and surface, surface properties, atmospheric temperatures, specific humidities and wind speeds. Details on the input data used in our experiments are provided in Section 3.4.

2.2 Radiation schemes

In this study three shortwave radiation schemes were applied in MUSC: 1) The IFS radiation scheme based on cycle 25r1 (Morcrette, 1991; White, 2004), 2) A new version of the HIRLAM radiation scheme called hlradia (Savijärvi 1990) containing aerosol parametrizations, 3) The acraneb2 scheme (Mašek et al., 2015). Table 1 summarises the main characteristics of these radiation schemes.

Each scheme treats the atmosphere as a 1D column consisting of a set of plane-parallel homogeneous layers. The grid box is split into a cloudy fraction and a clear-sky fraction and does not allow lateral exchanges between them. Atmospheric composition (i.e. aerosols, clouds and atmospheric gases) and the radiative properties of the surface are required as input to the radiation schemes. MUSC was run under clear-sky conditions for the experiments and sensitivity studies presented in this paper. Thus, details on cloud particles and cloud cover are not included. Further information on the basic differences between the radiation schemes is given in the following sub-sections.

2.2.1 IFS

The IFS SW radiation scheme (ECMWF, 2004; IFS cycle 25R1) is used by default in MUSC and is the most detailed of the three schemes applied in our experiments. It contains six SW spectral bands (0.185–0.25–0.44–0.69–1.19–2.38–4.00 μm), three in the ultraviolet/visible spectral range and three in the solar infrared range (Mascart and

Bougeault, 2011; White, 2004), and 14 LW bands. The IFS clear-sky SW radiative transfer is calculated using the Fouquart and Bonnel (1980) two-stream equations in Mascart and Bougeault (2011) where the reflectance, absorption and transmittance of the atmospheric layers are calculated in a similar manner to that outlined in Coakley and Chylek (1975). These calculations use the IOPs of aerosols and atmospheric molecules. Monthly climatologies of vertically integrated AOD550 for six aerosol categories - continental, sea, urban, desert, volcanic and background stratospheric are used (Tegen et al., 1997). These aerosols are distributed among the model levels using Tanre et al. (1984) climatological vertical profiles for each type as described in Section 3.1; these profiles are used in the calculations of model-level fluxes and heating rates. The aerosol IOPs are parametrized following Hess et al. (1998). Monthly climatologies of ozone and a fixed composition mixture of CO₂, N₂O, CH₄ and O₂ are also used. Further details on the aerosols are given in Section 3.1.

2.2.2 Hlradia

Hlradia, the simplest of the three schemes, considers 1 SW and 1 LW spectral band. Clear-sky transmittance, reflectance and absorptance of SW flux are taken into account at each model-level to obtain the radiative heating (vertical divergence of the net SW flux) and net SW fluxes. The radiative transfer is parametrized rather than solved explicitly, in order to make the scheme very fast for NWP use (Savijärvi, 1990). The impact of ozone, oxygen and carbon dioxide on SW irradiance is assumed to be constant over time and space. In older versions of the scheme, aerosols were accounted for using constant coefficients. However, the scheme has recently been modified to include parametrizations of the direct and semi-direct effects of aerosols, calculated using the two-stream approximation equations for anisotropic non-conservative scattering described by Thomas & Stamnes (2002).

Hlradia uses the GADS/OPAC aerosols of Koepke et al. (1997) and includes the following species: soot, minerals (nucleation, accumulation, coarse and transported modes), sulphuric acid, sea salt (accumulation and coarse modes), water soluble, and water insoluble aerosols. The IFS aerosols types described in Section 2.2.1 are mapped to GADS/OPAC species in accordance with ECMWF (2004). The aerosol IOPs are averaged over the entire SW spectrum using spectral weightings calculated, using the libRadtran/DISORT software package (Mayer & Kylling 2005; Stamnes et al., 1988), at a height of 2 km and a solar

zenith angle of 45 degrees for a standard mid-latitude summer atmosphere (Anderson et al., 1986). These IOPs are referred to as broadband IOPs hereafter in the paper. Hlradia uses the same vertical distributions of aerosols as the IFS scheme.

2.2.3 Acraneb2

The acraneb2 scheme (Mašek et al., 2015), which is more complex than hlradia but simpler than IFS, was developed as part of the ALARO-1 suite of physics parametrizations. Similar to hlradia, it is a broadband scheme using a single SW radiation interval. However, it uses the Ritter and Geleyn (1992) delta two-stream system for the clear-sky radiative transfer calculations with coefficients computed according to Räisänen (2002) i.e. by averaging the coefficients of all of the radiatively active species, weighted by their optical thicknesses. Acraneb2 uses the same climatologies of ozone and fixed composition mixture of CO₂, N₂O, CH₄ and O₂ as IFS. It also uses the same aerosol climatology as IFS but where the IOPs are spectrally averaged over the six IFS bands as is done for hlradia. One of the strengths of acraneb2 is that it possesses selective intermittency where slowly-varying gaseous transmissions are updated on a time-scale of hours (using a simple correction to the actual sun elevation).

3 Input and validation data

3.1 Aerosol climatology

The direct SW radiative effect of aerosols in MUSC is calculated using vertically integrated AOD550 and the following aerosol IOPs: AOD spectral scaling coefficients, spectral SSA and g . The spectral scaling coefficients are particularly important in the wildfire case study experiments because of the high dependence of the AOD of biomass burning aerosols on wavelength. The indirect radiative effect of aerosols is not included in the current version of the ALADIN-HIRLAM NWP system.

Monthly climatologies of Tegen et al. (1997) vertically integrated AOD550 for six aerosol categories (see Section 2.2.1) are used by default in MUSC. The aerosol IOPs for each spectral band and aerosol type are parametrized following Hess et al. (1998). The default aerosol types in MUSC are translated to GADS aerosol species before being used by hlradia

as outlined in Section 2.2.2. Spectrally averaged IOPs are used in both hlradia and acraneb2 as outlined in Sections 2.2.2 and 2.2.3.

Each radiation scheme uses Tanre et al. (1984) climatological vertical profiles to distribute the AODs for each aerosol type on model levels. In these schemes the surface-normalized vertical distribution of AOD, $\alpha(z)$, is described using the following exponential form which results in a decrease in aerosol attenuation with height

$$\alpha(z) = \exp(-z/h)$$

where z is the height above the ground and h is a vertical scale height.

It is also possible to replace the monthly Tegen climatology available in MUSC with other datasets such as the Max-Planck-Institute Aerosol Climatology version 1 (MACv1, Kinne et al., 2013) or the MACC reanalysis (Inness et al., 2013) dataset, which includes assimilated AOD measurements. For comparison, the MACC and Tegen aerosol datasets for August 2010 are shown in Figure 1 (see Section 5.1 for further details).

3.2 Aerosol observations

In the Russian wildfire case study (see Section 4.1 for details) we ran some of the simulations using AOD and IOPs derived from CIMEL sun/sky radiometer measurements recorded at the Aerosol Robotic Network (AERONET, Holben et al., 1998) station in Tõravere (58.3°N; 26.5°E), Estonia. Quality controlled level 2 AERONET data (Smirnov et al., 2000) were used. The AOD was derived (by AERONET) from measurements of the direct SW radiation flux at various wavelengths (Holben et al., 1998). SSA and g were calculated from diffuse irradiance measurements using an inversion algorithm by Dubovik and King (2000).

AOD550 and AOD scaling coefficients for the six IFS SW bands, assumed valid for the land (soot) aerosol type, were derived from the spectral AERONET measurements. These measurements range from 340 to 1020 nm whereas the IFS radiation scheme includes wavelengths from 185 to 4000 nm in the SW. Therefore, the first and sixth SW bands in the IFS scheme (1.19–2.38 μm and 2.38–4.00 μm) were extrapolated from the AERONET measurements. This may result in an error in these bands but the majority of the SW flux is contained in the remaining bands. As in the case of climatological aerosols, spectral

averages of the IOPs are derived for use in the hlradia and acraneb2 schemes.

3.3 BSRN radiative flux measurements

Global SWD radiation measurements recorded at Tõravere were compared to simulated fluxes for the August 2010 wildfire case study (see Section 4.1). These measurements are independent of the AERONET network but are part of the Baseline Surface Radiation Network (BSRN, Kallis, 2010) measurements described by Ohmura et al. (1998).

3.4 Atmospheric and surface input for MUSC

The input atmospheric and surface fields for the severe wildfire experiments at Tõravere were generated from hourly output snapshots from a 3D HARMONIE-AROME simulation. The simulation was carried out on a 2.5 km grid over Estonia for 08 August 2010 as described in Toll et al. (2015a) and the outputs were interpolated to the geographical coordinates of Tõravere for use by MUSC. As the experiments in this paper were run assuming clear-sky conditions, model level cloud water and cloud ice values were manually removed from each of the hourly atmospheric profile files generated for MUSC. These values were small but needed to be removed to allow direct comparison with observations.

4 Experiments

Three sets of experiments were conducted in this study (short names for each experiment have been included in brackets): 1) a case study of the summer 2010 Russian wildfires where smoke plumes affected Estonia (WFEXP) and the global SWD irradiance from MUSC was compared to observations, 2) the sensitivity of SWD fluxes to AOD (AODEXP), the aerosol vertical profile (VPEXP) and relative humidity (RHEXP) and 3) aerosol radiative transfer (transmittances) compared to the accurate DISORT scheme (RTEXP). 2) and 3) are sensitivity experiments and do not simulate the summer 2010 wildfires. A summary of these experiments in terms of the aerosols and radiation schemes used is given in Table 2.

4.1 Russian wildfire case study (WFEXP)

One of the worst cases of atmospheric pollution over Estonia in recent decades (Witte et al., 2011; Huijnen et al., 2012) occurred on 08 August 2010 when forest fires in the Baltic region coincided with severe thunderstorms (Toll and Männik, 2015). To study this extreme pollution event, we focussed MUSC single-column experiments on the Tõravere location in Estonia. This location was selected for three reasons: 1) the smoke plume had a strong impact on the area, 2) measurements of aerosol IOPs were available from a local AERONET station and 3) radiation flux measurements were available from the BSRN archive. We ran a series of 12 experiments using MUSC; 4 aerosol scenarios for each of the three radiation schemes (see Table 2 for summary). In particular, the following aerosol treatments were considered: 1) aerosol-free, 2) climatological AOD550 and parametrized IOPs, 3) observed AOD550 and parametrized IOPs and 4) aerosol observations (AOD550 and IOPs). In the experiments using observations (either AOD550 or both AOD550 and IOPs) the aerosols were assigned to the land/continental aerosol category while the remaining 5 categories of IFS aerosols (see Section 2.2.1) were set to zero. Accordingly, the climatological vertical distribution of IFS land aerosols was assumed.

In each experiment, a single time-step diagnostic MUSC simulation was run using the relevant input file (see Section 3.4) as the starting point and repeated for each hour between 00 UTC and 24 UTC. Thus, a series of single time-step simulations were run starting from the 00 UTC input file, 01 UTC input file and so on up to 24 UTC. The model was run in diagnostic mode in order to focus on the radiative properties when the state of the atmosphere and surface had not yet evolved from the initial values.

4.2 Aerosol sensitivity experiments (AODEXP, RHEXP and VPEXP)

In the aerosol sensitivity experiments outlined below, the 10 UTC atmospheric and surface files generated for the wildfire case study were used as input. In each of the experiments the relative effect of a different aerosol characteristic (AOD550, relative humidity and the vertical distribution of aerosols) on SWD fluxes was investigated. In each case single time-step diagnostic MUSC simulations were conducted for a range of values of each aerosol characteristic (see Table 2 for the summary).

Six experiments were carried out to investigate the effect of AOD550 on SWD fluxes

(AODEXP). In particular, two aerosol IOP configurations (observed and parametrized) were used with the IFS, hlradia and acraneb2 radiation schemes. In each case we varied AOD550 from 0 (no aerosols) to 5 (extremely polluted) in steps of 0.1 to investigate its influence on SW radiation fluxes at the surface.

The aerosol radiative transfer algorithms in the IFS and acraneb2 radiation schemes in the current version of the ALADIN-HIRLAM system assume a constant RH of 80%. In this regard, the hlradia scheme is more advanced as the RH dependence has been incorporated into the calculation of the radiative effect of aerosol IOPs. Four RH experiments (RHEXP) were carried out: 2 using hlradia and 2 using IFS where the latter were used to normalise the results from hlradia. As in AODEXP, the 10 UTC atmospheric and surface input files were used for the RHEXP experiments. Parametrized aerosol IOPs were employed in each case. Using hlradia, a series of single time-step diagnostic MUSC experiments were run for RH in the range 0 to 1.0 in increments of 0.1. The input atmospheric file was not edited to achieve the required RH. Instead, we hard-coded RH only for the aerosol transmission calculations. The series of RH simulations were run for an AOD550 of 0.1 and 1.0, which covers average and extreme aerosol quantities. Using IFS, it was only necessary to run one diagnostic MUSC simulation for each AOD550 because the IFS aerosol calculations were formulated using an assumed RH of 80%.

In the VP EXP experiments we tested the sensitivity of SWD fluxes and the SW heating rate to the vertical scale height h (see Section 3.1) using MUSC with the IFS radiation scheme. The experiments were initialised using the 10 UTC input files from the wildfire experiment and observed AOD550 and IOPs assigned to the land aerosol category. In MUSC, h has a default value of 1000 m for land aerosols. We ran single time-step diagnostic experiments using the following values of h for land aerosols: 527 m, 1000 m (default), 2109 m and 8343 m. For smaller h the aerosols are concentrated closer to the ground while larger values of h spread the aerosols higher into the atmosphere. The acraneb2 and hlradia schemes use the same vertical distribution of aerosols as IFS and exhibit a similar sensitivity in terms of SWD fluxes and the SW heating rate (results not included).

4.3 Aerosol radiative transfer (RTEXP)

Accurate aerosol radiative transfer is of equal importance to accurate aerosol IOPs. To examine the performance of the aerosol radiative transfer algorithms in MUSC, we

1 extracted the relevant subroutines from the IFS, hlradia and acraneb2 radiation scheme
2 codes and ran these as stand-alone formulations. These calculations require optical
3 thickness, SSA, g and the cosine of the solar zenith angle as input.

4 We ran experiments using the IFS Fouquart and Bonnel (1980) clear sky formulation, the
5 two-stream approximation (Thomas & Stamnes, 2002) used in hlradia and the acraneb2
6 Ritter and Geleyn (1992) two-stream approximation to calculate SW transmission through a
7 homogeneous atmospheric layer with optical properties resembling those of aerosols. In
8 particular, we used $SSA = 0.95$, $g = 0.7$ and cosine of the solar zenith angle of 0.6 while
9 varying the optical depth between 0.1 and 5. Additionally, we used the accurate DISORT
10 radiative transfer scheme (Stamnes et al., 1988) with 30 streams and the same input as the
11 IFS, hlradia and acraneb2 radiative transfer calculations.

13 **5 Results and Discussion**

14 **5.1 Russian wildfire case study (WFEXP)**

15 The results presented in this section include a comparison of AOD550 for the Tegen and
16 MACC reanalysis climatologies, time-series of spectral AOD, SSA and g from AERONET
17 and experiments run using MUSC with observed and climatological aerosol data and the
18 IFS, hlradia and acraneb2 radiation schemes. Figure 1 shows the AOD550 over
19 northwestern Europe on 08 August 2010 in the Tegen climatology used in MUSC and the
20 MACC reanalysis dataset (Inness et al., 2013). It is clear that the Tegen climatology greatly
21 underestimates aerosols when pollution is heavy, as was the case over Estonia and eastern
22 Russia on 08 August 2010. Overall, over northwestern Europe the values of the realistic
23 MACC AOD550 (maximum 3.5) are an order of magnitude higher than in the Tegen
24 climatology (maximum 0.33) for August which highlights a drawback of using the Tegen
25 dataset.

26 Figure 2 shows a time-series of AOD at Tõravere on 08 August 2010 for 7 wavelengths
27 (measurements from the AERONET archive). The strong spectral dependence of AOD is
28 clear from the figure; AOD is higher for shorter wavelengths. This notable wavelength
29 dependence is characteristic of biomass burning aerosols (Eck et al., 1999). The AOD550,

also shown in Figure 2 (*black dashed line*), and used in the experiments involving observations rather than the Tegen climatology, was calculated using the AERONET AOD at 500 nm (*cyan line*) and the Ångström exponent in the 440-675 nm spectral interval. For comparison, the significantly lower AOD550 from the Tegen climatology (*red dashed line*) is also included in the figure.

The remaining aerosol IOPs, SSA and g , from the AERONET inversion products database are shown in Figure 3 where daily averages are plotted as a function of wavelength. Although daily averages are shown, the time dependence of SSA and g on 08 August was small. The asymmetry factor, g , varies from 0.56 to 0.7 across the wavelength range and has an average value of 0.634. The latter was used in the wildfire experiments run using hlradia or acraneb2 with aerosol observations. The spectral values of g were interpolated to the six SW bands of IFS for experiments using this scheme.

The aerosol scattering per extinction ratio, represented by SSA, is high (close to 0.96 with a spectral average of 0.955) at each wavelength with little SW spectral dependence (Figure 3). This is similar to results by Dubovik et al. (2002) who showed that the typical SSA of smoke from biomass burning in Boreal forests is high. However, the scattering of smoke particles from this Russian wildfire event was higher than that of plumes from typical biomass burning in Boreal forests (Chubarova et al., 2012). As in the case of g , the average SSA was used in the hlradia and acraneb2 wildfire experiments involving aerosol observations while the spectral SSAs were interpolated to the IFS SW bands before use in the corresponding IFS experiments.

Figure 4a shows the global SWD radiative flux at the Earth's surface simulated using MUSC with the IFS radiation scheme for 08 August 2010 at Tõravere. We ran an experiment for each of the following 4 aerosol scenarios (also summarised in Table 2): 1) aerosol-free (*red curve*), 2) climatological AOD550 and parametrized IOPs (*black curve*), 3) observed AOD550 and parametrized IOPs (*green curve*) and 4) observed AOD550 and IOPs (*cyan curve*) and compared the global SWD fluxes to BSRN observations (*blue curve*). The discrepancy between simulated and observed SWD irradiance after 14 UTC is due to the development of convective clouds (Toll et al., 2015a) which are not accounted for in the MUSC clear-sky simulations.

The biases in global SWD flux (relative to observations) for the experiments using each

1 radiation scheme (and not just IFS) and the 4 different aerosol scenarios are depicted in
2 Figure 4b (*IFS dotted continuous lines, hlradia continuous lines, acraneb2 dashed lines*; the
3 aerosol scenario colour scheme is the same as in Figure 4a). Overall, the results for the three
4 schemes are similar (mostly to within 10-20 W/m² of each other for global SWD irradiance
5 which can be seen by comparing each group of three curves of the same colour),
6 particularly in their response to the different aerosol scenarios. The largest discrepancies
7 occur in the early morning; in hlradia this occurs because the sphericity of the atmosphere is
8 not taken into account. When the direct radiative effect of aerosols was excluded, global
9 SWD fluxes were overestimated by ~120 W/m² or 19 % (*red curves*) at midday compared
10 to BSRN observations. Accounting for the climatological average effect of aerosols using
11 the Tegen et al. (1997) dataset and Hess et al. (1998) IOP parametrizations (*black curves*)
12 improves the simulation of global SWD flux compared to the aerosol-free simulation.
13 However, there is still an overestimation of 60 W/m² or 10 % at noon, because the observed
14 AOD is higher than the climatological average (see Figures 1 and 2).

15 The use of AOD550 and IOPs derived from AERONET observations gives very good
16 agreement between the modelled and observed global SWD fluxes for each of the three
17 radiation schemes (*cyan curves*, bias < 20 W/m² or 4 % at noon). SWD flux was
18 underestimated by ~70 W/m² or 11 % (*green curves*) at noon when the direct radiative effect
19 of aerosols was accounted for using the observed AOD550 combined with parametrized
20 SSA, *g* and spectral scaling factors of land aerosols. This underestimation can be explained
21 by two factors.

22 Firstly, using the climatological AOD scaling factors for Tõravere in August leads to AOD
23 values which are 60% higher than those estimated from the AERONET spectral
24 measurements. Secondly, using a delta-Eddington optical depth scaling factor ($1-SSA \cdot g^2$;
25 Joseph et al. 1976) based on the climatological SSA and *g* values causes this factor to be
26 approximately 7% lower than the corresponding scaling factor based on the observed SSA
27 and *g* values. The combination of these two factors results in a scaled SW AOD of land
28 aerosols that is 48% larger than the observation-based value.

5.2 Aerosol sensitivity tests

5.2.1 AOD (AODEXP)

The sensitivity of global SWD fluxes to AOD550 is shown in Figure 5a for MUSC experiments run using the IFS, hlradia and acraneb2 radiation schemes. The results depicted by the cyan curves are for the case where observed IOPs (i.e. 10 UTC observation at Tõravere from the AERONET archive) were used in the simulations. Parametrized IOPs were used where the experiment results shown in green. For an AOD550 of 1, global SWD irradiance is $\sim 100 \text{ W/m}^2$ lower when parametrized rather than observed IOPs were used, regardless of the radiation scheme. The reason for this difference has been discussed in Section 5.1.

The effect of AOD550 on direct SWD flux is shown in Figure 5b for the three radiation schemes and observed (*cyan curves*) and parametrized (*green curves*) IOPs. For example, when parametrized IOPs were used an increase in AOD550 from 0 to 1 (i.e. no aerosols to heavy pollution) reduced the global SWD irradiance by $\sim 200 \text{ W/m}^2$ or 27 % (Figure 5a) but had a greater effect on the direct SWD flux ($\sim 330 \text{ W/m}^2$ or 49 %, Figure 5b). Direct SWD flux is very sensitive to AOD because it is extinguished by both absorption and scattering. Due to forward scattering, the global SWD flux is less affected as the forward scattered irradiance reaches the surface as diffuse SWD irradiance. The influence of AOD550 on global SWD flux is similar for each radiation scheme. Global SWD fluxes are lower when the parametrized IOPs are used; this is because of the combination of a higher AOD scaling factor and lower delta-Eddington optical depth scaling factor than based on observations as discussed in Section 5.1. This sensitivity test clearly illustrates the large effect AOD550 and the IOPs have on SW radiative fluxes at the surface and emphasizes the importance of using the correct aerosol IOPs in NWP.

5.2.2 Relative humidity (RHEXP)

The impact of relative humidity (RH), accounted for in the aerosol radiative transfer calculations in the hlradia scheme, on global and direct SWD fluxes is shown in Figure 6a. RH was varied from 0 to 1 in steps of 0.1; the AOD550 was set to 0.1 in the grey curves and to 1.0 (significant pollution) in the black curves. IFS land aerosol parametrized IOPs at Tõravere were used. Increasing RH from 0 to 1 increases global SWD flux by 1.5 % when

AOD550 = 0.1 (*grey line with filled circles*) and by 12 % when AOD550 = 1.0 (*black line with filled circles*). The effect on the corresponding direct SWD fluxes is greater as expected (increases of 2.5 % and 24 % for AOD550 = 0.1 and 1.0 respectively, *grey and black continuous lines*). Figure 6b shows global SWD irradiance from the experiments run using hlradia, an AOD550 of 0.1 (*grey*) and 1.0 (*black*) and RH varying from 0 to 1.0 as before. In this case, the global SWD fluxes are normalised using global SWD flux from a corresponding experiment run using the IFS radiation scheme. In the IFS scheme a constant RH of 80 % is assumed for “land” aerosols. For AODs close to the climatological value (i.e. 0.1 here) the relative differences between global SWD fluxes for experiments using hlradia and IFS are small, and negligible at a RH of 0.8. As the humidity deviates from 0.8, particularly for larger AODs, the differences between the global SWD fluxes from experiments using hlradia and IFS grow to ± 6 %. When the AODs are close to the climatological average (of the order of 0.1), the influence of RH on aerosol radiative transfer is less than 1%. In such cases, the assumption of a constant RH by the IFS and acraneb2 schemes is acceptable and is not a major source of error. On the other hand, for cases where pollution is high, the influence of RH on global SWD flux is $\sim \pm 6$ % and could be important, particularly for solar energy applications.

5.2.3 Vertical distribution of aerosols (VPEXP)

The inherent uncertainties associated with assuming fixed vertical profiles were investigated by modifying the assumptions about the shape of the IFS climatological vertical aerosol profile. Figure 7a shows normalised net SW fluxes on pressure levels for MUSC experiments run with the IFS radiation scheme and vertical scale heights (h , of land aerosols) of 527 m (*red*), 1000 m (*default, green*), 2109 m (*cyan*) and 8343 m (*blue*). AERONET aerosol observations at 10 UTC over Tõravere were used as input. The net SW is normalised relative to the aerosol-free case and varies by no more than 4% at each pressure level for the range of h plotted. For example, at 1000 hPa net SW varies by less than 3% with h , and by less than 4% at 800 hPa.

Figure 7b shows the SW heating rates for the same experimental set-up as in Figure 7a, where the heating rate is normalised relative to the corresponding aerosol-free simulations. The heating rates in the boundary layer changed by up to a factor of 2 in response to the aerosol vertical distribution (e.g. when h is halved (*red curve*) compared to the default

(*green curve*) the SW heating rate approximately doubles). Figure 7b also clearly illustrates that large values of h spread the aerosol higher in the atmosphere (*blue curve*) while for smaller values, the aerosols are more concentrated closer to the ground (*red curve*).

5.3 Aerosol radiative transfer (RTEXP)

Figure 8a shows transmission as a function of optical depth through a homogeneous atmospheric layer containing aerosol ($SSA = 0.95$, $g = 0.7$, cosine of the solar zenith angle set to 0.6). Transmittances calculated with the IFS Fouquart and Bonnel (1980) clear-sky radiative transfer formulation, the Thomas & Stamnes (2002) two-stream approximation used for aerosol transmittance in *hrradia* and the Ritter and Geleyn (1992) two-stream approximation used in *acraneb2* are compared to the transmittance from the accurate DISORT radiative transfer scheme (Stamnes et al., 1988) run using 30 streams. Figure 8b shows the relative differences in transmittance between these radiative transfer schemes and DISORT. As can be seen from Figures 8a and 8b the IFS, *hrradia* and *acraneb2* radiative transfer approximations give transmittances that are within a few percent of the accurate 30-stream DISORT calculations for optical thicknesses less than 1. Even for optical thicknesses of up to 5, the approximations remain within ± 13 % of the DISORT results.

6 Conclusions and future work

We carried out single column diagnostic experiments using the MUSC model and three radiation schemes (*IFS*, *hrradia* and *acraneb2*) to examine the influence of the direct radiative effects of aerosols on SW radiative flux. In particular, we focused on the effect that AOD550, aerosol IOPs, the relative humidity, vertical profile of AOD and the radiative transfer formulations has on SW fluxes.

In the wildfire case study we showed that the bias in modelled global SWD flux relative to observations was lowest when observed AOD550 and IOPs were included in the simulations (within ± 4 % at midday). This was true irrespective of the radiation scheme and its spectral resolution. Global SWD flux was greatly overestimated, by more than 15 % at midday, when aerosols were excluded and by +10 % at midday when the climatological aerosols were used (Tegen et al., 1997). On the other hand global SWD irradiance was underestimated by 11 % at noon, when observed AOD550 and parametrized IOPs, as

opposed to observed IOPs, were used in the experiments. This highlights the need for accurate information on both aerosol concentration and aerosol IOPs in order to improve the simulated radiation budget in the model. The importance of all of the aerosol IOPs, and not just AOD₅₅₀, in the direct radiative effect of aerosols on solar radiation was clearly demonstrated. The over- and under-estimation of global SWD flux leads to errors in model temperatures and energy fluxes. Therefore, during heavy pollution episodes the use of real-time aerosols would greatly improve the radiation budget and meteorological forecasts. The wildfire experiments also illustrate that the performance of the broadband hlradia and acraneb2 schemes is comparable to that of the spectral IFS scheme. The results attained for the three schemes were similar, with simulated global SWD fluxes mostly within 10-20 W/m² of each other for each aerosol scenario.

The dependency of the direct radiative effect of aerosols on relative humidity was up to ± 6 % for an AOD of 1.0. As a first approximation, assuming a constant relative humidity is acceptable but we suggest that relative humidity dependent parametrizations of aerosol IOPs should be used. The effect of the vertical profile of IFS land aerosols (via the vertical scale height) on net SW irradiance near the surface was found to be up to 4%. This is consistent with the finding of Meloni et al. (2015). The influence of the vertical profile on model-level SW heating rates was large, changing by up to a factor of 2 in the boundary layer in response to the aerosol vertical distribution. This highlights the need for using realistic vertical profiles of aerosols. In reality aerosols are distributed in discrete rather than continuous layers. The IFS, hlradia and acraneb2 radiative transfer approximations were tested for a range of optical depths and found to be accurate to within ± 13 % even for large aerosol loads compared to the DISORT model.

The influence of improvements in the representation of the direct radiative effect of aerosols on meteorological forecasts needs further study using 3D simulations. We plan to upgrade the aerosol climatology in the HARMONIE-AROME configuration of the ALADIN-HIRLAM system to the more realistic MACC reanalysis dataset. We will also investigate the option of acquiring real-time aerosol input, including the vertical profile of the aerosol properties, from 3D aerosol IOP estimates from the C-IFS model or chemical transport model simulations, possibly coupled to the NWP model.

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Figure 1. (a) AOD550 for August in the default climatology in the ALADIN-HIRLAM NWP system (Tegen et al., 1997) (b) AOD for 08 August 2010 from the MACC reanalysis (Inness et al., 2013). The location of Tõravere (58.3°N; 26.5°E), for which the MUSC single column experiments are run, is shown as a red dot in both panels. Note the factor of 10 difference in AOD between the Tegen climatology and the MACC reanalysis.

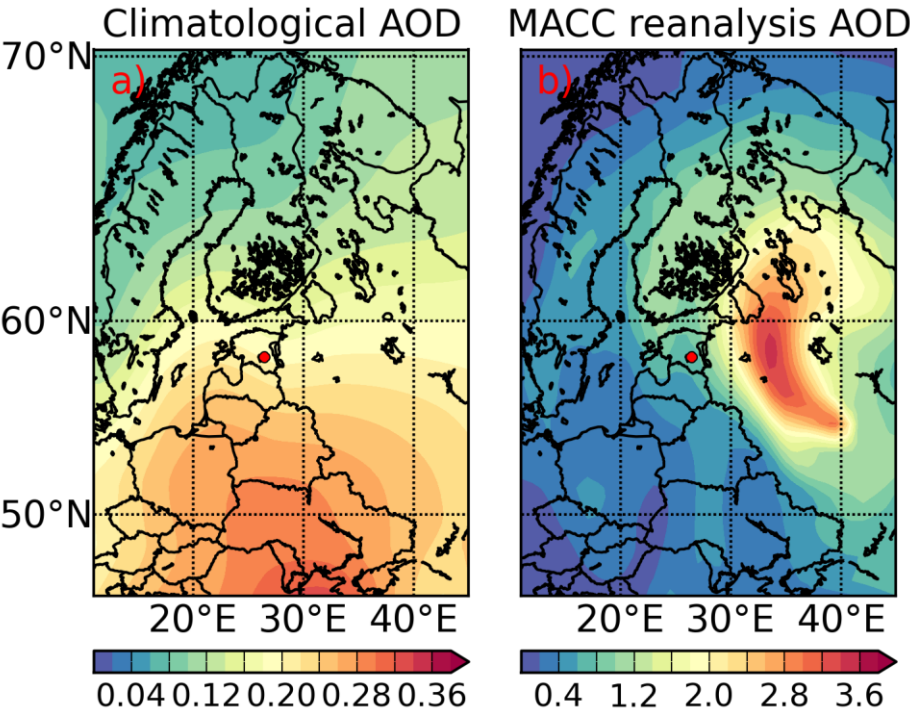


Figure 2. AERONET measurements of AOD at Tõravere on 08 August, 2010 for 7 SW wavelengths (nm). The AOD550 derived from these measurements (*black dashed line*) and the default climatological AOD550 at Tõravere (*red dashed line*) are also shown in the figure. Data are not available after 14UTC due to the presence of clouds.

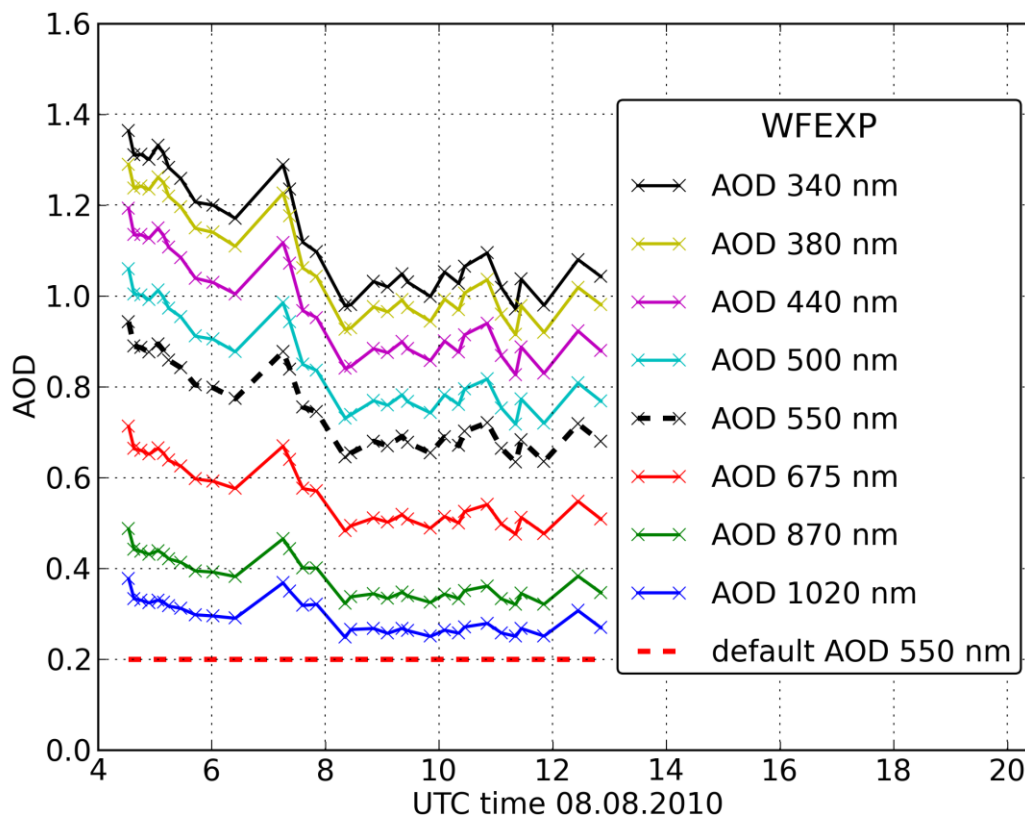


Figure 3. Single scattering albedo (SSA) and asymmetry factor (g) at Tõravere on 08 August 2010, as a function of wavelength, attained from the AERONET inversion products database. The data are averaged over the day – data at three different times were available and the time dependence was small.

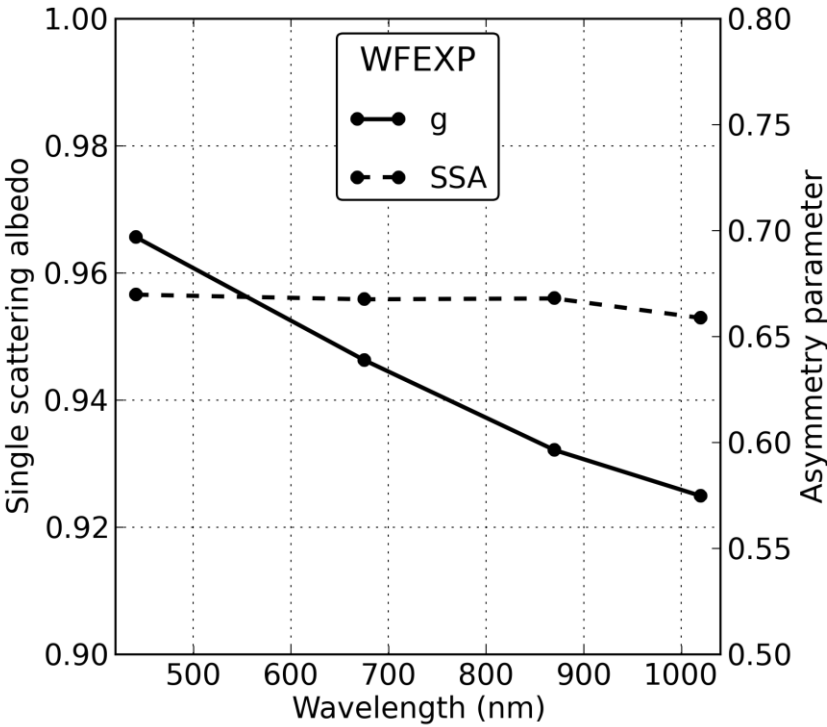
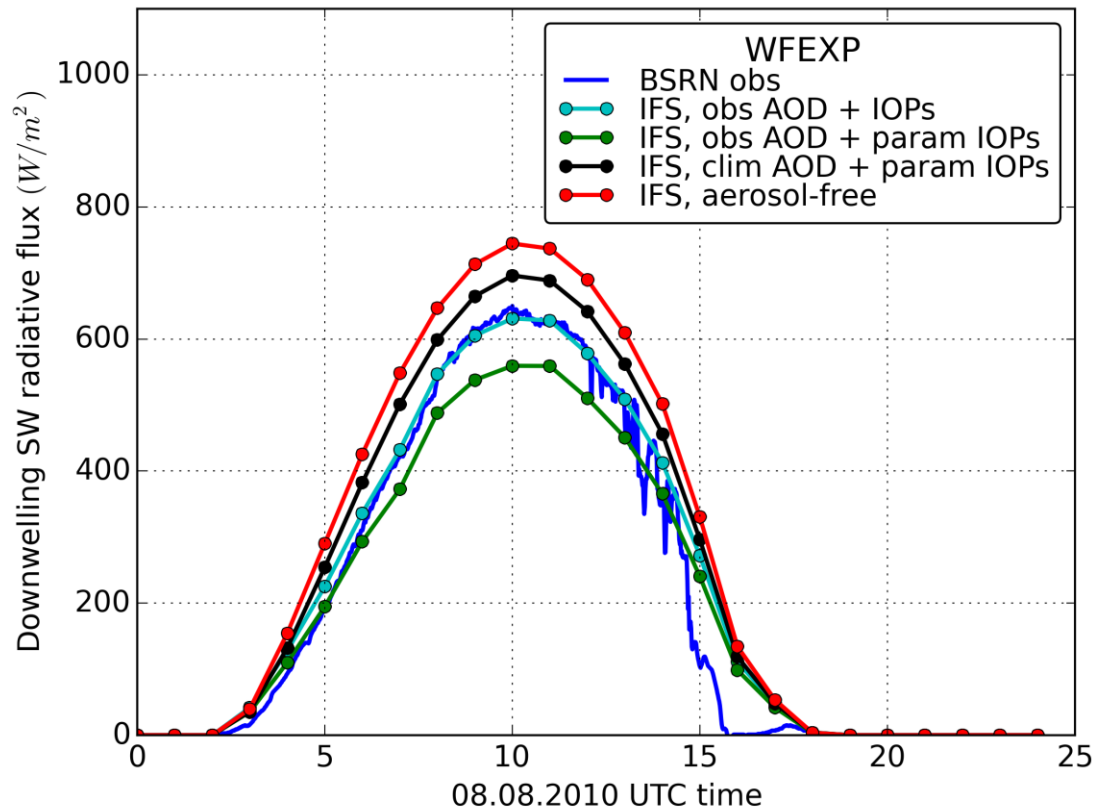
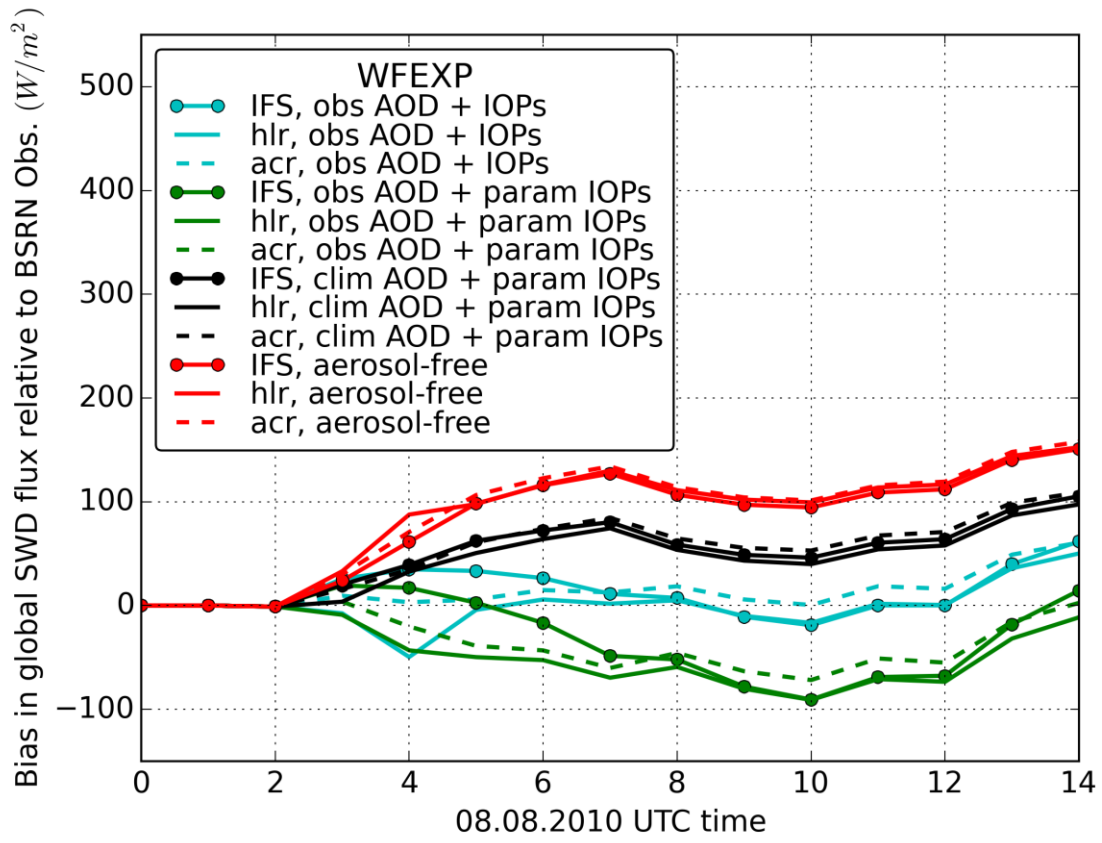


Figure 4. (a) Time-series of global SWD radiation flux (W/m^2) at Tõravere on 08 August 2010 simulated using MUSC (with the IFS radiation scheme) for four aerosol scenarios (*red*: aerosol free; *black*: climatological AOD550 and parametrized IOPs; *green*: observed AOD550 and parametrized IOPs; *cyan*: observed AOD550 and IOPs). BSRN global SWD flux measurements are shown in *blue* (b) Time-series of the bias in global SWD flux relative to BSRN observations for the same four aerosol scenarios as in (a). The results when the IFS radiation scheme was used are depicted by a dotted continuous line; the hlradia (hlr) and acraneb2 (acr) biases are shown using continuous and dashed lines respectively. Only data up to 14 UTC are shown because clouds developed after that, which were not included in the MUSC simulations.



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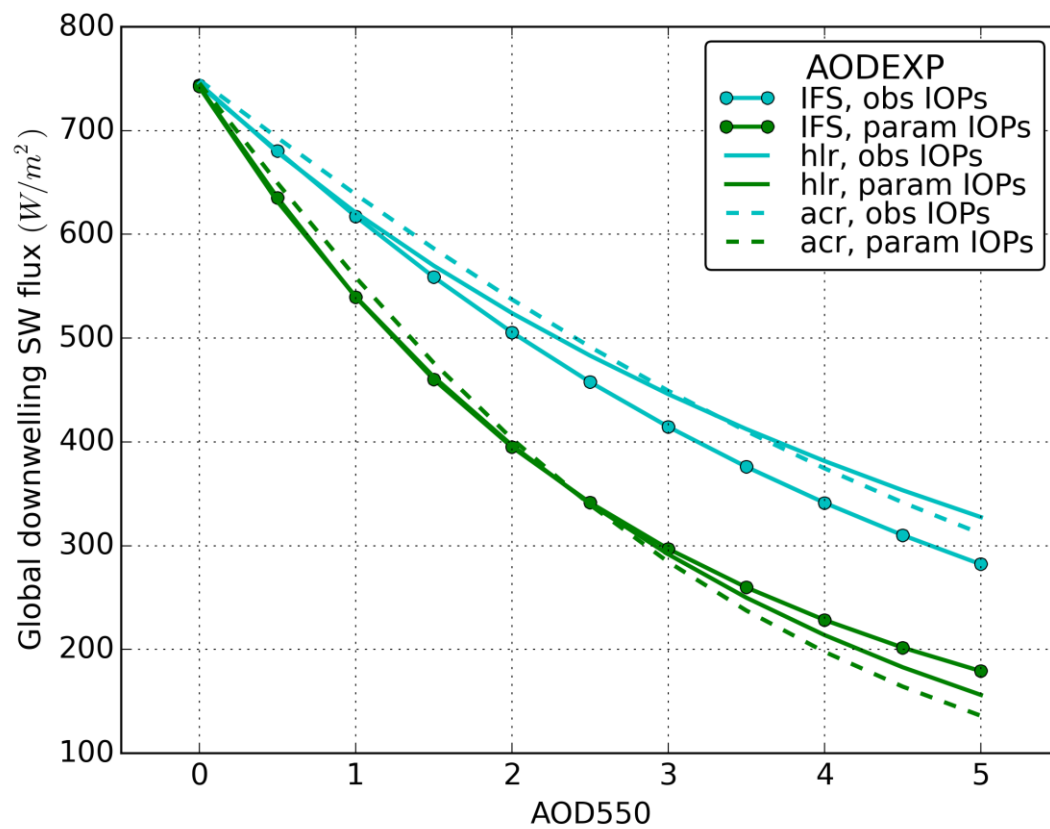
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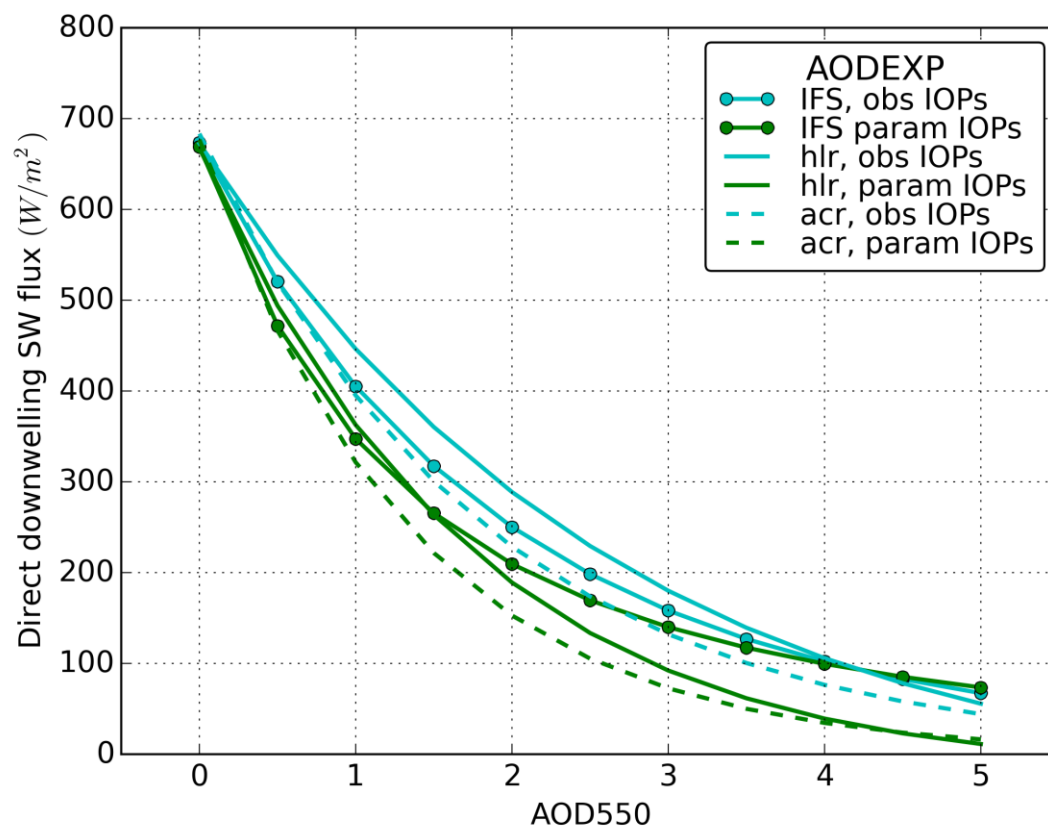
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Figure 5. Global SWD flux as a function of AOD550 for MUSC experiments run using the IFS (*dotted continuous curve*), hlradia (*continuous curve*) and acraneb2 (*dashed curve*) radiation schemes. The results depicted by the *cyan curves* are for the cases where observed IOPs (SSA, g and AOD scaling) were used. Parametrized IOPs were used for those shown in green. Direct SWD flux for MUSC experiments using the same aerosol scenarios and radiation schemes as in (a) is shown in (b).

(a)



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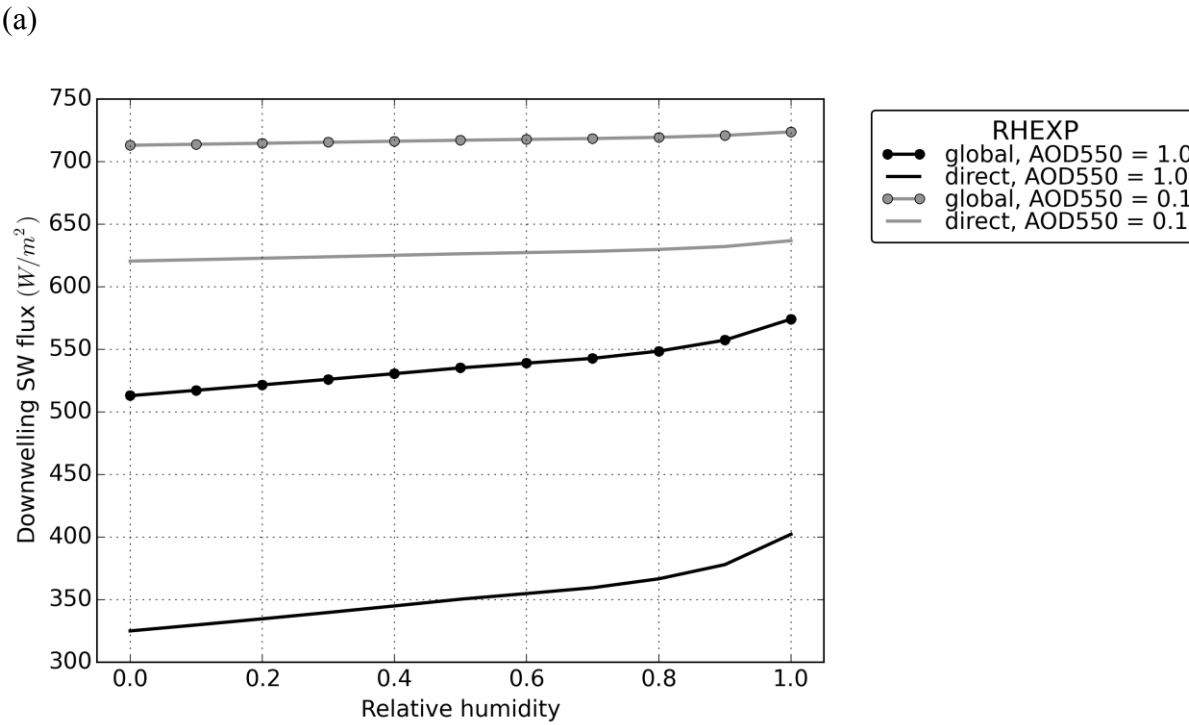
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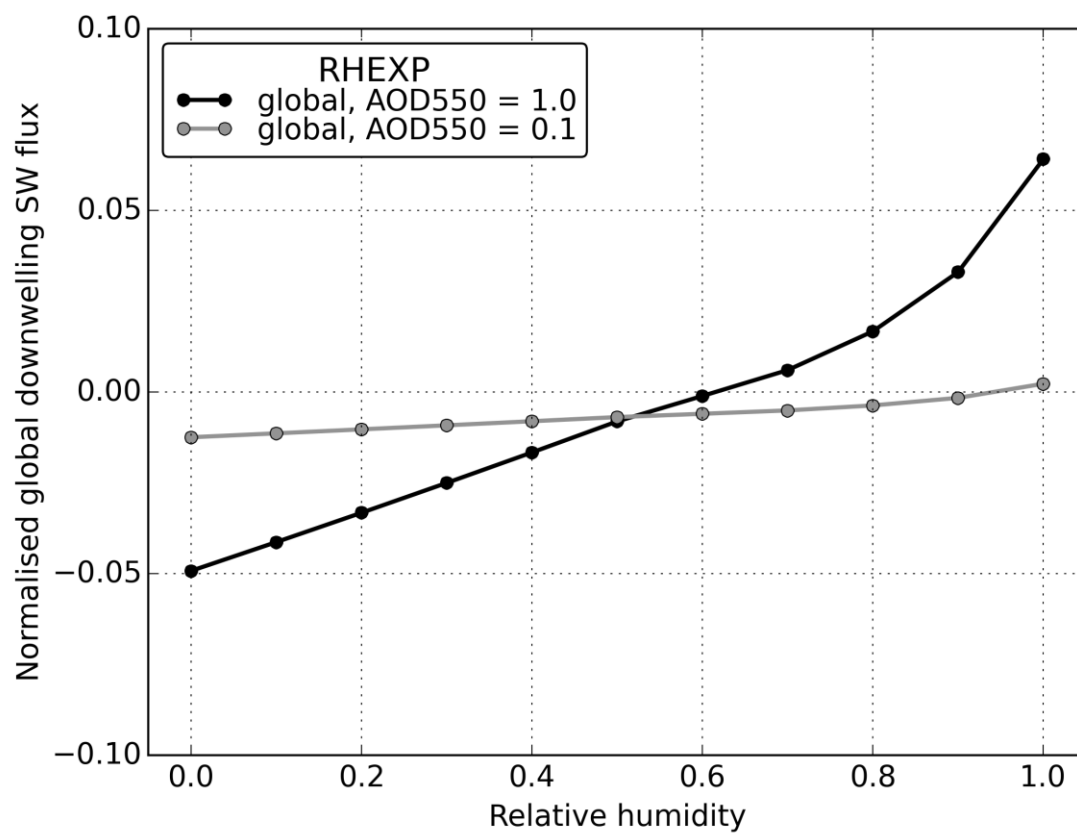
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Figure 6. (a) Global (*dotted continuous curves*) and direct (*continuous curves*) SWD fluxes, as a function of relative humidity, simulated using MUSC with the hlradia radiation scheme for AOD550 values of 0.1 (*grey*) and 1.0 (*black*) (b) similar to (a) but shows global SWD flux normalised relative to a corresponding experiment run using the IFS radiation scheme. The IFS scheme assumes a RH of 80% for land aerosols.



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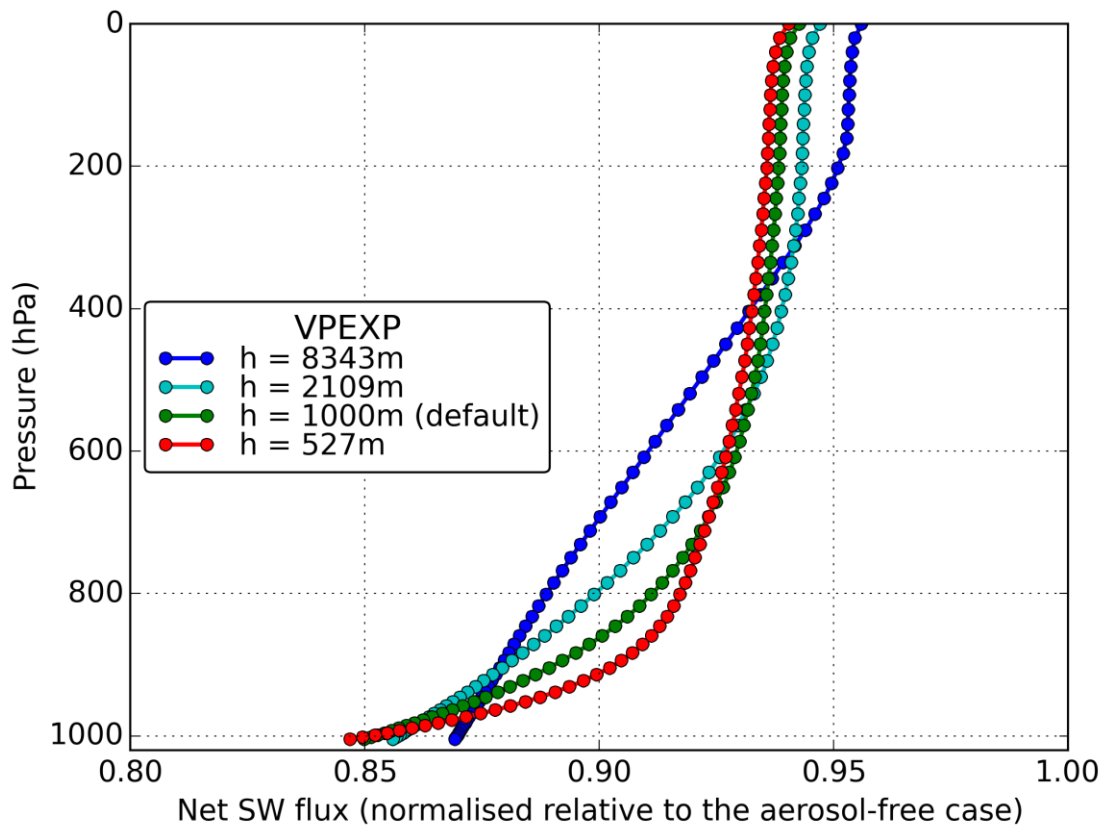
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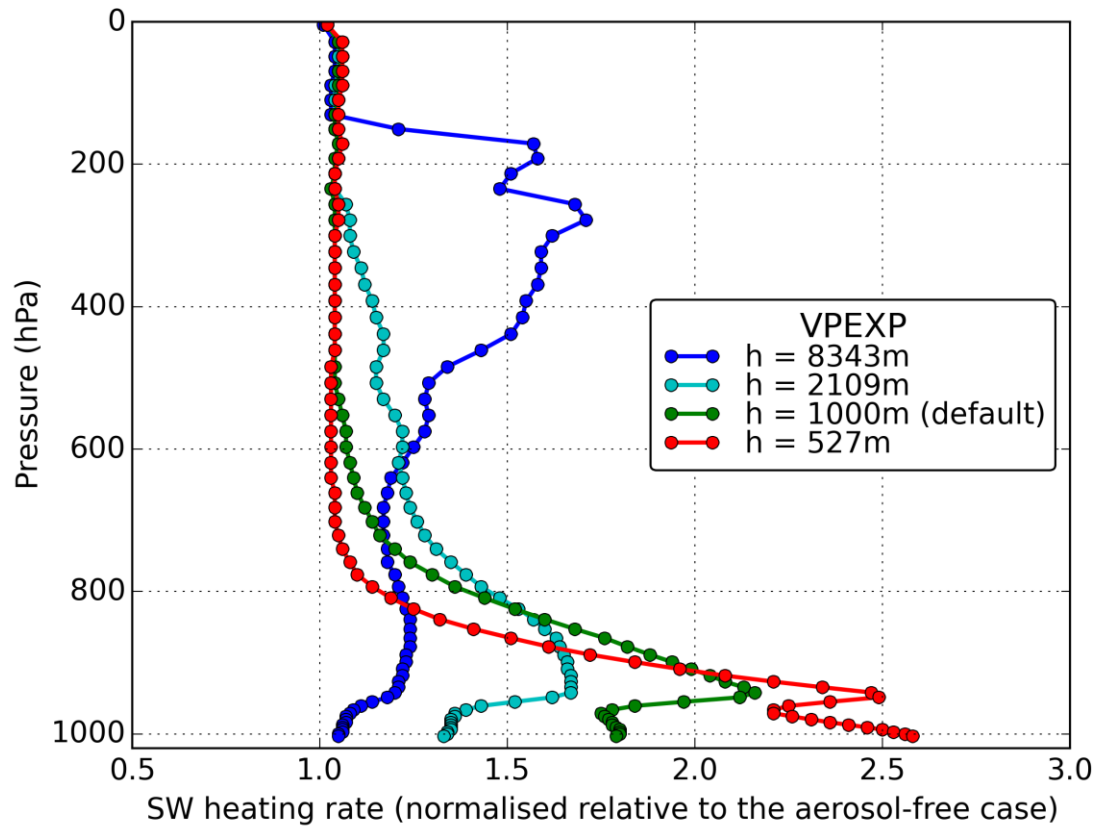
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Figure 7. (a) Net SW radiation fluxes (normalised relative to the aerosol-free case) as a function of atmospheric pressure for four vertical scale heights (h) of IFS land aerosols (b) similar to (a) but shows the normalised SW heating rate. The experiments were carried out using the IFS radiation scheme.

(a)



1 (b)



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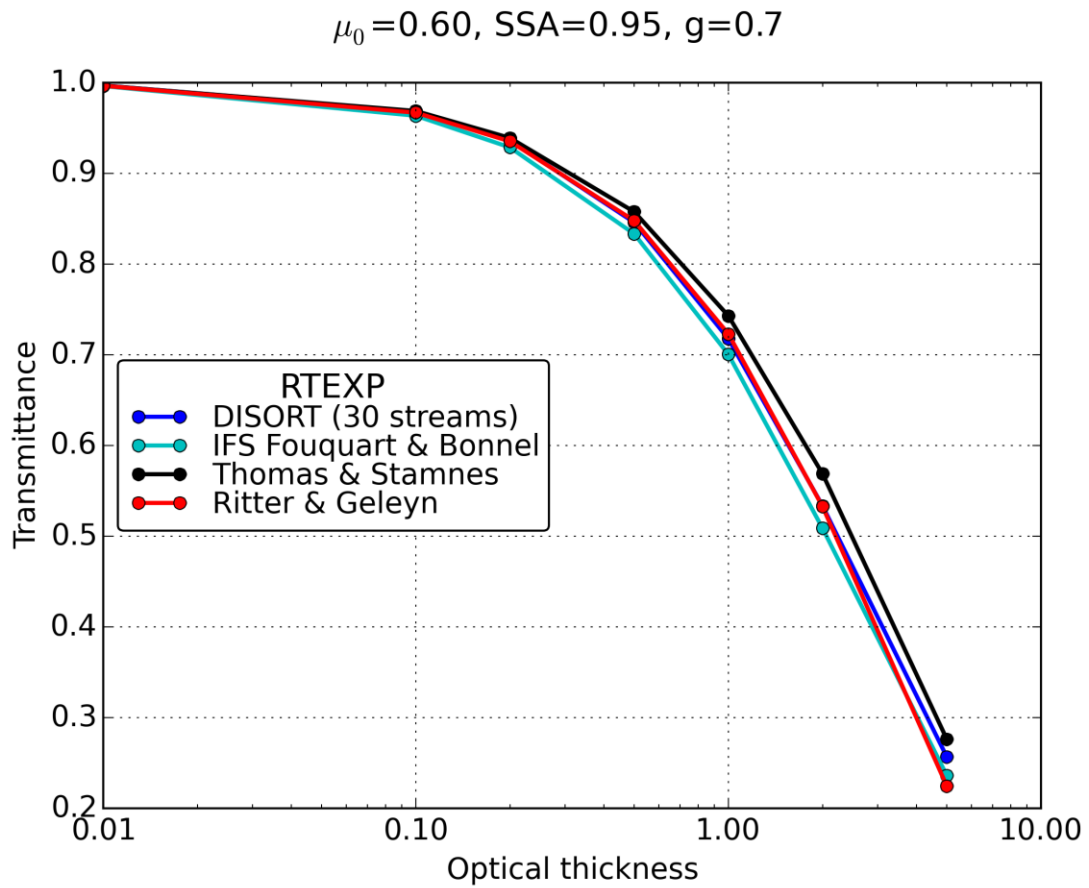
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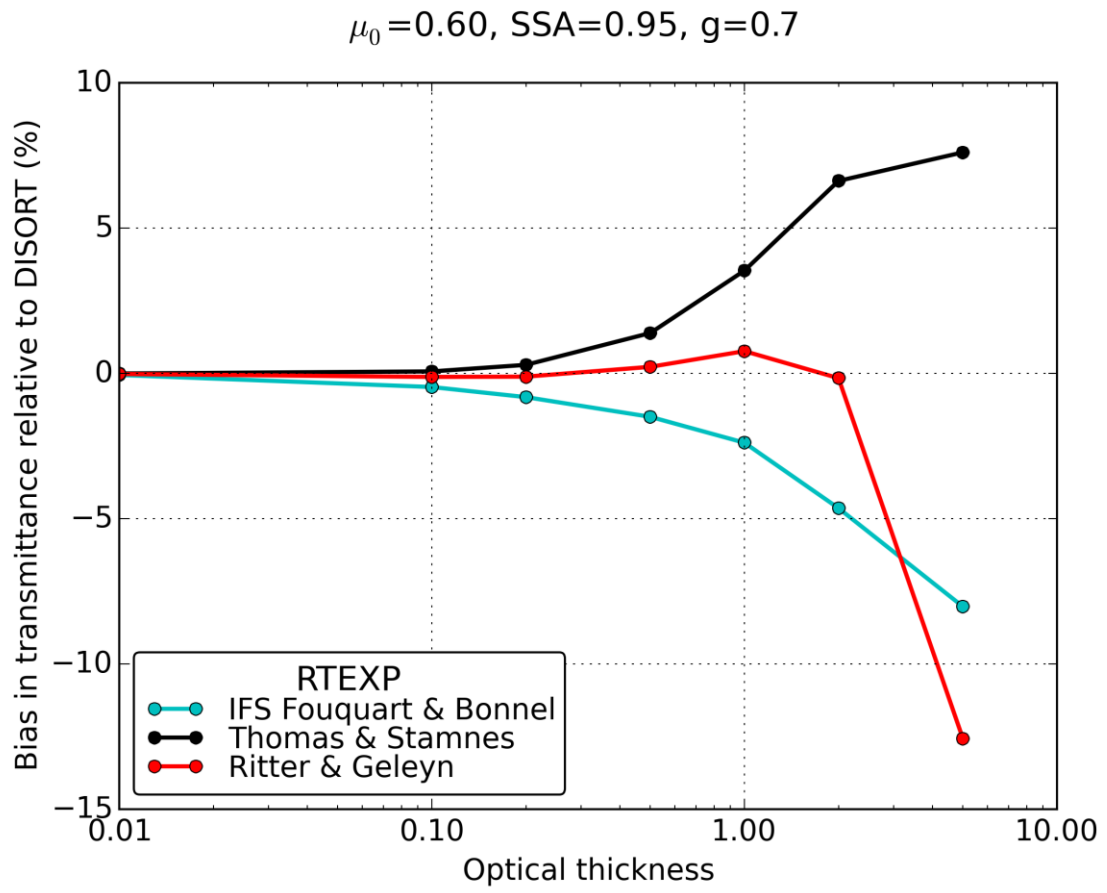
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Figure 8. (a) Transmission as a function of optical depth through a homogeneous atmospheric layer containing aerosol ($SSA = 0.95$, $g = 0.7$, cosine of the solar zenith angle set to 0.6) for the following radiative transfer algorithms: IFS Fouquart and Bonnel (1980) clear-sky radiative transfer formulation (*cyan curve*), the Thomas & Stamnes (2002) two-stream approximation used for aerosol transmittance in hlradia (*black curve*) and the Ritter and Geleyn (1992) two-stream approximation used in acraneb2 (*red curve*) and the accurate 30-stream DISORT (Stamnes et al., 1988) radiative transfer scheme (*blue curve*). (b) As for (a) but shows the transmission differences relative to the accurate DISORT scheme.



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Table 1. Summary of aerosol-radiation experiments including details of the radiation schemes and aerosol datasets used.

	IFS	hrradia	acraneb2
SW bands	6	1	1
LW bands	14	1	1
Ozone	Monthly climatology	Impact of O ₃ constant over time and space	Monthly climatology
Radiatively active gases	Fixed composition mixture of CO ₂ , N ₂ O, CH ₄ and O ₂	Impact of CO ₂ and O ₂ constant over time and space	Fixed composition mixture of CO ₂ , N ₂ O, CH ₄ and O ₂
Radiative transfer	Fouquart and Bonnel (1980) two-stream equations	Savijärvi, 1990 and the Thomas & Stamnes (2002) two stream equations	Ritter and Geleyn (1992) delta two-stream system

Table 2. Summary of aerosol-radiation experiments including details of the radiation schemes and aerosol datasets used.

Experiment	SW radiation scheme	AOD550	Other IOPs (AOD scaling, SSA, g)
Russian wildfire experiments: WFEXP	IFS	Aerosol-free	Aerosol-free
		Tegen climatology	Hess parametrizations
		Observed	Hess parametrizations
		Observed	Observed
	hrradia	Aerosol-free	Aerosol-free
		Tegen climatology	Hess parametrizations
		Observed	Hess parametrizations
		Observed	Observed
	acraneb2	Aerosol-free	Aerosol-free
		Tegen climatology	Hess parametrizations
		Observed	Hess parametrizations
		Observed	Observed
AOD550 experiments: AODEXP	IFS	Range [0,5] in steps of 0.1	Hess parametrizations
			Observed
	hrradia	Range [0,5] in steps of 0.1	Hess parametrizations
			Observed
	acraneb2	Range [0,5] in steps of 0.1	Hess parametrizations
			Observed
Vertical profile of aerosols experiments: VPEXP	IFS	Observed (10 UTC)	Observed (10 UTC)
Relative humidity experiments: RHEXP	hrradia	0.1	Hess parametrizations
		1.0	Hess parametrizations
	IFS	0.1	Hess parametrizations
		1.0	Hess parametrizations
Radiative transfer experiments: RTEXP	IFS	total AOD (not 550 nm) in the range [0.01,5]	SSA=0.95, g=0.7
	hrradia	total AOD in the range [0.01,5]	SSA=0.95, g=0.7
	acraneb2	total AOD in the range [0.01,5]	SSA=0.95, g=0.7
	DISORT	total AOD in the range [0.01,5]	SSA=0.95, g=0.7