Response to referees' comments

General: We are very much grateful to both the referees for appreciating our work and giving very helpful

suggestions and comments, which have significantly improved the manuscript. We have revised the

manuscript by carefully taking into account all the comments point by point. Text in red and blue colour

show the questions and answers, respectively.

Anonymous Referee #1

Received and published: 27 January 2016

This paper presents a year of data on CO and CO₂ concentrations from a site in Ahmedabad. High quality

concentration data from urban areas in general are sparse, and such data from the large urban areas in

rapidly developing regions are especially limited. These observations can contribute to understanding

emission patterns in a poorly studied region that is critically important to the global carbon budget. The

experimental methods are excellent and include decent calibration scheme. The text provides a good

summary of the methods and defines precision and accuracy. However, the discussion needs to be more

focused and strive to present a consistent set of key findings. As noted in detailed comments, some of

the observed variations in concentrations may not contribute to interpreting emissions patterns. The

results will be more convincing by focusing on the key aspects of the data. It is important to distinguish

between patterns with information about atmospheric dynamics (vertical mixing and transport) and

patterns that have information about emission sources.

Comments and suggestions for revised analysis.

Page: 32200

With respect to the evolution of CO₂ during night time. Even in cold regions there soils approach 0C

respiration continues throughout the night. At this site I don't think you can attribute lack of increasing

CO2 during night in some seasons to respiration being dormant. There is certainly no evidence included

in the text for this. In this site I would only expect respiration to be suppressed by very dry soils, so it could

be a reason in the spring, but temperatures are probably not cold enough to suppress respiration. You

don't show any data for night time winds. Differences in depth and strength of the nocturnal inversion and

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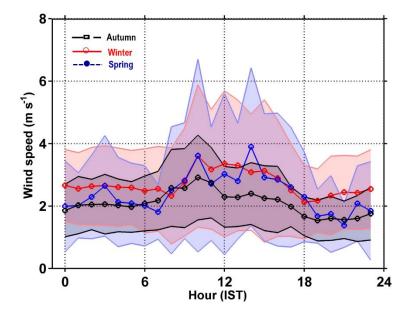
whether winds persist at night are factors that would impact whether trace gases accumulate at the surface during night. In subsequent section you show that night time concentrations of CO decline continuously in the winter and spring season, which indicates that there is enough vertical mixing of low CO air from above that once the CO source is turned off its concentration drops. Thus, the constant CO2 at night is evidence of a continued source in order to offset dilution by mixing of low CO2 air from aloft. The dynamics of CO2 is not just the depth of mixing. You can note that because there is active CO2 uptake during seasons when vegetation is active the entire mixed layer is depleted during daytime and when residual layer mixes to the surface in morning, low-CO2 air is mixed down.

Response: We are very much grateful for these wonderful explanation and suggestions. We have included above suggestions in the explanation of the main text as well as we have also revised respective sections, as per suggestions of the second referee. The revised text is given below.

Diurnal variation: CO₂

"Figure 5a shows the mean diurnal cycles of atmospheric CO₂ and associated 1-σ standard deviation (shaded region) during all the four seasons. All times are in Indian Standard Time (IST), which is 5.5 hrs ahead of the Universal Time (UT). Noticeable differences are observed in the diurnal cycle of CO2 from season to season. In general, maximum concentration has been observed during morning (0700-0800 hrs) and evening (1800-2000 hrs) hours, when the ABL is shallow, traffic is dense and vegetation respiration dominant due to absence of photosynthesis activity. The minimum of the cycles occurred in the afternoon hours (1400-1600 hrs), when the PBL is deepest and well mixed as well as when the vegetation photosynthesis is active. There are many interesting features in the period of 0000-0800 hrs. CO₂ concentrations start decreasing from 0000 to 0300 hrs and increases slightly afterwards till 0600-0700 hrs during summer and autumn. Respiration of CO₂ from the vegetation is mostly responsible for this night time increase. During winter and spring seasons CO₂ levels are observed constant during night hours and small increase is observed only from 0600 to 0800 hrs during the winter season. While in contrary to this, subsequent section shows a continuous decline in the night time concentrations of main anthropogenic tracer CO, which indicates that there is enough vertical mixing of low CO air from above once CO source is turned off, its concentration drops. Hence, constant levels of CO₂ at night hours during these seasons give the evidence of a continued but weak source (such as respiration) in order to offset dilution of mixing of low CO₂ air from aloft. Dry soil conditions could be one of the possible cause for weak respirations. Further, distinct timings have been observed in the morning peak of CO₂ during different seasons. It is mostly related to the sunrise time, which decides the evolution time of PBL height and beginning of vegetation photosynthesis. The sunrise occur at 0555-0620 hrs, 0620-0700 hrs, 0700-0723 hrs and 0720-0554 hrs during summer, autumn, winter and spring respectively. During spring and summer, rush hour starts after sunrise, so the vehicular emissions occur when the PBL is already high and photosynthetic activity has begun. The CO₂ concentration is observed lowest in the morning during the summer season as compared to other seasons. This is because CO₂ uptake by active vegetation deplete the entire mixed layer during day time and when residual layer mixes to the surface in the morning, low-CO₂ air is mixed down. In winter and autumn, rush hour starts parallel with the sunrise, so the emissions occur when the PBL is low and concentration build up is much stronger in these seasons than in spring and summer seasons".

We have also plotted the diurnal variations of mean wind speed, which shows relatively calm winds in the autumn, winter and spring seasons. This information is mentioned in the text, when we discussed the night time ratio of CO and CO₂. The wind data for monsoon period is not available due to break down of wind sensor in the high thunderstorm. The following figure is for your reference only.



Page: 32201

This paragraph about comparison to a model ought to come later as discussion and not be in the results. Also, keep in mind that the magnitude of concentration variation is not directly proportional to the magnitude of a flux. In a simple sense the amplitude of concentration changes are proportional to flux divided by mixed layer depth and strength of vertical mixing. In order to use the observed concentrations to evaluate the validity of modelled CO₂ fluxes you need to consider what the influence region for the concentration is and convert the observations and model to comparable units. Either combine

concentration data and typical mixing depth evolution to estimate a change in column density, or merge the CASA fluxes with a transport model to predict concentrations. The claim that model and observations are inconsistent is not convincing. The greatest magnitude of net daytime uptake and difference between CASA fluxes in day and night is in September through November, consistent with the peaks in amplitude of mixing ratio diel cycle (day/night difference of CO₂ concentration increases from 20 ppm in August to 50 ppm in October). So I don't see where the observations suggest productivity is higher in August than Sept-October.

Response: Thank you very much for your kind suggestion. Now, we have moved this figure in Section 4.7.1

It is clear from Figure 6 that the CO₂ flux diurnal cycle as modelled by CASA shows minimum day-night variation amplitude during the summer monsoon time (Jun-July-Aug). Given that the biosphere over Ahmedabad is water stressed for all other three seasons (except the summer monsoon time, Fig. 1A3), the behaviour of CASA model simulated diurnal variation is not in line with biological capacity of the plants to assimilate atmospheric CO₂.

Due this underestimation of CO₂ uptake in the summer monsoon season, we also find very large underestimation of the seasonal trough by ACTM in comparison with observations (Fig. 11). The variations in transport, PBL ventilation and horizontal winds are included in the ACTM simulation, therefore we do include "proportional to flux divided by mixed layer depth and strength of vertical mixing" in our model results.

For these reasons, we propose that summer time underestimation of CO_2 flux diurnal simulation by CASA is a clearly convincing case.

Page: 32202

The statement here on pg32302, line 26 about respiration contributing to CO₂ is inconsistent with the previous section suggesting that respiration was dormant.

Response: According the first suggestion, we have modified both explanations for CO₂ and CO. In CO₂ section, we have added following explanations.

"CO₂ concentrations start decreasing from 0000 to 0300 hrs and increases slightly afterwards till 0600-0700 hrs during summer and autumn. Respiration of CO₂ from the vegetation is mostly responsible for

this night time increase. During winter and spring seasons CO₂ levels are observed constant during night hours and small increase is observed only from 0600 to 0800 hrs during the winter season. While in contrary to this, subsequent section shows a continuous decline in the night time concentrations of main anthropogenic tracer CO, which indicates that there is enough vertical mixing of low CO air from above that once CO source is turned off, its concentration drops. Hence, constant levels of CO₂ at night hours during these seasons give the evidence of a continued but weak source (such as respiration) in order to offset dilution of mixing of low CO₂ air from aloft. Dry soil conditions could be one of the possible causes for weak respirations".

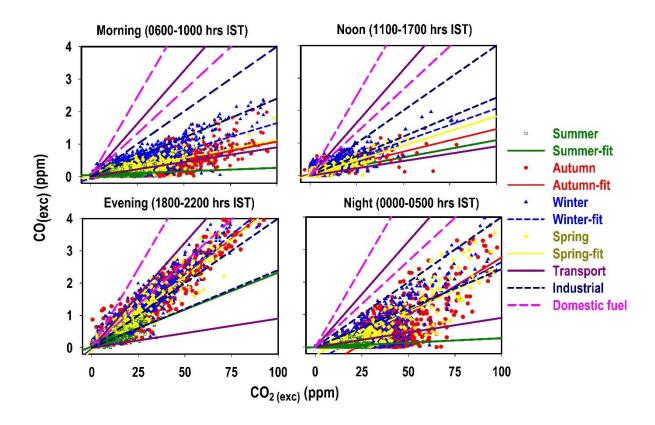
For CO section, we have modified the statement on pg32302, line 26.

"The third noticeable difference is that the CO levels decrease very fast after evening rush hours in all the seasons while this feature is not observed in case of CO₂, since respiration during night hours contributes to the levels of CO₂. The continuous drop of night time concentrations of CO indicates that there is enough vertical mixing of low CO air from above once the CO source is turned off".

Page: 32204

The regression slopes for CO: CO₂ are not credible estimates of the emission ratio. The difference between actual background CO₂ and the assumed constant value that is used to compute excess is correlated with time of day and thus with CO, so the slope of CO: CO₂ will be corrupted I do notice that the upper edge in all the figures appears to have a similar slope. That edge represents the air that is most strongly influenced by CO emission sources. Although I think it would be better to split up the data into groups that actually show a decent correlation, if you want to stick with the overall regression those lines should be shown on the figure and for comparison include some lines that show the slopes for a few representative emission sources.

Response: Thank you very much for raising this important point. As per the suggestion, we have removed the old diagram and added a following new one diagram, which shows the correlation at different time windows during different seasons and including the range of emission ratios of different sources from the available literature. We have included the following table, which gives the summary of the new diagram. According to diagram, we have significantly modified the text of the whole section.



	Slope in ppb/ppm (Coefficient of determination (r²)			
	Morning	Afternoon	Evening	Night
Seasons	(0600-1000 hrs)	(1100-1700 hrs)	(1800-2200 hrs)	(0000-0500 hrs)
Summer (JA)	0.9 (0.15)	10.0 (0.17)	19.5 (0.67)	0.5 (0.16)
Autumn (SON)	8.3 (0.48)	14.1 (0.75)	45.2 (0.90)	35.3 (0.71)
Winter (DJF)	14.3 (0.51)	20.0 (0.68)	47.2 (0.90)	30.0 (0.65)
Spring (MAM)	9.3 (0.68)	18.0 (0.80)	43.7 (0.93)	26.0 (0.80)

Note that in previous section examining diel cycles you made a convincing argument that CO emissions were shut down at night so concentrations declined but CO₂ from respiration continued. Thus, night time data should not be useful for finding an emission ratio.

Response: We totally agree with you that night time CO₂ levels are additionally affected by the respiration sources. Although there are influence of respiration sources in the levels of CO₂, but not very strong. We

have discussed it previously in the revised manuscript also. Since the wind condition is calm during this period due to no turbulence and most of dominant sources are shut off, the ratios during this period can be useful for broadly understanding about the emission characteristic of dominating sources over Ahmedabad.

I would suggest trying something similar to the analysis of Potosnak et al 1999 that seeks to extract the influence from biosphere and mean diel cycle. (J. Geophys. Res., 104(D8), 9561–9569, doi:10.1029/1999JD900102.)...

Response: It seems to be a good suggestion. However, it requires complete reanalysis of the data in different time bins and beyond the scope of present work. We are extremely sorry.

Page: 32205

In the end the CO: CO₂ ratios have such a wide range as to not be very useful at all. Unless you can reanalyse them to bring a narrower estimate it is not worthwhile to show this section. It is curious that the night time data have such a good correlation when the diel cycle analysis suggested that combustion emissions of both CO and CO₂ together were shut down. It would help to illustrate the relationship between CO and CO₂ in night by colouring the symbols for night-time data differently for time of day in Figure 8a I suspect the daytime values, with low correlation coefficients are not reliable, as you suggest by indicating the importance of CO₂ uptake. When biospheric influence influences the CO₂ mixing ratio you shouldn't bother to try to analyze the CO: CO₂ ratio.

Response: Thank you very much for the suggestion. As per suggestion, we have removed old figure and added a new one, in which the data are segregated in different time windows and coloured according to different seasons. The modified diagram is already shown previously (Page:32204).

As we discussed previously also the CO₂ levels are additionally affected by respiration sources during night time, but not very strongly. We also discussed that CO levels drop very fast during night time, which indicate that there is enough vertical mixing of low CO air from above that once the CO sources are turned off. Hence this mixing will enhance the correlation during night time, since there are no significant sources, which disturb their levels greatly. Correlation during day time is low only during monsoon season, since biospheric productivity play a large role in influencing the levels of CO₂. But for making the comparisons, we have included the day time values. While during other seasons, correlation is pretty good during day time due to significant atmospheric mixing of all emissions and comparatively lower biospheric

productivity. It concludes that during other seasons CO and CO2 levels are mostly dominated by common emission sources. This whole section is now modified according to previous comments.

Page: 32206

The previous section about CO: CO₂ slopes is rather muddled. It would be more convincing focussing on demonstrating the validity of just the night-time and rush-hour periods that you are using here. Showing the data for entire day just confuses things.

Response: As per the suggestion, now Section 4.5 includes the validity of EDGAR CO emissions from the night-time and rush hour periods measurements only. We have discussed previously also the measurements during these period will show combine influence from all anthropogenic sources mostly. Hence, the estimated slopes for these period will be helpful to validate the anthropogenic CO emissions of EDGAR inventory. According to that, we have modified our conclusion. We have replaced the fossil fuel emission term by the anthropogenic emissions. It includes all emissions such as vehicular emission, industrial emission as well as cooking sector emissions.

Assuming the discussion of ratios just for the relevant periods is more convincing you can also include some calculation of the uncertainty, which then feeds into providing estimates of uncertainty in the emissions you compute from those ratios and the CO₂ inventory. Uncertainty estimates are critical to include here.

Response: We are highly thankful for the suggestion. The possible causes for uncertainty are the uncertainty in estimated slopes, uncertainty in CO and CO2 emissions used for EDGAR inventory and uncertainty in the interpolation of the emission of both the gases. We have included the following text in the section due to the uncertainty in slopes. However, due to unavailability of uncertainty information in the emissions of EDGAR inventory, it is not possible to include these uncertainty in the calculation. We are very sorry for that.

"Further the uncertainty in total estimated emission of CO due to uncertainty associated with used slope is also calculated. Using this slope and based on CO₂ emissions from EDGAR inventory, the estimated fossil fuel emission for CO is observed to be 69.2±0.7 Gg (emission ± uncertainty) for the year of 2014."

Page: 32235

Consider plotting actual CO and CO₂ mixing ratios to see if the intercepts match the values chosen for background. In the active growing season the biospheric influence will impart a wide range of CO2 for

given values of CO, which is what shows for most seasons. A meaningful slope is difficult to extract in this case. A better estimate of CO₂: CO could be derived by using information from the mean diel cycle analysis to subtract a variable background, or restrict the analysis to just a fixed time of day, or analyze night and daytime separately...

Response: We are very much thankful for your suggestion. According to previous and this suggestion, we have modified the figure significantly, in which we have removed the diurnal variation of slopes and restrict our analysis at different time windows separately.

Minor editing

32197 line 25

There must be a missing word in the sentence; 'resulting in concentrations at the surface in the summer compared to the winter.

Response: Thank you very much for pointing out this slip. Yes, it was missing and we have corrected it now.

Anonymous referee #2

General comments:

This paper addresses temporal variations of atmospheric CO₂ and CO in an urban site in western India. There are not so many studies on greenhouse gases in urban environments. Furthermore, such study is rare in countries in development. This work is interesting to be published and is fully within the scope of ACP.

The authors address the seasonal and the diurnal scales, as well as CO/CO₂ ratios from which they infer information on the anthropogenic vs natural sources of CO and CO₂. They also propose a calculation of CO emissions for the studied city using such ratio and CO₂ emissions from inventory, as well as a short model/observations comparison. I acknowledge the large amount of work provided by the authors and interesting information issued from this study.

Response: We are very much grateful to the referee for the positive comments and appreciation of our work. We have tried our best to give all the answers point by point.

However, there are also some major issues to be addressed and reviewed before publication in ACP. These issues concern:

1. The form: The text is quite difficult to read and needs to be synthetized, especially the introduction, the seasonal study and the diurnal study. Some sentences are even repeated twice.

Response: According to the suggestion, we have synthetized all mentioned sections very carefully after removing, adding and rearranging sentences. Please note that, rather than adding, removing and rearranging sentences in between the written text, we have added modified text of whole section in new paragraph. The modified sections are given below.

Introduction:

Carbon dioxide (CO₂) is the most important anthropogenically emitted greenhouse gas (GHG) and has increased substantially from 278 to 390 parts per million (ppm) in the atmosphere since the beginning of the industrial era (circa 1750). It has contributed to more than 65% of the radiative forcing increase since 1750 and hence has significant impact on the climate system (Ciais et al., 2013). Major causes of CO₂ increase are anthropogenic emissions, especially fossil fuel combustion, cement production and land use change. Land and oceans are the two important sinks of atmospheric CO₂, which remove about half of the anthropogenic emissions (Le Quéré et al., 2014). The prediction of future climate change and its feedback rely mostly on our ability to quantify fluxes of greenhouse gases, especially CO₂, at regional (100-1000 km²) and global scales. Though the global fluxes of CO₂ can be estimated fairly well, the

regional scale (fluxes are associated with quite high uncertainty, especially over the South Asian region and the estimation uncertainty being larger than the value itself (Patra et al., 2013; Peylin et al., 2013). Detailed scientific understanding of the flux distributions is also needed for formulating effective mitigation policies.

Along with the need for atmospheric measurements for predicting the future levels of CO₂, quantifying the components of anthropogenic emissions of CO₂ is similarly important for providing independent verification of mitigation strategies as well as understanding the biospheric component of CO₂. Only CO₂ measurements cannot be helpful in making such study due to the larger role of biospheric fluxes in its atmospheric distributions. The proposed strategy for quantification of the anthropogenic component of CO₂ emissions is to measure simultaneously anthropogenic tracers,. CO can be used as a surrogate tracer for detecting and quantifying anthropogenic emissions from burning processes, since it is a major product of incomplete combustion (Turnbull et al., 2006; Wang et al., 2010). The vehicular as well as industrial emissions contribute large fluxes of CO₂ and CO to the atmosphere in urban regions. Several ground and aircraft based simultaneous studies of CO andCO₂ have been made in the past from different parts of the world (Turnbull et al., 2006; Wunch et al., 2009; Wang et al., 2010; Newman et al., 2013) but such a study is lacking in India except recently reported results from weekly samples for three Indian sites by Lin et al. (2015).

Measurements at different regions (eg. rural, remote, urban) and at different frequency (eg., weekly, daily, hourly etc) have their own advantage and limitations. For example, the measurements at remote locations at weekly interval can be useful for studying seasonal cycle, growth rate, and estimating the regional carbon sources and sinks after combining their concentrations with inverse modelling and atmospheric tracer transport models. However, some important studies, like diurnal variations, temporal covariance...etc are not possible from these measurements. Analysis of temporal covariance of atmospheric mixing processes and variation of flux on shorter time scales, e.g., sub-daily, is essential for understanding local to urban scale CO₂ flux variations (Ahmadov et al., 2007; Pérez-Landa et al., 2007; Briber et al., 2013; Lopez et al., 2013; Ammoura et al., 2014; Ballav et al., 2015). Urban regions contribute about 70% of global CO₂ emission from anthropogenic sources and further projected to increase over the coming decades. Therefore, measurements from these regions are very helpful for understanding emissions, growth as well as verifying the mitigation policies. The first observations of CO₂, CO and other greenhouse gases started in February 1993 from Cape Rama (CRI: a coastal site) on the south-west coast of India using flask samples (Bhattacharya et al., 2009). After that, several other groups have initiated the measurements of surface level greenhouse gases (Mahesh et al., 2014; Sharma et al., 2014; Tiwari et al., 2014; Lin et al., 2015). Most of these measurements are made at weekly or fortnightly time intervals. Two aircraft based measurement programs, namely, Civil Aircraft for the regular Investigation of the atmosphere Based on an Instrument Container (CARIBIC) (Brenninkmeijer et al., 2007) and Comprehensive Observation Network for TRace gases by AlrLiner (CONTRAIL) (Machida et al., 2008) have provided important first look on the South Asian CO₂ budget, but these data have their own limitations (Patra et al., 2011; Schuck et al., 2010, 2012). It is pertinent to mention here that till now, there are no reports of CO2 measurements over the urban locations of the Indian subcontinent, which could be an important player in the global carbon budget as well as mitigation purpose due to strong growing anthropogenic activities specifically fast growing traffic sector and sinks (large areas of forests and croplands). Hence, the present study is an attempt to reduce this gap by understanding the CO₂ mixing ratios in light of its sources and sinks at an urban region in India.

In the view from above, simultaneous continuous measurements of CO₂ and CO have been made since November 2013 from an urban site Ahmedabad located in the western India using very highly sensitive laser based technique. The preliminary results of these measurements for one month period have been reported in (Lal et al., 2015). These detailed measurements are utilized for studying the temporal variations (diurnal and seasonal) of both gases, their emission characteristics on diurnal and seasonal scale using their mutual correlations, estimating the contribution of vehicular and biospheric emission components in the diurnal cycle of CO₂ using the ratios of CO to CO₂ and rough estimate of the annual CO emissions from study region. Finally, the measurements of CO₂ have been compared with simulations using an atmospheric chemistry-transport model to discuss roles of various processes contributing to CO₂ concentration variations.

Seasonal cycle of CO₂ and CO:

The seasonal cycles of CO₂ and CO are mostly governed by the strength of emission sources, sinks and transport patterns. Although they follow almost identical seasonal patterns, but the factors responsible for their seasonal behaviours are distinct as for the diurnal variations. We calculate the seasonal cycle of CO₂ and CO using two different approaches. In the first approach, we use the monthly mean of all measurements and in the second approach we use monthly mean of the afternoon period (1200-1600 hrs) measurements only. The seasonal cycle from first approach will present the overall variability in both gases. On the other hand, second approach removes the auto covariance by excluding CO₂ and CO data mainly affected by local emission sources and represent seasonal cycles at the well mixed volume of the atmosphere. The CO₂ time series is detrended by subtracting a mean growth rate of CO₂ observed at Mauna Loa (MLO), Hawaii, i.e., 2.13 ppm yr⁻¹ or 0.177 ppm/month (www.esrl.noaa.gov/gmd/ccgg/trends/) for clearly depicting the seasonal cycle amplitude.

In general, total mean values of CO₂ and CO are observed lower in July having concentration 398.78±2.8 ppm and 0.15±0.05 ppm respectively. During summer monsoon months predominance of south-westerly winds which bring cleaner air from the Arabian Sea and the Indian Ocean over to Ahmedabad and high VC (Figure 1) are mostly responsible for the lower concentration of total mean of both the gases. CO₂ and CO concentrations are also at their seasonal low in the northern hemisphere due to net biospheric uptake and seasonally high chemical loss by reaction with OH, respectively. In addition, deep convections in the summer monsoon season efficiently transport the Indian emission (for CO, hydrocarbons) or uptake (for CO₂) signals at the surface to the upper troposphere, resulting lower concentrations at the surface in the as compared to the winter months (Kar et al., 2004; Randel and Park, 2006; Park et al., 2009; Patra et al., 2011; Baker et al., 2012). During autumn and early winter (December), lower VC values cause trapping of anthropogenically emitted CO₂ and CO and is the major cause for high concentrations of both gases during this period. In addition to this, wind changes from cleaner marine region to polluted continental region, especially from IGP region and hence could be additional factor for higher levels of CO₂ and CO during these seasons (autumn and winter). Elevated levels during these seasons are also examined in several other pollutants over Ahmedabad as discussed in previous studies (Sahu et al., 2006, Mallik et

al., 2016). Maximum concentrations of CO_2 and CO are observed to be 424.85±17 ppm and 0.83±0.53 ppm, respectively, during November. From January to May the total mean concentration of CO_2 decreases from 415.34±13.6 to 406.14±5.0 ppm and total mean concentration of CO decreases from 0.71±0.22 to 0.22±0.10 ppm. Higher VC and predominance of comparatively less polluted mixed air masses from oceanic and continental region results in lower total mean concentrations of both gases during this period.

There are some clear differences which are observed in the afternoon mean concentrations of CO_2 as compared to daily mean. The first distinct feature is that significant difference of about 5 ppm is observed in the afternoon mean of CO_2 concentration from July to August as compared to the difference in total mean concentration about ~ 0.38 ppm for the same period. Significant differences in the afternoon concentrations of CO_2 from July to August is mainly due to the increasing sink by net biospheric productivity after the Indian summer monsoonal rainfall. Another distinct feature is that the daily mean concentration of CO_2 is found higher in November while the afternoon mean concentration of CO_2 attain maximum value (406 ± 0.4 ppm) in April. Prolonged dry season combined with high daytime temperature (about $41^{\circ}C$) during April-May make the tendency of ecosystem to become a moderate source of carbon exchange (Patra et al., 2011) and this could be responsible for the elevated mean noon time concentrations of CO_2 . Unlike CO_2 , seasonal patterns of CO from total and afternoon mean concentrations are identical, although levels are different. It shows that the concentrations of CO is mostly governed by identical sources during day and night time throughout the year.

The average amplitude (max - min) of the annual cycle of CO₂ is observed around 13.6 and 26.07 ppm from the afternoon mean and total mean respectively. Different annual cycles and amplitudes have been observed from other studies conducted over different Indian stations. Similar to our observations of the afternoon mean concentrations of CO₂, maximum values are also observed in April at Pondicherry (PON) and Port Blair (PBL) with amplitude of mean seasonal cycles about 7.6±1.4 and 11.1±1.3 ppm respectively (Lin et al., 2015). Cape Rama (CRI), a costal site on the south-west coast of India show the seasonal maxima one month before than our observations in March annual amplitude about 9 ppm (Bhattacharya et al., 2009). The Sinhagad (SNG) site located over the Western Ghats Mountains, show very larger seasonal cycle with annual amplitude about 20 ppm (Tiwari et al., 2014). The amplitude of mean annual cycle at the free tropospheric site Hanle at altitude of 4500 m is observed to be 8.2±0.4 ppm, with maxima in early May and the minima in mid-September (Lin et al., 2015). Distinct seasonal amplitudes and patterns are due to differences in regional controlling factors for the seasonal cycle of CO₂ over these locations, e.g., the Hanle is remotely located from all continental sources, Port Blair site sampled predominantly marine air, Cape Rama observes marine air in the summer monsoon season and Indian flux signals in the winter, and Sinhagad represents a forested ecosystem. These comparisons show the need for CO₂ measurements over different ecosystems for constraining its budget.

The annual amplitude in afternoon and daily mean CO concentration is observed to be about 0.27 and 0.68 ppm, respectively. The mean annual cycles of CO over PON and PBL show the maxima in the winter months and minima in summer months same as our observations with annual amplitudes of 0.078±0.01 and 0.144±0.016 ppm, respectively. So the seasonal levels of CO are affected by large scale dynamics which changes air masses from marine to continental and vice versa and by photochemistry. The

amplitudes of annual cycle at these locations differ due to their climatic conditions and sources/sinks strengths.

4.3.1. Diurnal variation of CO₂

Figure 5a shows the mean diurnal cycles of atmospheric CO₂ and associated 1-σ standard deviation (shaded region) during all the four seasons. All times are in Indian Standard Time (IST), which is 5.5 hrs ahead of Greenwich Mean Time (GMT). Noticeable differences are observed in the diurnal cycle of CO2 from season to season. In general, maximum concentration has been observed during morning (0700-0800 hrs) and evening (1800-2000 hrs) hours, when ABL is shallow, traffic is dense and vegetation respiration dominate due to absence of photosynthetic activity. The minimum of the cycles occurred in the afternoon hours (1400-1600 hrs), when PBL is deepest and well mixed as well as when the vegetation photosynthesis is active. There are many interesting features in the period of 0000-0800 hrs. CO2 concentrations start decreasing from 0000 to 0300 hrs and increases slightly afterwards till 0600-0700 hrs during summer and autumn. Respiration of CO₂ from the vegetation is mostly responsible for this night time increase. During winter and spring seasons CO2 levels are observed constant during night hours and small increase is observed only from 0600 to 0800 hrs during the winter season. While in contrary to this, subsequent section shows a continuous decline in the night time concentrations of main anthropogenic tracer CO, which indicates that there is enough vertical mixing of low CO air from above that once the CO source is turned off, its concentration drops. Hence, constant levels of CO₂ at night hours during these seasons gives the evidence of a continued but weak sources (such as respiration) in order to offset dilution of mixing of low CO₂ air from aloft. Dry soil conditions could be one of the possible caused for weak respirations. Further, distinct timings have been observed in the morning peak of CO₂ during different seasons. It is mostly related to the sunrise time, which will decide the evolutions time of PBL height and beginning of vegetation photosynthesis. The sunrise occurs at 0555-0620 hrs, 0620-0700 hrs, 0700-0723 hrs and 0720-0554 hrs during summer, autumn, winter and spring, respectively. During spring and summer, rush hour starts after sunrise, so the vehicular emissions occur when the PBL has been already high and photosynthetic activity has begun. The CO₂ concentration is observed lowest in the morning during the summer season as compared to other seasons. This is because CO2 uptake by active vegetation deplete entire mixed layer during day time and when the residual layer mixes to the surface in the morning, low-CO₂ air is mixed down. In winter and autumn, rush hour starts parallel with the sunrise, so the emissions occur when the PBL is low and concentration builds up is much stronger in these seasons than in spring and summer seasons.

The diurnal amplitude is defined as the difference between the maximum and minimum concentrations of CO₂ in the diurnal cycle. The amplitudes of monthly averaged diurnal cycle of CO₂ from July, 2014 to May 2015 are shown in Figure 5b. The diurnal amplitude shows large month to month variation with increasing trend from July to October and decreasing trend from October onwards. The lowest diurnal amplitude of about 6 ppm is observed in July while highest amplitude about 51 ppm is observed in October. The amplitude does not change largely from December to March and is observed in the range of 25-30 ppm. Similarly from April to May the amplitude also varies in a narrow range from 12 to 15 ppm. The jump in the amplitude of the CO₂ diurnal cycle is observed highest (around 208%) from July to August. This is mainly due to significant increase of biospheric productivity from July to August after the rains in

Ahmedabad. It is observed that during July the noon time CO₂ levels are found in the range of 394-397 ppm while in August the noon time levels are observed in the range of 382-393 ppm. The lower levels could be due to the higher PBL height during afternoon and cleaner air, but in case of CO (will be discussed in next section), average day time levels in August are observed higher than in July. It rules out that the lower levels during August are due to the higher PBL height and presence of cleaner marine air, and confirms the higher biospheric productivity during August.

Near surface diurnal amplitude of CO₂ has been also documented in humid subtropical Indian station Dehradun and a dry tropical Indian station Gadanki (Sharma et al., 2014). In comparison to Ahmedabad, both these stations show distinct seasonal change in the diurnal amplitude of CO₂. The maximum CO₂ diurnal amplitude of about 69 ppm is observed during the summer season at Dehradun (30.3°N, 78.0°E, 435m), whereas maximum of about 50 ppm during autumn at Gadanki 400 (13.5°N, 79.2°E, 360 m).

Diurnal variation of CO:

Figure 7a shows seasonally averaged diurnal variation of CO. In general, the mean diurnal cycle of CO show lower concentration during noon (1200-1700 hrs) and two peaks in the morning (0800 to 1000 hrs) and in the evening (1800 to 2200 hrs) hours. This cycle exhibits the same pattern as the mean diurnal cycle of traffic flow, with maxima in the morning and at the end of the afternoon, which suggests the influence of traffic emissions on CO measurements. Along with the traffic flow, PBL dynamics also play a critical role in governing the diurnal cycle of CO. The amplitudes of the evening peak in diurnal cycles of CO are always greater than the morning peaks. It is because the PBL height evolves side by side with the morning rush hours traffic and hence increased dilution, while on the other hand, during evening hours, PBL height decrease along with evening dense traffic and favours accumulation of pollutants until the late evening under the stable PBL conditions. The noon time minimum of the cycle is mostly associated with the deepest and well mixed PBL.

In general, the average diurnal cycle patterns of both gases (CO₂ and CO) are similar, but having few noticeable differences. The first difference is observed in the timing of the occurrence of morning peaks: CO₂ peaks occur slightly before than CO peak due to the triggering photosynthesis process by the sunrise. On the other hand, the morning peaks of CO mostly depend on the rush hour traffic and are consistent at 0800-1000 hrs in all seasons. The second difference is that the afternoon concentrations of CO show little seasonal spread as compared to the afternoon concentrations of CO₂. Again, this is due to the biospheric control on the levels of CO₂ during the afternoon hours of different seasons while CO levels are mainly controlled by the dilution during these hours. The third noticeable difference is that the levels of CO decrease very fast after evening rush hours in all seasons while this feature is not observed in the case of CO₂ since respiration during night hours contributes to the levels of CO₂. The continuous drops of night time concentrations of CO indicates that there is enough vertical mixing of low CO air from above that once the CO source is turned off. The average morning (0800-0900 hrs) peak values of CO is observed minimum (0.18±0.1 ppm) in summer and maximum (0.72±0.16 ppm) in winter while evening peak shows minimum value (0.34±0.14 ppm) in summer and maximum (1.6±0.74 ppm) in autumn. The changes in

CO concentrations show large fluctuations from morning peak to afternoon minima and from afternoon minima to evening peak. From early morning maxima to noon minima, the changes in CO concentrations are found in the range of 20 -200%, while from noon minima to late evening maxima the changes in CO concentrations are found in the range of 85% to 680%. Similar diurnal variations with two peaks have also been observed in earlier measurements of CO as well as NOx at this site (Lal et al., 2000).

The evening peak contributes significantly to the diurnal amplitude of CO. The largest amplitude in CO cycle is observed in autumn (1.36 ppm) while the smallest amplitude is observed in summer (0.24 ppm). The diurnal amplitudes of CO are observed to be about 1.01 and 0.62 ppm, respectively during winter and spring. Like CO₂, the diurnal cycle of CO (Figure 7b) shows the minimum (0.156 ppm) amplitude in July and maximum (1.85 ppm) in October. After October the diurnal amplitude keeps on decreasing till summer. The evening peak contributes significantly to the diurnal amplitude of CO. The largest amplitude in CO cycle is observed in autumn (1.36 ppm) while the smallest amplitude is observed in summer (0.24 ppm). The diurnal amplitudes of CO are observed to be about 1.01 and 0.62 ppm, respectively during winter and spring. The monthly diurnal cycle of CO (Figure 7b) shows the minimum (0.156 ppm) amplitude in July and maximum (1.85 ppm) in October. After October the diurnal amplitude keeps on decreasing till summer.

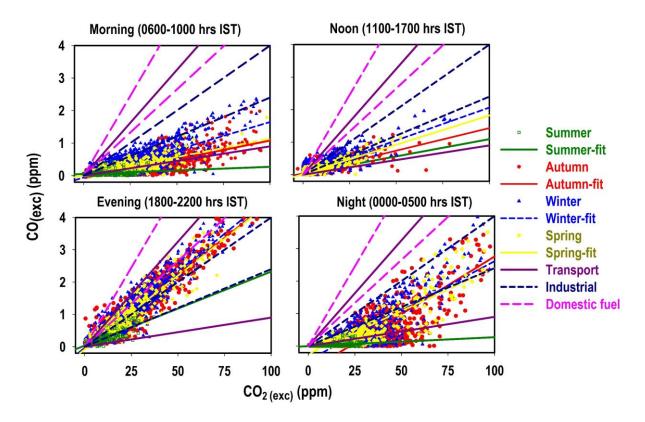
2. The content given on the emissions and the conclusions on the traffic sector vs the cooking and industrial one: there is a lack of information on the studied region and on the relative role of the different emission sectors that should be given quantitatively, with proper references. Especially, the part of emissions due to residential and slum cooking is almost not discussed, while the available literature explains that this emission sector is responsible for a large amount of atmospheric CO (less in CO₂).

Response: As per suggestion, we have included following more information about the relative contributions of different emission sectors in Section 2.

"An emission inventory for this city, which is developed for all known sources, shows the annual emissions (for year 2010) of CO₂ and CO about 22.4 million tons and 707,000 tons respectively (http://www.indiaenvironmentportal.org.in/files/file/Air-Pollution-in-Six-Indian-Cities.pdf). Out of these emissions, transport sector contribute about 36% in CO₂ emission and 25% in CO emissions, power plants contribute about 32% in CO₂ emissions and 30% in CO emission, industries contribute about 18% in CO₂ emissions and 12% in CO emissions and domestic sources contribute about 6% in CO₂ emissions and 22% in CO emissions."

The conclusion on the strong influence of traffic given in sector 4.4 is not convincing according to Table1. Response: Thank you very much for raising these points. This is also pointed by the Referee #1. Now we have modified our conclusions as per further suggestions. As per suggestions, we have removed old figure and added a following new figure, in which the data are segregated seasonally in different time

windows. After carefully analysing the CO: CO₂ ratios, now we have included the domestic and industrial emission sectors along with the transport sector. After this modification, the crux of this modified study is that, during evening hours, emission from transport and domestic sources mostly dominate while rest of periods transport and industrial emission sources mostly dominate mostly over the study region.

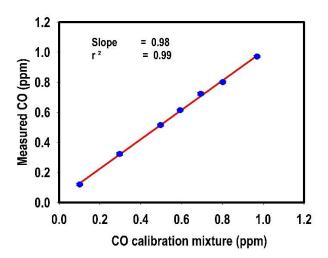


But it is then used as acquired in the following sections (e.g. 4.5 and 4.6), that should also be reviewed in function, taking into account especially the cooking sector, which is another anthropogenic source (nor a natural nor a fossil one), as well as the industrial one (as the authors mentioned there are several industries in the studied area).

Response: As per your and the first referee suggestions, now Section 4.5 includes the validity only from the night-time and rush-hours periods. The estimated slopes for these periods will show the combined influence from all anthropogenic sources and hence will be helpful to validate the anthropogenic emissions of CO. Rest of the results corresponding for different period, we have removed from the section for presenting results more clearly. Now we have modified our conclusion accordingly. We have replaced the fossil fuel emission terms from the anthropogenic emissions. It includes all emissions such as vehicular emission, industrial emission as well as cooking sector emissions.

3/. The validation of the data: the quality of the CO data set is poorly explain, and the data treated with a single calibration standard while no test is given on the instrumental linearity for this species. The data treatment should thus be better reported and these specific points precisely addressed.

Response: Thank you very much for your suggestion. We have calibrated this instrument with one calibration mixture from Linde, UK. As per suggestion, we have checked the instrument linearity. We have diluted the calibration mixture of CO from 970 ppb to 100 ppb. The calibration mixture is diluted in a glass made dead volume using highly precise and accurate pure air (air free from water vapour, particles, carbon monoxide (CO), sulphur dioxide (SO₂), oxides of nitrogen (NOx), ozone (O₃), and hydrocarbons (HC)) from ECO Physics generator. The flows of calibration mixtures and pure air were regulated using two separate mass flow controllers from Aalborg. For increasing the interaction time both gases (zero air and calibration mixture) to insure a homogeneous mixing, the spring shaped dead volume is used. Each diluted mixture is passed for 30 minutes in the instrument and average of last 10 minute concentration is used for the test. The instrument shows excellent linearity for CO and slope is observed 0.98. The calibration graph is given below for the reference only. The x-axis and y-axis show the calibration values and measured values by the analyzer, respectively. Accordingly we have modified the text in the section.



4/. The question of entertainment of air on top of the PBLH is not addressed, and the PBLH seems to be considered as mixed at any time of the day.

Response: We have added the information regarding the entertainment of air on top of the PBLH in the diurnal variation of both CO₂ and CO sections.

"During winter and spring seasons CO2 levels are observed constant during night hours and small increase is observed only from 0600 to 0800 hrs during the winter season. While in contrary to this, subsequent section shows a continuous decline in the night time concentrations of main anthropogenic tracer CO, which indicates that there is enough vertical mixing of low CO air from above that once the CO source is turned off, its concentration drops. Hence, constant levels of CO₂ at night hours during these seasons gives the evidence of a continued but weak sources (such as respiration) in order to offset dilution of mixing of low CO2 air from aloft. Dry soil conditions could be one of the possible caused for weak respirations. Further, distinct timings have been observed in the morning peak of CO₂ during different seasons. It is mostly related to the sunrise time, which will decide the evolutions time of PBL height and beginning of vegetation photosynthesis. The sunrise occurs at 0555-0620 hrs, 0620-0700 hrs, 0700-0723 hrs and 0720-0554 hrs during summer, autumn, winter and spring, respectively. During spring and summer, rush hour starts after sunrise, so the vehicular emissions occur when the PBL has been already high and photosynthetic activity has begun. The CO₂ concentration is observed lowest in the morning during the summer season as compared to other seasons. This is because CO2 uptake by active vegetation deplete entire mixed layer during day time and when the residual layer mixes to the surface in the morning, low-CO₂ air is mixed down. In winter and autumn, rush hour starts parallel with the sunrise, so the emissions occur when the PBL is low and concentration builds up is much stronger in these seasons than in spring and summer seasons".

The CO/CO₂ ratio diurnal variability should take this into account, a point that carefully needs to be studied at different time windows.

Response: As per the suggestions in further comments, we have modified this section, which discuss about the diurnal variability of CO/CO₂ ratios. Instead of using whole measurements of both gases for studying the emission characteristics, now the modified section discuss the emission characteristics of dominant emission sources at different time windows during all the four seasons. Most of details about this section are given in the following comments.

After these major revisions, I am convinced that this work will be of very good quality for publication in ACP.

Specific comments:

Abstract:

A sentence on your objectives should be given after the first sentence. What is the reliability of the CO₂ emissions inventory?

Response: We thank you very much for the suggestions. We have added and rearranged following lines in the abstract section.

"In order to draw effective emission mitigation policies for combating future climate change as well as independently validate the emission inventories for constraining their large range of uncertainties, especially over major metropolitan areas of developing countries, there is an urgent need for greenhouse gases measurements over representative urban regions. India is a fast developing country, where fossil fuel emissions have increased dramatically in the last three decades and predicted further to continue to grow by at least 6% per year through 2025. In the absence of systematic CO₂ measurements over the Indian urban locations, CO₂ along with an anthropogenic emission tracer carbon monoxide (CO) are being measured at Ahmedabad, a major urban site in western India, using a state-of-the-art laser based cavity ring down spectroscopy technique from November 2013 to May 2015 with a break during March to June 2014. These measurements enable us to understand the diurnal and seasonal variation in atmospheric CO₂ with respect to its sources (both anthropogenic and biospheric) and biospheric sinks."

Introduction

Much too long. Remove detailed information.

Response: Referee #1 also pointed out this issue. So now we have very carefully modified, rearranged and synthesized the introductory section.

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p. 32187
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Remove lines 8-11 (too long) ------Removed

Line 14: a country can be very small or very large so give rather km (100-1000 km² for the regional scale generally) ------Modified the text accordingly

Line 21: different... = this sentence is very unprecise ----- Removed

Line 28: not only traffic but also industry etc. ---- Changed accordingly

p. 32188

Lines 2-8: too long ------

Response: According to comment #1 we have synthesized the introductory section. So as a part of this change, we have removed some information from this para and modified text by adding the following lines.

"Along with the need for atmospheric measurements for predicting the future levels of CO₂, quantifying the components of anthropogenic emissions of CO₂ is similarly important for providing independent verification of mitigation strategies as well as understanding the biospheric component of CO₂. Only CO₂ measurements can not be helpful for making such study due to the large role of biospheric fluxes in its atmospheric distributions. The proposed strategy for quantification of the anthropogenic component of CO₂ emissions is to measure simultaneously the anthropogenic tracers. CO can be used as a surrogate tracer for detecting and quantifying anthropogenic emissions from burning processes, since it is a major product of incomplete combustion (Turnbull et al., 2006; Wang et al., 2010). The vehicular as well as industrial emissions contribute large fluxes of CO₂ and CO to the atmosphere in urban regions. Several ground based and aircraft based simultaneous studies of CO andCO₂ have been done in the past from different parts of the world (Turnbull et al., 2006; Wunch et al., 2009; Wang et al., 2010; Newman et al., 2013) but such study is lacking in India except recently reported results from weekly samples for three Indian sites by Lin et al. (2015)".

"Measurements at different regions (e.g., rural, remote, urban) and at different frequency (e.g., weekly, daily, hourly etc) have their own advantage and limitations. For example, measurements at remote locations at weekly interval can be useful for studying seasonal cycle, growth rate, and estimating the regional carbon sources and sinks after combining their concentrations with inverse modelling and atmospheric tracer transport models. However, some important studies like their diurnal variations, temporal covariance, etc are not possible from these measurements due to their limitations. Analysis of temporal covariance of atmospheric mixing processes and variation of flux on shorter timescales, e.g., sub-daily, is essential for understanding local to urban scale CO₂ flux variations (Ahmadov et al., 2007; Pérez-Landa et al., 2007; Briber et al., 2013; Lopez et al., 2013; Ammoura et al., 2014; Ballav et al., 2015). Urban regions contribute about 70% of global CO₂ emission from anthropogenic sources and further projected to increase further over the coming decades. Therefore, measurements over these regions are very helpful for understanding emissions growth as well as verifying the mitigation policies. "

p.32190

Objectives not clear, reformulate please.

Response: We have added following lines before the paragraph for highlighting the importance of the

study.

"It is pertinent to mention here that till now there are no reports of CO₂ measurements over the urban

locations of the Indian subcontinent, which could be an important player in the global carbon budget as

well as mitigation purpose due to strong growing anthropogenic activities specifically fast growing traffic

sector and sinks (large areas of forests and croplands). Hence, the present study is an attempt to reduce

this gap by understanding the CO2 mixing ratios in light of its sources and sinks at an urban region in

India".

Section 2

Lines 15-27

What is the height of the sampling height above ground level?

Response: The sampling height is about 25 meter above the ground level. This information is already

given in the line #6 of p. 31293.

The information given on the emission sectors should be improved. It is a key point of your argumentation

next. Please quantify here and give numbers on the relative role of the different CO2 and CO emission

(there Ahmedabad sectors in are several sources compare, here one:

http://www.indiaenvironmentportal.org.in/files/file/Air-Pollution-in-Six-IndianCities.pdf).

Response: We thank you very much for bringing out these points. We have added following information

in Section 2.

"An emission inventory for this city, which is developed for all known sources, shows the annual emissions

(for year 2010) of CO₂ and CO about 22.4 million tons and 707,000 tons respectively

(http://www.indiaenvironmentportal.org.in/files/file/Air-Pollution-in-Six-Indian-Cities.pdf). Out of these

emissions, transport sector contributes about 36% in CO₂ emission and 25% in CO emissions, power

plants contribute about 32% in CO₂ emissions and 30% in CO emission, industries contribute about 18%

in CO₂ emissions and 12% in CO emissions and domestic sources contribute about 6% in CO₂ emissions

and 22% in CO emissions".

p.32191

Line 17 and line 21: check months consistency.Checked

22

Section 3

This section generally lacks of precision on the procedures.

p.32193

Lines 1-2: do you mean the CRDS instruments in general, or yours? Your instruments should be discussed here, each CRDS instrumental is specific and needs to be validated (although this is right that they are usually within WMO recommendations).

Response: In this section we have discussed mostly about our instrument. We have modified respective sentences accordingly.

p.32194

Lines 1-12: this part is critical. The CO Dataset is calibrated with one single tank from Linde UK. Is it linked to the WMO scale? Despite this single cal tank, no linearity tests are reported for CO. How can you make sure your CO data set is not biased by an instrumental drift? Also, you need to report the accuracy of your measurements (both CO₂ and CO).

Response: The calibration tank of CO does not follow the WMO scale. Its traceability is based on NIST by weight. The accuracy of both gases is included in the respective section. As per suggestion and discussed previously, we have checked instrument linearity for CO and found that the instrument is highly linear for CO. The graph is already given in the previous discussion for the reference only. The modified text is given below.

"The measurement system is equipped with three high pressure aluminium cylinders containing gas mixtures of CO₂ (350.67±0.02, 399.68±0.002 and 426.20±0.006 ppm) in dry air from NOAA, Bolder USA, and one cylinder of CO (970 ppb) from Linde UK. These tanks were used to calibrate the instrument for CO₂ and CO. An additional gas standard tank (CO₂: 338 ppm, CO: 700 ppb), known as the "target", is used to monitor the instrumental drift and to assess the dataset accuracy and repeatability. The target tank values are calibrated against the CO₂ and CO calibration mixtures. The target tank and calibration gases were measured mostly in the mid of every month (Each calibration gas is passed for 30 minute and target tank for 60 minutes). The target gas is introduced into the instrument for a period of 24 hours also once in a six month, for checking the diurnal variability of instrument drift. Maximum drift for 24 hours has been calculated by subtracting the maximum and minimum value of 5 minute average, which were found

to be 0.2 ppm and 0.015 ppm respectively for CO₂ and CO. For all calibration mixtures, the measured concentration is calculated as the average of the last 10 minutes. The linearity of the instrument for CO₂ measurements has been checked by applying the linear fit equation of the CO₂ concentration of the calibration standards (350.67 ppm, 399.68 ppm and 426.20 ppm) measured by the analyzer. The slope is found in the range of 0.99 - 1.007 ppm with a correlation coefficient (r) of about 0.999. Further, linearity of the instrument for CO is also checked by diluting the calibration mixture from 970 ppb to 100 ppb. The calibration mixture is diluted in a glass made dead volume using highly precise and accurate pure air (air free from water vapour, particles, carbon monoxide (CO), sulphur dioxide (SO₂), oxides of nitrogen (NOx), ozone (O₃), and hydrocarbons (HC)) from ECO Physics generator and two mass flow controller. For increasing the interaction time of both the gases (zero air and calibration mixture) and to insure a homogeneous mixing, the spring shaped dead volume is used. Each diluted mixture is passed for 30 minutes in the instrument and average of last 10 minute data are used for the test. The instrument shows excellent linearity for CO and slope is observed to be 0.98. The accuracy of the measurements is calculated by subtracting the mean difference of measured CO2 and CO concentration from the actual concentration of both gases in target gas. The accuracies of CO₂ and CO are found in the range of 0.05-0.2 ppm and 0.01-0.025 ppm respectively. The repeatability of both gases are calculated by the standard deviation of the mean concentration of target gas measured by the analyser over the period of observations and found 0.3 ppm and 0.04 ppm for CO₂ and CO respectively."

Section 4

p.32195

This part are interesting but too long.

Response: We have shorten this part after removing some sentences and rearranging them. The modified text is given following.

"Figures 3a and 3c show the time series of 30 minute average CO₂ and CO concentrations for the period of November, 2013 - February, 2014 and July, 2014 to May, 2015. Large and periodic variations indicate the stronger diurnal dependence of both gases. Overall, the concentrations and variability of both gases are observed lowest in the month of July and August, while maximum scatter in the concentrations and several plumes of very high levels both gases have been observed from October, 2014 until mid-March 2015. Almost all plumes of CO₂ and CO are one to one correlated and are found mostly during evening rush hours and late nights. Figures 3e and 3f show the variations of CO₂ and CO concentrations with wind speed and direction for the study period except July, August and September due to non-availability of

wind data. Most of the high and low concentrations of both gases are found to be associated with low and high wind speeds. There is no specific direction for high levels of these gases. This probably indicates the transport sector is an important contributor to the local emissions since the measurement site is surrounded by city roads".

Lines 23-26: please reformulate

Response: We have reformulate the sentences, which are given below.

"Figures 3b and 3d show the probability distributions or frequency distributions of CO₂ and CO concentration during the study period. Both gases show different distributions from each other. This difference could be attributed due to the additional role of biospheric cycle (photosynthesis and respiration) on the levels of CO₂, apart from the common controlling factors (local sources, regional transport, PBL dynamics etc) responsible for distributions of both gases. The control of the boundary layer is common for the diurnal variations of these species because of their chemical lifetimes are longer (> months) than the timescale of PBL height variations (~ hrs)".

p.32196

Lines 22-23: what is the demonstration for this argument?

Response: We have modified the sentence. The modified part is given below.

"The seasonal cycle from first approach will present the overall variability in both gases. On the other hand, second approach removes the auto-covariance by excluding CO₂ and CO data mainly affected by local emission sources and represent seasonal cycles at the well mixed volume of the atmosphere".

p.32197

Remove lines 1-2 ----- Removed.

Lines 9-11: not clear ---- As mentioned in general comments, this section is synthesise and hence these lines are removed.

Lines 11-18: synthesise ---- As mentioned in previous comments also, after synthesising the respective section, some of mentioned lines are rearranged, while some are removed.

Lines 17-20: how much of the data coverage does this step represent? ----- Mean of every month ~24 hrs*30 = 720 hrs data.

p.32200

Lines 1-4: reformulate ------ As per previous suggestion for synthesizing this section, these sentences

are already modified. The modified text is given below.

"CO2 concentrations start decreasing from 0000 to 0300 hrs and increases slightly afterwards till 0600-

0700 hrs during summer and autumn. Respiration of CO₂ from the vegetation is mostly responsible for

this night time increase. During winter and spring seasons CO₂ levels are observed constant during night

hours and small increase is observed only from 0600 to 0800 hrs during the winter season. In contrary to

this, subsequent section shows a continuous decline in the night time concentrations of the main

anthropogenic tracer CO, which indicates that there is enough vertical mixing of low CO air from above

that once the CO source is turned off, its concentration drops. Hence, constant level of CO₂ at night hours

during these seasons give the evidence of a continued but weak sources (such as respiration) in order to

offset dilution of mixing of low CO2 air from aloft. Dry soil conditions could be one of the possible cause

for weak respirations."

p.32202

Lines: 15-20: reformulate

Response: We have reformulate the text by the following lines.

"The first difference is observed in the timing of the occurrence of morning peaks: CO₂ peaks occur slightly

before the CO peak due to the triggering photosynthesis process by the sunrise. On the other hand, the

morning peaks of CO mostly depend on the rush hour traffic and are consistent at 0800-1000 hrs in all

seasons. The second difference is that the afternoon concentrations of CO show little seasonal spread as

compared to the afternoon concentrations of CO₂. Again, this is due to the biospheric control on the levels

of CO₂ during the afternoon hours of different seasons while CO levels are mainly controlled by the dilution

during these hours".

p.32203

Define the baseline and background terms.

Response: Both words are used for same purpose: "the least affected levels of CO2 from local sources".

But since there may be some misunderstanding from different terminology, we have removed baseline

word from the text. We have included the following lines for defining the background.

26

"The measurements are generally affected by the dilution due to the boundary layer dynamics, but considering their ratios will cancel this effect. Further, the interpretation of correlation ratios in terms of their dominant emission sources needs to isolate first the local urban signal. For this, the measurements have to be corrected from their background influence. The background concentrations are generally those levels which have almost negligible influence from the local emission sources".

How sensitive is your 5th percentile method? This was for example assessed against MACC fields in Ammoura et al ACPD 2015 (a new method...). Give clues.

Response: We have added following lines for clarity about using 5th percentile method for calculating the background.

"It is observed that the mixing ratios of both gases at low wind speed, which show the influence of local urban signal, are significantly higher than background levels and hence confirm that the definition of background will not significantly affect the derived ratios (Ammoura et al., 2015). This technique of measuring the background is extensively studied by Ammoura et al (2015) and found suitable for both the gases CO and CO₂, even having the role of summer uptake on the levels of CO₂".

p.32204

The role of cooking (poor combustion => large co/co2), other FF sources etc should be considered here. Response: We have modified whole section according to previous comments and Referee #1 comments.

p.32205

Be careful here at hours when the PBLH evolves (see general)

Lines 17-20: this is critical. I do not agree with your argumentation. Table 1 does not show that the observed ratios (30-50 ppb/ppm) are much lower than the domestic sources (52.99 ppb/ppm). You cannot conclude that this is driven by gasoline emissions. And several solutions exist. You could have a mix of emissions from traffic and domestic sources for example. At what time do people have diner in Ahmedabad? Same time than rush hours or not? Etc. This section needs to be thought more and the different options argued to drive to a solid conclusion.

Response: The average dinner time is 1900-2100 hrs. So these ratios could be influenced by the mix emission sources such as vehicular emissions, domestic sources. Hence rather than fossil fuel emission

now we have included anthropogenic emission. We have added following explanations in the modified text.

"Except monsoon, the Δ CO $_{(exc)}$ / Δ CO2 $_{(exc)}$ ratios and their correlations are fairly comparable in other seasons in the evening rush hours, which indicate stronger influence of common emission sources. Ratios during this time can be considered as fresh emissions since dilution and chemical loss of CO can be considered negligible for this time. Most of these data fall in the domestic and transport sector emission ratio lines, which indicate that during this time intervals these sources mostly dominate (Table 1). On the other hand, during other time intervals most of the data are scattered between industrial and transport sectors emission ratio lines. Hence, from this we can conclude that during evening hours, transport and domestic sources mostly dominate while during other periods transport and industrial emission sources mostly dominate".

p.32206

Remind the question of the entrainment pb in the morning for example (check my general comments).

Response: We have already discussed the modified sentence previously according the general comments.

p.32207

It would be interesting to try to give an explanation about this. What emissions is EDGAR missing then? Is it a sector or is it underestimated on all sectors? What about emissions from slum /residential cooking for example? You might found this paper interesting on the CO emissions from New Delhi: http://aaqr.org/VOL15 No3 June2015/36 AAQR-14-07-TN0132 1137-1144.pdf

Response: Thank you very much for the reference and wonderful suggestion. Accordingly we have added following explanation in the text.

"The EDGAR inventory estimate the relative contributions of CO from industrial, transport and slum/residential sectors to be about 42%, 42% and 10%, respectively. The possible cause for underestimation of CO by the EDGAR inventory could be the underestimation of residential emission, inventory since other particularly for major urban Indian cities (http://www.indiaenvironmentportal.org.in/files/file/Air-Pollution-in-Six-Indian-Cities.pdf) show large relative contributions from residential sector. The uncertainty associated with the emission factors for different sectors could be another cause for the underestimation of CO emissions, since these are important parameters for developing the inventory (Sahu et al., 2015)".

Lines 14-15: following my remarks above, I do not agree with your argument on the large role you attribute

to CO emissions from fossil fuels incomplete combustion only. Other sectors are still on the race as long

as you did not demonstrate the contrary.

Response: We have modified the sentence. We replace the fossil fuel term from the anthropogenic

emission. Accordingly text is modified.

"CO has virtually no natural source in an urban environments except oxidation of hydrocarbons and hence

can help to disentangle the relative contributions of anthropogenic (from transport, power plant, industrial

etc) and the biospheric sources (mainly from respiration) of CO₂, by serving as a tracer of combustion

activity on shorter timescale (Duren and Miller, 2012)."

Lines 27-28: this was not clearly demonstrated as well.

Response: We have modified/removed the text by following lines.

"Figure 9a shows the excess diurnal variations of CO2 above the background levels during different

seasons. The observed concentrations of both gases can also be directly used for calculating the emission

ratio, provided the measured levels are not highly affected from natural sources as well as share the same

origin. We have used the evening time (1900-2100 hrs) data of CO_{2exc} and CO_{exc} for whole study period

to calculate the emission ratio of CO/CO₂ from predominant anthropogenic emission sources, since the

correlation (r = 0.95) for this period is very high and can hence, be considered that the levels of both gases

for this period are mostly affected by same types of anthropogenic sources. Also there can be considered

negligible contribution of biospheric sources in CO₂".

p.32208

Lines 4-6: very surprising, aren't people cooking at this time? 47 ppb/ppm is more than gasoline and in

between gasoline and biofuels/coal.

Response: We have modified the sentence. As discussed previously also, this ratio is mostly dominating

by vehicular and domestic fuel emissions. So we have included anthropogenic emission source in place

of gasoline and biofuels.

Line 13: same, no solid argumentation given for this

Response: We have removed the sentence from the text.

29

Lines 26-27: I do not agree here as well. I do not think this is true to say that the other sources do not emit

CO. What about wood burning, cooking etc. again. These are not natural but anthropogenic.

Response: This sentence is already modified as per previous suggestions.

p. 32209

Lines 1-13: this part should be fully rewritten according to the comments above.

Response: We have modified this section by adding following lines in the text.

"The average diurnal cycles of CO₂ above its background for each season are shown in (Figure 9a). In Section 4.3.1, we have discussed qualitatively the role of different sources in the diurnal cycle of CO₂. With the help of the above method, now the contributions of anthropogenic (CO_{2 (ant)}) and biospheric sources (CO_{2 (bio)}) are discussed quantitatively. Due to unviability of PBL measurements, we cannot disentangle the contributions of boundary layer dynamics. The diurnal pattern of CO_{2 (ant)} (Figure 9c) reflects the pattern like CO, because we are using constant Rco/co2 (ant) for all seasons. Overall, this analysis suggests that the anthropogenic emissions of CO₂ mostly from transport and industrial sectors during early morning during 0600-1000 hrs varied from 15 to 60% (4-15 ppm). During afternoon hours (1100-1700 hrs), the anthropogenic originated (transport and industrial sources, mainly) CO₂ varied between 20 and 70% (1-11 ppm). During evening rush hours (1800-2200 hrs), highest contributions of combined emissions of anthropogenic sources (mainly transport and domestic) are observed. During this period the contributions vary from 50 to 95% (2-44 ppm. During night/early morning hours (0000-0700 hrs) non-anthropogenic sources (mostly biospheric respiration) contribute from 8 to 41 ppm of CO₂ (Figure 9d). The highest contributions from 18 to 41 ppm are observed in the autumn from the respiration sources during night hours, since there is more biomass during this season after the South Asian summer monsoon. During the afternoon hours, lower biospheric component of CO₂ could be due to a combination of the effects of afternoon anthropogenic emissions, biospheric uptake of CO₂ and higher PBL height".

Line 19: rewrite with 1 time « component », not 3 times; remove « diurnal amplitude»; add « s » to

observation.

Response: Done as per suggestion.

Lines 24-26: give reference

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Response: We have included the following reference.

Law, R. M., Peters, W., Rödenbeck, C., Aulagnier, C., Baker, I., Bergmann, D. J., Bousquet, P., Brandt, J., Bruhwiler, L., Cameron-Smith, P. J., Christensen, J. H., Delage, F., Denning, A. S., Fan, S., Geels, C., Houweling, S., Imasu, R., Karstens, U., Kawa, S. R., Kleist, J., Krol, M. C., Lin, S.-J., Lokupitiya, R., Maki, T., Maksyutov, S., Niwa, Y., Onishi, R., Parazoo, N., Patra, P. K., Pieterse, G., Rivier, L., Satoh, M., Serrar, S., Taguchi, S., Takigawa, M., Vautard, R., Vermeulen, A. T., and Zhu, Z.: TransCom model simulations of hourly atmospheric CO₂: Experimental overview and diurnal cycle results for 2002, Global Biogeochemical Cycles, 22, doi:10.1029/2007GB003050, gB3009, 2008.

p.32211

Line 23: this part is vague about the tracers you used. Please give more clues to the reader.

Response: We have modified the explanations by adding informations in detail and following reference.

"Separate correlations of each CO₂ tracer with the observations are helpful to determine relative importance of each flux component in the CO₂ variation (Patra et al., 2008). Hence we separately studied the correlation between biospheric component of CO₂ and anthropogenic component of CO₂, estimated by model using CASA-3hr fluxes and EDGAR v4.2 inventory, and the measurements. The correlation coefficients give the hint about major controlling factors on the levels of CO₂. Figure 11b shows the resulting correlations for separate flux component with respect to measurements".

Patra, P. K., Law, R. M., Peters, W., Rödenbeck, C., Takigawa, M., Aulagnier, C., Baker, I., Bergmann, D. J., Bousquet, P., Brandt, J., Bruhwiler, L., Cameron-Smith, P. J., Christensen, J. H., Delage, F., Denning, A. S., Fan, S., Geels, C., Houweling, S., Imasu, R., Karstens, U., Kawa, S. R., Kleist, J., Krol, M. C., Lin, S.-J., Lokupitiya, R., Maki, T., Maksyutov, S., Niwa, Y., Onishi, R., Parazoo, N., Pieterse, G., Rivier, L., Satoh, M., Serrar, S., Taguchi, S., Vautard, R., Vermeulen, A. T., and Zhu, Z.: TransCom model simulations of hourly atmospheric CO₂: Analysis of synoptic-scale variations for the period 2002–2003, Global Biogeochemical Cycles, 22, doi:10.1029/2007GB003081, gB4013, 2008.

p.32212

5 Conclusions

Lines 2-4 Transition with CO should be improved.Improved

Lines 7-10: yes, this is an excellent remark!Modified

Lines 20: here again you mix seasons and climate features. Please modify. ------Done

p.32214

Line 2: remove (fossil fuel) unless you manage to demonstrate it correctly......

Response: We have modified the text after including the following lines according to the change made in the respective sections previously.

"The availability of simultaneous and continuous measurements of CO₂ and CO have made it possible to study their correlations at different time windows (during morning (0600-1000 hrs), noon (1100-1700 hrs), evening (1800-2200) and night (0000-0500 hrs) hours) of distinct seasons. Using the correlation slopes and comparing them with the emission ratios of different sources, contributions of distinct sources is discussed qualitatively. It is observed that during the evening hours, measurements over the study region are mostly affected by transport and domestic sources, while during other periods the levels of both gases are mostly dominated by the emissions from transport and industrial sources".

Lines 7-9: These lines should be changed (see remarks p.32208)Changed

Technical corrections:

Define ppm as part per millionDefined in the beginning of text

CO should be expressed in ppb (to be defined) and not in ppm, for consistency with the tables.

Do not mix noon (12h) and afternoon. Changes

Add « the » before model or before CASA model. You can also replace model by modeling framework to avoid repetitions.Done

Title:

« In CO₂ and CO » should be replaced by « of atmospheric ».Replaced

The short title « CO₂ over urban region » requires as well the keywords CO and India.

Abstract

Line 25: replace variations by ones......Replaced

Introduction

P.32157 Line 18: remove (such as the Kyoto Protocol)Removed

Section 2

P.32191 line 9/10: already said p.32190 line 17.This sentence is removed.

Section 3

Lines 12-26: check English style please (First... the second...).....We have checked the style and suitably modified the text wherever needed. The modified text is the following.

"The ambient measurements of CO₂ and CO concentrations have been made using the wavelength scanned cavity ring down spectroscopic (CRDS) technique based analyser (Picarro-G2401) at 0.5 Hz. CRDS offers highly sensitive and précised measurements of trace gases in the ambient air, due to its three main characteristics (Bitter et al., 2005; Chen et al., 2010; Karion et al., 2013). (1) Longer interaction path length (around 20 km). (2) Low operating pressure (~140 Torr) of cell to avoid pressure broadening which ensures that the peak height or area is linearly proportional to the concentration. (3) Decay time of light intensity is measured instead of the absorption of light intensity, making it independent of the fluctuation in the incident light intensity. "

P.32194 Line 23/24: remove for the model simulationsDone

Section 4

P.32196 Line 19: In the first approachDone

(Fig.4a)... 5 ConclusionsDone

p.32212

Line 12: remove « The unique flow of ».Removed

Line 20: is « transported » the right term?

Response: Thank you very much for pointing this. The appropriate term is emitted. We have removed the sentence "representing CO2 and CO transported from anthropogenic sources" for more clear explanation.

Lines 23-24: remove « seasonal... season) » and replace by activity (do not explain such process in your

conclusions)We have modified the sentence accordingly

p.32213

Line 12: replace effects of by undergoReplaced

p.32215

Line 3: validity is a bit strong... You could say we assessed independently, which is more neutral.

.....Done

Tables

Table 4: the legend does not seem appropriate to the table. The legend has been changed by the following text.

"Seasonal mean concentrations and diurnal amplitudes (max-min) of CO2 and CO over Ahmedabad."

Figures

Figure 1: « from the EDGAR... » Corrected

Date: 26 April 2016

Temporal variations in of atmospheric CO₂ and CO at Ahmedabad in western India

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Correspondence to: Shyam Lal (shyam@prl.res.in)

Abstract. About 70% of the anthropogenic carbon dioxide (CO₂) is emitted from the mega-cities and urban areas of the world. In order to draw effective emission mitigation policies for combating future climate change as well as independently validate the emission inventories for constraining their large range of uncertainties, especially over major metropolitan areas of developing countries, there is an urgent need for greenhouse gases measurements over representative urban regions. India is a fast developing country, where fossil fuel emissions have increased dramatically in the last three decades and predicted further to continue to grow by at least 6% per year through 2025. In the absence of systematic CO₂ measurements over the Indian urban locations, CO₂ along with an anthropogenic emission tracer carbon monoxide (CO) are being measured at Ahmedabad, a major urban site in western India, using a state-of-the-art laser based cavity ring down spectroscopy technique from November 2013 to May 2015 with a break during March to June 2014. These measurements enable us to understand the diurnal and seasonal variation in atmospheric CO₂with respect to its sources (both anthropogenic and biospheric) and biospheric sinks. In-situ simultaneous measurements of carbon dioxide (CO₂) and carbon monoxide (CO) have been made using a state-of-the-art laser based cavity ring 15 down spectroscopy technique at Ahmedabad, an urban site in western India, from November 2013 to May 2015 with a break during March to June 2014. These measurements enable us to understand the diurnal and seasonal variation in atmospheric CO₂ with respect to its anthropogenic and biospheric sources and sinks. The observed annual average concentrations of CO₂ and CO have been found to be are 413.0±13.7 ppm and 0.50±0.37 ppm respectively. Both the species CO₂ and CO show strong seasonality, with lower concentrations (400.3±6.8 ppm and 0.19±0.13 ppm), respectively during the south-west monsoon and higher values of concentrations (419.6 \pm 22.8 ppm and 0.72 \pm 0.68 ppm), respectively in during the autumn (SON) season. Strong diurnal variations are also observed for both

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the species. The common factors for diurnal cycles of CO_2 and CO are the vertical mixing and rush hour traffic, while the influence of biospheric fluxes is also seen in CO_2 diurnal cycle. Using CO and CO_2 covariation, we differentiate the anthropogenic and biospheric components of CO_2 and found that significant contributions of biospheric respiration and anthropogenic emission in the late night (0000 - 0500 hrs) and evening rush hours (1800-2200 hrs) respectively. We compute total yearly emission of CO to be 69.2 ± 0.07 Gg for the study region using the observed $CO:CO_2$ correlation slope and bottom-up CO_2 emission inventory. This calculated emission of CO is 52% larger than the estimated emission of CO by the EDGAR inventory. The observations of CO_2 have been compared with an atmospheric chemistry transport model (i.e., ACTM), which incorporates various components of CO_2 fluxes. ACTM is able to capture the basic variabilities, but both diurnal and seasonal amplitudes are largely underestimated compared to the observations. We attribute this underestimation by model to uncertainties in terrestrial biosphere fluxes and coarse model resolution. The fossil fuel signal from the model shows fairly good correlation with observed CO_2 variations, which supports the overall dominance of fossil fuel emissions over the biospheric fluxes in this urban region.

1 Introduction

Carbon dioxide (CO₂) is the most important anthropogenically emitted greenhouse gas (GHG) and has increased substantially from 278 ppm to 390 parts per million (ppm) in the atmosphere since the beginning of the industrial era (circa 1750). It has contributed to more than 65% of the radiative forcing increase since 1750 and hence leads to thehas significant impact on the climate system (Ciais et al., 2013). Major causes of CO2 increase are anthropogenic emissions, especially fossil fuel combustion, cement production and land use change. The cumulative anthropogenic CO₂ emissions from the preindustrial era to 2011, are estimated to be 545±85 PgC, out of which fossil fuel combustion and cement production contributed 365±30 PgC and land use change (including deforestation, afforestation and reforestation) contributed 180±80 PgC. Land and oceans are the two important sinks of atmospheric CO₂, which remove about half of the anthropogenic emissions (Le Quéré et al., 2014). The prediction of future climate change and its feedback rely mostly on our ability to quantifying fluxes of greenhouse gases, especially CO₂ at regional (100-1000 km²) and global scale. Though the global fluxes of CO₂ can be estimated fairly well, the regional scale (e.g. sub-continent and country level) fluxes are associated with quite high uncertainty especially over the South Asian region; and the estimation uncertainty being larger than the value itself (Patra et al., 2013; Peylin et al., 2013). Detailed scientific understanding of the flux distributions is also needed for formulating effective mitigation policies. (such as Kyoto Protocol). In inverse modelling, CO2 flux is estimated from atmospheric CO₂ observations and using an atmospheric transport model. Therefore, it is necessary to measure CO₂ concentrations covering different ecosystems and geographical areas

Along with the need of atmospheric measurements for predicting the future levels of CO₂, quantifying the components of anthropogenic emissions of CO2 is similarly important for providing independent verification of mitigation strategies as well as understanding the biospheric component of CO₂. Alone CO₂ measurements could not be helpful in making such study due to the large role of biospheric fluxes in its atmospheric distributions. The proposed strategy for quantification of the anthropogenic component of CO2 emissions is to measure simultaneously the anthropogenic tracers (Duren and Miller, 2012). Although, carbon monoxide (CO) is not a direct GHG but it affects elimate and air quality through the formation of CO_2 and ozone (O_3) . It affects the oxidizing capacity of atmosphere through reaction with the free OH radicals. Additionally, CO can be used as a surrogate tracer for detecting and quantifying anthropogenic emissions from burning processes, since it is a major product of incomplete combustion (Turnbull et al., 2006; Wang et al., 2010; Duren and Miller, 2012). The vehicular as well as industrial emissions contribute large fluxes of CO₂ and CO to the atmosphere in urban regions. The verification of future mitigation activities demand for quantifying the spatiotemporal distribution of these emissions. The CO emissions have large uncertainty as compared to CO₂, because its emission strongly depends on the combustion efficiency, the vehicle engine and their adopted technology as well as driving conditions. The correlation slope between the atmospheric variations of CO and CO₂ can be used to quantify the fossil fuel contribution, distinguish between different burning processes or to determine the burning efficiency and overall trend of anthropogenic emissions of CO in that city. The CO:CO2 ratios are higher for low combustion sources (e.g. forest fires) and lower for good or efficient combustion sources. Further, the CO:CO₂ ratio can be used for estimating the total emission of CO over an urban area provided the total CO₂ emission is known for that area. Hence, the information about CO:CO₂ ratio will be helpful to understand the effects on the CO emissions after adopting the newer vehicular technologies and new cleaner emission norms and finally will be beneficial for reducing the uncertainties in CO emission inventories. Several grounds based and aircraft based eorrelation simultaneous studies of CO; and CO₂ have been done in the past from different parts of the world (Turnbull et al., 2006; Wunch et al., 2009; Wang et al., 2010; Newman et al., 2013) but such study has not been done in India except recently reported results from weekly samples for three Indian sites by Lin et al. (2015).

Measurements at different regions (eg., rural, remote, urban) and at different frequency (eg., weekly, daily, hourly etc) have their own advantage and limitations. For example, the measurements at remote locations at weekly interval can be useful for studying seasonal cycle, growth rate, and estimating the regional carbon sources and sinks after combining their concentrations with inverse modelling and atmospheric tracer transport models. However, some important studies, like diurnal variations, temporal covariance, etc are not possible from these measurements. Analysis of temporal

covariance of atmospheric mixing processes and variation of flux on shorter time scales, e.g., subdaily, is essential for understanding local to urban scale CO₂ flux variations (Ahmadov et al., 2007; Pérez-Landa et al., 2007; Briber et al., 2013; Lopez et al., 2013; Ammoura et al., 2014; Ballav et al., 2015). Urban regions contribute about 70% of global CO₂ emission from anthropogenic sources and further projected to increase over the coming decades (Duren and Miller, 2012). Therefore, measurements from these regions are very helpful for understanding emissions, growth as well as verifying the mitigation policies. The first observations of CO2, CO and other greenhouse gases started in February 1993 from Cape Rama (CRI: a coastal site) on the south-west coast of India using flask samples (Bhattacharya et al., 2009). After that, several other groups have initiated the measurements of surface level greenhouse gases (Mahesh et al., 2014; Sharma et al., 2014; Tiwari et al., 2014; Lin et al., 2015). Most of these measurements are made at weekly or fortnightly time intervals. Two aircraft based measurement programs, namely, Civil Aircraft for the regular Investigation of the atmosphere Based on an Instrument Container (CARIBIC) (Brenninkmeijer et al., 2007) and Comprehensive Observation Network for TRace gases by AIrLiner (CONTRAIL) (Machida et al., 2008) have provided important first look on the South Asian CO₂ budget, but these data have their own limitations (Patra et al., 2011; Schuck et al., 2010, 2012). It is pertinent to mention here that till now there are no reports of CO₂measurements over the urban locations of the Indian subcontinent, which could be an important player in the global carbon budget as well as mitigation purpose due to strong growing anthropogenic activities specifically fast growing traffic sector and sinks (large areas of forests and croplands). Hence, the present study is an attempt to reduce this gap by understanding the CO₂ mixing ratios in light of its sources and sinks at an urban region in India.

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India is the second largest populous country in the world having about 1.3 billion inhabitants. Rapid socioeconomic developments and urbanization have made it the third largest CO₂ emitter next to China and USA since 2011 but it ranks at 137th level based on the per capita emission rate of CO₂ (EDGAR v4.2; CDIAC - . For example, in 2010 India's emission rate was 2.2 ton CO₂ eq/capita while the developed countries like USA, Russia and UK had emission rates of about 21.6, 17.6 and 10.9 ton CO₂ eq/capita respectively (EDGAR v4.2). The budgets of these gases on regional as well as global scales can be estimated by bottom-up and top down approaches. Large uncertainties are associated in the GHGs budgets over South Asia, especially over India than for other continents. Based on the atmospheric CO₂ inversion using model calculations, found that the biosphere in South Asia acted as the net CO₂ sink during 2007-2008 and estimated CO₂ flux of about -104±150 TgC yr⁻¹. Further, based on the bottom-up approach, gave an estimate of biospheric flux of CO₂ of about -191±193 TgC yr⁻¹ for the period of 2000-2009. Both these approaches show the range of uncertainty 100-150%. One of the major sources of these large uncertainties is the lack of spatial and temporal observations of these gases . The first observations of CO₂, CO and other greenhouse gases started in February 1993 from Cape Rama (CRI) on the south-west coast of India

using flask samples . After that, several other groups have initiated the measurements of surface level greenhouse gases . Most of these measurements are made at weakly or fortnightly time intervals or at lower accuracy. These data are very useful for several studies like analyzing seasonal cycle, growth rate, and estimating the regional (subcontinental) carbon sources and sinks after combining their concentrations with inverse modelling and atmospheric tracer transport models. However, some important studies like their diurnal variations, temporal covariance...etc are not possible from these measurements due to their limitations. Analysis of temporal covariance of atmospheric mixing processes and variation of flux on shorter timescales, e.g., sub-daily, is essential for understanding local to urban scale CO₂ flux variations. Two aircraft based measurements programs namely, Civil Aircraft for the regular Investigation of the atmosphere Based on an Instrument Container (CARIBIC) and Comprehensive Observation Network for TRace gases by AIrLiner (CONTRAIL) have provided important first look on the South Asian CO₂ budget, but these data have their own limitations . The focus of the Indigenous research is lacking in terms of making the continuous and simultaneous measurements of CO₂ and CO over the urban regions, where variety of emission sources influence the level of these gases.

In the view of above, Ssimultaneous continuous measurements of CO_2 and CO have been made since November 2013 from an urban site Ahmedabad located in the western India using very highly sensitive laser based technique. The preliminary results of these measurements for one month period have been reported in (Lal et al., 2015). These detailed measurements are utilized for studying the temporal variations (diurnal and seasonal) of both gases, their emissions characteristics on diurnal and seasonal scale using their mutual correlations, estimating the contribution of vehicular anthropogenic and biospheric emission components in the diurnal cycle of CO_2 using the ratios of CO to CO_2 and rough estimate of the annual CO emissions from study region. Finally, the measurements of CO_2 have been compared with simulations using an atmospheric chemistry-transport model to discuss roles of various processes contributing to CO_2 concentrations variations.

2 Site description, local emission sources and meteorology

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The measurement facility is housed inside the campus of the Physical Research Laboratory (PRL), situated in the western part of Ahmedabad (23.03°N, 72.55°E, 55 m AMSL) in the Gujarat state of India (Fig.1). It is a semi-arid, urban region in western India, having a variety of large and small scale industries (Textile mills and pharmaceutical production facilities) in east and north outskirts. The institute is situated about 15-20 km away from these industrial areas and surrounded by tree on all sides. The western part is dominated by the residential areas. The city has a population of about 5.6 million (Census India, 2011) and has large number of automobiles (about 3.2 million), which are increasing at the rate of about 10% per year. Most of the city buses and auto-rickshaws (three-wheelers)

use compressed natural gas (CNG) as a fuel. The transport-related activities are the major contributors of various pollutants (Mallik et al., 2015). An emission inventory for this city, which is developed for all known sources, shows the annual emissions (for year 2010) of CO₂ and CO about 22.4 million tons and 707,000 tons respectively (http://www.indiaenvironmentportal.org.in/files/file/Air-Pollution-in-Six-Indian-Cities.pdf). Out of these emissions, transport sector contributes about 36% in CO₂ emission and 25% in CO emissions, power plants contribute about 32% in CO₂ emissions and 30% in CO emission, industries contribute about 18% in CO₂ emissions and 12% in CO emissions and domestic sources contribute about 6% in CO₂ emissions and 22% in CO emissions. The Indo-Gangetic Plain (IGP) is situated in the northeast of Ahmedabad, which is very densely populated region and has high levels of pollutants emitted from various industries and power plants along with anthropogenic emissions from burning of fossil fuels and traditional biofuels (wood, cow-dung cake etc). The Thar Desert and the Arabian Sea are situated in the northwest and southwest of Ahmedabad respectively.

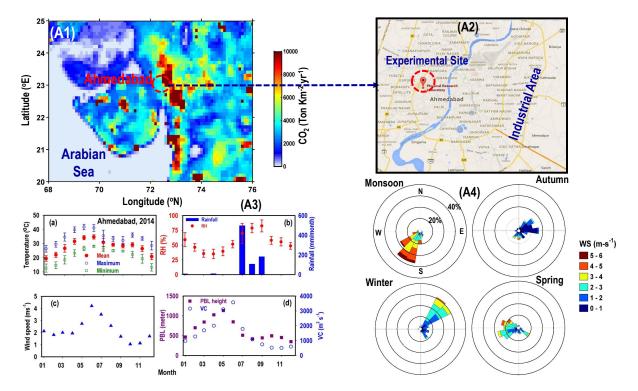


Figure 1: (A1) Spatial distribution of total anthropogenic CO₂ emissions from the EDGARv4.2 inventory over Ahmedabad and surrounding regions. (A2) The Ahmedabad city map showing location of the experimental site (PRL).(A3: a-d) Monthly average temperature with monthly maximum and minimum value, relative humidity (RH), rainfall, wind speed, PBL height and ventilation coefficient (VC) over Ahmedabad during the year 2014. Temperature, RH and wind speed are taken from Wunderground weather (www.wunderground.com) while rainfall and PBLH data are used from Tropical Rainfall Measuring Mission (TRMM) satellite and MEERA reanalysis data. (A4) Wind rose plots for Ahmedabad for the four seasons of 2014 using daily average data from Wunderground.

Figure 1shows average monthly variability of temperature, relative humidity (RH), wind speed based on data taken from Wunderground (http://www.wunderground.com), rainfall from Tropical Rainfall Measuring Mission (TRMM) and planetary boundary layer height (PBLH) from the Modern-Era Retrospective Analysis for Research and Applications (MEERA). The wind rose plot shows the surface level wind speed and direction during different seasons over Ahmedabad in 2014. This place is known for its semi-arid climate. Large seasonal variations are observed in the wind speed and direction over Ahmedabad. During monsoon the summer monsoon season (June-July-August), the inter-tropical convergence zone (ITCZ) moves northward across India. It results in the transport of moist and cleaner marine air from the Arabian Sea and the Indian Ocean to the study location by south westerly winds, or the so-called southwest monsoon (summer monsoon). The first shower due to the onset of the southwest monsoon occurs in July and it retreats in the mid of September over

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Ahmedabad. Due to heavy rain and winds mostly from oceanic region, RH shows higher values in July, August and September. Highest RH of about 83% is observed in September. The long-range transport of air masses from the northeast part of the Asian continent starts to prevail over Indian region when ITCZ moves back southward in September and October. These months are regarded as transition period for the monsoon. During autumn (September-October-November), the winds are calm and undergo a change in their direction from south west to north east. When the transition of winds takes placed from oceanic to continental region in October, the air gets dryer and RH decreases until December. The winds are north easterly during winter (December- January - February) and transport pollutants mostly from continental region (IGP region). During pre-monsoon the spring season (March - April - May), winds are north westerly and little south westerly which transport mixed air masses from continent and oceanic regions. The average wind speed is observed higher in June and July while lower in October and March when transition of wind starts from oceanic to continental and continental to oceanic respectively. The monthly averaged temperature starts increasing from January and attains maximum $(34.6\pm1.4^{\circ}\text{C})$ in June, followed by a decrease until September and temperature is slightly warmer in October compared to the adjacent months. The monthly variation in planetary boundary layer height (PBLH) closely resembles with the temperature pattern. Maximum PBLH of about 1130 m is found in June and it remains in the lower range at about 500 m during July to January. The ventilation coefficient (VC) is obtained by multiplying wind speed and PBL height which gradually increases from January to attain the maximum value in June and the lowest values of VC are observed in October and November.

210 3 Experiment and model details

3.1 Experimental method

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The measurements of ambient measurements of CO₂ and CO are performed concentrations have been made using a Picarro-G2401 instrument, which is based on the wavelength scanned cavity ring down spectroscopic (CRDS) technique based analyser (Picarro-G2401) at 0.5 Hz. CRDS offersis a now a well-established technique for making highly sensitivety and high precision precised measurements of trace gases in the ambient air, due to its three main characteristics (Bitter et al., 2005; Chen et al., 2010; Karion et al., 2013). (1). It provides very long Longer interaction path length (around 20 km) between the sample and the incident wavelength, by utilizing a 3-mirror configuration, which enhances its sensitivity over other conventional techniques like Non-dispersive Infrared Spectroscopy (NDIR) and Fourier Transform Infrared Spectroscopy (FTIR). The second is its ability to (2). Low operating pressure (140 Torr) of cell to avoid pressure broadening The operating low pressure (140 Torr) of cell allows to isolate a single spectral feature with a resolution of 0.0003 cm⁻¹, which ensures that the peak height or area is linearly proportional to the concentration. The third advantage is that (3) Decay time of light intensity is measured instead of

the absorption of light intensity, making it independent of the fluctuation in the incident light intensity. The measurements of trace gases using this technique are achieved by measuring the decay time of light intensity inside the cavity while the conventional optical absorption spectroscopy technique is based on absorption of light intensity. Hence, it increases the accuracy of measurements because it is insensitive to the fluctuations of incident light. The cell temperature of analyzer was maintained at 45°C throughout the study period. The precision and accuracy of these measurements follow the WMO compatibility goals of ±0.1 ppm CO₂ and ±2 ppb CO.

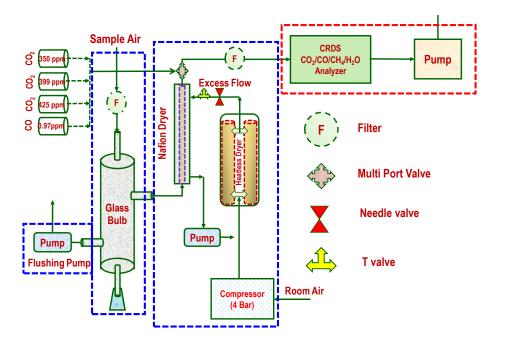


Figure 2: Schematic diagram of the experimental set-up analysis system. We introduce additionally a Naffion dryer in upstream of the inlet of instrument for removing the water vapour from the ambient air. Three calibration mixtures (three) from NOAA, USA are used to calibrate CO₂ measurements and one calibration mixture from Linde, UK is used to calibrate CO measurements. The red color box covers the analyser system received from the company, while two blue colours box cover the two stage moisture removing systems designed at our lab in PRL.

Figure 2 shows the schematic diagram of the measurement system experimental set-up, which consists an analyser system, a glass bulb, a Nafion dryer, a heat less dryer, other associated pumps and a set of calibration mixtures. Atmospheric air is sampled continuously from the terrace of the building (250 meter above the ground level) through 1/4 inch PFA Teflon tube via a glass manifold. An external pump is attached on one side of the glass manifold to flush the sample line. Water vapor affects the measurements of CO_2 by diluting its mole fractions in the air and by broadening

the spectroscopic absorption lines of other gases. Although, the instrument has ability to correct for the water vapour interferes by using an experimentally derived water vapor correction algorithms (Crosson, 2008), but it has an absolute H_2O uncertainty of $\sim 1\%$ (Chen et al., 2010) and can introduces a source of error using a single water vapor correction algorithm (Welp et al., 2013). This error can be minimized by either generating the correction coefficients periodically in the laboratory or by removing the water vapour from the sample air. Conducting the water vapor correction experiment is bit tricky and need extra care as discussed by Welp et al. (2013). Hence, we prefer to remove water vapour from the sample air by introducing a 50-strand Nafion dryer (Perma Pure, p/n PD-50T-24MSS) in the upstream of the analyser. Nafion dryer contains a bunch of semi-permeable membrane tubing separating an internal sample gas stream from a counter sheath flow of dry gas in stainless steel outer shell. The partial pressure of water vapour in the sheath air should be lower than the sample air for effectively removing the water vapour from the sample air. A heatless dryer generates dry air using a 4 bar compressor (KNF, MODEL: NO35ATE) which is used as a seath flow in Naffion dryer. After drying, sample air passes through the PTFE filter (polytetrafluoroethylene; 5μ m Sartorius AG, Germany) before entering the instrument cavity. This setup dries the ambient air near to 0.03% (300 ppm) concentration of H_2O .

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The measurement system is equipped with three high pressure aluminum cylinders containing gas mixtures of CO_2 (350.67 \pm 0.02, 399.68 \pm 0.01 and 426.20 \pm 0.01 ppm) in dry air from NOAA, Bolder USA, and one cylinder of CO (970 parts per billion (ppb)) from Linde UK. These tanks were used to calibrate the instrument for CO₂ and CO. An additional gas standard tank (CO₂: 338 ppm, CO: 700 ppb), known as the "target", is used to monitor the instrumental drift and to assess the dataset accuracy and repeatability. The target tank values are calibrated against the CO2 and CO calibration mixtures. The target tank and calibration gases were measured mostly in the mid of every month (Each calibration gas is passed for 30 minute and target tank for 60 minutes). The target gas is introduced into the instrument for a period of 24 hours also once in a six month, for checking the diurnal variability of instrument drift. Maximum drift for 24 hours has been calculated by subtracting the maximum and minimum value of 5 minute average, which were found to be 0.2 ppm and 0.015 ppm respectively for CO₂ and CO. For all calibration mixtures, the measured concentration is calculated as the average of the last 10 minutes. The linearity of the instrument for CO₂ measurements has been checked by applying the linear fit equation of the CO₂ concentration of the calibration standards (350.67 ppm, 399.68 ppm and 426.20 ppm) measured by the analyzer. The slope is found in the range of 0.99 - 1.007 ppm with a correlation coefficient (r) of about 0.999. Further, linearity of the instrument for CO is also checked by diluting the calibration mixture from 970 ppb to 100 ppb. The calibration mixture is diluted in a glass made dead volume using highly precise and accurate pure air (air free from water vapor, particles, carbon monoxide (CO), sulphur dioxide (SO₂), oxides of nitrogen (NOx), ozone (O₃), and hydrocarbons (HC)) from ECO

Physics generator. and two mass flow controllers. The flows of calibration mixture and pure air were regulated using two separate mass flow controllers from Aalborg. For increasing the interaction time of both the gases (zero air and calibration mixture) and to insure a homogeneous mixing, the spring shaped dead volume is used. Each diluted mixture is passed for 30 minutes in the instrument and average of last 10 minute data are used for the test. The instrument shows excellent linearity for CO and slope is observed to be 0.98. The accuracy of the measurements is calculated by subtracting the mean difference of measured CO_2 and CO concentration from the actual concentration of both gases in target gas. The accuracies of CO_2 and CO are found in the range of 0.05-0.2 ppm and 0.01-0.025 ppm respectively. The repeatability of both gases are calculated by the standard deviation of the mean concentration of target gas measured by the analyser over the period of observations and found 0.3 ppm and 0.04 ppm for CO_2 and CO respectively.

The CO_2 concentrations are reported on the WMO scale, using the three calibration mixtures of CO_2 (350.67±0.02, 399.68±0.002 and 426.20±0.006 ppm) from NOAA, Bolder USA, while the concentration of CO is reported against a calibration mixture of CO (970 ppb) from Linde UK. An additional gas standard tank (CO_2 : 338 ppm, CO: 700 ppm), known as the "target", is used to determine the precision of the instrument. The target tank values are calibrated against the CO_2 and CO calibration mixtures. The target gas is introduced in the instrument for a period of 24 hours. For CO_2 and CO, the 5 minute precisions were found to be 0.015 ppm and 0.005 ppm respectively within 1σ . Maximum drift for 24 hours has been calculated by subtracting the maximum and minimum value of 5 minute average which were found to be 0.2 ppm and 0.015 ppm respectively for CO_2 and CO. The linearity of the instrument for CO_2 measurements has been checked by using three calibration standards (350.67 ppm, 399.68 ppm and 426.20 ppm) of CO_2 . The linearity tests are conducted very frequently and the slope is found in the range of 0.99 – 1.007 ppm with correlation coefficient (r) of about 0.999.

3.2 Description of AGCM-based Chemistry Transport Model (ACTM)

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This study uses the Center for Climate System Research/National Institute for Environmental Studies/Frontier Research Center for Global Change (CCSR/NIES/FRCGC) atmospheric general circulation model (AGCM)-based chemistry-transport model (ACTM). The model is nudged with reanalysis meteorology using Newtonian relaxation method. The U and V components of horizontal winds are used from the Japan Meteorological Agency Reanalysis (JRA-25) (Onogi et al., 2007). The model has $1.125^{\circ} \times 1.125^{\circ}$ horizontal resolution (T106 spectral truncation) and 32 vertical sigma-pressure layers up to about 50 km. Three components namely anthropogenic emissions, monthly varying ocean exchange with net uptake and terrestrial biospheric exchange of surface CO₂ fluxes are used in the model. The fossil fuel emissions for the model simulations are taken from EDGAR inventory for the year of 2010. Air-sea fluxes from Takahashi et al. (2009) have been used for the oceanic CO₂ tracer. The oceanic fluxes are monthly and are linearly interpolated between mid-months. The

terrestrial biospheric CO₂ tracers are provided from the Carnegie-Ames-Stanford-Approach (CASA) process model (Randerson et al., 1997), after introducing a diurnal variability using 2 m air temperature and surface short wave radiation from the JRA-25 as per Olsen and Randerson (2004). The ACTM simulations has been extensively used in TransCom CO₂ model inter-comparison studies (Law et al., 2008; Patra et al., 2008).

4 Results and Discussion

4.1 Time series and general statistics

Figures 3a and 3c show the time series of 30 minute average CO₂ and CO concentrations for the period of November, 2013 - February, 2014 and July, 2014 to May, 2015. The concentrations of both gases exhibit ILarge and periodic variations indicate stronger diurnal dependence of both the gases.synoptic variability because the site is close to anthropogenic sources. Overall, the concentrations and variability of both the gases are observed lowest in the month of July and August, maximum scatter in the concentrations and several plumes of very high levels both gases have been observed from October, 2014 until mid-March 2015. Almost all plumes of CO₂ and CO are one to one correlated and are mostly found during evening rush hours and late nights. Figures 3e and 3f show the variations of CO₂ and CO concentrations with wind speed and direction for the study period except July, August and September due to non-availability of wind data. Most of the high and low concentrations of both these gases are found to be associated with low and high wind speeds. There is no specific direction for high levels of these gases. This probably indicates the transport sector is an important contributor to the local emissions since the measurement site is surrounded by the city roads.

Figures 3b and 3d show the probability distributions or frequency distributions of CO₂ and CO concentrations during the study period. Both gases show different distributions from each other. This difference could be attributed to the additional role of biospheric cycle (photosynthesis and respiration) on the levels of CO₂ apart from the common controlling factors (local sources, regional transport, PBL dynamics etc) responsible for distributions of both gases. The frequency distribution of CO₂ shows almost normal distribution while CO shows skewed towards right (lower concentrations). This is because, natural cycle of the biosphere (photosynthesis and respiration) along with some common controlling factors (local meteorology and anthropogenic sources), affects significantly the levels of CO₂. The control of the boundary layer is common for the diurnal variations of these species because of their chemical lifetimes are longer (> months) than the timescale of PBL height variations (~ hrs). However, biospheric fluxes of CO₂ can have strong hourly variations. During the study period the CO₂ concentrations varied between 382 - 609 ppm, with 16% of data lying below 400 ppm, 50% lying in the range 400-420 ppm, 25% between 420-440 ppm and 9% in the range of

440-570 ppm. Maximum frequency of CO_2 is observed at 402.5 ppm during the study period. The CO concentrations lies in the range of 0.071-8.8 ppm with almost 8% data lies below the most probable frequency of CO at 0.2 ppm, while 70% data lies between the concentrations of 0.21 ppm and 0.55 ppm. Only 8% data lies above the concentration of 1.6 ppm and rest of 14% data lies between 0.55 and 1.6 ppm. The annual mean concentrations of CO_2 and CO are found to be 413.0±13.7 ppm and 0.50±0.37 ppm respectively, after removing outliers beyond 2σ values.

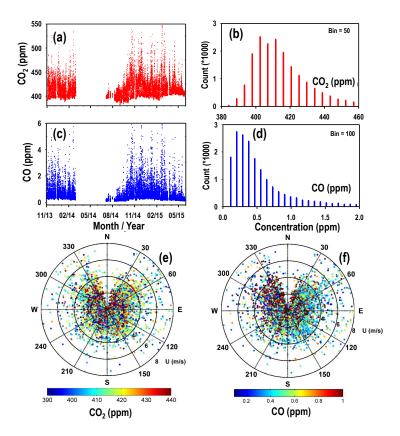


Figure 3: (a and c) Time series of 30 minute average values CO_2 and CO measured at Ahmedabad for the study period. (b and d) The frequency distribution in CO_2 and CO concentrations for the study period using 30 minute mean of both gases. (e and f) The polar plots show the variation of 30 minute averaged CO_2 and CO at this site with wind direction and speed during the study period except July, August and September due to unavailability of meteorology data.

4.2 Seasonal variations of CO₂ and CO

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The seasonal cycles of CO_2 and CO are mostly governed by the strength of emission sources, sinks and transport patterns. Although they follow almost identical seasonal patterns but the factors responsible for their seasonal behaviors are distinct as for the diurnal variations. We calculate the seasonal cycle of CO_2 and CO using two different approaches. In the first approach, we use the monthly

mean of all data measurements and in the second approach we use monthly mean for of the afternoon period (1200-1600 hrs) measurements only. The seasonal cycle from first approach will depiet the combined influence of local emissions (mostly) as well as that of large scale circulation present the overall variability in both gases. On the other hand, the second approach removes the autocovariance by excluding CO_2 and CO data mainly affected by local emission sources and represent seasonal cycles at the well mixed volume of the atmosphere. The CO_2 time series is de-trended by subtracting a mean growth rate of CO_2 observed at Mauna Loa (MLO), Hawaii, i.e., 2.13 ppm yr⁻¹ or 0.177 ppm/month (www.esrl.noaa.gov/gmd/ccgg/trends/) for clearly depicting the seasonal cycle amplitude. Figure 4a and Figure 4b show the variations of monthly average concentrations of CO_2 and CO using all daily (0-24 hrs) data and afternoon (1200-1600 hrs) data.

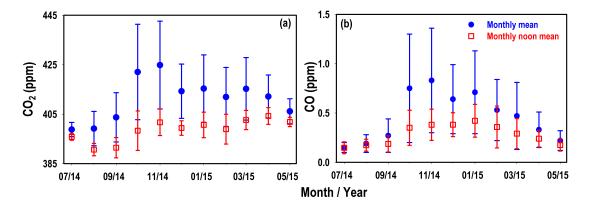


Figure 4: The seasonal variation of CO_2 and CO from July, 2014 to May, 2015 using their monthly mean concentrations. The blue dots and red rectangles show the monthly average concentrations of these gases for the total (0-24 hrs) and noon time (1200-1600 hrs) data respectively with 1σ spread.

Both average concentrations (total and noon time) of CO_2 exhibit strong seasonal cycle, but show distinct patterns (occurrence of maxima and minima) to each other. This difference occurs because seasonal cycle of CO_2 from all data is mostly governed by the PBL ventilation and large scale circulation while the seasonal cycle from noon time mean concentration is mostly related to the seasonality of vegetation activity. The total and noon time mean concentrations of CO show almost similar pattern and evince that the seasonal cycle of CO_2 from the afternoon mean is mostly controlled by the biospheric productivity, since biospheric cycle does not influence CO concentration directly. In general, total mean values of CO_2 and CO are observed lower in July having concentration 398.78 \pm 2.8 ppm and 0.15 \pm 0.05 ppm respectively. During summer monsoon months predominance of south-westerly winds which bring cleaner air from the Arabian Sea and the Indian Ocean over to Ahmedabad and high VC (Figure 1) are mostly responsible for the lower concentration of total mean of both the gases. CO_2 and CO concentrations are also at their seasonal low in the northern hemisphere due to net biospheric uptake and seasonally high chemical loss by reaction with

OH, respectively. In addition, deep convections in the summer monsoon season efficiently transport the Indian emission (for CO, hydrocarbons) or uptake (for CO₂) signals at the surface to the upper troposphere, resulting lower concentrations at the surface in the summer compared to the winter months (Kar et al., 2004; Randel and Park, 2006; Park et al., 2009; Patra et al., 2011; Baker et al., 2012). During autumn and early winter (December), lower VC values cause trapping of anthropogenically emitted CO₂ and CO and is the major cause for high concentrations of both gases during this period. In addition to this, wind changes from cleaner marine region to polluted continental region, especially from IGP region and hence could be additional factor for higher levels of CO₂ and CO during these seasons (autumn and winter). Elevated levels during these seasons are also examined in several other pollutants over Ahmedabad as discussed in previous studies (Sahu and Lal, 2006; Mallik et al., 2016). Maximum concentrations of CO₂ and CO are observed to be 424.8±17 ppm and 0.83±0.53 ppm, respectively, during November. From January to May the total mean concentration of CO₂ decreases from 415.3±13.6 to 406.1±5.0 ppm and total mean concentration of CO decreases from 0.71 ± 0.22 to 0.22 ± 0.10 ppm. Higher VC and predominance of comparatively less polluted mixed air masses from oceanic and continental region results in lower total mean concentrations of both gases during this period. A sudden increase in the total mean of both gases is observed from September to October and maximum concentrations of CO2 and CO are observed to be 424.8±17 ppm and 0.83±0.53 ppm, respectively, during November. From January to May the total mean concentration of CO_2 decreases from 415.3 ± 13.6 to 406.1 ± 5.0 ppm and total mean concentration of CO decreases from 0.71 ± 0.22 to 0.22 ± 0.10 ppm. During monsoon months predominance of south-westerly winds which bring cleaner air from the Arabian Sea and the Indian Ocean over to Ahmedabad and high VC (Figure 1) are responsible for the lower concentration of total mean of both the gases. CO₂ and CO concentrations are also at their seasonal low in the northern hemisphere due to net biospheric uptake and seasonally high chemical loss by reaction with OH, respectively. In addition, deep convections in the southwest monsoon season efficiently transport the Indian emission (for CO, hydrocarbons) or uptake (for CO₂) signals at the surface to the upper troposphere, resulting lower concentrations at the surface in the summer compared to the winter months. During autumn and early winter (December), lower VC values cause trapping of anthropogenically emitted CO₂ and CO. This is the major cause for high CO₂ and CO concentrations during this period. The north-easterly winds bring very high levels of pollutants from IGP region and could additionally enhance the levels of CO₂ and CO during these seasons (autumn and winter). Higher VC and predominance of comparatively less polluted mixed air masses from oceanic and continental region results in the lower total mean concentrations of both gases.

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There are some clear differences which are observed in the afternoon mean concentrations of CO_2 as compared to daily mean. The first distinct feature is that significant difference of about 5 ppm is observed in the afternoon mean of CO_2 concentration from July to August as compared to the dif-

ference in total mean concentration about \sim 0.38 ppm for the same period. Significant differences in the afternoon concentrations of CO₂ from July to August is mainly due to the increasing sink by net biospheric productivity after the Indian summer monsoonal rainfall. Another distinct feature is that the daily mean concentration of CO₂ is found highest in November while the afternoon mean concentration of CO₂ attains maximum value (406 ± 0.4 ppm) in April. Prolonged dry season combined with high daytime temperature (about 41° C) during April-May make the tendency of ecosystem to become moderate source of carbon exchange (Patra et al., 2011) and this could be responsible for the elevated mean noon time concentrations of CO₂. Unlike CO₂, seasonal patterns of CO from total and afternoon mean concentrations are identical, although levels are different. It shows that the concentrations of CO is mostly governed by identical sources during day and night time through out the year.

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430 The average amplitude (max - min) of the annual cycle of CO₂ is observed around 13.6 and 26.07 ppm from the afternoon mean and total mean respectively. Different annual cycles and amplitudes have been observed from other studies conducted over different Indian stations. Similar to our observations of the afternoon mean concentrations of CO2, maximum values are also observed in April at Pondicherry (PON) and Port Blair (PBL) with amplitude of mean seasonal cycles about 7.6±1.4 and 11.1±1.3 ppm respectively (Lin et al., 2015). Cape Rama (CRI), a costal site on the south-west coast 435 of India show the seasonal maxima one month before than our observations in March annual amplitude about 9 ppm (Bhattacharya et al., 2009). The Sinhagad (SNG) site located over the Western Ghats Mountains, show very larger seasonal cycle with annual amplitude about 20 ppm (Tiwari et al., 2014). The amplitude of mean annual cycle at the free tropospheric site Hanle at altitude of 4500 m is observed to be 8.2±0.4 ppm, with maxima in early May and the minima in mid-September (Lin et al., 2015). Distinct seasonal amplitudes and patterns are due to differences in regional controlling factors for the seasonal cycle of CO₂ over these locations, e.g., the Hanle is remotely located from all continental sources, Port Blair site is sampling predominantly marine air, Cape Rama observes marine air in the summer and Indian flux signals in the winter, and Sinhagad represents a forested ecosystem. These comparisons show the need for CO₂ measurements over different ecosystems for constraining its budget.

The annual amplitude in afternoon and daily mean CO concentration is observed to be about 0.27 and 0.68 ppm, respectively. The mean annual cycles of CO over PON and PBL show the maxima in the winter months and minima in monsoon summer months same as our observations with annual amplitudes of 0.078 ± 0.01 and 0.144 ± 0.016 ppm, respectively. So the seasonal levels of CO are affected by large scale dynamics which changes air masses from marine to continental and vice versa and by photochemistry. The amplitudes of annual cycle at these locations differ due to their climatic conditions and sources/sinks strengths.

455 4.3 Diurnal variation

The diurnal patterns for all months and seasons are produced by first generating the time series from the 15 min averages and then averaging the individual hours for all days of the respective month and season after removing the values beyond 2σ - standard deviations for each month as outliers.

4.3.1 Diurnal variation of CO₂

Figure 5a shows the mean diurnal cycles of atmospheric CO_2 and associated 1- σ standard deviation 460 (shaded region) during all the four seasons. All times are in Indian Standard Time (IST), which is 5.5 hrs ahead of Greenwich Mean Time (GMT)the Universal Time (UT). Noticeable differences are observed in the diurnal cycle of CO2 from season to season. In general, maximum concentration has been observed during morning (0700-0800 hrs) and evening (1800-2000 hrs) hours, when the ABL is shallow, traffic is dense and vegetation respiration dominate due to absence of photosynthe-465 sis activity. The minimum of the cycles occurred in the afternoon hours (1400-1600 hrs), when the PBL is deepest and well mixed as well as when the vegetation photosynthesis is active. There are many interesting features in the period of 0000-0800 hrs. CO₂ concentrations start decreasing from 0000 to 0300 hrs and increases slightly afterwards till 0600-0700 hrs during summer and autumn. Respiration of CO₂ from the vegetation is mostly responsible for this night time increase. During 470 winter and spring seasons CO2 levels are observed constant during night hours and small increase is observed only from 0600 to 0800 hrs during the winter season. In contrary to this, subsequent section shows a continuous decline in the night time concentrations of the main anthropogenic tracer CO, which indicates that there is enough vertical mixing of low CO air from above once the CO source is turned off, its concentration drops. Hence, constant levels of CO₂ at night hours during these seasons give the evidence of a continued but weak sources (such as respiration) in order to offset dilution of mixing of low CO₂ air from aloft. Dry soil conditions could be one of the possible causes for weak respiration. Further, distinct timings have been observed in the morning peak of CO₂ during different seasons. It is mostly related to the sunrise time, which decides the evolution time of PBL height and beginning of vegetation photosynthesis. The sunrise occurs at 0555-0620 hrs, 0620-0700 hrs, 0700-480 0723 hrs and 0720-0554 hrs during summer, autumn, winter and spring respectively. During spring and summer, rush hour starts after sunrise, so the vehicular emissions occur when the PBL has been already high and photosynthetic activity has begun. The CO₂ concentration is observed lowest in the the morning during the summer monsoon season as compared to other seasons. This is because CO₂ 485 uptake by active vegetation deplete the entire mixed layer during day time and when the residual layer mixes to the surface in the morning, low-CO2 air is mixed down. In winter and autumn, rush hour starts parallel with the sunrise, so the emissions occur when the PBL is low and concentration

builds up is much stronger in these seasons than in spring and summer monsoon seasons.

The diurnal amplitude is defined as the difference between the maximum and minimum concentrations of CO₂ in the diurnal cycle. The amplitudes of monthly averaged diurnal cycle of CO₂ from July, 2014 to May 2015 are shown in Figure 5b. The diurnal amplitude shows large month to month variation with increasing trend from July to October and decreasing trend from October onwards. The lowest diurnal amplitude of about 6 ppm is observed in July while highest amplitude about 51 ppm is observed in October. The amplitude does not change largely from December to March and is observed in the range of 25-30 ppm. Similarly from April to May the amplitude also varies in a narrow range from 12 to 15 ppm. The jump in the amplitude of the CO₂ diurnal cycle is observed highest (around 208%) from July to August. This is mainly due to significant increase of biospheric productivity from July to August after the rains in Ahmedabad. It is observed that during July the noon time CO₂ levels are found in the range of 394-397 ppm while in August the noon time levels are observed in the range of 382-393 ppm. The lower levels could be due to the higher PBL height during afternoon and cleaner air, but in case of CO (will be discussed in next section), average day time levels in August are observed higher than in July. It rules out that the lower levels during August are due to the higher PBL height and presence of cleaner marine air, and confirms the higher biospheric productivity during August.

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shows bi-modal feature in the diurnal cycle of CO₂ during the four seasons with morning and evening peaks. Both peaks are associated mostly with the vehicular emissions and PBL height during rush hours. There are many interesting features in the 0000-0800 hrs period. All times are in Indian Standard Time (IST), which is 5.5 hrs ahead of GMT. Concentrations of CO₂ start decreasing from 0000 hrs to 0300 hrs and afterwards increases until 0600 and 0700 hrs during monsoon and autumn. It could be mostly due to the accumulation of CO₂ emitted from respiration by the biosphere in the nocturnal boundary layer. During winter and spring the concentrations during night hours are almost constant and increase is observed only from 0600 hrs to 0800 hrs during winter. Dormant of respiration during these two seasons due to lower temperature could be one of the possible factors for no increase in CO₂ concentrations during night. No peak during morning hours is observed in spring. Distinct timings for the occurrence of the morning peak during different seasons is generally related to the sunrise time and consequently the evolutions of PBL height. The sunrise occur at 0555-0620 hrs, 0620-0700 hrs, 0700-0723 hrs and 0720-0554 hrs during monsoon, autumn, winter and spring, respectively. During spring and monsoon, rush hour starts after sunrise, so the vehicular emissions occur when the PBL is already high and photosynthetic activity has begun. But in winter and autumn rush hour starts parallely with the sunrise, so the emissions occur when the PBL is low and concentration build up is much strong in these seasons than in spring and monsoon seasons. CO₂ starts decreasing fast after these hours and attains minimum value around 1600 hrs. This quick drop of CO₂ after sunrise is linked to the dominance of photosynthesis over the respiration processes in addition to the higher atmospheric mixing height. CO₂ levels start increasing after 1600 hrs peak around 2100 hrs. Higher concentrations of CO₂ during these hours are mainly due to the rush hour

vehicular emissions and less dilution due to the lower PBL height. Comparative levels of CO₂ during evening rush hours except monsoon confirm separately the major influence from the same type of sources (vehicular emission) in its levels which do not show large variability as in post-midnight hours.

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The diurnal amplitude is defined as the difference between the maximum and minimum concentrations of CO₂ in the diurnal cycle. The amplitudes of monthly averaged diurnal cycle of CO₂ from July, 2014 to May 2015 are shown in Figure 5b. The diurnal amplitude shows large month to month variation with increasing trend from July to October and decreasing trend from October onwards. Lowest diurnal amplitude of about 6 ppm is observed in July while highest amplitude about 51 ppm is observed in October. The amplitude does not change much from December to March and is observed in the range of 25-30 ppm. Similarly from April to May the amplitude also varies in a narrow range from 12 to 15 ppm. The jump in the amplitude of CO₂ diurnal cycle is observed highest (around 208%) from July to August. This is mainly due to significant increase of biospheric productivity from July to August after the rains in Ahmedabad. It is observed that during July the noon time CO₂ levels are found in the range of 394-397 ppm while in August the noon time levels are observed in the range of 382-393 ppm. The lower levels could be due to the higher PBL height during afternoon and cleaner air, but in case of CO, average day time levels in August are observed higher than July. It rules out that the lower levels during August are due to the higher PBL height and presence of cleaner marine air, and confirms the higher biospheric productivity during August.

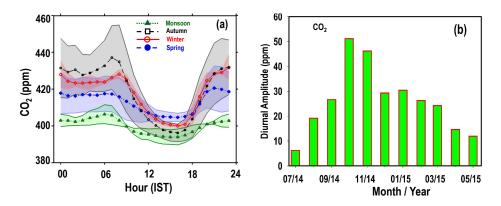


Figure 5: (a) Average diurnal variation of CO_2 over Ahmedabad during all the four seasons. (b) Monthly variation of average diurnal amplitude of CO_2 during from July, 2014 to May, 2015. All times are in Indian Standard Time (IST), which is 5.5 hrs ahead of Universal Time (UT).

The monthly average diurnal cycles of the biospheric net primary productivity from the CASA model for Ahmedabad and for the year of 2014 are shown Figure 10. The details of CASA flux are given in the Section 3.2. It is observed that the model shows higher biospheric productivity in September and October while the observations are suggesting higher productivity in August. This

indicates that the CASA model is not able to capture the signal of higher biospheric productivity for Ahmedabad and need to be improved. Similar discrepancy in the timing of maximum biospheric uptake is also discussed earlier by Patra et al. (2011) using inverse model CO₂ fluxes and CASA biospheric fluxes.

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Near surface diurnal amplitude of CO_2 has been also documented in humid subtropical Indian station Dehradun and a dry tropical Indian station Gadanki (Sharma et al., 2014). In comparison to Ahmedabad, both these stations show distinct seasonal change in the diurnal amplitude of CO_2 . The maximum CO_2 diurnal amplitude of about 69 ppm is observed during the monsoon summer season at Dehradun (30.3°N, 78.0°E, 435m), whereas maximum of about 50 ppm during autumn at Gadanki (13.5°N, 79.2°E, 360 m).

4.3.2 Diurnal variation of CO

Figure 6a shows seasonally averaged diurnal variation of CO. In general, the mean diurnal cycle of CO show lower concentration during noon (1200-1700 hrs) and two peaks in the morning (0800 to 1000 hrs) and in the evening (1800 to 2200 hrs) hours. This cycle exhibits the same pattern as the mean diurnal cycle of traffic flow, with maxima in the morning and at the end of the afternoon, which suggests the influence of traffic emissions on CO measurements. Along with the traffic flow, PBL dynamics also play a critical role in governing the diurnal cycle of CO. The amplitudes of the evening peak in diurnal cycles of CO are always greater than the morning peaks. It is because the PBL height evolves side by side with the morning rush hours traffic and hence increased dilution, while on the other hand, during evening hours, PBL height decrease along with evening dense traffic and favors accumulation of pollutants until the late evening under the stable PBL conditions. The noon time minimum of the cycle is mostly associated with the deepest and well mixed PBL.In general, the average diurnal cycle patterns of both gases (CO₂ and CO) are similar, but having few noticeable differences. The first difference is observed in the timing of the occurrence of morning peaks: CO₂ peaks occur slightly before the CO peak due to the triggering photosynthesis process by the sunrise. On the other hand, the morning peaks of CO mostly depend on the rush hour traffic and are consistent at 0800-1000 hrs in all seasons. The second difference is that the afternoon concentrations of CO show little seasonal spread as compared to the afternoon concentrations of CO₂. Again, this is due to the biospheric control on the levels of CO₂ during the afternoon hours of different seasons while CO levels are mainly controlled by the dilution during these hours. The third noticeable difference is that the levels of CO decrease very fast after evening rush hours in all the seasons while this feature is not observed in the case of CO₂ since respiration during night hours contributes to the levels of CO₂. The continuous drop of nighttime concentrations of CO indicates that there is enough vertical mixing of low CO air from above once the CO source is turned off. The average morning (0800-0900 hrs) peak values of CO is observed minimum (0.18±0.1 ppm) in summer and maximum

 $(0.72\pm0.16 \text{ ppm})$ in winter while evening peak shows minimum value $(0.34\pm0.14 \text{ ppm})$ in summer and maximum $(1.6\pm0.74 \text{ ppm})$ in autumn. The changes in CO concentrations show large fluctuations from morning peak to afternoon minima and from afternoon minima to evening peak. From early morning maxima to noon minima, the changes in CO concentrations are found in the range of 20 -200%, while from noon minima to late evening maxima the changes in CO concentrations are found in the range of 85% to 680%. Similar diurnal variations with two peaks have also been observed in earlier measurements of CO as well as NOx at this site (Lal et al., 2000).

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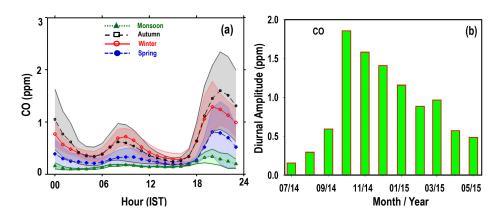


Figure 6: (a) Diurnal variation of CO over Ahmedabad during all the four seasons. (b) Monthly variation of the diurnal amplitude of CO.

shows seasonally averaged diurnal variation of CO. In general, the mean diurnal cycles of CO during all the seasons show lower concentration during noon (1200-1700 hrs) and two peaks, one in the morning (0800 to 1000 hrs) and other in the evening (1800 to 2200 hrs). This cycle exhibits the same pattern as the mean diurnal cycle of traffic flow, with maxima in the morning and at the end of the afternoon, which suggests the influence of traffic emissions on CO measurements. Along with the traffic flow, PBL dynamics also plays a critical role in governing the diurnal cycle of CO. The amplitudes of the evening peaks in diurnal cycles of CO are always greater than the morning peaks. It is because the PBL height evolve side by side with the morning rush hours traffic and hence increased dilution while during evening hours PBL height decrease along with evening time rush hours traffic and favors accumulation of pollutants until the late evening under the stable PBL conditions. The noon time minima is associated with the combined influence of boundary layer dilution and loss of CO due to OH radicals. The peaks during morning and evening rush hours, minima during afternoon hours in CO diurnal cycle during all seasons are similar as in CO₂. However, there are a few noticeable differences in the diurnal cycles of both the gases. The first noticeable difference is that the CO morning peak appears later than CO₂ peak. This is because as discussed earlier with sunrise time, PBL height starts evolve and same time photosynthesis process also get started and hence CO₂ morning peak depends on the sunshine time. But in case of CO,

timing of the morning peak mostly depends on the rush hour traffic and is consistent at 0800-1000 hrs in all seasons. The second noticeable difference is the afternoon concentrations of CO show little seasonal spread as compared to the afternoon concentrations of CO₂. Again, this is due to the biospheric control on the concentration of CO₂ during the afternoon hours of different seasons while CO levels are mainly controlled by the dilution during these afternoon hours. The third noticeable difference is that the levels of CO decrease very fast after evening rush hour in all seasons while this feature is not observed in case of CO₂ since respiration during night hours contributes to the levels of CO₂. The average morning (0800-0900 hrs) peak values of CO are observed minimum (0.18±0.1 ppm) in monsoon and maximum (0.72±0.16 ppm) in winter while its evening peak shows minimum value $(0.34\pm0.14 \text{ ppm})$ in monsoon and maximum $(1.6\pm0.74 \text{ ppm})$ in autumn. The changes in CO concentrations show large fluctuations from morning peak to afternoon minima and from afternoon minima to evening peak. From early morning maxima to noon minima, the changes in CO concentrations are found in the range of 20 -200% while from noon minima to late evening maxima the changes in CO concentrations are found in the range of 85% to 680%. Similar diurnal variations with two peaks have also been observed in earlier measurements of CO as well as NOx at this site.

The evening peak contributes significantly to the diurnal amplitude of CO. The largest amplitude in CO cycle is observed in autumn (1.36 ppm) while the smallest amplitude is observed in summer (0.24 ppm). The diurnal amplitudes of CO are observed to be about 1.01 and 0.62 ppm, respectively during winter and spring. Like CO₂, the diurnal cycle of CO (Figure 6b) shows the minimum (0.156 ppm) amplitude in July and maximum (1.85 ppm) in October. After October the diurnal amplitude keeps on decreasing till summer. The evening peak contributes significantly to the diurnal amplitude of CO. The largest amplitude in CO cycle is observed in autumn (1.36 ppm) while the smallest amplitude is observed in summer (0.24 ppm). The diurnal amplitudes of CO are observed to be about 1.01 and 0.62 ppm, respectively during winter and spring. The monthly diurnal cycle of CO (Figure 6b) shows the minimum (0.156 ppm) amplitude in July and maximum (1.85 ppm) in October. After October the diurnal amplitude keeps on decreasing till summer monsoon. The monthly diurnal cycle of CO (Figure 6b) shows minimum (0.156 ppm) amplitude in July and maximum (1.85 ppm) in October. After October the diurnal amplitude keep on decreasing till monsoon. The evening peak contributes significantly to the diurnal amplitude of CO. The largest amplitude in CO cycle is observed in autumn (1.36 ppm) while smallest amplitude is observed in monsoon (0.24 ppm).

4.4 Correlation between CO and CO₂

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The relationships of CO to CO₂ can be useful for investigating the CO source types and their combustion characteristics in the city region of Ahmedabad. The measurements are generally affected by the dilution due to the boundary layer dynamics, but considering their ratios will cancel this

effect. Further, the interpretation of correlation ratios in terms of their dominant emission sources needs to isolate first the local urban signal. For this, the measurements have to be corrected from their background influence. The background concentrations are generally those levels which have almost negligible influence from the local emission sources. For correlations study, in principle the baseline levels to be removed from the measured concentrations. Although, the most ideal case of determining the background levels are the continuous measurement of respective gases at a cleaner site. But due to unavailability of such measurements for our study period at cleaner sites, we use the 5^{th} percentile value of CO_2 and CO for each day as a background for corresponding day. It is observed that the mixing ratios of both gases at low wind speed, which show the influence of local urban signal, are significantly higher than background levels and hence confirm that the definition of background will not significantly affect the derived ratios (Ammoura et al., 2014). This technique of measuring the background is extensively studied by Ammoura et al. (2014) and found suitable for both the gases CO and CO₂, even having the role of summer uptake on the levels of CO₂. The excess CO_2 ($CO_{2(exc)}$) and CO ($CO_{(exc)}$) above the background for Ahmedabad city, are determined for each day after subtracting the background concentrations from the hours of each day $(CO_{2(exc)})$ $= CO_{2(obs)} - CO_{2(bg)}, CO_{(exc)} = CO_{(obs)} - CO_{(bg)}$.

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665 We use robust regression method for the correlation study. It is an alternative to least squares regression method and more applicable for analysing time series data with outliers arising from extreme events (http://www.ats.ucla.edu/stat/stata/dae/rreg.htm). Figure 7 illustrates the correlations between $CO_{(exc)}$ and $CO_{2(exc)}$ for the four seasons at different time windows of the day. Based on the dominance of different atmospheric processes and different emission sources as discussed in Section 4.3, the measurements are divided into the group of four different time windows: (1) Morning period (0600-1000 hrs), when PBL height is slowly evolving and rush hour traffic is there, (2) Afternoon period (1100-1700 hrs), when atmospheric is well mixed and traffic is relatively less, (3) Evening period (1800-2200 hrs), when influence of rush hour traffic is significantly high, (4) Night period (0000-0500 hrs), when atmospheric is calm and the anthropogenic sources of both gases are switched off. The measured slopes values for these time intervals are given in Table 1. The ranges of the emission ratios of CO/CO₂ for transport, industrial and domestic sources, as given in Table 2, are also plotted in the figures for broadly showing the dominance of different sources. The $\Delta CO_{(exc)}/\Delta CO_{2(exc)}$ ratios are observed lowest during summer, with a range varying from 0.9 ppb/ppm in morning to 19.5 ppb/ppm in evening period. Lowest coefficient of determination is also observed during this season, which suggest that the levels of CO and CO2 are controlled by 680 different factors. As discussed previously, higher biospheric productivity during this season mostly controls the CO₂ concentrations while CO concentrations are mostly controlled by the long range transport. During the winter season $\Delta CO_{(exc)}/\Delta CO_{2(exc)}$ ratios are observed highest and varies from 14.3 ppb/ppm in morning to 47.2 ppb/ppm in evening period. Relatively higher ratios dur-

ing winter than other three seasons indicates contribution of CO emission from additional biofuel 685 burning sources. From day to night, highest coefficient of determination is observed during spring season. As illustrated by the diurnal cycle, the CO2 is not significantly removed by the biosphere during spring with lower draw down in daily CO2. Along with this, higher VC during this season will result in very fast mixing. Therefore, very fast mixing will mostly regulate their relative varia-690 tion and will result in higher correlation in this season. Other factors like soil and plant respiration during this period may also control CO₂ concentrations due to which the correlation coefficient is not equal to 1. Except monsoon, the $\Delta CO_{(exc)}/\Delta CO_{2(exc)}$ ratios and their correlations are fairly comparable in other seasons in the evening rush hours, which indicate stronger influence of common emission sources. Ratios during this time can be considered as fresh emissions since dilution and chemical loss of CO can be considered negligible for this time. Most of these data fall in the 695 domestic and transport sector emission ratio lines, which indicate that during this time intervals these sources mostly dominate (Table 2). On the other hand, during other time intervals most of the data are scattered between industrial and transport sectors emission ratio lines. Hence, from this we can conclude that during evening hours, transport and domestic sources mostly dominate while during other periods transport and industrial emission sources mostly dominate. The impact of the possible sources of CO and CO₂ varies from month to month and hence season to season. The lowest correlation (r = 0.62, p =0.0001) is observed during monsoon, with a $\Delta CO_{(exc)}/\Delta CO_{2(exc)}$ ratio of 0.6±0.1 ppb/ppm. Lowest correlation suggest that different mechanisms control the levels of CO and CO₂ during the monsoon season. As discussed previously, higher biospheric productivity during this season mostly controls the CO₂ concentrations while CO concentrations are mostly controlled by the long range transport and higher loss due to OH. Highest correlation (r = 0.87, p<0.0001) with ΔCO_(exc)/ΔCO_{2(exc)} ratio of 8.4±0.17 ppb/ppm is observed during spring season. As illustrated by the diurnal cycle, the CO₂ is not significantly removed by the biosphere during spring with lower draw down in daily CO₂. Along withthis, higher VC during this season will result in very fast mixing. Therefore, very fast mixing will mostly regulate their relative variation and will result in higher correlation in this season. Other factors like soil and plant respiration during this period may also control CO₂ concentrations due to which the correlation coefficient is not equal to 1. The ratio of $\Delta CO_{(exc)}/\Delta CO_{2(exc)}$ is estimated to be 8.5 ± 0.15 ppb/ppm (r = 0.72) and 12.7 ± 0.17 ppb/ppm (r = 0.74) in autumn and winter respectively. Relatively higher ratios during winter than other three seasons indicates contribution of CO emission from additional biofuel burning sources. The winter time observed ratios are similar to the airmass influenced by both fossil fuel and biofuel emissions as discussed by Lin et al. (2015) over Pondicherry. Using CARIBIC observations, Lai et al. (2010) also reported the $\Delta CO/\Delta CO_2$ ratio in the range of 15.6-29.3 ppb/ppm from the airmass influenced by both biofuel and fossil fuel burning in the Indo-Chinese Peninsula. Further, $\Delta CO/\Delta CO_2$ ratio is also observed of about 13 ppb/ppm in South-east Asian outflow in February-April, 2001 during the TRACE-P campaign and suggest the combined influence of fossil fuel and biofuel burning (Russo

et al., 2003). The narrow range of the overall ratios (using all data) from autumn to spring (8.4 - 12.7 ppb/ppm) suggest the dominance of local emission sources during these seasons, and this range is correspond to the range of anthropogenic combustion sources (10-15 ppb/ppm) in developed countries (Suntharalingam et al., 2004; Takegawa et al., 2004; Wada et al., 2011). This suggest that the overall emissions of CO over Ahmedabad are mostly dominated by the anthropogenic combustion during these seasons.

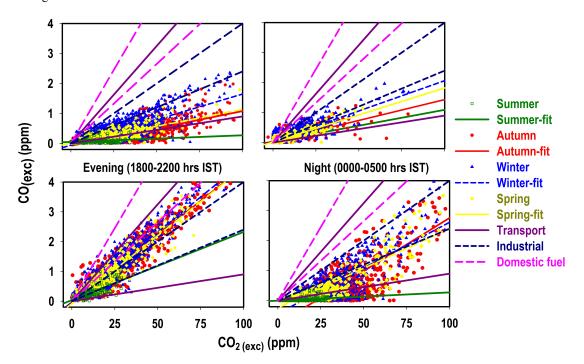


Figure 7: Scatter plots and regression fits of excess CO $(CO_{(exc)})$ vs. excess CO_2 $(CO_{2(exc)})$ during morning (0600-1000 hrs), noon (1100-1700 hrs), evening (1800-2200) and night (0000-0500 hrs) hours for all the four different seasons. Excess values of both species are calculated after subtracting their background concentrations. Each data points are averaged for 30 minutes. Emission ratios range of CO/CO_2 for different sources from the literature are also plotted in each figure.

Table 1: Correlation slopes $(\Delta CO_{(exc)}/\Delta CO_{2(exc)})$ in ppb/ppm) measured during different time intervals of distinct seasons. Coefficient of determination (r^2) is given inside the bracket.

Seasons	Slope in ppb/ppm (Coefficient of determination (r ²)				
	Morning	Afternoon	Evening	Night	
	(0600-1000 hrs)	(1100-1700 hrs)	(1800-2200 hrs)	(0000-0500 hrs)	
Summer (JA)	0.9 (0.15)	10.0 (0.17)	19.5 (0.67)	0.5 (0.16)	
Autumn (SON)	8.3 (0.48)	14.1 (0.75)	45.2 (0.90)	35.3 (0.71)	
Winter (DJF)	14.3 (0.51)	20.0 (0.68)	47.2 (0.90)	30.0 (0.75)	
Spring (MAM)	9.3 (0.68)	18.0 (0.80)	43.7 (0.93)	26.0 (0.80)	

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The $\Delta CO_{(exc)}/\Delta CO_{2(exc)}$ slope and their correlation may depend on the time of the day due to the variation in different controlling factors on their levels. Hence, we computed the diurnal cycle of $\Delta CO_{(exc)}/\Delta CO_{2(exc)} \ slope \ for \ all \ the \ seasons \ by \ binning \ the \ data \ for \ both \ hour \ and \ month$ × 24 hrs) as shown in Figure 7b. The colours indicate the correlation coefficients (r) for respective hour. These ratios do not reflect the diurnally varying PBL height, but rather the diurnally varying mix of fossil fuels and biogenic sources. The $\Delta CO_{(exc)}/\Delta CO_{2(exc)}$ slopes show very distinctive diurnal variation, being higher (30-50 ppb/ppm) in the evening rush hours with very good correlation (r>0.85) and lower (5-20 ppb/ppm) in the afternoon hours with lower correlation (r=0.5-0.6) during all the four seasons. Negative and lower slopes in afternoon hours during monsoon season indicate the higher biospheric productivity during this period. The slopes and their correlations are fairly comparable for all the four seasons in the evening rush hours which indicate stronger influence of common emission sources. Slopes during this time can be considered as fresh emissions since dilution and chemical loss of CO can be considered negligible for this time. These observed ratios are much lower than ratios related to domestic sources but are similar transport sector mostly dominated from gasoline combustion (Table 2). Except monsoon, the overall ratios in all four seasons were found in the range of 10-25 ppb/ppm during the daytime and 10-50 ppb/ppm during night-time.

Table 2: Emission ratios of CO/CO₂ (ppb/ppm), derived from emission factors (gram of gases emitted per kilogram of fuel burned).

Biomass burning	Transport		Industry	Domestic	
Crop-residue a,b,c	$Diesel^{d,e,f}$	Gasoline d,f	Coal	$Coal^{d,f}$	$Biofuel^{c,d}$
45.7-123.6	8.6-65.2	33.5	23.5-40.4	53.3-62.2	52.9-98.5

^aDhammapala et al. (2007); ^bCAO et al. (2008); ^cAndreae and Merlet (2001); ^dStreets et al. (2003); ^eSánchez-Ccoyllo et al. (2009); ^fWesterdahl et al. (2009)

4.5 Top-down CO emissions from observations

If the emissions of CO₂ are known for study locations, the emissions of CO can be estimated by multiplying the correlation slopes and molecular mass mixing ratios (Wunch et al., 2009; Wong et al., 2015). Final emissions of CO will depend on choosing the values of correlation slopes. The slopes should not be biased from particular local sources, chemical processing and PBL dynamics. We exclude monsoon the summer monsoon season data as the CO₂ variations mainly depend on the 750 biospheric productivity during this season. As discussed previously, the morning and evening rush hours data are appropriate for tracking vehicular emissions, while the afternoon data are affected by the other environmental factors, e.g., the PBL dynamics, biospheric activity and chemical process. The stable, shallow night-time PBL accumulates emissions since the evening and hence the correlation slope for this period can be used as a signature of the city's emissions. Hence, we calculate the slopes from the data corresponding to the period of night time (2300-0500 hrs) and evening 755 rush hour (1900-2200 hrs) 2300-0500 hrs. Additionally, slopes for morning hours (0600-1000 hrs), afternoon hours (1100-1700 hrs), and night hours (1800-0600 hrs) are also used for estimating the CO emissions to study the difference in the estimation of CO emissions due to choosing different times for slopes. The CO emission (E_{CO}) for Ahmedabad is calculated using the following formula.

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$$E_{CO} = \left(\alpha_{CO} \frac{M_{CO}}{M_{CO_2}}\right) E_{CO_2}$$
 (1)

Where, α_{CO} is the correlation slope of $CO_{(exc)}$ to $CO_{2(exc)}$ ppb ppm⁻¹, M_{CO} is the molecular mass of CO in g mol⁻¹, M_{CO_2} is the molecular mass of CO₂ in g mol⁻¹ and E_{CO_2} is the CO₂ emission in Gigagram (Gg) over Ahmedabad. The EDGARv4.2 emission inventory reported an annual emissions of CO_2 at $0.1^o \times 0.1^o$ for the period of 2000-2008 (http://edgar.jrc.ec.europa.eu/overview.php?v=42). It reported an annual CO_2 emission of 6231.6 Gg CO_2 yr⁻¹ by EDGARv4.2 inventory over the box (72.3<longitude<72.7°E, 22.8<latitude<23.2°N) which contain Ahmedabad coordinates in center of the box. We assume that the emissions of CO_2 are linearly changing with time and using increasing rate of emission from 2005 to 2008, we extrapolate the emission of CO_2 for 2014 over same area. The bottom-up CO_2 emission for the Ahmedabad is thus estimated of about 8368.6 Gg for the year of 2014. Further, for comparing the estimated emission with inventory emissions we extrapolated the CO emissions also for the year of 2014 using same method applied as for CO_2 . Further, we assumed same slopes for the year of 2008 and calculate CO emission for that year also. The slope values for different time period, estimated and inventory emissions of CO using different values of slope are given in Table 3.

Table 3: Estimates of emissions of CO using the the CO_2 emission from EDGAR inventory over the box (72.3<longitude<72.7°E, 22.8<latitude<23.2°N) and observed $CO_{(exc)}$: $CO_{2(exc)}$ slopes for different time periods. The correlation coefficient for corresponding slopes are given inside the bracket in slope column. Monsoon Data for the summer monsoon season are not included for calculating slopes.

Time (IST)	Slope (ppb/ppm) Correlation coefficient	EDGAR Emissions (Gg/yr)		Estimated Emssions Gg(yr)
	(r)	CO_2	CO	
2300 - 0500 hrs	13±0.14 (0.84)			69.2±0.7
0600 - 1000 hrs	11.4±0.19(0.75)			60.7±1.0
1100 - 1600 hrs	14.9±0.19(0.78)	8368.6	45.3	79.3±1.0
1800 - 0500 hrs 1900 - 2100 hrs	34.6±0.3747±0.27 (0.58)(0.95)			184.2±1.9 250.2±1.5
Full day (24 hr)	10.8±0.09(0.73)			57.5±0.5

The correlation between $CO_{(exc)}$ and $CO_{2(exc)}$ for the period of 2300-0500 hrs is tight and emission are mostly representative of the local anthropogenic sources. Hence, slope for this period can be considered for estimating the fossil fuel CO emissions for Ahmedabad. Further the uncertainty is total emission due to uncertainty associated with used slope is also calculated. Using this slope and based on CO₂ emissions from EDGAR inventory, the estimated fossil fuel emission for CO is observed 69.2±0.7 Gg (emission±uncertainty) for the year of 2014. The EDGAR inventory underestimates the emission of CO as they give the estimate about 45.3 Gg extrapolated for 2014. The slope corresponding to the night hours (1800-0600 hrs) give the highest estimate of CO. Using all combinations of slopes for other periods also, the derived CO emissions are larger than the bottom-up EDGAR emission inventory. The EDGAR inventory estimate the relative contributions of CO from industrial, transport and slum/residential sector to be about 42%, 42% and 10%, respectively. The possible cause for underestimation of CO by the EDGAR inventory could be the underestimation of residential emission, since other inventories particularly for major urban Indian cities (http://www.indiaenvironmentportal.org.in/files/file/Air-Pollution-in-Six-Indian-Cities.pdf) show large relative contributions from residential sector. The uncertainty associated with the emission factors for different sectors could be another cause for the underestimation of CO emissions, since these are important parameters for developing the inventory (Sahu et al., 2015).

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4.6 Diurnal tracking of CO₂ emissions

CO has virtually no natural source in an urban environments except oxidation of hydrocarbons and hence can help to disentangle the relative contributions of anthropogenic (from transport, power plant, industrial etc) and the biospheric sources (mainly from respiration) of CO₂, by serving as a tracer of combustion activity attribute CO2 enhancements to fossil fuel combustion on shorter timescale (Duren and Miller, 2012). As we discussed earlier that incomplete combustion of fossil fuels is the main sources of CO in urban environments and therefore can be used as a surrogate tracers to attribute CO₂ enhancements to fossil fuel combustion on shorter timescale. Several studies have used the simultaneous time-series measurements of CO₂ and CO to to segregate the fraction of CO₂ from anthropogenic sources and natural biospheric sources from it atmospheric concentrations demonstrated that the ratio of the excess concentrations of CO and CO2 in background eoneentrations can be used to determine the fraction of CO2 from fossil fuels and validated this method using carbon isotope (Δ ¹⁴CO₂) measurements (Levin et al., 2003; Turnbull et al., 2006, 805 2011; Lopez et al., 2013; Newman et al., 2013). This quantification technique is more practical, less expensive and less time consuming in comparison to the ¹⁴CO₂ method (Vogel et al., 2010). For performing this analysis, the background concentrations of CO and CO2 and the emission ratio of CO/CO₂ from anthropogenic emissions are required. The methods for calculating the background concentrations of CO₂ and CO are already discussed in Section 4.4. Figure 8a shows the excess diurnal variations of CO2 above the background levels during different seasons. As discussed in the previous section, the vehicular emissions are major emission sources over the study locations. The observed concentrations of both gases can also be directly used for calculating the emission ratio, provided the measured levels are not highly affected from natural sources and share the same origin. We have used the evening time (1900-2100 hrs) data of $CO_{2(exc)}$ and $CO_{(exc)}$ for whole study period to calculate the emission ratio of CO/CO₂ from predominant anthropogenic emission sources. The emission ratio for this time is calculated to be 47 ± 0.27 ppb CO/ppm CO₂ with very high correlation (r = 0.95) (Figure 8b), after excluding those data points, corresponding for which the mean wind speed is greater than 3 ms⁻¹ for avoiding the effect of fast ventilation. The tight correlation imply 820 that there is not a substantial difference in the emission ratio of these gases during the measurement period from November, 2013 to May, 2015. CO_{2(exc)} and CO_(exc) will be poorly correlated with each other if their emission ratio varies largely with time, assuming the correlation is mainly driven by emissions. Since anthropogenic emissions are very high for this period, contribution of respiration sources in the levels of CO₂ can be considered negligible during this period. This ratio can be 825 considered as the representative of anthropogenic sources, as discussed in previous section also. For ealculating the emission ratio of CO/CO₂ from the vehicular emissions, we used the evening time (1900-2200 hrs) concentrations of $CO_{2(exc)}$ and $CO_{(exc)}$ for whole study period since correlation for this period is very high. The other reason for choosing this time is that there is insignificant contribution of biospheric in CO₂ and no chemical loss of CO. We assume that negligible influence

of other sources (open biomass burning, oxidation of hydrocarbons) during this period. The emission ratio for this time is calculated to about 47 ± 0.27 ppb CO/ppm CO₂ with very high correlation (r = 0.95) (Figure 8b) after excluding those data points, corresponding for which the mean wind speed is greater than 3 ms⁻¹ for avoiding the effect of fast ventilation and transport from other sources. The tight correlation imply that there is not a substantial difference in the emission ratio of these gases during the measurement period from November, 2013 to May, 2015. CO_{2(exc)} and CO_(exc) will be poorly correlated with each other if their emission ratio varies largely with time, assuming the correlation is mainly driven by emissions. Since this ratio is mostly dominated by the transport sector, this analysis will give mainly the fraction of CO₂ from the emissions of transport sector. We define it as R_{CO/CO_{2(ant)}}. The standard deviation shows the uncertainty associated with slope which is very small. The contribution of transport sector (CO_{2(ant)}) in the diurnal cycle of CO₂ is calculated using following formula.

$$CO_{2Veh} = \frac{CO_{obs} - CO_{bg}}{R_{CO/CO_{2(gpt)}}} \tag{2}$$

where $CO_{(obs)}$ is the observed CO concentration and $CO_{(bg)}$ is a background CO value. Uncertainty in the $CO_{2(ant)}$ is dominated by the uncertainty in the $R_{CO/CO_{2(ant)}}$ and by the choice of $CO_{(bg)}$. The uncertainty in $CO_{2(ant)}$ due to the uncertainty in the $R_{CO/CO_{2(ant)}}$ is about 0.5% or 0.27 ppm and can be considered negligible. As discussed in Section 3, the uncertainty in the measurements of $CO_{(bg)}$ is very small and also can be considered negligible. Further, the contributions of CO_2 from other major sources are calculated by subtracting the $CO_{2(ant)}$ from the excess concentrations of CO_2 . These sources are those sources which do not emit significant amount of CO and can be considered mostly as natural sources (respiration), denoted by $CO_{2(bio)}$.

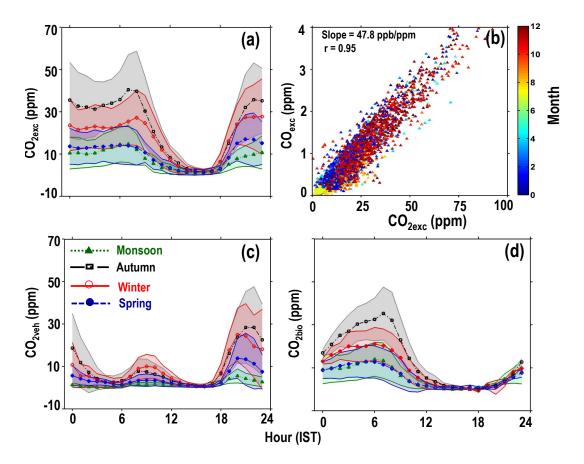


Figure 8: (a) Diurnal cycle of excess CO_2 over background levels during all the four seasons. (b) Correlation between excess CO and CO_2 for evening hours (1800-2100 hrs) during the study period. Contributions of fossil fuel (c) and biosphere (d) in the diurnal variation of excess CO_2 in all the four seasons.

The average diurnal cycles of CO_2 above its background for each season are shown in (Figure 8a). In Section 4.3.1, we have discussed qualitatively the role of different sources in the diurnal cycle of CO_2 . With the help of the above method, now the contributions of anthropogenic $(CO_{2(ant)})$ and biospheric sources $(CO_{2(bio)})$ are discussed quantitatively. Due to unavailability of PBL measurements, we cannot disentangle the contributions of boundary layer dynamics. The diurnal pattern of $CO_{2(ant)}$ (Figure 8c) reflects the pattern like CO, because we are using constant $R_{CO/CO_2(ant)}$ for all seasons. Overall, this analysis suggests that the anthropogenic emissions of CO_2 , mostly from transport and industrial sectors during early morning during 0600-1000 hrs varied from 15 to 60% (4-15 ppm). During afternoon hours (1100-1700 hrs), the anthropogenic originated (transport and industrial sources, mainly) CO_2 varied between 20 and 70% (1-11 ppm). During evening rush hours (1800-2200 hrs), highest contributions of combined emissions of anthropogenic sources (mainly transport and domestic) are observed. During this period the contributions vary from 50 to 95%

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(2-44 ppm. During night/early morning hours (0000-0700 hrs) non-anthropogenic sources (mostly biospheric respiration) contribute from 8 to 41 ppm of CO₂ (Figure 8d). The highest contributions from 18 to 41 ppm are observed in the autumn from the respiration sources during night hours, since there is more biomass during this season after the South Asian summer monsoon. During the afternoon hours, lower biospheric component of CO₂ could be due to a combination of the effects of afternoon anthropogenic emissions, biospheric uptake of CO₂ and higher PBL height.

The average diurnal cycles of CO_2 above its background for each seasons are shown in (Figure 8a). The diurnal pattern of $CO_{2(ant)}$ (Figure 8c) reflects the pattern like CO, because we are using constant $R_{CO/CO_{2(ant)}}$ for all seasons. Overall, this analysis suggests that the anthropogenic emissions of CO_2 from transport sectors during early morning from 0600-1000 hrs varied from 15 to 60% (4-15 ppm). During afternoon hrs (1100-1700 hrs), the vehicular emitted CO_2 varied between 20 and 70% (1-11 ppm) and during evening rush hours (1800-2200 hrs), it varies from 50 to 95% (2-44 ppm). During night/early morning hours (0000-0700 hrs) respiration contributes from 8 to 41 ppm of CO_2 (Figure 8d). The highest contributions from 18 to 41 ppm are observed in the autumn from the respiration sources during night hours, since there is more biomass during this season after the South Asian summer monsoon. During afternoon hours, lower biospheric component of CO_2 could be due to a combination of the effects of afternoon anthropogenic emissions, biospheric uptake of CO_2 and higher PBL height.

4.7 Model – Observations comparison

4.7.1 Comparison of diurnal cycle of CO₂

We first evaluate the ACTM in simulating the mean diurnal cycle of CO_2 over Ahmedabad by comparing the model simulated surface layer mean diurnal cycle of CO_2 . The atmospheric concentrations of CO_2 are calculated by adding the anthropogenic eomponent, oceanic eomponent and biospheric component from CASA process model. Figure 9a and Figure 9b show the residuals (Hourly mean-daily mean) of diurnal cycles of CO_2 based on the observations and the model simulations respectively. Model shows very little diurnal amplitude as compared to observationsal diurnal amplitude. Larger differences and discrepancies in night time and morning CO_2 concentrations between the model and observations might be contributed by diurnal cycle of the anthropogenic fluxes from local emissions and biospheric fluxes, and uncertainties in the estimation of PBLH by the model (Law et al., 2008). Hence, there is a need for efforts in improving the regional anthropogenic emissions as well as module for estimating the PBL height. It may be pointed out that the model's horizontal resolution $(1.125^o \times 1.125^o)$ is coarse for analysing local scale observations. However, the model is able to capture the trend of the diurnal amplitude, highest in autumn and lowest in monsoon the summer monsoon season. Figure 9c shows better agreement (r = 0.75) between the monthly change in modelled and observational diurnal amplitude of CO_2 from monthly mean diurnal cycle however

slope (m = 0.17) is very poor. We include the diurnal amplitudes of CO_2 for November and December, 2013 also for improving the total number of data points. The model captured the spread in the day time concentration of CO_2 from monsoonsummer to spring with a difference that the model shows lower concentration of CO_2 during noon hours in autumn while observations show lowest in monsoonthe summer monsoon season.

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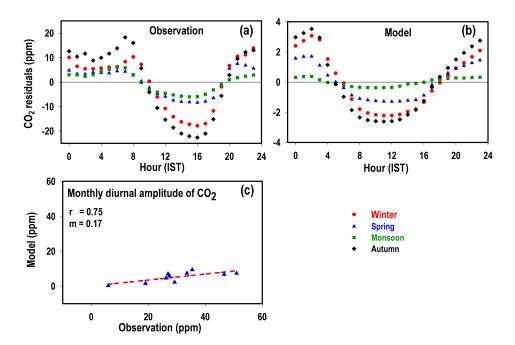


Figure 9: Residual of the diurnal cycle of CO₂ (in ppm) for (a) observations and (b) modal simulation over Ahmedabad in all the four seasons. Please note that the scales of the model and observational diurnal cycles are different. (c) Correlation between observed and the model simulated monthly mean diurnal cycle amplitudes.

The monthly average diurnal cycles of the biospheric net primary productivity from the CASA model for Ahmedabad and for the year of 2014 are shown Figure 10. The details of CASA flux are given in the Section 3.2. It is clear from Figure 10 that the CO₂ flux diurnal cycle as modelled by CASA shows minimum day-night variations amplitude during the summer monsoon time (Jun-July-August). Given that biosphere over Ahmedabad is water stressed for all other three seasons (except the summer monsoon time, Figure 1A3), the behaviour of CASA model simulated diurnal variation is not in line with biological capacity of the plants to assimilate atmospheric CO₂. Due to this underestimation of CO₂ uptake in the summer monsoon season, we also find very large underestimation of the seasonal through by ACTM in comparison with observations (Figure 9). Hence, there is a discrepancy in the diurnal flux of CO₂ simulated by CASA model. Similar discrepancy in the timing of maximum biospheric uptake is also discussed earlier by Patra et al. (2011) using inverse model CO₂ fluxes and CASA biospheric fluxes. Most of the atmospheric CO₂ uptake occur following the

Southwest monsoon season during July–September (Patra et al., 2011) and as a consequences we observe the lowest CO₂ concentration during the from the measurements during this season. But the model is not able to capture this feature since CASA biospheric flux (Figure 10) shows highest productivity in autumn and hence lowest concentrations of CO₂ in autumn during daytime. This It clearly suggest that there is a need for improving the biospheric flux for this region. It should be mentioned here that the CASA model used a land use map corresponding the late 1980s and early 1990s, which should be replaced by rapid growth in urbanised area in Ahmedabad (area and population increased by 91% and 42%, respectively, between 1990 and 2011). The model resolutions may be another factor for discrepancy. As Ballav et al. (2012) show that a regional model WRF-CO₂ is able to capture both diurnal and synoptic variations at two closely spaced stations within 25 km. Hence the regional models could be helpful for capturing these variabilities.

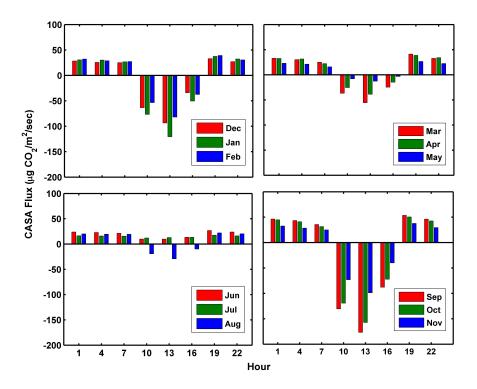


Figure 10: Diurnal variation of biospheric fluxes from the CASA ecosystem model.

4.7.2 Comparison of seasonal cycle of CO₂

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Figure 11a shows the performance of ACTM simulating mean seasonal cycle of CO_2 over Ahmedabad by comparing the model simulated mean surface seasonal cycle of CO_2 . Due to unavailability of data from March, 2014 to June, 2014 we plotted the monthly average of the year 2015 for same periods for visualizing the complete seasonal cycle of CO_2 . The seasonal cycles are calculated after subtracting the annual mean from each month, and corrected for growth rate using the observations at

MLO. For comparison, we used the seasonal cycle calculated from afternoon time average monthly concentrations, since the model is not able to capture the local fluctuations and produce better agreements when boundary layer is well mixed. In Table 4 we present the summary of the comparisons of the model and observations. The model reproduces the observed seasonal cycle in CO2 fairly well but with low seasonal amplitude about 4.15 ppm compared to 13.6 ppm observed. Positive bias during monsoon the summer monsoon season depicts the underestimation of biospheric productivity by the CASA model. The root mean square error is observed highest to be 3.21% in monsoonthe summer monsoon season. For understanding the role of biosphere, we also compared the seasonal cycle of CO₂ from noon time mean data with the seasonal cycle of CO₂ fluxes over South Asia region which is taken from the Patra et al. (2011) where they calculated it using a inverse model with including CARIBIC data and shifted a sink of 1.5Pg C yr⁻¹ sink from July to August and termed it as "TDI64/CARIBIC-modified". Positive and negative values of flux show the net release and net sink by the land biosphere over the South Asia. This comparison shows almost one to one correlation in the monthly variation of CO₂ and suggest that the lower levels of CO₂ during July, August and higher level in April are mostly due to the moderate source and sink of South Asian ecosystem during these months respectively. Significant correlation (r = 0.88) between South Asian CO₂ fluxes and monthly mean CO2 data for day time only suggest that the day time levels of CO2 are mostly controlled by the seasonal cycle of biosphere (Figure 11b).

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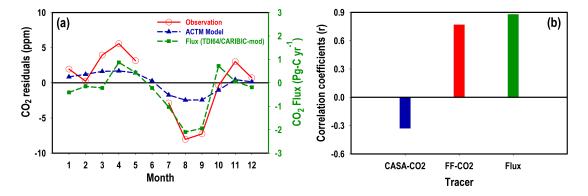


Figure 11: (a) The red circles and blue triangles show the mean seasonal cycles of CO_2 (in ppm) using afternoon values only, calculated from measurements and model over Ahmedabad. The green triangles show the seasonal cycles of CO_2 flux over South Asia, calculated from TDI64/CARIBIC-modified inverse model as given in Patra et al. (2011) (Figure 3d). (b) Blue bar and red bar shows the correlation coefficient (r) of model CO_2 concentration of biospheric tracer and fossil fuel tracer component with observed concentrations of CO_2 taking the entire annual time series of daily mean data, respectively. The green bar shows the correlation coefficient between the monthly residuals of afternoon mean only and the CO_2 flux over South Asia.

Separate correlations of each CO₂ tracer with the observations are helpful to determine relative importance of each flux component in the CO₂ variation (Patra et al., 2008). Hence, we perform separate correlation study between the measurements and biospheric, anthropogenic and oceanic component of CO₂, estimated by model using CASA-3hr fluxes (Randerson et al., 1997; Olsen and Randerson, 2004), EDGAR v4.2 inventory and air-sea fluxes from Takahashi et al. (2009) respectively. The correlation coefficient give the hint about dominating controlling factors of deriving the levels of CO₂. Separate correlation between individual tracers of model and observed data has been studied to investigate the relative contribution of individual tracer component in the CO2 variation (Figure 11b shows the resulting correlations for separate flux component with respect to measurements. We did not include the oceanic tracer and observed CO₂ correlation result, since no correlation has been observed between them. The comparison is based on daily mean of entire time series. Correlation between biospheric tracers and observed CO₂ have been found negative. This is because during growing season biospheric sources act as a net sink for CO2. Correlation of observed CO2 with fossil fuel tracer has been observed fairly well (r = 0.75). Hence, individual tracers correlation study also give the evidence of the overall dominance of fossil flux in overall concentrations of CO₂ over Ahmedabad for entire study period, and by assuming fossil fuel CO2 emission we can derive meaningful information on biospheric uptake cycle.

Table 4: Performance matrices used to quantify the level of agreement between the model simulations and observations. These statistics are based on hourly values in each day.

Parameter	Winter	Autumn	MonsoonSummer	All months
MB (ppm)	-2.72	12.64	-2.45	2.27
FGE (%)	0.96	3.12	2.0	1.76
RMSE (ppm)	5.21	12.82	9.14	8.60
RMSE (%)	1.27	3.21	2.20	2.09

970 This study suggest that the model is able to capture seasonal cycle with lower amplitude for Ahmedabad. However, the model fails to capture the diurnal variability since local transport and hourly daily flux play important roles for governing the diurnal cycle and hence there is a need for improving these features of the model.

5 Conclusions

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We report simultaneous in-situ measurements of CO₂ and CO concentrations in the ambient air at Ahmedabad, a semiarid urban region in western India using laser based CRDS technique during 2013-2015. Atmospheric concentrations of CO₂ were measured along with an anthropogenic tracer

CO at Ahmedabad, a semiarid urban region in western India using laser based CRDS technique during 2013-2015. The unique flow of air masses, originateding from both polluted continental regions as well as cleaner marine regions over the study location during different seasons, make this study most important for studying the characteristics of both type of air masses (polluted and relatively cleaner). Several key results are presented in this study. The observations show the range of CO₂ concentrations from 382 to 609 ppm and CO concentrations from 0.07 to 8.8 ppm, with the average of CO_2 and CO to be 416 ± 19 ppm and 0.61 ± 0.6 ppm respectively. The higher concentrations of both the gases are recorded for lower ventilation and for winds from north-east direction, representing CO₂ and CO transported from anthropogenic sources., while the lowest concentrations of both the gases are observed for higher ventilation and for the cleaner south-west winds from the Indian Oceanic region south-west direction, where air travels from the Indian Ocean. Along with these factors, the biospheric activity seasonal cycle (photosynthesis outweighs respiration during growing season and reverse during fall season) also controls the seasonal cycle of CO₂. Lowest day time CO₂ concentrations ranging from 382-393 ppm in August, suggest for the stronger biospheric productivity during this month over the study region, in agreement with an earlier inverse modelling study. This is in contrast to the terrestrial flux simulated by the CASA ecosystem model, showing highest productivity in September and October months. Hence, the seasonal cycles of both the gases reflect the seasonal variations of natural sources/sinks, anthropogenic emissions and seasonally varying atmospheric transport. The annual amplitudes of CO₂ variation after subtracting the growth rate based on the Mauna Loa, Hawaii data are observed to be about 26.07 ppm using monthly mean of all the data and 13.6 ppm using monthly mean of the afternoon period (1200-1600 hrs) data only. Significant difference between these amplitudes suggests that the annual amplitude from afternoon monthly mean data only does not give true picture of the variability. It is to be noted that most of the CO₂ measurements in India are based on day time flask samplings only.

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Significant differences in the diurnal patterns of CO₂ and CO are also observe, even though both the gases have major common emission sources and effects of undergo PBL dynamics and advection. Differences in their diurnal variability is probably the effect of terrestrial biosphere on CO₂ and chemical loss of CO due to reaction with OH radicals. The morning and evening peaks of CO are affected by rush hours traffic and PBL height variability and occur almost same time throughout the year. However, the morning peaks in CO₂ changes its time slightly due to shift in photosynthesis activity according to change in sun rise time during different seasons. The amplitudes of annual average diurnal cycles of CO₂ and CO are observed about 25 and 0.48 ppm respectively (Table 5). Both gases show highest amplitude in the autumn and lowest in monsoon the summer monsoon season. This shows that major influencing processes are common for both the gases, specific to this city and the monsoon India.

Table 5: Seasonal mean concentrations and diurnal amplitudes (max-min) of CO₂ and CO over Ahmedabad. Summary of results for study period. The total refers to the average of all the 24 hours data while noon time values refer to the average for 1200-1600 hrs. The 'exc' refer to the excess concentrations of CO₂ and CO after subtracting the background concentrations.

Period	Mean (ppm)		Diurn	al amplitude (ppm)	Correlatio	nCO _(exc) :CO _{2(exc)}	
	CO_2	СО	CO_2			Correlationcoefficient (r)	
Monsoon	400.3±6.8	0.19±0.13	12.4	0.24	0.6±0.1	0.62	
Autumn	419.6±22.8	0.72±0.71	40.9	1.36	8.5±0.2	0.72	
Winter	417.2±18.5	0.73±0.68	31.7	1.01	12.7±0.2	0.71	
Spring	415.4±14.8	0.41±0.40	15.9	0.62	8.4±0.1	0.87	
Annual	413.0±13.7	0.50±0.37	25.0	0.48	8.3±0.7	0.79	

The availability of simultaneous and continuous measurements of CO₂ and CO have made it possible to study their correlations at different time windows (during morning (0600-1000 hrs), noon (1100-1700 hrs), evening (1800-2200) and night (0000-0500 hrs) hours) of distinct seasons times of the day and during different seasons. Using the correlation slopes and comparing them with the emission ratios of different sources, contributions of distinct sources is discussed qualitatively. It is observed that during the evening hours, measurements over the study region are mostly affected by transport and domestic sources, while during other periods the levels of both gases are mostly dominated by the emissions from transport and industrial sources. The minimum value of slope and correlation coefficient of 0.8±0.2 ppb/ppm and 0.62 respectively are observed in monsoon. During other three seasons, the slopes vary in narrow range (Table 5) and indicate about the common emission sources of CO during these seasons. These slopes lie in the range (10 -15 ppb/ppm) of anthropogenic sources in developed countries, e.g., Japan. This suggest that the overall emissions of CO over Ahmedabad are mostly dominated by the anthropogenic (fossil fuel) combustion. These slopes also show significant diurnal variability having lower values (about 5-20 ppb/ppm) during noon hours and higher values (about 30-50 ppb/ppm) during evening rush hours with highest correlation (r>0.9). This diurnal pattern is similar to the traffic density and indicate the strong influence of vehicular emissions in the diurnal pattern of CO. Further, using the slope from the evening rush hours (1800-2200 hrs) data as vehicular anthropogenic emission ratios, the relative contributions of vehicular dominant anthropogenic emissions and biospheric emissions have been disentangle in the from the diurnal cycle of CO₂segregated. At rush hours, this analysis suggests that 90-95% of the total emissions of CO₂ are contributed by vehicular anthropogenic emissions. Using the relationship, the CO emission from Ahmedabad has been estimated. In this estimation, fossil fuel derived emission of CO₂ from EDGAR v4.2 inventory is extrapolated linearly from 2008

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to 2014 and it is assumed that there are no year-to-year variations in the land biotic and oceanic CO_2 emissions. The estimated annual emission CO for Ahmedabad is estimated to be 69.2 \pm 0.7 Gg for the year of 2014. The extrapolated CO emission from EDGAR inventory for 2014 shows a value smaller than this estimate by about 52%.

The observed results of CO_2 are also compared with an atmospheric general circulation model based chemistry transport model simulated CO_2 concentrations. The model captures some basic features like the trend of diurnal amplitude, seasonal amplitude etc, qualitatively but not quantitatively. The model captures the seasonal cycle fairly good but the amplitude is very less as compared to the observations. Similarly, performance of the model capturing the change in monthly averaged diurnal amplitude is quiet good (r = 0.72), however the slope is very poor. We also examined the correlation between the hourly averaged observed CO_2 and tracer of fossil fuel from model simulation and found fairly good correlation between them. However, no significant correlation has been observed between observed CO_2 and biospheric tracer. It suggests that the levels of CO_2 over Ahmedabad are mostly controlled by the fossil fuel combustion throughout the year.

This work demonstrate the usefulness of simultaneous measurements of CO_2 and CO in an urban region. The anthropogenic and biospheric component of CO_2 have been studied from its temporally varying atmospheric concentrations, and validity of "bottom-up" inventory have been assessed independently. Use of $CO_{(exc)}$: $CO_{2(exc)}$ ratios avoid some of the problems with assumptions that have to be made with modelling. These results represent a major urban region of India and will be helpful in validating emission inventories, chemistry-transport and terrestrial ecosystem models. However, a bigger network of sites is needed to elucidate more accurate distribution of emissions and their source regions, and run continuously over multiple years for tracking the changes associated with anthropogenic activities and emission mitigation policies.

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References

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- Ahmadov, R., Gerbig, C., Kretschmer, R., Koerner, S., Neininger, B., Dolman, A. J., and Sarrat, C.:
 Mesoscale covariance of transport and CO2 fluxes: Evidence from observations and simulations using the
 WRF-VPRM coupled atmosphere-biosphere model, Journal of Geophysical Research: Atmospheres, 112, doi:10.1029/2007JD008552, d22107, 2007.
 - Ammoura, L., Xueref-Remy, I., Gros, V., Baudic, A., Bonsang, B., Petit, J.-E., Perrussel, O., Bonnaire, N., Sciare, J., and Chevallier, F.: Atmospheric measurements of ratios between CO₂ and co-emitted species from traffic: a tunnel study in the Paris megacity, Atmospheric Chemistry and Physics, 14, 12871–12882, doi:10.5194/acp-14-12871-2014, 2014.
 - Andreae, M. O. and Merlet, P.: Emission of trace gases and aerosols from biomass burning, Global Biogeochemical Cycles, 15, 955–966, doi:10.1029/2000GB001382, 2001.
 - Baker, A. K., Schuck, T. J., Brenninkmeijer, C. A. M., Rauthe-Schöch, A., Slemr, F., van Velthoven, P. F. J., and Lelieveld, J.: Estimating the contribution of monsoon-related biogenic production to methane emissions from South Asia using CARIBIC observations, Geophysical Research Letters, 39, doi:10.1029/2012GL051756, 110813, 2012.
 - Ballav, S., PATRA, P. K., TAKIGAWA, M., GHOSH, S., DE, U. K., MAKSYUTOV, S., MURAYAMA, S., MUKAI, H., and HASHIMOTO, S.: Simulation of CO2 Concentration over East Asia Using the Regional Transport Model WRF-CO2, Journal of the Meteorological Society of Japan. Ser. I, 90, 959–976, doi:10.2151/jmsj.2012-607, 2012.
 - Ballav, S., Patra, P. K., Sawa, Y., Matsueda, H., Adachi, A., Onogi, S., Takigawa, M., and De, U.: Simulation of CO₂ concentrations at Tsukuba tall tower using WRF–CO₂ tracer transport model, J. Earth System Sci., in press, 2015.
- Bhattacharya, S. K., Borole, D. V., Francey, R. J., Allison, C. E., Steele, L. P., Krummel, P., Langenfelds, R.,

 Masarie, K. A., Tiwari, Y. K., and Patra, P.: Trace gases and CO₂ isotope records from Cabo de Rama, India,

 Current Science, 97, 2009.
 - Bitter, M., Ball, S. M., Povey, I. M., and Jones, R. L.: A broadband cavity ringdown spectrometer for in-situ measurements of atmospheric trace gases, Atmospheric Chemistry and Physics, 5, 2547–2560, doi:10.5194/acp-5-2547-2005, 2005.
- Briber, B. M., Hutyra, L. R., Dunn, A. L., Raciti, S. M., and Munger, J. W.: Variations in Atmospheric CO₂ Mixing Ratios across a Boston, MA Urban to Rural Gradient, Land, 2, 304, doi:10.3390/land2030304, 2013.
 - CAO, G., ZHANG, X., GONG, S., and ZHENG, F.: Investigation on emission factors of particulate matter and gaseous pollutants from crop residue burning, Journal of Environmental Sciences, 20, 50 55, doi:http://dx.doi.org/10.1016/S1001-0742(08)60007-8, 2008.
- 1100 Chen, H., Winderlich, J., Gerbig, C., Hoefer, A., Rella, C. W., Crosson, E. R., Van Pelt, A. D., Steinbach, J., Kolle, O., Beck, V., Daube, B. C., Gottlieb, E. W., Chow, V. Y., Santoni, G. W., and Wofsy, S. C.: High-accuracy continuous airborne measurements of greenhouse gases (CO₂ and CH₄) using the cavity ring-down spectroscopy (CRDS) technique, Atmospheric Measurement Techniques, 3, 375–386, doi:10.5194/amt-3-375-2010, 2010.
- 1105 Ciais, P., Sabine, C., Bala, G., Bopp, L., Brovkin, V., Canadell, J., Chhabra, A., DeFries, R., Galloway, J., Heimann, M., Jones, C., Quere, C., Myneni, R., Piao, S., and Thornton, P.: Carbon and Other Biogeochemical

- Cycles, book section 6, p. 465–570, Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, doi:10.1017/CBO9781107415324.015, 2013.
- Crosson, E.: A cavity ring-down analyzer for measuring atmospheric levels of methane, carbon dioxide, and water vapor, Applied Physics B, 92, 403–408, doi:10.1007/s00340-008-3135-y, 2008.
 - Dhammapala, R., Claiborn, C., Simpson, C., and Jimenez, J.: Emission factors from wheat and Kentucky blue-grass stubble burning: Comparison of field and simulated burn experiments, Atmospheric Environment, 41, 1512 1520, doi:http://dx.doi.org/10.1016/j.atmosenv.2006.10.008, 2007.
- Duren, R. M. and Miller, C. E.: Measuring the carbon emissions of megacities, Nature Clim. Change, 2, 560–1115 562, doi:http://dx.doi.org/10.1038/nclimate1629, 2012.
 - Kar, J., Bremer, H., Drummond, J. R., Rochon, Y. J., Jones, D. B. A., Nichitiu, F., Zou, J., Liu, J., Gille, J. C., Edwards, D. P., Deeter, M. N., Francis, G., Ziskin, D., and Warner, J.: Evidence of vertical transport of carbon monoxide from Measurements of Pollution in the Troposphere (MOPITT), Geophysical Research Letters, 31, doi:10.1029/2004GL021128, 2004.
- 1120 Karion, A., Sweeney, C., Wolter, S., Newberger, T., Chen, H., Andrews, A., Kofler, J., Neff, D., and Tans, P.: Long-term greenhouse gas measurements from aircraft, Atmospheric Measurement Techniques, 6, 511–526, doi:10.5194/amt-6-511-2013, 2013.
 - Lai, S. C., Baker, A. K., Schuck, T. J., van Velthoven, P., Oram, D. E., Zahn, A., Hermann, M., Weigelt, A., Slemr, F., Brenninkmeijer, C. A. M., and Ziereis, H.: Pollution events observed during CARIBIC flights in
- the upper troposphere between South China and the Philippines, Atmospheric Chemistry and Physics, 10, 1649–1660, doi:10.5194/acp-10-1649-2010, 2010.
 - Lal, S., Naja, M., and Subbaraya, B.: Seasonal variations in surface ozone and its precursors over an urban site in India, Atmospheric Environment, 34, 2713 2724, doi:http://dx.doi.org/10.1016/S1352-2310(99)00510-5, 2000.
- 1130 Lal, S., Chandra, N., and Venkataramani, S.: A study of CO₂ and related trace gases using a laser based technique at an urban site in western India., Current Science, 109, 2015.
 - Law, R. M., Peters, W., Rödenbeck, C., Aulagnier, C., Baker, I., Bergmann, D. J., Bousquet, P., Brandt, J., Bruhwiler, L., Cameron-Smith, P. J., Christensen, J. H., Delage, F., Denning, A. S., Fan, S., Geels, C., Houweling, S., Imasu, R., Karstens, U., Kawa, S. R., Kleist, J., Krol, M. C., Lin, S.-J., Lokupitiya, R., Maki, T., Maksyu-
- tov, S., Niwa, Y., Onishi, R., Parazoo, N., Patra, P. K., Pieterse, G., Rivier, L., Satoh, M., Serrar, S., Taguchi, S., Takigawa, M., Vautard, R., Vermeulen, A. T., and Zhu, Z.: TransCom model simulations of hourly atmospheric CO₂: Experimental overview and diurnal cycle results for 2002, Global Biogeochemical Cycles, 22, doi:10.1029/2007GB003050, gB3009, 2008.
 - Le Quéré, C., Moriarty, R., Andrew, R. M., Peters, G. P., Ciais, P., Friedlingstein, P., Jones, S. D., Sitch, S.,
- Tans, P., Arneth, A., Boden, T. A., Bopp, L., Bozec, Y., Canadell, J. G., Chevallier, F., Cosca, C. E., Harris, I., Hoppema, M., Houghton, R. A., House, J. I., Jain, A., Johannessen, T., Kato, E., Keeling, R. F., Kitidis, V., Klein Goldewijk, K., Koven, C., Landa, C. S., Landschützer, P., Lenton, A., Lima, I. D., Marland, G., Mathis, J. T., Metzl, N., Nojiri, Y., Olsen, A., Ono, T., Peters, W., Pfeil, B., Poulter, B., Raupach, M. R., Regnier, P., Rödenbeck, C., Saito, S., Salisbury, J. E., Schuster, U., Schwinger, J., Séférian, R., Segschneider,
- 1145 J., Steinhoff, T., Stocker, B. D., Sutton, A. J., Takahashi, T., Tilbrook, B., van der Werf, G. R., Viovy, N.,

- Wang, Y.-P., Wanninkhof, R., Wiltshire, A., and Zeng, N.: Global carbon budget 2014, Earth System Science Data Discussions, 7, 521–610, doi:10.5194/essdd-7-521-2014, 2014.
- Levin, I., Kromer, B., Schmidt, M., and Sartorius, H.: A novel approach for independent budgeting of fossil fuel CO₂ over Europe by ¹⁴CO₂ observations, Geophysical Research Letters, 30, doi:10.1029/2003GL018477, 2194, 2003.

1155

1165

1170

1175

- Lin, X., Indira, N. K., Ramonet, M., Delmotte, M., Ciais, P., Bhatt, B. C., Reddy, M. V., Angchuk, D., Balakrishnan, S., Jorphail, S., Dorjai, T., Mahey, T. T., Patnaik, S., Begum, M., Brenninkmeijer, C., Durairaj, S., Kirubagaran, R., Schmidt, M., Swathi, P. S., Vinithkumar, N. V., Yver Kwok, C., and Gaur, V. K.: Long-lived atmospheric trace gases measurements in flask samples from three stations in India, Atmospheric Chemistry and Physics, 15, 9819–9849, doi:10.5194/acp-15-9819-2015, 2015.
- Lopez, M., Schmidt, M., Delmotte, M., Colomb, A., Gros, V., Janssen, C., Lehman, S. J., Mondelain, D., Perrussel, O., Ramonet, M., Xueref-Remy, I., and Bousquet, P.: CO, NO_x and ¹³CO₂ as tracers for fossil fuel CO₂: results from a pilot study in Paris during winter 2010, Atmospheric Chemistry and Physics, 13, 7343–7358, doi:10.5194/acp-13-7343-2013, 2013.
- Machida, T., Matsueda, H., Sawa, Y., Nakagawa, Y., Hirotani, K., Kondo, N., Goto, K., Nakazawa, T., Ishikawa, K., and Ogawa, T.: Worldwide Measurements of Atmospheric CO₂ and Other Trace Gas Species Using Commercial Airlines, J. Atmos. Oceanic Technol., 25, 1744–1754, doi:10.1175/2008JTECHA1082.1, 2008.
 - Mahesh, P., Sharma, N., Dadhwal, V., Rao, P., Apparao, B., Ghosh, A., Mallikarjun, K., and Ali, M.: Impact of Land-Sea Breeze and Rainfall on CO₂ Variations at a Coastal Station, Journal of Earth Science and Climatic Change, 5, –, doi:10.4172/2157-7617.1000201, 2014.
 - Mallik, C., Lal, S., and Venkataramani, S.: Trace gases at a semi-arid urban site in western India: variability and inter-correlations, Journal of Atmospheric Chemistry, pp. 1–22, doi:10.1007/s10874-015-9311-7, 2015.
 - Mallik, C., Chandra, N., Venkataramani, S., and Lal, S.: Variability of atmospheric carbonyl sulfide at a semi-arid urban site in western India, Science of The Total Environment, 551–552, 725 737, doi:http://dx.doi.org/10.1016/j.scitotenv.2016.02.014, 2016.
 - Newman, S., Jeong, S., Fischer, M. L., Xu, X., Haman, C. L., Lefer, B., Alvarez, S., Rappenglueck, B., Kort, E. A., Andrews, A. E., Peischl, J., Gurney, K. R., Miller, C. E., and Yung, Y. L.: Diurnal tracking of anthropogenic CO₂ emissions in the Los Angeles basin megacity during spring 2010, Atmospheric Chemistry and Physics, 13, 4359–4372, doi:10.5194/acp-13-4359-2013, http://www.atmos-chem-phys.net/13/4359/2013/, 2013.
 - Olsen, S. C. and Randerson, J. T.: Differences between surface and column atmospheric CO₂ and implications for carbon cycle research, Journal of Geophysical Research: Atmospheres, 109, doi:10.1029/2003JD003968, http://dx.doi.org/10.1029/2003JD003968, d02301, 2004.
- Onogi, K., Tsutsui, J., Koide, H., Sakamoto, M., Kobayashi, S., Hatsushika, H., Matsumoto, T., Yamazaki, N., Kamahori, H., Takahashi, K., Kadokura, S., Wada, K., Kato, K., Oyama, R., Ose, T., Mannoji, N., and Taira, R.: The JRA-25 Reanalysis, Journal of the Meteorological Society of Japan. Ser. II, 85, 369–432, doi:10.2151/jmsj.85.369, 2007.
 - Park, M., Randel, W. J., Emmons, L. K., and Livesey, N. J.: Transport pathways of carbon monoxide in the Asian summer monsoon diagnosed from Model of Ozone and Related Tracers (MOZART), Journal of Geophysical Research: Atmospheres, 114, doi:10.1029/2008JD010621, 2009.

- Patra, P. K., Law, R. M., Peters, W., Rödenbeck, C., Takigawa, M., Aulagnier, C., Baker, I., Bergmann, D. J., Bousquet, P., Brandt, J., Bruhwiler, L., Cameron-Smith, P. J., Christensen, J. H., Delage, F., Denning, A. S., Fan, S., Geels, C., Houweling, S., Imasu, R., Karstens, U., Kawa, S. R., Kleist, J., Krol, M. C., Lin, S.-J., Lokupitiya, R., Maki, T., Maksyutov, S., Niwa, Y., Onishi, R., Parazoo, N., Pieterse, G., Rivier, L., Satoh, M.,
- 1190 Serrar, S., Taguchi, S., Vautard, R., Vermeulen, A. T., and Zhu, Z.: TransCom model simulations of hourly atmospheric CO₂: Analysis of synoptic-scale variations for the period 2002–2003, Global Biogeochemical Cycles, 22, doi:10.1029/2007GB003081, gB4013, 2008.

- Patra, P. K., Niwa, Y., Schuck, T. J., Brenninkmeijer, C. A. M., Machida, T., Matsueda, H., and Sawa, Y.: Carbon balance of South Asia constrained by passenger aircraft CO₂ measurements, Atmospheric Chemistry and Physics, 11, 4163–4175, doi:10.5194/acp-11-4163-2011, 2011.
- Patra, P. K., Canadell, J. G., Houghton, R. A., Piao, S. L., Oh, N.-H., Ciais, P., Manjunath, K. R., Chhabra, A., Wang, T., Bhattacharya, T., Bousquet, P., Hartman, J., Ito, A., Mayorga, E., Niwa, Y., Raymond, P. A., Sarma, V. V. S. S., and Lasco, R.: The carbon budget of South Asia, Biogeosciences, 10, 513–527, doi:10.5194/bg-10-513-2013, 2013.
- 1200 Pérez-Landa, G., Ciais, P., Gangoiti, G., Palau, J. L., Carrara, A., Gioli, B., Miglietta, F., Schumacher, M., Millán, M. M., and Sanz, M. J.: Mesoscale circulations over complex terrain in the Valencia coastal region, Spain Part 2: Modeling CO₂ transport using idealized surface fluxes, Atmospheric Chemistry and Physics, 7, 1851–1868, doi:10.5194/acp-7-1851-2007, 2007.
- Peylin, P., Law, R. M., Gurney, K. R., Chevallier, F., Jacobson, A. R., Maki, T., Niwa, Y., Patra, P. K., Peters, W.,
 Rayner, P. J., Rödenbeck, C., van der Laan-Luijkx, I. T., and Zhang, X.: Global atmospheric carbon budget: results from an ensemble of atmospheric CO₂ inversions, Biogeosciences, 10, 6699–6720, doi:10.5194/bg-10-6699-2013, 2013.
 - Randel, W. J. and Park, M.: Deep convective influence on the Asian summer monsoon anticyclone and associated tracer variability observed with Atmospheric Infrared Sounder (AIRS), Journal of Geophysical Research: Atmospheres, 111, doi:10.1029/2005JD006490, 2006.
 - Randerson, J. T., Thompson, M. V., Conway, T. J., Fung, I. Y., and Field, C. B.: The contribution of terrestrial sources and sinks to trends in the seasonal cycle of atmospheric carbon dioxide, Global Biogeochemical Cycles, 11, 535–560, doi:10.1029/97GB02268, 1997.
- Russo, R. S., Talbot, R. W., Dibb, J. E., Scheuer, E., Seid, G., Jordan, C. E., Fuelberg, H. E., Sachse, G. W.,
 Avery, M. A., Vay, S. A., Blake, D. R., Blake, N. J., Atlas, E., Fried, A., Sandholm, S. T., Tan, D., Singh,
 H. B., Snow, J., and Heikes, B. G.: Chemical composition of Asian continental outflow over the western
 Pacific: Results from Transport and Chemical Evolution over the Pacific (TRACE-P), Journal of Geophysical
 Research: Atmospheres, 108, doi:10.1029/2002JD003184, 2003.
- Sahu, L. and Lal, S.: Distributions of C2–C5 {NMHCs} and related trace gases at a tropical urban site in India, 1220 Atmospheric Environment, 40, 880 – 891, doi:http://dx.doi.org/10.1016/j.atmosenv.2005.10.021, 2006.
 - Sahu, S., Beig, G., and Parkhi, N.: High Resolution Emission Inventory of NOx and CO for Mega City Delhi, India, Aerosol and Air Quality Research, 15, 1137–1144, doi:10.4209/aaqr.2014.07.0132, 2015.
 - Schuck, T. J., Brenninkmeijer, C. A. M., Baker, A. K., Slemr, F., von Velthoven, P. F. J., and Zahn, A.: Greenhouse gas relationships in the Indian summer monsoon plume measured by the CARIBIC passenger aircraft,
- 1225 Atmospheric Chemistry and Physics, 10, 3965–3984, doi:10.5194/acp-10-3965-2010, 2010.

- Schuck, T. J., Ishijima, K., Patra, P. K., Baker, A. K., Machida, T., Matsueda, H., Sawa, Y., Umezawa, T., Brenninkmeijer, C. A. M., and Lelieveld, J.: Distribution of methane in the tropical upper troposphere measured by CARIBIC and CONTRAIL aircraft, Journal of Geophysical Research: Atmospheres, 117, doi:10.1029/2012JD018199, d19304, 2012.
- 1230 Sharma, N., Dadhwal, V., Kant, Y., Mahesh, P., Mallikarjun, K., Gadavi, H., Sharma, A., and Ali, M.: Atmospheric CO₂ Variations in Two Contrasting Environmental Sites Over India, Air, Soil and Water Research, 7, 61-68, doi:10.4137/ASWR.S13987, 2014.

1240

1255

2003.

- Streets, D. G., Bond, T. C., Carmichael, G. R., Fernandes, S. D., Fu, Q., He, D., Klimont, Z., Nelson, S. M., Tsai, N. Y., Wang, M. Q., Woo, J.-H., and Yarber, K. F.: An inventory of gaseous and primary aerosol emissions in Asia in the year 2000, Journal of Geophysical Research: Atmospheres, 108, doi:10.1029/2002JD003093,
- Suntharalingam, P., Jacob, D. J., Palmer, P. I., Logan, J. A., Yantosca, R. M., Xiao, Y., Evans, M. J., Streets, D. G., Vay, S. L., and Sachse, G. W.: Improved quantification of Chinese carbon fluxes using CO₂/CO correlations in Asian outflow, Journal of Geophysical Research: Atmospheres, 109, doi:10.1029/2003JD004362, 2004.
- Sánchez-Ccoyllo, O., Ynoue, R., Martins, L., Astolfo, R., Miranda, R., Freitas, E., Borges, A., Fornaro, A., Freitas, H., Moreira, A., and Andrade, M.: Vehicular particulate matter emissions in road tunnels in Sao Paulo, Brazil, Environmental Monitoring and Assessment, 149, 241-249, doi:10.1007/s10661-008-0198-5, 2009.
- Takahashi, T., Sutherland, S. C., Wanninkhof, R., Sweeney, C., Feely, R. A., Chipman, D. W., Hales, B., Friederich, G., Chavez, F., Sabine, C., Watson, A., Bakker, D. C., Schuster, U., Metzl, N., Yoshikawa-Inoue, H., Ishii, M., Midorikawa, T., Nojiri, Y., Körtzinger, A., Steinhoff, T., Hoppema, M., Olafsson, J., Arnarson, T. S., Tilbrook, B., Johannessen, T., Olsen, A., Bellerby, R., Wong, C., Delille, B., Bates, N., and de Baar, H. J.: Climatological mean and decadal change in surface ocean pCO2, and net sea-air {CO2}
- 1250 flux over the global oceans, Deep Sea Research Part II: Topical Studies in Oceanography, 56, 554 - 577, doi:http://dx.doi.org/10.1016/j.dsr2.2008.12.009, 2009.
 - Takegawa, N., Kondo, Y., Koike, M., Chen, G., Machida, T., Watai, T., Blake, D. R., Streets, D. G., Woo, J.-H., Carmichael, G. R., Kita, K., Miyazaki, Y., Shirai, T., Liley, J. B., and Ogawa, T.: Removal of NOx and NOy in Asian outflow plumes: Aircraft measurements over the western Pacific in January 2002, Journal of Geophysical Research: Atmospheres, 109, doi:10.1029/2004JD004866, 2004.
 - Tiwari, Y. K., Vellore, R. K., Kumar, K. R., van der Schoot, M., and Cho, C.-H.: Influence of monsoons on atmospheric CO₂ spatial variability and ground-based monitoring over India, Science of The Total Environment, 490, 570 – 578, doi:http://dx.doi.org/10.1016/j.scitotenv.2014.05.045, 2014.
- Turnbull, J. C., Miller, J. B., Lehman, S. J., Tans, P. P., Sparks, R. J., and Southon, J.: Comparison of 14CO₂, 1260 CO, and SF₆ as tracers for recently added fossil fuel CO₂ in the atmosphere and implications for biological CO₂ exchange, Geophysical Research Letters, 33, doi:10.1029/2005GL024213, 2006.
 - Turnbull, J. C., Karion, A., Fischer, M. L., Faloona, I., Guilderson, T., Lehman, S. J., Miller, B. R., Miller, J. B., Montzka, S., Sherwood, T., Saripalli, S., Sweeney, C., and Tans, P. P.: Assessment of fossil fuel carbon dioxide and other anthropogenic trace gas emissions from airborne measurements over Sacramento, California in spring 2009, Atmospheric Chemistry and Physics, 11, 705–721, doi:10.5194/acp-11-705-2011, 2011.
- 1265

- Vogel, F. R., HAMMER, S., STEINHOF, A., KROMER, B., and LEVIN, I.: Implication of weekly and diurnal 14C calibration on hourly estimates of CO-based fossil fuel CO2 at a moderately polluted site in southwestern Germany, Tellus B, 62, 512–520, doi:10.1111/j.1600-0889.2010.00477.x, 2010.
- Wada, A., Matsueda, H., Sawa, Y., Tsuboi, K., and Okubo, S.: Seasonal variation of enhancement ratios of trace gases observed over 10 years in the western North Pacific, Atmospheric Environment, 45, 2129 2137, doi:http://dx.doi.org/10.1016/j.atmosenv.2011.01.043, 2011.
 - Wang, Y., Munger, J. W., Xu, S., McElroy, M. B., Hao, J., Nielsen, C. P., and Ma, H.: CO₂ and its correlation with CO at a rural site near Beijing: implications for combustion efficiency in China, Atmospheric Chemistry and Physics, 10, 8881–8897, doi:10.5194/acp-10-8881-2010, 2010.
- 1275 Welp, L. R., Keeling, R. F., Weiss, R. F., Paplawsky, W., and Heckman, S.: Design and performance of a Nafion dryer for continuous operation at CO₂ and CH₄ air monitoring sites, Atmospheric Measurement Techniques, 6, 1217–1226, doi:10.5194/amt-6-1217-2013, 2013.

- Westerdahl, D., Wang, X., Pan, X., and Zhang, K. M.: Characterization of on-road vehicle emission factors and microenvironmental air quality in Beijing, China, Atmospheric Environment, 43, 697 705, doi:http://dx.doi.org/10.1016/j.atmosenv.2008.09.042, 2009.
- Wong, K. W., Fu, D., Pongetti, T. J., Newman, S., Kort, E. A., Duren, R., Hsu, Y.-K., Miller, C. E., Yung, Y. L., and Sander, S. P.: Mapping CH₄: CO₂ ratios in Los Angeles with CLARS-FTS from Mount Wilson, California, Atmospheric Chemistry and Physics, 15, 241–252, doi:10.5194/acp-15-241-2015, 2015.
- Wunch, D., Wennberg, P. O., Toon, G. C., Keppel-Aleks, G., and Yavin, Y. G.: Emissions of greenhouse gases from a North American megacity, Geophysical Research Letters, 36, doi:10.1029/2009GL039825, 2009.