1	
2	
3	
4	
5	
6	Theoretical analysis of mixing in liquid clouds. Part 3: Inhomogeneous mixing
7	
8	M. Pinsky(1), A. Khain(1), and A. Korolev(2)
9	
10	(1) Department of Atmospheric Sciences, The Hebrew University of Jerusalem, Israel
11	(2) Environment Canada, Cloud Physics and Severe Weather Section, Toronto, Canada
12	
13	
14	Submitted to
15	Atmospheric Chemistry and Physics
16	Revised: April 2016
17	Second revision: May 2016
18	
19	
20	Communicating author: Alexander Khain, The Hebrew University of Jerusalem,
21	khain@vms.huji.ac.il
22	
23	
24	
25	
26	

27 Abstract

28 An idealized diffusion-evaporation model of time-dependent mixing between a cloud volume and a droplet-free volume is analyzed. The initial droplet size distribution (DSD) in 29 30 the cloud volume is assumed to be monodisperse. It is shown that evolution of the microphysical variables and the final equilibrium state are unambiguously determined by two 31 non-dimensional parameters. The first one is the potential evaporation parameter R, 32 33 proportional to the ratio of the saturation deficit to the liquid water content in the cloud volume, that determines whether the equilibrium state is reached at 100% relative humidity, or 34 is characterized by a complete evaporation of cloud droplets. The second parameter Da is the 35 Damkölher number equal to the ratio of the characteristic mixing time to the phase relaxation 36 time. Parameters R and Da determine the type of mixing. 37

The results are analyzed within a wide range of values of R and Da. It is shown that there is no pure homogeneous mixing, since the first mixing stage is always inhomogeneous. The mixing type can change during the mixing process. Any mixing type leads to formation of a tail of small droplets in DSD and, therefore, to DSD broadening that depends on Da. At large Da, the final DSD dispersion can be as large as 0.2. The total duration of mixing varies from several to one hundred phase relaxation time periods, depending on R and Da.

The definitions of homogeneous and inhomogeneous types of mixing are reconsidered and clarified, enabling a more precise delimitation between them. The paper also compares the results obtained with those based on the classic mixing concepts.

47 Keywords: homogeneous and inhomogeneous mixing, turbulent diffusion, droplet48 evaporation

49

50

1. Introduction

51 Cloud physics typically investigates two types of turbulent mixing: homogeneous and extremely inhomogeneous (e.g. Burner and Brenguier, 2006; Andrejczuk et al., 2009; Devenish 52 53 et al., 2012; Kumar et al., 2012). The concept of extremely inhomogeneous mixing in clouds was introduced by Latham and Reed (1977); Baker and Latham (1979), Baker et al. (1980) and 54 Blyth et al. (1980). According to this concept, mixing of cloud air and sub-saturated air from 55 56 cloud surrounding results in complete evaporation of a fraction of cloud droplets, whereas size of other droplets remain unchanged. The studies of extremely inhomogeneous mixing were 57 closely related to investigation of different mechanisms underlying enhanced growth of cloud 58 59 droplets and warm precipitation formation (Baker et al., 1980; Baker and Latham, 1982). The concept of homogeneous mixing suggests that all the droplets partially evaporate, so the liquid 60 water content decreases while the droplet concentration remains unchanged (Lehmann et al., 61 2009; Pt1). The significance of the concepts of homogeneous and inhomogeneous mixing goes 62 far beyond formation of large-sized droplets. In fact, these concepts are closely related to the 63 mechanisms involved in formation of droplet size distributions (DSD) in clouds and to the 64 description of this formation in numerical cloud models. A detailed analysis of the classical 65 66 concepts of homogeneous and extremely inhomogeneous mixing is given by Korolev et al. 67 (2016, hereafter Pt1).

Mixing in clouds includes two processes: mechanical mixing caused by turbulent diffusion 68 and droplet evaporation accompanied by increasing relative humidity. The relative contribution 69 of these processes can be evaluated by comparison of two characteristic time scales: the 70 characteristic mixing time scale $\tau_{mix} \sim L^{2/3} \varepsilon^{-1/3}$ (where *L* is the characteristic linear scale of an 71 entrained volume and ε is the dissipation rate of turbulent kinetic energy) and the time of 72 phase relaxation $\tau_{pr} = (4\pi \mathscr{D}\overline{r}N)^{-1}$ (where N is droplet concentration in a cloud volume, \overline{r} is 73 the mean droplet radius and \mathcal{D} is the diffusivity of water vapor) characterizing the response of 74 the droplet population to changes in humidity (the list of notations is given in Appendix). The 75

choice of the phase relaxation time as the characteristic time scale of mixing is discussed byPinsky et al. (2016) (hereafter referred to as Pt. 2) and will be further elaborated below.

Mixing is considered homogeneous if $\tau_{mix}/\tau_{pr} \ll 1$. At the first stage of mixing, the initial gradients of the microphysical and thermodynamic variables rapidly decrease to zero. By the end of this stage, the fields of temperature, humidity (hence, the relative humidity, RH) and droplet concentration are spatially homogenized and all the droplets within the mixing volume experience the same saturation deficit. During the relatively lengthy second stage, droplets evaporate and increase the relative humidity in the volume. It was shown that homogeneous mixing takes place at scales below about 0.5 m (Pt. 2)

At spatial scales larger than ~0.5 m, $\tau_{mix}/\tau_{pr} > 1$ and the spatial gradients of RH remain for a long time. Consequently, droplets within the mixing volume experience different subsaturations, thus the mixing is considered inhomogeneous. At $\tau_{mix}/\tau_{pr} >> 1$, the mixing is considered extremely inhomogeneous.

According to the classical conceptual scheme, during the first stage of extremely 89 inhomogeneous mixing a fraction of droplets is transported into the droplet-free entrained 90 91 volume and evaporates completely. The evaporation continues until the evaporating droplets saturate the initially droplet-free volume. At the second stage, turbulent mixing between the 92 cloud volume and the initially droplet-free (but already saturated) volume homogenizes the 93 94 gradients of droplet concentration and other quantities. Since both volumes are saturated, mixing does not affect droplet sizes. As a result, the final (equilibrium) state is characterized by 95 the relative humidity RH=100% and the DSD shape similar to that before mixing, but with a 96 lower droplet concentration. The same result (a decrease in droplet concentration but 97 unchanged droplet size) is expected in cases of both monodisperse and polydisperse initial 98 99 DSD. Since the DSD shape does not change, the characteristic droplet sizes (i.e. the mean square radius, the mean volume radius and the effective radius) do not change either in thecourse of extremely inhomogeneous mixing.

Thus, according to the classical concepts, the final equilibrium state with RH=100% is reached either by a partial evaporation of all droplets (homogeneous mixing) or a total evaporation of a certain portion of droplets that does not affect the remaining droplets (extremely inhomogeneous mixing) (Lehmann et al., 2009; Pt1).

106 In analyses of in-situ measurements, the observed data are usually compared with those 107 expected at the final state of mixing as assumed by the classical mixing concepts. If droplet concentration decreases without a corresponding change in the characteristic droplet radius, 108 the mixing is considered "extremely inhomogeneous." If the characteristic droplet radius 109 decreases with an increase of the dilution level while droplet concentration decreases 110 insignificantly, the mixing is identified as "homogeneous." If both the characteristic droplet 111 radius and the droplet concentration change, the mixing is considered as "intermediate". 112 Quantitative evaluations of the microphysical processes specific for intermediate mixing 113 114 remain largely uncertain.

As was discussed in Pt 2, the final states of mixing suggested by the classical concepts are only hypothetical. To understand the essence of the final equilibrium states of mixing and evaluate the time needed to reach them, it is necessary to consider the time evolution of DSD in the course of mixing process. Time-dependent process of homogeneous mixing was analyzed in Pt. 2. It was shown that in important cases of wide polydisperse initial DSDs, the final state substantially differs from that hypothesized by the classical concepts.

In this study, which is a Pt 3 of the set of studies, we analyze the time-dependent process of inhomogeneous mixing. The structure of the paper is as follows. The main concept and the basic equations for time-dependent inhomogeneous mixing are described in Section 2. Analysis of non-dimensional diffusion-evaporation equations is presented in Section 3. The design and the results of simulations of non-homogeneous mixing are outlined in Sections 4 and 5. A discussion clarifying the concepts of homogeneous and inhomogeneous mixing is presented inconcluding Section 6.

128

129 **2.** The main concept and the basic equations

During mixing of cloud volume and entrained air volume, the following two processes 130 determine the change of the microphysical and thermodynamical variables: turbulent diffusion 131 resulting in mechanical smoothening of the gradients of temperature, water vapor and droplet 132 133 concentration, and droplet evaporation accompanied by phase transformation. In this study, inhomogeneous mixing is investigated based on the analysis and solution of a 1D diffusion-134 evaporation equation. To our knowledge, the idea of using a diffusive model of turbulent 135 136 mixing to describe the mixing process was first proposed by Baker and Latham (1982). A 137 diffusion-evaporation equation was also analyzed by Jeffery and Reisner (2006). In order to get a more precise understanding of the physics of mixing process the analysis is performed under 138 139 the following main simplifying assumptions:

a) turbulent mixing is analyzed neglecting vertical motions of mixing volumes, dropletcollisions and droplet sedimentation.

b) the total mixing volume is assumed adiabatic.

143 c) mixing is assumed to take place only along the *x*-direction, i.e. a 1D task is considered;

d) the initial DSD in the cloud volume is assumed monodisperse.

145 Other assumptions and simplifications are discussed below.

146

147 A schematic illustration of the initial conditions used in the study is shown in **Figure 1**. 148 Two air volumes are assumed to mix: a cloud volume (left) and a droplet-free volume (right), 149 each having the linear size of L/2. The value of L is assumed within the range of several tens 150 to a few hundred meters. The mixing starts at t = 0. The cloud volume is initially saturated 151 $S_1 = 0$, the initial droplet concentration is N_1 and the initial liquid water mixing ratio is

152
$$q_1 = \frac{4\pi\rho_w}{3\rho_a} N_1 r_0^3$$
. In the droplet-free volume the initial conditions are $RH_2 < 100\%$ (i.e.

153 $S_2 < 0$), $N_2 = 0$ and $q_2 = 0$. Therefore, the initial profiles of these quantities along the x-axis

are step functions

155
$$N(x,0) = \begin{cases} N_1 & \text{if } 0 \le x < L/2 \\ 0 & \text{if } L/2 \le x < L \end{cases}$$
(1a)

156
$$S(x,0) = \begin{cases} 0 & \text{if } 0 \le x < L/2 \\ S_2 & \text{if } L/2 \le x < L \end{cases}$$
 (1b)

157
$$q(x,0) = \begin{cases} q_1 & \text{if } 0 \le x < L/2 \\ 0 & \text{if } L/2 \le x < L \end{cases}$$
(1c)

The initial profile of droplet concentration is shown in Fig. 1. In this study, averaged equations are used. We do not consider mixing at scales below several millimeters. At the scales of averaging, there exist clear definitions of droplet concentration, supersaturation and other "macro scale" quantities. The mixing is assumed to be driven by isotropic turbulence within the inertial sub-range where the Richardson's law is valid. Accordingly, turbulent diffusion (turbulent mixing) is described by a 1D equation of turbulent diffusion with a turbulent coefficient *K*. The turbulent coefficient is evaluated as proposed by Monin and Yaglom (1975)

166
$$K(L) = C\varepsilon^{1/3}L^{4/3}$$
 (2)

167 In Eq. (2), C is a constant. Eq. (2) is valid in case turbulent diffusion is considered, i.e. at scales 168 where molecular diffusion can be neglected.

Since the total mixing volume is adiabatic, the fluxes of different quantities through the leftand right boundaries of the volume are equal to zero at any time instance, i.e.

171
$$\frac{\partial N(0,t)}{\partial x} = \frac{\partial N(L,t)}{\partial x} = 0; \quad \frac{\partial q(0,t)}{\partial x} = \frac{\partial q(L,t)}{\partial x} = 0; \quad \frac{\partial q_v(0,t)}{\partial x} = \frac{\partial q_v(L,t)}{\partial x} = 0 \quad (3)$$

172 where q_v is the water vapor mixing ratio.

During mixing, droplets in the mixing volume experience different subsaturations,
therefore, the initially monodisperse DSD will become polydisperse. The droplets that were

transported into the initially droplet-free volume will undergo either partial or complete

- evaporation. The evaporation leads to a decrease in both droplet size and droplet concentration.
- 177 The basic system of equations that describes the processes of diffusion and of evaporation 178 which occur simultaneously is to be derived. The first equation is written for value Γ defined 179 as

$$\Gamma = S + A_2 q \tag{4}$$

181 This value is conservative in a moist adiabatic process, i.e. it does not change during phase

182 transitions (Pinsky et al., 2013, 2014). In Eq. (4), the coefficient $A_2 = \frac{1}{q_v} + \frac{L_w^2}{c_p R_v T^2}$ is a weak

function of temperature that changes by ~10 % when temperatures change by ~10 $^{\circ}$ C (Pinsky et

184 al., 2013). In this study, it is assumed that $A_2 = constant$. In Eq. (4), $q = \frac{4\pi\rho_w}{3\rho_a} \int_0^\infty r^3 f(r) dr$ is the

185 liquid water mixing ratio and f(r) is the DSD. The quantity Γ obeys the diffusion equation

186
$$\frac{\partial \Gamma(x,t)}{\partial t} = K \frac{\partial^2 \Gamma(x,t)}{\partial x^2}$$
(5)

187 with the boundary conditions $\frac{\partial \Gamma(0,t)}{\partial x} = \frac{\partial \Gamma(L,t)}{\partial x} = 0$ and the initial profile at t = 0

188
$$\Gamma(x,0) = \begin{cases} A_2 q_1 & \text{if } 0 \le x < L/2 \\ S_2 & \text{if } L/2 \le x < L \end{cases}$$
(6)

189 Therefore, function $\Gamma(x,0)$ is positive in the left volume, and negative in the right volume.

190

Since Γ does not depend on phase transitions, Eq. (5) can be solved independently of other

equations. The solution of Eq. (5) with initial conditions (6) is (Polyanin and Zaitsev, 2004)

192

$$\Gamma(x,t) = \sum_{n=0}^{\infty} a_n \exp\left(-\frac{Kn^2 \pi^2 t}{L^2}\right) \cos\left(\frac{n\pi x}{L}\right) = \frac{1}{2} \left(S_2 + A_2 q_1\right) + (A_2 q_1 - S_2) \sum_{n=1}^{\infty} \frac{\sin(n\pi/2)}{n\pi/2} \exp\left(-\frac{Kn^2 \pi^2 t}{L^2}\right) \cos\left(\frac{n\pi x}{L}\right)$$
(7)

193 where the Fourier coefficients of expanding the step function (6) are

194
$$a_0 = \frac{1}{2} (A_2 q_1 + S_2)$$
 (8a)

195
$$a_n = (A_2 q_1 - S_2) \frac{\sin(n\pi/2)}{n\pi/2}, \ n = 1, 2, ...$$
 (8b)

196 An example of spatial dependencies of $\Gamma(x,t)$ at different time instances during the mixing is 197 shown in **Figure 2**. One can see a decrease in the initial gradients and a tendency to 198 establishing a horizontally uniform value of Γ . Since the initial volume was divided into two 199 equal parts, the diffusion leads to formation of a constant limit value of function Γ

200
$$\Gamma(x,\infty) = \frac{1}{2} (\Gamma(0,0) + \Gamma(L,0)).$$

The second basic equation is the equation for diffusional droplet growth, taken in the following form (Pruppacher and Klett, 2007)

$$203 \qquad \frac{d\sigma}{dt} = \frac{2S}{F} \tag{9}$$

204 where $\sigma = r^2$ is the square of droplet radius and $F = \frac{\rho_w L_w^2}{k_a R_v T^2} + \frac{\rho_w R_v T}{e_s(T)\mathcal{P}}$. The value of

coefficient F is considered constant in this study. The solution of Eq. (9) is

206
$$\sigma(t) = \frac{2}{F} \int_{0}^{t} S(t') dt' + \sigma_0$$
(10)

The third main equation describes the evolution of DSD. In the following discussion, the DSD will be presented in the form $g(\sigma)$ which is the distribution of the square of the radius. This formulation directly utilizes the property of the diffusion growth equation (9) according to which the time changes of DSD are reduced to shifting the distributions in the space of square radii, while the shape of the distribution remains unchanged. The standard DSD f(r) is related

212 to
$$g(\sigma)$$
 as $f(r) = 2r \cdot g(r^2)$.

213 The normalized condition for
$$g(\sigma)$$
 is

214
$$N = \int_{0}^{\infty} g(\sigma) d\sigma$$
 (11)

where *N* is the droplet concentration. Using DSD $g(\sigma)$, the liquid water mixing ratio can be presented as integral

217
$$q = \frac{4\pi\rho_w}{3\rho_a} \int_0^\infty \sigma^{3/2} g(\sigma) d\sigma$$
(12)

The 1D diffusion-evaporation equation for the non-conservative function $g(\sigma)$ can be written in the form (Rogers and Yau, 1989)

220
$$\frac{\partial g(\sigma)}{\partial t} = K \frac{\partial^2 g(\sigma)}{\partial x^2} - \frac{\partial}{\partial \sigma} \left(\frac{d\sigma}{dt} g(\sigma) \right)$$
(13)

where the first term on the right-hand side of Eq. (13) describes changes in the DSD due to
spatial diffusion, while the second term on the right-hand side describes changes in the DSD
due to evaporation. Substitution of Eq. (9) into Eq. (13) leads to the following equation

224
$$\frac{\partial g(x,t,\sigma)}{\partial t} = K \frac{\partial^2 g(x,t,\sigma)}{\partial x^2} - \frac{2S(x,t)}{F} \frac{\partial g(x,t,\sigma)}{\partial \sigma}$$
(14)

To close Eq. (14), Eq. (4) should be used in the form

226
$$S(x,t) = \Gamma(x,t) - A_2 q(x,t)$$
 (15)

where q(x,t) is calculated according to Eq. (12). Eqs. (12, 14, 15) constitute a closed set of equations allowing calculation of $g(x,t,\sigma)$.

To proceed to the equations for DSD moments, let us define a moment of DSD $g(\sigma)$ of order α as

231
$$m_{\alpha} = \overline{\sigma^{\alpha}} = \int_{0}^{\infty} \sigma^{\alpha} g(\sigma) d\sigma$$
 (16)

Multiplying Eq. (14) by σ^{α} , integrating within limits $[0...\infty]$ and assuming that $\sigma^{\alpha}g(\sigma) \rightarrow 0$ when $\sigma \rightarrow \infty$, yield a recurrent formula for the DSD moments

234

235
$$\frac{\partial m_{\alpha}(x,t)}{\partial t} = K \frac{\partial^2 m_{\alpha}(x,t)}{\partial x^2} + \alpha \frac{2S}{F} m_{\alpha-1}(x,t)$$
(17)

Eq. (17) provides a recurrent relationship between the DSD moments of different orders. This relationship was discussed by Pinsky et al.'s (2014) while analyzing diffusion growth in an ascending adiabatic parcel.

In particular, the equation for the liquid water mixing ratio that is a moment of the order of $\alpha = \frac{3}{2}$ can be written as

241
$$\frac{\partial q(x,t)}{\partial t} = K \frac{\partial^2 q(x,t)}{\partial x^2} + \frac{4\pi \rho_w N(x,t)\overline{r}(x,t)}{F \rho_a} S(x,t)$$
(18)

242 where the mean radius
$$\overline{r}(x,t) = \frac{m_{1/2}}{m_0}$$
.

In the general case, Eq. (18) is not closed, since concentration N(x,t) and $\overline{r}(x,t)$ are unknown functions of time and spatial coordinates.

The characteristic time of evaporation and of supersaturation change is the phase relaxation
time (Korolev and Mazin, 2003)

247
$$\tau_{pr} = \frac{\rho_a F}{4\pi \rho_w A_2 N \overline{r}}$$
(19)

248 Using Eq. (19), Eq. (18) can be rewritten as

249
$$\frac{\partial q(x,t)}{\partial t} = K \frac{\partial^2 q(x,t)}{\partial x^2} + \frac{1}{\tau_{pr}(x,t)} \left[\frac{1}{A_2} \Gamma(x,t) - q(x,t) \right] =$$
$$= K \frac{\partial^2 q(x,t)}{\partial x^2} + \frac{1}{A_2 \tau_{pr}(x,t)} S(x,t)$$
(20)

From Eqs. (20) and (15), the equation for supersaturation can be written in the following simpleform

252
$$\frac{\partial S(x,t)}{\partial t} = K \frac{\partial^2 S(x,t)}{\partial x^2} - \frac{S(x,t)}{\tau_{pr}(x,t)}$$
(20a)

Eqs. (20) and (20a) show that changes in the microphysical variables are determined by the rate of spatial diffusion (the first term on the right-hand side of these equations) and of evaporation (the second term on the right-hand side).

256

257

3. Analysis of non-dimensional equations

258 Spatial diffusion and evaporation depend on many parameters. It is the best to start the 259 analysis from the basic equation system presented in a non-dimensional form. A time scale 260 corresponding to the initial phase relaxation time in a cloud volume can be defined as

261
$$\tau_0 = \frac{\rho_a F}{4\pi \rho_w A_2 N_1 r_0}$$
(21)

and the non-dimensional time is $\tilde{t} = t/\tau_0$. Other non-dimensional parameters to be used are:

the non-dimensional phase relaxation time

264
$$\tilde{\tau}_{pr} = \tau_{pr} / \tau_0 = \frac{N_1 r_0}{N(\tilde{x}, \tilde{t}) \overline{r}(\tilde{x}, \tilde{t})}, \qquad (22a),$$

the normalized liquid water mixing ratio which is equal to the normalized liquid water content

266
$$\tilde{q} = \frac{q}{q_1}$$
, (22b),

the normalized supersaturation

268
$$\tilde{S} = \frac{S}{A_2 q_1}$$
(22c)

the non-dimensional conservative function

270
$$\tilde{\Gamma} = \frac{\Gamma}{A_2 q_1} \quad , \tag{22d},$$

the normalized square of droplet radius

272
$$\tilde{\sigma} = \frac{\sigma}{r_0^2},$$
 (22e),

the normalized droplet concentration

274
$$\tilde{N} = N/N_1$$
 (22f)

and the non-dimensional DSD

276
$$\tilde{g}(\tilde{\sigma}) = \frac{r_0^2}{N_1} g(\sigma)$$
(22g)

277 with normalization $\tilde{N} = \int_{0}^{1} \tilde{g}(\tilde{\sigma}) d\tilde{\sigma}$. The definition (22g) means that the integral of a non-

dimensional initial size distribution over the normalized square radius is equal to unity.

279 The non-dimensional distance and the non-dimensional time are defined as

280
$$\tilde{x} = x/L; \quad \tilde{t} = t/\tau_0$$
 (22h)

A widely used non-dimensional parameter showing the comparative rates of diffusion and evaporation is the Damkölher number:

283
$$Da = \frac{\tau_{mix}}{\tau_0} = \frac{L^2}{K\tau_0}$$
 (23)

where

$$\tau_{mix} = \frac{L^2}{K}$$
(24)

is the characteristic time scale of mixing. Using the non-dimensional parameters listed above,
Eq. (20) can be rewritten in a non-dimensional form as

288

289

$$\frac{\partial \tilde{q}(\tilde{x},\tilde{t})}{\partial \tilde{t}} = \frac{1}{Da} \frac{\partial^2 \tilde{q}(\tilde{x},\tilde{t})}{\partial \tilde{x}^2} + \frac{1}{\tilde{\tau}_{pr}(\tilde{x},\tilde{t})} \Big[\tilde{\Gamma}(\tilde{x},\tilde{t}) - \tilde{q}(\tilde{x},\tilde{t}) \Big] = \frac{1}{Da} \frac{\partial^2 \tilde{q}(\tilde{x},\tilde{t})}{\partial \tilde{x}^2} + \frac{1}{\tilde{\tau}_{pr}(\tilde{x},\tilde{t})} \tilde{S}(\tilde{x},\tilde{t})$$
(25)

290 where

291
$$\tilde{q}(\tilde{x},\tilde{t}) = \frac{N(\tilde{x},\tilde{t})\overline{\sigma^{3/2}}}{N_1 r_0^3} = \int_0^\infty \tilde{\sigma}^{3/2} \tilde{g}(\tilde{x},\tilde{t},\tilde{\sigma}) d\tilde{\sigma}$$
(26)

The initial conditions and the boundary conditions should be rewritten in a non-dimensional form as well. For instance, the normalized initial condition for the non-dimensional function $\tilde{q}(\tilde{x}, 0)$ can be derived from Eqs. (1c) and (22b)

295
$$\tilde{q}(\tilde{x},0) = \begin{cases} 1 & \text{if} \quad 0 \le \tilde{x} < 1/2 \\ 0 & \text{if} \quad 1/2 \le \tilde{x} < 1 \end{cases}$$
 (27)

296 The solution for $\tilde{\Gamma}(\tilde{x}, \tilde{t})$ obtained by a normalization of solution (7) is

297
$$\tilde{\Gamma}(\tilde{x},\tilde{t}) = \frac{1}{2} (1+R) + (1-R) \sum_{n=1}^{\infty} \frac{\sin(n\pi/2)}{n\pi/2} \exp\left(-\frac{n^2 \pi^2 \tilde{t}}{Da}\right) \cos(n\pi \tilde{x}), \quad (28)$$

298 where

299
$$R = \frac{S_2}{A_2 q_1}$$
 (29)

is a non-dimensional parameter referred to, hereafter, as a potential evaporation parameter (PEP). The PEP is proportional to the ratio of the amount of water vapour that should evaporate in order to saturate the initially droplet-free volume (that is determined by S_2) to the initial available liquid water q_1 in the cloud volume. The solution of Eq. (28) at $t \rightarrow \infty$ depends only on parameter *R*.

305
$$\tilde{\Gamma}(\tilde{x},\infty) = \frac{1}{2}(1+R)$$
(30)

The importance of PEP that determines a possible final state was illustrated in Pt. 1. PEP is also the sole parameter enabling calculation of the normalized mixing diagram for homogeneous mixing (Pt. 2). In this study, we consider cases when R < 0 since $S_2 < 0$, i.e. when droplets can only evaporate in the course of mixing.

310 The solution of Eq. (25) and the type of mixing depends on the values of two non-

dimensional parameters, namely, *Da* and *R*. Thus, when
$$R = \frac{S_2}{A_2 q_1} < -1$$
, $\tilde{\Gamma}(\tilde{x}, \infty) < 0$. It means

that the initially droplet-free volume V_2 is too dry and all the droplets in the mixing volume evaporate completely. At the final equilibrium state RH < 100%, i.e. $S(x, \infty) < 0$. If

314
$$R = \frac{S_2}{A_2 q_1} > -1$$
, $\tilde{\Gamma}(\tilde{x}, \infty) > 0$. This means that the mixed volume in the final state contains

droplets, i.e. the mixing leads expands the volume with droplets, i.e. the cloud volume. At the

final equilibrium state,
$$RH = 100\%$$
 (i.e. $S(x, \infty) = 0$). The case when $|R| = \left|\frac{S_2}{A_2q_1}\right| = \left|\tilde{S}_2\right| <<1$

corresponds to either *RH* close to 100% (i.e. S_2 is close to zero) (this case corresponds to the degenerated case considered in Pt. 1), and/or to the case when the liquid water mixing ratio in the cloud volume is large. In case $|R| \ll 1$, the second term on the right-hand side of Eq. (25) is much smaller than the first term, and the mixing is driven by turbulent diffusion only.

In case $Da \rightarrow 0$ (often considered as homogeneous mixing), at the beginning of the mixing 321 322 the diffusion term is much larger than the evaporation term, the second term on the right-hand side of Eq. (25). As mixing proceeds, within a short time period the total homogenization of all 323 324 the variables in the mixing volume is established and all the spatial gradients become equal to zero. At this time instance, the first term on the right-hand side becomes equal to zero, and the 325 326 second term on the right-hand side of Eq. (25), describing droplet evaporation, becomes 327 dominant. Thus, the analysis of the Eq. (25) shows that mixing consists of two stages. The first 328 mixing stage is a short stage of inhomogeneous mixing and the longer second stage of homogeneous mixing. The evolution of the microphysical variables during homogeneous 329 330 mixing is described in detail in Pt. 2.

 $Da \rightarrow \infty$ corresponds to extremely inhomogeneous mixing, according to the classic 331 332 concept. In this case, the diffusion term is much smaller than the evaporation term, so evaporation takes place under significant spatial gradients of RH. At $Da = \infty$, the adjacent 333 volumes do not mix at all and remain separated. This equivalent to existence of two 334 independent adiabatic volumes. Another interpretation of the limiting case $Da = \infty$ is an 335 infinite fast droplet evaporation. Both scenarios at $Da \rightarrow \infty$ indicate simplifications in the 336 337 definition of the extremely inhomogeneous mixing. At intermediate values of Da, mixing is inhomogeneous, when both turbulent diffusion and evaporation contribute simultaneously to 338 formation of the DSD. 339

340 Using Eq. (14) and normalization (22f), the equations for the non-dimensional size 341 distribution can be written as

342
$$\frac{\partial \tilde{g}(\tilde{x},\tilde{t},\tilde{\sigma})}{\partial \tilde{t}} = \frac{1}{Da} \frac{\partial^2 \tilde{g}(\tilde{x},\tilde{t},\tilde{\sigma})}{\partial \tilde{x}^2} + \frac{2}{3} \Big[\tilde{\Gamma}(\tilde{x},\tilde{t}) - \tilde{q}(\tilde{x},\tilde{t}) \Big] \frac{\partial \tilde{g}(\tilde{x},\tilde{t},\tilde{\sigma})}{\partial \tilde{\sigma}}$$
(31)

343

Eq. (31) is solved with the following initial conditions

345

346
$$\tilde{g}(\tilde{x},0,\tilde{\sigma}) = \begin{cases} \delta(\tilde{\sigma}-1) & \text{if } 0 \le \tilde{x} < 1/2 \\ 0 & \text{if } 1/2 \le \tilde{x} \le 1 \end{cases}$$
(32)

347 where $\delta(\tilde{\sigma}-1)$ is a delta function.

348

Table 1 presents the list of all the non-dimensional variables used in this study and the ranges of their variation. It is shown that six parameters determining the geometrical and microphysical properties of mixing can be reduced to two non-dimensional parameters, which enables a more efficient analysis of mixing. The ranges of parameter variations in Tab. 1 correspond to the simplifications used in the study (the initial DSD is monodisperse and *RH* \leq 100%).

355

4. Design of simulations

357 **Damkölher number** Da in clouds

The characteristic mixing time τ_{mix} can be evaluated using Eqs. (2) and (24)

359
$$au_{mix} = \frac{1}{C} \varepsilon^{-1/3} L^{2/3}$$
 (33)

There is significant uncertainty regarding the evaluation of τ_{mix} and Da in clouds, which is largely related to the choice of coefficient *C* in expression (33). These values differ in different studies: C = 10 (Jeffery and Reisner, 2006); C = 1 (Lehmann et al., 2009) and $C \approx 0.2$ (Monin and Yaglom, 1975) and Boffetta and Sokolov (2002). According to Lehmann et al. (2009), the values of Da in clouds of different types range from to 0.1 to several hundred. Thus, estimation of Da in clouds may vary within a wide range up to a few orders of magnitude. Da values of in stratocumulus clouds can be similar or even higher than those in cumulus clouds, since both τ_{mix} and τ_{pr} in stratiform clouds are larger than in cumulus clouds.

In our simulations, we compare the evolution of the microphysical parameters within a wide range of Da (from 1 up to 500) and of R (from -1.5 up to -0.1). Da = 1 represents the case closest to homogeneous mixing, while Da = 500 indicates extremely inhomogeneous mixing.

373

374 *Numerical method*

Calculations were performed using MATLAB solver PDEPE. We solve the equation system (31) for normalized DSD $\tilde{g}(\tilde{x}, \tilde{t}, \tilde{\sigma}_j)$ with the initial condition (32) and the Neumann boundary conditions

378

379
$$\frac{\partial \tilde{g}(0,\tilde{t},\tilde{\sigma}_j)}{\partial \tilde{x}} = \frac{\partial \tilde{g}(1,\tilde{t},\tilde{\sigma}_j)}{\partial \tilde{x}} = 0$$
(34)

380 where j = 1...24 are the bin numbers on a linear grid of square radii. The number of grid points 381 along the \tilde{x} axis was set equal to 81.

In calculation of the last term on the right-hand side of Eq. (31), the normalized supersaturation \tilde{S} was calculated first using the normalized conservative equation

384
$$\tilde{S}(\tilde{x},\tilde{t}) = \tilde{\Gamma}(\tilde{x},\tilde{t}) - \tilde{q}(\tilde{x},\tilde{t})$$
(35)

where $\tilde{\Gamma}(\tilde{x}, \tilde{t})$ is calculated using Eq. (28). Then, this term was formulated using Eq. (9) as

386
$$\frac{2}{3}\tilde{S}(\tilde{x},\tilde{t})\frac{\partial\tilde{g}(\tilde{x},\tilde{t},\tilde{\sigma}_{j})}{\partial\tilde{\sigma}_{j}} \approx \frac{\tilde{g}\left(\tilde{x},\tilde{t},\tilde{\sigma}_{j}+\frac{2}{3}\tilde{S}\Delta\tilde{t}\right)-\tilde{g}\left(\tilde{x},\tilde{t},\tilde{\sigma}_{j}\right)}{\Delta\tilde{t}}$$
(36)

Therefore, at each time step, the DSD \tilde{g} first was shifted to the left to the value $\frac{2}{3}\tilde{S}\Delta\tilde{t}$, where

388 $\Delta \tilde{t}$ is a small time increment chosen so that $\left|\frac{2}{3}\tilde{S}_{\max}\Delta \tilde{t}\right| \leq \frac{\Delta \tilde{\sigma}}{2}$. Next, the shifted DSD was 389 remapped onto the fixed square radius grid $\tilde{\sigma}_{j}$. We used the remapping method proposed by 390 Kovetz and Olund (1969), which conserves droplet concentration and LWC. After remapping, 391 the differences between the new and old DSDs were recalculated. The new values of LWC then 392 were determined using new values of DSD and Eq. (26). MATLAB utility PDEPE 393 automatically chooses the time step needed to provide stability of calculations.

394

395 **5. Results of simulations**

5.1 Full evaporation case

First, we consider the case R = -1.5, when all the cloud water evaporates completely. This process corresponds to the cloud dissipation caused by mixing with the entrained dry air. At the final state, RH is expected to be uniform and negative over the entire mixing volume.

Figure 3 shows spatial and time changes of \tilde{S} for Da = 1, 50 and 500. At the final state 400 for all the three cases \tilde{S} =-0.25, which is in agreement with the analytical solution of Eq. (30). 401 The final negative value indicates that all the droplets completely evaporated during mixing. At 402 403 Da = 1 (Fig.3ab), two stages of supersaturation evolution can be identified. The first short stage with $t < 0.4\tau_{pr}$ is the period of inhomogeneous mixing, when the gradients of RH 404 persist. By end of the second stage of about $14\tau_{pr}$, the equilibrium state is reached. Thus, at 405 small Da both types of mixing take place. In the cases of Da = 50 and Da = 500, the spatial 406 gradients exit during the entire period of mixing until the equilibrium state is reached 407 (approximately $50\tau_{pr}$ and $300\tau_{pr}$, respectively) (Fig.3cdef). Therefore, at these *Da* mixing is 408 inhomogeneous during entire mixing. 409

Analysis of Figs. 3 and 4 allows to introduce two characteristic time periods: (1) period T_{min} 414 during which the spatial gradients of the microphysical parameters persist, and mixing is 415 inhomogeneous, and (2) period T_{ev} during which droplet evaporation takes place. Both time 416 periods are dimensionless and normalized using τ_0 . Time period T_{ev} is equal either to the time 417 of complete droplet evaporation (when R < -1.0) or to the time period during which the 418 saturation deficit in the mixing volume becomes equal to zero (or close to zero if R > -1.0), 419 i.e. evaporation is actually terminated. Quantitative evaluations of T_{mix} and T_{ev} will be given in 420 Section 5.3. At $\tilde{t} < T_{mix}$, droplets in the mixing volume experience different saturation deficits. 421 Toward the end of time T_{mix} the saturation deficit becomes uniform over the entire mixing 422 423 volume because of mechanic mixing. At Da = 1, the homogenization of the saturation deficit and all the microphysical variables takes place during a very short time of about $0.5\tau_{\rm pr}$, and 424 then the evaporation of droplets is assumed to take place under the same subsaturation 425 426 conditions, so $T_{mix} \ll T_{ev}$.

Figs. 4a,b show that at $\tilde{t} \approx 0.35$, normalized LWC drops down from 1 to 0.4. Since the average value of the normalized LWC in the mixing volume is equal to 0.5 (see the initial condition in Eq. (27)), 20% of the droplet mass evaporates during this short inhomogeneous period. Thus, despite being quite short, inhomogeneous mixing stage plays an important role even at Da = 1.

Since at t = 0 the mixing volume is not spatially homogeneous by definition, there is always a period while spatial inhomogeneity exists. With increasing Da, the duration of the inhomogeneous stage increases and the duration of the homogeneous stage decreases. At

Da = 500, homogenization of the saturation deficit requires $250\tau_{pr}$, which is twice as long as 435 the time of complete droplet evaporation, i.e. $T_{mix} \approx 2T_{ev}$. This means that at Da = 500, droplet 436 evaporation takes place in the presence of the spatial gradients of supersaturation. After 437 complete evaporation of droplets, spatial gradients of the water vapour mixing ratios remain. 438 439 This kind of mixing is regarded as inhomogeneous.

440 At Da = 50, the time of complete evaporation is approximately equal to the time of supersaturation homogenization, i.e. $T_{mix} \approx T_{ev}$. In this case, as at Da = 500, the droplets 441 experience different saturation deficit within the mixing volume, so mixing is inhomogeneous 442 at Da = 50. 443

444 The differences in droplet evaporation at different *Da* can be seen in **Figure 5**., showing the relationships between \tilde{N} and \tilde{q} plotted with a certain time increment, so that each symbol 445 in the diagrams corresponds to a particular time instance. These symbols form curves. Each 446 panel of Fig. 5 shows three curves corresponding to different \tilde{x} : the centre of the initially cloud 447 volume ($\tilde{x} = 1/4$); the centre of the mixing volume ($\tilde{x} = 1/2$) and the centre of the initially 448 droplet-free volume ($\tilde{x} = 3/4$). The directions of the time increase are shown by arrows along 449 the corresponding curves. The initial points of the curves corresponding to $\tilde{t} = 0$ are 450 characterized by values $\tilde{q} = 1$ and $\tilde{N} = 1$ at $\tilde{x} = 1/4$, and by values $\tilde{q} = 0$ and $\tilde{N} = 0$ at $\tilde{x} = 3/4$. 451 The behaviour of the $\tilde{N} - \tilde{q}$ relationship provides important information about mixing 452 process. At $\tilde{t} < T_{mix}$, there are spatial gradients of \tilde{N} and \tilde{q} , i.e. \tilde{N} and \tilde{q} are different at 453 different \tilde{x} . This means that the three curves at $\tilde{t} < T_{mix}$ do not coincide. At $\tilde{t} > T_{mix}$, the spatial 454 gradients of \tilde{N} and \tilde{q} disappear and the three curves coincide. When the curves do not 455 coincide, mixing is inhomogeneous, and the coincidence of the curves indicates that the mixing 456 becomes homogeneous. In Fig. 5a and 5b (Da = 1 and Da = 5, respectively), the curves 457 coincide at point A corresponding to time $\tilde{t} = T_{mix}$.

458

Figs. 5a,b show that at Da = 1 and Da = 5, mixing consists of two stages: inhomogeneous 459 and homogeneous. The time instance $\tilde{t} = T_{mix}$ separates these two stages. In turn, the period of 460 homogeneous mixing (when evaporation is spatially homogeneous) can be separated into two 461 sub-periods. During the first sub-period, droplets evaporate only partially and \tilde{q} decreases at 462 the same droplet concentration. This sub-period is very pronounced at Da = 1, when \tilde{q} 463 decreases from about 0.4 to 0.1 at the unchanged droplet concentration. At the second sub-464 period, when $\tilde{q} < 0.1$, droplets evaporate completely, beginning with smaller ones, so both the 465 droplet concentration and \tilde{q} rapidly drop to zero. At Da = 5 (Fig. 5b), at the stage of 466 homogeneous evaporation (that begins at point "A") the decrease in \tilde{q} is accompanied by a 467 decrease in \tilde{N} . 468

At Da = 50 (Fig 5c), curves corresponding to different values of \tilde{x} do not coincide, except 469 at the final point "F", where $\tilde{N} = 0$ and $\tilde{q} = 0$. This means that horizontal gradients exist during 470 the entire mixing process and mixing is inhomogeneous till the final equilibrium state is 471 reached. Droplets penetrating into the initially droplet-free volume begin evaporating, so only a 472 small fraction of droplets reaches the centre of the droplet-free volume, as seen in Fig. 5c, 473 $\tilde{x} = 3/4$ (black curve). Accordingly, at $\tilde{x} = 3/4$ the droplet concentrations and \tilde{q} reach their 474 maxima (of 0.1 and 0.05, respectively) and then decrease to zero. At Da = 500 (Fig 5d), all 475 476 the droplets evaporate before reaching the centre of the dry volume, indicating an extremely high spatial inhomogeneity of droplet evaporation. Hence, only two curves for $\tilde{x} = 1/4$ and 477 $\tilde{x} = 1/2$ are seen in Fig.5d. 478

Fig. 5 also shows that the slopes of the curves describing the $\tilde{N} - \tilde{q}$ relationships are different at different values of \tilde{x} and change over time. At large Da, the slopes of the curves describing the dependencies $\tilde{N} - \tilde{q}$ in the initially cloud volume are close to linear. However, the slope at a high value of \tilde{q} is still flatter than that at a low value of \tilde{q} . This can be attributed to the fact that when \tilde{q} is large, it decreases faster than the concentration \tilde{N} because some 484 fraction of droplets evaporate only partially. At the end of the mixing when \tilde{q} is small, \tilde{N} 485 decreases faster than \tilde{q} , because the droplet concentration is determined by the smallest 486 droplets, while \tilde{q} is determined by larger droplets.

As was discussed in Pt. 1, according to the classical concept of extremely inhomogeneous mixing, the ratio q/N remains constant. For dimensionless \tilde{N} and \tilde{q} , the scattering points should be aligned along the 1:1 line. Therefore, the closeness of particular cases to the classical extremely inhomogeneous mixing can be evaluated by the deviation of the $\tilde{N} - \tilde{q}$ curve from the 1:1 line. One can see that at Da = 500 the $\tilde{N} - \tilde{q}$ relationship is closer to linear.

Despite the fact that at R < -1 all the droplets within the mixing volume evaporate, it is 492 interesting to follow the DSD evolution during this process. Figure 6 shows the time evolution 493 of a normalized DSD at Da = 1 and Da = 50. One can see a substantial difference in the DSD 494 evolutions at different Da. At Da = 1, different DSDs are formed very rapidly at different 495 values of \tilde{x} (panel a). The widest DSD occurs at $\tilde{x} = 1$, i.e. at the outer boundary of the initially 496 droplet-free volume. This is natural, because the supersaturation deficit is the highest at $\tilde{x} = 1$. 497 At $\tilde{t} > T_{mix} \approx 0.4$, DSD become similar at all values of \tilde{x} (Fig.6b). The DSD width continues to 498 increase due to partial droplet evaporation. This time period corresponds to the horizontal 499 segment of the $\tilde{N} - \tilde{q}$ relationship in Fig. 5a. Fig. 6c shows the DSD at the stage when a 500 decrease in LWC is accompanied by a decrease in number droplet concentration. The 501 corresponding point in the $\tilde{N} - \tilde{q}$ diagram at this time instance is quite close to the point "F" at 502 which $\tilde{N} = 0$ and $\tilde{q} = 0$. 503

At Da = 50, DSD are different at different \tilde{x} during the entire period of mixing. While DSD at $\tilde{x} > 0.5$ are wide and droplet evaporation is accompanied by a shift of DSD maximum to smaller droplet radii (this feature is typically attributed to homogeneous mixing), the DSD maximum at $\tilde{x} < 0.5$ (the initially cloud volume) shifts toward smaller radii only slightly until $\tilde{t} = 3.17$ (Fig. 6e). Further droplet evaporation either leads to a complete evaporation (at 509 $\tilde{x} \ge 0.5$) or shifts the DSDs to smaller droplet sizes (panel f). The maximum droplet 510 concentration takes place at $\tilde{x} = 0$. Fig. 6 shows that DSD shapes evolve substantially over 511 time, although the final state is characterized by complete droplet evaporation.

- 512
- 513 **5.2 Partial evaporation case**

514 **5.2.1** Evolution of the microphysical parameters at different values of *Da* and *R*

Here we consider the process of mixing at R > -1, i.e. when not all the droplets evaporate 515 completely. Figure 7 shows the horizontal profiles of a normalized supersaturation at different 516 517 Da and R. One can see that in all cases, the final state occurs when the equilibrium supersaturation $\tilde{S} = 0$ (RH=100%). However, this final value is reached quite differently 518 519 depending on Da. At Da = 1, rapid mixing leads to formation of spatially homogeneous humidity and supersaturation during a time period of a fraction of τ_{pr} . Then, supersaturation 520 within the mixing volume grows by evaporation of droplets, which are uniformly distributed 521 over the entire mixing volume. This process of homogeneous mixing was analyzed in detail in 522 Pt. 2. 523

524 At Da = 500, changes in supersaturation take place largely within the initially droplet-free volume. RH in the initially cloud volume undergoes only small changes. This process agrees 525 well with the classical concept of extremely inhomogeneous mixing. However, a strong 526 gradient of supersaturation remains within the initially drop-free volume for a long time (tens 527 of τ_{nr}). At Da = 50, the situation is intermediate. Mixing is intensive enough to decrease RH 528 in the initially cloud volume, but spatially uniform RH is established within about 5-10 τ_{pr} , 529 increasing with an increase in |R|. After this time instance, mixing takes place according to the 530 homogeneous scenario. 531

Figure 8 shows the horizontal profiles of normalized LWC at different Da and R. At the same R, the final equilibrium values of LWC are identical, as follows from Eq. (30); LWC decreases with an increase in |R|. At any Da, the decrease in the LWC in the cloud volume is caused largely by diffusion of droplets from the cloud volume into the initially droplet-free volume .

At Da = 500, evaporation in the cloud volume is small because \tilde{S} in these volumes is high 537 in cloud volumes during mixing (Fig. 7). At Da=1, the process of spatial homogenization 538 takes place during fractions of τ_{pr} , i.e. $T_{mix} < 1$. Then, during a relatively lengthy period of 539 $10\tau_{mr}$, evaporation decreases LWC over the entire mixing volume, which is characteristic of 540 homogeneous mixing. At Da = 50, spatial homogenization takes place during about $T_{mix} \approx 15$. 541 This is a slightly shorter time than it takes to establish the final equilibrium stage T_{tot} . Different 542 Da's cases reach equilibrium at different times. The process of reaching a final uniform LWC 543 lasts for $100\tau_{pr}$ at Da = 500 and for about τ_{pr} at Da = 1. 544

Figure 9 shows the profiles of the normalized droplet concentrations at different Da and 545 R. In contrast to LWC, the final concentration depends both on Da and R. Hence, profiles at 546 different *Da* can have different shapes at the same value of *R*. At R = -0.1 (which corresponds 547 to high RH in the initially dry volume) none of the droplets evaporate, so the final normalized 548 droplet concentration is equal to $\tilde{N} = 1/2$. This means that all the droplets in the initially cloud 549 volume are now uniformly distributed between both mixing volumes. At larger |R|, i.e., at 550 lower RH in an initially droplet-free volume, some droplets evaporate completely. The final 551 552 concentration decreases with an increase in Da.

The physical interpretation of this dependence is clear. At low Da, fast mixing leads to formation of a uniform RH throughout the entire mixing volume, and this affects all the droplets. At high Da, RH in the initially droplet -free volume remains low for a long time, and droplets that penetrate can evaporate. Therefore, the fraction of completely evaporated droplets increases with Da: at R = -0.1 there are no completely evaporated droplets at any Da. 558 At R = -0.3 a decrease in the droplet concentration takes place only at Da = 500, and at 559 R = -0.5 the droplet concentration decreases already at $Da \ge 50$.

The comparative contributions of different factors in establishing the final states of mixing are well seen in **Figure 10** presenting the relationships between normalized concentration and normalized LWC at three values of \tilde{x} : 1/4 (centre of the cloudy volume), 1/2 and 3/4 (centre of the initially dry volume) at R = -0.5 and different values of *Da*. Fig. 10 is analogous to Fig. 5, but plotted for R > -1.

At Da = 1 the mixing is very fast, which leads to a rapid decrease in LWC and in the 565 566 droplet concentration in the initially cloud volume and to an increase of these quantities in the initially droplet-free volume. As a result of the rapid mixing and homogenization, all the curves 567 coincide at point "A" (left panel). After this time instance, spatial homogeneous evaporation 568 takes place. Since at Da = 1 only partial, but not total, droplet evaporation occurs, the droplet 569 concentration remains unchanged even while LWC decreases. At Da = 50 and Da = 500, the 570 three curves coincide at the final point "F" only. At Da = 500, the relationship between the 571 droplet concentration and the mass becomes more linear (blue curve). The linear dependence is 572 consistent with the concept of extremely inhomogeneous mixing (see Pt1). Considerations 573 regarding the closeness of the $\tilde{N} - \tilde{q}$ relationship to the line 1:1 as a measure of 574 inhomogeneity of mixing made at R < -1 are also valid for R > -1. 575

576

577

5.2.2 Evolution of DSDs and the DSD parameters

Figure 11 presents examples of the DSD evolution at the center of the initially cloud volume ($\tilde{x} = 1/4$) (upper row) and of the initially droplet-free volume ($\tilde{x} = 3/4$) at R = -0.5and different values of Da. Several specific features of the DSD are notable. As a result of the rapid mixing at Da = 1 (left column), DSD become similar in both volumes already at $t = 0.317\tau_{pr}$ (black lines). Further evolution is similar in both volumes and is characterized by broadening of the DSD and its shifting and of the DSD toward smaller droplet sizes. This shift means a decrease in the mass at constant droplet concentration, which is typical ofhomogeneous mixing.

The initially monodisperse DSDs become polydisperse. The mechanism of the DSD broadening at Da = 1 is illustrated in **Figure 12**, showing the DSD at the earlier, inhomogeneous stage at different \tilde{x} . One can see that within very short periods when the spatial gradient of saturation deficit exists, droplets entering the initially droplet-free volume partially evaporate, reaching their minimal size at $\tilde{x} = 1$. In this way, a polydisperse DSD forms. As the mixing proceeds, DSD become spatially homogenized, as seen in the right panel of Fig. 12.

593 At Da = 50 and Da = 500, the DSD shapes substantially differ from those at Da = 1. There are two main differences: the peak of the distribution shifts only slightly (at Da = 50) or 594 595 does not shift at all (at Da = 500). At the same time, the DSD develops a long tail of small droplets. Since the mixing rate at these values of *Da* is slow, droplets penetrating deeper into 596 the initially dry volume remain there for long time and get smaller. As a result, at moderate 597 598 and large *Da*, a polydisperse DSDs form with droplet sizes ranging from zero to 1. Formation 599 of a long tail of small droplets in case of inhomogeneous mixing was simulated in direct 600 numerical simulation (DNS) by Kumar et al. (2012), as well as by means of "the explicitmixing parcel model" (EMPM) (Krueger et al., 1997; Su et al., 1998; Schlüter, 2006). 601

602 Figure 13 shows the spatial dependencies of the DSD dispersion (ratio of DSD r.m.s. width and the mean radius) at different time instances and different values of Da and R. One can see 603 that the dispersion increases with an increase in Da and in |R|. This behavior can be accounted 604 605 for by the fact that the DSD broadening toward smallest droplet size increases with the increase 606 in Da and in |R|. The DSD dispersion increases with time and with an increase in \tilde{x} , i.e. 607 further into the initially droplet free volume. At the same time, spatial homogenization takes 608 place, so at the final state at R = -0.5 the DSD dispersion reaches 0.11 at Da = 1 and about 0.2 609 at Da = 50 and Da = 500.

Observed DSD dispersion in different clouds typically ranges from 0.1 to 0.4 (Khain et al., 2000; Martin et al., 2004; Prabha et al., 2012) and can be caused the following factors: incloud nucleation (e.g. Khain et al., 2000; Pinsky and Khain, 2002), spatial averaging along aircraft traverses (Korolev, 1995) and non-symmetry in droplet nucleation/denucleation (Korolev, 1995). As seen in Fig. 13, this dispersion may be also caused by mixing at cloud edges at moderate and large *Da*. Hence, inhomogeneous mixing leads to DSD broadening.

The effective radius, r_{eff} , is an important DSD characteristic. According to the classical concept, r_{eff} remains unchanged during extremely inhomogeneous mixing, whereas decreases during homogeneous mixing. **Figure 14** shows spatial dependencies of r_{eff} at different time instances and different values of *Da* and *R*. At R = -0.1 (high RH in the surrounding volume) r_{eff} is similar for all values of *Da*. So, at high *R* (i.e., close to zero), the behaviour of r_{eff} does not allow to distinguish between mixing types.

At a given *R*, the final r_{eff} increases with increasing *Da*. For instance, at R = -0.5, r_{eff} at the final state differs from the initial r_{eff} value by less than 6% at Da = 500, while at Da = 1 r_{eff} decreases by 20%. At moderate and high *Da*, large gradients of r_{eff} exist during the mixing process. However, the gradient is high only in the initially droplet-free volume where r_{eff} decreases significantly due to the intense evaporation of droplets. Besides, r_{eff} growth very rapidly in the initially droplet free volume, so at high *Da* during most of the mixing time r_{eff} within the mixing volume becomes close to the initial r_{eff} value in the cloudy volume.

629

630

5.3 Delimitation between mixing types

Typically, the Da value is used as a criterion for delimitation between mixing types. Da = 1 is usually used as a boundary value separating homogeneous and inhomogeneous mixing. As shown in Section 4, mixing always starts as inhomogeneous. In the course of

mixing, the initial spatial gradients decrease and the air volumes either become identical or 634 635 remain different. In the former case, the second mixing stage is homogeneous. If inhomogeneity persists until the equilibrium state is established, mixing remains 636 637 inhomogeneous during the entire period. Both mixing stages can be characterized by duration, change in the droplet concentrations or LWCs, and other quantitative characteristics. These 638 characteristics are functions of two non-dimensional parameters R and Da, which can be 639 calculated and used for delimitation between mixing types. Since mixing between volumes 640 may turn from inhomogeneous into homogeneous before reaching the equilibrium state, it is 641 necessary to use some quantitative criteria to delimit mixing types. Below, delimitation is 642 performed for R > -1 which corresponds to partial evaporation of droplets by the end of 643 644 mixing.

645

646

5.3.1. Characteristic time periods of mixing

Three characteristic time periods of mixing are distinguished: a) mixing period T_{mix} , during which spatial gradients are smoothening (may be also called the homogenization period); b) period T_{ev} during which S < 0 and droplets evaporate until saturation is reached and c) the total mixing period T_{tot} that lasts until the final equilibrium stage is reached. In our analysis, all the three periods are assumed dimensionless quantities.

652 We use solution (28) for conservative function $\tilde{\Gamma}(\tilde{x}, \tilde{t})$ to define quantitatively time period 653 T_{mix} . The deviation of the solution from its final value $\Delta \tilde{\Gamma} = \tilde{\Gamma}(\tilde{x}, \tilde{t}) - \tilde{\Gamma}(\tilde{x}, \infty)$ at $\tilde{t} \to \infty$ can be 654 approximately estimated using the first term of the series expansion as

655
$$\begin{aligned} \left| \Delta \tilde{\Gamma} \right|_{\max} \approx \left| (1-R) \frac{\sin(\pi/2)}{\pi/2} \exp\left(-\frac{\pi^2 \tilde{t}}{Da} \right) \cos(\pi \tilde{x}) \right|_{\max} = \\ \left((1-R) \frac{2}{\pi} \exp\left(-\frac{\pi^2 \tilde{t}}{Da} \right) \end{aligned}$$
(37)

From Eq. (37) the estimation of T_{mix} can be written as

658
$$T_{mix} = -\frac{Da}{\pi^2} \ln \left[\frac{\pi}{2(1-R)} \left| \Delta \tilde{\Gamma} \right|_{max} \right]$$
(38a)

Suppose the value of the maximum deviation is $|\Delta \tilde{\Gamma}|_{\text{max}} = 0.02$. This is a small value compared to the initial leap of function $\tilde{\Gamma}$, which is equal to 1-R. At $|\Delta \tilde{\Gamma}|_{\text{max}} = 0.02$ the duration of the non-homogeneous stage is evaluated as

662
$$T_{mix} = -\frac{Da}{\pi^2} \ln \left[\frac{0.01\pi}{1-R} \right]$$
 (38b)

663

Several studies evaluate the evaporation time for droplets of a particular size using the equation for diffusion growth (e.g. Lehmann et al., 2009). In our study, the evaporation time duration T_{ev} is defined as the period during which the maximum deviation of supersaturation from zero exceeds the small value chosen as $\left|\Delta \tilde{S}\right|_{max} = 0.02$:

668
$$\left| \tilde{S}(\tilde{x}, T_{ev}) \right| \le \left| \Delta \tilde{S} \right|_{\max} = 0.02$$
 (39)

669

Although criterion (39) is rather subjective, it has an advantage over the criterion used by Lehmann et al. (2009), as Eq. (32) characterizes evaporation of the droplet population taking into account the simultaneous increase in supersaturation, but not of individual droplets of particular size at constant *S* as in Lehmann et al. (2009).

At the end of the mixing, both the thermodynamic equilibrium and the diffusion equilibrium are reached. Accordingly, the total time of mixing T_{tot} is evaluated as the maximum of the two time periods needed to achieve equilibrium $T_{tot} = \max\{T_{mix}, T_{ev}\}$. All the three characteristic time periods are normalized on the phase relaxation time, and, therefore, depend on the two non-dimensional parameters R and Da. The contours of the characteristic time durations T_{mix} , T_{ev} and T_{tot} in the Da - R diagrams are shown in **Figure 15**.

As follows from Eq. (38b), T_{mix} is proportional to Da. The dependence of T_{mix} on R is not 680 very strong, so T_{mix} slightly decreases with increasing R. This can be attributed to the fact that 681 the lower R, the smaller the initial inhomogeneity of function $\tilde{\Gamma}$ and the shorter the time to 682 align this inhomogeneity is. At small Da (high rate of homogenization of the volume), T_{ev} 683 depends largely on R. At large Da, T_{ev} depends substantially on Da, since the evaporation 684 685 rate depends on the number of droplets that diffuse to drier parts of the mixing volume. A comparison of Fig. 15c with Figs. 15a and 15b shows that at small Da, time T_{tot} is determined 686 by T_{ev} , while at large Da, T_{tot} is determined by T_{mix} . 687

688

689 5.3.2. Determination of boundaries between the mixing types on the R-Da plane

Several criteria can be proposed for delimitation between mixing types. We consider these criteria for R > -1. As discussed above, mixing always starts as inhomogeneous and late either become homogeneous or remains inhomogeneous till the final equilibrium state is established. At small *Da*, the homogenization takes place during $T_{mix} < T_{tot}$. The value of time fraction λ_1 of the inhomogeneous stage can serve as a criterion for definition of homogeneous mixing. This formula for the fraction can be written as

$$\delta 96 \qquad \lambda_1 = \frac{T_{mix}}{T_{tot}} \tag{40}$$

The case $\lambda_1 \le 0.5$, most time the mixing takes place according the homogeneous scenario and such regime is reasonable to regard as homogeneous mixing. If $\lambda_1(R, Da)$ changes within the range of $0.5 < \lambda_1 \le 1$, mixing appears to be intermediate. The criterion (40) depends on the nondimensional parameters *R* and *Da*. Figure 16a shows the boundaries separating mixing types _ on the *Da*-*R* plane. These boundaries separate all plane into several zones. At very small *R*, the duration of the phase transition is negligibly small. According to criterion (40), in this case mixing should be considered inhomogeneous, irrespective of the *Da* value. comparison of LWC variation rates due to different mechanisms. The mean normalized LWC (which is equal to the mean normalized liquid water mixing ratio) can be written as integral $\langle \tilde{q}(\tilde{t}) \rangle = \int_{0}^{1} \tilde{q}(\tilde{x}, \tilde{t}) d\tilde{x}$. The initial mean LWC is equal to $\langle \tilde{q}(t=0) \rangle = \frac{1}{2}$. The final equilibrium

Another criterion of delimitation between mixing types can be determined from a

LWC is equal to
$$\langle \tilde{q}(t=\infty) \rangle = \frac{1}{2}(1+R)$$
 (Eq. (30)). The total amount of liquid water that
evaporates in the course of mixing can be quantified by the difference between these two
values $\langle \tilde{q}(t=0) \rangle - \langle \tilde{q}(t=\infty) \rangle = -\frac{1}{2}R$. The amount of liquid water evaporated in the course of the
first inhomogeneous mixing stage is calculated by the equation
 $\langle \tilde{q}(t=0) \rangle - \langle \tilde{q}(T_{mix}) \rangle = \frac{1}{2} - \langle \tilde{q}(T_{mix}) \rangle$. Hence, parameter λ_2 which is a ratio of

704

714
$$\lambda_{2} = \frac{\left\langle \tilde{q}(t=0) \right\rangle - \left\langle \tilde{q}(T_{mix}) \right\rangle}{\left\langle \tilde{q}(t=0) \right\rangle - \left\langle \tilde{q}(t=\infty) \right\rangle} = \frac{2\left\langle \tilde{q}(T_{mix}) \right\rangle - 1}{R}$$
(41)

can serve as another possible criterion for delimitation between mixing types. This ratio characterizes the fraction of liquid water that evaporates at the initial inhomogeneous stage. Condition $\lambda_2 < 0.5$ in this case corresponds to homogeneous mixing, while condition $0.5 \le \lambda_2 < 1$ corresponds to intermediate mixing. We regard the case $\lambda_2 = 1$ as inhomogeneous mixing. Certainly, criterion λ_2 depends on the non-dimensional parameters *R* and *Da*. Fig. 16b illustrates delimitation between mixing types on the *Da*-*R* plane according to criterion λ_2 .

Comparison of Figs. 16a and 16b shows that both criteria lead to nearly similar separation of the Da - R plane into three zones corresponding to homogeneous, intermediate and inhomogeneous mixing. At the same time, the boundaries separating these zones are different depending on the delimitation criterion used. Nevertheless, it can be concluded that mixing can be considered homogeneous at Da below 4-10 and R < -0.1 and inhomogeneous at Da exceeding several tens.

"inhomogeneous mixing" (Burner and Brenguier, 2007) and "extremely 728 Terms 729 inhomogeneous mixing" (Lehmann et al., 2009; Gerber et al., 2008; Pt1) are used to denote the mixing regime when the relationship between the normalized values \tilde{N} and \tilde{q} is represented 730 by a straight 1:1 line, which is equivalent to the constant mean volume radius (in some studies, 731 732 the effective radius is used instead of the mean volume radius. According to the definition used 733 in the present study, extremely inhomogeneous mixing is the limiting case of inhomogeneous mixing when $Da \rightarrow \infty$. Despite the fact that the extremely inhomogeneous mixing is only an 734 idealization our approach allows to determine to what extent mixing can be considered to be 735 736 close to this limiting case. The measure of inhomogenity of mixing is the closeness of the $\tilde{N} - \tilde{q}$ relationship to the 1:1 straight line (see discussion above related to Figs. 5 and 10). 737

Figure 17a shows r.m.s. distance between the $\tilde{N} - \tilde{q}$ relationship and the 1:1 straight line, depending on *Da* and *R*. These dependences were calculated using the set of points \tilde{N}_i, \tilde{q}_i uniformly distributed over spatial interval [0 1] and time interval [0 T_{tot}]. The equation for

741 estimation is
$$\delta = \sqrt{\frac{1}{2M} \sum_{i=1}^{M} (\tilde{N}_i - \tilde{q}_i)^2}$$
, where *M* is the total number of points. This distance

corresponds to r.m.s. deviation of the normalized mean volume radius from 1. The dependences of the last deviation on Da and R and estimated as $\delta/3$ are shown in Fig. 17b. This estimation is based on the fact that the total mass of droplets is proportional to the cube of the mean volume radius. As expected, the distance decreases with increasing in Da. At large R, all the curves coincide indicating a degenerative case when type of mixing becomes indistinguishable.

We choose the value $\delta/3$ equal to 0.02 to determine the boundary of the extremely inhomogeneous mixing zone. The value of 0.02 corresponds to droplet radii deviation of a few tenths of a micron, which is so low that in in-situ measurements this case would always be attributed to extremely inhomogeneous mixing. In Fig.16 this boundary is marked by broken

line. The boundary shows that the mixing at *Da* exceeding several hundred can be attributed 751 to the extremely inhomogeneous. Between the boundary separating inhomogeneous mixing 752 from the intermediate one and the boundary separated inhomogeneous mixing from extremely 753 inhomogeneous there exists a wide zone of inhomogeneous mixing where the mean volume (or 754 755 the effective) radius may drop by 10% and more (Fig. 14), and where the DSD dispersion is substantial and the tail of small droplets is long enough (Fig. 11). Mixing diagrams currently 756 used for analysis of observed data (N-q dependences in the final equilibrium state of mixing) 757 758 do not contain this zone which, therefore, has remained unrecognized and uninvestigated.

759

760

6. Summary and conclusions

In this study, inhomogeneous turbulent mixing is investigated using a simple a 1D model of mixing between a saturated cloud volume and an undersaturated droplet-free volume. The mixing is simulated by solving a diffusion-evaporation equation written in the non-dimensional form. For simplicity, the initial volumes of cloudy and droplet-free air were assumed to be equal, and the initial DSD in the cloudy volume was assumed monodisperse.

766 Analysis of the diffusion-evaporation equation shows that the time-dependent process of 767 mixing and the final equilibrium state depend on two non-dimensional parameters. The first 768 parameter R, referred in this paper as potential evaporation parameter (PEP) is proportional to 769 the ratio between the saturation deficit in the initially droplet-free volume and the initial liquid water content in the cloudy volume. At R < -1, the final state is characterized by complete 770 771 droplet evaporation and a spatially homogeneous saturation deficit, which indicates dissipation of the cloudy volume. At R > -1, the final state is characterized by existence of droplets and 772 773 zero saturation deficit (RH=100%). In this case, the cloud volume expands after mixing with the entrained air. At small values of |R| (e.g., when RH in the entrained volume is close to 100%), 774 the effect of droplet evaporation on microphysics is small, and, formally, this kind of mixing 775 should be regarded as extremely inhomogeneous. Strictly speaking, this is a degenerate case, 776

when homogeneous and inhomogeneous mixing cannot be distinguished (see also Pt. 1). At R = 0, the droplet population turns into a passive admixture and its turbulent diffusion will be the same as different thermodynamic parameters.

The second parameter is the *Damkölher* number (*Da*) which is the ratio between the 780 characteristic mixing time and the phase relaxation time. This parameter compares the rates of 781 782 spatial diffusion and evaporation. Parameter Da (Eq. (23)) is logically appears in the nondimensional form of the diffusion-evaporation equation showing that Da is the ratio of the 783 mixing time defined as $\tau_{mix} = \frac{L^2}{K}$, to the initial drop relaxation time. The expression for this non-784 dimensional parameter clearly shows that since we consider an ensemble of evaporating droplets, 785 the drop relaxation time evaluated just before the mixing is the characteristic time scale of 786 787 inhomogeneous mixing process. In several studies (e.g., Baker and Latham, 1979; Burnet and Brenguier, 2007; Andejchuk et al., 2009) a question was raised as to which time scale should be 788 789 used in formulation of the Damkölher number: the time of an individual droplet evaporation at constant saturation deficit, or the phase relaxation time. This study, as well Pt. 2 show that the 790 phase relaxation time is the answer. The mixing time is introduced via the turbulent diffusion 791 792 coefficient which is a natural measure characterizing the diffusion rate and, in particular, determines the propagation rate of the fronts in the fields of droplet concentration and other 793 microphysical parameters. The turbulent diffusion coefficient is widely used to describe mixing 794 795 in cloud models at resolved scales.

The analysis was performed within a wide range of Da (from 1 to 500) and of R (from -1.5 to -0.1). The final LWC and the humidity in the mixing volume are determined by the mass conservation and do not depend on Da (see also Pt. 1 and Pt. 2). At the same time, the droplet concentration, as well as the shape of DSD and their parameters strongly depend on Da.

800 It is shown that the mixing of air volumes with initially different thermodynamical and 801 microphysical parameters consists of two stages characterized by two time periods: the time

during which microphysical characteristics become uniform over the total mixing volume T_{mix} , 802 and the time during which zero saturation deficit is reached (at R > -1), T_{ev} . At $\tilde{t} < T_{mix}$, the 803 spatial gradients of the microphysical values remain and the mixing regime can be regarded as 804 inhomogeneous. At $\tilde{t} > T_{mix}$, droplet evaporation, if it occurs at all, takes place within a 805 spatially homogeneous medium, so all the droplets in the mixing volume experience equal 806 saturation deficit. This regime can be regarded as homogeneous. It is shown, therefore, that at 807 808 small Da mixing between two volumes that starts as inhomogeneous can become 809 homogeneous towards the end of mixing.

810 This finding allows to delimit between mixing types. We presented two quantitative criteria on the Da-R plane that allow to delimit three mixing regimes: homogeneous, intermediate 811 812 and inhomogeneous. These criteria are based on comparison of the characteristic duration mixing and the evaporation rates. According to the criteria, at *Da* below about 5, mixing can 813 814 be regarded as homogeneous, i.e. the main microphysical changes take place during the homogeneous stage. At 5 < Da < 50, the changes in the microphysical parameters are more 815 significant at the inhomogeneous stage than at the homogeneous stage. In this case, the mixing 816 817 can be regarded as intermediate. Finally, at Da exceeding several tens, the spatial microphysical gradients remain until the final equilibrium stage is reached. In this case, the 818 819 mixing can be regarded as inhomogeneous. At Da exceeding a few hundred the deviations from predictions based of the classical concept of extremely inhomogeneous become relatively 820 821 small, which justifies attribute regarding this mixing as extremely inhomogeneous.

On the whole, the results of the present study are in line with the classic concepts defining homogeneous and inhomogeneous mixing types. However, several important points emerge from our work show serious limitations of classical concepts. A comparison of the classical concepts and the present study is presented in **Table 2**. Analysis of Tab. 2 shows the following. a) In contrast to many studies that analyze only the hypothetical final (equilibrium) state of mixing (Barnet and Brenguier, 2007; Gerber et al., 2008; Morrison and Grabowski, 2008; Hill et al., 2009), we consider the entire time-dependent processes of mixing and evaporation. At moderate and high Da, the mixing can last several minutes. In in-situ observations, we see mostly non-equilibrium stages which may account for a rather wide scattering of mixing diagrams even at the same values of Da (e.g., Lehmann et al., 2009).

Note that time dependent mixing was also considered in several studies (e.g. Baker et al.,
1980; Baker and Latham, 1982; Jeffery and Reisner, 2006; Krueger et al., 1997; Kumar et al.,
2012) using different approaches and numerical models. These studies, however, do not contain
analysis on non-dimensional diffusion-evaporation equation.

b) It is also shown in the study that the slopes of the $\tilde{N} - \tilde{q}$ relationship (between the 836 normalized droplet concentration and LWC) tends to the 1:1 line with increasing Da. The 837 838 closeness can be considered as a measure of extremely inhomogeneous mixing in terms of the classical concept (see Pt. 1). It has been found that the slope of the $\tilde{N} - \tilde{q}$ relationship depends 839 on the LWC and, accordingly, on time. At large LWC, \tilde{q} changes with time faster than \tilde{N} , 840 841 while at low LWC the concentration changes faster. Although mixing types are usually separated into homogeneous and extremely inhomogeneous, we have shown that there are wide 842 843 ranges of Da and R at which mixing should be considered intermediate or inhomogeneous, 844 but not extremely inhomogeneous. Within these ranges the effective radius can change by more than 10-15%. Standard mixing diagrams do not include this range that, to our knowledge, has 845 846 never been investigated despite the fact that multiple in-situ measurements indicate its 847 existence (e.g. Lu et al. 2014)

c) Many studies assume the existence of pure homogeneous mixing during which the initially monodisperse DSD remains monodisperse. Our study shows that at the very beginning, mixing is always inhomogeneous. This inhomogeneous stage leads to formation of a polydisperse DSD that broadens in the course of droplet evaporation. Hence, even at Da = 1the initially monodisperse spectrum becomes polydisperse. d) It is shown that at small *Da*, mixing includes both inhomogeneous and homogeneous stages, which means that type of mixing can change during the mixing process.

e) The classical concept assumes that the effective radius always decreases during 855 homogeneous mixing. Assuming an initially monodisperse DSD, we have found this 856 857 conclusion largely valid, with the exception small R. At the same time, it was shown in Pt. 2 that during homogeneous mixing, the effective radius can decrease, remain constant or increase 858 depending of the initial DSD shape. Thus, a decrease in the effective radius during mixing 859 cannot always be considered an indication of homogeneous mixing. Similarly, the invariability 860 the effective radius during mixing in the process cannot always be considered an indication of 861 862 extremely inhomogeneous mixing.

63 f) It is generally assumed that during homogeneous mixing droplet concentration remains 64 unchanged. In the present study, as well as in Pt. 2, it is shown that since mixing leads to a 65 polydisperse DSD, the smallest droplets may completely evaporate. At R < -1, the DSD 66 becomes very wide and all the droplets, the smallest ones first, evaporate.

g) It is generally assumed that inhomogeneous mixing does not alter DSD shape, but only 867 decreases droplet concentration. The present study showed that inhomogeneous mixing 868 significantly changes the DSD shape. DSD were found to be quite different in different regions 869 870 of mixing volumes. The main feature is the DSD broadening toward small droplet size, so the relative dispersion grows up to 0.2-0.3. These values are quite close to those observed in 871 atmospheric clouds (Khain et al., 2000). Elongated tails of small droplets during mixing were 872 873 simulated by Schlüter (2006) who described turbulent diffusion following to Kruger et al., (1997) and Su et al., (1998) as well as by Kumar et al. (2012) using DNS. We see that 874 formation of a polydisperse DSD is a natural result of inhomogeneous mixing and, therefore, 875 inhomogeneous mixing is an important mechanism of DSD broadening. A significant impact 876 877 of mixing on DSD shape was found identified in multiple studies, beginning with Warner 878 (1973).

h) The effective radius has been assumed to remain constant during extremely inhomogeneous mixing. Our results indicate that, indeed, at the final equilibrium stage at comparatively high RH the effective radius is close to that in the initially cloudy volume (especially at high Da). At the same time, we found that the effective radius varies in size and is smaller in the initially droplet-free volumes.

The results obtained in parts Pt1 and Pt 2, and especially in the current study (Pt 3) dedicated to analysis of turbulent mixing mechanisms in clouds determine the directions for future work. Since the widely used mixing diagrams show only a hypothetical equilibrium state, but not the instantaneous state of mixing that likely correspond to transition periods, the efficiency of the standard mixing diagrams is questionable. Moreover, the standard diagrams miss a very important mixing regime, namely, inhomogeneous mixing that occurs between two limiting cases of homogeneous and extremely inhomogeneous mixing (Fig. 16).

891 We believe that the results obtained will help to improve understanding and interpretation 892 of mixing process both in in-situ measurements and modeling. The approach allows to investigate the relationship between the main microphysical parameters typical of 893 inhomogeneous mixing, that differ from those in the limiting cases of extremely 894 inhomoheneous mixing. In addition, utilization of polydisperse DSD when solving diffusion-895 896 evaporation equation allows to investigate the role of the initial DSD shape in mixing. In-situ 897 measurements (e.g., Burnet and Brenguier, 2007; Gerber et al., 2008; Lehmann et al., 2009) and numerical models (Magaritz-Ronen et al., 2016) show a wide scattering of data on the 898 scattering diagrams. We expect location of various points on the diagrams (e.g. r_v^3 vs. dilution 899 rates) depends on the shape of the initial DSDs and characterizes the stage of mixing. The 900 method applied in the study allows investigation of evolution of DSD moments over space and 901 902 time.

Recently, there has been vigorous discussions concerning the possible existence of high
humidity layer near cloud edges that might affect mixing of cloud with its surrounding (Gerber

905	et al., 2008; Lehmann et al., 2009). In our opinion, this layer does exist and forms as a result of		
906	turbulent mixing of cloud with surrounding dry air, accompanied by complete droplet		
907	evaporation. The approach developed in the present paper allows to analyze formation of such		
908	humid layers.		
909	We believe that the results obtained in this study will foster the development of physically		
910	grounded parameterization of mixing in cloud models.		
911			
912	Acknowledgements		
913	This research was supported by the Israel Science Foundation (grant 1393/14), the Office of		

Science (BER), the US Department of Energy Award DE-SC0006788 and the Binational US-

- Israel Science Foundation (grant 2010446). Dr. Korolev's participation was supported by Environment Canada.

- Appendix. List of symbols
- **Table A here**

References 931 Andrejczuk M., W. Grabowski, S. P. Malinovski, P. K. Smolarkiewicz, 2009: Numerical 932 Simulation of Cloud-Clear Air Interfacial Mixing: Homogeneous versus Inhomogeneous 933 Mixing. J. Atmos. Sci., 66, 2993-2500. 934 935 Baker, M., and J. Latham: The evolution of droplet spectra and the rate of production of embryonic raindrops in small cumulus clouds. J. Atmos. Sci., 36, 1612–1615, 1979. 936 Baker, M., R. G. Corbin, and J. Latham: The influence of entrainment on the evolution of 937 938 cloud drop spectra: I. A model of inhomogeneous mixing. Quart. J. Roy. Meteor. Soc., 106, 581-939 598, 1980. Baker M. B. and J. Latham: A diffusive model of the turbulent mixing of dry and cloudy 940 air Quart. J. R. Met. Soc., 108, 871-898, 1982 941 Burnet, F., and J.-L Brenguier, Observational study of the entrainment-mixing process in 942 943 warm convective cloud, J. Atmos. Sci., 64, 1995–2011, 2007. Blyth, A. M., Choularton, T. W., Fullarton, G., Latham, J., Mill, C. S., Smith, M. H., and 944 945 Stromberg, I. M.: The Influence of entrainment on the evolution of cloud droplet spectra. 2. 946 Field experiments 5 at Great Dun Fell, Q. J. Roy. Meteor. Soc., 106, 821-840, 1980. Boffetta G. and I. M. Sokolov: relative dispersion in fully developed turbulence: The 947 Richardson's law and intermittency correction. Phys. Rev. Let., 88, 094501, 2002. 948 949 Denvich B. J., P. Bartello, J-L. Brenguier, L.R. Collins, W.W. Grabowski, R.H.A. Ijzermans, S.P. Malinovski, M.W. Reeks, J.C. Vassilicos, L-P. Wang, and Z. Warhaft: Droplet 950 growth in warm turbulent clouds. Q. J. Roy. Meteorol. Soc., 138, 1401-1429, 2012 951 Gerber H, Frick G, Jensen J.B, and Hudson J.G.: Entrainment, mixing, and microphysics in 952 trade-wind cumulus. J. Meteorol. Soc. Jpn., 86A. 87-106, 2008. 953

954	Hill, A. A., G. Feingold, and H. Jiang: The Influence of Entrainment and Mixing		
955	Assumption on Aerosol-Cloud Interactions in Marine Stratocumulus. J. Atmos. Sci., 66, 1450-		
956	1464, 2009.		
957	Jeffery, C.A., and J.M. Reisner: A study of cloud mixing and evolution using PDF methods.		
958	Part I: Cloud front propagation and evaporation. J. Atmos. Sci., 63, 2848-2864, 2006.		
959	Khain, A. P., M. Ovchinnikov, M. Pinsky, A. Pokrovsky, and H. Krugliak: Notes on the		
960	state-of-the-art numerical modeling of cloud microphysics. Atmos. Res. 55, 159-224, 2000.		
961	Korolev, A.V.: The influence of suresaturation fluctuations on droplet size spectra		
962	formation. J. Atmos. Sci., 52, 3620-3634, 1995.		
963	Korolev A., and I. Mazin: Supersaturation of water vapor in clouds, J. Atmos. Sci., 60,		
964	2957-2974, 2003.		
965	Korolev A., A. Khain, M. Pinsky, and J. French: Theoretical study of mixing in liquid		
966	clouds. Part 1: classical concept. Atmos. Chem. Phys., 2015 (submitted)		
967	Kovetz, A., and B. Olund: The effect of coalescence and condensation on rain formation in		
968	a cloud of finite vertical extent. J. Atmos. Sci., 26, 1060–1065, 1969.		
969	Krueger S., CW. Su and P. McMurtry: Modeling entrainment and finescale mixing in		
970	cumulus clouds. J. Atmos. Sci., 54, 2697-2712, 1997		
971	Kumar B., F. Janetzko, J. Schumacher, and R. A. Shaw: Extremely responses of a coupled		
972	scalar-particle system during turbulent mixing. New J. of Phys. 14, 115020, 2012		
973	Latham, J. and Reed, R. L.: Laboratory studies of effects of mixing on evolution of cloud		
974	droplet spectra, Q. J. Roy. Meteor. Soc., 103, 297-306, 1977.		
975	Lehmann, K., H. Siebert, and R. A. Shaw: Homogeneous and inhomogeneous mixing in		
976	cumulus clouds: Dependence on local turbulence structure. J. Atmos. Sci., 66, 3641-3659, 2009.		
977	Lu, C., Y. Liu, S. Niu, and S. Endo (2014), Scale dependence of entrainment-mixing		
978	mechanisms in cumulus clouds, J. Geophys. Res. Atmos., 119, 13,877-13,890, doi:10.1002/		
979	2014JD022265.		

980	Magaritz-Ronen L., A. Khain and M. Pinsky, 2016: About the Horizontal Variability of
981	Effective Radius in Stratocumulus Clouds. J. Geophys. Res. (in revision)
982	Martin G.M., D. W. Johnson and A. Spice: The measurements and parameterization of
983	effective radius of droplets in warm stratocumulus clouds. J. Atmos. Sci., 51, 1823-1842, 1994.
984	Monin, A.S. and Yaglom, A.M.: "Statistical Fluid Mechanics: Mechanics of Turbulence",
985	vol. 2, MIT Press. 1975
986	Morrison, H., and W. W. Grabowski: Modeling supersaturation and subgrid-scale mixing
987	with two-moment bulk warm microphysics. J. Atmos. Sci., 65, 792-812, 2008.
988	Pinsky, M. and A. P. Khain: Effects of in-cloud nucleation and turbulence on droplet
989	spectrum formation in cumulus clouds. Quart. J. Roy. Meteorol. Soc., 128, 1-33, 2002.
990	Pinsky M., I. P. Mazin, A. Korolev, and A. Khain: Supersaturation and diffusional droplet
991	growth in liquid clouds. J. Atmos. Sci., 70, 2778-2793, 2013.
992	Pinsky M., I. P. Mazin, A. Korolev and A. Khain: Supersaturation and diffusional droplet
993	growth in liquid clouds: Polydisperse spectra. J. Geophys. Res., Atmospheres, 119, 12,872-
994	12,887, 2014.
995	Pinsky, M., Khain, A., Korolev, A., and Magaritz-Ronen, L.: Theoretical study of mixing in
996	liquid clouds. Part 2: Homogeneous mixing. Atmos. Chem. Phys., , 2016 (submitted)
997	Polyanin A. D. and V. F. Zaitsev: Handbook of nonlinear partial differential equations.
998	Chapman & Hall/CRC, 809 pp., 2004
999	Prabha V. T., S. Patade, G. Pandithurai, A. Khain, D. Axisa, P. Pradeep Kumar, R. S.
1000	Maheshkumar, J. R. Kulkarni, and B. N. Goswami: Spectral width of premonsoon and
1001	monsoon clouds over Indo-Gangetic valley during CAIPEEX, J. Geop. Res. 117, D20205,
1002	doi:10.1029/2011JD016837 ', 2012
1003	Pruppacher, H.R., Klett, J.D Microphysics of Clouds and Precipitation. 2nd edn. Oxford
1004	Press, 914 p., 1997

1005	Rogers R. R. and Yau M. K: A Short Course in Cloud Physics, Pergamon press. 293pp.,
1006	1989
1007	Schlüter M. H.: The effects of entrainment and mixing process on the droplet size
1008	distribution in cumuli. A thesis submitted to the faculty of The University of Utah in partial
1009	fulfillment of the requirements for the degree of Master of Science, Department of
1010	Meteorology, The University of Utah, 92 pp., 2006
1011	Su CW., S.K. Krueger, P.A. McMurtry and P.H. Austin: Linear eddy modeling of droplet
1012	spectral evolution during entrainment and mixing in cumulus clouds. Atmos. Res., 47-48, 41-
1013	58, 1998.
1014	Telford, J.W., and S. K. Chai: A new aspect of condensation theory. Pageoph, 118, 720-
1015	742,1980
1016	Warner, J.: The microstructure of cumulus cloud. Pt. I, General features of the droplet
1017	spectrum, J. Atmos. Sci., 26, 1049-1059, 1969.
1018	Warner, J The microstructure of cumulus cloud: Part 4: The effect on the droplet spectrum
1019	of mixing between cloud and environment. J. Atmos. Sci. 30, 256–261, 1973.
1020	
1021	
1022	
1023	
1024	
1025	
1026	
1027	
1028	
1029	
1030	

1031 Table 1. Main parameters of the problem and their non-dimensional forms*

Quantity	Symbol	Non-dimensional form	Range of normalized
			values
Time	t	$\tilde{t} = \frac{t}{\tau_0}$	[o0]
Distance	x	$\tilde{x} = \frac{x}{L}$	[01]
Square of drop radius	σ	$\tilde{\sigma} = \frac{\sigma}{r_0^2}$	[01]
Droplet concentration	N	$\tilde{N} = \frac{N}{N_1}$	[01]
Liquid water mixing ratio	q	$\tilde{q} = \frac{q}{q_1}$	[01]
Distribution of square of drop	$g(\sigma)$	$\tilde{g}(\tilde{\sigma}) = \frac{r_0^2}{N_1} g(\sigma)$	
radius			
Conservative function	Γ	$\tilde{\Gamma} = \frac{\Gamma}{A_2 q_1}$	[<i>−∞</i> 1]
Supersaturation	S	$\tilde{S} = \frac{S}{A_2 q_1}$	[<i>−∞</i> 0]
Relaxation time	$ au_{pr}$	$\tilde{\tau}_{pr} = \frac{\tau_{pr}}{\tau_0}$	[1∞]
Damkölher number	Da	$Da = \frac{\tau_{mix}}{\tau_0} = \frac{L^2}{K\tau_0}$	[0∞]
Potential evaporation	R	$R = \frac{S_2}{A_2 q_1}$	[<i>−∞</i> 0]
parameter (PEP)			

*All normalized values depend on the initially given values of L, N_1 , r_0 , A_2 , S_2 and K

-	The present study	
Only the final equilibrium state is typically	The mixing period can last several minutes and	
analyzed; results of in-situ observations are	more. The microphysical structure of the	
interpreted assuming the equilibrium state.	mixing volumes during this period can differ	
	substantially from that at the final state	
Types of mixing are separated into	There are the wide ranges of Da and R values,	
homogeneous and extremely	at which mixing can be regarded as	
inhomogeneous.	intermediate or inhomogeneous (but not	
	extremely inhomogeneous).	
Mixing can start as purely homogeneous	Any mixing starts with the inhomogeneous	
	stage	
Homogeneous mixing leads to a DSD shift	Homogeneous mixing does not always lead to	
to small droplet sizes	the DSD shift to small droplet sizes (Pt. 2). The	
	shift depends on the DSD shape.	
Mixing can be analyzed within the	Mixing always leads to formation of	
framework of a monodisperse DSD	polydisperse DSD	
In the course of homogeneous mixing,	In the course of homogeneous mixing, droplet	
droplet concentration remains constant	concentration does not always remain constant	
	(Pt. 2)	
Extremely inhomogeneous mixing does not	Inhomogeneous mixing (including extremely	
change the DSD shape	inhomogeneous) leads to broadening of the	
	DSD towards small sizes	
In the course of inhomogeneous mixing, the	The effective radius varies only slightly (5-20	

	effective radius remains constant	%) in the initially cloud volume. The effective radius rapidly increases in the initially droplet- free volume, approaching the value of effective radius in the cloud volume. With increasing <i>Da</i> , the difference between the values of the
		effective radius in the initially cloud volume
		and that at the final state decreases in agreement with the classic concept.
1040		agreement with the classic concept.
1041		
1042		
1043		
1044		
1045		
1046		
1047		
1048		
1049		
1050		
1051		
1052		
1053		
1054 1055		
1055		
1057		

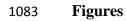
1058Tab. A. List of symbols

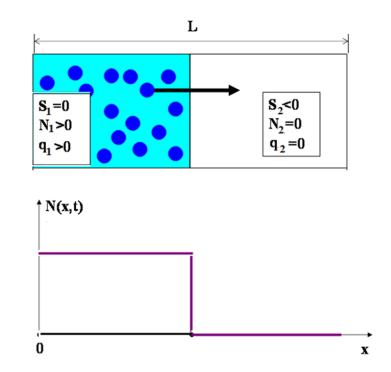
1059 ("nd" means non-dimensional)

Symbol	Description	Units
A_2	$\frac{1}{q_v} + \frac{L_w^2}{c_p R_v T^2} \text{, coefficient}$	nd
a_0, a_n	the Fourier series coefficients	nd
С	the Richardson's law constant	nd
C _p	specific heat capacity of moist air at constant pressure	$J kg^{-1}K^{-1}$
\mathcal{D}	coefficient of water vapour diffusion in the air	$m^2 s^{-1}$
Da	the Damkölher number	nd
е	water vapor pressure	N m ⁻²
e_s	saturation vapour pressure above a flat water surface	N m ⁻²
F	$F = \frac{\rho_w L_w^2}{k_a R_v T^2} + \frac{\rho_w R_v T}{e_s(T)\mathcal{P}} , \text{ coefficient}$	m ⁻² s
f(r)	droplet size distribution	m ⁻⁴
$g(\sigma)$	distribution of square radius	m-5
$\widetilde{g}(ilde{\sigma})$	normalized distribution of square radius	nd
k _a	coefficient of air heat conductivity	$J m^{-1}s^{-1}K^{-1}$
K	turbulent diffusion coefficient	m^2s^{-1}
L	characteristic spatial scale of mixing	m
L_{w}	latent heat for liquid water	J kg ⁻¹
m_{α}	moment of DSD of order α	m ⁻³
N	droplet concentration	
\tilde{N}	normalized droplet concentration	nd
N_1	Initial droplet concentration in a cloud volume	m ⁻³
р	pressure of moist air	N m ⁻²
q	liquid water mixing ratio	kg/kg
$q_{_1}$	Initial liquid water mixing ratio in a cloudy volume	kg/kg
$q_{_{V}}$	water vapor mixing ratio	kg/kg
\tilde{q}	normalised liquid water mixing ratio equal to	nd

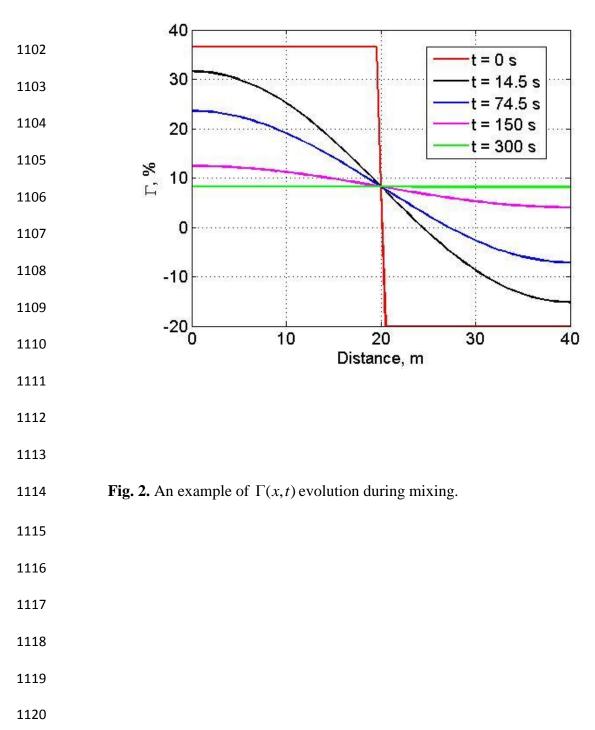
	normalized LWC	
r	droplet radius	m
r ₀	initial droplet radius	m
r ₀	mean droplet radius	m
r _v	mean volume radius	m
R	$\frac{S_2}{A_2q_1}$, potential evaporation parameter (PEP)	nd
R_a	specific gas constant of moist air	J kg ⁻¹ K ⁻¹
R_{ν}	specific gas constant of water vapor	J kg ⁻¹ K ⁻¹
S	$e/e_w - 1$, supersaruration over water	nd
$ ilde{S}$	normalized supersaturation	nd
S ₂	Initial supersaturation in a dry volume	nd
${ ilde S}_{ m max}$	maximal normalized supersaturation	nd
Т	temperature	K
T _{mix}	normalized duration of inhomogeneous stage	nd
T_{ev}	normalized duration of evaporation	nd
T_{tot}	normalized duration of mixing	nd
t	time	S
ĩ	non-dimensional time	nd
x	distance	m
ĩ	non-dimensional distance	nd
λ_1, λ_2	criteria of delimitation between the types of mixing	nd
ε	turbulent dissipation rate	m ² s ⁻³
$\Gamma(x,t)$	conservative function	nd
Γ	normalized conservative function	nd
$ ho_a$	air density	kg m ⁻³
ρ_w	density of liquid water	kg m ⁻³
σ	square of droplet radius	m ²
$ ilde{\sigma}$	normalized square of droplet radius	nd
$ au_{pr}$	phase relaxation time	s

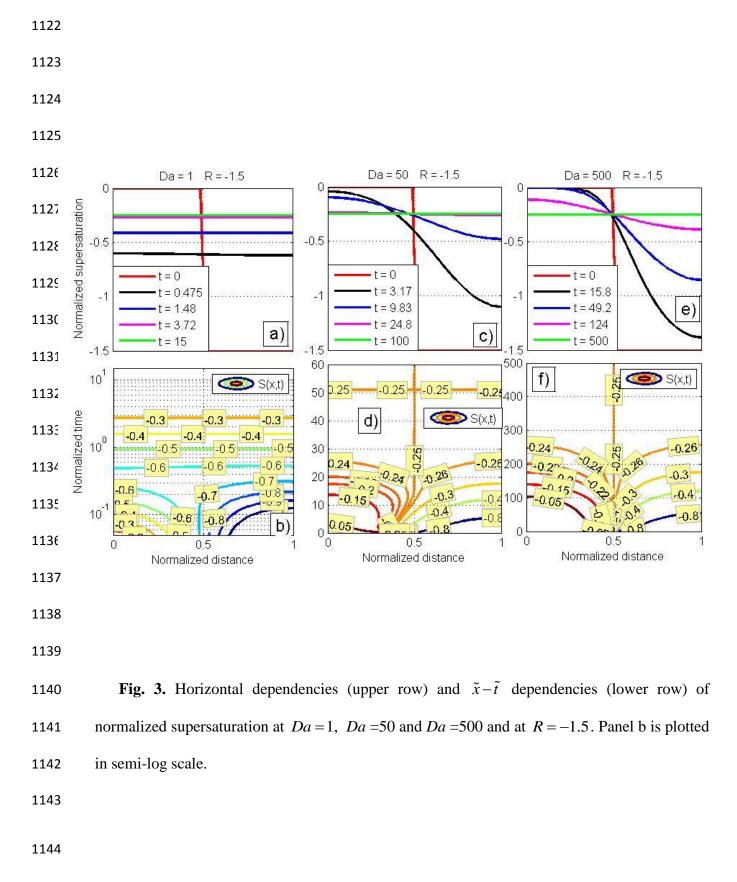
	$ ilde{ au}_{pr}$	normalized phase relaxation time	nd
	$ au_{mix}$	characteristic time of mixing	S
	$ au_0$	Initial time scale	S
1060			I
1061			
1062			
1063			
1064			
1065			
1066			
1067			
1068			
1069			
1070			
1071			
1072			
1073			
1074			
1075			
1076			
1077			
1078			
1079			
1080			
1081			
1082			





1089 Fig.1. The schematic illustration of the 1D mixing problem considered in the study. The 1090 initial state at t = 0 is illustrated. The left volume of length L/2 is a saturated cloudy volume; 1091 the right volume is a non-saturated air volume from the cloud environment.





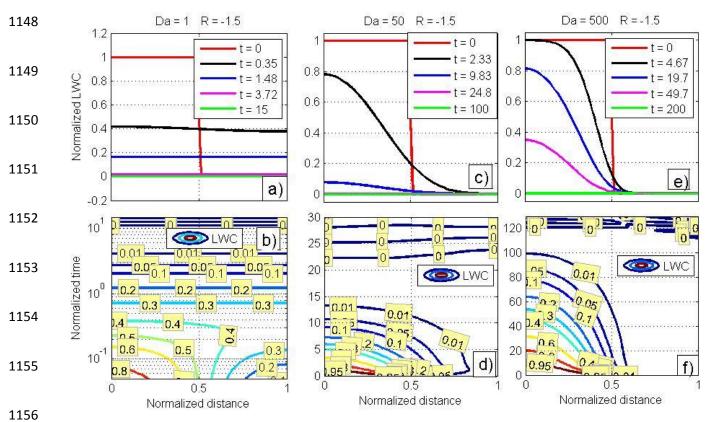


Fig. 4. The same as in Fig. 3, but for normalized LWC. Left bottom panel is plotted insemi-log scale.

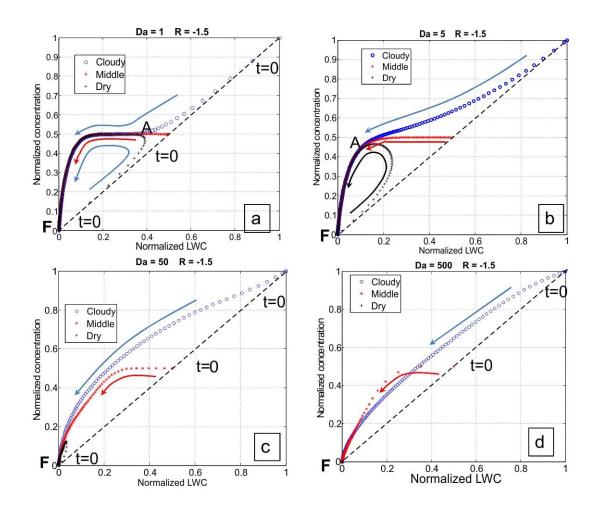




Fig. 5. Dependencies of normalized values of droplet concentration on normalized LWC 1165 at different Da and R = -1.5. Blue symbols mark the centre of the cloudy volume ($\tilde{x} = 1/4$), 1166 red symbols mark the interface between the cloudy volume and the dry volume ($\tilde{x} = 1/2$), and 1167 black crosses mark the centre of the initially droplet-free volume ($\tilde{x} = 3/4$). Symbols are 1168 plotted at different time instances. Symbols at t=0 show initial values of droplet concentration 1169 1170 and LWC at the three values of \tilde{x} . Arrows show the direction of movement of the points at the diagram with time. Point "A" marks the beginning of the spatially homogeneous stage, $\tilde{t} = T_{mir}$. 1171 Point "F" marks the final state. The dashed line indicates the relationship between \tilde{N} and \tilde{q} in 1172 extremely inhomogeneous mixing (according to the classical concept). 1173

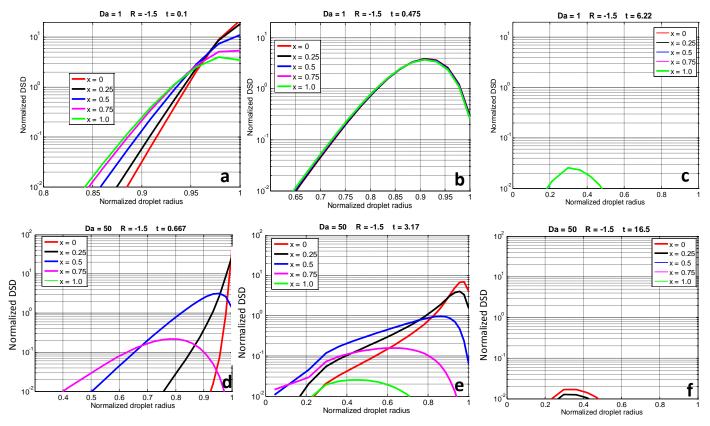
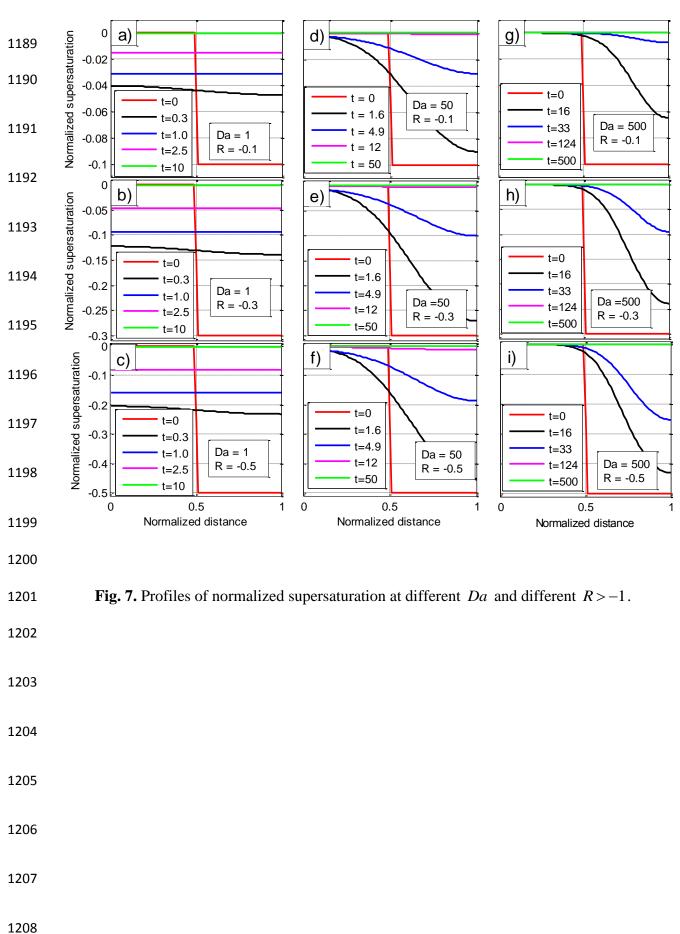
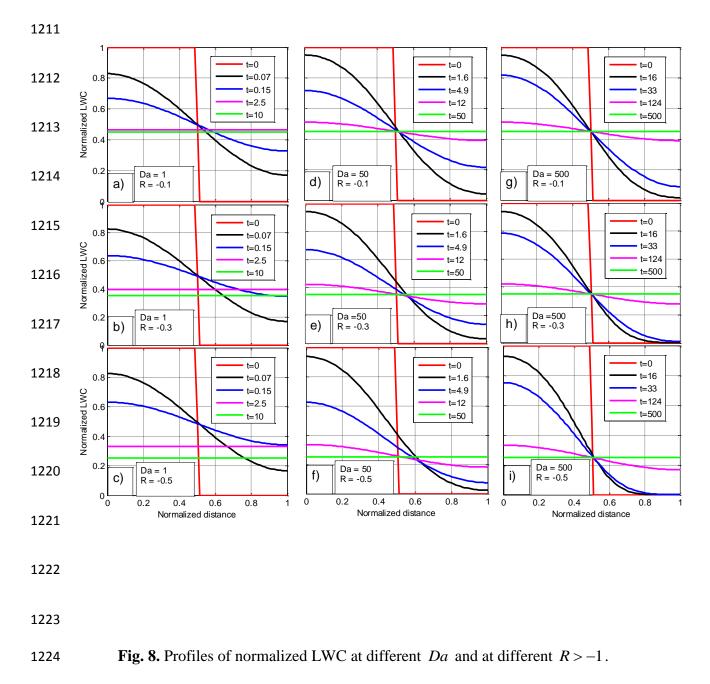


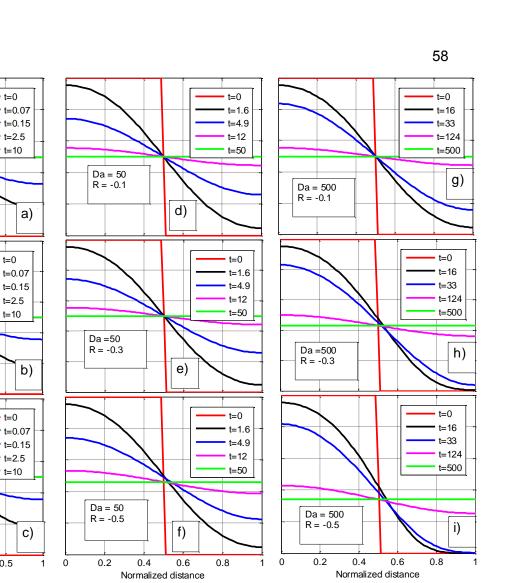
Fig. 6 Time evolution of DSD during droplet evaporation at Da = 1 (upper row) and 1180 Da = 50 (bottom row). In each panel, the normalized DSD are shown at different values of 1181 horizontal coordinate \tilde{x} . Different panels show DSD at different time instances.











Normalized concentration 9.0 9.0 7.0 8.0

q

Normalized concentration 9.0 9.0 7.0 7.0

9

Normalized concentration 9.0 9.0 7.0 7.0

0 L 0

Da = 1 R = -0.1

Da = 1 R = -0.3

Da = 1 R = -0.5

0.5

1

Normalized distance

t=0

t=2.5

t=10

t=0

t=2.5

t=10

t=0

t=10

0.5

0

1230

1231

1232

1233

1234

1235

1236

1237

1238

1239

1240

1241

1242

1243

- 1245
- 1246
- 1247

Fig. 9. Profiles of normalized droplet concentration at different Da and at different R > -1. 1248

- 1249
- 1250
- 1251
- 1252
- 1253

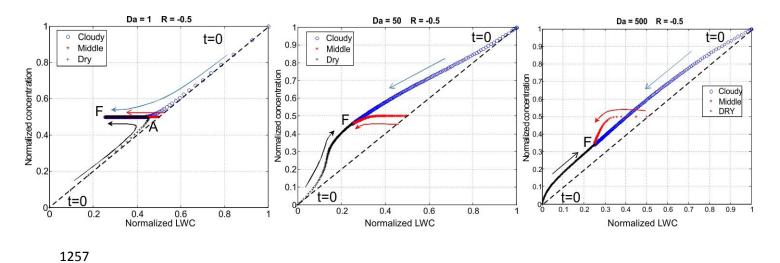
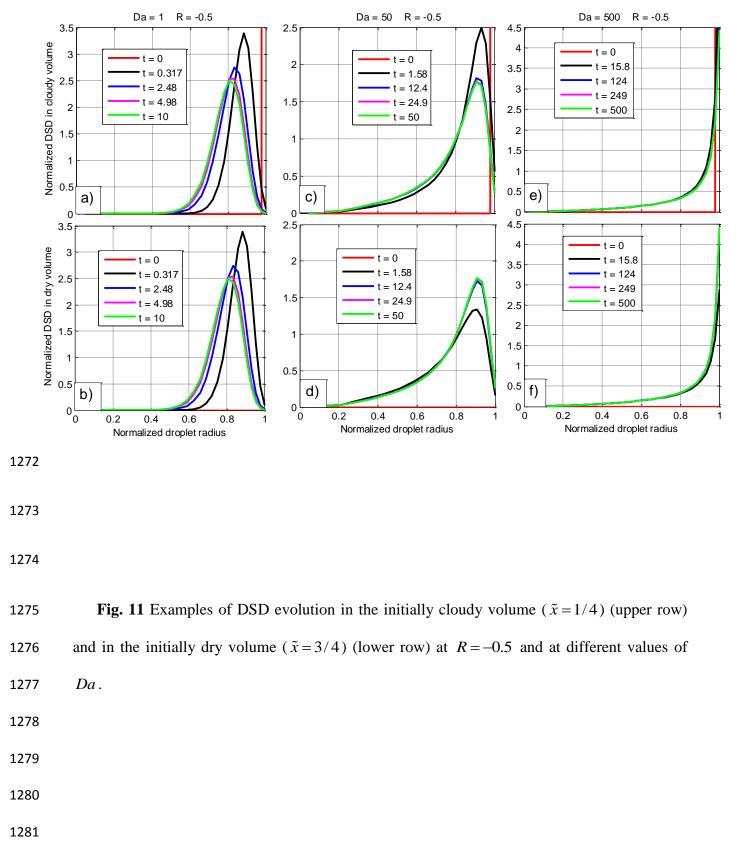
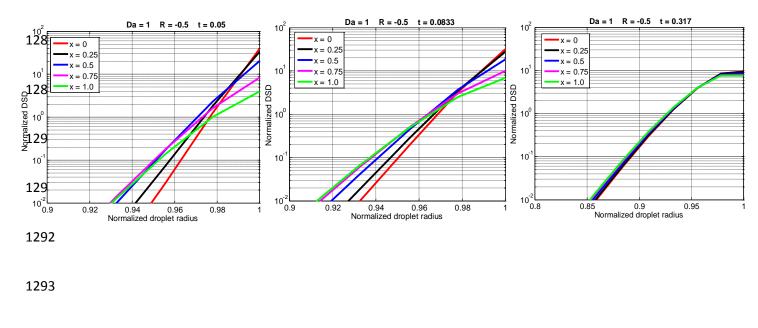




Fig. 10. Dependencies of normalized values of droplet concentration on normalized LWC at different *Da* and at R = -0.5. Blue circles mark the centre of the cloudy volume ($\tilde{x} = 1/4$), red symbols mark the initial interface ($\tilde{x}=1/2$) and black crosses mark the centre of the initially dry volume ($\tilde{x} = 3/4$). Arrows show the direction of movement of the points with time. Point "F" marks the final stationary state of the system. The dashed line indicates the relationship between \tilde{N} and \tilde{q} in extremely inhomogeneous mixing (according to the classical concept).





 1296
 Fig. 12. DSD at different \tilde{x} at the beginning of the mixing process for Da = 1 and

 1297
 R = -0.5.

 1298
 1299

 1300
 1301

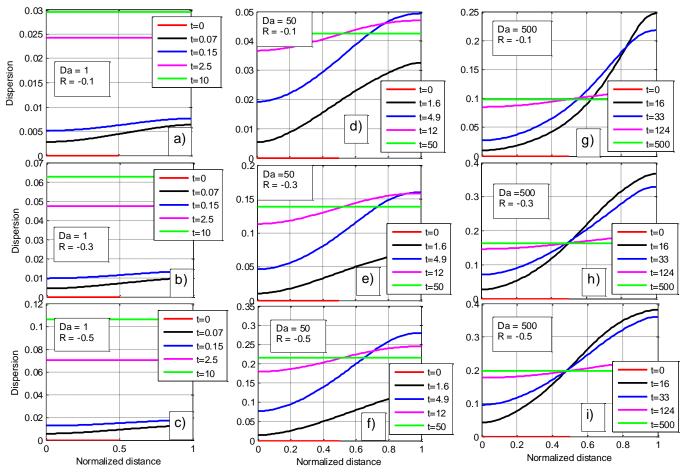
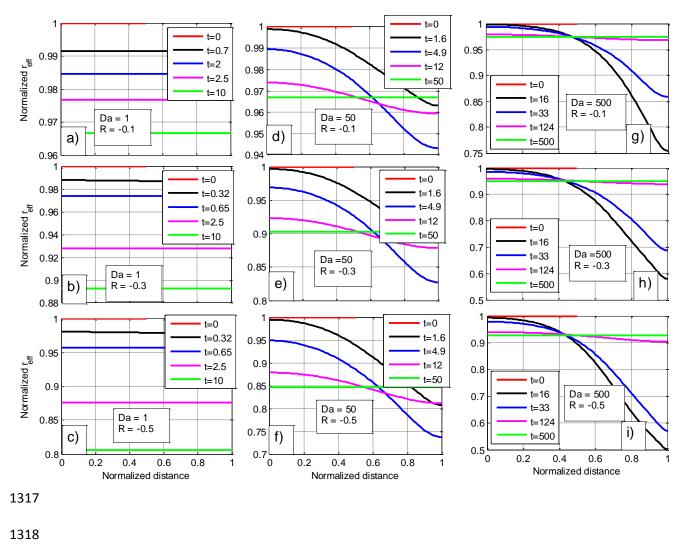
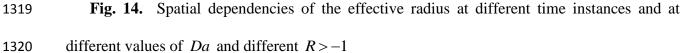
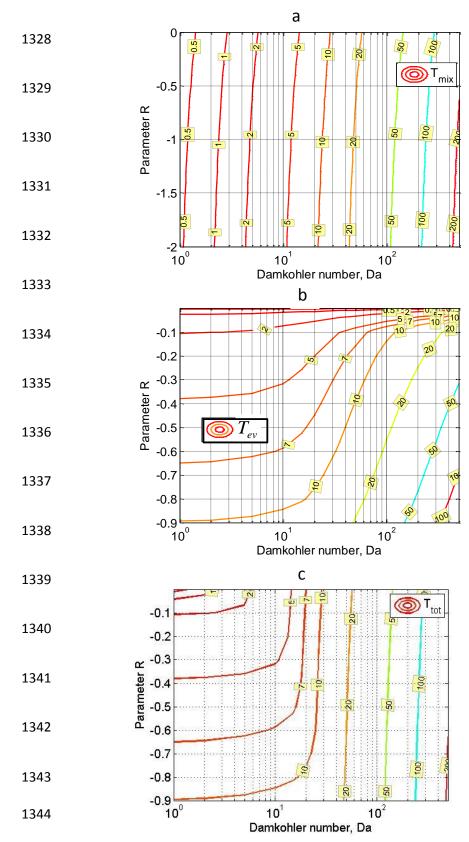


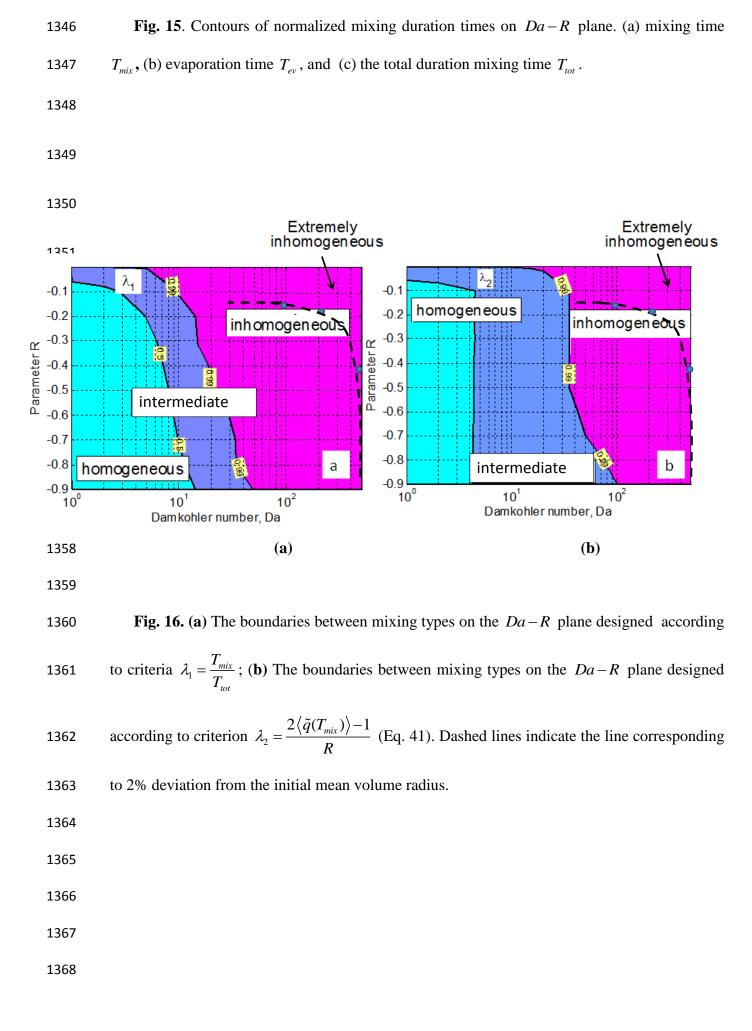
Fig. 13. Spatial dependencies of the relative DSD dispersion at different time instances and at different values of *Da* and different R > -1

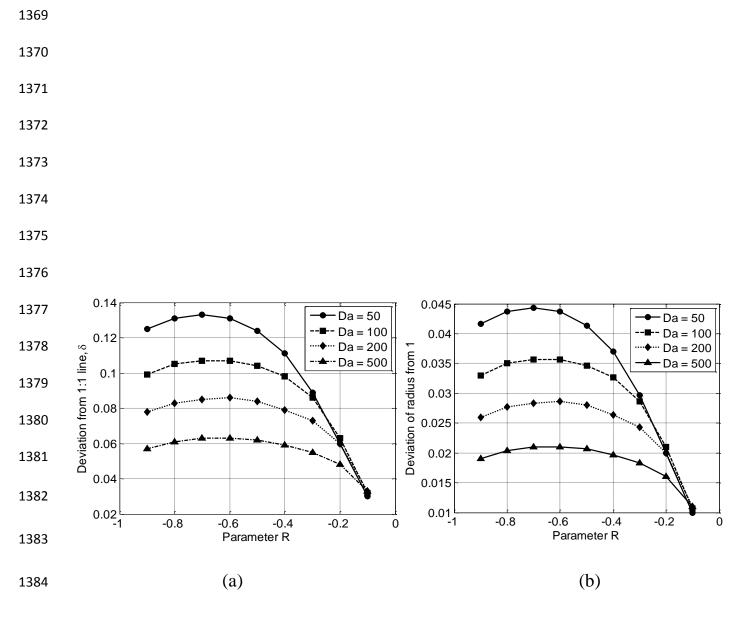












1385

Fig 17. (a) Dependencies of the r.m.s. distance of the $\tilde{N} - \tilde{q}$ relationship curve from straight line 1:1 suggested by classical concept of extremely inhomogeneous mixing. The dependencies are plotted for different values of *Da* and *R*. (b) The same as to the left panel but for r.m.s. deviations of the mean volume radius curve from that initial constant value assumed in the classical concept.