# RESOLVING ICE CLOUD OPTICAL THICKNESS BIASES BETWEEN CALIOP AND MODIS USING INFRARED RETRIEVALS

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### 19 Abstract20

21 Despite its importance as one of the key radiative properties that determines the impact of upper 22 tropospheric clouds on the radiation balance, ice cloud optical thickness (IOT) has proven to be 23 one of the more challenging properties to retrieve from space-based remote sensing measurements. 24 In particular, optically thin upper tropospheric ice clouds (cirrus) have been especially challenging 25 due to their tenuous nature, extensive spatial scales, and complex particle shapes and light 26 scattering characteristics. The lack of independent validation motivates the investigation presented 27 in this paper, wherein systematic biases between MODIS Collection 5 (C5) and CALIOP Version 28 3 (V3) unconstrained retrievals of tenuous IOT (< 3) are examined using a month of collocated A-29 Train observations. An initial comparison revealed a factor of two bias between the MODIS and 30 CALIOP IOT retrievals. This bias is investigated using an infrared (IR) radiative closure approach 31 that compares both products with MODIS IR cirrus retrievals developed for this assessment. The 32 analysis finds that both the MODIS C5 and the unconstrained CALIOP V3 retrievals are biased 33 (high and low, respectively) relative to the IR IOT retrievals. Based on this finding, the MODIS 34 and CALIOP algorithms are investigated with the goal of explaining and minimizing the biases 35 relative to the IR. For MODIS we find that the assumed ice single scattering properties used for the 36 C5 retrievals are not consistent with the mean IR COT distribution. The C5 ice scattering database

37 results in the asymmetry parameter (g) varying as a function of effective radius with mean values 38 that are too large. The MODIS retrievals have been brought into agreement with the IR by 39 adopting a new ice scattering model for Collection 6 (C6) consisting of a modified gamma 40 distribution comprised of a single habit (severely roughened aggregated columns); the C6 ice cloud 41 optical property models have a constant  $g \approx 0.75$  in the mid-visible spectrum, 5-15% smaller than 42 C5. For CALIOP, the assumed lidar ratio for unconstrained retrievals is fixed at 25 sr for the V3 43 data products. This value is found to be inconsistent with the constrained (predominantly 44 nighttime) CALIOP retrievals. An experimental data set was produced using a modified lidar ratio 45 of 32 sr for the unconstrained retrievals (an increase of 28%), selected to provide consistency with 46 the constrained V3 results. These modifications greatly improve the agreement with the IR and 47 provide consistency between the MODIS and CALIOP products. Based on these results the 48 recently released MODIS C6 optical products use the single habit distribution given above, while 49 the upcoming CALIOP V4 unconstrained algorithm will use higher lidar ratios for unconstrained 50 retrievals.

52 **1. Introduction** 

53 While clouds represent one of the largest modulators of Earth's radiation, with their impact 54 dependent on a variety of cloud physical and radiative properties, they remain one of the more 55 difficult components to represent in global climate models (Jiang et al. 2012). Passive satellite 56 observational datasets such as those from MODIS (Moderate Resolution Imaging 57 Spectroradiometer), AVHRR (Advanced Very High Resolution Radiometer), HIRS (High-spectral 58 Infrared Sounder), and ISCCP (International Satellite Cloud Climatology Project) provide long-59 term, global cloud observations (Wylie et al. 2005) (Heidinger et al. 2013; King et al. 2013; King 60 et al. 2003; Rossow 1991; Rossow and Schiffer 1999). However assessing the uncertainties in the 61 cloud radiative properties retrieved by these sensors has proved to be a complex and difficult task. 62 Until recently, validation of these retrievals was limited to ground and aircraft inter-comparisons. 63 But with the successful launch of CALIPSO (Cloud-Aerosol Lidar and Infrared Pathfinder Satellite 64 Observations) and CloudSat in April 2006 as part of the NASA-led Afternoon Constellation (A-65 Train) (Stephens et al. 2002; Winker et al. 2010), researchers now have access to a near-continuous 66 global record of vertically resolved observations of cloud and aerosol properties with nearly 67 coincident observations from MODIS Aqua.

Since launch, the CALIPSO lidar (the Cloud Aerosol Lidar with Orthogonal Polarization, or CALIOP) has proven to be a valuable tool for developing and evaluating passive cloud retrievals (Ackerman et al. 2008; Delanoë and Hogan 2010; Holz et al. 2008; Jin and Nasiri 2013; Kahn et al. 2014). CALIOP can directly measure cloud-top height with sensitivities that are significantly greater than the passive retrievals, while the CALIOP depolarization and attenuated

73 backscatter measurements provide vertically resolved cloud phase discrimination (Hu et al. 2009) 74 for cloud layers up to a cumulative optical depth of about 3.

75 Ice Optical Thickness (IOT) has also proved to be one of the more difficult properties to 76 retrieve from space-based passive sensor measurements and challenging to validate. In particular, it 77 is difficult to infer the microphysical and radiative properties of optically thin upper tropospheric 78 ice clouds (cirrus) from observations made by passive space-borne instruments due to their tenuous 79 nature, extensive spatial scales, complex particle shapes, and a wide range of particle sizes. Active 80 sensors such as CALIOP have the advantage that they directly measure the vertical structure of 81 clouds and aerosols however similar to the passive retrievals, assumptions regarding the ice 82 scattering properties (ie lidar ratio and multiple scattering) are necessary to invert the lidar signal 83 and retrieve the ice cloud extinction. These lack of constraints in both the MODIS and CALIOP ice 84 cloud retrievals result in considerable uncertainty and potential bias in the IOT which is the focus 85 of the manuscript. The manuscript begins by presenting an inter-comparison between the MODIS 86 C5 and CALIOP V3 IOT retrievals for optical tenuous cirrus (IOT < 3.0). A factor of two bias is 87 found between MODIS and CALIOP unconstrained retrievals (presented in Figure 1) and 88 described in section 4, raising a major question regarding the utility of these data records to study 89 ice cloud radiative processes. We next investigate the bias using an infrared (IR) radiative closure 90 experiment using collocated the MODIS 11 µm observations. Based on these results 91 modifications to the MODIS optical property retrievals with a focus on the ice scattering models is 92 investigated. For CALIOP experimentation with the value of the assumed lidar ratio used in the 93 unconstrained retrieval is evaluated. The result from this study provides the basis for the change in 94 the ice scattering models used by the recently released MODIS C6 ice cloud products and for

95 CALIOP the results provide one of the key studies motivating the changes to the CALIOP ice96 cloud extinction retrievals in the upcoming V4 product.

#### 97 MODIS and CALIOP Retrieval Background

98 Both MODIS and CALIOP IOT retrievals require a priori information concerning the ice 99 particle scattering properties that relate the measured reflectance (MODIS) or attenuated 100 backscatter (CALIOP) to the cloud's IOT and potentially the effective particle size. MODIS ice 101 cloud forward radiative calculations in the visible/near-infrared (VNIR) depend directly on the ice 102 particle phase function assumption, and to a first order on the associated asymmetry parameter (g). 103 For CALIOP, an assumed extinction-to-backscatter ratio is required for the unconstrained 104 retrievals where the algorithm is unable to make reliable estimates of cirrus IOT by measuring the 105 attenuated backscatter coefficients in some clear air region immediately above and below cloud 106 base (Young and Vaughan, 2009). Because solar background signals greatly reduce the signal-to-107 noise ratio (SNR) of the CALIOP daytime measurements, the vast majority of CALIOP daytime 108 IOT estimates are derived from unconstrained retrievals.

109 Uncertainties in the ice scattering property assumptions of either MODIS and/or CALIOP 110 could account for the biases found in Fig. 1. As will be discussed, an infrared (IR) cirrus IOT 111 retrieval is relatively insensitive to ice particle size and scattering details compared to MODIS and 112 CALIOP VNIR measurements, and thus provides an independent means to assess thin to 113 moderately optically thick ice cloud retrievals (IOT  $\sim 0-3$ ). In addition, an IR retrieval provides 114 radiative closure with solar reflectance based on MODIS IOT retrievals in the sense that 115 consistency in the two retrieved IOTs also implies forward model consistency with the respective 116 top-of-atmosphere (TOA) VNIR and IR observations.

117 Using the NASA-funded SSEC Atmosphere Product Evaluation and Test Element 118 (PEATE), now re-named the Suomi-NPP Atmosphere Science Investigator Processing System 119 (SIPS), the sensitivity of MODIS retrievals to ice single scattering properties are investigated by 120 repeated analyses of collocated January 2010 CALIOP and MODIS observations using a variety of 121 ice crystal habits (Yang et al. 2012) and size distributions. Based on comparisons against IR 122 retrievals, the MODIS MYD06 Collection 6 (C6) ice cloud optical property algorithm uses a single 123 habit - severely roughened aggregated columns (Yang et al. 2012) - instead of the size-dependent 124 multi-habit model (Baum et al. 2005) used for C5. The MYD06 C6 results compare well with a 125 new CALIOP version that uses a modified (larger) extinction-to-backscattering ratio for 126 unconstrained IOT retrievals.

127 The manuscript is organized as follows: Section 2 presents a detailed description of the 128 algorithms and data sets used in the analysis of the ice cloud optical depths with a focus on the IR 129 retrievals. Section 3 introduces the global inter-comparison between the MODIS collection 5 and 130 CALIPSO V3 ice cloud optical depths with section 4 presenting the comparison with the 131 collocated IR retrievals (ocean only). Section 5 discusses the impact of the ice model selection 132 (MODIS) and the assumed lidar ratio and multiple scattering correction (CALIOP) on the ice cloud 133 optical depth and then presents an inter-comparison of the MODIS and CALIOP retrievals 134 processed using a modified single scatter look up table (severely roughened aggregated columns) 135 and a modified of unconstrained lidar ratio of 31 (instead of 25 for V3). Section 6 summarizes the 136 results and with a focus on the rational for the selection of a single habit for the new single 137 scattering properties for the MODIS C6 ice cloud retrievals.

#### **138 2. Ice Cloud Optical Thickness Retrieval Datasets**

An overview of the relevant retrieval methodologies is presented here with a focus on the forward cloudy radiative transfer modeling assumptions and IR IOT retrievals developed specifically for this study.

#### 142 **2.1** *IR retrievals and radiative closure*

The MODIS channel suite includes a range of IR channels extending well into the  $CO_2$ absorption region (13-15  $\mu$ m). The calibration of the IR channels has been extensively validated and proven to have high accuracy, with uncertainties less than 0.5 K across a broad temperature range (Tobin et al. 2006). For ice clouds, the IR radiative transfer is dominated by absorption, and thus is less complex than for the VNIR retrieval. In this section we discuss the IR radiative transfer methodology that is used both to retrieve the IR IOT as well as evaluate the MODIS and CALIOP retrievals.

150 The goal of radiative closure study is to relate the differences in the CALIOP and 151 MODIS retrieved IOT to the measured TOA channel radiance or Brightness Temperature (BT) 152 in the MODIS 11 µm channel. To calculate the TOA cloudy radiances requires an accurate 153 radiative transfer model, knowledge of the cloud boundaries, and well-characterized surface 154 temperature/emissivity and atmospheric thermodynamic profiles. LBLDIS (Turner et al. 2003), a 155 cloudy radiative transfer model, is used for this analysis. The model elegantly combines the clear 156 sky Line By Line Radiative Transfer Model (LBLRTM) (Clough and Moncet 1992) with the 157 Discrete Ordinates Radiative Transfer (DISORT) (Stamnes et al. 1988), a proven and accurate 158 cloudy radiative transfer model. The inputs required for LBLRTM are surface temperature and 159 emissivity, vertically resolved temperature and water vapor profiles, and information regarding 160 trace gas concentrations such as CO<sub>2</sub> and O<sub>3</sub>. For this analysis the surface temperature and

161 thermodynamic profiles are extracted from the NOAA Global Data Assimilation System 162 (GDAS) files that provide profiles at 1° spatial resolution every 6 hours. For each MODIS and 163 CALIOP FOV, the closest (in both time and space) GDAS profile is selected. A fixed  $CO_2$ 164 concentration of 380 ppm and a climatological  $O_3$  profile is used. Given these inputs LBLRTM is 165 run on the selected FOV filtered using the collocated CALIOP V3 5km cloud layer products 166 (described in Section 3). The results of the clear sky validation are discussed in Section 4.

167 The cloud microphysics and thermodynamics are defined with a vertical resolution of 500 168 meters within the cloud boundaries defined by the CALIOP layer product. Only FOV where the 169 CALIOP is not attenuated at the surface are used greatly reducing uncertainties in the cloud base 170 determination. For example a cloud with a geometrical thickness of 1.5 km is divided into 3 171 layers with each layer defined by an optical thickness, effective radius and ice scattering model. 172 For example for a cloud with a total optical thickness of 1.5 each layer will have an optical 173 thickness of 0.5. Using this methodology the vertical temperature profile is accounted for in the 174 radiative transfer. For daytime IR forward model calculations, the effective radius from the 175 MODIS optical property retrieval is used for all cloud layers. For nighttime CALIOP 176 comparisons, a fixed effective radius of 40  $\mu$ m is used in the IR calculations. It is important to 177 note that at 11  $\mu$ m the OT retrieval is relatively insensitive to the assumed effective radius.

The last remaining variable needed to calculate the TOA IR radiance is IOT. LBLDIS is run independently using either the MODIS or CALIOP retrieved IOT, resulting in high spectral resolution TOA radiances with the only differences being the assumed IOT (i.e., MODIS or CALIOP). The spectrally resolved radiances are then integrated over the MODIS Aqua 11  $\mu$ m channel (band 31) spectral response function resulting in a simulated TOA radiance that can be directly compared to the measured MODIS 11  $\mu$ m observations. In addition to LBLDIS spectral calculations, TOA longwave fluxes are calculated using the Rapid Radiative Transfer Model (RRTM) (Mlawer et al. 1997) that is also based on DISORT and LBLRTM and utilizes a correlated-*k* method for gas absorption along with broadband ice cloud parameterizations from (Fu et al. 2000). Identical inputs are used for RRTM and the LBLDIS TOA calculations with the only variable being IOT. The TOA fluxes are subsequently used to quantify the impact of the IOT biases on the global characterization of ice cloud radiative forcing.

191 IR observations provide the independent reference to understand differences between 192 MODIS and CALIOP IOT retrievals. While radiance closure provides valuable information 193 regarding TOA radiances and fluxes it does not provide a direct assessment of the individual 194 CALIOP and MODIS IOT biases. To convert observed IR TOA radiance to IOT, two different 195 retrieval approaches were used. First, we developed an IR window IOT retrieval that uses the 196 collocated MODIS and CALIOP observations. This "reference" retrieval uses cloud boundary 197 information from CALIOP coupled with the LBLDIS forward model and then retrieves the IR 198 IOT using the MODIS 11  $\mu$ m window channel observations that are coincident and collocated 199 with CALIOP. A second method uses the spectral emissivity retrieved from the MODIS CO<sub>2</sub> 200 emissive cloud-top pressure retrieval that is then related to the IOT and effective radius using a 201 pre-computed lookup table (Heidinger et al. 2015). This method has the advantage of being 202 computationally very efficient, not requiring the CALIOP cloud boundaries, and providing IOT 203 for the entire MODIS swath. Both IR retrieval methods are discussed in more detail in the 204 following sub-sections.

#### 205 **2.1.1. Combined MODIS IR Window and CALIOP Retrievals**

A single channel IR window IOT retrieval was developed for this study using combined 206 207 CALIOP and MODIS observations and the LBLDIS forward radiative transfer modeling 208 discussed in the previous section. The method constrains the cloud boundaries using the 209 collocated CALIOP 5km layer products and uses surface and atmospheric temperatures 210 information from GDAS. TOA radiances are simulated using LBLDIS with IOT retrieved by 211 minimizing the measured MODIS channel 31 (11 µm) and calculated BT differences. The 212 retrieval assumes the cloud extinction is evenly distributed in the vertical throughout the cloud. 213 This simplification has the potential to bias the retrieval for FOV where the IOT is distributed 214 non-uniformly in the vertical (Maestri and Holz 2009). The cloud geometric thickness is thus 215 limited to no greater than 4 km to reduce IOT biases that can be introduced by non-homogeneous 216 layers.

#### 217 **2.1.2. MODIS IR Spectral Emissivity Retrievals**

The MODIS C6 CO<sub>2</sub> slicing algorithm provides retrieved spectral emissivity for the 8.5, 11, and 12 µm channels (channels 29, 31, 32) that have sensitivity to both the IOT and effective radius. As described in (Parol et al. 1991),  $\beta$  ratios can be approximated based on these emissivities and are related to the asymmetry parameter (g), single-scattering albedo ( $\omega_0$ ), and extinction efficiency ( $Q_e$ ) as follows:

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224 (1) 
$$\boldsymbol{\beta}_{\lambda_1\lambda_2} = \frac{\boldsymbol{Q}_{e,\lambda_1}(1-\boldsymbol{\omega}_{o,\lambda_1}\boldsymbol{g}_{\lambda_1})}{\boldsymbol{Q}_{e,\lambda_2}(1-\boldsymbol{\omega}_{o,\lambda_2}\boldsymbol{g}_{\lambda_2})}$$

Thus  $\beta$  is the ratio of the scaled absorption extinction in two spectral channels ( $\lambda_1$  and  $\lambda_2$ ). The effective radius is first retrieved by matching simulated ice single-scattering calculations of g(r),  $\omega_0(r)$ , and  $Q_e(r)$ , each integrated over the appropriate MODIS spectral response functions, to the retrieved MODIS  $\beta$  ratios which are calculated for both the 8.5-11  $\mu$ m and 11 – 12  $\mu$ m pairs. For this analysis the scattering properties of severely roughened aggregated columns (Yang et al. 2012) are used to be consistent with the MODIS C6 cloud optical property retrievals.

Using the effective radius to define g(r),  $\omega_0(r)$ , and  $Q_e(r)$ , the extinction optical thickness is then retrieved by relating the 11 µm emissivity to the extinction optical thickness in the form ((Van de Hulst 1974)):

235 (2) 
$$\tau_{vis} = \frac{2}{Q_e} \left( \frac{\tau_{abs}}{(1 - \omega_o g)} \right) ,$$

where  $\tau_{abs}$  is the IR absorption optical thickness and  $\tau_{vis}$  is the extinction optical thickness at 236 237 532 nm. This derivation assumes that the ratio between the absorption and extinction optical 238 thickness is a factor of 2 in the IR. Based on ice cloud single-scattering calculations (Yang et al. 239 2012) and assuming that the majority of ice clouds have an effective radius greater than 10  $\mu$ m, 240 this assumption is expected to have introduce no more than 10% uncertainty. (Heidinger et al. 241 2015) provides a more detailed discussion of the retrieval methodology. This approach can be 242 applied without the need for the CALIOP cloud boundaries, and provides full swath IR IOT 243 retrievals. We leverage this capability to investigate the MODIS IOT retrieval biases as a 244 function of view angle.

#### 245 **2.2 CALIOP Ice Cloud Optical Thickness Retrievals**

CALIOP is a two wavelength elastic backscatter lidar that measures attenuated backscattercomponents polarized parallel and perpendicular to the transmitted laser light at 532 nm and total

248 attenuated backscatter at 1064 nm (Hunt et al. 2009). Once the received signals have been 249 background-subtracted and calibrated (Powell et al. 2009), a tightly integrated suite of retrieval 250 algorithms is used to detect layer boundaries (Vaughan et al. 2009) and classify layers as either 251 clouds or aerosols (Liu et al. 2009). Layers classified as clouds are further classified according to 252 thermodynamic phase as either ice clouds or water clouds (Hu et al. 2009). Layer optical 253 thickness (including IOT) is then retrieved using one of two techniques: constrained or 254 unconstrained retrievals (Young and Vaughan 2009). Constrained retrievals are applied 255 whenever the effective two-way transmittance of a layer,

256

257 (3) 
$$\mathbf{T}_{eff}^{2} = exp(-2\eta\tau) = exp\left(-2\eta\int_{layer top}^{layer base} \sigma_{c}(r)dr\right)$$

- 258
- 259

260 can be directly and reliably measured. In this expression  $\tau$  is the layer optical depth (IOT for ice 261 clouds),  $\sigma_c(r)$  is the range-resolved cloud extinction coefficient, and  $\eta$  is a multiple scattering 262 correction factor whose value depends on the lidar sensing geometry and the scattering characteristics of the particulates being measured. While  $T_{\mbox{\scriptsize eff}}^2$  estimates can be obtained from 263 measurements of clear air, opaque water clouds, and ocean surfaces (see (Josset et al. 2012; 264 265 Yongxiang et al. 2007; Young 1995), respectively), the CALIOP V3 algorithm only implements the clear air technique, in which  $T_{eff}^2$  can be obtained directly from the ratio of the mean 266 267 attenuated scattering ratios calculated in regions of clear air located immediately above cloud top and below cloud base (Vaughan et al. 2005). Retrieving IOT from measurements of  $T_{e\!f\!f}^2$  requires 268 269 knowledge of the appropriate multiple scattering factor (Winker 2003). For CALIOP 270 measurements of cirrus clouds, (Josset et al. 2012) determined the mean multiple scattering factor to be  $0.61 \pm 0.15$ . In the CALIOP V3 algorithm,  $\eta$  is fixed at 0.6 for all cirrus clouds. More recent results suggest that the multiple scattering factor is dependent on cloud temperature (Garnier et al. 2015b)which is being considering for the upcoming version 4 products.

274 Constrained retrievals are the preferred method for retrieving IOT from CALIOP 275 measurements. However, because solar background light significantly degrades the CALIOP 276 SNR during daylight operations, V3 constrained retrievals occur almost exclusively during 277 nighttime observations, thus severely limiting direct comparisons with MODIS IOT retrievals 278 derived from VNIR solar reflectance. For the vast majority of daytime observations, CALIOP 279 IOT retrievals use an unconstrained technique that requires *a priori* knowledge of the cirrus 280 extinction-to-backscatter ratio (i.e., lidar ratio),

281 (4) 
$$S_c = \frac{\sigma_c(r)}{\beta_c(r)}$$

282 where  $\sigma_c(\mathbf{r})$  and  $\beta_c(\mathbf{r})$  are, respectively, the cloud extinction and backscatter coefficients. IOT is then obtained by solving the lidar equation using specified values of n and S<sub>c</sub> (Young and 283 284 Vaughan 2009). Note that while the cloud extinction and backscatter coefficients are explicitly 285 range-dependent, their ratio is assumed to be range-invariant. Although S<sub>c</sub> for ice clouds most 286 likely varies depending on crystal habit and size distribution, the CALIOP V3 unconstrained 287 retrievals use a globally constant default value of  $S_c = 25$  sr. Based on ground based lidar 288 observations there can be significant variability in the lidar ratio. The constant value is 289 considered one of the primary sources of uncertainty in the V3 ice cloud extinction retrievals. 290 This value was determined prior to launch from the best information available from numerous 291 ground-based and airborne data sets e.g., (Holz 2002; Sassen 2001; Yorks et al. 2011).

Errors in lidar ratio selection for unconstrained retrievals generate corresponding errors in the resultant estimates of IOT. In particular, an underestimate of S<sub>c</sub> will result in CALIOP underestimating IOT. The selection of the default CALIOP lidar ratio is thus one of the potential
major sources of bias in the CALIOP unconstrained retrievals that can be investigated using IR
observations from either MODIS or the CALIPSO IIR (Imaging Infrared Radiometer) instrument
(Garnier et al. 2015a).

- 298
- 299 **2.3 MODIS Ice Cloud Optical Thickness Retrievals**

The MODIS imager provides measurements in 36 spectral channels, covering the Visible Near Infrared (VNIR), Shortwave Infrared (SWIR), Midwave Infrared (MWIR), and thermal IR portions of the spectrum. Spatial resolution is 250 m in two VNIR channels, 500 m in the 5 VIS/SWIR channels, and 1 km in the remaining channels.

304 The MODIS cloud optical/microphysical property algorithm is used to generate a single 305 cloud product designated by the NASA Earth science data type (ESDT) names MOD06 and 306 MYD06 for Terra and Aqua MODIS, respectively (hereafter referred to as MYD06 since the 307 algorithms are essentially identical and this study is focused on MODIS Aqua observations). For 308 daytime measurements, the 1 km cloud retrieval algorithm uses multiple spectral channels 309 (primarily six VNIR, SWIR and MWIR channels, as well as several thermal channels) to 310 simultaneously retrieve cloud optical thickness, effective radius (and derived water path) and 311 thermodynamic phase for liquid and ice phase clouds. In addition to the 1 km MODIS Level-1B 312 calibrated radiance product, the algorithm requires the following input: MODIS cloud mask 313 (MYD35) including 250 m mask information (Ackerman et al., 1998), the cloud-top pressure 314 portion of MYD06 (Ackerman et al. 2008; Holz et al. 2008), and a variety of ancillary datasets. 315 Heritage algorithm work is discussed in (King et al. 2003; Nakajima and King 1990; Platnick 316 and Twomey 1994; Platnick et al. 2001).

317 C5 algorithm-related publications include ice radiative models (Ackerman et al. 2008; 318 Baum et al. 2005; Yang et al. 2007) multilayer detection (Wind et al. 2010), Clear Sky Restoral 319 filtering (Pincus et al. 2012; Zhang and Platnick 2011), pixel-level uncertainties, and L3 global 320 gridded statistics (King et al. 2013). An online list of the recent C6 algorithm updates is available 321 from the MODIS Atmosphere Team web site (Platnick 2014). The most relevant update for the 322 current discussion is the adoption of new ice cloud radiative models having an overall smaller 323 asymmetry parameter, as will be discussed in Sect. 5.1. Note for consistency with the spherical 324 droplet definition, as well as for use in deriving ice water path, the effective radius of a non-325 spherical ice particle is defined as <sup>3</sup>/<sub>4</sub> times the ratio of the average volume of the size distribution 326 to the average cross sectional area (Yang et al. 2007).

- 327
- 328 **2.4 Collocation and the Merged Dataset**

In this section we present the methods used to collocate and merge the CALIOP and MODIS observations providing the foundation for the inter-comparisons and analysis presented in the results of Sect. 4.

332 The analysis is based on one month (January 2010) of physically collocated CALIOP and 333 MODIS observations. MODIS is an imaging radiometer while CALIOP is a near-nadir viewing 334 lidar. Because each instrument has a unique viewing geometry with different spatial resolutions, 335 accurate inter-comparisons require collocating the observation FOVs. This analysis uses tools 336 that provide computationally efficient and accurate collocation (Nagle and Holz 2009). The 337 methodology defines master and follower instruments, with the master typically being the larger 338 FOV and the follower FOV collocated within the master footprint. In this investigation MODIS 339 is defined as the master with CALIOP the follower. The MODIS spatial resolution can be approximated as a rectangular box with a  $1 \times 2$  km resolution at nadir. The CALIOP IOT retrieval can be performed over horizontal averaging distances ranging from 5 km to 80 km, depending on the magnitude of the cloud signal relative to the background noise (Yongxiang et al. 2007). The CALIOP surface footprint is therefore approximated as an 80-meter wide swath with the alongtrack length depending on the amount of spatial averaging. The majority of observations used in this analysis are the 5 km averaged IOT. A more detailed description of the CALIOP and MODIS collocation is presented in (Holz et al. 2008).

Leveraging the UW Atmospheric Science Investigator-led Processing System (SIPs) processing capabilities, a month of collocated MODIS and CALIOP collocated observations were processed using the CALIOP and MODIS IOT retrievals with the only difference being incremental changes to the ice cloud parameterizations used in the retrieval algorithms. This approach isolates the impact of the parameterization changes and/or algorithm modifications and provides a direct assessment of the changes in IOT.

#### 353 **3. CALIOP V3 and MODIS C5 Cirrus Optical Thickness Inter-Comparisons**

354 Figure 1 presents the MODIS C5 IOT retrievals compared with CALIOP V3 IOT for one month 355 (January 2010) of non-polar (±60 degrees latitude) daytime ocean observations. The CALIOP 356 5 km layer products are used to select only single layer ice clouds were both the CALIOP phase 357 retrieval (Hu et al. 2009) and the MODIS optical property phase retrieval identify ice clouds. The 358 CALIOP phase detection is sensitive to scattering from oriented ice (specular reflection), and 359 such cases are excluded from the data set. Because the CALIOP layer detection algorithm 360 employs a nested, multi-resolution spatial averaging scheme (Vaughan et al. 2009), the CALIOP 361 5 km layer products can report distinct layers in cases where the base of the upper layer is

362 separated from the top of the lower layer by as little as a single range bin (60m). For a passive 363 retrieval such as from MODIS, a 60 m vertical separation will have little impact on the retrieval 364 results assuming both layers are ice. To improve the comparison yield and provide a more 365 representative distribution of single layer ice clouds for inter-comparing the passive 366 observations, CALIOP 5 km ice cloud layers with a vertical separation of 3 km or less are 367 merged to form single, vertically contiguous layers. The CALIOP extinction profile is then 368 integrated for each profile using the redefined layer boundaries, thus providing an aggregated 369 IOT. Ice clouds with total geometrical thickness greater than 4 km using this single layer 370 definition are excluded from the comparison.

371 The MODIS IOT retrievals are filtered using the C5 MODIS Quality Assurance (QA) 372 parameters and a horizontal heterogeneity threshold. MODIS IOT retrievals, (i.e., with the QA 373 usefulness flag set to 1 and the QA confidence flag set to 3, are used in the comparison). This 374 filtering provides all ice cloud retrieval where both the IOT and effective radius successfully 375 converged within the lookup table. Unlike liquid water clouds QA values of 2 and 1 are not used 376 for C5 retrievals. Using this filtering provides the highest quality MODIS retrievals and removes 377 all cloud edges from the comparison. To reduce uncertainties resulting from spatial sampling 378 differences between MODIS and CALIOP, the standard deviation of a 5x5 pixel box centered 379 over the collocated pixel is computed. Only collocated pixels where the MODIS IOT standard 380 deviation is less than 0.5 are used; we find, however, that the comparison results are relatively 381 insensitive to this threshold.

Figure 1 reveals a systematic bias between the MODIS C5 and CALIOP IOT's, with MODIS approximately a factor of two larger than the CALIOP unconstrained retrievals. An independent methodology is needed to assess this difference since both retrievals depend on ice 385 scattering property assumptions. As discussed in the methodology section, the IR observations 386 provide sensitivity to the IOT given well-constrained cloud boundaries with uncertainties that are 387 independent of the CALIOP and MODIS VNIR retrievals. Spectrally resolved TOA radiances 388 are calculated for the three different retrieval methods - MODIS, CALIOP unconstrained 389 (daytime measurements), and CALIOP constrained (nighttime measurements) – using LBLRTM 390 and LBLDIS. All three calculations use identical cloud boundaries defined by the merged 391 CALIOP 5 km layer heights and the same thermodynamic profiles and ocean surface 392 temperatures (GDAS), with the only difference being the IOT used in the calculation. The 393 spectrally resolved TOA radiances are then integrated over the MODIS channel 31 (11 µm) 394 spectral response function. To investigate the accuracy of the combined GDAS and TOA clear 395 sky LBLRTM calculations, simulated TOA 11 µm BT for clear sky FOVs identified using both 396 the MODIS and CALIOP cloud masks were compared to the measured MODIS 11 µm channel 397 BTs. The mean bias between the simulated and observed BT is less than 0.2 K, which is within 398 the expected calibration uncertainty of MODIS (Tobin et al. 2006).

399 Figure 2a presents the MODIS C5 and CALIOP V3 BT closure results. The figure 400 reveals a sobering finding which is that neither the MODIS C5 nor the CALIOP V3 401 unconstrained IOT retrievals provide radiative closure in the window IR. Furthermore, the 402 respective retrievals are biased in opposite directions. For MODIS C5, the calculated TOA BT is 403 colder than the measured BT with a mean bias of -8.7K, implying the MODIS IOT is on average 404 biased high. In contrast, the TOA BT calculated using the CALIOP V3 unconstrained IOT has a 405 mean bias of +12.1 K, suggesting the CALIOP retrieval is biased low. The CALIOP V3 406 constrained retrievals, which do not require an assumed lidar ratio but only an estimate of the 407 multiple scattering correction, demonstrate much better agreement with a mean bias of +1.4 K.

408 To put the biases into a radiative context, the cloudy IR TOA fluxes are computed for 409 each collocation using RRTM. The calculations use the CALIOP cloud boundaries, the surface and atmospheric profiles from GDAS, and the MODIS retrieved effective radius. For each 410 411 collocation RRTM calculations are computed with the only difference being the IOT used 412 (MODIS or CALIOP) with the results presented in Figure 2b. The mean TOA flux difference between MODIS and CALIOP unconstrained retrievals is +23 W m<sup>-2</sup> with a standard deviation of 413 21W m<sup>-2</sup>. For the tenuous ice clouds being investigated, the sensitivity of the TOA flux to IOT is 414 415 primarily driven by the thermal contrast between the surface and the mean emitting temperature 416 of the cloud (Corti and Peter 2009). The very large differences in the wings of the distribution in 417 Fig. 1b occur primarily near the tropics where the thermal contrast is greatest between the cloud and the surface. For this region TOA differences as large as 50 W m<sup>-2</sup> are found in Figure 2b. 418

#### 419 **4. IR Retrievals as a Reference Optical Thickness**

Because the sensitivity of IR IOT retrievals to ice crystal habit selection is minimal, these retrievals provide an independent means to evaluate the CALIOP and MODIS solar reflectance retrievals. As discussed in Sect. 2, the main sources of uncertainty in the IR IOT originate from characterizing the surface temperature and having an accurate determination of the cloud emitting temperature. To reduce the surface temperature uncertainty, the results of this section are restricted to non-polar ( $\pm$  60 degrees) ocean-only cases.

The comparisons with IR window IOT retrievals shown in Figure 3 reveal biases in both the MODIS (a) and daytime CALIOP unconstrained (b) retrievals (high and low, respectively) that are consistent with the radiative closure results presented in Figure 2. The magnitude of the bias relative to the IR is approximately +40% for MODIS. For CALIOP there is a non-linear dependence between the IOT and the negative bias relative to the IR, with the bias increasing 431 substantially for IR IOTs greater than unity; the CALIOP results are discussed further in Section432 5.2.

433 A limitation of the IR window IOT data set is that only a small subset of the MODIS 434 across track swath can be assessed due to the very close coordination between the MODIS and 435 CALIOP orbits. To investigate MODIS IOT scan angle dependencies we use the MODIS 436 spectral IR IOT retrieval described in Sect. 2.1.2. Figure 4a shows the MODIS C5 liquid (warm 437 colors) and ice (cool colors) phase cloud optical thickness for an example MODIS data granule 438 (January 11 2010, 06:25 UTC). Fig. 4b presents the histogram of the ratio between the MODIS 439 IOT and the full swath IR IOT (described in section 2.1.2) separated by viewing angle ranges as 440 indicated by the colored lines overlaid on the IOT image. A ratio of unity would suggest good 441 agreement between the spectral IR and VNIR IOT retrievals. However, as illustrated in the 442 following section, for the MODIS C5 retrievals (solid lines) the modes of the distributions vary 443 with scan angle, and the bias is seen to be an increasing as a function of scan angle. The 444 dependence on the scan angle (ie scattering angle) for the C5 retrievals results from the strong 445 angular variation in the C5 phase functions. This is an important result, as it demonstrates 446 necessity that this scattering angle dependence can provide an additional constraint on ice 447 radiative model selection. In addition, because CALIPSO and Aqua have similar orbits, only a 448 small range of MODIS viewing angles are included in the collocated inter-comparison, thus the 449 possible strong dependence on viewing angle implies the collocated analysis is representative 450 only of the view angle ranges sampled. Finally, given the lack of significant scattering in the IR, 451 the scan dependent bias further suggests the issue is with the MODIS C5 VNIR retrievals. This is 452 investigated in the next section.

## 453 5. Investigating the sensitivity of ice scattering model selections for MODIS and CALIOP ice454 cloud retrievals

#### 455 **5.1** Ice Radiative Model Sensitivities in MODIS

456 Though a primary focus of this investigation is on optimizing C6 ice models to improve 457 IOT inter-comparisons, it is understood that ice model crystal habits also affect the particle 458 single scattering albedo retrieved using the SWIR and MWIR channels that provide effective 459 particle size information. Figure 5a and Figure 5b show the 2.13 µm and 3.7 µm channel co-460 albedo, respectively, as a function of Cloud Effective Radius (CER) for four habit realizations, 461 namely the C5 habit mixture (black line) and the three severely roughened habits solid aggregate 462 plates (green line), solid bullet rosettes (red line), and aggregate columns (blue line) (Yang et al. 463 2012). To the extent that CER retrievals of an asymptotically thick cloud in the SWIR/MWIR are 464 essentially a retrieval of co-albedo, the difference between the aggregated column and C5 model 465 co-albedo implies an effective radius difference of  $+2 \,\mu\text{m}$  and  $-8 \,\mu\text{m}$  at the 2.1  $\mu\text{m}$  and 3.7  $\mu\text{m}$ 466 wavelengths, respectively, for a C5 effective radius of about 35 µm; smaller C5 retrieved sizes 467 would result in larger differences.

468 Figure 6 shows the asymmetry parameter sensitivity to habit for the same four habits 469 shown in Figure 5. Evidently the habit-sensitivity of the asymmetry parameter is also strong in 470 both the 2.1µm and 3.7µm MODIS channels. While the asymmetry parameters of three severely 471 roughened habits are not constant with effective size (though at 2.1 µm the aggregate plates and 472 aggregate columns are nearly constant), the C5 model has much larger size sensitivity at both 473 wavelengths (Cole et al. 2014; van Diedenhoven et al. 2014; Yang et al. 2008). Aggregated 474 columns, with smaller asymmetry parameters relative to C5, will result in a larger retrieved CER 475 estimates. This is because the resulting increase in modeled SWIR reflectance for a given 476 effective size causes the measured reflectance to be associated with a more absorbing (i.e.,

477 larger) particle. Therefore, the effect of both co-albedo and asymmetry parameter differences
478 between the severely aggregated column habit and the C5 model act to increase retrieved
479 effective radii at 2.1 µm, while at 3.7 µm some cancellation of effects can be expected.

480 The single habit radiative models shown in Figure 5 and Figure 6 are used to build look-481 up tables that were integrated into the MODIS C6 cloud retrieval development code. A month of 482 data was processed for each habit. It was found that the habit that provided the best consistency 483 with the IR window retrievals (Sect. 2.1.1) is the severely roughened aggregated column model. 484 The IOT retrieval comparison with the IR window retrievals using this model is shown in Figure 485 7a, where the MODIS reflectance-based retrievals using the severely roughened aggregated 486 column model are now clustered around the 1-to-1 line. In addition, this aggregated column 487 model was used to assess the MODIS retrieval swath dependence previously shown in Figure 4b. 488 The improvement of the aggregated column model (dashed lines) relative to the C5 model (solid 489 lines) is significant. Both results led to the decision to use the severely roughened aggregated 490 column radiative model for the MODIS C6 cloud optical/microphysical property retrievals.

491 Figure 8 shows an example of ice cloud retrievals for C5 and C6 for typhoon Fung-492 Wong. The typhoon was located south of Taiwan at the time of the MODIS Aqua data granule 493 acquisition on September 20, 2014 (0530 UTC). The C5 and C6 ice (cool colors) and liquid 494 (warm colors) cloud optical thickness retrievals are shown in the middle and right panels, 495 respectively. In addition to ice radiative model differences, MYD06 C5 and C6 have different 496 schemes for the cloud thermodynamic phase yielding different ice and liquid phase pixel 497 populations, though the optical thickness spatial patterns are similar for regions having the same 498 phase. Because of the different phase assignments made by these two scheme, quantifying ice 499 model retrieval sensitivities requires the comparisons be restricted to only those pixels for which

both algorithms generate successful retrievals that identify identical cloud phases. With this pixel filtering, the left panel of Figure 8b shows the normalized IOT distribution for the optical thickness range of the plot. The C6 IOT mode is roughly 27% smaller than the C5 mode, while the mean is decreased by about 15%, from 4.16 for C5 to 3.55 for C6. The 2.1  $\mu$ m ice cloud effective particle radius retrievals are shown in the right panel, with the C6 mode and mean both increasing by about 4  $\mu$ m (+15%) for C6 relative to C5.

506

#### 5.2 MODIS C6 model selection methodology

507 The MODIS IOT retrieval depends strongly on assumed ice scattering properties that are 508 needed to relate the measured reflectance to the retrieved IOT. The MODIS C5 retrieval used 509 empirically derived habit and size distributions with asymmetry parameters ranging between 0.79 510 and 0.88 depending on the ice cloud effective radius (Baum et al., 2005). By conducting an 511 infrared closure analysis, we have shown that the C5 parameterization is not representative of the 512 globally averaged ice scattering properties. More recent investigations of the ice cloud asymmetry 513 parameter suggest that most ice clouds have values around 0.75 in the visible spectrum (Cole et al. 514 2012; van Diedenhoven et al. 2013). Additionally, use of the C5 ice cloud radiative model results 515 in MODIS retrieval biases are strongly dependent on the viewing angle, as demonstrated in Figure 516 4. These findings motivated the investigation of new ice scattering models that have lower 517 asymmetry parameters and weaker dependence on ice effective radius.

518 Since the MODIS C5 algorithms were finalized, new ice scattering models that incorporate 519 roughened ice crystal parameterizations have been developed (Yang et al. 2012). Experimentation 520 with these new models demonstrates that a modified gamma distribution of severely roughened 521 aggregated columns provides a significantly lower visible asymmetry parameter (~0.75) that shows 522 very little dependence on ice effective radius. For testing purposes, the MODIS cloud retrieval 523 algorithm team implemented these new scattering properties in the MYD06 retrieval algorithm. 524 The updated algorithm was then run on the Atmospheric PEATE and the resulting data was 525 collocated with CALIOP measurements. Simulated TOA cloudy MODIS 11 µm brightness 526 temperatures (BT) were then computed using the reprocessed MODIS IOT retrievals and are 527 compared to the MODIS measured BT, These new results are presented in Figure 10b. The 528 updated ice scattering models generate greatly improved IOT estimates that show very close to a 529 one-to-one correspondence with the independently derived IR IOT values (Figure 7a) and is 530 consistent with the findings of (Baum et al. 2014). Additionally the view angle dependent bias is 531 largely removed, as presented in Figure 4b. Based on these results, the recently reprocessed 532 MODIS C6 cloud optical/microphysical property product (now in forward production) uses a 533 modified gamma distribution consisting of a single habit of severely roughened aggregated 534 columns for ice cloud retrievals. An additional benefit of the single habit is that it simplifies the 535 retrieval and increases the reproducibility of the scattering properties by the research community. It 536 is important to note that the selection of the single habit modified gamma distribution was to 537 provide a radiative consistency with the IR, not a microphysical model.

Figure 10a presents the same filtered 2-D histogram comparing CALIOP and MODIS as Figure 1 but using the ice radiative model modifications made for MODIS and the updated lidar ratio (32 sr) for CALIOP. Figure 10b presents the IR radiative closure for the updated IOT retrievals for January 2010. Notice the large bias between the MODIS and CALIOP un-constrained IOT is significantly reduced and the IR radiative closure shows very good agreement for all three IOT retrievals. There is still a tendency for the MODIS IOT to be larger than CALIOP in Figure 10a. The MODIS C6 IR closure in Figure 10b also demonstrates this bias, with the tail of the 545 distribution weighted to negative BT differences suggesting the remaining bias is specific to546 MODIS.

547

#### 548 **5.3** Ice Lidar Ratio Sensitivities in CALIOP

549 As previously discussed, CALIOP uses one of two methods, (i.e., constrained and unconstrained) 550 to retrieve IOT. The constrained method requires high SNR in clear air regions immediately 551 above and below the cloud. This SNR requirement limits the constrained retrieval primarily to 552 nighttime FOVs, because solar background light severely degrades the clear air SNR during the 553 daytime. This precludes direct comparison of the constrained retrievals with the MODIS daytime 554 optical property retrievals. The IR retrieval, being day/night independent, allows for direct inter-555 comparisons between the MODIS IR IOT retrievals and both the constrained and un-constrained 556 CALIOP IOT retrievals providing a means to evaluate the two retrieval methods against a 557 consistent reference.

558 Figure 3b presents the joint histogram between the unconstrained CALIOP and the 559 MODIS window IR IOT for January 2010 for single layer cirrus. The filtering criteria are the 560 same as in Figure 1, except both day and night observations are included. The CALIOP layer 561 optical thickness is filtered using the extinction quality control (QC) flags provided as part of the 562 L2 products. Only QC values of 0 (unconstrained solution, no lidar ratio adjustment), 2 563 (unconstrained solution, lidar ratio decreased) and 4 (unconstrained solution, lidar ratio 564 increased) were selected. Consistent with the findings of (Garnier et al. 2015a), Figure 3b shows 565 CALIOP unconstrained IOT is significantly low-biased with respect to the IR IOT, with a non-566 linear dependence as a function of IOT. Figure 9 compares the CALIOP constrained retrievals 567 (QC=1) to the MODIS IR COT for the same filtering criteria. This comparison reveals a distinct 568 difference between the CALIOP constrained and unconstrained retrievals (Figure 3b), as the 569 constrained retrievals demonstrate a significantly smaller bias relative to the IR IOT. While the 570 CALIOP IOT retrieval requires estimates of the multiple scattering contributions for both the 571 constrained and unconstrained retrievals, the un-constrained method also requires an assumed 572 lidar ratio whereas the constrained retrieval does not. Because both retrievals use an identical 573 fixed multiple scattering factors, the difference between the constrained and unconstrained 574 retrievals relative to the IR can be attributed to the use of an assumed lidar ratio in the 575 unconstrained retrieval.

576 To investigate the sensitivity of the CALIOP IOT retrievals to the lidar ratio, a month of 577 CALIOP L2 products was processed (January 2010) with the default lidar ratio increased to 578 32 sr. This revised value is the mean of all V3 constrained solutions of ice clouds with randomly 579 oriented ice crystals (3,091,952 cases) measured between 28 November 2007 (when CALIPSO permanently changed its pointing angle to 3° off nadir) and 30 June 2012. It is important to note 580 581 that the selection of this new default lidar ratio was based on on-going quality assurance analyses 582 conducted by the CALIOP algorithm team that were wholly independent of the IR inter-583 comparisons with the final value dependent on change to the multiple scattering correction and 584 calibration. In addition the CALIOP team is currently investigating more complex multiple 585 scatteirng parameterization that depends on the cloud temperature (Garnier et al. 2015b). The 586 modified CALIOP product was ingested by the Atmospheric PEATE and collocated with both 587 the MODIS C5 and C6 products and the MODIS IR retrievals. The modified CALIOP 588 unconstrained retrievals compared to the reference IR IOT is presented in Figure 7b. Compared 589 to the standard V3 products (Figure 3b) the change in the lidar ratio significantly reduced the 590 bias compared to IR IOT, and the non-linear behavior at large IOT is almost completely

removed. This is because optical depth is a nonlinear function of lidar ratio, thus weakly scattering layers show minimal changes in IOT while the changes in strongly scattering layers are much more substantial. This result strongly suggests that the current V3 unconstrained lidar ratio of 25 sr should be increased in future versions of the CALIOP data products.

#### 595 **6.** Conclusions

596 MODIS Collection 5 (C5) ice optical thickness (IOT) retrievals are compared to the version 3 (V3) 597 CALIOP IOT for one month (January 2010) of collocated single layer ice clouds. The comparison 598 reveals a factor of two differences between the retrievals as presented in Figure 1. Using IR 599 observations from MODIS as an independent means of assessing the CALIOP and MODIS IOT 600 clearly demonstrates that both retrievals have significant biases, but in opposite directions: MODIS 601 C5 systematically overestimates IOT while CALIOP V3 systematically underestimates IOT.

602 The decision to use the single severely roughened aggregate column habit as the MODIS 603 C6 ice cloud radiative model was made solely to achieve closure with IR retrievals in a global 604 sense. Our use of this model for this purpose does not imply that it is a suitable microphysical 605 model for use in understanding ice particle physical processes (e.g., size distribution evolution, fall 606 speed distribution, etc). Furthermore, the IR comparisons were done in conjunction with collocated 607 CALIOP observations that that allow for the filtering of multi-layer ice phase clouds from the 608 statistical study; the data set used here is clearly a subset of actual scenes and so may not be 609 reflective of the full distribution of ice clouds observed by the sensors. It is important to note that 610 this investigation focuses on ice clouds with optical depths less then 3.0 as this is the limit for 611 CALIOP. Finally, it is recognized that using a fixed ice radiative model for global retrievals is 612 only meaningful in a climatological sense and may be expected to breakdown in instantaneous

and/or regional studies with variations of inferred ice models with region, cloud type, dynamicsand cloud top height shown by (Cole et al. 2014; van Diedenhoven et al. 2014).

615 The severely roughened aggregated column model adopted for the MODIS C6 ice cloud 616 algorithm has a fixed aspect ratio with an asymmetry parameter of about 0.75 in the visible for 617 all effective sizes. This produces results that are quite consistent with those generated using the 618 Inhomogeneous Hexagonal Mono-crystal (IHM) model derived by (C.-Labonnote et al. 2001) 619 (asymmetry parameter of about 0.77) that provided a good match with observed POLDER view 620 angle-dependent VNIR reflectance. Other studies have also suggested that featureless (i.e., 621 smooth) phase functions indicative of roughened or highly asymmetric aggregated habits with 622 relatively small asymmetry parameters are needed to match aircraft and satellite observations 623 e.g., (Baran et al. 2001; C. -Labonnote et al. 2000; van Diedenhoven et al. 2013).

624 The Generalized Habit Model (GHM) (Baum et al. 2010) was also tested but did not 625 result in the same level of radiative closure with the IR IOT retrievals compared to the severely 626 roughened aggregated columns (comparison shown in Fig. 7a). While there was an improvement 627 with respect to the C5 ice model (comparison shown in Fig. 3a), the GHM model resulted in IOT 628 retrievals that were still significantly larger than the IR because of larger asymmetry parameters 629 in the visible relative to the severely roughened aggregated column model (about 0.77 at an 630 effective radius of 5 µm up to 0.82 at 60 µm). (Cole et al. 2012) also tested the GHM as well as 631 single habit models from (Yang et al. 2012) and (Yang et al. 2003) against POLDER polarized 632 and total reflectance observations across a range of scattering angles.

633 Polarized angular observations agreed well with a severely roughened version of the 634 GHM. However, it was concluded that there was no single habit/model that is best in all respects 635 for the reflectance (derived spherical albedo) consistency tests, though the severely roughened aggregated column model was not included in the analysis. Similarly, (Baran and Labonnote 2007) also noted that though the IHM model provided good consistency with POLDER directional reflectance distributions, it was less successful in matching the angular distribution of polarized reflectances. Due to vertical size stratification in ice clouds it is possible that different models are needed to match polarized observations (weighted towards the uppermost portion of the cloud-top) with total reflectance observations (weighted deeper into the cloud), e.g., (Platnick 2000) and (Zhang et al. 2010).

643 Given that MODIS retrievals are based on total reflectance, it is expected that directional 644 reflectance consistency with POLDER is the more relevant metric. Further, the study of (Zhang 645 et al. 2010; Zhang et al. 2009) shows there is little difference between IOT retrieved from 646 reflectance and IR observations for the model case study considered. (Fauchez et al. 2014) 647 demonstrated that for 1km IR observations, sensitivities to 3-D effects are limited to horizontal 648 heterogeneity (plane-parallel approximation or PPA bias) and the effect of vertical heterogeneity 649 is small. Though more extensive heterogeneity studies are needed, these studies do suggest the 650 utility of using IR IOT retrievals to assess MODIS reflectance-based ice radiative models. 651 Finally, we note that recent comparisons have demonstrated consistency between Aqua MODIS 652 C6 IOT retrievals and those from AIRS Version 6 (Kahn 2015).

For CALIOP it is found that the bias relative to the IR for the V3 IOT retrievals depends on the retrieval method used. While CALIOP can make direct measurements of the effective two-way transmittance of the layer, the retrieved optical thickness depends only on an estimate of the multiple scattering factor and the accuracy of the molecular attenuated backscatter profile (calculated from a temperature and pressure profile using Rayleigh scattering theory). However, daytime solar background noise limits the applicability of this constrained retrieval technique to 659 mostly nighttime observations, thus prohibiting direct comparisons to the MODIS daytime optical 660 retrievals. For the constrained retrieval we find good agreement with the IR radiative closure 661 (Figure 2) and the IR IOT in Figure 9. However, the majority of the daytime CALIOP retrievals 662 use the unconstrained method that requires an *a priori* specification of the cloud extinction-to-663 backscatter ratio. It is these unconstrained retrievals that are directly compared to the MODIS C5 664 IOT in Figure 1 and to the IR in Figure 2 and Figure 3. The CALIOP V3 unconstrained IOT 665 retrievals show a significant low bias relative to both the IR and the constrained CALIOP 666 retrievals. Since both CALIOP methods assume an identical multiple scattering correction, this 667 suggests that the default lidar ratio (25 sr) used in the V3 CALIOP unconstrained retrievals is too 668 low. As part of this investigation the CALIOP algorithm team processed a month of retrievals 669 using a lidar ratio of 32 sr for the unconstrained retrievals with results presented in Figure 7b. It is 670 important to note that the selection of a lidar ratio of 32 sr was not based on the IR inter-671 comparison studies, but instead was derived from independent analyses of the nighttime 672 constrained retrievals conducted by the CALIOP algorithm team in order to improve the accuracy 673 of the CALIOP unconstrained retrievals and increase the consistency of IOTs reported by the 674 constrained and unconstrained retrievals.

675

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**Figure 1:** A two dimensional histogram comparing MODIS C5 and CALIOP V3 single layer ice cloud daytime optical thickness retrievals for January 2010 (ocean surfaces,  $\pm 60^{\circ}$  latitude). Notice the color scale is logarithmic.



906 Calculated - Measured 11um (K)
 907 Figure 2 presents the radiative closure results (A) for 1 month (January 2010) of collocated
 908 single layer ice cloud observations using LBLRTM and DISORT to calculate the TOA 11 μm
 909 radiance that are compared to MODIS channel 31 observations. The only difference in the
 910 calculations is the IOT retrieval method. The differences in TOA fluxes resulting from using the

- 911 MODIS or CALIOP daytime IOT retrievals in the calculation are presented the right histogram912 (B).
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914 915 **Figure 3** The 2-D histogram comparing the MODIS C5 (a) and CALIOP V3 (b) retrievals to the

- 916 reference IR IOT retrieval.
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**Figure 4** The MODIS IOT retrievals dependence on scan angle is investigated in the above figures. The image presents the MODIS C5 OT retrievals on January 11 2010 at 06:25 UTC. The right figure presents a histogram of the ratio between the MODIS IOT for both C5 (solid line) and C6 (dashed line) and full swath IR retrieval for only those FOVs which were identified as ice by MODIS. The histograms are separated by view angle the approximate regions for each color marked by the associated color lines on the left image. Notice the significant scan dependent bias relative to the IR IOT for the MODIS C5 retrievals.

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Figure 5. The relationship between effective radius and single scattering co-albedo in the 930 931 MODIS (a) 2.13 and (b) 3.7 µm channels for different ice particle radiative models. See Fig. 6 932 for model details. Since effective radius retrievals for an optically thick cloud are a retrieval of 933 co-albedo, the difference between the C5 and aggregated column model co-albedo implies a retrieved effective radius difference of  $+2 \mu m$  and  $-8 \mu m$ , respectively, for a C5 effective radius 934 935 retrieval of about 35 µm.





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939 Figure 6. The relationship between effective radius and single scatter asymmetry parameter in 940 the MODIS (a) 0.67 and (b) 2.13 µm channels for different ice particle radiative models. Notice 941 the strong dependence of the MODIS C5 model asymmetry parameter on effective size. The

other models consist of a single habit with severely roughened surfaces. The single habitcalculations are made for a modified gamma size distribution and an effective variance of 0.10.





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**Figure 7** The joint histogram comparing the MODIS C6 IOT with the reference IOT retrieval (A). Notice the significant improvement in the agreement resulting from the change to severely roughened aggregated columns. The CALIOP non-constrained IOT using a modified lidar ratio of 32 is compared to collocated IR MODIS retrieved IOT in the right figure (B). Notice the significant improvement in the no-linear bias compared to **Figure 3**b.

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- **Figure 8a** Example retrieval results for an Aqua MODIS data granule (MYD06 2014, 20
- 955 September, 0530 UTC). The RGB composite is shown in the left panel while IOP retrievals for
- 956 Collections 5 and 6 are shown in the center and right panels, respectively. Note the difference in
- 957 the phase determination between the two collections.



**Figure 8b.** Collections 5 and 6 distributions of ice cloud optical thickness and effective radius derived from a combination of the MODIS 0.86 and 2.1  $\mu$ m channels for the data granule of Fig. 8a (or 8). The distributions are limited to common pixels for which both collections agree that the pixel has an ice phase and the retrievals were successful. The IOT modes are at about 1.5 and 1.1 for C5 and C6, respectively, representing about a 27% reduction in the most recent Collection; the effective radius modes increase by about 15%. The mean for the range shown in the plots is given in the legends.

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970 Figure 9 The CALIOP V3 constrained IOT retrieval for single layer clouds is compared to the

BLDIS reference IOT retrieval. Due to single to noise limitations the comparison is limited tonighttime only FOV.







976 Figure 10 (=> Fig. 10) The CALIOP unconstrained IOT but processed using a modified lidar ratio of 32 is compared to the new single habit ice scattering LUT used in the updated MODIS C6 IOT retrievals in the left (A) figure. Notice the improved bias relative to the MODIS C5 and V3 CALIOP retrievals presented in Figure 1. The radiative closure analysis using the updated retrievals is presented in the right (B) figure. The modifications have greatly improved agreement with the measured MODIS 11 µm channel compared to MODIS C5 and the current V3 CALIOP retrievals presented in Figure 2.