

Reply to Referee 1

We are grateful to the referee for the encouraging comments and careful reviews which helped to improve the quality of our paper. In the following we quoted each review question in the square brackets and added our response after each paragraph.

[Review Comment: In this paper, the observational and modeling studies show a relationship between the haze events over CEC and Tibetan Plateau (TP). I think it is a very interesting paper. It provides a new perspective to investigate the causes of the haze pollution in China. I believe this manuscript is appropriate for publication in Atmospheric Chemistry and Physics and would recommend publication subject to primarily minor revisions outlined below.]

Reply 1: Thank you for the encouraging comments.

[1. In this paper, which are the simulation results? Which are the observed results? Please comment on these.]

Reply 2: Following this comment, we have added a paragraph at the beginning of the revised Section 3 (Results and discussion) with following sentences:

“In this Section, we identify the contributions of pollutant emissions and climate change to interannual haze variations (in Sect. 3.1), reveal a climatological “susceptible region” for haze formation in China (in Sect, 3.2), analyze the relationships between TP’s thermal forcing and haze over CEC (in Sect, 3.3) and investigate the TP-warming inducing favorable meteorology for CEC’s haze (in Sect. 3.4) based on the meteorological observations. In order to more convincingly demonstrate the observed results, Section 3.5 presents the results of a sensitivity simulation experiment about impacts of the TP’s thermal forcing on CEC’s aerosol variations.”

[2. How to separate the time series into three phases in Fig.1? How to define the haze event in the figure? Please comment on these.]

Reply 3: In response to the referee's questions, we have clarified the statements in the revised manuscript as follows:

“In accompany with an unceasing increase in the Chinese pollutant emissions during recent decades, the significant interannual variations of haze occurrences in CEC over recent decades could be separated into three interdecadal phases with the trends of slow ascending (4.6d/10a) from the 1960s to 70s, less changing (1.7 d/10a) during the 1980s–1990s and sharply rising with a trend reaching 13.0d/10a going into the 21st century (upper panel of Fig. 1). Although of a continuous increasing trend in the pollutant emissions over the recent decades, the haze variations in CEC have evolved with the different trends of slow, less and sharply ascending over three interdecadal periods,” (Please see the second paragraph of Sect. 3.1).

“This study adopts a widely-used comprehensive haze definition using surface in-situ observations of visibility, relative humidity and weather phenomenon. The observed relative humidity of less than 90% is used to distinguish haze from fog under the visibility <10km, (Schichtel et al., 2001; Doyle and Dorling, 2002; Ding and Liu, 2014).” (Please see the first paragraph of Sect. 2).

References:

Schichtel, B. A., Husar, R. B., Falke, S. R. and Wilson, W. E.: Haze trends over the United States 1980–1995, *Atmos Environ*, 35,5205–5210, 2001.

Doyle. M. and Dorling, S.: Visibility trends in the UK 1950–1997, *Atmos Environ*, 36,3161–3172, 2002.

[3.Please give us the more detailed information about CO2 emission.]

Reply 3: In the revised manuscript, we have added “The Chinese CO₂ emission data over 1961-2012 are downloaded online from the website (http://cdiac.ornl.gov/CO2_Emission/timeseries/national) ” in the first paragraph of Section 2, and “the Chinese CO₂ emission data source: http://cdiac.ornl.gov/CO2_Emission/timeseries/national” in the caption of Fig. 1.

[4. “In accompany with an unceasing increase in the Chinese pollutant emissions in recent decades, the significant interannual variations of haze occurrences in CEC have evolved with the trends of slow ascending from the 1960s to 70s, less changing during the 1980s–1990s and sharply rising with a trend reaching 13.0d/10a going into the 21st century (upper panel of Fig. 1), implying that climate change could also play an important role in the variations of haze events in CEC apart from the anthropogenic dimension of pollutant emission sources related to the rapid industrialization of China.” I think this conclusion looks a little messy. You need to give us more evidence to prove this point.]

Reply 4: Following the suggestions, we have clarified that conclusion in the revised manuscript as follows (in Sect. 3.1):

“In accompany with an unceasing increase in the Chinese pollutant emissions during recent decades, the significant interannual variations of haze occurrences in CEC over recent decades could be separated into three interdecadal phases with the trends of slow ascending (4.6d/10a) from the 1960s to 70s, less changing (1.7 d/10a) during the 1980s–1990s and sharply rising with a trend reaching 13.0d/10a going into the 21st century (upper panel of Fig. 1). Although of a continuous increasing trend in the pollutant emissions over the recent decades, the haze variations in CEC have evolved with the different trends of slow, less and sharply ascending over three interdecadal periods, implying that climate change could also play an important role in the variations of haze events in CEC apart from the anthropogenic dimension of pollutant emission sources related to the rapid industrialization of China. A steady decline of East Asian monsoon winds is negatively correlated to haze occurrences in the CEC with the coefficient of

determination, $R^2=0.6419$ passing the confidence level of 99.9% (lower panel of Fig.1), indicating a consequence of East Asian monsoon climate change to CEC haze pollution.”

[5. Fig.3: Surface wind speed or wind speed at 10 m? Please check it.]

Reply: Yes, it is wind speed at 10 m. It has been corrected in the caption of Fig.3.

[6. Could you give us the information about pollutant emission in lower TP's Q1 (1996,2002) and higher TP's Q1 (1998,2003)?]

Reply: The haze variations are generally controlled by pollutant emissions and meteorological conditions. In order to more convincingly demonstrate the contribution of the TP-warming to the haze frequency over CEC, a sensitivity simulation by employing the global air quality model GEM-AQ/EC is designed to isolate the emission influence on interannual variations of aerosols, where the monthly data of anthropogenic emissions by fossil fuel and biomass burning as well as the sulfate emissions compiled using EDGAR2.0 (Gong et al., 2012) are introduced without any interannual changes from 1995 to 2004. As designed in the sensitivity simulation, the pollutant emissions in lower TP's Q₁ (1996,2002) and higher TP's Q₁ (1998,2003) are same used in the simulation with the emission inventory dataset EDGAR2.0.

In the revised manuscript, we have added the above-discussions in Sect. 2 (the last paragraph) and in Sect. 3.5 (lines 348-349; lines 360-362) .

[7. As we known, dust is one of the absorbing aerosols in the atmosphere, which can influence the climate directly by modulating the radiation budget, affect the microphysical properties of clouds, and alter the surface albedo of the ground covered by snow or glacier. TP dust could impact on regional and global climate (e.g., Lau et al., 2006,2010; Huang et al., 2007; Chen et al., 2013). Could you consider the climatic

effects of the TP dust in this paper? I am wondering whether there is a relationship between TP dust and haze over CEC.

References:

1. Lau, K. M., M. K. Kim, and K. M. Kim (2006), Asian summer monsoon anomalies induced by aerosol direct forcing: The role of the Tibetan Plateau, *Clim. Dyn.*, 26(7-8), 855–864, doi:10.1007/s00382-006-0114-z.

2. Lau, W. K. M., M. K. Kim, K. M. Kim, and W. S. Lee (2010), Enhanced surface warming and accelerated snow melt in the Himalayas and Tibetan Plateau induced by absorbing aerosols, *Environ. Res. Lett.*, 5, 025204, doi:doi:10.1088/1748-9326/5/2/025204.

3. Huang, J., P. Minnis, Y. Yi, Q. Tang, X. Wang, Y. Hu, Z. Liu, K. Ayers, C. Trepte, and D. Winker (2007), Summer dust aerosols detected from CALIPSO over the Tibetan Plateau, *Geophys. Res. Lett.*, 34, L18805, doi:10.1029/2007GL029938, 2007.

4. Chen, S., J. Huang, C. Zhao, Y. Qian, L. R. Leung, and B. Yang, Modeling the transport and radiative forcing of Taklimakan dust over the Tibetan Plateau: A case study in the summer of 2006, *J. Geophys. Res. Atmos.*, 118, 2013.]

Reply: Many thanks for the referee's discussion. We agree with all the suggestions. We have added these in Section 1: Introduction (lines 120-130) as follows:

“Aerosol transport and deposition have been increasingly dirtying and even melting the snow- and ice-dominated wintertime TP (Ramanathan and Carmichael, 2008; Xu et al., 2009). This process leads to decreases in the snow and ice albedos, which could be largely responsible for climate change in the TP region (Hansen and Nazarenko, 2004). As one of the absorbing aerosols in the atmosphere, dust can influence the climate directly by modulating the radiation budget, affect the microphysical properties of clouds, and alter the surface albedo of the ground covered by snow or glacier TP dust could impact on regional and global climate (Lau et al., 2006, 2010; Huang et al., 2007; Chen et al., 2013; Liu et al., 2008). “

We have accordingly cited the following articles in the revised manuscript:

Chen, S., Huang, J., Zhao, C., Qian, Y., Leung, L. R., and Yang, B.: Modeling the transport and radiative forcing of Taklimakan dust over the Tibetan plateau: a case study in the summer of 2006, *J. Geophys. Res.-Atmos.*, 118, 797–812, doi:10.1002/jgrd.50122, 2013.

Lau, K. M., Kim, M. K., and Kim, K. M.: Asian summer monsoon anomalies induced by aerosol direct forcing: the role of the Tibetan Plateau, *Clim. Dynam.*, 26, 855–864, doi:10.1007/s00382-006-0114-z, 2006.

Lau, W. K. M., Kim, M. K., Kim, K. M. and Lee, W. S.: Enhanced surface warming and accelerated snow melt in the Himalayas and Tibetan Plateau induced by absorbing aerosols, *Environ. Res. Lett.*, 5, 025204, doi:10.1088/1748-9326/5/2/025204, 2010.

Liu, Z., Liu, D., Huang, J., Vaughan, M., Uno, I., Sugimoto, N., Kittaka, C., Treppe, C., Wang, Z., Hostetler, C., and Winker, D.: Airborne dust distributions over the Tibetan Plateau and surrounding areas derived from the first year of CALIPSO lidar observations, *Atmos. Chem. Phys.*, 8, 5045–5060, doi:10.5194/acp-8-5045-2008, 2008.

Huang, J., Minnis, P., Chen, B., Huang, Z., Liu, Z., Zhao, Q., Yi, Y., and Ayers, J. K.: Long-range transport and vertical structure of Asian dust from CALIPSO and surface measurements during PACDEX, *J. Geophys. Res.*, 113, D23212, doi:10.1029/2008JD010620, 2008.

Wang, H.-J., Chen, H.-P. and Liu J.-P.: Arctic sea ice decline intensified haze pollution in Eastern China, *Atmos. Oceanic Sci. Lett.*, 8(1), 1-9, doi: 10.3878/AOSL20140081, 2015.

Zhang, R. H., Li, Q. and Zhang, R. N.: Meteorological conditions for the persistent severe fog and haze event over eastern China in January 2013, *Science China: Earth Sciences*, 57, 26–35, doi: 10.1007/s11430-013-4774-3, 2014.

Yan, L. and Liu, X.: Has climatic warming over the Tibetan Plateau paused or continued in recent years?, *Journal of Earth, Ocean and Atmospheric Sciences*, 1, 13-28, 2014.

Duan, A. , Wu, G. X., Zhang, Q., Liu, Y.: New proofs of the recent climate warming over the Tibetan Plateau as a result of the increasing greenhouse gases emissions, *Chinese Science Bulletin*, 51(11), 1396-1400, 2006.

Liu, X. and Chen, B.: Climatic warming in the Tibetan Plateau during recent decades, *Int. J. Climatol.* 20, 1729-1742, 2000.

Reply to Referee 2 (Prof. Dr. Xuhui Lee)

We are grateful to Prof. Dr. Xuhui Lee for the encouraging comments and careful reviews which helped to improve the quality of our paper. In the following we quoted each review question in the square brackets and added our response after each paragraph.

[The authors of this paper propose that meteorological conditions on the Tibetan Plateau are partly responsible for the deteriorating air quality in eastern China. They found a robust correlation between wintertime visibility in eastern China with the heating rate (Q1) on the plateau derived from a reanalysis model. An air quality model was used to attribute changes in air quality to emission changes and to changes in meteorology. They found that during years of negative Q1 anomalies, the surface wind was weakened, the air in the boundary layer became more stable, and more moisture was transported from the ocean. All these factors contributed to unfavorable dispersion conditions. I support publication of the paper in ACP.]

Reply 1: Thank you for the encouraging comments

[The current manuscript style is not suited for the readership of ACP. Because literature review is mixed with methods, results and discussion, it is difficult to determine which statements describe past research by other people and which describe new results from this study. They should restructure the manuscript according to the standard template used by a technical journal, by separating the information clearly into different sections (introduction/literature review, methods, results, discussion, conclusions). Data and sources should be clearly explained. Model scenarios and their rationale should also be explained in the methods section.]

Reply 2: Following the comments, the revised manuscript has been restructured based on the original manuscript with the following outline:

1 Introduction

In this Sect., we present all the literature review with moving the descriptions of past research from the original Sect. 4 (please see lines 120-130 in the revised manuscript with track changes) and removed the new results and analysis from this study, as a part of results and discussion, to Sect. 3.1 (Contributions of pollutant emissions and climate change to interannual haze variations in China).

2 Data and methods

In this Sect., the design of sensitivity simulation is moved from the original Sect. 6 (please see lines 193-202 in the revised manuscript with track changes).

3 Results and discussion

3.1 Contributions of pollutant emissions and climate change to interannual haze variations in China

This Sect. is built with the sentences of the original introduction .

3.2 A climatological “susceptible region” for haze formation in China

3.3 Relationships between TP’s thermal forcing and haze over CEC

we removed the literature review to lines 120-130 (in the revised manuscript with track changes) in Sect. 1 (Introduction).

3.4 The TP-warming inducing favorable meteorology for CEC’s haze

3.5 A sensitivity simulation experiment on effect of the TP-warming

The design of sensitivity simulation is moved to lines 193-202 (in the revised manuscript with track changes) in the Sect. 2(Data and methods).

4 Conclusions

[The first name of all the authors should be spelled out completely. It is very difficult to distinguish Chinese names, and using initials only makes the job even more difficult.]

Reply 3: We have checked that the first names of all the authors were spelled out completely in our submitted PDF-file of manuscript, and after the submission the ACP changed them to the initials for the ACP discussion.

[Quality of the graphs should be improved. Some legends and axis labels are too small or too faint.]

Reply 4: We are sorry for the poor figure quality. All the Figures have been redrawn with high quality for ACP-publication (Please see the new Figs. In the revised manuscript).

[I understand that the authors have plan to continue to elucidate mechanisms underlying the Q1-air quality “teleconnection”. Here I present several questions that arose from my reading of the manuscript. They do not need a response at this time but may be helpful to the authors in their future research:

1) Q1 is a quantity derived from atmospheric reanalysis. Given that meteorological observations are sparse on the Tibetan Plateau, I wonder how robust this quantity is. The current research relies solely on one reanalysis product (NCEP?). Do other reanalyses show similar Q1 variabilities (both inter-annual and long-term trend)?

2) It has been known for some time that the eastern Asian monsoon has been weakening in the past three decades. As far as I know, there is no accepted theory for explaining why this is occurring. This study presents empirical evidence suggesting that the weakening trend is tied to changes in land cover on the plateau, as implicated by the strong correlation between Q1 and wind. In this regard, two questions are worth pursuing: a) what types of change on the surface (snow albedo change, glacier melt, desertification, shrinking of shrub land, etc) are responsible for the changing Q1 signal in the atmosphere? d) How do you test the TP hypothesis versus other competing hypotheses regarding the monsoon trend?

3) Related to my point 2 above, the author suggests that higher Q1 values are caused by contamination of snow cover by pollutants transported to the plateau, due to the fact that dirtier snow absorbs more solar energy. Can you find direct evidence for a declining trend in the snow albedo? If this proves true, the snow albedo effect constitutes a positive feedback on air quality at the regional scale, which may have contributed to the observed decline in visibility. However, the feedback mechanism cannot explain the large year-to-year variabilities in Q1. It is intriguing

that Q1 and air quality are correlated both in terms of the decadal trend and in terms of interannual variabilities. Why does Q1 change rapidly from one year to the next?]

Reply 4: Your constructive suggestions are greatly appreciated and very helpful for our further study. In this preliminary study based on long-term observational data and a sensitivity simulation experiment, we investigate a relationship between the haze pollution in China and TP's environment and climate changes. It should be emphasized that considering the quality of reanalysis data over and around the TP, a comparison between NCEP/NCAR and other reanalysis data sets such as JRA-25, ERA-Interim, or MERRA is necessary in further work. The understandings of TP's thermal forcing changes and East Asian monsoon declines are challenging topics. The impacts of TP's climate change on air quality in China could be further studied on the shifts in weather patterns, pollutant emissions, depositions and chemical reactions in the atmosphere to comprehensively understand the meteorological drivers of air quality in a changing climate and also to consider the ocean-related impacts of climate change.

We have accordingly added the above discussions at the end of Sect. 4 (lines 402-410 of Conclusions in revised manuscript with track changes).

1 **Climate modulation of the Tibetan Plateau on haze in China**

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Key Points

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25 1) A large-scale "susceptible region" for haze occurrences is climatologically
26 identified over central-eastern China (CEC) harbored by the Tibetan Plateau (TP).

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28 2) Thermal anomalies of the TP induce the change in meteorological drivers
29 downstream for frequent haze events in CEC.

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31 3) Implications of the TP-topography for atmospheric environment will be having
32 potential utility for development planning in China.

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Abstract

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46 Rapid increases in pollutant emissions in conjunction with stagnant meteorological
47 conditions result in haze pollution in China. Recent frequent haze in China has
48 attracted worldwide attention. Here we show a relationship between the haze events
49 and Tibetan Plateau (TP)'s environment and climate changes. Based on
50 observational data taken over recent decades, we identify central-eastern China (CEC)
51 as a climatological large-scale "susceptible region" of frequent haze, which is
52 harbored by the TP with its impact on mid-latitude westerly winds. The
53 observational and modeling studies demonstrate that the interannual variations in the
54 thermal forcing of TP are positively correlated with the incidences of wintertime haze
55 over CEC. Further analysis indicates that the TP-climate warming induced changes in
56 atmospheric circulation driving frequent haze events in CEC. The frequent haze
57 occurrences in CEC are consistent with decreasing winter monsoon winds,
58 intensifying downward air flows and increasing atmospheric stability in the lower
59 troposphere over the CEC in association with upstream plateau's thermal anomalies.
60 Therefore, variations of haze in China are related to mechanical and thermal forcing
61 by the TP. Our results also suggest that implications of the large TP-topography for
62 environment and climate changes should be taken into account for air pollution
63 mitigation policies in China.

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67 **1 Introduction**

68 Haze in poor visibility with high particulate matter (PM) or aerosol levels is a
69 pervasive air quality problem facing China, posing a major challenge for public health
70 (Huang et al., 2014;Zhang et al., 2015). The frequent haze pollution has been notable
71 for hitting record high levels of PM pollution over central-eastern China (CEC) in
72 history since 1961 (Ding and Liu, 2014). In January 2013, extremely severe and
73 persistent haze events swept over much of CEC-region. A large area of CEC from
74 the North China Plain, including Beijing, across the Lower Yangtze Valley Plain to the
75 Sichuan Basin, was blanketed in thick haze and smog for almost one month. It is
76 estimated by the Chinese government that this wintertime haze covered a quarter of
77 the total land area in China with 600 million people, half of the Chinese population,
78 exposed to the haze air pollution (NDRC, 2013;Gu, 2013). China's National
79 Meteorological Center released its first ever "haze" orange alert (CMA, 2010) in
80 response to the air quality index frequently reaching hazardous levels for this regional
81 haze event. The PM_{2.5} (PM with an aerodynamic diameter less than 2.5 micrometers)
82 concentrations at 33 cities in the CEC region were more than 300 $\mu\text{g m}^{-3}$ for longer
83 than half a month, and some monitors reported hourly peak PM_{2.5} levels of 900 $\mu\text{g m}^{-3}$,
84 which is classified as "Beyond Index" (NDRC, 2013;Gu, 2013). The suffering of
85 those in China from haze and poor air quality has attracted worldwide attention(Wang
86 et al., 2014;Arden Pope III and Dockery, 2013;Chen et al., 2013b;Kan et al.,
87 2011;Park et al., 2013;Zhao et al., 2013).

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89 China has been experiencing the increased air pollution, commonly attributed to the
90 large increases in pollutant emissions associated with the rapid economic
91 development. However, air quality is modulated by changes in meteorology and
92 climate (Tagaris et al., 2009;Zhang et al., 2014;Wang et al., 2015). In accompany
93 with an unceasing increase in the Chinese pollutant emissions in recent decades, the
94 CEC region had observed the significant interannual variations of haze occurrences
95 (Ding and Liu, 2014). The changing East Asian monsoon climate could also play an
96 important role in the variations of haze events in CEC apart from the anthropogenic
97 dimension of pollutant emission sources related to the rapid industrialization of China.
98 The surface wind speed associated with East Asian monsoons has significantly
99 weakened in both winter and summer in the recent three decades (Xu et al., 2006;Oey
100 et al., 2013). The weakening of the East Asian monsoons could increase air pollutants
101 mainly by the changes in atmospheric circulation and weather conditions (Zhu et al.,
102 2012;Niu et al., 2010). Weak advection of cold air, in conjunction with strong
103 subsidence and stable atmospheric stratification, can easily produce a stagnation area
104 in the lower troposphere resulting in regional pollutant accumulations, which are
105 favorable for the development of CEC haze events (Zhao et al., 2013). In addition,
106 in the presence of high soil moisture, strong surface evaporation results in increases in
107 the near-surface relative humidity, which is also conducive to haze formation(Xiao et
108 al., 2011). The contribution of the meteorological factors to the variance of the daily
109 haze evolution was estimated to reach 0.68, which could explain more than 2/3 of the
110 variance for the persistent severe haze events over CEC in January 2013 (Zhang et al.,

111 2014). By changing East Asian winter monsoon climate, the Arctic sea ice decline
112 could intensify haze pollution in CEC (Wang et al., 2015).

113

114 It is generally accepted that meteorological conditions in China are closely connected
115 to the large topography of Tibetan Plateau (TP) (Yanai et al., 1992; Xu et al., 2010; Wu
116 et al., 2012; Ye and Wu, 1998). Precipitation, land surface temperature and surface
117 air temperature have increased on the TP over the past decades (Zhong et al., 2011).

118 The TP has exhibited the largest surface radiative flux changes induced by aerosols
119 (e.g. black carbon and dust) contaminating snow and ice compared to any other
120 snow- and ice-covered regions in the world (Qian et al., 2011). Aerosol transport and

121 deposition have been increasingly dirtying and even melting the snow- and
122 ice-dominated wintertime TP (Ramanathan and Carmichael, 2008; Xu et al., 2009).

123 This process leads to decreases in the snow and ice albedos, which could be largely
124 responsible for climate change in the TP region (Hansen and Nazarenko, 2004). As
125 one of the absorbing aerosols in the atmosphere, dust can influence the climate

126 directly by modulating the radiation budget, affect the microphysical properties of
127 clouds, and alter the surface albedo of the ground covered by snow or glacier TP. Dust

128 transport and depositions could impact on regional climate and environment over the
129 TP (Lau et al., 2006; Lau et al., 2010; Huang et al., 2008; Chen et al., 2013a; Liu et al.,

130 2008). The question remains whether the rapid changes in climate experienced by

131 the TP could exert an influence on the haze variations in the downstream CEC region,

132 the lower flatlands harbored by the large TP-topography. The consequent processes

133 linking the TP-climate change with the CEC haze pollution should be highly possible
134 and worth investigating, even though it is obvious that increasing anthropogenic
135 pollutant emissions contribute to high haze frequency (Zhang et al., 2013). In this
136 study, we attempt to determine the physical connection between climate change in the
137 TP and haze occurrences in the CEC region to more comprehensively understand the
138 large-area haze formation in China especially under the background of global
139 warming with the TP's environment and climate changes.

140

141 **2 Data and methods**

142 In this study, we used the observational records of visibility, weather phenomenon,
143 relative humidity and 10m wind from 1961 to 2012 archived at the China
144 Meteorological Administration (CMA) , and the meteorological variables of air
145 temperature, winds and relative humidity from the reanalysis data generated by the
146 US National Center for Environmental Prediction–National Center for Atmospheric
147 Research (NCEP/NCAR). This study adopts a widely-used comprehensive haze
148 definition using surface in-situ observations of visibility, relative humidity and
149 weather phenomenon. The observed relative humidity of less than 90% is used to
150 distinguish haze from fog under the visibility <10km, (Schichtel et al., 2001;Doyle
151 and Dorling, 2002;Ding and Liu, 2014). The Chinese CO₂ emission data during
152 1961-2012 are downloaded online from the website
153 (http://cdiac.ornl.gov/CO2_Emission/timeseries/national).

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155 Following the studies of Yanai (1961), Yanai and Johnson (1993) , Yanai and Tomita
156 (1998), the apparent heat source (Q_1) and apparent moisture sink (Q_2) are calculated.
157 Atmospheric heat sources and moisture sinks are respectively gauged with the Q_1 and
158 Q_2 . As Q_1 includes Q_2 and radiative heating, here we concentrate only on the
159 collective effect of apparent heating (Q_1) over the TP. The heat source column (in
160 units of $W m^{-2}$) over the TP is obtained with both horizontal and vertical integration of
161 Q_1 over the TP-area of 78°E-103°E and 28°N-38°N covering the most region with the
162 altitude of higher than 3000m (see the large TP-rectangle in upper panel of Fig. 1) to
163 form a one-dimensional variable representing the TP-thermal forcing. The correlation
164 coefficients between the TP-heat source column and the meteorological variables (U-,
165 V- and W-components of wind and air temperature) are calculated to build their
166 horizontal and vertical distributions of correlations. Zonal, meridional and vertical
167 components of the correlation vector are respectively derived through the correlation
168 coefficients of the TP-heat source column to U-, V- and W-components of vector of
169 wind and air temperature, indicating the changes in wind and air temperature induced
170 by the TP-thermal forcing.

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172 In the modeling study, we used the global air quality model system GEM-AQ/EC,
173 which is an integration of gas phase chemistry and aerosol modules in the
174 meteorological model GEM (Global Environmental Multiscale weather prediction
175 model of Environment Canada). Full details of the development and evaluation of
176 GEM-AQ/EC are given by Gong et al. (2012). The validations of 10 year (1995-2004)

177 GEM-AQ/EC modeling prove that the model provided satisfactory simulations of the
178 distribution and variation of global and regional aerosols (Gong et al., 2012;Zhao et
179 al., 2012). Regional variations of aerosols in East Asia are reasonably captured by the
180 GEM-AQ/EC modeling compared to the observed aerosol concentrations and aerosol
181 optical depth.

182

183 Air quality change is generally driven by two factors: pollutant emissions and
184 meteorology. In order to exclude the emission influence on interannual variations of
185 aerosols, a sensitivity simulation with GEM-AQ/EC is designed without year-to-year
186 changes in anthropogenic aerosol emissions from 1995 to 2004 for an assessment on
187 the impact of the TP-warming on air quality change in China. The sensitivity
188 simulation experiment of GEM-AQ/EC was configured with 28 hybrid vertical levels
189 and the model top at 10 hPa as well as the horizontal model grids in a global uniform
190 resolution of $1^{\circ} \times 1^{\circ}$. The GEM-AQ/EC was run with the fully nudged variables of
191 wind, temperature, pressure and water vapor of NCEP-reanalysis meteorology every
192 24 hours from 1995 to 2004.

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194 **3 Results and discussion**

195 In this Section, we identify the contributions of pollutant emissions and climate
196 change to interannual haze variations (in Sect. 3.1), reveal a climatological
197 "susceptible region" for haze formation in China (in Sect, 3.2), analyze the
198 relationships between TP's thermal forcing and haze over CEC (in Sect, 3.3) and

199 investigate the TP-warming inducing favorable meteorology for CEC's haze (in Sect.
200 3.4) based on the meteorological observations. In order to more convincingly
201 demonstrate the observed results, Section 3.5 presents the results of a sensitivity
202 simulation experiment about impacts of the TP's thermal forcing on CEC's aerosol
203 variations.

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205 **3.1 Contributions of pollutant emissions and climate change to interannual** 206 **haze variations in China**

207 China has experienced the huge increases in CO₂ emissions from fossil fuel
208 combustion with the certain attendant pollutant emissions and aerosol loading over
209 recent decades (upper panel of Fig. 1), which has a direct physical link to more
210 frequent haze occurrences in situ in China. The regional emissions of air pollutants
211 contribute largely to the haze pollution in CEC with a high coefficient of
212 determination, $R^2=0.9025$ between interannual variations of haze frequency and CO₂
213 emission in China (upper panel of Fig. 1), reflecting that the frequent haze events are
214 strongly associated with the large increases in anthropogenic pollutant emissions in
215 recent decades.

216

217 In accompany with an unceasing increase in the Chinese pollutant emissions during
218 recent decades, the significant interannual variations of haze occurrences in CEC over
219 recent decades could be separated into three interdecadal phases with the trends of
220 slow ascending (4.6d/10a) from the 1960s to 70s, less changing (1.7 d/10a) during the

221 1980s–1990s and sharply rising with a trend reaching 13.0d/10a going into the 21st
222 century (upper panel of Fig. 1). Although of a continuous increasing trend in the
223 pollutant emissions over the recent decades, the haze variations in CEC have evolved
224 with the different trends of slow, less and sharply ascending over three interdecadal
225 periods, implying that climate change could also play an important role in the
226 variations of haze events in CEC apart from the anthropogenic dimension of pollutant
227 emission sources related to the rapid industrialization of China. A steady decline of
228 East Asian monsoon winds is negatively correlated to haze occurrences in the CEC
229 with the coefficient of determination, $R^2=0.6419$ passing the confidence level of
230 99.9% (lower panel of Fig.1), indicating a consequence of East Asian monsoon
231 climate change to CEC haze pollution.

232

233 **3.2 A climatological “susceptible region” for haze formation in China**

234 Examination of ground-based observations of the frequency of haze events from 1961
235 to 2012 (CMA, 2010) reveals that the haze air pollution in China typically has the
236 highest levels in the CEC region covering a vast area from the eastern edges of the TP
237 and the Loess Plateau to China’s Pacific coast, and haze occurrences in CEC oscillate
238 seasonally between the peak in winter and the low in summer (Fig. 2). Based upon
239 these climate data, we could climatologically regard the CEC, with the lowlands
240 harboured by the upstream plateaus of western China, as a large-scale “susceptible
241 region” of frequent haze events in China (left panel of Fig. 2). Upper panel of
242 Figure 3 shows that low average wind speeds tend to be coincident with the centers of

243 pollutant haze events over the CEC (left panel of Fig. 2), reflecting the climatological
244 “susceptible region” of haze occurrences in connection with a stagnation area in the
245 lower troposphere in China.

246

247 Due to the influence of the TP terrain on the typical westerly winds in this region, the
248 air flowing from the windward plateaus descends in a north-south oriented zone
249 between about 110°E and 125°E (middle panel of Fig. 3). Accompanying this strong
250 downward current are weak winds in the near-surface layers that lie in the lee side of
251 the plateaus. These air flow and wind condition lead to development of a “harbor”
252 that accumulates air pollutants in the CEC region. The weak wind and downward
253 current areas coincide well with the centers of frequent haze events in China (middle
254 and lower panels of Fig. 3). The “susceptible region” of haze events over the CEC
255 region from the eastern edge of the plateaus to the lower flatlands is associated with
256 the “harbor” effect of the unique TP topography under specific meteorological
257 conditions that trap air pollutants.

258

259 Because haze is climatologically mostly a winter phenomenon in the CEC (right panel
260 of Fig. 2), the following analysis on the TP’s climate effect on haze pollution in CEC
261 and the related mechanisms is focused on the winter season.

262

263 **3.3 Relationships between TP’s thermal forcing and haze over CEC**

264 As a vast elevated landmass, the TP acts thermodynamically as a synoptic-scale

265 wintertime cooling source protruding into the free atmosphere(Qiu, 2008;Liu and
266 Chen, 2000b;Ruddiman and Kutzbach, 1989;Yeh et al., 1957). The TP region, as a
267 wintertime cooling source (negative values of apparent heat source Q_1), has been
268 experiencing a warming trend over recent decades, especially since 2001 (upper left
269 panel of Fig. 4). A striking climate warming over the TP during the last decades has
270 been revealed by many studies (Liu and Chen, 2000a;Duan et al., 2006;Yan and Liu,
271 2014).

272

273 Against the backdrop of global climate change, the question may be posed: Does the
274 warming of the TP region cause changes in the atmospheric environment in China
275 resulting in more frequent haze events in the CEC? The historical data analysis
276 indicates that a significant correlation exists between the wintertime cooling source
277 represented by the apparent heat source column Q_1 integrated over the TP and the
278 number of haze days averaged regionally in the CEC over recent decades (upper panel
279 of Fig. 4). It is also found in upper panel of Fig. 4 that the changes of wintertime Q_1
280 over the TP were reversed from cooling to warming in the late 1990s, which could be
281 connected with the trends in haze occurrences with less changing over the 1980s-90s
282 and sharp increasing during the 21st century in China under the increases in pollutant
283 emission levels (upper panel of Fig. 1). Based on the composite analysis on the haze
284 frequencies in winter with positive and negative anomalies in wintertime cooling
285 source of the TP, the haze increasing and decreasing incidences over the CEC are
286 found in good agreement with the positively and negatively anomalous TP-cooling

287 sources (lower panel of Fig. 4). The frequency of haze events over the CEC region is
288 positively correlated with climatic warming over the TP.

289

290 **3.4 The TP-warming inducing favorable meteorology for CEC's haze**

291 Further analyses provide information on the mechanisms relating climatic warming of
292 the TP and enhancement in haze occurrence in the CEC. The favorable meteorology
293 for haze occurrences is well known to be lower wind speeds, weaker vertical mixing,
294 stronger subsidence, higher air humidity and more stable low-level stratification. We
295 are still pondering the question whether climatic warming of the TP could strengthen
296 the aforementioned meteorological conditions downstream for frequent haze events in
297 the CEC to reveal the mechanism how thermal anomalies of the TP in climate change
298 influence the incidence of haze over the CEC

299

300 The East Asian winter monsoon, which climatologically prevails over the CEC,
301 typically maintains near-surface northeastern winds(Ding, 1994). In upper panel of
302 Figure 5, two horizontal components of the correlation vector are derived through two
303 correlation coefficients of Q_1 to U- and V-surface wind components, respectively,
304 where the arrow length denotes the combined correlation with a longer arrow
305 implying a better correlation, and the arrow direction means the direction of
306 anomalous wind induced by the TP-thermal effect. The correlation vector over the
307 CEC in upper panel of Figure 5 indicates that the variations of thermal forcing over
308 the TP could give rise to the weakening winter monsoon winds (southwest wind

309 anomalies) induced by changes of Q_1 over the TP. Furthermore, the anomalous south
310 wind components resulting from climate change in the TP (positive correlations of Q_1
311 to V-wind components in upper panel of Fig. 5) can enhance transport of water vapor
312 from the oceans to the CEC(Niu et al., 2010). By increasing the moisture in the lower
313 troposphere driven by the strong vapor transport(Hung and Kao, 2010) (lower panel
314 of Fig. 5), in addition to decline in the East Asian winter monsoon with weak
315 advection of cold air, haze formation can be enhanced(Zhao et al., 2013;Xiao et al.,
316 2011).

317

318 Upper panels of Figure 6 present the results of composite analysis on vertical
319 variations in air temperature in five winters respectively with the most positive and the
320 most negative anomalies in the TP cooling source. The air temperature changes with
321 upper warmer and lower cooler are found in the middle and lower troposphere over
322 the CEC region in winter with positive Q_1 anomalies on the TP, and an inverse
323 structure of the CEC air temperature changes in winter with negative Q_1 anomalies on
324 the TP (upper panel of Fig. 6). The TP's warming and cooling anomalies could lead to
325 a "warm shield" and a "cool shield" in the atmosphere over the CEC. The correlation
326 analyses of observation data over 1981-2012 confirm that the vertical structure of
327 anomalous air temperature similar to that induced the TP's positive thermal effect
328 (upper-left panel of Fig.6) with a "warm shield" intensifying the subsidence in the
329 lower troposphere is responsible for the frequent haze occurrences over the CEC
330 (lower panel of Fig. 6). Associated with the warming TP, the vertical variations of air

331 temperature with upper warmer and lower cooler could easily build an inversion layer
332 in the atmosphere over the polluted CEC, which results in more stably stratified
333 atmosphere in this region (Fig. 6). Heavy haze pollution processes in winter are
334 highly related with the existence of atmospheric inversion layer(Xu et al., 2003).

335

336 The cumulative consequences of weakening winter monsoon winds, intensifying
337 downward air flows, a more humid and more stable atmosphere as the favorable
338 meteorological conditions for haze formation would be expected to strengthen the air
339 pollutant “harbor” effect of the TP and increase the number and severity of haze
340 events in the CEC. Therefore, the haze formation over CEC is significantly modulated
341 by the TP’s climate change under the increase and even without changes in the current
342 levels of anthropogenic pollutant emissions.

343

344 **3.5 A sensitivity simulation experiment on effect of the TP-warming**

345 In order to more convincingly demonstrate the connection of the TP-warming to the
346 haze frequency over CEC, a sensitivity simulation by employing the global air quality
347 model GEM-AQ/EC is designed to isolate the emission influence on interannual
348 variations of aerosols, where the monthly data of anthropogenic emissions by fossil
349 fuel and biomass burning as well as the sulfate emissions compiled using EDGAR2.0
350 (Gong et al., 2012) are introduced without any interannual changes from 1995 to 2004.

351 The results of this sensitivity simulation are used to assess the impact of climate
352 change on interannual change of air quality over CEC in this study.

353

354 Haze and aerosol changes are determined by both pollutant emission and meteorology,
355 and the effects of meteorology are difficult to separate from aerosol observations. The
356 10-year GEM-AQ/EC simulation without interannual changes in the anthropogenic
357 emissions provides a possibility to identify the meteorological effect on the
358 interannual variations of aerosols. To investigate the implications of TP's climate
359 change for interannual aerosol variations in CEC's haze, a composite analysis of
360 surface aerosol concentrations over CEC (Fig. 7) were performed for two winters with
361 lower TP's Q_1 (1996, 2002) and two winters with higher TP's Q_1 (1998, 2003) during
362 the simulation period of 1995-2004 according to the interannual Q_1 changes over the
363 TP (Fig. 4). As designed in the sensitivity simulation, the pollutant emissions in lower
364 TP's Q_1 (1996,2002) and higher TP's Q_1 (1998,2003) are same in the simulation with
365 the emission inventory dataset EDGAR2.0 (Gong et al., 2012). Because the effect of
366 emissions was singled out in the interannual aerosol variations modeled in the
367 sensitivity simulation experiment, the simulated variations in aerosol concentrations
368 over CEC could be purely attributed to the changes of meteorological drivers in the
369 context of changing climate. The analysis results show that the TP heating
370 anomalies could lead to enhancements of 30-45% in wintertime surface aerosol
371 concentrations over the CEC-region compared to the winters with the TP cooling
372 anomalies (Fig. 7). Because changes of aerosol levels in the surface atmosphere
373 determine haze formation, this sensitivity simulation confirmed that the frequent haze
374 in China with the significantly interannual variations is closely related to thermal

375 forcing by the TP, and climate change of the TP could intensify pollutant haze in
376 China even without increases in the current anthropogenic pollutant emissions.

377

378 **4. Conclusions**

379 Based on observational data over the recent decades, we identify the CEC region, the
380 lower flatlands along the eastern plateau edges in China as a climatological
381 large-scale “susceptible region” of pollutant haze in connected with downward
382 currents and weak near-surface winds in consequence of the “harbor” impact of large
383 TP topography on mid-latitude westerlies. The climate analysis reveals that the
384 increasingly frequent haze in the CEC region is related with decreasing winter
385 monsoon winds, intensifying descending air and increasing atmospheric stability in
386 the lower-troposphere over the CEC in association with plateau’s thermal anomalies.
387 Climate impact of the TP’s mechanical and thermal forcing driving changes in
388 atmospheric circulation and meteorological conditions downstream is potentially
389 contributing to the increasing trend in haze events in China. A sensitivity simulation
390 also confirmed that the frequent haze in CEC with the significantly interannual
391 variations is closely connected with thermal forcing by the TP.

392

393 The Chinese government has been making great strides in reducing emissions and
394 mitigating air pollution. However, the interplay of China’s unique landform
395 distribution with climate change and its associated more extreme weather events
396 could impair the effectiveness of air pollution control measures in China. With the

397 influence of the TP climate change, the CEC region is facing a bigger challenge to
398 realize air quality maintenance plan. The TP “harbor effect” and climate change
399 should be considered in making decisions on the locations of new industrial facilities
400 for development planning in China in order to preferentially reduce anthropogenic
401 emissions in the “susceptible region” of haze and in turn reduce the number and
402 severity of haze events in the central-eastern region of China.

403

404 In this preliminary study based on long-term observational data and a sensitivity
405 simulation experiment, we investigate a relationship between the haze pollution in
406 China and TP’s environment and climate changes. It should be emphasized that
407 considering the quality of reanalysis data over and around the TP, a comparison
408 between NCEP/NCAR and other reanalysis data sets such as JRA-25, ERA-Interim,
409 or MERRA is necessary in further work. The understandings of TP’s thermal forcing
410 changes and East Asian monsoon declines are challenging topics. The impacts of TP’s
411 climate change on air quality in China could be further studied on the shifts in
412 weather patterns, pollutant emissions, depositions and chemical reactions in the
413 atmosphere to comprehensively understand the meteorological drivers of air quality in
414 a changing climate and also to consider the ocean-related impacts of climate change.

415

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424 **References**

- 425 Arden Pope III, C., and Dockery, D. W.: Air pollution and life expectancy in China and beyond PNAS, ,
426 doi: 10.1073/pnas.1310925110 2013.
- 427 Chen, S., Huang, J., Zhao, C., Qian, Y., Leung, L. R., and Yang, B.: Modeling the transport and radiative
428 forcing of Taklimakan dust over the tibetan plateau: a case study in the summer of 2006, *J. Geophys.*
429 *Res.*, 118, 797–812, doi:10.1002/jgrd.50122, 2013a.
- 430 Chen, Y. Y., Ebenstein, A., Greenstone, M., and Li, H. B.: Evidence on the impact of sustained exposure
431 to air pollution on life expectancy from China's Huai River policy, PNAS doi:
432 10.1073/pnas.1300018110, 2013b.
- 433 CMA: China Meteorological Administration, National meteorological standard of the People's
434 Republic China: Observation and forecasting levels of haze (in Chinese), QX-T 113-2010, 2010.
- 435 Ding, Y. H.: Monsoons over China, Kluwer Academic Publishers, Dordrecht/Boston/London, 1994.
- 436 Ding, Y. H., and Liu, Y. J.: Analysis of long-term variations of fog and haze in China in recent 50 years
437 and their relations with atmospheric humidity., 57: , *Science China: Earth Sciences*, 57, 36-46, doi:
438 10.1007/s11430-11013-14792-11431, 2014.
- 439 Doyle, M., and Dorling, S.: Visibility trends in the UK 1950–1997, *Atmos Environ*, 36, 3161–3172, 2002.
- 440 Duan, A., Wu, G. X., Zhang, Q., and Liu, Y.: New proofs of the recent climate warming over the Tibetan
441 Plateau as a result of the increasing greenhouse gases emissions, *Chinese Science Bulletin*, 51,
442 1396-1400, 2006.
- 443 Gong, S. L., Lavoué, D., Zhao, T. L., Huang, P., and Kaminski, J. W.: GEM-AQ/EC, an on-line global
444 multi-scale chemical weather modeling system: model development and evaluation of global aerosol
445 climatology,, *Atmos. Chem. Phys.*, 12, 8237-8256, doi: 10.5194/acp-12-8237-2012, 2012.
- 446 Gu, L.: Half of Chinese live in haze: report, *China News*, at
447 <http://www.ecns.cn/cns-wire/2013/07-12/72889.shtml> 2013.
- 448 Hansen, J., and Nazarenko, L.: Soot climate forcing via snow and ice albedos, PNAS, 101, 423-428,
449 2004.
- 450 Huang, J., Minnis, P., Chen, B., Huang, Z., Liu, Z., Zhao, Q., Yi, Y., and Ayers, J. K.: Long-range transport
451 and vertical structure of Asian dust from CALIPSO and surface measurements during PACDEX, *J.*
452 *Geophys. Res.*, 113, D23212, doi:10.1029/2008JD010620, 2008.
- 453 Huang, R.-J., Zhang, Y., Bozzetti, C., Ho, K.-F., Cao, J.-J., Han, Y., Daellenbach, K. R., Slowik, J. G., Platt, S.
454 M., Canonaco, F., Zotter, P., Wolf, R., Pieber, S. M., Brun, E. A., Crippa, M., Ciarelli, G., Piazzalunga, A.,
455 Schwikowski, M., Abbazade, G., Schnelle-Kreis, J., Zimmermann, R., An, Z., Szidat, S., Baltensperger,
456 U., Haddad, I. E., and Prevot, A. S. H.: High secondary aerosol contribution to particulate pollution

457 during haze events in China, *Nature*, 514, 218-222, 10.1038/nature13774
458 [http://www.nature.com/nature/journal/v514/n7521/abs/nature13774.html#supplementary-informa](http://www.nature.com/nature/journal/v514/n7521/abs/nature13774.html#supplementary-information)
459 [tion](http://www.nature.com/nature/journal/v514/n7521/abs/nature13774.html#supplementary-information), 2014.

460 Hung, C.-W., and Kao, P.-K.: Weakening of the winter monsoon and abrupt increase of winter rainfalls
461 over northern taiwan and southern china in the early 1980s, *J. Climate*, 23, 2357 - 2367,
462 10.1175/2009JCLI3182.1, 2010.

463 Kan, H., Chen, R., and Tong, S.: Ambient air pollution, climate change, and population health in China,
464 *Environmental international*, doi: 10.1016/j.envint.2011.1003.1003, 2011.

465 Lau, K. M., Kim, M. K., and Kim, K. M.: Asian summer monsoon anomalies induced by aerosol direct
466 forcing: the role of the Tibetan Plateau, *Clim. Dynam.*, 26, 855–864, doi:10.1007/s00382-006-0114-z,
467 2006.

468 Lau, W. K. M., Kim, M. K., Kim, K. M., and Lee, W. S.: Enhanced surface warming and accelerated snow
469 melt in the Himalayas and Tibetan Plateau induced by absorbing aerosols, *Environ. Res. Lett.*, 5,
470 025204, doi:10.1088/1748 9326/5/2/025204, 2010.

471 Liu, X., and Chen, B.: Climatic warming in the Tibetan Plateau during recent decades, *Int. J.*
472 *Climatol.* , 20, 1729-1742, 2000a.

473 Liu, X., and Chen, B.: Climatic warming in the Tibetan Plateau during recent decades, *Int. j Climatol.*,
474 20, 1729-1742, 2000b.

475 Liu, Z., Liu, D., Huang, J., Vaughan, M., Uno, I., Sugimoto, N., Kittaka, C., Trepte, C., Wang, Z., Hostetler,
476 C., and Winker, D.: Airborne dust distributions over the Tibetan Plateau and surrounding areas
477 derived from the first year of CALIPSO lidar observations, *Atmos. Chem. Phys.*, 8, 5045-5060,
478 doi:10.5194/acp-8-5045-2008, 2008.

479 NDRC: Chinese National Development and Reform Commission Report; available at
480 http://www.sdpc.gov.cn/jjxsfx/t20130710_549549.htm, 2013.

481 Niu, F., Li, Z., Li, C., Lee, K.-H., and Wang, M.: Increase of wintertime fog in China: Potential impacts of
482 weakening of the Eastern Asian monsoon circulation and increasing aerosol loading, *Journal of*
483 *Geophysical Research: Atmospheres*, 115, D00K20, 10.1029/2009jd013484, 2010.

484 Oey, L.-Y., Chang, M.-C., Chang, Y.-L., Lin, Y.-C., and Xu, F.-H.: Decadal warming of coastal China Seas
485 and coupling with winter monsoon and currents, *Geophys. Res. Lett.*, 40 5, 10.1002/2013GL058202,
486 2013.

487 Park, S.-U., Cho, J.-H., and Park, M.-S.: Analyses of high aerosol concentration events (dense haze/mist)
488 occurred in East Asia during 10-16 January 2013 using the data simulated by the Aerosol Modeling
489 System, *International Journal of Chemistry*, 03, 10-26, 2013.

490 Qian, Y., Flanner, M. G., Leung, L. R., and Wang, W.: Sensitivity studies on the impacts of Tibetan
491 Plateau snowpack pollution on the Asian hydrological cycle and monsoon climate, *Atmos. Chem. Phys.*,
492 11, 1929-1948, 10.5194/acp-11-1929-2011, 2011.

493 Qiu, J.: China: The third pole, *Nature*, 454, 393-396, 2008.

494 Ramanathan, V., and Carmichael, G.: Global and regional climate changes due to black carbon, *Nature*,
495 1, 221-227, 2008.

496 Ruddiman, W. F., and Kutzbach, J. E.: Forcing of late Cenozoic northern hemisphere climate by plateau
497 uplift in southern Asia and the American west, *J. Geophys. Res.*, D15, 18409-18427,
498 doi:18410.11029/JD18094iD18415p18409., 1989.

499 Schichtel, B. A., Husar, R. B., Falke, S. R., and Wilson, W. E.: Haze trends over the United States
500 1980-1995, *Atmos Environ*, 35, 5205–5210, 2001.

501 Tagaris, E., Liao, K. J., Delucia, A. J., Deck, L., Amar, P., and Russell, A. G.: Potential impact of climate
502 change on air pollution-related human health effects, *Environ Sci Technol*, 43(13), 4979-4988, 2009.

503 Wang, H.-J., Chen, H.-P., and Liu, J.-P.: Arctic sea ice decline intensified haze pollution in Eastern China,
504 *Atmos. Oceanic Sci. Lett.*, 8, 1-9, doi: 10.3878/AOSL20140081, 2015.

505 Wang, Y., Zhang, R. Y., and Saravanan, R.: Asian pollution climatically modulates midlatitude
506 cyclones following hierarchical modeling and observational analysis, *Nature Commun.*
507 , <http://dx.doi.org/10.1038/ncomms4098>, 2014.

508 Wu, G. X., Liu, Y., He, B., Q. Bao, Duan, A., and Jin, F.-F.: Thermal Controls on the Asian Summer
509 Monsoon, *Scientific Reports* 2, 404, 10.1038/srep00404, 2012.

510 Xiao, Z.-M., Zhang, Y.-F., Hong, S.-M., Bi, X.-H., Jiao, L., Feng, Y.-C., and Wang, Y.-Q.: Estimation of the
511 main factors influencing haze, based on a long-term monitoring campaign in Hangzhou, China,
512 *Aerosol and Air Quality Research*, 11, 873-882, doi: 10.4209/aaqr.2011.4204.0052, 2011.

513 Xu, B., Cao, J. J., Hansen, J., Yao, T., Joswila, D. R., Wang, N., Wu, G., Wang, M., Zhao, H., Yang, W., Liu,
514 X., and He, J.: Black soot and the survival of Tibetan glaciers, *PNAS*, 106, 22114-22118, 2009.

515 Xu, M., Chang, C.-P., Fu, C., Qi, Y., Robock, A., Robinson, D., and Zhang, H.: Steady decline of east
516 Asian monsoon winds, 1969-2000: Evidence from direct ground measurements of wind speed *J.*
517 *Geophys. Res.*, 111, D24111, 10.1029/2006JD007337, 2006.

518 Xu, X., Ding, G., Zhou, L., Zheng, X., Bian, L., Qiu, J., Yang, L., and Mao, J.: Localized 3D-structural
519 features of dynamic-chemical processes of urban air pollution in Beijing winter, *Chinese Science*
520 *Bulletin*, 48(8), 819-825, 2003.

521 Xu, X., Lu, C., Shi, X., and Ding, Y.: Large-scale topography of China: A factor for the seasonal
522 progression of the Meiyu rainband?, *J. Geophys. Res.*, 115, D02110,
523 doi:10.1029/2009JD012444, 2010.

524 Yan, L., and Liu, X.: Has climatic warming over the Tibetan Plateau paused or continued in recent
525 years?, *Journal of Earth, Ocean and Atmospheric Sciences*, 1, 13-28, 2014.

526 Yanai, M.: A detailed analysis of typhoon formation, *J. Meteor. Soc. Japan*, 39, 187-214, 1961.

527 Yanai, M., Li, C. F., and Song, Z. S.: Seasonal heating of the Tibetan Plateau and its effects on the
528 evolution of the Asian summer monsoon, *J. Meteor. Soc. Japan*, 70, 319-351, 1992.

529 Yanai, M., and Johnson, R. H.: Impacts of cumulus convection on thermodynamic fields. In *The*
530 *Representation of Cumulus Convection*
531 *in Numerical Models of the Atmosphere*, Emanuel KA, Raymond DJ(eds), Vol. 24. AMS Monograph;
532 39-62., 1993.

533 Yanai, M., and Tomita, T.: Seasonal and Interannual Variability of Atmospheric Heat Sources and
534 Moisture Sinks as Determined from NCEP-NCAR Reanalysis, *J. Climate*, 11, 463-482, 1998.

535 Ye, D. Z., and Wu, G. X.: The role of the heat source of the Tibetan Plateau in the general circulation,
536 *Met. Atmos. Phys.*, 67, 181-198, 1998.

537 Yeh, T. C., Luo, S. W., and Chu, P. C.: The wind structure and heat balance in the lower troposphere
538 over Tibetan Plateau and its surrounding, *Acta Meteor. Sin.*, 28, 108-121, 1957.

539 Zhang, R. H., Li, Q., and Zhang, R. N.: Meteorological conditions for the persistent severe fog and haze
540 event over eastern China in January 2013 *Science China: Earth Sciences*, 57, 26-35, doi:
541 10.1007/s11430-013-4774-3, 2014.

542 Zhang, X. Y., Sun, J., Wang, Y., Li, W., Zhang, Q., Wang, W., Quan, J., Cao, G., Wang, J., Yang, Y., and
543 Zhang, Y.: Factors contributing to haze and fog in China (in Chinese), *Chinese Sci. Bull.*, 58, 1178-1187,
544 doi:10.1360/972013-150, 2013

545 Zhang, X. Y., Wang, J. Z., Wang, Y. Q., Liu, H. L., Sun, J. Y., and Zhang, Y. M.: Changes in chemical
546 components of aerosol particles in different haze regions in China from 2006 to 2013 and contribution
547 of meteorological factors, *Atmos. Chem. Phys.*, 15, 12935-12952, 10.5194/acp-15-12935-2015, 2015.
548 Zhao, T. L., Gong, S. L., Huang, P., and Lavoué, D.: Hemispheric transport and influence of meteorology
549 on global aerosol climatology, *Atmos. Chem. Phys.*, 12, 7609-7624, doi:10.5194/acp-12-7609-2012,
550 2012.
551 Zhao, X. J., Zhao, P. S., Xu, J., Meng, W., , P., W. W., Dong, F., He, D., and Shi, Q. F.: Analysis of a winter
552 regional haze event and its formation mechanism in the North China Plain, *Atmos. Chem. Phys.*, 13,
553 5685-5696, doi: 5610.5194/acp-5613-5685-2013, 2013.
554 Zhong, L., Su, Z., Ma, Y., Salama, M. S., and Sobrino, J. A.: Accelerated Changes of Environmental
555 Conditions on the Tibetan Plateau Caused by Climate Change, *J. Climate*, 24, 6540-6550, 2011.
556 Zhu, J., Liao, H., and Li, J.: Increases in aerosol concentrations over eastern China due to the
557 decadal-scale weakening of the East Asian summer monsoon, *Geophys. Res. Lett.*, 39, L09809,
558 10.1029/2012GL051428, 2012.

559

560 **Figure Captions**

561

562 Figure 1. Interannual variations in the total CO₂ emissions in mainland China and the
563 haze event frequency averaged in the CEC region over 1961-2012 with their
564 relationship in the inserted figure (upper panel). The Chinese CO₂ emission data
565 source: http://cdiac.ornl.gov/CO2_Emission/timeseries/national. Two blue dotted
566 lines separate the time series into three phases of the 1960s-70s, the 1980s-90s and the
567 21st century. Interannual variations in wind at 10m (blue line) and the number of
568 days with haze (red line) in the CEC over 1961-2012 (lower-left panel) and their
569 scatter plot (lower-right panel).

570

571 Figure 2. Distribution of annual haze event frequency (days per year) averaged over
572 1961-2012 in China, and Chinese topography of the TP and the Loess Plateau with
573 altitudes is shown with yellow shades (left panel). Monthly variation of haze
574 frequency averaged from 1961-2012 over the CEC region (right panel).

575

576 Figure 3. **Near-surface** wind speed distribution (m s^{-1}) averaged over 1961-2012 in
577 China with the red rectangle marking the region for cross sections in the middle and
578 lower panels (upper panel). Cross sections of horizontal wind speed (m s^{-1} ; filled color
579 contours) and vertical circulations illustrated by stream lines (middle panel) and zonal
580 variations of annual haze event frequency (lower panel) at 27°N - 41°N averaged over
581 1961-2012. Note that near-surface vertical and horizontal winds are not illustrated
582 well here due to north-south variations in the terrain and approximation of the
583 location of the TP in this figure. All fields are for the annual-averages.

584

585 Figure 4. Interannual variability in the apparent heat source Q_1 (the negative values
586 denote cooling) integrated vertically over the TP and haze event frequency averaged
587 in the CEC in winter (December, January and February) over 1980-2012 and their
588 correlation (upper panel). The differences of haze frequencies (days) averaged in
589 five winters with most positive (lower left panel) and most negative Q_1 anomalies
590 (lower right panel) on the TP relative to the mean haze frequency from 1980 to 2012.

591

592 Figure 5. The distributions of the interannual correlations of the apparent heat
593 source Q_1 over the TP to the local V-component of surface wind in winter over
594 1961- 2012 (color shading). Arrows denote correlation vectors (showing both
595 correlation coefficients of Q_1 to U- and V-surface wind components) in China. The

596 correlation coefficients of 0.12 (-0.12), 0.14 (-0.14) and 0.19 (-0.19) respectively
597 passing the significance levels of 90%, 95% and 99%. A vertical section of the
598 trends in vapour content (g kg^{-1} per 10 years) in winter over 1961-2012 averaged
599 along 27°N - 41°N (lower panel).

600

601 Figure 6. Vertical sections of the anomalous air temperature ($^{\circ}\text{C}$) averaged along
602 27°N - 41°N in five winters with most positive (upper left panel) and most negative Q_1
603 anomalies (upper right panel) on the TP from 1980 to 2012, and vertical sections of
604 the correlations of the number of haze days with air temperature (lower left panel) and
605 vertical circulations (lower right panel) in winter from 1980 to 2012.

606

607 Figure 7. The percentages (%; contour lines) of differences of surface aerosol
608 concentrations between winters of 1998 and 2003 with positive and winters of 1996
609 and 2002 with negative Q_1 anomalies on the TP relative to the surface aerosol levels
610 averaged over winters of 1996, 1998, 2002 and 2003 ($\mu\text{g m}^{-3}$; color contours) modeled
611 by the sensitivity simulation experiment with GEM-AQ/EC.

612