1

2

Secondary formation of nitrated phenols: insights from observations during the Uintah Basin Winter Ozone Study (UBWOS) 2014

3 Bin Yuan^{1,2}, John Liggio³, Jeremy Wentzell³, Shao-Meng Li³, Harald Stark^{2,4}, James M.

4 Roberts¹, Jessica Gilman^{1,2}, Brian Lerner^{1,2}, Carsten Warneke^{1,2}, Rui Li^{1,2}, Amy

5 Leithead³, Hans D. Osthoff⁵, Robert Wild^{1,2}, Steven S. Brown^{1,6}, and Joost A. de

6 Gouw^{1,2,6}

7 1. NOAA Earth System Research Laboratory (ESRL), Chemical Sciences Division,

- 8 Boulder, CO, USA
- 9 2. Cooperative Institute for Research in Environmental Sciences, University of Colorado
- 10 at Boulder, Boulder, CO, USA
- 11 3. Environment Canada, Science and Technology Branch, Toronto, ON, Canada
- 12 4. Aerodyne Research Inc., Billerica, MA, USA
- 13 5. Department of Chemistry, University of Calgary, Calgary, Canada

14 6. Department of Chemistry and Biochemistry, University of Colorado at Boulder, CO,

15 USA

16

18 Abstract

19 We describe the results from online measurements of nitrated phenols using a time 20 of flight chemical ionization mass spectrometer (ToF-CIMS) with acetate as reagent ion 21 in an oil and gas production region in January and February of 2014. Strong diurnal 22 profiles were observed for nitrated phenols, with concentration maxima at night. Based 23 on known markers (CH₄, NO_x, CO₂), primary emissions of nitrated phenols were not 24 important in this study. A box model was used to simulate secondary formation of 25 phenol, nitrophenol (NP) and dinitrophenols (DNP). The box model results indicate that 26 oxidation of aromatics in the gas phase can explain the observed concentrations of NP 27 and DNP in this study. Photolysis was the most efficient loss pathway for NP in the gas 28 phase. We show that aqueous-phase reactions and heterogeneous reactions were minor 29 sources of nitrated phenols in our study. This study demonstrates that the emergence of 30 new ToF-CIMS (including PTR-TOF) techniques allows for the measurement of 31 intermediate oxygenates at low levels and these measurements improve our 32 understanding on the evolution of primary VOCs in the atmosphere.

34 1. Introduction

35 Nitrated phenols are a family of aromatic compounds with both nitro (-NO₂) and 36 hydroxyl groups (-OH) connected to a benzene ring. Nitrated phenols have been detected 37 in the gas phase, aerosol, cloud water and rainwater (Harrison et al., 2005a). Many 38 studies have shown that nitrated phenols are one of the important components of brown 39 carbon in aerosol (Desyaterik et al., 2013;Mohr et al., 2013;Zhang et al., 2013;Lin et al., 40 2015), as they absorb light in the atmosphere (Bejan et al., 2007). Photolysis of some 41 nitrated phenols was reported to produce nitrous acid (HONO) (Bejan et al., 2006) and 42 hydroxyl (OH) radicals (Cheng et al., 2009), while the oxidation and photolysis of them 43 contribute to secondary organic aerosol (SOA) formation, especially in biomass burning 44 plumes (Mohr et al., 2013;Kitanovski et al., 2012;Lauraguais et al., 2014). There is also 45 evidence that nitrated phenols are phytotoxic and contribute to forest decline (Rippen et 46 al., 1987; Natangelo et al., 1999). Some nitrated phenols are known to be mutagenic and 47 are of concern to human health (Fernandez et al., 1992).

48 Sources of nitrated phenols in the atmosphere include emissions from vehicle 49 exhaust (Inomata et al., 2013; Tremp et al., 1993; Sekimoto et al., 2013) and biomass 50 burning (Mohr et al., 2013). Nitrated phenols are also produced from photooxidation of 51 aromatic hydrocarbons in the atmosphere: for example benzene oxidizes to 2-nitrophenol 52 (2-NP) and 4-nitrophenol (4-NP), and toluene oxidizes to methylnitrophenols (MNP) 53 (Harrison et al., 2005a). Figure 1 shows the reactions leading to secondary formation of 54 NP and dinitrophenols (DNP) in the atmosphere (Jenkin et al., 2003). Oxidation of 55 benzene by OH radicals forms phenol and further reactions of phenol with either OH or 56 NO_3 radicals yield phenoxy (C_6H_5O) radicals, which react further with NO_2 to generate 57 NP. In addition to benzene oxidation, C₆H₅O radicals are also generated from the reaction 58 of NO with phenyl peroxy ($C_6H_5O_2$) radicals, a product from reactions of some other 59 aromatic precursors, e.g. benzaldehyde (Caralp et al., 1999). Further oxidation of nitrated 60 phenols by obtaining another nitro- group produces dinitrophenols (DNP). The yields of 61 NP from phenol oxidation by OH radicals (Atkinson et al., 1992;Olariu et al., 62 2002;Berndt and Boge, 2003) and NO₃ radicals (Atkinson et al., 1992;Bolzacchini et al., 63 2001) have been reported. Berndt and Boge (2003) also showed that the NP yield from

OH oxidation of phenol increases at higher NO₂ concentrations. In addition to gas phase
reactions, nitrated phenols are formed from aqueous-phase reactions in aerosol or cloud
water (Vione et al., 2001, 2005). The importance of the aqueous reactions compared to
gas-phase reactions is highly dependent on liquid water content in the atmosphere
(Harrison et al., 2005b).

69 The sinks of nitrated phenols in the gas phase include reactions with OH radicals 70 (Atkinson et al., 1992;Bejan et al., 2007), with NO₃ radicals (Atkinson et al., 1992), with 71 chlorine atoms (Bejan et al., 2015) and photolysis (Bejan et al., 2007; Chen et al., 2011). 72 It has been proposed that photolysis is the dominant gas phase atmospheric loss for 73 nitrated phenols (Bejan et al., 2007). Despite the importance of photolysis of nitrated 74 phenols, the photolysis frequency of nitrated phenols under ambient conditions was only 75 reported in a single non-peer reviewed publication (1.4% of photolysis frequency of NO₂) 76 (Bardini, 2006). The chemical products from photolysis of nitrated phenols were 77 proposed, but the proposed products were not fully evaluated against experimental results 78 (Bejan et al., 2006). Nitrated phenols are also removed by various processes in the 79 aqueous phase, including reactions with OH, NO₃ and photolysis (Vione et al., 2009).

80 Measurements of nitrated phenols have been mainly conducted using offline 81 methods (Harrison et al., 2005a). Air samples are usually collected on filters or cartridges 82 and then analyzed by liquid chromatography (LC) methods (Rubio et al., 2012;Harrison 83 et al., 2005a; Delhomme et al., 2010). These detection methods are time-consuming and 84 measurements as a function of the time of day are not usually possible (Delhomme et al., 85 2010). The lack of fast-response online measurements has prevented, at least partially, a 86 thorough investigation of sources and sinks of nitrated phenols. Recently, Mohr et al. 87 (2013) deployed a chemical ionization mass spectrometer (CIMS) using acetate as the 88 reagent ion to measure nitrated phenols online in the particle phase in winter of London 89 and based on their measurements the authors concluded that nitrated phenols were mainly 90 from wood burning in this region of the atmosphere.

In this study, we conducted high time resolution measurements of nitrated phenols
in the gas phase at a site in an oil and gas production region in winter. High
concentrations of ozone and secondary products (Edwards et al., 2014) were observed at

94 this site, as the result of photochemical degradation of large amounts of alkanes and

aromatics emitted from oil and gas production in this region (Warneke et al., 2014).

96 Using the present dataset, we investigate diurnal variations, sources and sinks of nitrated

97 phenols. We use a box model to analyze the budget of nitrated phenols in the atmosphere,

98 and provide insights into the formation mechanism of nitrated phenols.

99 2. Measurements

The Uintah Basin Winter Ozone Study (UBWOS 2014) was conducted in January
and February of 2014 at the Horse Pool site in the Uintah Basin, where over 10,000
active oil and gas wells are located.

103 2.1 Acetate ToF-CIMS

104 2.1.1 Instrument operation

105 An Aerodyne time-of-flight (ToF) CIMS (Lee et al., 2014) that uses acetate 106 $(CH_3C(O)O)$ as the reagent ion was deployed at the Horse Pool site during the UBWOS 107 2014 to measure organic acids, inorganic acids and nitrated phenols. These compounds 108 are ionized in the ion-molecule reaction region (IMR, 61.8±0.3 mbar) via proton 109 abstraction (Veres et al., 2008) or by a sequence of clustering-declustering/deprotonation 110 reactions (Brophy and Farmer, 2015) in the reaction with acetate ions. Acetate ions were 111 produced by introducing saturated acetic anhydride/N₂ mixture (5 mL/min) mixed with another flow of N₂ (2.5 L/min) into a polonium-210 (210 Po) radioactive source. The 112 113 instrument was operated at strong de-clustering conditions by applying voltages in the 114 first quadrupole ion guide (i.e. SSQ, 2.50±0.01 mbar) during UBWOS 2014, with the 115 ratio of acetate cluster (CH₃C(O)O⁻•CH₃C(O)OH)/acetate (CH₃C(O)O⁻) at $0.4\%\pm0.1\%$. 116 Under such de-clustering conditions, the conjugate anions were usually observed as the 117 product ions, with little contribution from cluster ions. The reagent ions and product ions 118 are analyzed using a high-resolution time of flight mass spectrometer (Tofwerk AG, Switzerland). The signals of acetate ion were approximately $1-2 \times 10^6$ count per second 119 (cps) during the campaign (ToF extraction frequency=25 kHz). The mass resolution of 120 121 the ToF analyzer during UBWOS 2014 was approximately 3200 for ions of m/z > 200.

122 Background signals associated with the instrument were measured every 2 hours 123 for 15 min by passing ambient air through three stages of zero air generation: a platinum 124 catalytic converter heated to 350 °C, nylon wool coated with sodium bicarbonate 125 (NaHCO₃), and activated charcoal, which were used in series to remove acidic gases 126 from the sample air and determine instrument backgrounds. During the UBWOS 2014 127 study, two CIMS inlets constructed from Teflon tubing heated to ~40 °C with similar 128 lengths (~10 m) placed at heights of 1 m and 18.5 m above ground were switched 129 automatically every 30 min during the period of January 24- February 1, to measure the 130 vertical concentration gradient of nitrated phenols and other acidic gases. Inlet switching 131 between a long-heated and a short-unheated inlet was conducted during February 1-5, to 132 explore possible inlet interferences to CIMS measurements of nitrated phenols from the 133 long-heated inlet. We did not observe differences in signals between the long and short 134 inlets for nitrated phenols except DNP (Figure S1), indicating that potential loss in the 135 sampling line was minimal for the reported single nitrated phenols in this study. The inlet 136 issues for DNP will be discussed in Section 2.1.2.

137 2.1.2 Data processing

138 The ToF-CIMS data was processed using the Tofware software package 139 (www.tofwerk.com/tofware) written in Igor Pro (Wavemetrics Inc., USA). The detailed 140 data processing procedures are presented in recent studies (Yatavelli et al., 2014;Stark et 141 al., 2015). Post-measurement mass calibrations were performed using nine isolated ions, 142 including m/z 31.9904 (O₂⁻), m/z 34.9694 (Cl⁻), m/z 44.9982 (CHO₂⁻), m/z 59.0139 143 $(C_2H_3O_2)$, m/z 61.9884 (NO_3) , m/z 143.9840 $(C_3F_4O_2)$, m/z 162.9824 $(C_3F_5O_2)$, m/z144 193.9808 (C₄F₆O₂⁻) and m/z 243.9776 (C₅F₈O₂⁻). The four fluorine-containing ions in the 145 list were released from the Teflon inlet during UBWOS 2014 and their persistent 146 presence was used for mass calibration. The accuracy of mass calibration was 4.7 ± 1.9 147 ppm for the whole campaign and the errors of mass calibration for individual ions were 148 usually within 10 ppm (average+ 3σ). The fitted raw signals for the targeted compounds 149 were normalized using an acetate signal at the level of 1×10^6 cps. 150

150 The fitted m/z used for quantification of concentrations of nitrated phenols in the 151 acetate CIMS are m/z 138.0197 (C₆H₄NO₃⁻) for NP, m/z 152.0353 (C₇H₆NO₃⁻) for MNP, 152m/z 166.0510 (C₈H₈NO₃⁻) for dimethylnitrophenol+ ethylnitrophenol (DMNP) and m/z153183.0047 (C₆H₃N₂O₅⁻) for DNP. Compounds with the same molecular formulas as154nitrated phenols include phenyl nitrates/benzyl nitrates, methoxynitrobenzenes,155nitrobenzyl alcohols and hydroxycarboxylic acids derived from pyridine. The first three156groups of compounds have lower acidities than acetic acid (Bartmess, 2015) and hence157they are unlikely to be observed in acetate CIMS, while hydroxycarboxylic acids derived158from pyridine are expected to be small in the atmosphere.

159 High-resolution (HR) peak fitting to m/z 138, m/z 152 and m/z 183 in the averaged 160 mass spectra of ToF-CIMS on a typical day (January 25, 2014) are shown as examples in 161 Figure 2. Isotope signals from lower masses (dark green lines) accounted for small 162 fractions of the m/z signals. Multiple overlapping ion peaks were identified in the m/z163 channels. In addition to nitrated phenols, several ions without deprotonation were also 164 present in the even m/z (e.g. C₈H₁₀O₂⁻ at m/z 138), possibly due to electron transfer 165 reactions and/or fragmentation in the quadrupole ion guides (Stark et al., 2015). The 166 signals of NP and MNP were either the largest or significantly larger than their neighboring peaks at their respective m/z, whereas the signal of DNP was much smaller 167 168 than its neighboring peaks on January 25, 2014. Smaller ratios of the signals between the 169 targeted peak and its neighboring peaks have been shown to deteriorate the precision of 170 the fitted signals for the targeted peak (Cubison and Jimenez, 2015;Müller et al., 171 2011; Corbin et al., 2015). Based on the provided equations in Cubison and Jimenez 172 (2015), the imprecision arising from mass calibration (not including counting error) for the signals of NP, MNP and DNP are 3.2%, 1.8% and 47% based on the mass spectra of 173 174 January 25, respectively. Imperfect mass calibration can also affect fitted magnitudes of 175 ion signals. Figure S2 shows the sensitivity of the fitted signals of various masses as a 176 function of the errors in mass calibration. The signal changes at 10 ppm (average+ 3σ) 177 error of mass calibration relative to the perfect mass calibration (error=0 ppm) for NP, 178 MNP and DNP signals are as high as 14%, 5% and 81%, respectively. The results from 179 both precision calculation and sensitivity of fitted magnitudes indicate that the peak 180 signals of NP and MNP can be fitted well with low uncertainties. The peak fitting at m/z181 166 for DMNP shows similar results as m/z 138 for NP and m/z 152 for MNP. However, 182 large uncertainties are associated with the peak signals of DNP on January 25, 2014,

which is mainly affected by the $C_3F_6HO_2^-$ ions (*m/z* 182.9886) as indicated by the opposite behaviors of the DNP ion and $C_3F_6HO_2^-$ ion in Figure S2C.

185 The C₃F₆HO₂⁻ ion (m/z 182.9886) was released from the heated Teflon inlet along 186 with other fluorine-containing ions that were used for mass calibration. The release of 187 $C_{3}F_{6}HO_{2}$ ion was supported by much higher signals from the long-heated inlet compared 188 to the short-unheated inlet, when inlet-switching experiments were conducted in February 189 2-5 (Figure S1). Long-heated inlets were used for most of the time during UBWOS 2014 190 (January 23- February 13), except during January 18-22, when a short-unheated inlet was 191 used. The averaged mass spectra of m/z 183 measured on January 18 is shown in Figure 192 2D. Compared to the mass spectra on January 25, $C_3F_6HO_2^-$ signals on January 18 were 193 lower and the signals of DNP were larger than of $C_3F_6HO_2^-$ ions. As a result, the 194 uncertainty from peak fitting for the DNP ion was much lower on January 18 (Figure 195 S2D). Thus, we will only use measured DNP data in the beginning of the campaign 196 (January 18-22), when the long heated inlet was not connected to the acetate CIMS and 197 no inlet switching was performed.

198 The response of the CIMS instrument for nitrated phenols, including 2-NP, 4-NP, 2-199 methyl-4-nitrophenol and 2,5-dinitrophenol, was calibrated using a Liquid Calibration 200 Unit (LCU, IONICON Analytik). In the LCU, a water solution with known 201 concentrations of the targeted compounds is nebulized and diluted by another gas stream 202 at different flow rates to produce a gas standard at various concentrations (Kaser et al., 203 2013). The results of the calibrations to various nitrated phenols are shown in Table 1. 204 The sensitivity of 4-NP in our instrument was determined to be higher than that of 2-NP 205 by a factor of 2.1. A higher sensitivity of 4-NP in acetate CIMS was reported in Mohr et 206 al. (2013), but in that study the difference was significantly larger at three orders of 207 magnitude (Mohr et al., 2013). The different sensitivity ratios of 4-NP/2-NP can be 208 caused by many different instrumental conditions between our instrument and that in 209 Mohr et al. (2013), such as the amount of acetic anhydride introduced into the instrument, 210 IMR and SSQ pressures, and de-clustering settings in the quadrupole ion guides, all of 211 which affect sensitivities of acetate CIMS significantly (Stark et al., 2012). The main 212 reagent ions in IMR were shown to be acetic acid-acetate clusters rather than acetate 213 (Bertram et al., 2011) and the cluster distributions in IMR may depend on operated

214 pressure in IMR and the amount of acetic anhydride introduced into the ion source. While 215 de-clustering in SSQ helps the interpretation of recorded mass spectra, de-clustering also 216 obscures a precise understanding of cluster distributions in IMR and hence accurate 217 prediction of sensitivities in acetate CIMS. This result also emphasizes the importance of 218 instrument calibrations in deriving concentration from acetate CIMS. We note that 3-219 nitrophenol (3-NP) is not usually present in the atmosphere (Harrison et al., 2005a). 220 Thus, the average of the sensitivities of 2-NP and 4-NP was used for calculating 221 concentrations of NP. DMNP was not calibrated in this study and we assumed the same 222 sensitivity as determined for MNP.

223 The accuracies of nitrated phenols measurements by the CIMS are conservatively 224 estimated to be around 40% for NP and 50% for other nitrated phenols, mainly arising 225 from uncertainties in the concentration output of the LCU (~10%), uncertainties 226 associated with calibration procedures (\sim 5%), errors in high-resolution (HR) peak fittings 227 to mass spectra (see above and Figure S2), and the representativeness of the calibrated 228 species to other isomers (0-30% for NP and 0-40% for other nitrated phenols). Assuming 229 random errors in the observed ion counts follow a Poisson distribution, detection limits of 230 nitrated phenols, i.e. concentrations with a signal to noise ratio (S/N) of 3, are calculated 231 to be 0.1-0.3 ppt for 1-min average data (Table 1). Following the discussions in Bertram 232 et al. (2011), the measured background ion counts in ToF-CIMS drift over time and thus 233 detection limits are more appropriately calculated as the concentrations at three times the 234 standard deviation of the measurement background counts. The determined detection 235 limits of nitrated phenols increase to the range of 0.3-0.5 ppt based on this approach 236 (Bertram et al., 2011) (Table 1).

237 **2.2 Other measurements**

Volatile organic compounds (VOCs), including hydrocarbons and oxygenates, were
measured using an online gas chromatograph-mass spectrometer (GC-MS) (Gilman et al.,
2013). A commercial proton transfer reaction time of flight mass spectrometer (PTRTOF) (IONICON Analytik, Austria) was also deployed at the Horse Pool site to measure
various VOC species (Warneke et al., 2015). Measurements of phenol, cresols and
dimethylphenols+ethylphenols (DMP) were accomplished by the PTR-TOF at *m/z*

- 244 95.0491 ($C_6H_6OH^+$), *m/z* 109.0648 ($C_7H_8OH^+$) and *m/z* 123.0804 ($C_8H_{10}OH^+$),
- respectively. An example of high-resolution peak fitting to m/z 95 in the mass spectra of
- the PTR-TOF is shown in Figure S3. The sensitivities to these phenols are estimated here
- from the calibrated sensitivities of m/z 93.0699 (toluene), m/z 107.0855 (C8-aromatics)
- and m/z 121.1012 (C9-aromatics) and the ratio of proton-transfer rate coefficients (k) of
- the phenols versus the aromatic hydrocarbons (Cappellin et al., 2012) (see details in the
- SI). Considering the uncertainties in the rate coefficients k, the accuracies of the
- determined concentrations of phenols can be up to 50% (de Gouw and Warneke, 2007).
- 252 Measurements of NO₃ and N₂O₅ were conducted by a cavity ring-down
- spectroscopy instrument (Dubé et al., 2006). NO_x (NO+NO₂), NO_y and O₃ were
- 254 measured with another cavity ring-down spectroscopy instrument (Wild et al., 2014).
- 255 Measurements of methane (CH₄) and carbon dioxide (CO₂) were performed with a
- commercial cavity ring-down spectrometry instrument (Picarro G2301). A pair of
- 257 commercial spectral radiometers (Metcon Inc.) was used to measure photolysis
- 258 frequencies of ozone and NO₂.
- 259

9 **3. Results and Discussions**

260 **3.1 Diurnal variations**

261 Measured diurnal profiles of NP, MNP and DMNP during the UBWOS 2014 are 262 shown in Figure 3. Very strong diurnal variations in concentrations of these nitrated 263 phenols were observed. Concentrations of nitrated phenols were higher at night and lower 264 in the daytime. The ratios between the concentrations in the two hours around midnight 265 (23:00-1:00 MST) and in the two hours around noon (11:00-13:00 MST) are 2.9, 3.9 and 266 4.7 for NP, MNP and DMNP, respectively. This indicates that the substituted alkyl 267 groups enhance the diurnal variations of nitrated phenols, either through larger source at 268 night or stronger loss in the daytime.

Primary emissions of VOCs and NO_x at the Horse Pool site are predominantly due to oil and gas production activities, as the Horse Pool site is surrounded by oil and gas production wells. VOCs and NO_x emitted from nearby oil and gas wells led to periodic concentration spikes during the UBWOS campaigns (Warneke et al., 2014; Yuan et al., 2015). Figure 4 shows two types of episodes encountered during UBWOS 2014. The first 274 was associated with high concentrations of methane and benzene, as an example of 275 fugitive emissions from oil and gas wells. No enhancement of nitrated phenols was 276 observed for the first emission episode. The second episode was associated with high 277 concentrations of NO_v , NO_x and CO_2 , as an example of either vehicular emissions or 278 other fuel combustion activities related to oil and gas extractions (e.g., compressors, 279 dehydrators and pump jacks). High NO_x/NO_y ratios (0.96±0.01) indicate that a fresh 280 combustion plume was encountered. We observed small enhancement of NP during the 281 second emission episode. The enhancement ratio of NP/NO_v in this plume is determined to be $4.6\pm0.7\times10^{-3}$ ppt/ppb, which is comparable with the reported NP/NO_x emission 282 ratios $(1-50 \times 10^{-3} \text{ ppt/ppb})$ from gasoline and diesel vehicles (Inomata et al., 283 284 2013;Sekimoto et al., 2013). Using the obtained enhancement ratio of NP/NO_v, we determine that primary emissions from combustion sources only account for less than 2% 285 286 of NP concentrations during UBWOS 2014. In addition to these primary sources, 287 biomass burning was not observed in the UBWOS campaigns, based on the absence of 288 any enhancement of biomass burning markers like acetonitrile. We conclude that primary 289 emissions of nitrated phenols were not significant during UBWOS 2014.

In addition to primary emissions, secondary formation from oxidation of phenols is an important source for nitrated phenols (Harrison et al., 2005a). Phenol exhibited a concentration maximum in the afternoon (Figure 3B). The diurnal profile of phenol is more similar to that of secondary acetaldehyde, than that of primary emitted benzene. It suggests that secondary formation was the most important source of phenol. Substituted phenols (cresols and DMP) also had similar diurnal variations as phenol.

296 **3.2 Modeling analysis for NP**

297 3.2.1 Box model results

We will focus on NP to understand the budget of nitrated phenols, because NP had higher concentrations than the substituted nitrated phenols (MNP and DMNP) and there is more information on sources and sinks of NP in the literature. A series of zerodimensional box model simulations on the formation of phenol and NP were conducted using the online AtChem tool (https://atchem.leeds.ac.uk). The MCM v3.2 (Jenkin et al., 2012) was used as the chemical mechanism in the box model. We note that ambient

304 temperature (-5±5 °C) during UBWOS 2014 was much lower than the temperature 305 (around 25 °C) at which rate constants of many reactions are usually measured. Rate 306 constants as a function of temperature are only available for the reactions of OH radical 307 with benzene and phenol among those shown in Figure 1 and they were already included 308 in the MCM v3.2. The model ran in a time-dependent mode and a 48-hour spin-up time 309 was applied in the box model. Measured concentrations of various hydrocarbons, NO_x, 310 O₃, NO₃ and photolysis frequencies (Table S2) were used as constraints in the box model. 311 The simulation period of the model was chosen to be January 18-27, a period associated 312 with several buildup episodes of ozone and other secondary products, with high measured 313 concentrations of NP and without precipitation. Following previous box model studies 314 (Yuan et al., 2015;Edwards et al., 2014), dilution and deposition processes were 315 represented together using a diurnally varying first-order physical loss parameter in the box model. The physical loss rate at night $(5.8 \times 10^{-6} \text{ s}^{-1})$ was calculated from the decrease 316 317 rate of NP concentration between 0:00-6:00 am when the chemical loss was expected to be low (see Section 3.3.4). A higher physical loss rate $(2.0 \times 10^{-5} \text{ s}^{-1})$ during daytime was 318 319 used to account for larger turbulent mixing during daytime (Edwards et al., 2014), which 320 results in the decrease of concentrations of inert tracers in the afternoon, e.g. benzene 321 (Figure 3) and methane. Based on sensitivity tests of the box model, increase and 322 decrease of the physical loss rate terms by a factor of two resulted in -48% and +39% of 323 changes in the modeled NP concentrations.

324 As shown in the introduction section, photolysis has been recognized as an 325 important sink for nitrated phenols. However, the photolysis of NP (and other nitrated 326 phenols) is not included in the MCM v3.2. We added the photolysis frequency of NP 327 from Bardini (2006) (1.4% of photolysis frequency of NO₂) into the MCM v3.2 and this 328 model run is referred to as the base simulation. Here, we assume that photolysis of NP 329 produces 2-phenoxy biradicals and HONO, as proposed in Bejan et al. (2006) (Figure 1 330 Route1). There are other possible chemical routes for photolysis of NP: producing 331 phenoxy radicals (C_6H_5O) by losing NO₂ (Route2 in Figure 1) and producing 332 nitrophenoxy radical by hydrogen abstraction (Route3 in Figure 1). The simulation test in 333 Figure S4 indicates that the pathway forming C₆H₅O radicals and NO₂ is an ineffective 334 sink for NP, since C₆H₅O radical will re-form NP by reacting with NO₂. However, we

cannot exclude this pathway to occur along with that producing 2-phenoxy biradicals and

HONO. The photolysis frequency determined in Bardini (2006) based on concentration

337 changes of 2-nitrophenol in a chamber may not include this pathway as well. As a result,

attributing the photolysis rates determined in Bardini (2006) to other pathways other than

- Route2 is reasonable. The route producing nitrophenoxy radical will be discussed in
- 340 Section 3.3.

The simulated results for phenol and NP from the base case of the box model are shown in Figure 5. The modeled diurnal variations agreed reasonably well with the observation for both NP and phenol in the base simulation, except for the phenol nighttime levels that will be discussed below. Although modeled NP concentrations are higher than the measurements for both daytime and nighttime, the agreement between measurements and model results is still within their combined uncertainties.

347 The average measured concentrations of phenol at night are higher than 10 ppt, but 348 the modeled phenol concentrations are usually less than 2 ppt. At night, the production of phenol from benzene oxidation halts, and the fast reaction with NO₃ $(2.8 \times 10^{-11} \text{ cm}^3)$ 349 molecule⁻¹ s⁻¹ at 298 K) removes phenol quickly (Figure 7 and discussion in Section 350 351 3.2.3). Measured nighttime NO₃ radicals were quite low during UBWOS 2014 (1.4 ± 2.4 352 ppt). As a check on the possible uncertainties in measurements of NO_3 at these low 353 levels, simulations by varying NO₃ concentrations by a factor of 2 result in little 354 improvement for the modeled concentrations of phenol (Figure S5). Another simulation 355 using calculated NO₃ concentrations from the equilibrium between NO₃ and N_2O_5 356 (Figure S5) also indicates that uncertainties in NO₃ measurements cannot account for the 357 discrepancies between measured and modeled phenol at night. The high phenol 358 concentrations measured at night might be a result of primary emissions. Indeed, the 359 measured phenol concentration was slightly enhanced in the plume with high methane 360 concentrations (see Figure 4A). However, a simulation using the measured phenol 361 concentrations as a constraint in the box model, predicted much higher NP concentrations 362 than measurements (Figure S4). Perhaps a more likely explanation for the enhanced 363 phenol concentrations at night is that the measurements of phenol by PTR-TOF suffer 364 from chemical interferences at night. Vinylfuran might be a candidate (Karl et al.,

365 2007;Stockwell et al., 2015). Thus, the modeled concentrations of phenol shown in366 Figure 5 will be used in the following discussions.

367 3.2.2 NO₂ dependence of NP yields

368 As shown in Figure 1, NP is generated from the reaction of NO_2 with phenoxy 369 radicals (C_6H_5O) (Berndt and Boge, 2003), which is an intermediate from the reactions 370 of OH and NO₃ radicals with phenol and the reaction of phenylperoxy radicals ($C_6H_5O_2$) 371 with NO. In addition to NO₂, C_6H_5O radicals also react with NO and O₃ (Platz et al., 372 1998) (see Figure 1). Thus, the yield of NP has been reported to depend on NO₂ 373 concentrations in the atmosphere (Berndt and Boge, 2003). 374 In the MCM v3.2, only the reactions of C_6H_5O radical with O_3 and NO_2 are included and here we added the reaction of C_6H_5O with NO ($k=1.88\times10^{-12}$ cm³ molecule⁻¹ 375 376 1 s⁻¹) for a new simulation. Compared to the base simulation, the modeled concentrations 377 of NP are lower, especially for the period of 11:00-17:00, as the effective yield of NP is 378 reduced (Figure 5). The small enhancement during the period of 11:00-17:00 in NP 379 concentrations from the base simulation is absent in the simulation with the reaction of 380 C_6H_5O with NO. The variations of modeled NP concentrations in the daytime from the 381 new simulation are in better agreement with the measurements (Figure 5). This indicates 382 the reaction of NO and C₆H₅O radical should be considered to account for the NO₂

383 dependence of NP formation.

384 Another simulation using fixed NP yields from phenol oxidation reported in 385 Atkinson et al. (1992) (6.7% for OH oxidation and 25.1% for NO₃ oxidation) is also 386 performed. This simulation neglects any dependence of NP yield from phenol oxidation 387 on concentrations of NO₂, O₃ and NO. We observed lower concentrations during both the 388 day and night compared to the base simulation (Figure 5). However, the enhancement of 389 modeled NP in the period of 11:00 -17:00 is distinctly observed with the simulation using 390 the fixed yields in Atkinson et al. (1992), which is in contrast to the lowest concentration 391 in the afternoon from our observations. This, again, indicates there must be a dependence 392 of NP yield from phenol oxidation on NO₂ concentrations in the atmosphere.

393 3.2.3 Gas-particle partitioning of NP

394 NP formed in the gas phase can partition into particles (Harrison et al., 2005a). 395 Measurements in several studies demonstrated that 2-NP and MNP were mainly found in 396 the gas phase (Herterich and Herrmann, 1990;Cecinato et al., 2005;Morville et al., 2006). 397 However, the reported particle fractions of 4-NP and DNP exhibit a broad range in 398 values: the particle fractions of 4-NP and DNP reported in Herterich and Herrmann 399 (1990) were both lower than 15%, whereas most of the concentrations of 4-NP (>75%) 400 (Cecinato et al., 2005) and DNP (>95%) (Morville et al., 2006) were detected in particles 401 in these two studies.

402 The concentrations of NP and other nitrated phenols in particles were not measured 403 in this study. We acknowledge that some fractions of nitrated phenols in particles may 404 evaporate into gas phase in the heated inlets. If it holds true, the measured concentrations 405 of nitrated phenols in this study would fall somewhere between concentrations in gas-406 phase and the total gas+particle concentrations. Here, the gas-particle partitioning of NP 407 as a function of time was estimated using the equilibrium absorption partitioning theory 408 (Pankow, 1994; Donahue et al., 2006) (see details in the SI), based upon pure-compound 409 liquid vapor pressures of 2-NP and 4-NP (Schwarzenbach et al., 1988) (Table 2) and 410 organic aerosol (OA) concentrations measured with an aerosol mass spectrometer 411 (AMS). The dependence with temperature was accounted for using the Clausius-412 Clapeyron relationship with reported enthalpies of evaporation (H_{vap}) (Schwarzenbach et 413 al., 1988). Although vapor pressures from Schwarzenbach et al. (1988) might have 414 significant uncertainties, Schwarzenbach et al. (1988) provided the only comprehensive 415 measurements of sub-cooled liquid vapor pressures of nitrated phenols in the literature. 416 The calculated fractions in particle (F_p) for 2-NP were generally very low (campaign average: $1.1\pm0.9\times10^{-4}$, max: 7.3×10^{-4}), whereas F_p for 4-NP were higher (average: 417

418 0.053 \pm 0.048, max: 0.38). The variability of the determined F_p values is the result of

419 variations of both OA concentrations ($12.5\pm8.7 \ \mu g/m^3$, min: $<1 \ \mu g/m^3$, max; $42.6 \ \mu g/m^3$)

420 and ambient temperature (-5±5 °C, min: -17 °C, max: 10 °C) during the campaign. The

421 higher F_p for 4-NP is expected, as 4-NP (1.4×10⁻³ Torr at 298 K) has much lower vapor

422 pressure than 2-NP (0.20 Torr at 298 K) (Table 2). In addition to absorption, partitioning

423 of NP into the aqueous phase of particles is another possible pathway affecting F_p . This

424 mechanism is estimated using Henry's law constants (Sander, 2015) and the liquid water

425 content (LWC) $(8.4\pm7.5 \,\mu\text{g/m}^3)$ in aerosol determined using the ISORROPIA model

426 (Fountoukis and Nenes, 2007). The estimated F_p values based on aqueous phase

427 partitioning for 2-NP (1.3×10^{-7}) and 4-NP (3.0×10^{-5}) are both much lower than F_p

428 estimated from the equilibrium absorption partitioning theory, indicating partitioning of

429 NP into the particle aqueous phase was not important during UBWOS 2014.

The gas-particle partitioning of 2-NP and 4-NP determined above was incorporated into the box model by constraining the estimated F_p in the determination of gas-particle mass transport rates. The mass transport rates (R_{in} and R_{out}) of a species into and out of particles with radius *r* are approximated by (Jacob, 2000):

$$R_{in} = \left(\frac{r}{D_g} + \frac{4}{v\alpha}\right)^{-1} A \times c_g = \frac{1}{\tau} \times c_g \tag{1}$$

$$R_{out} = \left(\frac{r}{D_g} + \frac{4}{\upsilon\alpha}\right)^{-1} A \times \frac{c_p}{K_{ep}} = \frac{1}{\tau} \times \frac{c_p}{K_{ep}}$$
(2)

434 where c_q and c_p are concentrations of the species in the gas and particle phase. D_q is the gas-phase molecular diffusion coefficient ($m^2 s^{-1}$), v is the mean molecular speed 435 (m s⁻¹), α is the mass accommodation coefficient. A is the aerosol surface area per unit 436 volume of air (m² m⁻³). K_{ep} is the equilibrium constant, i.e. c_p/c_g , or $/(1 - F_p)$. The 437 438 characteristic time scale of mass transfer (τ) is estimated to be on the order of minutes for 439 particles in the troposphere (Bowman et al., 1997; Jacob, 2000). Thus, rather than 440 determining the characteristic time scale explicitly, we assume that the equilibrium is 441 maintained at each model step (τ =5 min). After entering into particles, no further reaction 442 of NP was prescribed in the model. The modeled diurnal profiles of NP associated with 443 the inclusion of gas-particle partitioning are shown in Figure 6. Compared to the base 444 simulation, the modeled NP concentrations in the gas phase using the estimated F_p from 4-NP are lower (4-8%) for most of the day and slightly higher (2-3%) in the morning 445 when NP concentrations decreased quickly. Since the predicted F_p from 2-NP is very 446 447 small, the modeled NP concentrations in the gas phase using the estimates from 2-NP 448 were almost identical to the base simulation. In contrast with the modeled concentrations 449 of NP in the gas phase, the modeled total concentrations of NP in gas and particle phase 450 are consistently higher than the base simulation that does not consider gas-particle

451 partitioning. In summary, we observe relatively small changes of the modeled gaseous

452 NP concentrations after the inclusion of gas-particle partitioning in the box model.

453 Further measurements of the gas/particle partitioning of nitrated phenols are needed to

454 explain the variety of F_p values observed in different studies and/or the potential

differences between measurements and prediction from the equilibrium absorption

456 partitioning model.

457 3.2.4 Budget analysis of phenol and nitrophenol

458 Diurnal profiles of formation and loss rates derived from the base simulation of the 459 box model for both phenol and NP are shown in Figure 7. Production of phenol only 460 occurs in the daytime from OH oxidation of benzene. The magnitudes of losses of phenol 461 due to OH oxidation (21 ppt/day) and NO₃ oxidation (19 ppt/day) are comparable on a 462 daily basis. From morning to afternoon (8:00-15:00), production of phenol is larger than 463 the losses, resulting in continuous growth of phenol concentrations in this period. After 464 15:00 pm, the losses start to surpass the production and phenol concentrations decrease 465 quickly. With fast reduction of phenol concentrations in the evening, phenol loss from the 466 reaction with NO₃ mainly occurred before midnight.

467 As shown in Figure 7, NP is produced during both daytime and night, with more 468 production in the daytime. As mentioned earlier, the only formation pathway of NP is 469 from the reaction of phenoxy radicals (C_6H_5O) with NO₂, so the contribution from 470 different pathways to NP formation can be derived from source analysis of C₆H₅O 471 radicals (Figure 7C). The production of C_6H_5O radicals is dominated by the reaction of 472 $C_6H_5O_2$ radicals with NO during daytime (71% for 24-hour average) and the reaction of 473 phenol with NO₃ radicals at night (27% for 24-hour average). There are several sources 474 contributing to the formation of $C_6H_5O_2$ radicals in the MCM, including photolysis of 475 benzaldehyde and peroxybenzoic acid, OH oxidation of benzoic acid and phenyl 476 hydroperoxide, and degradations of other radicals (e.g., phenylperoxyacyl radical 477 $C_6H_5CO_3$), suggesting a wide range of aromatic compounds as the precursors of NP in 478 the daytime. The reaction of phenol with OH radicals only accounts for a small fraction 479 of the production of C₆H₅O radicals (2% for 24-hour average), due to the small yield of 480 NP (6%) from the reaction of OH with phenol in the MCM. This indicates phenol is not

an important precursor for NP during daytime. The destruction of NP is mainly due to 481 482 photolysis (17 ppt/day), with some contributions from NO₃ reaction (1.7 ppt/day). The 483 reaction with OH radicals is not important for the losses of NP in the UBWOS 2014. 484 Dilution/deposition processes account for 20% of the total loss of NP in the box model. 485 Our results are consistent with a previous proposal on photolysis as the dominant 486 chemical loss pathway for nitrated phenols (Bejan et al., 2007). Based on the conditions 487 at the Horse Pool site, the lifetime of NP due to photolysis at noontime is calculated to be 488 \sim 80 min. As the result of the short lifetime of NP during daytime, the production (23.6 489 ppt/day) and loss rates (23.1 ppt/day) of NP maintain a balanced budget of NP on a daily 490 basis. The inclusion of the reaction of phenoxy radicals (C₆H₅O) with NO discussed in 491 Section 3.2.2 would mainly affect the NP budget at midday, with smaller production and 492 loss in this period.

The different diurnal variations of production and loss rates of NP explain the measured diurnal profile of NP concentrations shown in Figure 3. The increase of loss rates from photolysis result in the quick decline of NP concentrations in the morning. The formation of NP from NO₃ oxidation of phenol in the evening exceeds the destruction of NP, which accounts for the enhancement of NP in this period. The formation and loss rates of NP are comparable in the afternoon and relatively constant concentrations of NP were observed.

500 A previous study showed that photolysis of nitrated phenols contributes to HONO 501 formation (Bejan et al., 2006). If we assume photolysis of nitrated phenols at rates of 502 1.4%×J(NO₂) yields HONO at a 100% yield (upper limit), photolysis of NP, MNP and 503 DMNP together accounted for a formation rate of HONO of 1.5 ± 1.9 ppt/hour around 504 noontime (9:00-15:00) during UBWOS 2014. This photolysis source would increase the 505 steady state concentrations of HONO by 0.5 ppt in early morning (7:00-8:30) and 0.2 ppt 506 during the noontime period (9:00-15:00), which are only small fractions of measured 507 HONO concentrations (50-100 ppt) during UBWOS 2014 (Edwards et al., 2014).

508 **3.3 Dinitrophenol**

Further oxidation of NP in the presence of NO_x produces DNP. The measured time
series of DNP in January 18-22 is shown in Figure 8. A similar diurnal profile was

observed for DNP as other nitrated phenols, with higher concentrations at night and lower
in the daytime. We also notice that the peak time of DNP concentrations at night was
somewhat later than NP, consistent with further oxidation of NP as the source of DNP.

514 In the MCM v3.2, reactions of NP with OH or NO_3 radicals generate nitrophenoxy 515 radicals (NO₂C₆H₅O·), which react further with NO₂ to form DNP. Here, we assume 516 DNP has the same photolysis rate as NP (1.4% of photolysis frequency of NO₂) (Bardini, 517 2006) and we added the photolysis into the MCM v3.2. The simulated concentrations of 518 DNP from the box model are also displayed in Figure 8. The agreement between 519 measurements and simulation is quite good from the base simulation. DNP has also been 520 observed in the particle phase at significant fractions (Morville et al., 2006). Using the 521 equilibrium absorption partitioning theory described in Section 3.3.3 and vapor pressures 522 of two different DNP isomers (2,4-DNP and 2,5-DNP) (Table 2), we incorporated the 523 calculated particle fractions of DNP (Table 2) into the box model as a sensitivity 524 simulation. The predicted DNP concentrations from this simulation are around 5% lower 525 than the base simulation at night. Considering the limited information on DNP formation, 526 the agreement between measured and modeled concentrations of DNP from both 527 simulations is encouraging. This degree of agreement implies that DNP concentrations 528 measured in UBWOS 2014 are explainable by known chemical reactions in the gas 529 phase.

530 As described in Section 3.2, photolysis is the dominant sink for NP and box 531 model results indicate that photolysis of NP may not generate phenoxy radical (by losing 532 NO₂). The other possible product from photolysis of NP is nitrophenoxy radical (Figure 1 533 Route 3), which would act as a secondary source of DNP. This assumption is evaluated 534 as a new simulation. The simulation predicted concentration peaks of DNP in the 535 morning (the orange line in Figure 8), which are not observed in our measurements. 536 Thus, we exclude nitrophenoxy radical as the main products of photolysis of NP. 537 However, the product and exact chemical mechanisms for photolysis of NP remain 538 unclear and thus the photolysis of NP warrants further detailed studies. 539 **3.4 Non-gas phase reactions**

540 The box model only considers gas phase reactions that produce nitrated phenols. In 541 addition to gas phase reactions, aqueous reactions in particles and heterogeneous 542 reactions are other potential sources of nitrated phenols (Harrison et al., 2005a). As 543 shown in Section 3.3.3, using chemical compositions of aerosol at the Horse Pool site and 544 the ISORROPIA model (Fountoukis and Nenes, 2007), the liquid water content (LWC) in aerosol during UBWOS 2014 was estimated to be $8.4\pm7.5 \text{ µg/m}^3$ (whole campaign 545 average), or $8.4\pm7.5\times10^{-12}$ expressed as the volume fraction. Based on the box modeling 546 547 results in Harrison et al. (2005b), aqueous reactions contribute less than 2% of NP production at 3×10^{-9} volume fraction LWC. Thus, aqueous reactions in UBWOS 2014 548 549 should not be a significant source for nitrated phenols compared to gas phase reactions.

550 The Uintah Basin was covered by snow during UBWOS 2014. The importance of 551 heterogeneous reactions on the snow surface to formation of nitrated phenols is evaluated 552 using measurements of the vertical gradients of these species. Here, the concentration 553 gradient is defined as the concentrations measured at 18.5 m subtracted from those 554 measured at 1 m. As shown in Figure 9, we observed negative concentration gradients for 555 nitrated phenols at night, indicating that deposition was playing a role and consequently 556 ground snow was a net sink for nitrated phenols at night. A previous study suggested 557 heterogeneous reaction of N_2O_5 with phenol in the aqueous phase produces NP (Heal et 558 al., 2007). Strong deposition of N_2O_5 to the snow surface was observed at night during 559 UBWOS 2014, but as discussed no production of nitrated phenols near the snow surface 560 was detected at night (Figure 9). The vertical gradients for nitrated phenols in the daytime 561 fluctuated around zero with large variations, which might be a result of their low 562 concentrations during daytime. The analysis of vertical gradients implies that 563 heterogeneous reactions on snow surface may not be important for formation of nitrated 564 phenols in the atmosphere during UBWOS 2014.

565 **4.** Conclusions

566 In this study, nitrated phenols in the gas phase were measured using an online 567 acetate ToF-CIMS in an oil and gas production region during winter. Strong diurnal 568 profiles were observed for nitrated phenols, with concentration maxima at night. As the 569 dominant sink for nitrated phenols, photolysis accounted for lower concentrations of

570 nitrated phenols during daytime. We determined that the photolysis of nitrated phenols 571 was not an important source of HONO during UBWOS 2014. Based on box model 572 results, NP was mainly formed in the daytime (73%) from a wide range of precursors, 573 with significant contribution from the reaction of phenol with NO₃ radicals at night 574 (27%). Box model results also indicated that gas phase oxidation of aromatics was able to 575 explain the measured concentrations of NP and DNP. We demonstrated that box model 576 results provided valuable information on the detailed chemical mechanisms in the 577 formation and destruction of NP, e.g., the NO₂ dependence of NP yields from phenol 578 oxidations and chemical products of NP photolysis. We determined that aqueous-phase 579 reactions and heterogeneous reactions were minor sources of nitrated phenols in this 580 study. Although the dataset of nitrated phenols was collected in an oil and gas production 581 region, the chemistry in secondary formation of nitrated phenols and the dynamics of the 582 budget of nitrated phenols in other regions, e.g., urban areas, should behave similarly to 583 those shown in this study.

584 Biomass burning activity did not affect the UBWOS 2014 measurements, and the 585 concentrations of phenols and nitrated phenols were mainly from oxidations of aromatics 586 in the atmosphere. The measurements during UBWOS 2014 provided a great opportunity 587 to study secondary formation of nitrated phenols in the absence of other confounding 588 sources. The UBWOS 2014 campaign also represented the first coincident and high time-589 resolution measurements of aromatic hydrocarbons, phenols and nitrated phenols in 590 ambient air. The measurements of phenol and nitrated phenols provided a better 591 understanding of their sources, budgets and roles in atmospheric chemistry and for the 592 evaluation of the oxidation mechanisms of aromatics. This is achieved by the emergence 593 of the new ToF-CIMS and PTR-TOF techniques. We envision that these techniques will 594 provide the ability to detect many other intermediate compounds in the atmosphere and 595 that the measurements will advance the understanding of atmospheric oxidation 596 processes.

597

598 Acknowledgement

599 The Uintah Basin Winter Ozone Studies were a joint project led and coordinated by the

- 600 Utah Department of Environmental Quality (UDEQ) and supported by the Uintah Impact
- 601 Mitigation Special Service District (UIMSSD), the Bureau of Land Management (BLM),
- 602 the Environmental Protection Agency (EPA) and Utah State University. This work was
- 603 funded in part by the Western Energy Alliance, and NOAA's Atmospheric Chemistry,
- 604 Climate and Carbon Cycle program. We thank Questar Energy Products for site
- 605 preparation and support. Chemical compositions of aerosol were provided by Tim Bates
- and James Johnson from NOAA Pacific Marine Environmental Laboratory (PMEL) and
- 607 the Joint Institute for the Study of the Atmosphere and Ocean (JISAO) at University of
- 608 Washington.
- 609

610 **References:**

- 611 Atkinson, R., Aschmann, S. M., and Arey, J.: Reactions of hydroxyl and nitrogen trioxide 612 radicals with phenol, cresols, and 2-nitrophenol at 296 ± 2 K. Environmental Science &
- 613 Technology, 26, 1397-1403, 10.1021/es00031a018, 1992.
- Bardini, P.: Atmospheric Chemistry of Dimethylphenols & Nitrophenols, PhD,
- 615 University College Cork, Corcaigh, 2006.
- 616 Bartmess, J. E.: Negative Ion Energetics Data, NIST Chemistry WebBook, NIST
- 617 Standard Reference Database Number 69, edited by: Linstrom, P. J., and Mallard, W.
- 618 G., National Institute of Standards and Technology, Gaithersburg MD, 20899
- 619 <u>http://webbook.nist.gov</u>, (retrieved August 28, 2015). 2015.
- Bejan, I., Abd El Aal, Y., Barnes, I., Benter, T., Bohn, B., Wiesen, P., and Kleffmann, J.:
 The photolysis of ortho-nitrophenols: a new gas phase source of HONO, Phys Chem
- 622 Chem Phys, 8, 2028-2035, 2006.
- Bejan, I., Barnes, I., Olariu, R., Zhou, S., Wiesen, P., and Benter, T.: Investigations on
- the gas-phase photolysis and OH radical kinetics of methyl-2-nitrophenols, Phys Chem
 Chem Phys, 9, 5686-5692, 10.1039/B709464G, 2007.
- 626 Bejan, I., Duncianu, M., Olariu, R., Barnes, I., Seakins, P. W., and Wiesen, P.: Kinetic
- 627 Study of the Gas-Phase Reactions of Chlorine Atoms with 2-Chlorophenol, 2-
- 628 Nitrophenol, and Four Methyl-2-nitrophenol Isomers, The Journal of Physical
- 629 Chemistry A, 119, 4735-4745, 10.1021/acs.jpca.5b02392, 2015.
- Berndt, T., and Boge, O.: Gas-phase reaction of OH radicals with phenol, Phys Chem
 Chem Phys, 5, 342-350, 10.1039/B208187C, 2003.
- 632 Bertram, T. H., Kimmel, J. R., Crisp, T. A., Ryder, O. S., Yatavelli, R. L. N., Thornton, J.
- A., Cubison, M. J., Gonin, M., and Worsnop, D. R.: A field-deployable, chemical
 ionization time-of-flight mass spectrometer, Atmos. Meas. Tech., 4, 1471-1479,
- 635 10.5194/amt-4-1471-2011, 2011.
- 636 Bolzacchini, E., Bruschi, M., Hjorth, J., Meinardi, S., Orlandi, M., Rindone, B., and
- Rosenbohm, E.: Gas-Phase Reaction of Phenol with NO3, Environmental Science &
 Technology, 35, 1791-1797, 10.1021/es001290m, 2001.
- Bowman, F. M., Odum, J. R., Seinfeld, J. H., and Pandis, S. N.: Mathematical model for
 gas-particle partitioning of secondary organic aerosols, Atmospheric Environment, 31,
 3921-3931, http://dx.doi.org/10.1016/S1352-2310(97)00245-8, 1997.
- Brophy, P., and Farmer, D. K.: A switchable reagent ion high resolution time-of-flight
 chemical ionization mass spectrometer for real-time measurement of gas phase
- 644 oxidized species: characterization from the 2013 southern oxidant and aerosol study,
- 645 Atmospheric Measurement Techniques, 8, 2945-2959, 10.5194/amt-8-2945-2015,
- 646 2015.
- 647 Cappellin, L., Karl, T., Probst, M., Ismailova, O., Winkler, P. M., Soukoulis, C., Aprea,
- 648 E., Mark, T. D., Gasperi, F., and Biasioli, F.: On quantitative determination of volatile 649 organic compound concentrations using proton transfer reaction time-of-flight mass
- 650 spectrometry, Environ Sci Technol, 46, 2283-2290, 10.1021/es203985t, 2012.
- 651 Caralp, F., Foucher, V., Lesclaux, R., J. Wallington, T., and D. Hurley, M.: Atmospheric
- 652 chemistry of benzaldehyde: UV absorption spectrum and reaction kinetics and
- mechanisms of the C6H5C(O)O2 radical, Phys Chem Chem Phys, 1, 3509-3517,
- 654 10.1039/A903088C, 1999.

- 655 Cecinato, A., Di Palo, V., Pomata, D., Tomasi Scianò, M. C., and Possanzini, M.:
- 656 Measurement of phase-distributed nitrophenols in Rome ambient air, Chemosphere, 59, 657 679-683, http://dx.doi.org/10.1016/j.chemosphere.2004.10.045, 2005.
- 658 Chen, J., Wenger, J. C., and Venables, D. S.: Near-Ultraviolet Absorption Cross Sections
- of Nitrophenols and Their Potential Influence on Tropospheric Oxidation Capacity,
 The Journal of Physical Chemistry A, 115, 12235-12242, 10.1021/jp206929r, 2011.
- 661 Cheng, S.-B., Zhou, C.-H., Yin, H.-M., Sun, J.-L., and Han, K.-L.: OH produced from o-
- 662 nitrophenol photolysis: A combined experimental and theoretical investigation, The
 663 Journal of Chemical Physics, 130, 234311,
- 664 doi:http://dx.doi.org/10.1063/1.3152635, 2009.
- 665 Corbin, J., Othman, A., D. Allan, J., R. Worsnop, D., D. Haskins, J., Sierau, B.,
- Lohmann, U., and A. Mensah, A.: Peak-fitting and integration imprecision in the
- 667 Aerodyne aerosol mass spectrometer: effects of mass accuracy on location-constrained
- 668 fits, Atmospheric Measurement Techniques, 8, 4615-4636, 10.5194/amt-8-4615-2015,
 669 2015.
- Cubison, M. J., and Jimenez, J. L.: Statistical precision of the intensities retrieved from
 constrained fitting of overlapping peaks in high-resolution mass spectra, Atmospheric
 Measurement Techniques, 8, 2333-2345, 10.5194/amt-8-2333-2015, 2015.
- 673 de Gouw, J., and Warneke, C.: Measurements of volatile organic compounds in the 674 earth's atmosphere using proton-transfer-reaction mass spectrometry, Mass
- 675 Spectrometry Reviews, 26, 223-257, 2007.
- Delhomme, O., Morville, S., and Millet, M.: Seasonal and diurnal variations of
 atmospheric concentrations of phenols and nitrophenols measured in the Strasbourg
- area, France, Atmospheric Pollution Research, 16-22, 10.5094/APR.2010.003, 2010.
- 679 Desyaterik, Y., Sun, Y., Shen, X., Lee, T., Wang, X., Wang, T., and Collett, J. L.:
- Speciation of "brown" carbon in cloud water impacted by agricultural biomass burning
 in eastern China, Journal of Geophysical Research: Atmospheres, 118, 7389-7399,
- 682 10.1002/jgrd.50561, 2013.
- Donahue, N. M., Robinson, A. L., Stanier, C. O., and Pandis, S. N.: Coupled Partitioning,
 Dilution, and Chemical Aging of Semivolatile Organics, Environmental Science &
 Technology, 40, 2635-2643, 10.1021/es052297c, 2006.
- 686 Dubé, W. P., Brown, S. S., Osthoff, H. D., Nunley, M. R., Ciciora, S. J., Paris, M. W.,
- 687 McLaughlin, R. J., and Ravishankara, A. R.: Aircraft instrument for simultaneous, in
- situ measurement of NO3 and N2O5 via pulsed cavity ring-down spectroscopy, Review
- 689 of Scientific Instruments, 77, 034101, doi:<u>http://dx.doi.org/10.1063/1.2176058</u>,
 690 2006.
- 691 Edwards, P. M., Brown, S. S., Roberts, J. M., Ahmadov, R., Banta, R. M., deGouw, J. A.,
- Dube, W. P., Field, R. A., Flynn, J. H., Gilman, J. B., Graus, M., Helmig, D., Koss, A.,
- 693 Langford, A. O., Lefer, B. L., Lerner, B. M., Li, R., Li, S. M., McKeen, S. A., Murphy,
- 694 S. M., Parrish, D. D., Senff, C. J., Soltis, J., Stutz, J., Sweeney, C., Thompson, C. R.,
- Trainer, M. K., Tsai, C., Veres, P. R., Washenfelder, R. A., Warneke, C., Wild, R. J.,
- 696 Young, C. J., Yuan, B., and Zamora, R.: High winter ozone pollution from carbonyl
- 697 photolysis in an oil and gas basin, Nature, 514, 351-354, 10.1038/nature13767, 2014.
- 698 Fernandez, P., Grifoll, M., Solanas, A. M., Bayona, J. M., and Albaiges, J.: Bioassay-
- directed chemical analysis of genotoxic components in coastal sediments,
- 700 Environmental Science & Technology, 26, 817-829, 10.1021/es00028a024, 1992.

- 701 Fountoukis, C., and Nenes, A.: ISORROPIA II: a computationally efficient
- 702 thermodynamic equilibrium model for K+ Ca2+ Mg2+ NH4+ Na+ SO42- NO3- Cl-
- 703 H2O aerosols, Atmos. Chem. Phys., 7, 4639-4659, 10.5194/acp-7-4639-2007, 2007.
- 704 Gilman, J. B., Lerner, B. M., Kuster, W. C., and de Gouw, J. A.: Source signature of
- 705 volatile organic compounds from oil and natural gas operations in northeastern 706 Colorado, Environ Sci Technol, 47, 1297-1305, 10.1021/es304119a, 2013.
- 707 Harrison, M. A. J., Barra, S., Borghesi, D., Vione, D., Arsene, C., and Iulian Olariu, R.:
- 708 Nitrated phenols in the atmosphere: a review, Atmospheric Environment, 39, 231-248, 709 http://dx.doi.org/10.1016/j.atmosenv.2004.09.044, 2005a.
- 710 Harrison, M. A. J., Heal, M. R., and Cape, J. N.: Evaluation of the pathways of 711 tropospheric nitrophenol formation from benzene and phenol using a multiphase
- 712 model, Atmos. Chem. Phys., 5, 1679-1695, 10.5194/acp-5-1679-2005, 2005b.
- 713 Heal, M. R., Harrison, M. A. J., and Neil Cape, J.: Aqueous-phase nitration of phenol by 714 N2O5 and CINO2, Atmospheric Environment, 41, 3515-3520,
- 715 http://dx.doi.org/10.1016/j.atmosenv.2007.02.003, 2007.
- 716 Herterich, R., and Herrmann, R.: Comparing the distribution of nitrated phenols in the 717 atmosphere of two German hill sites, Environ Technol, 11, 961-972,
- 718 10.1080/09593339009384948, 1990.
- 719 Inomata, S., Tanimoto, H., Fujitani, Y., Sekimoto, K., Sato, K., Fushimi, A., Yamada, H.,
- 720 Hori, S., Kumazawa, Y., Shimono, A., and Hikida, T.: On-line measurements of 721 gaseous nitro-organic compounds in diesel vehicle exhaust by proton-transfer-reaction
- 722 mass spectrometry, Atmospheric Environment, 73, 195-203, 723 http://dx.doi.org/10.1016/j.atmosenv.2013.03.035, 2013.
- 724 Jacob, D. J.: Heterogeneous chemistry and tropospheric ozone, Atmospheric
- 725 Environment, 34, 2131-2159, http://dx.doi.org/10.1016/S1352-2310(99)00462-8, 726 2000.
- 727 Jenkin, M. E., Saunders, S. M., Wagner, V., and Pilling, M. J.: Protocol for the 728 development of the Master Chemical Mechanism, MCM v3 (Part B): tropospheric 729 degradation of aromatic volatile organic compounds, Atmos. Chem. Phys., 3, 181-193, 730 10.5194/acp-3-181-2003, 2003.
- 731 Jenkin, M. E., Wyche, K. P., Evans, C. J., Carr, T., Monks, P. S., Alfarra, M. R., Barley, 732
- M. H., McFiggans, G. B., Young, J. C., and Rickard, A. R.: Development and chamber
- 733 evaluation of the MCM v3.2 degradation scheme for β -caryophyllene, Atmos. Chem.
- 734 Phys., 12, 5275-5308, 10.5194/acp-12-5275-2012, 2012.
- 735 Karl, T. G., Christian, T. J., Yokelson, R. J., Artaxo, P., Hao, W. M., and Guenther, A.:
- 736 The Tropical Forest and Fire Emissions Experiment: method evaluation of volatile 737 organic compound emissions measured by PTR-MS, FTIR, and GC from tropical
- 738 biomass burning, Atmospheric Chemistry and Physics, 7, 5883-5897, 2007.
- 739 Kaser, L., Karl, T., Guenther, A., Graus, M., Schnitzhofer, R., Turnipseed, A., Fischer,
- 740 L., Harley, P., Madronich, M., Gochis, D., Keutsch, F. N., and Hansel, A.: Undisturbed
- 741 and disturbed above canopy ponderosa pine emissions: PTR-TOF-MS measurements 742 and MEGAN 2.1 model results, Atmos. Chem. Phys., 13, 11935-11947, 2013.
- 743 Kitanovski, Z., Grgić, I., Yasmeen, F., Claeys, M., and Čusak, A.: Development of a
- 744 liquid chromatographic method based on ultraviolet-visible and electrospray ionization
- 745 mass spectrometric detection for the identification of nitrocatechols and related tracers

- in biomass burning atmospheric organic aerosol, Rapid Communications in Mass
- 747 Spectrometry, 26, 793-804, 10.1002/rcm.6170, 2012.
- Lauraguais, A., Coeur-Tourneur, C., Cassez, A., Deboudt, K., Fourmentin, M., and
 Choël, M.: Atmospheric reactivity of hydroxyl radicals with guaiacol (2-
- 750 methoxyphenol), a biomass burning emitted compound: Secondary organic aerosol
- formation and gas-phase oxidation products, Atmospheric Environment, 86, 155-163,
- 752 http://dx.doi.org/10.1016/j.atmosenv.2013.11.074, 2014.
- Lee, B. H., Lopez-Hilfiker, F., Mohr, C., Kurtén, T. C., Worsnop, D., and Thornton, J.
- A.: An Iodide-Adduct High-Resolution Time-of-Flight Chemical-Ionization Mass
 Spectrometer: Application to Atmospheric Inorganic and Organic Compounds,
- Figure 2017 Spectrometer. Application to Atmospheric morganic and Organic Compo 756 Environmental Science & Technology, 10.1021/es500362a, 2014.
- 757 Lin, P., Liu, J., Shilling, J. E., Kathmann, S., Laskin, J., and Laskin, A.: Molecular
- Characterization of Brown Carbon (BrC) Chromophores in Secondary Organic Aerosol
 Generated From Photo-Oxidation of Toluene, Phys Chem Chem Phys,
- 760 10.1039/C5CP02563J, 2015.
- Mohr, C., Lopez-Hilfiker, F. D., Zotter, P., Prévôt, A. S. H., Xu, L., Ng, N. L., Herndon,
 S. C., Williams, L. R., Franklin, J. P., Zahniser, M. S., Worsnop, D. R., Knighton, W.
- 763 B., Aiken, A. C., Gorkowski, K. J., Dubey, M. K., Allan, J. D., and Thornton, J. A.:
- 764 Contribution of Nitrated Phenols to Wood Burning Brown Carbon Light Absorption in
- 765 Detling, United Kingdom during Winter Time, Environmental Science & Technology,
- 766 47, 6316-6324, 10.1021/es400683v, 2013.
- Morville, S., Scheyer, A., Mirabel, P., and Millet, M.: Spatial and Geographical
 Variations of Urban, Suburban and Rural Atmospheric Concentrations of Phenols and
 Nitrophenols (7 pp), Environ Sci Pollut R, 13, 83-89, 10.1065/espr2005.06.264, 2006.
- 770 Müller, M., George, C., and D'Anna, B.: Enhanced spectral analysis of C-TOF Aerosol
- 771 Mass Spectrometer data: Iterative residual analysis and cumulative peak fitting,
- 772 International Journal of Mass Spectrometry, 306, 1-8,
- 773 <u>http://dx.doi.org/10.1016/j.ijms.2011.04.007</u>, 2011.
- Natangelo, M., Mangiapan, S., Bagnati, R., Benfenati, E., and Fanelli, R.: Increased
 concentrations of nitrophenols in leaves from a damaged forestal site, Chemosphere,
 20, 1405, 1502, http://dxia.org/10.1016/J00045_0525(00000270.1, 1000)
- 77638, 1495-1503, http://dx.doi.org/10.1016/S0045-6535(98)00370-1, 1999.
- Olariu, R. I., Klotz, B., Barnes, I., Becker, K. H., and Mocanu, R.: FT–IR study of the
 ring-retaining products from the reaction of OH radicals with phenol, o-, m-, and p-
- cresol, Atmospheric Environment, 36, 3685-3697,
- 780 <u>http://dx.doi.org/10.1016/S1352-2310(02)00202-9</u>, 2002.
- 781 Pankow, J. F.: An absorption model of the gas/aerosol partitioning involved in the
- formation of secondary organic aerosol, Atmospheric Environment, 28, 189-193, Doi:
 10.1016/1352-2310(94)90094-9, 1994.
- 784 Platz, J., Nielsen, O. J., Wallington, T. J., Ball, J. C., Hurley, M. D., Straccia, A. M.,
- 785 Schneider, W. F., and Sehested, J.: Atmospheric Chemistry of the Phenoxy Radical,
- C6H5O(•): UV Spectrum and Kinetics of Its Reaction with NO, NO2, and O2, The
 Journal of Physical Chemistry A, 102, 7964-7974, 10.1021/jp9822211, 1998.
- 788 Rippen, G., Zietz, E., Frank, R., Knacker, T., and Klöpffer, W.: Do airborne nitrophenols
- 789 contribute to forest decline?, Environmental Technology Letters, 8, 475-482,
- 790 10.1080/09593338709384508, 1987.

- Rubio, M. A., Lissi, E., Herrera, N., Pérez, V., and Fuentes, N.: Phenol and nitrophenols
 in the air and dew waters of Santiago de Chile, Chemosphere, 86, 1035-1039,
- 793 <u>http://dx.doi.org/10.1016/j.chemosphere.2011.11.046</u>, 2012.
- Sander, R.: Compilation of Henry's law constants (version 4.0) for water as solvent,
 Atmospheric Chemistry and Physics, 15, 4399-4981, 10.5194/acp-15-4399-2015, 2015.
- Schwarzenbach, R. P., Stierli, R., Folsom, B. R., and Zeyer, J.: Compound properties
- relevant for assessing the environmental partitioning of nitrophenols, Environmental
 Science & Technology, 22, 83-92, 10.1021/es00166a009, 1988.
- 799 Sekimoto, K., Inomata, S., Tanimoto, H., Fushimi, A., Fujitani, Y., Sato, K., and
- Yamada, H.: Characterization of nitromethane emission from automotive exhaust,
 Atmospheric Environment, 81, 523-531, 10.1016/j.atmosenv.2013.09.031, 2013.
- 802 Stark, H., Yatavelli, R. L. N., Kimmel, J. R., Bertram, T. H., Thornton, J. A., Jimenez, J.
- L., and Worsnop, D. R.: Cluster Formation and Ion Chemistry in the High Pressure
- 804 Inlet of a Chemical Ionization Mass Spectrometer: Lessons learned from Field and
- Laboratory Studies, 22nd International Symposium on Gas Kinetics, June 18-22, 2012,
 Boulder, CO, 2012.
- Stark, H., Yatavelli, R. L. N., Thompson, S. L., Kimmel, J. R., Cubison, M. J., Chhabra,
 P. S., Canagaratna, M. R., Jayne, J. T., Worsnop, D. R., and Jimenez, J. L.: Methods to
- 809 extract molecular and bulk chemical information from series of complex mass spectra
- 810 with limited mass resolution, International Journal of Mass Spectrometry,
- 811 10.1016/j.ijms.2015.08.011, 2015.
- Stockwell, C. E., Veres, P. R., Williams, J., and Yokelson, R. J.: Characterization of
 biomass burning emissions from cooking fires, peat, crop residue, and other fuels with
 high-resolution proton-transfer-reaction time-of-flight mass spectrometry, Atmos.
 Cham. Phys. 15, 845, 865, 10,5104/com 15, 845, 2015, 2015.
- 815 Chem. Phys., 15, 845-865, 10.5194/acp-15-845-2015, 2015.
- 816 Tremp, J., Mattrel, P., Fingler, S., and Giger, W.: Phenols and nitrophenols as
- tropospheric pollutants: Emissions from automobile exhausts and phase transfer in the
 atmosphere, Water Air Soil Pollut, 68, 113-123, 10.1007/BF00479396, 1993.
- 819 Veres, P., Roberts, J. M., Warneke, C., Welsh-Bon, D., Zahniser, M., Herndon, S., Fall,
 820 R., and de Gouw, J.: Development of negative-ion proton-transfer chemical-ionization
 821 mass spectrometry (NI-PT-CIMS) for the measurement of gas-phase organic acids in
- the atmosphere, International Journal of Mass Spectrometry, 274, 48-55, DOI
- 823 10.1016/j.ijms.2008.04.032, 2008.
- Vione, D., Maurino, V., Minero, C., and Pelizzetti, E.: Phenol photonitration upon UV
 irradiation of nitrite in aqueous solution I: Effects of oxygen and 2-propanol.
- 826 Chemosphere, 45, 893-902, <u>http://dx.doi.org/10.1016/S0045-6535(01)00035-2</u>,
 827 2001.
- 828 Vione, D., Maurino, V., Minero, C., and Pelizzetti, E.: Aqueous Atmospheric Chemistry:
- Formation of 2,4-Dinitrophenol upon Nitration of 2-Nitrophenol and 4-Nitrophenol in
 Solution, Environmental Science & Technology, 39, 7921-7931, 10.1021/es050824m,
- 831 2005.
- 832 Vione, D., Maurino, V., Minero, C., Duncianu, M., Olariu, R.-I., Arsene, C., Sarakha, M.,
- and Mailhot, G.: Assessing the transformation kinetics of 2- and 4-nitrophenol in the

atmospheric aqueous phase. Implications for the distribution of both nitroisomers in the atmosphere, Atmospheric Environment, 43, 2321-2327,

836 http://dx.doi.org/10.1016/j.atmosenv.2009.01.025, 2009.

- 837 Warneke, C., Geiger, F., Edwards, P. M., Dube, W., Pétron, G., Kofler, J., Zahn, A.,
- Brown, S. S., Graus, M., Gilman, J., Lerner, B., Peischl, J., Ryerson, T. B., de Gouw, J.
- A., and Roberts, J. M.: Volatile organic compound emissions from the oil and natural
- gas industry in the Uinta Basin, Utah: point sources compared to ambient air
- composition, Atmos. Chem. Phys., 14, 10977-10988, 10.5194/acpd-14-11895-2014,
- 842 2014.
- 843 Warneke, C., Veres, P. R., Murphy, S. M., Soltis, J., Field, R. A., Graus, M. G., Koss, A.,
- Li, S. M., Li, R., Yuan, B., Roberts, J. M., and de Gouw, J. A.: PTR-QMS vs. PTR-
- 845 TOF comparison in a region with oil and natural gas extraction industry in the Uintah
- Basin in 2013, Atmos. Meas. Tech., 8, 411-420, 10.5194/amtd-7-6565-2014, 2015.
 Wild, R. J., Edwards, P. M., Dube, W. P., Baumann, K., Edgerton, E. S., Quinn, P. K.,
- 848 Roberts, J. M., Rollins, A. W., Veres, P. R., Warneke, C., Williams, E. J., Yuan, B.,
- and Brown, S. S.: A Measurement of Total Reactive Nitrogen, NOy, together with
- NO2, NO, and O3 via Cavity Ring-down Spectroscopy, Environ Sci Technol, 48,
- 850 NO2, NO, and O3 via Cavity Ring-down Spectroscopy, Environ Sci Technol, 2 851 9609-9615, 10.1021/es501896w, 2014.
- 852 Yatavelli, R. L. N., Stark, H., Thompson, S. L., Kimmel, J. R., Cubison, M. J., Day, D.
- 853 A., Campuzano-Jost, P., Palm, B. B., Hodzic, A., Thornton, J. A., Jayne, J. T.,
- Worsnop, D. R., and Jimenez, J. L.: Semicontinuous measurements of gas-particle partitioning of organic acids in a ponderosa pine forest using a MOVI-HRToF-CIMS,
- 856 Atmos. Chem. Phys., 14, 1527-1546, 10.5194/acp-14-1527-2014, 2014.
- Yuan, B., Veres, P. R., Warneke, C., Roberts, J. M., Gilman, J. B., Koss, A., Edwards, P.
- 858 M., Graus, M., Kuster, W. C., Li, S. M., Wild, R. J., Brown, S. S., Dubé, W. P., Lerner,
- B. M., Williams, E. J., Johnson, J. E., Quinn, P. K., Bates, T. S., Lefer, B., Hayes, P. L.,
- Jimenez, J. L., Weber, R. J., Zamora, R., Ervens, B., Millet, D. B., Rappenglück, B.,
- and de Gouw, J. A.: Investigation of secondary formation of formic acid: urban
- 862 environment vs. oil and gas producing region, Atmos. Chem. Phys., 15, 1975-1993,
 863 10.5194/acp-15-1975-2015, 2015.
- Zhang, X., Lin, Y.-H., Surratt, J. D., and Weber, R. J.: Sources, Composition and
- 865 Absorption Ångström Exponent of Light-absorbing Organic Components in Aerosol
- Extracts from the Los Angeles Basin, Environmental Science & Technology, 47, 36853693, 10.1021/es305047b, 2013.
- 868



Figure 1. Formation of phenol, nitrophenol (NP) and dinitrophenol (DNP) from the
photo-oxidation of benzene in the atmosphere (Jenkin et al., 2003). Reactions in blue are

included in the MCM v3.2, whereas reactions in red are added or evaluated in this study.

- 875 For NP, DNP and the intermediate radicals, other isomers are expected but not shown for
- the sake of clarity.
- 877
- 878



879m/Q (Th)m/Q (Th)880Figure 2. High-resolution peak fitting to the averaged mass spectra of acetate ToF-CIMS881for m/z 138 (A), m/z 152 (B) and m/z 183 (C) on January 25, 2014 and m/z 183 (D) on882Jan. 18, 2014 during UBWOS 2014. The dark green lines indicate the calculated isotope883signals from lower masses.



885MST886Figure 3. (A) Diurnal profiles of measured NP, MNP, DMNP. (B) Diurnal profiles of887benzene, acetaldehyde, phenol, cresol and DMP. Photolysis frequencies of NO2 are

shown in both A and B for reference.



890 891 Figure 4. (A) An episode with high concentrations of methane and benzene on January

892 27, 2014 during UBWOS 2014. The source for this episode was fugitive emissions from

893 oil and gas activities. (B) An episode with high concentrations of NO_v and CO₂ on

894 February 3, 2014 during UBWOS 2014. The source for this episode was fuel combustion

895 (e.g., vehicle exhaust and/or other combustion sources for oil and gas extraction).



Figure 5. (A and C) Comparison of measured and modeled time series of phenol (A) and
NP (C). (B and D) Diurnal profiles of measured and modeled concentrations of phenol
(B) and NP (D). Photolysis frequencies of NO₂ are shown in B and D for reference. Error
bars in (B) and (D) indicate the accuracies of measured concentrations of phenol (50%)
and NP (40%), respectively.



903 MST
904 Figure 6. Diurnal profiles of measured and modeled concentrations of NP from the base

905 simulation and the simulations considering gas/particle partitioning. Photolysis

906 frequencies of NO_2 are shown for reference. Error bars indicate the accuracies of

907 measured concentrations of NP (40%).



909MSTMSTMST910Figure 7. Diurnal profiles of production and loss rates from different pathways for phenol

911 (A) and NP (B) derived from the base simulation of the box model. (C) Diurnal profiles
912 of production rates from different pathways for C₆H₅O radicals. The inserted pie in C

913 shows contributions from three different pathways to formation of C_6H_5O radicals on a

- 914 daily basis.
- 915



916 917 Figure 8. Comparison of measured and modeled time series of DNP. Measured time

918 series of NP is also shown for comparison. Error bars indicate the accuracies of measured

- 919 DNP concentrations (50%).
- 920
- 921



923

Figure 9. Diurnal profiles of vertical gradients for nitrated phenols measured in January

- 22 - February 1. The measured vertical gradient of N_2O_5 measured in February 6- 14 is
- also shown for comparison.

927 Tables

Species	Abbreviation	Ion	m/z	Sensitivity		Detection Limit, ppt ^f	
				Value,	Ratio to	Method	Method
				ncps/ppt	HCOOH ^e	1	2
Nitrophenol	NP	C ₆ H ₄ NO ₃ ⁻	138.0197	13.2 ^a	2.6	0.18	0.45
Methylnitrophenol	MNP	C ₇ H ₆ NO ₃ ⁻	152.0353	16.6 ^b	3.3	0.24	0.36
Dimethylnitrophenol + ethylnitrophenol	DMNP	C ₈ H ₈ NO ₃ ⁻	166.0510	16.6 ^c	3.3	0.14	0.36
Dinitrophenol	DNP	$C_6H_3N_2O_5$	183.0047	10.3 ^d	2.0	0.23	0.58

928 Table 1. Sensitivities and detection limits of nitrated phenols in acetate ToF-CIMS

929 a: Average from calibrations of 2-NP (8.4 ncps/ppt) and 4-NP (18.0 ncps/ppt).

- b: Calibration of 2-methyl-4-nitrophenol.
- 931 c: Using the same value as MNP.
- d: Calibration using 2,5-dinitrophenol.

e: Based on the determined sensitivity of formic acid (HCOOH) at 5.0 ncps/ppt during UBWOS

934 2014.

f: Method 1 is based on the random errors of observed counts follow Poisson distribution,

whereas method 2 is calculated as the concentrations with counts at three times of standard

937 deviations of measured background counts (see discussions in text and in Bertram et al. (2011)).

- 938
- 939
- 940
- 941 942
- 943

Table 2. Vapor pressure, enthalpy of evaporation (ΔH_{vap}) and calculated concentration

945

fractions in particle phase (F_p) for several nitrated phenols

Species	Vapor pressure at 298 K, Torr ^a	$\Delta H_{vap},\mathrm{kJ/mol}^{\mathrm{a}}$	F_p ,
2-NP	0.20	53.1	$1.1\pm0.9\times10^{-4}$
4-NP	1.4×10^{-3}	80.0	5.3±4.8×10 ⁻²
2,4-DNP	8.4×10^{-3}	70.4	$4.6 \pm 4.4 \times 10^{-3}$
2,5-DNP	1.2×10^{-3}	68.5	$2.8 \pm 2.7 \times 10^{-3}$

a: Calculated from data in Schwarzenbach et al. (1988).