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## Tropospheric mercury vertical profiles between 500 and 10000 m in central Europe

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### Abstract

Measurements of the vertical distribution of atmospheric mercury (Hg) are rare, because airborne measurements are expensive and labour intensive. Consequently, only a few vertical Hg profile measurements have been reported since the 1970s. Be-

- <sup>5</sup> sides the CARIBIC passenger aircraft observations, the latest vertical profile over Europe was measured in 1996. Within the Global Mercury Observation System (GMOS) project four vertical profiles were taken on board research aircraft (CASA-212) in August 2013 in background air over different locations in Slovenia and Germany. Each vertical profile consists of at least seven 5 min horizontal flight sections from 500 m
- <sup>10</sup> above ground to 3000 m a.s.l. Gaseous elemental mercury (GEM) was measured with a Tekran 2537X analyser and a Lumex RA-915-AM. Total gaseous mercury (TGM) was measured using a Tekran 2537B analyser and gaseous oxidized mercury (GOM) was sampled onto 8 denuders for post flight analysis (one for each profile, three during the transfer flights, and two blanks). In addition to the mercury measurements, SO<sub>2</sub>,
- <sup>15</sup> CO, O<sub>3</sub>, NO, NO<sub>2</sub>, as well as basic meteorological parameters (pressure, temperature, relative humidity) have been measured. Additional ground based speciated mercury measurements at the GMOS master site in Waldhof (Germany) were used to extend the profile to the ground.

No vertical gradient was found inside the well mixed boundary layer (variation by less than 0.1 ngm<sup>-3</sup>) at different sites with GEM varying from location to location between 1.4 and 1.6 ngm<sup>-3</sup> (STP; standard conditions: p = 1013.25 hPa, T = 273.15 K). At all locations GEM dropped to 1.3 ngm<sup>-3</sup> (STP) when entering the free troposphere and remained constant at higher altitudes. The combination of the vertical profile, measured on 21 August 2013, over Leipzig (Germany) with the CARIBIC measurements during ascent and descent to Frankfurt airport (Germany) at approximately the same time provide a unique central European vertical profile from inside the boundary layer (550 ma.s.l.) to the upper free troposphere (10 500 ma.s.l.) and shows a fairly constant free tropospheric TGM concentration of 1.3 ngm<sup>-3</sup> (STP). The highest GOM concen-



trations of up to  $60 \text{ pg m}^{-3}$  (STP, denuder samples) were found above the boundary layer during the transfer flights.

## 1 Introduction

Mercury and its compounds are very toxic and therefore hazardous for human health
 and the environment (Selin, 2009). Therefore it is on the priority list of many international agreements and conventions dealing with environmental protection and human health, including the United Nations Environment Program (UNEP) Minamata convention on mercury (www.mercuryconvention.org). Mercury is emitted to the atmosphere from a variety of anthropogenic (e.g. coal and oil combustion) and natural sources (e.g.
 evaporation from ocean and lakes) (Pirrone et al., 2010). The most efficient transport

- pathway for mercury is the atmosphere (Filzgerald et al., 1998). However, measurements of the vertical distribution of atmospheric mercury are rare, because airborne measurements are time consuming and expensive. Between 1978 and 2014 only seven campaigns performed airborne mercury measurements over Europe. Apart from the
- <sup>15</sup> CARIBIC dataset (Civil Aircraft for the Regular Investigation of the atmosphere Based on an Instrument Container, www.caribic-atmospheric.com), the last European vertical profile of mercury was measured in June 1996. Table 1 summarises all European airborne mercury measurements known to us together with their key findings (including this study).
- The GMOS 2012 measurement campaign at Mt. Etna (Global Mercury Observation System; www.gmos.eu; Weigelt et al., 2015b) focused on volcanic emissions and therefore no vertical profile was measured. CARIBIC measurements focus on the tropopause region and measures vertical profiles only above 6 km during ascent and descent from/to airports. During the four measurement campaigns over Europe be-
- tween 1978 and 1996 a vertical gradient was found neither in the planetary boundary layer (PBL) nor in the free troposphere. This was expected, because most of the atmospheric mercury is in its elemental state Hg(0) with a long atmospheric life time of six



months to one year (Lindberg et al., 2007). Due to the long lifetime, Hg is well mixed in the atmosphere. All known vertical profile measurements of Hg were summarized by Swartzendruber et al. (2009) (data are shown in Fig. 7 for comparison to this study). Hg vertical profiles were measured by Radke et al. (2007), Talbot et al. (2008), and
<sup>5</sup> Swartzendruber et al. (2006, 2008) in different locations over the Pacific Ocean and

- the US between 2002 and 2008. Vertical profiles over Canada were reported by Banic et al. (2003) for the period between 1995 and 1998. Friedli et al. (2004) report vertical profiles measured over Japan/Korea and China in Spring 2001. In the Swartzendruber et al. (2009) summary, a paper by Ebinghaus and Slemr (2000) represents the only
- <sup>10</sup> European vertical profile. Recently, Brooks et al. (2014) reported speciated mercury vertical profiles measured over USA over a period of almost one year from August 2012 to June 2013.

Except for large vertical GEM gradients reported by Radke et al. (2007), no pronounced GEM vertical gradients were observed by other researchers (Swartzendruber

- et al., 2009; Brooks et al., 1014). Usually the GEM concentrations in the planetary boundary layer (PBL; 0–1 to 3 km) were found to be the same as in the lower free troposphere (FT). As mercury is emitted from the underlying surface, we would expect at least a slightly higher concentration inside the PBL compared to the FT. The absence of a vertical gradient inside the PBL and the FT is caused by the "fast" mixing velocity
- of Hg (hours to days), compared to the atmospheric life time (6 to 12 month) and the insufficient precision of the available mercury analysers to detect concentration gradients of less than  $0.1 \text{ ng m}^{-3}$ .

The European Tropospheric Mercury Experiment (ETMEP) was carried out in July/August 2012 (ETMEP-1) and August 2013 (ETMEP-2) to measure local emis-

sions and to perform vertical profile measurements from inside the boundary layer to the lower free troposphere. In total 10 measurement flights were performed over Italy, Slovenia, and Germany with two small, flexible aircraft. The ETMEP-1 campaign focused on volcanic emissions as such and not on the investigation of vertical profiles.



We report here the results of the ETMEP-2 campaign, which focused on vertical profile measurements over central Europe.

## 2 Measurement location and methodology

From 19 to 22 August 2013, five ETMEP-2 measurement flights were carried out over
<sup>5</sup> central Europe (Fig. 1). After take-off on 19 August at the aircraft's home base in Parma (northern Italy) the first vertical profile was measured in the early afternoon over the GMOS Master site "Iskraba" in Slovenia. Thereafter the second vertical profile was flown over Idrija (Slovenia), a former mercury mining area. On 21 August, in the morning the transfer flight from Ronchi dei Legionari (north-east Italy) to Leipzig (central Germany) was used as the second measurement flight to obtain a central European horizontal profile inside or slightly above the boundary layer (flight 2). During this flight no vertical profile was flown. After refuelling at Leipzig airport, the third flight was carried out on the same day. Within this flight, two vertical profiles were flown; the first one at noon downwind of a coal-fired power plant south of Leipzig (Lippendorf) and

- the second one at early afternoon over the Leipzig city-centre. With the fourth measurement flight on 22 August (take-off in Leipzig), the fifth vertical profile was flown in the forenoon over the GMOS master site "Waldhof" (northern Germany), representing central European rural background air. Thereafter, the aircraft was refuelled at Leipzig airport and flown back to Parma on the same day. This last transfer flight (flight 5) was used to obtain a second central European horizontal profile slightly above the bound-
- ary layer. Here we present and discuss the vertical profiles over Iskraba, Idrija, Leipzig, and Waldhof. The Lippendorf vertical profile downwind the coal fired power plant will be discussed in a separate paper (Weigelt et al., 2015c).

Each vertical profile consists of at least seven horizontal flight legs, lasting five minutes each. The altitude for the flight legs was chosen, starting inside the boundary layer at about 400 m above ground. For each vertical profile the highest flight level was 3000 m a.s.l. Each flight-level-change was performed within 2.5 min. Consequently,



each vertical profile took 50 min. The measurement campaign described above was performed with a CASA 212 two engine turboprop aircraft (Fig. 2a). This aircraft is operated by Compagnia Generale Ripreseaeree (http://www.terraitaly.it/). The CASA 212 has a maximum payload of 2.7 tons, allowing to carry the measurement instru-

- <sup>5</sup> ments, different service instruments, the power supply, two pilots, and 5 operators. The aircraft normal cruising speed is 140 kn ( $\sim 260 \text{ km h}^{-1}$ ). At this speed the maximum flight distance is  $\sim 1600 \text{ km}$ . The maximum flight level of the unpressurized aircraft is 8500 m. As it was not possible to fly with oxygen masks, the maximum flight level for the ETMEP-2 campaign was limited to 10 000 ft ( $\sim 3000 \text{ ma.s.l.}$ ).
- Previously, the CASA 212 was used as a research aircraft to carry remote sensing LIDAR systems (light detection and ranging), but not for in situ measurements. Therefore, the aircraft had no gas inlet. To transfer unbiased ambient air from outside the aircraft boundary layer to the measurement instruments, a gas inlet system has been developed and manufactured at the Helmholtz-Zentrum Geesthacht (Fig. 2b). The gas
- <sup>15</sup> inlet was specially designed for the cruising speed of the CASA 212. The air enters the inlet with a speed of about  $260 \text{ km h}^{-1}$  (~  $72 \text{ m s}^{-1}$ ). By expansion, the air velocity is reduced to about  $15 \text{ km h}^{-1}$  (~  $5 \text{ m s}^{-1}$ ). At  $260 \text{ km h}^{-1}$  about  $120 \text{ Lmin}^{-1}$  (ambient conditions) enters the inlet. In the centre of the expansion area the main sampling line starts. All instruments pull their measurement air from this main sampling line (all
- <sup>20</sup> together about  $25 \text{ Lmin}^{-1}$ ). The remaining  $95 \text{ Lmin}^{-1}$  are directed to the back of the inlet where the air speed is increased by a nozzle and the air exits. By replacing the inlet- and outlet nozzle with smaller or larger ones, this inlet system is also suited for other aircraft types with different cruising speed. In the expanded area (behind the main sample line) the air temperature (*T*), static pressure (*p*), and relative humidity
- (RH) are measured. To optimize for trace gas measurements and to avoid contamination, the whole inside of the inlet was coated with Teflon and only Teflon tubes (PFA) were used for the sampling line. The outside of the inlet was copper coated to avoid electrostatic charging. The inlet body was mounted onto a 6 cm wide and 90 cm long telescope tube. This telescope tube was flexibly mounted into the aircraft fuselage. Af-



ter take-off, the telescope tube was pushed down by  $\sim 40$  cm from inside the aircraft, to ensure the inlet nozzle is outside the aircraft boundary layer. Before landing the telescope tube was pulled back into the aircraft fuselage. Inlet and telescope tube were equipped with controllable heaters to prevent icing. However, because the measure-

- ment flights were carried out in summer at altitudes below 3000 ma.s.l., it was never necessary to switch on the heating system. Inside the cabin the tubing from telescope tube to instruments (~ 2.5 m long 3/8" main sample tube with PFA manifolds to instruments) was not heated. The temperature inside the cabin was 18 to 30 °C. Aerosol particles were filtered out at the instrument individual inlets by using PTFE membrane
   filter (pore size 0.2 μm). All data were corrected for individual instrument response time
  - due to sampling tube length and instrument internal analysis.

For the campaign the aircraft was equipped with three mercury measurement instruments, one Lumex RA-915AM, a Tekran 2537B, and a Tekran 2537X (cf. Table 2). The high resolution Lumex RA-915 AM measures gaseous elemental mercury (GEM)

- with a raw signal temporal resolution of only 1 s. The measurement principle is based on atomic absorption spectroscopy (AAS) with Zeeman background correction. Due to the limited sensitivity however, the raw signal is noisy (about  $\pm 4 \text{ ngm}^{-3}$ ) and is dependent on pressure and temperature. As the aircraft cruising speed is about 72 m s<sup>-1</sup>, this highly resolved raw signal is very useful to detect small scale highly concentrated
- <sup>20</sup> mercury plumes (concentration change > noise level). Therefore, Lumex data were used for data analysis but are not sown in the vertical profile plots. The Tekran 2537B and 2537X analysers are based on cold vapour atomic fluorescence spectroscopy (CVAFS) and can measure total gaseous mercury (TGM). Because the CVAFS needs pre-concentrated samples, the Tekran analysers pre-amalgamate Hg from the sam-
- <sup>25</sup> ple air on solid gold cartridges and achieve a minimum temporal resolution of 150 s. For the ETMEP-2 flights a quartz wool trap was installed upstream the Tekran 2537X analyser, removing only gaseous oxidized mercury (GOM) and aerosol particles with particle bound mercury (PBM) but no GEM from the air stream (cf. Lyman and Jaffe, 2011).



The Tekran 2537B analyser was operated as backup instrument without a quartz wool trap. Although the Teflon made (PFA and PTFE) aircraft gas inlet system was not tested for GOM transmission efficiency, we expect nearly quantitative GOM transmission from the inlet to the instrument. Therefore, the Tekran 2537B measurement are believed to represent total gaseous mercury (TGM = GEM + GOM) concentrations. To estimate the concentration of GOM, additionally 8 manual denuder samples were taken (sampling time 1 h or longer, sampling flow controlled using a mass flow controller). In parallel to the denuder samples two blank tests were performed by handling the denuders exactly the same way the samples were handled (denuder preparation, installation

- <sup>10</sup> to sampling setup, storage, analysis), but without pulling sample air thru. After all flights had been finished, the denuders were analysed for their total GOM loadings in the laboratory. This method has a relatively high uncertainty of about  $\pm 5 \text{ pg m}^{-3}$ . Nevertheless, additional information on the amount of GOM is obtained by that approach. During each vertical profile one denuder was used, whereas one denuder was loaded during <sup>15</sup> the transfer flight Ronchi dei Legionari – Leipzig and two denuders were used along the
- transfer flight Leipzig Parma. The blank tests were performed during the first vertical profile over Iskraba and the last vertical profile over Waldhof (see flight information and Fig. 1).

For the identification and characterization of different air masses carbon monoxide (CO), ozone (O<sub>3</sub>), sulphur dioxide (SO<sub>2</sub>), nitric oxide (NO), nitric dioxide (NO<sub>2</sub>), and the basic meteorological parameters temperature (*T*), pressure (*p*), and relative humidity (RH) were measured simultaneously at high temporal resolution. Instrument details are summarised in Table 2. CO and SO<sub>2</sub> can be used for the identification of city plumes and plumes of power stations, respectively. O<sub>3</sub> can be used to characterize upper tropospheric/lower stratospheric air or to explain oxidation processes. A ratio of NO / NO<sub>2</sub> provides information about the age of polluted air masses. Usually FT air is much dryer than PBL air and, therefore, the RH measurements can distinguish these two air masses. Model meteorological data like potential vorticity, equivalent potential temperature, relative- and specific humidity, cloud cover, cloud water content, 3 dimen-



sional wind vector, as well as five day backward trajectories were calculated every 150 s along the aircraft flight tracks for additional information. These calculations are based on meteorological analysis data from the European Centre for Medium-Range Weather Forecasts (ECMWF) and the TRAJKS trajectory model (Scheele et al., 1996).

- <sup>5</sup> Before take-off all instruments were warmed up for at least 45 min, using an external ground power supply. During the starting of the engines the power was interrupted for less than 3 min. Since 45 min were too short to stabilize the Tekran 2537 internal permeation source, these instruments were calibrated directly after each measurement flight before the engine shut down. All data were recalculated, using the post flight
- calibration. The pressure in the fluorescent cells of both Tekran instruments was kept constant using upstream pressure controllers at the exits of the cells. This eliminated the known pressure dependence of the response signal (Ebinghaus and Slemr, 2000; Talbot et al., 2007). The Lumex analyser has a much shorter warm up time of less than 10 min and was, therefore, calibrated before take-off. The CO instrument calibration
- <sup>15</sup> takes 60 s and was, therefore, performed during the measurement flights every 20 min. The O<sub>3</sub>, SO<sub>2</sub>, NO, NO<sub>2</sub> instruments have a fairly constant signal response and were thus calibrated before and after the ETMEP-2 measurement campaign with external calibration gases. The factory calibration was used for the pressure, temperature and relative humidity sensors. The measurements were synchronized using their individual <sup>20</sup> response times. Please note that all mercury (TGM, GEM, and GOM) concentrations are reported at standard conditions (p = 1013.25 hPa, T = 273.15 K). At these standard
- are reported at standard conditions (p = 1013.25 hPa, T = 273.15 K). At these standard conditions 1 ng m<sup>-3</sup> corresponds to a mixing ratio of 112 ppqv (parts per quadrillion by volume).

#### 3 Results

The first vertical profile was measured on 19 August 2013 from 11:15 to 12:15 UTC over the GMOS Master site "Iskraba" (Fig. 1). As Iskraba is located in mountainous terrain, the lowermost flight leg was performed at 1000 m a.s.l. The measurements are



summarised in Fig. 3. The squares represent the constant flight level measurement points (5 min each). The stars represent the measurements while climbing between two flight levels (2.5 min average). The data, represented by squares are thus more significant and the data illustrated as stars do provide additional information of the ver-

- <sup>5</sup> tical structure. Please note that the RH and the air temperature (T) are plotted with high temporal resolution (1 s) in the right most panel. RH increases with increasing altitude and shows no step change to lower RH which would identify the top of PBL. Hence, the whole profile in Fig. 3 was flown within the PBL. The measurements indicate a very constant mercury concentration without any vertical gradient for TGM and GEM. With
- 1.44 ng m<sup>-3</sup> the whole column average TGM concentration was somewhat below the northern hemispheric background concentrations of 1.5–1.7 ng m<sup>-3</sup> (Lindberg et al., 2007) but was comparable with the August 2013 monthly median of 1.41 ng m<sup>-3</sup> at Mace Head/Ireland (Weigelt et al., 2015a) and a median concentration of 1.40 ng m<sup>-3</sup> all vertical profiles over Tennessee, USA, in 2012–2013 (Brooks et al., 2014). With
- <sup>15</sup> 1.38 ng m<sup>-3</sup> the column averaged GEM concentration was only slightly lower than TGM but this difference is smaller than the combined uncertainties of both instruments and thus insignificant. The manual denuder sampling (integral value, representative for the whole boundary layer over Iskraba) gave a GOM concentration between 1.9 and 13.2 pg m<sup>-3</sup>, which is smaller than the GOM concentration which can be resolved by
- the TGM GEM differential measurement. This was the second lowest GOM concentration measured on all ETMEP-2 flights with the manual denuder sampling technique (Table 3). No ground-based reference data for the GMOS Iskraba site were available due to technical reasons. Furthermore, Lumex instrument data were not available due the instrument failure. Besides mercury, neither CO, nor O<sub>3</sub>, NO, and NO<sub>2</sub> mixing ratios
- <sup>25</sup> indicate a significant vertical gradient. Only the SO<sub>2</sub> mixing ratio increased from 1000 to 1500 m a.s.l. and remained constant there above. In general the measurements thus showed that the air over Iskraba was well mixed within the PBL.

After the experimental flights over Iskraba were completed, the second vertical profile was flown on the same day about 80 km northwest over the former mercury mining



area "Idrija". Until the 1990s, Idrija was the second largest mercury mine in operation worldwide (Grönlund et al., 2005). This profile was measured between 12:25 and 13:25 UTC (Fig. 4). Due to the mountainous terrain the seven horizontal fight legs were performed within the altitude range 1350 m to 3150 m a.s.l. On the contrary to Iskraba,

<sup>5</sup> the uppermost flight leg over Idrija was flown above the PBL, in FT air. This is clearly indicated by a significantly reduced RH (the right most panel in Fig. 4). The boundary layer top was found at 2850 to 2900 m a.s.l.

Compared to Iskraba, the mercury concentration over Idria was with 1.5 to 1.6 ng m<sup>-3</sup> (GEM) and 1.6 to 1.7 ng m<sup>-3</sup> (TGM) about 10 to 15% higher. With 18.0 to 28.8 pg m<sup>-3</sup>,

- <sup>10</sup> also GOM (manual denuder sampling) was found to be significantly higher than over Iskraba. The elevated mercury concentrations might be caused by increased emission from the soil around Idrija due to the former mining activity. However, as over Iskraba no vertical GEM or TGM concentration gradient was observed inside the PBL. It should be noted that above the PBL the GEM and TGM concentrations were found to be <sup>15</sup> significantly lower (GEM: 1.23 ngm<sup>-3</sup>; TGM: 1.32 ngm<sup>-3</sup>). Ozone, CO, NO<sub>2</sub>, and SO<sub>2</sub>
- <sup>15</sup> significantly lower (GEW. 1.23 fight ), read. 1.32 fight ). O2016, CO,  $NO_2$ , and  $SO_2$ mixing ratios behave similarly, although  $NO_2$  and  $SO_2$  show a small gradient within the PBL with slightly decreasing mixing ratios with increasing altitude. At 2700 m a.s.l. near the top of the PBL, all trace gas mixing ratios start to decrease and the mixing ratios at 3150 m a.s.l. in the FT are the lowest of the whole profile.  $O_3$  and CO mixing
- ratios decrease by about 20 % when entering FT, NO<sub>2</sub> by about 60 % and SO<sub>2</sub> drops essentially to the detection limit. This step in mixing ratio at the PBL top indicates that FT air is separated from the PBL air due to slow air mass exchange. Nitrogen oxide (NO) shows no vertical gradient from inside the PBL to the FT. It should be noted the NO mixing ratios are close to the instrument's detection limit and might be not representative or have at least a large uncertainty. The stars at 2900 ma.s.l. represent

a mixture of the PBL and FT air, explaining the concentrations are between the PBL and FT air concentration (e.g. GEM 1.3 ngm<sup>-3</sup> and TGM 1.4 ngm<sup>-3</sup>).

On 21 August 2013, two vertical profiles were measured over central Germany in the area of the city of Leipzig (Fig. 1). The first profiling was carried out downwind of



a coal-fired power plant and is the subject of another paper (Weigelt et al., 2015c). Thereafter, the second profile was flown between 11:10 UTC and 12:10 UTC over the city centre of Leipzig (population 500 000). The Leipzig profile was flown upwind of the power plant and was taken as a reference for the profile downwind of the power plant measurements. The profile is shown in Fig. 5. The lowermost flight level over Leipzig was 450 m above ground (600 m a.s.l.) and the highest one was 3020 m a.s.l.

From 21 to 23 August 2013, additionally four CARIBIC measurement flights were performed aboard a passenger aircraft (Lufthansa airbus A340-600) from Frank-furt/Germany to Caracas/Venezuela and Vancouver/Canada and back. Among other

- instruments (Brenninkmeijer et al., 2007), the CARIBIC system carries a Tekran 2537A mercury analyser, measuring TGM along the flight track with a temporal resolution of 600 s (Ebinghaus et al., 2007; Slemr et al., 2014). On 21 to 23 August 2013, a high pressure system dominated the weather over Germany and Western Europe when the ETMEP-2 and the CARIBIC measurements were carried out. The wind direction in the
- free troposphere (3–10 km) was west to northwest and the forward- and backward trajectory analysis showed that both the ETMEP-2 and CARIBIC aircraft sampled about the same air mass (see Fig. S1 in the Supplement). This allows supplementing and comparing the ETMEP-2 Leipzig vertical profile with the independent CARIBIC measurements during ascent/descent from/to Frankfurt airport, only some 350 km apart.
- <sup>20</sup> For this extension only free tropospheric CARIBIC measurements east of 0° E are additionally plotted in Fig. 5, providing a vertical profile extending from 600 to 10500 m a.s.l. Stratospheric CARIBIC measurements (with  $O_3 > 80$  ppb) are not shown.

The ETMEP-2 measured RH vertical profile identified the PBL top over the city centre of Leipzig at 2200 to 2250 ma.s.l. While the first five ETMEP-2 horizontal flight legs were flown inside the PBL, the last two legs were performed in FT air. Again, inside and above the PBL no vertical gradient was apparent for GEM, TGM, O<sub>3</sub>, CO, NO, and SO<sub>2</sub>, indicating well mixed air masses. Only for NO<sub>2</sub> a negative vertical gradient was found inside the PBL, but not above. Inside the PBL the average GEM and TGM concentration was 1.50 and 1.55 ngm<sup>-3</sup>, which is in between the concentrations found



inside the PBL over Iskraba and Idrija. The FT GEM and TGM concentration over Leipzig was measured to be 1.2 to  $1.3 \text{ ng m}^{-3}$ . Similar concentrations were also found in the FT air over Idrija (Fig. 4), Waldhof (Fig. 6, flight leg five and six), as well as during the transfer flights Ronchi dei Legionari – Leipzig and Leipzig – Parma (not shown).

<sup>5</sup> The GOM concentration from denuder samples along the Leipzig profile was 1 pg m<sup>-3</sup> (lower detection limit) to 10.6 pg m<sup>-3</sup>, representing the lowest measured concentration along all flights (cf. Table 3).

The CARIBIC and ETMEP-2 FT data match very well. The average TGM concentration is 1.23 ng m<sup>-3</sup> for the ETMEP-2 and 1.30 ng m<sup>-3</sup> for CARIBIC dataset. This means
 that no vertical GEM gradient is apparent in the entire FT over Central Europe. Inside the PBL the GEM and TGM concentration is about 20% higher. Furthermore trace gases CO, O<sub>3</sub>, and NO measured aboard CARIBIC match the ETMEP-2 measurements very well, supporting the notion that the same FT air mass was sampled during the CARIBIC and ETMEP-2 flights. Consequently, the combined ETMEP-2 and CARIBIC data set provides to the best of our knowledge the first complete vertical mercury profiles from inside the PBL to the upper FT.

The last vertical profile was flown on 22 August 2013, over the GMOS master site "Waldhof" (Fig. 6). Since this profile was measured in the forenoon (08:15 to 09:15 UTC; 10:15 to 11:15 local time), the PBL was with the top at 1750–1850 m a.s.l.

- rather shallow when compared to the previous profiles. Thus only the first four flight legs were flown inside the PBL and the remaining three were above. As measured during all previous vertical profiles, again a significant difference between PBL and FT air was apparent for GEM and TGM concentrations, and CO, NO, and SO<sub>2</sub> mixing ratios. The two lower FT flight legs indicated typical GEM and TGM concentrations of 1.27 and
- <sup>25</sup> 1.19 ng m<sup>-3</sup> (1950 m a.s.l.) and 1.22 and 1.22 ng m<sup>-3</sup> (2490 m a.s.l.), respectively. However, in the uppermost flight level at 3030 m a.s.l. GEM and TGM concentrations were 0.99 and 0.98 ng m<sup>-3</sup>, respectively, i.e. about 25 % lower. Furthermore, in that layer not only the GEM and TGM concentrations, but also the CO and O<sub>3</sub> mixing ratios were about ~ 25 % lower. At the same time RH was with 66.6 % substantially higher and SO<sub>2</sub>



with 1.1 ppb slightly higher. Five day backward trajectories (Fig. S2 in the Supplement) suggest that the air from this uppermost flight leg originated from the subtropical east Atlantic (about 30° N, 25° W). On the contrary, the air measured during all lower flight legs (in PBL and FT air) came from north Canada (north of 60° N, west of 50° W).

- Inside the PBL the GEM and TGM concentrations were with 1.93 and 1.95  $\text{ngm}^{-3}$ , respectively, the highest in the uppermost flight leg (1470 m a.s.l.). Similarly, the CO mixing ratio was also elevated and the SO<sub>2</sub> raw signal indicated some short peaks to 1.5 ppb (not shown). The coincidence of elevated GEM and TGM concentrations with elevated CO and SO<sub>2</sub> mixing ratios was probably caused by a combustion plume.
- <sup>10</sup> Below this plume again a fairly constant profile was measured for GEM (1.66 ng m<sup>-3</sup>), TGM (1.73 ng m<sup>-3</sup>), CO (121.4 ppb), O<sub>3</sub> (52.4 ppb), and NO (at detection limit). Only NO<sub>2</sub> and SO<sub>2</sub> mixing ratios increased towards the ground from 1.1 and 1.1 ppb, at 962 m a.s.l. to 1.7 and 1.6 ppb at 429 m a.s.l., respectively.
- GEM concentration measured by a Tekran speciation unit at the ground at the Wald-<sup>15</sup> hof site was with 1.92 ng m<sup>-3</sup> somewhat elevated. Concurrently measured concentrations of GOM (3.6 pg m<sup>-3</sup>) and particle bound mercury (PBM, 7.8 pg m<sup>-3</sup>) were somewhat elevated too. The Waldhof three-year-average (2009–2011) is 1.0 pg m<sup>-3</sup> for GOM and 6.3 pg m<sup>-3</sup> for PBM (Weigelt et al., 2013). With 2.0 ppb the ground based NO<sub>2</sub> mixing ratio follows the increasing gradient toward the ground. On the contrary the Waldhof NO mixing ratio was significantly higher (1.0 ppb), and O<sub>3</sub> (36.4 ppb) and SO<sub>2</sub> (1.0 ppb)
- mixing ratios were somewhat lower than the airborne measurements. The measured air temperature and pressure however matched very well.

The GOM concentration from the manual denuder sampling was calculated to be between 24.6 and  $37.3 \text{ pgm}^{-3}$ , which is much higher than the ground based measured

<sup>25</sup> concentration. This mismatch is probably caused by the different air masses that have been sampled. While the ground based measurement represent only air from the lower PBL which are directly influenced by dry deposition, the airborne sampling represents the entire air column from inside the PBL to the lower FT. Vertical GOM profiles reported by Brooks et al. (2014) show clearly a tendency to lower GOM concentrations



at the lowest altitudes. Except Iskraba, all denuder samples were taken to more than 17% of the sampling time above the PBL. The three denuder samples taken on the transfer flights between Italy and Germany do represent mainly FT air. For all three denuders a relatively high GOM loading of  $18.4-65.6 \text{ pgm}^{-3}$  was found (Table 3). These

- <sup>5</sup> GOM concentrations are in reasonable agreement with ~ 20–110 pg m<sup>-3</sup> measured in August 2012 at altitudes up to 6 km by Brooks et al. (2014) over Tennessee, USA. It is assumed that above the PBL the GOM concentration is higher because less aerosol surface is available to condense the GOM onto, the RH is usually lower, and the radiation flux is higher (less humidity results in fewer clouds and less light scattering.
- <sup>10</sup> Furthermore, solar radiation is scattered and reflected at the PBL cloud top and is partly scattered back above the PBL). All these conditions favour elevated GOM concentrations with a maximum at altitudes between 2 and 5 km (Brooks et al., 2014). Within future studies more detailed GOM vertical profiles with a better vertical resolution should be carried out.

#### 15 4 Conclusions

Opposite to most of the previously reported vertical profiles, we always observed a significant difference between PBL and FT air (Fig. 7). While the FT GEM and TGM background concentration over central Europe was measured to ~1.3 ngm<sup>-3</sup>, 10–30 % higher GEM and TGM concentrations were found in the PBL. Besides this abrupt jump at the PBL top, at all sampling locations, neither in the boundary layer, nor in the free troposphere a clear vertical gradient was apparent. This is in agreement with most of the vertical profiles obtained elsewhere (Swartzendruber et al., 2009; Brooks et al., 2014). Vertical profiles with pronounced decreasing GEM concentrations with increasing altitude were reported by Radke et al. (2007) and Brooks et al. (2014), but only <sup>25</sup> for spring month April, May, and June. These are the months with the strongest strato-

for spring month April, May, and June. These are the months with the strongest stratosphere to troposphere ozone flux in the Northern Hemisphere (Olsen et al., 2004) and the anomalous vertical profiles with strong vertical GEM gradients are probably related



to it. In summer months GEM and TGM are homogeneously distributed inside the PBL and FT. The combination of ETMEP-2 measurements over Leipzig with CARIBIC measurement over Western Europe (Fig. 5) gives a unique vertical profile from 0.5 km (lower PBL) to 10.5 km (upper FT). From above the PBL to the FT top the TGM back-<sup>5</sup> ground concentration is on average 1.3 ngm<sup>-3</sup>.

During all vertical profiles, as well as during the transfer flights between Slovenia, Germany, and Italy denuder samples were taken in PBL- and FT air. The analysis of the denuders for GOM indicated an increased GOM concentration above the PBL. As this is a region favouring the generation of GOM (low particle surface (in comparison to PBL), low humidity, and high actinic fluxes from top and below (reflection at PBL top)), this finding is reasonable. The vertical distribution and the range of observed GOM concentrations reported here are in agreement with measurements by Brooks et al. (2014). Considering a GOM concentration between 1.0 and 65.5 pg m<sup>-3</sup> (Table 3) and an average FT GEM concentration of 1.3 ng m<sup>-3</sup>, the ratio of GOM to GEM is 0.1

to 5.0%. Note however the large uncertainty of the measured GOM concentrations of  $\pm 5 \text{ pg m}^{-3}$ . Therefore, more detailed GOM vertical profiles with a better vertical resolution should be carried out in the future.

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#### References

- 5 Banic, C. M.: Vertical distribution of gaseous elemental mercury in Canada, J. Geophys. Res., 108, 1–14, doi:10.1029/2002JD002116, 2003.
- Brenninkmeijer, C. A. M., Crutzen, P., Boumard, F., Dauer, T., Dix, B., Ebinghaus, R., Filippi, D., Fischer, H., Franke, H., Frieß, U., Heintzenberg, J., Helleis, F., Hermann, M., Kock, H. H., Koeppel, C., Lelieveld, J., Leuenberger, M., Martinsson, B. G., Miemczyk, S., Moret, H. P.,
- Nguyen, H. N., Nyfeler, P., Oram, D., O'Sullivan, D., Penkett, S., Platt, U., Pupek, M., Ramonet, M., Randa, B., Reichelt, M., Rhee, T. S., Rohwer, J., Rosenfeld, K., Scharffe, D., Schlager, H., Schumann, U., Slemr, F., Sprung, D., Stock, P., Thaler, R., Valentino, F., van Velthoven, P., Waibel, A., Wandel, A., Waschitschek, K., Wiedensohler, A., Xueref-Remy, I., Zahn, A., Zech, U., and Ziereis, H.: Civil Aircraft for the regular investigation of
- the atmosphere based on an instrumented container: The new CARIBIC system, Atmos. Chem. Phys., 7, 4953–4976, doi:10.5194/acp-7-4953-2007, 2007.
  - Brooks, S., Ren, X., Cohen, M., Luke, W. T., Kelley, P., Artz, R., Hynes, A., Landing, W., and Martos, B.: Airborne vertical profiling of mercury speciation near Tullahoma, TN, USA, Atmosphere (Basel), 5, 557–574, doi:10.3390/atmos5030557, 2014.
- <sup>20</sup> Brosset, C.: The behavior of mercury in the physical environment, Water Air Soil Pollut., 34, 145–166, 1987.
  - Ebinghaus, R. and Slemr, F.: Aircraft measurements of atmospheric mercury over southern and eastern Germany, Atmos. Environ., 34, 895–903, doi:10.1016/S1352-2310(99)00347-7, 2000.
- <sup>25</sup> Ebinghaus, R., Slemr, F., Brenninkmeijer, C. A. M., van Velthoven, P., Zahn, A., Hermann, M., O'Sullivan, D. A., and Oram, D. E.: Emissions of gaseous mercury from biomass burning in South America in 2005 observed during CARIBIC flights, Geophys. Res. Lett., 34, 1–5, doi:10.1029/2006GL028866, 2007.

Fitzgerald, W. F., Engstrom, D. R., Mason, R. P., and Nater, E. A.: The case for atmospheric mer-

<sup>30</sup> cury contamination in remote areas, Environ. Sci. Technol., 32, 1–7, doi:10.1021/es970284w, 1998.



- Friedli, H. R.: Mercury in the atmosphere around Japan, Korea, and China as observed during the 2001 ACE-Asia field campaign: measurements, distributions, sources, and implications, J. Geophys. Res., 109, 1–13, doi:10.1029/2003JD004244, 2004.
- Grönlund, R., Edner, H., Svanberg, S., Kotnik, J., and Horvat, M.: Mercury emissions from the Idrija mercury mine measured by differential absorption lidar techniques and a point monitoring absorption spectrometer, Atmos. Environ., 39, 4067–4074, doi:10.1016/j.atmosenv.2005.03.027, 2005.
  - Kvietkus, K.: Investigation of the gaseous and particulate mercury concen- trations along horizontal and vertical profiles in the lower troposphere, in: Proceedings of the 10th World Clean
- <sup>10</sup> Air Congress, Held at Espoo, Finland, 28 May–2 June 1995: Atmospheric Pollution, edited by: Anttiba, P., Kämäri, J., and Jolvanen, M., Finnish Air Pollution Prevention Society, 1995, Helsinki, 284–287, 1995.
  - Lindberg, S., Bullock, R., Ebinghaus, R., Engstrom, D., Feng, X., Fitzgerald, W., Pirrone, N., Prestbo, E., and Seigneur, C.: A synthesis of progress and uncertainties in attributing the sources of mercury in deposition. AMBIO A J. Hum, Environ., 36, 19–33, doi:10.1579/0044-
- sources of mercury in deposition, AMBIO A J. Hum. Environ., 36, 19–33, doi:10.1579/0044-7447(2007)36[19:ASOPAU]2.0.CO;2, 2007.
  - Lyman, S. N. and Jaffe, D. A.: Formation and fate of oxidized mercury in the upper troposphere and lower stratosphere, Nat. Geosci., 5, 114–117, doi:10.1038/ngeo1353, 2011.
  - Olsen, M. A., Schoeberl, M. R., and Douglass, A. R.: Stratosphere–troposphere exchange of mass and ozone, J. Geophys. Res., 109, D24114, doi:10.1029/2004JD005186, 2004.

20

- Pirrone, N., Cinnirella, S., Feng, X., Finkelman, R. B., Friedli, H. R., Leaner, J., Mason, R., Mukherjee, A. B., Stracher, G. B., Streets, D. G., and Telmer, K.: Global mercury emissions to the atmosphere from anthropogenic and natural sources, Atmos. Chem. Phys., 10, 5951–5964, doi:10.5194/acp-10-5951-2010, 2010.
- Radke, L. F., Friedli, H. R., and Heikes, B. G.: Atmospheric mercury over the NE Pacific during spring 2002: Gradients, residence time, upper troposphere lower stratosphere loss, and long-range transport, J. Geophys. Res., 112, 1–17, doi:10.1029/2005JD005828, 2007.
   Scheele, M. P., Siegmund, P. C., and Van Velthoven, P. F. J.: Sensitivity of trajectories to data resolution and its dependence on the starting point: in or outside a tropopause fold, Meteorol.
- Appl., 3, 267–273, doi:10.1002/met.5060030308, 2007.
   Selin, N. E.: Global biogeochemical cycling of mercury: a review, Annu. Rev. Environ. Resour., 34, 43–63, doi:10.1146/annurev.environ.051308.084314, 2009.



28236

- Slemr, F., Schuster, G., and Seiler, W.: Distribution, speciation, and budget of atmospheric mercury, J. Atmos. Chem., 3, 407–434, doi:10.1007/BF00053870, 1985.
- Slemr, F., Ebinghaus, R., Brenninkmeijer, C. A. M., Hermann, M., Kock, H. H., Martinsson, B. G., Schuck, T., Sprung, D., van Velthoven, P., Zahn, A., and Ziereis, H.: Gaseous mercury distri-
- <sup>5</sup> bution in the upper troposphere and lower stratosphere observed onboard the CARIBIC passenger aircraft, Atmos. Chem. Phys., 9, 1957–1969, doi:10.5194/acp-9-1957-2009, 2009.
  - Slemr, F., Weigelt, A., Ebinghaus, R., Brenninkmeijer, C., Baker, A., Schuck, T., Rauthe-Schöch, A., Riede, H., Leedham, E., Hermann, M., van Velthoven, P., Oram, D., O'Sullivan, D., Dyroff, C., Zahn, A., and Ziereis, H.: Mercury plumes in the global upper tropo-
- <sup>10</sup> sphere observed during flights with the CARIBIC observatory from May 2005 until June 2013, Atmosphere (Basel), 5, 342–369, doi:10.3390/atmos5020342, 2014.
  - Swartzendruber, P. C., Jaffe, D. A., Prestbo, E. M., Weiss-Penzias, P., Selin, N. E., Park, R., Jacob, D. J., Strode, S., and Jaeglé, L.: Observations of reactive gaseous mercury in the free troposphere at the Mount Bachelor Observatory, J. Geophys. Res. Atmos., 111, D24301, doi:10.1029/2006JD007415, 2006.
- Swartzendruber, P. C., Chand, D., Jaffe, D. A., Smith, J., Reidmiller, D., Gratz, L., Keeler, J., Strode, S., Jaeglé, L., and Talbot, R.: Vertical distribution of mercury, CO, ozone, and aerosol scattering coefficient in the Pacific Northwest during the spring 2006 INTEX-B campaign, J. Geophys. Res. Atmos., 113, D10305, doi:10.1029/2007JD009579, 2008.

15

25

- Swartzendruber, P. C., Jaffe, D. A., and Finley, B.: Development and first results of an aircraftbased, high time resolution technique for gaseous elemental and reactive (oxidized) gaseous mercury, Environ. Sci. Technol., 43, 7484–7489, doi:10.1021/es901390t, 2009.
  - Talbot, R., Mao, H., Scheuer, E., Dibb, J., and Avery, M.: Total depletion of Hg<sup>°</sup> in the upper troposphere-lower stratosphere, Geophys. Res. Lett., 34, L23804, doi:10.1029/2007GL031366, 2007.
- Talbot, R., Mao, H., Scheuer, E., Dibb, J., Avery, M., Browell, E., Sachse, G., Vay, S., Blake, D., Huey, G., and Fuelberg, H.: Factors influencing the large-scale distribution of Hg<sup>°</sup> in the Mexico City area and over the North Pacific, Atmos. Chem. Phys., 8, 2103–2114, doi:10.5194/acp-8-2103-2008, 2008.
- Weigelt, A., Temme, C., Bieber, E., Schwerin, A., Schuetze, M., Ebinghaus, R., and Kock, H. H.: Measurements of atmospheric mercury species at a German rural background site from 2009 to 2011 – methods and results, Environ. Chem., 10, 102–110, doi:10.1071/EN12107, 2013.



- Weigelt, A., Ebinghaus, R., Manning, A. J., Derwent, R. G., Simmonds, P. G., Spain, T. G., Jennings, S. G., and Slemr, F.: Analysis and interpretation of 18 years of mercury observations since 1996 at Mace Head, Ireland, Atmos. Environ., 100, 85–93, doi:10.1016/j.atmosenv.2014.10.050, 2015a.
- <sup>5</sup> Weigelt, A., Ebinghaus, R., Pirrone, N., Bödewadt, J., Esposito, G., Mannarino, V., Montagnoli, M., Slemr, F., Sprovieri, F., and van Velthoven, P. F. J.: Airborne measurements of the mercury emissions from Mt. Etna volcano (Italy) in July/August 2012, in preparation, 2015b. Weigelt, A., Ebinghaus, R., Pirrone, N., Bieser, J., Bödewadt, J., Esposito, G., and Slemr, F.: Airborne measurements of mercury emissions from a modern coal fired power plant in central
- <sup>10</sup> Europe, in preparation, 2015c.



**Table 1.** Summary of all known European airborne atmospheric mercury measurements untilDecember 2014.

Time	Location	Altitude	Key finding	Literature
1978–1981	Central Europe	6–12 km	<ul> <li>no vertical gradient</li> </ul>	Slemr et al. (1985)
1981	West of Göteborg	up to 3 km	• decrease with altitude proportional to pressure decrease $\rightarrow$ no vertical gradient when transferring to STP conditions	Brosset (1987)
Jun 1988	Eastern Lithuania	?	• concentration proportional to pressure at sampling altitude → no vertical gradient when transferring to STP conditions	Kvietkus et al. (1995)
Jun 1996	Eastern Germany	0.5–3.75 km	<ul> <li>no vertical gradient</li> <li>increased concentration observed near source region up to ~ 2 km altitude</li> </ul>	Ebinghaus and Slemr (2000)
since 2005	Europe and global (CARIBIC Project)	6–12 km	<ul> <li>long term monitoring in UT and LS (trend analysis)</li> <li>large scale plume identification</li> </ul>	Slemr et al. (2009, 2014) www.caribic- atmospheric.com
Jul/Aug 2012	Mt. Etna volcano (Southern Italy)	0–4 km	no/low gaseous mercury emis- sion from Mt. Etna volcano	Weigelt et al. (2015b)
Aug 2013	Central Europe (Slovenia and Germany)	0–3 km 6–11 km	• significant difference between boundary layer and free tropo- sphere, but no vertical gradient inside individual layers	this study

**ACPD** 15, 28217-28247, 2015 **Tropospheric** mercury vertical profiles between 500 and 10 000 m A. Weigelt et al. **Title Page** Abstract Introduction References Tables Figures 4 Close Back Full Screen / Esc **Printer-friendly Version** Interactive Discussion

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**Table 2.** List of instruments, installed into the CASA 212 research aircraft. The acronyms are: GEM = gaseous elemental mercury; GOM = gaseous oxidized mercury; CO = CarbonMonoxide;  $O_3 = Ozone$ ;  $SO_2 = Sulphur dioxide$ ; NO = Nitric oxide;  $NO_2 = Nitric dioxide$ .

Parameter	Instrument name	Temporal resolution	Uncertainty	Lower detection limit
GEM	Lumex RA-915AM (modified, T- stabilised by Lumex company)	1 s (raw signal)	±4 ng m <sup>-3</sup> (1 s raw signal) ±0.25 ng m <sup>-3</sup> (120 s average)	0.5 ng m <sup>-3</sup> (120 s average)
GEM	Tekran: 2537X (with upstream quartz wool trap)	150 s	$\pm 0.1  \text{ng}  \text{m}^{-3}$	0.1 ng m <sup>-3</sup>
GEM + unknown amount of GOM <sup>a</sup>	Tekran 2537B	150 s	$\pm 0.1  \text{ng}  \text{m}^{-3}$	0.1 ng m <sup>-3</sup>
GOM	manually denuder samples	2600 to 3600 s	±5 pg m <sup>-3,b</sup>	1 pg m <sup>-3</sup>
СО	Aero Laser AL5002	1 s	±1 ppb	1.5 ppb
O <sub>3</sub>	Teledyne API 400A	10 s	±0.5% of reading	0.6 ppb
SO <sub>2</sub>	Thermo: 43C Trace Level	10 s	±3% of reading	0.2 ppb
NO NO <sub>2</sub>	Teledyne API M200EU	10 s 10 s	±10% of reading	0.05 ppb
Pressure	Sensor Technics CTE7001	1 s	±1% of reading	0 mbar
Temperature	LKM Electronic DTM5080	1 s	±0.13°C	–50 °C
Relative Humidity (RH)	Vaisala HMT333	8s	±1.0 % RH (0–90 % RH) ±1.7 % RH (90–100 % RH)	0%
GPS data (3d position, speed, heading)	POS AV	1s	±5 m (horizontal) <sup>c</sup> ±15 (vertical) <sup>c</sup>	-

<sup>a</sup> The aircraft inlet system transmission efficiency for GOM was not tested.

<sup>b</sup> Difference of the two blank tests.

<sup>c</sup> The GPS accuracy is dependent on the number of satellites. The given numbers are estimated values.

<b>AC</b> 15, 28217–	<b>PD</b> 28247, 2015			
Tropospheric mercury vertical profiles between 500 and 10 000 m				
A. Weigelt et al.				
Title	Title Page			
Abstract	Introduction			
Conclusions	References			
Tables	Figures			
14	►I			
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Back	Close			
Full Screen / Esc				
Printer-friendly Version				
Interactive Discussion				

**Discussion Paper** 

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Discussion Paper

**Discussion** Paper

**Table 3.** Results of the manual denuder samples during all ETMEP-2 measurement flights in 2013 over central Europe. GOM data were corrected for denuder blank test, additionally performed over Iskraba/Slovenia and Waldhof/Germany.

Date	Location	Profile character (relative sam- pling time in PBL <sup>a</sup> and FT <sup>b</sup> air)	analysed GOM concentration [pg m <sup>-3</sup> ]
19 Aug 2013	Iskraba/Slovenia	vertical (100 % PBL; 0 %FT)	1.9–13.2
19 Aug 2013	Idrija/Slovenia	vertical (83 % PBL; 17 % FT)	18.0–28.8
21 Aug 2013	Ronchi dei Legionari/Italy to Leipzig/Germany	horizontal (20 % PBL; 80 % FT)	18.4–24.0
21 Aug 2013	Lippendorf/Germany	vertical (76 % PBL: 24 % FT)	7.0–15.7
21 Aug 2013	Leipzig/Germany	vertical (61 % PBL; 39 % FT)	1.0 <sup>c</sup> –10.6
22 Aug 2013	Waldhof/Germany	vertical (54 % PBL; 46 % FT)	24.6–37.3
22 Aug 2013	Leipzig/Germany to Parma/Italy 1 (central- and south Germany)	horizontal (0 % PBL; 100 % FT)	55.4–65.6
22 Aug 2013	Leipzig/Germany to Parma/Italy 2 (south Germany and Alps)	horizontal (0 % PBL; 100 % FT)	22.1–31.1

<sup>a</sup> Planetary boundary layer (PBL).

<sup>b</sup> Free troposphere (FT).

<sup>c</sup> If a concentration was found to be below the method lower detection limit of 1.0 pg m<sup>-3</sup>, the lower detection limit is given.





**Figure 1.** Flight tracks of the European Tropospheric Mercury Experiment part 2 (ETMEP-2) research flights in August 2013. Flights are separated by the flight track colour. The home base of the used aircraft was Parma/Italy. Over Waldhof, Leipzig, Lippendorf, Idrija, and Iskraba vertical profiles were flown. The underlying map was taken from Google Earth.





**Figure 2.** For the ETMEP-2 campaign in August 2013 the CASA 212 (a) from the Italian company Compagnia Generale Ripreseaeree (http://www.terraitaly.it/) was equipped with specially designed and manufactured trace gas inlet (b).





**Figure 3.** Vertical profile, measured on 19 August 2013 from 13:17:30 to 14:07:30 (local time) over the GMOS master site "Iskraba" (45.561° N, 14.858° E, elevation: 530 m a.s.l.; mountain terrain). Squares represent 300 s averages with horizontal flight leg; stars indicate 150 s averages during climbing between two neighbouring flight legs. GOM was sampled onto a denuder during the whole profile. Two blank measurements were performed at the beginning and at the end of the ETMEP-2 campaign. Therefore the given GOM concentration (high concentration with consideration of low blank and vice versa) is an average over the whole air column. The red dashed line indicates the planetary boundary layer (PBL) top, which is not representative here because all measurements were performed below the boundary layer top. GEM and TGM concentrations are given at standard conditions (p = 1013.25 hPa, T = 273.15 K).





**Figure 4.** Same as Fig. 3, but for the former mercury mining area "Idrija" (45.000° N, 14.022° E, elevation: 330 m; mountain terrain up to 800 m). The profile was measured on 19 August 2013 from 14:30:00 to 15:20:00 (local time). The PBL top (red dashed line) was determined to be at 2850 to 2900 m a.s.l. TGM and GEM concentrations are given at standard conditions (p = 1013.25 hPa, T = 273.15 K).





**Figure 5.** Vertical profile, measured within the ETMEP-2 campaign on 21 August 2013 from 13:15:00 to 14:07:30 (local time) over the city centre of Leipzig/Germany (51.353° N, 12.434° E, elevation: 125 m, flat terrain) and from 21–23 August 2013 over Western Europe (east of 0° W; CARIBIC). While the ETMEP-2 data were averaged for 300 s (squares) and 150 s (stars), the CARIBIC data (triangles) represent 600 s averages. The plots have the same structure as Fig. 3. The PBL top (red dashed line) was determined to be at 2200 to 2250 ma.s.l. Please note, *y* axis is broken between 3500 and 6000 m. TGM and GEM concentrations are given at standard conditions (*p* = 1013.25 hPa, *T* = 273.15 K).





**Figure 6.** Same as Fig. 3, but over the GMOS master site Waldhof/Germany (52.801° N, 10.756° E, elevation: 75 m, flat terrain). The profile was measured on 22 August 2013 from 10:22:30 to 11:17:30 (local time). The PBL top (red dashed line) was determined to be at 1750 to 1850 ma.s.l. Additionally the data measured at the same time at the ground based site "Waldhof" are plotted. TGM and GEM concentrations are given at standard conditions (p = 1013.25 hPa, T = 273.15 K).







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Figure 7. Comparison of known vertical gaseous mercury profiles (TGM and GEM). Data plotted in black were taken from Swatzendruber et al. (2009). Data in grey represent the August measurement from Brooks et al. (2014). Coloured data represent ETMEP-2 data (Fig. 3-6). The Waldhof 1.47 km fight leg average was removed for this plot, because of probably inside plume measurement (cf. discussion to Fig. 6).