Simulating Secondary Organic Aerosol in a Regional Air 1

Quality Model Using the Statistical Oxidation Model: 1. 2

Assessing the Influence of Constrained Multi-generational 3

#### Ageing 4

- Shantanu H. Jathar<sup>1,2</sup>, Christopher D. Cappa<sup>2</sup>, Anthony S. Wexler<sup>2</sup>, John H. Seinfeld<sup>3</sup>, and Michael J. Kleeman<sup>2</sup> 5
- 6
- 7 <sup>1</sup> Mechanical Engineering, Colorado State University, Fort Collins CO
- <sup>2</sup> Civil and Environmental Engineering, University of California, Davis CA 8
- 9 <sup>3</sup> Chemical Engineering, California Institute of Technology, Pasadena CA
- 10 Corresponding authors: Christopher D. Cappa (cdcappa@ucdavis.edu) and Michael J. Kleeman

Multi-generational oxidation of volatile organic compound (VOC) oxidation products can

- (mikleeman@ucdavis.edu). 11
- 12

14

#### 13 Abstract

15 significantly alter the mass, chemical composition and properties of secondary organic aerosol 16 (SOA) compared to calculations that consider only the first few generations of oxidation 17 reactions. However, the most commonly used state-of-the-science schemes in 3-D regional or 18 global models that account for multi-generational oxidation (1) consider only functionalization 19 reactions but do not consider fragmentation reactions; (2) have not been constrained to 20 experimental data; and (3) are added on top of existing parameterizations. The incomplete 21 description of multi-generational oxidation in these models has the potential to bias source 22 apportionment and control calculations for SOA. In this work, we used the Statistical Oxidation 23 Model (SOM) of Cappa and Wilson (2012), constrained by experimental laboratory chamber 24 data, to evaluate the regional implications of multi-generational oxidation considering both 25 functionalization and fragmentation reactions. SOM was implemented into the regional 26 UCD/CIT air quality model and applied to air quality episodes in California and the eastern US. 27 The mass, composition and properties of SOA predicted using SOM were compared to SOA 28 predictions generated by a traditional "two-product" model to fully investigate the impact of 29 explicit and self-consistent accounting of multi-generational oxidation. 30 Results show that SOA mass concentrations predicted by the UCD/CIT-SOM model are 31 very similar to those predicted by a two-product model when both models use parameters that 32 are derived from the same chamber data. Since the two-product model does not explicitly resolve 33 multi-generational oxidation reactions, this finding suggests that the chamber data used to 34 parameterize the models captures the majority of the SOA mass formation from multi-

generational oxidation under the conditions tested. Consequently, the use of low and high NO<sub>x</sub> 35 36 vields perturbs SOA concentrations by a factor of two and are probably a much stronger 37 determinant in 3-D models than multi-generational oxidation. While total predicted SOA mass is 38 similar for the SOM and two-product models, the SOM model predicts increased SOA 39 contributions from anthropogenic (alkane, aromatic) and sesquiterpenes and decreased SOA 40 contributions from isoprene and monoterpene relative to the two-product model calculations. The 41 SOA predicted by SOM has a much lower volatility than that predicted by the traditional model; 42 resulting in better qualitative agreement with volatility measurements of ambient OA. On 43 account of its lower-volatility, the SOA mass produced by SOM does not appear to be as 44 strongly influenced by the inclusion of oligomerization reactions, whereas the two-product 45 model relies heavily on oligomerization to form low volatility SOA products. Finally, an 46 unconstrained contemporary hybrid scheme to model multi-generational oxidation within the 47 framework of a two-product model in which "ageing" reactions are added on top of the existing 48 two-product parameterization is considered. This hybrid scheme formed at least three times more 49 SOA than the SOM during regional simulations as a result of excessive transformation of semi-50 volatile vapors into lower volatility material that strongly partitions to the particle phase. This 51 finding suggests that these "hybrid" multi-generational schemes should be used with great 52 caution in regional models.

53

#### 54 **1.** Introduction

Organic aerosol (OA) is generally the dominant component of submicrometer-sized atmospheric particulate matter (Jimenez et al., 2009), which plays an important role in the energy budget of the earth (IPCC, 2007) and the health effects of air pollution (Bernstein et al., 2004). Despite its prominence, OA is the least understood component of atmospheric aerosol. Largescale chemical transport models are the essential tool to simulate concentration distributions, which are needed to form strategies to mitigate, the climate and health impacts of atmospheric aerosols.

62 OA is a complex mixture of thousands of different compounds that have a wide range of 63 properties (Goldstein and Galbally, 2007). OA can be directly emitted to the atmosphere in 64 particulate form (so-called primary organic aerosol; POA) or it can be formed *in situ* by the 65 oxidation of volatile organic compounds (VOCs) to yield lower volatility products that condense 66 into the aerosol phase, so-called secondary organic aerosol (SOA). This latter route is generally 67 the predominant one to form OA. Continuous oxidation of VOCs and their oxidation products 68 yields a broad range of products, including those that have intermediate and low volatility. The 69 importance of such "multi-generational oxidation" on SOA production has been widely 70 established in laboratory chamber experiments (Chacon-Madrid et al., 2010;Chacon-Madrid et 71 al., 2013; Yee et al., 2013; Donahue et al., 2012; Chhabra et al., 2011; Henry and Donahue, 2012). 72 Multi-generational oxidation includes the initial formation of oxidized products of lower 73 volatility as well as the loss of SOA mass after initial formation owing to fragmentation 74 reactions. For example, experiments performed with the Potential Aerosol Mass (PAM) reactor, 75 which aims to simulate prolonged VOC oxidation, are always associated with formation 76 followed by destruction of OA mass (Lambe et al., 2012). Simulations that capture this behavior 77 require inclusion of multi-generational oxidation. In addition to altering predicted SOA mass, 78 inclusion of multi-generational oxidation is expected to alter the oxidation state of OA, which 79 has important repercussions for OA properties (e.g., water uptake, toxicity) (Jimenez et al., 80 2009).

81 Traditionally, models of SOA formation in chamber experiments have represented SOA 82 formation from VOCs using two to four surrogate products per VOC, the yields for which have 83 been parameterized to reproduce observed levels of SOA (Odum et al., 1996). These models 84 generally assume that the surrogate products are non-reactive (i.e., do not undergo multi-85 generational oxidation). These models, whether implemented in "two-product" or "volatility 86 basis set" (VBS) forms (Donahue et al., 2006), generally under-predict ambient concentrations of 87 SOA (Carlton et al., 2010). Some models have used simple chemical schemes to mimic the 88 effects of multi-generational oxidation. While these schemes differ in their details, in essence, 89 they assume that the vapors and the products of each surrogate traditional VOC species react 90 with the hydroxyl radical (OH) to form lower volatility products (Robinson et al., 2007; Pye and 91 Seinfeld, 2010; Baek et al., 2011). Such "ageing" schemes to account for multi-generational 92 oxidation of traditional VOC products share similarities with reaction schemes applied to the 93 oxidation of intermediate-volatility organic compounds (IVOCs) and POA vapors (Robinson et 94 al., 2007). Note that oxidation of IVOCs and POA vapors is assumed to proceed only through 95 these ageing-type reactions, whereas oxidation of the semi-volatile products of traditional VOC 96 precursors is an augmentation to the existing two-product or VBS parameterization. Models that

97 include these ageing schemes predict SOA mass concentrations that close the gap with measured 98 ambient concentrations of OA mass. As a result, over the past five years, both research and 99 regulatory groups have incorporated these schemes into their 3-D models (e.g., Environmental 100 Protection Agency's Community Multiscale Air Quality Model (CMAQ) (Koo et al., 2014), 101 PMCAMx (Murphy and Pandis, 2009; Tsimpidi et al., 2009), WRF-CHEM (Ahmadov et al., 102 2012;Lane et al., 2008;Tsimpidi et al., 2009)). These first order SOA schemes have three major 103 mechanistic drawbacks. First, they typically do not account for laboratory evidence of 104 fragmentation of oxygenated organic molecules that can lead to decreases in SOA concentrations 105 (Chacon-Madrid and Donahue, 2011;Henry and Donahue, 2012). Second, they assume that the 106 multi-generational oxidation of products of different anthropogenic VOCs (e.g., alkanes versus 107 aromatics) or different biogenic VOCs (e.g., isoprene versus monoterpenes) share the same 108 reaction mechanism. Finally (and most importantly), these schemes remain under-unconstrained 109 in that they have not been rigorously tested against measurements of multi-generational products 110 (or classes of products) under realistic ambient conditions, and they are typically added on top of 111 existing parameterizations. These concerns apply specifically to the multi-generational oxidation 112 schemes that are commonly applied to traditional VOCs, but these are also relevant to the 113 oxidation schemes associated with IVOCs and POA vapors. Chemically explicit models have 114 seldom been used in 3-D modeling (e.g. Johnson et al. (2006), Chen et al. (2006), Ying and Li 115 (2011)) due to their heavy computational burden, although some studies have used reduced complexity forms for 3-D modeling (e.g. Utembe et al. (2011), Lin et al. (2012)) or have 116 117 implemented them for box modeling studies (e.g. Lee-Taylor et al. (2011)). 118 In this work, we use the Statistical Oxidation Model (SOM) of Cappa and Wilson (2012) 119 to model the multi-generational oxidation reactions inherent in SOA formation. The SOM 120 provides an efficient framework to track the experimentally-constrained chemical evolution and 121 gas/particle partitioning of SOA using a carbon and oxygen grid. In Jathar et al. (2015), we 122 detailed the coupling of the SOM with the gas-phase chemical mechanism SAPRC-11 (Carter 123 and Heo, 2013) within the UCD/CIT regional air quality model and used the new model to make 124 predictions over the South Coast Air Basin (SoCAB) in California and the eastern United States (US). Here, we use the UCD/CIT-SOM model to investigate the influence of constrained multi-125

126 generational oxidation on the mass concentrations and properties of SOA and contrast those

results against predictions from a traditional two-product model and an unconstrained multi-generational oxidation model.

129

# 130 2. Model Description and Simulations

131 **2.1. Air Quality Model** 

132 The UCD/CIT air quality model is a regional chemical transport model (CTM) (Kleeman 133 and Cass, 2001) used here to simulate SOA formation for two geographically-distinct domains 134 and time periods: (1) the state of California simulated at a grid resolution of 24 km followed by a 135 nested simulation over the SoCAB at a grid resolution of 8 km from July 20 to August 2, 2005, and (2) the eastern half of the US simulated at a grid resolution of 36 km from August 20<sup>th</sup> to 136 September 2<sup>nd</sup>, 2006. Details about the latest version of the UCD/CIT model are provided in 137 138 Jathar et al. (2015) and summarized in Table S.1. Briefly, anthropogenic emissions for California 139 were based on the California Regional PM10/PM2.5 Air Quality Study (CRPAQS) inventory of 140 2000 but scaled to match conditions in 2005. FINN (Fire Inventory for National Center for 141 Atmospheric Research) (Wiedinmyer et al., 2011) and MEGAN (Model of Emissions of Gases 142 and Aerosols from Nature) (Guenther et al., 2006) were used to calculate wildfire and biogenic 143 emissions in California. Anthropogenic and wildfire emissions for the eastern US were based on 144 the 2005 National Emissions Inventory (NEI), and biogenic emissions were estimated using 145 BEIS (Biogenic Emissions Inventory System) version 3. Hourly meteorological fields were 146 generated using the Weather Research and Forecasting (WRF) v3.4 model (www.wrf-147 model.org). National Center for Environmental Protection's NAM (North American Mesoscale) 148 analysis data were used to set the initial and boundary conditions for WRF. Gas- and particle-149 phase initial and hourly-varying boundary conditions were based on the results from the global 150 model MOZART-4/NCEP (Emmons et al., 2010). Gas-phase chemistry was modeled using 151 SAPRC-11. In all simulations, POA was treated as non-volatile, yet absorptive, as per the 152 treatment in the regulatory Community Multiscale Air Quality (CMAQ) version 4.7 model 153 (Carlton et al., 2010). As such, contributions of semi-volatile and intermediate volatility organic 154 compound emissions (which are commonly assumed to originate from the evaporation of and co-155 emitted with POA) to the SOA burden were not considered in this study.

156

#### 157 **2.2.** SOA Models

158 Four types of SOA models are compared in this work: (1) A "Base" two-product model 159 that is equivalent to the SOA model used in CMAQ and representative of SOA models used in 160 most chemical transport (Carlton et al., 2010) and global climate models (Henze et al., 2008); (2) 161 A modified version of the Base model, "BaseM", which uses the two-product framework, but in 162 which the SOA formation parameters were determined using newer chamber data; (3) A "SOM" 163 model (Cappa and Wilson, 2012) in which multi-generational oxidation is accounted for through 164 semi-explicit representation of progressive generations of gas-phase oxidation of the products 165 and precursors of SOA, and that was parameterized based on the same dataset as the BaseM 166 model; (4) A "cascading" oxidation model, wherein ageing of semi-volatile products was 167 accounted for a posteriori using ageing rates derived from separate experiments. All of the SOA 168 models utilize fully dynamic gas/particle partitioning for OA species as in Kleeman and Cass 169 (2001). The following sub-sections describe the four SOA models. To aid the reader, a 170 conceptual schematic comparing various SOA models (e.g. 2-product, SOM, VBS) is provided 171 in Figure. S.1.

172

# 173 **2.2.1. Base**

174 The Base model simulated SOA formation as per the pathways and parameters in CMAQ 175 model version 4.7 (Carlton et al., 2010) from the following gas-phase precursors: long alkanes 176 (ALK5), benzene (BENZENE), low-yield aromatics (ARO1), high-yield aromatics (ARO2), 177 isoprene, monoterpenes (TRP1) and sesquiterpenes (SESQ). The species in parentheses are the 178 model species representing those compounds in SAPRC-11 (the gas-phase chemical mechanism 179 used here). The pathways considered include: (1) oxidation of the above-mentioned precursors to 180 form non-reactive semi-volatile products that partition into the particle-phase (Odum et al., 1996) 181 (the so-called two-product model, where model parameters were previously determined from 182 fitting chamber data); (2) acid enhancement of isoprene SOA (Surratt et al., 2007). SOA 183 formation from aromatics is NO<sub>x</sub> dependent; low levels of NO<sub>x</sub> result in higher SOA formation 184 and vice-versa. The Base model was extended to include particle-phase oligomerization 185 (Kalberer et al., 2004), for which particle-phase semi-volatile components were converted to non-volatile components with  $k_{\text{oligomer}} = 9.6 \times 10^{-6} \text{ s}^{-1}$ . In summary, the Base model was run in 186 187 two configurations, with and without oligomerization reactions: Base and Base-OLIG.

188

#### 189 **2.2.2. Base Modified**

The "modified" version of the Base model, termed "BaseM" was created to facilitate a 190 191 true evaluation of multi-generational oxidation in a two-product model framework. The BaseM 192 model: (1) used recent chamber data (Jathar et al., 2015) from California Institute of Technology 193 to determine alternate two-product model parameters; and (2) did not include acid-catalyzed 194 enhancement of isoprene SOA and oligomerization reactions. The two-product fit parameters 195 and data sources are listed in Table S.2. Note that the "long alkane" BaseM parameterization has 196 been developed using experimental results for SOA formation from *n*-dodecane (Loza et al., 197 2014).

198

199

# 99 **2.2.3. Statistical Oxidation Model**

200 The SOM parameterizes multi-generational oxidation using a two-dimensional carbon-201 oxygen grid to track the evolution of gas- and particle-phase organic products arising from the 202 oxidation of SOA precursors (Cappa and Wilson, 2012;Cappa et al., 2013;Zhang et al., 2014). 203 This evolution through the SOM grid is VOC-specific and defined by six parameters: (P1-P4) 204 yields of the four products that add 1, 2, 3, and 4 oxygen atoms, respectively, without 205 fragmentation; (P5) the probability of fragmentation; and (P6) the decrease in vapor pressure (or 206 volatility) of the species per addition of oxygen atom. Details of the implementation and 207 parameterization of the SOM model in the UCD-CIT are presented in (Jathar et al., 2015). Briefly, six SOM grids with precursor-specific parameter sets were used to represent SOA 208 209 formation from the same precursor classes in the Base model. Parameter sets were separately 210 determined from high NO<sub>x</sub> (low yield) and low NO<sub>x</sub> (high yield) chamber data as the SOM in its 211 current configuration cannot yet account for continuous variation in NO<sub>x</sub>. The SOM parameters 212 were completely determined from explicit fitting to chamber data where the number of fit data 213 points greatly exceeded the number of fitting parameters (6). Thus, the SOM model will be 214 referred to as "constrained" multi-generational oxidation. The SOM parameters and data sources 215 are listed in Table S.3.

The SOM model parameters used in the present study were determined without accounting for losses of vapors to chamber walls, which can lead to a substantial underestimation of the actual SOA formation potential of a given precursor (<u>Matsunaga and Ziemann</u>, 219 <u>2010;Zhang et al., 2014</u>). A companion paper evaluates vapor wall-loss effects on the SOM

- results (<u>Cappa et al., 2015</u>). The SOM parameter fits were derived using dynamic gas-particle
- 221 partitioning assuming an accommodation coefficient of unity, which tends to minimize the
- 222 influence of vapor wall loss (<u>McVay et al., 2014</u>), and thus represents a conservative lower
- bound of SOA formation. The SOM model was additionally extended to consider the influence
- of oligomerization reactions by allowing irreversible conversion of particle-phase SOM species
- into a single non-volatile species using the same  $k_{\text{oligomer}}$  as in the Base model, referred to as
- 226 SOM-OLIG. Oligomerization reactions were added *a posteriori* to the SOM model, i.e.
- 227 oligomerization reactions were not included as part of the data fitting and parameter
- determination and are included in the present study only as a sensitivity case.
- 229

# 230 2.2.4. Cascading Oxidation Model

231 Additional simulations were performed using a contemporary multi-generational 232 oxidation scheme, the Cascading Oxidation Model (COM). The COM builds on the two-product 233 Base model but allows for additional reaction of the semi-volatile products using the scheme of 234 Baek et al. (2011). Briefly, the two semi-volatile products from a given precursor react with OH, 235 with the highest volatility product converted into the lowest volatility product and the lowest 236 volatility product converted to a non-volatile product (see SI Section on Cascading Oxidation 237 Model). Like most other schemes that have thus far been used to represent multi-generational 238 oxidation of SOA from traditional VOCs in 3-D models (Lane et al., 2008), COM does not 239 consider fragmentation reactions, is not fit or constrained to experimental data, and adds these 240 ageing reactions on top of an existing parameterization. The COM model will be referred to as 241 "unconstrained" multi-generational oxidation.

242

| Simulation               | Description   |  |  |  |
|--------------------------|---|--|--|--|
| Base                     | Equivalent to Carlton et al. (2010) without oligomerization       |  |  |  |
| Base-OLIG                | Equivalent to Carlton et al. (2010)                               |  |  |  |
| BaseM (low yield)        | two-product model using new high NO <sub>x</sub> data (low yield) |  |  |  |
| BaseM (high yield)       | two-product model using new low NO <sub>x</sub> data (high yield) |  |  |  |
| SOM (low yield)          | New high NO <sub>x</sub> data, no vapor wall losses               |  |  |  |
| SOM (high yield)         | New low NO <sub>x</sub> data, no vapor wall losses                |  |  |  |
| SOM-OLIG (low yield) and | SOM with inclusion of alignmenization                             |  |  |  |
| SOM-OLIG (high yield)    | SOW with inclusion of ongomenzation                               |  |  |  |

243 Table 1: Simulations performed in this work.

 COM
 Base-OLIG model with added ageing reactions

 244

### 245

# 246 2.3. Simulations

247 Table 1 lists the simulations performed in this work. We performed two simulations with 248 the Base model (with and without oligomerization), two with the BaseM model (low and high 249 yield), four with the SOM model (low and high yield and with oligomerization accounted for) 250 and one with the COM model. These nine simulations were performed for both domains: SoCAB 251 and the eastern US. Simulations were performed for 19 days with the first 5 days used for spin 252 up. For the SoCAB, each simulated day using the SOM required approximately 4 h of elapsed 253 time (on 40 Intel i5-3570 processor cores) so a 19-day episode was simulated in less than 4 days. 254 For the eastern US, each simulated day required approximately 9 h of elapsed time so a 19-day 255 episode was simulated in about 8 days. The SOM simulations were approximately four times 256 slower than the BaseM simulations on account of the large number of model species.

257

#### 258 **3. Results**

# 259 **3.1. Base versus BaseM**

260 Although the main focus of the present study is on understanding the role of multi-261 generational oxidation in SOA models, it is useful to begin by considering differences between 262 the predictions from Base and BaseM (two-product parameters fit to more recent data sets). The 263 14-day averaged, precursor-resolved SOA concentrations at two sites in the SoCAB (Los 264 Angeles: urban, Riverside: urban outflow) and at two sites in the eastern US (Atlanta: urban, 265 Smoky Mountains: remote) from Base and BaseM are compared in Figure 1. Base model 266 predictions of total semi-volatile SOA concentrations (i.e. SOA exclusive of oligomers) at all 267 four sites are similar to the BaseM (low yield) model predictions that were parameterized using 268 high-NO<sub>x</sub> chamber data. This outcome is perhaps not surprising at Los Angeles, Riverside and 269 Atlanta since these urban areas have higher  $NO_x$  levels and, correspondingly, the Base 270 simulations effectively used high-NO<sub>x</sub> parameters. While there are slight increases in SOA from 271 some precursors and decreases from others, BaseM, in comparison to Base, predicted negligible 272 contributions from alkane SOA. The general agreement between Base and BaseM (low yield) in 273 rural/remote areas like the Smoky Mountains (where more than three-quarters of the SOA comes 274 from terpene oxidation) also resulted from increases in SOA from some precursors and decreases

275 from others. These precursor-specific differences are a result of slight differences between the 276 two-product yields for these species in Base (Carlton et al., 2010) and BaseM. The comparison 277 between Base and BaseM suggests that while the newer data might not dramatically affect the 278 SOA concentrations in high-NO<sub>x</sub> (or urban) areas — at least those that still have marginal 279 biogenic contributions — the newer data could increase SOA concentrations (factor of ~2) in 280 low-NO<sub>x</sub> (or rural/remote) areas. One important difference is that the BaseM parameterizations 281 for mono- and sesquiterpenes indicate a NO<sub>x</sub> dependence, whereas the Base parameterizations 282 have no  $NO_x$  dependence for these compounds. This has implications for the assessment of 283 anthropogenic influences on biogenic SOA and whether biogenic SOA can, to some extent, be 284 controlled (Carlton et al., 2007). Further, the substantial decrease in alkane SOA concentrations 285 in BaseM compared to Base suggests that the Base alkane parameterization might be over-286 predicting SOA formation from alkanes, at least those that make up ALK5, making it an even 287 smaller fraction of the total SOA mass.

288

# 289 **3.2. Effect of Constrained Multi-Generational Oxidation**

### **3.2.1. SOA Concentrations**

291 Predictions from BaseM and SOM, which were parameterized using the same data, were 292 used to investigate the influence of multi-generational oxidation. Domain-wide, 14-day averaged 293 SOA concentrations from BaseM and SOM for the SoCAB and for the eastern US, along with 294 the ratio of the SOA concentrations between SOM and BaseM, are shown in Figure 2. The SOA 295 concentrations presented are averages of the low-yield and high-yield simulations. 296 Consideration of either the low-yield or high-yield simulations individually affects the details, 297 but not the general conclusions about multi-generational oxidation below, even though the SOA 298 mass concentrations from the high-yield simulations are typically 2-4 times larger than from the 299 low-yield simulations (see Figure S.2). In both the SoCAB and the eastern US, the predicted 300 spatial distribution of SOA is generally similar between BaseM and SOM, with only minor 301 differences evident in some locations. For the SoCAB, the SOA concentrations in SOM are 302 somewhat lower everywhere compared to BaseM, by 10-20% in the Los Angeles metropolitan 303 area (marked by a red box) and by about 20-30% in regions dominated by biogenic SOA (e.g., 304 Los Padres National Forest located in the northwest corner of the simulated domain). Similarly, 305 the SOM predictions for SOA concentrations in the eastern US are 0-20% lower than BaseM

306 predictions over most of the domain. The urban versus biogenic difference was not evident,

307 probably owing to a coarser grid resolution (36 km for the eastern US versus 8 km for the

- 308 SoCAB). It appears that multi-generational oxidation does not dramatically increase (from
- 309 additional functionalization reactions) or decrease (from additional fragmentation reactions) the
- total SOA concentrations formed from the precursor compounds considered in either region.
- In Figure 1, at all sites, the SOM SOA concentrations are roughly the same or slightly higher than the BaseM SOA concentrations for the low-yield simulations but consistently lower for the high-yield simulations, by 18-25%. When averaged, the SOM SOA concentrations are slightly lower than the BaseM simulations, largely due to the lower predictions of SOA from mono-terpene and sesquiterpenes in the SOM high yield simulations. The low- versus high-yield distinction suggests that the SOM-predicted SOA is probably similar to BaseM-predicted SOA in urban areas (low yield or high NO<sub>x</sub>) but lower in rural/remote areas (high yield or low NO<sub>x</sub>).
- 318 The seemingly limited influence of multi-generational oxidation on total SOA 319 concentrations runs counter to the findings from previous work that suggests multi-generational 320 oxidation is an important source of SOA (Robinson et al., 2007; Murphy and Pandis, 2009; Baek 321 et al., 2011; Fast et al., 2014; Dzepina et al., 2009). However, these previous efforts accounted for 322 multi-generational VOC oxidation by adding ageing reactions for semi-volatile products on top 323 of an existing parameterization, similar to the COM model, and thus may suffer from "double 324 counting" to some extent (we will return to this point later). These results also indicate that the 325 two-product model parameterization inherently captures some of the influence of multi-326 generational oxidation, at least over the timescales and conditions relevant for the SoCAB and 327 the eastern US. This can be understood by considering that, although the two-product model 328 assumes non-reactive products, the chamber-observed SOA formation is dependent on 329 production from all reaction generations, even at short oxidation lifetimes (half to a full day of 330 photochemistry); the extent to which multi-generational oxidation influences the two-product fit 331 parameters will depend on the extent to which later generation products are responsible for the 332 actual SOA formation in a given experiment. In summary, it is possible that the chamber-333 observed SOA formation accounts for the majority of the multi- generational oxidation reactions 334 that contribute to SOA mass and hence, a two-product approach to model SOA formation would 335 already include the mass-enhancement associated with multi-generational oxidation. However,

such a two-product model may not necessarily accurately represent the chemical composition ofSOA

338 The behavior of SOM vs. BaseM predictions is similar in the SoCAB and the eastern US, 339 with minor differences likely related to the size of the domain and the average atmospheric 340 lifetime of the simulated SOA, differences in the evolution of SOA from the various precursors, 341 and the dominance of certain precursors in different domains. These precursor-specific SOA 342 concentrations are visualized in Figure 1 and listed as domain-wide averages in Table S.4. These 343 results indicate that SOM typically produced more SOA from alkanes (although very little 344 overall) but less from terpenes and isoprene in both the SoCAB and the eastern US, compared to 345 BaseM. For aromatics and sesquiterpenes the concentrations are generally similar between the 346 two models, although slightly greater for sesquiterpenes for the eastern US SOM simulations. 347 The use of the SOM model that inherently accounts for multi-generational oxidation leads to 348 more SOA mass for some compounds (due to enhanced functionalization) but less SOA mass for 349 others (due to fragmentation) compared to a static representation of the semi-volatile products. 350 SOA concentrations in chamber photooxidation experiments have been observed to decrease at 351 longer times for some VOCs, notably isoprene (<u>Chhabra et al., 2011</u>) and alpha-pinene (<u>Henry</u> 352 and Donahue, 2012). Such behavior is captured by SOM but not by BaseM, which does not 353 account for fragmentation. Consequently, SOA concentrations in BaseM can never decrease 354 from reactions. The general similarity in the total simulated SOA from BaseM and SOM results 355 in large part from offsetting trends associated with different SOA precursors. This suggests that 356 the use of constrained multi-generational oxidation SOA models, such as SOM, over two-357 product models may help to provide a clearer picture of the sources of SOA in a given region, 358 even if the different modeling approaches lead to similar total SOA mass concentrations. 359 The simulated total OA concentrations (POA+SOA) are compared to ambient OA 360 measurements made at the STN (Speciated Trends Network) and IMPROVE (Interagency 361 Monitoring of Protected Visual Environments) air quality monitoring sites in the SoCAB and the 362 eastern US. (IMPROVE sites tend to be remote and with lower OA concentrations compared to 363 STN sites, which tend to be more urban.) Table 2 lists statistical metrics of fractional bias and 364 fractional error that capture model performance for OA for all simulations for both domains at 365 the STN and IMPROVE sites. The simulated SOA fraction of total OA differs greatly between

the SoCAB (~10%) and the eastern US (~80%). Consequently, changes in the amount of SOA

367 simulated will have a larger influence on the total OA in the eastern US, and thus on the

- 368 comparison with observations. Despite these differences, there is no substantial change in model
- 369 performance between Base, BaseM and SOM in either domain, with all simulations under-
- 370 predicting the total OA. In contrast, COM, which leads to substantial increases in the simulated
- 371 SOA mass concentrations within both domains (see Section 3.3), improved model performance
- at the STN and IMPROVE sites for the SoCAB and at the STN sites for the eastern US.
- 373

374 Table 2: Fractional bias and fractional error at STN and IMPROVE sites for the SoCAB and the

area eastern US for the Base, BaseM (average of low- and high-yield), COM and SOM (average of

low- and high-yield) simulations. Green, yellow, and orange shading represent 'good', 'average'and 'poor' model performance (Boylan and Russell, 2006).

|            |       |       |         |       |            |       | <u> </u> |       |
|------------|-------|-------|---------|-------|------------|-------|----------|-------|
| Simulation | SoCAB |       |         |       | Eastern US |       |          |       |
|            | STN   |       | IMPROVE |       | STN        |       | IMPROVE  |       |
|            | Frac. | Frac. | Frac.   | Frac. | Frac.      | Frac. | Frac.    | Frac. |
|            | Bias  | Error | Bias    | Error | Bias       | Error | Bias     | Error |
| Base       | -62   | 62    | -34     | 43    | -78        | 89    | -10      | 57    |
| BaseM      | -61   | 62    | -30     | 41    | -78        | 87    | -8       | 57    |
| SOM        | -62   | 63    | -33     | 43    | -80        | 89    | -12      | 55    |
| COM        | -28   | 43    | 27      | 50    | 3          | 61    | 85       | 92    |

378

## **379 3.2.2. SOA Volatility**

380 The effective volatility of the SOA was characterized for the Base, BaseM and SOM 381 simulations. SOA volatility influences the sensitivity of the SOA to dilution and temperature 382 changes. Since Base, BaseM and SOM use model species that have very different volatilities, as 383 characterized by the species saturation concentration,  $C^*$ , volatility distributions were developed 384 in which individual species are grouped into logarithmically spaced bins of effective  $C^*$ , referred to as volatility basis set-equivalent (VBS<sub>eq</sub>) distributions (<u>Donahue et al., 2006</u>). In Figure 3(a,c), 385 386 we show the normalized, episode-averaged VBSeq distributions of SOA at Los Angeles and 387 Atlanta for the Base, BaseM and SOM simulations. Qualitatively, the SOA VBS<sub>eq</sub> distributions for Base and BaseM are similar, with the bulk of the gas+particle mass being in the  $C^* = 1$  to 388 1000 µg m<sup>-3</sup> range. In sharp contrast, the SOA volatility distribution for the SOM simulation had 389 a substantial fraction of SOA mass in the  $C^* = 0.0001$  to 1 µg m<sup>-3</sup> range, much lower than the 390 391 Base/BaseM simulations. At atmospherically-relevant OA concentrations (1-10 µg m<sup>-3</sup>), the 392 mass in these low  $C^*$  bins would be exclusively in the particle-phase.

393 It is not possible to compare the simulated volatility distributions to ambient observations 394 since direct measurement of volatility distributions has not been demonstrated for such low  $C^*$ 395 species. However, the effective volatility of SOA particles has been experimentally assessed by 396 considering the response of particles to heating in a thermodenuder (Cappa and Jimenez, 397 2010;Huffman et al., 2009). High volatility species generally evaporate at lower temperatures 398 than low volatility species. The theoretical response of the predicted SOA mass, expressed as the 399 mass fraction remaining (MFR), to heating in a thermodenuder over the range 25 to 105 °C was 400 simulated using the model of Cappa (2010). The  $C^*$  values varied with temperature according to 401 the Clausius-Clapeyron equation and the enthalpy of vaporization was assumed to be a function of  $C^*$  with  $\Delta H_{vap} = 131 - 11 \times log_{10}C^*$ . (See SI section Thermodenuder Model.) We plot the 402 results in Figure 3(b,d). At both Los Angeles and Atlanta, differences in the predicted SOA 403 404 volatility are quite evident. In general, the effective SOA volatility was higher in the Base and 405 BaseM simulations than in the SOM simulations. The SOA from the Base and BaseM 406 simulations is almost entirely evaporated when heated to 70 °C, and some evaporation occurs even at 25 °C as a response to vapor stripping in the denuder. In contrast, the SOA from the 407 408 SOM simulations did not entirely evaporate until 100 °C and exhibits a more gradual decrease 409 with temperature. The SOM-simulated SOA TD evaporation is much more similar to the behavior observed in both laboratory experiments and field assessments of SOA volatility 410 411 (Cappa and Jimenez, 2010;Huffman et al., 2009;Lee et al., 2010). This suggests that SOM is 412 producing SOA with more physically realistic properties even though the Base/BaseM and SOM 413 simulations produced similar SOA concentrations.

414

### 415 **3.2.3. Influence of Oligomerization**

416 The Base-OLIG model includes an oligomerization pathway in which semi-volatile, 417 condensed-phase material is converted to a non-volatile, yet absorptive material on a fixed 418 timescale. This effectively "pumps" semi-volatile vapors to the particle phase and leads to 419 increased SOA concentrations. It has the additional effect of making the SOA less sensitive to 420 dilution and changes in temperature. To examine the influence of oligomerization, Figure 4 421 shows predictions of the precursor-resolved SOA concentrations from the Base, Base-OLIG, 422 SOM and SOM-OLIG simulations for Los Angeles and Riverside, CA. The total SOA 423 concentrations in Base-OLIG are ~60% higher than Base but the SOA concentrations in SOM- 424 OLIG were only ~14% higher than SOM. This difference can be understood through the 425 differences between the SOM and Base volatility distributions for semi-volatile species. For the 426 Base model, a large fraction of the oxidation products have  $C^* > 1 \mu \text{g m}^{-3}$ , and thus a sizable fraction is in the gas-phase. This gas-phase material can be viewed as potential SOA, and as 427 428 oligomers are formed this material is converted to actual SOA. For SOM, much of the material has  $C^* \leq 1 \mu \text{g m}^{-3}$ , and thus most of it is already in the particle phase. Consequently, when it is 429 converted to oligomers only a marginal influence on the total SOA concentration results. 430 431 Overall, it is evident that the influence of oligomerization on simulated SOA concentrations is 432 tightly linked to the semi-volatile product distribution. This may influence the timescales of SOA 433 formation, since in SOM production of lower volatility material is related to the timescales of 434 gas-phase oxidation, whereas in Base, the specified oligomerization rate coefficient, which is 435 largely under-constrained, controls the timescale of low (essentially non-) volatile material. 436

# 437 **3.3.** Comparing Constrained Multi-generational Oxidation to

438 Unconstrained Schemes

439 The 14-day averaged SOA concentrations from the COM, Base, and SOM simulations 440 for the SoCAB and the eastern US are compared in Figure 5. Recall that COM allows for 441 conversion of the semi-volatile products in the Base model to lower-volatility products on top of 442 the original 2-product parameterization. The COM simulations predict a factor of 4 to 8 increase 443 in SOA concentrations over the Base and SOM simulations, attributable to the production of 444 low-volatility and non-volatile SOA from the added oxidation reactions. Because COM, like 445 many ad hoc ageing schemes (Simon and Bhave, 2011; Robinson et al., 2007; Pye and Seinfeld, 446 2010; Baek et al., 2011), lacks fragmentation and adds ageing reactions on top of an existing 447 parameterization, and with sufficient oxidation all semi-volatile products will be converted into 448 non-volatile SOA. This means that the ultimate SOA mass yield is equal to the sum of the mass 449 vields of the individual products, independent of their vapor pressures. Given that SOM 450 inherently accounts for multi-generational oxidation as part of the model parameterization, this 451 comparison clearly suggests that the unconstrained schemes used in the COM simulations form 452 too much SOA and that such schemes are not truly representative of multi-generational oxidation 453 in the atmosphere.

454 Some previous studies have defended the use of a COM-type model because its 455 implementation improved model performance (Lane et al., 2008; Murphy and Pandis, 456 2009; Shrivastava et al., 2008), as was also observed here (Table 2). However, given that COM-457 type models remain generally unconstrained and have been inconsistently applied to different 458 VOC precursor types (e.g. ageing of anthropogenics but not biogenics) (Farina et al., 2010;Lane 459 et al., 2008; Murphy and Pandis, 2009), and since recent testing of a COM-type scheme in the 460 laboratory demonstrated that such schemes do, indeed, lead to over-prediction of SOA mass 461 concentration (Zhao et al., 2015), we suggest that this apparently improved agreement is more 462 likely fortuitous than a true indication of improved representation of atmospheric chemistry. It 463 should be noted that the current study specifically assesses the performance of a COM-type 464 model on the SOA production from traditional VOCs only, exclusive of potential contributions 465 of IVOCs and semi-volatile POA vapors to the SOA burden. Previous studies that have 466 examined the influence of multi-generational oxidation of traditional VOCs using COM-type 467 models have typically combined the effects of VOC ageing and IVOC and POA vapor oxidation 468 (e.g. Murphy and Pandis (2009); Jathar et al. (2011)) together and have not investigated the role 469 of each process separately. Consequently, our results, which isolate the influence of using a 470 COM-type oxidation scheme, suggest COM-type models may be inappropriate for use in 471 regional air quality models even though they can lead to improved model/measurement 472 comparison (Table 2). They also imply that models that employed COM-like schemes have 473 potentially underplayed the role of other important OA formation pathways such as aqueous 474 (aerosol, fog, cloud) processing of water-soluble organics (Ervens et al., 2011) and particle-475 surface reactions (Liggio et al., 2005; Shiraiwa et al., 2013). Future work to integrate semi-476 volatile POA treatments with constrained multi-generational ageing schemes like SOM is 477 needed.

478

## 479 **4. Discussion**

When constrained using the same chamber data, the BaseM (traditional two-product model that does not resolve multi-generational oxidation) and SOM models predict roughly the same SOA mass concentrations and spatial distribution for regional air pollution episodes in the SoCAB and the eastern US. This suggests that the chamber data used to constrain the BaseM and SOM parameterizations presumably already includes a majority of the SOA mass that would be 485 attributable to multi-generational oxidation. The extent to which multi-generational oxidation 486 influences the production of SOA in a given chamber experiment depends on both the volatility 487 and reactivity of the first-generation products and the time-scale of the experiment (Wilson et al., 488 2012). If SOA formation is dominated by first-generation products, then explicit accounting for 489 multi-generational ageing will not be important. Alternatively, if most SOA is formed from 490 second-generation products with little direct contribution from first-generation products, than a 491 static representation (such as with the 2-product model) might be sufficient even when multi-492 generational ageing is, in fact, dominant. But if SOA formation is balanced between 493 contributions from first, second and later generation products, then the extent to which a static 494 representation will capture the influence of multi-generational ageing may be highly variable and 495 sensitive to the experimental conditions and number of oxidation lifetimes. Consequently, the 496 appropriateness of extrapolating static model parameterizations to longer (global atmospheric) 497 timescales remains unclear. The results presented here indicate that the 2-product model does 498 capture the influence of multi-generational ageing as part of the parameterization in terms of 499 mass concentration, at least for the regional episodes considered, but it is also apparent that the 500 simulated SOA properties (e.g. volatility) and the explicit contributions of various SOA types are 501 not fully captured by such simple models.

502 The BaseM and SOM simulations show that the SOA concentrations in the SoCAB and 503 eastern US vary by a factor of two when using parameterizations developed from low vs. high 504  $NO_x$  chamber experiments. Hence, we can argue that for the present simulations  $NO_x$ 505 dependence is a much more important factor for SOA production than multi-generational 506 oxidation. While most 3-D models include schemes to simulate the NO<sub>x</sub> dependence of SOA 507 formation, these schemes remain *ad hoc* as they are based on limited experimental measurements 508 and also rely on the ability of the model to accurately predict radical concentrations (RO<sub>2</sub>, HO<sub>2</sub>) 509 or VOC-to-NO<sub>x</sub> ratios. In this work, the model predictions from the low- and high-yield 510 simulations bound the NO<sub>x</sub>-dependent uncertainty in SOA concentrations and we recommend 511 that future work examine this issue in much more detail.

512 SOM predicts a modestly different composition of SOA than BaseM despite similar total 513 mass concentrations of SOA. The composition predicted by SOM has a slightly higher 514 contribution from alkanes, aromatics (anthropogenic) and sesquiterpenes and a lower 515 contribution from isoprene and monoterpenes. These modest differences in the predicted 516 composition of SOA have implications for understanding the sources of ambient aerosol and 517 eventually the regulation of these sources to achieve compliance with National Ambient Air 518 Quality Standards (NAAQS). These more accurate SOA predictions resolved by chemical 519 families should be tested in epidemiological studies to determine if they are associated with 520 adverse health effects. Additionally, SOM predicted a much lower-volatility SOA than BaseM, 521 and SOM predictions are in better qualitative agreement with ambient thermodenuder 522 measurements of OA volatility. Since the SOA has a much lower volatility, there is very little 523 enhancement (10-15%) with the inclusion of oligomerization reactions, implying that while 524 oligomerization might affect composition, it may not be a source of additional SOA formation as 525 the Base model suggests.

526 In this work, we consider POA as non-volatile and non-reactive and do not consider SOA contributions from IVOCs or semi-volatile POA vapors. Oxidation of IVOCs and semi-volatile 527 528 POA vapors (i.e. SVOCs) can lead to the production of new SOA mass, but evaporation of POA 529 leads to a decrease in the total OA mass. To some extent, these effects are offsetting (especially 530 for SVOCs, which do not contribute new carbon mass to a model). To the extent that the loss of 531 POA is balanced exactly by the formation of SOA from IVOCs and 'recycling' of semi-volatile 532 POA vapors, the simulations here represent a scenario in which the total OA mass is conserved, 533 although possibly with the wrong spatial distribution (Robinson et al., 2007). Most efforts to 534 incorporate SOA formation from IVOCs and SVOCs have simulated their oxidation using a version of the VBS model in which multi-generational ageing is implicit, but highly 535 536 underconstrained and structured in such a way that the ultimate (long time) SOA yield is greater 537 than unity because all mass is converted to low-volatility products and oxygen addition is 538 assumed. The SOM framework provides a way to explicitly account for the influence of multi-539 generational chemistry in SOA formation experiments that include semi-volatile POA vapors 540 and IVOCs (Gordon et al., 2014a; Gordon et al., 2014b; Gordon et al., 2013; Grieshop et al., 541 2009a; Grieshop et al., 2009b; Hennigan et al., 2011; Miracolo et al., 2011; Miracolo et al., 542 2012;Platt et al., 2013;Platt et al., 2014;Nordin et al., 2013;Chirico et al., 2010;Heringa et al., 543 <u>2011; Tkacik et al., 2014</u>), and thus should be useful for constraining the contribution of these 544 compound classes to the ambient OA budget. In addition, the simulations here do not consider 545 the influence of vapor wall losses on SOA formation. Such losses can influence SOA yields in 546 chambers, and consequently the parameterizations that result from fitting of such chamber data.

- 547 The influence of vapor wall losses on simulated ambient SOA and OA concentrations within the
- 548 SOM framework is examined in a companion paper (<u>Cappa et al., 2015</u>). Ultimately, models like
- the SOM can be applied to chamber experiments to better understand the role and contribution of
- 550 POA, IVOCs and vapor wall-losses to total OA.
- 551 Finally, the comparison between the constrained SOM and the unconstrained COM
- 552 (commonly used in large-scale models) suggests that COM may be double counting SOA
- 553 formation. These simple ageing schemes should be refit to chamber data where all parameters
- can be matched to observed trends in a self-consistent manner.
- 555

# 556 **5.** Acknowledgements

- 557 This work was supported by the California Air Resources Board (CARB) under contracts
- 558 11-755 and 12-312. Although this work was funded by the CARB, the statements and
- conclusions are those of the authors and not necessarily those of the CARB.
- 560

# 561 **6.** References

Ahmadov, R., McKeen, S. A., Robinson, A. L., Bahreini, R., Middlebrook, A. M., de Gouw, J.
A., Meagher, J., Hsie, E. Y., Edgerton, E., Shaw, S., and Trainer, M.: A volatility basis set model
for summertime secondary organic aerosols over the eastern United States in 2006, Journal Of
Geophysical Research-Atmospheres, 117, D06301, 2012.

Baek, J., Hu, Y., Odman, M. T., and Russell, A. G.: Modeling secondary organic aerosol in
CMAQ using multigenerational oxidation of semi-volatile organic compounds, Journal of
Geophysical Research: Atmospheres, 116, D22204, 2011.

Bernstein, J. A., Alexis, N., Barnes, C., Bernstein, I. L., Bernstein, J. A., Nel, A., Peden, D.,
Diaz-Sanchez, D., Tarlo, S. M., and Williams, P. B.: Health effects of air pollution, The Journal
of Allergy and Clinical Immunology, 114, 1116-1123, 2004.

Boylan, J. W., and Russell, A. G.: PM and light extinction model performance metrics, goals,
and criteria for three-dimensional air quality models, Atmospheric Environment, 40, 4946-4959,
2006.

Cappa, C., and Wilson, K.: Multi-generation gas-phase oxidation, equilibrium partitioning, and
the formation and evolution of secondary organic aerosol, Atmospheric Chemistry and Physics,
12, 9505-9528, 2012.

578 Cappa, C. D.: A model of aerosol evaporation kinetics in a thermodenuder, Atmospheric
579 Measurement Techniques, 3, 579-592, 2010.

- 580 Cappa, C. D., and Jimenez, J. L.: Quantitative estimates of the volatility of ambient organic 581 aerosol, Atmospheric Chemistry & Physics, 10, 5401-5924, doi:10.5194/acp-10-5409-2010,
- 581 acrossit, Autospheric Chemistry & Thysics, 10, 5401-5724, doi:10.5174/acp-10-5407-2010, 582 2010.
- Cappa, C. D., Zhang, X., Loza, C. L., Craven, J. S., Yee, L. D., and Seinfeld, J. H.: Application
  of the Statistical Oxidation Model (SOM) to Secondary Organic Aerosol formation from
  photooxidation of C 12 alkanes, Atmospheric Chemistry and Physics, 13, 1591-1606, 2013.
- Cappa, C. D., Jathar, S. H., Wexler, A. S., Seinfeld, J., and Kleeman, M. J.: Simulating
  secondary organic aerosol in a regional air quality model using the statistical oxidation model –
  Part 2: Assessing the influence of vapor wall losses, Atmospheric Chemistry & Physics
  Discussions, 15, 30081-30126, 2015.
- 590 Carlton, A. G., Pinder, R. W., Bhave, P. V., and Pouliot, G. A.: To what extent can biogenic 591 SOA be controlled?, Environmental Science & Technology, 44, 3376-3380, 2007.
- 592 Carlton, A. G., Bhave, P. V., Napelenok, S. L., Edney, E. O., Sarwar, G., Pinder, R. W., Pouliot,
- 593 G. A., and Houyoux, M.: Model representation of secondary organic aerosol in CMAQv4. 7,
- 594 Environmental Science & Technology, 44, 8553-8560, 2010.
- 595 Carter, W. P., and Heo, G.: Development of revised SAPRC aromatics mechanisms, 596 Atmospheric Environment, 77, 404-414, 2013.
- 597 Chacon-Madrid, H., and Donahue, N.: Fragmentation vs. functionalization: chemical aging and 598 organic aerosol formation, Atmos. Chem. Phys, 11, 10553-10563, doi:10.5194/acp-11-10553-599 2011 2011.
- Chacon-Madrid, H. J., Presto, A. A., and Donahue, N. M.: Functionalization vs. fragmentation:
  n-aldehyde oxidation mechanisms and secondary organic aerosol formation, Phys. Chem. Chem.
  Phys., 12, 13975–13982, doi: 10.1039/c0cp00200c, 2010.
- 603 Chacon-Madrid, H. J., Henry, K. M., and Donahue, N. M.: Photo-oxidation of pinonaldehyde at 604 low NO x: from chemistry to organic aerosol formation, Atmospheric Chemistry and Physics, 605 13, 3227-3236, 2013.
- 606 Chen, J., Mao, H., Talbot, R. W., and Griffin, R. J.: Application of the CACM and MPMPO
  607 modules using the CMAQ model for the eastern United States, Journal of Geophysical Research:
  608 Atmospheres (1984–2012), 111, 2006.
- Chhabra, P. S., Ng, N. L., Canagaratna, M. R., Corrigan, A. L., Russell, L. M., Worsnop, D. R.,
  Flagan, R. C., and Seinfeld, J. H.: Elemental composition and oxidation of chamber organic
  aerosol, Atmospheric Chemistry and Physics, 11, 8827-8845, 2011.
- 612 Chirico, R., DeCarlo, P., Heringa, M., Tritscher, T., Richter, R., Prevot, A., Dommen, J., 613 Weingartner, E., and Wehrle, G.: Impact of aftertreatment devices on primary emissions and 614 secondary organic aerosol formation potential from in-use diesel vehicles: results from smog 615 chamber experiments, Atmospheric Chemistry and Physics, 10, 11545-11563, 2010.

- Donahue, N., Robinson, A., Stanier, C., and Pandis, S.: Coupled partitioning, dilution, and
  chemical aging of semivolatile organics, Environ. Sci. Technol, 40, 2635-2643,
  doi:10.1021/es052297c, 2006.
- Donahue, N. M., Henry, K. M., Mentel, T. F., Kiendler-Scharr, A., Spindler, C., Bohn, B.,
  Brauers, T., Dorn, H. P., Fuchs, H., and Tillmann, R.: Aging of biogenic secondary organic
  aerosol via gas-phase OH radical reactions, Proceedings of the National Academy of Sciences,
  109, 13503-13508, 2012.
- Dzepina, K., Volkamer, R., Madronich, S., Tulet, P., Ulbrich, I., Zhang, Q., Cappa, C., Ziemann,
  P., and Jimenez, J.: Evaluation of recently-proposed secondary organic aerosol models for a case
  study in Mexico City, Atmospheric Chemistry and Physics, 9, 5681-5709, doi:10.5194/acp-95681-2009, 2009.
- Emmons, L., Walters, S., Hess, P., Lamarque, J.-F., Pfister, G., Fillmore, D., Granier, C.,
  Guenther, A., Kinnison, D., and Laepple, T.: Description and evaluation of the Model for Ozone
  and Related chemical Tracers, version 4 (MOZART-4), Geoscientific Model Development, 3,
  43-67, 2010.
- Epstein, S. A., Riipinen, I., and Donahue, N. M.: A semiempirical correlation between enthalpy
  of vaporization and saturation concentration for organic aerosol., Environmental science &
  technology, 44, 743-748, 2010.
- Ervens, B., Turpin, B., and Weber, R.: Secondary organic aerosol formation in cloud droplets
  and aqueous particles (aqSOA): a review of laboratory, field and model studies, Atmospheric
  Chemistry and Physics, 11, 11069-11102, 2011.
- Farina, S. C., Adams, P. J., and Pandis, S. N.: Modeling global secondary organic aerosol
  formation and processing with the volatility basis set: Implications for anthropogenic secondary
  organic aerosol, Journal of Geophysical Research, 115, D09202, doi:10.1029/2009JD013046,
  2010.
- 641 Fast, J. D., Allan, J., Bahreini, R., Craven, J., Emmons, L., Ferrare, R. A., Hayes, P. L., Hodzic, 642 A., Holloway, J., Hostetler, C. A., Jimenez, J. L., Jonsson, S., Liu, Y., Metcalf, C., Middlebrook, 643 A. M., Novak, J., Pekour, M., Perring, A. E., Russell, L., Sedlacek, A., Seinfeld, J., Setyan, A., Shilling, J., Shrivastava, M., Springston, S., Song, C., Subramanian, R., Taylor, J. W., Vinoj, V., 644 Yang, Q., Zaveri, R. A., and Zhang, Q.: Modeling regional aerosol and aerosol precursor 645 646 variability over California and its sensitivity to emissions and long-range transport during the 647 2010 CalNex and CARES campaigns, Atmospheric Chemistry and Physics, 14, 10013-10060, 648 doi:10.5194/acp-14-10013-2014, 2014.
- Goldstein, A. H., and Galbally, I. E.: Known and unexplored organic constituents in the earth's
  atmosphere, Environmental Science & Technology, 41, 1514-1521, doi:10.1021/es072476p,
  2007.
- Gordon, T. D., Tkacik, D. S., Presto, A. A., Zhang, M., Jathar, S. H., Nguyen, N. T., Massetti, J.,
  Truong, T., Cicero-Fernandez, P., Maddox, C., Rieger, P., Chattopadhyay, S., Maldonado, H.,
  Maricq, M. M., and Robinson, A. L.: Primary Gas- and Particle-Phase Emissions and Secondary

Organic Aerosol Production from Gasoline and Diesel Off-Road Engines, Environmental 655 656 Science & Technology, 47, 14137-14146, 2013.

657 Gordon, T. D., Nguyen, N. T., May, A. A., Presto, A. A., Lipsky, E. M., Maldonado, S., 658 Chattopadhyay, S., Gutierrez, A., Maricq, M., and Robinson, A. L.: Secondary Organic Aerosol 659 Formed from Light Duty Gasoline Vehicle Exhaust Dominates Primary Particulate Matter Emissions, Atmospheric Chemistry & Physics, 14, 4461-4678, doi:10.5194/acpd-13-23173-660 661 2013, 2014a.

- 662 Gordon, T. D., Nguyen, N. T., Presto, A. A., Lipsky, E. M., Maldonado, S., Maricq, M., and Robinson, A. L.: Secondary organic aerosol production from diesel vehicle exhaust: impact of 663 aftertreatment, fuel chemistry and driving cycle, Atmospheric Chemistry & Physics, 14, 4643-664 4659, doi:10.5194/acp-14-4643-2014, 2014b. 665
- 666 Grieshop, A., Donahue, N., and Robinson, A.: Laboratory investigation of photochemical oxidation of organic aerosol from wood fires 2: analysis of aerosol mass spectrometer data, 667 Atmospheric Chemistry and Physics, 9, 2227-2240, 2009a. 668
- 669 Grieshop, A. P., Logue, J. M., Donahue, N. M., and Robinson, A. L.: Laboratory investigation of 670 photochemical oxidation of organic aerosol from wood fires 1: measurement and simulation of 671 organic aerosol evolution, Atmospheric Chemistry and Physics, 9, 1263-1277, 10.5194/acp-9-672 1263-2009, 2009b.
- 673 Guenther, A., Karl, T., Harley, P., Wiedinmyer, C., Palmer, P., and Geron, C.: Estimates of global terrestrial isoprene emissions using MEGAN (Model of Emissions of Gases and Aerosols 674 675 from Nature), Atmospheric Chemistry & Physics, 6, 3181-3210, 2006.
- 676 Hennigan, C. J., Miracolo, M. A., Engelhart, G. J., May, A. A., Presto, A. A., Lee, T., Sullivan, 677 A. P., McMeeking, G. R., Coe, H., Wold, C. E., Hao, W. M., Gilman, J. B., Kuster, W. C., 678 Gouw, J., Schichtel, B. A., Collett, J. L., Kreidenweis, S. M., and Robinson, A. L.: Chemical and 679 physical transformations of organic aerosol from the photo-oxidation of open biomass burning 680 emissions in an environmental chamber, Atmospheric Chemistry and Physics, 11, 7669-7686, 681 doi:10.5194/acp-11-7669-2011 2011.
- 682 Henry, K. M., and Donahue, N. M.: Photochemical aging of α-pinene secondary organic aerosol: effects of OH radical sources and photolysis, The Journal of Physical Chemistry A, 116, 5932-683 684 5940, 2012.
- 685 Henze, D., Seinfeld, J., Ng, N., Kroll, J., Fu, T., Jacob, D., and Heald, C.: Global modeling of secondary organic aerosol formation from aromatic hydrocarbons: high- vs. low-yield pathways, 686 687 Atmospheric Chemistry and Physics, 8, 2405--2420, 2008.
- 688 Heringa, M., DeCarlo, P., Chirico, R., Tritscher, T., Dommen, J., Weingartner, E., Richter, R., 689 Wehrle, G., Prevot, A., and Baltensperger, U.: Investigations of primary and secondary 690 particulate matter of different wood combustion appliances with a high-resolution time-of-flight 691
- aerosol mass spectrometer, Atmospheric Chemistry and Physics, 11, 5945-5957, 2011.

- Huffman, J., Docherty, K., Mohr, C., Cubison, M., Ulbrich, I., Ziemann, P., Onasch, T., and Jimenez, J.: Chemically-resolved volatility measurements of organic aerosol fom different
- 694 sources., Environmental science & technology, 43, 5351–5357, 2009.
- 695 IPCC: Climate change 2007: The physical science basis, Agenda, 6, 07, 2007.
- 696 Jathar, S., Farina, S., Robinson, A., and Adams, P.: The influence of semi-volatile and reactive
- 697 primary emissions on the abundance and properties of global organic aerosol, Atmospheric
- 698 Chemistry and Physics, 11, 7727-7746, doi:10.5194/acp-11-7727-2011 2011.
- 699 Jathar, S. H., Cappa, C. D., Wexler, A. S., Seinfeld, J. H., and Kleeman, M. J.: Multi-
- generational oxidation model to simulate secondary organic aerosol in a 3-D air quality model,
- 701 Geosci. Model Dev., 8, 2553-2567, 10.5194/gmd-8-2553-2015, 2015.
- Jathar, S. H., Cappa, C. D., Wexler, A. S., Seinfeld, J. H., and Kleeman, M. J.: Multi generational Oxidation Model to Simulate Secondary Organic Aerosol in a 3D Air Quality
   Model, Geophysical Model Development, accepted for publication.
- Jathar, S. H., Mahmud, A., Barsanti, K. C., Asher, W., Pankow, J. F., and Kleeman, M. J.: Water
   uptake and its influence on gas/particle partitioning of secondary organic aerosol in the United
   States, Atmospheric Environment, submitted.
- Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H.,
  DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. M.,
  Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. A., Hueglin,
  C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara, P., Ehn, M.,
  Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J., E, Dunlea, J., Huffman, J. A.,
- 713 Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo, Y., Schneider, J., Drewnick, F.,
- 714 Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D., Cottrell, L., Griffin, R., Takami, A.,
- 715 Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J. Y., Zhang, Y. M., Dzepina, K., Kimmel, J.
- 716 R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M., Williams, L. R., Wood, E. C.,
- 717 Middlebrook, A. M., Kolb, C. E., Baltensperger, U., and Worsnop, D. R.: Evolution of Organic
- 718 Aerosols in the Atmosphere, Science, 326, 1525-1529, 10.1126/science.1180353, 2009.
- Johnson, D., Utembe, S. R., Jenkin, M. E., Derwent, R. G., Hayman, G. D., Alfarra, M. R., Coe,
  H., and McFiggans, G.: Simulating regional scale secondary organic aerosol formation during
  the TORCH 2003 campaign in the southern UK, Atmospheric Chemistry and Physics, 6, 403-
- 722 418, 2006.
  - Kalberer, M., Paulsen, D., Sax, M., Steinbacher, M., Dommen, J., Prevot, A. S. H., Fisseha, R.,
    Weingartner, E., Frankevich, V., Zenobi, R., and Baltensperger, U.: Identification of polymers as
    major components of atmospheric organic aerosols, Science, 303, 1659-1662,
    10.1126/science.1092185, 2004.
  - Kleeman, M. J., and Cass, G. R.: A 3D Eulerian source-oriented model for an externally mixed
    aerosol, Environmental Science & Technology, 35, 4834-4848, 2001.

Koo, B., Knipping, E., and Yarwood, G.: 1.5-Dimensional volatility basis set approach for modeling organic aerosol in CAMx and CMAQ, Atmospheric Environment, 95, 158-164, 2014.

Lambe, A. T., Onasch, T. B., Croasdale, D. R., Wright, J. P., Martin, A. T., Franklin, J. P.,
Massoli, P., Kroll, J. H., Canagaratna, M. R., and Brune, W. H.: Transitions from
functionalization to fragmentation reactions of laboratory secondary organic aerosol (SOA)
generated from the OH oxidation of alkane precursors, Environmental science & technology, 46,
5430-5437, 2012.

- Lane, T. E., Donahue, N. M., and Pandis, S. N.: Simulating secondary organic aerosol formation
  using the volatility basis-set approach in a chemical transport model, Atmospheric Environment,
  42, 7439-7451, 2008.
- Lee-Taylor, J., Madronich, S., Aumont, B., Baker, A., Camredon, M., Hodzic, A., Tyndall, G. S.,
  Apel, E., and Zaveri, R. A.: Explicit modeling of organic chemistry and secondary organic
  aerosol partitioning for Mexico City and its outflow plume, Atmospheric Chemistry and Physics,
  11, 13219-13241, 2011.
- Lee, B. H., Kostenidou, E., Hildebrandt, L., Riipinen, I., Engelhart, G. J., Mohr, C., DeCarlo, P.
  F., Mihalopoulos, N., Prevot, A. S. H., and Baltensperger, U.: Measurement of the ambient
  organic aerosol volatility distribution: application during the Finokalia Aerosol Measurement
  Experiment (FAME-2008), Atmospheric Chemistry and Physics, 10, 12149-12160, 2010.
- Liggio, J., Li, S. M., and McLaren, R.: Reactive uptake of glyoxal by particulate matter, Journal
  of Geophysical Research: Atmospheres (1984–2012), 110, D10304, doi:10.1029/2004JD005113,
  2005.
- Lin, G., Penner, J. E., Sillman, S., Taraborrelli, D., and Lelieveld, J.: Global modeling of SOA formation from dicarbonyls, epoxides, organic nitrates and peroxides, Atmospheric Chemistry and Physics, 12, 4743-4774, 2012.
- Loza, C. L., Craven, J. S., Yee, L. D., Coggon, M. M., Schwantes, R. H., Shiraiwa, M., Zhang,
  X., Schilling, K. A., Ng, N. L., Canagaratna, M. R., Ziemann, P. J., Flagan, R. C., and Seinfeld,
  J. H.: Secondary organic aerosol yields of 12-carbon alkanes, Atmospheric Chemistry and
  Physics, 14, 1423-1439, 2014.
- Matsunaga, A., and Ziemann, P. J.: Gas-wall partitioning of organic compounds in a Teflon film
   chamber and potential effects on reaction product and aerosol yield measurements, Aerosol
   Science and Technology, 44, 881-892, 2010.
- McVay, R. C., Cappa, C. D., and Seinfeld, J. H.: Vapor–Wall Deposition in Chambers:
   Theoretical Considerations, Environmental science & technology, 48, 10251-10258, 2014.
- Miracolo, M., Hennigan, C., Ranjan, M., Nguyen, N., Gordon, T., Lipsky, E., Presto, A.,
  Donahue, N., and Robinson, A.: Secondary aerosol formation from photochemical aging of
  aircraft exhaust in a smog chamber, Atmos. Chem. Phys, 11, 4135-4147, doi:10.5194/acp-114135-2011, 2011.

- Miracolo, M. A., Drozd, G. T., Jathar, S., Presto, A. A., Lipsky, E., Corporan, E., and Robinson,
  A.: Fuel composition and secondary organic aerosol formation: gas-turbine exhaust and
  alternative aviation fuels, Environmental Science & Technology, 46, 8493-8501,
  doi:10.1021/es300350c, 2012.
- Murphy, B., and Pandis, S.: Simulating the formation of semivolatile primary and secondary
  organic aerosol in a regional chemical transport model., Environmental science & technology,
  43, 4722-4728, doi:10.1021/es803168a, 2009.
- Nordin, E. Z., Eriksson, A. C., Roldin, P., Nilsson, P. T., Carlsson, J. E., Kajos, M. K., Hellén,
  H., Wittbom, C., Rissler, J., Löndahl, J., Swietlicki, E., Svenningson, B., Bohgard, M., Kulmala,
  M., Hallquest, M., and Pagels, J. H.: Secondary organic aerosol formation from idling gasoline
  passenger vehicle emissions investigated in a smog chamber, Atmospheric Chemistry and
  Physics, 13, 6101-6116, 2013.
- Odum, J. R., Hoffmann, T., Bowman, F., Collins, D., Flagan, R. C., and Seinfeld, J. H.:
  Gas/particle partitioning and secondary organic aerosol yields, Environmental Science &
  Technology, 30, 2580-2585, 1996.
- Platt, S. M., El Haddad, I., Zardini, A. A., Clairotte, M., Astorga, C., Wolf, R., Slowik, J. G.,
  Temime-Roussel, B., Marchand, N., Ježek, I., Drinovec, L., Močnik, G., Möhler, O., Richter, M.,
  Barmet, P., Bianchi, F., Baltensperger, U., and Prévôt, A. S. H.: Secondary organic aerosol
  formation from gasoline vehicle emissions in a new mobile environmental reaction chamber,
  Atmospheric Chemistry and Physics, 13, 9141-9158, 2013.
- Platt, S. M., Haddad, I. E., Pieber, S. M., Huang, R. J., Zardini, A. A., Clairotte, M., SuarezBertoa, R., Barmet, P., Pfaffenberger, L., Wolf, R., Slowik, J. G., Fuller, S. J., Kalberer, M.,
  Chirico, R., dommen, J., astorga, C., Zimmermann, R., Marchand, N., Hellebust, S., TemimeRoussel, B., Baltensperger, U., and Prévôt, A. S. H.: Two-stroke scooters are a dominant source
  of air pollution in many cities, Nature communications, 5, 2014.
- Pye, H., and Seinfeld, J.: A global perspective on aerosol from low-volatility organic
  compounds, Atmos. Chem. Phys, 10, 4377-4401, doi:10.5194/acp-10-4377-2010, 2010.
- Robinson, A. L., Donahue, N. M., Shrivastava, M. K., Weitkamp, E. A., Sage, A. M., Grieshop,
  A. P., Lane, T. E., Pierce, J. R., and Pandis, S. N.: Rethinking organic aerosols: Semivolatile
  emissions and photochemical aging, Science, 315, 1259-1262, 2007.
- 796 Shiraiwa, M., Yee, L. D., Schilling, K. A., Loza, C. L., Craven, J. S., Zuend, A., Ziemann, P. J.,
- and Seinfeld, J. H.: Size distribution dynamics reveal particle-phase chemistry in organic aerosol
- formation, Proceedings of the National Academy of Sciences, 110, 11746-11750, 2013.
- Shrivastava, M. K., Lane, T. E., Donahue, N. M., Pandis, S. N., and Robinson, A. L.: Effects of
  gas particle partitioning and aging of primary emissions on urban and regional organic aerosol
  concentrations, Journal of Geophysical Research-Atmospheres, 113, D18301,
  doi:10.1029/2007JD009735, 2008.

803 Simon, H., and Bhave, P. V.: Simulating the degree of oxidation in atmospheric organic 804 particles, Environmental science & technology, 46, 331-339, 2011.

Surratt, J. D., Lewandowski, M., Offenberg, J. H., Jaoui, M., Kleindienst, T. E., Edney, E. O.,
and Seinfeld, J. H.: Effect of acidity on secondary organic aerosol formation from isoprene,
Environmental science & technology, 41, 5363-5369, 2007.

- 808 Tkacik, D. S., Lambe, A. T., Jathar, S., Li, X., Presto, A. A., Zhao, Y., Blake, D., Meinardi, S.,
- 809 Jayne, J. T., Croteau, P. L., and Robinson, A. L.: Secondary Organic Aerosol Formation from in-
- 810 Use Motor Vehicle Emissions Using a Potential Aerosol Mass Reactor, Environmental Science
- 811 & Technology, 48, 11235-11242, 10.1021/es502239v, 2014.
- 812 Tsimpidi, A., Karydis, V., Zavala, M., Lei, W., Molina, L., Ulbrich, I., Jimenez, J., and Pandis,
- 813 S.: Evaluation of the volatility basis-set approach for the simulation of organic aerosol formation
- 814 in the Mexico City metropolitan area, Atmos. Chem. Phys, 10, 525-546, doi:10.5194/acp-10-
- 815 525-2010, 2009.
- 816 Utembe, S. R., Cooke, M. C., Archibald, A. T., Shallcross, D. E., Derwent, R. G., and Jenkin, M.
- 817 E.: Simulating secondary organic aerosol in a 3-D Lagrangian chemistry transport model using
- 818 the reduced Common Representative Intermediates mechanism (CRI v2-R5), Atmospheric
- 819 Environment, 45, 1604-1614, 2011.
- Wiedinmyer, C., Akagi, S., Yokelson, R. J., Emmons, L., Al-Saadi, J., Orlando, J., and Soja, A.:
  The Fire INventory from NCAR (FINN): A high resolution global model to estimate the
  emissions from open burning, Geoscientific Model Development, 4, 625, 2011.
- emissions from open burning, Geoscientific Model Development, 4, 625, 2011.
- Wilson, K. R., Smith, J. D., Kessler, S. H., and Kroll, J. H.: The statistical evolution of multiple generations of oxidation products in the photochemical aging of chemically reduced organic aerosol, Physical Chemistry Chemical Physics, 14, 1468-1479, 2012.
- Yee, L. D., Kautzman, K. E., Loza, C. L., Schilling, K. A., Coggon, M. M., Chhabra, P. S.,
  Chan, M. N., Chan, A. W. H., Hersey, S. P., and Crounse, J. D.: Secondary organic aerosol
  formation from biomass burning intermediates: phenol and methoxyphenols, Atmospheric
  Chemistry and Physics, 13, 8019-8043, 2013.
- Ying, Q., and Li, J.: Implementation and initial application of the near-explicit Master Chemical
  Mechanism in the 3D Community Multiscale Air Quality (CMAQ) model, Atmospheric
  Environment, 45, 3244-3256, 2011.
- Zhang, X., Cappa, C. D., Jathar, S. H., McVay, R. C., Ensberg, J. J., Kleeman, M. J., and
  Seinfeld, J. H.: Influence of vapor wall loss in laboratory chambers on yields of secondary
  organic aerosol, Proceedings of the National Academy of Sciences, 111, 5802-5807, 2014.
- Zhao, B., Wang, S., Donahue, N. M., Chuang, W., Hildebrandt Ruiz, L., Ng, N. L., Wang, Y.,
  and Hao, J.: Evaluation of One-Dimensional and Two-Dimensional Volatility Basis Sets in
  Simulating the Aging of Secondary Organic Aerosol with Smog-Chamber Experiments,
  Environmental Science & Technology, 49, 2245-2254, 10.1021/es5048914, 2015.
- 840

- 842 **Figure Captions**
- 843

844 Figure 1: 14-day averaged SOA concentrations at Los Angeles (a), Riverside (b), Atlanta (c) and 845 Smoky Mountains (d) for the Base, BaseM, and SOM simulations resolved by the precursor/ 846 pathway.

847

848 Figure 2: (a-b) 14-day averaged SOA concentrations in SoCAB for the BaseM and SOM 849 simulations. (c) Ratio of the 14-day averaged SOA concentration from the SOM simulation to 850 that from the BaseM simulation. The BaseM and SOM results are averages of the low yield and 851 high yield simulations. Red box indicates urban areas surrounding Los Angeles.

852

853 Figure 3: Volatility distributions of the 14-day averaged gas+particle SOA mass at Los Angeles 854 (a) and Atlanta (c) for the Base, BaseM and SOM simulations. Thermograms that capture the 855 volatility of the 14- day averaged gas+particle SOA mass at Los Angeles (b) and Atlanta (d) for 856 the Base, BaseM and SOM simulations.

857

858 Figure 4: 14-day averaged SOA concentrations at (a) Los Angeles and (b) Riverside for the Base, 859 Base- OLIG, SOM, SOM-OLIG simulations resolved by the precursor/pathway.

860

861 Figure 5: 14-day averaged SOA concentrations in SoCAB (a-c) and the eastern US (d-f) for the 862 Base, COM and SOM simulations. The SOM results are averages of the low-yield and high-yield

- 863 simulations.
- 864

865 Figure S.1: Schematic illustrating the differences between some of the different ways of

866 modeling SOA. From top to bottom: the 2-product model; the COM-type model, i.e. 2-product

867 with ageing; the VBS as applied to VOCs with no ageing; the VBS as applied to VOCs with

868 additional ageing; the VBS as applied to S/IVOCs; and the SOM. The black arrows indicate the

869 production of products directly from the parent VOC and the orange arrows indicate ageing

870 reactions, i.e. reactions involving product species. For the SOM, all species are reactive and both

871 functionalization and fragmentation are possible. In the other models that include ageing, only functionalization reactions are included, i.e. reactions that decrease compound vapor pressures.

872 873

874 Figure S.2:14-day averaged SOA concentrations in SoCAB for the BaseM and SOM simulations 875 for the low-yield and high-yield parameterizations.













0

Low Yield



High Yield