

This discussion paper is/has been under review for the journal Atmospheric Chemistry and Physics (ACP). Please refer to the corresponding final paper in ACP if available.

# Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø, Denmark: insight into sources of iodine isotopes and their species transformations

L. Y. Zhang<sup>1,2</sup>, X. L. Hou<sup>1,2</sup>, and S. Xu<sup>3</sup>

Received: 29 June 2015 - Accepted: 31 August 2015 - Published: 15 September 2015

Correspondence to: X. L. Hou (xiho@dtu.dk)

Published by Copernicus Publications on behalf of the European Geosciences Union.

Discussion

Discussion Paper

Discussion Paper

Printer-friendly Version

Interactive Discussion

**ACPD** 

15, 25139–25173, 2015

Speciation of <sup>127</sup>I and 129 I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page

**Abstract** Introduction Conclusions References **Tables Figures** 





Full Screen / Esc

<sup>&</sup>lt;sup>1</sup>Center for Nuclear Technologies, Technical University of Denmark, Risø Campus, Roskilde 4000, Denmark

<sup>&</sup>lt;sup>2</sup>SKLLQG, Shaanxi Key Laboratory of Accelerator Mass Spectrometry Technology and Application, Xi'an AMS Center, Institute of Earth Environment CAS, Xi'an 710061, China <sup>3</sup>Scottish Universities Environmental Research Center, East Kilbride, G75 0QF, UK

Discussion Paper

Discussion Paper

Discussion Paper

# **ACPD**

15, 25139-25173, 2015

Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

I 

I 

Back Close

Printer-friendly Version

Full Screen / Esc

Interactive Discussion



25140

The atmospheric chemistry of iodine has attracted increasing attention in recent years, in part because of its central role in governing the geochemical cycling of iodine, and also for its influences on ozone depletion, primary particle formation, and global climate (Carpenter, 2003; O'Dowd et al., 2002; Saiz-Lopez et al., 2012). Most atmospheric iodine studies to date have concentrated on laboratory smog chamber experiments and modeling studies, with a paucity of field measurements of atmospheric samples. Apart from gaseous iodine (e.g. molecular iodine, iodocarbons) in the atmosphere, aerosols contain a significant fraction of iodine, and play an important role for iodine transportation from ocean to land. In particular, speciation analysis of iodine in aerosols is expected to speculate the reactions and atmospheric processes of iodine, and develop a comprehensive understanding of biogeochemical cycles of iodine in nature.

Most of the work on aerosol iodine speciation has focused on water-soluble iodine (WSI) (Baker, 2004, 2005; Gilfedder et al., 2008). Early modeling studies suggested that aerosol iodate should be the only stable iodine specie as a by-product of iodine oxides production, with negligible iodide due to transformation into gaseous iodine (Mc-Figgans et al., 2000; Vogt et al., 1999). However, this prediction went against several previous observations that showed significant iodide and soluble organic iodine, accounting for 10-100% of WSI in aerosols (Baker, 2004, 2005; Gabler and Heumann, 1993). An improved model was then developed, primarily by increasing the proportion of iodide in the aerosols (Pechtl et al., 2007). Few studies have considered waterinsoluble iodine in aerosols (Gilfedder et al., 2010; Tsukada et al., 1987), which might hardly be converted to gaseous iodine before being deposited on the Earth's surface. Existing observational data suggest that water-insoluble iodine is abundant in aerosol particles, representing 17-53% of the total iodine (Gilfedder et al., 2010). However, there is no comprehensive investigation of aerosol iodine species that encompasses aqueous-soluble and insoluble iodine species. This constitutes a significant gap in our understanding of the interaction among atmospheric iodine species.

ACPD

Paper

Discussion Paper

Discussion Paper

Discussion Paper

15, 25139-25173, 2015

Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

I ✓ ▶I

Back Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



25141

Discussion Paper

**ACPD** 

15, 25139-25173, 2015

Speciation of 127 I and 129 I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page

Abstract Introduction Conclusions References **Tables Figures** 

Full Screen / Esc

Close

Back

Printer-friendly Version

Interactive Discussion



Due to the difficulty of source identification of natural  $^{127}$ I,  $^{129}$ I ( $T_{1/2} = 15.7$  Ma) has proven be an especially useful geochemical tracer both in the marine environment and in the atmosphere (Hou et al., 2007; Jabbar et al., 2012; Michel et al., 2012). A large amount of <sup>129</sup>I has been released to the environment by nuclear reprocessing plants (NRPs). These dominate all other anthropogenic sources, with 6000 kg released as of 2009, from Sellafield (United Kingdom) and La Hague (France) (He et al., 2013b; Hou et al., 2007; Raisbeck et al., 1995). Anthropogenic <sup>129</sup>I is currently distributed unevenly around the globe in water, atmosphere and terrestrial systems (Hou et al., 2009b). Above-ground nuclear explosions also produced <sup>129</sup>I, injecting it into the troposphere and stratosphere. The proportion that entered to the stratosphere became mixed globally before re-entering the troposphere. Larger particles remained in the troposphere for about 20 days before being deposited on the earth's surface (Tölgyessy, 1993). Gaseous releases from reprocessing plants and accidents are primary sources of 129 I to the local environment (Hou et al., 2009a; Xu et al., 2013). Re-emission of iodine from the marine boundary layer plays a significant role in the atmospheric budget of <sup>129</sup>I (Englund et al., 2010b). As with stable iodine (1271), 1291 also travels in the atmosphere in particulate form. Particulate <sup>129</sup>I has been used successfully in monitoring studies, and as an atmospheric tracer (Englund et al., 2010b; Jabbar et al., 2012; Santos et al., 2005; Tsukada et al., 1987). While, speciation analysis of <sup>129</sup>I in aerosols is extremely scarce with only one our previous study (Xu et al., 2015).

Here, we present the results of speciation analysis of stable <sup>127</sup>I and radioactive <sup>129</sup>I in aerosols collected in Denmark. Our goal is to understand how the iodine is carried in the atmosphere and to investigate transformation mechanisms between iodine species in aerosols, as well as gas-aerosol exchange processes involving iodine. To our knowledge, this is the first paper to address the conversion among overall iodine species in aerosols.

#### 2.1 Aerosol sampling

The aerosol samples were collected on 0.45 µm polypropylene filters (Type G-3, PTI, Germany), attached to an in-house aerosol collector at Risø campus, Technical University of Denmark, Denmark (55°41.77′ N, 12°05.39′ E) (Fig. 1). The sampling filters are normally replaced every week, but during the Fukushima nuclear accident, the changing frequency was increased to every 3–4 days. Seven aerosol samples were collected from 31 March–2 May 2011 (local time, the same below), shortly after the Fukushima accident on 11 March 2011. One sample was collected from 8–15 December 2014. The samples were put into plastic bags and stored in the dark before analysis. The sample information is listed in Table 1.

# 2.2 Speciation analysis of <sup>127</sup>I and <sup>129</sup>I in aerosols

Separation of iodine species from aerosol. The aerosol samples were analyzed using a newly developed method (Zhang et al., 2015). In brief, iodine was extracted sequentially using deionized water and sodium hydroxide solutions for water-soluble and NaOH-soluble iodine (WSI and NSI). Total iodine (TI) and residual insoluble iodine (RII) were separated by alkaline ashing from the original air filters and the residual filters after NaOH solution leaching, respectively. Iodide and iodate in the water leachate were separated using anion exchange chromatography. Figure 2 shows a diagram of the separation procedure for speciation analysis of iodine isotopes.

Determination of <sup>127</sup>I by ICP-MS and <sup>129</sup>I by AMS. <sup>127</sup>I in all fractions were diluted by a factor of 1–20 using ammonium to 1 % NH<sub>3</sub>·H<sub>2</sub>O. Cesium was added as internal standard to a final concentration of 2 ng g<sup>-1</sup> to monitor the ionization efficiency of iodine in the ICP-MS measurement. Standard addition method was employed for the measurement of <sup>127</sup>I in ash leachates of total iodine and the insoluble fraction. An iodine standard solution (NaI, CPI International, California, USA) was spiked into the

Discussion Paper

Discussion Paper

Discussion Paper

Discussion Paper

**ACPD** 

15, 25139-25173, 2015

Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page

Full Screen / Esc
Printer-friendly Version

Close

Back



reserved aliquots and diluted with 1 %  $NH_3 \cdot H_2O$  to a final concentration of 2 ng g<sup>-1</sup>. <sup>127</sup>I in the diluted solution was measured by ICP-MS (Thermo Fisher, X Series II) using an Xt cone under normal mode, as described elsewhere (Hou et al., 2007).

 $^{129}\text{I}$  was measured using a 5 MV accelerator mass spectrometry (NEC, Wisconsin, USA) at Scottish University Environmental Research Center, UK (Xu et al., 2013). The prepared AgI precipitates were mixed with silver powder (Sigma-Aldrich Co., USA) with a AgI : Ag mass ratio of 1 : 2 and pressed into copper holders. A terminal voltage of 3.5 MV and I $^{5+}$  ions were chosen for  $^{129}\text{I}$  detection. The measured  $^{129}\text{I}$  /  $^{127}\text{I}$  ratios were corrected against a standard with an  $^{129}\text{I}$  /  $^{127}\text{I}$  ratio of 1.138  $\times$  10 $^{-10}$ , prepared from a NIST 4949C standard and an  $^{127}\text{I}$  carrier. The analytical precision of standards and samples are within 5 %.  $^{129}\text{I}$  /  $^{127}\text{I}$  ratios in the procedural blanks are about 5  $\times$  10 $^{-13}$ , 1–3 orders of magnitude lower than those measured for the samples. All results were corrected for procedural blanks.

#### 3 Results

The concentrations of total iodine in aerosols from Risø, Denmark ranged at 1.04–2.48 ng m $^{-3}$  (average of  $1.79\pm0.52$  ng m $^{-3}$ ) for  $^{127}I$  and  $11.31–72.98\times10^5$  atoms m $^{-3}$  (average of  $(43.65\pm18.88)\times10^5$  atoms m $^{-3}$ ) for  $^{129}I$ , during March–May 2011 (Table 2). Total  $^{127}I$  concentration of 2.36 ng m $^{-3}$  during 8–15 December 2014 fell within the range in 2011, while  $^{129}I$  concentration of 97.00  $\times$  10 $^5$  atoms m $^{-3}$  was about two times higher than the average value in 2011. Much lower values of  $^{129}I$  and  $^{127}I$  concentrations were observed during 18 April–2 May compared to those before 18 April (Fig. 3), but  $^{129}I$  levels didn't show a synchronous variation with  $^{127}I$  concentrations. The results of  $^{127}I$  concentrations in the studied aerosols are compatible with those from an inland city, Regensburg, Germany and from some Pacific islands (Gabler and Heumann, 1993, and references therein).  $^{129}I$ /  $^{127}I$  atomic ratios in the investigated aerosols were (17.84–86.84)  $\times$  10 $^{-8}$ , which are consistent with those found at Foehr island, north of Germany in April 2002 (Michel et al., 2012) and in southern Sweden

ACPD

15, 25139-25173, 2015

Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

I 

▶ I

Back Close

Printer-friendly Version

Full Screen / Esc



Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

I ✓ ▶I

Back Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



during 1983–2008 (Englund et al., 2010b; Michel et al., 2012). However, the measured <sup>129</sup>I concentrations and <sup>129</sup>I / <sup>127</sup>I ratios are 1–2 orders of magnitude higher than those observed in Vienna, Austria during 2001–2002, and in the high altitude eastern Alps (2700 m) during 2001 (Jabbar et al., 2011, 2012).

For  $^{127}$ I, the WSI fraction of aerosols collected in 2011 occurred predominantly as iodide, ranging from 0.12–0.33 ng m $^{-3}$ . Minor iodate of 0.02–0.03 ng m $^{-3}$  was detected only in two samples (AE11–12 during 4–7 April and AE11–17 during 26 April–2 May) (Table 1 and Fig. 4). The highest measured  $^{127}$ I $^-$  concentration (0.74 ng m $^{-3}$ ) was observed in aerosol collected in 2014, and there was no measurable water-soluble organic iodine. Water-soluble  $^{129}$ I shows a similar species pattern as water-soluble  $^{127}$ I except that iodate-129 was below the detection limit in all the analyzed samples. The concentrations of  $^{129}$ I $^-$  varied in the range of (3.26–5.91) × 10 $^5$  atoms m $^{-3}$  with an average value of 4.41 × 10 $^5$  atoms m $^{-3}$  in the aerosols from 2011. This value is about 8 times lower than that (30.12 × 10 $^5$  atoms m $^{-3}$ ) in the sample collected in 2014. High NSI concentrations were observed in all of the aerosol samples, with an average of 0.64 ± 0.21 ng m $^{-3}$  for  $^{127}$ I and (13.55 ± 10.12) × 10 $^5$  atoms m $^{-3}$  for  $^{129}$ I. RII is the dominant species in aerosol samples for both iodine isotopes, with concentration range of 0.34–1.66 ng m $^{-3}$  for  $^{127}$ I and (4.27–39.94) × 10 $^5$  atoms m $^{-3}$  for  $^{129}$ I.

There is no available standard reference material for iodine isotopes in aerosols which can be used for validation of the analytical method. While, the sums of all of iodine species in the analyzed samples are comparable with the measured total iodine in the corresponding samples for both <sup>127</sup>I and <sup>129</sup>I in a range of 85–110%, indicating the reliability of our analytical results (Fig. 5).

For the samples collected before 18 April 2011, the proportion of <sup>129</sup>I and <sup>127</sup>I species follows an order of RII > NSI > iodide. For the samples collected later in April 2011 and in the winter 2014, the three iodine fractions for both <sup>129</sup>I and <sup>127</sup>I account for almost equivalent proportion of total iodine (Fig. 5). Compared to the former sampling period (31 March–18 April 2011), the most apparent feature of the latter sampling period is a dramatic diminution of iodine in the RII fraction, by 38.7 % for <sup>127</sup>I and 26.8 % for <sup>129</sup>I.

This was compensated by a corresponding increase in the iodide fraction of the aerosol particles. Iodate, as the least abundant iodine specie (< 3%), was only observed in two aerosol samples for  $^{127}$ I, and no  $^{129}$ IO $_3^-$  was detected.

An range of  $^{129}I/^{127}I$  atomic ratios of  $(15.56-102.36) \times 10^{-8}$  was determined for the different iodine species in aerosols. The variations of  $^{129}I/^{127}I$  ratio in iodide, NSI and RII are similar to those in total  $^{129}I/^{127}I$ , with high values observed prior to 18 April 2011 and December 2014 (Table 2 and Fig. 3).

#### 4 Discussion

#### 4.1 Sources of <sup>127</sup>I and <sup>129</sup>I in aerosols

Concentration variations of <sup>127</sup>I and <sup>129</sup>I in aerosols against meteorological parameters (i.e. wind direction, wind speed and temperature) during the sampling period show that wind direction has a dominant influence (Fig. 3). Back trajectory model analysis shows that <sup>127</sup>I and <sup>129</sup>I in the aerosols was directly controlled by the sources and pathways of air masses (Figs. 6, S1 and S2 in the Supplement). The relatively high <sup>127</sup>I and <sup>129</sup>I concentrations was observed in in aerosols collected early April 2011 and December 2014, when the air mass was mainly transported from the Atlantic Ocean and the North Sea by prevailing westerly winds. Relatively low concentrations of iodine isotopes were observed in the aerosols collected later in April, when the air masses were dominated by prevailing easterly winds and passed over the European continent and the Baltic Sea.

Marine emission of volatile iodine species is a major source of iodine in the atmosphere. It results in relatively elevated <sup>127</sup>I concentrations in the marine atmosphere, as compared to the terrestrial atmosphere (Saiz-Lopez et al., 2012). During the sampling period of 11–14 April, the sampled air mass was transported by westerly winds from a vast area over the northern Atlantic Ocean. This caused an elevated <sup>127</sup>I concentration. Except for the Norwegian Sea, most areas of the northern Atlantic have low seawater

**ACPD** 

15, 25139–25173, 2015

Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

I ◀ ▶I

Back Close

Full Screen / Esc

Printer-friendly Version



<sup>129</sup>I concentrations (He et al., 2013a), which is what we observed for the 11–14 April aerosol sample.

The highest measured  $^{129}I$  concentrations,  $(72.98-97.00) \times 10^5$  atoms m<sup>-3</sup>, were observed in the aerosol samples collected in 4-7 April 2011 and 8-15 Decem-5 ber 2014. These samples reflect that air masses were transported from two directions, west/northwest and south/southwest (Figs. 6, S1 and S2). The dominant south/southwesterly wind passed over regions of elevated <sup>129</sup>I, including the North Sea and reprocessing sites at Cap de La Hague and Sellafield. In contrast to the marinederived <sup>127</sup>I, more than 95% of environmental <sup>129</sup>I is discharged from Sellafield and La Hague, west and southwest of Denmark, respectively (Fig. 1). Only a small fraction of <sup>129</sup>I in gaseous form (about 2–5 kg year<sup>-1</sup> in the past 20 years) has been released to the atmosphere from two reprocessing plants at La Hague and Sellafield, and this has become diluted over a large area, including southern Europe (Ernst et al., 2003; Hou et al., 2007). However, a large fraction of <sup>129</sup>I (200–300 kg year<sup>-1</sup> since 1995) has been discharged directly to the sea in liquid form from these two reprocessing plants. After being introduced into the English Channel and Irish Sea, the <sup>129</sup>I follows surface ocean currents to the North Sea, Kattegat, and Baltic Sea, and continues to the Arctic along the Norwegian coast (Alfimov et al., 2004a; Buraglio et al., 1999; Hou et al., 2007; Raisbeck et al., 1995; Yi et al., 2012). Remarkably, elevated <sup>129</sup>I concentrations of up to  $10^{10}$ – $10^{11}$  atoms L<sup>-1</sup> have been found in the North Sea,  $10^9$ – $10^{10}$  atoms L<sup>-1</sup> in Norwegian coastal waters and the Kattegat, and  $10^8 - 10^9$  atoms L<sup>-1</sup> in the Baltic Sea (Aldahan et al., 2007; Alfimov et al., 2004b; He et al., 2014; Hou et al., 2007; Michel et al., 2012; Yi et al., 2011). In addition to direct releases to the atmosphere from the two reprocessing plants, volatile iodine is also emitted from seawater, especially in the North Sea and Kattegat. This process is facilitated by sea-spray and biological activity of macroalgae and microalgae (McFiggans, 2005; O'Dowd et al., 2002). This secondary iodine source is known to significantly increase atmospheric <sup>129</sup>I concentrations and may be expected to contribute to the iodine measured in our aerosol samples.

#### **ACPD**

15, 25139–25173, 2015

Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Printer-friendly Version



Printer-friendly Version



The lowest measured  $^{129}$ I concentrations,  $(11-13) \times 10^5$  atoms m<sup>-3</sup>, were observed in aerosol samples collected in 18-26 April and 26 April-2 May 2011. Back trajectory analysis (Fig. S1) shows that in this period the air masses at the sampling site were mainly transported by easterly or northwesterly winds. Terrestrial emissions of 5 iodine occur through vegetation and terrestrial microorganisms, and are much lower than marine emissions (Bewers and Haysom, 1974). This is reflected in these two aerosol samples by their relatively low <sup>129</sup>I concentrations. An elevated <sup>210</sup>Pb level (249–253 μBq m<sup>-3</sup>) (Table 1) for this period is also consistent with a continental origin (210 Pb in the air is a decay product of 222 Rn which is mainly released from the soil in the continental area). The <sup>129</sup>I concentrations of aerosol samples collected in 11-14 and 14-18 April 2011 ranged of  $(43.81-47.27) \times 10^5$  atoms m<sup>-3</sup>, which is lower than the 7 April sample  $(72.98 \times 10^5 \text{ atoms m}^{-3})$  by a factor of 1.6, but about 4 times higher than the aerosol samples that passed over the European continent from 18 April to 2 May 2011. A back trajectory analysis clearly shows that the dominant air masses during the sampling periods were westerly, i.e. from the Atlantic Ocean but cross a narrow area of the northern North Sea (Figs. 6 and S1). The secondarily high <sup>129</sup>I concentrations in these aerosol samples should be attributed to the re-emission of 129 I from the highly contaminated seawater in the North Sea. It can therefore be concluded that besides the direct air releases of <sup>129</sup>I from the two European reprocessing plants, secondary emission of <sup>129</sup>I from the highly contaminated North Sea water is the dominant source of <sup>129</sup>I in the aerosols collected in Denmark. This is also supported by <sup>129</sup>I measurements in aerosols from high altitude European sites (Jabbar et al., 2012).

It should be noted that such a result was not be observed for precipitation samples collected in central Sweden during 1998-1999 (Buraglio et al., 2001). This might be attributed to different mechanisms that iodine incorporates into particles and precipitation. Gaseous iodine species, e.g. I<sub>2</sub> and CH<sub>2</sub>I<sub>2</sub>, emitted by iodine-rich seaweed are important precursors for the formation of new particles, due to their relatively rapid photolysis to active iodine oxides (e.g. IO, OIO) (McFiggans, 2005; O'Dowd et al., 2002; Saiz-Lopez et al., 2012). Hence, iodine in marine aerosols directly participates in

## **ACPD**

15, 25139-25173, 2015

Speciation of <sup>127</sup>I and 129 I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page **Abstract** Introduction

Conclusions References

**Tables Figures** 

Back Close

Full Screen / Esc

aerosol formation. In contrast, iodine in precipitation originates from washout process of gaseous iodine and air-borne particulate iodine, primarily in the form of polar iodine compounds (e.g. HI, HOI, IO, I<sup>-</sup>, IO<sub>3</sub><sup>-</sup>, I<sub>2</sub> and unidentified organic iodine species) in H<sub>2</sub>O molecules during precipitation events (Buraglio et al., 2001; Gilfedder et al., 2007). In addition, aerosol samples were collected over a relatively long time scale (3–7 days), reflecting an accumulated signal during the sampling period, while precipitation events are normally much shorter. Therefore, the source terms of iodine isotopes in precipitation might not necessarily follow back trajectory analysis.

#### 4.2 Species of <sup>129</sup>I and <sup>127</sup>I in aerosols

WSI is virtually pure iodide in the aerosols investigated, with iodate and water-soluble organic iodine accounting for less than 3% of total iodine, and these are only measureable in two samples. Iodate was once considered to be the only WSI species in aerosol (Vogt et al., 1999). This was supported by earlier field observations demonstrating that iodate was dominant in size-segregated aerosols from the tropical Atlantic (Wimschneider and Heumann, 1995). However, this iodate-dominant feature was not found in other aerosol samples, e.g. in the northwest Atlantic Ocean and in tropical atmospheric aerosols (Baker, 2004, 2005). In these cases, iodide was the dominant aerosol phase, as observed in this study. Significant amounts of soluble organic iodine, accounting for 83-97% of WSI, has been reported in aerosols collected at the Mace Head atmospheric research station on the west coast of Ireland (Gilfedder et al., 2008). Water-soluble organic iodine accounting for 4-75 % of WSI were also measured in aerosols collected from a cruise from the UK to the Falkland Islands in 2003 (Baker, 2005). This suggests that the proportion of soluble organic iodine in aerosols varies regionally and depends on particular aerosol sources and formation processes. Some of this variability might also be related to the analytical methods employed for speciation analysis (Zhang et al., 2015).

It is not clear how iodide is formed in the atmosphere, in an oxidizing environment containing oxygen and ozone. Early models predict a negligible iodide concentration in

ACPD

15, 25139–25173, 2015

Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

I 

I 

Back Close

Full Screen / Esc
Printer-friendly Version

Interactive Discussion



25149

Back

Printer-friendly Version

Interactive Discussion



particle phases based on an assumption that the iodide in aerosols only origins from the low levels of gaseous HI (McFiggans et al., 2000; Vogt et al., 1999). This work in combination with other previous results (Baker, 2004; Xu et al., 2015) suggests that there must be alternative primary pathways that contribute to iodide formation at the 5 observed levels. It's generally accepted that iodine atoms are photochemically produced by photolysis of gaseous iodinated compounds, and oxidized by ozone to form reactive iodine oxides (Carpenter, 2003; Saiz-Lopez et al., 2012; Vogt et al., 1999). The formation of iodide from iodine atoms and other reactive iodine compounds must rely on electron-donors that are capable of reducing high valence iodine species to iodide. One possibility is the involvement of sulfur compounds (Chatfield and Crutzen, 1990). Possible reaction pathways are given in Table 3. Gaseous SO<sub>2</sub> can be formed in nature by reactions of dimethyl sulfate (DMS) with hydroxide and nitrate. By associating with  $H_2O$ , these reactions produce  $HSO_3^-$  and  $SO_3^{2-}$  (Eqs. 1 and 2). Native iodine and other reactive species (not shown) can be reduced to I on gas-aerosol interfaces (Eg. 3). Other iodine species in aerosols can also reduced by reductive sulfur compounds to form iodide (Eq. 4). The electron-donors are not limited to sulfur compounds either, for example, nitrogen in the form of gaseous ammonia  $(NH_3 \rightarrow NO_2/NO_3)$  (McFiggans et al., 2000; Saiz-Lopez and Plane, 2004) and elemental mercury (Hg<sup>0</sup>  $\rightarrow$  HgO/HgX, where X is a halogen,  $I^-$ ,  $Br^-$  or  $CI^-$ ) (Lindberg et al., 2002; Simpson et al., 2007) are also candidates responsible for iodide formation.

We note that relatively low WSI 129 I and 127 I was measured in marine-sourced aerosols from the North Sea, as compared to the European continent-sourced aerosols (Figs. 4 and 5). This is consistent with the findings drawn from an iodine speciation study of coastal aerosol samples from England (Baker et al., 2001), where the concentrations of total water-soluble iodine from continental aerosols were significantly higher than those from marine aerosols.

A large proportion of <sup>129</sup>I and <sup>127</sup>I in our aerosol samples were NaOH-soluble, which is consistent with the results of aerosol from Tsukuba, Japan, collected shortly after the Fukushima nuclear accident in March 2011 (Xu et al., 2015). Abundant NaOH-soluble

### **ACPD**

15, 25139–25173, 2015

Speciation of <sup>127</sup>I and 129 I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page **Abstract** Introduction

Conclusions References

**Tables Figures** 

Close

Full Screen / Esc

Discussion Paper

Discussion

**ACPD** 

15, 25139-25173, 2015

Speciation of <sup>127</sup>I and 129 I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page

**Abstract** Introduction Conclusions References **Tables Figures** 

Full Screen / Esc

Close

Back

Printer-friendly Version

Interactive Discussion



129 (32-44 % of total 129 l) in Fukushima-derived aerosols was attributed to coarse vegetation-related organic particles concentrated during spring. The measured NaOHsoluble iodine (NSI) fractions of <sup>129</sup>I and <sup>127</sup>I during the entire sampling period in the spring of 2011 and winter of 2014 are similar. This indicates that NSI is relatively stable 5 and less affected by the source and pathways of air masses than WSI. NaOH leaching is often used to extract organic substance in fractionation analysis of soil and sediment (Englund et al., 2010a; Hou et al., 2003) based on the high solubility of organic matter, such as humic substances. Organic compounds are important contributors to aerosols, such as lipidic, saccharides and proteinaceous materials (O'Dowd et al., 2004; Quinn et al., 2014). A significant portion of atmospheric aerosols was found to be humic-like substances (HULIS), named for their strong structural similarity to humic and fulvic acids (Havers et al., 1998). Most of these organic compounds are water-soluble, but a significant water-insoluble fraction of the HULIS material is hydrophobic and acidic in character, and can be dissolved in an alkaline solvent, like NaOH (Feczko et al., 2007; Havers et al., 1998). Therefore, NaOH-soluble iodine is suggested to be likely associated with HULIS in aerosols.

RII in aerosols has received less attention than WSI (Gilfedder et al., 2010; Tsukada et al., 1987). The early report on water-insoluble iodine fraction in aerosol particles showed that water-insoluble iodine accounted for 27-58% of total iodine bound in aerosols from Tokyo, Japan, collected in 1983–1984 (Tsukada et al., 1987). Another similar result of 17-53% of total iodine as insoluble species was reported for aerosols from the west coast of Ireland in 2007, and from a ship transect from China to Antarctica in 2005-2006 (Gilfedder et al., 2010). Taking the alkaline-leachable iodine in aerosols into account, these results are very consistent with our observations from Risø (Fig. 5). The residual insoluble <sup>129</sup>I fractions were reported to be 4–23% of total <sup>129</sup>I in Fukushima-derived aerosol particles (Xu et al., 2015), less than the proportion in the aerosols collected in Denmark. This discrepancy reveals different formation processes of RII species for the NRPs-derived <sup>129</sup>I in this study as compared to those from Fukushima. A significant difference is the timing of the <sup>129</sup>I releases. NRPs have been

Printer-friendly Version

releasing <sup>129</sup>I into the European environment for about 50 years, allowing it to follow geochemical pathways on timescales ranging from days to decades. In contrast, RII from Fukushima-derived aerosols had only days to react with their environment prior to sampling, 15-22 March 2011.

The origin of the RII fraction is not well understood at present. It's possible that the RII fraction is derived from suspended soil particles (Xu et al., 2013). However, speciation analysis of <sup>129</sup>I in soils show that residual iodine after NaOH leaching accounts for less than 10% of the total, and this minor component is assumed to be associated with minerals (Hou et al., 2003; Qiao et al., 2012). Another possibility is a connection with soot and black carbon from combustion processes, found in anthropogenic aerosols (Kondratyev et al., 2006; Rose et al., 2006), However, the aerosols collected for this study in early April 2011 and winter of 2014 were mainly marine aerosols, with high RII. Alternatively, the RII fraction in aerosols might be associated with metal oxides that originated by suspension of fine inorganic particles. A relatively large fraction of iodine in soil and sediment has been observed in metal oxides associated form (Hou et al., 2003). Gaseous iodine in the atmosphere could interact with these inorganic particles and become firmly bound. The relatively high RII in the marine aerosols (Figs. 4 and 5) suggests that some marine components facilitate the association of gaseous iodine with inorganic particles.

A significant positive correlation between <sup>127</sup> lodide and <sup>7</sup>Be in aerosol samples (R = 0.76, p = 0.05) is observed (Fig. 7). <sup>7</sup>Be  $(t_{1/2} = 54 \text{ days})$  is produced in the stratosphere by cosmic ray reactions with oxygen and nitrogen and rapidly transported to the troposphere by vertical mixing processes to be finally deposited onto the earth. The production of Be decreases with decreasing altitude in the troposphere because of the attenuation of cosmic rays by interactions with atomic targets in the atmosphere (Turekian et al., 1983). This positive correlation might indicate that the formation of iodide occurs in air masses at higher altitude. However, no significant correlation between <sup>129</sup>I<sup>-</sup> and <sup>7</sup>Be is observed (Fig. 7). This is consistent with the process of secondary emission of <sup>129</sup>I from contaminated seawater, which would only increase <sup>129</sup>I in

### **ACPD**

15, 25139-25173, 2015

Speciation of <sup>127</sup>I and 129 I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page **Abstract** Introduction Conclusions References **Tables Figures** 

Back Close

Full Screen / Esc

the troposphere. The significant negative correlation of NSI of  $^{129}$ I with  $^{7}$ Be (R = 0.73, p = 0.06) likely reflects the association of iodine with organic matter at low altitude, where organic matter is liberated by biological activity in the ocean and on land.

# 4.3 Fukushima-derived <sup>129</sup>I signal in the European atmosphere

The Fukushima Dai-ichi nuclear power plant accident on 11 March 2011 released radioiodine to the atmosphere, primarily as <sup>131</sup>I and <sup>129</sup>I, which was mainly transported eastwards by prevailing winds. Based on <sup>129</sup>I levels in the offshore Fukushima seawater, the released 129 amount from this accident was estimated to be 1.2 kg (Hou et al., 2013). Gaseous <sup>129</sup>I travelled across the Pacific Ocean, American continent and Atlantic Ocean, and some fraction arrived on the European continent after 1-2 weeks (Clemenza et al., 2012; Leon et al., 2011; Manolopoulou et al., 2011; Pittauerová et al., 2011). Anthropogenic <sup>129</sup>I has been reported from a variety of environmental samples in Japan, including soil, seawater, precipitation and aerosols (Buesseler et al., 2012; Hou et al., 2013; Muramatsu et al., 2015; Xu et al., 2013, 2015). The level of <sup>129</sup>I in aerosols collected in Tsukuba, about 170 km from the Fumushima Dai-ichi NPP, reached  $5 \times 10^8$  atoms m<sup>-3</sup> (Xu et al., 2015). While the Fukushima-derived <sup>129</sup>I in environmental samples outside of Japan was less well-characterized. 131 in the aerosol samples collected at Risø, Denmark, 10 days after the Fukushima accident have been observed (Fig. 8) (Nielsen et al., 2011). The <sup>131</sup>I reached the peak on 24-30 March 2011 (763 µBq m<sup>-3</sup> in aerosol), then fell to below detection limits for <sup>131</sup>I in the middle of May. Based on the measured <sup>131</sup>I radioactivity in the aerosol samples and an  $^{129}I/^{131}I$  atomic ratio of  $16.0 \pm 2.2$  deduced from the aerosol samples collected at Tsukuba, Japan shortly after the Fukushima accident (Xu et al., 2015), the Fukushima-derived <sup>129</sup>I signal in Denmark can be reconstructed (Fig. 8). The highest <sup>129</sup>I concentration in the aerosols in Denmark from the Fukushima accident is estimated to be  $6.3 \times 10^4$  atoms m<sup>-3</sup> on 30–31 March 2011, which accounts for less than  $\sim 6\%$ of total  $^{129}I$  (1.1–9.7 × 10<sup>6</sup> atoms m<sup>-3</sup>) in Denmark when the Fukushima  $^{131}I$  peak was

**ACPD** 

15, 25139-25173, 2015

Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page

Full Screen / Esc

Close

Back

Printer-friendly Version



measured. Considering the rapid decline of <sup>129</sup>I levels in aerosols and precipitation in Japan to nearly pre-accident levels within two years (Xu et al., 2013), the contribution of Fukushima-derived <sup>129</sup>I to the <sup>129</sup>I level and inventory in Europe is now negligible in comparison to NRPs-derived <sup>129</sup>I in the European atmosphere.

### 4.4 Dry deposition flux of iodine isotopes

Dry deposition fluxes of iodine can be estimated as  $F_d = C_d \cdot v_d$  (Baker et al., 2001), where  $F_d$  is the flux ( $\mu g \, m^{-2} \, year^{-1}$  for  $^{127}I$  and atoms  $m^{-2} \, year^{-1}$  for  $^{129}I$ ),  $C_d$  is the aerosol iodine concentration in the atmosphere (ng m<sup>-3</sup> for <sup>127</sup>I and atoms m<sup>-3</sup> for  $^{129}I$ ), and  $v_d$  is the deposition velocity (cm s<sup>-1</sup>). The deposition velocities are size-dependent, at  $0.1 \, \text{cm s}^{-1}$  for fine particles (Aerodynamic diameter  $< 1 \, \mu \text{m}$ ) and 1.5 cm s<sup>-1</sup> for larger particles (Duce et al., 1991). The aerosol samples in this work were collected on a PP filter with a pore size of approximately 0.45 μm. According to the particle size distribution at Spieka-Neufeld, Germany, whose atmospheric input of iodine is also from the North Sea, as in this study, 67% of aerosol iodine accumulated in particles larger than 1 µm and 33% in 0.45-1 µm particles (Gabler and Heumann, 1993). Employing these parameters, dry deposition fluxes of iodine in Denmark can be estimated to be 342-815 µg m<sup>-2</sup> year<sup>-1</sup> for <sup>127</sup>I, and  $(0.4-3.2) \times 10^{12}$  atoms m<sup>-2</sup> year<sup>-1</sup> for <sup>129</sup>I during 2011–2014. If the average values of  $1.8 \,\mathrm{ng}\,\mathrm{m}^{-3}$  and  $43.7 \times 10^5 \,\mathrm{atoms}\,\mathrm{m}^{-3}$  are used as the representative concentrations of <sup>127</sup>I and <sup>129</sup>I in aerosol, respectively for 2011, average dry deposition fluxes of iodine in 2011 would be  $591 \, \mu \mathrm{g \, m^{-2} \, year^{-1}}$  and  $1.4 \times 10^{12} \, \mathrm{atoms \, m^{-2} \, year^{-1}}$  for  $^{127} \mathrm{I}$  and 129 I, respectively. The 127 I deposition flux falls within the range of natural stable iodine deposition, at 460-830 µg m<sup>-2</sup> year<sup>-1</sup>, observed at Norfolk, on the north coast of the United Kingdom (Baker et al., 2001). It is also comparable with measured values from southern Sweden, but higher than northern Sweden (Englund et al., 2010b). The dry deposition flux of <sup>129</sup>I for aerosol samples collected in Sweden between 1983 and 2008 shows a range of 0.33-6.6 × 10<sup>11</sup> atoms m<sup>-2</sup> year<sup>-1</sup> in southern Sweden, and

**ACPD** 

15, 25139-25173, 2015

Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

I 

I 

I 

I 

Back Close

Printer-friendly Version

Full Screen / Esc



 $0.008-1.6 \times 10^{11}$  atoms m<sup>-2</sup> year<sup>-1</sup> in northern Sweden (Englund et al., 2010b), which are 1–3 orders of magnitude lower than the deposition fluxes determined here. This is attributed to the distant sampling locations in the Swedish study, much further away from the major point sources of aerosol <sup>129</sup>I than Denmark. The wet deposition of <sup>129</sup>I at the Risø site was calculated to be  $(1.25\pm0.30)\times10^{12}$  atoms m<sup>-2</sup> year<sup>-1</sup> during 2001–2006 (Hou et al., 2009a), which is consistent with our calculated dry deposition flux of <sup>129</sup>I, indicating that subequal contributions of <sup>129</sup>I by dry and wet deposition. It should be noted, however, that the estimated dry deposition flux was based on only one month of data in the spring and on a modeled deposition velocity, which could potentially introduce significant uncertainty due to temporal variations of <sup>129</sup>I in aerosols and particulate-size dependent velocity.

#### 5 Conclusions

Based on the analytical results on speciation analysis of <sup>129</sup>I and <sup>127</sup>I in aerosols collected in Denmark immediately after the Fukushima accident and the discussion above, the following conclusions can be drawn:

- Iodide is the dominant form (> 97 %) of the water-soluble iodine in aerosol, its formation might be related to atmospheric reductants, such as reductive SO<sub>3</sub> and disulfites. The most dominant species of iodine in aerosols are NSI and RII, accounting for up to 80 % of total iodine, NSI is predominantly bound to organic matter, such as HULIS, while RII might be associated with inorganic components, such as metal oxides.
- 2. Westerly-dominated air masses from the Atlantic Ocean contribute high <sup>127</sup>I concentrations compared to easterly air masses from the northeastern European continent. <sup>129</sup>I in aerosols in Denmark is primarily derived from European nuclear reprocessing plants. This includes secondary emission of <sup>129</sup>I from seawater in

**ACPD** 

15, 25139–25173, 2015

Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page

Back

Printer-friendly Version

Full Screen / Esc

Close



- 3. Fukushima-derived  $^{129}$ I contributed less than  $\sim 6\%$  to the total  $^{129}$ I in the European atmosphere immediately after the Fukushima accident, a negligible quantity when compared to  $^{129}$ I levels released from Sellafield and La Hague nuclear reprocessing plants.
- 4. Water-insoluble iodine (NSI and RII) in aerosols that is primarily associated with organic compounds is crucial for investigation of geochemical cycling of iodine in the atmosphere due to its large fraction. This requires more intensive studies to extend our observations on both temporal and spatial scales.

# The Supplement related to this article is available online at doi:10.5194/acpd-15-25139-2015-supplement.

Acknowledgements. L. Y. Zhang is grateful for the support from all of the colleagues in the Division of Radioecology (headed by Sven P. Niensen), Center for Nuclear Technologies, Technical University of Denmark, for her PhD project. This work is partly supported by the projects of Innovation Methodology (No. 2012IM030200) and Fundamental Scientific Research (2015FY110800) from the Ministry of Science and Technology of China. We acknowledge George Burr for polishing English presentation.

#### References

- Aldahan, A., Possnert, G., Alfimov, V., and Cato, K.: Anthropogenic <sup>129</sup>I in the Baltic Sea, Nucl. Instrum. Meth. B. 259, 491–495, 2007.
  - Alfimov, V., Aldahan, A., Possnert, G., and Winsor, P.: Anthropogenic iodine-129 in seawater along a transect from the Norwegian coastal current to the North Pole, Mar. Pollut. Bull., 49, 1097–1104. 2004a.

.\_

---

Discussion Paper

Discussion Paper

Discussion Pape

#### **ACPD**

15, 25139–25173, 2015

Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

Full Screen / Esc

Close

Printer-friendly Version

Back



15, 25139–25173, 2015

Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.

- Title Page

  Abstract Introduction

  Conclusions References

  Tables Figures

  I 

  I 

  I 

  Back Close
- Printer-friendly Version

Full Screen / Esc

- Interactive Discussion
  - © 0 BY

- Alfimov, V., Aldahan, A., Possnert, G., Kekli, A., and Meili, M.: Concentrations of <sup>129</sup>I along a transect from the North Atlantic to the Baltic Sea, Nucl. Instrum. Meth. B, 223–224, 446–450, 2004b.
- Baker, A. R.: Inorganic iodine speciation in tropical Atlantic aerosol, Geophys. Res. Lett., 31, L23S02, doi:10.1029/2004GL020144, 2004.
- Baker, A. R.: Marine Aerosol Iodine Chemistry: The Importance of Soluble Organic Iodine, Environ. Chem., 2, 295–298, 2005.
- Baker, A. R., Tunnicliffe, C., and Jickells, T. D.: lodine speciation and deposition fluxes from the marine atmosphere, J. Geophys. Res., 106, 28743–28749, 2001.
- Bewers, J. M. and Haysom, H. H.: The terrigenous dust contribution to fluoride and iodide in atmospheric precipitation, J. Rech. Atmos., 8, 689–697, 1974.
- Buesseler, K. O., Jayne, S. R., Fisher, N. S., Rypina, I. I., Baumann, H., Baumann, Z., Breier, C. F., Douglass, E. M., George, J., Macdonald, A. M., Miyamoto, H., Nishikawa, J., Pike, S. M., and Yoshida, S.: Fukushima-derived radionuclides in the ocean and biota off Japan, P. Natl. Acad. Sci. USA, 109, 5984–5988, 2012.
- Buraglio, N., Aldahan, A., and Possnert, G.: Distribution and inventory of <sup>129</sup>I in the central Arctic Ocean, Geophys. Res. Lett., 26, 1011–1014, 1999.
- Buraglio, N., Aldahan, A., Possnert, G., and Vintersved, I.: <sup>129</sup>I from the nuclear reprocessing facilities traced in precipitation and runoff in northern Europe, Environ. Sci. Technol., 35, 1579–1586, 2001.
- Carpenter, L. J.: Iodine in the Marine Boundary Layer, Chem. Rev., 103, 4953-4962, 2003.
- Chatfield, R. B. and Crutzen, P. J.: Are there interactions of iodine and sulfur species in marine air photochemistry?, J. Geophys. Res.-Atmos., 95, 22319–22341, 1990.
- Clemenza, M., Fiorini, E., Previtali, E., and Sala, E.: Measurement of airborne <sup>131</sup>I, <sup>134</sup>Cs and <sup>137</sup>Cs due to the Fukushima reactor incident in Milan (Italy), J. Environ. Radioact., 114, 113–118, 2012.
- Duce, R. A., Liss, P. S., Merrill, J. T., Atlas, E. L., Buat-Menard, P., Hicks, B. B., Miller, J. M., Prospero, J. M., Arimoto, R., Church, T. M., Ellis, W., Galloway, J. N., Hansen, L., Jickells, T. D., Knap, A. H., Reinhardt, K. H., Schneider, B., Soudine, A., Tokos, J. J., Tsunogai, S., Wollast, R., and Zhou, M.: The atmospheric input of trace species to the world ocean, Global Biogeochem. Cy., 5, 193–259, 1991.
- Englund, E., Aldahan, A., Hou, X., Petersen, R., and Possnert, G.: Speciation of iodine (I-127 and I-129) in lake sediments, Nucl. Instrum. Meth. B, 268, 1102–1105, 2010a.

- Englund, E., Aldahan, A., Hou, X., Possnert, G., and Soderstrom, C.: Iodine (I-129 and I-127) in aerosols from northern Europe, Nucl. Instrum. Meth. B, 268, 1139–1141, 2010b.
- Ernst, T., Szidat, S., Handl, J., Jakob, D., Michel, R., Schnabel, Ch., Synal, H. A., Arevalo, F. J. S., Benne, I., Boess, J., Gehrt, E., Capelle, A., Schneider, J., Schafer, W., and Bottcher, J.: Migration of iodine-129 and iodine-127 in soils, Kerntechnik (Germany), 68, 155–167, 2003.
- Feczko, T., Puxbaum, H., Kasper-Giebl, A., Handler, M., Limbeck, A., Gelencsér, A., Pio, C., Preunkert, S., and Legrand, M.: Determination of water and alkaline extractable atmospheric humic-like substances with the TU Vienna HULIS analyzer in samples from six background sites in Europe, J. Geophys. Res.-Atmos., 112, 1–9, 2007.
- Gabler, H. E. and Heumann, K. G.: Determination of particulate iodine in aerosols from different regions by size fractionating impactor sampling and IDMS, Intern. J. Anal. Chem., 50, 129–146, 1993.
  - Gilfedder, B. S., Petri, M., and Biester, H.: Iodine speciation in rain and snow: Implications for the atmospheric iodine sink, J. Geophys. Res.-Atmos., 112, D07301, doi:10.1029/2006JD007356, 2007.
- Gilfedder, B. S., Lai, S., Petri, M., Biester, H., and Hoffmann, T.: Iodine speciation in rain, snow and aerosols and possible transfer of organically bound iodine species from aerosol to droplet phases, Atmos. Chem. Phys. Discuss., 8, 7977–8008, doi:10.5194/acpd-8-7977-2008, 2008.
- Gilfedder, B. S., Chance, R., Dettmann, U., Lai, S. C., and Baker, A. R.: Determination of total and non-water soluble iodine in atmospheric aerosols by thermal extraction and spectrometric detection (TESI), Anal. Bioanal. Chem., 398, 519–526, 2010.
  - Havers, N., Burba, P., Lambert, J., and Klockow, D.: Spectroscopic Characterization of Humic-Like Substances in Airborne Particulate Matter, J. Atmos. Chem., 29, 45–54, 1998.
- He, P., Hou, X., Aldahan, A., Possnert, G., and Yi, P.: Iodine isotopes species fingerprinting environmental conditions in surface water along the northeastern Atlantic Ocean, Scientific Reports, 3, 2685, 1–8, 2013a.
  - He, P., Aldahan, A., Possnert, G., and Hou, X. L.: A summary of global <sup>129</sup>I in marine waters, Nucl. Instrum. Meth. B, 294, 537–541, 2013b.
- He, P., Aldahan, A., Possnert, G., and Hou, X.: Temporal Variation of Iodine Isotopes in the North Sea, Environ. Sci. Technol., 48, 1419–1425, 2014.

**ACPD** 

15, 25139-25173, 2015

Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

4

Full Screen / Esc

Close

Back

Printer-friendly Version



- Hou, X., Fogh, C. L., Kucera, J., Andersson, K. G., Dahlgaard, H., and Nielsen, S. P.: Iodine-129 and Caesium-137 in Chernobyl contaminated soil and their chemical fractionation, Sci. Total
- Hou, X., Aldahan, A., Nielsen, S. P., Possnert, G., Nies, H., and Hedfors, J.: Speciation of I-129 and I-127 in seawater and implications for sources and transport pathways in the North Sea, Environ. Sci. Technol., 41, 5993-5999, 2007.

Environ., 308, 97–109, 2003.

- Hou, X., Aldahan, A., Nielsen, S. P., and Possnert, G.: Time Series of I-129 and I-127 Speciation in Precipitation from Denmark, Environ. Sci. Technol., 43, 6522-6528, 2009a.
- Hou, X., Hansen, V., Aldahan, A., Possnert, G., Lind, O. C., and Lujanienė, G.: A review on speciation of iodine-129 in the environmental and biological samples, Anal. Chim. Acta, 632, 181-196, 2009b.
- Hou, X., Povinec, P. P., Zhang, L., Shi, K., Biddulph, D., Chang, C., Fan, Y., Ješkovský, M., Liu, Q., Luo, M., Steier, P., Zhou, W. J., Hou, Y., and Golser, R.: Iodine-129 in Seawater Offshore Fukushima: Distribution, Inorganic Speciation, Sources, and Budget, Environ, Sci. Technol., 47. 3091–3098. 2013.
- Jabbar, T., Steier, P., Wallner, G., Kandler, N., and Katzlberger, C.: AMS analysis of iodine-129 in aerosols from Austria, Nucl. Instrum. Meth. B, 269, 3183-3187, 2011.
- Jabbar, T., Steier, P., Wallner, G., Priller, A., Kandler, N., and Kaiser, A.: Iodine Isotopes (127) and <sup>129</sup>I) in Aerosols at High Altitude Alp Stations, Environ. Sci. Technol., 46, 8637–8644, 2012.
- Kondratyev, K. Y., Ivlev, L. S., Krapivin, V. F., and Varotsos, C. A.: Aerosol formation processes, in: Atmospheric Aerosol Properties: Formation, Processes and Impacts, edited by: Kondratyev, K. Y., Ivlev, L. S., Krapivin, V. F., and Varotsos, C. A. (Eds.), ISBN: 978-3-540-26263-3 (Print) 978-3-540-37698-9 (Online), Springer Berlin Heidelberg, 187–263, 2006.
- Leon, J. D., Jaffe, D. A., Kaspar, J., Knecht, A., Miller, M. L., Robertson, R. G. H., and Schubert, A. G.: Arrival time and magnitude of airborne fission products from the Fukushima, Japan, reactor incident as measured in Seattle, WA, USA, J. Environ. Radioact., 102, 1032-1038, 2011.
  - Lindberg, S. E., Brooks, S., Lin, C., Scott, K. J., Landis, M. S., Stevens, R. K., Goodsite, M., and Richter, A.: Dynamic Oxidation of Gaseous Mercury in the Arctic Troposphere at Polar Sunrise, Environ. Sci. Technol., 36, 1245-1256, 2002.

**ACPD** 

15, 25139–25173, 2015

Speciation of <sup>127</sup>I and 129 I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page

**Abstract** Introduction Conclusions References

**Tables Figures** 

Back Close

Full Screen / Esc

**Printer-friendly Version** 



Paper

- Manolopoulou, M., Vagena, E., Stoulos, S., Ioannidou, A., and Papastefanou, C.: Radioiodine and radiocesium in Thessaloniki, Northern Greece due to the Fukushima nuclear accident, J. Environ. Radioact., 102, 796–797, 2011.
- McFiggans, G.: Atmospheric science: Marine aerosols and iodine emissions, Nature, 433, E13–E13, 2005.
- McFiggans, G., Plane, J. M. C., Allan, B. J., Carpenter, L. J., Coe, H., and O'Dowd, C.: A modeling study of iodine chemistry in the marine boundary layer, J. Geophys. Res.-Atmos., 105, 14371–14385, 2000.
- Michel, R., Daraoui, A., Gorny, M., Jakob, D., Sachse, R., Tosch, L., Nies, H., Goroncy, I., Herrmann, J., Synal, H. A., Stocker, M., and Alfimov, V.: lodine-129 and iodine-127 in European seawaters and in precipitation from Northern Germany, Sci. Total Environ., 419, 151–169, 2012.
- Muramatsu, Y., Matsuzaki, H., Toyama, C., and Ohno, T.: Analysis of <sup>129</sup>I in the soils of Fukushima Prefecture: preliminary reconstruction of <sup>131</sup>I deposition related to the accident at Fukushima Daiichi Nuclear Power Plant (FDNPP), J. Environ. Radioact., 139, 344–350, 2015.
- Nielsen, S. P., Anderson, K. P., and Miller, A.: December 2011 Radioactivity in the Risø District January–June 2011, Risø Report, Risø-R-1800(EN), 2011.
- O'Dowd, C. D., Jimenez, J. L., Bahreini, R., Flagan, R. C., Seinfeld, J. H., Hämeri, K., Pirjola, L., Kulmala, M., Jennings, S. G., and Hoffmann, T.: Marine aerosol formation from biogenic iodine emissions, Nature, 417, 632–636, 2002.
- O'Dowd, C. D., Facchini, M. C., Cavalli, F., Ceburnis, D., Mircea, M., Decesari, S., Fuzzi, S., Yoon, Y. J., and Putaud, J.: Biogenically driven organic contribution to marine aerosol, Nature, 431, 676–680, 2004.
- Pechtl, S., Schmitz, G., and von Glasow, R.: Modelling iodide iodate speciation in atmospheric aerosol: Contributions of inorganic and organic iodine chemistry, Atmos. Chem. Phys., 7, 1381–1393, doi:10.5194/acp-7-1381-2007, 2007.
  - Pittauerová, D., Hettwig, B., and Fischer, H. W.: Fukushima fallout in Northwest German environmental media, J. Environ. Radioact., 102, 877–880, 2011.
- Qiao, J., Hansen, V., Hou, X., Aldahan, A., and Possnert, G.: Speciation analysis of <sup>129</sup>I, <sup>137</sup>Cs, <sup>232</sup>Th, <sup>238</sup>U, <sup>239</sup>Pu and <sup>240</sup>Pu in environmental soil and sediment, Appl. Radial. Isot., 70, 1698–1708, 2012.

**ACPD** 

15, 25139–25173, 2015

Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

I 

I 

Back Close

Printer-friendly Version

Full Screen / Esc



- Quinn, P. K., Bates, T. S., Schulz, K. S., Coffman, D. J., Frossard, A. A., Russell, L. M., Keene, W. C., and Kieber, D. J.: Contribution of sea surface carbon pool to organic matter enrichment in sea spray aerosol, Nat. Geosci., 7, 228–232, 2014.
- Raisbeck, G. M., Yiou, F., Zhou, Z. Q., and Kilius, L. R.: <sup>129</sup>I from nuclear fuel reprocessing facilities at Sellafield (U.K.) and La Hague (France); potential as an oceanographic tracer, J. Mar. Syst., 6, 561–570, 1995.
- Rose, D., Wehner, B., Ketzel, M., Engler, C., Voigtländer, J., Tuch, T., and Wiedensohler, A.: Atmospheric number size distributions of soot particles and estimation of emission factors, Atmos. Chem. Phys., 6, 1021–1031, doi:10.5194/acp-6-1021-2006, 2006.
- Saiz-Lopez, A. and Plane, J. M. C.: Novel iodine chemistry in the marine boundary layer, Geophys. Res. Lett., 31, L04112, doi:10.1029/2003GL019215, 2004.
- Saiz-Lopez, A., Gómez Martín, J. C., Plane, J. M. C., Saunders, R. W., Baker, A. R., Von Glasow, R., Carpenter, L. J., and McFiggans, G.: Atmospheric chemistry of iodine, Chem. Rev., 112, 1773–1804, 2012.
- Santos, F. J., López-Gutiérrez, J. M., García-León, M., Suter, M., and Synal, H. A.: Determination of <sup>129</sup>I/<sup>127</sup>I in aerosol samples in Seville (Spain), J. Environ. Radioact., 84, 103–109, 2005.
- Simpson, W. R., von Glasow, R., Riedel, K., Anderson, P., Ariya, P., Bottenheim, J., Burrows, J., Carpenter, L. J., Frieß, U., Goodsite, M. E., Heard, D., Hutterli, M., Jacobi, H.-W., Kaleschke, L., Neff, B., Plane, J., Platt, U., Richter, A., Roscoe, H., Sander, R., Shepson, P., Sodeau, J., Steffen, A., Wagner, T., and Wolff, E.: Halogens and their role in polar boundary-layer ozone depletion, Atmos. Chem. Phys., 7, 4375–4418, doi:10.5194/acp-7-4375-2007, 2007.
- Tölgyessy, J. (Ed.): Chemistry and Biology of Water, Air and Soil-Environmental Aspects, Studies in Environmental Science, Elsevier, Amsterdam-London-New York-Tokyo, 53, 1–858, 1993.

25

- Tsukada, H., Hara, H., Iwashima, K., and Yamagata, N.: The iodine content of atmospheric aerosols as determined by the use of a fluoropore filter for collection, Bull. Chem. Soc. Jpn., 60, 3195–3198, 1987.
- Turekian, K. K., Benninger, L. K., and Dion, E. P.: <sup>7</sup>Be and <sup>210</sup>Pb total deposition fluxes at New Haven, Connecticut and at Bermuda, J. Geophys. Res.-Oceans, 88, 5411–5415, 1983.
- Vogt, R., Sander, R., von Glasow, R., and Crutzen, P.: Iodine Chemistry and its Role in Halogen Activation and Ozone Loss in the Marine Boundary Layer: A Model Study, J. Atmos. Chem., 32, 375–395, 1999.

**ACPD** 

15, 25139–25173, 2015

Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

I 

I 

I 

Back Close

Printer-friendly Version

Full Screen / Esc



- Weather Wunderground: available at: http://www.wunderground.com/, last access: 26 May 2015.
- Wimschneider, A. and Heumann, K. G.: Iodine speciation in size fractionated atmospheric particles by isotope dilution mass spectrometry, Fresenius J. Anal. Chem., 353, 191–196, 1995.
- Xu, S., Freeman, S. P. H. T., Hou, X., Watanabe, A., Yamaguchi, K., and Zhang, L.: Iodine Isotopes in Precipitation: Temporal Responses to <sup>129</sup>I Emissions from the Fukushima Nuclear Accident, Environ. Sci. Technol., 47, 10851–10859, 2013.
  - Xu, S., Zhang, L., Freeman, S. P. H. T., Hou, X., Shibata, Y., Sanderson, D., Cresswell, A., Doi, T., and Tanaka, A.: Speciation of Radiocesium and Radioiodine in Aerosols from Tsukuba after the Fukushima Nuclear Accident, Environ. Sci. Technol., 49, 1017–1024, 2015.
  - Yi, P., Aldahan, A., Hansen, V., Possnert, G., and Hou, X.: Iodine Isotopes (<sup>129</sup>I and <sup>127</sup>I) in the Baltic Proper, Kattegat, and Skagerrak Basins, Environ. Sci. Technol., 43, 903–909, 2011.
  - Yi, P., Aldahan, A., Possnert, G., Hou, X., Hansen, V., and Wang, B.: <sup>127</sup>I and <sup>129</sup>I Species and Transformation in the Baltic Proper, Kattegat, and Skagerrak Basins, Environ. Sci. Technol., 46, 10948–10956, 2012.
  - Zhang, L., Hou, X., and Xu, S.: Speciation Analysis of <sup>129</sup>I and <sup>127</sup>I in Aerosols Using Sequential Extraction and Mass Spectrometry Detection, Anal. Chem., 87, 6937–6944, 2015.

15

**ACPD** 

15, 25139–25173, 2015

Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

I 

▶ I

Back Close

Full Screen / Esc



Printer-friendly Version

Interactive Discussion

**Table 1.** Sampling information of aerosols collected at Risø, Denmark in 2011 and 2014. Data of <sup>131</sup>I, <sup>7</sup>Be and <sup>210</sup>Pb in the aerosol samples are adopted from the DTU Nutech report (Nielsen et al., 2011). The reference time was the mid-point of the sampling period, and analytical uncertainties were 5 % for <sup>131</sup>I, and within 1 % for <sup>7</sup>Be and <sup>210</sup>Pb).

Sample No	Sampling date	Air flux, m <sup>3</sup>	Air flux, m <sup>3</sup> h <sup>-1</sup>	Weight, g	<sup>131</sup> l, μBq m <sup>-3</sup>	<sup>7</sup> Be, μBq m <sup>-3</sup>	<sup>210</sup> Pb, μBq m <sup>-3</sup>
AE11-1	31 Mar-04 Apr 2011	88 833	2757	72.5	205	1925	66
AE11-2	04-07 Apr 2011	64 339	2751	79.2	218	1482	47
AE11-3	07–11 Apr 2011	55 911	1744	79.5	147	1482	47
AE11-4	11–14 Apr 2011	27 083	1096	70.9	110	2750	172
AE11-5	14–18 Apr 2011	48 317	1505	77.9	58.3	2750	172
AE11-6	18–26 Apr 2011	101 400	1593	80.8	20.9	4528	249
AE11-7	26 Apr-02 May 2011	54 600	1117	77.7	14.8	4027	253
AE14-1	08-15 Dec 2014	37917	2727	21.7	< D.L.	1499	54.9

15, 25139–25173, 2015

Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Back Close

Full Screen / Esc

Printer-friendly Version



**Table 2.** Analytical results of chemical species of <sup>127</sup>I and <sup>129</sup>I in aerosols collected from Risø, Denmark during spring 2011 and winter 2014.

Sample	TI	WSI	lodate	lodide	NSI	RII
<sup>127</sup> l conc	<sup>127</sup> I concentration, ng m <sup>-3</sup>					
AE11-1	1.187 ± 0.062	$0.152 \pm 0.002$	ND	$0.158 \pm 0.008$	$0.340 \pm 0.019$	$0.606 \pm 0.042$
AE11-2	$1.797 \pm 0.116$	$0.141 \pm 0.01$	$0.022 \pm 0.012$	$0.119 \pm 0.006$	$0.556 \pm 0.037$	$0.977 \pm 0.049$
AE11-3	$1.927 \pm 0.115$	$0.264 \pm 0.004$	ND	$0.259 \pm 0.013$	$0.813 \pm 0.027$	$0.983 \pm 0.05$
AE11-4	$2.480 \pm 0.129$	$0.258 \pm 0.013$	ND	$0.276 \pm 0.014$	$0.825 \pm 0.049$	$1.664 \pm 0.085$
AE11-5	$2.027 \pm 0.104$	$0.221 \pm 0.011$	ND	$0.237 \pm 0.012$	$0.638 \pm 0.036$	$1.308 \pm 0.087$
AE11-6	$1.506 \pm 0.112$	$0.305 \pm 0.007$	ND	$0.327 \pm 0.017$	$0.624 \pm 0.033$	$0.585 \pm 0.03$
AE11-7	$1.041 \pm 0.055$	$0.316 \pm 0.019$	$0.033 \pm 0.024$	$0.283 \pm 0.014$	$0.377 \pm 0.027$	$0.343 \pm 0.018$
AE14-1	$2.356 \pm 0.127$	$0.618 \pm 0.019$	ND	$0.739 \pm 0.039$	$0.929 \pm 0.057$	$0.802 \pm 0.041$
$^{129}$ I concentration, $\times 10^5$ atoms m <sup>-3</sup>						
AE11-1	28.57 ± 1.11	$3.07 \pm 0.08$	ND	$3.60 \pm 0.43$	$8.33 \pm 0.84$	$14.14 \pm 0.64$
AE11-2	$72.98 \pm 5.64$	$4.72 \pm 0.41$	ND	$4.63 \pm 0.27$	$20.55 \pm 1.43$	$39.94 \pm 1.64$
AE11-3	$25.60 \pm 0.98$	$3.60 \pm 0.15$	ND	$4.11 \pm 0.72$	$10.80 \pm 0.40$	$13.49 \pm 2.83$
AE11-4	$47.27 \pm 1.55$	$4.78 \pm 0.23$	ND	$5.27 \pm 0.57$	$13.43 \pm 1.24$	$30.26 \pm 4.24$
AE11-5	$43.81 \pm 1.28$	$5.55 \pm 0.27$	ND	$5.91 \pm 0.74$	$12.58 \pm 0.46$	$27.51 \pm 1.28$
AE11-6	$12.73 \pm 0.42$	$3.26 \pm 0.10$	ND	$3.26 \pm 0.30$	$4.60 \pm 0.49$	$5.91 \pm 0.36$
AE11-7	$11.31 \pm 0.43$	$3.34 \pm 0.11$	ND	$4.08 \pm 0.24$	$3.36 \pm 0.25$	$4.27 \pm 0.74$
AE14-1	$97.00 \pm 3.01$	$26.85 \pm 0.65$	ND	$30.12 \pm 1.68$	$34.74 \pm 0.80$	$39.01 \pm 1.49$
$^{129}$ I / $^{127}$ I atomic ratio, × $10^{-8}$						
AE11-1	$50.78 \pm 3.31$	42.73 ± 1.28	ND	$48.04 \pm 6.23$	$51.70 \pm 5.98$	49.19 ± 4.07
AE11-2	$85.70 \pm 8.63$	$70.73 \pm 8.06$	ND	$81.93 \pm 6.36$	$78.03 \pm 7.50$	$86.21 \pm 5.62$
AE11-3	$28.03 \pm 1.99$	$28.82 \pm 1.26$	ND	$33.49 \pm 6.13$	$28.04 \pm 1.40$	$28.94 \pm 6.25$
AE11-4	$40.21 \pm 2.48$	$39.11 \pm 2.72$	ND	$40.29 \pm 4.84$	$34.34 \pm 3.78$	$38.37 \pm 5.72$
AE11-5	$45.60 \pm 2.70$	$52.85 \pm 3.66$	ND	$52.50 \pm 7.16$	$41.58 \pm 2.79$	$44.36 \pm 3.61$
AE11-6	$17.84 \pm 1.46$	$22.59 \pm 0.89$	ND	$21.08 \pm 2.24$	$15.56 \pm 1.85$	$21.32 \pm 1.69$
AE11-7	$22.92 \pm 1.48$	$22.30 \pm 1.55$	ND	$30.43 \pm 2.38$	$18.81 \pm 1.93$	$26.28 \pm 4.73$
AE14-1	$86.84 \pm 5.40$	$91.70 \pm 3.55$	ND	$86.03 \pm 6.63$	$78.84 \pm 5.20$	$102.63 \pm 6.51$

Analytical uncertainties are  $1\sigma$ .

# **ACPD**

15, 25139–25173, 2015

Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.





**Table 3.** Possible pathways of formation of iodide by reduction of sulfur compounds.

Phase	Reactions	Equations	References
Gas	$DMS + OH \rightarrow SO_2$	(1)	Chatfield and Crutzen (1990)
	$DMS + NO_3 \rightarrow SO_2$		
Gas/Aerosol	$SO_2 + H_2O \rightarrow HSO_3^-$	(2)	
	$SO_2 + H_2O \rightarrow SO_3^{2-3}$		
Gas-Aerosol interface	$I + HSO_3^- \rightarrow I^- + SO_4^{2-}$	(3)	
	$I + SO_3^{2-} \rightarrow I^- + SO_4^{2-}$		
Aerosol	$HOI + HSO_3^-/SO_3^{2-} \rightarrow I^- + SO_4^{2-}$	(4)	Saiz-Lopez et al. (2012)
	$HOI + SO_3^{2-} \rightarrow I^{-} + SO_4^{2-}$	. ,	
	$HOI_2 + HSO_3^-/SO_3^{2-} \rightarrow I^- + SO_4^{2-}$		
	$HOI_{2} + SO_{3}^{2-} \rightarrow I^{-} + SO_{4}^{2-}$		

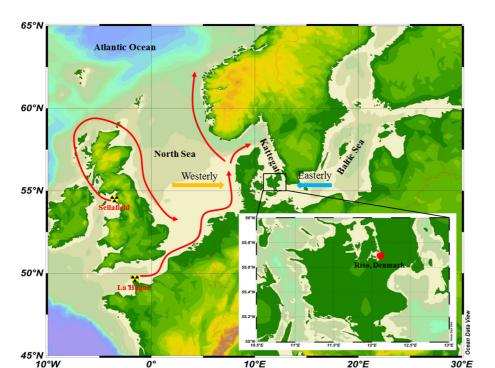
15, 25139-25173, 2015

Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Printer-friendly Version





**Figure 1.** Map showing the sampling site (red dot in the inset) at Risø, Denmark for aerosol collection. The two nuclear reprocessing plants (radioactivity labels), Sellafield (United Kingdom) and La Hague (France) have released a large quantity of <sup>129</sup>I to the marine and atmospheric environment. This anthropogenic <sup>129</sup>I discharged to marine systems has been transported to the North Sea, Baltic Sea, Norwegian Sea, and beyond, by ocean currents (red lines). Yellow and blue arrows show the westerly wind and easterly wind, respectively.

15, 25139-25173, 2015

Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures







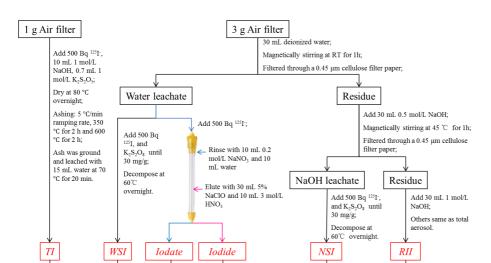




Full Screen / Esc

Printer-friendly Version





**Figure 2.** Schematic diagram of the analytical procedure for the determination of <sup>127</sup>I and <sup>129</sup>I species in aerosols. TI for total iodine, WSI for water-soluble iodine, NSI for NaOH-soluble iodine and RII for residual-insoluble iodine.

Add Cs+ as internal standard;

Dilute with 1% NH3·H3O

<sup>127</sup>I measurement by ICP-MS

Add iodine carrier:

Solvent extraction

129I measurement by AMS

**ACPD** 

15, 25139–25173, 2015

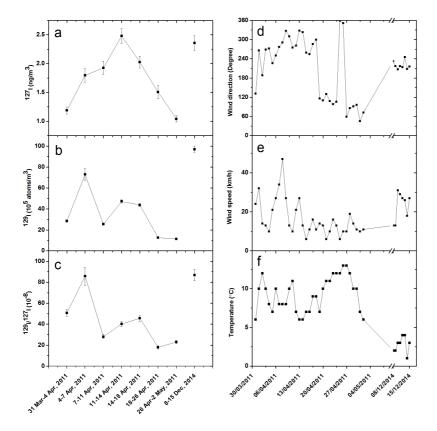
Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Printer-friendly Version

Full Screen / Esc





**Figure 3.** Variation of  $^{127}$ I concentrations **(a)**,  $^{129}$ I concentrations **(b)**,  $^{129}$ I  $/^{127}$ I ratios in aerosols **(c)**, wind direction **(d)**, wind speed **(e)** and temperature **(f)** during the sampling period. The historical meteorological data, including temperature, wind direction and wind speed were obtained from the observation station of Hangarvej in Roskilde, Denmark (55.594° N 12.128° E) based on 41 m a.s.l. (above sea level) (Weather Wunderground webpage, 2015).

15, 25139-25173, 2015

Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

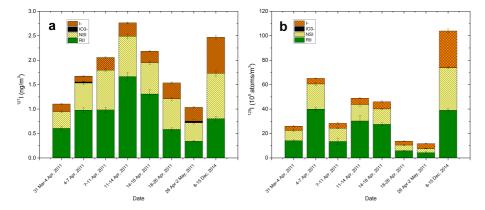
I ✓ ▶I

Back Close

Full Screen / Esc

Printer-friendly Version





**Figure 4.** Concentrations of iodine species in the aerosol samples for <sup>127</sup>I (a) and <sup>129</sup>I (b), indicating that NSI and RII are major iodine species and that iodide is the dominant fraction of water-soluble iodine.

15, 25139–25173, 2015

Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

I ◀ ▶I

■ Back Close

Full Screen / Esc

Printer-friendly Version





15, 25139–25173, 2015

Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

**ACPD** 

L. Y. Zhang et al.



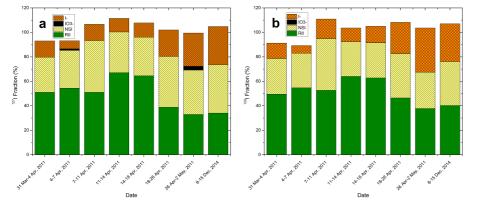


Figure 5. Distribution of iodine species in the aerosol samples for <sup>127</sup>I (a) and <sup>129</sup>I (b), compared with the measured concentrations of total <sup>127</sup>I and <sup>129</sup>I in the original aerosols.



15, 25139–25173, 2015

**ACPD** 

# Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.







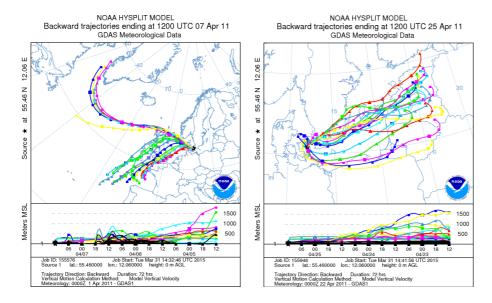


Figure 6. The 72 h (3 days) air mass back trajectories for starting altitudes of 0 m a.g.l. (above ground level) calculated from the FNL database of the National Ocean and Atmospheric Administration (NOAA) and simulated by using the Hybrid Single-Particle Langrangian Integrated Trajectory (HY-SPLIT) model. 4-7 April 2011 (left) and 21-25 April 2011 (right).



15, 25139-25173, 2015

**ACPD** 

# Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.





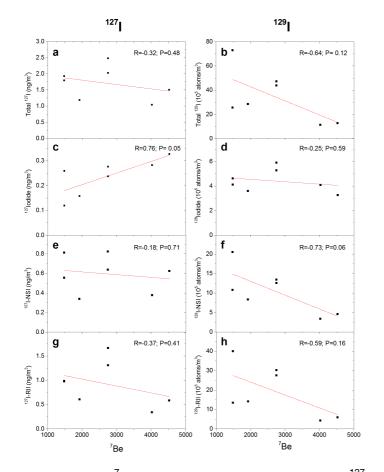
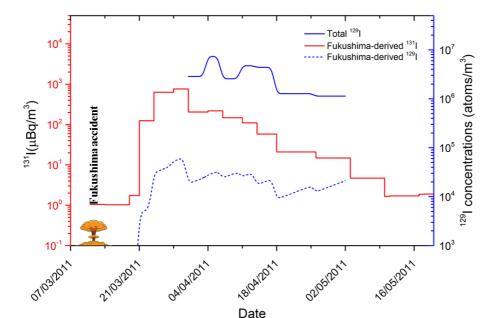


Figure 7. Correlation of aerosol <sup>7</sup>Be with iodine species including total <sup>127</sup>I (a) and <sup>129</sup>I (b),  $^{127}I^{-}$  (c),  $^{129}I^{-}$  (d), NSI for  $^{127}I$  (e) and  $^{129}I$  (f), as well as RII for  $^{127}I$  (g) and  $^{129}I$  (h).



**Figure 8.** <sup>131</sup>I radioactivity (red), <sup>129</sup>I concentrations (blue) in aerosols from Risø, Denmark after the Fukushima accident (Nielsen et al., 2011). The Fukushima-derived <sup>129</sup>I concentrations are calculated based on an <sup>129</sup>I / <sup>131</sup>I atomic ratio of  $16.0 \pm 2.2$  deduced from Fukushima-affected aerosol samples (Xu et al., 2015).

15, 25139–25173, 2015

Speciation of <sup>127</sup>I and <sup>129</sup>I in atmospheric aerosols at Risø

L. Y. Zhang et al.

Title Page

Full Screen / Esc

Back

Close

Printer-friendly Version

Interactive Discussion

