1	Mixing state and sources of submicron regional background aerosols in the
2	North Qinghai-Tibetan Plateau and the influence of biomass burning
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15 Abstract: Transmission electron microscopy (TEM) was employed to obtain morphology, size, composition, and mixing state of background aerosols with diameter less than 1 µm in the North 16 Qinghai-Tibetan Plateau (QTP) during 15 September to 15 October, 2013. Individual aerosol 17 18 particles mainly contained secondary inorganic aerosols (SIA-sulfate and nitrate) and organics during clean periods (PM_{2.5} mass concentration less than 2.5 μ g/m³). The presence of K-Na-Cl 19 20 associated with organics and an increase of soot particles suggest that an intense biomass burning event caused the highest PM_{2.5} concentrations (> 30 μ g/m³) during the study. A large number 21 22 fraction of the fly ash-containing particles (21.73%) suggests that coal combustion emissions in the QTP significantly contributed to air pollutants at the medium pollution level (PM2.5: 10-30 23 $\mu g/m^3$). We concluded that emissions from biomass burning and from coal combustion both 24 constantly contribute to anthropogenic particles in the QTP atmosphere. Based on size 25 26 distributions of individual particles in different pollution levels, we found that gas condensation on 27 existing particles is an important chemical process for the formation of SIA with organic coating. 28 TEM observations show that refractory aerosols (e.g., soot, fly ash, and visible organic particles) 29 likely adhere to the surface of SIA particles larger than 200 nm due to coagulation. Organic 30 coating and soot on surface of the aged particles likely influence their hygroscopic and optical properties in the QTP, respectively. To our knowledge, this study reports the first microscopic 31 32 analysis of fine particles in the background QTP air.

34 **1. Introduction**

With an immense area (about 2,400,000 km²) and mean elevation of more than 4000 m above 35 sea level, the Tibetan Plateau (TP), called the "ridge of the world and third polar", plays a key role 36 37 in Asian climatology, especially the formation of monsoons (Lau et al., 2006). Climate on the TP 38 has warmed 0.3 $^{\circ}$ C per decade over the past three decades, which is twice the rate of observed 39 global warming (Xu et al., 2009). Anthropogenic aerosols and their ice and cloud condensation 40 nuclei (CCN) directly or indirectly leaded to the brightening and dimming phenomenon before 41 and after the 1980s in the TP (You et al., 2010). However, light absorbing carbonaceous aerosol 42 particles (i.e., black carbon (BC) and brown carbon (BrC)) can warm the troposphere 43 (Ramanathan and Carmichael, 2008) and accelerate glacier retreat (Xu et al., 2009). Both the 44 radiative effects of aerosols and their role in cloud forming processes depend on their number, size, 45 chemical properties, and mixing state.

46 As a consequence, a better understanding of climate change can be achieved by characterizing 47 the TP aerosols. Quite a few aerosol measurements have been conducted in the TP. Ion 48 composition records from a shallow ice core (Zheng et al., 2010) and black soot in the Tibetan 49 glaciers (Xu et al., 2009) both showed that anthropogenic aerosols have increased significantly in 50 the most recent 50 years in the TP. Li et al. (2013a) obtained aerosol components such as 61% mineral, 3% ammonium, 4% nitrate, 18% sulfate, 2% black carbon, and 12% organic matter in 51 PM_{2.5} at a concentration of 21.5 µg m⁻³ during summer of 2010 at Qinghai Lake (36 59 N, 52 53 99 54'E; 3200 m a.s.l.) in the northeastern part of the TP. Coal burning and biomass burning were the major sources for anthropogenic aerosols. Xu et al. (2014) showed that the PM_{2.5} mass 54 concentration of 9.5 \pm 5.4 µg m⁻³ during a year-long study at the Qilian Shan Station (39.50 N, 55 56 96.51 E; 4180 m a.s.l.), a remote site on the northeast edge of the Tibetan Plateau, and their water soluble ionic species were dominated by SO_4^{2-} (39%), CO_3^{2-} (19%), Ca^{2+} (16%), NO_3^{-} (10%), and 57 NH4⁺ (6%). The study suggests anthropogenic aerosol and natural mineral dust from the Gobi 58 59 desert together contribute to the particle loading in this remote air. Li et al. (2007) also found 60 anthropogenic ions from residential combustion emissions in precipitation samples at Nam Co station of the central TP. In addition, long-range transport of pollutants from eastern and 61 northwestern China and northern India can contribute black carbon and other air pollutants to the 62

TP region (Cao et al., 2010;Wang et al., 2010;Engling et al., 2011;Kopacz et al., 2011;Lu et al., 2012;Xu et al., 2013;Zhao et al., 2013;Cong et al., 2015;Duo et al., 2015). Although anthropogenic sources make but a minor contribution to the background TP atmosphere, these anthropogenic aerosols significantly enhanced aerosol optical properties (AOD) in the central and northeastern TP in summer (Cong et al., 2009b;Che et al., 2011;Xia et al., 2011). Therefore, study of the composition and sources of aerosol particles in the TP is necessary to understand their effects on the optical, CCN or IN activity.

70 The TP has various geographic and natural environmental ecosystems such as mountains, lake 71 basins, deserts, forests, and grasslands. Previous sampling sites have included mountain forests, 72 lake basins, and urban areas in the TP (Li et al., 2007; Che et al., 2011; Engling et al., 2011; Li et al., 73 2013a;Xu et al., 2014). Grassland is one of the largest geomorphology in the TP. There are only a 74 few herdsmen and farmers living in the vast grasslands of the norther TP. Air pollutants from 75 anthropogenic and natural sources can be easily transported over low bushes in the grasslands 76 under high wind speed in north TP (Figure S1). However, fine aerosols in the troposphere have not 77 been studied over the vast grassland in the northern TP. In this study, we collected samples at a 78 national station of background atmospheric monitoring (NSBAM) on the top of Moshidaban 79 mountain of Menyuan county in Qinghai province (37°35.370'N; 101°17.329'E; elevation: 3295 80 m), which is within the northern part of the Qinghai-Tibet Plateau (QTP) (Figure 1). The NSBAM 81 is 180 km north of Qinghai Lake (Li et al., 2013a) and 160 km from Waliguan station (Che et al., 82 2011). Only a few herdsmen live in the grassland but many agricultural activities (e.g., growing 83 hulless barley and rape) mainly occur around Qinghai Lake and Menyuan County.

84 Although many atmospheric scientists have commented on the probable aerosol impacts on 85 climate and monsoon (Lau et al., 2006), atmospheric observations are still very limited because of 86 the unique geographic and natural environment, electric supply problems, high maintenance costs 87 for instruments, and lack of skilled operators. Because of the immensity of the TP, these previous 88 studies, quite scattered in diverse locations, are insufficient to adequately characterize the aerosols 89 throughout this vast region. In addition, the mixing state of individual particles has not been 90 examined, and only a few studies of particle types and soot have been conducted through electron 91 microcopy (Zhang et al., 2001a;Cong et al., 2009a). Understanding the mixing state of individual 92 particles sheds light on their source, ageing processes, optical, and hygroscopic properties (Posfai

93 and Buseck, 2010:Li et al., 2015). In the present study, high-resolution transmission electron 94 microscopy (TEM) is employed to study the mixing state and composition of individual submicron particles with diameters $< 1 \ \mu m$. The pollution levels have been evaluated and 95 identified through continuous gaseous and particulate instruments at the sampling site. The 96 anthropogenic sources were further identified based on particle types in the QTP. 97

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2. Experimental methods 99

100 2.1 Aerosol sampling

101 Aerosol particles were collected onto copper TEM grids coated with carbon film (carbon type-B, 300-mesh copper, Tianld Co., China) by a two-stage impactor with a 1-mm-diameter jet 102 nozzle and a 0.5-mm-diameter jet nozzle and an air flow of 1.0 L min⁻¹. Both stages have a 50% 103 104 collection efficiency, the first at 0.80 μ m and the second at 0.20 μ m, with an atmospheric pressure of 69 kpa, a temperature of 283.5K, and an assumed particle density of 2 g/cm³. Sampling times 105 106 varied from 30 s to 15 min, depending on particle loading. After collection, each sample was 107 placed in a sealed, dry plastic tube and stored in a desiccator at 25 °C and 20 ± 3% RH to minimize 108 exposure to ambient air and preserve it for analysis. Altogether, 70 individual samples were 109 collected at the NABAM with the elevation at 3295 m, of which we analyzed the fine particles 110 collected on only the second stage.

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2.2 Individual particle analysis

112 21 aerosol particle samples collected on TEM grids were analyzed with a JEOL JEM-2100 TEM operated at 200 kV. The details about the aerosol collection were marked in Figure 2. 113 114 Elemental composition was determined semi-quantitatively by using an energy-dispersive X-ray 115 spectrometer (EDS) that can detect elements heavier than C. Cu was excluded from the analyses 116 because the TEM grids are made of Cu. The distribution of aerosol particles on TEM grids was not 117 uniform, with coarser particles occurring near the center and finer particles occurring on the 118 periphery. Therefore, to ensure that the analyzed particles were representative, five areas were 119 chosen from the center and periphery of the sampling spot on each grid. Every particle in the selected area was analyzed. EDS spectra were collected for 15 s in order to minimize radiation 120 exposure and potential beam damage. To better understand the properties of internally mixed 121

122 aerosol particles, we also analyzed the composition of different components of individual particles, 123 such as coatings, inclusions, and aggregations. The sampling was controlled to avoid coagulation 124 on the substrate during sampling. The projected areas of individual particles were determined by 125 the iTEM software (Olympus soft imaging solutions GmbH, Germany), the standard image 126 analysis platform for electron microscopy. Altogether 4218 particles in these samples were 127 measured for statistical analyses.

128 2.3 Particle size measurement

129 Atomic force microscopy (AFM) with a tapping mode analyzed aerosol particles under ambient conditions. AFM, a digital NanoscopeIIIa Instrument, can detect the 130 three-dimensional morphology of particles. The AFM settings contain imaging forces 131 between 1 and 1.5 nN, scanning rates between 0.5 and 0.8 Hz, and scanning range sizes at 10 132 133 µm with a resolution of 512 pixels per length. After the AFM analysis, composition of the same particles was confirmed by TEM, with 194 fine aerosol particles analyzed by this 134 method. The NanoScope analysis software can automatically obtain bearing area (A) and 135 bearing volume (V) of each analyzed particle according to the following formula. 136

$$A = \pi r^2 = \pi \times \left(\frac{d}{2}\right)^2 = \frac{\pi d^2}{4} \to d = \sqrt{\frac{4A}{\pi}}$$
(1)

$$V = \frac{4}{3}\pi r^{3} = \frac{4}{3} \times \frac{\pi D^{3}}{8} = \frac{\pi D^{3}}{6} \to D = \sqrt[3]{\frac{6V}{\pi}}$$
(2)

Where d is the equivalent circle diameter (ECD) and D is the equivalent volume diameter (EVD). By plotting the ECD against the EVD (Figure 3), we also obtain the relationship between them as EVD=0.64ECD As a result, ECD (d) of individual aerosol particles measured from the iTEM software can be further converted into EVD (D) based on this relationship. In this study, we only considered fine aerosol particles with equivalent volume diameter smaller than 1 μ m, where the correlation between the ECD and EVD is especially good (Figure 3).

144 2.4 FLEXPART particle dispersion model

Air mass history was determined using the Lagrangian particle dispersion model FLEXPART (version 9.02; (Stohl et al., 1998)). FLEXPART simulates the release of thousands of passive tracer air parcels at the specific location, advecting them backwards in

time, providing a representation of the spatial distribution of the air mass at an upwind time 148 referred to as a "retroplume". FLEXPART was driven with 6 h meteorology data from NCEP 149 Climate Forecast System Version 2 (NCEP - CFSv2), including land cover, temperature, 150 relative humidity, and three- dimensional wind, in 37 levels with a resolution of $0.5^{\circ} \times 0.5^{\circ}$. 151 In this study, the modeling periods were 72h each simulations and it simulated 4 times each 152 153 day (beginning at 00:00, 06:00, 12:00, and 18:00, respectively) from 0000 UTC 10 September 2013 to 0000UTC 16 October 2013. Every simulation containing 10000 particles released at 154 155 the beginning over an altitude range of 3395 m a.s.l to 3995 m a.s.l and the model outputs were recorded every 3-h. The output data of each simulation were combined together to make 156 157 the figure 1.

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2.5 PM_{2.5}, trace gases, BC, and meteorological measurements

159 Thermo TEOM 1405 $PM_{2.5}$ and PM_{10} continually measure the particulate mass concentrations 160 in one-hour averages. Gaseous air pollutants were measured continuously from 1 September to 15 161 October, 2013: O₃ by a UV photometric analyzer (Teledyne Instruments, Model 400EU); SO₂ by a pulsed UV fluorescence analyzer (M100EU), CO by a non-dispersive infrared analyzer (M300EU), 162 163 and NO and NO₂ by a commercial chemiluminescence analyzer (M200EU), with the 164 concentrations being recorded in five-minute averages. BC concentrations were measured by an 165 Aethalometer and were recorded in one hour averages. In addition, the non-refractory submicron 166 aerosol species including organics, sulfate, nitrate, ammonium, and chloride were measured in-situ 167 by an Aerodyne Aerosol Chemical Speciation Monitor (ACSM) (Du et al., 2015). Wind direction, 168 wind speed, relative humidity (RH), and temperature were measured and recorded in one hour 169 averages. The time-series meteorological data, shown in Figure S1, and the $PM_{2.5}$ and gaseous 170 concentrations, were provided by the NSBAM.

171 The average CO mixing ratio is 44.78 ppb at the NSBAM, much lower than the 149 ppb at 172 Waliguan in the summer of 2006, that is the site of the observation station of the World 173 Meteorological Organization's (WMO) Global Atmospheric Watch (GAW) (Xue et al., 2011). This 174 contrast shows that the NSBAM adequately represents background conditions in the expansive 175 grasslands of the northern TP. Time-series concentration variations of six pollutants (i.e., PM_{2.5}, 176 BC, SO_2 , NO_x , CO, and O_3) show that their highest concentrations occurred from 15 September to 177 25 September, 2013, and from 11 October to 15 October, 2013, (Figure 2). Therefore, we 178 considered these two periods as typical high-pollution events. Table 1 shows that five pollutants' concentrations (i.e., PM_{2.5}, BC, SO₂, NO_x, and CO) were higher during these pollution events than 179 180 in the intervening cleaner period; O_3 concentration was close. When the combustion-tracing CO and NO_x concentrations increase, O_3 mixing ratios generally decrease in the QTP due to 181 photochemical consumption (Xue et al., 2011). The primary BC concentrations during the two 182 pollution events were 17% and 81% higher than the intervening cleaner period, respectively 183 (Table 1). PM_{2.5} concentration at 17.06 μ g/m³ at the NSBAM is slightly lower than the 21.5 μ g/m³ 184 185 in the cleaner Qinghai-lake area in the summer of 2010 (Li et al., 2013a). The air mass back trajectories during the sampling period commonly came from the northwestern TP and crossed the 186 187 Qinghai-lake area into the northern TP (Figure 1). The air masses during the pollution events 188 adequately represented highly aged and processed long-range transported ambient aerosols in the TP. Figure 2 further displays aerosol collection time under three different PM_{2.5} levels, e.g., PM_{2.5} 189 \geq 30 µg/m³, PM_{2.5} between 10-30 µg/m³, PM_{2.5} < 10 µg/m³, which represent high pollution level, 190 191 medium pollution level, and clean pollution level.

192 **3. Results**

193 **3.1** Major fine aerosol particles and mixing states

194 Based on elemental composition and morphology, aerosol particles were classified into six 195 major categories: mineral dust, K-Na-Cl, fly ash, secondary inorganic aerosol (SIA) containing 196 ammoniated sulfates and nitrates, organics, and soot (i.e., BC) (Figures 4-5). For example, mineral 197 dust particles normally display irregular shapes and fly ash particles are spherical, although they 198 both have similar compositions such as Si and Al. This detailed particle classification scheme is 199 described in our previous studies (Li et al., 2014b;Li et al., 2015). The nanosized metal particles 200 which have been frequently detected in ambient aerosols in East China (Li et al., 2013b;Li et al., 201 2013c) were absent in the Qinghai-Tibet plateau. Mixing properties among the six types of 202 particles were characterized in detail. TEM observations indicate that SIA and organics coexisted 203 in individual fine particles and that organic carbon (OC) coated (e.g., Figures 4d and 5a) or 204 homogeneously mixed (e.g., Figure 6) with these SIA particles. In other words, OC occurred on 205 surfaces of the SIA particles. In addition, we found that many SIA-OC particles had visible 206 inclusions such as mineral dust, fly ash, OC, and soot particles. Identification of the refractory 207 inclusions in internally mixed particles enables one to trace particle sources and their history in the aging air mass. Their mixing properties consist mostly of SIA-soot-OC (e.g., Figure 4c), SIA-fly

ash-soot (e.g., Figure 5d), SIA-fly ash-OC (visible) (e.g., Figure 5d), SIA-fly ash (e.g., Figure 5c),

SIA-mineral, SIA-visible OC (e.g., Figure 5a). Therefore, SIA and OC in aerosol particles in the
Qinghai-Tibet Plateau were the two most important influences of the mixing state of primary
particles.

213 **3.2** Size distribution of aerosol particles

In this section, we describe the size distribution of individual particles with their diameters from 40 nm to 1 µm in different pollution levels. 684 particles collected during clean periods show a median diameter of 230 nm. Figure 7a shows that the size distribution of particles without inclusions determines ambient particle size.

1214 particles during high pollution levels have a median diameter of 260 nm (Figure 7b). 218 219 Particles with inclusions and particles without inclusions have median diameters of 300 nm and 220 230 nm, respectively. Figure 7b shows particles with and without inclusions jointly determine the 221 particle size distribution during high pollution levels. We noticed that the size distribution of 222 inclusions in SIA displays a median diameter of 150 nm. A similar pattern of size distribution of 223 2355 particles occurred in medium pollution levels (Figure 7c). The median diameters of total 224 individual particles, particles with inclusions, particle without inclusions, and inclusions are 290 225 nm, 340 nm, 250 nm, and 150 nm, respectively. Therefore, the inclusions (e.g., mineral, fly ash, 226 soot, and spherical OC particles) significantly enhanced particle sizes in the background air once 227 they were internally mixed with sulfates.

228

229 4. Discussion

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4.1 Identification of the pollution events

TEM observations show that individual particle types display large differences under three different $PM_{2.5}$ levels: $\geq 30 \ \mu g/m^3$ (high pollution level), 10-30 $\mu g/m^3$ (medium pollution level), and $< 10 \ \mu g/m^3$ (clean pollution level) (Figure 3). Aerosol particles collected in clean periods mainly contained SIA and OC. Figure 6 shows that individual SIA particles were commonly coated by OC. Consistently, the ACSM measurement showed that SO_4^{2-} and OC were the main components in PM₁, accounting for 33-36% and 34-48% in mass at the sampling site, respectively (Du et al., 2015). Figure 8 presents the composition of all the analyzed individual particles in the three pollution levels. In the clean period, we found only a few anthropogenic particles such as flyash, soot, or their mixed particles with their contributions being less than 10% (Figure 8a).

240 The increase of KCl and soot particles is suggestive of an intense biomass burning event at the 241 background site (Li et al., 2003). In this study, abundant K-Na-Cl particles and soot-containing particles (e.g., Figure 4) only occurred in pollution period 1 and period 2 which have been 242 243 indicated as biomass burning (Du et al., 2015). Because the K-Na-Cl particles associated with OC occurred only in the short pollution periods and are smaller than typical sea-salt or soil 244 245 particles (mostly $>1 \mu m$), it is unlikely that they originated from natural sources such as saline Qinghai Lake and desert. Our field experimental investigations showed that a few farmers burned 246 247 cole flowers and highland barley during the autumn harvest season, which are main season crops in the QTP. In addition, the burning of Yak dung for residential heating likely caused the high 248 PM_{2.5} in 11-15 October (Du et al., 2015). 249

250 The distribution of different particle types in the medium pollution level is similar to the high 251 pollution level, except for the absence of K-Na-Cl particles. Figure 8c shows that a large amount 252 of fly ash-containing particles occurred in medium pollution level. Fly ash is generally considered 253 as a reliable fingerprint of coal combustion in residential cooking, power plants, and industrial activities (Li and Shao, 2009). The fly ash-containing particles increase from 11.40% in the high 254 255 pollution level to 21.73% in the medium one, but soot-containing particles decrease from 38.40% 256 to 25.87%. This result indicates that coal combustion emissions in the QTP significantly affected 257 the background air quality. Compared to individual particles in polluted East China, absence of nanometer metal particles in the QTP suggests that there are no large heavy metal-related 258 259 industrial emissions in the area under air mass back trajectories (Figure 1). The China Energy 260 Statistical Yearbook of 2013 shows coal combustion occurs in power plants (48.5%), heavy 261 industries (36.4%), and house cooking/heating in rural areas (8.6%) in Qinghai province, 262 particularly nearby the large cities such as Xi'ning (Wen et al., 2013). Although we didn't find any 263 K-Na-Cl particle in the samples under medium pollution level, 50% of SIA (OC coating) and SIA-soot particles containing minor K (Figure 4c-d) frequently occurred in the samples. During 264 265 long-range transport, once coagulation and condensation of ammoniated sulfates and sulfuric acid 266 in biomass burning particles increase, K-rich particles may transform into sulfur-rich particles 267 with certain amounts of K (Li et al., 2014b). Therefore, SIA particles containing minor K suggest

that biomass burning emissions likely contributed to the QTP aerosols on a more or less constant basis. A similar result has been obtained through the analysis of organic species in $PM_{2.5}$ at Qinghai-lake (Li et al., 2013a). Therefore, we conclude that emissions from biomass burning and from coal combustion significantly contribute to the formation of anthropogenic fine particles in the atmosphere over the QTP.

273 4.2 Regional effects of biomass burning and industrial emissions

274 The previous studies proved that trace gases such as SO₂, NO_x, and volatile organic 275 compounds (VOCs) from anthropogenic and natural sources had been transported long distances in the QTP and were transformed into secondary aerosol particles (Xue et al., 2011;Li et al., 276 277 2013a;Du et al., 2015;Xu et al., 2015). For example, Du et al. (2015) suggested that oxygenated 278 organic aerosols from anthropogenic sources and biomass burning transported over a long distance 279 to the sampling site in the QTP. Because of measurement limitations, there is no research about 280 refractory aerosol particles (e.g., mineral, fly ash, and soot) in fine particles and primary organic 281 particles. In contrast, TEM observations can adequately characterize these refractory particles 282 internally mixed with SIA based on their unique morphology and composition (Figures 4-5). We 283 identified abundant refractory particles at the three pollution levels at the regional background site. 284 Therefore, the nanosized refractory particles and trace gases from various anthropogenic sources 285 including biomass burning can together be transported long distances. The FLEXPART simulation 286 shows that these anthropogenic particles mainly originated from biomass burning between the 287 Qinghai lake and Menyuan county and heavy industries and coal-fired power plants in western 288 areas of Xining city (Figure 1).

289 4.3 Mixing mechanisms of aged aerosol particles

290 In this study, individual particle clearly containing more than two types of aerosol components 291 (e.g., mineral dust, K-Na-Cl, fly ash, SIA, organics, and soot) has been defined as aged particle. 292 More than 90% of particles at the background site were highly aged. SIA with OC coating were 293 the dominant particles and could determine the hygroscopic properties of the ambient aerosol 294 particles. OC coatings on inorganic particles can induce an early deliquescence of particle surface 295 compared to that of the pure inorganic compounds (Li et al., 2014a). Recently, Mikhailov et al. 296 (2015) found that the semi-solid state of the OC coating can lead to kinetic limitations of water 297 uptake and release during hydrate and dehydrate cycles in the background area. OC dominated in 298 fine particles, accounting for 43% of mass on average, followed by sulfate (28%) and nitrate (1%) 299 (Du et al., 2015). TEM observations further indicated that OC can heterogeneously and homogeneously be mixed with all the fine SIA particles (Figures 6, 9-10). This finding is in 300 301 agreement with the study of new particle formation and growth events during the sampling period, in which oxygenated organics significantly contributed into particle growth in the QTP (Du et al., 302 303 2015). SIA with OC coating (i.e., particle without inclusions in Figure 7) shift to one smaller size 304 than the total individual particles. Gas condensation on the existing particles is an important 305 chemical process for formation of SIA with OC coating. Therefore, OC coating of the aged 306 aerosol particles is likely an important factor to determine particle hygroscopic growth and phase 307 transitions in the QTP.

Inclusions within SIA particles increase their size by 36-42% (Figure 7). The size distribution 308 309 of individual particles shows that particles without inclusions have a median size of 200 nm - 250 310 nm at the background site. Figure 7 shows that the number of particles with inclusions increases 311 substantially with diameters above 200 nm. In addition, Figures 9-10 show soot, fly ash, and 312 visible OC particles likely adhere to the surface of SIA particles, which is different from many 313 refractory particles embedded within SIA particles in East China (Li and Shao, 2009;Li et al., 2011b;Li et al., 2014b). Therefore, the coagulation process between primary refractory particles 314 315 and SIA particles with diameters > 200 nm could be dominant in the atmosphere. The results are 316 different from the background aerosol particles in Siberia where these soot, fly ash and visible OC 317 particles are absent (Mikhailov et al., 2015). In particular, the mixing structure of soot on the 318 surface of SIA particles is different from the previous studies (Li et al., 2003;Adachi et al., 319 2010;Li et al., 2015). Therefore, sulfates cannot act as the lens to enhance optical absorption of 320 soot particles before individual particles totally deliquesces in humid air (Ramanathan and 321 Carmichael, 2008). Light absorption of soot on surfaces of sulfate particles have 30% lower than soot centered within sulfate particles (Adachi et al., 2010). We found that number of 322 soot-containing particles increase during the biomass burning periods (Figure 8b-c). Also, Figure 3 323 shows that the BC concentrations exceeded $1.0 \ \mu g \ m^{-3}$ during biomass burning periods which is 324 325 two times higher than the average value during the sampling period. As a result, large amounts of 326 soot particles from biomass burning in background atmosphere likely change atmospheric optical 327 absorption and modify optical feedback of ice/snow after their deposition in the QTP (Che et al.,

2011;Ming et al., 2012). The microstructure of soot particles can improve understanding of the
optical properties of fine aerosol particles and better evaluate their climate impacts using climate
models.

We also notice that a large number of particles without inclusions can occur with diameters > 331 332 300 nm, although the particle numbers decrease (Figure 7). TEM observations reveal distinct rims 333 on some larger particles, as the examples shown in Figures 5a, 5d, 9b, 10a. Our previous studies 334 showed that the cloud and fog residues on the substrate can display distinct rims (Kojima et al., 335 2004;Li et al., 2011a;Li et al., 2011b). Therefore, these large particles with distinct rims probably undergo complicated atmospheric transformation such as cloud/fog processing during their growth. 336 337 Briefly, our study indicates that aerosol particles in different size regimes have different 338 atmospheric chemical or physical processes in the background air over the QTP.

4.4 Further considerations about fine particles over Qinghai-Tibetan plateau

340 Emissions from coal combustion and biomass burning contribute fine particles into the 341 background air over the QTP. Because the complex aerosol particles from different anthropogenic 342 sources intruded into pristine background air, the suspended aerosols became highly aged. 343 Because of the sensitive feedback of climate in the TP, these aged aerosol particles in the plateau 344 become particularly interesting. Firstly, transport and sources of aerosol particles should be 345 evaluated, and, indeed, most studies in the TP have accomplished this (Cong et al., 2009a;Cong et 346 al., 2009b;Engling et al., 2011;Lu et al., 2012;Du et al., 2015;Xu et al., 2015). These studies all 347 suggested that long-range transport of fine particles from biomass burning and other 348 anthropogenic sources (cooking and vehicular emissions) often reach the TP. However, the 349 emissions of coal combustion from power plants or other industrial sources have a decided 350 regional influence. The statement has not been reported. Our studies provide new evidence that fly 351 ash particles serve as a reliable fingerprint of coal-combustion at the background site. Following 352 economic development in western China, coal combustion increases, chiefly for electrical power 353 generation and other industrial activities (Figure S2). Secondly, highly aged particles such as 354 ambient aerosols and CCN in the atmosphere and sediment in ice/snow can directly or indirectly 355 impact on climate in the TP (Cong et al., 2009b;You et al., 2010;Che et al., 2011;Lu et al., 356 2012; Ming et al., 2012; He et al., 2014; Wang et al., 2015; Yang et al., 2015). At the background sampling site the mean BC concentration was $0.54 \ \mu g/m^3$. Interestingly, we found that most fine 357

358 soot (BC) particles adhere to individual SIA (Figures 9-10) in the Qinghai plateau while many soot particles were embedded within SIA in polluted areas of East China (Li and Shao, 2009;Li et 359 360 al., 2011b). The detailed physical properties (e.g., mixing structure and size) of soot in air and ice/snow should be further studied in the TP, which can improve the current climate models. 361 Thirdly, fine aerosol particles in the TP mainly contain OC and sulfates with minor nitrates. The 362 363 result is largely different from fine particles with high nitrate in more polluted areas (Li et al., 364 2013a;Du et al., 2015;Xu et al., 2015). Although regional transport from anthropogenic sources or 365 biomass burning significantly increase particle concentrations, mineral dust from surrounding deserts and organics from plants are still dominant in the TP throughout the year (Zhang et al., 366 367 2001b;Wang et al., 2010;Li et al., 2013a;Xu et al., 2014).

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369 5. Conclusions

Time-series of six pollutants (PM_{2.5}, BC, SO₂, NO_x, CO, and O₃) were obtained at a national 370 371 station of background atmospheric monitoring (NSBAM) on the top of Moshidaban Mountain in the North QTP during 15 September – 15 October, 2013. The mean concentrations of PM_{2.5}, BC, 372 SO₂, NO_x, CO, and O₃ were 17.06 μ g/m³, 0.54 μ g/m³, 1.27 ppb, 2.05 ppb, 44.78 ppb, and 50.00 373 374 ppb, respectively. TEM was employed to study individual fine particles with the diameter less 375 than 1 µm that were classified into six major particle types: mineral dust, K-Na-Cl, fly ash, 376 secondary inorganic aerosol (SIA) containing ammoniated sulfates and nitrates, organics, soot (i.e., 377 BC). Individual fine particle types display large differences under three different $PM_{2.5}$ levels: $PM_{2.5} \ge 30 \ \mu g/m^3$ (high pollution level), $10 \le PM2.5 < 30 \ \mu g/m^3$ (medium pollution level), and < 378 $10 \ \mu g/m^3$ (clean period). Individual fine particles in clean periods mainly contained SIA and 379 380 organics. The presence of K-Na-Cl coated by organics and increased soot particles during high 381 pollution levels suggests an intense biomass burning event near the background site. Large 382 amounts of fly ash-containing particles occurred in medium pollution level. The fly ash-containing 383 particles increased from 11.40% at the medium pollution level to 21.73% at the high pollution 384 level, but soot-containing particles decreased from 38.40% to 25.87%. This result indicates that 385 coal combustion emissions in the QTP significantly affected the background air quality. In 386 addition, SIA particles containing minor K suggest that biomass burning emissions were a constant contributor to aerosol particles in the QTP. We concluded that the emissions from 387

biomass burning and coal-used anthropogenic activities contribute anthropogenic particles into the

389 QTP atmosphere.

390 Aerosol particles containing SIA core and OC coating display smaller median size than the total particles. Gas condensation on the particles is an important chemical process for their 391 formation. The number concentration of particles with inclusions increased markedly above 200 392 393 nm. TEM observations show that refractory aerosols (e.g., soot, fly ash, and visible organic particles) likely adhere to the surface of SIA particles, suggesting physical coagulation could be 394 395 dominant in background air. These results notably improve our understanding of sources and aging processes of long-range transported aerosols in the QTP. The transport of these aerosol 396 397 particles, as well as their, hygroscopic, and optical properties and atmospheric chemistry, require 398 further study in the TP.

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412 References

- Adachi, K., Chung, S. H., and Buseck, P. R.: Shapes of soot aerosol particles and implications for their
 effects on climate, J. Geophys. Res., 115, doi:10.1029/2009JD012868, 2010.
- Cao, J., Tie, X., Xu, B., Zhao, Z., Zhu, C., Li, G., and Liu, S.: Measuring and modeling black carbon (BC)
 contamination in the SE Tibetan Plateau, J. Atmos. Chem., 67, 45-60, 2010.
- Che, H., Wang, Y., and Sun, J.: Aerosol optical properties at Mt. Waliguan Observatory, China, Atmos.
 Environ., 45, 6004-6009, 2011.
- Cong, Z., Kang, S., Dong, S., and Zhang, Y.: Individual particle analysis of atmospheric aerosols at Nam
 Co, Tibetan Plateau, Aerosol Air Qual. Res., 9, 323-331, 2009a.
- 421 Cong, Z., Kang, S., Smirnov, A., and Holben, B.: Aerosol optical properties at Nam Co, a remote site in
 422 central Tibetan Plateau, Atmos. Res., 92, 42-48, 2009b.
- 423 Cong, Z., Kawamura, K., Kang, S., and Fu, P.: Penetration of biomass-burning emissions from South Asia
 424 through the Himalayas: new insights from atmospheric organic acids, Sci. Rep., 5, DOI:
 425 10.1038/srep09580, 2015.
- 426 Du, W., Sun, Y. L., Xu, Y. S., Jiang, Q., Wang, Q. Q., Yang, W., Wang, F., Bai, Z. P., Zhao, X. D., and Yang, Y.
- 427 C.: Chemical characterization of submicron aerosol and particle growth events at a national
 428 background site (3295 m a.s.l.) in the Tibetan Plateau, Atmos. Chem. Phys. Discuss., 15, 13515-13550,
 429 2015.
- Duo, B., Zhang, Y., Kong, L., Fu, H., Hu, Y., Chen, J., Li, L., and Qiong, A.: Individual particle analysis of
 aerosols collected at Lhasa City in the Tibetan Plateau, J. Environ. Sci., 29, 165-177, 2015.
- 432 Engling, G., Zhang, Y.-N., Chan, C.-Y., Sang, X.-F., Lin, M., Ho, K.-F., Li, Y.-S., Lin, C.-Y., and Lee, J. J.:
- 433 Characterization and sources of aerosol particles over the southeastern Tibetan Plateau during the
- 434 Southeast Asia biomass-burning season, Tellus B, 63, 117-128, 10.1111/j.1600-0889.2010.00512.x,
 435 2011.
- He, C., Li, Q., Liou, K.-N., Takano, Y., Gu, Y., Qi, L., Mao, Y., and Leung, L. R.: Black carbon radiative
 forcing over the Tibetan Plateau, Geophy. Res. Lett., 41, 2014GL062191, 2014.
- Kojima, T., Buseck, P. R., Wilson, J. C., Reeves, J. M., and Mahoney, M. J.: Aerosol particles from tropical
 convective systems: Cloud tops and cirrus anvils, J. Geophys. Res., 109, 12201-12201, 2004.
- 440 Kopacz, M., Mauzerall, D. L., Wang, J., Leibensperger, E. M., Henze, D. K., and Singh, K.: Origin and
- radiative forcing of black carbon transported to the Himalayas and Tibetan Plateau, Atmos. Chem.Phys., 11, 2837-2852, 2011.
- Lau, K. M., Kim, M. K., and Kim, K. M.: Asian summer monsoon anomalies induced by aerosol direct
 forcing: the role of the Tibetan Plateau, Clim Dyn, 26, 855-864, 2006.
- Li, C., Kang, S., Zhang, Q., and Kaspari, S.: Major ionic composition of precipitation in the Nam Co region, Central Tibetan Plateau, Atmos. Res., 85, 351-360, 2007.
- Li, J., Posfai, M., Hobbs, P. V., and Buseck, P. R.: Individual aerosol particles from biomass burning in
 southern Africa: 2, Compositions and aging of inorganic particles, J. Geophys. Res., 108,
 doi:10.1029/2002JD002310, 2003.
- 450 Li, J. J., Wang, G. H., Wang, X. M., Cao, J. J., Sun, T., Cheng, C. L., Meng, J. J., Hu, T. F., and Liu, S. X.:
- Abundance, composition and source of atmospheric PM 2.5 at a remote site in the Tibetan Plateau,
 China, Tellus B, 65, 10.3402/tellusb.v65i0.20281, 2013a.
- 453 Li, W., Li, P., Sun, G., Zhou, S., Yuan, Q., and Wang, W.: Cloud residues and interstitial aerosols from
- 454 non-precipitating clouds over an industrial and urban area in northern China, Atmos. Environ., 45,
- 455 2488-2495, 2011a.

- Li, W., Wang, T., Zhou, S., Lee, S., Huang, Y., Gao, Y., and Wang, W.: Microscopic Observation of
 Metal-Containing Particles from Chinese Continental Outflow Observed from a Non-Industrial Site,
 Environ. Sci. Technol., 47, 9124-9131, 2013b.
- Li, W., Wang, Y., Collett, J. L., Chen, J., Zhang, X., Wang, Z., and Wang, W.: Microscopic Evaluation of
 Trace Metals in Cloud Droplets in an Acid Precipitation Region, Environ. Sci. Technol., 47, 4172-4180,
 2013c.
- Li, W., Chi, J., Shi, Z., Wang, X., Chen, B., Wang, Y., Li, T., Chen, J., Zhang, D., Wang, Z., Shi, C., Liu, L.,
 and Wang, W.: Composition and hygroscopicity of aerosol particles at Mt. Lu in South China:
 Implications for acid precipitation, Atmos. Environ., 94, 626-636,
 <u>http://dx.doi.org/10.1016/j.atmosenv.2014.06.003</u>, 2014a.
- Li, W., Shao, L., Shi, Z., Chen, J., Yang, L., Yuan, Q., Yan, C., Zhang, X., Wang, Y., Sun, J., Zhang, Y., Shen,
 X., Wang, Z., and Wang, W.: Mixing state and hygroscopicity of dust and haze particles before leaving
 Asian continent, J. Geophys. Res., 119, 1044-1059, 2014b.
- Li, W., Shao, L., Zhang, D., Ro, C.-U., Hu, M., Bi, X., Geng, H., Matsuki, A., Niu, H., and Chen, J.: A review
 of single aerosol particle studies in the atmosphere of East Asia: morphology, mixing state, source, and
- 471 heterogeneous reactions, J. Clean. Prod., DOI:10.1016/j.jclepro.2015.1004.1050 (in press), 2015.
- Li, W. J., and Shao, L. Y.: Transmission electron microscopy study of aerosol particles from the brown
 hazes in northern China, J. Geophys. Res., 114, doi:10.1029/2008JD011285, 2009.
- Li, W. J., Zhou, S. Z., Wang, X. F., Xu, Z., Yuan, C., Yu, Y. C., Zhang, Q. Z., and Wang, W. X.: Integrated
 evaluation of aerosols from regional brown hazes over northern China in winter: Concentrations,
 sources, transformation, and mixing states, J. Geophys. Res., 116, doi:10.1029/2010JD015099, 2011b.
- 477 Lu, Z., Streets, D. G., Zhang, Q., and Wang, S.: A novel back-trajectory analysis of the origin of black
 478 carbon transported to the Himalayas and Tibetan Plateau during 1996-2010, Geophy. Res. Lett., 39,
 479 L01809, 10.1029/2011gl049903, 2012.
- Mikhailov, E. F., Mironov, G. N., Pöhlker, C., Chi, X., Krüger, M. L., Shiraiwa, M., Förster, J. D., Pöschl, U.,
 Vlasenko, S. S., Ryshkevich, T. I., Weigand, M., Kilcoyne, A. L. D., and Andreae, M. O.: Chemical
 composition, microstructure, and hygroscopic properties of aerosol particles at the Zotino Tall Tower
- 483 Observatory (ZOTTO), Siberia, during a summer campaign, Atmos. Chem. Phys., 15, 8847-8869,
 484 10.5194/acp-15-8847-2015, 2015.
- 485 Ming, J., Du, Z., Xiao, C., Xu, X., and Zhang, D.: Darkening of the mid-Himalaya glaciers since 2000 and
 486 the potential causes, Environ. Res. Lett., 7, 014021, 10.1088/1748-9326/7/1/014021, 2012.
- 487 Posfai, M., and Buseck, P. R.: Nature and Climate Effects of Individual Tropospheric Aerosol Particles,
 488 Annu. Rev. Earth Pl. Sc., 38, 17-43, 2010.
- Ramanathan, V., and Carmichael, G.: Global and regional climate changes due to black carbon, Nature
 Geosci, 1, 221-227, 2008.
- Stohl, A., Hittenberger, M., and Wotawa, G.: Validation of the lagrangian particle dispersion model
 FLEXPART against large-scale tracer experiment data, Atmos. Environ., 32, 4245-4264,
 <u>http://dx.doi.org/10.1016/S1352-2310(98)00184-8</u>, 1998.
- 494 Wang, Q. Y., Huang, R. J., Cao, J. J., Tie, X. X., Ni, H. Y., Zhou, Y. Q., Han, Y. M., Hu, T. F., Zhu, C. S., Feng,
- 495 T., Li, N., and Li, J. D.: Black carbon aerosol in winter northeastern Qinghai-Tibetan Plateau, China: the
- effects from South Asia pollution, Atmos. Chem. Phys. Discuss., 15, 14141-14169, 2015.
- Wang, X., Huang, J., Zhang, R., Chen, B., and Bi, J.: Surface measurements of aerosol properties over
 northwest China during ARM China 2008 deployment, J. Geophys. Res., 115,
 doi:10.1029/2009JD013467, 2010.

- Wen, J., Meng, H., Wang, X., and coauthors: China Energy Statistical Yearbook, Department of Energy
 Statistics, National Bureau of Statistics, People's Republic of China, Beijing, 1-355, 2013.
- Xia, X., Zong, X., Cong, Z., Chen, H., Kang, S., and Wang, P.: Baseline continental aerosol over the
 central Tibetan plateau and a case study of aerosol transport from South Asia, Atmos. Environ., 45,
 7370-7378, 2011.
- Xu, B., Cao, J., Hansen, J., Yao, T., Joswia, D. R., Wang, N., Wu, G., Wang, M., Zhao, H., Yang, W., Liu, X.,
 and He, J.: Black soot and the survival of Tibetan glaciers, Proc. Natl. Acad. Sci. U.S.A., 106,
 22114-22118, 2009.
- Xu, J., Zhang, Q., Li, X., Ge, X., Xiao, C., Ren, J., and Qin, D.: Dissolved Organic Matter and Inorganic Ions
 in a Central Himalayan Glacier—Insights into Chemical Composition and Atmospheric Sources, Environ.
- 510 Sci. Technol., 47, 6181-6188, 2013.
- Xu, J., Wang, Z., Yu, G., Qin, X., Ren, J., and Qin, D.: Characteristics of water soluble ionic species in fine
 particles from a high altitude site on the northern boundary of Tibetan Plateau: Mixture of mineral
 dust and anthropogenic aerosol, Atmos. Res., 143, 43-56, 2014.
- 514 Xu, J. Z., Zhang, Q., Wang, Z. B., Yu, G. M., Ge, X. L., and Qin, X.: Chemical composition and size
- 515 distribution of summertime PM2.5 at a high altitude remote location in the northeast of the Qinghai–
- 516 Xizang (Tibet) Plateau: insights into aerosol sources and processing in free troposphere, Atmos. Chem.
 517 Phys., 15, 5069-5081, 2015.
- 518 Xue, L. K., Wang, T., Zhang, J. M., Zhang, X. C., Deliger, Poon, C. N., Ding, A. J., Zhou, X. H., Wu, W. S.,
- Tang, J., Zhang, Q. Z., and Wang, W. X.: Source of surface ozone and reactive nitrogen speciation at
 Mount Waliguan in western China: New insights from the 2006 summer study, J. Geophys. Res., 116,
 doi:10.1029/2010JD014735, 2011.
- Yang, S., Xu, B., Cao, J., Zender, C. S., and Wang, M.: Climate effect of black carbon aerosol in a Tibetan
 Plateau glacier, Atmos. Environ., 111, 71-78, 2015.
- You, Q., Kang, S., Flügel, W.-A., Sanchez-Lorenzo, A., Yan, Y., Huang, J., and Martin-Vide, J.: From
 brightening to dimming in sunshine duration over the eastern and central Tibetan Plateau (1961–
- 526 2005), Theor. Appl. Climatol., 101, 445-457, 2010.
- 527 Zhang, D., Iwasaka, Y., and Shi, G.: Soot particles and their impacts on the mass cycle in the Tibetan
 528 atmosphere, Atmos. Environ., 35, 5883-5894, 2001a.
- Zhang, X. Y., Arimoto, R., Cao, J. J., An, Z. S., and Wang, D.: Atmospheric dust aerosol over the Tibetan
 Plateau, J. Geophys. Res., 106, 18471-18476, 2001b.
- 531 Zhao, Z., Cao, J., Shen, Z., Xu, B., Zhu, C., Chen, L. W. A., Su, X., Liu, S., Han, Y., Wang, G., and Ho, K.:
- Aerosol particles at a high-altitude site on the Southeast Tibetan Plateau, China: Implications for pollution transport from South Asia, J. Geophys. Res., 118, 11,360-311,375, 2013.
- 534 Zheng, W., Yao, T., Joswiak, D. R., Xu, B., Wang, N., and Zhao, H.: Major ions composition records from
- a shallow ice core on Mt. Tanggula in the central Qinghai-Tibetan Plateau, Atmos. Res., 97, 70-79,2010.
- 537



Figure 1 FLEXPART retroplume simulations during 10 September-15 October. Topographical map showing the sampling location and surrounding regions in the Tibetan Plateau. Xi'ning is the caption city of Qinghai province. Menyuan represents sampling site. The black line shows the major back trajectories of air mass during 10 September - 15 October, 2013 based on the spatial distribution of the air mass. The main air mass mainly passed through rural areas, grasslands, and desert.



Figure 2 Time-series concentration of air pollutants (i.e., CO, NO_x, SO₂, BC, and PM_{2.5}) during 11 September – 15 October, 2013. The sampling time was marked by grey column ($PM_{2.5} \ge 30 \mu g/m^3$), pink column ($PM_{2.5}$ between 10-30 $\mu g/m^3$), blue column ($PM_{2.5} < 10 \mu g/m^3$).



Figure 3 Correction of equivalent circle diameters (ECD) vs equivalent volume diameter (EVD) of 194 aerosol particles.



Figure 4 TEM images of (a-b) K-Na-Cl particle with organic coating on 22 September and 13 October. (c) SIA-soot with organic coating on 22 September. (d) SIA with organic coating on 11 October. EDS spectra shows elemental compositions of each particle type in each TEM image.



Figure 5 TEM images of (a) SIA-soot-OC (visible) with organic coating on 16 September. (b) SIA-soot-OC on 14 September. (c) SIA-fly ash-OC on 5 October. (d) SIA-fly ash-soot-OC (visible) with organic coating on 16 September. EDS spectra shows elemental composition of each particle type in each TEM image.



Figure 6 TEM image of individual particles collected in clean period with $PM_{2.5}$ mass concentration less than 10 μ g/m³. SIA particles tend to homogeneously mix with organics.



Figure 7 Size distributions of individual particles, inclusions, particles with inclusions, and particles without inclusions. (a) Clean periods under $PM_{2.5}$ at 10 µg/m³. (b) The high pollution level under $PM_{2.5}$ larger than 30 µg/m³. (c) The medium pollution level under $PM_{2.5}$ among 10 µg/m³ - 30 µg/m³.



Figure 8 Identification of the pollution events based on individual particle analysis (a) 684 individual particles and 108 SIA-inclusion particles. (b) 1214 individual particles and 307 SIA-inclusion particles. Eight samples were collected in $PM_{2.5}$ larger than 30 µg m⁻³ induced by biomass burning emission. (c) 2355 individual particles and 750 SIA-inclusion particles. Nine samples were collected in $PM_{2.5}$ at the range of 10-30 µg m⁻³ induced by biomass burning and industrial emissions. Four samples were collected in $PM_{2.5}$ smaller than 10 µg m⁻³ which indicates clean period.



Figure 9 Individual particles during biomass burning periods with $PM_{2.5}$ mass concentration larger than 30 µg/m³. (a) OC and SIA-soot-(OC coating) particles on 12 October. (b) SIA-soot-(visible OC) on 19 October. OC in particles 1, 3, 6, 10, 13, 14 are heterogeneously mixed with SIA particles and particles 2, 4, 5, 8 homogeneously mixed with minor SIA.



Figure 10 Individual particles collected under $PM_{2.5}$ mass concentration among 10-30 µg/m³. (a) Mixture of SIA and soot, fly ash particles collected on 14 September. (b) Mixture of SIA and soot, fly ash particles collected on 18 September. (c) Mixture of SIA and fly ash, soot, organics collected on 29 September. (c) Mixing of SIA and fly ash, soot, organics collected on 10 October. Organics are heterogeneously mixed with SIA.

	All data		polluted period-1		polluted period-2		other period	
Pollutants	mean ±SD	n	mean ±SD	n	mean ±SD		mean ±SD	n
	Max, Min		Max, Min		Max, Min	n	Max, Min	
DM	17.06±11.39	715	17.6±11.46	152	24.45±15.12	99	15.32±10.41	464
PM _{2.5}	68.70, 0.20		59.10, 0.20		68.70, 0.30		62.80, 0.20	
DC	0.54±0.42	805	0.55±0.52	176	0.85±0.50	119	0.47 ±0.40	510
BC	3.73, 0.02		3.73, 0.04		2.04, 0.02		3.73, 0.03	
80	1.27±1.34	8822	1.2±0.99	1981	2.73±3.09	1063	1.03±0.65	5778
SO ₂	13.93, 0.02		8.43, 0.20		13.93, 1.41		8.43, 0.02	
NO	2.05 ± 1.96	8842	2.37±1.33	2001	3.41 ± 1.70	1063	1.69±0.97	5778
NOX	9.86, 0.31		9.33, 0.65		9.59, 0.55		9.33, 0.31	
60	44.78±48.03	7822	63.45±55.59	1939	104.23±54.69	1030	24.68±39.91	4853
CO	318.00, 0.20		318.00, 0.20		272.40, 0.60		318.00, 0.20	
0	50±7.86	8817	47.87±7.70	2000	49.01±10.00	1020	50.53±7.56	5778
03	98.63, 20.43		67.70, 26.66		98.63, 20.43	1039	96.77, 26.66	

Table 1 Concentrations of six air pollutants during the sampling period, two pollution periods, and clean period

All data period: 10 Sept.-15 Oct. 2013; Polluted period-1: 18 Sept.-25 Sept. 2013; Polluted period-2: 11 Oct.-15 Oct. 2013