Charging and Coagulation of Radioactive and Nonradioactive

Particles in the Atmosphere

Yong-ha Kim, ^a Sotira Yiacoumi, ^a Athanasios Nenes, ^{b,c} Costas Tsouris ^{a,d,*}
^a School of Civil and Environmental Engineering, Georgia Institute of Technology, Atlanta,
Georgia 30332-0373, USA
^b School of Earth and Atmospheric Sciences, Georgia Institute of Technology, Atlanta, Georgia
30332-0340, USA
^c School of Chemical and Biomolecular Engineering, Georgia Institute of Technology, Atlanta,
Georgia 30332-0100, USA
^d Oak Ridge National Laboratory, Oak Ridge, Tennessee 37831-6181, USA
*Email: <u>tsourisc@ornl.gov</u> ; Telephone: 865-241-3246; Fax: 865-241-4829
Submitted for publication in
Atmospheric Chemistry and Physics
July 2015
Revised manuscript submitted in February 2016
Notice: This manuscript has been authored by UT-Battelle, LLC under Contract No. DE-AC05-00OR22725 with the US Department of Energy. The United States Government retains and the publisher, by accepting the article for publication, acknowledges that the United States Government retains a non-exclusive, paid-up, irrevocable, world-wide license to publish or reproduce the published form of this manuscript, or allow others to do so, for United States Government purposes. The Department of Energy will provide public access to these results of federally sponsored research in accordance with the DOE Public Access Plan (http://energy.gov/downloads/doe-public-access-plan).

1

Charging and Coagulation of Radioactive and Nonradioactive

25

26

Particles in the Atmosphere

27 Yong-ha Kim, a Sotira Yiacoumi, Athanasios Nenes, b,c Costas Tsouris Athanasios Nenes, b,c Costas Tsouris 28 29 30 ^aSchool of Civil and Environmental Engineering, Georgia Institute of Technology, Atlanta, 31 Georgia 30332-0373, USA 32 ^bSchool of Earth and Atmospheric Sciences, Georgia Institute of Technology, Atlanta, Georgia 33 30332-0340, USA 34 ^cSchool of Chemical and Biomolecular Engineering, Georgia Institute of Technology, Atlanta, 35 Georgia 30332-0100, USA ^dOak Ridge National Laboratory, Oak Ridge, Tennessee 37831-6181, USA 36 37 *Email: tsourisc@ornl.gov; Telephone: 865-241-3246; Fax: 865-241-4829 38 39 ABSTRACT 40 Charging and coagulation influence one another and impact the particle charge and size 41 distributions in the atmosphere. However, few investigations to date have focused on the 42 coagulation kinetics of atmospheric particles accumulating charge. This study presents three 43 approaches to include mutual effects of charging and coagulation on the microphysical evolution 44 of atmospheric particles such as radioactive particles. The first approach employs ion balance, 45 charge balance, and a bivariate population balance model (PBM) to comprehensively calculate 46 both charge accumulation and coagulation rates of particles. The second approach involves a

much simpler description of charging, and uses a monovariate PBM and subsequent effects of charge on particle coagulation. The third approach is further simplified assuming that particles instantaneously reach their steady-state charge distributions. It is found that compared to the other two approaches, the first approach can accurately predict time-dependent changes in the size and charge distributions of particles over a wide size range covering from the free molecule to continuum regimes. The other two approaches can reliably predict both charge accumulation and coagulation rates for particles larger than about 0.04 micrometers and atmospherically relevant conditions. These approaches are applied to investigate coagulation kinetics of particles accumulating charge in a radioactive neutralizer, the urban atmosphere, and an atmospheric system containing radioactive particles. Limitations of the approaches are discussed.

1. INTRODUCTION

58

59

60

61

62

63

64

65

66

67

68

69

70

71

72

73

74

75

76

77

78

79

Atmospheric particles play an important role in airborne transport of contaminants, such as radionuclides. Contaminants emitted from anthropogenic sources (e.g., nuclear plant accidents) can be captured by background aerosols and then are transported together with pre-existing particles. Contaminant-laden particles can be deposited onto the ground by dry and wet deposition (primary contamination) and subsequently resuspended by wind or heat-driven convection and moved to other areas (secondary contamination). For instance, due to these atmospheric dispersion patterns, radioactive particles (e.g., ¹³⁷Cs) released during the Fukushima accident were sampled in-situ 150 km away from the emission site (Yamauchi et al., 2012), and also found in many places around the world (Hoeve and Jacobson, 2012). Similar dispersion patterns of radioactive particles were observed after the Chernobyl accident (Yoshenko et al., 2006a, 2006b). These examples suggest that particle deposition, which is largely affected by aerosol microphysics, can determine the fate of contaminants during atmospheric transport. Thus, accurate understanding of the microphysical behavior of atmospheric particles is necessary to more accurately predict transport of contaminants (especially long-lived ones, such as ¹³⁷Cs), as well as their potential environmental impacts. The microphysical behavior of atmospheric particles is driven by such properties as charge and size (Fuchs, 1989; Pruppacher and Klett, 1997). Atmospheric particles can acquire charge via self-charging and diffusion charging. Self-charging refers to charge accumulation caused by radioactive decay which typically leads to emission of electrons from particle surfaces. Diffusion charging is attributed to diffusion of ions from the surrounding atmosphere onto the surface of particles. Radioactive particles can be charged through these two charging mechanisms

(Greenfield, 1956; Yeh et al., 1976; Clement and Harrison, 1992; Clement et al., 1995; Gensdarmes et al., 2001; Walker et al., 2010; Kweon et al., 2013; Kim et al., 2014; 2015) while natural atmospheric particles are typically charged by diffusion charging (Hoppel, 1985; Yair and Levin, 1989; Renard et al., 2013). Particle charging can modify not only the charge but also the size distribution because charge on the particles modifies their coagulation rate coefficients by generating electrostatic interactions (Fuchs, 1989; Tsouris et al., 1995; Chin et al., 1998). Coagulation of atmospheric particles can influence their charging because the concentration of atmospheric ions is affected by the particle size distribution (Yair and Levin, 1989), thereby modifying the diffusion charging rates of the particles. Also, particle coagulation can result in charge neutralization or accumulation on atmospheric particles (Alonso et al., 1998). These effects imply that particle coagulation can influence the particle charge distribution. Thus, particle charging and coagulation can mutually affect each other and simultaneously affect both charge and size distributions in the atmosphere. Theoretical and experimental investigations have been performed to examine the charging of radioactive particles and background aerosols in the atmosphere. However, the effects of coagulation of particles on the charge distribution have been frequently neglected by assuming that the size distribution is constant while they are charged (Greenfield, 1956; Hoppel, 1985; Yair and Levin, 1989). The assumption may be valid if the particle concentration is low or the steady-state charge distribution is instantaneously attained (Hoppel, 1985; Renard et al., 2013; Kim et al., 2015). If the timescale for particle charging is longer than that for particle coagulation, the assumption may no longer be valid (Yair and Levin, 1989). Also, the effects of charging on the particle size distribution are frequently neglected in aerosol transport models involving microphysics of atmospheric particles. A possible reason for neglecting the charging effects may

80

81

82

83

84

85

86

87

88

89

90

91

92

93

94

95

96

97

98

99

100

101

be that the steady-state mean charge of atmospheric particles rarely may affect their coagulation rates (Seinfeld and Pandis, 2006). However, neglecting electrostatic particle-particle interactions may increase uncertainty of prediction results if particles can acquire multiple elementary charges (e.g., radioactive particles). The simplified assumption of omitting electrostatic particle interactions may create uncertainty in transport predictions of radioactive particles. Hence, it may be necessary to take into account the mutual effects of particle charging and coagulation processes in predicting the behavior of atmospheric particles carrying radioactive contaminants. Previous attempts to consider charging effects include Oron and Seinfeld (1989a, 1989b), who developed sectional approaches to simultaneously predict the behavior of charged and uncharged atmospheric particles. Laakso et al. (2002) developed a general dynamic equation, including charging and coagulation kinetics of atmospheric particles. Yu and Turco (2001) presented a kinetic approach that involves diffusion charging and particle coagulation. However, the validity of these approaches was not evaluated using analytical solutions. Alonso (1999) and Alonso et al. (1998) developed analytical and numerical approaches to estimate time-dependent changes in the size distributions of singly charged and neutral particles; thus, these approaches cannot be used to investigate the coagulation kinetics of particles acquiring multiple elementary charges. Also, none of these approaches considered self-charging; therefore, the aforementioned approaches may be subject to error when they are used to simulate atmospheric dispersion of radioactive plumes. Our study presents three approaches to simultaneously predict time-dependent changes of the charge and size distributions of radioactive and nonradioactive particles over a wide size range.

103

104

105

106

107

108

109

110

111

112

113

114

115

116

117

118

119

120

121

122

123

124

Development, validity, application, and limitations of these approaches are discussed.

2. MODEL DEVELOPMENT

126 2.1. Ion balance model

125

Many atmospheric processes can generate and remove ions in air. Typical ion sources in the atmosphere involve natural and artificial radioactivity, as well as cosmic rays. Ions are generally removed by ion-ion recombination and ion-particle attachment. Changes in ion concentrations by these processes can be given by (Kim et al., 2015):

131
$$\frac{dn_{\text{ion}}^{+}}{dt} = -n_{\text{ion}}^{+} \sum_{k} \sum_{j} \beta_{k,j}^{+} N_{k,j} - \alpha_{rc} n_{\text{ion}}^{+} n_{\text{ion}}^{-} + q, \qquad (1)$$

132
$$\frac{dn_{\text{ion}}^{-}}{dt} = -n_{\text{ion}}^{-} \sum_{k} \sum_{j} \beta_{k,j}^{-} N_{k,j} - \alpha_{rc} n_{\text{ion}}^{+} n_{\text{ion}}^{-} + q + q_{e}, \qquad (2)$$

where n_{ion}^{\pm} refers to the number concentrations of positive or negative ions, the indices k and j represent the size and number of elementary charges of particles, respectively, $\beta_{k,j}^{\pm}$ is the attachment coefficient between a particle and an ion, $N_{k,j}$ is the number concentration of particles, α_{rc} is the recombination coefficient of ions, and t is time. The first two terms of the right-hand-side (RHS) of equations 1 and 2 represent the loss rate of ions due to ion-particle attachment and ion-ion recombination, respectively. The third term denotes the production rate of ion pairs, q:

$$139 q = q_b + q_I, (3)$$

where q_b is the ion production rate by cosmic rays and natural radioactivity, and q_I is the ion production rate by radionuclides released by nuclear events. Electrons released by radioactive decay are taken into account via the last term of eq 2. Changes in the ion concentrations may affect the electrical conductivity of the atmosphere, σ_{air} (Harrison and Carslaw, 2003):

144
$$\sigma_{\text{air}} = e(\mu_{+} n_{\text{ion}}^{+} + \mu_{-} n_{\text{ion}}^{-}),$$
 (4)

145 where e is the electrical charge and μ_{\pm} is the mobility of positive or negative ions. In eq 4, the 146 terms in the parentheses of the RHS represent polar air conductivities.

147

148

155

156

157

158

159

2.2. Charge balance models

Self-charging generally accumulates positive charge on the surface of particles, while diffusion charging adds both positive and negative charges, indicating that the charging mechanisms can compete with one another. For radioactive particles involved in beta decay, time-dependent changes in their charge distributions due to competition of the charging mechanisms can be expressed by (Clement and Harrison, 1992; Kim et al., 2015):

154
$$\frac{dN_{kj}}{dt} = A_{k,j-1}N_{k,j-1} - A_{k,j}N_{k,j} + \beta_{k,j-1}^{+}n_{\text{ion}}^{+}N_{k,j-1} - \beta_{k,j}^{+}n_{\text{ion}}^{+}N_{k,j} + \beta_{k,j+1}^{-}n_{\text{ion}}^{-}N_{k,j+1},$$

$$-\beta_{k,j}^{-}n_{\text{ion}}^{-}N_{k,j}$$
(5)

where A is the radioactivity of the radioactive particles. In eq 5, self-charging is represented by terms that include A, while diffusion charging is represented by terms with $\beta_{k,j}^{\pm}$. If the terms for self-charging are removed, eq 5 becomes identical to the charge balance model presented by Renard et al. (2013) who predicted electrification phenomena of aerosols in the real atmosphere. The mean value of the particle charge distributions can be given by:

$$160 \qquad \frac{dJ_k}{dt} = \frac{d}{dt} \left(\frac{\sum j N_{kj}}{\sum N_{kj}} \right). \tag{6}$$

- The mean charge J of the radioactive particles of size k can also be approximated using a simple
- charge balance equation (Kim et al., 2015):

163
$$\frac{dJ_k}{dt} = A_k + \beta_{k,J}^+ n_{\text{ion}}^+ - \beta_{k,J}^- n_{\text{ion}}^-, \tag{7}$$

- where A_k is the radioactivity of the size k particles. Similarly to eq 5, the mean charge
- accumulation rate of the radioactive particles (eq 7) depends on the competition between self-
- and diffusion charging. Eqs 5 and 7 indicate that the net charge of beta-emitting radioactive
- particles converges to a steady state where self-charging balances diffusion charging. The
- timescale, τ, needed to reach a steady state can be given by (Clement and Harrison, 1992):

$$\tau = \frac{1}{\beta^- n_{\text{ion}}^-},\tag{8}$$

- where $\overline{\beta}$ is the mean attachment coefficient between negative ions and particles. At steady
- state, the mean charge of the radioactive particles can be approximated using (Clement et al.,
- 172 1995):

173
$$J_{k} = \begin{cases} y - \left(\frac{y(X-1)}{\exp(2\lambda y) - 1}\right) & \lambda y > 0.22 & (9a) \\ y + \frac{X-1}{2\lambda} & \lambda y \le 0.22 & (9b) \end{cases}, \tag{9}$$

174 with
$$\lambda = \frac{e^2}{8\pi\varepsilon_0\varepsilon r_k k_B T}$$
, $y = \frac{\varepsilon_0 A_k}{e\mu_L n_{\text{ion}}^0}$, $n_{\text{ion}}^0 = \sqrt{\frac{q}{\alpha}}$, $X = \frac{\mu_+ n_{\text{ion}}^+}{\mu_L n_{\text{ion}}^-}$,

where ε_0 is the vacuum permittivity, ε is the dielectric constant of the air, r is the radius of particles, k_B is the Boltzmann constant, T is the temperature, $n_{\rm ion}^0$ is the mean ion concentration, and I is the ionization rate coefficient of beta-emitting radionuclides. Eq 9 suggests that the steady-state mean charge of beta-emitting radioactive particles is highly influenced by their size and decay rates, as well as the concentrations and mobilities of ions in air. The second term of the RHS of eq 9(b) represents charge accumulated only by diffusion charging; thus, it can be used to approximate the steady-state mean charge of nonradioactive particles, such as background aerosols that are externally mixed with radionuclides.

183 2.3. Population balance models

175

176

177

178

179

180

181

- 184 2.3.1. Bivariate population balance model
- A bivariate population balance model, expressed in terms of particle volume *x* and charge *j*, can be used to predict effects of coagulation on time-dependent changes in the particle size and charge distributions. In the bivariate population balance model, the time-evolution of the number densities of charged and uncharged particles, *n*, due to coagulation, can be given by (Zebel, 1958; Oron and Seinfeld, 1989a&b):

190
$$\frac{\partial n(x,j)}{\partial t} = \frac{1}{2} \sum_{j'=-\infty}^{\infty} \int_{0}^{x} F_{j',j-j'}(x',x-x') n(x',j') n(x-x',j-j') dx' - \sum_{j'=-\infty}^{\infty} \int_{0}^{\infty} F_{j,j'}(x,x') n(x,j) n(x',j') dx'$$
(10)

where *F* is the coagulation frequency (m³ s⁻¹), which is also called the coagulation (rate) coefficient (Jacobson, 2005; Seinfeld and Pandis, 2006), and which can be obtained by multiplying the collision frequency and the collision efficiency. The two terms on the RHS of eq 10 represent the production and loss rates of charged and uncharged particles by coagulation, respectively. A numerical solution of eq 10 can be obtained through the discretization of the integral terms, respectively (Oron and Seinfeld, 1989a&b). Vanni (2000) tested several sectional approaches and showed that the approach of Kumar and Ramkrishna (1996) is simpler and more accurate than other tested approaches, and also preserves mass and number of particles. Thus, the sectional approach of Kumar and Ramkrishna (1996) was used in this study to discretize the integral terms of eq 10, leading to the following discretized form:

$$201 \qquad \frac{dN_{kj}}{dt} = \sum_{j'=-\infty}^{\infty} \sum_{\substack{l,m \\ x_{k-1} \le x_l + x_m \le x_{k+1}}}^{l \ge m} \left(1 - \frac{1}{2} \delta_{l,m}\right) \eta_{l,m} F_{l,m,j-j',j'} N_{l,j-j'} N_{m,j'} - \sum_{j'=-\infty}^{\infty} \sum_{l=1}^{M} F_{k,l,j,j'} N_{k,j} N_{l,j'} , \qquad (11)$$

where indices l and m refer to the size bins, δ is the Kronecker delta, $\eta_{l,m}$ is a property distribution factor between two size bins given by Kumar and Ramkrishna (1996), and M is the total number of the size bins. If coagulation is induced by thermal energy (i.e., Brownian coagulation), the Brownian collision frequency β^{Br} (m³ s⁻¹), which is also called the Brownian collision kernel (Jacobson, 2005), is given by (Fuchs, 1989):

$$\beta_{kl}^{Br} = 4\pi \left(r_k + r_l\right) \left(D_{p,k} + D_{p,l}\right) \left(\frac{r_k + r_l}{r_k + r_l + \sqrt{g_k^2 + g_l^2}} + \frac{4\left(D_{p,k} + D_{p,l}\right)}{(r_k + r_l)\sqrt{v_{p,k}^2 + v_{p,l}^2}}\right)^{-1},$$
(12)

where D_p is the particle diffusion coefficient, g is the particle mean traveling distance, and \overline{v}_p is the particle thermal speed in air. Coagulation of charged particles is influenced by electrostatic

- 210 particle-particle interactions. This effect can be accounted for by multiplying the collision
- frequency with the collision efficiency, α^{Br} , which is also called the coalescence efficiency
- 212 (Fuchs, 1989; Jacobson, 2005; Seinfeld and Pandis, 2006):

$$213 \qquad \alpha_{kl}^{Br} = \frac{u}{e^u - 1} \tag{13}$$

214 with
$$u = \frac{j_k j_l e^2}{4\pi\varepsilon_0 \varepsilon (r_k + r_l) k_B T}$$

- 215 In eq 13, u indicates the relative importance between electrostatic potential energy and thermal
- 216 energy in coagulation.
- 2.3.2. Monovariate population balance model
- The time-evolution of the size distribution of particles can be estimated using a monovariate
- 219 population balance model, with only the particle volume as the variable (Kumar and
- 220 Ramkrishna, 1996):

$$221 \qquad \frac{dN_k}{dt} = \sum_{\substack{l,m\\x_{k-1} \le x_l + x_m \le x_{k+1}}}^{l \ge m} \left(1 - \frac{1}{2} \delta_{l,m}^{Kr}\right) \eta_{l,m} F_{l,m} N_l N_m - \sum_{l=1}^M F_{k,l} N_k N_l \ . \tag{14}$$

- The coagulation frequency can be simply corrected using the mean charge of particles. However,
- the collision efficiency computed with the mean charge can be different from that with the
- particle charge distributions (Matsoukas, 1997). To include effects of the particle charge
- distributions on the coagulation frequency, eq (13) can be replaced by the average collision

efficiency $\bar{\alpha}$ (Clement et al., 1995), which involves interaction of all charged particles of size k with any charged particles of size l.

228
$$\overline{\alpha}_{kl}^{Br} = 1 + \frac{\sum_{j_k j_l \neq 0} N_{k, j_k} N_{l, j_l} \left(\alpha_{kl}^{Br} - 1 \right)}{\sum_{j_k} N_{k, j_k} \sum_{j_l} N_{l, j_l}}.$$
 (15)

- The particle charge distributions needed to calculate $\bar{\alpha}^{Br}$ can be obtained by assuming a
- 230 Gaussian distribution:

231
$$N_{kj} = \frac{\sum_{j} N_{k,j}}{\sqrt{2\pi}\sigma_k} \exp\left(-\frac{(j-J_k)^2}{2\sigma_k^2}\right)$$
 (16)

232 with

235

236

237

238

239

240

241

242

$$233 \qquad \sigma^2 = y + \frac{1}{2\lambda}.$$

2.4. Approaches to couple particle charging with coagulation kinetics

Figure 1 shows three approaches which can be used to predict the time-evolution of the charge and size distributions of particles in the atmosphere. All the approaches can be used to simulate charging and coagulation kinetics of atmospheric particles carrying contaminants, including radioactive particles. Approach 1 is a rigorous scheme that simultaneously computes both charge accumulation and coagulation rates of particles using the ion balance model (eqs 1 and 2), the charge balance model (eq 5), and the bivariate population balance model (eq 11). Approach 2 is a simplified scheme of Approach 1, which can be used to predict the particle charge distribution using the mean charge of particles (eq 7) and the Gaussian distribution (eq 16). In order to easily

simulate the coagulation of charged particles, Approach 2 employs the monovariate population balance model (eq 14) that corrects the collision frequency using the average collision efficiency (eq 15). Approach 2 can be simplified to Approach 3 by assuming that charge accumulation rates of particles instantaneously reach a steady state, with a timescale based on 5 times larger than τ from eq 8. The steady-state particle charge distribution can be approximated by eqs 9 and 16. In Approach 3, the collision frequency is multiplied by the average collision efficiency to include the influence of electrostatic forces on coagulation. In this work, polarization of particles is not taken into account.

3. RESULTS AND DISCUSSION

3.1. Methods to simulate particle charging

The three approaches attained above employ different methods to simulate charging of particles. These methods were evaluated by comparing their prediction results with measurements obtained using radioactive charge neutralizers (Liu and Pui, 1974; Wiedensohler and Fissan, 1991; Alonso et al., 1997) and radioactive particles (Gensdarmes et al., 2001). Initial conditions for the simulations were determined from the measurements. The properties of ions observed during the measurements are shown in Table 1. For the measurements providing the values of ion mass, $\beta_{k,j}^{\pm}$ was calculated using Fuchs (1963) and Hoppel and Frick (1986). However, the mass of ions was not measured during the experiments performed by Gensdarmes et al. (2001). In these experiments, $\beta_{k,j}^{\pm}$ was estimated using analytical equations given by Gunn (1954) and Harrison and Carslaw (2003).

3.1.1. Diffusion-Charging Mechanism

264

265

266

267

268

269

270

271

272

273

274

275

276

277

278

279

280

281

282

283

284

Figure 2 shows the steady-state charge distributions of nonradioactive particles over a wide size range. Here, the particles were charged by the diffusion charging mechanism. For particles larger than approximately 0.04 µm in diameter, the prediction results of all approaches were in good agreement with the measurements [Figures 2 (a) and (b)]. Below 0.04-µm particle size, Approach 1 accurately forecasted the particle charge distributions, but Approaches 2 and 3 underestimated the number concentrations of the negatively charged particles [Figure 2 (a)] although the mean charge values of the particles given by all approaches were comparable. Similar discrepancies were observed for the number concentrations of the positively charged particles smaller than about 0.025 µm (not shown). Analysis of the discrepancies suggests that they originate from the standard deviation involved in the Gaussian distribution (eq 16). At a given temperature, the width of the particle charge distributions can be significantly influenced by three parameters: the particle size, ion mass, and ion mobility (Wiedensohler and Fissan, 1991). In Approaches 2 and 3, however, the effects of the ion properties are not involved, so particle size primarily drives the standard deviation, which can differ from what Approach 1 gives. When Approaches 2 and 3 used the standard deviation values obtained by Approach 1, their simulation results became closer to the measurements, although the discrepancies are still seen for negatively and positively charged particles smaller than about 0.02 µm.

3.1.2. Competition of Self-Charging and Diffusion-Charging Mechanisms

285

286

287

288

289

290

291

292

293

294

295

296

297

298

299

300

301

302

303

304

305

In our previous work (Kim et al., 2014; 2015), it has been shown that Approaches 1 and 3 can reliably simulate charging of radioactive particles. Thus, in this study, we focused on evaluating the validity of Approach 2 with the experiments of Gensdarmes et al. (2001) who measured the charge distributions of ¹³⁷Cs particles under various ionizing conditions. Ionizing rates of air molecules were estimated using a linear energy transfer equation for energetic electrons emitted by beta decay (Kim et al., 2015). Results of Approach 1 were included as a reference. Figure 3 shows the charge accumulation on radioactive particles under two ionizing conditions: $q_I = 3.7 \times 10^8 \text{ m}^{-3} \text{ s}^{-1}$ and $q_I = 7.1 \times 10^6 \text{ m}^{-3} \text{ s}^{-1}$. Approach 2 predictions are in good agreement with observations and Approach 1 values. During the measurements, the self-charging rate of the radioactive particles was constant because of the long half-life of ¹³⁷Cs (approximately 30 years), suggesting that changes in their charge accumulation rates may be dominated by diffusion charging rates. The ion concentrations in air can rapidly increase at the high ionizing rate considered, suggesting that the diffusion charging rate of the ¹³⁷Cs particles quickly increased and then became comparable to their self-charging rate (eq 7). The charge accumulation on the radioactive particles promptly reached a steady-state value, and the particle charge distribution was similar to the initial condition (Figure 4). In contrast, the time required to reach the steadystate value was much longer at the low ionizing rate considered; hence, the particle charge distribution shifted to the right (see Figure 4), i.e., more positive charge. The agreement observed in Figure 4 between simulation results by Approach 2 and experimental data by Gensdarmes et al. (2001) suggests that Approach 2 can accurately forecast the competition between self- and

diffusion charging on submicron particles carrying radionuclides, and precisely predict the particle charge distributions.

3.1.3. Timescale to Reach Steady-State Charge Accumulation Rate

306

307

308

309

310

311

312

313

314

315

316

317

318

319

320

321

322

323

324

325

326

327

To evaluate the steady-state assumption of particle charging for atmospheric conditions, the timescale for reaching steady-state (eq 8) is evaluated with Approaches 1 and 2. Figure 5 shows time-dependent changes in the concentrations of negatively charged particles under two different initial conditions of Alonso et al. (1997) who measured the charge distributions of particles of a few nanometers. All particles were initially uncharged or negatively charged. Because the particle size was very small, Approach 1 was used to predict the time-evolution of the particle concentrations. As time elapsed, the initially uncharged particles became negatively charged by capturing negative ions. The diffusion of positive ions led to the discharging of the initially negatively charged particles. For the initial conditions used, the charging and discharging rates of the particles reached a steady state after approximately 0.2 s, respectively. This charging/discharging behavior predicted by Approach 1 is in good agreement with the measurements of Alonso et al.(1997). However, the timescales obtained from eq 8 are shorter than the prediction results, as well as the measurements, because Approach 1 and the observations provided exact timescales, while τ in eq 8 is a scaling parameter. Similar results were also obtained for different initial conditions, as well as for the ¹³⁷Cs particles (see Table 2). Eq 8 is based on the assumptions that (i) $X \approx 1$, (ii) all particles are initially uncharged, and (iii) the ion concentrations are constant (Clement and Harrison, 1992). As seen in Figure 6, when all the assumptions were applied, the diffusion charging rate of the radioactive particles became - 8.1×10^{-3} s⁻¹ at the timescale provided by eq 8, which corresponds to approximately 63% of the

steady-state self-charging rate. If the timescale is increased by a factor of five, the diffusion charging rate reaches about 99 % of the steady-state self-charging rate. Similar results were obtained for other cases with radioactive particles. Thus, eq 8 is valid if the assumptions can be used, and a reliable timescale to reach a steady state (e.g., 99%) can be obtained by multiplying the equation by a factor of 5. However, because the assumptions cannot be used in typical atmospheric conditions, such as $X \neq 1$ (Harrison and Carslaw, 2003), the steady-state assumption of Approach 3 should be evaluated using Approaches 1 or 2.

So far, we have evaluated the validity of the methods used in the three approaches to predict charge accumulation on atmospheric particles. The evaluation results suggest that the method employed in Approach 1 can accurately simulate charging of particles in the free molecule (d_p < 0.01 μ m), transition (d_p = 0.01 – 0.2 μ m), and continuum (d_p > 0.2 μ m) regimes. The methods used in Approaches 2 and 3 can reliably forecast charging rate of atmospheric particles larger than 0.04 μ m.

3.2. Validity of the three approaches to couple particle charging with coagulation

3.2.1. Bivariate Population Balance Model for Approach 1

Based on the numerical approach of Alonso et al.(1998), Alonso (1999) suggested an analytical approach to simultaneously investigate charging and coagulation kinetics of nonradioactive particles, smaller than $0.02~\mu m$ in diameter. Results of the analytical approach agreed well with those of the numerical approach, but the applicability of both analytical approaches may be limited as discussed in section Introduction. The analytical approach, however, was found to be

useful to validate numerical solutions of population balance equations including diffusion charging and coagulation (Alonso, 1999). In this study, the analytical approach of Alonso (1999) was used to evaluate Approach 1 because Approaches 2 and 3 are not applicable to particles smaller than 0.02 µm, as shown in Figure 2a. In integrating eq 11, an equidistant diameter grid was used for discretization. Particles were assumed to be initially uncharged, monodispersed (d_n) = 0.003 μ m, N_t = 10¹⁵ m⁻³), with a constant ion concentration. The mobility of negative ions was slightly greater than that of positive ions, according to the properties taken from Alonso et al. (1997) (See Table 1). Figure 7 shows the time dependent concentration of particle charge classes. The concentration of uncharged 0.003-um particles decreased with time because of loss due to coagulation and charging by captured ions. However, the loss of particle concentration caused by diffusion charging was negligible because the ion-particle attachment coefficient was small, suggesting that the time-evolution of the charged-particle concentration depends on the uncharged-particle concentration. More negative than positive ions were captured by the uncharged 0.003-um particles and, thus, the concentration of the negatively charged 0.003-um particles is slightly higher than that of the positively charged 0.003-um particles. The concentrations of larger particles (e.g., charged and uncharged 0.006-um particles) increased over time because of the size growth of small particles due to coagulation, as well as the diffusion-charging mechanism. These evolution patterns predicted by Approach 1 were in good agreement with the prediction results given by the analytical approach. As can be seen in Figures 2 and 5, the ion balance and charge balance models of Approach 1 accurately predicted the diffusion charging of nanoparticles, suggesting that the numerical solution of the bivariate population balance model (eq 11) reliably predicts coagulation of particles acquiring charge.

349

350

351

352

353

354

355

356

357

358

359

360

361

362

363

364

365

366

367

368

369

3.2.2. Average Collision Efficiency of Approaches 2 and 3

372

373

374

375

376

377

378

379

380

381

382

383

384

385

386

387

388

389

390

391

392

393

Approaches 2 and 3 employ an average collision efficiency and are coupled to the monovariate, instead of the bivariate, population balance model. These approaches provided accurate particle charge distributions for various cases (e.g., Figures 2 and 4; Kim et al., 2014), suggesting that their validity may be highly influenced by the accuracy of the average collision efficiency. Thus, we compared simulation results of Approach 2 with those of Approach 1 to check if the average collision efficiency (eq 15) can appropriately account for the influence of the charge distributions of particles on their size growth via coagulation. For comparison, simulation results of Approach 2 using the mean charge (eq 13), as well as those for uncharged particles, were included. Similarly to Oron and Seinfeld (1989 a&b), we assumed monodispersed initial size distributions $(d_p = 0.1 \text{ } \mu\text{m}, \ 0.5 \text{ } \mu\text{m}, \ 1 \text{ } \mu\text{m}, \ N_t = 10^{13} \text{ } \text{m}^{-3}, \ \text{and} \ n_{\text{ion}}^0 = 10^{16} \text{ } \text{m}^{-3}).$ The geometrical grids $(x_{k+1} = 2x_k)$ were used to cover a wide particle-size range. Other basic assumptions were similar to those considered for the validation test of Approach 1. Figure 8 shows the time-evolution of the particle size distributions induced by particle charging and coagulation. The simulation conditions led to the accumulation of more negative than positive charges on the particles. At t = 1 min, the number fraction of the negatively charged particles was 0.7, while that of the positively charged and uncharged particles was 0.16 and 0.14, respectively. Thus, the size growth of the particles by coagulation was suppressed due to the generation of strong repulsive electrostatic forces (Approaches 1 and 2 vs Uncharged). While most particles were negatively charged, some particles captured positive ions. Owing to electrostatic attractive forces, the positively charged particles can more frequently coagulate with the negatively charged particles and grow. Therefore, the coagulation rates predicted by

Approach 2 with the average collision efficiency were slightly higher than those for the case assuming that all particles were negatively charged [Approach 2 (eq 13) vs Approach 2 (eq 15)]. These coagulation patterns predicted by Approach 2 using the average collision efficiency were in good agreement with those given by Approach 1 [Approach 1 vs Approach 2 (eq 15)], as well as the particle charge distributions in various size ranges (Figure S1). Similar results were obtained for different initial particle size distributions ($d_p = 0.1 \mu m$; $d_p = 1 \mu m$). Thus, the monovariate population balance model using the average collision efficiency can be used to simultaneously investigate charging and coagulation kinetics of atmospheric particles. These validation tests suggest that all the approaches developed can be used to reliably couple particle charging with coagulation kinetics of atmospheric particles.

3.3. Applications

3.3.1. Radioactive Neutralizer

Radioactive neutralizers are typically used to control the charge of atmospheric particles in many laboratory-scale experiments. The applicability of the three approaches to studies using radioactive neutralizers was evaluated using the experiments of Alonso et al. (1998) who measured the size distribution of nanometer-size particles passing by a ²⁴¹Am radioactive neutralizer under various residence times.

Figure 9 shows the size distribution of negatively charged particles when the residence time, t = 0.318 s. All particles were initially uncharged, but some particles became charged by capturing positive and negative ions in the neutralizer. Approach 1 accurately predicted the size

distribution of the negatively charged particles, while the predictions by Approaches 2 and 3 were different from the measurements. The particle size distributions predicted by all the approaches were similar (not shown). It can be concluded that the Gaussian distribution used in Approaches 2 and 3 cannot accurately predict the charge distributions of very small particles (See Figure 2). Thus, Approaches 2 and 3 should not be used for particles smaller than 0.04 μm. As shown in Figures 2-5, Approach 1 can accurately predict the charge accumulation rate of radioactive and nonradioactive particles in the free-molecule, transition, and continuum regimes. Approach 1 employs the interpolation formula of Fuchs that can be used to compute the collision frequency of the particles in these regimes, revealing that this approach can also precisely predict the charge distribution of larger particles undergoing coagulation. These results suggest that Approach 1 can be a reasonable option to simultaneously simulate charging and coagulation of particles of any size in laboratory-scale experiments. 3.3.2. Charging and Coagulation of Nonradioactive Particles in Urban Atmosphere Hoppel (1985) simulated charging of 0.06-µm urban aerosols by diffusion charging; however, effects of coagulation on their steady-state charge distribution were excluded from the simulation. Changes in the particle charge and size distributions by charging and coagulation were investigated in this work on the basis of the simulation of Hoppel (1985) for comparison. The simulation time was approximately 100 min, but for completeness, we repeated and extended the

giving $q_b \approx 10^7 \,\text{m}^{-3} \,\text{s}^{-1}$ (Hoppel, 1985).

415

416

417

418

419

420

421

422

423

424

425

426

427

428

429

430

431

432

433

434

435

436

Hoppel (1985) simulation to 1 day. The extended results were compared with prediction results

of Approach 1, which involves the effects of coagulation on the particle charge distribution. It

was assumed that cosmic rays and natural radioactivity generate ion pairs in the atmosphere,

Figure 10 presents changes in the particle charge and size distributions vs time. The simulation results performed by Hoppel (1985) showed that the particle charge distribution approached its steady-state value after approximately 90 min. However, as time elapsed, the particles grew in size due to coagulation. The size growth led to the generation of large particles capturing many ions, thereby modifying the particle charge distribution.

The simulation results of Hoppel (1985) also indicated that the ion concentrations became unchanged after reaching a steady state. However, coagulation reduced the particle number concentrations which can affect the loss rate of ions by diffusion charging (see eqs 1 and 2). The reduction in the particle concentrations increased the ion concentrations, thereby enhancing the electrical conductivity of the postulated atmosphere (Figure 11). The ion concentrations and electrical conductivity are expected to increase until ion-ion recombination becomes the major ion removal mechanism. These results suggest that coagulation can affect the electrical properties in the atmosphere, as well as the particle charge distribution.

- 3.3.3. Charging and Coagulation of Radioactive Particles in the Atmosphere
- 451 3.3.3.1. Comparison with results by Greenfield (1956)

Nuclear events can release particles carrying radionuclides. Greenfield (1956) simulated time-evolution of the charge distribution of 0.1 µm radioactive particles emitting energetic electrons at 6-km altitude. Because Greenfield (1956) assumed that the particle size distribution is constant for 4 hours, the influence of coagulation on the particle charge distribution was evaluated using the simulation conditions postulated by Greenfield (1956).

Figure 12 shows changes in the particle charge distributions vs time. Both self-charging and diffusion charging influenced the charge accumulation on radioactive particles. Due to many ion pairs produced by beta radiation (Figure S2), positive charge accumulated on the particles by self-charging was rapidly neutralized by capturing negative ions. Thus, the particle charge distribution given by Approach 1 was slightly shifted to the right of the zero elementary charge in Figure 12(a), although large particles with a high level of radioactivity were generated by coagulation. As time elapsed, the particle charge distribution was slightly moved to the left. Because the decay rates of the highly radioactive particles were reduced over time, their selfcharging rates also decreased, and this led to the slight movement of the charge distribution to the left in Figure 12(a). The discrepancies between the predictions of Approach 1 and Greenfield (1956) result mainly from the ion-particle attachment coefficient used in the simulation. The values assumed by Greenfield (1956) were beyond the ion-particle attachment coefficient found by other researchers (e.g., Hoppel and Frick, 1986), leading to the discrepancies observed (Kim et al., 2015). Beta radiation caused by radioactive decay rapidly increased the ion concentrations, thereby enhancing the electrical conductivity in the atmosphere (Figure S2). In contrast to the case shown in Figure 11, the ion concentrations and air conductivity significantly decreased with time because the ionization rate of air molecules decreased considerably, and ion-ion recombination was responsible for the change in the concentrations. Nevertheless, the air conductivity enhancement by beta radiation was much higher than that by cosmic rays and natural radioactivity.

457

458

459

460

461

462

463

464

465

466

467

468

469

470

471

472

473

474

475

476

After the Chernobyl and Fukushima accidents, short- and long-range transport of particles carrying radionuclides, such as ¹³⁷Cs and ¹³⁴Cs, affected the electrical properties of the local atmosphere in many places (Israelsson and Knudsen, 1986; Yamauchi et al., 2012). In particular, beta radiation led to significant changes in the electrical conductivity and potential gradient in the local atmosphere (Kim et al., 2015). Israelsson et al. (1987) suggested that an increase in the electrical conductivity led to enhancement of lightning activities at radioactively contaminated sites in Sweden. These observations reveal that the approaches developed in this study can be employed to investigate the influence of radionuclides on electrification phenomena in the atmosphere.

3.3.3.2. Steady-state assumption of radioactive particle charging

The steady-state assumption of particle charging can be useful to simulate coagulation of radioactive particles in model studies of radioactivity transport. Charging and coagulation kinetics of radioactive particles were investigated using Approaches 2 and 3 to evaluate the validity of the steady-state assumption of radioactive particle charging. For comparison, the size growth of particles by coagulation was simulated by assuming the Boltzmann charge distribution. We used the simulation condition employed to validate the average collision efficiency, but additionally presumed that radioactive decay of ¹³⁴Cs is responsible for the ionization of air molecules. The specific radioactivity of ¹³⁴Cs was obtained from Clement and Harrison (1992). The ionization rate of ¹³⁴Cs was estimated according to Kim et al. (2015). Under these conditions, eq 8 revealed that $5\tau \approx 4.3$ ms. Thus, we assumed that charge accumulation rates of ¹³⁴Cs particles instantaneously reach steady state, and evaluated this assumption for the simulation conditions of $X \approx 0.7$ and ion concentration given by $q_I = I_{Cs-134} \times A_{Cs-134} \times N_t$.

Figure 13 shows the charge and size distributions of the 134 Cs particles after 2 hours of evolution. The prediction results of Approach 3 were different from those of the case assuming the Boltzmann charge distribution, but agreed well with those of Approach 2, suggesting that the steady-state assumption of radioactive particle charging can be valid if τ is small. We also tested the assumption of Approach 3 using different initial conditions (e.g., $d_p = 0.3 \mu m$), and the agreement was still maintained (not shown).

3.4. Computational costs

The computational costs to predict transport of particles containing contaminants depends on the number of ordinary differential equations (ODEs) solved during simulation. Thus, the number of ODEs involved in the three approaches was evaluated by assuming 30 size bins, which corresponds to those used in the two-moment aerosol sectional microphysics model, covering particle diameters from 0.01 µm to 10 µm (Adams and Seinfeld, 2002).

Table 3 shows an example of the computational costs of the three approaches. For Approach 1, we assumed that atmospheric particles can acquire up to fifteen elementary charges regardless of their sign, thereby resulting in 932 ODEs. Because Approaches 2 and 3 employed the monovariate population balance model, a fewer number of ODEs were involved in these Approaches than in Approach 1, suggesting that they are computationally more efficient. For instance, compared to Approach 1, Approaches 2 and 3 more quickly computed the charge

accumulation and coagulation rates of urban aerosols.

A simple way to reduce the number of ODEs included in Approach 1 is to assume that atmospheric particles acquire only a few electrical charges. For example, Laakso et al. (2002) assumed that submicron particles can acquire elementary charges from -5 to +5. This assumption can be valid if the particle size is small [see Figure 2 (a)]. However, when the simulation conditions were used, we observed loss of submicron particles because they can acquire more elementary charges (Figure S3). To preserve mass and charge, one may optimize the minimum number of elementary charges using the charge balance model of Approach 1, and then begin the simulation of charging of particles undergoing coagulation.

Approach 3 includes all the physics of charging and coagulation to predict the particle size/charge distribution, but, compared to Approaches 1 and 2, it is computationally more suitable for use in a 3-D global transport model to predict the transport of radioactivity in the environment after a radiological event such as a nuclear plant accident.

4. CONCLUSIONS

Understanding the behavior of atmospheric particles is important to accurately predict short- and long-range transport of contaminants. Particle charging and coagulation processes can strongly affect the behavior of atmospheric particles because these processes can change their important physical and electrical properties, such as size and charge. This study has shown three approaches with a wide range of complexity and applications to involve the mutual effects of charging and coagulation processes in the simulation of particle charge and size distributions vs time. Depending on the initial conditions, these approaches can be employed to accurately

predict the behavior of atmospheric particles carrying radioactive contaminants. We have shown the approaches to be applicable to a wide variety of atmospheric (laboratory and field) applications. The accuracy of the approaches depends on the assumptions made to reduce computational cost. The developed approaches can be readily incorporated into microphysical and transport models of any scale to account for charging phenomena of atmospheric particles.

ACKNOWLEDGMENTS. This work was supported by the Defense Threat Reduction Agency under grant number DTRA1-08-10-BRCWMD-BAA. The manuscript has been co-authored by UT-Battelle, LLC, under Contract No. DEAC05-00OR22725 with the U.S. Department of Energy.

REFERENCES

- Adams, P. J. and Seinfeld, J. H.: Predicting global aerosol size distributions in general circulation
- 553 models, J. Geophys. Res., 107(D19), 4370, doi:10.1029/2001JD001010, 2002.
- Alonso, M., Kousaka, Y., Nomura, T., Hashimoto, N., and Hashimoto, T.: Bipolar charging and
- neutralization of nanometer-sized aerosol particles, J. Aerosol Sci., 28, 1479-1490,
- doi:10.1016/S0021-8502(97)00036-0, 1997.
- Alonso, M., Hashimoto, T., Kousaka, Y., Higuchi, M., and Nomura, T.: Transient bipolar
- charging of a coagulating nanometer aerosol. J. Aerosol Sci., 29, 263-270,
- 559 doi:10.1016/S0021-8502(97)10007-6, 1998.
- Alonso, M.: Simultaneous charging and Brownian coagulation of nanometre aerosol particles, J.
- Phys. A: Math. Gen., 32, 1313-1327, doi:10.1088/0305-4470/32/8/003, 1999.
- 562 Chin, C.-J., Yiacoumi, S., and Tsouris, C: Shear-induced flocculation of colloidal particles in
- stirred tanks, J. Colloid Interface Sci., 206, 532-545, doi:10.1006/jcis.1998.5737, 1998.
- Clement, C. F., and Harrison, R. G.: The charging of radioactive aerosols, J. Aerosol Sci., 23,
- 565 481-504, doi:10.1016/0021-8502(92)90019-R, 1992.
- Clement, C.F., Clement, R.A., and Harrison, R.G: Charge distributions and coagulation of
- radioactive aerosols, J. Aerosol Sci., 26, 1207-1225, doi:10.1016/0021-8502(95)00525-0,
- 568 1995.
- 569 Fuchs, N. A.: On the stationary charge distribution on aerosol particles in a bipolar ionic
- atmosphere. Geofis. Pura. Appl., 56, 185–193, doi: 10.1007/BF01993343, 1963
- Fuchs, N.A.: The Mechanics of Aerosols; Dover Publications, 1989.

572 Gensdarmes, F., Boulaud, D., and Renoux, A.: Electrical charging of radioactive aerosols— 573 comparison of the Clement-Harrison models with new experiments, J. Aerosol Sci. 32, 574 1437-1458, doi:10.1016/S0021-8502(01)00065-9, 2001. 575 Greenfield, S. M.: Ionization of radioactive particles in the free air, J. Geophys. Res., 61, 27-33, 576 doi:10.1029/JZ061i001p00027, 1956. 577 Gunn, R.: Diffusion charging of atmospheric droplets by ions, and the resulting combination 578 coefficients, J. Meteor., 11, 339-347, doi: http://dx.doi.org/10.1175/1520-579 0469(1954)011<0339:DCOADB>2.0.CO;2, 1954. 580 Harrison, R. G., and Carslaw, K. S.: Ion-aerosol-cloud processes in the lower atmosphere, Rev. 581 Geophys., 41, 1012, doi: 10.1029/2002RG000114, 2003. 582 Hoeve, J. E. T., and Jacobson, M. Z.: Worldwide health effects of the Fukushima Daiichi nuclear 583 accident. Energy Environ. Sci., 5, 8743-8757, doi: 10.1039/C2EE22019A, 2012. 584 Hoppel, W. A.: Ion-aerosol attachment coefficients, ion depletion, and the charge distribution on 585 aerosols, J. Geophys, Res., 90, 5917-5923, doi: 10.1029/JD090iD04p05917, 1985. 586 Hoppel, W. A., and Frick, G. M.: Ion-aerosol attachment coefficients and the steady-state charge 587 distribution on aerosols in a bipolar ion environment, Aerosol Sci. Technol., 5, 1-21, 588 doi:10.1080/02786828608959073, 1986. 589 Jacobson, M. Z.: Fundamentals of atmospheric modeling, Cambridge University Press, New 590 York, 2005. 591 Israelsson, S., and Knudsen, E.: Effects of radioactive fallout from a nuclear power plant 592 accident on electrical parameters, J. Geophys. Res., 91, 11909-11910, doi: 593 10.1029/JD091iD11p11909, 1986.

594 Israelsson, S., Schütte, T., Pisler, E., and Lundquist, S.: Increased occurrence of lightning flashes 595 in Sweden during 1986, J. Geophys. Res., 92, 10996-10998, doi: 596 10.1029/JD092iD09p10996, 1987. 597 Kim, Y.-H., Yiacoumi, S., Lee, I., McFarlane, J., and Tsouris, C.: Influence of radioactivity on 598 surface charging and aggregation kinetics of particles in the atmosphere, Environ. Sci. 599 Technol., 48, 182-189, doi: 10.1021/es4047439, 2014. 600 Kim, Y.-H., Yiacoumi, S., and Tsouris, C.: Surface charge accumulation of particles containing 601 radionuclides in open air, J. Environ. Radioactivity, 143, 91-99, doi: 602 10.1016/j.jenvrad.2015.02.017, 2015. 603 Kumar, S. and Ramkrishna, D.: On the solution of population balance equations by 604 discretization—I. A fixed pivot technique, Chem. Eng. Sci., 51, 1311-1332, 605 doi:10.1016/0009-2509(96)88489-2, 1996. 606 Kweon, H., Yiacoumi, S., Lee, I., McFarlane, J., and Tsouris, C.: Influence of surface potential 607 on the adhesive force of radioactive gold surfaces. Langmuir, 29, 11876-11883, doi: 608 10.1021/la4008476, 2013. 609 Laakso, L, Mäkelä, J. M., Pirjola, L., and Kulmala, M.: Model studies on ion-induced nucleation 610 in the atmosphere. J. Geophys. Res., 107, 4427, doi:10.1029/2002JD002140, 2002. 611 Liu, B. Y. and Pui, D. Y.: Equilibrium bipolar charge distribution of aerosols. J. Colloid 612 Interface Sci., 49, 305-312, doi:10.1016/0021-9797(74)90366-X, 1974. 613 Matsoukas, T.: The coagulation rate of charged aerosols in ionized gases. J. Colloid Interface 614 Sci., 187, 474-483, doi:10.1006/jcis.1996.4723, 1997.

615 Ooe, H., Seki, R., and Ikeda, N.: Particle-size distribution of fission products in airborne dust 616 collected at Tsukuba from April to June 1986, J. Environ. Radioactivity, 6, 219-223, 617 doi:10.1016/0265-931X(88)90078-1, 1988. 618 Oron, A., and Seinfeld, J. H.: The dynamic behavior of charged aerosols: II. Numerical solution 619 by the sectional method, J. Colloid Interface Sci., 133, 66-79, doi:10.1016/0021-620 9797(89)90282-8, 1989a. 621 Oron, A., Seinfeld, J. H.: The dynamic behavior of charged aerosols: III. Simultaneous charging 622 and coagulation, J. Colloid Interface Sci., 133, 80-90, doi:10.1016/0021-9797(89)90283-623 X, 1989b. 624 Pruppacher, H. R. and Klett, J. D.: Microphysics of Clouds and Precipitation, Kluwer Acad., 625 Norwell, Mass., 1997. 626 Renard, J.-B., Tripathi, S. N., Michael, M., Rawal, A., Berthet, G., Fullekrug, M., Harrison, R. 627 G., Robert, C., Tagger, M., and Gaubicher, B.: In situ detection of electrified aerosols in 628 the upper troposphere and stratosphere, Atmos. Chem. Phys., 13, 11187–11194, 629 doi:10.5194/acp-13-11187-2013, 2013. 630 Seinfeld, J. H., and Pandis, S.N.: Atmospheric chemistry and physics: from air pollution to 631 climate change, John Wiley and Sons, New Jersey, 2006. 632 Taboada-Serrano, P., Chin, C., Yiacoumi, S., and Tsouris, C.: Modeling aggregation of colloidal 633 particles, Curr. Opin. Colloid Interface Sci., 10, 123-132, 634 doi:10.1016/j.cocis.2005.07.003, 2005. 635 Tsouris, C., Yiacoumi, S., and Scott, T.: Kinetics of heterogeneous magnetic flocculation using a 636 bivariate population-balance equation, Chem. Eng. Commun., 137, 147-159, 637 doi:10.1080/00986449508936373, 1995.

638 Vanni, M.: Approximate population balance equations for aggregation-breakage processes. J. 639 Colloid Interface Sci., 221, 143-160, doi:10.1006/jcis.1999.6571, 2000. 640 Walker, M. E., McFarlane, J., Glasgow, D. C., Chung, E., Taboada-Serrano, P., Yiacoumi, S., 641 and Tsouris, C.: Influence of radioactivity on surface interaction forces, J. Colloid 642 Interface Sci., 350, 595-598, doi:10.1016/j.jcis.2010.06.042, 2010. 643 Wiedensohler, A., and Fissan, H.J.: Bipolar charge distributions of aerosol particles in high-644 purity argon and nitrogen, Aerosol Sci. Technol., 14, 358-364, doi: 645 10.1080/02786829108959498, 1991. 646 Yair, Y., and Levin, Z.: Charging of polydispersed aerosol particles by attachment of 647 atmospheric ions, J. Geophys. Res., 94, 13085-13091, doi: 10.1029/JD094iD11p13085, 648 1989. 649 Yamauchi, M., Takeda, M., Makino, M., Owada, T., and Miyagi, I.: Settlement process of 650 radioactive dust to the ground inferred from the atmospheric electric field measurement, 651 Ann. Geophys., 30, 49–56, doi:10.5194/angeo-30-49-2012, 2012. Yeh, H. C., Newton, G. J., Raabe, O. G., and Boor, D. R.: Self-charging of ¹⁹⁸Au-labeled 652 653 monodisperse gold aerosols studied with a miniature electrical mobility spectrometer. J. 654 Aerosol Sci., 7, 245-253, doi:10.1016/0021-8502(76)90039-2, 1976. 655 Yoshenko, V.I., Kashparov, V.A., Protsak, V.P., Lundin, S.M., Levchuk, S.E., Kadygrib, A.M., 656 Zvarich, S.I., Khomutinin, X.V., Maloshtan, I.M., Lanshin, V.P., Kovtun, M.V., and 657 Tschiervsch, J.: Resuspension and redistribution of radionucleotides during grassland and 658 forest fires in the Chernobyl exclusion zone. Part I: fire experiments, J. Environ.

Radioactivity, 86, 143-163, doi:10.1016/j.jenvrad.2005.08.003, 2006a.

660	Yoshenko, V.I. Kashparov, V.A., Levchuk, S.E., Glukhovskiy, A.S,. Khomutni, Y.V., Protsak,
661	V.P., Lundin, S.M., and Tschiersch, J.: Resuspension and redistribution of
662	radionucleotides during grassland and forest fires in the Chernobyl exclusion zone. Part
663	II: modeling the transport process, J. Environ. Radioactivity, 87, 260-278,
664	doi:10.1016/j.jenvrad.2005.12.003, 2006b.
665	Yu, F., and Turco, R. P.: From molecular clusters to nanoparticles: Role of ambient ionization in
666	tropospheric aerosol formation, J. Geophys. Res., 106, 4797-4814, doi:
667	10.1029/2000JD900539, 2001.
668	Zebel, G.: Zur theorie des verhaltens elektrisch geladener aerosole, Colloid Polym. Sci., 157, 37
669	50, doi: 10.1007/BF01734032, 1958.

Table 1. Properties of ions used for experimental observations

	Posit	ive ions	Negative ions	
References	Mass (AMU)	Mobility $(cm^2 V^{-1} s^{-1})$	Mass (AMU)	Mobility $(cm^2 V^{-1} s^{-1})$
Alonso et al. (1997)	150	1.15	80	1.65
Liu and Pui (1974) ^a	140	1.4	101	1.6
Wiedensohler and Fissan (1991)	140	1.4	101	1.6
Gensdarmes et al. (2001)	-	1.19	-	1.54

^a The properties of ions were obtained from Wiedensohler and Fissan (1991), who used a radioactive neutralizer similar to that employed by Liu and Pui (1974).

671672

Table 2. Timescales required for particles to reach steady-state charge

Charging mechanism	Particle diameter (µm)	Steady-state mean ion concentration (m ⁻³)	τ (s) (eq 8)	Numerical timescale (s) (Approach 1 or 2)	Timescale from measurments (s)	Reference
Diffusion charging	0.0071	8.0×10^{12}	0.04	$0.2 \sim 0.4$	0.3	
	0.0070	2.0×10^{13}	0.02	$0.2 \sim 0.4$	0.3	Alonso et al.
	0.0027	8.0×10^{12}	0.04	$0.2 \sim 0.4$	0.3	(1997)
	0.0026	2.0×10^{13}	0.02	$0.2 \sim 0.4$	0.3	
	0.82	2.1×10^{9}	171.1	> 1,000	-	
Diffusion charging	0.82	1.5×10^{10}	23.6	> 200	-	Gensdarmes
and self-charging	1.05	7.6×10^8	471.6	> 3,000	-	et al.(2001)
	1.05	7.0×10^9	51	> 500	-	

Table 3. Computational costs of the approaches used

Example: 30 size bins		Approach 1	Approach 2	Approach 3
	Ion balance model	2	2	-
The number	Charge balance model	20 × 21 = 020	30	-
of ODEs	Population balance model	$30 \times 31 = 930$	30	30
	Total	932	62	30
Computational time for urban aerosol ^{a,b} (s)		12724.2	302.6	8.5

^a Simulation conditions: $d_g = 0.116$ μm, $\sigma_g = 1.46$, $N_t = 6.718 \times 10^9$ m⁻³ and $q = 10^7$ m⁻³ s⁻¹ (Kim et al., 2015). The simulation time is 6 hours. ^b Computational resources: Intel(R) Core(TM)2 Duo CPU E6850 @ 3.00 GHz with 4GB RAM and

Matlab ODE solver.

676 Figure captions

- Figure 1. Three approaches to predict time-dependent changes in the particle size and charge
- distributions in the atmosphere.
- Figure 2. Steady-state charge distributions of particles capturing positive and negative ions. The
- symbols represent the measurements of the charge distributions of particles.
- Figure 3. Charge accumulation on 137 Cs particles under two ionizing conditions: $q_1 = 7.1 \times 10^6 \text{ m}^{-3}$
- 682 s⁻¹ and $q_I = 3.7 \times 10^8$ m⁻³ s⁻¹ ($d_p = 0.82$ μm; $A_{Cs-137} = 12.8$ mBq). The symbols represent the mean
- value of the particle charge distributions measured by Gensdarmes et al.(2001).
- Figure 4. Charge distributions of 137 Cs particles under two ionizing conditions: $q_1 = 7.1 \times 10^6 \text{ m}^{-3}$
- 685 s⁻¹ and $q_I = 3.7 \times 10^8 \text{ m}^{-3} \text{ s}^{-1}$ ($d_p = 0.82 \text{ } \mu\text{m}$; $A_{\text{Cs-137}} = 12.8 \text{ } \text{mBq}$). The prediction results of
- Approach 2 were compared with the measurements of Gensdarmes et al.(2001).
- Figure 5. Timescale to reach steady-state charge accumulation rates of 0.0071 μm nanoparticles.
- The lines are the simulation results of Approach 1. The symbols are the measurements of Alonso
- et al.(1997). Charging timescales were estimated using eq 8 ($\tau_{charging} = 0.042$ s and $\tau_{discharging} =$
- 690 0.017 s), as well as Approach 1 and the measurements.
- Figure 6. Charge accumulation rate of 137 Cs particles for each charging mechanism ($d_p = 0.82$
- 692 µm; $A_{Cs-137} = 12.8 \text{ mBq}$; $q_I = 7.1 \times 10^6 \text{ m}^{-3} \text{ s}^{-1}$). The assumptions used in eq 8 were applied to
- evaluate the validity of the equation. η_{Cs-137} corresponds to the self-charging rate of the
- 694 radioactive particles.
- Figure 7. Validation of the numerical solution for the bivariate population balance model under a
- monodispersed initial condition ($d_p = 0.003 \, \mu \text{m}$, $N_t = 10^{15} \, \text{m}^{-3}$, and $n_{\text{ion}}^0 = 10^{18} \, \text{m}^{-3}$). The lines
- and symbols represent the results of the analytical solution of Alonso (1999) and eq 11,
- 698 respectively.

- Figure 8. Time-evolution of the particle size distributions predicted by the monovariate
- population balance model with the average collision efficiency (eq 15) under a monodispersed
- 701 initial condition ($d_p = 0.5 \mu \text{m}$, $N_t = 10^{13} \text{ m}^{-3}$; $n_{\text{ion}}^0 = 10^{16} \text{ m}^{-3}$). Approach 1 was used as a reference
- that includes the mutual effect of surface charging and coagulation on the particle size and
- 703 charge distributions.
- Figure 9. Evolution of the size distribution of negatively charged particles in a ²⁴¹Am radioactive
- neutralizer. For the initial condition, $d_g = 0.0055 \, \mu \text{m}$, $\sigma_g = 1.23$, and $N_t = 5 \times 10^{15} \, \text{m}^{-3}$. The lines
- represent the simulation result of Approach 1. The symbols are the measurements of Alonso et
- 707 al.(1998).
- Figure 10. Time-evolution of the charge (a) and size (b) distributions of atmospheric particles in
- 709 the postulated atmosphere of Hoppel (1985) ($d_p = 0.06 \, \mu \text{m}$, $N_t = 2.3 \times 10^{10} \, \text{m}^{-3}$; $q = 10^7 \, \text{m}^{-3} \, \text{s}^{-1}$).
- Approach 1 was used to involve the effects of coagulation on the Hoppel (1985) simulation.
- Figure 11. Time-evolution of the mean ion concentration, n_0 and air conductivity, σ_{air} in the
- 712 postulated atmosphere of Hoppel (1985) ($d_p = 0.06 \, \mu \text{m}$, $N_t = 2.3 \times 10^{10} \, \text{m}^{-3}$; $q = 10^7 \, \text{m}^{-3} \, \text{s}^{-1}$).
- Figure 12. Time-evolution of the charge (a) and size (b) distributions of monodispersed
- radioactive particles at 6 km altitude ($d_p = 0.1 \, \mu \text{m}$, $N_t = 3.55 \times 10^{10} \, \text{m}^{-3}$; $I = 1.5 \times 10^4 \, \text{s}^{-1}$). Approach
- 715 1 was used to simultaneously simulate surface charging and coagulation of radioactive particles.
- Figure 13. The charge (a) and size (b) distributions of initially monodispersed 134 Cs particles (d_p
- 717 = 0.5 μ m, A_{Cs-134} = 14.5 Bq; N_t = 10¹³ m⁻³). The simulation time is 2 hours.

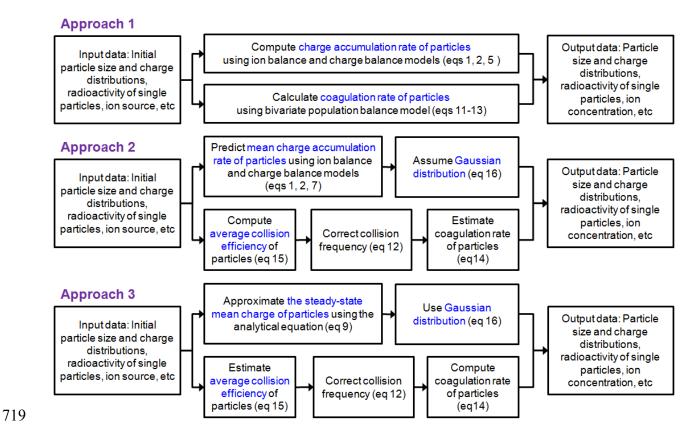


Figure 1. Three approaches to predict time-dependent changes in the particle size and charge distributions in the atmosphere.

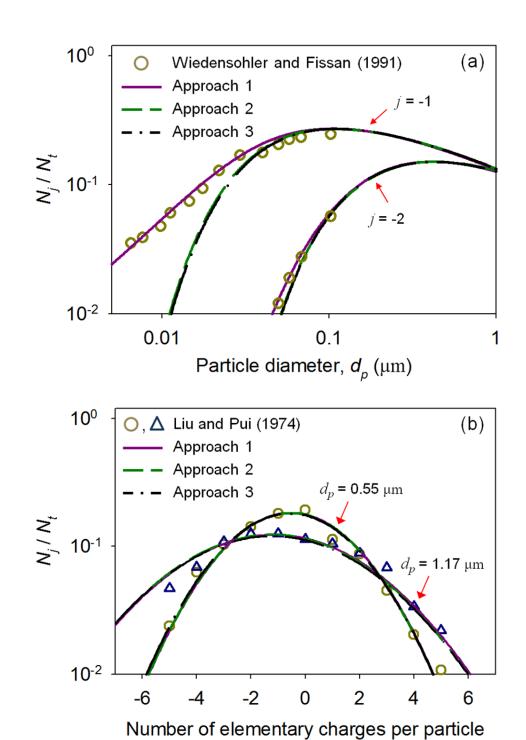


Figure 2. Steady-state charge distributions of particles capturing positive and negative ions. The symbols represent the measurements of the charge distributions of particles.

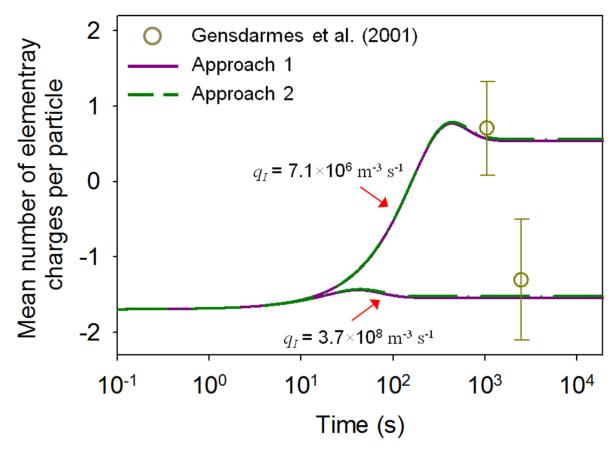


Figure 3. Charge accumulation on 137 Cs particles under two ionizing conditions: $q_I = 7.1 \times 10^6 \text{ m}^{-3}$ s⁻¹ and $q_I = 3.7 \times 10^8 \text{ m}^{-3} \text{ s}^{-1}$ ($d_p = 0.82 \text{ } \mu\text{m}$; $A_{\text{Cs-}137} = 12.8 \text{ mBq}$). The symbols represent the mean value of the particle charge distributions measured by Gensdarmes et al.(2001).

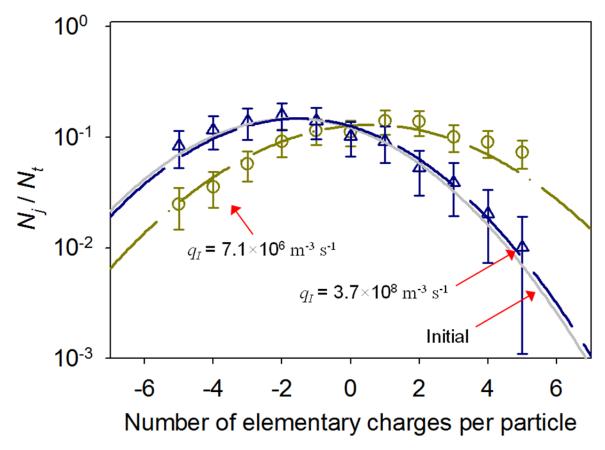


Figure 4. Charge distributions of 137 Cs particles under two ionizing conditions: $q_I = 7.1 \times 10^6 \text{ m}^{-3}$ s^{-1} and $q_I = 3.7 \times 10^8 \text{ m}^{-3} \text{ s}^{-1}$ ($d_p = 0.82 \text{ }\mu\text{m}$; $A_{\text{Cs-}137} = 12.8 \text{ }m\text{Bq}$). The prediction results of Approach 2 were compared with the measurements of Gensdarmes et al.(2001).

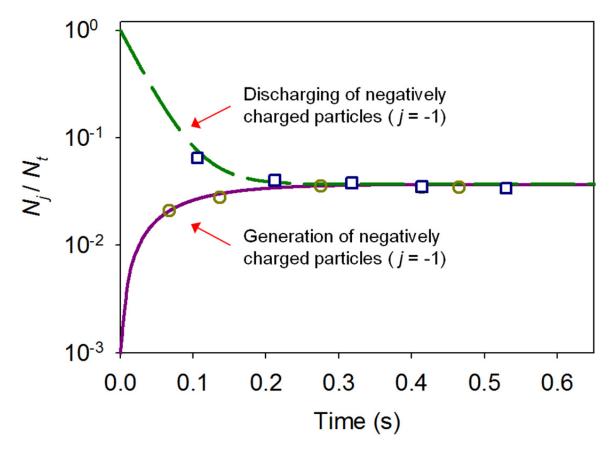


Figure 5. Timescale to reach steady-state charge accumulation rates of 0.0071 μ m nanoparticles. The lines are the simulation results of Approach 1. The symbols are the measurements of Alonso et al.(1997). Charging timescales were estimated using eq 8 ($\tau_{charging} = 0.042$ s and $\tau_{discharging} = 0.017$ s), as well as Approach 1 and the measurements.

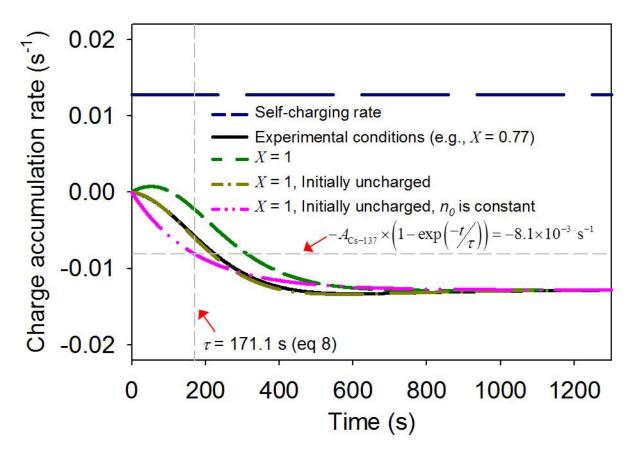


Figure 6. Charge accumulation rate of 137 Cs particles for each charging mechanism ($d_p = 0.82$ µm; $A_{Cs-137} = 12.8$ mBq; $q_I = 7.1 \times 10^6$ m⁻³ s⁻¹). The assumptions used in eq 8 were applied to evaluate the validity of the equation. A_{Cs-137} corresponds to the self-charging rate of the radioactive particles.

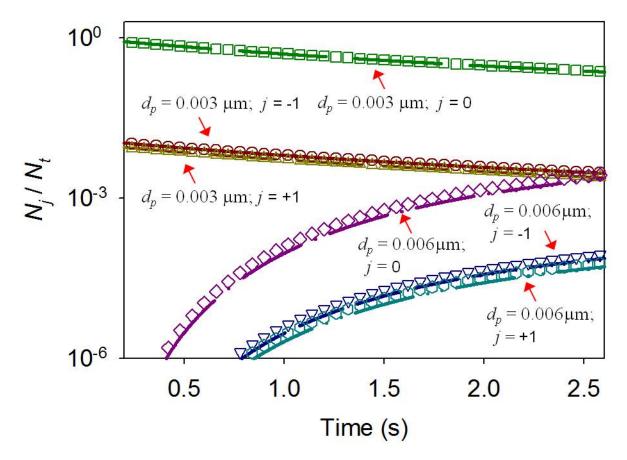


Figure 7. Validation of the numerical solution for the bivariate population balance model under a monodispersed initial condition ($d_p = 0.003 \, \mu \text{m}$, $N_t = 10^{15} \, \text{m}^{-3}$, and $n_{\text{ion}}^0 = 10^{18} \, \text{m}^{-3}$). The lines and symbols represent the results of the analytical solution of Alonso (1999) and eq 11, respectively.

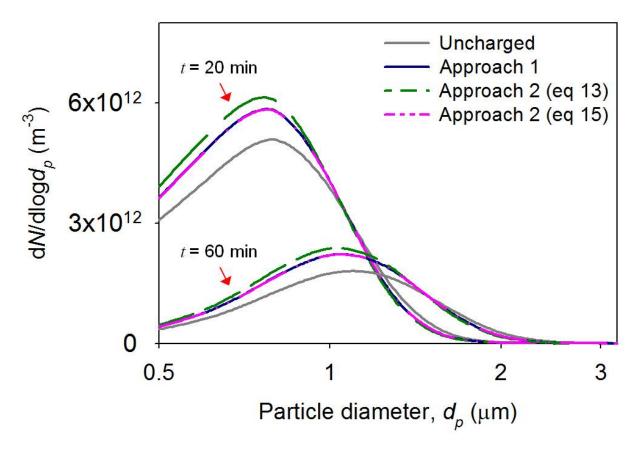


Figure 8. Time-evolution of the particle size distributions predicted by the monovariate population balance model with the average collision efficiency (eq 15) under a monodispersed initial condition ($d_p = 0.5 \, \mu m$, $N_t = 10^{13} \, m^{-3}$; $n_{\rm ion}^0 = 10^{16} \, m^{-3}$). Approach 1 was used as a reference that includes the mutual effect of surface charging and coagulation on the particle size and charge distributions.

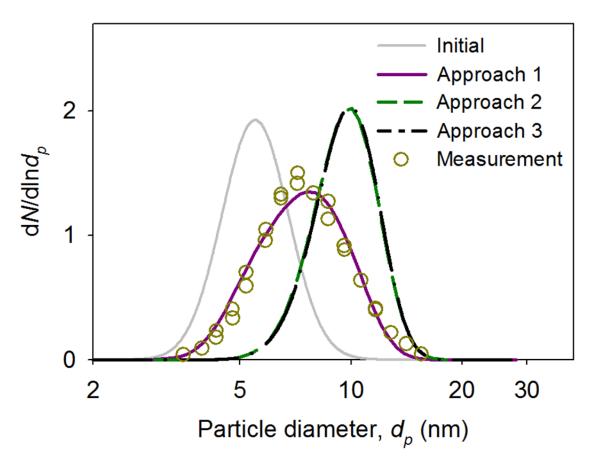


Figure 9. Evolution of the size distribution of negatively charged particles in a 241 Am radioactive neutralizer. For the initial condition, $d_g = 0.0055 \, \mu \text{m}$, $\sigma_g = 1.23$, and $N_t = 5 \times 10^{15} \, \text{m}^{-3}$. The simulation time is 0.318s. The measurements were taken from Alonso et al.(1998).

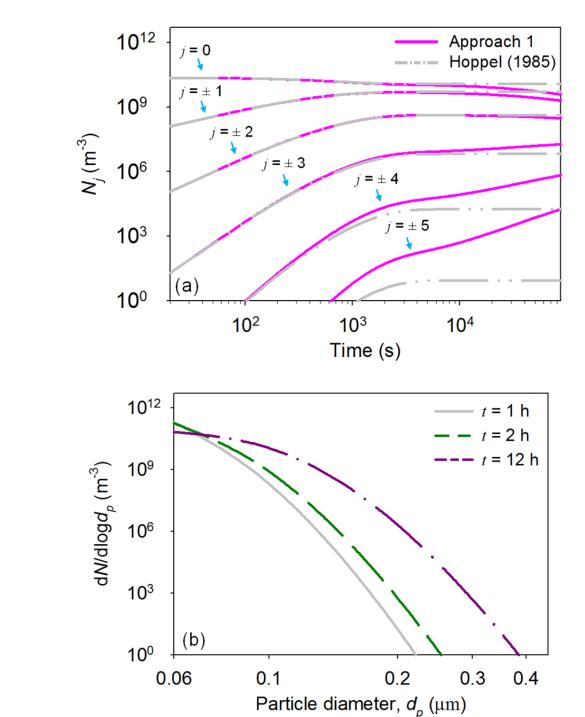


Figure 10. Time-evolution of the charge (a) and size (b) distributions of atmospheric particles in the postulated atmosphere of Hoppel (1985) ($d_p = 0.06 \, \mu \text{m}$, $N_t = 2.3 \times 10^{10} \, \text{m}^{-3}$; $q = 10^7 \, \text{m}^{-3} \, \text{s}^{-1}$). Approach 1 was used to involve the effects of coagulation on the Hoppel (1985) simulation.

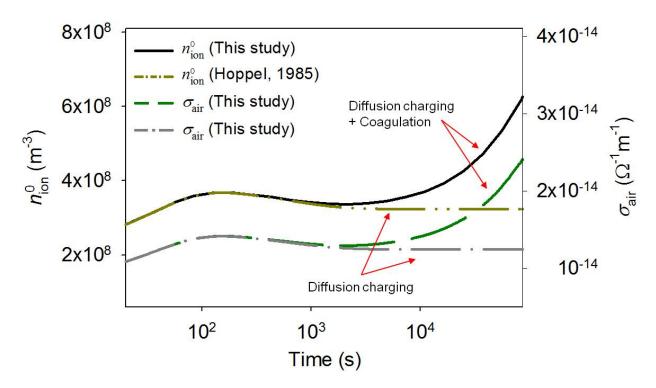


Figure 11. Time-evolution of the mean ion concentration, n_0 and air conductivity, $\sigma_{\rm air}$ in the postulated atmosphere of Hoppel (1985) ($d_p = 0.06 \, \mu \text{m}$, $N_t = 2.3 \times 10^{10} \, \text{m}^{-3}$; $q = 10^7 \, \text{m}^{-3} \, \text{s}^{-1}$).

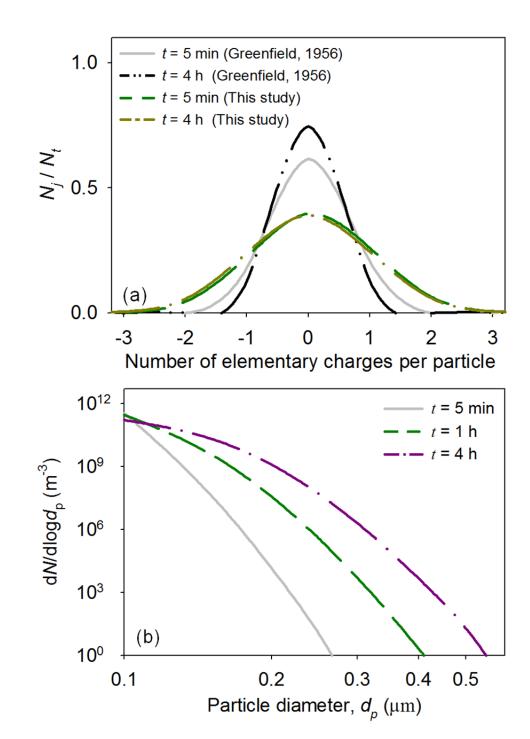
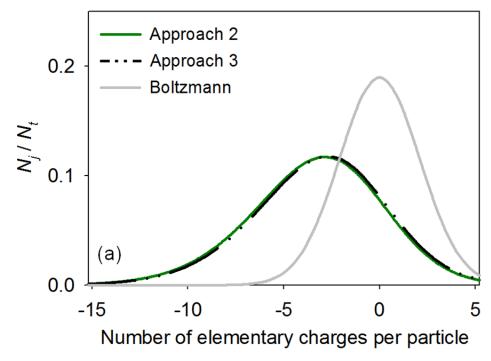


Figure 12. Time-evolution of the charge (a) and size (b) distributions of monodispersed radioactive particles at 6 km altitude ($d_p = 0.1 \, \mu m$, $N_t = 3.55 \times 10^{10} \, m^{-3}$; $I = 1.5 \times 10^4 \, s^{-1}$). Approach 1 was used to simultaneously simulate surface charging and coagulation of radioactive particles.



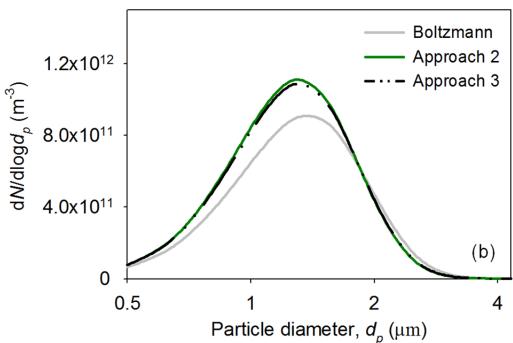


Figure 13. The charge (a) and size (b) distributions of initially monodispersed ¹³⁴Cs particles (d_p 781 = 0.5 μ m, A_{Cs-134} = 14.5 Bq; N_t = 10¹³ m⁻³). The simulation time is 2 hours.