#### 1 Non-OH chemistry in oxidation flow reactors for the study of atmospheric chemistry

### 2 systematically examined by modeling

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18 Abstract. Oxidation flow reactors (OFRs) using low-pressure Hg lamp emission at 185 and 254 nm 19 produce OH radicals efficiently and are widely used in atmospheric chemistry and other fields. However, 20 knowledge of detailed OFR chemistry is limited, allowing speculation in the literature about whether 21 some non-OH reactants, including several not relevant for tropospheric chemistry, may play an important role in these OFRs. These non-OH reactants are UV radiation,  $O({}^{1}D)$ ,  $O({}^{3}P)$ , and  $O_{3}$ . In this 22 23 study, we investigate the relative importance of other reactants to OH for the fate of reactant species in 24 OFR under a wide range of conditions via box modeling. The relative importance of non-OH species is 25 less sensitive to UV light intensity than to water vapor mixing ratio (H<sub>2</sub>O) and external OH reactivity 26 (OHRext), as both non-OH reactants and OH scale roughly proportionally to UV intensity. We show that 27 for field studies in forested regions and also the urban area of Los Angeles, reactants of atmospheric 28 interest are predominantly consumed by OH. We find that  $O(^{1}D)$ ,  $O(^{3}P)$ , and  $O_{3}$  have relative 29 contributions to VOC consumption that are similar or lower than in the troposphere. The impact of O 30 atoms can be neglected under most conditions in both OFR and troposphere. We define "riskier OFR 31 conditions" as those with either low H<sub>2</sub>O (<0.1%) or high OHR<sub>ext</sub> ( $\geq$ 100 s<sup>-1</sup> in OFR185 and >200 s<sup>-1</sup> in 32 OFR254). We strongly suggest avoiding such conditions as the importance of non-OH reactants can be 33 substantial for the most sensitive species, although depending on the species present OH may still 34 dominate under some riskier conditions. Photolysis at non-tropospheric wavelengths (185 and 254 nm) 35 may play a significant (>20%) role in the degradation of some aromatics, as well as some oxidation 36 intermediates, under riskier reactor conditions, if the quantum yields are high. Under riskier conditions, 37 some biogenics can have substantial destructions by O<sub>3</sub>, similarly to the troposphere. Working under 38 low O<sub>2</sub> (volume mixing ratio of 0.002) with the OFR185 mode allows OH to completely dominate over 39 O<sub>3</sub> reactions even for the biogenic species most reactive with O<sub>3</sub>. Non-tropospheric VOC photolysis may 40 have been a problem in some laboratory and source studies, but can be avoided or lessened in future 41 studies by diluting source emissions and working at lower precursor concentrations in laboratory studies, 42 and by humidification. Photolysis of SOA samples is estimated to be significant (>20%) under the upper 43 limit assumption of unity quantum yield at medium  $(1x10^{13} \text{ and } 1.5x10^{15} \text{ photons cm}^{-2} \text{ s}^{-1} \text{ at } 185 \text{ and}$ 44 254 nm, respectively) or higher UV flux settings. The need for quantum yield measurements of both 45 VOC and SOA photolysis is highlighted in this study. The results of this study allow improved OFR 46 operation and experimental design, and also inform the design of future reactors.

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### 47 **1** Introduction

48 For decades, environmental chambers have been employed to study atmospheric chemical processes, 49 particularly, volatile organic compound (VOC) oxidation processes in the atmosphere (Cocker et al., 2001; 50 Carter et al., 2005; Presto et al., 2005; Wang et al., 2011; Platt et al., 2013), without the interference of 51 some transport processes (e.g., advection and wet deposition). These oxidation processes are the key 52 to secondary organic aerosol (SOA) formation (Odum et al., 1996; Hoffmann et al., 1997; Hallquist et al., 53 2009), and air pollutant removal (Levy II, 1971). Atmospheric simulation chambers usually have volumes 54 on the order of several m<sup>3</sup>, and use light sources longer than 300 nm (e.g., sunlight or UV blacklights) to 55 generate oxidants (mainly OH). These settings leads to OH concentrations (10<sup>6</sup>-10<sup>8</sup> molecules cm<sup>-3</sup>) that are not much higher than the typical ambient values ( $10^6-10^7$  molecules cm<sup>-3</sup>; Mao et al., 2009). 56 57 Relatively low OH concentrations require long residence/simulation times (generally hours) and limit 58 the ability of those setups to reach very high photochemical ages that are atmospherically-relevant 59 (George et al., 2007; Kang et al., 2007; Carlton et al., 2009; Seakins, 2010; Wang et al., 2011). Residence 60 times are ultimately limited due to losses of gases and particles to Teflon walls with timescales of tens 61 of minutes to several hours (Cocker et al., 2001; Matsunaga and Ziemann, 2010; Zhang et al., 2014) as 62 well as by the limited volume of the bag relative to the sampling instrumentation (Nguyen et al., 2014). 63 Besides, large sizes and support systems (e.g. clean air generators) make it difficult to use large 64 chambers in field or source studies.

Oxidation flow reactors (OFR) are an alternative that offers some advantages over environmental chambers, especially for rapid changes of experimental conditions, and/or for field experiments. They generally have a smaller size (on the order of 10 L), and typically use low-pressure Hg lamps as light sources for producing OH in large amounts via O<sub>3</sub> and/or H<sub>2</sub>O photolysis. These design choices lead to good portability, short experimental timescales, ability to reach long photochemical ages, and potentially reduced wall losses.

71 Due to these advantages, OFRs have been employed in many recent field and laboratory studies 72 in atmospheric chemistry, particularly in SOA-related research (George et al., 2007; Kang et al., 2007, 73 2011; Smith et al., 2009; Massoli et al., 2010; Cubison et al., 2011; Lambe et al., 2011a, 2011b, 2012, 74 2013; Bahreini et al., 2012; Saukko et al., 2012; Wang et al., 2012; Ortega et al., 2013; Li et al., 2013). 75 OFRs are also used in related applied fields, such as scrubbing of pollution from air (Johnson et al., 2014). 76 In contrast to their popularity, the chemistry occurring in OFRs is still incompletely characterized, 77 although the formation and interconversion reactions of most oxidants in OFRs have been well 78 characterized (Sander et al., 2011; Ammann et al., 2015). To our knowledge, there are only three studies 79 of OFR radical oxidation chemistry up to date: Ono et al. (2014) focused on the dependence of O<sub>3</sub> 80 destruction on  $H_2O$  concentration. We have recently made progress on the characterization of  $HO_x$ 81 radical chemistry in OFRs (Li et al., 2015; Peng et al., 2015). We have developed a kinetic model for OFRs, 82 which provides predictions in good agreement with laboratory experiments. This model has also shown 83 that OH exposure (OH<sub>exp</sub>, i.e., OH concentration integrated over the reactor residence time) increases 84 with H<sub>2</sub>O concentration and UV intensity, and decreases with external OH reactivity [OHR<sub>ext</sub>= $\sum k_i c_i$ , i.e., the sum of the products of concentrations of externally introduced OH-consuming species (*ci*) and rate constants of their reactions with OH (*ki*)]. The OH<sub>exp</sub> decrease due to OHR<sub>ext</sub> was defined as "OH suppression," and can reach two orders-of-magnitude in some cases (Li et al., 2015; Peng et al., 2015). We also showed that relative uncertainties of the outputs of our box model (e.g., OH<sub>exp</sub>) due to uncertain kinetic parameters are typically only 20% (Peng et al., 2015). However, none of these studies directly address the fate of VOCs (including oxygenated VOCs), simply regarded as external OH reactants in the prior studies.

92 The primary reason for the use of the OFRs studied here is for the study of reactions of species or 93 mixtures of atmospheric relevance with the OH radical. However, other highly reactive species are also 94 present at very elevated concentrations, including the radicals O(<sup>1</sup>D) and O(<sup>3</sup>P), 185 and 254 nm photons, 95 and O<sub>3</sub>. If a substantial fraction of the species of interest reacted with those non-OH reactants, then the 96 chemistry in the OFR would deviate from the OH radical chemistry intended to investigate. The absence 97 of systematic research on VOC fate in OFRs leaves room for some speculation that non-OH or non-98 tropospheric chemistry can play a major role in OFRs: for example, Johnson et al. (2014) suggested that 99 O(<sup>1</sup>D) and O(<sup>3</sup>P) significantly consumed VOCs. Klems et al. (2015) concluded that photons at 254nm 100 from Hg-lamp emission played an important role in their OFR experiment, especially for downstream 101 chemistry. Lack of clarity about these types of questions and of clear guidelines about how to apply 102 OFRs to avoid such problems have limited the application of OFRs for years. In this paper, we apply the 103 model in Peng et al. (2015) to systematically investigate whether significant non-tropospheric or non-104 OH chemistry occurs in OFRs, and what experimental conditions make it more important. Considering 105 the enormous complexity of organic radical (particularly organic peroxy) chemistry, we only examine 106 the non-OH fate of stable species in the present work. The fate of organic radicals should be the subject 107 of future studies. The results allow improved OFR operation and experimental design, as well as 108 guidance for the design of future reactors.

#### 109 2 Methods

The OFR and the model used here have been described in detail elsewhere (Kang et al., 2007; Li etal., 2015; Peng et al., 2015). Here, we only present a brief introduction for each.

# 112 2.1 Potential Aerosol Mass flow reactor

113 Kang et al. (2007) first introduced the Potential Aerosol Mass (PAM) flow reactor. Although there 114 were earlier versions of the PAM reactor, the version of cylindrical geometry with a volume of ~13 L has 115 been widely used and is currently in use in many SOA research groups (Massoli et al., 2010; Cubison et 116 al., 2011; Kang et al., 2011; Lambe et al., 2011a, 2011b, 2012, 2013; Bahreini et al., 2012; Saukko et al., 117 2012; Wang et al., 2012; Li et al., 2013; Ortega et al., 2013). The reactor is made of aluminum or of glass 118 and aluminum, and equipped with 1-4 low-pressure Hg lamps (model no. 82-9304-03, BHK Inc.) located 119 inside the flow tube. The Hg lamps produce UV emissions at 185 and 254 nm, whose intensity can be 120 rapidly computer-controlled. The operation mode using both 185 and 254 nm emissions is called 121 "OFR185". In this mode, photons at 185 nm dissociate H<sub>2</sub>O and O<sub>2</sub> molecules to produce OH+H and 122  $O(^{3}P)$ , respectively. Recombination of  $O(^{3}P)$  with  $O_{2}$  forms  $O_{3}$ . UV light at 254 nm then photolyzes  $O_{3}$  to 123 produce  $O(^{1}D)$ , which reacts with H<sub>2</sub>O and produces additional OH. OFR can also be operated in another 124 mode where photons at 185 nm are filtered by quartz sleeves around the lamps. In this case, only 254 125 nm UV light is active to generate OH ("OFR254" mode), and injection of externally generated O<sub>3</sub> into 126 the reactor is required for OH production. The amount of injected O<sub>3</sub> plays a critical role in the OFR 127 chemistry (Peng et al., 2015). For this reason this amount (X ppm) is also included in OFR operation 128 mode notation in the form of OFR254-X. For example, OFR254-70 and OFR254-7 denote experiments 129 with 70 and 7 ppm  $O_3$  injected, respectively. We use the PAM as the basic OFR design. Other designs 130 will be specified below if needed. Rapid computer-controlled UV lamp setting allows rapidly scanning 131 UV lamp settings during an experiment, and has unique applications to OFR experiments in field studies 132 (Hu et al., 2015; Ortega et al., 2015; Palm et al., 2015). In these experiments, OFRs enable the 133 exploration of a very large range of photochemical age during a short period (~2 hr) when ambient 134 conditions often do not significantly change.

135 2.2 Model description

136 We use the same model as in Peng et al. (2015), a standard chemical-kinetic model under plug-137 flow conditions. The effect of non-plug flow residence time distributions (RTD) was also investigated in 138 that study. Non-plug-flow models result in similar OH<sub>exp</sub> than plug-flow in most cases, except under 139 specific conditions with very high UV, H2O, and OHRext (Peng et al., 2015). Therefore, plug-flow OHexp is 140 used in this study, as a proxy of OHexp estimated from OHRext decay and to avoid the much increased 141 computational expense for complex RTDs. All O<sub>x</sub> and HO<sub>x</sub> reactions available in the JPL Chemical Kinetic 142 Data Evaluation (Sander et al., 2011) are taken into account. Reactions of some external OH reactants 143 (externally introduced reactants destructing OH), such as SO<sub>2</sub>, CO and NO<sub>x</sub>, are also included. SO<sub>2</sub> is used 144 as a proxy of other external OH reactants (e.g., VOCs). We believe that this is a realistic approximation 145 in terms of OHR<sub>ext</sub> decay vs. OH<sub>exp</sub> for many precursors, as discussed in Peng et al. (2015).

146 When studying the OFR, we assume a residence time of 180 s, and use typical temperature (295 147 K) and atmospheric pressure (835 mbar) in Boulder, CO, USA. H<sub>2</sub>O mixing ratio (abbr. H<sub>2</sub>O hereafter) 148 ranges from 0.07 to 2.3% (equivalent to relative humidity (RH) of 3–92%). According to Li et al. (2015), 149 UV photon fluxes (abbr. UV hereafter) at 185 and 254 nm are estimated to be in the ranges  $1.0 \times 10^{11}$ -1.0x10<sup>14</sup> and 4.2x10<sup>13</sup>-8.5x10<sup>15</sup> photons cm<sup>-2</sup> s<sup>-1</sup>, respectively. Four levels of OHR<sub>ext</sub>, 0, 10, 100, and 1000 150 151 s<sup>-1</sup> covering the range of most field and laboratory studies are investigated. In the explored parameter 152 space, the same 3-character labels as in Peng et al. (2015) are used to denote typical cases (Table 1). For 153 OFR254, we study OFR254-70 and OFR254-7, representing OFR254 experiments with high (Palm et al., 154 2015) and low O<sub>3</sub> (Kang et al., 2011; Liu et al., 2015), respectively. To model literature OFR studies, we 155 adopt corresponding parameters (reactor volume, H<sub>2</sub>O, residence time, etc.), and estimate parameters 156 that are specified or measured (e.g., UV) as needed. In particular, for some field studies, where long 157 time-series of experimental data [42 d in BEACHON-RoMBAS (Palm et al., 2015), 42 d in SOAS (Hu et al., 158 2015), and 15 d in CalNex-LA (Ortega et al., 2015)] were recorded, we model all valid datapoints and 159 present outputs in the form of histograms. Note that the outputs for field studies, i.e., histograms, have a complex dependence on ambient temperature, H<sub>2</sub>O, and OHR<sub>ext</sub>, as well as UV steps used. The specific
 histogram shapes for different field campaigns are influenced by both ambient and experimental
 parameters.

163 3 Results and Discussions

In the following sections we explore the relative importance of five non-OH pathways (photons at 185 and 254 nm, O(<sup>1</sup>D), O(<sup>3</sup>P), and O<sub>3</sub>) vs. the OH reaction for species of atmospheric interest, including a variety of typical biogenic and anthropogenic VOCs and a few important inorganic species (e.g., SO<sub>2</sub> and NO<sub>2</sub>). Because of the huge complexity of VOC oxidation mechanisms, only consumption/oxidation of specific VOCs is investigated. In such an investigation, a large amount of kinetic data is required. We collected the required data (Tables S1–S3) according to the principles in Section S1. Photolysis of SOA is also investigated.

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### 3.1 Fractional loss of VOCs to non-OH reactants

172 As shown in Peng et al. (2015), OHexp in OFRs depends on various physical conditions, e.g., H<sub>2</sub>O and 173 OHRext. However, the non-OH reactants are much less dependent on these parameters. H<sub>2</sub>O and 174 external OH reactants only contribute less than 1% to absorption at 185 and 254 nm. Therefore, they 175 have almost no impact on effective UV. O<sub>3</sub> can absorb a fraction of the 254 nm radiation, but the optical 176 depth due to 70 ppm O<sub>3</sub> in the reactor is ~0.11, and thus the attenuation of 254 nm photons by O<sub>3</sub> 177 absorption is a minor effect. The dominant fates of  $O({}^{1}D)$  and  $O({}^{3}P)$  are the quenching by air and the 178 recombination with  $O_2$ , respectively, with which no reactions involving  $H_2O$  or external OH reactants can 179 compete. Thus the concentrations of  $O(^{1}D)$  and  $O(^{3}P)$  in the reactor depend on UV intensity and H<sub>2</sub>O, 180 but not on OHRext. Since OH can be strongly modulated by OHRext, changing input conditions may result 181 in very different relative importance of OH to other reactive species. To fully evaluate this issue, it is 182 necessary to explore a very wide range of conditions as in Peng et al. (2015).

183 **3.1.1** Common features of all the non-OH reactants

184 Figs. 1–5 show the relative consumption of several species vs. ratio of exposures to individual non-185 OH species (X) to OH exposure (Xexp/OHexp) for 185 and 254 nm photons, O(<sup>1</sup>D), O(<sup>3</sup>P), and O<sub>3</sub>, 186 respectively. Fig. S1-S5 present the same information in an alternative format that may be useful to 187 evaluate the fate of species not included in our study. They show ratios of rate constants of species with 188 OH to those with non-OH species (X), mathematically equivalent to the X<sub>exp</sub>/OH<sub>exp</sub> where corresponding 189 X/(OH+X) is 50%. X/(OH+X) denotes the fractional importance of X in the sum of the contribution of OH 190 and X to VOC. If a VOC has a higher ratio of rate constant with OH to that with a non-OH reactant (X) 191 than with another reactant (Y), the relative contribution of X, X/(OH+X), should be smaller than that of 192 Y. In these figures, we also show the X<sub>exp</sub>/OH<sub>exp</sub> range for OFR254-70, OFR254-7, and OFR185, under 193 different conditions, including key laboratory and field studies, and identify the X<sub>exp</sub>/OH<sub>exp</sub> ranges where 194 non-OH contribution to species fate is significant. A dissociation quantum yield of unity is assumed for 195 the photolysis reactions, which results in upper limits for the relative importance of those pathways. 196 In these figures, the relationships of all non-OH reactive species to OH are similar for certain

197 common conditions. We define three types of conditions to help guide experimental design and

198 evaluation in terms of the relative importance of non-OH reactants. Under "riskier conditions" of 199 high/very high  $OHR_{ext}$  ( $\geq 100 \text{ s}^{-1}$  in OFR185 and  $> 200 \text{ s}^{-1}$  in OFR254 (-7 to -70)) and/or low  $H_2O$  (<0.1%), 200 non-OH reactions can be significant depending on the species. Conversely, under "safer conditions" with 201 relatively low OHR<sub>ext</sub> (<30 s<sup>-1</sup> in OFR185 and <50 s<sup>-1</sup> in OFR254), and high H<sub>2</sub>O (>0.8% in OFR185 and >0.5% in OFR254), and moderate or higher UV (>1x10<sup>12</sup> photons cm<sup>-2</sup> s<sup>-1</sup> at 185 nm) in OFR185, reaction with 202 203 OH is dominant (Figs. 1–5 and S1–5). We denote all other conditions as "transition conditions." High 204 H<sub>2</sub>O and zero/low OHR<sub>ext</sub> lead to strong OH production and no/weak OH suppression, respectively. Thus, 205 OH is more abundant and dominates species consumption under those conditions. In the case of low 206 H<sub>2</sub>O and high OHR<sub>ext</sub>, OH is generally lower because of less production and more suppression. These 207 conditions increase the relative contribution of non-OH species. UV light intensity is generally less 208 influential on non-OH VOC fate than H<sub>2</sub>O and OHR<sub>ext</sub>, although OH production is nearly proportional to 209 UV (Peng et al., 2015), because the non-OH reactive species also scale (nearly) proportional to UV. As a 210 result, UV generally has smaller effects on exposure ratios between OH and the non-OH reactants. 211 However, under a UV near the lower bound of the explored range in this study (<1x10<sup>12</sup> photons cm<sup>-2</sup> s<sup>-</sup> 212 <sup>1</sup> at 185 nm) in OFR185, OH production is so small that the effect of  $OHR_{ext}$  on OH suppression can be 213 amplified and hence some exposure ratios may be affected. In OFR254 OH is more resilient to 214 suppression even at low UV because of the OH-recycling by initially injected O<sub>3</sub> (Peng et al., 2015). Note 215 that we call these conditions "riskier" and "safer" mainly in terms of non-tropospheric VOC fate, but 216 not of VOC fate by all non-OH reactants, as some of the non-OH reactant studied in this work may also 217 play a role under some tropospheric conditions (see Sections 3.1.4 and 3.1.5). In addition to the 218 common features above, individual non-OH reactants have their own features as well as a few 219 exceptions to the above mentioned general observations, which we will detail below.

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#### 3.1.2 Reactions with OH vs. photolysis at 185 nm

Under riskier reactor conditions, photolysis at 185 nm of several aromatic compounds, such as toluene, benzene, and p-xylene, are estimated to be significant and even dominant vs. the OH reactions (Fig. 1). This results from their aromatic ring, which is not only highly efficient as a chromophore, but also relatively resistant to OH attack.

225 It is not common to perform field studies for SOA at H<sub>2</sub>O as low as 0.1% or OHR<sub>ext</sub>≥100 s<sup>-1</sup> (Table 226 1). According to F185<sub>exp</sub>/OH<sub>exp</sub> calculated from the field studies where OFR185 was deployed, i.e., 227 BEACHON-RoMBAS (Palm et al., 2015), SOAS (Hu et al., 2015), and CalNex-LA (Ortega et al., 2015) all 228 these studies are generally under safer conditions (infrequent low H<sub>2</sub>O mixing ratio and ambient OH 229 reactivity estimated to be ~15-25 s<sup>-1</sup>). For instance, none of these field studies fell into the conditions 230 where the fractional importance of photolysis at 185 nm was significant for aromatic species. However, 231 in some source studies using OFRs, e.g. when sampling biomass burning smoke (FLAME-3; Ortega et al. 232 2013) or air in a traffic tunnel (Tkacik et al., 2014), OHRext can be very high reaching values of several 233 100 s<sup>-1</sup> (Table S4). Especially on the smoke study, photolysis of aromatics may have played a role. 234 However, it has long been known that excited aromatic molecules may undergo various deactivation 235 pathways (e.g., vibronic coupling, intersystem crossing, and collisional quenching) without molecular fragmentation (Beddard et al., 1974; Nakashima, 1982; Nakashima and Yoshihara, 1983; Fang and
Phillips, 2002), preventing unity quantum yields. Therefore, the photolysis of aromatics at 185 nm in the
above mentioned source studies may not be as significant as estimated in Fig. 1.

239 Under riskier conditions, some organic peroxy nitrates and nitrates (e.g., peroxyacetyl nitrate and 240 2-propyl nitrate in Fig. 1) have an estimated contribution from photolysis at 185 nm to their fate that is 241 comparable to or even larger than that of reaction with OH. Nevertheless, this does not mean that we 242 need to make extra efforts to avoid the photolysis of organic nitrates and peroxy nitrates at 185 nm. 243 Although they have cross-sections at 185 nm ~10-100 times smaller than those of aromatics, these 244 organic compounds react remarkably slowly with OH (~2 order of magnitude slower than reactions of 245 aromatics with OH), so photolysis appears substantial with respect to reaction with OH, but really two 246 slow rates are being compared. Thus, absolute photolyzed amounts of these species are not substantial. 247 For example, only ~10% of peroxyacyl nitrate is photolyzed by 185 nm photons at the highest OFR185 248 lamp setting. Even if photolysis of nitrates and peroxy nitrates by low-pressure Hg lamp emission 249 proceeds to a significant extent, it may still not be a problem, as it generally leads to the same products 250 as their ambient photolysis (Renlund and Trott, 1984; Roberts and Fajer, 1989). Nitrate and peroxy 251 nitrate photolysis is actually much more important in the atmosphere than in OFRs for the same 252 photochemical age (see below and Fig. 6).

SO<sub>2</sub> has been used in some studies to calibrate OH<sub>exp</sub> (Lambe et al., 2015; Li et al., 2015). It does undergo significant photolysis at 185 nm under riskier conditions. However, this photolysis does not lead to an overestimation of OH<sub>exp</sub>, since the S-bearing product of SO<sub>2</sub> photolysis at 185 nm, SO, converts back to SO<sub>2</sub> very rapidly through its reaction with O<sub>2</sub>.

Oxidation intermediates may also photolyze at 185 nm. However, their photolysis is unlikely to be
 significant when OFR is operated under safer conditions. To clarify this issue, a detailed discussion about
 the photolysis of oxidation intermediates at 254 nm is required as a premise. We thus discuss oxidation
 intermediate photolysis at both 185 and 254 nm in Section 3.1.3.

261 3.1.3 Reactions with OH vs. photolysis at 254 nm

262 The photon flux in the reactor at 254 nm is 80–250 times larger than at 185 nm (Li et al., 2015). 263 Although absorption cross sections of all molecules investigated in this study are significantly lower at 264 254 than 185 nm, the higher photon flux compensates, at least partially, for this effect, so that in OFR185 265 photolysis of many species at 254 nm is of similar relative importance as photolysis at 185 nm, with 266 potentially important effects at low H<sub>2</sub>O and/or high OHRext (Figs. 2 and S2). As for 185 nm, 254 nm 267 photolysis is a concern mainly for aromatic compounds, because of the high light absorptivity and low 268 OH reactivity of aromatic rings as previously discussed. Again, note that this concern may be less serious 269 than shown in Fig. 2 because of possible lower quantum yields. Photolysis of organic nitrates and peroxy 270 nitrates at 254 nm also appears to be important relative to reactions with OH, and is still not a concern 271 for the same reasons as photolysis at 185 nm. SO<sub>2</sub> can absorb efficiently at 254 nm, but it is still not a 272 problem for SO<sub>2</sub>-based OH<sub>exp</sub> calibration, since photons at 254 nm are not sufficiently energetic to 273 dissociate SO<sub>2</sub> molecules.

High UV generally appears to be more problematic than low UV in OFR254 (Fig. S2). This is in contrast to the trend of OFR185. In OFR254, O<sub>3</sub> is the only primary OH source, and a substantial fraction of O<sub>3</sub> can be photolyzed at the highest lamp settings, leading to a substantial reduction of OH production (compared with proportional scaling with UV). OFR254-70 appears to be less prone to riskier conditions than OFR254-7, since higher O<sub>3</sub> favors HO<sub>2</sub>-to-OH recycling, making OH more resilient to suppression (Peng et al., 2015).

280 Under highly risky conditions (H<sub>2</sub>O<0.1% and OHR<sub>ext</sub>≥100 s<sup>-1</sup> for OFR185, and H<sub>2</sub>O<0.1% and 281 OHRext>200 s<sup>-1</sup> for OFR254), some saturated carbonyl compounds (e.g., pyruvic acid, methyl ethyl ketone 282 and hydroxyacetone) have significant photolysis at 254 nm relative to reactions with OH. This significant 283 relative contribution of photolysis also results from remarkably slow reactions of saturated carbonyl 284 compounds with OH. Although secondary species without C=C double bond, e.g., saturated carbonyls, 285 hydroperoxides, and nitrates, can be photolyzed in OFRs at low H<sub>2</sub>O and/or high OHR<sub>ext</sub>, their photolysis 286 only proceeds to a ~10-1000 times smaller extent than ambient photolysis at the same photochemical 287 age (Fig. 6).

288 Unsaturated carbonyls may have much higher absorption cross-sections if their C=C bonds are 289 conjugated with carbonyls. However, according to our following analysis, conjugated unsaturated 290 carbonyl compounds do not often cause problems of non-tropospheric photolysis at 254 nm. Carbonyls 291 have  $\pi - \pi^*$  and  $n - \pi^*$  transitions. The former corresponds to high cross-section (typically >10<sup>-18</sup> cm<sup>2</sup>) and 292 typically occurs around or below 200 nm. The latter is forbidden, and thus has weak absorption (crosssection on the order of 10<sup>-19</sup> cm<sup>2</sup> or lower), and typically occurs around or above 300 nm (Turro et al., 293 294 2009). Conjugation usually does not substantially enhance the absorption of  $n-\pi^*$  transition but it does 295 for  $\pi$ - $\pi$ \* transitions (Turro et al., 2009). As a result, through conjugation, the only reason why cross-296 sections of carbonyls at 254 nm may be elevated above 10<sup>-18</sup> cm<sup>2</sup> is the red-shift of the maximum 297 absorption wavelength of their  $\pi$ - $\pi$ \* transitions due to conjugation. According to Woodward's rules 298 (Pretsch et al., 2009) and available cross-section data of  $\alpha$ , $\beta$ -unsaturated carbonyls in Keller-Rudek et al. 299 (2015), a conjugation of at least 3–4 double bonds is required for the excitation at 254 nm to dominantly 300 correspond to  $\pi - \pi^*$  transition. Conjugated oxidation intermediates containing at least 3–4 double 301 bonds including C=C bond(s) are virtually impossible to form from aliphatic hydrocarbon oxidation in 302 OFRs. Nevertheless, such intermediates may form via ring-opening pathways of aromatic oxidation (Calvert et al., 2002; Atkinson and Arey, 2003; Strollo and Ziemann, 2013). E,E-2,4-hexadienedial may 303 304 be regarded as an example of this type of intermediates. Even under assumption of a unity quantum 305 yield, its fraction of photolysis at 254 nm is not much higher than that of aromatic precursors (Fig. 2). 306 Therefore, 254 nm photolysis of conjugated intermediates should not be problematic as long as safer 307 experimental conditions are adopted.

308 To our knowledge, the only exception that has strong absorption at 254 nm due to conjugation 309 with <2 double bonds are  $\beta$ -diketones, which may be formed in aliphatic hydrocarbon oxidation, 310 particularly that of long-chain alkanes (Ziemann and Atkinson, 2012). The peculiarity of  $\beta$ -diketones is 311 that their enol form may have a highly conjugated ring structure due to very strong resonance (Scheme

S1), and hence cross-sections on the order of  $10^{-17}$  cm<sup>2</sup> at 254 nm (Messaadia et al., 2015). However, 312 even under the assumption of unity quantum yield, the fractional contribution of 254 nm photolysis of 313 314 acetylacetone (representative of  $\beta$ -diketones) is only slightly larger than for aromatic VOCs (Fig. 2), since 315 its enol form also contains a C=C bond leading to very high reactivity toward OH. Furthermore, we argue 316 that the actual probability that a concrete structural change (in number and type of functional groups, 317 O/C ratio, average C oxidation state etc.) of  $\beta$ -diketones resulting from photoexcitation at 254 nm may 318 be low. As their excitation at 254 nm corresponds to  $\pi$ - $\pi$ \* transition, their rigid ring structure likely 319 hinders cyclic structural change at the 1<sup>st</sup> singlet excited state ( $S_1(\pi,\pi^*)$ ) while the biradical structure of 320 the 1<sup>st</sup> triplet state ( $T_1(\pi,\pi^*)$ ) may favor H-shift between two O atoms, which ends up with the 321 same/similar structure than prior to the H-shift (Scheme S1). Also, the excitation of  $\beta$ -diketones at 254 322 nm may also lead to charge transfer complex formation via direct excitation and/or radiationless 323 transition from a local excited state (Phillips and Smith, 2015), which is likely to result in low quantum 324 yields, as discussed in detail below.

325 In addition to conjugated species, Phillips and Smith (2014, 2015) reported a new type of highly 326 absorbing species that may be formed from VOC oxidation. Although their studies were conducted in 327 the condensed phase, it is likely that the main conclusions of these studies are generally transferable to 328 the gas-phase conditions, since no long-range interactions, which do not exist in normal gases, were 329 involved in these studies. Phillips and Smith (2014, 2015) investigated the photoabsorption enhancement of multifunctional oxygenated species in SOA and found that the high absorptivity of 330 331 these species can largely be explained by transitions toward the electronic states of charge transfer 332 complex formed between hydroxyl groups (donor) and neighboring carbonyl groups (acceptor). They 333 also pointed out that charge transfer complexes of this kind have a continuum of states whose energy 334 levels range from that of local excited states (radiative transition wavelength <300 nm) to very low levels 335 (radiative transition wavelength >600 nm). The latter are insufficient to cause common photochemical 336 reactions. Relaxation through a continuum of states is usually ultrafast according to Fermi's golden rule 337 (Turro et al., 2009), likely leading to low quantum yields of chemical reactions. The low quantum yields 338 may be seen even from species with only one hydroxyl and one carbonyl: the photolysis of 3-hydroxy-339 3-methyl-2-butanone and 4-hydroxy-2-butanone at wavelengths down to 270 nm has quantum yields 340 around only 0.1 (Bouzidi et al., 2014, 2015). Although measurements of photolysis quantum yield for 341 multifunctional species are challenging and rare, it is reasonable to expect even lower quantum yields 342 for larger and/or highly substituted (by hydroxyl and carbonyl) species, since larger species have more 343 degrees of freedom for relaxation of excited molecules, and more and/or larger complex sites generally 344 lead to more efficient relaxation through a continuum of states, in accordance with common 345 photophysical sense (Sharpless and Blough, 2014). Therefore, even though species with a number of 346 hydroxyls and carbonyls are formed in VOC oxidation and can absorb >1 order of magnitude more 347 efficiently at 254 nm than mono- and difunctional species, they may still have low effective photolysis 348 rates because of low quantum yields.

349

For this type of species, we estimate an upper limit of the fractional importance of their photolysis

350 at 254 nm. Molar absorption coefficients of charge transfer transitions of organic molecules are usually 351 ~10<sup>3</sup>–1x10<sup>4</sup> L mol<sup>-1</sup> cm<sup>-1</sup>, i.e., cross-sections of ~3.9x10<sup>-18</sup>–3.9x10<sup>-17</sup> cm<sup>2</sup> (Foster, 1969). Based on that, it 352 is reasonable to estimate an upper limit of absorption cross-sections of charge transfer transitions of 353 5x10<sup>-17</sup> cm<sup>2</sup>. On the other hand, photolysis quantum yields of multifunctional species are unlikely to be 354 larger than that of species with only one carbonyl and one hydroxyl, i.e., ~0.1 (see discussion above). We thus take 0.1 as an upper limit of photolysis quantum yields. Besides, 6x10<sup>-12</sup> cm<sup>3</sup> molecule<sup>-1</sup> s<sup>-1</sup> can 355 356 be a conservative estimate of rate constants of multifunctional oxygenated species with OH, as it is 357 roughly an average value for ketones (Atkinson and Arey, 2003), and the enhancement of H-abstraction 358 by hydroxyl groups (Kwok and Atkinson, 1995; Ziemann and Atkinson, 2012) and the fast abstraction of 359 aldehydic H atoms (Atkinson and Arey, 2003) are completely neglected. With the three estimates 360 combined, the estimated maximum fractional contribution from photolysis at 254 nm to the fate of 361 multifunctional species (Fig. 2) is close to that of E,E-2,4-hexadienedial and acetylacetone.

The problem of photolysis of oxidation intermediates at 185 nm is unlikely to be worse than at 254 nm. According to available UV spectra of carbonyl compounds in Keller-Rudek et al. (2015), 185 nm is almost always located within the  $\pi$ - $\pi$ \* transition band, whose maximum cross-section is on the order of 10<sup>-17</sup> cm<sup>2</sup>. Even if all types of radiative transitions at normal radiation intensity are considered, an approximate upper limit of absorption cross-sections is ~10<sup>-16</sup> cm<sup>2</sup> (Evans et al., 2013). However, UV intensity at 185 nm in the OFR185 mode is ~100 times lower than that at 254 nm (Li et al., 2015). The photolysis rate of oxidation intermediates at 185 nm should thus be generally smaller than at 254 nm.

369 Therefore, in summary, photolysis of oxidation intermediates are, to our knowledge, 370 conservatively estimated to be of limited importance relative to their reactions with OH, as long as the 371 experimental conditions are in the safer range. Although studies on photolysis quantum yields of 372 oxidation intermediates are very sparse, we reason, based on the existing studies on this topic and 373 common photophysical and photochemical rules, that the photolysis quantum yields of these species 374 may be lower than the values assumed in this study (e.g., 1 for E,E-2,4-hexadienedial and acetylacetone 375 and 0.1 for multifunctional species). As a result, actual rates and relative importance of photolysis might 376 be significantly smaller than the upper limits estimated in our study.

As discussed for photolysis at 185 nm, in all ambient OFR field studies (BEACHON-RoMBAS, SOAS, and CalNex-LA), reactions with OH dominate over photolysis at 254 nm (Fig. 2). The fractional consumption of several anthropogenic aromatic VOCs, such as benzene and naphthalene, in the urban CalNex-LA campaign by 254 nm photolysis is estimated as a few percent under most conditions and at most ~15%. At the BEACHON-RoMBAS and SOAS forested sites, photolysis at 254 nm should be a negligible contributor to the fate of biogenic VOCs such as isoprene and monoterpenes.

Some laboratory and source studies may have had an appreciable contribution to aromatic species fate from 254 nm at low H<sub>2</sub>O and/or high OHR<sub>ext</sub>. F254<sub>exp</sub>/OH<sub>exp</sub> in the biomass smoke and urban tunnel source studies (source OHR<sub>ext</sub> up to ~300 s<sup>-1</sup>; Ortega et al., 2013; Tkacik et al., 2014) and the Kang et al. (2011) laboratory study (H<sub>2</sub>O down to ~0.1%) can be as high as  $10^6$ – $10^7$  cm/s. In this range, photolysis of a few aromatic VOCs (e.g., benzene and naphthalene) at 254 nm could account for ~20–80% of their 388 destruction.

389 Note that photolysis of oxidation intermediates also needs to be taken into account. If 390 multifunctional species,  $\beta$ -diketones, and extensively conjugated species are photolyzed as shown in Fig. 391 2, these photolyses would be significant in some previous source and laboratory studies examined here, 392 as they were conducted at relatively low H<sub>2</sub>O and/or high OHR<sub>ext</sub>. To our knowledge, none of these 393 studies reported a significant photolysis of oxidation intermediates. Klems et al. (2015) attributed large 394 amounts of fragmentation products detected in their OFR experiments with dodecanoic acid to 395 photolysis of peroxy radicals. However, these products may also be at least partially accounted for by 396 photolysis of carbonyls leading to carbon-chain cleavage via Norrish reactions (Laue and Plagens, 2005). 397 The OFR used by Klems et al. (2015) has a different design from the PAM, which is regarded as the base 398 design in this study. Their reactor employs a light source stronger than the PAM's highest lamp setting, with UV at 254 nm estimated to be  $\sim 3x10^{16}$  photons cm<sup>-2</sup> s<sup>-1</sup> ( $\sim 4$  times the value at the highest lamp 399 400 setting of the PAM OFR) based on the lamp power and the reactor geometry. Such high UV may even 401 result in significant photolysis of saturated carbonyl intermediates, which are very likely formed in the 402 oxidation of long-chain alkane-like dodecanoic acid.

403 **3.1.4** Reactions with OH vs. reactions with O(<sup>1</sup>D) and O(<sup>3</sup>P)

404 The results for these two radicals are shown in Figs. 3-4. The potential impact of O(<sup>1</sup>D) is smaller 405 than for 185 and 254 nm photons, due to the low concentration of O(<sup>1</sup>D) in the reactor. Only for methane 406 may reaction with  $O(^{1}D)$  be significant, because the reaction of methane with  $O(^{1}D)$  is close to the 407 collision rate, while CH4 is the most resistant VOC to H-abstraction by OH. This could be important if CH4 408 was used for OH<sub>exp</sub> calibration under riskier reactor conditions, or if the fate of CH<sub>4</sub> is important to the 409 experiment for other reasons. Other VOCs react more slowly with  $O(^{1}D)$  and much faster with OH. As a 410 result, reactions of VOCs (other than CH<sub>4</sub>) with O(<sup>1</sup>D) in all laboratory, field, and source studies 411 previously discussed are almost always negligible. We also note that the ratio of  $O(^{1}D)_{exp}/OH_{exp}$  in the 412 OFR is actually much lower than in the troposphere (Monks, 2005), except under some riskier conditions. 413 It is believed that the contribution of  $O(^{1}D)$  to VOC destruction in the atmosphere should be negligible 414 (Calvert et al., 2002), and their relative importance is even lower under most OFR conditions.

415 Reactions with O(<sup>3</sup>P) are small or negligible contributors to VOC consumption except under 416 extreme riskier conditions. Unless at low  $H_2O$  (<0.1%) and very high external OH reactivity (~1000 s<sup>-1</sup>), 417 VOC consumption by O(<sup>3</sup>P) cannot be larger than 10% of that by OH (Figs. 4 and S4). This results from 418 both very low concentrations of O(<sup>3</sup>P) and relatively low reactivity compared to that of OH. Among the 419 species that we examine, biogenic VOC consumption may have some contribution from O(<sup>3</sup>P) under the 420 abovementioned riskier conditions, due to the higher reactivity of double bonds in these species with 421  $O(^{3}P)$ . For example  $\alpha$ -pinene in the mixture experiments in Kang et al. (2011) may have had a ~5% 422 contribution from  $O(^{3}P)$ . Similarly to  $O(^{1}D)$ ,  $O(^{3}P)_{exp}/OH_{exp}$  in the troposphere (Calvert et al., 2002) is 423 higher than in the OFR except for riskier conditions. Thus the relative importance of both  $O(^{1}D)$  and  $O(^{3}P)$ 424 to OFR chemistry is typically lower than in the troposphere.

425 3.1.5 Reactions with OH vs. reactions with O<sub>3</sub>

Reaction with O<sub>3</sub> is a major, even dominant pathway of the consumption of many biogenic VOCs in the troposphere. However, it is of interest to quantify the relative importance of OH vs. O<sub>3</sub> across OFR experiments (Figs. 5 and S5). This allows comparison with the relative importance in the troposphere, as well as potentially designing experiments where the relative influence of O<sub>3</sub> is minimized or adjusted as desired.

A large amount of  $O_3$  is injected into OFR254, and  $O_3$  concentration in that type of reactor does not change much with UV flux (negligibly under most conditions and up to a factor of ~2 at high H<sub>2</sub>O and UV, Peng et al. 2015). Since  $OH_{exp}$  is proportional to UV flux, as UV decreases,  $OH_{exp}$  is lowered and the fractional species destruction by  $O_3$  increases. In OFR185,  $O_3$  production is almost linearly dependent on UV, while a significant portion of OH production has a quadratic relationship with UV. Thus OH increases faster with increasing UV than  $O_3$ . Therefore, lower UV in OFR185 also leads to a higher relative importance of  $O_3$  for VOC consumption.

438 The distribution of O<sub>3exp</sub>/OH<sub>exp</sub> expected for the troposphere was obtained from the GISS ModelE2 439 climate model (Schmidt et al., 2014) and is estimated as the ratio of the simulated daily mean 440 concentration of O<sub>3</sub> and OH on a horizontal grid of 2 degrees in latitude and 2.5 degrees in longitude for 441 the year 2000. Interestingly, the simulated relative importance of  $O_3$  to OH in the troposphere is higher 442 than when OFRs are operated under safer conditions, and similar to OFRs when they are operated under 443 riskier conditions (Figs. 5 and S5). In those cases, a number of biogenic VOCs can be significantly 444 consumed by O<sub>3</sub>. In particular, some monoterpenes (e.g.,  $\alpha$ -terpinene) and sesquiterpenes (e.g.,  $\beta$ -445 caryophyllene) have a fractional reaction with  $O_3$  close to 100% in the troposphere. In contrast to 446 biogenics, reactions with  $O_3$  do not play any role in the consumption of most anthropogenic VOCs, e.g., 447 benzene, toluene, and alkanes. Besides, ozonolysis of saturated oxidation intermediates (e.g. carbonyls 448 and alcohols) is minor or negligible in both OFRs and the atmosphere, since they react with OH at  $^{-10^{-1}}$ <sup>13</sup>-10<sup>-11</sup> cm<sup>3</sup> molecule<sup>-1</sup> s<sup>-1</sup> while their ozonolysis rate constants are <10<sup>-20</sup> cm<sup>3</sup> molecule<sup>-1</sup> s<sup>-1</sup> (Atkinson 449 450 and Arey, 2003). However, unsaturated oxidation intermediates may have larger contributions from O<sub>3</sub> 451 because of C=C bonds. In particular, dihydrofurans, possible intermediates of saturated hydrocarbon 452 oxidation (Ziemann and Atkinson, 2012; Aimanant and Ziemann, 2013), may be predominantly oxidized 453 by O<sub>3</sub> in the troposphere. In OFR254, they can still have significant contributions from O<sub>3</sub> even outside 454 the low-H<sub>2</sub>O and/or high-OHR<sub>ext</sub> conditions.

455 An experimentalist may be interested in obtaining an  $O_{3exp}/OH_{exp}$  in an OFR close to ambient values, 456 which requires lower H<sub>2</sub>O and higher OHR<sub>ext</sub> conditions, although care should be taken to avoid other 457 non-tropospheric reactions under those conditions. On the other hand, one may want to study OH-458 dominated chemistry and thus want to avoid significant ozonolysis of VOCs to reduce the complexity of 459 VOC fate. This is analogous to the addition of excess NO to suppress O<sub>3</sub> in some chamber experiments. 460 In this case the OFRs should be operated under opposite conditions, i.e., high H<sub>2</sub>O, high UV, and low 461 OHR<sub>ext</sub>. This strategy enhancing OH<sub>exp</sub> is effective for most VOCs, except those with the highest  $k_{O3}/k_{OH}$ 462 ratios, e.g.,  $\alpha$ -terpinene and  $\beta$ -caryophyllene. To further decrease the importance of reactions of VOC 463 with O<sub>3</sub>, it is necessary to lower the O<sub>3</sub> concentration. For OFR254, one can inject less O<sub>3</sub> into the reactor

and increase the UV lamp setting. The comparison between OFR254-70 and OFR254-7 in Fig. S5
demonstrates this approach. For OFR185, we propose another strategy, i.e., lowering O<sub>2</sub> concentration
in the reactor. This decreases O<sub>3</sub> production but affects OH production to a much lesser extent, thanks
to the major OH production by H<sub>2</sub>O photolysis. We simulate the OFR185 cases with 2‰ O<sub>2</sub> and observe
that VOC ozonolysis can be excluded at high H<sub>2</sub>O and high UV (Fig. S5).

469 Among the literature OFR studies, the field studies employing OFRs in urban and forested areas all 470 operated under O<sub>3exp</sub>/OH<sub>exp</sub> values 100 times lower than in the atmosphere. In these field studies 471 reaction of almost all VOCs with O<sub>3</sub> can be neglected, except for the most reactive biogenics with O<sub>3</sub>, 472 e.g.,  $\alpha$ -terpinene and  $\beta$ -caryophyllene. The source study in an urban tunnel of Tkacik et al. (2014) 473 operated under similar conditions. Some laboratory studies using OFR254 (Kang et al., 2011; Lambe et 474 al., 2011b) as well as the biomass smoke source study (Ortega et al., 2013) operated at O<sub>3exp</sub>/OH<sub>exp</sub> close 475 to tropospheric values, because the injected O<sub>3</sub> plays a key role for OFR254 studies and the biomass 476 smoke experiments were conducted at high  $OHR_{ext}$ . Nevertheless, only  $\alpha$ -pinene and  $\beta$ -pinene, both 477 biogenics, are significantly consumed by O<sub>3</sub>. Another OFR254 study, Klems et al. (2015), had O<sub>3exp</sub>/OH<sub>exp</sub> 478 significant lower than tropospheric values, since the initial O<sub>3</sub> in their experiment was only 2 ppm, and 479 the UV light in their experiment was stronger than our lamps' highest setting and further reduced 480 effective O<sub>3</sub>.

481 3.1.6 Reactions with <sup>1</sup>O<sub>2</sub> and HO<sub>2</sub>

482 Singlet oxygen, <sup>1</sup>O<sub>2</sub>, can be produced in various ways in OFRs (Calvert et al., 2002; Ono et al., 2014) 483 and react with alkenes at appreciable rate constants (~10<sup>-17</sup>-10<sup>-14</sup> cm<sup>3</sup> molecule<sup>-1</sup> s<sup>-1</sup>; Huie and Herron, 484 1973; Eisenberg et al., 1986). We estimate  ${}^{1}O_{2}$  concentration by the expression proposed by Ono et al. 485 (2014). Only at the lowest H<sub>2</sub>O, the highest lamp setting, and the highest OHR<sub>ext</sub> in this study (Table 1), 486 may the most reactive alkene (endo-cyclic conjugated dienes, e.g., cyclopentadiene,  $\alpha$ -terpinene, and 487  $\alpha$ -phellandrene) have >10% contribution from <sup>1</sup>O<sub>2</sub> to their fate. For all other species and under all other 488 conditions, reactions of VOCs with  ${}^{1}O_{2}$  are negligible. Thus, this reactant is not discussed further in the 489 present work.

490 HO<sub>2</sub> is a major radical in the OFR chemistry. However, it is much less reactive than OH toward VOCs. 491 Typically, the rate constants of reactions of HO<sub>2</sub> with alkenes are smaller than 10<sup>-20</sup> cm<sup>3</sup> molecule<sup>-1</sup> s<sup>-1</sup> at 492 room temperature, and those with almost all saturated VOCs (except aldehydes and ketones) are even 493 smaller (Tsang, 1991; Baulch et al., 1992, 2005). Therefore, we briefly discuss reactions of HO<sub>2</sub> with 494 aldehydes and ketones, and neglect those with all other VOCs in this study. Ketones react with  $HO_2$  at rate constants on the order of 10<sup>-16</sup> cm<sup>3</sup> molecule<sup>-1</sup> s<sup>-1</sup> or lower (Gierczak and Ravishankara, 2000; Cours 495 496 et al., 2007). Therefore, only at low H<sub>2</sub>O, low UV and high OHR<sub>ext</sub>, the reaction of acetone with HO<sub>2</sub> may 497 compete with that with OH. The same is likely true for the reactions of acetaldehyde and larger 498 aldehydes with HO<sub>2</sub>, as their rate constants are likely to be around or less than 1x10<sup>-14</sup> cm<sup>3</sup> molecule<sup>-1</sup> s<sup>-1</sup> 499 <sup>1</sup> (da Silva and Bozzelli, 2009). Formaldehyde is the only stable carbonyl compound that may react with 500  $HO_2$  (rate constant: 7.9x10<sup>-14</sup> cm<sup>3</sup> molecule<sup>-1</sup> s<sup>-1</sup>; Ammann et al., 2015) at a rate competing with that 501 with OH under conditions that are not low-H<sub>2</sub>O, low-UV, and high-OHR<sub>ext</sub>. Note that the reaction of

502 formaldehyde with HO<sub>2</sub> is also significant in the atmosphere (Pitts and Finlayson, 1975; Gäb et al., 1985). 503 However, its product, hydroxymethylperoxy radical, dominantly undergoes decomposition via thermal 504 reaction and photolysis (Kumar and Francisco, 2015), compared to the hydroxymethylhydroperoxide 505 formation pathway via a further reaction with HO<sub>2</sub> (Ziemann and Atkinson, 2012). Even if 506 hydroxymethylhydroperoxide is produced in appreciable amounts, in the high-OH environment of OFRs, 507 this species can be easily predicted to convert into formic acid (Francisco and Eisfeld, 2009) and 508 eventually CO2. All these products have very few interactions with other VOCs, and hence should not 509 significantly perturb the reaction system of OFRs.

510

## 3.1.7 Overall contribution of non-OH reactants to gas-phase chemistry

In this section we summarize the combined effect of all non-OH reactants to VOC consumption.
 However, we can no longer use X<sub>exp</sub>/OH<sub>exp</sub> to express total non-OH VOC consumption as for individual
 reactants. Total non-OH VOC consumption is thus discussed case by case.

514 In the explored range of conditions (i.e., H<sub>2</sub>O, UV, OHR<sub>ext</sub>, and initial O<sub>3</sub> for OFR254), there are 515 obviously conditions where all non-OH fates of VOCs are negligible. Most simply, the highest H<sub>2</sub>O and 516 UV in this study and a very small non-zero OHRext result in a VOC consumption nearly 100% by OH, 517 regardless of the VOC type (Table S6). Lowering UV can make non-OH contribution even smaller for 518 OFR185, but not for OFR254. This difference occurs because OH production is reduced while O<sub>3</sub> roughly 519 remains at the same level in OFR254, leading to enhanced relative contribution from O<sub>3</sub> to the fate of 520 biogenics. At the lowest non-zero UV in Li et al. (2015)'s PAM reactor (7.9x10<sup>11</sup> photons cm<sup>-2</sup> s<sup>-1</sup> at 185 nm;  $2.0 \times 10^{14}$  photons cm<sup>-2</sup> s<sup>-1</sup> at 254 nm), the fractional destructions of  $\alpha$ -pinene (representative of 521 522 biogenic VOCs) by O<sub>3</sub> are 21% and 4.2% in OFR254-70 and OFR254-7, respectively. Other OFR designs 523 may reach lower UV, e.g., the "low UV" case defined in this study (1.0x10<sup>11</sup> photons cm<sup>-2</sup> s<sup>-1</sup> at 185 nm; 4.2x10<sup>13</sup> photons cm<sup>-2</sup> s<sup>-1</sup> at 254 nm). At these UV levels, the fate of  $\alpha$ -pinene by O<sub>3</sub> further increases to 524 525 44% and 13% in OFR254-70 and OFR254-7, respectively.

526 On the other hand, non-OH reactants can dominate VOC fate under opposite conditions that lead 527 to low OH production and strong OH suppression. At the lowest H<sub>2</sub>O and UV and the highest OHR<sub>ext</sub> in 528 this study, >95% of  $\alpha$ -pinene and ~80% of toluene are consumed by non-OH reactants in OFR185. In 529 OFR254, almost all α-pinene has non-OH fate while non-OH fate of toluene is still minor or negligible. If 530 OFRs are operated at the low UV setting from Li et al. (2015)'s PAM, ~8 times higher the lowest UV in 531 this study, the situation hardly changes, as a very large OH suppression persists. Nevertheless, if UV is 532 increased to the highest level, non-OH fate of  $\alpha$ -pinene is lowered to ~60–70% and that of toluene in 533 OFR185 decreases to 24%. When H<sub>2</sub>O is at the highest level, non-OH fate is systematically lower than at 534 low H<sub>2</sub>O in all types of OFRs. In particular, Case HHV has non-OH fates of  $\alpha$ -pinene only up to ~10% and 535 negligible non-OH fates of toluene, despite very high OHR<sub>ext</sub>.

536 We also summarize VOC fate for key laboratory, source, and field studies examined in the present 537 work in Fig. 7. For each case, the fate of one or a few typical VOCs is investigated. In laboratory studies, 538 Kang et al. (2011) performed experiments with a mixture of  $\alpha$ -pinene, m-xylene, and p-xylene, and one 539 of Lambe et al. (2011b)'s experiments used biogenics ( $\alpha$ -pinene and isoprene, respectively) as 540 precursors. In both cases, O<sub>3</sub> plays a role in the fate of biogenics when H<sub>2</sub>O is low (Kang et al.), or OHR<sub>ext</sub> 541 is high and UV is low (Lambe et al.), as shown in Fig. 7. The fate of isoprene by O<sub>3</sub> is less significant 542 despite the very high  $OHR_{ext}$ , because isoprene, compared to  $\alpha$ -pinene, is more reactive with OH and 543 less reactive with  $O_3$ . Besides,  $O(^{3}P)$  contributes up to a few percent to the fate of biogenics. In the 544 literature experiments performed at a higher H<sub>2</sub>O (Kang et al.) or a higher UV (Lambe et al.), non-OH 545 fate of both VOCs significantly decreases because of increases in OH<sub>exp</sub>. About 20% of p-xylene in Kang et al.'s mixture experiment at very low H<sub>2</sub>O may be destroyed by 254 nm photons, under the assumption 546 547 of unity quantum yield. Other laboratory study cases with aromatics have lower non-OH fates because 548 of higher H<sub>2</sub>O. n-decane in one of Lambe et al.'s experiments and dodecanoic acid in Klems et al. (2015)'s 549 study are consumed ~100% by OH, as these alkane(-like) species neither react rapidly with O<sub>3</sub>, nor 550 absorb UV efficiently. However, as previously discussed, some carbonyl compounds may be formed and 551 significantly photolyzed at 254 nm in Klems et al.'s experiments, although the huge complexity of 552 intermediates and limited knowledge of reaction mechanisms prevents a quantitative assessment of 553 the fate of carbonyl intermediates by photons at 254 nm.

554 Source and field studies usually have highly complex precursors. For the urban tunnel study (Tkacik 555 et al., 2014) and the CalNex-LA study in the Los Angeles basin (Ortega et al., 2015), we choose toluene 556 as a representative of aromatic species, as these are major anthropogenic VOCs and SOA precursors in 557 urban environment (Dzepina et al., 2009; Borbon et al., 2013; Hayes et al., 2015; Jathar et al., 2015). 558 Although alkanes are also major anthropogenic VOCs, their non-OH fate is not quantitatively assessed 559 for the same reason as discussed for dodecanoic acid in Klems et al. (2015)'s experiment. For the smoke 560 aging study, FLAME-3 (Ortega et al., 2013), we select benzene and  $\alpha$ -pinene, which are important VOCs 561 in biomass burning emissions (Warneke et al., 2011). For the BEACHON-RoMBAS and SOAS studies at 562 forested sites,  $\alpha$ -pinene and isoprene are chosen, respectively, as they are major emitted biogenic VOCs 563 at those corresponding sites. Both the urban tunnel and FLAME-3 studies have aromatic precursors 564 significantly photolyzed at 185 and 254 nm (assuming quantum yield = 1) under the conditions of high 565 source concentrations (Fig. 7). The toluene fate by UV in the tunnel study is less substantial than that of 566 benzene in FLAME-3, since NO<sub>x</sub>, the largest fraction of external OH reactant in the tunnel study, is 567 converted into HNO<sub>3</sub> very rapidly (Li et al., 2015) and does not further suppress OH. In the cases of lower 568 OHRext (e.g., the tunnel experiments with low source concentration and CalNex), toluene is dominantly 569 consumed by OH. It also holds for biogenic VOCs that non-OH fate decreases with decreasing OHRext 570 due to less OH suppression. The non-OH fate of  $\alpha$ -pinene in FLAME-3, dominated by reaction with O<sub>3</sub>, 571 is larger than 20%, while the non-OH fates of  $\alpha$ -pinene in BEACHON-RoMBAS and of isoprene in SOAS 572 are both negligible, since OHRext in the former study is >10 times higher than in the latter two studies.

573 3.2 SOA photolysis

Recently, photolysis in the UV range has been found to be a potentially significant sink of some types SOA in the troposphere (Updyke et al., 2012; Lambe et al., 2013; Liu et al., 2013, 2015; Hodzic et al., 2015; Wong et al., 2015; Romonosky et al., 2016). It is necessary to also investigate SOA photolysis in OFRs, as photons used in OFRs are highly energetic and non-tropospheric. UV extinction due to aerosol optical scattering and in-particle absorption under OFR conditions is generally negligible (Hodzic
et al., 2015). For simplicity, we estimate photodegradation ratios of various SOA component surrogates
as well as several SOA samples whose absorptivity was measured in previous studies (Updyke et al.,
2012; Lambe et al., 2013; Liu et al., 2013; Romonosky et al., 2016) (Fig. 8) under the assumption of unity
quantum yield to obtain upper limits of photodegradation ratios, and also under the assumption of
lower (0.1 and 0.01) quantum yields.

584 Most SOA functional groups are oxygenated (e.g., peroxides, carbonyls, carboxylic acids, alcohols). 585 The absorption cross sections of most of these functional groups are too low at 185 and 254 nm given 586 the OFR residence time and UV light intensity, leading to a small contribution of photolysis to SOA 587 degradation (Fig. 8). For example, glycolaldehyde has a negligible fractional contribution of photolysis 588 except at the highest lamp setting, when only ~5% of this species photolyzes at each wavelength. 589 Species (e.g., isoprene) with conjugated double bonds as efficient chromophores will not be present in 590 SOA because of their high reactivity with OH and O<sub>3</sub>. Nitrate groups may have a ~30% contribution from 591 photolysis at the highest UV settings, and a negligible contribution at intermediate or low UV settings.

592 Aromatic rings are more resistant to OH attack and usually strongly absorb UV light. Under our 593 assumptions, the photolysis of some aromatic SOA components (e.g., o-cresol at 185 nm and 594 naphthalene at 254 nm) is already important at medium UV flux. At the highest lamp setting, most 595 aromatics in SOA would be destroyed if the quantum yields are indeed near unity. However, as 596 previously discussed, photolysis quantum yields of aromatics may be significantly lower than 1. This is 597 more probable in the condensed phase (Damschen et al., 1978; Baker et al., 2015) than in the gas phase, 598 as quenching processes in the condensed-phase matrix are usually much more efficient than through 599 gas-phase molecular collisions. It has recently been reported that photolysis quantum yields of 600 aromatics in SOA were low under UVB irradiation (Romonosky et al., 2015). Although the range and 601 relevance of the species investigated in that study are limited, it is reasonable to assume low quantum 602 yield for aromatic photolysis in SOA at 185 and 254 nm.

603 Wong et al. (2015) conducted  $\alpha$ -pinene-derived SOA photolysis experiments in a chamber under 604 UVB irradiation (down to 284 nm). They observed at 85% RH ~30% SOA photolyzed after >30 min 605 irradiation and a photolysis quantum yield of ~1 during the first 10 min. However, in OFRs such a high 606 SOA photodegradation percentage would not occur, since Wong et al. (2015)'s experiments had a high 607 photon flux ( $\sim$ 4x10<sup>15</sup> photons cm<sup>-2</sup> s<sup>-1</sup>) and a long irradiation time, and hence a photon flux exposure 608 that is ~5 times that at the highest lamp setting in the OFRs modeled in our work. According to the 609 measurements of Wong et al. (2015), a photolysis fraction of ~6% would be expected for this type of 610 SOA in our OFRs under the highest UV flux, with lower percentages at lower UV settings. In addition, 611 the approximate unity quantum yield observed in Wong et al. (2015) may be due to (hydro)peroxides in 612  $\alpha$ -pinene-derived SOA, since peroxides have high photolysis quantum yields (Goldstein et al., 2007; 613 Epstein et al., 2012), while other functional groups (i.e., mainly hydroxyl and carbonyl) in oxygenated 614 species in SOA are unlikely to have for reasons discussed below.

615 Note that a simple addition of absorptivities of different functional groups may not explain SOA

616 absorptivity (Phillips and Smith, 2015). According to the absorption data of SOA samples from Lambe et 617 al. (2013) and Romonosky et al. (2015a), real SOA absorbs ~1-3 orders of magnitude more than non-618 aromatic component surrogate species shown in Fig. 8 at 254 nm. As discussed for multifunctional 619 oxidation intermediates (with carbonyls and hydroxyls), SOA absorption enhancement may be largely 620 due to transitions of charge transfer complexes formed between carbonyls and hydroxyls in 621 multifunctional oxygenated SOA components (Phillips and Smith, 2014, 2015). These complexes 622 between carbonyls and hydroxyls also have continua of states likely leading to ultrafast relaxation and 623 hence low photolysis quantum yields. Charge transfer transitions have been extensively shown in 624 measurements (Alif et al., 1991; Gao and Zepp, 1998; Johannessen and Miller, 2001; O'Sullivan et al., 625 2005; Zhang et al., 2006; Osburn et al., 2009; Sharpless and Blough, 2014) to have very low quantum 626 yields in the condensed phase. Sharpless and Blough (2014) compiled quantum yields of various 627 products of humic-like matter photolysis down to 280 nm. No quantum yields except those of the 628 product <sup>1</sup>O<sub>2</sub>, which is generally unimportant for OFRs (see Section 3.1.6), are higher than 0.01. If the 629 photolysis quantum yields of the SOA samples in Fig. 8 at 254 nm are no more than 0.01, no SOA samples 630 will be photolyzed by 20% even at the highest OFR lamp setting, and photolysis of most SOA samples at 631 254 nm will be minor or negligible in OFRs. Thus, to our current knowledge, lack of solid information on 632 quantum yields of SOA components with multiple carbonyls and hydroxyls at 254 nm prevents a clear 633 assessment of SOA photolysis in OFRs at the medium and high UV. On the other hand, direct 634 measurements are desirable for this issue and caution should still be exercised for OFR experiments at 635 relatively high UV.

SOA photolysis at 185 nm may be lower compared to that at 254 nm. SOA absorptivity data at 185 nm are not available. According to SOA mass-specific absorption cross-section (MAC) data between 250 and 300 nm in Romonosky et al. (2015a), there is a linear relationship between the logarithm of MAC and wavelength for most SOA samples: MAC increases by a factor of ~3 per 50 nm decrease in wavelength. We thus extrapolate this relationship to 185 nm, where MAC is estimated to be ~3.5 times higher than that at 254 nm. However, the UV flux at 185 nm in our OFR is ~100 times lower than at 254 nm.

643 Based on the discussion above, the SOA photodegradation ratio of ~30% in Wong et al. (2015)'s 644 non-OFR setup may be explained.  $\alpha$ -pinene-derived SOA has ~20–50% weight fraction of peroxides 645 (Docherty et al., 2005; Epstein et al., 2014), which may undergo photolysis in SOA to convert into 646 carbonyls (and hydroxyls) (Epstein et al., 2014). We speculate that after the formation of carbonyls from 647 peroxides, SOA materials cannot proceed significantly further with photolysis as discussed for charge 648 transfer between carbonyl and hydroxyl above. In the experiments of Wong et al. (2015), as well as 649 Epstein et al. (2014), effective photolysis rate constants/quantum yields decreased as SOA photolysis 650 proceeded. Photolysis rates were substantially reduced after a ~30% mass loss due to photolysis in 651 Wong et al. (2015)'s experiments. This mass loss ratio is consistent with the mass percentage of 652 peroxides in  $\alpha$ -pinene-derived SOA. Again, we note that, according to the extrapolation from Wong et al. (2015)'s results, the mass loss percentage expected in our OFR under the highest UV flux is ~6% for 653

 $\alpha$ -pinene-derived SOA. This value is much lower than that shown in Fig. 8 under the assumption of unity quantum yield (~40%) because of a substantially decreasing quantum yield in the real photolysis experiments. Therefore, in OFRs, even if (hydro)peroxides in SOA may be photolyzed in appreciable amounts, SOA mass is unlikely to be largely destroyed by photons in OFRs, as (hydro)peroxides may convert into carbonyls and hydroxyls, which may substantially lower subsequent photolysis quantum yields.

According to the discussion above, measurements of quantum yields and/or products of SOA photolysis are highly desirable, especially for the photolysis of SOA containing dominantly carbonyl and hydroxyl groups, as (hydro)peroxides, which are likely to form in OFRs, may convert into hydroxyls and carbonyls. With more data on quantum yields of SOA photolysis, a clearer strategy for including or excluding SOA photolysis in OFRs can be made.

665 Even though SOA photolysis can be significant in OFRs, it only proceeds to a much lesser extent 666 compared to ambient SOA photolysis. We calculate the numbers of e-fold decay of SOA photolysis in 667 OFR254-70 and the troposphere according to the effective ambient photolysis lifetime of SOA from 668 Romonosky et al. (2015a). Under the condition of 70% RH ( $H_2O = 1.4\%$ ) and OHR<sub>ext</sub> = 25 s<sup>-1</sup> (typical of ambient conditions), SOA samples are estimated to undergo ~0.01–0.5 e-fold photolysis timescales (i.e., 669 670  $\sim$ 1–35% OA photolyzed) in OFR254-70 at an equivalent photochemical age of 1 week under the upper 671 limit assumption of unity quantum yields (Table S8). However, in the atmosphere, those samples may proceed with  $10^2-10^4$  e-fold decays of photolysis (i.e., virtually complete destruction) at the same 672 673 photochemical age, if ambient SOA photolysis quantum yields are assumed to be those of H<sub>2</sub>O<sub>2</sub> (unity 674 below 400 nm). Even if the quantum yield of acetone (non-zero below 320 nm, see Romonosky et al., 675 2015a) is taken as a surrogate for SOA, most types of SOA would still be completely or nearly completely 676 photolyzed under ambient conditions. These results demonstrate that ambient SOA photolysis is likely 677 to be much more important than in OFRs. On the other hand, they also highlight the need for studies 678 of ambient SOA photolysis quantum yields and photolytic aging, as ambient SOA is unlikely to be 679 completely destroyed by photons within only 1 week. Either their quantum yields are much lower than 680 used in this study, or the photolabile groups are destroyed and leave behind others that are not (or less) 681 photolabile during photolytic aging.

### 682 3.3 Guidelines for OFR experimental design and operation

It is necessary to avoid significant non-tropospheric chemistry in OFRs in order to more accurately simulate tropospheric aging. Only photolysis at 185 and 254 nm are important non-tropospheric pathways in OFRs and reactions with O atoms are generally unimportant. Ozonolysis is also a major VOC sink in the troposphere, and the desirability of including or excluding its effects depends on the goals of each experiment.

In the cases where the exclusion of VOC ozonolysis is desired, there is no dilemma for the experimental design, as the exclusion of both VOC ozonolysis and non-tropospheric VOC consumption requires similar conditions, i.e., safer conditions. As shown above, all examined field studies do not have significant non-tropospheric contribution to VOC fate, while some past laboratory and source studies 692 do because of low  $H_2O$  and/or very high OHR<sub>ext</sub> in those experiments. It is possible to improve the latter 693 experiments by increasing H<sub>2</sub>O and/or lowering OHR<sub>ext</sub>. In detail, source humidification and dilution can 694 be feasible measures to increase H<sub>2</sub>O and decrease OHR<sub>ext</sub>, respectively. For example, increasing RH 695 from 3% to 60% (H<sub>2</sub>O from ~0.06% to ~1.2%) lowers the percentage of non-tropospheric consumption 696 of p-xylene in Kang et al. (2011)'s mixture experiment from ~20% to 1.5%. Humidifying the average 697 condition of the BEACHON-RoMBAS (Palm et al., 2015) campaign from  $H_2O = 1.6\%$  (RH = 63%) to  $H_2O =$ 698 2.3% (RH = 92%) leads to significant (from ~20% for 185 nm photon flux to a factor of ~3 for  $O(^{3}P)$ ) 699 decreases in all exposure ratios between non-OH reactants and OH (Table S4). Also, a 5-fold source 700 dilution in FLAME-3 reduces the non-tropospheric fate of benzene from >60% to ~15% (Table S4 and 701 "Improved" cases in Fig. 7). Injecting less precursor is a simple way to keep a reasonably low OHRext in 702 laboratory studies. The comparison between the cases with high and low concentrations in the urban 703 tunnel study (Tkacik et al., 2014) is a good example (Table S4 and Fig. 7). Note that when taking the 704 measures above to limit non-tropospheric VOC fate, one generally reduces the contribution from all 705 non-OH reactants. Increasing H<sub>2</sub>O and source dilution also significantly lower the relative importance of 706 ozonolysis in the fate of  $\alpha$ -pinene in Kang et al.'s mixture experiments and the FLAME-3 study (Table S4 707 and Fig. 7). Although increasing UV may increase OH production, OH reactant destruction, and hence 708 the relative contribution of OH to VOC fate in some cases, one has to be cautious when taking this 709 measure to reduce effective OHRext, as high UV may cause non-tropospheric photolysis of aromatic SOA 710 components.

711 In laboratory experiments, running OFRs under safer conditions ensures a minor contribution of 712 non-tropospheric photolysis, based on the current knowledge of oxidation intermediate photolysis (Fig. 713 2). This also reduces the relative contribution of ozonolysis to VOC fate. However, when more 714 information becomes available about photoysis quantum yields of oxidation intermediates (vs. the 715 upper limits assumed here), there may be additional flexibility to include ozonolysis while excluding 716 non-tropospheric VOC consumption. As the precursor composition is usually relatively simple in 717 laboratory experiments, it is sufficient to ensure the insignificance of non-tropospheric consumption of 718 only the precursor(s) and possible intermediates (usually oxygenated species), rather than for a large 719 variety of VOC precursors and intermediates. For example, in the case of quantum yields significantly 720 lower than used in the present work, we may perform OFR254-70 experiments with a large amount of 721 biogenics at medium H<sub>2</sub>O and UV. In this case, a tropospheric O<sub>3exp</sub>/OH<sub>exp</sub> ratio can be obtained without 722 major side effects, because the fractional contribution of photolysis of possible intermediates is still 723 minor due to low quantum yields. On the other hand, OFR experiments with some anthropogenic VOCs, 724 such as alkanes, can just be conducted at high H<sub>2</sub>O and low OHR<sub>ext</sub> to avoid the contribution of all non-725 OH reactants, since ozonolysis of alkanes is negligible even at a tropospheric O<sub>3exp</sub>/OH<sub>exp</sub>.

OFR experiments can be simply conducted under safer conditions to avoid non-tropospheric VOC fate, while riskier conditions can lead to significant non-tropospheric VOC fate, depending on the species under study. The conditions in between, i.e., "transition" conditions, are explicitly discussed above. However, one may want to be able to more quantitatively estimate the relative importance of 730 non-OH reactants under different conditions so that a more detailed experimental planning becomes 731 possible that simultaneously ensures insignificant non-tropospheric VOC fate and specific experimental 732 goals. For this purpose, we provide a series of estimation equations for non-OH reactant exposures 733 (Section S3, Table S9, and Fig. S6, as well as Excel file). With these equations, the relative contribution 734 of non-OH reactants under all conditions explored in this study can be easily estimated. In OFR studies 735 where a different OFR design is adopted and/or chemistry beyond the approximations in our model is 736 involved, a new model may need to be established, which can be done in similar manner as Peng et al. 737 (2015), to obtain the relative importance of non-OH VOC fate and then perform experimental design.

### 738 4 Conclusions

739 We used a kinetic model to investigate non-OH contribution (from 185 and 254 nm photons,  $O(^{1}D)$ , 740 O(<sup>3</sup>P), and O<sub>3</sub>) to VOC destruction, as well as to SOA photolysis at 185 and 254 nm in OFRs. We assessed 741 the relative significance of the VOC consumption by non-OH reactants to that by OH in OFRs and the 742 troposphere. The only non-tropospheric reaction that can play a major role under OFR conditions is 743 photolysis, especially at 254 nm. The relative importance of photolysis is largest under riskier OFR 744 conditions where OH is low due to low H<sub>2</sub>O and/or high OHR<sub>ext</sub>. Due to lack of quantum yield data, we 745 estimated upper limits of the relative importance of photolysis for the few most susceptible oxidation 746 intermediates, which are comparable to those from aromatic precursors. Reactions of O atoms are not 747 competitive and are actually of lower relative importance (vs. OH) in OFRs than in the troposphere. VOC 748 ozonolysis is much less important than in the troposphere under typical OFR conditions and of similar 749 importance under riskier OFR conditions. Photolysis of SOA in OFRs could be significant at medium and 750 high UV, but only if corresponding quantum yields are high. If SOA photolysis quantum yields are of the 751 order of 0.01 or lower, as measured for many humic-like substances (Sharpless and Blough, 2014), SOA 752 photolysis in OFRs may be minor or negligible under most conditions. Although the reaction fates may 753 be different, numbers of e-fold decays of photolysis for a given OHexp are at least an order-of-magnitude 754 lower in the OFRs compared to the troposphere.

755 We examined some past field, laboratory, and source studies using OFRs. In the field studies of 756 aged urban and forest ambient air, non-OH VOC fate was not important because of relatively high H<sub>2</sub>O 757 and moderate OHRext. However, some laboratory and source studies were conducted at low H<sub>2</sub>O and/or 758 high OHRext, and have significant non-tropospheric VOC consumption. Humidification and/or dilution 759 are recommended in these cases to reduce the importance of non-tropospheric reactants. We proposed 760 different approaches to avoid non-OH VOC consumption, as well as strategies to employ insignificant 761 non-tropospheric photolysis and significant tropospheric ozonolysis simultaneously in laboratory 762 experiments. Our work has implications for the interpretation of past OFR studies, and should be useful 763 for designing and conducting future OFR experiments for atmospheric research, as well as in related 764 applied fields.

The need for systematic measurements of photolysis quantum yields, for both VOC and SOA, and both at actinic wavelengths and at 185 and 254 nm, was highlighted in this study. When quantum yield data become available, photolysis of oxidation precursors, oxidation intermediates, and SOA in OFRs can be much better quantified, its relative importance compared to OH oxidation, ambient photolysis
etc. can be better evaluated, and experimental planning might then be able to be less conservative and
have more freedom to avoid non-tropospheric photolysis and realize specific experimental objective(s).

771

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### 782 References

- Aimanant, S. and Ziemann, P. J.: Chemical Mechanisms of Aging of Aerosol Formed from the Reaction
  of n-Pentadecane with OH Radicals in the Presence of NOx, Aerosol Sci. Technol., 47(9), 979–990,
  doi:10.1080/02786826.2013.804621, 2013.
- Alif, A., Pilichowski, J. and Boule, P.: photochemistry and environment XIII phototransformation of 2nitrophenol in aqueous solution, J. Photochem. Photobiol. A Chem., 59, 209–219, doi:10.1016/10106030(91)87009-K, 1991.
- Ammann, M., Cox, R. A., Crowley, J. N., Jenkin, M. E., Mellouki, A., Rossi, M. J., Troe, J., Wallington, T. J.,
  Cox, B., Atkinson, R., Baulch, D. L. and Kerr, J. A.: IUPAC Task Group on Atmospheric Chemical Kinetic
  Data Evaluation, [online] Available from: http://iupac.pole-ether.fr/#, 2015.
- Atkinson, R. and Arey, J.: Atmospheric degradation of volatile organic compounds., Chem. Rev., 103(12),
  4605–38, doi:10.1021/cr0206420, 2003.
- Bahreini, R., Middlebrook, A. M., Brock, C. A., de Gouw, J. A., McKeen, S. A., Williams, L. R., Daumit, K.
  E., Lambe, A. T., Massoli, P., Canagaratna, M. R., Ahmadov, R., Carrasquillo, A. J., Cross, E. S., Ervens, B.,
  Holloway, J. S., Hunter, J. F., Onasch, T. B., Pollack, I. B., Roberts, J. M., Ryerson, T. B., Warneke, C.,
  Davidovits, P., Worsnop, D. R. and Kroll, J. H.: Mass spectral analysis of organic aerosol formed downwind
  of the Deepwater Horizon oil spill: field studies and laboratory confirmations., Environ. Sci. Technol.,
  46(15), 8025–34, doi:10.1021/es301691k, 2012.
- Baker, L. a., Horbury, M. D., Greenough, S. E., Coulter, P. M., Karsili, T. N. V., Roberts, G. M., Orr-Ewing, A.
  J., Ashfold, M. N. R. and Stavros, V. G.: Probing the Ultrafast Energy Dissipation Mechanism of the
  Sunscreen Oxybenzone after UVA Irradiation, J. Phys. Chem. Lett., 6, 1363–1368,
  doi:10.1021/acs.jpclett.5b00417, 2015.
- Baulch, D. L., Bowman, C. T., Cobos, C. J., Cox, R. A., Just, T., Kerr, J. A., Pilling, M. J., Stocker, D., Troe, J.,
  Tsang, W., Walker, R. W. and Warnatz, J.: Evaluated kinetic data for combustion modeling: Supplement
  II, J. Phys. Chem. Ref. Data, 34(3), 757–1397, doi:10.1063/1.1748524, 2005.
- Baulch, D. L., Cobos, C. J., Cox, R. A., Esser, C., Frank, P., Just, T., Kerr, J. A., Pilling, M. J., Troe, J., Walker,
  R. W. and Warnatz, J.: Evaluated Kinetic Data for Combustion Modelling, J. Phys. Chem. Ref. Data, 21(3),
  411, doi:10.1063/1.555908, 1992.
- Beddard, G. S., Fleming, G. R., Gijzeman, O. L. J. and Porter, G.: Vibrational Energy Dependence of
  Radiationless Conversion in Aromatic Vapours, Proc. R. Soc. A Math. Phys. Eng. Sci., 340(1623), 519–533,
  doi:10.1098/rspa.1974.0168, 1974.
- Borbon, A., Gilman, J. B., Kuster, W. C., Grand, N., Chevaillier, S., Colomb, A., Dolgorouky, C., Gros, V.,
  Lopez, M., Sarda-Esteve, R., Holloway, J., Stutz, J., Petetin, H., McKeen, S., Beekmann, M., Warneke, C.,
  Parrish, D. D. and De Gouw, J. A.: Emission ratios of anthropogenic volatile organic compounds in
  northern mid-latitude megacities: Observations versus emission inventories in Los Angeles and Paris, J.
  Geophys. Res. Atmos., 118(4), 2041–2057, doi:10.1002/jgrd.50059, 2013.
- Bouzidi, H., Aslan, L., El Dib, G., Coddeville, P., Fittschen, C. and Tomas, A.: Investigation of the Gas-Phase
  Photolysis and Temperature-Dependent OH Reaction Kinetics of 4-Hydroxy-2-butanone, Environ. Sci.
  Technol., 49(20), 12178–12186, doi:10.1021/acs.est.5b02721, 2015.
- Bouzidi, H., Laversin, H., Tomas, a., Coddeville, P., Fittschen, C., El Dib, G., Roth, E. and Chakir, a.:
  Reactivity of 3-hydroxy-3-methyl-2-butanone: Photolysis and OH reaction kinetics, Atmos. Environ.,
  98(3), 540–548, doi:10.1016/j.atmosenv.2014.09.033, 2014.
- Calvert, J. G., Atkinson, R., Becker, K. H., Kamens, R. M., Seinfeld, J. H., Wallington, T. H. and Yarwood,
  G.: The Mechanisms of Atmospheric Oxidation of the Aromatic Hydrocarbons, Oxford University Press,
  USA. [online] Available from: https://books.google.com/books?id=P0basaLrxDMC, 2002.
- 827 Carlton, A. G., Wiedinmyer, C. and Kroll, J. H.: A review of Secondary Organic Aerosol (SOA) formation
  828 from isoprene, Atmos. Chem. Phys., 9(14), 4987–5005, doi:10.5194/acp-9-4987-2009, 2009.
- Carter, W. P. L., Cocker, D. R., Fitz, D. R., Malkina, I. L., Bumiller, K., Sauer, C. G., Pisano, J. T., Bufalino, C.
  and Song, C.: A new environmental chamber for evaluation of gas-phase chemical mechanisms and
  secondary aerosol formation, Atmos. Environ., 39(40), 7768–7788,
  doi:10.1016/j.atmosenv.2005.08.040, 2005.

- Cocker, D. R., Flagan, R. C. and Seinfeld, J. H.: State-of-the-Art Chamber Facility for Studying Atmospheric
   Aerosol Chemistry, Environ. Sci. Technol., 35(12), 2594–2601, doi:10.1021/es0019169, 2001.
- Cours, T., Canneaux, S. and Bohr, F.: Features of the potential energy surface for the reaction of HO2
  radical with acetone, Int. J. Quantum Chem., 107(6), 1344–1354, doi:10.1002/qua.21269, 2007.
- 837 Cubison, M. J., Ortega, a. M., Hayes, P. L., Farmer, D. K., Day, D., Lechner, M. J., Brune, W. H., Apel, E.,
- Diskin, G. S., Fisher, J. a., Fuelberg, H. E., Hecobian, a., Knapp, D. J., Mikoviny, T., Riemer, D., Sachse, G.
  W., Sessions, W., Weber, R. J., Weinheimer, a. J., Wisthaler, a. and Jimenez, J. L.: Effects of aging on organic aerosol from open biomass burning smoke in aircraft and laboratory studies, Atmos. Chem.
  Phys., 11(23), 12049–12064, doi:10.5194/acp-11-12049-2011, 2011.
- Bamschen, D. E., Merritt, C. D., Perry, D. L., Scott, G. W. and Talley, L. D.: Intersystem Crossing Kinetics,
  J. Phys. Chem., 82(21), 2268–2272, doi:10.1021/j100510a002, 1978.
- Bocherty, K. S., Wu, W., Lim, Y. Bin and Ziemann, P. J.: Contributions of Organic Peroxides to Secondary
  Aerosol Formed from Reactions of Monoterpenes with O3, Environ. Sci. Technol., 39(11), 4049–4059,
  doi:10.1021/es050228s, 2005.
- B47 Dzepina, K., Volkamer, R. M., Madronich, S., Tulet, P., Ulbrich, I. M., Zhang, Q., Cappa, C. D., Ziemann, P.
  B48 J. and Jimenez, J. L.: Evaluation of recently-proposed secondary organic aerosol models for a case study
  B49 in Mexico City, Atmos. Chem. Phys., 9(15), 5681–5709, doi:10.5194/acp-9-5681-2009, 2009.
- Eisenberg, W. C., Taylor, K. and Murray, R. W.: Gas-Phase Kinetics of the Reaction of Singlet Oxygen with
  Olefins at Atmospheric Pressure, J. Phys. Chem., 90(20), 1945–1948, doi:10.1021/j100400a041, 1986.
- Epstein, S. A., Blair, S. L. and Nizkorodov, S. A.: Direct Photolysis of α-Pinene Ozonolysis Secondary
  Organic Aerosol: Effect on Particle Mass and Peroxide Content, Environ. Sci. Technol., 48(19), 11251–8,
  doi:10.1021/es502350u, 2014.
- Epstein, S. A., Shemesh, D., Tran, V. T., Nizkorodov, S. A. and Gerber, R. B.: Absorption Spectra and
  Photolysis of Methyl Peroxide in Liquid and Frozen Water, J. Phys. Chem. A, 116(24), 6068–6077,
  doi:10.1021/jp211304v, 2012.
- Evans, R. C., Douglas, P. and Burrow, H. D., Eds.: Applied Photochemistry, Springer Netherlands,
  Dordrecht., 2013.
- Fang, W. H. and Phillips, D. L.: The crucial role of the S1/T2/T1 intersection in the relaxation dynamics of aromatic carbonyl compounds upon  $n \rightarrow \pi^*$  excitation, ChemPhysChem, 3(10), 889–892, doi:10.1002/1439-7641(20021018)3:10<889::AID-CPHC889>3.0.CO;2-U, 2002.
- 863 Foster, R.: Organic Charge-Transfer Complexes, Academic Press, New York., 1969.
- Francisco, J. S. and Eisfeld, W.: Atmospheric Oxidation Mechanism of Hydroxymethyl Hydroperoxide <sup>+</sup>,
  J. Phys. Chem. A, 113(26), 7593–7600, doi:10.1021/jp901735z, 2009.
- Gäb, S., Hellpointner, E., Turner, W. V. and Koŕte, F.: Hydroxymethyl hydroperoxide and
  bis(hydroxymethyl) peroxide from gas-phase ozonolysis of naturally occurring alkenes, Nature,
  316(6028), 535–536, doi:10.1038/316535a0, 1985.
- Gao, H. and Zepp, R. G.: Factors Influencing Photoreactions of Dissolved Organic Matter in a Coastal
  River of the Southeastern United States, Environ. Sci. Technol., 32(19), 2940–2946,
  doi:10.1021/es9803660, 1998.
- George, I. J., Vlasenko, A., Slowik, J. G., Broekhuizen, K. and Abbatt, J. P. D.: Heterogeneous oxidation of
  saturated organic aerosols by hydroxyl radicals: uptake kinetics, condensed-phase products, and particle
  size change, Atmos. Chem. Phys., 7(16), 4187–4201, doi:10.5194/acp-7-4187-2007, 2007.
- 875 Gierczak, T. and Ravishankara, A. R.: Does the HO2 radical react with ketones?, Int. J. Chem. Kinet., 32(9),
  876 573–580, doi:10.1002/1097-4601(2000)32:9<573::AID-KIN7>3.0.CO;2-V, 2000.
- Goldstein, S., Aschengrau, D., Diamant, Y. and Rabani, J.: Photolysis of Aqueous H 2 O 2 : Quantum Yield
  and Applications for Polychromatic UV Actinometry in Photoreactors, Environ. Sci. Technol., 41(21),
  7486–7490, doi:10.1021/es071379t, 2007.
- Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen, J., Donahue,
  N. M., George, C., Goldstein, A. H., Hamilton, J. F., Herrmann, H., Hoffmann, T., Iinuma, Y., Jang, M.,

- Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr, A., Maenhaut, W., McFiggans, G., Mentel, T. F., Monod, A.,
  Prevot, A. S. H., Seinfeld, J. H., Surratt, J. D., Szmigielski, R. and Wildt, J.: The formation, properties and
  impact of secondary organic aerosol: current and emerging issues, Atmos. Chem. Phys., 9(14), 5155–
  5236 [online] Available from: <Go to ISI>://WOS:000268535500040, 2009.
- Hayes, P. L., Carlton, a. G., Baker, K. R., Ahmadov, R., Washenfelder, R. a., Alvarez, S., Rappenglück, B.,
  Gilman, J. B., Kuster, W. C., de Gouw, J. a., Zotter, P., Prévôt, a. S. H., Szidat, S., Kleindienst, T. E., Offenberg,
  J. H., Ma, P. K. and Jimenez, J. L.: Modeling the formation and aging of secondary organic aerosols in Los
  Angeles during CalNex 2010, Atmos. Chem. Phys., 15(10), 5773–5801, doi:10.5194/acp-15-5773-2015,
  2015.
- Hodzic, A., Madronich, S., Kasibhatla, P. S., Tyndall, G., Aumont, B., Jimenez, J. L., Lee-Taylor, J. and
  Orlando, J.: Organic photolysis reactions in tropospheric aerosols: effect on secondary organic aerosol
  formation and lifetime, Atmos. Chem. Phys., 15(16), 9253–9269, doi:10.5194/acp-15-9253-2015, 2015.
- Hoffmann, T., Odum, J. R., Bowman, F., Collins, D., Klockow, D., Flagan, R. C. and Seinfeld., J. H.:
  Formation of Organic Aerosols from the Oxidation of Biogenic Hydrocarbons, J. Atmos. Chem., 26(2),
  189–222, doi:10.1023/A:1005734301837, 1997.
- 897 Hu, W. W., Campuzano-Jost, P., Palm, B. B., Day, D. A., Ortega, A. M., Hayes, P. L., Krechmer, J. E., Chen, 898 Q., Kuwata, M., Liu, Y. J., de Sá, S. S., McKinney, K., Martin, S. T., Hu, M., Budisulistiorini, S. H., Riva, M., 899 Surratt, J. D., St. Clair, J. M., Isaacman-Van Wertz, G., Yee, L. D., Goldstein, A. H., Carbone, S., Brito, J., 900 Artaxo, P., de Gouw, J. A., Koss, A., Wisthaler, A., Mikoviny, T., Karl, T., Kaser, L., Jud, W., Hansel, A., 901 Docherty, K. S., Alexander, M. L., Robinson, N. H., Coe, H., Allan, J. D., Canagaratna, M. R., Paulot, F. and 902 Jimenez, J. L.: Characterization of a real-time tracer for isoprene epoxydiols-derived secondary organic 903 aerosol (IEPOX-SOA) from aerosol mass spectrometer measurements, Atmos. Chem. Phys., 15(20), 904 11807–11833, doi:10.5194/acp-15-11807-2015, 2015.
- Huie, R. E. and Herron, J. T.: Kinetics of the reactions of singlet molecular oxygen (O21Dg) with organic
  compounds in the gas phase, Int. J. Chem. Kinet., 5(2), 197–211, doi:10.1002/kin.550050204, 1973.
- Jathar, S. H., Cappa, C. D., Wexler, a. S., Seinfeld, J. H. and Kleeman, M. J.: Multi-generational oxidation
  model to simulate secondary organic aerosol in a 3-D air quality model, Geosci. Model Dev., 8(8), 2553–
  2567, doi:10.5194/gmd-8-2553-2015, 2015.
- Johannessen, S. C. and Miller, W. L.: Quantum yield for the photochemical production of dissolved
   inorganic carbon in seawater, Mar. Chem., 76(4), 271–283, doi:10.1016/S0304-4203(01)00067-6, 2001.
- Johnson, M. S., Nilsson, E. J. K., Svensson, E. A. and Langer, S.: Gas-Phase Advanced Oxidation for
  Effective, Efficient in Situ Control of Pollution, Environ. Sci. Technol., 48(15), 8768–8776,
  doi:10.1021/es5012687, 2014.
- Kang, E., Root, M. J., Toohey, D. W. and Brune, W. H.: Introducing the concept of Potential Aerosol Mass
  (PAM), Atmos. Chem. Phys., 7(22), 5727–5744, doi:10.5194/acp-7-5727-2007, 2007.
- Kang, E., Toohey, D. W. and Brune, W. H.: Dependence of SOA oxidation on organic aerosol mass
  concentration and OH exposure: experimental PAM chamber studies, Atmos. Chem. Phys., 11(4), 1837–
  1852, doi:10.5194/acp-11-1837-2011, 2011.
- Keller-Rudek, H., Moortgat, G. K., Sander, R. and Sörensen, R.: The MPI-Mainz UV/VIS Spectral Atlas of
  Gaseous Molecules of Atmospheric Interest, [online] Available from: www.uv-vis-spectral-atlasmainz.org, 2015.
- Klems, J. P., Lippa, K. a and McGivern, W. S.: Quantitative Evidence for Organic Peroxy Radical
  Photochemistry at 254 nm, J. Phys. Chem. A, 119(2), 344–351, doi:10.1021/jp509165x, 2015.
- Kumar, M. and Francisco, J. S.: Red-Light-Induced Decomposition of an Organic Peroxy Radical: A New
  Source of the HO 2 Radical, Angew. Chemie Int. Ed., doi:10.1002/anie.201509311, 2015.
- Kwok, E. and Atkinson, R.: Estimation of hydroxyl radical reaction rate constants for gas-phase organic
  compounds using a structure-reactivity relationship: An update, Atmos. Environ., 29(14), 1685–1695,
  doi:10.1016/1352-2310(95)00069-B, 1995.
- Lambe, A. T., Ahern, A. T., Williams, L. R., Slowik, J. G., Wong, J. P. S., Abbatt, J. P. D., Brune, W. H., Ng, N.
  L., Wright, J. P., Croasdale, D. R., Worsnop, D. R., Davidovits, P. and Onasch, T. B.: Characterization of
- 932 aerosol photooxidation flow reactors: heterogeneous oxidation, secondary organic aerosol formation

- and cloud condensation nuclei activity measurements, Atmos. Meas. Tech., 4(3), 445–461,
  doi:10.5194/amt-4-445-2011, 2011a.
- Lambe, A. T., Cappa, C. D., Massoli, P., Onasch, T. B., Forestieri, S. D., Martin, A. T., Cummings, M. J.,
  Croasdale, D. R., Brune, W. H., Worsnop, D. R. and Davidovits, P.: Relationship between Oxidation Level
  and Optical Properties of Secondary Organic Aerosol, Environ. Sci. Technol., 47(12), 6349–6357,
  doi:10.1021/es401043j, 2013.
- Lambe, A. T., Chhabra, P. S., Onasch, T. B., Brune, W. H., Hunter, J. F., Kroll, J. H., Cummings, M. J., Brogan,
  J. F., Parmar, Y., Worsnop, D. R., Kolb, C. E. and Davidovits, P.: Effect of oxidant concentration, exposure
  time, and seed particles on secondary organic aerosol chemical composition and yield, Atmos. Chem.
  Phys., 15(6), 3063–3075, doi:10.5194/acp-15-3063-2015, 2015.
- Lambe, A. T., Onasch, T. B., Croasdale, D. R., Wright, J. P., Martin, A. T., Franklin, J. P., Massoli, P., Kroll, J.
  H., Canagaratna, M. R., Brune, W. H., Worsnop, D. R. and Davidovits, P.: Transitions from
  Functionalization to Fragmentation Reactions of Laboratory Secondary Organic Aerosol (SOA)
  Generated from the OH Oxidation of Alkane Precursors, Environ. Sci. Technol., 46(10), 5430–5437,
  doi:10.1021/es300274t, 2012.
- Lambe, A. T., Onasch, T. B., Massoli, P., Croasdale, D. R., Wright, J. P., Ahern, A. T., Williams, L. R., Worsnop,
  D. R., Brune, W. H. and Davidovits, P.: Laboratory studies of the chemical composition and cloud
  condensation nuclei (CCN) activity of secondary organic aerosol (SOA) and oxidized primary organic
  aerosol (OPOA), Atmos. Chem. Phys., 11(17), 8913–8928, doi:10.5194/acp-11-8913-2011, 2011b.
- Laue, T. and Plagens, A.: Named Organic Reactions, 2nd ed., John Wiley & Sons, Chichester, England,
  New York. [online] Available from: http://www.wiley.com/WileyCDA/WileyTitle/productCd047001041X.html, 2005.
- Levy II, H.: Normal atmosphere: large radical and formaldehyde concentrations predicted., Science,
  173(3992), 141–143, doi:10.1126/science.173.3992.141, 1971.
- Li, R., Palm, B. B., Borbon, A., Graus, M., Warneke, C., Ortega, a M., Day, D. a, Brune, W. H., Jimenez, J.
  L. and de Gouw, J. a: Laboratory Studies on Secondary Organic Aerosol Formation from Crude Oil Vapors,
  Environ. Sci. Technol., 47(21), 12566–12574, doi:10.1021/es402265y, 2013.
- Li, R., Palm, B. B., Ortega, A. M., Hu, W., Peng, Z., Day, D. A., Knote, C., Brune, W. H., de Gouw, J. and
  Jimenez, J. L.: Modeling the radical chemistry in an Oxidation Flow Reactor (OFR): radical formation and
  recycling, sensitivities, and OH exposure estimation equation, J. Phys. Chem. A, 119(19), 4418–4432,
  doi:10.1021/jp509534k, 2015.
- Liu, P. F., Abdelmalki, N., Hung, H.-M., Wang, Y., Brune, W. H. and Martin, S. T.: Ultraviolet and visible
  complex refractive indices of secondary organic material produced by photooxidation of the aromatic
  compounds toluene and m-Xylene, Atmos. Chem. Phys., 15(3), 1435–1446, doi:10.5194/acp-15-14352015, 2015.
- Liu, P., Zhang, Y. and Martin, S. T.: Complex refractive indices of thin films of secondary organic materials
  by spectroscopic ellipsometry from 220 to 1200 nm, Environ. Sci. Technol., 47, 13594–13601,
  doi:10.1021/es403411e, 2013.
- Mao, J., Ren, X., Brune, W. H., Olson, J. R., Crawford, J. H., Fried, a., Huey, L. G., Cohen, R. C., Heikes, B.,
  Singh, H. B., Blake, D. R., Sachse, G. W., Diskin, G. S., Hall, S. R. and Shetter, R. E.: Airborne measurement
  of OH reactivity during INTEX-B, Atmos. Chem. Phys., 9(1), 163–173, doi:10.5194/acp-9-163-2009, 2009.
- 974 Massoli, P., Lambe, A. T., Ahern, A. T., Williams, L. R., Ehn, M., Mikkilä, J., Canagaratna, M. R., Brune, W. 975 H., Onasch, T. B., Jayne, J. T., Petäjä, T., Kulmala, M., Laaksonen, A., Kolb, C. E., Davidovits, P. and Worsnop, 976 D. R.: Relationship between aerosol oxidation level and hygroscopic properties of laboratory generated 977 secondary organic aerosol (SOA) particles, Geophys. Res. Lett., 37(24), L24801, 978 doi:10.1029/2010GL045258, 2010.
- Matsunaga, A. and Ziemann, P. J.: Gas-Wall Partitioning of Organic Compounds in a Teflon Film Chamber
  and Potential Effects on Reaction Product and Aerosol Yield Measurements, Aerosol Sci. Technol., 44(10),
  881–892, doi:10.1080/02786826.2010.501044, 2010.
- 982 Messaadia, L., El Dib, G., Ferhati, A. and Chakir, A.: UV–visible spectra and gas-phase rate coefficients 983 for the reaction of 2,3-pentanedione and 2,4-pentanedione with OH radicals, Chem. Phys. Lett., 626,

- 984 73–79, doi:10.1016/j.cplett.2015.02.032, 2015.
- 985 Monks, P. S.: Gas-phase radical chemistry in the troposphere, Chem. Soc. Rev., 34, 376–395, doi:10.1039/b307982c, 2005.
- 987 Nakashima, N.: Laser photolysis of benzene. V. Formation of hot benzene, J. Chem. Phys., 77(12), 6040,
   988 doi:10.1063/1.443847, 1982.
- Nakashima, N. and Yoshihara, K.: Laser flash photolysis of benzene. VIII. Formation of hot benzene from
  the S[sub 2] state and its collisional deactivation, J. Chem. Phys., 79(6), 2727–2735,
  doi:10.1063/1.446176, 1983.
- Nguyen, T. B., Crounse, J. D., Schwantes, R. H., Teng, a. P., Bates, K. H., Zhang, X., St. Clair, J. M., Brune,
  W. H., Tyndall, G. S., Keutsch, F. N., Seinfeld, J. H. and Wennberg, P. O.: Overview of the Focused Isoprene
  eXperiment at the California Institute of Technology (FIXCIT): mechanistic chamber studies on the
  oxidation of biogenic compounds, Atmos. Chem. Phys., 14(24), 13531–13549, doi:10.5194/acp-1413531-2014, 2014.
- O'Sullivan, D. W., Neale, P. J., Coffin, R. B., Boyd, T. J. and Osburn, C. L.: Photochemical production of
  hydrogen peroxide and methylhydroperoxide in coastal waters, Mar. Chem., 97(1-2), 14–33,
  doi:10.1016/j.marchem.2005.04.003, 2005.
- Odum, J. R., Hoffmann, T., Bowman, F., Collins, D., Flagan Richard, C. and Seinfeld John, H.: Gas particle
  partitioning and secondary organic aerosol yields, Environ. Sci. Technol., 30(8), 2580–2585,
  doi:10.1021/es950943+, 1996.
- 1003 Ono, R., Nakagawa, Y., Tokumitsu, Y., Matsumoto, H. and Oda, T.: Effect of humidity on the production
  1004 of ozone and other radicals by low-pressure mercury lamps, J. Photochem. Photobiol. A Chem., 274, 13–
  1005 19, doi:10.1016/j.jphotochem.2013.09.012, 2014.
- Ortega, A. M., Day, D. A., Cubison, M. J., Brune, W. H., Bon, D., de Gouw, J. A. and Jimenez, J. L.: Secondary
  organic aerosol formation and primary organic aerosol oxidation from biomass-burning smoke in a flow
  reactor during FLAME-3, Atmos. Chem. Phys., 13(22), 11551–11571, doi:10.5194/acp-13-11551-2013,
  2013.
- Ortega, A. M., Hayes, P. L., Peng, Z., Palm, B. B., Hu, W., Day, D. A., Li, R., Cubison, M. J., Brune, W. H.,
  Graus, M., Warneke, C., Gilman, J. B., Kuster, W. C., de Gouw, J. A. and Jimenez, J. L.: Real-time
  measurements of secondary organic aerosol formation and aging from ambient air in an oxidation flow
  reactor in the Los Angeles area, Atmos. Chem. Phys. Discuss., 15(15), 21907–21958, doi:10.5194/acpd15-21907-2015, 2015.
- 1015 Osburn, C. L., Retamal, L. and Vincent, W. F.: Photoreactivity of chromophoric dissolved organic matter
  1016 transported by the Mackenzie River to the Beaufort Sea, Mar. Chem., 115(1-2), 10–20,
  1017 doi:10.1016/j.marchem.2009.05.003, 2009.
- Palm, B. B., Campuzano-Jost, P., Ortega, A. M., Day, D. A., Kaser, L., Jud, W., Karl, T., Hansel, A., Hunter, J.
  F., Cross, E. S., Kroll, J. H., Peng, Z., Brune, W. H. and Jimenez, J. L.: In situ secondary organic aerosol formation from ambient pine forest air using an oxidation flow reactor, Atmos. Chem. Phys. Discuss., 15(21), 30409–30471, doi:10.5194/acpd-15-30409-2015, 2015.
- Peng, Z., Day, D. A., Stark, H., Li, R., Lee-Taylor, J., Palm, B. B., Brune, W. H. and Jimenez, J. L.: HOx radical
  chemistry in oxidation flow reactors with low-pressure mercury lamps systematically examined by
  modeling, Atmos. Meas. Tech., 8(11), 4863–4890, doi:10.5194/amt-8-4863-2015, 2015.
- Phillips, S. M. and Smith, G. D.: Light Absorption by Charge Transfer Complexes in Brown Carbon Aerosols,
   Environ. Sci. Technol. Lett., 1(10), 382–386, doi:10.1021/ez500263j, 2014.
- Phillips, S. M. and Smith, G. D.: Further Evidence for Charge Transfer Complexes in Brown Carbon
  Aerosols from Excitation–Emission Matrix Fluorescence Spectroscopy, J. Phys. Chem. A, 119(19), 4545–
  4551, doi:10.1021/jp510709e, 2015.
- Pitts, J. N. and Finlayson, B. J.: Mechanismen der photochemischen Luftverschmutzung, Angew. Chemie,
  87(1), 18–33, doi:10.1002/ange.19750870103, 1975.
- Platt, S. M., El Haddad, I., Zardini, a. a., Clairotte, M., Astorga, C., Wolf, R., Slowik, J. G., Temime-Roussel,
  B., Marchand, N., Ježek, I., Drinovec, L., Močnik, G., Möhler, O., Richter, R., Barmet, P., Bianchi, F.,

- Baltensperger, U. and Prévôt, a. S. H.: Secondary organic aerosol formation from gasoline vehicle
  emissions in a new mobile environmental reaction chamber, Atmos. Chem. Phys., 13(18), 9141–9158,
  doi:10.5194/acp-13-9141-2013, 2013.
- Presto, A. A., Huff Hartz, K. E. and Donahue, N. M.: Secondary Organic Aerosol Production from Terpene
  Ozonolysis. 1. Effect of UV Radiation, Environ. Sci. Technol., 39(18), 7036–7045, doi:10.1021/es050174m,
  2005.
- Pretsch, E., Bühlmann, P. and Badertscher, M.: Structure Determination of Organic Compounds, Springer
   Berlin Heidelberg, Berlin, Heidelberg., 2009.
- 1042 Renlund, A. M. and Trott, W. M.: ArF Laser-induced decomposition of simple energetic niolecules, Chem.
  1043 Phys. Lett., 107(6), 555–560, doi:10.1016/S0009-2614(84)85155-6, 1984.
- 1044 Roberts, J. M. and Fajer, R. W.: UV absorption cross sections of organic nitrates of potential atmospheric 1045 importance and estimation of atmospheric lifetimes, Environ. Sci. Technol., 23(8), 945–951, 1989.
- Romonosky, D. E., Ali, N. N., Saiduddin, M. N., Wu, M., Lee, H. J. (Julie), Aiona, P. K. and Nizkorodov, S.
  A.: Effective absorption cross sections and photolysis rates of anthropogenic and biogenic secondary
  organic aerosols, Atmos. Environ., 130, 172–179, doi:10.1016/j.atmosenv.2015.10.019, 2016.
- Romonosky, D. E., Laskin, A., Laskin, J. and Nizkorodov, S. A.: High-Resolution Mass Spectrometry and
   Molecular Characterization of Aqueous Photochemistry Products of Common Types of Secondary
   Organic Aerosols, J. Phys. Chem. A, 119(11), 2594–2606, doi:10.1021/jp509476r, 2015.
- Sander, S. P., Friedl, R. R., Barker, J. R., Golden, D. M., Kurylo, M. J., Wine, P. H., Abbatt, J. P. D., Burkholder,
  J. B., Kolb, C. E., Moortgat, G. K., Huie, R. E. and Orkin, V. L.: Chemical Kinetics and Photochemical Data
  for Use in Atmospheric Studies Evaluation Number 17, 2011.
- Saukko, E., Lambe, A. T., Massoli, P., Koop, T., Wright, J. P., Croasdale, D. R., Pedernera, D. A., Onasch, T.
  B., Laaksonen, A., Davidovits, P., Worsnop, D. R. and Virtanen, A.: Humidity-dependent phase state of
  SOA particles from biogenic and anthropogenic precursors, Atmos. Chem. Phys., 12(16), 7517–7529,
  doi:10.5194/acp-12-7517-2012, 2012.
- 1059 Schmidt, G. a, Kelley, M., Nazarenko, L., Ruedy, R., Russell, G. L., Aleinov, I., Bauer, M., Bauer, S. E., Bhat, 1060 M. K., Bleck, R., Canuto, V., Chen, Y., Cheng, Y., Clune, T. L., Del Genio, A., de Fainchtein, R., Faluvegi, G., 1061 Hansen, J. E., Healy, R. J., Kiang, N. Y., Koch, D., Lacis, A. a, LeGrande, A. N., Lerner, J., Lo, K. K., Matthews, 1062 E. E., Menon, S., Miller, R. L., Oinas, V., Oloso, A. O., Perlwitz, J. P., Puma, M. J., Putman, W. M., Rind, D., 1063 Romanou, A., Sato, M., Shindell, D. T., Sun, S., Syed, R. A., Tausnev, N., Tsigaridis, K., Unger, N., 1064 Voulgarakis, A., Yao, M.-S. and Zhang, J.: Configuration and assessment of the GISS ModelE2 1065 contributions to the CMIP5 archive, J. Adv. Model. Earth Syst., 6(1), 141–184, 1066 doi:10.1002/2013MS000265, 2014.
- Seakins, P. W.: A brief review of the use of environmental chambers for gas phase studies of kinetics,
  chemical mechanisms and characterisation of field instruments, EPJ Web Conf., 9, 143–163,
  doi:10.1051/epjconf/201009012, 2010.
- Sharpless, C. M. and Blough, N. V: The importance of charge-transfer interactions in determining
  chromophoric dissolved organic matter (CDOM) optical and photochemical properties, Environ. Sci.
  Process. Impacts, 16(4), 654, doi:10.1039/c3em00573a, 2014.
- Smith, J. D., Kroll, J. H., Cappa, C. D., Che, D. L., Liu, C. L., Ahmed, M., Leone, S. R., Worsnop, D. R. and
  Wilson, K. R.: The heterogeneous reaction of hydroxyl radicals with sub-micron squalane particles: a
  model system for understanding the oxidative aging of ambient aerosols, Atmos. Chem. Phys., 9(9),
  3209–3222, doi:10.5194/acp-9-3209-2009, 2009.
- Strollo, C. M. and Ziemann, P. J.: Products and mechanism of secondary organic aerosol formation from
  the reaction of 3-methylfuran with OH radicals in the presence of NOx, Atmos. Environ., 77, 534–543,
  doi:10.1016/j.atmosenv.2013.05.033, 2013.
- Tkacik, D. S., Lambe, A. T., Jathar, S., Li, X., Presto, A. A., Zhao, Y., Blake, D., Meinardi, S., Jayne, J. T.,
  Croteau, P. L. and Robinson, A. L.: Secondary Organic Aerosol Formation from in-Use Motor Vehicle
  Emissions Using a Potential Aerosol Mass Reactor, Environ. Sci. Technol., 48(19), 11235–11242,

- 1085 doi:10.1021/es502239v, 2014.
- Tsang, W.: Chemical kinetic data base for combustion chemistry part V. Propene, J. Phys. Chem. Ref. data,
  20(2), 221–274, doi:10.1063/1.555880, 1991.

Turro, N. J., Ramamurthy, V. and Scaiano, J. C.: Principles of Molecular Photochemistry: An Introduction,
University Science Books, Sausalito, CA, USA. [online] Available from:
http://www.uscibooks.com/turro2.htm, 2009.

- 1091 Updyke, K. M., Nguyen, T. B. and Nizkorodov, S. a.: Formation of brown carbon via reactions of ammonia
  1092 with secondary organic aerosols from biogenic and anthropogenic precursors, Atmos. Environ., 63, 22–
  1093 31, doi:10.1016/j.atmosenv.2012.09.012, 2012.
- Wang, B., Lambe, A. T., Massoli, P., Onasch, T. B., Davidovits, P., Worsnop, D. R. and Knopf, D. a.: The
  deposition ice nucleation and immersion freezing potential of amorphous secondary organic aerosol:
  Pathways for ice and mixed-phase cloud formation, J. Geophys. Res., 117(D16), D16209,
  doi:10.1029/2012JD018063, 2012.
- Wang, J., Doussin, J. F., Perrier, S., Perraudin, E., Katrib, Y., Pangui, E. and Picquet-Varrault, B.: Design of
  a new multi-phase experimental simulation chamber for atmospheric photosmog, aerosol and cloud
  chemistry research, Atmos. Meas. Tech., 4(11), 2465–2494, doi:10.5194/amt-4-2465-2011, 2011.
- Warneke, C., Roberts, J. M., Veres, P., Gilman, J., Kuster, W. C., Burling, I., Yokelson, R. and de Gouw, J.
  a.: VOC identification and inter-comparison from laboratory biomass burning using PTR-MS and PIT-MS,
  Int. J. Mass Spectrom., 303(1), 6–14, doi:10.1016/j.ijms.2010.12.002, 2011.
- Wong, J. P. S., Zhou, S. and Abbatt, J. P. D.: Changes in Secondary Organic Aerosol Composition and Mass
  due to Photolysis: Relative Humidity Dependence, J. Phys. Chem. A, 119(19), 4309–4316,
  doi:10.1021/jp506898c, 2015.
- Zhang, X., Cappa, C. D., Jathar, S. H., McVay, R. C., Ensberg, J. J., Kleeman, M. J. and Seinfeld, J. H.:
  Influence of vapor wall loss in laboratory chambers on yields of secondary organic aerosol., Proc. Natl.
  Acad. Sci. U. S. A., 111(16), 5802–7, doi:10.1073/pnas.1404727111, 2014.
- 1110 Zhang, Y., Xie, H. and Chen, G.: Factors Affecting the Efficiency of Carbon Monoxide Photoproduction in
  1111 the St. Lawrence Estuarine System (Canada), Environ. Sci. Technol., 40(24), 7771–7777,
  1112 doi:10.1021/es0615268, 2006.
- 1113 Ziemann, P. and Atkinson, R.: Kinetics, products, and mechanisms of secondary organic aerosol 1114 formation, Chem. Soc. Rev., 41(19), 6582, doi:10.1039/c2cs35122f, 2012.

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Figure 1. Fractional importance of the photolysis rate at 185 nm of several species of interest vs. the reaction rate with OH, as a function of the ratio of exposure to 185 nm photons (F185) and OH. F185 exposure (in photons cm<sup>-2</sup>) is the product of 185 nm photon flux (in photons cm<sup>-2</sup> s<sup>-1</sup>) and time (in s). F185 exposure / OH exposure is thus in cm s<sup>-1</sup>. The modeled frequency distributions of ratios of 185 nm photon exposure to OH exposure under riskier, safer, and transition conditions for OFR185 are also shown. The curves of aromatics and inorganic gases are highlighted by solid dots and upward triangles, respectively. The lower inset shows histograms of model-estimated F185/OH exposures for three field studies where OFR185 was used to process ambient air. Their ordinate is the fractional occurrence of a given condition (X<sub>exp</sub>/OH<sub>exp</sub>). All histograms are normalized to be of identical total area (i.e., total probability of 1). The upper inset (black and blue markers) shows similar information for source studies of biomass smoke

(FLAME-3; Ortega et al., 2013) and an urban tunnel (Tkacik et al., 2014). All curves, markers, and histograms in this figure share the same abscissa.



1127 Figure 2. Same format as Fig. 1, but for the fractional importance of the photolysis rate at 254 nm vs. the reaction rate with OH as a function of the ratio of exposure to 254 nm

- 1128 (F254) and OH. The modeled frequency distributions of ratios of 254 nm photon exposure to OH exposure under riskier, safer, and transition conditions for OFR185 and OFR254
- 1129 (-7 to-70) are also shown. The curves of saturated carbonyl compounds and possible highly absorbing oxidation intermediates are highlighted by downward triangles and
- 1130 squares, respectively. The insets show histograms of model-estimated F254/OH exposures for three field studies where OFR185 was used to process ambient air. In addition to
- source studies of biomass smoke (FLAME-3) and urban tunnel (Tkacik et al., 2014), F254 exposure/OH exposure ratios in two laboratory studies (Kang et al., 2011; Lambe et al.,
- 1132 2011b) are shown in the upper inset. Colored tags indicate species used in the laboratory experiments. The lower and upper limits of F254 exposure/OH exposure ratios in the
- 1133 experiments with a certain source in a certain study are denoted by tags below and above the markers, respectively.



**Figure 3**. Same format as Fig. 2, but for the ratio of the reaction rate with  $O(^{1}D)$  vs. OH as a function of the relative exposure to  $O(^{1}D)$  and OH. A typical value of the relative exposure of  $O(^{1}D)$  and OH in the troposphere estimated according to Monks (2005) is also shown.



**Figure 4**. Same format as Fig. 3, but for the ratio of the reaction rate with  $O(^{3}P)$  vs. OH as a function of the relative exposure of  $O(^{3}P)$  and OH. A typical value of the relative exposure of  $O(^{3}P)$  and OH in the troposphere from Calvert et al. (2002) is also shown.



**Figure 5**. Same format as Fig. 2, but for the fractional importance of the reaction rate with O<sub>3</sub> vs. OH as a function of the relative exposure of O<sub>3</sub> and OH. The curves of biogenics

- 1142 are highlighted by squares. Also shown are modeled distributions of the relative exposure of O<sub>3</sub> and OH at the Earth's surface (solid line) and throughout the column from the
- 1143 surface to a height with a pressure of 150 hPa (dashed line). The distributions were calculated from the mean daily concentrations of O<sub>3</sub> and OH as simulated by the GISS
- 1144 ModelE2.



**Figure 6.** Ambient photolysis fractions of secondary species in a week (calculated from photolysis rates reported in Hodzic et al. (2015)) vs. photolysis fractions of those species in OFR185 and OFR254-70 when reaching the same photochemical age (ambient OH concentration of 1.5x10<sup>6</sup> molecules cm<sup>-3</sup> assumed) under conditions of 70% relative humidity (water vapor mixing ratio of 1.4%) and 25 s<sup>-1</sup> initial OHR<sub>ext</sub>. If the points of a certain species for both OFR185 and OFR254-70 are available, the species name is tagged on the OFR185 point (downward arrow), otherwise on the OFR254-70 point (upward arrow). The 1:1, 1:10, 1:100, and 1:1000 lines are also shown for comparison.





**Figure 7.** VOC fate in several representative cases of the laboratory, source, and field studies examined in this work. More details on VOC fate in these studies can be found in Table S4.



1156Absorption cross-section at 254 nm (cm /molecule)1157Figure 8. Percentage of SOA photodegradation at (upper panel) 185 and (lower panel) 254 nm at different1158UV levels as a function of absorption cross-section under the assumptions of quantum yields of 1, 0.1,1159and 0.01. Absorption cross-sections of some SOA component surrogates (black tag) and SOA samples1160(orange tag; calculated from data in Lambe et al. (2013) and Romonosky et al. (2015a)) are also shown.

1161 Table 1. Code of the labels of typical cases. A case label is composed of three characters denoting the1162 water mixing ratio, the photon flux, and the external OH reactivity, respectively.

	Water mixing ratio	Photon flux	External OH reactivity
Options -	L=low (0.07%)	L=low (10 <sup>11</sup> photons cm <sup>-2</sup> s <sup>-1</sup> at 185 nm; 4.17x10 <sup>13</sup> photons.cm <sup>-2</sup> s <sup>-1</sup> at 254 nm)	0
	M=medium (1%)	M=medium (10 <sup>13</sup> photons cm <sup>-2</sup> s <sup>-1</sup> at 185 nm; 1.45x10 <sup>15</sup> photons cm <sup>-2</sup> s <sup>-1</sup> at 254 nm)	L=low (10 s <sup>-1</sup> ) typically for remote or clean urban air
	H=high (2.3%)	H=high (10 <sup>14</sup> photons cm <sup>-2</sup> s <sup>-1</sup> at 185 nm; 8.51x10 <sup>15</sup> photons cm <sup>-2</sup> s <sup>-1</sup> at 254 nm)	H=high (100 s <sup>-1</sup> ) typically for polluted urban air
		"L"=lowest in Li et al. (2015)'s PAM (7.9x10 <sup>11</sup> photons cm <sup>-2</sup> s <sup>-1</sup> at 185 nm; 2.04 x10 <sup>14</sup> photons.cm <sup>-2</sup> s <sup>-1</sup> at 254 nm)	V=very high (1000 s <sup>-1</sup> only for laboratory experiments
Example	LHO:	low water mixing ratio, high photon flux, no external OH reactivity	