1	Possible influence of atmospheric circulations on winter haze
2	pollution in Beijing-Tianjin-Hebei region, northern China
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12 Abstract:

Using the daily records derived from the synoptic weather stations and the 13 14 NCEP/NCAR and ERA-Interim reanalysis data, the variability of the winter haze pollutions (indicated by the mean visibility and number of hazy days) in Beijing-15 Tianjin-Hebei (BTH) region during the period 1981 to 2015 and its relationship to the 16 atmospheric circulations in middle-high latitude were analyzed in this study. The winter 17 haze pollution in BTH had distinct inter-annual and inter-decadal variabilities without 18 a significant long-term trend. According to the spatial distribution of correlation 19 coefficients, six atmospheric circulation indices (I₁ to I₆) were defined from the key 20 areas in sea level pressure (SLP), zonal and meridional winds at 850 hPa (U850, V850), 21 geopotential height field at 500 hPa (H500), zonal wind at 200 hPa (U200), and air 22 temperature at 200 hPa (T200), respectively. All of the six indices have significant and 23 stable correlations with the winter visibility and number of hazy days in BTH. In the 24 raw (unfiltered) correlations, the correlation coefficients between the six indices and 25 the winter visibility (number of hazy days) varied from 0.57 (0.47) to 0.76 (0.6) with 26 an average of 0.65 (0.54); in the high-frequency (<10 yr) correlations, the coefficients 27 varied from 0.62 (0.58) to 0.8 (0.69) with an average of 0.69 (0.64). The six circulation 28 indices together can explain 77.7% (78.7%) and 61.7% (69.1%) variances of the winter 29 30 visibility and number of hazy days in the year-to-year (inter-annual) variability, respectively. The increase of I_c (a comprehensive index derived from the six individual 31

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circulation indices) can cause a shallowing of the East Asian trough at the middle 32 troposphere and a weakening of the Siberian high pressure field at sea level, and then 33 accompanied by a reduction (increase) of horizontal advection and vertical convection 34 (relative humidity) in the lowest troposphere and a reduced boundary layer height in 35 BTH and its neighboring areas, which are favorable for the formation of haze pollutions 36 37 in BTH winter, and vice versa. The high level of the prediction statistics and the reasonable mechanism suggested that the winter haze pollutions in BTH can be 38 forecasted or estimated credibly based on the optimized atmospheric circulation indices. 39 Thus it is helpful for government decision-making departments to take actions in 40 advance in dealing with probably severe haze pollutions in BTH indicated by the 41 42 atmospheric circulation conditions.

43 Key word: haze pollution, visibility, atmospheric circulation, Beijing-Tianjin-Hebei,

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45 **1 Introduction**

Beijing-Tianjin-Hebei (BTH) region is located in northern China, with approximately 46 110 million residents and 216,000 km² in size. As the rapid progress of urbanization 47 and industrial development over the past three decades, the BTH region has become 48 one of China's most economically developed regions and the third economic engine in 49 China. Recently, the Chinese government has been promoting the integration of the 50 three neighboring regions to optimize the industrial layout and improve the allocation 51 of resources. Undoubtedly, the BTH region is becoming more and more important in 52 China or even the world economy in the future. However, the rapid economic growth 53 and urbanization have increased the level of air pollution in recent decades (Streets et 54 al., 2007; Chan and Yao, 2008; Wang et al., 2009; Wang et al., 2010; Gao et al., 2011). 55 56 Most of eastern China has frequently suffered from severe haze or smog days in recent years, especially in the BTH region. For example, the continuously haze pollutions in 57 January 2013 greatly threatened human health and traffic safety (Kang et al., 2013; 58 Wang et al., 2013). Roughly speaking, the haze pollution can be attributed to two 59 aspects: pollutant emissions to the lower atmosphere from fossil fuel combustion or 60 61 construction and favorable meteorological conditions. Meteorological conditions are controlling the occurrence of haze pollution (Wu, 2012; Zhang et al., 2013). 62 Specifically, weather conditions play an essential role in the daily fluctuation of air 63 pollutant concentrations (Zhang et al., 2015). 64

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At present, many studies have focused on the physical and chemical properties of

pollutants in Beijing and other cities (Feng et al., 2006; Yu et al., 2011; Xu et al., 2013; 66 Zhao et al., 2013). And also studies demonstrated the influence of weather conditions 67 or synoptic situations upon air pollutions (Zhao et al., 2009; Zhang et al., 2015). They 68 elucidated clearly the formation and chemical composition of air pollutants and the 69 dominant meteorological factors during heavy pollution in the BTH region. On the 70 71 other hand, some studies demonstrated that the haze pollution occurring in the BTH region could be strongly affected by the local atmospheric circulations including sea-72 land and mountain-valley breeze circulations and the planetary boundary layer height 73 (Lo et al., 2006; Liu et al., 2009; Chen et al., 2009; Miao et al., 2015). Recently, Wang 74 et al. (2015) suggested that the reduction of autumn Arctic sea ice leads to anomalous 75 76 atmospheric circulation changes which favor less cyclone activity and more stable atmosphere and leading to more hazy days in eastern China. Moreover, Wang et al. 77 (2013) showed that east China suffered from severe haze pollutions in January 2013 78 may be due to a sudden stratospheric warming over the mid-high latitude of Northern 79 Hemisphere, which lead to an anomalous steady atmosphere dominated in northern 80 81 China. Thus, it is interesting to examine whether the winter haze pollution in BTH has 82 been influenced by other known or unknown atmospheric circulations or teleconnections in the mid-high latitude of the Northern Hemisphere and whether there 83 are some potential circulations that can be used for the forecast or evaluation of the 84 85 winter haze pollution in BTH. To date, it is not clear about these questions, and a few 86 studies have been performed to explore these issues.

Owing to a lack of long-term instrumental records for air pollutant concentration, 87 the understanding of the evolution of air pollution and their relations to atmospheric 88 circulations is limited. In this paper, we intend to use the atmospheric visibility and the 89 number of hazy days derived from the synoptic meteorological stations to denote the 90 91 evolution of haze pollution in the BTH region since 1980s. Many studies demonstrated that, in the absence of certain weather conditions (e.g., rain, fog, dust and snowstorm), 92 the visibility is an excellent indicator of air quality because its degradation results from 93 light scattering and absorption by atmospheric particles and gases that can originate 94 from natural or anthropogenic sources (Baumer et al., 2008; Chang et al., 2009; 95 96 Sabetghadam et al., 2012; Baddock et al., 2014), although visibility was influenced comprehensively by airborne pollutants and meteorological parameters such as relative 97 humidity, wind speed, temperature, pressure and solar radiation (Wen and Yeh, 2010; 98 Deng et al., 2014; Zhang et al., 2015). 99

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The main purpose of this study is to examine the possible relations between the

atmospheric circulations and the winter haze pollution (the mean visibility and mean 101 number of hazy days) over the BTH region and investigate the possible physical 102 mechanism, which could be useful for a prediction of the winter haze pollution and 103 could provide a scientific support to the government to take effective measures in 104 advance to reduce or control the pollutant emission in case of an anomalous circulations 105 106 leading to a serious haze pollution in the region. This paper is organized as follows. Section 2 describes the data and method used. Section 3 shows major results and 107 discussions. Conclusion is summarized in section 4. 108

2 Data and methods

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110 2.1 Research area and station data

The atmospheric visibility recorded at the 19 synoptic meteorological stations 111 located in the research area from 1 January 1980 to 28 February 2015 were used (Figure 112 1). The visibility by human observers is recorded by four times (02:00, 08:00, 14:00 113 and 20:00, Beijing local time) or three times (08:00, 14:00 and 20:00, Beijing local time) 114 per day. A good continuous monitoring operation was maintained throughout the entire 115 period, with the missing data rates for the 19 stations varying from a minimum of 1.7% 116 to a maximum of 2.1%, with a mean 1.9%. On the other hand, the distribution of the 117 stations is relatively uniform, indicating that the mean visibility or hazy days is a good 118 representative for the whole BTH region. 119

In the present study, the days with visibility ≤ 5 km and relative humidity < 90%at 14:00PM (local time) were defined as hazy days, except the special weather phenomena occurred at this moment including rain, fog, dust and snow (Schichtel et al., 2010; Wu et al., 2014;). The mean number of hazy days (*NHD*) of each winter in the BTH region can be calculated by:

$$\overline{NHD} = \frac{1}{n} \sum_{i=1}^{n} N_i \tag{1}$$

where *n* is the number of stations (here *n*=19), *N* denotes the number of hazy days in a station in each winter (December, January and February). The mean visibility (\overline{Vis}) of each winter in the BTH region can be calculated by:

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$$\overline{Vis} = \frac{1}{n} \sum_{i=1}^{n} \left(\frac{1}{m} \sum_{j=1}^{m} V_{ij} \right)$$
(2)

where *n* is the number of stations (here n=19), *m* is the number of valid days in winter. It should be noted that the winter in 1981 consists of December 1980, January

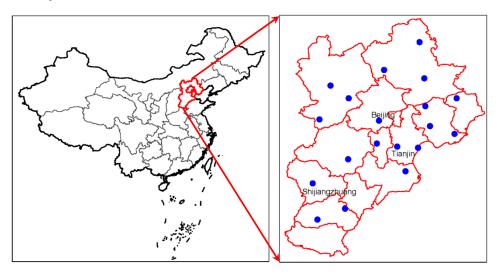




Figure 1 Research area and locations of the 19 synoptic meteorological stations

135 **2.2 Reanalysis data**

The global NCEP/NCAR reanalysis data of the monthly sea level pressure (SLP), 136 zonal and meridional winds at 850 hPa (U850, V850), geopotential height field at 500 137 hPa (H500), zonal wind at 200 hPa (U200) and air temperature at 200, 150, 100 and 70 138 hPa (T200, T150, T100, T70) with a 2.5°×2.5° spatial resolution from January 1980 to 139 February 2015 were used (Kalnay et al., 1996). Moreover, in order to obtain a higher 140 spatial resolution in the BTH region, the ERA-Interim reanalysis data of the monthly 141 relatively humidity (RH), vertical speed (W), zonal (U) and meridional (V) winds from 142 1000 to 500 hPa (16 pressure levels in total) and the boundary layer height (BLH) with 143 a 0.125°×0.125° spatial resolution confined to the area 33-45°N and 110-122°E were 144 also used (Dee et al., 2011). 145

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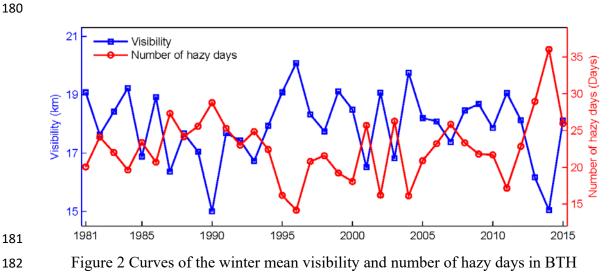
147 2.3 Analysis method

For the statistical and atmospheric circulation analyses carried out in the study, the 148 common statistical methods such as the composite analyses and the least square 149 regression and the Pearson correlation analyses with a two-tailed Student's t-test were 150 applied in this research. A principal component analysis (PCA) was also used to extract 151 the principal mode of multiple time series. Moreover, in order to reduce the possible 152 effects of low-frequency variation or long-term trends and to examine whether or not 153 the correspondence between the two time series on inter-annual time-scale is stable, the 154 high-frequency (< 10yr) correlation of the high-pass filtered time series was also tested 155 for time series analyses (Gong and Luterbacher, 2008; Zhang et al., 2010). 156 157

158 **3 Results and discussions**

159 3.1 Evolution of the winter visibility and hazy days in the BTH region

The regional mean visibility and number of hazy days in winter in BTH were 160 presented in Figure 2. As expected, the visibility was negatively correlated to the 161 number of hazy days with the raw and high-frequency (< 10yr) correlation coefficients 162 between them of -0.91 and -0.93, respectively. Both of them are significant at the 0.01 163 level (p < 0.01 for short). More hazy days generally denote lower mean visibility in 164 winter due to the light scattering and absorption effects of air pollutants (Baumer et al., 165 2008; Sabetghadam et al., 2012). There are intense inter-annual fluctuations in both the 166 visibility and the number of hazy days over the entire period of 1981 to 2015. The 167 168 decadal fluctuations can be also distinguished for both the visibility and the number of hazy days throughout the entire period. A significant reducing trend of visibility 169 (p<0.05) and increasing trend of number of hazy days (p<0.01) dominated in the 1980s. 170 And then, the visibility experienced an increasing trend in 1990s and a decreasing trend 171 since 2001, and the hazy days showed an anti-phase changes, but none of them are 172 statistically significant with exception of the number of hazy days trend in 1990s 173 (p<0.05). The mean visibility maximum in 1990s reached to 18.3 km (larger than the 174 mean values of 17.9 km over the entire period); and the minimum number of hazy days 175 in 1990s reached to 20.6 days (less than the mean values of 22.7 days over the entire 176 period). However, the long-term trends of them are not statistically significant, although 177 a weak reducing and increasing trends can be founded in the curves of winter visibility 178 and number of hazy days, respectively. 179



184 **3.2** Relationship between haze pollution and atmospheric circulations

We first examined the correlation coefficients between the visibility and number of 185 hazy days and the most common atmospheric teleconnection or oscillation indices over 186 the mid-high latitude of Northern hemisphere (see Table 1), which could affect the 187 winter climate variability over China, such as the Arctic Oscillation (AO), the Northern 188 Atlantic Oscillation (NAO), the Pacific/North American pattern (PNA), the Eurasian 189 190 pattern (EU), the Western Pacific pattern (WP) and the Siberian High (SBH) (Wallace and Gutzler, 1981; Zhang et al., 2009; Gong and Ho, 2012). It can be seen that both of 191 the raw (r1) and high-frequency (r2) correlations show that the visibility and number 192 of hazy days are correlated weakly with the winter AO, NAO and PNA. However, the 193 visibility is highly positively correlated with EU, WP and SBH; and the number of hazy 194 195 days is highly negatively correlated with EU, WP and SBH, most of them are significant at the 0.01 or 0.05 level. 196

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Table 1 Correlation coefficients of visibility and hazy days and circulation indices

		AO	NAO	PNA	EU	WP	SBH
Visibility	<i>r</i> 1	-0.11	0.00	0.16	0.61**	0.40^{*}	0.39*
	<i>r</i> 2	0.05	0.22	0.16	0.71**	0.37*	0.36*
Number of	<i>r</i> 1	0.13	0.13	-0.10	-0.51**	-0.47**	-0.32
hazy days	<i>r</i> 2	-0.01	-0.11	-0.10	-0.70**	-0.56**	-0.37*

199 200 ** Significant at the 0.01 level, * Significant at the 0.05 level. The *r*1 and *r*2 terms indicate the raw correlation and high-frequency (< 10yr) correlation, respectively.

201 Furthermore, the general characteristics of spatial distribution of the correlation coefficients between visibility and number of hazy days in BTH and the major 202 meteorological fields from surface to tropopause in Northern Hemisphere including 203 SLP, U850, V850, H500, U200, T200, T150 and T70 were also examined (Figure 3 and 204 4). Owning to a generally anti-pattern for the number of hazy days, thus only the 205 206 correlation maps with visibility were analyzed for simplicity. In SLP (Figure 3a), a positive correlation center dominated most of East Asian continent, while a negative 207 correlation center dominated the area from northeast Asia to northwest Pacific, 208 respectively. This spatial pattern may reflect the effects of land-sea thermal contrast on 209 the lower troposphere condition over BTH region. The pressure increasing in East 210 211 Asian continent and decreasing in area from northeast Asia to northwest Pacific suggest that they favor the visibility increase in the BTH region in winter, and vice versa. In 212 UV850 (Figure 3b), an anomalously anti-cyclonic and northerly pattern are 213 predominant most of Siberia and eastern China. This suggests that an anomalous 214

northerly advection from Siberian to eastern China improve the winter visibility in the 215 BTH region. In H500 (Figure 3c), there exist a "-+-" wave train pattern along the 216 Eurasia-west Pacific in the mid-high latitude, extending from the central-eastern 217 Europe through Siberia to north China-Korean peninsula-Japan-northwest Pacific 218 Ocean, similar to the EU pattern (Wallace and Gutzler, 1981). This pattern implies that 219 220 a deepening of East Asian trough and a weakening of blocking will favor the winter visibility increase in the BTH region. In U200 (Figure 3d), there also exist a wave train 221 pattern from northwest Russia through Siberia to northwest Pacific Ocean. This pattern 222 may imply that the south (north) of East Asian Jet stream strengthened (weakened) 223 coincided with the anomalous ascending (sinking) motions occurred in the south (north) 224 225 of the Jet stream entrance at the upper troposphere, which will lead to a strengthening northerly appeared in the lower troposphere. Hence it is not conducive to the 226 accumulation of pollutants over BTH region in the winter. 227

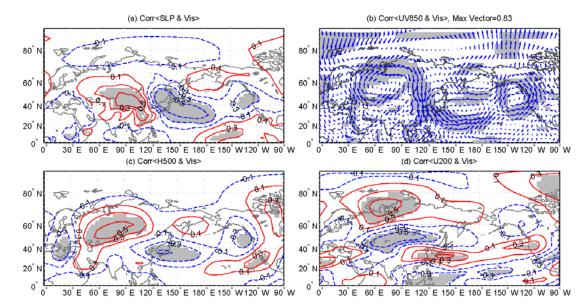


Figure 3, Spatial distribution of correlation coefficients between visibility and SLP (a), UV850
(b), H500 (c) and U200 (d) (Area significant at the 0.05 level are shaded; either U850 or V850
significant at the 0.05 level are shaded in b)

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Besides the lower troposphere, previous studies suggested that the anomalous stratospheric warming over the Northern Hemisphere led to the severe haze pollutions in east China in January 2013 (Wang et al., 2013). Here, the spatial distribution of the correlation coefficients between visibility and the temperature from the upper troposphere to lower stratosphere at 200 hPa (T200), 150 hPa (T150), 100 hPa (T100) and 70 hPa (T70) were checked. Negative correlations are found from eastern Siberia to the northern North Pacific including Alaska in T200, T150, T100 and T70,
respectively (Figure 4), with the biggest correlation in T200 (Figure 4a). The
significantly negative correlation suggest that the warming at 200 hPa over eastern
Siberia to the northern North Pacific would indicate a decreasing of winter visibility,
namely a worsening of haze pollutions in the BTH region.

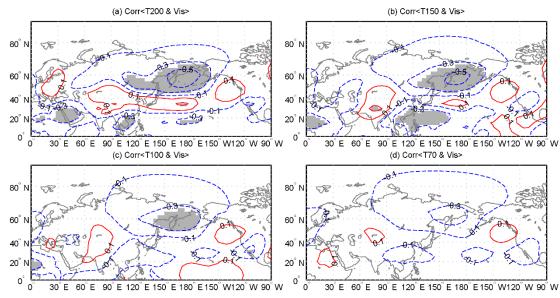


Figure 4 Spatial distribution of correlation coefficients between visibility and T200 (a), T150 (b),
T100 (c) and T70 (d) (Area significant at the 0.05 level are shaded)

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Based on the above analyses, we wonder whether the meteorological variables in 250 the significant correlation areas can be used to predict or evaluate the variability of the 251 winter visibility and haze pollutions in the BTH region. Thus, the six indices for 252 atmospheric circulations or teleconnections were defined based on the key regions 253 shown in the previous correlation maps as listed in Table 2. We computed the raw and 254 255 high-frequency correlation coefficients of the winter visibility and number of hazy days in BTH and the six atmospheric circulation indices. All of the six indices (I₁ to I₆) show 256 highly positive or negative correlations with the winter visibility and number of hazy 257 days, with significance at the 0.01 level (Table3). Moreover, we note that most of the 258 high-frequency correlations are larger than the raw correlations except the correlations 259 260 between visibility and I₁. This suggests that the links between the air quality in BTH and the circulations indices are very stable from year to year. The significantly positive 261 or negative correlations should be a reflection of the physical response mechanisms 262 between them, which will be discussed in the latter section. 263

Table 2 List of the definition for the six circulation indices

Index	Variable	Expression
I ₁	SLP	SLP (38~50N, 84~108E) – SLP (36~52N, 126~150E; 24~40N,150~184E)
I ₂	U _{850hPa}	U ₈₅₀ (55~75N, 40~110E) –U ₈₅₀ (40~50N, 45~75E)
I ₃	V _{850hPa}	V ₈₅₀ (32~64N, 104~120E)
I ₄	H _{500hPa}	H ₅₀₀ (46~64N, 50~92E) –H ₅₀₀ (28~44N, 16~28E; 28~42N, 120~156E)
I ₅	U _{200hPa}	U ₂₀₀ (42~52N,60~110E) –U ₂₀₀ (64~76N,50~96E; 28~36N, 120~152E)
I ₆	T _{200hPa}	T ₂₀₀ (46~66N, 146~196E)

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Table 3 Correlation coefficients of visibility and number of hazy days and circulation indices

		I_1	I_2	I ₃	I_4	I_5	I ₆	
Visibility	<i>r</i> 1	0.73**	0.57**	-0.76**	0.62**	-0.59**	-0.61**	
	<i>r</i> 2	0.70**	0.68**	-0.80**	0.72**	-0.62**	-0.62**	
Number of	<i>r</i> 1	-0.60**	-0.47**	0.60**	-0.47**	0.52**	0.60**	
hazy days	r2	-0.61**	-0.65**	0.69**	-0.67**	0.58**	0.64**	
Same as Table 1								

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3.3 Predictions for visibility and number of hazy days based on the circulation indices

In order to assess the prediction capability of the six circulation indices for the 273 winter haze pollutions in BTH, the winter mean visibility and number of hazy days 274 were estimated by applying a multivariate regression method with the least square 275 estimate. The estimated curves by the fitting and the cross-validation with a leave-one-276 out method were displayed in Figure 5. Intuitively, both of the fitting curves and the 277 cross-validation curves are fairly consistent with the observed winter mean visibility 278 279 and number of hazy days over the last three decades. The raw and high-frequency correlation coefficients between the observed and the fitting visibility (number of hazy 280 days) are 0.88 (0.78) and 0.86 (0.77), respectively. All of them are significantly at 0.01 281 level. The six circulation indices together can explain 77.7% (78.7%) and 61.7% 282 (69.1%) variances of the winter visibility and number of hazy days over the BTH region 283 284 in the year to year (inter-annual) variability, respectively. A good fitting does not mean that there must be stable relationships between the dependent variable and explanatory 285 variables. Thus we emphasized testing the stability of the statistic models by means of 286 the Leave-N-out cross-validations. The statistics for the cross-validation estimations 287

were listed in Table 4, including the explained variance (r^2) , the standard error (SE), and reduction of error (RE). Previous studies suggested that RE is an extremely rigorous verification statistic because it has no lower bound, RE > 0 indicating the skillful estimation, RE > 0.2 indicating the reliable estimation and RE = 1.0 indicating a perfect estimation (Fritts, 1976; Gong and Luterbacher, 2008; Zhang et al., 2010).

293 The statistics for both the visibility and number of hazy days are generally stable (no sharply increase or decrease) when N increased from 1 to 11 (more than 30% sample 294 removed in regression models), although the r^2 and RE (SE) slightly decreased 295 (increased) with the increasing of N. For the visibility, the r^2 varied from 52.5% to 62.7% 296 with an average of 57.6%, the SE varied from 0.74 to 0.84 with an average of 0.79, the 297 298 RE varied from 0.49 to 0.61 with an average of 0.55. For the number of hazy days, the r^2 varied from 31.1% to 41.5% with an average of 35.2%, the SE varied from 3.37 to 299 3.66 with an average of 3.54, the RE varied from 0.23 to 0.38 with an average of 0.30. 300 The mildly changes of these statistics suggest that the statistic models between the given 301 atmospheric circulations and the haze pollution indicators are stably even in the case of 302 parts of sample missed. On the other hand, we noted the statistics for the visibility 303 estimations are generally better than that for the number of hazy days estimations in all 304 tests. However, the minimum values of r^2 and RE for the number of hazy days 305 estimations are still lager than 30% and 0.2, respectively. Based on these statistics, it 306 can be concluded that the predictions for the winter visibility and number of hazy days 307 308 in the BTH region based on the circulation indices are overall reliable during the entire 309 period, especially for the mean visibility. That is to say, the winter haze pollutions in BTH can be evaluated or estimated well by the optimized atmospheric circulations. 310

The relatively larger errors for the estimated values referred to the observed 311 visibility and number of hazy days have been found since the winter in 2009 (Figure 5). 312 313 We re-computed all the statistics for the period 1981 to 2008, the results displayed that all the values of r^2 and RE (SE) for visibility and number of hazy days predictions 314 increased (decreased) much more than the entire period (Table 4), suggesting that the 315 316 statistic estimation models are much more stable and reliable before 2009. Why did the prediction efficiency of the statistic estimation models decrease relatively in the last 317 318 few years? It can be distinguished that the estimations for the winter mean visibility are distinctly lower (higher) than the observed in the winters of 2009 and 2010 (2014), and 319 vice versa for the number of hazy days. We speculated that these phenomena can be 320 attributed partly to the fluctuations of pollutant emissions because the pollutant 321 emissions over northern China around 2008 were controlled strictly by the Chinese 322

government associated with the 2008 Olympic Games in Beijing (An et al., 2007; 323 Zhang et al., 2010; Gao et al., 2011). The decrease of pollutant emissions led to the 324 improvement of air quality (increasing visibility and decreasing hazy days) in 2009 and 325 2010, although the atmospheric conditions remained the same and did not contributed 326 to the spread and elimination of air pollutants. However, pollutant emissions especially 327 328 in the areas of BTH rebounded after the Olympic Games, with the decrease in visibility and increase in hazy days in the BTH region around 2012 to 2014 to some extent (Zhang 329 et al., 2015). Generally, the errors between the observed visibility (haze days) and the predicted 330 331 could be attributed to the natural variability of atmospheric circulation and the changes of pollutant emissions. However, the contribution rates of each factor are not clear now, thus further studies will 332 333 be necessary to unravel these issues.

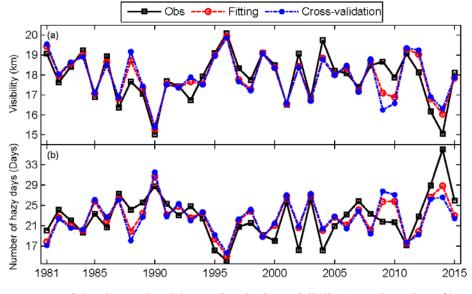


Figure 5 Curves of the observed and the predicted winter visibility (a) and number of hazy days
(b) in the BTH region since 1981



Table 4 List of the statistics for the Leave-N-out cross-validation estimations

N	Period covering	Visibility			Number of hazy days		
		r^{2} (%)	SE	RE	r^{2} (%)	SE	RE
1	1981-2015	62.7	0.74	0.61	41.5	3.37	0.38
	1981-2008	87.1	0.42	0.87	53.9	2.56	0.52
3	1981-2015	56.8	0.80	0.54	34.3	3.57	0.28
	1981-2008	86.8	0.42	0.87	52.6	2.59	0.51
5	1981-2015	59.2	0.78	0.57	35.3	3.54	0.30
	1981-2008	86.8	0.42	0.87	46.7	2.75	0.43

7 —	1981-2015	59.0	0.78	0.56	37.5	3.48	0.33
/ _	1981-2008	86.4	0.43	0.86	44.7	2.80	0.41
9 –	1981-2015	56.2	0.80	0.54	32.5	3.62	0.27
9 –	1981-2008	84.2	0.46	0.84	40.8	2.90	0.36
11	1981-2015	52.5	0.84	0.49	31.1	3.66	0.23
11 -	1981-2008	84.4	0.46	0.84	48.2	2.71	0.44

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(N denotes the number of sample removed in the cross-validation regressions; only the odd numbers of N were listed for short)

342 **3.4 Possible mechanism of the circulations related to the winter haze pollutions**

In order to explore the possible mechanism and role of the investigated circulation 343 344 indices on the winter visibility and number of hazy days in the BTH region, the links between the given large-scale atmospheric circulations and the local meteorological 345 conditions, which have close relations with the haze pollutions, were examined. For 346 simplicity, a comprehensive index labeled as Ic was synthesized from the six individual 347 circulation indices (I₁ to I₆) by applying a PCA method, namely the first principal 348 component (PC1). The high values of the explained variance (64.4% in PC1) indicated 349 that the comprehensive index of I_c roughly reflect the integrated features of all the six 350 indices. Thus, we used the Ic instead of the six individual indices in the following 351 analysis. Generally, the positive (negative) Ic indicate the lower (higher) visibility and 352 more (less) hazy days in the BTH region in winter. 353

First we examined the links between the I_c and the meteorological fields of SLP 354 and H500 respectively. Based on the NCEP/NCAR reanalysis data, Figure 6(a) and (b) 355 present the climatological mean of SLP and H500 in winter averaged from 1981 to 2010, 356 respectively. The changes of SLP and H500 in winter in association with a one-357 358 standard-deviation positive Ic during the winters 1981 to 2015 are shown in Figure 6(c) and (d), respectively. In the climatological mean fields, the BTH region were located in 359 the trough of East Asian trough at the middle troposphere and in the ridge of Siberian-360 Mongolia high in SLP field, which indicate the northerly dominated the BTH region in 361 winter. The regression maps show that the SLP decreased in the Siberian-Mongolia high 362 363 areas and increased in the western Pacific in SLP and the geopotential height decreased in the most areas of Siberia and increased in the northern China to western Pacific. 364 These patterns suggest that both the East Asian trough and Siberian high weaken with 365 increasing I_c, that further implies that the winter cold air activity will be weaken and 366 then lead to an anomalous steady atmospheric conditions in BTH and its adjacent areas 367

in winter. Namely, the less strong Siberian high and East Asian trough and associated 368 northerly winds in the low and middle troposphere will lead to a severe haze pollution 369 (lower visibility and more hazy days) due to the favorable meteorological conditions 370 for the accumulation and chemical reaction of pollutants. Anyway, we wonder whether 371 it is true as we speculated. We further examined the links between the comprehensive 372 373 index of Ic and the local meteorological conditions which play direct roles in the formation of haze pollutions, including the wind fields (Figure 7), relative humidity 374 (Figure 8) and vertical velocity (Figure 9) at the lowest troposphere (averaged from 375 1000 hPa to 900 hPa with an interval of 25 hPa) and the boundary layer height (Figure 376 10) based on the ERA-Interim reanalysis data. 377

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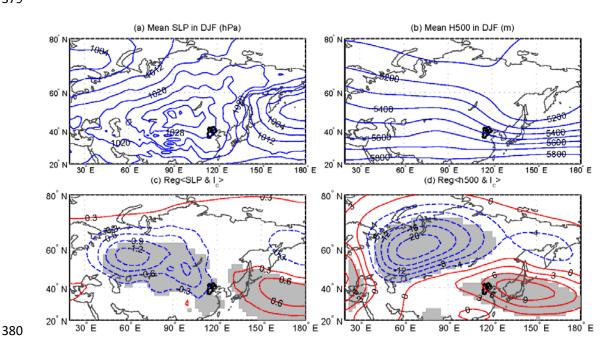


Figure 6 The climatological mean fields of SLP (a) and H500 (b) averaged in winter 1981 to 2010, 381 and the spatial distribution of the regression coefficients of SLP (c) and H500 (d) upon the Ic over 382 the period 1981 to 2015 (Area significant at the 0.05 level are shaded) 383

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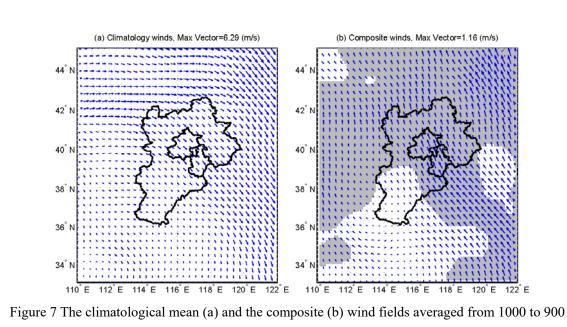
385

Figure 7(a) displays the climatological mean wind field averaged from 1000 to 900 hPa over the winter 1981 to 2010. At lower level, the northwesterly winds dominated 386 the BTH, and the wind speed in Beijing, Tianjin and north of Hebei province was larger 387 than that in the south of Hebei province. Figure 7(b) shows the composite (positive I_c 388 winters minus negative Ic winters) wind field averaged from 1000 to 900 hPa over the 389 winter 1981 to 2015. In the composite wind field, the anomalous southeasterly winds 390 dominated the BTH region instead of the northwesterly in the climatological mean wind 391

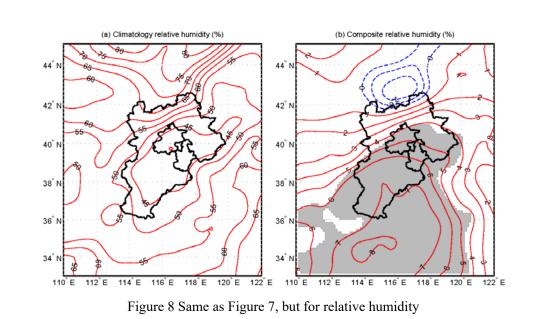
field, indicating the weakening of the northwesterly significantly over BTH and its neighboring areas when I_c increased. Previous studies (Zhang et al., 2015) demonstrated the decreasing of wind speed is not conducive to the diffusion of air pollutants and easily lead to haze pollutions in Beijing. It may be true for the whole BTH region. Thus, the increasing of I_c will lead to a decrease in the visibility and increase in the number of hazy days in winter over the BTH region.

Same as Figure 7, Figure 8(a) and (b) present the climatology and composite 398 relative humidity averaged from the lowest troposphere respectively. In the composite 399 map, all the areas of BTH are covered by the positive values and most of them are 400 significant at the 0.05 level. They indicate that the winter relative humidity was 401 402 anomalous higher in the positive I_c years than that in the negative I_c years. As pointed in the Introduction, a high relative humidity is one of the important reasons for visibility 403 degradation. This is because that the high relative humidity is favorable for the accumulation and 404 405 hygroscopic growth of pollutants, which can strengthen the light scattering and absorption by 406 atmospheric particles and gases and then cause the visibility degradation directly (Baumer et al., 407 2008; Zhang et al., 2015). Thus a positive Ic imply that a decreasing of visibility accompanied by the increasing number of hazy days may occur in the winter over BTH 408 region. Figure 9(a) and (b) present the climatology and composite vertical speeds 409 averaged from the lowest troposphere respectively. The positive (negative) values of 410 411 vertical speed in unit of Pa/s denote sinking (ascending) motion. The climatological 412 vertical speeds show that the downward air motions dominated the BTH region in the winter. In the composite vertical speed field, the most areas of BTH were covered by 413 the significantly negative values, which suggested a less vertical exchanges of air 414 occurred in this areas in the positive I_c winters. In other words, the increased I_c may 415 result in a weaker vertical convection and forcing the lowest troposphere more stable. 416 417 It's easy to understand the anomalous stabilization will lead to much haze pollutions. Moreover, a similar result can be found in the planetary boundary layer height, which 418 was reduced significantly in the most of BTH and its adjacent areas in the positive I_c 419 winters (Figure 10). The decreased boundary layer height will depress the air pollutants 420 into a narrower air column in a certain area and then lead to an increasing of the 421 422 pollutants concentration. Thus, a winter with the lower visibility and more hazy days in the BTH region would be expected in the case of the lower boundary layer height 423 caused by the anomalously high Ic. 424

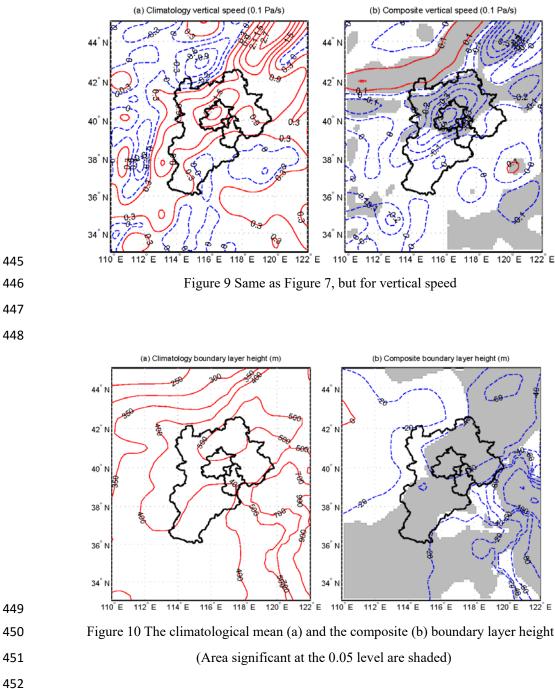
In view of the responses of the local surface winds, relative humidity, vertical motion and boundary layer to the comprehensive index of I_c mentioned above, the close relationships between the winter mean visibility and number of hazy days over BTH region and the given six atmospheric circulations are generally feasible in the physical mechanism. It is reasonable and reliable to estimate the winter haze pollutions in the BTH region based on the seasonal forecast fields derived from climate simulation. Thus it will be helpful to provide scientific references for the governmental decisions in advance about the reducing or controlling of pollutants emission to deal with the probably severe haze pollutions in the BTH region.



hPa (Area significant at the 0.05 level are shaded)







+JZ

453 **4 Conclusions**

Using the daily visibility and number of hazy days recorded in the 19 meteorological stations and the NCEP/NCAR and ERA-Interim reanalysis data, the evolution of the winter haze pollutions in the BTH region since 1981 and their possible relations to atmospheric circulations were examined in this study.

The results showed that the winter mean visibility has a significant negative correlation with the number of hazy days and both of them show distinctly inter-annual

variability during the entire period 1981 to 2015. The correlation coefficients between 460 the winter haze pollutions (the visibility and number of hazy days) and the most 461 common atmospheric circulations over the mid-high latitude of northern hemisphere 462 were re-examined. Results showed that the relations between the haze pollutions in 463 464 BTH and the winter AO, NAO and PNA were very weak, but they correlated 465 significantly with EU, WP and SBH. Furthermore, the six new indices (I1 to I6) derived from the key areas in the fields of SLP, U&V850, H500, U200 and T200 were closely 466 related to the winter haze pollutions in BTH. We can estimate the visibility and number 467 of hazy days by using the six indices and the fitting and the leave-N-out cross-validation 468 methods, respectively. In general, the high level of the estimation statistics suggested 469 470 the winter haze pollutions in BTH can be estimated or predicted in a reasonable degree based on the optimized atmospheric circulation indices. However, we also noted that 471 the statistic estimation models for the visibility and number of hazy days may be 472 influenced in part by a prominent change of the pollutants emission artificially. Thus, it 473 is valuable and significant for government decision-making departments to take actions 474 475 in advance in dealing with the probably severe haze pollutions in BTH indicated by the circulation conditions, such as to control the pollutants discharge. 476

In order to investigate the link processes between the haze pollutions and the given 477 atmospheric circulations more simply, a comprehensive index (I_c) was synthesized from 478 479 the six individual circulation indices by applying a PCA method. The winter Ic increase 480 appear to cause a shallowing of the East Asian trough at the middle troposphere and a weakening of the Siberian high pressure field at sea level, and then accompanied by a 481 482 reduction (increase) of horizontal advection and vertical convection (relative humidity) 483 in the lowest troposphere and a reduced boundary layer height in BTH and its neighboring areas, which are not conducive to the spread and elimination of air 484 485 pollutants but favor the formation of haze pollutions in BTH winter. In short, the reasonable link processes and the stable statistic relationships suggested that the 486 atmospheric circulation indices can be used to predict or evaluate generally the haze 487 pollutions in BTH winter to some extent. 488

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