1	Possible influence of atmospheric circulations on winter hazy						
2	pollution in Beijing-Tianjin-Hebei region, northern China						
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4	Ziyin Zhang ¹ *, Xiaoling Zhang ¹ , Daoyi Gong ² , Seong-Joong Kim ³ , Rui Mao ² ,						
5	Xiujuan Zhao ¹						
6	¹ Environmental Meteorology Forecast Center of Beijing-Tianjin-Hebei, Chinese						
7	Meteorological Administration, Beijing 100089, China						
8	² State Key Laboratory of Earth Surface Processes and Resource Ecology, Beijing						
9	Normal University, Beijing 100875, China						
10	³ Korea Polar Research Institute, Incheon 406-840, Korea						
11							

12 Abstract:

Using the daily records derived from the synoptic weather stations and the 13 14 NCEP/NCAR and ERA-Interim reanalysis data, the variability of the winter hazy pollutions (indicated by the mean visibility and number of hazy days) in Beijing-15 Tianjin-Hebei (BTH) region during the period 1981 to 2015 and its relationship to the 16 atmospheric circulations in middle-high latitude were analyzed in this study. The winter 17 hazy pollution in BTH had distinct inter-annual and inter-decadal variabilities without 18 19 a significant long-term trend. According to the spatial distribution of correlation coefficients, six atmospheric circulation indices (I₁ to I₆) were defined from the key 20 areas in sea level pressure (SLP), zonal and meridional winds at 850 hPa (U850, V850), 21 geopotential height field at 500 hPa (H500), zonal wind at 200 hPa (U200), and air 22 temperature at 200 hPa (T200), respectively. All of the six indices have significant and 23 stable correlations with the winter visibility and number of hazy days in BTH. In the 24 raw (unfiltered) correlations, the correlation coefficients between the six indices and 25 the winter visibility (number of hazy days) varied from 0.57 (0.47) to 0.76 (0.6) with 26 an average of 0.65 (0.54); in the high-frequency (<10 yr) correlations, the coefficients 27 varied from 0.62 (0.58) to 0.8 (0.69) with an average of 0.69 (0.64). The six circulation 28 indices together can explain 77.7% (78.7%) and 61.7% (69.1%) variances of the winter 29 30 visibility and number of hazy days in the year-to-year (inter-annual) variability, respectively. The increase of I_c (a comprehensive index derived from the six individual 31

^{*} Correspondence to: Ziyin Zhang, Environmental Meteorology Forecast Center of Beijing-Tianjin-Hebei, Chinese Meteorological Administration, Beijing 100089, China. E-mail: zzy_ahgeo@163.com

circulation indices) can cause a shallowing of the East Asian trough at the middle 32 troposphere and a weakening of the Siberian high pressure field at sea level, and then 33 accompanied by a reduction (increase) of horizontal advection and vertical convection 34 (relative humidity) in the lowest troposphere and a reduced boundary layer height in 35 BTH and its neighboring areas, which are favorable for the formation of hazy pollutions 36 37 in BTH winter, and vice versa. The high level of the prediction statistics and the reasonable mechanism suggested that the winter hazy pollutions in BTH can be 38 forecasted or estimated credibly based on the optimized atmospheric circulation indices. 39 However, we also noted that the statistic estimation models would be largely influenced 40 by the artificial control of a pollutant discharge. Thus it is helpful for government 41 42 decision-making departments to take actions in advance in dealing with probably severe hazy pollutions in BTH indicated by the atmospheric circulation conditions. 43

44 **Key word**: hazy pollution, visibility, atmospheric circulation, Beijing-Tianjin-Hebei,

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46 **1 Introduction**

Beijing-Tianjin-Hebei (BTH) region is located in northern China, with approximately 47 110 million residents and 216,000 km² in size. As the rapid progress of urbanization 48 and industrial development over the past three decades, the BTH region has become 49 one of China's most economically developed regions and the third economic engine in 50 China. Recently, the Chinese government has been promoting the integration of the 51 three neighboring regions to optimize the industrial layout and improve the allocation 52 of resources. Undoubtedly, the BTH region is becoming more and more important in 53 China or even the world economy in the future. However, the rapid economic growth 54 and urbanization have increased the level of air pollution in recent decades (Streets et 55 56 al., 2007; Chan and Yao, 2008; Wang et al., 2009; Wang et al., 2010; Gao et al., 2011). Most of eastern China has frequently suffered from severe haze or smog days in recent 57 years, especially in the BTH region. For example, the continuously hazy pollutions in 58 January 2013 greatly threatened human health and traffic safety (Kang et al., 2013; 59 Wang et al., 2013). Roughly speaking, the hazy pollution can be attributed to two 60 61 aspects: pollutant emissions to the lower atmosphere from fossil fuel combustion or construction and favorable meteorological conditions. Meteorological conditions are 62 controlling the occurrence of hazy pollution (Wu, 2012; Zhang et al., 2013). 63 Specifically, weather conditions play an essential role in the daily fluctuation of air 64 pollutant concentrations (Zhang et al., 2015). 65

At present, many studies have focused on the physical and chemical properties of 66 pollutants in Beijing and other cities (Feng et al., 2006; Yu et al., 2011; Xu et al., 2013; 67 Zhao et al., 2013). And also studies demonstrated the influence of weather conditions 68 or synoptic situations upon air pollutions (Zhao et al., 2009; Zhang et al., 2015). They 69 elucidated clearly the formation and chemical composition of air pollutants and the 70 71 dominant meteorological factors on hazy days or during heavy pollution in Beijing and its neighboring areas. On the other hand, some studies demonstrated that the hazy 72 pollution occurring in the BTH region could be strongly affected by the local 73 atmospheric circulations including sea-land and mountain-valley breeze circulations 74 and the planetary boundary layer height (Lo et al., 2006; Liu et al., 2009; Chen et al., 75 76 2009; Miao et al., 2015). Recently, Wang et al. (2015) suggested that the reduction of autumn Arctic sea ice leads to anomalous atmospheric circulation changes which favor 77 less cyclone activity and more stable atmosphere and leading to more hazy days in 78 eastern China. Moreover, Wang et al. (2013) showed that east China suffered from 79 severe hazy pollutions in January 2013 may be due to a sudden stratospheric warming 80 81 over the mid-high latitude of Northern Hemisphere, which lead to an anomalous steady 82 atmosphere dominated in northern China. Thus, it is interesting to examine whether the winter hazy pollution in BTH has been influenced by other known or unknown 83 atmospheric circulations or teleconnections in the mid-high latitude of the Northern 84 85 Hemisphere and whether there are some potential circulations that can be used for the 86 forecast or evaluation of the winter hazy pollution in BTH. To date, it is not clear about these questions, and a few studies have been performed to explore these issues. 87

Owing to a lack of long-term instrumental records for air pollutant concentration, 88 the understanding of the evolution of air pollution and their relations to atmospheric 89 circulations is limited. In this paper, we intend to use the atmospheric visibility and the 90 91 number of hazy days derived from the synoptic meteorological stations to denote the evolution of hazy pollution in the BTH region since 1980s. Many studies demonstrated 92 that, in the absence of certain weather conditions (e.g., rain, fog, dust and snowstorm), 93 the visibility is an excellent indicator of air quality because its degradation results from 94 light scattering and absorption by atmospheric particles and gases that can originate 95 96 from natural or anthropogenic sources (Baumer et al., 2008; Chang et al., 2009; Sabetghadam et al., 2012; Baddock et al., 2014), although visibility was influenced 97 comprehensively by airborne pollutants and meteorological parameters such as relative 98 humidity, wind speed, temperature, pressure and solar radiation (Wen and Yeh, 2010; 99 Deng et al., 2014; Zhang et al., 2015). 100

The main purpose of this study is to examine the possible relations between the 101 atmospheric circulations and the winter hazy pollution (the mean visibility and mean 102 number of hazy days) over the BTH region and investigate the possible physical 103 mechanism, which could be useful for a prediction of the winter hazy pollution and 104 could provide a scientific support to the government to take effective measures in 105 106 advance to reduce or control the pollutant emission in case of an anomalous circulations leading to a serious hazy pollution in the region. This paper is organized as follows. 107 Section 2 describes the data and method used. Section 3 shows major results and 108 discussions. Conclusion is summarized in section 4. 109

110 **2 Data and methods**

111 **2.1 Research area and station data**

The atmospheric visibility recorded at the 19 synoptic meteorological stations 112 located in the research area from 1 January 1980 to 28 February 2015 were used (Figure 113 1). The visibility by human observers is recorded by four times (02:00, 08:00, 14:00 114 and 20:00, Beijing local time) or three times (08:00, 14:00 and 20:00, Beijing local time) 115 per day. A good continuous monitoring operation was maintained throughout the entire 116 period, with the missing data rates for the 19 stations varying from a minimum of 1.7% 117 to a maximum of 2.1%, with a mean 1.9%. On the other hand, the distribution of the 118 stations is relatively uniform, indicating that the mean visibility or hazy days is a good 119 representative for the whole BTH region. 120

In the present study, the days with visibility ≤ 5 km and relative humidity < 90%at 14:00PM (local time) were defined as hazy days, except the special weather phenomena occurred at this moment including rain, fog, dust and snow (Schichtel et al., 2010; Wu et al., 2014;). The mean number of hazy days (*NHD*) of each winter in the

125 BTH region can be calculated by:

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$$\overline{NHD} = \frac{1}{n} \sum_{i=1}^{n} N_i$$
(1)

where *n* is the number of stations (here *n*=19), *N* denotes the number of hazy days in a station in each winter (December, January and February). The mean visibility (\overline{Vis}) of each winter in the BTH region can be calculated by:

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$$\overline{Vis} = \frac{1}{n} \sum_{i=1}^{n} \left(\frac{1}{m} \sum_{j=1}^{m} V_{ij} \right)$$
 (2)

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where n is the number of stations (here n=19), m is the number of valid days in

winter. It should be noted that the winter in 1981 consists of December 1980, January







Figure 1 Research area and locations of the 19 synoptic meteorological stations

136 2.2 Reanalysis data

The global NCEP/NCAR reanalysis data of the monthly sea level pressure (SLP), 137 zonal and meridional winds at 850 hPa (U850, V850), geopotential height field at 500 138 hPa (H500), zonal wind at 200 hPa (U200) and air temperature at 200, 150, 100 and 70 139 hPa (T200, T150, T100, T70) with a 2.5°×2.5° spatial resolution from January 1980 to 140 February 2015 were used (Kalnay et al., 1996). Moreover, in order to obtain a higher 141 spatial resolution in the BTH region, the ERA-Interim reanalysis data of the monthly 142 143 relatively humidity (RH), vertical speed (W), zonal (U) and meridional (V) winds from 1000 to 500 hPa (16 pressure levels in total) and the boundary layer height (BLH) with 144 a 0.125°×0.125° spatial resolution confined to the area 33-45°N and 110-122°E were 145 also used (Dee et al., 2011). 146

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148 **2.3 Analysis method**

For the statistical and atmospheric circulation analyses carried out in the study, the 149 common statistical methods such as the composite analyses and the least square 150 regression and the Pearson correlation analyses with a two-tailed Student's t-test were 151 applied in this research. A principal component analysis (PCA) was also used to extract 152 the principal mode of multiple time series. Moreover, in order to reduce the possible 153 effects of low-frequency variation or long-term trends and to examine whether or not 154 the correspondence between the two time series on inter-annual time-scale is stable, the 155 high-frequency (< 10yr) correlation of the high-pass filtered time series was also tested 156 for time series analyses (Gong and Luterbacher, 2008; Zhang et al., 2010). 157

159 **3 Results and discussions**

160 3.1 Evolution of the winter visibility and hazy days in the BTH region

The regional mean visibility and number of hazy days in winter in BTH were 161 presented in Figure 2. As expected, the visibility was negatively correlated to the 162 163 number of hazy days with the raw and high-frequency (< 10yr) correlation coefficients between them of -0.91 and -0.93, respectively. Both of them are significant at the 0.01 164 level (p < 0.01 for short). More hazy days generally denote lower mean visibility in 165 winter due to the light scattering and absorption effects of air pollutants (Baumer et al., 166 2008; Sabetghadam et al., 2012). There are intense inter-annual fluctuations in both the 167 168 visibility and the number of hazy days over the entire period of 1981 to 2015. The decadal fluctuations can be also distinguished for both the visibility and the number of 169 hazy days throughout the entire period. A significant reducing trend of visibility 170 (p<0.05) and increasing trend of number of hazy days (p<0.01) dominated in the 1980s. 171 And then, the visibility experienced an increasing trend in 1990s and a decreasing trend 172 since 2001, and the hazy days showed an anti-phase changes, but none of them are 173 statistically significant with exception of the number of hazy days trend in 1990s 174 (p<0.05). The mean visibility maximum in 1990s reached to 18.3 km (larger than the 175 mean values of 17.9 km over the entire period); and the minimum number of hazy days 176 in 1990s reached to 20.6 days (less than the mean values of 22.7 days over the entire 177 period). However, the long-term trends of them are not statistically significant, although 178 a weak reducing and increasing trends can be founded in the curves of winter visibility 179 and number of hazy days, respectively. 180

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Figure 2 Curves of the winter mean visibility and number of hazy days in BTH

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185 **3.2** Relationship between hazy pollution and atmospheric circulations

We first examined the correlation coefficients between the visibility and number of 186 hazy days and the most common atmospheric teleconnection or oscillation indices over 187 the mid-high latitude of Northern hemisphere (see Table 1), which could affect the 188 winter climate variability over China, such as the Arctic Oscillation (AO), the Northern 189 190 Atlantic Oscillation (NAO), the Pacific/North American pattern (PNA), the Eurasian pattern (EU), the Western Pacific pattern (WP) and the Siberian High (SBH) (Wallace 191 and Gutzler, 1981; Zhang et al., 2009; Gong and Ho, 2012). It can be seen that both of 192 the raw (r1) and high-frequency (r2) correlations show that the visibility and number 193 of hazy days are correlated weakly with the winter AO, NAO and PNA. However, the 194 195 visibility is highly positively correlated with EU, WP and SBH; and the number of hazy days is highly negatively correlated with EU, WP and SBH, most of them are significant 196 at the 0.01 or 0.05 level. 197

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Table 1 Correlation coefficients of visibility and hazy days and circulation indices

		AO	NAO	PNA	EU	WP	SBH
Visibility	<i>r</i> 1	-0.11	0.00	0.16	0.61**	0.40^{*}	0.39*
	<i>r</i> 2	0.05	0.22	0.16	0.71**	0.37*	0.36 [*]
Number of	<i>r</i> 1	0.13	0.13	-0.10	-0.51**	-0.47**	-0.32
hazy days	<i>r</i> 2	-0.01	-0.11	-0.10	-0.70**	-0.56**	-0.37*

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** Significant at the 0.01 level, * Significant at the 0.05 level. The *r*1 and *r*2 terms indicate the raw correlation and high-frequency (< 10yr) correlation, respectively.

Furthermore, the general characteristics of spatial distribution of the correlation 202 coefficients between visibility and number of hazy days in BTH and the major 203 meteorological fields from surface to tropopause in Northern Hemisphere including 204 SLP, U850, V850, H500, U200, T200, T150 and T70 were also examined (Figure 3 and 205 206 4). Owning to a generally anti-pattern for the number of hazy days, thus only the correlation maps with visibility were analyzed for simplicity. In SLP (Figure 3a), a 207 positive correlation center dominated most of East Asian continent, while a negative 208 correlation center dominated the area from northeast Asia to northwest Pacific, 209 respectively. This spatial pattern may reflect the effects of land-sea thermal contrast on 210 211 the lower troposphere condition over BTH region. The pressure increasing in East Asian continent and decreasing in area from northeast Asia to northwest Pacific suggest 212 that they favor the visibility increase in the BTH region in winter, and vice versa. In 213 UV850 (Figure 3b), an anomalously anti-cyclonic and northerly pattern are 214

predominant most of Siberia and eastern China. This suggests that an anomalous 215 northerly advection from Siberian to eastern China improve the winter visibility in the 216 BTH region. In H500 (Figure 3c), there exist a "-+-" wave train pattern along the 217 Eurasia-west Pacific in the mid-high latitude, extending from the central-eastern 218 Europe through Siberia to north China-Korean peninsula-Japan-northwest Pacific 219 220 Ocean, similar to the EU pattern (Wallace and Gutzler, 1981). This pattern implies that a deepening of East Asian trough and a weakening of blocking will favor the winter 221 visibility increase in the BTH region. In U200 (Figure 3d), there also exist a wave train 222 pattern from northwest Russia through Siberia to northwest Pacific Ocean. This pattern 223 may imply that the south (north) of East Asian Jet stream strengthened (weakened) 224 225 coincided with the anomalous ascending (sinking) motions occurred in the south (north) of the Jet stream entrance at the upper troposphere, which will lead to a strengthening 226 northerly appeared in the lower troposphere. Hence it is not conducive to the 227 accumulation of pollutants over BTH region in the winter. 228





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Figure 3, Spatial distribution of correlation coefficients between visibility and SLP (a), UV850
(b), H500 (c) and U200 (d) (Area significant at the 0.05 level are shaded; either U850 or V850 significant at the 0.05 level are shaded in b)

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Besides the lower troposphere, previous studies suggested that the anomalous stratospheric warming over the Northern Hemisphere led to the severe hazy pollutions in east China in January 2013 (Wang et al., 2013). Here, the spatial distribution of the correlation coefficients between visibility and the temperature from the upper troposphere to lower stratosphere at 200 hPa (T200), 150 hPa (T150), 100 hPa (T100) and 70 hPa (T70) were checked. Negative correlations are found from eastern Siberia
to the northern North Pacific including Alaska in T200, T150, T100 and T70,
respectively (Figure 4), with the biggest correlation in T200 (Figure 4a). The
significantly negative correlation suggest that the warming at 200 hPa over eastern
Siberia to the northern North Pacific would indicate a decreasing of winter visibility,
namely a worsening of hazy pollutions in the BTH region.





Figure 4 Spatial distribution of correlation coefficients between visibility and T200 (a), T150 (b),
 T100 (c) and T70 (d) (Area significant at the 0.05 level are shaded)

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Based on the above analyses, we wonder whether the meteorological variables in 251 the significant correlation areas can be used to predict or evaluate the variability of the 252 winter visibility and hazy pollutions in the BTH region. Thus, the six indices for 253 atmospheric circulations or teleconnections were defined based on the key regions 254 255 shown in the previous correlation maps as listed in Table 2. We computed the raw and high-frequency correlation coefficients of the winter visibility and number of hazy days 256 in BTH and the six atmospheric circulation indices. All of the six indices (I_1 to I_6) show 257 highly positive or negative correlations with the winter visibility and number of hazy 258 days, with significance at the 0.01 level (Table3). Moreover, we note that most of the 259 260 high-frequency correlations are larger than the raw correlations except the correlations between visibility and I₁. This suggests that the links between the air quality in BTH 261 and the circulations indices are very stable from year to year. The significantly positive 262 or negative correlations should be a reflection of the physical response mechanisms 263 between them, which will be discussed in the latter section. 264

Table 2 List of the definition for the six circulation indices

Index	Variable	Expression
I ₁	SLP	SLP (38~50N, 84~108E) - SLP (36~52N, 126~150E; 24~40N,150~184E)
I ₂	U _{850hPa}	U ₈₅₀ (55~75N, 40~110E) –U ₈₅₀ (40~50N, 45~75E)
I ₃	V _{850hPa}	V ₈₅₀ (32~64N, 104~120E)
I4	H _{500hPa}	H ₅₀₀ (46~64N, 50~92E) -H ₅₀₀ (28~44N, 16~28E; 28~42N, 120~156E)
I ₅	U _{200hPa}	U ₂₀₀ (42~52N,60~110E) –U ₂₀₀ (64~76N,50~96E; 28~36N, 120~152E)
I ₆	T _{200hPa}	T ₂₀₀ (46~66N, 146~196E)

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Table 3 Correlation coefficients of visibility and number of hazy days and circulation indices

		I_1	I_2	I_3	I_4	I_5	I_6
Visibility	rl	0.73**	0.57**	-0.76**	0.62**	-0.59**	-0.61**
	r2	0.70**	0.68**	-0.80**	0.72**	-0.62**	-0.62**
Number of	<i>r</i> 1	-0.60**	-0.47**	0.60^{**}	-0.47**	0.52**	0.60**
hazy days	r2	-0.61**	-0.65**	0.69**	-0.67**	0.58**	0.64**
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Same as Table 1

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3.3 Predictions for visibility and number of hazy days based on the circulation indices

In order to assess the prediction capability of the six circulation indices for the 274 winter hazy pollutions in BTH, the winter mean visibility and number of hazy days 275 were estimated by applying a multivariate regression method with the least square 276 estimate. The estimated curves by the fitting and the cross-validation with a leave-one-277 out method were displayed in Figure 5. Intuitively, both of the fitting curves and the 278 279 cross-validation curves are fairly consistent with the observed winter mean visibility and number of hazy days over the last three decades. The raw and high-frequency 280 correlation coefficients between the observed and the fitting visibility (number of hazy 281 days) are 0.88 (0.78) and 0.86 (0.77), respectively. All of them are significantly at 0.01 282 level. The six circulation indices together can explain 77.7% (78.7%) and 61.7% 283 284 (69.1%) variances of the winter visibility and number of hazy days over the BTH region in the year to year (inter-annual) variability, respectively. A good fitting does not mean 285 that there must be stable relationships between the dependent variable and explanatory 286 variables. Thus we emphasized testing the stability of the statistic models by means of 287

- the Leave-N-out cross-validations. The statistics for the cross-validation estimations were listed in Table 4, including the explained variance (r^2) , the standard error (SE), and reduction of error (RE). Previous studies suggested that RE is an extremely rigorous verification statistic because it has no lower bound, RE > 0 indicating the skillful estimation, RE > 0.2 indicating the reliable estimation and RE = 1.0 indicating a perfect estimation (Fritts, 1976; Gong and Luterbacher, 2008; Zhang et al., 2010).
- The statistics for both the visibility and number of hazy days are generally stable 294 (no sharply increase or decrease) when N increased from 1 to 11 (more than 30% sample 295 removed in regression models), although the r^2 and RE (SE) slightly decreased 296 (increased) with the increasing of N. For the visibility, the r^2 varied from 52.5% to 62.7% 297 with an average of 57.6%, the SE varied from 0.74 to 0.84 with an average of 0.79, the 298 RE varied from 0.49 to 0.61 with an average of 0.55. For the number of hazy days, the 299 r^2 varied from 31.1% to 41.5% with an average of 35.2%, the SE varied from 3.37 to 300 3.66 with an average of 3.54, the RE varied from 0.23 to 0.38 with an average of 0.30. 301 The mildly changes of these statistics suggest that the statistic models between the given 302 atmospheric circulations and the hazy pollution indicators are stably even in the case of 303 304 parts of sample missed. On the other hand, we noted the statistics for the visibility estimations are generally better than that for the number of hazy days estimations in all 305 tests. However, the minimum values of r^2 and RE for the number of hazy days 306 estimations are still lager than 30% and 0.2, respectively. Based on these statistics, it 307 308 can be concluded that the predictions for the winter visibility and number of hazy days 309 in the BTH region based on the circulation indices are overall reliable during the entire period, especially for the mean visibility. That is to say, the winter hazy pollutions in 310 BTH can be evaluated or estimated well by the optimized atmospheric circulations. 311

The relatively larger errors for the estimated values referred to the observed 312 313 visibility and number of hazy days have been found since the winter in 2009 (Figure 5). We re-computed all the statistics for the period 1981 to 2008, the results displayed that 314 all the values of r^2 and RE (SE) for visibility and number of hazy days predictions 315 increased (decreased) much more than the entire period (Table 4), suggesting that the 316 statistic estimation models are much more stable and reliable before 2009. Why did the 317 prediction efficiency of the statistic estimation models decrease in the last few years? 318 It can be distinguished that the estimations for the winter mean visibility are distinctly 319 lower (higher) than the observed in the winters of 2009 and 2010 (2014), and vice versa 320 for the number of hazy days. We speculated that these phenomena can be attributed to 321 the fluctuations of pollutant emissions in part because the pollutant emissions over 322

northern China around 2008 were controlled strictly by the Chinese government 323 associated with the 2008 Olympic Games in Beijing (An et al., 2007; Zhang et al., 2010; 324 Gao et al., 2011). The decrease of pollutant emissions led to the improvement of air 325 quality (increasing visibility and decreasing hazy days) in 2009 and 2010, although the 326 atmospheric conditions remained the same and did not contributed to the spread and 327 328 elimination of air pollutants. However, pollutant emissions especially in the areas of BTH rebounded after the Olympic Games, with the decrease in visibility and increase 329 in hazy days in the BTH region around 2012 to 2014 to some extent (Zhang et al., 2015), 330 although the atmospheric conditions remained relatively the same as before. From this 331 result, it can be assumed the statistic estimation models for the winter mean visibility 332 333 and number of hazy days would be largely influenced by an artificial control of pollutant discharge. 334

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Figure 5 Curves of the observed and the predicted winter visibility (a) and number of hazy days
(b) in the BTH region since 1981

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Table 4 List of the statistics for the Leave-N-out cross-validation estimations

N	Period covering	Visibility			Number of hazy days		
		r^{2} (%)	SE	RE	r^{2} (%)	SE	RE
1	1981-2015	62.7	0.74	0.61	41.5	3.37	0.38
	1981-2008	87.1	0.42	0.87	53.9	2.56	0.52
3	1981-2015	56.8	0.80	0.54	34.3	3.57	0.28
	1981-2008	86.8	0.42	0.87	52.6	2.59	0.51
5	1981-2015	59.2	0.78	0.57	35.3	3.54	0.30

	1981-2008	86.8	0.42	0.87	46.7	2.75	0.43
_	1981-2015	59.0	0.78	0.56	37.5	3.48	0.33
/	1981-2008	86.4	0.43	0.86	44.7	2.80	0.41
9	1981-2015	56.2	0.80	0.54	32.5	3.62	0.27
	1981-2008	84.2	0.46	0.84	40.8	2.90	0.36
11	1981-2015	52.5	0.84	0.49	31.1	3.66	0.23
	1981-2008	84.4	0.46	0.84	48.2	2.71	0.44

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3.4 Possible mechanism of the circulations related to the winter hazy pollutions

(N denotes the number of sample removed in the cross-validation regressions; only the odd

numbers of N were listed for short)

In order to explore the possible mechanism and role of the investigated circulation 345 indices on the winter visibility and number of hazy days in the BTH region, the links 346 between the given large-scale atmospheric circulations and the local meteorological 347 conditions, which have close relations with the hazy pollutions, were examined. For 348 349 simplicity, a comprehensive index labeled as I_c was synthesized from the six individual 350 circulation indices (I₁ to I₆) by applying a PCA method, namely the first principal component (PC1). The high values of the explained variance (64.4% in PC1) indicated 351 that the comprehensive index of I_c roughly reflect the integrated features of all the six 352 indices. Thus, we used the Ic instead of the six individual indices in the following 353 analysis. Generally, the positive (negative) Ic indicate the lower (higher) visibility and 354 more (less) hazy days in the BTH region in winter. 355

First we examined the links between the Ic and the meteorological fields of SLP 356 and H500 respectively. Based on the NCEP/NCAR reanalysis data, Figure 6(a) and (b) 357 present the climatological mean of SLP and H500 in winter averaged from 1981 to 2010, 358 359 respectively. The changes of SLP and H500 in winter in association with a onestandard-deviation positive Ic during the winters 1981 to 2015 are shown in Figure 6(c) 360 and (d), respectively. In the climatological mean fields, the BTH region were located in 361 the trough of East Asian trough at the middle troposphere and in the ridge of Siberian-362 Mongolia high in SLP field, which indicate the northerly dominated the BTH region in 363 364 winter. The regression maps show that the SLP decreased in the Siberian-Mongolia high areas and increased in the western Pacific in SLP and the geopotential height decreased 365 in the most areas of Siberia and increased in the northern China to western Pacific. 366 These patterns suggest that both the East Asian trough and Siberian high weaken with 367 increasing Ic, that further implies that the winter cold air activity will be weaken and 368

then lead to an anomalous steady atmospheric conditions in BTH and its adjacent areas 369 in winter. Namely, the less strong Siberian high and East Asian trough and associated 370 northerly winds in the low and middle troposphere will lead to a severe hazy pollution 371 (lower visibility and more hazy days) due to the favorable meteorological conditions 372 for the accumulation and chemical reaction of pollutants. Anyway, we wonder whether 373 374 it is true as we speculated. We further examined the links between the comprehensive index of I_c and the local meteorological conditions which play direct roles in the 375 formation of hazy pollutions, including the wind fields (Figure 7), relative humidity 376 (Figure 8) and vertical velocity (Figure 9) at the lowest troposphere (averaged from 377 1000 hPa to 900 hPa with an interval of 25 hPa) and the boundary layer height (Figure 378 379 10) based on the ERA-Interim reanalysis data.

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Figure 6 The climatological mean fields of SLP (a) and H500 (b) averaged in winter 1981 to 2010,
and the spatial distribution of the regression coefficients of SLP (c) and H500 (d) upon the I_c over
the period 1981 to 2015 (Area significant at the 0.05 level are shaded)

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Figure 7(a) displays the climatological mean wind field averaged from 1000 to 900 hPa over the winter 1981 to 2010. At lower level, the northwesterly winds dominated the BTH, and the wind speed in Beijing, Tianjin and north of Hebei province was larger than that in the south of Hebei province. Figure 7(b) shows the composite (positive Ic winters minus negative Ic winters) wind field averaged from 1000 to 900 hPa over the winter 1981 to 2015. In the composite wind field, the anomalous southeasterly winds dominated the BTH region instead of the northwesterly in the climatological mean wind field, indicating the weakening of the northwesterly significantly over BTH and its neighboring areas when I_c increased. Previous studies (Zhang et al., 2015) demonstrated the decreasing of wind speed is not conducive to the diffusion of air pollutants and easily lead to hazy pollutions in Beijing. It may be true for the whole BTH region. Thus, the increasing of I_c will lead to a decrease in the visibility and increase in the number of hazy days in winter over the BTH region.

Same as Figure 7, Figure 8(a) and (b) present the climatology and composite 400 relative humidity averaged from the lowest troposphere respectively. In the composite 401 map, all the areas of BTH are covered by the positive values and most of them are 402 403 significant at the 0.05 level. They indicate that the winter relative humidity was anomalous higher in the positive I_c years than that in the negative I_c years. As pointed 404 in the Introduction, a high relative humidity is one of the important reasons for visibility 405 degradation. This is because that the high relative humidity is favorable for the accumulation and 406 407 hygroscopic growth of pollutants, which can strengthen the light scattering and absorption by 408 atmospheric particles and gases and then cause the visibility degradation directly (Baumer et al., 409 2008; Zhang et al., 2015). Thus a positive Ic imply that a decreasing of visibility accompanied by the increasing number of hazy days may occur in the winter over BTH 410 region. Figure 9(a) and (b) present the climatology and composite vertical speeds 411 412 averaged from the lowest troposphere respectively. The positive (negative) values of 413 vertical speed in unit of Pa/s denote sinking (ascending) motion. The climatological vertical speeds show that the downward air motions dominated the BTH region in the 414 winter. In the composite vertical speed field, the most areas of BTH were covered by 415 the significantly negative values, which suggested a less vertical exchanges of air 416 occurred in this areas in the positive I_c winters. In other words, the increased I_c may 417 418 result in a weaker vertical convection and forcing the lowest troposphere more stable. It's easy to understand the anomalous stabilization will lead to much hazy pollutions. 419 Moreover, a similar result can be found in the planetary boundary layer height, which 420 was reduced significantly in the most of BTH and its adjacent areas in the positive Ic 421 winters (Figure 10). The decreased boundary layer height will depress the air pollutants 422 423 into a narrower air column in a certain area and then lead to an increasing of the pollutants concentration. Thus, a winter with the lower visibility and more hazy days in 424 the BTH region would be expected in the case of the lower boundary layer height 425 caused by the anomalously high Ic. 426

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In view of the responses of the local surface winds, relative humidity, vertical

motion and boundary layer to the comprehensive index of Ic mentioned above, the close 428 relationships between the winter mean visibility and number of hazy days over BTH 429 region and the given six atmospheric circulations are generally feasible in the physical 430 mechanism. It is reasonable and reliable to estimate the winter hazy pollutions in the 431 BTH region based on the seasonal forecast fields derived from climate simulation. Thus 432 433 it will be helpful to provide scientific references for the governmental decisions in advance about the reducing or controlling of pollutants emission to deal with the 434 probably severe hazy pollutions in the BTH region. 435





438 Figure 7 The climatological mean (a) and the composite (b) wind fields averaged from 1000 to 900
439 hPa (Area significant at the 0.05 level are shaded)

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Figure 8 Same as Figure 7, but for relative humidity

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Figure 9 Same as Figure 7, but for vertical speed







Conclusions

Using the daily visibility and number of hazy days recorded in the 19 meteorological stations and the NCEP/NCAR and ERA-Interim reanalysis data, the evolution of the winter hazy pollutions in the BTH region since 1981 and their possible relations to atmospheric circulations were examined in this study.

The results showed that the winter mean visibility has a significantly negative correlation with the number of hazy days and both of them show distinctly inter-annual variability during the entire period 1981 to 2015. The correlation coefficients between the winter hazy pollutions (the visibility and number of hazy days) and the most

common atmospheric circulations over the mid-high latitude of northern hemisphere 460 were re-examined. Results showed that the relations between the hazy pollutions in 461 BTH and the winter AO, NAO and PNA were very weak, but they correlated 462 significantly with EU, WP and SBH. Furthermore, the six new indices (I₁ to I₆) derived 463 from the key areas in the fields of SLP, U&V850, H500, U200 and T200 were closely 464 465 related to the winter hazy pollutions in BTH. We can estimate the visibility and number of hazy days by using the six indices and the fitting and the leave-N-out cross-validation 466 methods, respectively. In general, the high level of the estimation statistics suggested 467 the winter hazy pollutions in BTH can be estimated or predicted in a reasonable degree 468 based on the optimized atmospheric circulation indices. However, we also noted that 469 470 the statistic estimation models for the visibility and number of hazy days may be influenced by a prominent change of the pollutants emission artificially. Thus, it is 471 valuable and significant for government decision-making departments to take actions 472 in advance in dealing with the probably severe hazy pollutions in BTH indicated by the 473 circulation conditions, such as to control the pollutants discharge. 474

475 In order to investigate the link processes between the hazy pollutions and the given atmospheric circulations more simply, a comprehensive index (I_c) was synthesized from 476 the six individual circulation indices by applying a PCA method. The winter Ic increase 477 appear to cause a shallowing of the East Asian trough at the middle troposphere and a 478 479 weakening of the Siberian high pressure field at sea level, and then accompanied by a 480 reduction (increase) of horizontal advection and vertical convection (relative humidity) in the lowest troposphere and a reduced boundary layer height in BTH and its 481 482 neighboring areas, which are not conducive to the spread and elimination of air pollutants but favor the formation of hazy pollutions in BTH winter. In short, the 483 reasonable link processes and the stable statistic relationships suggested that the 484 485 atmospheric circulation indices can be used to predict or evaluate generally the hazy pollutions in BTH winter to some extent. 486

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