



Annual variations of
carbonaceous PM_{2.5}
in Malaysia

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This discussion paper is/has been under review for the journal Atmospheric Chemistry and Physics (ACP). Please refer to the corresponding final paper in ACP if available.

Annual variations of carbonaceous PM_{2.5} in Malaysia: influence by Indonesian peatland fires

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Received: 16 June 2015 – Accepted: 4 August 2015 – Published: 21 August 2015

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Published by Copernicus Publications on behalf of the European Geosciences Union.

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Abstract

In this study, we quantified carbonaceous PM_{2.5} in Malaysia through annual observations of PM_{2.5}, focusing on organic compounds derived from biomass burning. We determined organic carbon (OC), elemental carbon (EC) and concentrations of solvent-extractable organic compounds (biomarkers derived from biomass burning sources and *n*-alkanes). We observed seasonal variations in the concentrations of pyrolyzed OC (OP), levoglucosan (LG), mannosan (MN), galactosan, syringaldehyde, vanillic acid (VA) and cholesterol. The average concentrations of OP, LG, MN, galactosan, VA and cholesterol were higher during the southwest monsoon season (June–September) than during the northeast monsoon season (December–March), and these differences were statistically significant. Conversely, the syringaldehyde concentration during the southwest monsoon season was lower. The PM_{2.5} OP/OC₄ mass ratio allowed distinguishing the seven samples, which have been affected by the Indonesian peatland fires (IPFs). In addition, we observed significant differences in the concentrations between the IPF and other samples of many chemical species. Thus, the chemical characteristics of PM_{2.5} in Malaysia appeared to be significantly influenced by IPFs during the southwest monsoon season. Furthermore, we evaluated two indicators, the vanillic acid/syringic acid (VA/SA) and LG/MN mass ratios, which have been suggested as indicators of IPFs. The LG/MN mass ratio ranged from 14 to 22 in the IPF samples and from 11 to 31 in the other samples. Thus, the respective variation ranges partially overlapped. Consequently, this ratio did not satisfactorily reflect the effects of IPFs in Malaysia. In contrast, the VA/SA mass ratio may serve as a good indicator, since it significantly differed between the IPF and other samples. However, the OP/OC₄ mass ratio provided more remarkable differences than the VA/SA mass ratio, offering an even better indicator. Finally, we extracted biomass burning emissions' sources such as IPF, softwood/hardwood burning and meat cooking through varimax-rotated principal component analysis.

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als) and See et al. (2007) (water-soluble ions, metals and polycyclic aromatic hydrocarbons).

Several studies exist on the chemical characteristics of haze ambient particulates, which have been potentially affected by IPFs in Malaysia and Singapore (e.g., Abas et al., 2004a, b; Betha et al., 2014; Engling et al., 2014; Fang et al., 1999; Fujii et al., 2015b; He et al., 2010; Keywood et al., 2003; Narukawa et al., 1999; Okuda et al., 2002; See et al., 2006; Yang et al., 2013). In most cases, the field observation periods were short. Even when long-term observations have been obtained, however, only typical chemical species such as ions and metals have been analysed. Nevertheless, organic compounds significantly contribute to the IPF aerosols (Fujii et al., 2014). In Malaysia especially, there are no available quantitative data regarding variations of several organic compound concentrations based on long-term observations of PM_{2.5}.

The three major sources of air pollution in Malaysia are mobile, stationary and open burning sources including the burning of solid wastes and forest fires (Afroz et al., 2003). The annual burned biomass in Malaysia has been estimated to be 23 Tg on average (Streets et al., 2003). Therefore, it is necessary to distinguish the effects of IPFs from those of other sources, particularly local biomass burning. Fujii et al. (2015b) reported the total suspended particulate matter (TSP) concentrations in the different carbon fractions (OC1, OC2, OC3, OC4 and pyrolysed OC (OP)) defined by the IMPROVE_A protocol (Chow et al., 2007) in Malaysia during the haze periods affected by IPFs. They proposed the OP/OC4 mass ratio as a useful indicator of transboundary haze pollution from IPFs at receptor sites even in light haze; the ratio during the haze periods were higher (> 4) than during the non-haze periods (< 2).

In the present study, the carbonaceous PM_{2.5} components are quantitatively characterised using annual PM_{2.5} observations in Malaysia, with special regard to the organic compounds resulting from biomass burning. Furthermore, the OP/OC4 mass ratio is used as an indicator to investigate the effects of IPFs on carbonaceous PM_{2.5} species in this area. In addition, other indicators that potentially record the effects of

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Figure 2 presents the monthly hotspot counts in the Sumatra Island detected by the NOAA-18 satellite (Indofire). During the southwest monsoon season on September 2011 and June 2012, hotspots exceeded 3000 on several occasions. The hotspot counts in September 2011 and June 2012 mainly derived from the South Sumatra (60 % of the hotspot counts) and the Riau (42 %) provinces, respectively. The sampling sites are dominantly downwind regions in the Sumatra Island during the southwest monsoon season. Thus, some samples have probably been affected by IPFs. The three-day backward air trajectories for the sampling periods (Fig. S1) support this conclusion.

3.2 PM_{2.5} chemical characteristics and seasonal variations

3.2.1 OC and EC

The annual concentrations of OC and EC are 7.0 ± 5.4 and $3.1 \pm 1.1 \mu\text{gCm}^{-3}$, respectively. The OC and EC concentrations' statistical results for each monsoon season appear in Table 1. The average OC concentration during the southwest monsoon season (June–September) is higher than that during other seasons. In particular, an extremely high OC concentration ($> 25 \mu\text{gCm}^{-3}$) is observed on 12 September 2011 and on 15 and 16 June 2012. There is no statistically significant difference in the EC concentration between the southwest and northeast (December–March) monsoon seasons according to the two-sided Wilcoxon rank sum test (p value: $p > 0.05$) with R-software. In Bangi (~ 30 km southeast of the sampling site), the OC concentration was $11 \pm 3.2 \mu\text{gCm}^{-3}$ in September 2013 (Fujii et al., 2015c), in good agreement with the present results for the southwest monsoon season. The OC/EC mass ratios during the southwest monsoon, post-monsoon (October–November), northeast monsoon and pre-monsoon (April–May) season range among 1.2–6.5, 1.4–2.4, 0.99–3.0 and 1.2–2.3, respectively. A high OC/EC mass ratio value (> 4) is found only for some samples collected on September 2011 and June 2012. These values have probably

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been affected by biomass burning, because aerosols emitted from biomass burning usually present higher OC/EC mass ratios (Cong et al., 2015).

The daily variations of the OC fractions' mass concentrations during the sampling periods are presented in Fig. 3. The annual concentrations of OC1, OC2, OC3, OC4 and OP are 0.51 ± 0.80 , 1.9 ± 1.1 , 2.3 ± 1.4 , 1.2 ± 0.36 and $1.1 \pm 2.2 \mu\text{g m}^{-3}$, respectively. Statistically significant differences among the OP concentrations during the southwest and northeast monsoon seasons are observed according to the two-sided Wilcoxon rank sum test ($p < 0.001$). In particular, high OP concentrations are clearly observed in September 2011 and June 2012, in addition to the higher OC/EC mass ratios described above. Fujii et al. (2015b) supported that the enhanced OP concentrations in TSP, which are observed in Malaysia during the haze periods, are affected by the IPFs. The enhanced OP concentrations in $\text{PM}_{2.5}$ during the southwest monsoon season, which are observed in the present study, are also probably affected by IPFs from the Sumatra Island. The increased number of hotspots recorded (Fig. 2) and backward air trajectories (Fig. S1) further support this conclusion.

3.2.2 Biomarkers

Ten biomarkers are identified in this study, which have been suggested as indicators of biomass burning processes such as wood burning and meat cooking. The annual concentrations of levoglucosan (LG), mannosan (MN), galactosan, *p*-hydroxybenzoic acid, vanillic acid (VA) and syringic acid (SA) are 86 ± 95 , 4.8 ± 5.7 , 1.2 ± 1.6 , 1.1 ± 1.3 , 0.19 ± 0.28 and $0.25 \pm 0.28 \text{ ng m}^{-3}$, respectively; notably, they exhibit great variability. The annual concentrations of vanillin, syringaldehyde, dehydroabietic acid and cholesterol are 1.2 ± 0.80 , 0.51 ± 0.42 , 1.3 ± 1.0 and $1.3 \pm 0.72 \text{ ng m}^{-3}$, respectively. The biomarker statistical results for each monsoon season are listed in Table 1.

LG is a specific indicator for cellulose burning emissions and generally formed during cellulose pyrolysis at temperatures above 300°C (Fujii et al., 2015b; Lin et al., 2010; Shafizadeh, 1984; Simoneit et al., 1999). The MN and galactosan are derived from hemicellulose pyrolysis products; they can also be used as tracers of biomass burn-

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ing besides LG (e.g., Engling et al., 2014; Fujii et al., 2014, 2015b; Zhu et al., 2015). Statistically significant differences are observed among the concentrations of LG, MN and galactosan obtained during the southwest and northeast monsoon seasons on the basis of the two-sided Wilcoxon rank sum test ($p < 0.001$); high concentrations of these compounds are mostly observed during the southwest monsoon season (especially September 2011 and June 2012; Fig. S2). In Singapore, Engling et al. (2014) suggested that the enhanced concentrations of these compounds during the haze periods were due to the IPFs during the southwest monsoon season. Thus, the presently observed enhanced concentrations of these compounds may also be attributed to the IPFs.

In a previous report, PM_{2.5} lignin unit-originating compounds in samples collected at IPF source were quantified (Fujii et al., 2015a). Lignin is an aromatic polymer consisting of phenylpropane units linked through many ether and C–C linkages. Its aromatic structure varies depending on the species; softwood lignins exclusively contain guaiacyl (G) types, hardwood lignins include both G and syringyl (S) types, whereas herbaceous plants include G, S and *p*-hydroxyphenyl (H) types (Fujii et al., 2015a, b). The composition of these aromatic nuclei within the lignin pyrolysis products resulting from biomass burning may be useful in identifying the biomass type (Fujii et al., 2015a; Simoneit et al., 1993). In the present study, vanillin and VA (compounds derived from G units), syringaldehyde and SA (compounds derived from S units) as well as *p*-hydroxybenzoic acid (compounds derived from H units or the secondary decomposition of G and S units) (Fujii et al., 2015b) have been quantified. There are significant differences between the concentrations of syringaldehyde and VA derived from lignin pyrolysis during the southwest and northeast monsoon seasons on the basis of the two-sided Wilcoxon rank sum test ($p < 0.001$), corresponding to seasonal variations. The average VA concentration during the southwest monsoon season is 5.3 times greater than that during the northeast monsoon season. In contrast, the average concentration of syringaldehyde during the northeast monsoon season is 2.6 times greater than that during the southwest monsoon season. This may be due to the transboundary pollution

by prevailing winds from the Chinese region including Thailand and Vietnam during the northeast monsoon season (Fig. S1; Khan et al., 2015).

Dehydroabietic acid and cholesterol are quantified as indicators of softwood burning and meat cooking, respectively (Fujii et al., 2015b; Lin et al., 2010). The two-sided Wilcoxon rank sum test indicates that the difference between the cholesterol concentration during the southwest and northeast monsoon seasons is statistically significant ($p < 0.001$). The dehydroabietic acid and cholesterol concentrations recorded in the interval between June and July 2014 in Bangi, which is located ~ 30 km southeast of the sampling site, range between 2.6–8.7 and 1.5–5.7 ng m^{-3} , respectively (Fujii et al., 2015b). The PJ industrial area's concentrations of these compounds are lower than those in the Bangi suburban area owing to the decreased impact of softwood burning and meat cooking in PJ.

3.2.3 *N*-alkanes

The total annual concentration of *n*-alkanes is $79 \pm 63 \text{ ng m}^{-3}$. The total *n*-alkanes concentration during the southwest monsoon, post-monsoon, northeast monsoon and pre-monsoon season is 110 ± 93 , 57 ± 20 , 67 ± 18 and $55 \pm 41 \text{ ng m}^{-3}$, respectively. The highest concentration is observed during the southwest monsoon season. Figure 4 illustrates the molecular distribution of *n*-alkanes during the southwest and northeast monsoon seasons. There are no significant differences among the concentrations of C_{22} – C_{26} , C_{29} , C_{30} and C_{32} in the two seasons ($p > 0.05$). High concentrations of $> C_{24}$ are mainly observed in September 2011 and June 2012 when many hotspots are detected in the Sumatra Island (Fig. 2). Fujii et al. (2015a) suggested that IPFs increase the C_{27} , C_{28} and C_{29} concentrations in $\text{PM}_{2.5}$ at the receptor site relative to other sources such as vehicle and biomass burning. Thus, the enhanced *n*-alkanes concentrations in $\text{PM}_{2.5}$ during the southwest monsoon season may be mainly attributed to IPFs.

The carbon number maximum (C_{max}) in *n*-alkanes during the southwest and northeast monsoon seasons is C_{27} (in 83% of the samples) and C_{26} (75%), respectively

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(Fig. 5). Reported C_{\max} values range from 27 to 33, characteristic of biogenic sources (higher plant-wax), whereas lower C_{\max} values may indicate major petrogenic input (Abas et al., 2004a; Gogou et al., 1996; He et al., 2010). The C_{\max} during the southwest monsoon season (C_{27}) suggests primarily biogenic sources and is in perfect agreement with the measured value for the IPF source (Fujii et al., 2015b).

The carbon preference index (CPI) has been widely used to roughly estimate the effects of anthropogenic or biogenic sources (e.g., Chen et al., 2014; He et al., 2010). The CPI is defined as the sum of the concentrations of the odd carbon-number n -alkanes divided by that of the even carbon-number n -alkanes. The n -alkanes from terrestrial vegetation typically exhibit high CPI values (> 2), whereas those from anthropogenic sources present CPI values close to one (Chen et al., 2014; He et al., 2010). Here, the CPI values are calculated by the following equation.

$$\text{CPI} = \frac{C_{23} + C_{25} + C_{27} + C_{29} + C_{31} + C_{33}}{C_{22} + C_{24} + C_{26} + C_{28} + C_{30} + C_{32}} \quad (2)$$

The CPI values during the southwest and northeast monsoon seasons are 1.2 ± 0.15 and 0.96 ± 0.12 , respectively; these values are close to one for both seasons, indicating an anthropogenic n -alkane source. Thus, the CPI value is not susceptible to IPF influence, since the CPI value at IPF source is 1.4 ± 0.088 (Fujii et al., 2015a), which is not high. Consequently, the CPI cannot be used to identify IPFs sources at a receptor site.

3.3 Indonesian peatland fire effect

The hotspot data and backward air trajectories suggest that IPFs strongly modify many chemical species concentrations mostly during the southwest monsoon season. However, IPFs do not always occur during the southwest monsoon season. Therefore, significant differences in some chemical species concentrations among samples affected by IPF and others should be observed. To distinguish IPF samples from other samples obtained during the southwest monsoon season, the OP/OC4 mass ratio is used,

which is a useful indicator for IPF (Fujii et al., 2015b). The ratio value is > 4 for seven samples (11–13 September 2011 and 14–17 June 2012); these samples are regarded as the IPF samples. The OP/OC4 mass ratio for the IPF and other samples is 7.4 ± 3.4 and 0.44 ± 0.49 , respectively, exhibiting significant differences among them according to the two-sided Wilcoxon rank sum test ($p < 0.001$). Figure 6 shows the p values used to determine the statistical significance in a hypothesis test of the differences between the IPF and other samples for all the quantified species. Significant differences ($p < 0.001$) are recorded for many chemical species. Thus, the chemical characteristics of PM_{2.5} in Malaysia are significantly influenced by IPFs.

Furthermore, the VA/SA and LG/MN mass ratios in the IPF source are investigated as potential indicators, as suggested in previous studies (Fujii et al., 2014, 2015a). The VA/SA mass ratio for IPF and other samples is 1.7 ± 0.36 and 0.59 ± 0.27 , respectively, providing a good indicator ($p < 0.001$). Although the VA/SA mass ratio at the IPF source is 1.1 ± 0.16 (Fujii et al., 2015a), the ratios for IPF samples are higher. Opsahl and Benner (1998) reported photochemical reactivity of VA and SA in the Mississippi River water. They demonstrated that the early degradation of SA in the water is mostly due to its higher photochemical reactivity compared with VA. Even though there are no reports of such degradations in air, SA is considered to be less stable than VA in air as well as in water, which leads to an increased VA/SA ratio after long-range transportation. On the other hand, the LG/MN mass ratio for the IPF and other samples ranges from 14 to 22 and 11 to 31, respectively (Fig. S3). Therefore, the LG/MN mass ratio is inappropriate to extract the effects of IPF in Malaysia, because its value's ranges in the IPF and other samples partially overlap.

The daily variability of the C₂₇ and LG concentration as well as the VA/SA and OP/OC4 mass ratios are presented in Fig. 7; similar trends are observed in all cases. However, the concentrations of LG, MN and galacotsan (Fig. S2) increase abruptly on 10 August 2011, although this sample is not categorised as an IPF sample. We hypothesised that this increase results from local biomass burning, since LG emissions are produced by several different biomass burning sources (Oros and Simoneit, 2001a, b;

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Oros et al., 2006). Therefore, LG levels are not directly indicative of the IPF contribution in Malaysia; instead, C_{27} may be a useful indicator (Fig. 7). Although the VA/SA mass ratio can be used as an IPF indicator, as we mentioned before, the OP/OC4 mass ratio highlights the differences between the IPF and other samples better than the VA/SA mass ratio (Fig. 7).

3.4 Carbonaceous $PM_{2.5}$ contributions

The possible sources of carbonaceous $PM_{2.5}$ are investigated through varimax-rotated PCA of the PJ_A and PJ_S datasets. Over 80% of the cumulative variance in the PJ_A and PJ_S datasets is explained by three and five factors, respectively (Table 2). For the PJ_A data (Table 2a), the total variance explained by the three factors is 80%. Factor A1, which explains 60% of the variance, is heavily loaded (loading factor: > 0.65) with OC, LG, MN, galactosan, *p*-hydroxybenzoic acid, VA and C_{25} – C_{33} , which direct towards an IPF source. Factor A2, which corresponds to 12% of the variance, is heavily loaded with C_{22} – C_{24} , suggesting a petrogenic source (Abas et al., 2004a; Gogou et al., 1996; He et al., 2010). Factor A3, which explains 8.0% of the variance in the data set, is heavily loaded with SA and dehydroabietic acid, indicating mixed (softwood and hardwood) biomass burning sources. For the PJ_S dataset (Table 2b), the total variance explained by five factors is 82%. Factor S1 explains 43% of the data's variance and is heavily loaded with C_{27} – C_{33} , which suggests tire wear emission (Rogge et al., 1993). Factor S2 explains 19% of the variance and is heavily loaded with LG, MN, galactosan, VA and SA, which correspond to biomass burning source. Factor S3, which explains 11% of the variance, is heavily loaded with C_{22} – C_{26} , which indicate a petrogenic source, similar to factor A2. Although heavy loading with only syringaldehyde is found in factor S4 (5.0% of the variance), its source could not be identified. Finally, factor S5 explains 4.5% of the variance and is heavily loaded with EC and cholesterol, which are produced when cooking meat.

Wahid et al. (2013) reported varimax-rotated PCA results on the distribution of inorganic ions within fine-mode aerosols (< 1.5 μ m) at Kuala Lumpur, which is close to the

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present study's sampling site (~ 10 km). They extracted three principal components from this analysis: (1) motor vehicles, (2) soil and earth's crust and (3) sea spray. Jamhari et al. (2014) applied varimax-rotated PCA on polycyclic aromatic hydrocarbon data in PM₁₀ at Kuala Lumpur. They extracted two factors, which were attributed to (1) natural gas emission and coal combustion and (2) vehicles and gasoline emissions. In the present study, only biomass burning could be identified as a factor through comparison with these previous analyses. Factors such as soil, sea spray and coal combustion could not be identified, because the key inorganic compounds produced from these sources were not determined.

4 Conclusions

Annual PM_{2.5} observations in Malaysia have been conducted to quantitatively characterise carbonaceous PM_{2.5}, especially focusing on organic compounds derived from biomass burning for the first time. The main conclusions are summarised as follows:

Concentrations of OP, LG, MN, galactosan, syringaldehyde, VA and cholesterol exhibit seasonal variability. The average concentrations of OP, LG, MN, galactosan, VA and cholesterol during the southwest monsoon season are higher than those during the northeast monsoon season, and the differences are statistically significant. In contrast, the syringaldehyde concentration during the southwest monsoon season is lower.

Seven IPF samples are distinguished on the basis of the PM_{2.5} OP/OC4 mass ratio. In addition, significant differences are observed for the concentrations of many chemical species between the IPF and other samples. Thus, the PM_{2.5} chemical characteristics in Malaysia are clearly influenced by IPFs during the southwest monsoon season. Furthermore, two previously suggested indicators of IPF sources have been evaluated, the VA/SA and LG/MN mass ratio. The LG/MN mass ratio ranges from 14 to 22 in the IPF samples and from 11 to 31 in the other samples. The two ratio distributions partial overlap. Thus, the LG/MN mass ratio is not considered appropriate for extracting the effects of IPFs in Malaysia. In contrast, significant differences among the VA/SA mass

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ratios in the IPF and other samples suggest that it may serve as a good indicator. However, the OP/OC4 mass ratio differentiates the IPF samples better than VA/SA mass ratio. Consequently, the OP/OC4 mass ratio is proposed as a better indicator than the VA/SA mass ratio. Finally, varimax-rotated PCA enabled to discriminate biomass burning components such as IPFs, softwood/hardwood burning and meat cooking.

The Supplement related to this article is available online at doi:10.5194/acpd-15-22419-2015-supplement.

Acknowledgements. This study was supported by JSPS Kakenhi Grant Number (15H02589, 15J08153).

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Table 1. Statistical results of chemical species concentrations. Av = Average. SD = Standard deviation.

Compounds	Southwest monsoon (Jun–Sep)		Post-monsoon (Oct–Nov)		Northeast monsoon (Dec–Mar)		Pre-monsoon (Apr–May)	
	Av ± SD	Range	Av ± SD	Range	Av ± SD	Range	Av ± SD	Range
OC and EC [$\mu\text{g m}^{-3}$]								
OC	10 ± 7.8	3.6–36	5.6 ± 2.4	2.5–11	5.2 ± 1.4	2.7–8.2	4.2 ± 1.4	2.8–7.3
EC	3.0 ± 0.95	1.0–5.6	3.2 ± 1.3	1.1–5.9	3.4 ± 1.1	1.6–6.1	2.6 ± 1.2	1.4–4.5
Biomarkers [ng m^{-3}]								
levoglucosan	160 ± 130	32–490	64 ± 39	19–130	40 ± 14	17–64	49 ± 21	23–86
mannosan	8.4 ± 8.2	1.5–30	3.4 ± 2.6	0.95–9.1	2.6 ± 1.2	0.84–5.3	2.5 ± 1.2	1.2–5.3
galactosan	2.3 ± 2.3	0.38–8.3	0.86 ± 0.72	0.29–2.8	0.60 ± 0.35	0.13–1.3	0.62 ± 0.34	0.33–1.5
<i>p</i> -hydroxybenzoic acid	1.9 ± 1.9	0.18–7.5	0.79 ± 0.67	0.036–2.2	0.64 ± 0.30	0.20–1.2	0.50 ± 0.25	0.24–1.0
vanillin	1.6 ± 1.1	0.54–5.5	1.2 ± 0.66	0.45–2.2	1.0 ± 0.38	0.21–1.7	0.96 ± 0.42	0.30–1.7
syringaldehyde	0.29 ± 0.22	0.085–1.0	0.59 ± 0.22	0.26–1.2	0.77 ± 0.54	0.074–2.2	0.36 ± 0.22	0.093–0.77
vanillic acid	0.39 ± 0.39	0.074–1.9	0.11 ± 0.070	0.031–0.22	0.073 ± 0.057	0.013–0.26	0.066 ± 0.027	0.034–0.12
syringic acid	0.35 ± 0.41	0.075–2.4	0.26 ± 0.21	0.058–0.59	0.17 ± 0.13	0.029–0.64	0.16 ± 0.084	0.049–0.28
dehydroabietic acid	1.7 ± 1.1	0.10–5.4	1.1 ± 0.69	0.31–2.4	1.1 ± 1.1	0.14–4.6	0.67 ± 0.24	0.16–0.98
cholesterol	1.8 ± 0.82	0.50–3.7	1.2 ± 0.51	0.57–2.0	0.98 ± 0.51	0.026–2.0	1.3 ± 0.56	0.51–2.0
<i>n</i> -alkanes [ng m^{-3}]								
docosane	3.2 ± 0.82	1.8–5.0	2.9 ± 0.61	2.0–4.0	3.0 ± 0.53	1.9–4.2	4.0 ± 4.8	2.1–19
tricosane	3.6 ± 1.2	2.0–7.2	3.2 ± 0.91	2.0–4.8	3.2 ± 0.65	1.8–4.4	5.0 ± 7.6	2.1–29
tetracosane	5.8 ± 3.2	2.5–19	5.7 ± 1.7	3.3–8.7	6.1 ± 2.3	2.9–15	6.3 ± 8.5	2.7–33
pentacosane	8.9 ± 6.7	3.5–34	5.7 ± 2.3	3.1–11	6.0 ± 1.6	3.7–9.2	5.8 ± 5.5	3.2–23
hexacosane	13 ± 9.8	4.3–49	8.6 ± 3.7	3.6–18	9.7 ± 2.8	5.0–16	7.1 ± 5.3	3.5–23
heptacosane	16 ± 14	4.7–64	7.2 ± 2.6	3.6–12	8.2 ± 2.4	3.7–14	5.8 ± 3.4	3.3–16
octacosane	12 ± 12	2.6–54	4.3 ± 1.8	1.7–7.9	5.9 ± 3.0	2.3–17	3.6 ± 1.7	2.3–8.2
nonacosane	13 ± 13	3.0–55	4.9 ± 2.1	1.5–8.7	6.3 ± 2.2	3.3–13	4.5 ± 1.4	2.6–7.8
triacontane	7.9 ± 7.8	2.0–36	3.8 ± 2.0	1.6–9.0	5.2 ± 2.7	2.0–16	3.3 ± 1.7	1.7–8.3
hentriacontane	14 ± 14	2.8–59	4.8 ± 1.9	1.8–8.4	5.7 ± 2.0	3.3–11	4.3 ± 1.2	2.9–6.9
dotriacontane	6.7 ± 5.5	1.6–27	3.4 ± 0.72	2.4–4.5	4.6 ± 1.3	2.8–7.8	3.1 ± 0.88	1.8–4.4
tritriacontane	6.8 ± 7.1	1.2–33	2.5 ± 0.97	1.1–4.2	2.8 ± 0.92	1.2–5.0	2.1 ± 0.72	1.5–3.8



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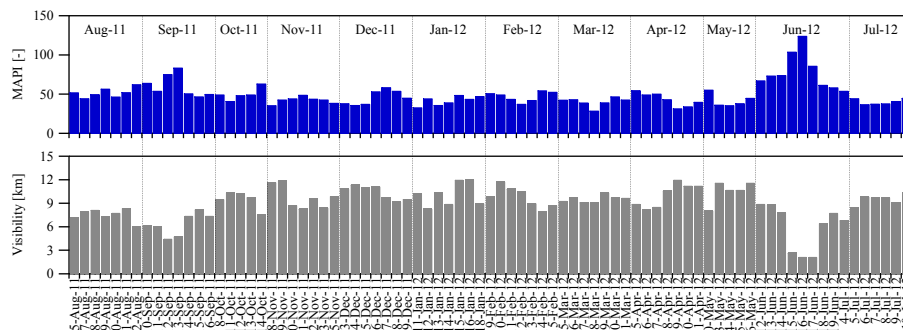


Figure 1. Daily variability of the MAPI and visibility during the sampling periods.

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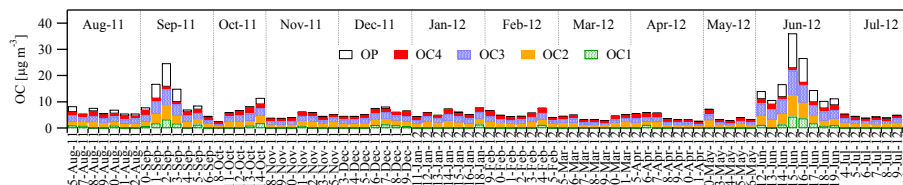


Figure 3. Daily variation of the OC fractions' mass concentrations during the sampling periods.

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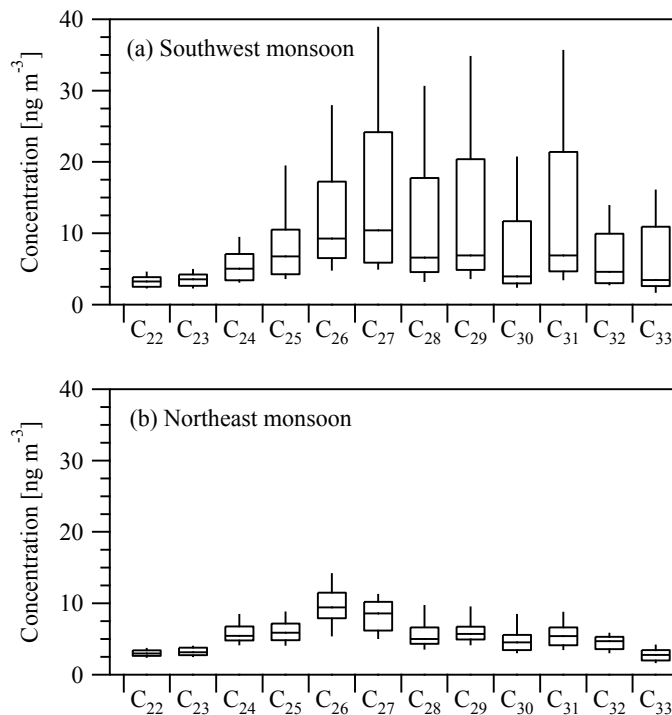


Figure 4. Box-whisker plots of molecular distributions of *n*-alkanes during the **(a)** southwest and **(b)** northeast monsoon seasons. The horizontal lines in the box represent the 25th, 50th, and 75th percentiles. The whiskers represent the 10th and 90th percentiles.

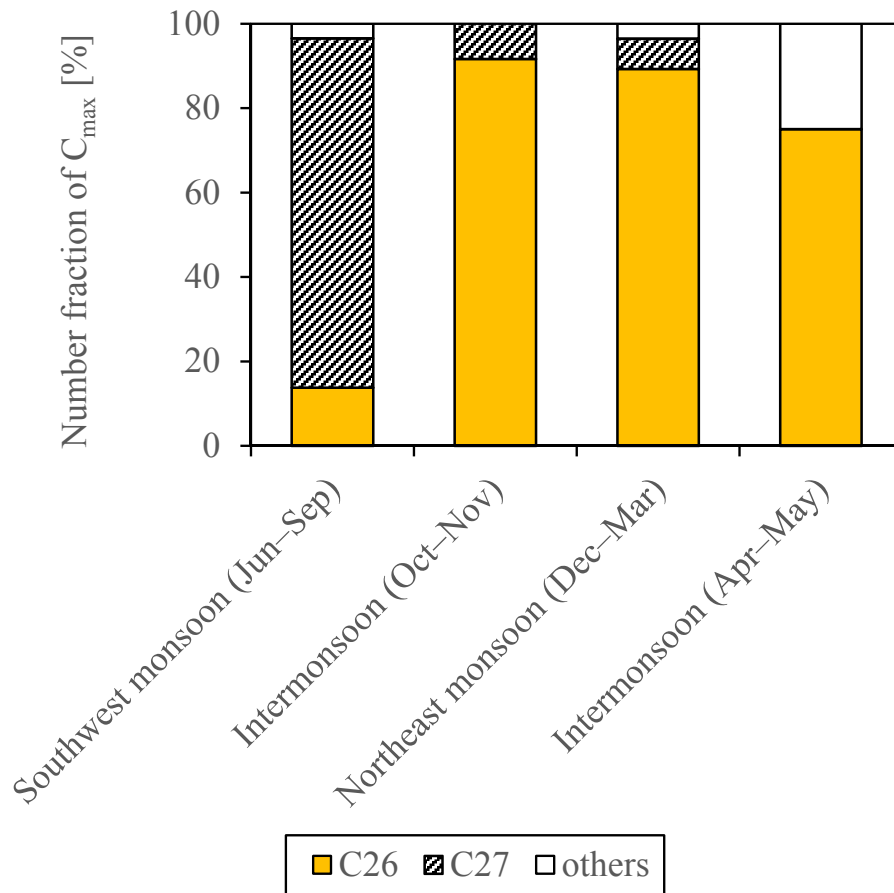


Figure 5. Number fraction of C_{\max} in the $PM_{2.5}$ samples for each monsoon season.

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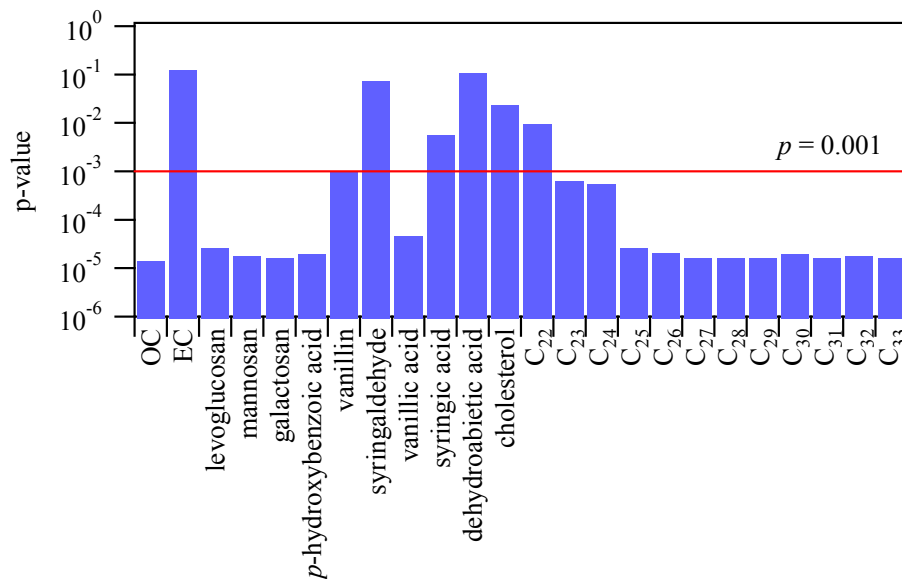


Figure 6. *P* values to determine significance in the two-sided Wilcoxon rank sum test between the IPF and other samples.

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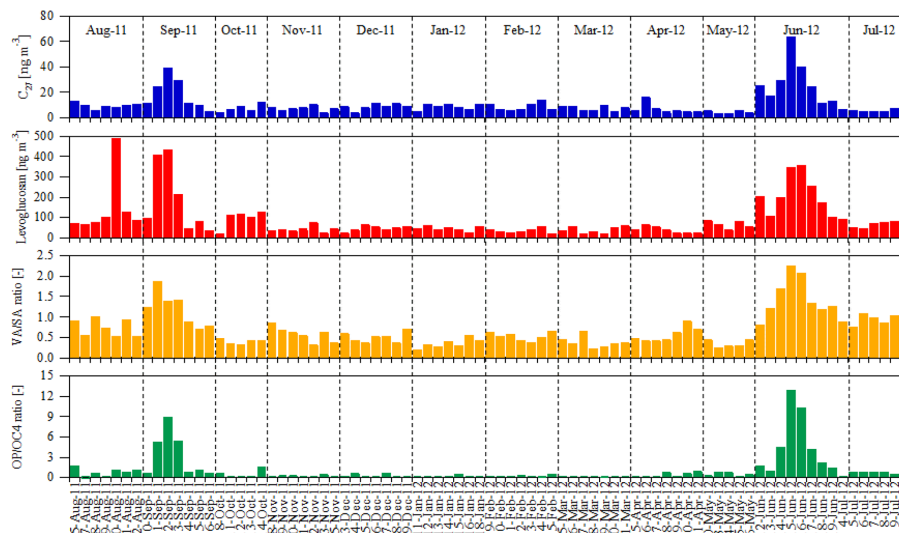


Figure 7. Daily variability of the C₂₇ and LG concentration as well as the VA/SA and OP/OC4 mass ratios during the sampling periods.

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