

1 **Global and regional emissions estimates of 1,1-difluoroethane (HFC-152a,**
2 **CH₃CHF₂) from in situ and air archive observations.**

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33
34 **Abstract**

35 High frequency, in situ observations from eleven globally-distributed sites for the period
36 1994–2014 and archived air measurements dating from 1978 onward have been used to
37 determine the global growth rate of 1,1-difluoroethane (HFC-152a, CH₃CHF₂). These
38 observations have been combined with a range of atmospheric transport models to derive
39 global emission estimates in a top-down approach. HFC-152a is a greenhouse gas with a
40 short atmospheric lifetime of about 1.5 years. Since it does not contain chlorine or bromine,
41 HFC-152a makes no direct contribution to the destruction of stratospheric ozone and is
42 therefore used as a substitute for the ozone depleting chlorofluorocarbons (CFCs) and
43 hydrochlorofluorocarbons (HCFCs). The concentration of HFC-152a has grown substantially
44 since the first direct measurements in 1994, reaching a maximum annual global growth rate
45 of 0.84 ± 0.05 ppt/yr in 2006, implying a substantial increase in emissions up to 2006.
46 However, since 2007, the annual rate of growth has slowed to 0.38 ± 0.04 ppt/yr in 2010 with

47 a further decline to an annual average rate of growth in 2013–2014 of -0.06 ± 0.05 ppt/yr.
48 The annual average Northern Hemisphere (NH) mole fraction in 1994 was 1.2 ppt rising to an
49 annual average mole fraction of 10.1 ppt in 2014. Average annual mole fractions in the
50 Southern Hemisphere (SH) in 1998 and 2014 were 0.84 and 4.5 ppt, respectively. We
51 estimate global emissions of HFC-152a have risen from 7.3 ± 5.6 Gg/yr in 1994 to a
52 maximum of 54.4 ± 17.1 Gg/yr in 2011, declining to 52.5 ± 20.1 Gg/yr in 2014 or 7.2 ± 2.8
53 Tg-CO₂ eq/yr. Analysis of mole fraction enhancements above regional background
54 atmospheric levels suggests substantial emissions from North America, Asia and Europe.
55 Global HFC emissions (so called ‘bottom up’ emissions) reported by the United Nations
56 Framework Convention on Climate Change (UNFCCC) are based on cumulative national
57 emission data reported to the UNFCCC, which in turn are based on national consumption
58 data. There appears to be a significant underestimate (> 20 Gg) of ‘bottom-up’ reported
59 emissions of HFC-152a, possibly arising from largely underestimated USA emissions and
60 undeclared Asian emissions.

61

62 **1. Introduction**

63 HFC-152a (CH₃CHF₂) is primarily sold as an aerosol and foam-blowing agent (Greally
64 et al., 2007) and as a component of some refrigerant blends (Ashford et al., 2004). Emissions
65 to the atmosphere show both temporal and regional variability depending on the specific
66 application in which HFC-152a is used. Incorporation of HFC-152a into aerosol propellants
67 results in prompt release, whereas when used as a single-component non-encapsulated
68 blowing agent, release occurs over a period of about 2 years (McCulloch et al., 2009).
69 Refrigerant use of HFC-152a results in release over longer periods, possibly up to 20 years.
70 Reported emissions of HFC-152a are likely to be incomplete as a consequence of a limited
71 number of producers and confidentiality considerations. Emissions of HFC-152a for some
72 countries are aggregated with other HFCs in a category reported to the UNFCCC as
73 “unspecified mix”. For example, emissions reported by the USA to the UNFCCC for HFC-
74 152a, 227ea, 245ca and 43-10mee are shown in the database as “commercially confidential”
75 and they constitute the aggregated “unspecified” emissions. HFC-152a emissions from the
76 USA are estimated to be the primary contributor to the total for this gas from Annex 1
77 countries (Lunt et al., 2015). Previous papers (Manning and Weiss, 2007; Millet et al., 2009;
78 Stohl et al., 2009; Barletta et al., 2011; Miller et al., 2012; Simmonds et al., 2015) have
79 reported major differences between USA HFC-152a emission estimates derived from
80 atmospheric measurements (top down) and emissions calculated from US reports to the
81 UNFCCC (bottom up). The apparent under-reporting of USA emissions to the UNFCCC
82 ranges from 20–60 Gg based on annual average estimates.

83 HFC-152a has the smallest 100-year global warming potential (GWP_{100} , 138) of all the
84 major HFCs (Forster et al., 2007; Myhre et al., 2013), with a short atmospheric lifetime of 1.5
85 years, due to efficient reaction with tropospheric hydroxyl (OH) radicals (SPARC Report No.
86 6, 2013). Unlike hydrocarbons, HFC-152a does not participate in the reaction to form ozone
87 in the troposphere. These desirable properties have made HFC-152a especially attractive as a
88 replacement, not only for CFCs and HCFCs, but also increasingly for HFC-134a in technical
89 aerosol applications and mobile air-conditioners (IPCC/TEAP, 2011).

90 Ryall et al. (2001) using observations from Mace Head, Ireland reported the
91 distribution of European HFC-152a emissions, concentrated in Germany, and estimated an
92 average European total emission of 0.48 Gg/yr for 1995-1998. Reimann et al. (2004) used a
93 3-year data set (2000–2002) of HFC-152a observations at the Swiss Alpine station
94 Jungfraujoch and trajectory modelling, also noting a predominantly German source for
95 European HFC-152a emissions. This group measured an atmospheric growth rate of 0.3
96 ppt/yr (ppt–parts per trillion, 10^{-12} , mol/ mol or pmol/mol) from 2000 to 2002 and a
97 December 2002 mole fraction at the Jungfraujoch station of 3.2 ppt, from which they
98 estimated a European emission strength of 0.8 Gg/yr for 2000–2002.

99 In the Southern hemisphere HFC-152a monthly means, annual means and trends have
100 been reported from observations at Cape Grim, Tasmania, for 1998–2004 (Sturrock et al.,
101 2001; Fraser et al., 2014a; Krummel et al., 2014;). The HFC-152a annual means have grown
102 from 0.8 ppt (0.1 ppt/yr) in 1998 to 1.8 ppt (0.4 ppt/yr) in 2004. More recent estimates of SE
103 Australian HFC-152a emissions (2005–2012) have been calculated by interspecies
104 correlation and model inversions and by extrapolation based on population (Fraser et al.,
105 2014a).

106 Here we further expand the HFC-152a record up to the end of 2014 using in situ
107 observations from eleven globally-distributed monitoring stations (9 Advanced Global
108 Atmospheric Gases Experiment (AGAGE) stations and 2 affiliated stations), together with
109 atmospheric transport models to independently estimate HFC-152a emissions on regional and
110 global scales. We then compare these with HFC-152a emission estimates compiled from
111 national reports to the United Nations Framework Convention on Climate Change
112 (UNFCCC) and Emissions Database for Global Atmospheric Research (EDGAR v4.2;
113 <http://edgar.jrc.ec.europa.eu/>), using the same techniques reported for other greenhouse gases
114 (O’Doherty et al., 2009, 2014; Miller et al., 2010; Vollmer et al., 2011; Krummel et al., 2014;
115 Rigby et al., 2014).

116

117 2. Experimental methods

118 2.1 Instrumentation and Calibration

119
120 High frequency, in situ measurements of HFC-152a were made by gas
121 chromatography-mass spectrometry (GC-Agilent 6890) coupled with quadrupole mass
122 selective detection (MSD-Agilent 5973/5975). Measurements commenced at Mace Head,
123 Ireland in 1994 and Cape Grim, Tasmania in 1998, using a custom-built automated pre-
124 concentration system (Adsorption Desorption System -ADS) to selectively and quantitatively
125 retain halogenated compounds from 2 L air samples. Based on a Peltier-cooled pre-
126 concentration microtrap cooled to -50°C during the adsorption phase, the ADS provided on
127 site calibrated air samples every 4 hours, i.e. 6 per day (Simmonds et al., 1995). In 2004 the
128 ADS-GC-MS was replaced with a more advanced custom-built pre-concentration system
129 (Medusa) with enhanced cooling to $\sim -180^{\circ}\text{C}$ and the relatively mild adsorbent HayeSep D
130 (Miller et al., 2008; Arnold et al., 2012). Agilent 5973 MSDs (mass selective detector) were
131 also upgraded to the Agilent 5975 MSDs over the course of the Medusa observations.
132 Analysis of each 2 L sample of ambient air was alternated with analysis of a 2 L reference
133 gas (designated as a working standard) to correct for short-term instrumental drift, resulting
134 in 12 (Medusa) individually calibrated air measurements per day. Working standards were
135 prepared for each station by compressing ambient air into 34 L electropolished stainless steel
136 canisters (Essex Industries, Inc., Missouri) using modified oil-free compressors (SA-6, RIX,
137 California). Exceptions to this were the Cape Grim and Zeppelin stations, where the working
138 standards were filled using a cryogenic filling technique. Research-grade helium, which was
139 used as a carrier gas in the Medusa systems, was further purified by passage through a heated
140 “getter” type purifier (Valco Instruments, Houston, TX). The carrier gas was analysed for
141 blanks on a regular basis and blank levels of HFC-152a were below the limit of detection at
142 all field stations.

143 Table 1 lists the geographical location and the time when routine ambient
144 measurements of HFC-152a began at each monitoring station. Stations with the longest
145 observational records that deployed both ADS and Medusa GC-MS instruments include
146 Mace Head (MHD), Jungfraujoch (JFJ), Ny-Ålesund (ZEP) and Cape Grim (CGO). Medusa
147 GC-MS instruments were installed at five other AGAGE stations Trinidad Head (THD),
148 Gosan (GSN), Ragged Point, (RPB), Shangdianzi (SDZ) and Cape Matatula (SMO) between
149 2003 and 2010. In addition two AGAGE affiliated stations Monte Cimone (CMN) and
150 Hateruma (HAT), which use comparable GC-MS instruments, but a different pre-

151 concentration design for sample enrichment, commenced HFC-152a measurements in 2001
152 and 2004, respectively. Importantly, all eleven stations listed in Table 1 report HFC-152a
153 measurements relative to the Scripps Institution of Oceanography (SIO-05) calibration scale
154 (as dry gas mole fractions in pmol/mol).

155
156 The estimated accuracy of the calibration scale for HFC-152a is 4%: a more detailed
157 discussion of the measurement technique and calibration procedure is reported elsewhere
158 (Miller et al., 2008; O'Doherty et al., 2009; Mühle et al., 2010). HFC-152a was determined
159 using the MS in selected ion monitoring mode (SIM) with a target ion CH_3CF_2^+ (m/z 65) and
160 qualifier ion CH_3CF^+ (m/z 46). To ensure that potential interferences from co-eluting species
161 did not compromise the analysis, the ratio of the target to qualifier ion was continuously
162 monitored. Measurement precision was calculated as the daily standard deviation (1σ) of the
163 ratios of each standard response to the average of the closest-in-time preceding and
164 subsequent standard responses. Typical daily precisions vary from station to station with a
165 range of 0.1–0.4 ppt. Individual station precisions were used to estimate the precision of each
166 in situ measurement.

167 168 *2.2 Northern and Southern Hemisphere archived air samples*

169
170 In order to extend the HFC-152a data record back before the commencement of high-
171 frequency measurements, analyses of Northern Hemisphere (NH) and Southern Hemisphere
172 (SH) archived air samples dating back to 1978, were carried out using three similar Medusa
173 GC-MS instruments at the Scripps Institution of Oceanography (SIO), La Jolla, California,
174 the Commonwealth Scientific and Industrial Research Organisation (CSIRO), Aspendale,
175 Australia and the Cape Grim Baseline Station, Tasmania. The SH samples are part of the
176 Cape Grim Air Archive (CGAA) described in Langenfelds et al., (1996); and Krummel et al.,
177 (2007). The NH samples analysed for this paper were filled during background conditions
178 mostly at Trinidad Head, but also at La Jolla, California; Cape Meares, Oregon; Ny Ålesund,
179 Svalbad and Point Barrow, Alaska (some samples are courtesy of the National Oceanic and
180 Atmospheric Administration (NOAA)).

181 In addition, eight SH samples were measured at SIO and compared with SH samples
182 of similar age measured at CSIRO (February 1995, July 1995, November 1995, June 1998,
183 July 2004, February 2006, August 2008, and December 2010, $\Delta x = 0.01$ – 0.07 ppt $\Delta t = 1$ – 33
184 days) and three NH samples were measured at CSIRO and compared with NH samples of the
185 same age measured at SIO (May 1989 and April 1999, $\Delta x = 0.02$ – 0.06 ppt, $\Delta t = 1$ – 11 days).

186 The good agreement between SIO and CSIRO archived air stored in different types of tanks
187 (stainless steel tanks, Essex Industries, Inc and Silcosteel treated tanks, Restek Corporation)
188 serves both as proof of the good consistency of the individual Medusa GC-MS instruments
189 and the integrity of the tanks used. Samples were analysed in replicate typically 3–6 times
190 each and several NH tanks were re-measured over a number of years.

191 192 *2.3 Selection of baseline data*

193
194 Baseline in situ monthly mean HFC-152a mole fractions were calculated by excluding
195 values enhanced by local and regional pollution influences, as identified by the iterative
196 AGAGE pollution identification algorithm, (see Appendix in O’Doherty et al., 2001).
197 Briefly, baseline measurements are assumed to have a Gaussian distribution around the local
198 baseline value, and an iterative process is used to filter out the points that do not conform to
199 this distribution. A second-order polynomial is fitted to the subset of daily minima in any
200 121-day period to provide a first estimate of the baseline and seasonal cycle. After
201 subtracting this polynomial from all the observations a standard deviation and median are
202 calculated for the residual values over the 121-day period. Values exceeding three standard
203 deviations above the baseline are thus identified as non-baseline (polluted) and removed from
204 further consideration. The process is repeated iteratively to identify and remove additional
205 non-baseline values until the new and previous calculated median values agree within 0.1%.
206 For the core AGAGE stations, in situ baseline data and archive air data, extending the record
207 to periods prior to the in situ measurement period, are then combined for each hemisphere,
208 and outliers are rejected by an iterative filter.

209

210 **3. Modelling studies**

211

212 We pursued several approaches to determine emissions at global, continental and
213 regional scales. The methodologies have been published elsewhere and are summarised
214 below. The global, continental and some regional estimates incorporate a priori estimates of
215 emissions, which were subsequently adjusted using the observations.

216 There are several sources of information on production and emissions of HFC-152a; none of
217 which, on their own, provides a complete database of global emissions. The more
218 geographically comprehensive source of information is provided by the parties to the
219 UNFCCC, but only includes Annex 1 countries (developed countries). The 2014 database
220 covers years 1990 to 2012 and are reported in Table 2(II) s1 in the Common Reporting

221 Format (CRF) available at [http://unfccc.int/national](http://unfccc.int/national_reports/annex_ghg_inventories/national) reports/annex ighg inventories/national
222 [inventories submissions/items/8108.php](http://unfccc.int/national_reports/annex_ghg_inventories/national_inventories_submissions/items/8108.php). An alternative inventory estimate was also obtained
223 from the Emissions Database for Global Atmospheric Research (EDGAR v4.2;
224 <http://edgar.jrc.ec.europa.eu/>), a database that estimates global emission inventories of
225 anthropogenic greenhouse gases (GHGs) on a country, region and grid basis up to 2008.

226 To infer “top-down” emissions we select observations from the various observing sites
227 listed in Table 1 and four chemical transport models. These eleven sites are sensitive to many
228 areas of the world in which HFC-152a emissions are reported, however other areas of the
229 globe that are not well monitored by this network are also likely to have significant emissions
230 (such as South Asia, South Africa, and South America).

231 232 *3.1 Global emissions estimates using the AGAGE two-dimensional 12-box model* 233

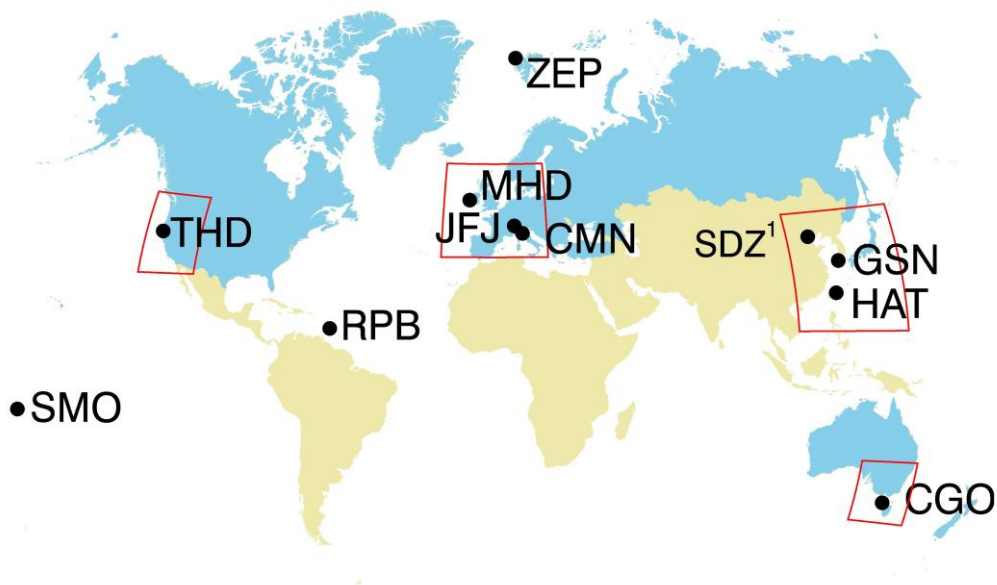
234 To estimate global-average mole fractions and derive growth rates, a two-dimensional
235 model of atmospheric chemistry and transport was employed. The AGAGE 12-box model
236 simulates trace gas transport in four equal mass latitudinal sections (divisions at 30-90°N, 0-
237 30°N, 30-0°S and 90-30°S) and at three heights (vertical divisions at 200, 500 and 1000 hPa).
238 The model was originally developed by Cunnold et al. (1983) (nine-box version), with
239 subsequent improvements by Cunnold et al. (1994) and Rigby et al. (2013, 2014). Emissions
240 were estimated between 1989 and 2014 using a Bayesian method in which an a priori
241 constraint (EDGAR v4.2) on the emissions growth rate was adjusted using the baseline-
242 filtered AGAGE observations (Rigby et al., 2011a, 2014). Global emissions were derived that
243 included estimates of the uncertainties due to the observations, the prior and the lifetime of
244 HFC-152a, as detailed in the supplementary material in Rigby et al. (2014). Note that
245 historically and here the 12-box model only uses observations from the core AGAGE sites,
246 Mace Head, Trinidad Head, Ragged Point, Cape Matatula, and Cape Grim.

247 248 *3.2 Global and continental emissions estimates using a combined Eulerian and Lagrangian* 249 *model* 250

251 We used the methodology outlined in Lunt et al. (2015) and Rigby et al. (2011b) to
252 derive emissions of HFC-152a from continental regions. The high-resolution, regional UK
253 Met Office Numerical Atmospheric dispersion Modelling Environment (NAME), Manning et
254 al. (2011) was used to simulate atmospheric HFC transport close to a subset of AGAGE
255 monitoring sites, which were strongly influenced by regional HFC sources (domains shown
256 by red boxes in Figure 1). Simultaneously, the influence of changes to the global emissions

257 field on all measurement stations was simulated using the global Model for OZone and
 258 Related Tracers, MOZART (Emmons et al., 2010). We estimated annual emissions for the
 259 period 2007-2012 and aggregated the derived emissions fields into continental regions,
 260 separating countries that either do (“Annex-1”), or do not (“non-Annex-1”) report detailed,
 261 annual emissions to the UNFCCC. Emissions were estimated using a hierarchical Bayesian
 262 inverse method (Ganesan et al., 2014, Lunt et al., 2015) and all high-frequency observations
 263 from 10 of the 11 sites listed in Table 1, excluding Shangdianzi due to the short time series.
 264 The hierarchical Bayesian method includes uncertainty parameters (e.g. model “mismatch”
 265 errors and a priori uncertainties) in the estimation scheme, reducing the influence of
 266 subjective choices on the outcome of the inversion.

267



291

292 Figure 1. Location of AGAGE and affiliated stations.
 293 Ny-Ålesund, Zeppelin, Norway (ZEP); Mace Head, Ireland (MHD); Jungfrauoch,
 294 Switzerland (JFJ); Monte Cimone, Italy (CMN);Trinidad Head, USA (THD); Shangdianzi,
 295 China (SDZ); Gosan, South Korea (GSN); Hateruma, Japan (HAT); Ragged Point, Barbados
 296 (RPB); Cape Matatula, American Samoa (SMO); and Cape Grim, Tasmania (GCO). Red
 297 boxes indicate “local regions” where the NAME model was used with increased resolution
 298 compared to the global MOZART model, Annex 1 countries are shaded blue and non-Annex
 299 1 countries are shaded yellow.
 300 Note: ¹ Shangdianzi (SDZ) was not used in any of the modelling studies due to the relatively
 301 short time series.

302 *3.3 High-resolution regional emissions estimates using InTEM*

303

304 A method for estimating emissions from observations and atmospheric transport
305 modelling with NAME referred to as InTEM, ‘Inversion Technique for Emission Modelling’
306 (Manning et al., 2011), uses a simulated annealing method (Press et al., 1992) to search for
307 the emission distribution that produces a modelled times-series that has the best statistical
308 match to the observations from certain AGAGE stations (e.g. Mace Head, Cape Grim).
309 NAME was driven with output from the operational analysis of the UK Met Office
310 Numerical Weather Prediction model, the Unified Model, at global horizontal resolution of
311 17-40 km (year dependent). InTEM estimates the spatial distribution of emissions across a
312 defined geographical area, and can either start from a random emission distribution or be
313 constrained by an inventory-defined distribution. Emission totals from specific geographical
314 areas are calculated by summing the derived emissions from each grid (non-uniform) in that
315 region.

316 The uncertainty estimation used within InTEM is described in detail elsewhere
317 (Manning et al 2011). Briefly, the uncertainty space was explored by a) solving the inversion
318 multiple times with a range of baseline mole fractions within the baseline uncertainty
319 estimated during the baseline fitting process and b) by altering the 3-year inversion time
320 window by one month throughout the data period thereby solving over a particular one year
321 period many times using different observations. In total for each annual estimate up to 111
322 inversions were performed, the median and 5th and 95th percentiles were used as the final total
323 and spread. For the Australian estimates data between 2002 through 2011 were used, for the
324 NW European estimates data between Nov. 1994 and Dec. 2013 were used.

325 *3.4 High resolution European emission estimates using the FLEXPART model*

326 A regional Bayesian inversion system using backward simulations of a Lagrangian
327 particle dispersion model FLEXPART (Stohl et al., 2005) was applied to the HFC-152a
328 observations from Mace Head, Jungfraujoch and Mt. Cimone for the period 2006 to 2014.
329 The inversion technique follows the description by Stohl et al. (2009) and was previously
330 applied to regional halocarbon emissions from Europe (Keller et al., 2012, Maione et al.,
331 2014) and China (Vollmer et al., 2009). For these emission estimates, the background was
332 determined by applying the Robust Extraction of Baseline Signal (REBS) filter described in
333 detail by Ruckstuhl et al, (2012). The transport model FLEXPART was driven with output
334 from the operational analysis of the Integrated Forecast System (IFS) of the European Centre

335 for Medium Range Weather Forecast (ECMWF) using a spatial resolution of $0.2^\circ \times 0.2^\circ$ for a
336 nested domain covering the larger area of the European Alps and a spatial resolution of $1^\circ \times$
337 1° elsewhere.

338 The FLEXPART model was applied to the HFC-152a observations from Mace Head,
339 Jungfraujoch and Mt. Cimone for the period 2006 to 2014. Prior to 2006, the model
340 resolution of Integrated Forecast System (IFS) was not sufficiently fine to realistically
341 simulate the transport to the two high altitude sites Jungfraujoch and Mt. Cimone. Therefore,
342 no attempt was made here to apply the inversion system to years before 2006. As prior
343 information of the HFC-152a emissions we used country totals as submitted to UNFCCC.
344 These were spatially disaggregated following the HFC-152a distribution given in EDGAR
345 (v4.2). For countries not reporting HFC-152a emissions to UNFCCC we used the values
346 given in EDGAR. The EDGAR inventory was only available up to the year 2008 beyond this
347 year the EDGAR 2008 distribution was used. The uncertainty of the prior emissions was set
348 so that the region total uncertainty equalled 20 % of the region total emissions. The regional
349 inversion grid covered a region similar to that shown in Figure 1.

350

351 *3.5 Regional emissions estimates using the inter-species correlation (ISC) methods*

352 We also present regional emissions estimates using inter-species correlation (ISC)
353 methods (Yokouchi et al., 2005). Emissions of a number of trace gases from the
354 Melbourne/Port Phillip region (CFCs, HCFCs, HFCs, carbon tetrachloride: Dunse et al.,
355 2001, 2002, 2005; O'Doherty et al., 2009; Fraser et al., 2014a, b), including HFC-152a
356 (Greally et al., 2007), have been estimated utilising in situ high frequency measurements
357 from Cape Grim and ISC with co-incident carbon monoxide (CO) measurements.

358 ISC works best for co-located sources – however extensive modelling has shown that
359 by the time the Melbourne/Port Phillip plume reaches Cape Grim (300 km from the source) it
360 is well mixed and the likely inhomogeneity of the source regions (for CO and HFC-152a in
361 this case) does not have a significant influence on the derived emissions. It should be noted
362 that in order to obtain a significant sampling of Port Phillip pollution episodes at Cape Grim,
363 data from 3 years (for example 2011-2013) are used to derive annual emissions (for 2012).
364 (InTEM also uses data from 3 years to derive annual emissions.) The ISC uncertainties given
365 in the paper include (1) the uncertainties in the estimates of CO emissions from
366 Melbourne/Port Phillip (2) the uncertainties in the overall correlation between CO and

367 HCFC-152a as seen in pollution episodes at Cape Grim (3) the uncertainties in the
368 geographic extent of the HFC-152a and CO source regions impacting on Cape Grim and their
369 entrained population.

370 Using HCFC-22 as the reference tracer, Li et al., (2011) reported that China is the
371 dominant emitter of halocarbons in East Asia. North American HFC-152a emissions have
372 been estimated from atmospheric data using interspecies correlation based techniques with
373 CO (Millet et al., 2009, Barletta et al., 2011) and fossil fuel CO₂ (Miller et al., 2012) as the
374 reference emissions.

375 **4. Results and Discussion**

376

377 *4.1 In situ observations*

378

379 The time series of HFC-152a in situ observations recorded at selected AGAGE and
380 affiliated monitoring stations are shown in Figure 2 (a-c). Data have been filtered into
381 baseline (black) and above baseline (red) using the AGAGE pollution algorithm, as discussed
382 in section 2.3. Figure 2a shows the mole fractions in ppt for the four stations that deployed
383 both ADS and Medusa GC-MS instruments (Mace Head, Zeppelin, Jungfraujoch, and Cape
384 Grim). Most notable are the substantial above baseline events at Mace Head and Jungfraujoch
385 that are influenced primarily by emissions from European sources. Conversely, the Zeppelin
386 Arctic station and the SH station at Cape Grim have relatively small above baseline events
387 implying smaller emissions from local or regional sources.

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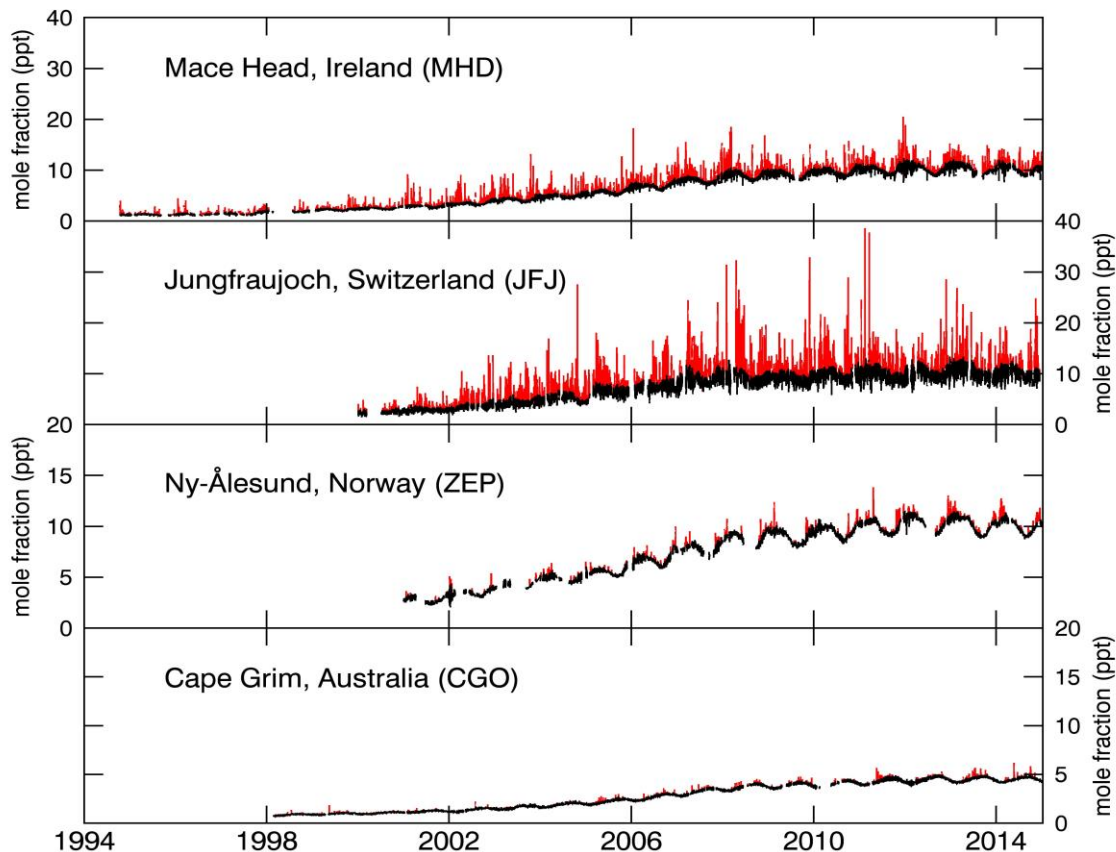
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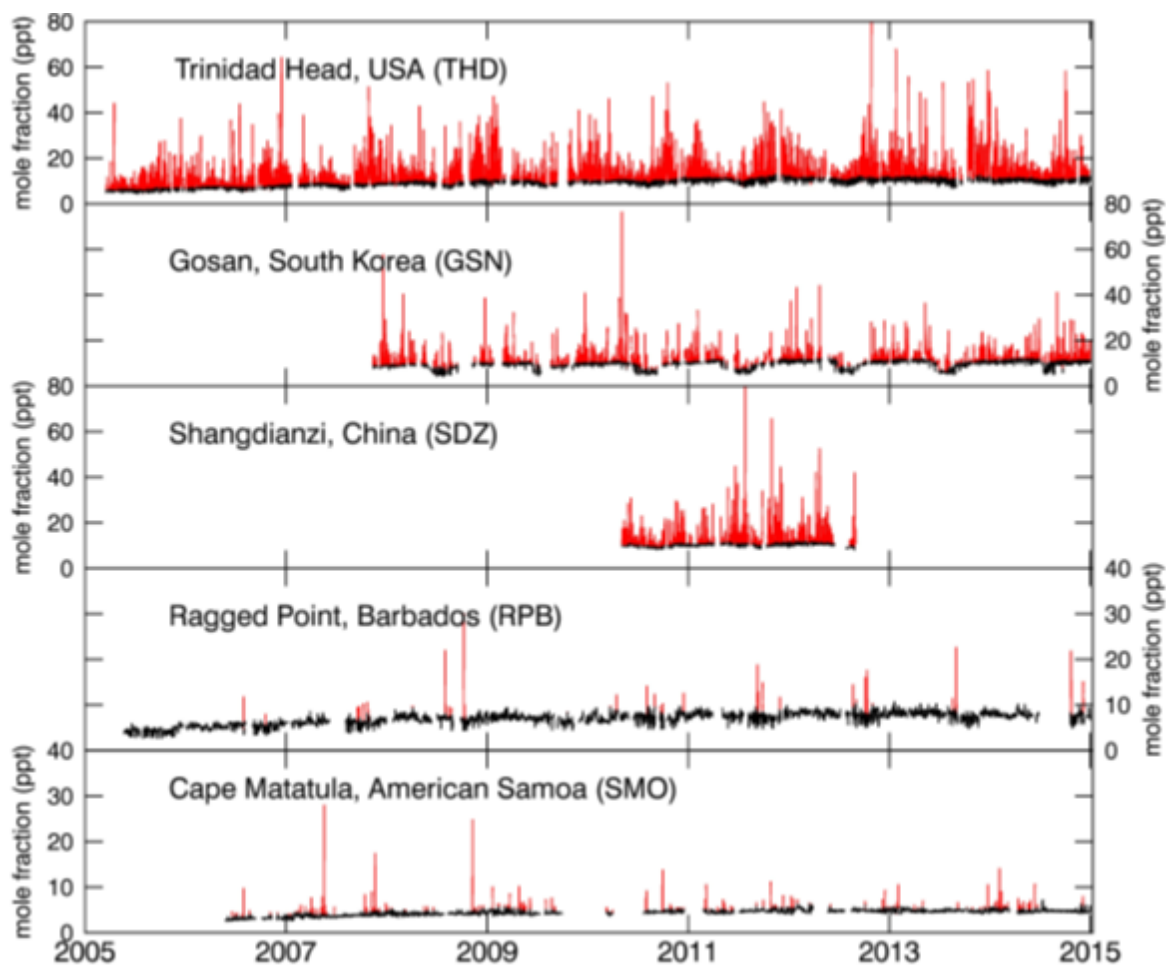
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427 Figure 2a. Time series of HFC-152a mole fractions (ppt) recorded at the four monitoring
428 stations with combined ADS and Medusa data. (MHD, JFJ, ZEP and CGO), (note the
429 different Y-axis scales). Data have been assigned as baseline (black) and non-baseline (red)
430 using the AGAGE pollution identification algorithm.
431

432 Figure 2b shows measurements at the five other AGAGE stations (Trinidad Head,
433 Gosan, Ragged Point, Shangdianzi, and Cape Matatula), which used only Medusa GC-MS
434 instruments. The North American site at Trinidad Head and the Asian sites at Shangdianzi
435 and Gosan are the most strongly influenced by regional emissions. The tropical sites at
436 Ragged Point, Barbados and Cape Matatula, American Samoa show very few enhancements
437 above the baseline and these are due mostly to local emissions occurring under night time
438 inversion conditions and occasional influences from regional emission sources (note the
439 different Y-axis scales). Although the Shangdianzi station was operational for only a short
440 period, the enhancements above baseline are significant due to the sensitivity of this site to
441 Chinese emissions, and comparable in magnitude to those at Gosan.
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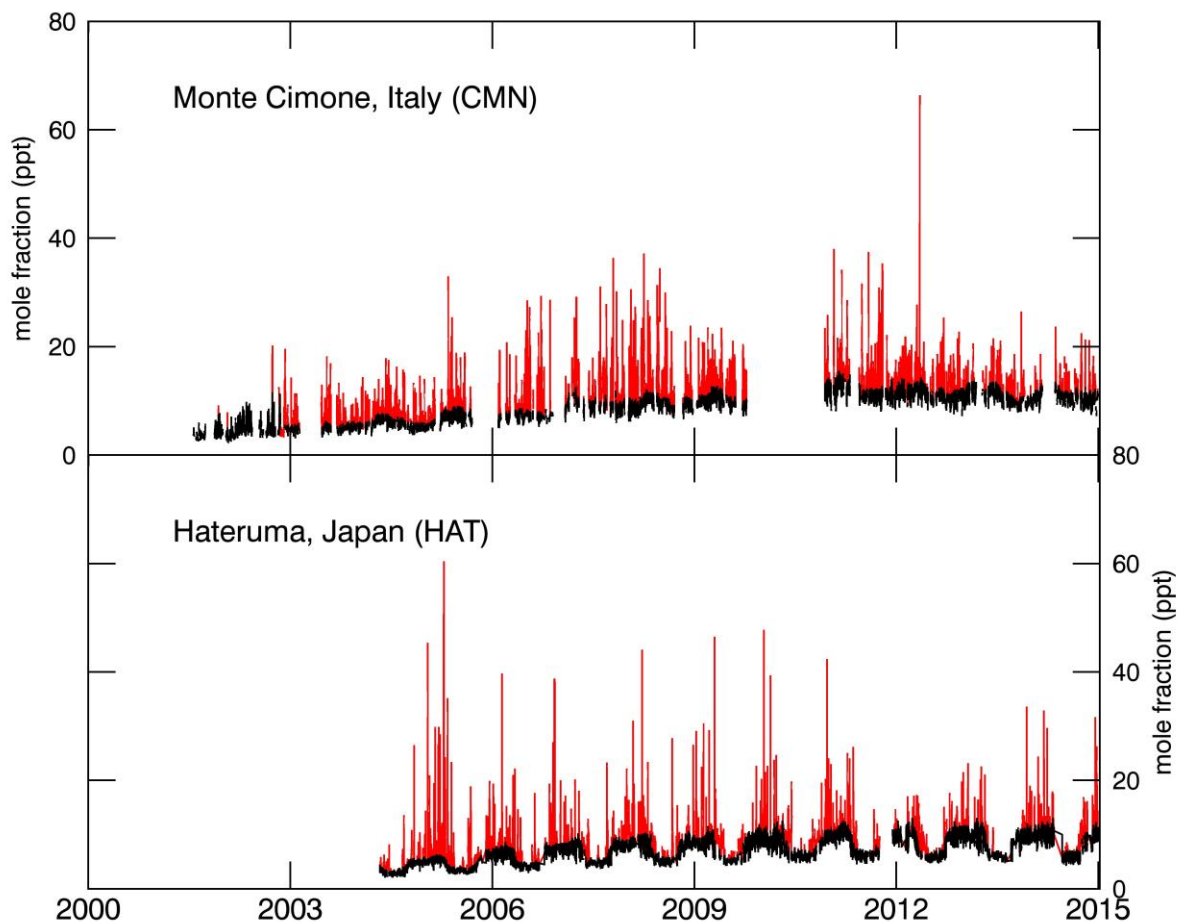


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Figure 2b. Time series of HFC-152a mole fractions (ppt), recorded with the Medusa GC-MS instruments at the five AGAGE monitoring stations (THD, GSN, SDZ, RPB, and SMO). Data have been assigned as baseline (black) and non-baseline (red) using the AGAGE pollution identification algorithm.

451 Figure 2c illustrates the time series from the two affiliated AGAGE stations (Monte
 452 Cimone and Haturuma) that used comparable GC-MS instruments but with different methods
 453 of pre-concentration. Monte Cimone, like the Jungfrauoch, is also influenced by substantial
 454 emissions from sources in continental Europe. Haturuma is influenced by sources in China,
 455 Korea, Taiwan and Japan (Yokouchi et al., 2006).

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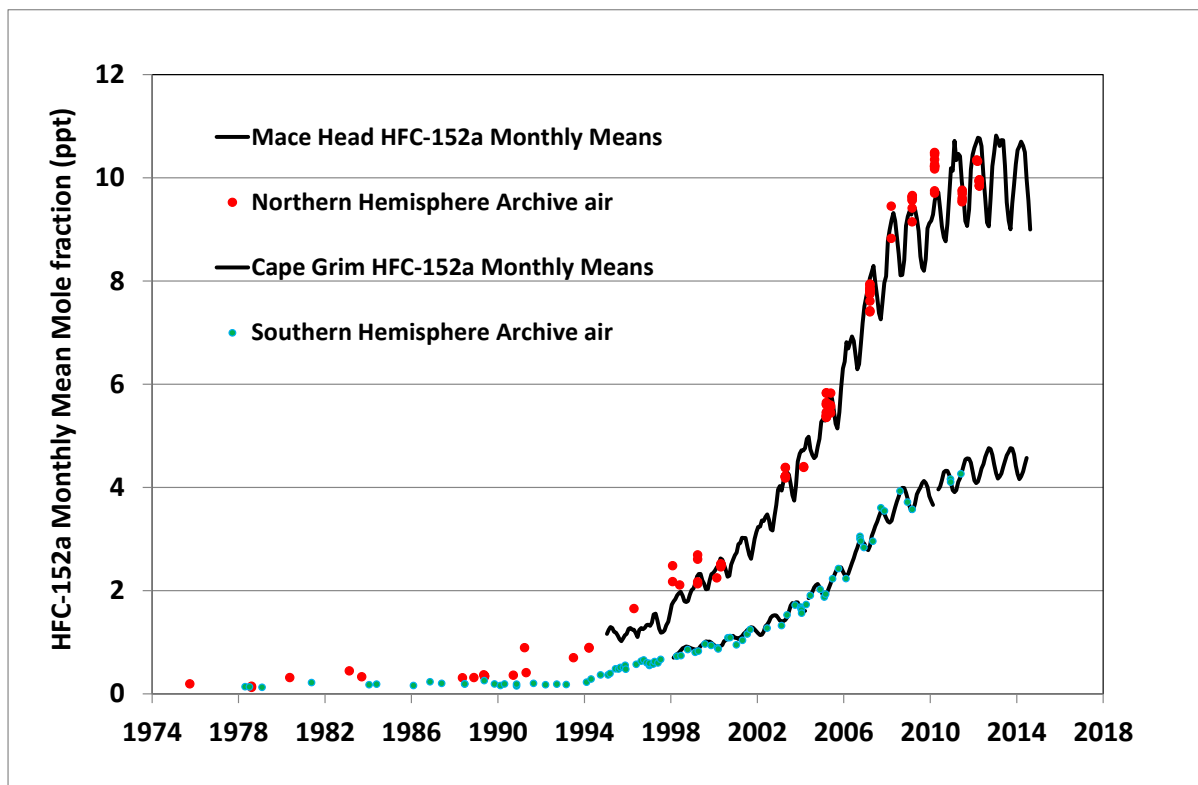


465
 466 Figure 2c. Time series of HFC-152a mole fractions (ppt) recorded with the GC-MS
 467 instruments at the two affiliated AGAGE stations CMN and HAT. Data have been assigned
 468 as baseline (black) and non-baseline (red) using the AGAGE pollution identification
 469 algorithm.

470
 471 *4.2 Atmospheric Trends and Seasonal Cycles*

472
 473 Figure 3 shows the in situ measurements of HFC-152a, as baseline monthly means
 474 (excluding pollution events), obtained from the two AGAGE stations Mace Head and Cape
 475 Grim with the longest time series that deployed both ADS and Medusa GC-MS instruments.
 476 Superimposed in Figure 3 are the NH and SH archived flask data extending back to 1978.
 477 Annual average mole fractions at Mace Head increased from 1.2 ppt in 1994 to 10.2 ppt by
 478 2014, Cape Grim annual average mole fractions increased from 0.84 ppt in 1998 when in situ
 479 measurements first began to 4.5 ppt in 2014. However, in the last few years the rates of
 480 growth at both sites have slowed to almost zero.

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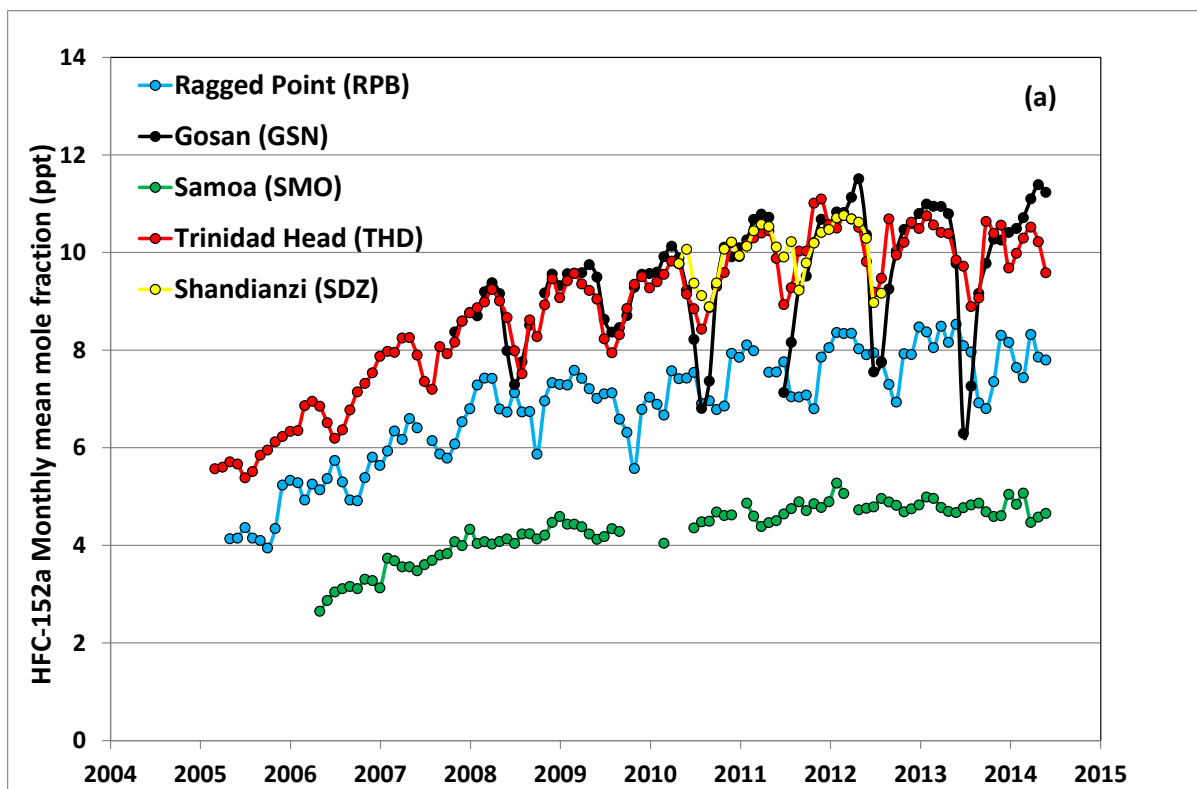


485
 486 Figure 3. HFC-152a baseline monthly mean mole fraction (ppt) recorded at Mace Head-
 487 MHD (ADS GC-MS, 1994-2003; Medusa GC-MS, 2004-2014) and at Cape Grim-CGO
 488 (ADS GC-MS, 1998-2003; Medusa GC-MS, 2004-2014) and from analysis of archived NH
 489 and SH air samples extending back to 1975: in situ (black), air archive NH (red) and SH
 490 (blue).

491
 492 The NH archived samples are more variable than the SH archived samples. The SH archive is
 493 collected only under strict baseline conditions (Southern Ocean air) and is far removed from
 494 the major sources of HFC-152a. Conversely in the NH, where most major sources of
 495 emissions are located, sampling under strict baseline conditions is more difficult to achieve.

496
 497 Figure 4a illustrates HFC-152a baseline monthly means obtained from the five other
 498 AGAGE observing sites (Ragged Point, Gosan, Cape Matatula, Trinidad Head and
 499 Shangdianzi using only the more advanced Medusa GC-MS. There is a large seasonal cycle
 500 at Gosan with a very deep minimum due to summertime transport from the Southern
 501 Hemisphere (Li et al., 2011). Barbados can also be influenced by southern hemispheric air
 502 during the hurricane season (Archibald et al., 2015).

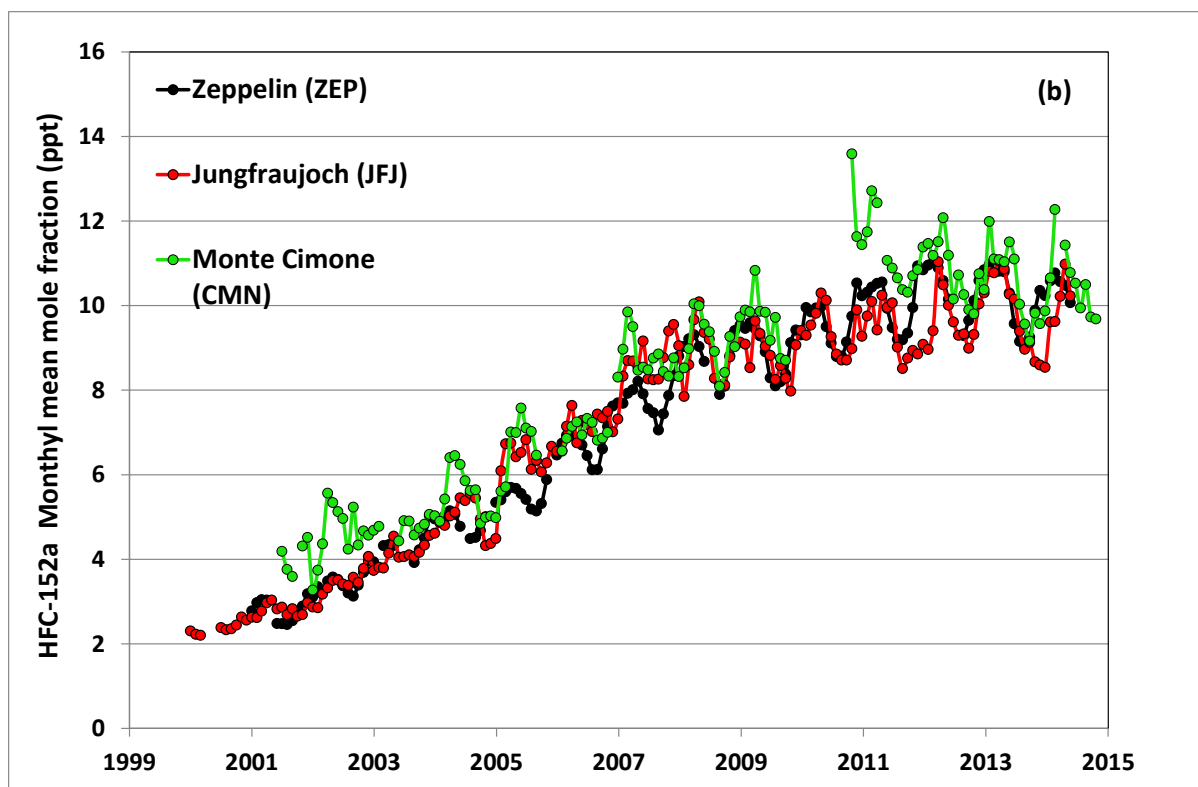
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 508 Figure 4a. Medusa GC-MS baseline monthly mean mole fractions (ppt) recorded at Ragged
 509 Point, Gosan, Cape Matatula, Trinidad Head, and Shangdianzi. Observations at Shangdianzi
 510 were discontinued in August 2012.

511
 512 Figure 4b shows the baseline monthly mean mole fractions for the three mountain
 513 stations. Ny-Ålesund and Jungfraujoch, using combined ADS and Medusa GC-MS
 514 measurements and Monte Cimone, which used a commercial pre-concentrator GC-MS. In
 515 most years Monte Cimone exhibits enhanced mole fractions during the NH spring months
 516 (March–May).

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532 Figure 4b. Combined ADS and Medusa GC-MS baseline monthly mean mole fraction
 533 recorded at Ny-Ålesund , Jungfraujoch and Monte Cimone.

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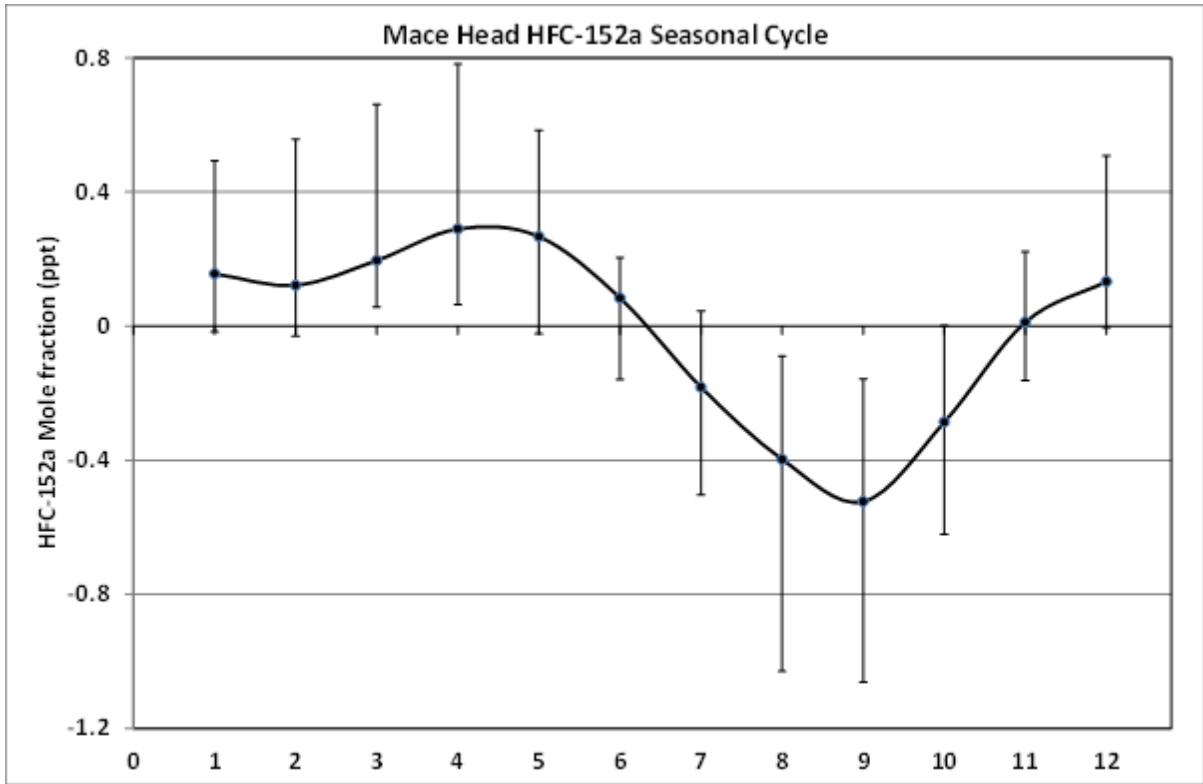
535 The HFC-152a seasonal cycles at Mace Head and Cape Grim shown in Figures 5a and
 536 5b, are broadly representative of the Northern and Southern Hemispheres, respectively. The
 537 seasonal cycle at Mace Head shows a NH spring maximum (April-May) and late summer
 538 minimum (August-October), while the SH seasonal cycle at Cape Grim exhibits a broad
 539 austral spring maximum (July-November) and a late summer minimum (January-April). The
 540 summer minimum at both locations is attributed to enhanced summertime loss (OH) with
 541 possibly a contribution from seasonally varying emissions in the NH that may be out-of-
 542 phase with the NH sink. At Cape Grim an additional source of seasonality is due to
 543 seasonally varying transport between the NH and SH, which is generally in phase with the
 544 sink induced seasonal cycle. This competition between OH summertime loss and seasonally
 545 varying transport has been observed at many other AGAGE locations (Prinn et al., 1992;
 546 Greally et al., 2007; O'Doherty et al., 2009, 2014 and Li et al., 2011).

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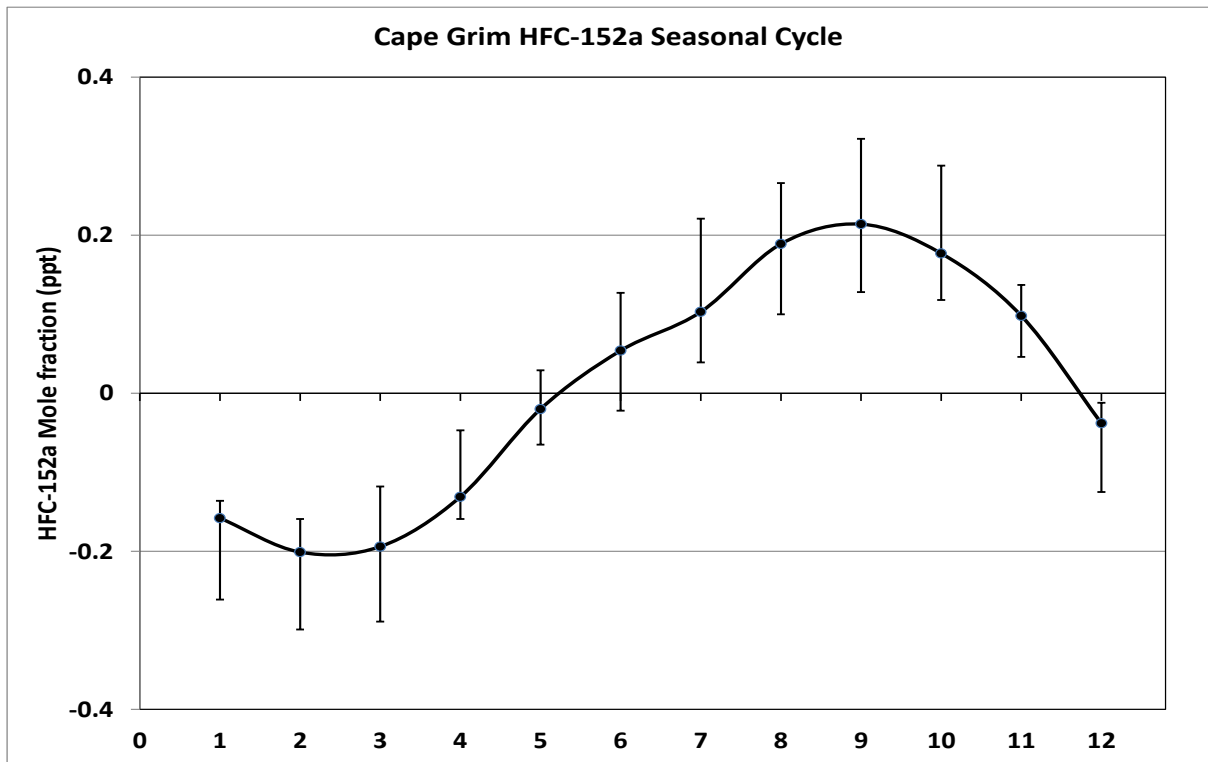
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552 Figure 5a. Average seasonal cycle at Mace Head, Ireland (2004-2014). Black line represents
 553 the average for each month of these individual years and the error bars represent the min to
 554 max range.

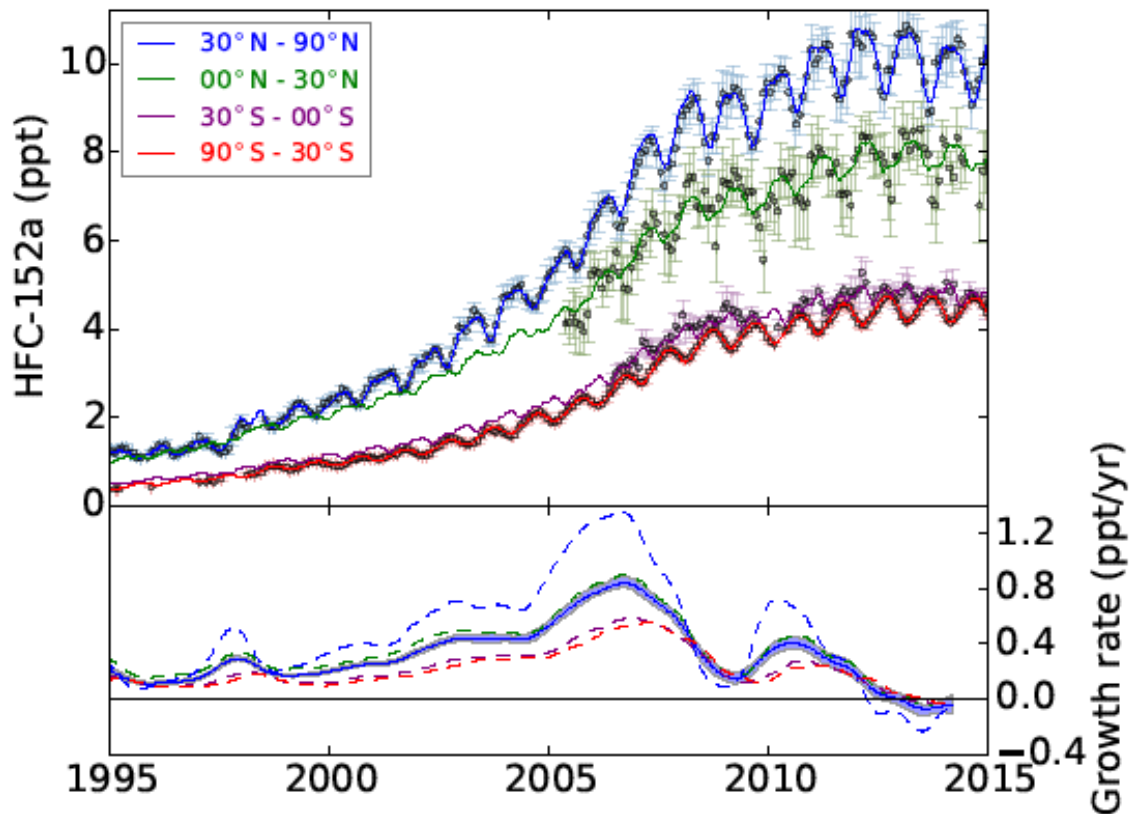


555

556 Figure 5b. Average seasonal cycle at Cape Grim, Tasmania (2004-2014). Black line
 557 represents the average for each month of these individual years and the error bars represent
 558 the min to max range.

559 Figure 6 shows the mole fractions output from the AGAGE global 12-box model,
560 along with the monthly-mean semi-hemispheric average observations used in the inversion.
561 The figure also shows the running mean growth rate, smoothed using a Kolmogorov–
562 Zurbenko filter with a window of approximately 12 months (Rigby et al., 2014). Most
563 notable is the positive growth rate from 1995 reaching a maximum of ~ 0.84 ppt/yr in 2006,
564 followed by a steady decline in the growth rate with a minimum in 2008-2009 during the time
565 of the economic recession. The positive growth rate then resumes increasing to ~ 0.4 ppt/yr in
566 2010 followed by a subsequent decrease with an annual average growth rate in 2013-2014 of
567 minus ~ 0.06 ppt/yr.

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582 Figure 6. Top panel: AGAGE 12-box model mole fractions (solid line) for the two
583 NH (30°N-90°N, MHD and THD and 00°N-30°N, RGP) and the two SH (30°S-00°S,
584 SMO and 90°S-30°S, CGO) latitudinal bands. The points show the semi-hemispheric
585 monthly mean observations from the 5 AGAGE stations used in the inversion (MHD,
586 THD, RPB, SMO, CGO). Lower panel: HFC-152a semi-hemisphere annualized
587 growth rates are shown as dashed lines (see Rigby et al., 2014 for smoothing method),
588 with the solid blue line and shading showing the global mean and its uncertainty.
589
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591 The strong inter-hemispheric gradient demonstrates that emissions are predominantly
592 in the NH, as has been illustrated for many other purely anthropogenic trace gases (Prinn et
593 al., 2000). The globally averaged mole fraction in the lower troposphere in 2014 is estimated
594 to be 6.84 ± 0.23 ppt and the annual rate of increase is -0.06 ± 0.05 ppt/yr. As reported by
595 Rigby et al. (2014) the major long lived synthetic greenhouse gases (SGHG) which include
596 CFCs, HCFCs, HFCs and perfluorocarbons (SF_6 and NF_3), as well as CH_3CCl_3 and CCl_4
597 were responsible for 350 ± 10 mW/m² of direct radiative forcing in 2012. The radiative
598 forcing of HFC-152a, determined from the AGAGE 12-box model in this study, was $0.61 \pm$
599 0.02 mW/m² in 2014, which represents only a tiny fraction (~0.2%) of the global radiative
600 forcing of the SGHG.

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602 **5. Top-down Emission Estimates**

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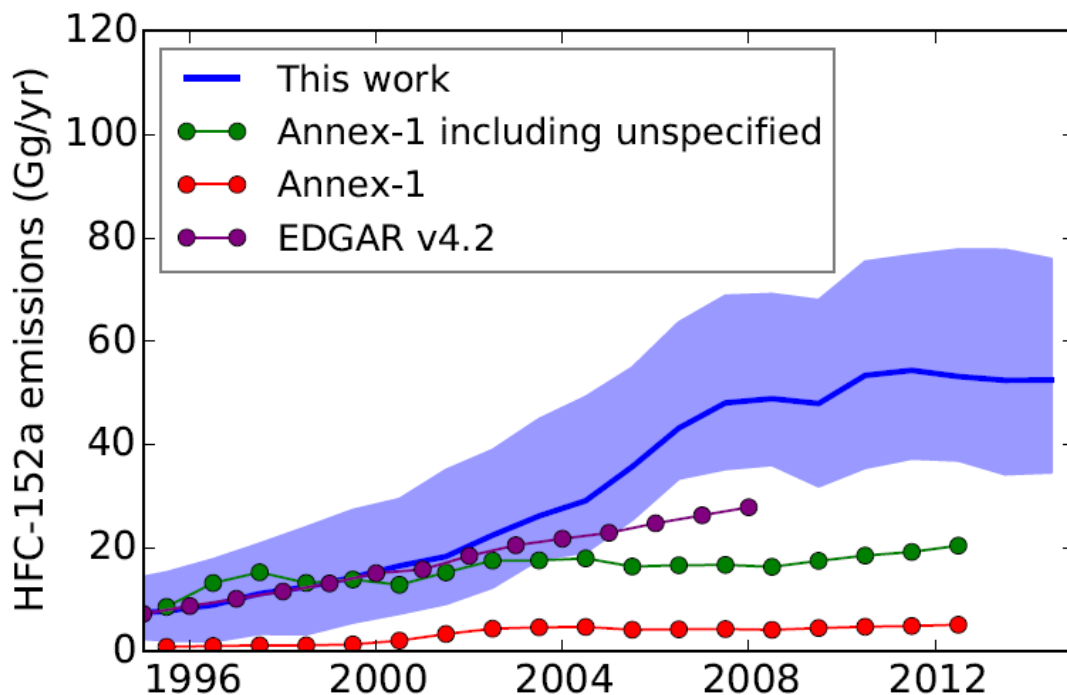
604 *5.1 Global estimates*

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606 Estimated global emissions of HFC-152a using the 12-box model and the reported
607 UNFCCC and EDGAR emission inventories are shown in Figure 7 and Table 2. The blue
608 solid line represents our model-derived emissions, with the 1 σ error band shown by the
609 shaded areas. Model derived emissions grew steadily from 1995-2007 with a non-
610 statistically-significant decrease in emissions in 2009 to 48 ± 16.4 Gg/yr, during the time of
611 the economic downturn in 2008-2009. The mean emission reached a maximum of 54.4 ± 17.1
612 Gg/yr in 2011, followed by a period of relatively stable emissions, the mean showing a slight
613 decline to 52.5 ± 20.1 Gg/yr (7.2 ± 2.8 Tg-CO₂ eq/yr) in 2014.

614

615



616 Figure 7. HFC-152a emissions estimates derived from observations (blue line and shading,
 617 1σ uncertainty) and inventories. The purple line shows the global emissions estimates from
 618 EDGAR (v4.2), the red line shows the emissions reported to the UNFCCC and the green line
 619 shows emissions calculated from all data reported to UNFCCC, including allowance for the
 620 HFC-152a component of unspecified emissions.

621
 622 The data shown in column 3 of Table 2 are the totals of submissions by the national
 623 governments to the UNFCCC (Rio Convention) as reported in Table 2(II) s1 in the Common
 624 Reporting Format (CRF), available on the UNFCCC website (http://unfccc.int/national-reports/annex_ghg_inventories/national_inventories_submissions/items/8108.php). The values
 625 were taken from the 2014 database and cover years 1995, the baseline year for submissions,
 626 to 2012. In addition to reporting calculated emissions of HFCs 23, 32, 125, 134a, 143a, 152a,
 627 227ea, 236fa, 245ca and 43-10mee, individually, many countries also included "unspecified"
 628 emissions in this database (as the sum of their CO₂ equivalents). Where the unspecified
 629 component was small in relation to the national specified emissions, it was disaggregated by
 630 assuming that it had the same fractional contribution of each HFC as reported in the specified
 631 components (adjusted for their CO₂ equivalence). However, in the US, although values of
 632 emissions of several HFCs are calculated specifically for the individual substances, HFCs
 633 152a, 227ea, 245ca and 43-10mee are shown in the database as "commercially confidential"
 634 and their emissions apparently constitute the substantial aggregated "unspecified" emissions
 635 reported. Hence, for the US, these unspecified annual emissions were divided only between
 636 HFCs 152a, 227ea, 245ca and 43-10mee, assuming the same ratio as their reported global
 637 emissions, all expressed as CO₂ equivalents. The values shown in column 4 of Table 2 are the
 638

639 global totals of HFC-152a after adjusting in these ways for the quantities included in
640 “unspecified” emissions.

641 The additional component of US emissions makes a substantial contribution to the
642 very large difference between the UNFCCC data as reported and the adjusted values. This is
643 partly due to the low global warming potential of HFC-152a (a factor of 10 lower than other
644 HFCs) which magnifies its mass component in the 8200 Gg CO₂ equivalent of US
645 “unspecified” emissions.

646 The AGAGE observation based global emissions are substantially higher than the
647 emissions calculated from the UNFCCC GHG reports (2014 submission). It is not
648 unreasonable that UNFCCC-reported emissions are lower than the AGAGE global emission
649 estimates, since countries and regions in Asia (e.g. China, Indonesia, Korea, Malaysia, the
650 Philippines, Taiwan, Vietnam), the Indian sub-continent (e.g. India, Pakistan), the Middle
651 East, South Africa and Latin America do not report to the UNFCCC. Where we include the
652 HFC-152a component of unspecified emissions (green line in Figure 7) results are consistent
653 within the error bars until approximately 2003 to 2005 when they start to diverge (UNFCCC
654 + ‘unspecified’ lower). From 1996-2002, estimated emissions from EDGAR (v4.2) are
655 generally consistent with AGAGE emissions, but then begin to diverge with EDGAR
656 emissions 22 Gg below 2008 AGAGE emissions, the last year for which EDGAR reports
657 emissions.

658

659 *5.2 Regional Emissions of HFC-152a Inferred for Europe, United States, Asia* 660 *and Australia*

661

662 Lunt et al. (2015) have reported global and regional emissions estimates for the most
663 abundant HFCs, based on inversions of atmospheric mole fraction data, aggregated into two
664 categories; those from Annex 1 countries and those from non-Annex 1 countries. The
665 inversion methodology used the NAME model to simulate atmospheric transport close to the
666 monitoring sites, and the Model for Ozone and Related chemical Tracers (MOZART,
667 Emmons et al., 2010) to simultaneously calculate the effect of changes to the global
668 emissions field on each measurement site. The model sensitivities were combined with a
669 prior estimate of emissions (based on EDGAR) and the atmospheric measurements, in a
670 hierarchical Bayesian inversion (Ganesan et al., 2014), to infer emissions.

671 Using this method we infer emissions estimates for the entire world, Europe, North
672 America and East Asia. Table 3 lists our estimated regional emissions in Gg/yr averaged

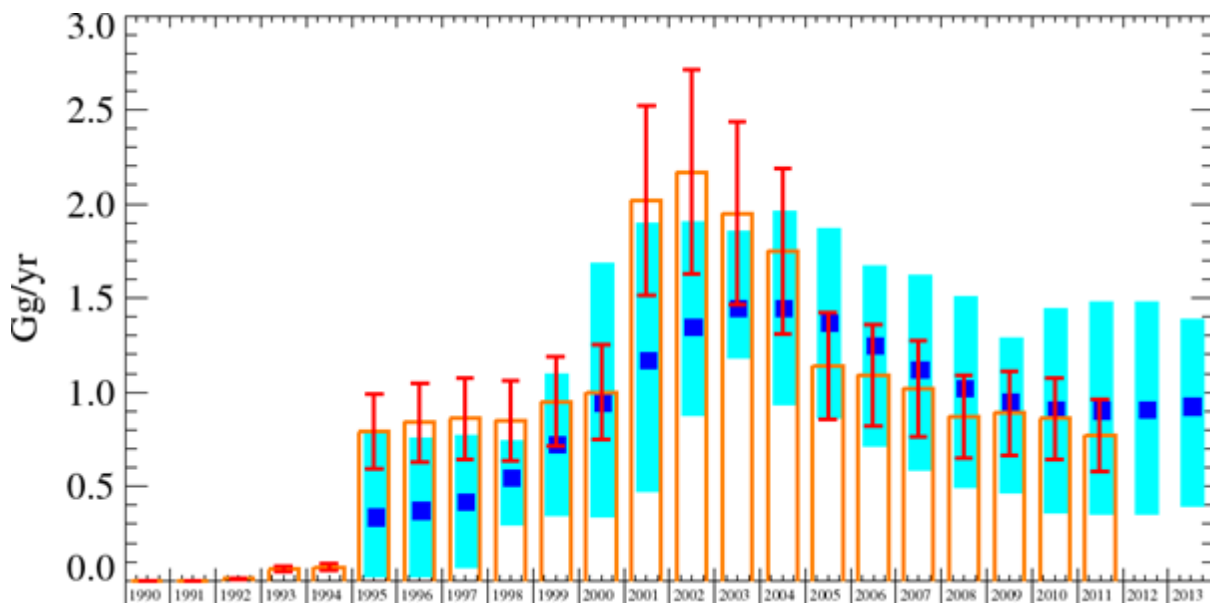
673 across two time periods: 2007–2009 and 2010–2012, together with our global emission
674 estimates averaged over the same time periods from the 12-box model. It is apparent that
675 North American average annual emissions (~30 Gg) are the major contributor to the global
676 total with Europe contributing annual average emissions from about 5–6 Gg/yr. East Asia and
677 Europe contribute emissions of ~7 Gg/yr and ~6 Gg/yr, respectively to the global total. The
678 2007–2009 North American emission estimate of 28 Gg/yr agrees within the uncertainties of
679 HFC-152a emission estimates reported in Barletta et al. (2011) and Simmonds et al. (2015).
680 The North American estimate indicates one reason why the UNFCCC reported amount
681 appears to be so low; more than half the global emissions appear to come from this
682 continental region, yet the UNFCCC reports do not include specific HFC-152a emissions
683 from the US.

684 *5.2.1 InTEM North-West Europe (NWEU) estimated emissions from Mace Head observations*

685
686 The HFC-152a perturbations above baseline, observed at Mace Head, are driven by
687 emissions on regional scales that have yet to be fully mixed on the hemisphere scale. The
688 Mace Head observations are coupled with NAME model air history maps using the inversion
689 system InTEM to estimate surface emissions across NWEU (Manning et al., 2011). NWEU is
690 defined as United Kingdom, Ireland, Germany, France, Benelux and Denmark.

691 As shown in Fig. 8, the NWEU emission estimates for HFC-152a from InTEM (rolling 3-yr
692 averages) agree to within inversion uncertainties with the UNFCCC data (2013 submission)
693 for most years. The estimates of NWEU emissions grew steadily from 1995 reaching a
694 maximum emission of 1.6 ± 0.21 Gg/yr in 2003 with a subsequent decline to 0.98 ± 0.34
695 Gg/yr in 2013.

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707 Figure 8: Emission (Gg/yr) estimates for HFC-152a from North-West Europe. The blue
 708 uncertainty bars represent the 5th and 95th percentiles of the InTEM estimates (rolling 3-yr
 709 averages). The orange bars and associated uncertainty are the UNFCCC inventory estimates
 710 for the NWEU region. (25% uncertainty is estimated by the UK in their National Inventory
 711 Report (NIR) submission to the UNFCCC, the same uncertainty was assumed for North-West
 712 Europe given the lack of additional information).

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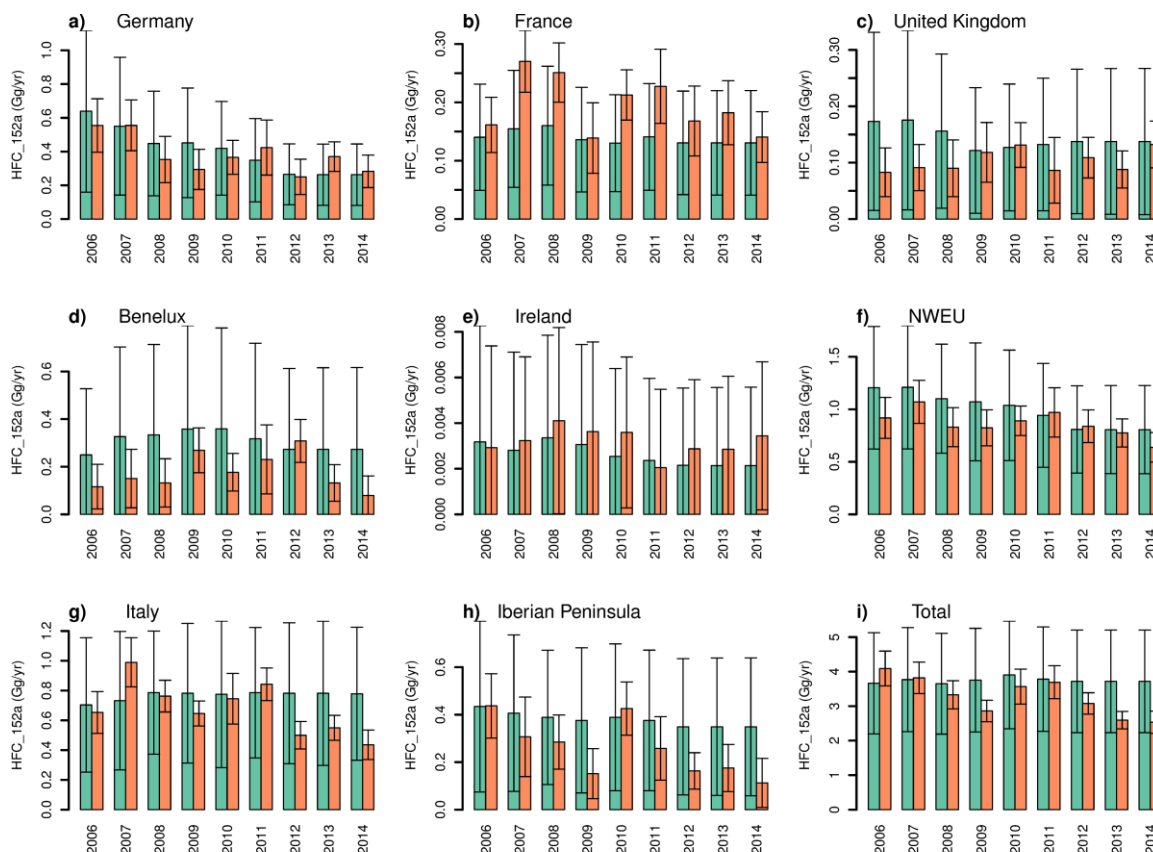
714 *5.2.2 European estimated emissions from European observations at Mace Head*
 715 *Jungfraujoch and Mt. Cimone*

716

717 The temporal evolution of emission estimates for different European regions are given
 718 in Figure 9. In contrast to the InTEM estimates the Bayesian inversion derived emissions in
 719 NWEU were slightly smaller than the UNFCCC estimate and showed a continued decrease
 720 until 2014. Total emissions in the inversion domain ranged from 4 ± 0.5 Gg/yr (2σ
 721 confidence range) for 2006 to only 2.5 ± 0.2 Gg/yr in 2014. This is considerably smaller than
 722 the European Annex I estimate given in section 5.2, but covers a significantly smaller
 723 geographical region. The estimate given in section 5.2 encompassed all countries in Europe
 724 extending beyond the bounds of the area indicated in Fig. 1 (red box). The steady decline in
 725 emissions was interrupted by a local maximum in the years 2010–2012, when emissions
 726 reached 3.6 ± 0.5 (Gg/yr). A minimum in the posterior emissions can be seen in 2009 and was
 727 most pronounced for the Iberian Peninsula, Italy, France and Germany, which might indicate
 728 the influence of the European recession in 2008–2009. For NWEU the emission estimate
 729 remains slightly below the UNFCCC estimates and those estimated by InTEM, but support
 730 the declining trend in European emissions. Despite the fact that Italy does not report HFC-
 731 152a emissions to the UNFCCC, the largest by country emissions were estimated for Italy

732 (up to 1 Gg/yr in 2007). However, a strong decline in these emissions after 2011 was
 733 established here. Similar values for Italian HFC-152a emissions were reported by Brunner et
 734 al. (2012) using observations from Jungfraujoch and Mace Head (but not Mt. Cimone) in an
 735 extended Kalman Filter inversion.

736



737
 738 Figure 9: HFC-152a emission estimates for different European regions using the Bayesian
 739 regional inversion (orange bars) and prior estimates as reported to UNFCCC (green bars).
 740 Error bars indicate 2σ confidence levels. Total prior uncertainties were set to 20% of the total
 741 domain emissions, which may result in different levels of relative uncertainty for each
 742 country/region. Note that prior estimates for Italy were taken from EDGAR instead. Prior
 743 values for 2012 were repeated for each region after 2012.

744

745 5.2.3 US estimated emissions.

746

747 Estimates of North American emissions have been reported by several groups (see
 748 also estimates from this study in Table 3). Millet et al. (2009) report average US emissions
 749 for 2004-2006 of 7.6 Gg (4.8-10 Gg) compared with the UNFCCC average 2005-2006
 750 estimate of 12.3 Gg calculated from UNFCCC data. Miller et al. (2012) provided HFC-152a
 751 emissions estimates averaged from 2004-2009 of 25 Gg (11-50 Gg). Barletta et al. (2011)
 752 reported a 2008 HFC-152a emission estimate of 32 ± 4 Gg. In a recent investigation of the
 753 surface-to-surface transport of HFC-152a from North America to Mace Head, Ireland, an

754 interspecies correlation method with HFC-125 as the reference gas was also used to estimate
755 North American emissions primarily from the eastern seaboard region. The average 2008
756 HFC-152a emission estimate was 31.3 ± 5.9 Gg (Simmonds et al., 2015); in very close
757 agreement with the estimate from Barletta et al. (2011). HFC-152a emission estimates for
758 2005 (10.1 Gg) and 2006 (12.5 Gg) reported by Stohl et al. (2009) are close to the
759 (recalculated) UNFCCC estimates in those years.

760 If the sources of emissions from the US were solely technical aerosols and
761 construction foam, emissions would be expected to be far lower. These were the historic
762 uses in Europe and Japan and resulted in emissions ten times less than those estimated
763 for the US. However, in the US, do-it-yourself (DIY) refilling of car air conditioners is not
764 only permitted but thriving (Zhan et al., 2014), with an estimated 24 million DIY
765 refilling operations attempted each year. The practice is banned in Europe (OJ., 2014).
766 Furthermore, there is ample evidence online that HFC-152a is extensively used in DIY
767 refilling on account of its lower cost. It is a technically suitable replacement for HFC-134a,
768 although there are safety concerns of importance to vehicle manufacturers (Hill., 2003). If the
769 quantities estimated by Zhan et al. 2014 were met using HFC-152a diverted from the retail
770 trade in technical aerosols, some 10 to 20 Gg/yr of HFC-152a could be released into the
771 atmosphere from this source alone.

772 773 *5.2.4 East Asian emissions* 774

775 Emissions of HFC-152a from China were estimated to be 4.3 ± 2.3 Gg/yr in 2004-
776 2005 (Yokouchi et al., 2006), 3.4 ± 0.5 Gg/yr in 2008 (Stohl et al., 2010) and 5.7 (4.3-7.6)
777 Gg/yr in 2008 (Kim et al., 2010). Li et al. (2011) using an interspecies correlation method
778 also reported emission estimates for East Asia (China, South Korea and Taiwan, with HCFC-
779 22 as the reference tracer) and Japan (reference tracer HFC-134a) for the period between
780 November 2007 and December 2008. For China, emissions were estimated to be 5.4 (4-7.4)
781 Gg/yr. In contrast the Taiwan region Korea and Japan had lower estimated emissions totalling
782 1.39 Gg/yr. These estimates are within the uncertainties of our East Asia emissions reported
783 in section 5.2 and Table 3.

784 Yao et al. (2012), using the interspecies correlation method with carbon monoxide as
785 the reference tracer, reported more recent Chinese emissions of 2 ± 1.8 Gg/yr in 2010-2011.
786 This would imply some reduction in Chinese emissions compared with earlier years.

787

788 5.2.5 Australian HFC-152a emissions from Cape Grim data

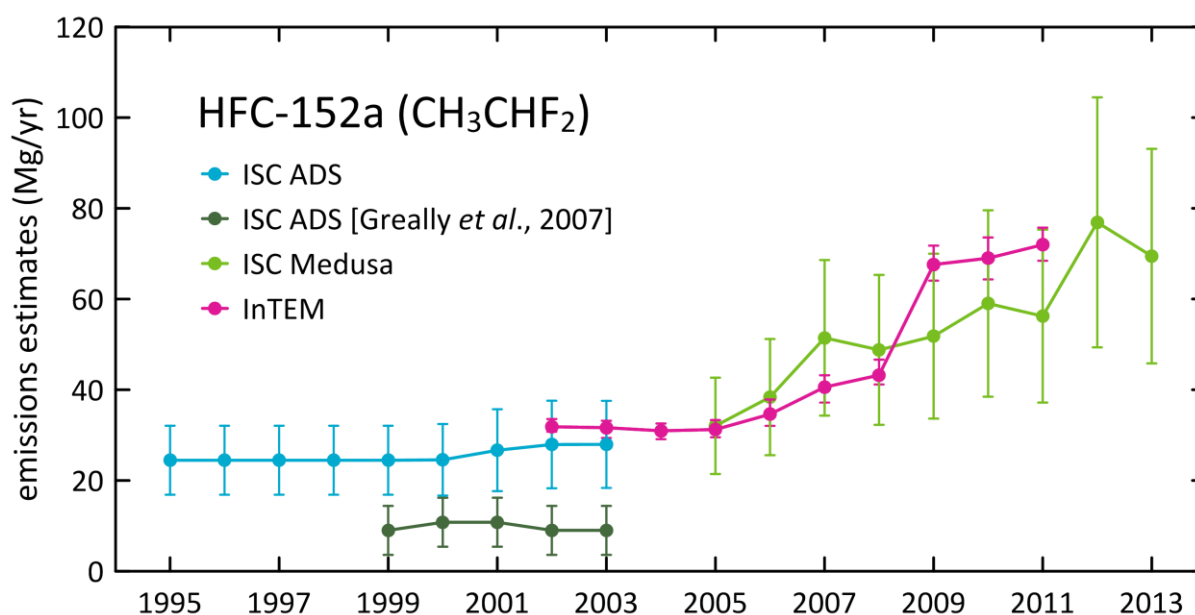
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790 SE Australian emissions of HFC-152a are estimated using the positive enhancements
791 above baseline or background concentrations observed at Cape Grim using interspecies
792 correlation with CO as the reference species (ISC: Dunse et al., 2005; Grealley et al., 2007)
793 and inverse modelling (InTEM: Manning et al., 2003, 2011). Figure 2a (CGO) shows an
794 overall increase in the magnitude of HFC-152a pollution episodes, presumably due to
795 increasing regional emissions. Detailed analysis of these pollution episodes using air mass
796 back trajectories shows clearly that the HFC-152a pollution seen at Cape Grim originates
797 largely from Melbourne and the surrounding Port Phillip region.

798 Australian HFC-152a emissions of 5-10 Mg/yr via interspecies correlation (ISC) have
799 been reported for the period 1998-004, although it was noted that these emission estimates
800 were near the detection limit of the ISC method (Grealley et al., 2007). Recently, significant
801 improvements have been made to this ISC method, including a revised (upward) CO
802 emissions inventory for the Melbourne/Port Phillip region, exclusion of high CO events in the
803 Cape Grim in situ CO record, resulting from CO emissions from biomass burning and coal
804 combustion in the Latrobe Valley (east of Port Phillip) and a revised (upward) population-
805 based scaling factor (5.4), used to convert Melbourne/Port Phillip emissions to Australian
806 emissions (Fraser et al., 2014a, b). Each of these changes to the ISC method resulted in
807 higher trace gas emission estimates. The revised (compared to Grealley et al., 2007) Australian
808 HFC-152a emission estimates from the ISC method are shown in the 2nd column of Table 4
809 and in Figure 10 as 3-year running averages.

810 The InTEM model (Manning et al., 2003, 2011) has been used to derive HFC-152a
811 emissions from Victoria/Tasmania (Fraser et al., 2014a). Annual Australian emissions are
812 calculated from Victoria/Tasmania emissions using a population based scale factor of 3.7 and
813 are shown in Figure 10 and the 3rd column of Table 4, interpolated from rolling 3-year
814 emission estimates. Over the period 2002-2011, the average Australian HFC-152a emissions
815 from ISC and InTEM agree to within 2%. The method for estimating the InTEM
816 uncertainties are discussed above. No additional uncertainty was applied to the estimates
817 through the process of up-scaling from Victoria/Tasmania to Australian totals. The
818 assumption was made that the use of HFC-152a per head of population was identical across
819 Australia as we have no more detailed information.

820



821

822 Figure 10. Australian HFC-152a emissions (Mg/yr) calculated from Cape Grim *in situ*
 823 observations via ISC, using ADS and Medusa data, and inverse modelling using InTEM
 824 (Medusa data). Australian emissions are derived from SE Australian emissions, scaled by
 825 population (see text). Uncertainties are 25th-75th percentiles (InTEM) and 1 s.d. (ISC).

826 Australian HFC-152a emissions have increased steadily from 25 Mg/yr in the late-
 827 1990s to over 60 Mg/yr in the late-2000s. The 2012 and 2013 emissions have been estimated
 828 from Cape Grim data by ISC at 77 and 69 Mg, respectively. Australian HFC-152a emissions
 829 (1998-2004) are 25-30 Mg, significantly higher than estimated (5-10 Mg/yr) in Greally et al.
 830 (2007), resulting from improvements in the ISC method (see above).

831 Compared to the global values derived above, Australian emissions are 0.1% of global
 832 emissions based on ISC/InTEM data. It is unusual for Australian emissions of an industrial
 833 chemical to be as low as 0.1% of global emissions. For other HFCs, CFCs and HCFCs (for
 834 example HFC-134a, CFC-12, HCFC-22), Australian emissions as fraction of global
 835 emissions are typically 1-2%, similar to Australia's fraction of global gross domestic product
 836 (GDP, 1.9%, 2014) but significantly larger than Australia's fraction of global population
 837 (0.33%, 2014) (Fraser et al., 2014b).

838 The possible reasons for the low Australian HFC-152a emissions (relatively low use
 839 in Australia compared to rest of world) are being investigated. One suggestion (M. Bennett,
 840 Refrigerant Reclaim Australia, personal communication, 2013) is that a significant major-
 841 volume use in other parts of the world for HFC-152a is as an aerosol propellant, a use not
 842 taken up to any significant degree in Australia.

843 6. Conclusions

844 Atmospheric abundances and temporal trends of HFC-152a have been estimated from
845 data collected at the network of eleven globally-distributed monitoring sites. The longest
846 continuous in situ record at Mace Head, Ireland covers a 20-year period from 1994-2014.
847 Other stations within the network have observational records from 9-16 years, with only a
848 short record (2010–2012) at Shangdianizi, China. From selected baseline in situ
849 measurements and measurements of archived air samples dating back to 1978 the long-term
850 growth rate of HFC-152a has been deduced. Analysing the enhancements above baseline
851 coupled with atmospheric transport models permitted us to estimate both regional and global
852 HFC-152a emissions. However, it should be noted that the various models use different
853 domains to obtain regional emissions estimates.

854 The annual average NH (Mace Head + Trinidad Head) baseline mole fraction in 1994
855 was 1.2 ppt reaching an annual average mole fraction of 10.1 ppt in 2014. In the SH (Cape
856 Grim) the annual average mole fraction increased from 0.84 ppt in 1998 to 4.5 ppt in 2014.
857 Using the global average mole fraction obtained from the AGAGE 12-box model we estimate
858 that the HFC-152a contribution to radiative forcing was 0.61 ± 0.02 mW/m² in 2014. Since
859 the first in situ measurements in 1994 the global annual growth rate of HFC-152a has
860 increased to a maximum annual growth rate in 2006 of 0.81 ± 0.05 ppt/yr. More recently the
861 average annual growth rate has slowed to 0.38 ± 0.04 ppt/yr in 2010, and become negative,
862 with a growth rate in 2013-2014 of minus 0.06 ± 0.05 ppt/yr.

863 Global HFC-152a emissions increased from 7.3 ± 5.6 Gg/yr in 1994 to 52.5 ± 20.15
864 Gg/yr in 2014. Global emissions are dominated by emissions from North America with this
865 region being responsible for ~67% of global emissions in our estimates. Estimates of north-
866 western European emissions of ~0.9 Gg/yr, (2010-2012 average) agree within the
867 uncertainties for the two regional models (see sections 3.3 and 3.4) and overlap with the
868 UNFCCC inventory. For the combined Eulerian and Lagrangian models (see 3.2 and Table
869 3), that encompass all European countries, we derive a 2010-2012 average emission of 5.2
870 Gg/yr. East Asian countries contribute 1 Gg/yr (Annex 1) and 6 Gg/yr (Non-Annex 1) to the
871 global total (2010-2012 averages). All of the models studies indicate a current declining trend
872 in European and Asian emissions.

873 Substantial differences in emission estimates of HFC-152a were found between this
874 study and those reported to the UNFCCC which we suggest arises from underestimated North
875 American emissions and undeclared Asian emissions; reflecting the incomplete global

876 reporting of GHG emissions to the UNFCCC and/or biases in the accounting methodology.
877 Ongoing, continuous, and accurate globally and regionally distributed atmospheric
878 measurements of GHGs, such as HFC-152a, are required for ‘top-down’ quantification of
879 global and regional emissions of these gases, thereby enabling improvements in national
880 emissions inventories, or ‘bottom-up’ emissions data collected and reported to the UNFCCC
881 (Weiss and Prinn, 2011).

882

883 **Data availability**

884 The entire ALE/GAGE/AGAGE data base comprising every calibrated measurement
885 including pollution events is archived on the Carbon Dioxide Information and Analysis
886 Center (CDIAC) at the U.S. Department of Energy, Oak Ridge National Laboratory.

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888

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1238 Table 1. Overview of the 11 Measurement Stations used in this Study, their Coordinates and
 1239 Periods for which Data are available.

Station	Latitude	Longitude	ADS Data*	Medusa Data**
Ny-Ålesund, Norway ¹	78.9° N	11.9° E	2001-2010	September 2010-present
Mace Head, Ireland ¹	53.3° N	9.9° W	1994-2004	June 2003-present
Jungfrauoch, Switzerland ¹	46.5° N	8.0° E	2000-2008	May 2008-present
Monte Cimone, Italy ²	44.2° N	10.7° E		June 2001-present ²
Trinidad Head, California ¹	41.0° N	124.1° W		March 2005 -present
Shangdianzi, China ^{1,3}	40.4° N	117.7° E		May 2010-August 2012
Gosan, Jeju Island, Korea ¹	33.2° N	126.2° E		November 2007-present
Hateruma, Japan ²	21.1° N	123.8° E		May 2004-present ²
Ragged Point, Barbados ¹	13.2° N	59.4° W		May 2005-present
Cape Matatula, Samoa ¹	14.2° S	170.6° W		May 2006-present
Cape Grim, Tasmania ¹	40.7° S	144.7° E	1998- 2004	Jan 2004-present

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¹ AGAGE stations

² Affiliated stations use a different pre-concentration system (non-Medusa) than the AGAGE stations, but comparable GC-MS analytical instruments (see Yokouchi et al., 2006, Maione et al., 2014).

³ Shangdianzi was only operational for a short period and is not included in the modelling studies.

* Period of HFC-152a data record using ADS-GC-MS.

** Period of HFC-152a data record using Medusa-GC-MS.

1265 Table 2. Estimates of global emissions of HFC-152a (Gg/yr \pm 1 σ) based on AGAGE in situ
 1266 measurements using the AGAGE 2-D 12-box model. Emission inventories as reported in
 1267 UNFCCC National Inventory Reports (2014 submission), EDGAR (v4.2) database and
 1268 recalculated from the UNFCCC data as described in the text.
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Year	AGAGE (Gg/yr)	UNFCCC as reported (Gg/yr)	EDGAR (4.2) (Gg/yr)	UNFCCC including “unspecified” contribution (Gg/yr)
1994	7.3 \pm 5.6			
1995	7.9 \pm 7.4	1.0	7.3	8.8
1996	9.1 \pm 8.4	1.1	8.9	13.3
1997	11.3 \pm 8.6	1.3	10.3	15.4
1998	12.5 \pm 10.9	1.2	11.7	13.3
1999	14.4 \pm 11.2	1.4	13.2	14.0
2000	16.6 \pm 12.2	2.2	15.2	13.0
2001	18.4 \pm 13.4	3.5	15.9	15.4
2002	22.5 \pm 14.7	4.5	18.6	17.6
2003	26.3 \pm 15.3	4.7	20.6	17.7
2004	29.2 \pm 15.6	4.8	21.7	18.1
2005	35.8 \pm 14.7	4.3	23.0	16.5
2006	43.3 \pm 14.9	4.4	24.9	16.7
2007	48.1 \pm 17.6	4.4	26.4	16.8
2008	48.9 \pm 16.7	4.3	28.0	16.4
2009	48.0 \pm 16.4	4.6		17.6
2010	53.4 \pm 17.5	4.9		18.6
2011	54.4 \pm 17.1	5.0		19.3
2012	53.2 \pm 18.5	5.2		20.5
2013	52.5 \pm 17.8			
2014	52.5 \pm 20.1			

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 1272 Table 3. Annex 1 and non-Annex 1 global and regional emissions in Gg/yr averaged over two
 1273 3-year periods. Values in the final column are from the 12-box model, all other values are
 1274 from the combined Eulerian and Lagrangian model of Lunt et al. (2015). The global
 1275 estimates from the 12-box model are not in exact agreement with the combined Annex I and
 1276 non-Annex I emissions reported in Lunt et al. 2015. However, this is not unexpected, given
 1277 the vastly different transport and inversion models used to estimate these terms. We note that
 1278 the uncertainty range of the combined Annex I and non-Annex I estimates does overlap with
 1279 the uncertainty range from the 12-box model, and a similar growth in emissions is seen across
 1280 the two averaging periods.
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3-year Averages	Europe Annex 1	North America Annex 1	East Asia Annex 1	East Asia Non-Annex 1	GLOBAL Annex 1	GLOBAL Non-Annex 1	GLOBAL 12-Box model
	2007–2009	6.4 (5.2-7.5)	28.0 (22.5-33.4)	0.4 (0.2-1.2)	5.8 (4.5-7.5)	35.2 (27.7-42.6)	6.6 (4.3-9.2)
2010–2012	5.2 (4.1-6.4)	31.6 (24.5-38.6)	1.0 (0.5-1.6)	6.0 (4.3-8.2)	40.2 (31.3-49.3)	6.6 (3.9-9.8)	53.9 (43.0-67.3)

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1283 Table 4. Australian HFC-152a emissions (Mg/yr, 3-year running averages) calculated from
 1284 Cape Grim *in situ* observations via ISC (ADS and Medusa data, uncertainty: ± 1 s.d.) and
 1285 inverse modelling using InTEM (Medusa data, range: 25th-75th percentiles); ISC, NAME
 1286 averages weighted by uncertainties, ISC InTEM average for 2004 is based only on InTEM
 1287 data.

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YEAR	ISC	InTEM	ISC and InTEM average	ISC/InTEM ratio
1999	24 \pm 7			
2000	25 \pm 8			
2001	27 \pm 9			
2002	28 \pm 10	32 (31-34)	31 \pm 2	
2003	28 \pm 10	32 (29-33)	31 \pm 4	0.88
2004	29 \pm 10	31 (29-33)	31 \pm 2	0.94
2005	32 \pm 10	31(30-33)	31 \pm 4	1.03
2006	38 \pm 10	35 (32-38)	35 \pm 6	1.09
2007	51 \pm 15	41 (37-43)	42 \pm 6	1.24
2008	49 \pm 15	43 (41-47)	44 \pm 5	1.14
2009	52 \pm 15	68 (64-72)	65 \pm 8	0.76
2010	59 \pm 20	69 (64-74)	67 \pm 10	0.86
2011	56 \pm 15	72 (68-76)	69 \pm 7	0.78
2012	77 \pm 25			
2013	69 \pm 24			

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