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Spring and summer contrast in new particle formation over nine forest areas in North America

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Abstract

Recent laboratory chamber studies indicate a significant role for highly oxidized low volatility organics in new particle formation (NPF) but the actual role of these highly oxidized low volatility organics in atmospheric NPF remains uncertain. Here, particle

- size distributions (PSDs) measured in nine forest areas in North America are used to characterize the occurrence and intensity of NPF and to evaluate model simulations using an empirical formulation in which formation rate is a function of the concentrations of sulfuric acid and low volatility organics from alpha-pinene oxidation (Nucl-Org), and using an ion-mediated nucleation mechanism (excluding organics; Nucl-IMN). On
- ¹⁰ average, NPF occurred on ~70% of days during March for the four forest sites with springtime PSD measurements, while NPF occurred on only ~10% of days in July for all nine forest sites. Both Nucl-Org and Nucl-IMN schemes capture the observed high frequency of NPF in spring, but the Nucl-Org scheme significantly over-predicts while the Nucl-IMN scheme slightly under-predicts NPF and particle number concentrations
- ¹⁵ in summer. Statistical analyses of observed and simulated ultrafine particle number concentrations and frequency of NPF events indicate that the scheme without organics agrees better overall with observations. The two schemes predict quite different nucleation rates (including their spatial patterns), concentrations of cloud condensation nuclei, and aerosol first indirect radiative forcing in North America, highlighting the need to raduae NPE uncertainties in radianal and global earth autom medale.

²⁰ to reduce NPF uncertainties in regional and global earth system models.

1 Introduction

Particle number concentration is an important factor in the health and climate impacts of atmospheric aerosols. High number concentrations of ultrafine particles may lead to adverse health effects (Knibbs et al., 2011). As cloud condensation nuclei (CCN),

²⁵ atmospheric particles modify cloud properties and precipitation and thus affect the hydrological cycle and climate indirectly. Aerosol indirect radiative forcing (IRF) remains



a major uncertainty in assessing climate change (IPCC, 2013; Carslaw et al., 2013). Secondary particles formed via nucleation dominate the global total particle number abundance (Spracklen et al., 2008; Pierce and Adams, 2009; Yu and Luo, 2009), contributing as high as 80–95% to total particle (>3 nm) and 50–80% to CCN number

- ⁵ concentrations in most parts of the lower troposphere (Yu and Luo, 2009). Global simulations also indicate that the aerosol IRF is quite sensitive to nucleation parameterizations (Wang and Penner, 2009; Kazil et al., 2010; Yu et al., 2012). Furthermore, as a result of its dependence on meteorological conditions, emissions, and chemistry, new particle formation (NPF) is an important physical process involved in a number of
- climate feedback loops. Laboratory experiments indicate that ammonia, amines, ions, certain organic compounds and oxidants can all contribute to NPF. But there are numerous chemical reactions and physical processes involved, so there is no single unified theory that accurately describes NPF and the levels of enhancements due to different species vary widely among various studies (e.g., Zhang et al., 2004; Erupe et al., 2011; Chen et al., 2012; Dawson et al., 2012; Almeida et al., 2013; Berndt et al., 2014;
- ¹⁵ 2011; Chen et al., 2012; Dawson et al., 2012; Almeida et al., 2013; Berndt et al., 2014 Riccobono et al., 2014; Glasoe et al., 2015).

The indication that nucleation is significantly enhanced by organic species formed via oxidation of biogenic volatile organic compounds (VOCs; e.g., Riccobono et al., 2014) is of interest as it may have important implications for the interactions of anthropogenic

- and biogenic emissions and the associated climate forcing. Nevertheless, the various roles of organics and their oxidation products in NPF in the real atmosphere remains an active research area. Particle size distributions (PSDs) have been measured in a variety of environments around the globe and many of these measurements have been used to study NPF events (e.g., Kulmala et al., 2004). To evaluate the potential role
- of oxidation products of biogenic VOCs in NPF, PSDs taken over forested areas are particularly useful. In the present study, we analyze NPF events and non-events based on PSDs measured over nine forest areas in North America (NA) and compare them to model simulations with and without including organics in the nucleation rate calculation. Since biogenic VOC emissions and their oxidation are strongest in the summer, we use



the observed spring and summer contrast in NPF events to study the possible role of organics in NPF in NA and evaluate our current understanding of NPF processes in the atmosphere.

2 Methods

5 2.1 Measurements

PSDs from ~ 3–10 nm to >~ 100 nm have been widely used to identify and to study particle nucleation and growth events in the atmosphere (e.g., Kulmala et al., 2004). In this study, we focus on PSDs observed in various years in eight US and one southern Canadian forest sites summarized in Table 1. The locations of these sites are marked
on a Google map (http://maps.google.com) in Fig. 1. Egbert (EGB) is surrounded by a mixture of forests and farmland which is subject to a strong biogenic influence under certain conditions (Slowik et al., 2010), while all other eight sites are directly located in forest or mountain areas. PSDs have been measured at these nine sites with Scanning Mobility Particle Sizers (SMPSs) during different time periods. Sites 1–4 have at least one year of continuous PSD data while sites 5–9 have PSD data only during the summer (for different years). PSDs for one spring and one summer month for Sites 1–4 and for one summer month for Sites 5–9 are given in Fig. 1. It is clear that nucleation in the spring is much more frequent and stronger (i.e., higher concentrations of nucleation mode particles) when compared to summer months. Detailed analysis of these data

 $_{\rm 20}$ $\,$ and comparisons with simulations are given in the Results section.

2.2 Model and simulations

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The modeling work in this study is based on the GEOS–Chem model (e.g., Bey et al., 2001) with an advanced particle microphysics (APM) model incorporated (Yu and Luo, 2009). GEOS-Chem is a state-of-the-art global 3-D model of atmospheric composition driven by assimilated meteorological observations from the Goddard Earth Observing



System (GEOS) of the NASA Global Modeling and Assimilation Office (GMAO). The model has been developed and used by many research groups and contains a number of state-of-the-art modules treating various chemical and aerosol processes (e.g., Bey et al., 2001; Chung and Seinfeld, 2002; Martin et al., 2003; Park et al., 2004; Evans
and Jacob, 2005; Liao et al., 2007; Henze et al., 2008). The APM module in GEOS-Chem is optimized to simulate the formation of secondary particles and their growth to CCN sizes, using high size resolution (30 bins) for particles in the dry size range of 1.2–120 nm diameter. There are an additional 10 bins for 0.12–12 µm particles (Yu and Luo, 2009). More detailed information about GEOS-Chem and updates can be found at the model website (http://geos-chem.org/).

To study the possible role of organics in NPF, we employ the organics-mediated nucleation parameterization (Nucl-Org) derived from a CLOUD chamber study as given in Riccobono et al. (2014),

$J_{1.7} = k_m \times [H_2 SO_4]^2 \times [BioOxOrg]$

¹⁵ where $J_{1.7}$ is the formation rate (# cm⁻³ s⁻¹) of particles of ~ 1.7 nm, k_m is the fitting pre-factor with a value of 3.27×10^{-21} cm⁶ s⁻¹ (90% confidence interval: 1.73×10^{-21} to 6.15×10^{-21} cm⁶ s⁻¹), [H₂SO₄] and [BioOxOrg] are the gas-phase concentrations (# cm⁻³) of H₂SO₄ and biogenic oxidized organic (BioOxOrg) vapors, respectively. In the chamber study reported in Riccobono et al. (2014), BioOxOrg molecules are organic compounds of relatively low volatility from the oxidation of pinanediol (a firstgeneration oxidation product of α -pinene) and represent later-generation oxidation products of biogenic monoterpenes.

In GEOS-Chem v8-03-02 on which this study is based, reactive VOCs are grouped into six categories (VOC_{*i*}, *i* = 1–6), with VOC₁ = α -pinene + β -pinene + sabinene + Δ -

²⁵ 3 carene + terpenoid ketones; VOC_2 = limonene; $VOC_3 = \alpha$ -terpinene + γ -terpinene + terpinolene; VOC_4 = myrcene + terpenoid alcohols + ocimene; VOC_5 = sesquiterpenes; and VOC_6 = isoprene. Yu (2011) extended the two-product secondary organic



(1)

aerosol (SOA) formation model (Chung and Seinfeld, 2002) to predict concentrations of extremely low volatile secondary organic gas ([LV-SOG]) from successive oxidation aging of the first generation oxidation products of various VOCs, i.e., semivolatile SOG (SV-SOG) and medium-volatile SOG (MV-SOG). LV-SOG has a saturation vapor con-⁵ centration in the range of ~ 0.0001–0.03 μ g m⁻³, corresponding to the extremely low volatile SOG reported in Ehn et al. (2014). The production rate of LV-SOG depends on the saturation vapor concentrations of MV-SOG and SV-SOG and temperature (Yu, 2011).

To use Eq. (1) to calculate organics-mediated nucleation, we separate LV-SOG from the oxidation products of VOC₁ (named LV-SOG_{*α*-pinene} thereafter) from those associated with VOC₂₋₆. LV-SOG_{*α*-pinene}, a major component of total LV-SOG, corresponds well to the BioOxOrg reported in the chamber study of Riccobono et al. (2014). For comparison, simulations based on the ion-mediated nucleation of the H₂SO₄-H₂O binary system (Nucl-IMN; Yu, 2010) are also presented. Nucleation rates based on the Nucl-Org scheme depend on [H₂SO₄] and [LV-SOG_{*α*-pinene}] only while those based

- on the Nucl-IMN scheme depend non-linearly on $[H_2SO_4]$, temperature (*T*), relative humidity (RH), ionization rate, and surface area of pre-existing particles (Yu, 2010). Based on the CLOUD chamber study (Almeida et al., 2013; Riccobono et al., 2014), IMN is significant or dominant when $J_{1.7}$ is below ~ 1 cm⁻³ s⁻¹ but neutral nucleation
- ²⁰ takes over when $J_{1.7} > \sim 1 \text{ cm}^{-3} \text{ s}^{-1}$. Nucl-IMN calculates formation rates of particles at wet sizes of ~ 1.5 nm. The nucleated particles are added into the first bin (wet size ~ 1.5 nm) of the secondary particles, and the processes to grow them to large sizes considered in GEOS-Chem/APM include kinetic condensation of H₂SO₄ and total LV-SOG, equilibrium uptake of nitrate and ammonium, partitioning uptake of SV-SOG and MV-SOG, and self-coagulation (Yu, 2011).

The horizontal resolution of GEOS-Chem employed for this study is $2^{\circ} \times 2.5^{\circ}$ and there are 47 vertical layers (with 14 layers from surface to ~ 2 km above the surface). The emission inventories and schemes used can be found at the GEOS-Chem website and have also been summarized in Yu and Hallar (2014). Biogenic VOC emissions from



terrestrial vegetation are based on the Model of Emissions of Gases and Aerosols from Nature (MEGAN; Guenther et al., 2006) which computes emissions for plant functional types as a function of temperature, solar radiation, leaf area index (LAI), and leaf age. MEGAN does not consider the effect of water stress (drought) which may suppress
 ⁵ biogenic emissions (Pryor et al., 2014).

This study focuses on the NA region, and values of all tracers and relevant parameters in all layers over the nine forest sites marked in Fig. 1 were output at each chemistry time step (30 min) for detailed analysis and comparison with measurements. Since PSD measurements at different sites were taken in different years, multiple-year simulations have been carried out and comparisons are for the same month/year for any given site. Since GEOS-Chem was driven by assimilated meteorology, the key meteorology fields are generally consistent with observations at sites of comparisons (e.g., Yu and Hallar, 2014).

2.3 NPF events and non-events

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- To assess the ability of different nucleation schemes to capture NPF events (or non-events) at various sites, an objective criterion is needed to decide from the evolution of PSDs whether a given day is an event day or not. An NPF event has generally been defined as the appearance of a clear new nucleation mode (<~25 nm) for a significant period of time (hours) accompanied by the growth of the mode diameter during its existence (e.g., Dal Maso et al., 2005). Similar to the method used in Hallar et al. (2011) and Venzac et al. (2008), in this study an NPF event is defined if the dN/dlogDp value (where dN is the number concentration of particles within diameter range dD_p) averaged within three bins near 10 nm maintains a level of 3000 cm⁻³ or
- higher continuously for more than two hours, and the mode diameter grows during the
 period. The growth rate of the nucleation mode was calculated by linearly fitting the
 change of mode diameters (Lehtinen and Kulmala, 2002). To simplify the comparison,
 all days with measurements available but not defined as NPF event days are counted
 as non-event days.



3 Results

VOC emissions in the summer are known to be much higher than VOC emissions in the spring. The high VOC emissions coupled with strong photochemistry and oxidation lead to significantly higher production rates and concentrations of oxidized SOGs of varied volatilities (medium volatile, semivolatile, and low-volatile) in the summer. Both [LV-SOG] and [LV-SOG_{α -pinene}] in the summer are much higher (by a factor of ~ 4–10) than those in the spring while $[H_2SO_4]$ has a similar concentration in the spring as in the summer (not shown). Figure 2 shows horizontal distributions of monthly mean nucleation rates (J) in the boundary layer (0-1 km above the surface) in March and July of 2006 based on two different nucleation schemes (i.e, Nucl-Org and Nucl-IMN). Two 10 nucleation schemes clearly predict significantly different spatial patterns and seasonal variations of nucleation rates. In March, $J_{\text{Nucl-Org}}$ is highest in the southeastern US as a result of high [LV-SOG_{α -pinene}] in the region, reaching 5–10 cm⁻³ s⁻¹ (Fig. 2a), while $J_{\rm IMN}$ has two separate easiern and western nucleation zones with monthly mean J up to ~ 1–2 cm⁻³ s⁻¹ (Fig. 2b). In July, boundary layer $J_{\text{Nucl-Org}}$ (Fig. 2c) is two orders of 15

- magnitude higher than $J_{\rm IMN}$ (Fig. 2d) although both indicate a high nucleation center in the northeastern US surrounding the Ohio valley with other scattered high nucleation zones in southern and western US along the coast. The low $J_{\rm IMN}$ is due to the high temperature in the summer. The shift of high $J_{\rm Nucl-Org}$ nucleation zone from southeastern in March to northeastern in July is mainly associated with the change in the spatial distributions of [LV-SOG_{α -pinene}].
 - Comparisons of model predictions with in situ measurements are critical to evaluate our current understanding of atmospheric nucleation processes and to identify areas requiring further research. The PSDs measured at the nine forest sites, as marked in
- Figs. 1–2, provide a reasonably representative dataset for this purpose. As shown in Fig. 1, all the PSDs have lower cut-off sizes of 10 nm or smaller with some as small as 3–5 nm. Due to wall losses and lower charging and counting efficiencies of small particles, PSDs for mobility diameters smaller than ~ 10 nm may have large uncertainties.



The nano-SMPSs generally have upper cutoff sizes of ~ 100 nm. The GEOS-Chem-APM model simulates PSDs from ~ 1.5 nm to larger than 10 µm. Figure 3 gives an example of the comparison between observed and simulated PSDs during two ten-day periods in March and July of 2006 in Duke Forest, along with time series of the con-⁵ centration of condensation nuclei (CN) between 10 and 100 nm (CN₁₀₋₁₀₀, integrated from PSDs). Similar plots for one spring month and one summer month for sites 1–4 and one summer month for sites 5–9 are given in Figs. S1–S9 in the Supplement. The diurnal profile of CN₁₀₋₁₀₀ is a good indication of new particle formation and growth. Statistical analyses of observed and simulated CN₁₀₋₁₀₀ for all sites are presented later in this section.

To assess the ability of different nucleation schemes to capture NPF events (or nonevents) at various sites, we use the criteria described in Section 2.3 to decide from the evolution of PSDs whether a given day is an event day or not. To illustrate this decision process, we mark the event (\checkmark) and non-event (\checkmark) days determined based on this criterion for the observed and simulated PSDs in Fig. 3. For the ten-day period in March, this criterion gives 8, 9, and 9 event days for the observed NPF in Duke forest (Fig. 3a), simulated with Nucl-Org (Fig. 3c) and simulated with Nucl-IMN (Fig. 3e), respectively. For the 10-day period in July, the corresponding numbers of event days are 0, 10, and 0. The requirement for the dN/dlogD_p at D_p = ~ 10 nm of 3000 cm⁻³ or higher is to filter out weak NPFs such as the one on Day 80 of observed and simulated PSD (Fig. 3a, c, e) and Day 84 of observed PSD (Fig. 3d). The two-hour duration

- requirement is intended to filter out the short and sharp spikes likely as a result of local plumes rather than regional nucleation, such as the one appearing on Days 201 and 202 of the observation. Our visual inspection indicates that this criterion captures most
- ²⁵ of major observed and simulated NPFs. While this criterion is not perfect, it enables us to evaluate objectively the ability of different nucleation schemes to capture NPF events (or non-events) at various sites during long observation periods. A slight change in this criterion (for example, increasing dN/dlogD_p at D_p =10 nm from 3000 to 5000 cm⁻³, or



the duration from two hours to three hours) does not impact the main conclusions of this study.

It is clear from Fig. 3a–b that NPF events observed in Duke Forest are much more frequent and concentrations of nucleation mode particles are much higher in the spring than in the summer. More frequent and stronger observed NPF events in the spring have been reported in many previous field measurement studies (e.g., Dal Maso et al., 2005; Pryor et al., 2010; Kanawade et al., 2011; Pillai et al., 2013; Pierce et al., 2014). Both schemes predict frequent NPF events in March (Fig. 3c, e), which is consistent with observations, although the Nucl-Org scheme generally gives stronger NPF events and higher CN_{10-100} (Fig. 3g). Model simulations based on Nucl-Org scheme (Eq. 1) suggest that frequent (almost daily, Fig. 3d) and stronger NPF (Fig. 3h) should occur in the summer in Duke Forest, which was not observed with measurements (Fig. 3b). The relatively less frequent or lack of events in the summer indicates that, at least, LV-SOG_{*α*-pinene} cannot be the dominant nucleation factor as [LV-SOG_{*α*-pinene]} should

¹⁵ be highest during the summer. However, the Nucl-IMN scheme is consistent with the observed lack of nucleation in the summer in Duke Forest. Note that there is no explicit temperature dependence in Eq. (1), an issue that is discussed later.

Figure 4 presents a statistical analysis and comparison of simulated NPF event days and CN_{10-100} based on two nucleation schemes with observations at four forest sites

- for the spring month and for all nine sites for the summer month. Table 2 gives the averaged values. For the statistical analysis, we look into the fraction (or frequency) of nucleation event days (F_{nucl} ; Fig. 4a spring; Fig. 4b summer), the proportion correct (PC; Fig. 4c–d), Pearson correlation coefficient (r) of the observed and simulated scatterplots of CN₁₀₋₁₀₀ (Fig. 4c–d), and monthly mean CN₁₀₋₁₀₀ (Fig. 4e–f). For the four sites with PSD data in the spring NDE system assured as 81 e.5. 87 and 42.% of days
- sites with PSD data in the spring, NPF events occurred on 81, 65, 87, and 43 % of days for sites DUK, MMSF, SPL, and EGB, respectively (Fig. 4a), with an average value of 69 %. Both nucleation schemes capture most of the NPF events in the spring month (four-site average of 76 % for both Nucl-Org and Nucl-IMN). In July, there are no obvious NPF events for sites 1, 5–9 (DUK, WFM, PSP, UMBS, TNF, OZA) and infrequent



events (~ 20 %) for sites 2–4 (MMSF, SPL, EGB). Nucl-Org significantly overpredicts the nucleation frequency (by up to ~ 50–97 %) for sites 1–2 (DUK, MMSF) and 4–6 (EGB, WFM, PSP) while Nucl-IMN slightly underpredicts F_{nucl} (by ~ 10 %) for sites 2–3 (MMSF, SPL). On average, Nucl-Org overpredicts F_{nucl} by 50 % while F_{nucl} based on Nucl-IMN (7 %) is close to the observed value (8 %) for the nine sites in the summer month.

 F_{nucl} alone is not adequate to assess the performance of the model. For example, for Site 2 (MMSF) in March, F_{nucl} based on Nucl-IMN is closer to the observed values than F_{nucl} based on Nucl-Org but r and PC are actually lower (Fig. 4c). Our analysis indicates that a large fraction of NPF event days predicted by Nucl-IMN for MMSF in March does not match the NPF events observed on specific days. In terms of r and PC, the values based on Nucl-Org and Nucl-IMN are close for the spring month (Fig. 4c and Table 2), but Nucl-IMN overall gives higher r and PC values for July (Fig. 4d and Table 2). The correlation was calculated for a whole month of hourly CN_{10-100} data

- (~ 700 data points). Due to clear diurnal variations associated with NPF in the spring (Fig. 3a) and the fact that both schemes capture ~ 70% of event and non-event days, the average r value reaches 0.37 for Nucl-Org and 0.49 for Nucl-IMN. As a result of the lack of obvious diurnal variations and inability of the global model to capture the subgrid CN plumes (such as the short-duration CN spikes in Fig. 3h), the r value is much
- ²⁰ lower for July (nine-site average of 0.16 for Nucl-Org and 0.22 for Nucl-IMN) although PC reaches 0.88 for Nucl-IMN. The lower average PC value for Nucl-Org (0.43) is due to the over-prediction of NPF events (i.e., F_{nucl}).

Particle number concentrations and CN_{10-100} in NA (on a regional scale) are largely dominated by secondary particles formed via nucleation (Yu and Luo, 2009). Our anal-

²⁵ ysis demonstrates that nucleation schemes have a strong impact on the simulated particle number concentrations. Compared to observations, CN_{10-100} averaged at the four sites is 40% higher based on Nucl-Org and 14% higher based on Nucl-IMN for the spring month (Table 2). For the summer month, CN_{10-100} averaged at the nine sites is 161% higher based on Nucl-Org and 27% lower based on Nucl-IMN (Table 2).



For specific sites (Fig. 4e–f), CN_{10–100} based on Nucl-Org are generally higher than observed values in the spring for all four sites (especially site 1 – DUK) and are much higher than those in the summer month for sites 1–2 (DUK, MMSF) and 3–5 (SPL, EGB, WFM). Nucl-IMN substantially overpredicted CN_{10–100} for site 4 (EGB) in the spring and underpredicted CN_{10–100} for sites 8–9 (TNF, OZA) in the summer month. Possible sources of the difference include sub-grid inhomogeneity as well uncertainties in emissions, chemistry, aerosol microphysics, and meteorology. It is noteworthy that the global model, with a horizontal resolution of 2° × 2.5°, is aimed to capture regional events. If the measurements at a given site during certain periods were affected by sub-grid scale topography and processes (emissions, plumes, etc.), the global model would not be able to capture these effects. Comparisons of high-resolution simulations with observations are needed to address the issue.

4 Implications and discussion

We showed in the last section that Nucl-Org and Nucl-IMN schemes predict guite different spatial distributions and significant differences between the spring and summer 15 seasons with respect to NPF events. Table 2 shows that Nucl-IMN predictions agree better with observations in all categories, especially during July. One logical question to ask is: what is the implication of such differences? Figure 5 gives the ratios of the CCN concentration in the lower troposphere (0-3 km) based on Nucl-Org to the CCN concentration based on Nucl-IMN as well as the difference of aerosol first indirect radiative 20 forcing (FIRF; $FIRF_{Nucl-Org} - FIRF_{Nucl-IMN}$). The CCN concentrations are calculated at a water supersaturation ratio of 0.4% (CCN0.4) from simulated PSDs and the calculation of aerosol FIRF in GEOS-Chem is based on the approach discussed in Yu et al. (2013). As a result of higher nucleation rates, CCN0.4 based on Nucl-Org are approximately a factor of two higher than CCN0.4 based on Nucl-IMN in July over most 25 parts of NA (Fig. 5a). Higher CCN leads to higher cloud droplet number concentrations and enhanced cloud albedo, resulting in an additional negative FIRF (cooling) of 1-



 $2.5 \,\mathrm{W}\,\mathrm{m}^{-2}$ in a large fraction of NA with a domain average of $-1.27 \,\mathrm{W}\,\mathrm{m}^{-2}$. Since CCN can also impact precipitation (second indirect impact) and cloud cover, it is important to reduce uncertainty in the new particle formation calculation in regional and global climate models.

- ⁵ The significant overprediction of NPF events and particle number concentrations in the summer by the Nucl-Org scheme cannot be explained by the uncertainty in the pre-factor (k_m) in the Nucl-Org parameterization (Eq. 1). It indicates that the organicsenhanced nucleation parameterization derived from laboratory chamber studies may not be directly applicable to the atmospheric conditions, at least in the summer months
- ¹⁰ when VOC emissions, photochemistry, and SOA formation are strongest. One possible explanation for this result is that the concentration of organic compounds and their oxidation products exist at much lower concentrations in the atmosphere than those used in chamber studies. While both LV-SOG_{*α*-pinene} in the model and BioOxOrg in the chamber studies are later-generation oxidation products of biogenic monoterpenes, it
- ¹⁵ is possible that only a subset of LV-SOG_{α -pinene} may act as BioOxOrg vapors that are involved in nucleation. The uncertainty in the predicted concentrations of organic compounds actively participating in nucleation will not only impact the calculated nucleation rates but also the growth rate of these particles. Nevertheless, this will not be able to explain the observed spring-summer contrast in NPF since the concentrations of these
- ²⁰ specific compounds (i.e., BioOXOrg) are much higher in the summer, but observations show more frequent and stronger NPF events in the spring (Table 2 and Fig. 4).

Another possible reason for the overprediction of the Nucl-Org scheme is the influence of temperature. According to the Nucl-IMN scheme, the main reason for the lack of nucleation in July in relatively clean environments is the high temperature that reduces the supersaturation ratio of sulfuric acid and NPF. It should be noted that the Nucl-Org parameterization (Eq. 1) was derived from laboratory chamber studies at T = 278 K and RH = 39% (Riccobono et al., 2014). Based on nucleation theory, temperature should influence the nucleation rate significantly unless the nucleation is barrierless. Higher summer temperatures may inhibit H₂SO₄-Organic nucleation and thus



explain the lack of NPF in forest areas. If this is the case, the thermodynamic data for the H_2SO_4 -Organics system is needed to derive a *T*-dependent nucleation scheme which is presently not available. Furthermore, water vapor and base molecules such as ammonia and amines may also affect nucleation barriers and thus nucleation rates.

⁵ On the other hand, ammonia, amines, and certain organics concentrations are typically much higher in the summer and can also enhance IMN when the nucleation barrier is higher. This may explain the July under-prediction of CN by binary IMN.

Further research is needed to understand the different mechanisms of NPF in the atmosphere in different environments and represent them properly in the models. Con-

- tinuous field measurements of PSDs down to nanometer sizes (with improved accuracy in smaller sizes) in various environments along with concentrations of key precursor gasses are important. Thermodynamic data with regard to the interactions and stability of multiple-component pre-nucleation clusters is essential to develop more robust nucleation schemes that can take into account major influencing factors. Finally, develenergy and publication environments along with regard to the interactions and stabilinteraction schemes that can take into account major influencing factors. Finally, devel-
- oped nucleation schemes should be validated against field measurements taken under a wide range of varying conditions (such as season, temperature, vegetation types, complexity of terrain).

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Table 1. A summary of names, location, information, and references of nine forest sites where particle size distribution (PSD) observations are available and used in the present study. PSD measurements at different sites cover different periods of time and this study focuses on detailed analysis of one spring month and one summer month.

Site Name	Site Loca- tion	Site Information	PSD Measure- ment Periods	Period of PSD data used in this study	References	
1. Duke Forest (DUK)	(35.98° N, 79.09° W)	Suburban forest in south- eastern US	Nov., 2005 to Sep, 2007.	March and July of 2006	Pillai et al. (2013)	
2. Morgan-Monroe State Forest (MMSF)	(39.32° N, 86.42° W)	Rural forest in Midwestern US	Dec., 2006 to Dec., 2008	March and July of 2007	Pryor et al. (2010)	
3. Storm Peak Labora- tory (SPL)	(40.46° N, 106.74° W)	West summit of Mt. Werner in northwestern Colorado	Various periods from 2001	March 15– April 15, and July of 2012	Hallar et al. (2011)	
4. Egbert (EGB)	(44.23° N, 79.78° W)	Mixture of forests and farmland in Ontario, Canada	May, 2007 to May, 2008	July, 2007; March, 2008	Pierce et al. (2014)	
5. Whiteface Mountain (WFM)	(44.4° N, 73.9° W)	Summit of the Adirondack Mountains	7/10/2002– 8/07/2002	7/10/2002– 8/07/2002	Bae et al. (2010)	
6. Pinnacle State Park (PSP)	(42.10° N, 77.21° W)	Mixed deciduous and conif- erous forests in southwest New York State	7/13/2004– 8/11/2004	7/13/2004– 8/11/2004	Bae et al. (2010)	
7. University of Michi- gan Biological Station (UMBS)	(45.56° N, 84.72° W)	Transition zone between the mixed hardwood and the boreal forest	1 July–3 August 2009	July, 2009	Kanawade et al. (2011)	
8. Talladega National Forest (TNF)	(32.94° N, 87.16° W)	Isoprene-dominated rural forest in southeastern US	June 1 to 15 July, 2013	6/15–7/15, 2013	Lee et al. (2015)	
9. Ozarks Forest (OZA)	(38.74° N, 92.2° W)	"isoprene volcano" zone in the Central US	May to October 2012.	July, 2012	Yu et al. (2014)	

ACPD 15, 21271-21298, 2015 Spring and summer contrast in new particle formation F. Yu et al. **Title Page** Abstract Introduction Conclusions References **Tables** Figures 4 Close Back Full Screen / Esc **Printer-friendly Version** Interactive Discussion

Discussion Paper

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Table 2. Observed and simulated fraction (or frequency) of nucleation event days (F_{nucl}) and CN_{10-100} , for one spring month (sites 1–4 average) and one summer month (sites 1–9 average). The Pearson correlation coefficient (r) of simulated and observed time series of CN_{10-100} as well as proportion correct (PC) of days.

	Spring (4 sites)				Summer (9 sites)			
	F _{nucl}	CN ₁₀₋₁₀₀	r	PC	F _{nucl}	CN ₁₀₋₁₀₀	r	PC
Observation	0.69	5441			0.08	2450		
Nucl-Org	0.76	7606	0.37	0.68	0.60	6385	0.16	0.43
Nucl-IMN	0.76	6198	0.49	0.63	0.07	1783	0.22	0.88



Figure 1. The locations of nine measurement sites in the forest areas in North America (see Table 1 for the names and references of the sites). Also shown are the particle size distributions measured in one spring month at sites 1–4 (with blue frames) and one summer month at all sites (with red frames).





Figure 2. Horizontal distributions of monthly mean nucleation rates (*J*) in the boundary layer (0-1 km above surface) in March **(a, b)** and July **(c, d)** of 2006 based on two different nucleation schemes, i.e, Nucl-Org **(a, c)** and Nucl-IMN **(b, d)** (see text for details). The locations of sites 1-9 (Table 1) are marked.





Figure 3. Particle size distributions (PSDs) observed (**a**, **b**) and simulated based on Nucl-Org (**c**, **d**) and Nucl-IMN (**e**, **f**) schemes during two ten-day periods in March (**a**, **c**, **e**) and July (**b**, **d**, **f**) of 2006 in Duke Forest (DUK), along with time series of the concentration of condensation nuclei between 10 and 100 nm (CN_{10-100}) (**g**, **h**). The event (\checkmark) and non-event (\checkmark) days determined based on the criterion given in Sect. 2.3 for the observed and simulated PSDs are marked in Fig. 3a–f.







Figure 4. Statistical analysis of simulated NPF events and CN_{10-100} based on two nucleation schemes (Nucl-Org, Nucl-IMN) and comparison with observations at four forest sites for the spring month (**a**, **c**, **e**) and nine sites for the summer month (**b**, **d**, **f**). F_{nucl} is the fraction (or frequency) of nucleation event days (**a**, **b**). The proportion correct (PC) (**c**, **d**) is defined as (# of modeled events that were events + # of modeled non-events that were non-events)/(total # of modeled events + total # of modeled non-events). Monthly mean CN_{10-100} (**e**, **f**) and correlation coefficient (*r*) (**c**, **d**) of hourly mean CN_{10-100} during the month are also given.



Figure 5. (a) Ratios of the concentration of CCN (at water supersaturation ratio of 0.4 %) in the lower troposphere (0–3 km) based on Nucl-Org scheme to those based on Nucl-IMN scheme, and **(b)** Difference of aerosol first indirect radiative forcing (FIRF; $FIRF_{Nucl-Org} - FIRF_{Nucl-IMN}$).

