# Gas-particle Partitioning and Hydrolysis of Organic Nitrates Formed from the Oxidation of α-Pinene in Environmental Chamber Experiments

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## Abstract

Gas-particle partitioning and hydrolysis of organic nitrates (ON) influences their role as sinks and sources of $NO_x$ and their effects on the formation of tropospheric ozone and organic aerosol (OA). In this work organic nitrates were formed from the photo-oxidation of α-pinene in environmental chamber experiments under different conditions. Particle-phase ON hydrolysis rates consistent with observed ON decay exhibited a nonlinear dependence on relative humidity (RH): an ON decay rate of 2 $day^{-1}$ was observed when the RH ranged between 20 and 60%, and no significant ON decay was observed at RH lower than 20%. In experiments when the highest observed RH exceeded the deliquescence RH of the ammonium sulfate seed aerosol, the particle-phase ON decay rate was as high as 7 $day^{-1}$ and more variable. The ON gas-particle partitioning was dependent on total OA concentration and temperature, consistent with absorptive partitioning theory. In a volatility basis set the ON partitioning was consistent with mass fractions of [0 0.11 0.03 0.86] at saturation mass concentrations (C*) of [1 10 100 1000] μg $m^{-3}$.

## 1 Introduction

Organic nitrates (ON) play an important role in atmospheric chemistry as they can act as sinks and sources of $NO_x$ ($NO + NO_2$) and thereby affect the formation of tropospheric ozone and organic aerosol. The sink reaction – addition of NO to a peroxy radical (R-O-O·) to form an organic nitrate (R-O-$NO_2$) – breaks the ·OH initiated oxidation cycle and reduces the formation of ozone (Seinfeld and Pandis, 2006). Most R-O-$NO_2$ molecules are semi-volatile and are therefore expected to partition between the gas and particle phases. They can be transported in either phase and can become a source of $NO_x$ when they are photolyzed or oxidized, contributing to the regional nature of $NO_x$ pollution. Attempts to implement organic nitrate decomposition reactions in a chemical transport model which did not account for gas-particle partitioning of organic nitrates resulted in over-prediction of $NO_x$ and ozone concentrations (Yarwood et al., 2012), consistent with an over-estimate of the strength of organic nitrates as $NO_x$ sources.

Recent studies have suggested that organic nitrates in the condensed phase may undergo hydrolysis, leading to the formation of $HNO_3$ (Day et al., 2010; Darer et al., 2011; Hu et al., 2011; Liu et al., 2012; Browne et al., 2013; Jacobs et al., 2014; Rindelaub et al., 2015). This is a more permanent sink for $NO_x$ and would decrease the regeneration of $NO_x$ from organic nitrates. While these studies have found

evidence for hydrolysis of aerosol-phase organic nitrates ($ON^{aer}$), it is not clear at which rate ON hydrolysis occurs. Correctly modeling organic nitrates and ozone formation depends on knowledge of the ON partitioning and hydrolysis rate.

While ON hydrolysis in the bulk phase has been studied for decades (Baker and Easty, 1950; Baker and Easty, 1952; Boschan et al., 1955), organic nitrate hydrolysis in atmospheric particles has only recently started to receive attention. Day et al. (2010) observed a decrease in particulate organic nitrates measured in coastal southern California under acidic conditions at high relative humidity and hypothesized hydrolysis as the cause. Browne et al. (2013) used ON hydrolysis to justify observations over the Boreal Forest of higher levels of $HNO_3$ despite higher production rates of organic nitrates. The chamber experiments (0 to >80% RH) performed by Liu et al. (2012) using trimethylbenzene (an anthropogenic volatile organic compound) and HONO as oxidant were the first to measure the hydrolysis of condensed organic nitrates. Rindelaub et al. (2015) observed ON hydrolysis while measuring partitioning of α-pinene SOA but did not directly quantify it. Boyd et al. (2015) measured hydrolysis of ON formed from nitrate radical oxidation of β-pinene.

The partitioning of organic nitrates to the particle phase is important to determine their fate as only condensed organic nitrates are expected to hydrolyze appreciably to $HNO_3$. Absorptive partitioning theory (Pankow, 1994; Donahue et al., 2006, Rollins et al., 2013; Rindelaub et al., 2015) has been used to describe the gas-particle partitioning of organic nitrates. Rollins et al. (2013) used partitioning data from the 2010 CalNex campaign to find a volatility basis set distribution for ON observed at ambient aerosol concentrations. Rindelaub et al. (2015) observed the partitioning of organic nitrates formed from the ·OH initiated oxidation of α-pinene at various levels of relative humidity. However, other work has suggested that the partitioning of organic nitrates to the particle phase is irreversible (Perraud et al., 2012). The goals of this work were to form organic nitrates in controlled environmental chamber experiments from the ·OH-dominated oxidation of α-pinene under high $NO_x$ conditions and various relative humidity levels and:

1. Quantify the hydrolysis rate of organic nitrates.

2. Verify whether the gas-particle partitioning of organic nitrates is reversible and can therefore be modeled by absorptive partitioning theory

3. Parameterize the gas-particle partitioning of organic nitrates

# 2 Methods

## 2. 1 Environmental Chamber Experiments

All experiments were performed in the Atmospheric Physicochemical Processes Laboratory Experiments (APPLE) chamber located at the University of Texas at Austin (UT-Austin). The APPLE chamber is a ~12 $m^3$ Teflon ® bag suspended inside of a temperature-controlled room. The walls of the room are lined with UV lights which can be used to induce photolysis reactions. The intensity of the UV lights has been characterized by the photolysis rate of $NO_2$, which was measured to be 0.4 $min^{-1}$, similar to ambient $NO_2$

photolysis rates (e.g., 0.46 $min^{-1}$ at a zenith angle of 40°, Carter et al., 2005). Before each experiment the bag was flushed for at least 12 hours with clean air from an Aadco clean air generator (Model 737-14A) at a flow rate exceeding 100 liters per minute (LPM). Ammonium sulfate ($(NH_4)_2SO_4$) particles (Fisher Scientific, 99.5%) were injected both to monitor wall loss rates (Hildebrandt et al., 2009) as well as to act as seed particles onto which organic vapors can condense. Gas phase NO was injected directly into the chamber from a cylinder (Airgas, 9.94 PPM ±2%) and liquid-phase α-pinene (Sigma Aldrich, 98%) was injected to a glass bulb and subsequently evaporated into the chamber with a steady stream of mildly heated air. $H_2O_2$, which photolyzes to 2·OH, was used as ·OH radical source and was either injected by bubbling air through an aqueous $H_2O_2$ solution (Fisher Scientific, 30% weight) or by injecting $H_2O_2$ solution into a glass bulb and subsequently evaporating it into the chamber with a steady stream of mildly heated air. Some experiments were performed under dry conditions (<5% relative humidity); in other experiments humidity was increased by passing air through clean water and then into the chamber. Experimental conditions and results are summarized in Table 1. Results are discussed in Sect. 3.

Reactions were allowed to proceed for at least 4 hours with continuous UV light. Experiments were run in a batch mode with no injections or dilution after the experiment was started; the bag volume of 12 $m^3$ allowed ample time for sampling. In some cases the temperature effects on gas-particle partitioning were observed by increasing temperature to 40 °C in the chamber after the UV lights had been turned off (see Sect. 3.2).

### 2.1.1 Instrumentation

The composition of $PM_1$ (particulate matter smaller than 1 micrometer in diameter) was measured using an Aerosol Chemical Speciation Monitor (ACSM) from Aerodyne Research Inc. (Ng et al., 2011). In the ACSM, particles are flash-vaporized on a heater at 600 °C, and the resulting gas molecules are ionized using electron-impact ionization. This harsh ionization method results in fragmentation of most molecules. The molecular fragments, which are measured by a quadrupole mass spectrometer, are attributed to four categories—organics, nitrate, sulfate, and ammonium - using a fragmentation table (Allan et al., 2004). The instrument alternates between normal sampling and sampling through a particle filter, enabling subtraction of a gas-phase background. During this study the ACSM was operated at a time resolution (filter/sample cycle length) of approximately 90 seconds. The size distribution of particles was measured using a Scanning Electrical Mobility System (SEMS) from Brechtel Manufacturing, Inc. The SEMS uses a Differential Mobility Analyzer (DMA) to size-select particles based on their electric mobility, which are then counted by a Condensation Particle Counter (CPC). The DMA continuously cycled between the voltages which select particles ranging from 5 to 1000 nm in diameter, resulting in a time resolution of the particle size distribution of approximately 60 seconds.

Gas phase reaction products were monitored using a High-Resolution Time-of-Flight Chemical Ionization Mass Spectrometer (HR-ToF-CIMS) from Aerodyne Research, Inc.. The HR-ToF-CIMS uses softer chemical ionization which results in minimal fragmentation of parent molecules. Mass spectra are derived from measurements of the ions' time-of-flight as they are pulsed through a low pressure chamber in a "V"

shape. Two chemical reagent ions were used—water clusters ($H_3O^+ \cdot (H_2O)_n$) and iodide-water clusters ($I^- \cdot (H_2O)_n$). Water cluster ionization is most sensitive towards detection of moderately oxidized hydrocarbons; the ability to ionize and thus sensitivity is based on the relative proton affinity between the water cluster and the parent molecule (Lindinger et al., 1998). This method was used to monitor α-pinene as well as early-generation oxidation products. Iodide-water cluster ionization is most sensitive towards detection of more highly oxidized hydrocarbons; this method was used to observe later-generation oxidation products as well as $HNO_3$ and $H_2O_2$. In the work presented here data from the HR-ToF-CIMS are only used qualitatively since, as it was later discovered, a partially clogged inlet may have interfered with instrument calibration and quantitative measurements.

Concentrations of NO and $O_3$ were measured using Teledyne chemiluminescence $NO_x$ and absorption $O_3$ monitors (200E and 400E, respectively); concentrations of $NO_2$ were measured via an $NO_2$ monitor from Environnement (Model AS32M), which uses a Cavity Attenuated Phase Shift (CAPS) method to directly measure $NO_2$ (Kebabian et al., 2008). The advantage of this direct $NO_2$ measurement is that it does not rely on $NO_2$ conversion to NO and therefore does not suffer from interference by other oxidized nitrogen compounds such as HONO and organic nitrates (Winer et al., 1974).

### 2.1.2 Data Analysis

Data from the ACSM were analyzed in Igor Pro using the software package "ACSM Local," which includes a correction for relative ion transmission efficiency as well as changes in the flow rate throughout the experiment. The SEMS volume concentration was converted to mass using the densities 1.77 g/cm$^3$ for ammonium sulfate and 1.4 g/cm$^3$ for organics and organic nitrates (Ng et al., 2007). The time series of particle mass concentration (not corrected for wall losses) during Expt. 7 is shown in Fig. S1; other experiments exhibited similar time series.

All PM nitrate was assumed to be organic because no inorganic nitrate was introduced in these controlled experiments. PM nitrate is measured by the ACSM as $NO^+$ and $NO_2^+$ fragments, and the standard fragmentation table is used to estimate the portion of the *m/z* 30 that is due to $NO^+$ (as opposed to $(^{15}N)_2^+$ or Organics). Nitric acid is formed in the gas phase as well as in the particle phase through hydrolysis, but it is assumed that nitric acid concentrations are negligible in the particle phase due to its high vapor pressure (Fry et al., 2009). A Henry's Law calculation suggests that the total amount of aqueous $HNO_3$ in particles is 3 orders of magnitude lower than that in the gas phase.

The ACSM does not detect all sampled particles, primarily due to particle bounce at the vaporizer, resulting in a collection efficiency (CE) smaller than 1. Collection efficiency and wall losses were accounted for simultaneously by multiplying the ACSM concentrations of organics and organic nitrates by the mass concentration ratio $C_{SEMS}^{t=0}/C_{ACSM}^{seed}(t)$ as has been done in previous work (Hildebrandt et al., 2009). Here, $C_{SEMS}^{t=0}$ is the mass concentration of ammonium sulfate seed just before the UV lights are turned on and organic aerosol formation commences and $C_{ACSM}^{seed}$ (t) is the time dependent mass concentration of $(NH_4)_2SO_4$ measured by the ACSM throughout the experiment. This correction assumes that particles on the chamber walls participate in gas-particle partitioning as though they are still in suspension and that the suspended ammonium sulfate concentration changes only due to wall losses. It

accounts for partitioning of organic vapors into wall-deposited particles (Hildebrandt et al., 2009) but does not account for losses of organic vapors onto the clean Teflon® walls (e.g. Matsunaga and Ziemann, 2010).

The ACSM standard fragmentation table was adjusted based on filter measurements taken in each experiment as described in the supplementary information. Data from the HR-ToF-CIMS were analyzed in Igor Pro (Wavemetrics) using Tofware, the software provided with the instrument. The data were first mass calibrated based on HR-ToF-CIMS reagent ions and other known ions. The baseline was subtracted and the average peak shape was found so it could be used for high resolution analysis, through which multiple ions can be identified at any given integer *m/z*. Ions up to *m/z* 300 were analyzed in high resolution mode. Only prominent ions were fit above *m/z* 200 because of the high number of possible ions at this high *m/z*. After ions were identified in the high resolution spectrum, the peaks were integrated to yield a time series of ions. Analyte ion concentrations were then normalized by the reagent ion concentrations – the sum of $H_3O^+$, $H_3O^+ \bullet (H_2O)$ and $H_3O^+ \bullet (H_2O)_2$ for water cluster ionization and the sum of $I^-$ and $I^- \bullet (H_2O)$ for iodide-cluster ionization. This correction accounts for changes in reagent ion concentrations and instrument sensitivity during and between experiments. Relative humidity can affect instrument sensitivity but this varied by less than 5% during each experiment.

The partitioning coefficient of a species is defined as the ratio of the species concentration in the particle phase to the total species concentration (gas and particle phase). For a single compound the partitioning coefficient is the same whether it is on a mass or mole basis. However, for a mix of compounds, such as those formed in ·OH-initiated oxidation, the mass and mole-basis partitioning coefficients will be different, with the coefficient expected higher on a mass basis since higher molecular weight compounds typically have lower vapor pressure. The partitioning coefficient in this work was calculated on a mole basis, in part because fragmentation in the ACSM makes it impossible to tell the original size and identity of ON molecules. This mole-basis partitioning coefficient is also more useful for most modeling efforts which group chemical species without knowledge of their exact molecular identity. The particle-phase ON concentration was quantified using data from the ACSM: the mass concentration of nitrate measured by the ACSM was converted to mixing ratio (ppb) using the molecular weight of the nitrate functional group (62 g/mol). This assumes that the ON have only one nitrate functional group. Conversion of the nitrate mass concentration to mixing ratio avoids the need to assume an ON molecular weight (needed to estimated ON mass concentrations from ACSM) and is therefore deemed to be a more accurate measure of ON from the ACSM. Quantification of all gas phase ON species would necessitate calibration and identification of all ON species which is not feasible. Instead, a chamber box model and nitrogen balance was employed to estimate total gas-phase ON as described below.

## 2.2 Chamber Modeling and Partitioning Coefficient

In these experiments only five major forms of oxidized nitrogen are present in significant concentrations—NO, $NO_2$, $HNO_3$, $ON^{gas}$ and $ON^{aer}$ (gas and aerosol-phase organic nitrates, respectively). Figure S2 shows that, based on the Statewide Air Pollution Research Center (SAPRC) model (http://www.engr.ucr.edu/~carter/SAPRC/), the concentrations of other forms of reactive nitrogen are orders of magnitude lower than the concentrations of these five forms. Concentrations of NO and $NO_2$

were measured using gas-phase monitors, $ON^{aer}$ was measured using the ACSM, and concentrations of $HNO_3$ were approximated using the SAPRC box model. The concentration of $ON^{gas}$ was then calculated based on a nitrogen mass balance ($ON^{gas} = NOx^{initial} - NO_2 - NO - ON^{aer} - HNO_3{}^{model}$), and the partitioning coefficient was calculated as a time series for each experiment."

The concentration of $H_2O_2$ used in the model was adjusted until the modeled NO, $NO_2$, and $O_3$ concentrations closely matched those measured throughout each experiment as shown in Fig. S3 for Expt. 7. In Expts. 1, 2, and 3 (which were used to calculate the volatility basis set parameters) $H_2O_2$ was injected by passing air through a solution of $H_2O_2$ and into the chamber, and the amount of $H_2O_2$ injected could not be estimated. In later experiments $H_2O_2$ was directly injected into the chamber through evaporation of a known liquid volume of solution, and the estimated amount of $H_2O_2$ injected was about half of the $H_2O_2$ concentrations consistent with observed NO, $NO_2$ and $O_3$ based on the model. This difference may be due to inefficient injection or uncertainties in the UV spectrum. The absorption cross section of $H_2O_2$ only minimally intersects with the emission spectrum of the UV lights used in this chamber. Thus, small changes in the UV spectrum (or errors in measurements of the spectrum) could cause significant errors in estimated concentrations of [OH], which drive the observed concentrations of NO, $NO_2$ and $O_3$.

SAPRC simulations were conducted with the reaction mechanism Carbon Bond 6 revision 2 (CB6r2), which includes organic nitrate hydrolysis through a rate estimated from a combination of the work of Liu et al. (2012) and Rollins et al. (2013); updates to CB6r2 were described by Hildebrandt Ruiz and Yarwood (2013) . Experiments were modeled with and without organic nitrate hydrolysis to see the effect this has on the predicted ON partitioning coefficient. The overall effect of this process was small, with a maximum effect being a 5% decrease to the partitioning coefficient by removing the hydrolysis mechanism from the model. This corresponded to a 17% decrease in $HNO_3$, which suggests that the partitioning coefficient estimated in this work is not very sensitive to changes in the modeled $HNO_3$ concentrations. For the results and analysis presented here the $HNO_3$ concentrations were taken from CB6r2 with the inclusion of the ON hydrolysis process for experiments above 20% RH and without the hydrolysis process for experiments below 20% RH.

According to absorptive partitioning theory (Pankow, 1994; Donahue et al., 2006), the gas-particle partitioning of an organic species depends on its vapor pressure and the concentration of organic material in the particle phase. The fraction of a compound i in the particle phase ($Y_i$) is given by (Donahue et al., 2006):

$$Y_i = \left(1 + \frac{C_i^*}{C_{OA}}\right)^{-1} \quad (1)$$

where $C_{OA}$ is the organic aerosol concentration and $C_i^*$ is the saturation mass concentration of species i (the saturation vapor pressure converted to concentration units). In the volatility basis set (VBS, Donahue et al., 2006), organic species are lumped by $C_i^*$ spaced logarithmically. This leads to an overall partitioning coefficient

$$Y_{tot} = \sum_{i=1}^{n} F_i \left(1 + \frac{C_i^*}{C_{OA}}\right)^{-1} \quad (2)$$

(Rollins et al., 2013), where $F_i$ is the fraction of organic species in the volatility bin described by $C_i^*$. In this work we used measurements of $C_{OA}$ and $Y_{tot}$ to fit the $F_i$ using a Matlab optimization routine. These VBS parameters can be used in models to represent the gas-particle partitioning of organic nitrates and account for changes in partitioning with temperature and $C_{OA}$.

# 3 Results and Discussion

A typical time series of compounds containing oxidized nitrogen is shown in Fig. 1 (Expt. 7). Initially the chamber contains only NO and a small amount of $NO_2$, in addition to α-pinene and inorganic seed aerosol. When the UV lights are activated at time = 0 the NO immediately begins to react with ·OH and other radicals to form $NO_2$ and additional $NO_y$ compounds such as organic nitrates. Ozone formation also starts and thus in these experiments ozone also plays a role in the oxidation of α-pinene. Based on the SAPRC model 15% of the total α-pinene reacts with ozone while the rest reacts with ·OH. Table 1 summarizes results from all experiments. Concentrations of $O_3$, $ON^{aer}$, PM organics, and $ON^{gas}$ are averaged over approximately 20 minutes of the time when PM organics and nitrates peak in concentration. This averaging period was chosen so that experiments with different $H_2O_2$ concentrations could be compared even though they reach their maximum concentrations at different rates. Higher initial loading of NOx, α-pinene, and $H_2O_2$ resulted in higher concentrations of ozone and PM.

Figure 2 shows time series of molecular ions identified using the HR-ToF-CIMS using water cluster ("positive mode") and iodide-water cluster ("negative mode") ionization. Many compounds are identified with the CIMS and a select few of the most prominent compounds were chosen for Fig. 2. In short time periods after switching reagent ions the sensitivity of the HR-ToF-CIMS slowly adjusts to a steady state value. Minor changes during these short time periods should be taken with caution but the overall trends over the 4.5 hour experiment are useful in viewing oxidation trends. The initial data collected in negative mode show that formation of organic nitrates begins immediately after oxidation has started. Later in the experiment the less-oxygenated compounds observed in positive mode begin to decrease while the more highly oxygenated compounds observed in negative mode continue to increase, consistent with oxidation and conversion of less oxidized compounds to more highly oxidized compounds continuing throughout the experiment. Highly oxidized compounds which still contain ten carbon atoms (as the precursor α-pinene) begin to decrease towards the end of the experiment while fragmented compounds (containing less than ten carbon atoms) continue to increase, consistent with fragmentation of the carbon backbone during oxidation. Molecular weights of the gas-phase compounds identified here range from 221 to 279 g $mol^{-1}$ and align well with the range of molecular weights estimated by Fry et al. (2009) for particle-phase organic nitrates formed from $NO_3$ oxidation of β-pinene (229±12 to 434±25 g $mol^{-1}$). Gas-phase organic nitrates identified here are therefore expected to be semi- volatile and to partition significantly to the particle phase.

## 3.1 Hydrolysis of Organic Nitrates

Concentrations of wall-loss corrected (normalized to sulfate) PM nitrate were observed to decrease at the end of most experiments. These decreases of PM nitrate are attributed to physical or chemical processes in the gas and aerosol phases, and an exponential decay was fit to the data to quantify the

decay. The exception was experiments 1 and 3 during which production of SOA was slow (primarily due to lower initial $H_2O_2$ and α-pinene) and continued throughout the experiment, so a decay could not be observed. Examples of the decay for a humid and dry experiment are shown in Fig. S4. The decay rates for each experiment are reported in Table 1 and appear to depend on relative humidity as shown in Fig. 3. When the RH ranged between 20 and 60%, an ON decay rate of 2 $day^{-1}$ was observed; no significant ON decay was observed at RH lower than 20%. Experiments conducted at an average RH of 67% or higher can exhibit a significantly higher decay rate, probably due to effects of being near the deliquescence relative humidity of the ammonium sulfate seed aerosol. In experiments 10 and 12, which have decay rates well above 2 $day^{-1}$, the chamber was initially cooled to 20 °C before the UV lights were turned on. Once the UV lights were activated the temperature then increased to 25 °C and the RH settled at the values indicated in Table 1. For these experiments the RH was above 80% (the deliquescence RH, DRH, of $(NH_4)_2SO_4$ for several minutes, potentially resulting in aqueous aerosol. Experiment 11 also reached a relative humidity above deliquescence, yet it shows a lower nitrate loss rate than Expts. 10 and 12. The ratio of organics and nitrates to sulfate (seed) particles was much lower in Expt. 11 compared to Expts. 10 and 12, but whether and why this would result in a different nitrate loss rate is currently unclear. The relative humidity in Expt. 4 did not reach the DRH of $(NH_4)_2SO_4$. Future work should focus on the fate of ON at higher (> 60%) relative humidity. The generally higher nitrate loss rate at higher RH makes hydrolysis of particulate nitrate functional groups the most plausible explanation for the observed decay.

PM organics also decreased in some experiments, but their loss rate was lower and more variable than that of nitrate. Based on the work by Chuang et al. (2015) the addition of a nitrate functional group decreases volatility of a compound by 2.5 orders of magnitude – slightly more than the substituted alcohol group from hydrolysis. Thus, the organic compound resulting from ON is more volatile than the original organic nitrate, and as a result could partition to the gas phase, resulting in a decrease in PM organics.

No direct observation of hydrolysis (conversion of the -$ONO_2$ group to an –OH group) has been made in this or previous work. The estimated hydrolysis lifetime of 12 hours (loss rate of 2 $day^{-1}$) for particulate organic nitrates is similar to hydrolysis rates suggested by other studies under humid conditions. Liu et al. (2012) observed a trend similar to that shown in Fig. 3 in chamber experiments in which ON were formed from the oxidation of tri-methyl benzene using HONO as the ·OH and $NO_x$ source. In those experiments, PM nitrate was found to have negligible loss rate below 20% RH but a lifetime of 6 hours at 40% RH and higher. Perring et al. (2009) estimated the lifetime of isoprene nitrates to be between 95 minutes and 16 hours depending on their branching ratio in isoprene ·OH oxidation. Boyd et al. (2015) measured a lifetime of 3-4.5 hours for 10% of ON formed from $NO_3$ oxidation of β-pinene, with a much longer lifetime for the remaining 90%. This suggests that 10% of the ON functional groups were tertiary with the rest being primary or secondary as those have been shown to hydrolyze much slower in the bulk phase (Darer et al., 2011; Hu et al., 2011). More tertiary ON groups are expected from α-pinene than β-pinene, based on the location of the double bond.

Similar VOC precursors such as α-pinene and β-pinene can form different fractions of primary/secondary and tertiary ON. When $NO_3$ reacts and bonds with the terminal double bond of β-pinene, an alkyl radical

is formed in either a primary or tertiary position (opposite of the carbon-nitrate bond). The tertiary alkyl radical is more stable, so primary organic nitrates are expected to be more abundant. The double bond in α-pinene is not terminal, so the $NO_3$ reaction produces either a secondary or tertiary ON and alkyl radical. $NO_3$ typically bonds with the less substituted carbon of a double bond so that a more highly substituted alkyl radical is formed. The reverse is true for OH+NO chemistry. In this case NO reacts with the peroxy-radical to form the nitrate group. The peroxy-radical, a product of $O_2$ and an alkyl radical, is likely to be on a more substituted carbon as this would have been the more stable alkyl radical. Thus, more highly substituted ON are expected from OH + NOx than from $NO_3$ chemistry. This has important implications for attempts to model ON and the resulting $NO_x$ recycling.

As Table 1 shows experiments were conducted at varying $NO_x$ and α-pinene concentrations, relative humidity, and hydrogen peroxide (·OH radical source) levels, which resulted in different final concentrations of PM nitrate and total OA. Liu et al. (2012) suggested that a lower PM nitrate / OA ratio at higher RH could be due to ON hydrolysis. In these experiments, the correlation between the ratio of PM nitrate/total OA (measured when total OA was highest) and RH was very low ($R^2$ = 0.02). Thus, based on these experiments, differences in the observed final PM nitrate / OA are due to experimental conditions other than relative humidity.

## 3.2 Gas-particle Partitioning of Organic Nitrates

In order to test the reversibility of ON partitioning the temperature of the chamber was increased after OA had formed (and when the UV lights were off) in some experiments. Figure 4 shows gas and particle-phase measurements taken from a representative experiment (Expt. 2). After the UV lights are turned off there is a 60 minute period in which the temperature stabilizes around 15 °C. This is followed by ~90 minutes of heating to a final temperature of 45 °C. After this the chamber is quickly cooled back to 15 °C. Figure 4b shows a time series of the Org/sulfate and $ON^{aer}$/sulfate ratios measured by the ACSM. Sulfate has a low vapor pressure and does not evaporate significantly at the temperatures investigated; therefore changes in the $ON^{aer}$/sulfate and Org/sulfate ratios with chamber temperature can be attributed to partitioning of organic nitrates and other organic species between the gas and particle phases or wall losses of gas-phase species. As Fig. 4b shows, Org/sulfate and $ON^{aer}$/sulfate decreased with increasing temperature and increased with decreasing temperature, suggesting evaporation of species at higher temperatures and their re-partitioning to the particle phase at lower temperatures.

Figure 4c shows the effects of temperature on various compounds measured in the gas phase. Several organic compounds – with and without ON functional groups - increase with increasing temperature. This suggests that these compounds are present in both the gas and particle phases and evaporate at higher temperature resulting in increased gas phase concentrations. As temperature is increased the percent change in the concentration of gas-phase $C_{10}H_{16}O_2$ is less than the change in $C_{10}H_{16}O_4$ and the percent change in the concentration of gas-phase $C_{10}H_{15}NO_4$ is less than the change in $C_{10}H_{15}NO_6$. This is consistent with the more highly oxidized compounds having a lower vapor pressure and evaporating less. As the temperature is decreased back to 15 °C the concentrations return to the pre-heating trends, suggesting that re-condensation to the particle-phase has occurred. These observations, as well as the trends seen in particle-phase measurements, are consistent with equilibrium partitioning and inconsistent with the irreversible partitioning of ON recently suggested by Perraud et al. (2012).

Other processes may influence particle and gas concentrations of organic compounds. Continuing reactions with $O_3$ and nitrate radicals (since $O_3$ and $NO_2$ are both present) limit the ability to stop all chemical activity. This is seen in the gas phase compounds, some of which appear to be changing in concentration after the UV lights are off. Despite this a clear change is seen in all compounds with a temperature increase. During the cooling phase (beginning at t = 320 minutes) the particle phase organic and nitrate concentrations do not return to the original levels. It is likely that some organic compounds are lost to the walls of the Teflon chamber, especially since they reach the coldest temperatures during active cooling, and thus Org/sulfate does not return to the values seen before temperature changes began. Despite these limitations it is clear that both the Org/sulfate and $ON^{aer}$/sulfate ratios decrease with heating, consistent with semi-volatile organics and organic nitrates.

Table 1 summarizes the ON partitioning coefficient averaged over approximately 20 minutes from when PM organics and nitrates peak in concentration. PM concentration and partitioning data are not calculated for experiments above 60% RH. As discussed, these experiments had higher and less consistent nitrate decay rates which may affect partitioning. There may also be issues in applying the wall-loss and CE corrections used here and in previous work (Hildebrandt et al., 2009), which assume that the particles measured by the SEMS before lights are turned on consist only of ammonium sulfate. In these high RH experiments the particles may contain a significant fraction of water. The wall loss correction also assumes that particles lost to the walls participate in partitioning as though they were still in suspension. This assumption may be poor if small amounts of water condense onto the walls of the chamber in these high RH experiments.

Data taken throughout the lower-concentration experiments (Expts 1, 2, and 3) with UV lights on were fit to a volatility basis set as these experiments were conducted under conditions which are more atmospherically relevant. Volatility basis set (VBS) parameters are static but are often used to approximate dynamic systems such as the one observed in these experiments. Experimental data were used after total PM organics (corrected for wall losses) had reached 2 μg $m^{-3}$ to avoid effects of noise and model uncertainty at the beginning of the experiments when concentrations of both gas- and particle-phase organic nitrates were low. Outlying points (for example, when PM organics temporarily rose above 2 μg $m^{-3}$ but subsequent data suggested that condensation had not begun) were removed as well. Figure 5 shows the data used to find the volatility basis set along with the fit. The $C^*$ values used for this were 1, 10, 100, and 1000 μg $m^{-3}$; the corresponding mass fractions ($F_i$ ) calculated to give the best fit for Eq. (2) (Sect. 2.2) are $F_i$ = 0, 0.11, 0.03, and 0.86.

As seen in Fig. 5 these results indicate that under typical ambient conditions (< 40 μg/$m^3$ of OA) 5-10% of organic nitrates formed from the photo-oxidation of α-pinene under high NOx conditions are expected to partition to the particle phase. This is significantly lower than the organic nitrate partitioning calculated by Rollins et al. (2013) for organic nitrates measured in Bakersfield, CA during the CalNex campaign in 2010. In those measurements >30% partitioning of ON was observed at organic aerosol concentrations of 10 μg/$m^3$. The difference could be attributed to differences in precursor molecules and levels of oxidation. Studies have shown that high $NO_x$ conditions can shift photochemical oxidation products of terpenes towards higher volatility compounds (Wildt et al. 2014). Rollins et al. determined using the SPARC model (Hilal et al., 2003) that precursor molecules (a mix of C5-C15 VOCs)

would need two stages of oxidative chemistry beyond the initial oxidation of the VOC to reach the point when 19-28% would partition to the particle phase for a $C_{OA}$ of 3 µg $m^{-3}$. This may suggest that the ON formed in our experiments have undergone fewer than three generations of oxidation as they are more volatile than the ON measured in Bakersfield during CalNex 2010. It should also be noted that the thermal dissociation-laser induced fluorescence (TD-LIF) instrument used by Rollins et al. (2013) has been shown in a recent study to measure PM ON a factor of two higher than the ON measured by aerosol mass spectrometers (Ayres et al., 2015) which utilize similar measurement and detection techniques as the ACSM used in this work. While the reasons for this difference are unknown it would result in a higher partitioning coefficient compared to the one calculated from the AMS (or ACSM) and explain part of the observed difference.

# 4 Conclusions

Organic nitrates formed during the oxidation of α-pinene decay in the particle phase at a rate of 2 $day^{-1}$ when RH is between 20 and 60%; no significant decay was observed below 20% RH. During experiments when the highest observed RH exceeded the deliquescence RH of the ammonium sulfate seed aerosol, the particle-phase ON decay was as high as 7 $day^{-1}$ and more variable. The dependence of observed decay rate on relative humidity suggests organic nitrate hydrolysis as the most plausible explanation. The gas-particle partitioning of ON determines their potential to hydrolyze. Partitioning of the ON is reversible and can be described by a volatility basis set.

The conversion of NOx to organic nitrates affects local ozone production. Partitioning and hydrolysis of organic nitrates affect regional concentrations of organic particulate matter and ozone. The organic nitrate partitioning coefficient and hydrolysis rates from this work can be used to include these processes in chemical transport models and more accurately represent the effect of organic nitrates on concentrations of ozone and particulate matter.

# 5 Acknowledgements

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Table 1. Experimental conditions and summary of results.

| Exp | initial α-pinene (ppb) | initial NO (ppb) | RH (%) | $H_2O_2$ conc in model (ppb)[a] | $O_3$ (ppb)[b] | $ON^{aer}$ ($\mu g/m^3$)[b,c] | PM Org ($\mu g/m^3$)[b,c] | $ON^{gas}$ (ppb)[b] | Part coeff[d] | Hyd. ($day^{-1}$) |
|---|---|---|---|---|---|---|---|---|---|---|
| 1 | 40 | 30 | 22 | 100 | 90 | 7 | 90 | 13 | 0.19 | NA[e] |
| 2 | 40 | 40 | 39 | 60 | 50 | 6 | 60 | 11 | 0.18 | 2.2 |
| 3 | 40 | 40 | 0 | 40 | 50 | 4 | 30 | 13 | 0.10 | NA[e] |
| 4 | 130 | 110 | 68 | 600 | 210 | | | | | 2.4 |
| 5 | 130 | 130 | 22 | 900 | 330 | 70 | 780 | 57 | 0.33 | 1.8 |
| 6 | 130 | 120 | 50 | 500 | 240 | 40 | 460 | 47 | 0.26 | 1.9 |
| 7 | 130 | 120 | 15 | 200 | 210 | 50 | 510 | 34 | 0.38 | 0.2 |
| 8 | 80 | 80 | 0 | 1000 | 300 | 30 | 310 | 32 | 0.26 | 0.6 |
| 9 | 80 | 80 | 0 | 1500 | 330 | 20 | 270 | 28 | 0.25 | 0.2 |
| 10 | 50 | 50 | 70 | 600 | 180 | | | | | 6.9 |
| 11 | 40 | 40 | 70 | 200 | 70 | | | | | 2.5 |
| 12 | 50 | 50 | 67 | 500 | 170 | | | | | 5.2 |

[a] $H_2O_2$ concentration for which SAPRC model most closely matched measurements of $NO_x$ and $O_3$

[b] Measured and averaged over a 20 minutes period when PM organics peaked

[c] Corrected for wall-losses as described in Sect. 2.1.2

[d] Molar basis

[e] Experimental conditions resulted in aerosol growth throughout the experiment

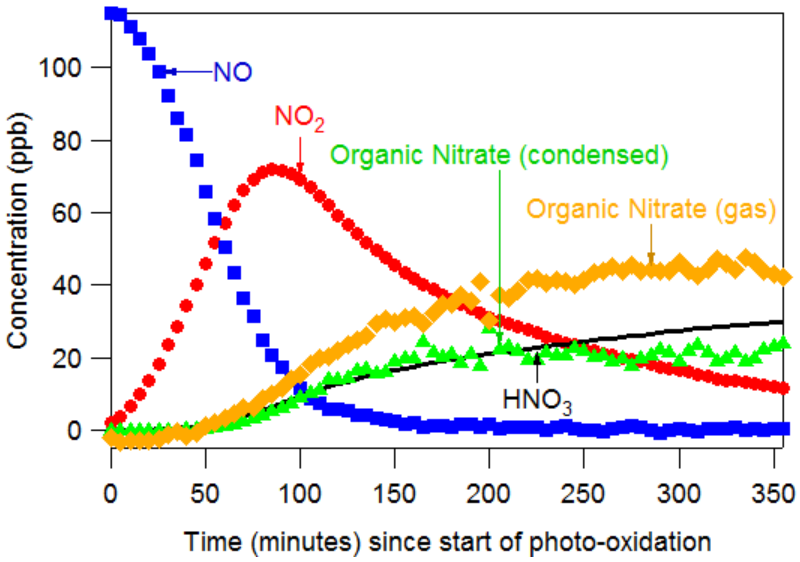



Figure 1 – Time series of oxidized-nitrogen species in Expt. 7. NO, $NO_2$, and $ON^{aer}$ are measured directly. $HNO_3$ is modeled using SAPRC. $ON^{gas}$ is calculated from a mass balance.

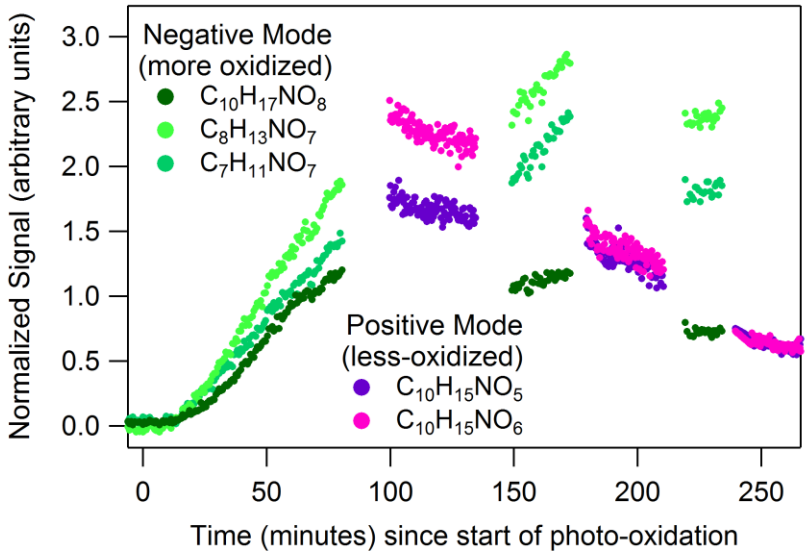



Figure 2 – Time series of selected organic nitrates identified by HR-ToF-CIMS (Expt. 10)

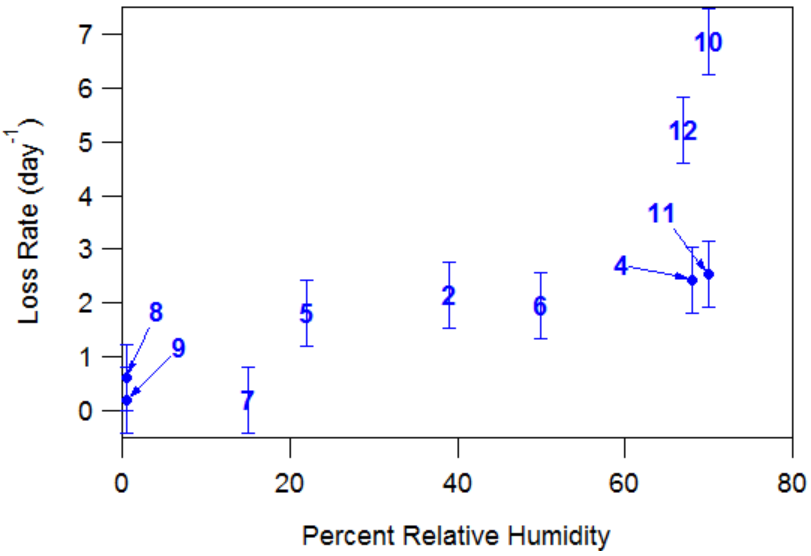



Figure 3. The organic nitrate loss rate as a function of relative humidity for Expts. 2, 4-12. Uncertainty (error bars) is estimated as 0.6 $day^{-1}$, the highest loss rate observed in experiments below 5% RH (Expt. 8).

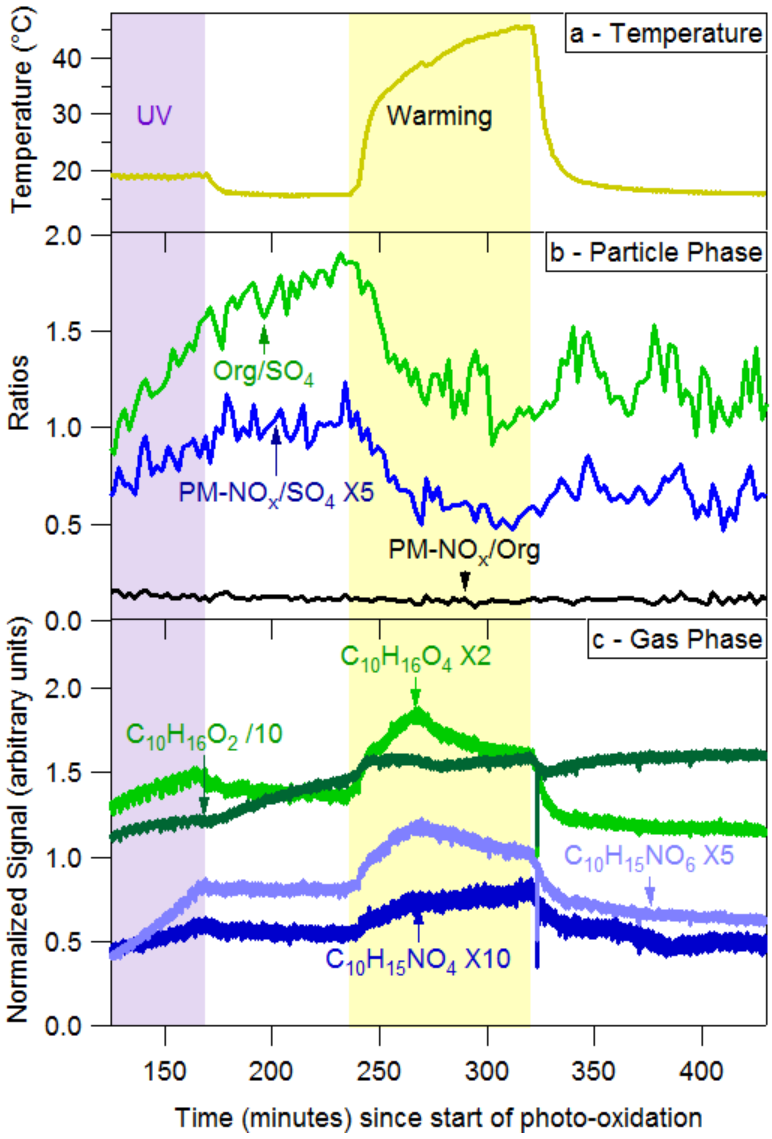
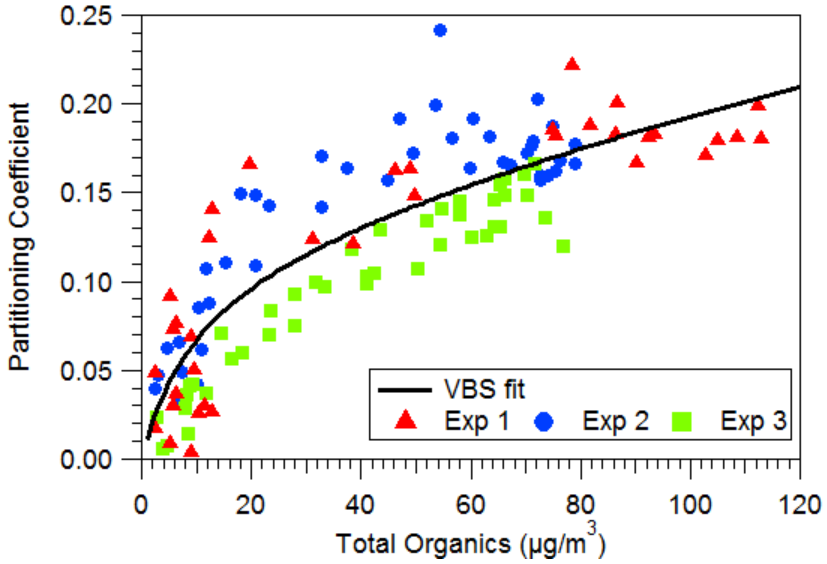



Figure 4 – Temperature effects on gas-particle partitioning of organic nitrates (Expt. 2).




604

605 Figure 5 – Volatility basis set fit from this work shown with data from Expts. 1, 2, and 3.