Yafang Cheng
Co-Editor of ACP

Dear Dr. Cheng,

We apologize that a few of the comments by referee #4 were either missed or not clearly answered in the initial author's response. Please find the below discussion which addresses these points.

[1] No ancillary data are presented, such as aerosol size distributions.

> ***[A1]*** Unfortunately, this data was not measured at all sites (e.g. UBC), and to prepare this information for publication would require a large amount of work. We hope the Editor agrees that this information is not required for the current publication.

[2] The aerosols were not brought to a standardized humidity so that the sizing would be more consistently comparable. Possible aging of the samples was not controlled for.

> ***[A2]*** To address the referees comment, in the revised manuscript we have added the caveat that the aerosol was not brought to a standardized humidity, and therefore small differences in the measured INP size between locations may result due to differences in hygroscopic growth. We have also added the caveat that the duration of sample storage varied in this study, and that additional studies are needed to quantify the effect of storage on INP activity.

Sincerely,

Allan Bertram
Professor of Chemistry
University of British Columbia

Title: Size-resolved measurements of ice nucleating particles at six locations in North America and one in Europe

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Abstract

Detailed information on the size of ice nucleating particles (INPs) may be useful in source identification, modeling their transport in the atmosphere to improve climate predictions, and determining how effectively or ineffectively instrumentation used for quantifying INPs in the atmosphere captures the full INP population. In this study we report immersion-mode INP number concentrations as a function of size at six ground sites in North America and one in Europe using the micro-orifice uniform deposit impactor-droplet freezing technique (MOUDI-DFT), which combines particle size-segregation by inertial impaction and a microscope-based immersion freezing apparatus. The lowest INP number concentrations were observed at Arctic and alpine locations and the highest at suburban and agricultural locations, consistent with previous studies of INP concentrations in similar environments. We found that 91 ± 9, 79 ± 17, and 63 ± 21 % of INPs had an aerodynamic diameter > 1 μm at ice activation temperatures of -15, -20, and -25 °C, respectively, when averaging over all sampling locations. In addition, 62 ± 20, 55 ± 18, and 42 ± 17 % of INPs were in the coarse mode (> 2.5 μm) at ice activation temperatures of -15, -20, and -25 °C, respectively, when averaging over all sampling locations. These results are consistent with six out of the nine studies in the literature that have focused on the size distribution of INPs in the atmosphere. Taken together, these findings strongly suggest that supermicron and coarse mode aerosol particles are a significant component of the INP population in many different ground-level environments. Further size-resolved studies of INPs as a function of attitude are required since the size distribution of INPs may be different at high altitudes due to size-dependent removal processes of atmospheric particles.

1 Introduction

Ice nucleating particles (INPs) are a unique class of aerosol particles that catalyze ice

46 formation under atmospheric conditions. A variety of particle types have been identified as INPs,
47 including mineral dust, black carbon, volcanic ash, glassy aerosols, and primary biological
48 particles such as bacteria, fungal spores, and pollen (see reviews by Szyrmer and Zawadzki,
49 1997; Möhler et al., 2007; Ariya et al., 2009; Després et al., 2012; Hoose and Möhler, 2012;
50 Murray et al., 2012; Yakobi-Hancock et al., 2013). Although only a small fraction of aerosol
51 particles nucleate ice (e.g. Rogers et al., 1998), INPs are important since they can lead to changes
52 in the properties and lifetimes of mixed-phase and ice clouds, ultimately affecting climate and
53 precipitation (Baker, 1997; Lohmann and Feichter, 2005; Baker and Peter, 2008; DeMott et al.,
54 2010; Creamean et al., 2013).

55 Vali et al. (2015) describes four modes of heterogeneous ice nucleation: deposition
56 nucleation, where ice forms on the INP directly from the gas phase; condensation freezing,
57 where ice nucleates during the condensing of water onto the INP; immersion freezing, where
58 crystallization is initiated by an INP within a supercooled liquid droplet; and contact freezing,
59 where the freezing of a supercooled liquid droplet is due to impaction by an INP. In this study
60 we focus on freezing via the immersion mode in dilute solution droplets, which is relevant to
61 mixed-phase cloud conditions.

62 Due to the importance of INPs for climate and precipitation, there has been a renewed
63 interest in measuring the concentrations of INPs in the atmosphere (DeMott et al., 2011). While
64 much of this work has focused on measurements of the total number concentration of INPs, there
65 has been less emphasis on determining their size distributions in the atmosphere. Information on
66 airborne INP size distributions may be particularly helpful in identifying the predominant INP
67 sources. For example, information on the size distribution of INPs may help rule out or support
68 the role of fungal spores in atmospheric ice nucleation since they are often in the supermicron

range (Graham et al., 2003; Elbert et al., 2007; Sesartic and Dallafior, 2011; Després et al., 2012; Huffman et al., 2012). A similar approach can be used with black carbon particles, since they are mainly in the submicron range (Clarke et al., 2004; Schwarz et al., 2008, 2013).

Previous modeling studies have shown that the transport and distribution of INPs, and aerosol particles in general, are sensitive to the size of the particles assumed in the models (Burrows et al., 2009; Wilkinson et al., 2011). Information on the size distributions of INPs are thus needed for accurate modeling of their transport and distributions in the atmosphere (Morris et al., 2004; Hoose et al., 2010a, 2010b; Sesartic et al., 2013; Haga et al., 2014; Spracklen and Heald, 2014).

Information on the size distribution of INPs is also needed to determine if techniques used to measure atmospheric INP concentrations capture the entire INP population. For example, the continuous flow diffusion chamber (Rogers et al., 2001b) is often used for measuring INPs (e.g. DeMott et al., 1998; Rogers et al., 2001a; Richardson et al., 2007; Pratt et al., 2009; Prenni et al., 2009; Eidhammer et al., 2010; Chou et al., 2011; Friedman et al., 2011; Hoyle et al., 2011; Corbin et al., 2012; Garcia et al., 2012; Tobo et al., 2013; McCluskey et al., 2014). This type of instrument has the advantage of providing real-time measurements of INPs with the ability to detect very large INP number concentrations, but the aerodynamic diameter of particles measured with it is limited, from $d_{50} \leq 2.4$ μm in some studies (e.g. Garcia et al., 2012) to $d_{50} \leq 0.75$ μm in others (e.g. DeMott et al., 2003). Such techniques may miss supermicron or coarse mode (i.e. larger than 2.5 μm) INPs. The exact proportion of INPs missed may depend on temperature. Such online instruments have typically focused on measurements below approximately -20 °C as sample volume considerations limit effective sampling of lower INP number concentrations at warmer temperatures. The exact proportion of INPs missed may also

depend on altitude since the removal of atmospheric particles by wet and dry deposition in the atmosphere is expected to be size dependent. As an example, supermicron particles have larger dry deposition loss rates than submicron particles.

Previous studies of INPs as a function of size have been carried out in the field (e.g. Vali, 1966; Rosinski et al., 1986; Mertes et al., 2007; Santachiara et al., 2010) and in the laboratory (e.g. Welti et al., 2009; O'Sullivan et al., 2015). These and additional studies are further discussed in Sect. 3.2. In the current study, we add to the existing body of size-resolved INP measurements by reporting ground-level INP size distributions from six locations in North America and one in Europe, covering a diverse set of environments and investigating immersion freezing at -15, -20, and -25 °C.

2 Methods

2.1 Sampling sites

The seven locations used in this study are detailed in Table 1 and shown in Fig. 1. All reported sampling periods are local times. Measurements using the sampling instrumentation described in the next section were made at five locations in Canada: Alert, Nunavut; the Labrador Sea near Newfoundland and Labrador; Whistler Mountain, British Columbia; the University of British Columbia (UBC) campus, British Columbia and Amphitrite Point, British Columbia. Measurements in Canada were conducted as part of the larger NETwork on Climate and Aerosols: addressing key uncertainties in Remote Canadian Environments project (NETCARE; http://netcare-project.ca/). Measurements were also made at Saclay, France and Colby, Kansas, USA.

2.1.1 Alert

Arctic sampling was conducted at the Dr. Neil Trivett Global Atmosphere Watch Observatory in Alert, Nunavut, Canada (labeled 1 in Fig. 1; Cobbett et al., 2007) between March 29 and July 23, 2014. This Arctic research station is part of a global network for measuring chemical and physical perturbations of the atmosphere. Aerosol particles were collected through a louvered total suspended particulate (TSP) inlet (Mesa Labs Inc., Butler, NJ, USA) and 0.9 m mast located on the upper level of an outdoor platform free of surrounding obstructions, and were stored in the dark at -15 or 4 °C for a period of 10–112 days prior to analysis.

2.1.2 Whistler Mountain

The Whistler Peak High Elevation Site is located at the summit of Whistler Mountain in Whistler, British Columbia, Canada (labeled 3 in Fig. 1) and operated by Environment Canada (Gallagher et al., 2011; Macdonald et al., 2011). Aerosol particle collection at this alpine site occurred between March 30 and April 23, 2014. The louvered TSP inlet was located approximately 10 m from a chairlift operating station. Although there are no continuous combustion sources at the site, sampled air may have been influenced by engine exhaust for short periods of time due to nearby snowmobile operation. Samples were stored in the dark at 4 °C for a period of 1–4 days prior to analysis.

2.1.3 Amphitrite Point

The coastal site at Amphitrite Point on Vancouver Island, British Columbia, Canada (labeled 5 in Fig. 1) is operated by Environment Canada, the BC Ministry of Environment, and Metro Vancouver for the continuous monitoring of aerosols and trace gases influenced by marine trajectories (McKendry et al., 2014; Yakobi-Hancock et al., 2014; Mason et al., 2015b). The mobile laboratory used during sampling was located approximately 100 m from the high tide line of the Pacific Ocean along a rocky shoreline, separated from the ocean by a narrow row of trees

and shrubs approximately 2–10 m in height. Sampling took place from August 6 to August 27, 2013 using a louvered TSP inlet and 3 m mast. Aerosol particles were stored at room temperature and analyzed within 1 day of collection.

2.1.4 The Labrador Sea

The Canadian Coast Guard Service vessel CCGS Amundsen serves as both an icebreaker for shipping lanes and an Arctic research vessel. One set of aerosol particle samples was collected from the top of the bridge of this vessel on July 11, 2014 while in the Labrador Sea off the coast of Newfoundland and Labrador, Canada (labeled 2 in Fig. 1). While sampling was within the marine boundary layer in the presence of sea spray aerosols, back trajectories (not included) calculated using the Hybrid Single-Particle Lagrangian Integrated Trajectory (HYSPLIT4) model of the National Oceanographic and Atmospheric Administration (Draxler and Rolph, 2014) indicate that the sampled air mass spent the majority of the previous 72-hour period over land. Air was passed through a louvered TSP inlet and 1.5 m mast during sampling, and collected aerosol particles were stored in the dark at 4 °C for a period of 45–46 days prior to analysis.

2.1.5 Saclay, France

Aerosol particle samples were collected at the Commissariat à l'Energie Atomique (CEA) Atmospheric Supersite (AS), CEA l'Orme des Merisiers. The CEA-AS Observatory is a suburban area located 30 km southeast of Paris in Saclay, France (labeled 7 in Fig. 1). The CEA-AS Observatory is surrounded by different sources of bioaerosols such as forest and agricultural fields, and is often influenced by marine or urban air masses (Baisnée et al., 2014). Measurements were made as part of the BIODETECT 2014 intensive campaign, an intercomparison of bioaerosol detection methods (Sarda-Estève et al., 2014). During this study

period, the site was heavily influenced by urban outflow. A large set of ancillary measurements was done to constrain all the particulate matter sources. Aerosol particles were sampled through a TSP inlet and 10 m mast between July 15 and August 4, 2014, and were stored in the dark at 4 °C for a period of 55–217 days prior to analysis.

2.1.6 UBC Campus

Four sets of aerosol particle samples were collected from a weather station on the roof of the five-story Earth Sciences Building on the UBC campus in British Columbia, Canada (labeled 4 in Fig. 1). The UBC campus is located on a peninsula and is surrounded by forest on three sides and ocean on the fourth. The site has been classified as suburban since it is less than 10 km from downtown Vancouver. Samples were collected through a TSP inlet and 0.5 m mast between May 12 and May 16, 2014. The aerosol particles were stored in the dark at 4 °C for a period of 21–23 days prior to analysis.

2.1.7 Colby, KS

Aerosol particles were collected at the soybean and sorghum fields of the Kansas State University Northwest Research Center in Colby, KS, USA (labeled 6 in Fig. 1). One sample was collected at each location during combine harvesting from a distance approximately 3–10 m downwind of the field. A third sample was also collected at the sorghum field the night following harvest. Sampling took place on October 14 and 15, 2014 and samples were stored in the dark at 4 °C for a period of 41–46 days prior to analysis.

2.2 Size-resolved INP number concentrations

INP number concentrations as a function of size and temperature were determined using the micro-orifice uniform deposit impactor-droplet freezing technique (MOUDI-DFT; Huffman

et al., 2013; Mason et al., 2015a). This technique combines aerosol particle collection by a cascade inertial impactor with sharp size-cutoff characteristics (the MOUDI; Marple et al., 1991) with an established droplet freezing apparatus (the DFT) for determining immersion-mode freezing properties (Koop et al., 1998; Iannone et al., 2011; Haga et al., 2013). A similar approach has also been used to study deposition nucleation by particles collected from the atmosphere (Wang et al., 2012; Knopf et al., 2014).

2.2.1 Aerosol particle sampling

Size-fractionated aerosol particle samples were collected onto hydrophobic glass cover slips (HR3-215; Hampton Research, Aliso Viejo, CA, USA) using a model 110R or 120R MOUDI (MSP Corp., Shoreview, MN, USA). Previous work has shown that these hydrophobic glass surfaces do not cause significant heterogeneous ice nucleation (e.g. Haga et al., 2013, 2014; Wheeler et al., 2015). Substrate holders were used on the impaction plates of the MOUDI to reproducibly position the hydrophobic glass cover slips in regions where aerosol deposit particle concentrations did not vary significantly (for details see Mason et al., 2015a). At most locations MOUDI stages 2–9 were used, corresponding to particle size bins of 10–5.6, 5.6–3.2, 3.2–1.8, 1.8–1.0, 1.0–0.56, 0.56–0.32, 0.32–0.18, and 0.18–0.10 μm (50 % cutoff aerodynamic diameter; Marple et al., 1991), respectively. The range in particle size collected at each location is given in Table 1.

Bounce within inertial impactors such as the MOUDI can occur during aerosol sampling, where particles impact the collection substrate but are not retained. This rebounding of particles from the surface could possibly alter the INP number concentrations and size distributions being measured. If composition is held constant, bounce is expected to increase with particle size because of their greater kinetic energy (Dahneke, 1971). Hence, INP number concentrations for

large particle sizes may be underestimated here. Bounce is also expected to increase with decreasing relative humidity (RH). Previous work has shown that having a sample RH of 70 % or greater can be effective in reducing particle bounce (e.g. Winkler, 1974; Fang et al., 1991; Stein et al., 1994; Vasiliou et al., 1999; Chen et al., 2011; Bateman et al., 2014), although its efficacy is dependent on particle type (Winkler, 1974; Lawson, 1980; Saukko et al., 2012). For six out of the seven sites investigated here, the average RH during sampling was 69% or greater. Recently, results from the MOUDI-DFT and the continuous flow diffusion chamber were compared during an ambient field campaign at Colorado State University (Mason et al., 2015a). For particle sizes < 2.4 µm, the INP number concentrations measured by the MOUDI-DFT were within experimental error of those measured by the continuous flow diffusion chamber technique, suggesting that bounce was not an issue during these previous ambient field measurements. Based on these previous measurements, in the current studies we do not consider the issue of particle bounce when calculating INP number concentrations and size distributions. Nevertheless, additional studies are warranted to better quantify the effect of bounce.

Ryan Mason 1/19/2016 7:38 PM
Deleted: relative humidity

During sampling, the aerosols were not brought to a standardized humidity, and the RH varied from site-to-site (see Table 1 for the average RH during sampling at each site). This variability in RH could lead to a small variability in INP size due to differences in hygroscopic growth. We further note that the duration of sample storage varied in the current study. Additional studies are needed to quantify the effect of storage on INP activity.

2.2.2 Freezing measurements

Samples were analyzed by the DFT to determine the number concentration of particles active in the immersion-freezing mode. Details of the experimental procedure can be found in Mason et al. (2015a). Briefly, samples were transferred to a temperature- and humidity-

controlled flow cell coupled to an optical microscope equipped with a 5× magnification objective (Axiolab; Zeiss, Oberkochen, Germany). Water droplets were condensed onto the sample and monitored using a CCD camera recording a digital video. Since the relative humidity of the gas flow during droplet condensation was held at approximately 120 %, water condensation occurred uniformly on the cover slip, and growing droplets coagulated as they grew to a final size of 97 ± 42 µm (mean diameter and 1 standard deviation (SD) uncertainty). The freezing temperature of each droplet was then determined during cooling at a rate of -10 °C min^{-1} using the video timestamp and a resistance temperature detector located within the flow cell. Note that droplet growth and coagulation occurs in the same manner for samples containing particles and clean hydrophobic glass cover slips (no particles deposited). In addition, based on an analysis of samples collected at Amphitrite Point, more than 99 % of particles become incorporated into the droplets prior to the freezing experiments. Here we regard ice nucleation as a singular process (i.e. strictly temperature-dependent) but note that the stochastic (i.e. time-dependent) component to immersion freezing (Vali, 2014) may alter the median freezing temperature of a droplet by 0.5–2 °C per decade change in cooling rate (Murray et al., 2011; Welti et al., 2012; Wright and Petters, 2013; Wright et al., 2013; Wheeler et al., 2015).

A potential issue with the droplet freezing technique is heterogeneous ice nucleation initiated by the hydrophobic glass cover slips used to collect atmospheric particles. To address this issue, experiments were conducted using new hydrophobic glass cover slips that were processed in the same manner as ambient samples except they were not exposed to atmospheric particles drawn into the MOUDI. For the five hydrophobic glass cover slips investigated, which contained 231 droplets generated during the freezing experiments, the average freezing temperature was -36.5 ± 0.5 ºC (1 SD). In addition, none of the droplets froze above -33.7 ºC in

these blank experiments. Since we only report INP number concentrations for temperatures from -15 ºC to -25 ºC in this study, heterogeneous ice nucleation by the substrate is unlikely to contribute to the reported INP number concentrations.

2.2.3 Calculating the number concentration of INPs

The number of INPs in the DFT, #INPs(T), was calculated using the following equation which accounts for the possibility of a droplet containing multiple INPs (Vali, 1971):

$$\#\mathrm{INPs}(T) = -\ln\left(\frac{N_\mathrm{u}(T)}{N_\mathrm{o}}\right) N_\mathrm{o} f_{0.25-0.10\mathrm{mm}} f_\mathrm{ne} \qquad (1)$$

where $N_\mathrm{u}(T)$ is the number of unfrozen droplets at temperature T, N_o is the total number of droplets, $f_{\mathrm{nu},0.25-0.10\mathrm{mm}}$ is a non-uniformity correction factor that takes into account non-uniformity at the 0.25-1 mm scale, and f_ne is a statistical uncertainty derived for a given number of detected nucleation events, with fewer nucleation events leading to greater statistical uncertainty (Koop et al., 1997). For the results reported, here f_ne is derived for a confidence level of 0.95.

Equation (1) assumes that droplets in a given freezing experiment have the same volume (Vali, 1971). Using droplet freezing experiments reported in Mason et al. (2015a), which are similar to experiments presented here, we explored if this assumption leads to uncertainties when applied to the DFT experiments. Number of INPs was first calculated from the DFT experiments as described above. Second, number of INPs were calculated by first separating the droplet freezing results into 2-4 bins based on droplet volume. After binning the data by droplet volume, droplets in each bin are more similar in volume. Equation 1 was then used to calculate the number of INPs in each bin. Finally, the total number of INPs was determined by summing the numbers of INPs calculated for each bin. We found that the total number of INPs determined both ways (with and without binning) agreed within the experimental uncertainty at freezing

274 temperatures of -25 °C and above for 97 % of freezing events. We also analyzed the same data
275 by first separating the droplet freezing results into 2-4 bins based on the maximum area the
276 droplets covered. Again, the number of INPs determined with and without binning was in good
277 agreement, being within the experimental uncertainty at freezing temperatures of -25 °C and
278 above for 98 % of freezing events. Based on this analysis we conclude that the application of
279 Equation (1), which assumes monodisperse droplets, to the DFT results at -25 °C and above does
280 not lead to large uncertainties.

281 In the DFT, once a droplet freezes it may grow by vapor diffusion and contact a
282 neighboring liquid droplet, causing the latter to freeze as well. To account for these non-
283 immersion freezing events, we calculated the #INPs(T) in the immersion mode using two
284 difference scenarios: (i) we calculated an upper limit by assuming that all droplets which
285 underwent the processes discussed above froze by immersion freezing; and (ii) we calculated a
286 lower limit to the INP concentration by assuming that all droplets which underwent the processes
287 discussed above remained liquid until the homogeneous freezing temperature of approximately -
288 37 °C (Wheeler et al., 2015).

289 After #INPs(T) were determined for a given freezing experiment, atmospheric INP
290 number concentrations, [INPs(T)], were calculated using the following equation:

291 $$[\mathrm{INPs}(T)] = \#\mathrm{INPs}(T)\left(\frac{A_{\mathrm{deposit}}}{A_{\mathrm{DFT}}V}\right) f_{\mathrm{nu,1mm}} \quad (2)$$

292 where A_{deposit} is the total area of the sample deposit on the MOUDI impaction plate, A_{DFT} is the
293 area of the sample analyzed by the DFT (1.2 mm^2 in all samples), V is the volume of air sampled
294 by the MOUDI, and $f_{\mathrm{nu,1mm}}$ is a correction factor to account for non-uniformity in particle
295 concentration of the sample deposit at the 1 mm scale. See Mason et al. (2015a) for details.

Values of $A_{deposit}$, $f_{nu,0.25–0.10mm}$, $f_{nu,1mm}$, and f_{ne} are given in Tables S1 and S2 and discussed in Section S1 of the Supplement. Reported INP number concentrations at each location are averaged over all samples and have been adjusted to standard temperature and pressure. In calculating averages over all sampling locations, measurements have not been weighted by sample number.

For the experimental conditions used in the current study the maximum number concentration of INPs that could be detected was roughly 20 L^{-1}. This maximum number concentration is greater than that reported in Huffman et al. (2013) because in the current studies shorter sampling times were used.

3 Results and discussion

3.1 INP number concentrations

The total number concentration of INPs active at -15, -20, and -25 °C are shown for each site in Fig. 2. Freezing events were rare at temperatures warmer than -15 °C, accounting for only 1.3 % of all cases, and are therefore not reported. Some of the DFT experiments proceeded such that all droplets froze at temperatures slightly below -25 °C. Since this scenario prohibits calculation of INP number concentrations, -25 °C is the lowest temperature reported. As expected, INP number concentrations were found to increase with decreasing freezing temperature with the average concentration at -25 °C (3.8 ± 2.9 L^{-1}) being more than an order of magnitude larger than at -15 °C (0.25 ± 0.15 L^{-1}).

INP number concentrations were relatively low at the Alert and Whistler Mountain sites with values of 0.05 and 0.10 L^{-1} at -15 °C, 0.22 and 0.16 L^{-1} at -20 °C, and 0.99 and 1.1 L^{-1} at -25 °C, respectively. These findings are consistent with previous measurements at similar locations.

For example, Arctic measurements of Bigg (1996) for particles < 10 μm and Fountain and Ohtake (1985) using filter samples found mean INP concentrations of 0.01 L^{-1} at -15 °C and 0.13 L^{-1} at -20 °C, respectively, with both deposition and condensation modes likely possible, and Prenni et al. (2007) measured an average INP number concentration for particles with and aerodynamic diameter < 1.5 μm of approximately 0.33 L^{-1} between -8 and -28 °C with deposition, condensation, and immersion modes likely possible. At high elevation sites, Bowers et al. (2009) at Mt. Werner in Colorado and Conen et al. (2012) at the research station Jungfraujoch in Switzerland measured mean immersion-mode INP number concentrations of approximately 0.02 L^{-1} at -10 and -12 ºC, respectively, using filter samples of 0.2–0.3 μm pore size. High elevation sites can receive large quantities of dust, which can act as efficient INPs at lower temperatures (Chou et al., 2011), but this was unlikely during our measurement period based on the low INP number concentrations.

INP number concentrations at Amphitrite Point were 0.23, 0.94, and 2.15 L^{-1} at droplet freezing temperatures of -15, -20, and -25 °C, respectively. Despite the predominance of marine air masses being sampled, the major source of INPs at Amphitrite Point during the study period was likely biological particles from local vegetation (Mason et al., 2015b). Similar values were measured in the Labrador Sea, where INP number concentrations at -15, -20, and -25 °C were 0.38, 1.3, and 2.8 L^{-1}, respectively. These concentrations are consistent with previous measurements within the marine boundary layer in regions influenced by air flow off of nearby coasts, for instance those of Schnell (1977) in the immersion freezing mode for sizes > 0.45 μm off the coast of Nova Scotia, roughly 1100–1500 km southwest of our sampling site in the Labrador Sea, and Rosinski et al. (1995) over the East China Sea in the deposition and condensation modes for sizes > 0.2 μm. However, INP number concentrations found during

marine studies can vary by several orders of magnitude with changing location as summarized by Burrows et al. (2013).

The highest concentrations of INPs at a freezing temperature of -25 °C were found at the Colby sites, where the average number concentration was 8.9 L^{-1}. Aerosol sampling was conducted adjacent to soya and sorghum fields during and following periods of combine operation. This high concentration of INPs is consistent with previous work of Garcia et al. (2012) that showed elevated concentrations of INPs downwind of corn fields during combine harvesting, and Bowers et al. (2011) who found greater INP concentrations in air above cropland than above suburban or forest sites.

The suburban sites of Saclay and the UBC campus also showed high INP concentrations, being 4.4 and 6.1 L^{-1}, respectively, at a freezing temperature of -25 °C. Both sites were likely influenced by multiple sources of INPs. For example, both are in close proximity to major metropolitan centers and forest vegetation, which are potential sources of anthropogenic INPs (e.g. Hobbs and Locatelli, 1970; Al-Naimi and Saunders, 1985; Knopf et al., 2010, 2014; Ebert et al., 2011; Corbin et al., 2012; Cziczo et al., 2013; Brooks et al., 2014) and biological INPs (e.g. Vali et al., 1976; Kieft and Ruscetti, 1990; Richard et al., 1996; Hirano and Upper, 2000; Diehl et al., 2002; Prenni et al., 2009; Iannone et al., 2011; Pummer et al., 2012; Huffman et al., 2013; Tobo et al., 2013; Haga et al., 2014; Wright et al., 2014), respectively. The sampling site at Saclay was also within 1 km of agricultural fields, an additional source of biological aerosols that may act as INPs (e.g. Lindow et al., 1982; Hirano et al., 1985; Georgakopoulos and Sands, 1992; Möhler et al., 2008; Bowers et al., 2011; Garcia et al., 2012; Haga et al., 2013; Morris et al., 2013; Hiranuma et al., 2015).

3.2 INP size distributions

Figure 3a shows the relative contribution of supermicron aerosol particles to the total measured INP population. Note that the same particle size range was not investigated at all locations with particles in the range of 0.10–0.18 μm not being measured at Whistler Mountain or Amphitrite Point, and no uncertainty is reported for the Labrador Sea measurement as only a single sample was available. Averaging over all sampling locations with a 1 SD uncertainty, 91 ± 9, 79 ± 17, and 63 ± 21 % of INPs had an aerodynamic diameter > 1 μm at ice activation temperatures of -15, -20, and -25 °C, respectively. At -15 °C, the percentage of supermicron INPs ranged from 78 % at Whistler Mountain up to 100 % at the Labrador Sea and Colby sites. At lower temperatures, there was more variation between samples: at -20 ºC the percentage of supermicron INPs ranged from 52 % at Whistler Mountain to 100 % over the Labrador Sea, and at -25 ºC the percentage of supermicron INPs ranged from 39 % at Whistler Mountain to 95 % over the Labrador Sea.

Figure 3b shows the fraction of INPs that are in the coarse mode, calculated by assuming that half of the INPs found in the 1.8–3.2 μm MOUDI size cut were larger than 2.5 μm (i.e. INPs are uniformly distributed over this size range). Measurements of the total particle size distribution were not available at all locations to test this assumption. Furthermore, it is not known if the INP size distribution follows the total particle size distribution *a priori*. Averaging over all sampling locations, the percentage of INPs in the coarse mode was 62 ± 20, 55 ± 18, and 42 ± 17 % (1 SD) at ice activation temperatures of -15, -20, and -25 °C, respectively. The percentage of INPs in the coarse mode was found to range from 38 % at Saclay to 91 % in Colby at -15 °C, from 26 % at Whistler Mountain to 73 % in Colby at -20 °C, and from 20 % at Alert to 64 % at the Labrador Sea at -25 °C. Despite great diversity in the studied locations, each had a significant contribution from coarse mode particles to the measured INP population.

387 The median sizes of INPs at ice activation temperatures of -15, -20, and -25 °C are shown
388 in Fig. 4 with the 25th and 75th percentile values. At -15 °C, the median INP size is relatively
389 large at all locations, varying from 2.1 μm at Saclay to 4.7 μm at Colby with an average of 3.2 ±
390 0.9 μm (1 SD). As droplet freezing temperature decreased, the median INP size also decreased,
391 with the exception of samples from the Labrador Sea and the UBC campus. At -25 °C, the
392 median size of INPs varied from 0.83 μm at Alert to 3.1 μm at the Labrador Sea site with an
393 average of 1.9 ± 1.0 μm. The median size of the INPs was > 1 μm in all cases with the exception
394 of the Alert and Whistler Mountain sites at a freezing temperature of -25 °C. Alert, in the polar
395 tundra at high latitude, and Whistler Mountain, at high elevation and periodically in the free
396 troposphere, are remote with fewer local sources of aerosols.

397 In Fig. 4, the difference between the 75th and 25th percentile sizes is relatively small at all
398 locations at -15 °C. A narrow INP size distribution at -15 °C is consistent with a single type or
399 class of particles dominating freezing at this temperature. With decreasing temperature, the
400 interquartile range significantly increased: the 75th and 25th percentile INP sizes decreased by an
401 average factor of 1.2 and 3.7, respectively, between -15 and -25 °C, corresponding to a 63 %
402 increase in the average interquartile range.

403 INP size distributions are further explored in Fig. 5, where the fraction of the measured
404 INP number concentration found in each MOUDI size bin is shown. Colors on the red end of the
405 scale illustrate that a large fraction of the INPs measured at a particular location belong to that
406 particle size bin. Total INP number concentrations as a function of temperature are given in Fig.
407 S1 and histograms of the INP size distributions are given in Figs. S2–S8 of the Supplement.
408 Figure 5a shows that most INPs active at -15 °C were 1–10 μm in size. In particular, when
409 averaged over all locations, 72 % of the INPs active at -15 °C were between 1.8 and 5.6 μm.

Furthermore, the major mode (i.e. the global maximum in a size distribution) was always larger than 1.8 μm. At lower freezing temperatures the INP size distributions broadened with increased contributions from smaller aerosol particles, evident by the more uniform intensity of Fig. 5b and c. By -25 °C, six of the seven locations had a submicron INP mode, and at Alert and Whistler Mountain this was the major mode. A general broadening of the INP size distribution with decreasing temperature would be expected if there were an increase in the number of particle types exhibiting ice activity with decreasing activation temperature.

Several previous studies have conducted size-resolved INP measurements. In most of these studies, the fraction of INPs larger than a given particle size was not reported, and in some cases the temperatures studied were different than in the current study. To better compare our data with these previous studies, we have used the literature data to calculate the fraction of INPs larger than either 1, 1.2, or 2.5 μm at temperatures as close as possible to the freezing temperatures we used. Details of the calculations are presented in the Supplement, and the results of the calculations are summarized in Table 2.

Table 2 shows that in six out of the nine previous studies, a large fraction (16–100 %) of the INPs was found to be supermicron in size, consistent with the current study. Here the work of Rosinski et al. (1986) is considered as two separate studies given the change in the investigated mode of ice nucleation. Although the condensation and immersion freezing measurements reported in Rosinski et al. (1986) appear contradictory, it is important to note that that the two freezing results for Rosinski et al. (1986) shown in Table 2 correspond to different temperature ranges (compare -5 °C to -6 °C for condensation and -10.8 °C for immersion freezing). One possibility is that small particles dominated the INP population at the warmest temperatures while larger particles dominated the INP population at the colder temperatures investigated by

433 Rosinski et al. (1986).

434 The ice nucleation efficiency of particles as a function of size have also been investigated
435 in laboratory experiments (e.g. Lüönd et al. (2010), Archuleta et al. (2005), Welti et al. (2009)).
436 In general this work has shown the ice nucleation efficiency increases as particle size increases.
437 Furthermore, several studies have investigated correlations between the concentrations of INPs
438 and aerosol particles above a certain size (Richardson et al., 2007; DeMott et al., 2010; Chou et
439 al., 2011; Field et al., 2012; Huffman et al., 2013; Prenni et al., 2013; Tobo et al., 2013; Ardon-
440 Dryer and Levin, 2014; Jiang et al., 2014, 2015). For example, using data from a variety of field
441 measurements DeMott et al. (2010) observed a correlation between INP concentrations and the
442 concentration of aerosol particles > 0.5 μm, and Ardon-Dryer and Levin (2014) found that INP
443 concentrations in Israel were better correlated to the concentration of aerosol particles 2.5–10 μm
444 in size than those < 2.5 μm. These results are also consistent with INPs being relatively large in
445 size.

446 **4 Summary and conclusions**

447 INP number concentrations in the immersion mode as a function of size and droplet
448 freezing temperature were determined at six locations across North America and one in Europe.
449 INP number concentrations varied by as much as an order of magnitude between locations, and
450 were generally found to be lowest at the remote sites of Alert and Whistler Mountain and highest
451 at the agricultural sites of Colby and the suburban sites of Saclay and the UBC campus,
452 consistent with previous studies. Several key findings indicate the potential importance of large
453 INPs at ground level: (1) 91 ± 9 and 62 ± 20 % of INPs measured at -15 °C across all locations
454 are supermicron or in the coarse mode, respectively; (2) at the lowest temperature analyzed, -25
455 °C, 63 ± 21 and 42 ± 17 % of INPs across all locations remained in the supermicron regime and

456 coarse mode, respectively; (3) at -15 °C, the median INP size was relatively large at all locations,
457 varying from 2.1 μm at Saclay to 4.7 μm at Colby with an average of 3.2 ± 0.9 μm; and (4) at -
458 25 °C, the median size INP varied from 0.83 μm at Alert to 3.1 μm above the Labrador Sea with
459 an average of 1.9 ± 1.0 μm.

460 Our measurements indicate that, when averaged over all studied locations and
461 temperatures, 78 ± 19 and 53 ± 20 % of immersion-mode INPs may be missed if either
462 supermicron particles or coarse mode particles are not sampled at ground sites. As noted in Sect.
463 1, some instrumentation for measuring ambient INP number concentrations restricts the upper
464 range of sampled aerosol particles. The data presented may be useful for estimating the fraction
465 of INPs not measured with these instruments at ground sites and in different environments.

466 All measurements used in this study were conducted at ground level and, apart from
467 those at Whistler Mountain, were also close to sea level. As the large contribution of
468 supermicron and coarse mode INPs to the overall INP population noted here may not necessarily
469 hold for higher altitudes, additional size-resolved INP measurements as a function of altitude are
470 needed. In obtaining such data, careful consideration will be needed toward sampling issues with
471 aerosol inlets and transfer through sample lines on aircraft platforms. It can be anticipated that
472 supermicron particle transfer will be strongly restricted without special inlets or care, and this
473 may especially impact INP sampling.

474 A caveat to this study is that our measurements were confined to aerosol particle sizes
475 greater than either 0.10 or 0.18 μm. If there were a significant contribution from INPs of smaller
476 sizes (Vali, 1966; Schnell and Vali, 1973; Pummer et al., 2012; Augustin et al., 2013; Fröhlich-
477 Nowoisky et al., 2015; O'Sullivan et al., 2015; Tong et al., 2015; Wilson et al., 2015), and these
478 smaller sizes did not coagulate or get scavenged by larger particles, the values presented here

would represent upper limits to the contribution of supermicron and coarse mode particles to the total INP population. Additional studies exploring the relative atmospheric abundance of INPs < 0.10 μm are necessary (Hader et al., 2014). Future studies of the size distribution of INPs should also include measurements of particle mixing state to determine if particles are internally or externally mixed at the locations where the size distribution of INPs are being measured.

Acknowledgements

The authors thank the three anonymous referees and G. Vali for helpful comments on the manuscript. The authors also thank R. B. Stull and R. Schigas for access to the UBC campus site and associated weather data, and L. A. Miller for assistance coordinating the measurements from the CCGS Amundsen. The authors also wish to thank Juniper Buller and Anton Horvath for access to the Whistler Mountain sampling site, and Freddie Lamm and the agricultural specialists at the Kansas State University Northwest Research Center for access to the Colby, KS sites, advice on harvesting time frames, and their help during harvesting. The sampling site at Amphitrite Point is jointly supported and maintained by Environment Canada, the British Columbia Ministry of Environment, and Metro Vancouver. We thank the Canadian Coast Guard and Department of Fisheries and Oceans staff from the Amphitrite Point site and the CCGS Amundsen for their help. The Natural Sciences and Engineering Research Council of Canada supported this research. K. J. Suski, P. J. DeMott, and T. C. J. Hill acknowledge support under U.S. National Science Foundation grant AGS 1358495, which also provided support for measurements at the Colby, KS site. W. M. Lassar and K. M. Pierce acknowledge funding support through the University of Denver undergraduate research center. J. A. Huffman, W. M. Lassar and K. M. Pierce acknowledge the CEA/DAM/CBRN-E research programs and particularly D. Baisnée (CEA) for support in the intensive campaign and J. Sciare (CNRS) for

the ancillary aerosol data.

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933 **Table 1.** The seven locations used in this study and conditions during sampling.

Location	Environment	Coordinates	Elevation (m)	Sampling period	Number of Samples	Average sampling time (h)	Average temperature (°C)	Average relative humidity (%)	Particle size range (μm)[a]
Alert, NU, Canada	Arctic	82.45° N 62.51° W	12 agl 200 asl	Mar. 29 - Jul. 23, 2014	9	17.6	-17.4	77	0.10–10
Whistler Mountain, BC, Canada	Alpine	50.06° N 122.96° W	2 agl 2182 asl	Mar. 30 - Apr. 23, 2014	4	6.7	0.8	83	0.18–10
Amphitrite Point, BC, Canada	Coastal	48.92° N 125.54° W	5.5 agl 25 asl	Aug. 6–27, 2013	34	7.8	13.8	97	0.18–10
The Labrador Sea, Canada	Marine	54.50° N 55.37° W	2 agl 15 asl	Jul. 11, 2014	1	6.2	10.9	75	0.10–10
CEA, Saclay, France	Suburban	48.70° N 2.14° E	10 agl 168 asl	Jul. 15 - Aug. 4, 2014	15	7.2	20.6	69	0.10–10
UBC campus, BC, Canada	Suburban	49.26° N 123.25° W	24 agl 120 asl	May 12–16, 2014	4	6.3	15.2	70	0.10–10
Colby, KS, USA	Agricultural	39.39° N 101.06° W and 39.39° N 101.08° W	2 agl 968 asl	Oct. 14–15, 2014	3	4.5	17.0	48	0.10–10

934 [a]Aerodynamic diameter based on the 50 % cutoff of the MOUDI (Marple et al., 1991).
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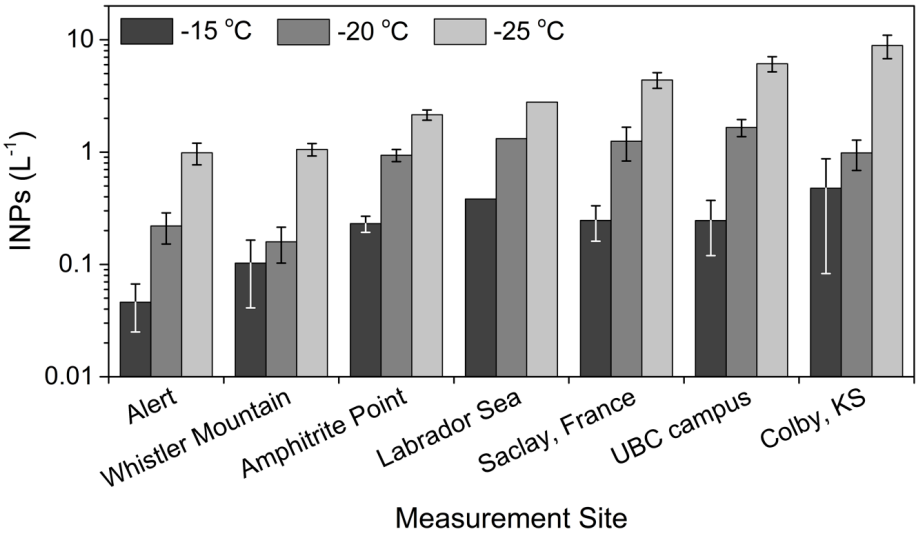
Table 2. Previous size-resolved INP measurements.

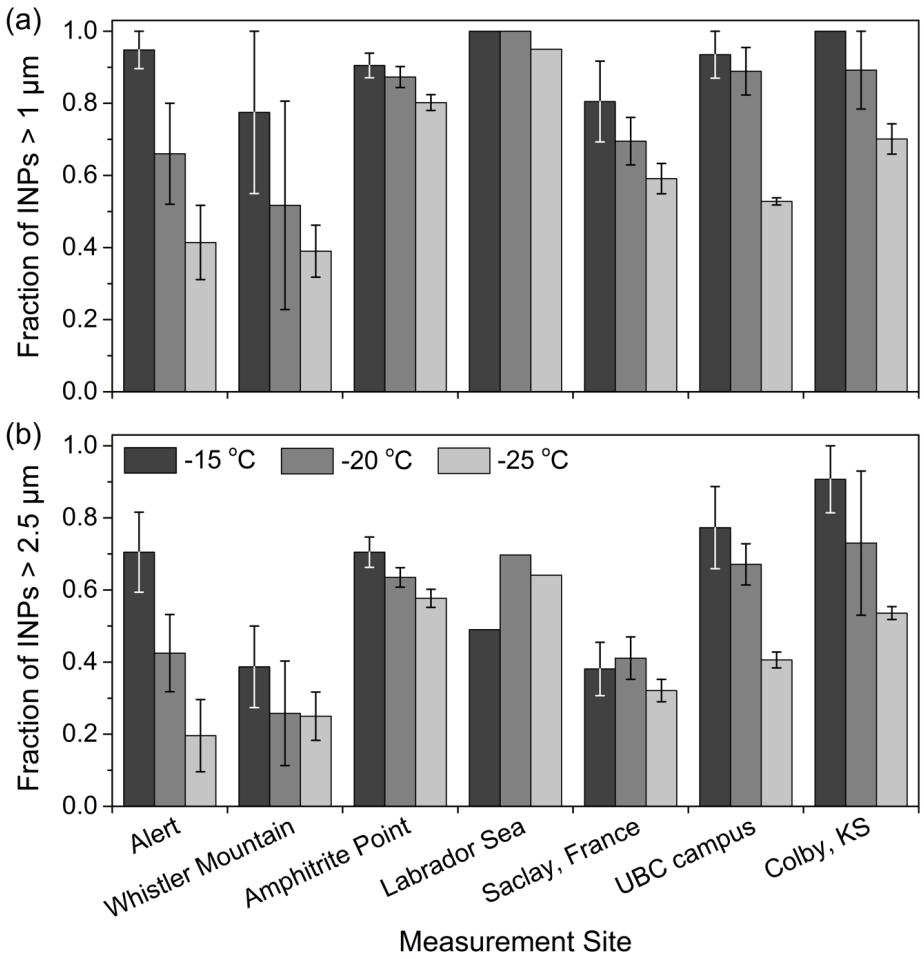
Study	Location	Geographical Description	Altitude	Particle sizes investigated (µm)	Mode of ice nucleation	Ice nucleation temperature (°C)[a]	Result
Rucklidge (1965)	West Plains, Missouri	Forest and pasture	4 m agl	TSP	Condensation and/or deposition	-12 to -25	≤ 14 % of INPs > 1 µm
Vali (1966)	Alberta, Canada	Western Canada	Hail melt water	TSP[b]	Immersion	-12.8	16 % of INPs > 1.2 µm[c]
Rosinski et al. (1986)	Central and western South Pacific Ocean	Marine	Near sea level	0.5 to > 8	Immersion	-10.8	100 % of INPs > 1 µm
				< 0.5 to > 8	Condensation	-5 to -6	1 % of INPs > 1 µm
Rosinski et al. (1988)	Gulf of Mexico	Marine	Near sea level	0.1 to > 4.5	Condensation	-15 to -16	45 % of INPs > 1 µm
Berezinski et al. (1988)	European territory of the former Soviet Union	Eastern Europe	100–500 m agl	0.1 to > 100	Condensation	-15 to -20	37 % of INPs > 1 µm
Mertes et al. (2007)	Jungfraujoch, Switzerland	Alpine	3580 m asl	0.02 to 5	Unknown (ice residual)	-17.4	< 1 % of ice residuals > 1 µm[c]
Santachiara et al. (2010)	S. Pietro Capofiume, Italy	Rural Italy	3 m agl	TSP	Condensation	-17 to -19	47 and 30 % of INPs > 1 and 2.5 µm, respectively
Huffman et al. (2013)	Manitou Experimental Forest, CO, USA	Forest during/after rainfall	4 m agl	0.32 to >18	Immersion and deposition	-15 to -20	89 % of INPs > 1 µm
		Forest during dry periods	4 m agl	0.32 to >18	Immersion and deposition	-15 to -20	46 % of INPs > 1 µm

[a]We used data at the temperatures of this study (-15, -20, and -25 °C) when available, otherwise the next closest temperature was used.
[b]TSP = total suspended particulate.
[c]Particle size reported for Vali (1966) is based on filter pore size, and particle size reported in Mertes et al. (2007) is based on electrical mobility and optical measurements . In all other studies, particle size is given as the aerodynamic diameter.

Figure 1. Sampling locations used in this study: (1) Alert, Nunavut, Canada; (2) Whistler Mountain, British Columbia, Canada; (3) Amphitrite Point, British Columbia, Canada; (4) the Labrador Sea, Canada; (5) CEA in Saclay, France; (6) the University of British Columbia campus, British Columbia, Canada; and (7) Colby, Kansas, USA. Site coordinates are given in Table 1 with details in Sect. 2.1. The image was modified from Bing Maps, 2014 (http://bing.com/maps).



966 **Figure 2.** Mean INP number concentrations at droplet freezing temperatures of -15 °C (dark
967 grey), -20 °C (intermediate grey), and -25 °C (light grey). Uncertainty is given as the standard
968 error of the mean, assuming a normal distribution. As only one sample was available from the
969 Labrador Sea, no uncertainty is reported.



970 **Figure 3.** The mean fraction of INPs larger than (a) 1 µm and (b) 2.5 µm. Uncertainty is the
971 standard error of the mean, assuming a normal distribution. Shading in the histogram
972 corresponds to INP activation temperature: -15 °C is dark grey, -20 °C is an intermediate grey,
973 and -25 °C is light grey. As 2.5 µm does not align with the size cut of a MOUDI stage, the
974 fraction of INPs larger than 2.5 µm was found by assuming that number concentration of INPs
975 1.8–3.2 µm in size was uniformly distributed over that size range. As only one sample was
976 available from the Labrador Sea, no uncertainty is reported.

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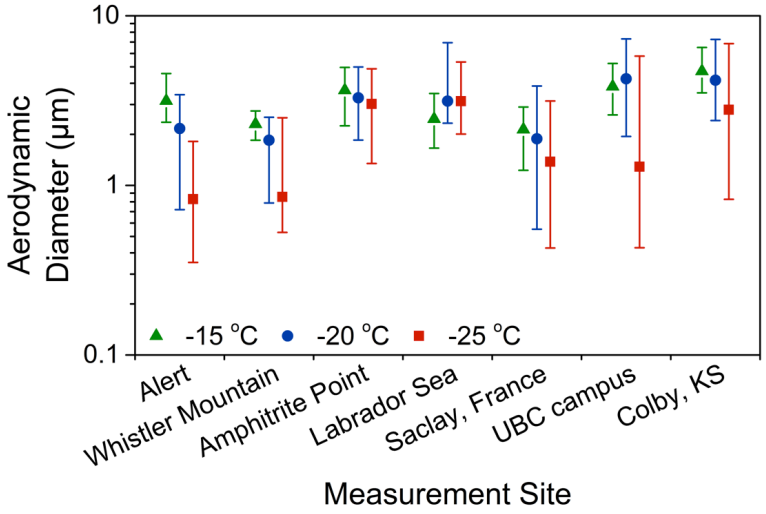
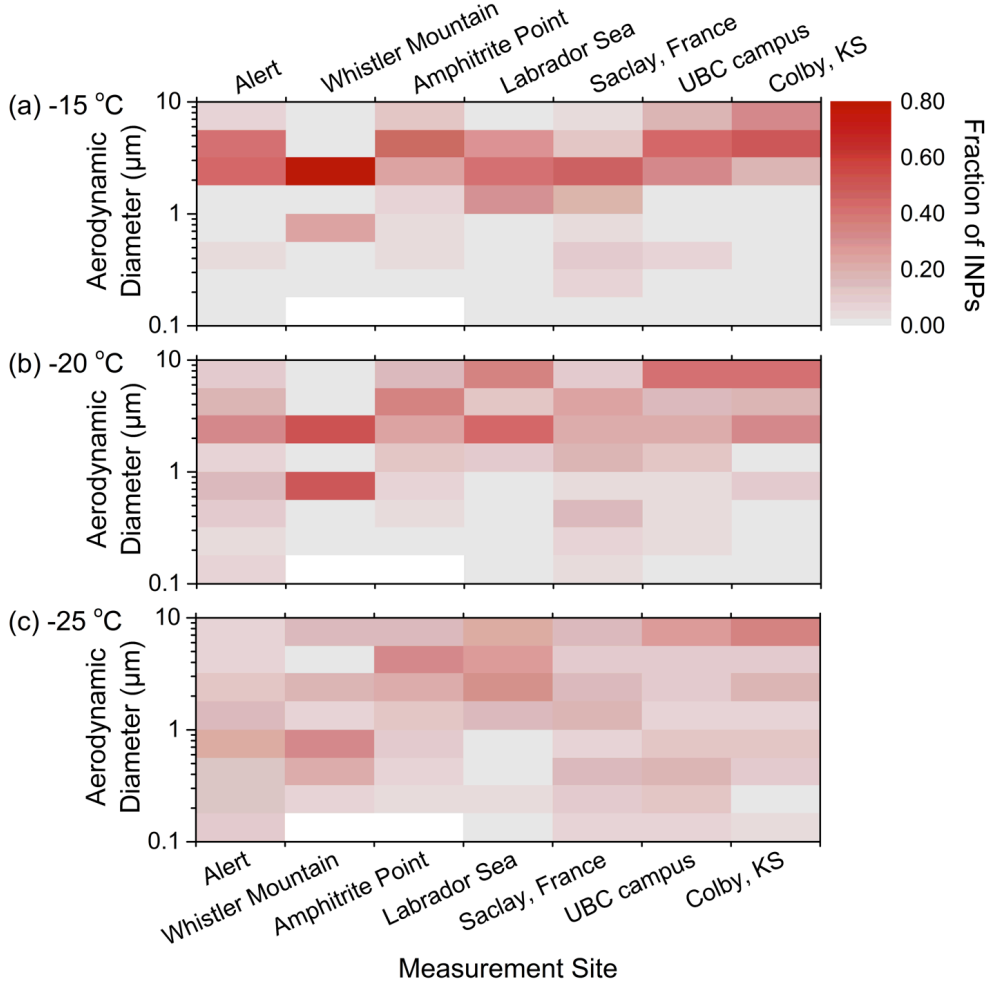


Figure 4. The median size of INPs at ice-activation temperatures of -15 °C (green), -20 °C (blue), and -25 °C (red) when averaged over all analyzed samples. Upper and lower uncertainties are the 75th and 25th percentiles, respectively.



984 **Figure 5.** Fractional INP concentrations as a function of aerosol particle size, location, and
985 activation temperature: (a) -15 °C; (b) -20 °C; and (c) -25 °C. The color bar indicates the fraction
986 of INPs measured in each particle size bin. Aerosol particle sizes correspond to the 50 % cutoff
987 aerodynamic diameters of the MOUDI stages (Marple et al., 1991). Missing sizes for the
988 Whistler Mountain and Amphitrite Point sites are uncolored.