



## Abstract

Interaction of biogenic volatile organic compounds (VOC) with anthropogenic VOC affects the physicochemical properties of secondary organic aerosol (SOA). We investigated cloud droplet activation (CCN activity), droplet growth kinetics, and hygroscopicity of mixed anthropogenic and biogenic SOA (ABSOA) compared to pure biogenic SOA (BSOA) and pure anthropogenic SOA (ASOA). Selected monoterpenes and aromatics were used as representative precursors of BSOA and ASOA, respectively.

We found that BSOA, ASOA, and ABSOA had similar CCN activity despite the higher oxygen to carbon ratio (O/C) of ASOA compared to BSOA and ABSOA. For individual reaction systems, CCN activity increased with the degree of oxidation. Yet, when considering all different types of SOA together, the hygroscopicity parameter,  $\kappa_{\text{CCN}}$ , did not correlate with O/C. Droplet growth kinetics of BSOA, ASOA, and ABSOA was comparable to that of  $(\text{NH}_4)_2\text{SO}_4$ , which indicates that there was no delay in the water uptake for these SOA in supersaturated conditions.

In contrast to CCN activity, the hygroscopicity parameter from hygroscopic tandem differential mobility analyzer (HTDMA) measurement,  $\kappa_{\text{HTDMA}}$ , of ASOA was distinctly higher (0.09–0.10) than that of BSOA (0.03–0.06), which was attributed to the higher degree of oxidation of ASOA. The ASOA components in mixed ABSOA enhanced aerosol hygroscopicity. Changing the ASOA fraction by adding BVOC to ASOA or vice versa AVOC to BSOA changed the hygroscopicity of aerosol, in line with the change in the degree of oxidation of aerosol. However, the hygroscopicity of ABSOA cannot be described by a simple linear combination of pure BSOA and ASOA systems. This indicates that additional processes, possibly oligomerization, affected the hygroscopicity.

Closure analysis of CCN and HTDMA data showed  $\kappa_{\text{HTDMA}}$  was lower than  $\kappa_{\text{CCN}}$  by 30–70%. Better closure was achieved for ASOA compared to BSOA. This discrepancy can be attributed to several reasons. ASOA seemed to have higher solubility in sub-

saturated conditions and/or higher surface tension at the activation point than that of BSOA.

## 1 Introduction

Secondary organic aerosol (SOA) is an important class of atmospheric aerosol with impacts on air quality, human health and climate change (Hallquist et al., 2009; Kanakidou et al., 2005; Jimenez et al., 2009; Zhang et al., 2011; Verma et al., 2014). Despite substantial improvements in the understanding of SOA formation mechanisms and properties, considerable uncertainties remain about the regional and global budget of SOA (e.g. Goldstein and Galbally, 2007). Models often do not correctly predict the ambient concentrations of organic aerosol (OA) (e.g., Spracklen et al., 2011; Heald et al., 2005), and usually the modeled concentrations underestimate the observed OA concentrations (Spracklen et al., 2011). Recent studies suggested that interactions between biogenic volatile organic compounds (VOC) and anthropogenic emissions can enhance SOA formation and often, ambient OA concentrations correlate with anthropogenic tracers such as CO or isopropyl nitrate (de Gouw et al., 2008, 2005; Weber et al., 2007; Shilling et al., 2013; Xu et al., 2015). However,  $^{14}\text{C}$  carbon isotopic analysis showed that the SOA material itself is composed mostly of modern carbon, i.e. arises from biogenic sources, including biomass burning (Weber et al., 2007; Zotter et al., 2014). These observations suggest that a consideration of an anthropogenic enhancement can reduce the discrepancies between models and observations (de Gouw et al., 2005; Goldstein et al., 2009; Hoyle et al., 2011; Worton et al., 2011; Glasius et al., 2011).

Anthropogenic VOCs (AVOC), such as aromatic compounds are possibly important factors that lead to enhanced SOA formation as their oxidation products can interact with biogenic VOC (BVOC) oxidation products during SOA formation, as shown by several studies (Hoyle et al., 2011; Emanuelsson et al., 2013; Flores et al., 2014). In a recent study, Emanuelsson et al. (2013) found that anthropogenic SOA (ASOA) com-

ponents reduce the volatility of biogenic SOA (BSOA) in a non-linear way with respect to the ASOA fraction, possibly by oligomerization or a phase change such as formation of a glassy state (Emanuelsson et al., 2013; Virtanen et al., 2010; Koop et al., 2011). The reduced volatility in the mixed SOA (anthropogenic-biogenic SOA, ABSOA) can enhance SOA persistence and concentrations in the atmosphere. Flores et al. (2014) investigated the optical properties of BSOA, SOA from simultaneous addition of BVOC and AVOC and SOA from sequential addition of BVOC and AVOC. They found that both SOA from mixed AVOC and BVOC show an increase of scattering component of the refraction index with aging (increase of the oxygen to carbon ratio (O/C)) and the increase is greater for SOA from simultaneous addition of BVOC and AVOC than SOA from sequential addition of VOC.

Besides the thermochemical and optical properties, cloud droplet activation (cloud condensation nuclei (CCN) activity) and hygroscopicity are important physicochemical properties that have critical implications for the impact of aerosol on climate. It is possible that enhanced oligomerization which happens in the mixed aerosol particles could modify its CCN activity and hygroscopicity (Xu et al., 2014). Given that CCN activity and hygroscopicity correlate with the aerosol O/C in many cases and given that ASOA tends to have a higher O/C (Chhabra et al., 2011; Emanuelsson et al., 2013), it is expected that ASOA components enhance the CCN activity and hygroscopicity of mixed SOA. However, to the best of our knowledge, the influence of the interaction of AVOC with BVOC on CCN activity and hygroscopicity of SOA has not been studied from the literature.

Several field studies found a delay in droplet growth kinetics of the aerosol from anthropogenic origin when compared with the aerosol from biogenic origin (Shantz et al., 2010, 2012). ASOA, as an important anthropogenic aerosol, may contribute to this delay. In addition, a recent laboratory study suggests limited mixing in SOA formed by sequentially mixing a biogenic precursor ( $\alpha$ -pinene) with an anthropogenic one (toluene) (Loza et al., 2013). It is possible that limited mixing could affect water uptake by ABSOA and delay droplet growth when the diffusion of water within aerosol

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particle was suppressed. Yet, no laboratory studies on the droplet growth kinetics of ASOA and ABSOA have been found.

In this study, we investigated the effect of the interaction of ASOA and BSOA on CCN activity and hygroscopicity of aerosol. We also studied the kinetics of droplet growth of ASOA, BSOA and mixed ABSOA.

## 2 Experimental

### 2.1 Experiment setup and instrumentation

The experiments were conducted in the atmosphere simulation chamber SAPHIR (Simulation of Atmospheric PHotochemistry In a large Reaction chamber). SAPHIR is a double-wall Teflon chamber with a volume of 270 m<sup>3</sup>. Details of the chamber have been previously described (Rohrer et al., 2005; Bohn et al., 2005). The chamber uses natural sunlight for illumination and is equipped with a louvre system which can be used to simulate dark processes. For the experiments described here, various instrumentations were used to characterize gas phase and the particulate phase species.

Chamber parameters like temperature, relative humidity, flow rate, and photolysis frequencies were also recorded. The actinic flux and the corresponding photolysis frequencies were provided from measurements using a spectral radiometer (Bohn et al., 2005; Bohn and Zilken, 2005).

The number concentration and size distributions of aerosol were measured by a scanning mobility particle sizer (SMPS, DMA model 3081/CPC model 3785, TSI Shoreview, USA) and separate condensation particle counter (CPC, model 3786, TSI) to allow detection of nucleation particles down to 3 nm.

The chemical composition of aerosol was measured by a High-Resolution Time-of-Flight Aerosol Mass Spectrometer (HR-ToF-AMS, Aerodyne Research Inc., USA). To characterize the degree of oxidation of aerosol, the oxygen to carbon ratio (O/C) and  $f_{44}$  (fractional contribution of  $m/z$  44 to the total organics signal) were obtained from the

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mass spectra. The O/C (corrected for the minor influence of gaseous components) was derived by the elemental analysis of mass spectra obtained in the high mass resolution mode (W-mode) of the mass spectrometer as described by Aiken et al. (2007) and Aiken et al. (2008). The contributions of gas phase CO<sub>2</sub> and water vapor to  $m/z$  44 and to  $m/z$  18, respectively, were characterized using a CO<sub>2</sub> and H<sub>2</sub>O analyzer (Picarro, Santa Clara, USA). The values were subtracted to obtain the particle signals before the elemental analysis (Allan et al., 2004).

Droplet activation and droplet growth were measured using a size scanning CCN method as described previously (Buchholz, 2010; Zhao et al., 2010). This method, also known as Scanning Mobility CCN analysis (SCMA, Moore et al., 2010), has been successfully used in a number of previous studies (Asa-Awuku et al., 2008, 2009, 2010; Padro et al., 2007; Engelhart et al., 2008, 2011). The measurement was done by coupling a differential mobility analyzer (DMA, model 3081, TSI Shoreview, USA) with a cloud condensation nuclei counter (CCNC, Droplet Measurement Technique, USA) and condensation particle counter (CPC3786, TSI). Particles first passed through the DMA and the outgoing air was split into two paths connecting to the CCNC and CPC, which measure the CCN and cloud nuclei (CN) concentrations, respectively. The DMA scanned over a size range between 10 and 450 nm while the supersaturation (SS) remained constant. And four to five different supersaturations in the range of 0.1–1.3% were used depending on the particle sizes. From the measurement, CCN activation fraction over size and the dry activation diameter (or critical dry diameter,  $D_{\text{crit}}$ ) was obtained using a method as described in Buchholz (2010). Briefly, for each particle size, the CN and CCN concentrations measured were used to calculate the activation fraction ( $a_f = \text{CCN}/\text{CN}$ ). Before  $a_f$  was calculated, the measured CN and CCN concentrations were corrected for multiple charged particles. To separate the single from the multiple charged particles, the fraction of multiple charged particles was calculated according to a Boltzmann charge distribution using the measured size distribution (Wiedensohler, 1988). Then,  $a_f$  was determined for each charge class separately and

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fitted with a Gaussian error function (Rose et al., 2008). The dry activation diameter at the set SS is the turning point of this function.

For each SS at least three full scans were performed and the resulting  $D_{\text{crit}}$  were averaged. For the calibration of SS,  $D_{\text{crit}}$  of ammonium sulfate at various SS was measured and compared to theoretical data in the literature (Rose et al., 2008). The set SS was corrected according to the theoretical data. From the CCN data, the hygroscopicity parameter  $\kappa_{\text{CCN}}$  was calculated according to the one parameter representation of the Köhler equation proposed by Petters and Kreidenweis (2007). A higher hygroscopicity parameter  $\kappa$  indicates a more hygroscopic material, i.e. cloud droplet activation at lower SS for particles of a given size or at smaller size for a given SS.

The hygroscopic growth of the aerosol was measured using a home-built hygroscopic tandem differential mobility analyzer (HTDMA). The details of the HTDMA were described previously (Buchholz, 2010; Zhao et al., 2010). Particles were selected using the first DMA and then were exposed to a prescribed relative humidity to measure the growth factor. Hygroscopic growth was measured at different RH. The sizes of the humidified particles were determined by the second DMA which was operated in a scanning mode in combination with a CPC (model 3022A, TSI). The size selected aerosol flow and the sheath air flow of the second DMA were humidified at room temperature (25–30 °C depending on the surroundings) to almost the same relative humidity (RH) with the sheath air at slightly higher RH. The second DMA was kept in a thermo-insulated box which was cooled to 20 °C. Both aerosol and sheath air flow were cooled down to the same temperature before entering the second DMA, and thus the RH increased to its final value. The residence time of particles at the final humidity is approximately 30 s before they entered the SMPS operated with sheath air of the same RH. The hygroscopic growth factor (GF) was calculated as the ratio of the size of the wet particle ( $D_{\text{wet}}$ ) to the selected dry size ( $D_{\text{dry}}$ ). The HTDMA was calibrated using ammonium sulfate aerosol by comparing with the theoretical growth curve (Rose et al., 2008). From hygroscopic growth factor at 90% RH, the hygroscopicity parameter,  $\kappa_{\text{HTDMA}}$ , was calculated according to Petters and Kreidenweis (2007).

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SOA samples were collected on PTFE filters at the end of different experiments to obtain detailed insight into the chemical composition of the aerosol particles. The details of sample collection and analysis are described in Emanuelsson et al. (2013) and Kristensen and Glasius (2011). Before the filters, the air passed through an annular denuder coated with XAD-4 resin to remove gaseous organic species. The filters were extracted and analyzed using a Dionex Ultimate 3000 HPLC system coupled through an electrospray (ESI) inlet to a q-TOF mass spectrometer (micro-TOFq, Bruker Daltonics GmbH, Bremen, Germany), which was operated in both positive and negative mode. Pinonic acid, cis-pinic acid, terpenylic acid, diaterpenylic acid acetate (DTAA) and 3-methyl butane tri-carboxylic acid (MBTCA) were quantified using authentic standards.

For SOA from part of the experiments (experiment #B3, AB4, AB6 as in Table 1), samples were also collected on quartz fiber filters and analyzed by ultra-high resolution mass spectrometry (UHRMS). In this analysis, the aerosol samples were extracted as described elsewhere (Kourtchev et al., 2013). The extracts were analyzed using an ultrahigh resolution LTQ Orbitrap Velos mass spectrometer (Thermo Scientific, Bremen, Germany) equipped with a TriVersa Nanomate robotic nanoflow chip-based ESI source (Advion Biosciences, Ithaca NY, USA). The Orbitrap MS instrument calibration, settings and mass spectral data interpretation are described in Kourtchev et al. (2014). The mass accuracy of the instrument was below 1.5 ppm and the instrument mass resolution was 100 000 at  $m/z$  400. The negative ionization mass spectra were collected in three replicates and in this study, only ions that appeared in all three analytical replicates were kept for evaluation.

The VOC were measured by a high resolution proton transfer reaction-mass spectrometer (HR-PTR-MS, Ionicon, Innsbruck, Austria) (Jordan et al., 2009) and gas chromatography coupled to a mass spectrometer (GC-MS, PerkinElmer, Waltham, USA) (Apel et al., 2008; Kaminiski, 2014).

The OH concentration was measured directly using laser-induced fluorescence (LIF) (Fuchs et al., 2012). The OH radicals inside the chamber are mainly formed by the pho-

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tolysis of HONO formed via a photolytic process on the chamber walls, and to a minor fraction by  $O_3$  photolysis (Rohrer et al., 2005). From the OH concentration, the OH dose was calculated and used as a common abscissa in order to better compare different experiments. The OH dose is the integral OH concentration over time that gives the accumulated OH concentrations to which gases and particles were exposed in the course of experiment. One hour exposure to a typical atmospheric OH concentration of  $2 \times 10^6$  molecules  $cm^{-3}$  is then equivalent to an OH dose of  $7.2 \times 10^9$  molecules  $cm^{-3}$  s.

## 2.2 Experimental procedure

The experimental procedures have been described elsewhere in details (Emanuelsson et al., 2013; Flores et al., 2014) and only a short description is given here. In the BSOA experiments, a monoterpene mixture of  $\alpha$ -pinene and limonene with a molar ratio of 1 : 1 was used as the representative BSOA precursors and its photooxidation induced BSOA formation. Ozone was added to initialize BVOC oxidation and particle formation. In the ASOA experiments, toluene or xylene was used as the representative ASOA precursors. In the mixed SOA experiments (ABSOA), AVOC and BVOC were added, either simultaneously or sequentially to investigate the potential effect of adding order.

In total, three BSOA experiments (including one using  $\alpha$ -pinene as precursor), seven ASOA experiments, and six mixing experiments (ABSOA) with biogenic and anthropogenic precursors were analyzed in detail and the summary for these experiments is given in Table 1. In two experiments (#AB1, AB2), AVOC was added 6.3 h before the BVOC. In two experiments (#AB3, AB4), BVOC was added 2.5 and 5 h before the AVOC, respectively. In the other two experiments (#AB5, AB6), BVOC and AVOC were added simultaneously into the chamber. In the ABSOA experiments, the fraction of ASOA in the total aerosol was estimated using a method based on the aerosol mass yield and VOC consumed as described by Emanuelsson et al. (2013).

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although ASOA components increased the O/C of ABSOA and decreased its volatility (Emanuelsson et al., 2013), it did not enhance its CCN activity. We conclude that the interaction of AVOC with BVOC during the photochemical oxidation does not affect CCN activity of the SOA. Decreasing volatility and invariant CCN activity at higher O/C induced by presence of ASOA components indicates different effects of the main aging processes, oligomerization and photochemistry, on volatility and CCN activity of aerosol. While both oligomerization and photochemical aging reduce the volatility (Emanuelsson et al., 2013), they have counterbalancing effects on CCN activity. Oligomerization increases the molar volume and decreases the molar concentration of solutes, which results in an increase of water activity and thus the saturation ratio. Therefore, oligomerization reduces CCN activity, while photochemical aging enhances the CCN activity by producing more oxygenated compounds via functionalization (Zhao et al., 2015b). Functionalization does not significantly change the molar volume of organics, but it increases the van't Hoff factor due to both its interaction of polar functional groups with water and additional dissociation effects by carboxylic and hydroxyl group (as in Eq. A6) (Petters et al., 2009; Tritscher et al., 2011). While the effects of oligomerization and photochemical aging on the surface tension of the droplets should also be considered, it is outside the scope of this study and we assume no or compensating effects. In any case, in this study we did not observe a significant difference in CCN activity between ABSOA and other SOA, indicating that the effects of possible oligomerization on the CCN activity of ABSOA are likely compensated by the effects of photochemical aging.

For BSOA and ABSOA,  $\kappa_{\text{CCN}}$  generally increases with O/C at a given SS (Fig. S2). For ASOA, the relationship between  $\kappa_{\text{CCN}}$  and O/C is not so apparent. The  $\kappa_{\text{CCN}}$  values of BSOA (0.11–0.19) are consistent with  $\kappa_{\text{CCN}}$  values obtained in previous studies (Massoli et al., 2010; Lambe et al., 2011; Frosch et al., 2011). An increase of  $\kappa_{\text{CCN}}$  with O/C of SOA has been reported by a number of previous studies (Massoli et al., 2010; Lambe et al., 2011). Yet, other studies reported that  $\kappa_{\text{CCN}}$  of SOA is largely independent of O/C or  $f_{44}$ , another proxy of the degree of oxidation of organic aerosol (Frosch et al.,

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2011; Alfara et al., 2013). The inconclusive results in the literature may be attributed to the differences in reaction conditions including the VOC type and concentration,  $\text{NO}_x$  concentration, photochemical aging (OH dose) and O/C range. In this study, the observation that  $\kappa_{\text{CCN}}$  is invariant on O/C in ASOA may be associated with the relatively high and narrow range of observed O/C for ASOA. Emanuelsson et al. (2013) proposed that first generation products of AVOC (mainly carbonyls) may have higher vapor pressure compared to BVOC products. Therefore, AVOC need a higher OH dose and more oxidation steps before a significant amount of material starts to condense on the particles. As a result, once SOA started to form, ASOA had already reached a relatively high O/C due to aging in the gas phase (Fig. 1b). As a consequence, a further increase of O/C in ASOA is likely slower and less pronounced.

Considering all types of SOA investigated here,  $\kappa_{\text{CCN}}$  did not correlate with O/C here (Fig. 1b), which is in agreement with the findings of Alfara et al. (2013) that water uptake is independent of  $f_{44}$  as a metric of the degree of oxidation for multiple biogenic precursors. This is because O/C affects the solubility of the solute but other effects such as molecular weight, surface tension and density are also important. Massoli et al. (2010) also showed, the correlation between  $\kappa_{\text{CCN}}$  derived from CCN and O/C is less significant as in individual systems, when considering all systems together including  $\alpha$ -pinene, trimethylbenzene, and m-xylene.

For ASOA systems, particle formation was studied for different aromatic precursors at low  $\text{NO}_x$  ( $\text{NO}_x < 1$  ppb) and high  $\text{NO}_x$  concentrations (10 ppb of NO added). The CCN activity of ASOA from different precursors including toluene, xylene, and benzene was found to be similar. Although  $\text{NO}_x$  levels affect the overall SOA particle composition and yields (Ng et al., 2007a, b; Hildebrandt et al., 2009; Zhang et al., 2014; Presto et al., 2005; Eddingsaas et al., 2012), they had no significant effect on the CCN activity of ASOA comparing the low  $\text{NO}_x$  ( $\text{NO}_x < 1$  ppb) with high  $\text{NO}_x$  (10 ppb NO added) condition (Fig. S3). The CCN activity of ASOA generally agreed with the range obtained from a number of previous studies (Massoli et al., 2010; Prenni et al., 2007; Lambe et al., 2011). Similarly,  $\text{NO}_x$  level has been found to not influence the CCN activity of

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SOA from  $\alpha$ -pinene ozonolysis and subsequent photochemical aging (Frosch et al., 2011).

### 3.1.2 Droplet growth kinetics

Droplet growth kinetics was investigated using the method of “threshold droplet growth analysis” (TDGA), which has been used successfully in many field and laboratory studies (Engelhart et al., 2008; Asa-Awuku et al., 2009, 2010; Bougiatioti et al., 2011). In this method, the droplet growth kinetics was assessed by comparing the droplet sizes from various SOA with that from ammonium sulfate, which is highly hygroscopic and rapidly grows under supersaturated condition. When two particles are exposed to the same SS, they will grow to droplets of similar size, if their critical SS and the mass transfer of water vapor are similar. In this study, the TDGA method was applied to the size-resolved CCNC data and droplet size was compared for activated particles with  $SS_{crit}$  equal to the instrument SS (i.e. particles with a size of  $D_{crit}$ ) (Asa-Awuku et al., 2009, 2010; Bougiatioti et al., 2011).

Droplet size as a function of SS for BSOA, ASOA and ABSOA were compared with that of ammonium sulfate (Fig. 3). The droplet sizes of BSOA, ABSOA or ASOA are similar with those of ammonium sulfate. This indicates the absence of a kinetic barrier for the water uptake of these SOA during droplet activation. Our study is in agreement with several previous studies showing comparable droplet growth kinetics of SOA from monoterpenes with that of ammonium sulfate (Engelhart et al., 2008; Frosch et al., 2011). For SOA from toluene or xylene, no report on droplet growth kinetics was found in the literature. The droplet growth of aerosol from anthropogenic sources in the field containing both organics and ammonium sulfate has been shown to be slower than that of the pure ammonium sulfate, using a static diffusion cloud condensation chamber (Shantz et al., 2010, 2012). Based on our study, ASOA from common aromatics, does not explain such delay and the observations by Shantz and co-workers must have been caused by other aerosol components (Shantz et al., 2010, 2012).

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A recent laboratory study of Loza et al. (2013) suggests limited mixing in the SOA formed from BSOA precursor  $\alpha$ -pinene and ASOA precursor toluene added sequentially. In contrast, the study of Hildebrandt et al. (2011) supported the pseudo-ideal mixing of BSOA and ASOA components according to the aerosol mass yield. Based on the droplet growth kinetics of the ABSOA (Fig. 3) observed in this study, either mixing did occur or if limited mixing of ASOA and BSOA components occurred but did not seem to affect the water uptake in the supersaturated conditions.

### 3.2 Hygroscopicity of SOA from HTDMA measurement

Figure 4 shows the hygroscopicity ( $\kappa_{HTDMA}$ ) of BSOA, ASOA, and ABSOA.  $\kappa_{HTDMA}$  of BSOA was between 0.03 and 0.06, and increased slightly with OH dose.  $\kappa_{HTDMA}$  of ASOA was around 0.09–0.1, significantly higher than that of BSOA. Subsaturated hygroscopic growth of ASOA and BSOA was obviously different whereas their CCN activity was basically similar. The influence of the SOA types on the hygroscopic growth is different from their influence on CCN activity. The comparison between the water uptake in the subsaturated conditions from hygroscopic growth and that in the supersaturated conditions from CCN activity is discussed in the Sect. 3.3.  $\kappa_{HTDMA}$  of ASOA did not change much with the OH dose and  $\kappa_{HTDMA}$  of ASOA from different aromatic precursors (toluene, xylene, and benzene) were similar (#A1, 2, 4, 5 in Fig. 4).

The higher  $\kappa_{HTDMA}$  of ASOA can be related to the chemical composition represented by O/C. The O/C of BSOA was about 0.3–0.5, distinctively lower than that of ASOA (0.7–0.8). O/C has been found to correlate with  $\kappa_{HTDMA}$  for various SOA systems (Jimenez et al., 2009; Massoli et al., 2010; Duplissy et al., 2008, 2011; Lambe et al., 2011). The same arguments as for CCN activity apply here. ASOA reached much higher O/C at the same OH dose compared to BSOA (Fig. S4) (Emanuelsson et al., 2013) because first generation products of AVOC have a smaller carbon number and higher vapor pressure compared to BVOC. Thus, first generation products of AVOC need thus more oxidation steps before starting to condense significantly on particles as discussed in Sect. 3.1. In addition, constituents of aromatic ASOA generally

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Morphological effect can also play a role. If the ASOA and BSOA components were not well mixed in the aerosol particles in the experiments with sequential VOC additions, there would be more BSOA components on SOA particle surface in the experiments #AB1 and #AB2. This could affect the  $\kappa_{\text{HTDMA}}$  and contributed to the non-linear effect. But this cannot explain the non-linear effect in the experiment with VOCs added simultaneously.

The ABSOA filter samples from experiment #AB4 and #AB6 were extracted and analyzed for oligomers. We did not find indications that ABSOA contained more dimers compared to BSOA. This can be attributed to the low ASOA fraction  $\leq 5\%$  in experiments #AB4 and #AB6 (estimated using the method as in Emanuelsson et al., 2013). The low ASOA fraction was caused by the low OH concentration and low chemical turnover of the aromatics in these experiments because high concentrations of VOC were used in order to generate enough particle mass for optical measurement (Flores et al., 2014). The low fraction of ASOA resulted in little oligomer formation by the interaction between the ASOA components and BSOA components.

### 3.3 Closure between the hygroscopicity parameter from CCN and HTDMA

The hygroscopicity parameter  $\kappa$  was obtained from CCN and HTDMA measurements in supersaturated and subsaturated conditions, respectively. For all SOA types studied here, there is a significant gap between  $\kappa_{\text{HTDMA}}$  and  $\kappa_{\text{CCN}}$ .  $\kappa_{\text{HTDMA}}$  was significantly lower than  $\kappa_{\text{CCN}}$  with  $\kappa_{\text{HTDMA}}/\kappa_{\text{CCN}}$  around 0.3–0.7 (Fig. 7). The ratio of  $\kappa_{\text{HTDMA}}/\kappa_{\text{CCN}}$  for BSOA and ABSOA was lower than that of ASOA, which is closer to 1. This means that there is a smaller gap between  $\kappa_{\text{HTDMA}}$  and  $\kappa_{\text{CCN}}$  for ASOA compared to BSOA and ABSOA.

The closure between  $\kappa_{\text{CCN}}$  and  $\kappa_{\text{HTDMA}}$  of SOA has been studied and discussed by a number of previous studies with varying results (Dusek et al., 2011; Alfarra et al., 2013; Good et al., 2010; Duplissy et al., 2008; Juranyi et al., 2009; Prenni et al., 2007; Massoli et al., 2010; Wex et al., 2009; Whitehead et al., 2014). The discrepancy between  $\kappa_{\text{CCN}}$  and  $\kappa_{\text{HTDMA}}$  found here can be attributed to several possible reasons as

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discussed in the literature (Prenni et al., 2007; Massoli et al., 2010; Frosch et al., 2011; Good et al., 2010; Alfarra et al., 2013; Wex et al., 2009; Whitehead et al., 2014; Petters et al., 2009; Dusek et al., 2011). An important reason is the presence of slightly soluble compounds. These compounds only dissolve partly in the subsaturated condition while they can dissolve completely in the supersaturated conditions due to more water available. Therefore,  $\kappa$  is underestimated to varied extent in the subsaturated condition. ASOA components here seemed to have higher solubility compared to BSOA components, and thus the gap between  $\kappa_{\text{CCN}}$  and  $\kappa_{\text{HTDMA}}$  was smaller than that of BSOA.

Surface tension can also play a role in this discrepancy.  $\kappa$  was calculated using the surface tension of pure water ( $0.073 \text{ N m}^{-1}$ ). If the surface tension of the droplets is lower than that of water,  $\kappa$  would be overestimated. While  $\kappa_{\text{HTDMA}}$  is not so sensitive to the change of surface tension,  $\kappa_{\text{CCN}}$  is quite sensitive due to an approximate cubic relationship between  $\kappa_{\text{CCN}}$  and surface tension (Petters and Kreidenweis, 2007). The surface tension under the subsaturated conditions is assumed to be lower than that under the supersaturated conditions due to the more concentrated organics in the droplets under the subsaturated condition (Prisle et al., 2008). If the surface tension effect for BSOA would be larger than for ASOA, i.e. surface tension in the supersaturated condition is lower, this could explain the larger discrepancy between  $\kappa_{\text{HTDMA}}$  and  $\kappa_{\text{CCN}}$  for BSOA. However, surface active organics can be enriched at the surface to such a high extent that the Raoult term is significantly diminished (Prisle et al., 2008). This difference can compensate the overestimation by using the surface tension of water and the compensating effects make using surface tension of water be a reasonable choice.

Furthermore, the  $\kappa$ -Köhler model does not account explicitly for changes in non-ideality of a solution as a function of saturation ratio i.e. water activity. Instead,  $\kappa$  might be not constant through the whole range of water activity (Petters and Kreidenweis, 2007). In addition, different aerosol behaviors such as evaporation and condensation of organics in HTDMA and CCN instrument and limited diffusivity of water in case of glassy particles can contribute to the discrepancy (Whitehead et al., 2014; Asa-Awuku

#### 4 Conclusions

We investigated the droplet activation, droplet growth kinetics and hygroscopicity of the BSOA, ASOA, and ABSOA formed from monoterpenes and aromatics used as representative BVOC and AVOC.

We found that BSOA, ASOA and ABSOA had similar CCN activity although ASOA had a higher O/C. Adding BVOC after ASOA formation, or adding AVOC after BSOA formation did not significantly change the CCN activity of SOA. The similar CCN activity of BSOA, ASOA and ABSOA suggests that ASOA components and the interaction of ASOA with BSOA did not significantly change the CCN activity of SOA. This was likely due to compensating effects of potential oligomerization (reducing effect) and photochemical aging (enhancing effect) on CCN activity.

$\kappa_{\text{CCN}}$  generally increased slightly with photochemical aging, using OH dose as a proxy of photochemical aging and increased with O/C of aerosol for individual reaction systems. But when taking all the SOA types into account,  $\kappa_{\text{CCN}}$  did not correlate with O/C.

Analysis of the droplet growth kinetics shows that the droplet sizes from BSOA, ASOA, and ABSOA in supersaturated conditions were similar to those obtained with ammonium sulfate, indicating that none of these SOA has a kinetic barrier for water uptake. The fast water uptake of ASOA indicates that ASOA formed by aromatic precursors is not responsible for the droplet growth delay found in field studies (Shantz et al., 2010, 2012). This finding also suggests that potentially limited mixing between BSOA and ASOA reported in the literature does not hinder the water uptake in supersaturated conditions.

In contrast to CCN activity, the hygroscopicity of ASOA was distinctively higher than that of BSOA. The higher hygroscopicity was related to the higher O/C of ASOA com-

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pared to BSOA. Therefore, the ASOA component in ABSOA enhanced the hygroscopicity of aerosol and the extent depended on the ASOA fraction. Adding BVOC to ASOA or AVOC to BSOA changed the aerosol hygroscopicity, which was consistent with co-occurring changes in the ASOA fraction and the degree of oxidation of the aerosol represented by  $f_{44}$ . However, the hygroscopicity of ABSOA cannot be described by a linear combination of pure BSOA and ASOA systems. This indicates that additional processes such as oligomerization suppressed the hygroscopicity, which is in agreement with the interpretations for CCN activity.

Comparing hygroscopicity parameter  $\kappa$  obtained from CCN and HTDMA measurement shows a discrepancy between  $\kappa$  from the two methods.  $\kappa_{\text{HTDMA}}$  was significantly lower than  $\kappa_{\text{CCN}}$  for all SOA types studied here, by 30–70%. This discrepancy could not be resolved but can be attributed to the presence of slightly soluble materials, possible surface tension effect, or non-ideality of solutions and different behaviors of aerosol in the instruments. Better closure between HTDMA and CCN was found for ASOA than BSOA and ABSOA. ASOA seemed to have higher solubility in the subsaturated condition and/or lower surface tension at the activation point compared to BSOA.

This study has important implications for assessing the impact of SOA formed by the interaction of biogenic VOC with anthropogenic VOC emissions on the radiative forcing and climate. Since the interaction of AVOC with BVOC reduces the volatility (Emanuelsson et al., 2013), it prolongs particle persistence, which further enhances the particle concentration. Yet, based on this study, the CCN activity is not significantly affected. Therefore, models to assess the climatic effects of SOA formed through the interaction of biogenic VOC with anthropogenic VOC emissions could use single series of hygroscopicity parameter  $\kappa_{\text{CCN}}$  for BSOA, ASOA, and ABSOA to predict CCN concentration. However, significant mixing of ASOA and BSOA components can change the hygroscopic growth factor of the particles, which further affects the optical properties of SOA.

Comparing emission rates of aromatic compounds and isoprenoids (Lamarque et al., 2010; Guenther et al., 2012) and considering the turnover rates with OH and O<sub>3</sub> sug-

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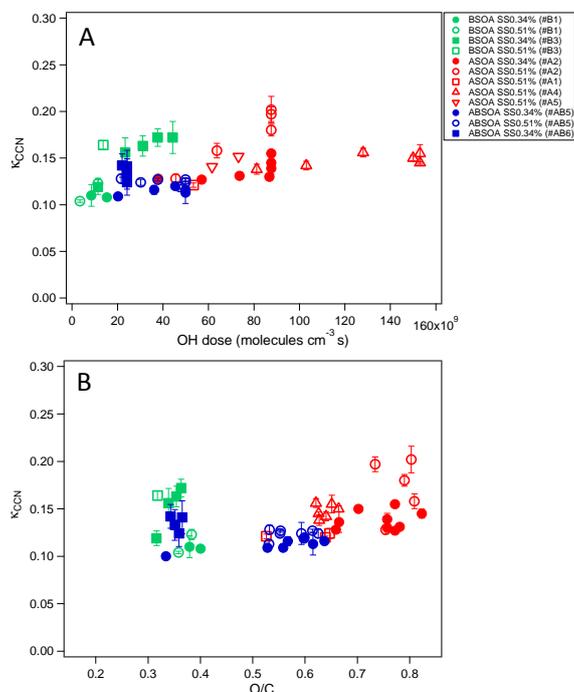
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**Table 1.** Summary of the experiments in this study.

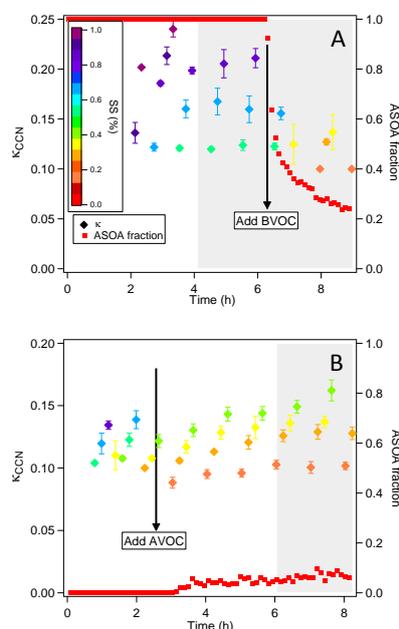
Experiment #	Experiment type	Experiment description	BVOC (ppb)	AVOC (ppb)	Ozone added (ppb)
B1	BSOA	BVOC only	$\alpha$ -pinene, limonene (20, 20)	N.A	50
B2	BSOA	BVOC only	$\alpha$ -pinene (40)	N.A	160
B3	BSOA	BVOC only	$\alpha$ -pinene, limonene (48, 48)	N.A	200
A1	ASOA	AVOC only	N.A	toluene (85)	0
A2	ASOA	AVOC only	N.A	toluene (85)	0
A3	ASOA	AVOC only	N.A	p-xylene (30)	0
A4	ASOA	AVOC only	N.A	p-xylene (30)	0
A5	ASOA	AVOC only	N.A	benzene (280)	0
A6	ASOA	AVOC only	N.A	benzene (280)	0
A7	ASOA	AVOC only	N.A	p-xylene-d <sub>10</sub> (200)	200
AB1	ABSOA	AVOC added 6.3h before BVOC	$\alpha$ -pinene, limonene (20, 20)	p-xylene (30)	0
AB2	ABSOA	AVOC added 6.3h before BVOC	$\alpha$ -pinene, limonene (20, 20)	toluene (85)	0
AB3	ABSOA	BVOC added 2.6h before AVOC	$\alpha$ -pinene, limonene (20, 20)	toluene (85)	50
AB4	ABSOA	BVOC added 5h before AVOC	$\alpha$ -pinene, limonene (39, 39)	p-xylene-d <sub>10</sub> (51)	200
AB5	ABSOA	AVOC and BVOC added together	$\alpha$ -pinene, limonene (4, 4)	toluene (85)	60
AB6	ABSOA	AVOC and BVOC added together	$\alpha$ -pinene, limonene (42, 42)	p-xylene-d <sub>10</sub> (90)	200

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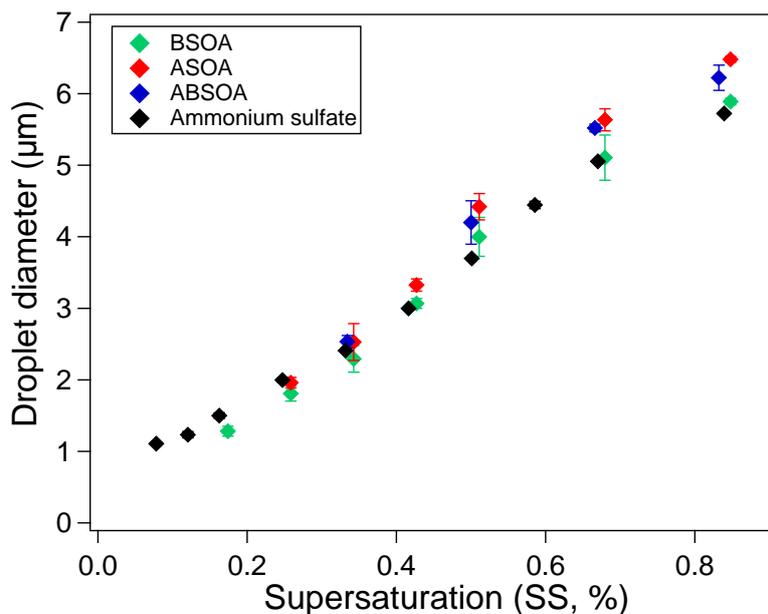
**Figure 1.** CCN activity of BSOA, ASOA, and ABSOA as a function of OH dose (a) and O/C of aerosol (b).  $\kappa_{CCN}$  at two supersaturations (SS) 0.34% (solid markers) and 0.51% (open markers) is shown. The green, red and blue colors indicate BSOA, ASOA and ABSOA, respectively.  $\kappa_{CCN}$  at similar SS is compared since  $\kappa_{CCN}$  depended on SS. BSOA, ASOA, and ABSOA showed no significant difference at a given SS. Although ASOA had higher degree of oxidation (O/C) than BSOA, ASOA components did not enhance  $\kappa_{CCN}$ .

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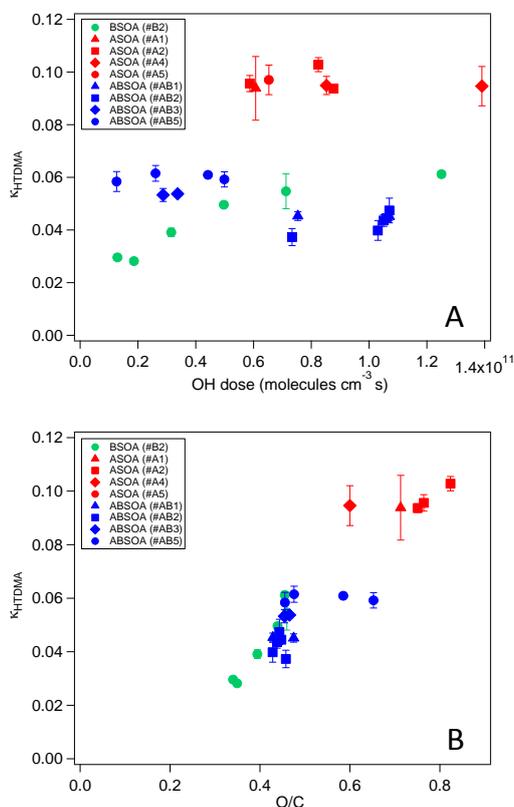
**Figure 2.** CCN activity of ABSOA from sequential VOC addition at various supersaturations (SS). (a) Toluene was added first and monoterpenes were added 6 h afterwards as indicated by the vertical arrow. The shaded areas show dark periods. ASOA fraction (red dots, right axis) decreased remarkably upon BVOC addition but  $\kappa_{CCN}$  showed no significant change. (b) Monoterpenes were added first and toluene was added 2.6 h afterwards as indicated by the vertical arrow.  $\kappa_{CCN}$  showed no significant change upon the addition of AVOC to BSOA. Note that time series of  $\kappa_{CCN}$  should be followed at similar SS because  $\kappa_{CCN}$  depended on SS and the SS range may change due to the particle size change.

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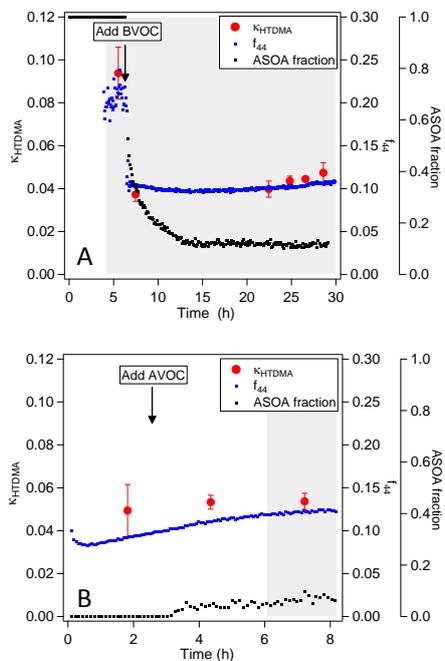
**Figure 3.** Droplet sizes of BSOA, ASOA, ABSOA and ammonium sulfate aerosols at various supersaturations (SS). In the CCNC, all SOA particles reached comparable droplet sizes compared to ammonium sulfate.

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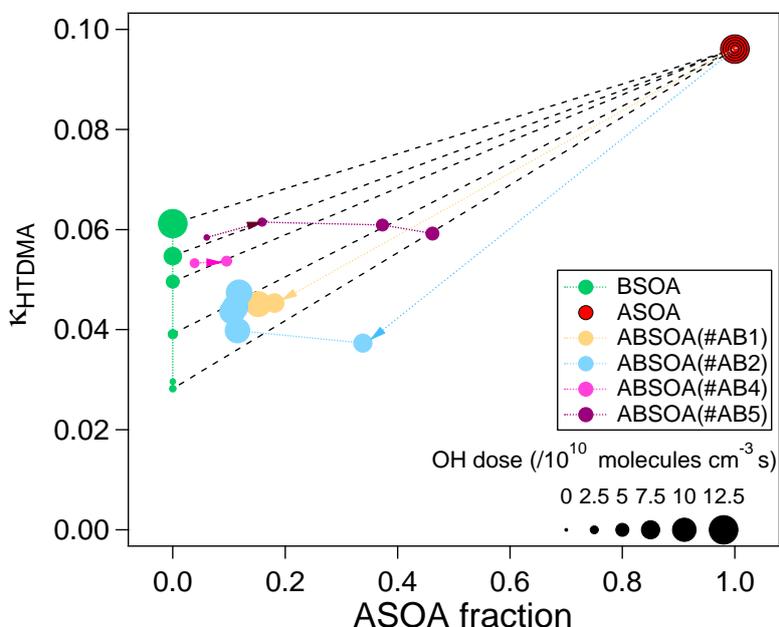
**Figure 4.**  $\kappa_{\text{HTDMA}}$  for BSOA, ASOA and ABSOA as a function of OH dose (a) and O/C of aerosol (b). ASOA had a distinctively higher  $\kappa_{\text{HTDMA}}$  and O/C than BSOA.

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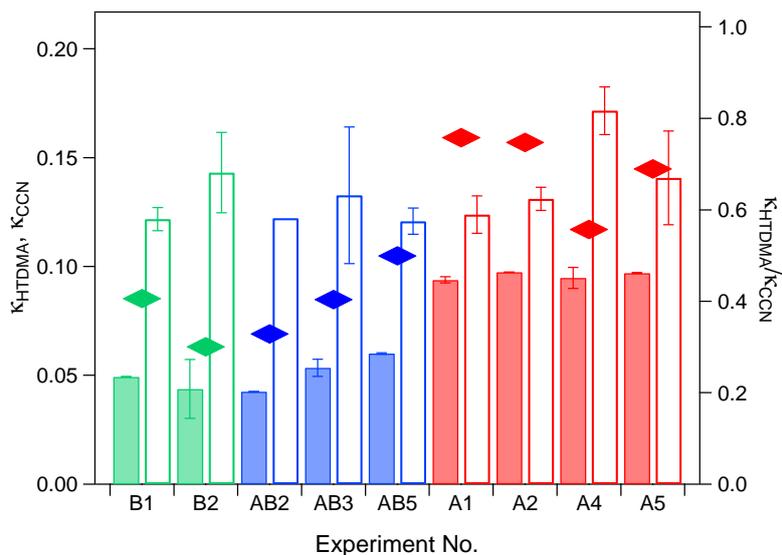
**Figure 5.**  $\kappa_{\text{HTDMA}}$  of ABSOA from experiments when AVOC and BVOC were added sequentially (same experiments as in Fig. 2). **(a)** Toluene was added first and the monoterpenes were added 6 h afterwards as indicated by the vertical arrow. The shaded areas indicate the dark periods.  $\kappa_{\text{HTDMA}}$  decreased significantly with the decrease of ASOA fraction due to the addition of BVOC. At the same time,  $f_{44}$  decreased distinctively. **(b)** Monoterpenes were added first and toluene was added 2.6 h afterwards as indicated by the vertical arrow.  $\kappa_{\text{HTDMA}}$  increased slightly as only a minor fraction (< 10 %) of ASOA was formed. At the same time,  $f_{44}$  also increased slightly.

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**Figure 6.**  $\kappa_{\text{HTDMA}}$  of ABSOA as a function of ASOA fraction. The size of the marker denotes the OH dose. For ASOA,  $\kappa_{\text{HTDMA}}$  is largely independent of OH dose and the average value was used, shown as the same point for different OH dose. The dashed lines connecting the pure BSOA and ASOA of the same OH doses stand for the linear combination of the  $\kappa_{\text{HTDMA}}$  of pure BSOA and pure ASOA at that given OH dose. A data point beyond the line corresponding to the OH dose of that point indicates a non-linear effect. The cases of non-linear effect can be found for ABSOA #AB1, #AB2 and #AB5. The dotted lines with arrow of ABSOA show the time order of data points.

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**Figure 7.** Comparison of  $\kappa_{\text{HTDMA}}$  and  $\kappa_{\text{CCN}}$  from HTDMA and CCN measurements for BSOA, ASOA, and ABSOA. The solid bar denotes the  $\kappa_{\text{HTDMA}}$  and the open bar denotes  $\kappa_{\text{CCN}}$  (left axis). The diamond denotes the ratio of  $\kappa_{\text{HTDMA}}$  to  $\kappa_{\text{CCN}}$  (right axis). The green, blue and red color indicates the BSOA, ABSOA and ASOA, respectively. For all SOA here, there was a discrepancy between  $\kappa_{\text{HTDMA}}$  and  $\kappa_{\text{CCN}}$ . For ASOA, the gap between  $\kappa_{\text{HTDMA}}$  and  $\kappa_{\text{CCN}}$  (low  $\kappa_{\text{HTDMA}}/\kappa_{\text{CCN}}$ ) was smaller compared to BSOA.