

Dear Editor,

Many thanks for forwarding the review report to us.

We have carefully read and replied to the reviewer's each comments. Particularly, we have addressed the reviewers' concern about the additivity of our synthetic inversions by including in Appendix B a mathematical derivation that demonstrates our approach is valid. We have also included an Appendix C that includes additional information about the bias correction that we use in our joint data assimilation, which was also requested by the reviewer.

Below are our detailed responses to each of the reviewer comments (marked in italics).

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*1. Going more into details: The major comment of reviewer one was essentially ignored, and no further information or changes were implemented.*

With respect, we disagree. We were asked to explain or comment why our results were different to those from Reuter et al. As described in our original reply, we have included additional quasi-regional inversions in Appendix A that represent a substantial amount of work. With the additional calculations we showed that even for regional flux inversions it is non-trivial to develop an on-line bias correction scheme to remove adverse effects from the erroneous description of boundary conditions without the risk of compromising the result regional flux estimates.

Admittedly, we are not using the same inversion system as Reuter et al (clearly acknowledged in the paper) and hence we are unable to conclude whether the inversions suffer from the same problems but it does highlight that valid issues remain in the interpretation of data over a geographical isolation with respect to the rest of the world. To address this, we have already added text in conclusion to stress:

'...we need to better quantify observation and model errors, and need to better understand the sensitivity of each inversion system to the assimilated observations as well as to their possible biases.'

*2. Reviewer 2 suggested changing not only the title, but also the abstract. While the title was changed, the text of the abstract was not significantly altered. This reviewer's comments about the non-additive nature of the flux adjustments is well taken, and not really addressed. Just because the adjustments appear to be additive in this case (at least at the scales considered) does not mean that the claim is generally true, or that the conclusions are valid. Rather than discussing this, they simply repeat what they did in the paper (rather confusingly, as the description of scenario 2 is erroneous, and is the same as scenario 3, which isn't true). This reviewer also requested a few specific plots, in order to assess the results.*

As stated in our original reply, we explained why reviewer 2 had unfortunately misunderstood the experiments we reported. We are surprised that this question has been raised again, as the additivity is obvious for our value-replacing testing with the same observation coverage and the same observation uncertainties. Below, we explain this in more detail for this reviewer.

We designed numerical experiments to focus on the sensitivity of the inferred European uptake to possible systematic errors associated with the values of current GOSAT XCO<sub>2</sub> retrievals by using the

45 same observation coverage and observation uncertainties. In these experiments, we replaced all or  
part of the GOSAT  $X_{\text{CO}_2}$  data with model values sampled at the time and location of the GOSAT data  
(taken from the reference INV\_TCCON). These model  $X_{\text{CO}_2}$  values are assigned with the same  
uncertainties associated with the GOSAT values that they replace. As a result, our inversions such as  
INV\_ACOS\_MOD\_ONLYEU and INV\_ACOS\_MOD\_NOEU share the same gain matrix as the GOSAT  
inversion INV\_ACOS. So the additivity of their individual contributions to the increased European  
uptake is obvious. In Appendix B, we show this mathematically that the consistency shown by our  
numerical results, despite approximations in solving posterior fluxes using the EnKF approach, is  
50 generally true.

3. *While I can understand that the authors may not to choose to include the figures in their  
manuscript, they should have at least included the figures in the response to the reviewer, as  
requested. Not doing so really suggests that they are ignoring the reviewer's comments in general. In  
55 a similar vein, this reviewer also asks that the explicit equation of the bias correction be provided,  
and the authors choose not to accede to this simple request.*

An egregious oversight for which we apologise profusely. As described previously, we use the on-  
line ‘bias correction’ as a systematic approach to infer the differences between the model and  
60 GOSAT retrievals at regional and monthly scales in order to show the effects of the small (sub-ppm)  
and varying sub-regional systematic differences over Europe on inferred European uptake estimates  
of  $\text{CO}_2$ .

As acknowledged in the manuscript Appendix A, we are acutely aware of the limits of our on-line  
bias correction, and have consequently treated our inferred fluxes from the joint data assimilation  
65 with caution. With this paper we stress the current approach as a method of inferring systematic  
differences at regional and monthly scales with no intention at this stage to suggest it as the final  
solution towards consistent flux inversions. We do not want to place emphasis on such a simple bias  
correction scheme but if the reviewer thinks the equation is important, we are happy to include in  
the revised manuscript (Appendix C):

$$70 \quad y^c = y - \text{bias}(m, i), \text{ (C1)}$$

where  $y$  represents GOSAT retrievals before the (extra) bias correction, and  $y^c$  is the bias-corrected  
 $X_{\text{CO}_2}$  data that we assimilate in our joint data assimilation experiments. For simplicity, we have  
assumed the regional (sub-regional) bias,  $\text{bias}(m, i)$ , is a function only of month ( $m$ ) and  
geographical region ( $i$ ).

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4. *Similarly, they replied to all of my comments with argumentation rather than accepting any  
changes (excluding typos and the title change), and chose to ignore the reasonable suggestion (also  
suggested by another reviewer in a previous round) of excluding TCCON from the reference inversion.*

80 We acknowledge that the reviewer provided many good suggestions, including, for example, the  
study focused on low-latitude regions. As discussed in our replies, limited by available observations  
we cannot add these additional studies in our manuscript, which represent a substantial body of  
work. For example, we are working on assimilating GOSAT and OCO-2 glint observations with dense  
coverage over tropical regions but as we are sure, the reviewer will appreciate that these data are

85 not properly validated and significant efforts are still required before they produce scientifically useful results.

Regarding the  $X_{CO_2}$  data collected by TCCON network: we appreciate the strong feelings of the reviewer for using these data as an independent validation. However, as we described in our previous reply, TCCON have been used in deriving bias correction for GOSAT XCO<sub>2</sub> retrievals. Also, 90 the TCCON column measurements are sensitive to CO<sub>2</sub> air masses transported from many different geographical regions. So that better (or worse) agreement with TCCON does not necessarily indicate a better (or worse) flux estimate at the regional scale (such as Europe). As other groups have found, the large amount of TCCON observations (as well as their sensitivity to long-range transport of air masses) are a useful asset for us to improve in-situ inversions, which is consistent with our aim to 95 build a reliable reference inversion for this study. This point has been stressed again in the revised manuscript (Section 2).

*5. Furthermore, I requested that they provide a map of the sub-regions that they used for their bias correction, along with the diagnosed errors. They acknowledged that this was a good suggestion, but that they'd like to present them in a future study based on the latest GOSAT XCO<sub>2</sub> retrievals. It is as 100 though they are trying to squeeze another publication out of the same approach rather than adapting the current manuscript to make it fit for publication. This is not an acceptable response, in my opinion.*

As detailed in our previous reply and above, we used a simple approach to estimate regional and sub-regional systematic differences between the model and GOSAT XCO<sub>2</sub> retrievals. We are fully 105 aware of the challenge to develop a robust on-line bias correction scheme. We have been experimenting with more elaborated bias-correction approaches, but have no intention to publish another paper just based on the simple scheme presented in this manuscript; this has never been the philosophy followed by the University of Edinburgh group.

We did not include such a map in our last revision because we believe such details are not critical for 110 the main discussion. But as one example, we have now included one figure in Appendix C to show the derived regional bias for March 2010 over 10 TransCom regions and 2 European (West and North) sub-regions. Although like other inter-comparisons between model and  $X_{CO_2}$  retrievals, the inferred systematic differences show interesting spatial structure, we are unable to validate these biases with independent observations. So our main goal is just limited to showing that the derived 115 small and varying biases over East and West Europe can affect the estimated European uptake.

*6. With respect to the title, if the paper is finally accepted for publication, the title needs to be changed again. While it is true that "outwith" is found in the dictionary, its use is restricted almost 120 exclusively to Scotland, and is unintelligible to many native English speakers, let alone international scientists. This could easily be substituted with "outside of".*

We follow the reviewer's suggestion, and change the title as its suggestion.



125 **Estimates of European uptake of CO<sub>2</sub> inferred from GOSAT X<sub>CO<sub>2</sub></sub> retrievals: sensitivity to measurement bias inside and outside Europe**

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140 Estimates of the natural CO<sub>2</sub> flux over Europe inferred from in situ measurements of atmospheric CO<sub>2</sub> mole fraction have been used previously to check top-down flux estimates inferred from space-borne dry-air CO<sub>2</sub> column (X<sub>CO<sub>2</sub></sub>) retrievals. Several recent studies have shown that CO<sub>2</sub> fluxes inferred from X<sub>CO<sub>2</sub></sub> data from the Japanese Greenhouse gases Observing SATellite (GOSAT) and the Scanning Imaging Absorption Spectrometer for Atmospheric CHartography (SCIAMACHY) have larger  
145 seasonal amplitudes and a more negative annual net CO<sub>2</sub> balance than those inferred from the in situ data. The cause of this elevated European uptake of CO<sub>2</sub> is still unclear, but some recent studies have suggested that this is a genuine scientific phenomenon. Here, we put forward an alternative hypothesis and show that realistic levels of bias in GOSAT data can result in an erroneous estimate of elevated uptake over Europe. We use a global flux inversion system to examine the  
150 relationship between measurement biases and estimates of CO<sub>2</sub> uptake from Europe. We establish a reference in situ inversion that uses an Ensemble Kalman Filter (EnKF) to assimilate conventional surface mole fraction observations and X<sub>CO<sub>2</sub></sub> retrievals from the surface-based Total Carbon Column Observing Network (TCCON). We use the same EnKF system to assimilate two independent versions of GOSAT X<sub>CO<sub>2</sub></sub> data. We find that the GOSAT-inferred European terrestrial biosphere uptake peaks  
155 during the summer, similar to the reference inversion, but the net annual flux is 1.40±0.19 GtC/a compared to a value of 0.58±0.14 GtC/a for our control inversion that uses only in situ data. To reconcile these two estimates, we perform a series of numerical experiments that assimilate observations with added biases or assimilate synthetic observations for which part or all of the GOSAT X<sub>CO<sub>2</sub></sub> data are replaced with model data. We find that for our global flux inversions, a large  
160 portion (60-90%) of the elevated European uptake inferred from GOSAT data in 2010 is due to retrievals outside the immediate European region, while the remainder can largely be explained by a sub-ppm retrieval bias over Europe. We use a data assimilation approach to estimate monthly

GOSAT  $X_{\text{CO}_2}$  biases from the joint assimilation of in situ observations and GOSAT  $X_{\text{CO}_2}$  retrievals. The inferred biases represent an estimate of systematic differences between GOSAT  $X_{\text{CO}_2}$  retrievals and the inversion system at regional or sub-regional scales. We find that a monthly varying bias of up to 0.5 ppm can explain an overestimate of the annual sink of up to 0.20 GtC/a. Our results highlight the sensitivity of  $\text{CO}_2$  flux estimates to regional observation biases, which have not been fully characterized by the current observation network. Without further dedicated measurements we cannot prove or disprove that European ecosystems are taking up a larger-than-expected amount of  $\text{CO}_2$ . More robust inversion systems are also needed to infer consistent fluxes from multiple observation types.

## 1. Introduction

Observed atmospheric variations of carbon dioxide ( $\text{CO}_2$ ) are due to atmospheric transport and surface flux processes. Using prior knowledge of the spatial and temporal distribution of these fluxes and atmospheric transport it is possible to infer (or invert for) the a posteriori estimate of surface fluxes from atmospheric concentration data. The geographical scarcity of such observations precludes robust flux estimates for some regions due to large uncertainties associated with meteorology and a priori fluxes. Arguably, our knowledge of top-down estimates of regional  $\text{CO}_2$  fluxes, particularly at tropical and high northern latitudes, have not significantly improved for over a decade [Gurney et al., 2002; Peylin et al., 2013], reflecting the difficulty of maintaining a surface measurement programme over vulnerable and inhospitable ecosystems. Atmospheric transport model errors compound errors introduced by poor observation coverage, resulting in significant differences between flux estimates on spatial scales  $< O(10,000 \text{ km})$  [e.g. Law et al., 2003; Yuen et al., 2005; Stephens et al., 2007]

The Greenhouse gases Observing SATellite (GOSAT), a space-borne mission launched in a sun-synchronous orbit in early 2009, was purposefully designed to measure  $\text{CO}_2$  columns using short-wave IR wavelengths. Validation of current  $X_{\text{CO}_2}$  column retrievals using co-located upward-looking FTS measurements of the Total Carbon Column Observing Network (TCCON) [Wunch et al., 2011] show a standard deviation of 1.6-2.0 ppm (e.g., Parker et al., 2012). Their global biases are typically smaller than 0.5 ppm [Oshchepkov et al., 2013]. The disadvantage of using the TCCON is that sites are mainly at northern extra-tropical latitudes with little or no coverage where our knowledge of the carbon cycle is weakest. Many surface flux estimation algorithms are particularly sensitive to systematic errors so that sub-ppm biases can still significantly change the patterns of regional flux estimates [Chevallier et al., 2010]. This is further complicated by the seasonal coverage of GOSAT data at high latitudes during winter months when solar zenith angles are too large to retrieve reliable values for  $X_{\text{CO}_2}$  [Liu et al., 2014].

Several independent studies have shown that regional flux distributions inferred from GOSAT  $X_{\text{CO}_2}$  retrievals are significantly different from those inferred from in situ data [Basu et al., 2013; Deng et al., 2013; Chevallier et al., 2014]. In particular, these studies report a larger-than-expected annual net emission over tropical continents and a larger-than-expected net annual uptake over Europe. While the GOSAT inversions suffer from larger observation errors, atmospheric transport errors and issues from the seasonal coverage of higher latitudes, the in-situ inversions are also unreliable over many regions due to poor coverage and atmospheric transport errors. Inter-comparisons revealed significant inconsistency in regional flux estimates inferred from in-situ observations by using

205 different inversion systems, over many regions important for global carbon cycle, including Europe [Peylin et al., 2013]. Consequently, there is an ongoing debate about whether a recent study that shows a large European uptake of CO<sub>2</sub> [Reuter et al., 2014] reflects a real phenomenon or is an artefact due to deficiencies both in the observations and in the inverse modelling.

210 We report the results from a small set of experiments that show systematic bias can introduce a large difference between European fluxes inferred from GOSAT and those inferred from in-situ data by using a global flux inversion approach. In the next section we provide an overview of the inverse model framework used to interpret data from the in-situ observation network (including both the conventional surface observation network and the relatively new TCCON network), and from the space-based GOSAT X<sub>CO<sub>2</sub></sub> data. In section 3, we present results from two groups of global inversion  
215 experiments that characterize the role of systematic bias in regional flux estimates. Further experiments for quasi-regional flux inversions are presented in Appendix A. In section 4, we use a modified version of the inverse model framework to estimate monthly biases by jointly assimilating all data. We conclude the paper in section 5.

## 2. Description and Evaluation of Control In-situ and GOSAT Experiments

220 We use the GEOS-Chem global chemistry transport model to relate surface fluxes to the observed variations of atmospheric CO<sub>2</sub> concentrations [Feng et al., 2009] at a horizontal resolution of 4°x5°, driven by GEOS-5 meteorological analyses from the Global Modeling and Assimilation Office Global Circulation Model based at NASA Goddard Space Flight Centre. We use an Ensemble Kalman Filter (EnKF) [Feng et al., 2009; 2011] to estimate regional fluxes from insitu or GOSAT observations for  
225 three years from 2009-2011, but we focus on 2010 to minimize error due to spin-up and edge effects. We estimate monthly fluxes on a spatial distribution that is based on TransCom-3 [Gurney et al., 2002] with each continental region further divided equally into 12 sub-regions and each ocean region further divided equally into 6 sub-regions. As a result, we estimate fluxes for 199 regions, compared to 144 regions we have used in previous studies [Feng et al, 2009; Chevallier et al., 2014].

230 In all global inversion experiments we assume the same set of a priori flux inventories, including: (1) monthly fossil fuel emissions [Oda and Maksyutov, 2011]; (2) weekly biomass burning emissions (GFED v3.0) [van der Werf et al., 2010]; (3) monthly oceanic surface CO<sub>2</sub> fluxes [Takahashi et al., 2009]; and (4) 3-hourly terrestrial biosphere-atmosphere CO<sub>2</sub> exchange [Olsen and Randerson, 2004]. We assume that the a priori uncertainty for each land sub-region is proportional to a  
235 combination of the net biospheric emission (70%) at the current month, and its annual variation (30%). We also assume that the a priori errors are correlated with each other with a spatial correlation length of 800 km, and a temporal correlation of 1 month [Chevallier et al., 2014]. We then determine the coefficient for the assumed a priori uncertainty by scaling the aggregated annual uncertainty over all 133 land sub-regions to 1.9 GtC/a. In particular, the resulting annual a priori  
240 uncertainty for European region is about 0.52 GtC/a, with the monthly uncertainty varying from 2.0 GtC/a for the summer months to about 0.8 GtC/a for winter months, which is generally larger than the a priori monthly uncertainty used by Deng et al. (2014). Prior uncertainties over oceans are determined under similar assumption but with a longer spatial correlation (1500 km), and a smaller aggregated annual error (0.6 Gt/a). Our experiments show that doubling the a priori uncertainty  
245 increases the European uptake inferred from GOSAT data by about 0.21 GtC/a (from 1.40 GtC/a to

1.61 GtC/a), compared to a smaller increase of 0.09 GtC/a for the in-situ inversion (from 0.58 GtC/a to 0.67 GtC/a).

Our control inversion experiment (INV\_TCCON, Table 1 and Figure 1) assimilates in-situ observations, including the conventional surface observations at 76 sites [Feng et al., 2011] and, in particular, the total column  $X_{CO_2}$  retrievals from all the TCCON sites of the GGG2014 dataset (see Wennberg et al., 2014, and <https://tcccon-wiki.caltech.edu> for more details) to improve observation constraints. In some studies, TCCON data was used to evaluate posterior fluxes. However TCCON data has been used to derive bias corrections for GOSAT  $X_{CO_2}$  retrievals [Cogan et al., 2012], and also the nature of total column measurements means that they are sensitive to air mass transported from other regions, which complicate the assessment of European flux estimates.

We use daytime (09:00 to 15:00 local time) mean TCCON retrievals, with the observation errors determined by the standard deviation about their daytime mean. To account for the inter-site biases as well as the model representation errors, we enlarge the TCCON observation errors by 0.5 ppm. Including TCCON observations increases the annual net uptake over Europe in 2010 from 0.49 GtC/a, as inferred from surface observations only, to 0.58 GtC/a. The increase is mainly due to a larger summer uptake. TCCON data also reduce the a posteriori uncertainty by about 15% from 0.16 GtC/a to 0.14 Gt/a. However considering the limited spatial resolution (only 12 sub regions for the whole TransCom European region), and unquantified model transport and representation errors, we anticipate that the complete a posteriori uncertainty is larger than the value estimated by the inversion system itself, as suggested by large inter-model variations found for in-situ inversions [e.g., Peylin et al., 2013].

For the two control GOSAT inversions (Figure 1), we use two independent data sets: (1)  $X_{CO_2}$  retrievals from JPL ACOS team (v3.3) [Osterman et al., 2013] (INV\_ACOS); and (2) the full-physics  $X_{CO_2}$  retrievals (v4.0) from the University of Leicester [Cogan et al., 2012] (INV\_UOL). For both data sets, we assimilate only the H-gain data over land regions, and apply the bias corrections recommended by the data providers. We double the reported observation errors, as suggested by the retrieval groups.

As a performance indicator for our ability to fit fluxes to observed  $X_{CO_2}$  concentrations, we compare a posteriori model concentrations with GOSAT  $X_{CO_2}$  retrievals and show that INV\_ACOS and INV\_UOL agree much better than INV\_TCCON. For example, the bias against ACOS  $X_{CO_2}$  retrievals is -0.45 ppm for INV\_TCCON and 0.02 ppm for INV\_ACOS with a corresponding reduction in the global standard deviation from 1.69 ppm to 1.57 ppm. However comparison of GOSAT a posteriori concentrations against independent HIPPO-3 measurements is worse than INV\_TCCON with a positive bias of 0.47 ppm and 0.66 ppm for INV\_ACOS and INV\_UOL, respectively, which are mainly caused by the overestimation of  $CO_2$  concentrations (~1.5-2.0 ppm) at low latitudes (Figure 2).

### 3. Results

Figure 1 and Table 1 shows the three inversion experiments, INV\_TCCON, INV\_ACOS, and INV\_UOL, have similar European uptake values in June 2010 (0.69 GtC/m for INV\_TCCON and ~0.72 GtC/m for GOSAT inversions), and are generally consistent with other GOSAT inversion experiments (e.g., Deng et al., 2014; Chevallier et al., 2014). But the GOSAT inversions have an annual net uptake of about  $1.40 \pm 0.19$  GtC/a compared to the in-situ inversion of  $0.58 \pm 0.14$  GtC/a. Figure 1 also shows



significant differences between their monthly flux estimates in early spring and winter when there is only sparse GOSAT observation coverage, particularly over northern Europe. Both INV\_UOL and INV\_ACOS have a cumulative total of about 0.51 GtC more uptake than INV\_TCCON during February-  
290 April of 2010, with a further 0.37 GtC uptake accumulated over the following summer and autumn. This larger uptake is partially cancelled out by larger emissions (0.17-0.08 GtC) at the end of 2010.

Figure 2 shows that INV\_TCCON a posteriori CO<sub>2</sub> mole fractions agree well with the independent HIAPER Pole-to-Pole Observations (HIPPO-3) aircraft measurements below 5 km over the Pacific Ocean in 2010 [Wofsy et al., 2010], with a small bias of 0.05 ppm, and a sub-ppm standard deviation  
295 of 0.87 ppm. Figure 3 shows further evaluation of a posteriori CO<sub>2</sub> mole fractions using descending and ascending profile observations over two European airports from the CONTRAIL experiment [Machida et al., 2008]. We calculate monthly mean CONTRAIL measurements during 2010 using data below 3 km, where there is greater sensitivity to local surface fluxes. Our current model resolution precludes small-scale sources (or sinks) so we expect model bias. We find that INV\_TCCON agrees  
300 best with CONTRAIL observations, in particular at the beginning of the 2010, partially reflecting the poor GOSAT X<sub>CO2</sub> coverage over Europe during the winter and early spring. However, we cannot conclude from the slightly degraded agreement with CONTRAIL (as well as with HIPPO-3) that the European uptake inferred from GOSAT data is incorrect, because unaccounted small local emissions/sinks, and model transport errors can affect the comparison against aircraft observations.

Figure 3 also presents an additional model simulation forced by a hybrid flux (denoted by the magenta broken line) where the INV\_TCCON a posteriori fluxes outside Europe are replaced by the results from INV\_ACOS. The resulting CO<sub>2</sub> concentrations from these hybrid fluxes are, as expected, higher than the a posteriori model concentrations for INV\_ACOS because of the larger European emissions (i.e., less uptake) inferred by INV\_TCCON. But they are also systematically higher than the  
305 INV\_TCCON simulation, in particular during spring months, despite the same European fluxes being used to force these two simulations. This suggests an overestimate of CO<sub>2</sub> transported into the European region by the GOSAT inversions. Further comparison of the INV\_TCCON simulation and the hybrid run reveals that systematic differences in the inflow into the European domain can affect the atmospheric X<sub>CO2</sub> gradient across this region. In the INV\_TCCON simulation, the mean X<sub>CO2</sub> difference  
310 between east (east of 20°E) and west (west of 20°E) Europe is ~0.04 ppm for May, 2010, which is increased to 0.16 ppm in the hybrid run (cf. E-W X<sub>CO2</sub> gradient of -0.20 ppm for GOSAT ACOS data).  
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To understand the differences between the INV\_TCCON and GOSAT inversions, we conducted two groups of sensitivity tests (Table 1 and Figure 4). First, we replaced all or part of the GOSAT X<sub>CO2</sub> retrievals assimilated in INV\_ACOS with those from a model simulation forced by the a posteriori  
320 fluxes from INV\_TCCON. In experiment INV\_ACOS\_MOD\_ALL (Figure 4), where we replace all GOSAT data with CO<sub>2</sub> concentrations inferred from INV\_TCCON, we reproduce INV\_TCCON with small exceptions at beginning of 2010, reflecting the seasonal variation in GOSAT coverage. In a related experiment INV\_ACOS\_MOD\_NOEU for which we only replace X<sub>CO2</sub> retrievals outside Europe with the model simulation, the differences between the GOSAT and in-situ inversions are  
325 significantly reduced, particularly over the period with limited observation coverage, although the actual X<sub>CO2</sub> retrievals are still assimilated over Europe. The simulated GOSAT data outside Europe reduces the estimate of European uptake from 1.40 GtC/a to 0.88 GtC/a. In other words, the GOSAT observations outside the European region are responsible for about 60% (0.52 GtC/a) of the total enhanced European sink (0.82 GtC/a) with the remainder (0.30 GtC/a) due to observations taken

330 directly over Europe. The large contribution from GOSAT retrievals outside Europe has also been  
confirmed by the high uptake (1.17 Gt/a) in a counterpart experiment (INV\_ACOS\_MOD\_ONLYEU)  
where only GOSAT retrievals within Europe are replaced by the model simulations. We show in  
Appendix B that theoretically the difference between INV\_ACOS and INV\_ACOS\_MOD\_ALL is equal  
335 to the sum of the individual uptake increases in the paired synthetic inversions of  
INV\_ACOS\_MOD\_NOEU and INV\_ACOS\_MOD\_ONLYEU.

For INV\_UOL, when we replace the  $X_{CO_2}$  data outside Europe by the a posteriori INV\_TCCON model  
simulations, European uptake is reduced to 0.67 GtC/a (INV\_UOL\_MOD\_NOEU, Table 1), indicating  
an external contribution of nearly 90% to the enhanced uptake of 0.82 GtC/a. Together with Figure  
3, these results suggest that GOSAT inversions result in an overestimated  $CO_2$  inflow. This will  
340 subsequently lead to the fitted European flux having to compensate, via mass balance, by being  
erroneously low even when un-biased GOSAT  $X_{CO_2}$  data are assimilated over the immediate  
European region. We find similar effects in the quasi-regional inversions (Figure A1), where only  
observations within European region are assimilated, with flux estimates from INV\_TCCON or from  
INV\_ACOS being used to provide lateral boundary conditions around Europe.

345 Second, we crudely demonstrate how regional bias could explain the remaining discrepancy of up to  
0.30 GtC/a between GOSAT and in situ inversions over Europe. In our experiment  
INV\_ACOS\_SPR\_0.5ppm, we add a bias of +0.5 ppm to the GOSAT ACOS retrievals within Europe  
taken in February-April, inclusively, which effectively reduces the uptake by 0.1 GtC/a from 1.40  
GtC/a to 1.30 GtC/a. Similarly, when the bias of +0.5 ppm is added to the GOSAT data taken in June-  
350 August we find a larger reduction of 0.15 GtC/a for the summer peak uptake  
(INV\_ACOS\_SUM\_0.5ppm), partially due to a larger a priori uncertainty and denser GOSAT coverage  
during the summer. These results emphasize the importance of characterizing sub-ppm regional  
bias to avoid erroneous flux estimates.

#### 4. Bias estimation.

355 Here we demonstrate a simple approach to quantify systematic bias in  $X_{CO_2}$  retrievals based on a  
simple on-line bias correction scheme. We assimilate the GOSAT  $X_{CO_2}$  retrievals together with the  
surface and TCCON observations in two experiments: INV\_ACOS\_INS and INV\_UOL\_INS (Table 1).  
We also include monthly GOSAT  $X_{CO_2}$  regional biases over 11 TransCom land regions [Gurney et al.,  
2002] as parameters to be inferred together with surface fluxes from the joint assimilation of in-situ  
360 and satellite observations. To investigate the spatial pattern of the  $X_{CO_2}$  biases within Europe, we  
split Europe into West Europe (west of 20°E) and East Europe (east of 20°E). We assume that a priori  
for monthly biases is  $0.0 \pm 0.5$  ppm. For simplicity, we have assumed the a priori errors for regional  
 $X_{CO_2}$  biases are not correlated. Compared to the off-line comparisons between GOSAT  $X_{CO_2}$  retrieval  
and model concentrations, the main advantage of the on-line bias estimation is that the  
365 uncertainties associated with error in flux estimates can be partially taken into account. However,  
biases derived by this approach reflect the systematic difference between the model simulation and  
GOSAT data over large (continental) regions, which also contain systematic model errors (such as the  
atmospheric transport and representation errors). In addition, the inversion results are affected by  
the relative weights assigned to different data sets, as well as by the relative prior uncertainty  
370 assumed for surface fluxes and for the observation bias. The seasonal variation of the mean  $CO_2$   
concentration is an important sign of the underlined biosphere seasonal cycle. We show in

Appendix A that when we inflate the a priori uncertainty for the assumed observation bias, the observation constraints on flux estimate will become weaker. Also, the on-line bias correction is only effective for detecting and correcting bias at specified patterns, which may increase the sensitivity to other uncharacterized systematic errors. Despite these weaknesses, a joint data assimilation approach can exploit complementary constraints from in situ and satellite  $X_{CO_2}$  data: for example there are few GOSAT observations over northern Europe during autumn and winter months, while Eastern Europe has few in-situ observations. We have also limited the a priori uncertainty for the monthly observation biases to 0.5 ppm. **Figure C1 (Appendix C) shows, for example, the inferred monthly mean bias for March 2010.**

In the joint inversions INV\_ACOS\_INS and INV\_UOL\_INS, the annual European uptake is estimated to be 0.62 GtC/a and 0.67 GtC/a, respectively (Table 1), which is close to the reference value of 0.58 GtC/a inferred from the in situ observations. To test the impact of the on-line bias correction, we set the a priori uncertainty of regional  $X_{CO_2}$  bias to be 0.01 ppm so that on-line bias correction is effectively turned off. As a result, the annual European uptake for INV\_ACOS\_INS is increased by 0.15 GtC to 0.77 GtC/a, which is close to INV\_ACOS\_MOD\_NOEU, but about 55% of the GOSAT only inversions (1.40 GtC/a).

Figure 5 shows the estimated monthly biases in ACOS and UOL  $X_{CO_2}$  retrievals over East and West Europe during 2010. Monthly biases are typically smaller than 0.5 ppm over the two regions, but have different seasonal cycles. Additional experiment shows that after ACOS  $X_{CO_2}$  data over Europe have been corrected for the inferred biases, the European annual uptake by INV\_ACOS is reduced by 0.20 GtC/a, representing more than half of the contribution from GOSAT observations within Europe. This result is consistent with our sensitivity tests. The effect of bias correction is much smaller for INV\_UOL (about 0.07 GtC/a), because of the different bias patterns. Differences in GOSAT  $X_{CO_2}$  retrievals and their effects on regional flux estimates have also been investigated in previous studies (e.g., Takagi et al., 2014).

## 5. Discussion and Conclusions

We used an ensemble Kalman Filter to infer regional  $CO_2$  fluxes from three different  $CO_2$  data sets: 1) surface in situ mole fraction observations and TCCON  $X_{CO_2}$  retrievals; 2) GOSAT  $X_{CO_2}$  retrievals from the JPL ACOS team; and 3) GOSAT  $X_{CO_2}$  retrievals from the University of Leicester. Our results, consistent with previous studies, show that these GOSAT data in a global flux inversion context result in a significantly larger European uptake than inferred from in situ data during 2010.

We showed using sensitivity experiments that a large portion (60-90%) of the elevated European uptake of  $CO_2$  is related to the systematically higher model  $CO_2$  mass being transported into Europe, due to the assimilation of GOSAT  $X_{CO_2}$  data outside the European region. We find some evidence using aircraft observations over the Pacific that GOSAT a posteriori fluxes result in higher  $CO_2$  concentration over lower latitudes. But limited observation coverage and unaccounted model errors prevent us from confidently concluding that GOSAT  $X_{CO_2}$  data are biased high or low. Our global and quasi-regional (Appendix A) flux inversion experiments show that the main consequence of the elevated  $CO_2$  inflow to the European domain is that the European uptake must increase because of mass balance, even when GOSAT  $X_{CO_2}$  retrievals within the European domain are not biased. A crude sensitivity test (INV\_ACOS\_OUT\_0.5ppm) shows that reducing ACOS  $X_{CO_2}$  data outside the European region by 0.5 ppm will reduce European annual uptake from 1.40 GtC/a to 0.98 GtC/a. Erroneous

415 interpretation of  $X_{CO_2}$  data can result from analyses if biased boundary conditions are not addressed. However, as shown in Appendix A, a gross mis-characterization and correction of bias may weaken observation constraints, which can also lead to erroneous flux estimates.

We also showed using sensitivity tests that sub-ppm bias can explain the remaining 0.30 GtC/a flux difference between the in situ inversion and INV\_ACOS after accounting for biased boundary conditions. By simultaneously assimilating the in situ and GOSAT observations to estimate surface fluxes and monthly  $X_{CO_2}$  biases, we infer a monthly observation bias that is typically less than 0.5 ppm over East and West Europe, but is able to cause an elevated sink of up to 0.20 GtC/a. The inferred monthly biases for UOL  $X_{CO_2}$  are also not the same as the ACOS  $X_{CO_2}$  data, particularly over West Europe during the summer months. This level of sensitivity of regional flux estimate to time-varying sub-ppm observation bias highlights the challenges we face as a community when evaluating  $X_{CO_2}$  retrievals using current observation networks.

Flux estimates are sensitive to a priori assumptions, idiosyncrasies of applied inversion algorithms, and the underlying model atmospheric transport [Chevallier et al., 2014; Peylin et al., 2014; Reuter et al., 2014]. The possible presence of regional observation biases further complicates the inter-comparisons of flux estimates based on different inversion approaches, as they may have different sensitivities to certain observation biases. In our assimilation of ACOS  $X_{CO_2}$  retrievals, we find that doubling the a priori flux error (INV\_ACOS\_DBL\_ERR) increases the estimated European uptake from 1.40 GtC/a to 1.61 GtC/a, consistent with the hypothesis on the increased vulnerability to the observation biases both within and outside Europe when using weak a priori constraints. In contrast, doubling the a priori flux errors only increases the uptake by 0.05 GtC/a to 0.67 GtC/a for the joint data assimilation (INV\_ACOS\_INS\_DBL\_ERR), with very little changes in the estimated biases (not shown). Examples in Appendix A also demonstrate different responses to regional and sub-regional biases before and after an on-line scheme is used to correct the systematic error across Europe. These differences emphasize the need for a closer examination of the responses of the inversion systems to the assimilated observations, as well as to their possible biases, to help understand the inter-model variations in estimated regional fluxes.

Complicated interactions between observations and the assimilation system also mean that our present study does not exclude other possible causes for the elevated European uptake reported by previous research from assimilation of GOSAT data. Instead, it highlights the adverse effects of possibly uncharacterized regional biases in current GOSAT  $X_{CO_2}$  retrievals that can attract erroneous interpretation of resulting regional flux estimates. A more thorough evaluation of the  $X_{CO_2}$  retrievals using independent and sufficiently accurate/precise observations is urgently required to increase the confidence of regional  $CO_2$  flux estimates inferred from space-based observations. Without additional observations, we cannot rule out either the lower European uptake estimate of around 0.6 GtC/a (inferred from the in situ inversion INV\_TCCON and the joint inversion INV\_ACOS\_INS and INV\_UOL\_INS) or the higher European uptake estimate of around 1.40 GtC/a (inferred from GOSAT data). There is also no sufficient reason to believe that the mean value among these diverse estimates is more reliable, because our study suggests that small systematic errors can result in significant differences in the estimated fluxes, and the influences of random errors have also not been fully quantified. The observational density required to infer flux estimates over a limited spatial domain such as Europe is crucial. For the time frame of this analysis, the TCCON network provided good coverage for Europe, North America, South East Asia and Australia/New Zealand. Great efforts

were also taken to reduce inter-station biases. In future the TCCON measurement network may be supported by smaller, more mobile FTIR instruments, which can be established, at least on a campaign basis, in tropical and high latitude locations where observational gaps are greatest.

460 **Our joint data assimilation approach assimilates in-situ and space-borne observations. It also provides estimates of systematic differences between XCO<sub>2</sub> retrievals and the inversion system at regional /sub-regional scales. However the resulting differences will include the observation biases and deficiencies in the underlying inversion approach.** To achieve consistent flux estimates inferred from assimilating multiple data sets using different inversion approaches, we need to better quantify observation and model errors, and need to better understand the sensitivity of each inversion system to the assimilated observations as well as to their possible biases. It is difficult to develop a robust bias correction scheme before properly characterizing observation biases and the responses by the inversion system.

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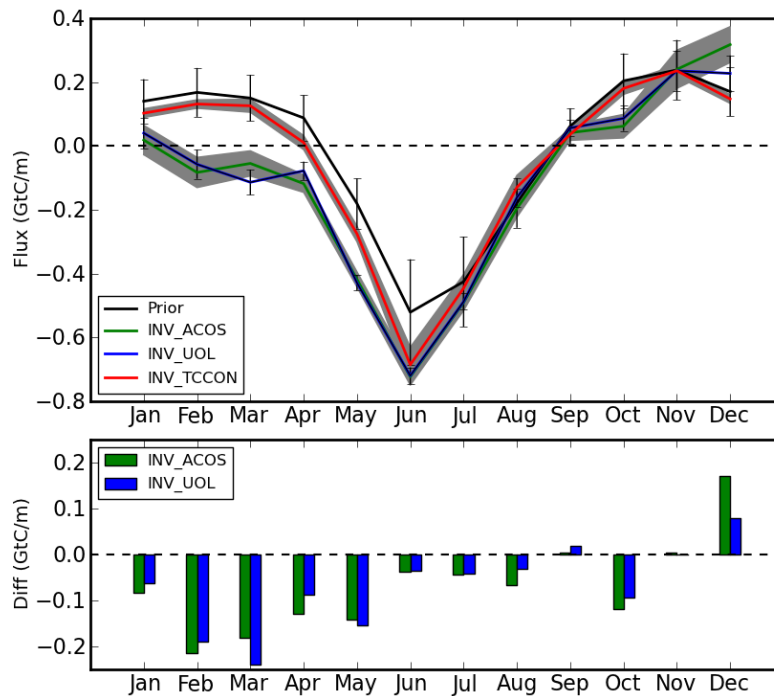
## Tables

Name	Data	Flux (GtC/a)	Uncertainty (GtC/a)
INV_TCCON	In-situ Flask and TCCON $X_{CO_2}$	-0.58	0.14
INV_ACOS	ACOS $X_{CO_2}$ retrievals	-1.40	0.19
INV_UOL	UOL $X_{CO_2}$ retrievals	-1.4	0.20
INV_ACOS_MOD_ALL	Model simulation of ACOS $X_{CO_2}$ by using INV_TCCON posterior fluxes	-0.64	0.19
INV_ACOS_MOD_NOEU	As INV_ACOS_MOD_ALL but the real ACOS $X_{CO_2}$ retrievals are assimilated within Europe.	-0.88	0.19
INV_UOL_MOD_NOEU	As INV_UOL, but outside the Europe, UOL $X_{CO_2}$ retrievals are replaced with INV_TCCON simulations.	-0.67	0.19
INV_ACOS_MOD_ONLYEU	As INV_ACOS, but $X_{CO_2}$ retrievals within EU are replaced by INV_TCCON simulations	-1.17	0.19
INV_ACOS_OUT_0.5ppm	As INV_ACOS, but a bias of -0.5 ppm has been added to $X_{CO_2}$ retrievals outside Europe.	-0.98	0.19
INV_ACOS_SPR_0.5ppm	As INV_ACOS, but 0.5 ppm bias has been added to the European data in February, March, and April.	-1.30	0.19
INV_ACOS_SUM_0.5ppm	As INV_ACOS, but 0.5 ppm bias has been added to the European data in June, July, and August.	-1.25	0.19
INV_ACOS_INS	ACOS $X_{CO_2}$ retrievals and In-situ flask and TCCON data	-0.62	0.13
INV_UOL_INS	UOL $X_{CO_2}$ retrievals and in-situ flask and TCCON data	-0.67	0.13
INV_ACOS_DBL_ERR	ACOS $X_{CO_2}$ retrievals, but the a priori uncertainties have been doubled	-1.61	0.27
INV_ACOS_INS_DBL_ERR	GOSAT ACOS $X_{CO_2}$ retrievals and In-situ flask and TCCON data but the a priori flux uncertainties have been doubled	-0.67	0.16

615 **Table 1:** The magnitude and uncertainty of the European annual CO<sub>2</sub> biosphere flux (GtC/a) from 14 global flux inversion experiments. Except INV\_ACOS\_INS\_DBL\_ERR and INV\_ACOS\_DBL\_ERR, the aggregated European annual uptake of the a priori fluxes is  $-0.1 \pm 0.52$  GtC/a.

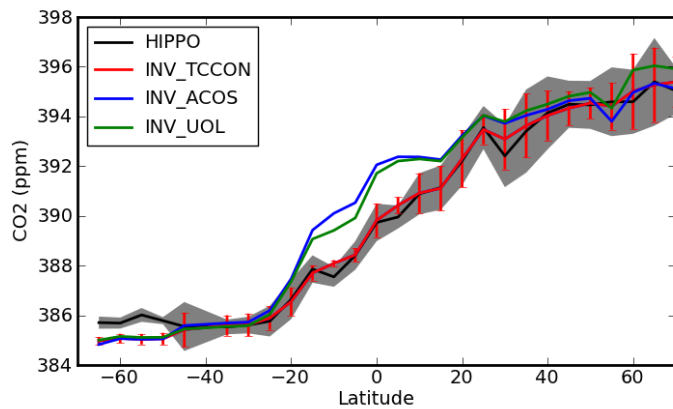
## Figures

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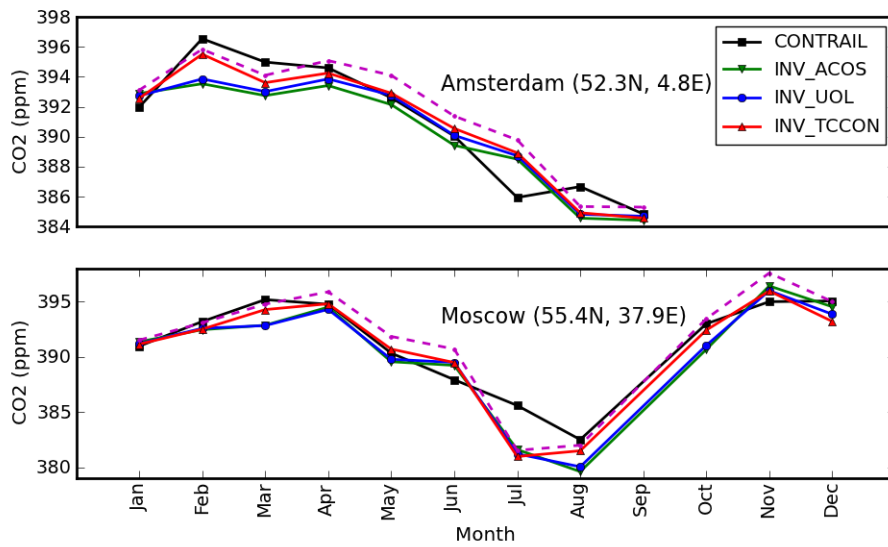


**Figure 1.** Monthly a posteriori estimates (GtC/m) for European biospheric CO<sub>2</sub> fluxes in 2010 using three inversion experiments (top panel): 1) INV\_TCCON (red line), 2) INV\_ACOS (green line), and INV\_UOL (blue line). The black line denotes a priori values. The vertical black lines and grey shading denotes the uncertainties of the corresponding a priori or a posteriori flux estimates, respectively. Differences in monthly CO<sub>2</sub> uptake (GtC/m) between INV\_TCCON and two GOSAT inversions (bottom panel): INV\_ACOS (green bars) and INV\_UOL (blue bars).

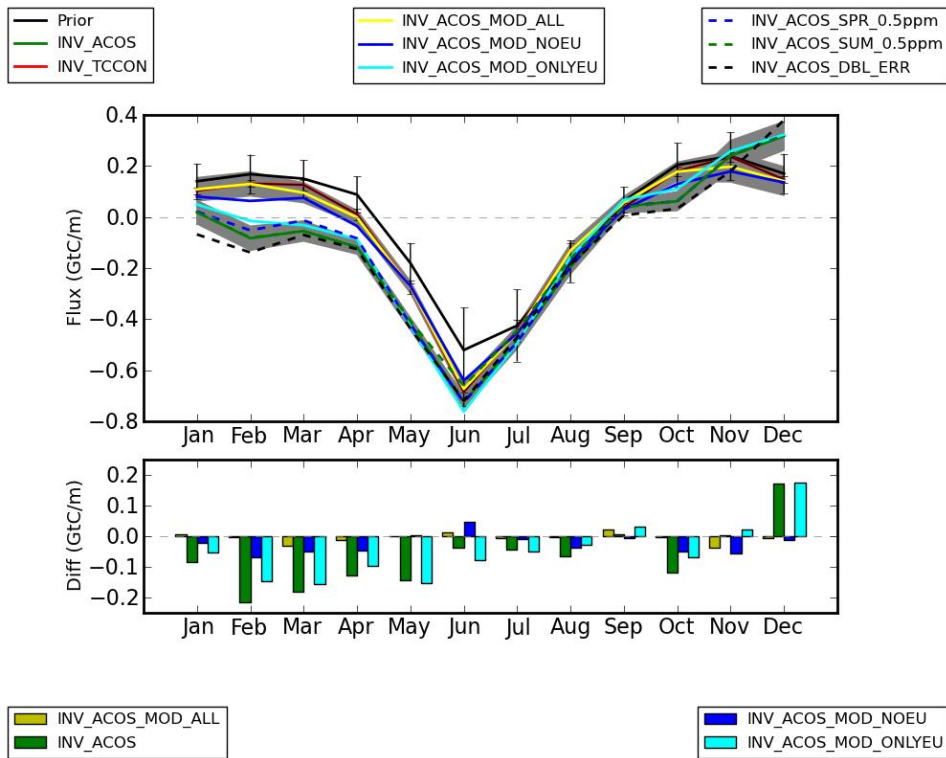
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**Figure 2:** HIPPO-3 and GEOS-Chem model atmospheric CO<sub>2</sub> mole fractions (ppm) over the Pacific Ocean below 5 km (black). GEOS-Chem is driven by different a posteriori flux estimates: 1) INV\_TCCON (red), 2) INV\_ACOS (blue), and 3) INV\_UOL (green). HIPPO-3 and model CO<sub>2</sub> mole fractions are binned into 5° latitude boxes. We calculate the mass-weighted average over these latitude boxes by assigning each HIPPO-3 and GEOS-Chem model value a weighting factor according to the observation altitude (air pressure). The grey envelope (red vertical lines) indicates the one standard deviation of HIPPO-3 measurements (INV\_TCCON model values) within each latitude box.



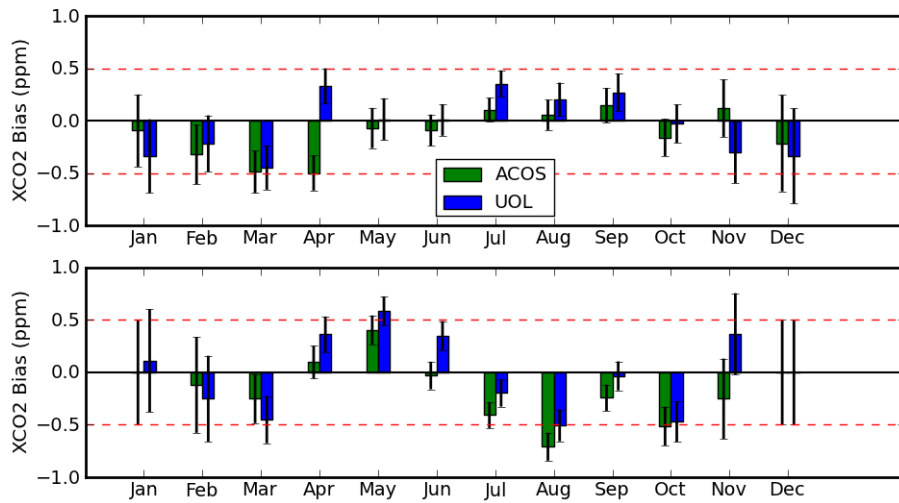
**Figure 3:** Monthly mean observed and model a posteriori model CO<sub>2</sub> mole fractions (ppm) below 3km above Amsterdam (the top panel) and Moscow (the bottom panel) airports during 2010, respectively [Machida et al., 2008]. The three sets of a posteriori model concentrations are inferred from three inversion experiments: INV\_TCCON (red line), INV\_ACOS (green line), and INV\_UOL (blue line). The broken magenta line represents a model simulation where the European fluxes from INV\_ACOS inversion are replaced by INV\_TCCON estimates.



655 **Figure 4:** Monthly European biospheric flux estimates (GtC/m) from two groups of sensitivity  
 experiments (top panel, Table 1). Black, green and red solid lines denote the a priori and the  
 INV\_ACOS and INV\_TCCON inversions, respectively. Differences between INV\_TCCON inversion and  
 sensitivity inversions (bottom panel): 1) INV\_ACOS\_MOD\_ALL (yellow), where all GOSAT retrievals  
 are replaced by the model simulations forced by INV\_TCCON a posteriori fluxes; 2) INV\_ACOS  
 660 (green), where original GOSAT ACOS retrievals are assimilated; 3) INV\_ACOS\_NOEU (blue) where all  
 the GOSAT retrievals outside the European region are replaced by the INV\_TCCON simulations; and  
 4) INV\_ACOS\_MOD\_ONLYEU (cyan) where only GOSAT retrievals within the European region are  
 replaced by the INV\_TCCON simulations.

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**Figure 5:** Estimates of monthly CO<sub>2</sub> biases (ppm) in GOSAT ACOS (green) and UOL (blue) X<sub>CO2</sub> retrievals over (top) West (West of 20°E) and (bottom) East (East of 20°E) Europe. The black vertical lines represent the uncertainty.

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## Appendix A: Quasi-regional flux inversion.

To further study the contributions from  $X_{\text{CO}_2}$  retrievals within and outwith Europe we have performed quasi-regional flux inversions to infer the European uptake of  $\text{CO}_2$  in 2010, based on the same EnKF approach as the global flux inversions. In contrast to the global experiments (Table 1), for the quasi-regional inversions we assimilate observations only over Europe, and assign a small a priori flux uncertainty to any region outside Europe in order to minimize the influence of observations taken over Europe on other regions. Consequently, a posteriori flux estimates outside of Europe are close to their a priori values. We use the a posteriori fluxes from INV\_TCCON as the a priori estimates for 12 sub-regions in Europe, and assume their uncertainty is two thirds of that we use for the global flux inversions. This is because the a posteriori estimates from INV\_TCCON have already been refined by in situ data.

To investigate the influence of lateral boundary conditions on the quasi-regional flux inversions, we use two different sets of a posteriori estimates to define fluxes outside Europe: 1) INV\_TCCON (INV\_BD\_TCCON) and 2) INV\_ACOS (INV\_BD\_ACOS). Figure A1 shows that INV\_BD\_ACOS has a higher annual uptake of 1.58 GtC/a than INV\_BD\_TCCON with an uptake of 0.79 GtC/a (Table A1), with differences larger during the first half of 2010. The estimate for INV\_BD\_ACOS is similar to its global inversion counterpart INV\_ACOS. Large differences between INV\_BD\_ACOS and INV\_BD\_TCCON highlight the importance of accurate lateral boundary conditions to a regional European inversion.

We use on-line bias correction schemes to reduce the adverse impacts from incorrect boundary conditions around Europe. Similar to Reuter et al. (2014), we estimate monthly observation biases across Europe using our quasi-regional flux inversion system. Here, we introduce a monthly bias to remove the systematic difference between model and GOSAT observations across the whole European region, and assume an associated a priori uncertainty of 100 pm [Reuter et al, 2014]. This is different from our previous bias assumption of 0.5 ppm over East and West Europe for INV\_ACOS\_INS. Compared to INV\_ACOS\_INS, we also do not assimilate any in situ observations as additional constraints. Figure A1 shows that such a bias correction scheme (INV\_BD\_ACOS\_BC) successfully reduces European uptake of  $\text{CO}_2$  during 2010 to 0.96 GtC/a from 1.58 GtC/a for INV\_BD\_ACOS. Table A1 shows that after the applying bias correction scheme, INV\_BD\_ACOS\_BC and INV\_BD\_TCCON\_BC are consistent (0.94 GtC/a vs 0.96 GtC/a) despite different lateral boundary conditions provided by INV\_ACOS and from INV\_TCCON. But INV\_BD\_TCCON\_BC (0.94 GtC/a) has 0.15 GtC/a more uptake than INV\_BD\_TCCON (0.79 GtC/a). We find a similar difference using UOL data (not shown), which infer an annual uptake of 0.71 GtC/a (0.56 GtC/a) with (without) the on-line bias correction.

We next examine the effectiveness of the inversion system that uses an on-line bias correction with large a priori uncertainty. Generally, large a priori uncertainty for biases will lead to the eventual loss of constraint by the observed mean  $\text{CO}_2$  concentration across Europe. The weakened constraint can be seen by the enlarged a posteriori error (by 0.04 GtC/a) for INV\_BD\_TCCON\_BC. In additional OSSEs (Table A2) we find that the loss of such a constraint can result in large systematic errors in estimated fluxes.

In these OSSEs, we assume the a priori estimates for 12 European sub-regions to be the same as the a priori used by INV\_TCCON. Similar to INV\_BD\_TCCON, we set the fluxes outside European region to

725 be the a posteriori estimates by INV\_TCCON. We assimilate the INV\_TCCON model ACOS  $X_{CO_2}$  retrievals over Europe, to test the ability of the system to recover the “true” European flux (defined by INV\_TCCON) from the assumed a priori that we define as the CASA model. Without the on-line bias correction, the quasi-regional inversion INV\_REG\_ENKF reproduces the truth for most months (Figure A2), and the associated annual uptake of 0.55 GtC/a compared to the true value of 0.58 GtC/a. If we also estimate monthly  $X_{CO_2}$  bias with a large a priori uncertainty of 100 ppm (INV\_REG\_BC), the a posteriori European uptake is systematically underestimated for almost all 730 months in 2010 (Figure A2). Consequently, the a posteriori annual uptake is about 0.38 GtC/a, which is 35% smaller than the truth (Table A2). Weakening the observation constraint also enlarges the a posteriori uncertainty from 0.22 GtC/a for INV\_REG\_ENKF to 0.27 for INV\_REG\_BC. But we find that increases in the estimated a posteriori uncertainty (by 0.05 GtC/a) are smaller than the increase in the systematic deviation from the true annual uptake (by 0.19 GtC/a).

735 More importantly, we find that the derived annual uptake is not linearly correlated to the assumed true fluxes. In experiment INV\_REG\_BC\_SP (Table A2) we replace the true fluxes (defined by INV\_TCCON) over the first 3 of 12 European sub-regions, which are at the south part of Europe (roughly south of 47° N), with values from CASA model. As a result, the new true fluxes have an annual uptake of about 0.48 GtC/a across Europe, which is about 18% (0.1 GtC/a) lower than the 740 original one defined by INV\_TCCON for INV\_REG\_BC. We then re-generate model ACOS  $X_{CO_2}$  data by running GEOS-Chem driven by the new hybrid true fluxes. However, after assimilating the new model  $X_{CO_2}$  data, INV\_REG\_BC\_SP infers an annual uptake of 0.37 GtC/a, which is almost the same as the posterior estimate (0.38 GtC/a) of INV\_REG\_BC, failing to reproduce the 18% decrease from the true value of 0.58 GtC/a assumed for INV\_REG\_BC to the 0.48 GtC/a assumed for INV\_REG\_BC\_SP. 745 In contrast, the quasi-inversion without on-line bias correction (INV\_REG\_ENKF\_SP) well reproduces such decrease.

The bias correction across Europe can also increase the sensitivity to sub-regional biases. To illustrate this we added 1 ppm bias to the simulated observations during June to August of 2010 over south-west Europe between 35°N to 42°N and 15°W to 20°E (mostly over Spain and Italy). Without 750 an on-line bias correction, adding the 1 ppm bias over the south-west strip leads to a small change (0.01GtC/a) in the annual uptake: a (slightly) reduced uptake in the first half of 2010 is largely compensated by a slightly enhanced uptake in the second half of 2010. Conversely, when we use an on-line bias correction with large prior errors (INV\_REG\_BC\_1ppm), the 1 ppm positive bias increases the uptake by about 0.24 GtC in June, July and August. This implies that without the 755 constraint from the mean concentration across the whole European region, the inversion system is free to interpret the higher concentrations over the small south-west strip as the signal of more uptakes over other larger part of Europe. As a result, the annual uptake changes from an underestimation of 35% by INV\_REG\_BC to an overestimation of 15% by INV\_REG\_BC\_1ppm (0.65 GtC/a) (Table A2).

760 In summary, our quasi-regional inversion experiments highlight the sensitivity of regional flux inversions to the accurate description of the boundary conditions around the domain. Using an on-line bias correction can be helpful when the bias has been properly characterized. Over-correcting the bias can weaken the observation constraints, and possibly increase sensitivity to other small-scale unknown biases. We have also tested bias correction schemes using a different inversion 765 algorithm (the Maximum A Posteriori (MAP) approach, Fraser et al., 2014), and found similar

770 deficiencies when the a priori uncertainty of the regional observation bias is assumed to be very large. Our studies cannot prove or disprove Reuter et al. (2014), but it does highlight previously unrecognized limitation to the approach. The diversity of results reached under different assumptions associated with observation biases and emission spatial patterns highlight the importance for us as a community to investigate the interaction between observation and the inversion system for achieving consistent flux estimates in the future from assimilation of the upcoming observations from OCO-2 satellite as well as from the improved in situ networks.

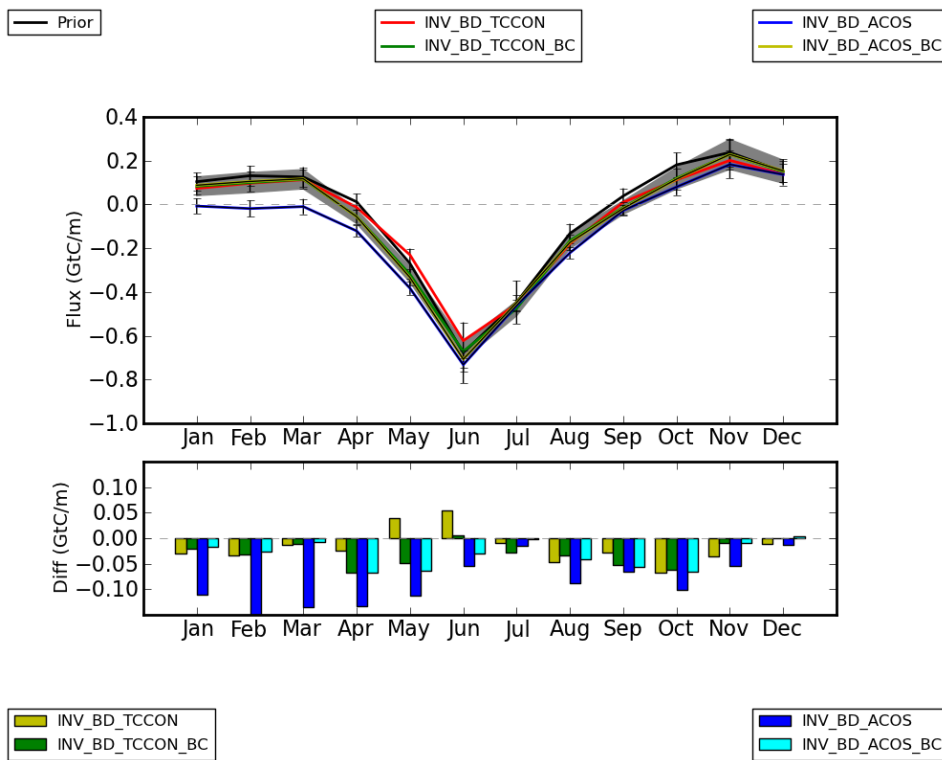


Name	Description	Flux (GtC/a)	Uncertainty (GtC/a)
INV_BD_TCCON	Only ACOS data over Europe are assimilated to infer monthly fluxes over 12 European sub-regions. Fluxes outside EU are fixed to INV_TCCON inversion.	-0.79	0.18
INV_BD_TCCON_BC	The same as INV_BD_TCCON, but monthly bias with an assumed prior uncertainty of 100 ppm are included as additional parameters to be estimated.	-0.94	0.22
INV_BD_ACOS	The same as INV_BD_TCCON, but external regional fluxes are fixed to INV_ACOS.	-1.58	0.18
INV_BD_ACOS_BC	The same as INV_BD_ACOS, but estimates for monthly observation bias included.	-0.96	0.22

**Table A1:** The same as Table 1 but for quasi-regional inversions where only ACOS  $X_{CO_2}$  are assimilated.

Name	Description	Flux (GtC/a)	Uncertainty (GtC/a)
INV_REG_ENKF	Synthetic ACOS data over Europe are assimilated to infer monthly fluxes over 12 European sub-regions, which prior estimates are assumed to be same as INV_ACOS (i.e., CASA model). Here we assume the true fluxes be a posteriori of INV_TCCON inversion.	-0.55	0.22
INV_REG_BC	The same as INV_REG_ENKF, but estimates for monthly bias are included as additional parameters.	-0.38	0.25
INV_REG_ENKF_1ppm	The same as INV_REG_ENKF, but 1ppm bias is added to the synthetic observations over a strip at south-west Europe for three months from June to August in 2010.	-0.54	0.22
INV_REG_BC_1ppm	The same as INV_REG_BC, 1ppm bias is added to the synthetic observations over a strip at south-west Europe for three months from June to August in 2010.	-0.65	0.25
INV_REG_ENKF_SP	The same as INV_REG_ENKF, but the 'true fluxes' over the first 3 of the 12 European sub-regions are replaced by CASA model values.	-0.47	0.22
INV_REG_BC_SP	The same as INV_REG_ENKF_SP, but with on-line bias correction with assumed prior uncertainty of 100 pm.	-0.37	0.25

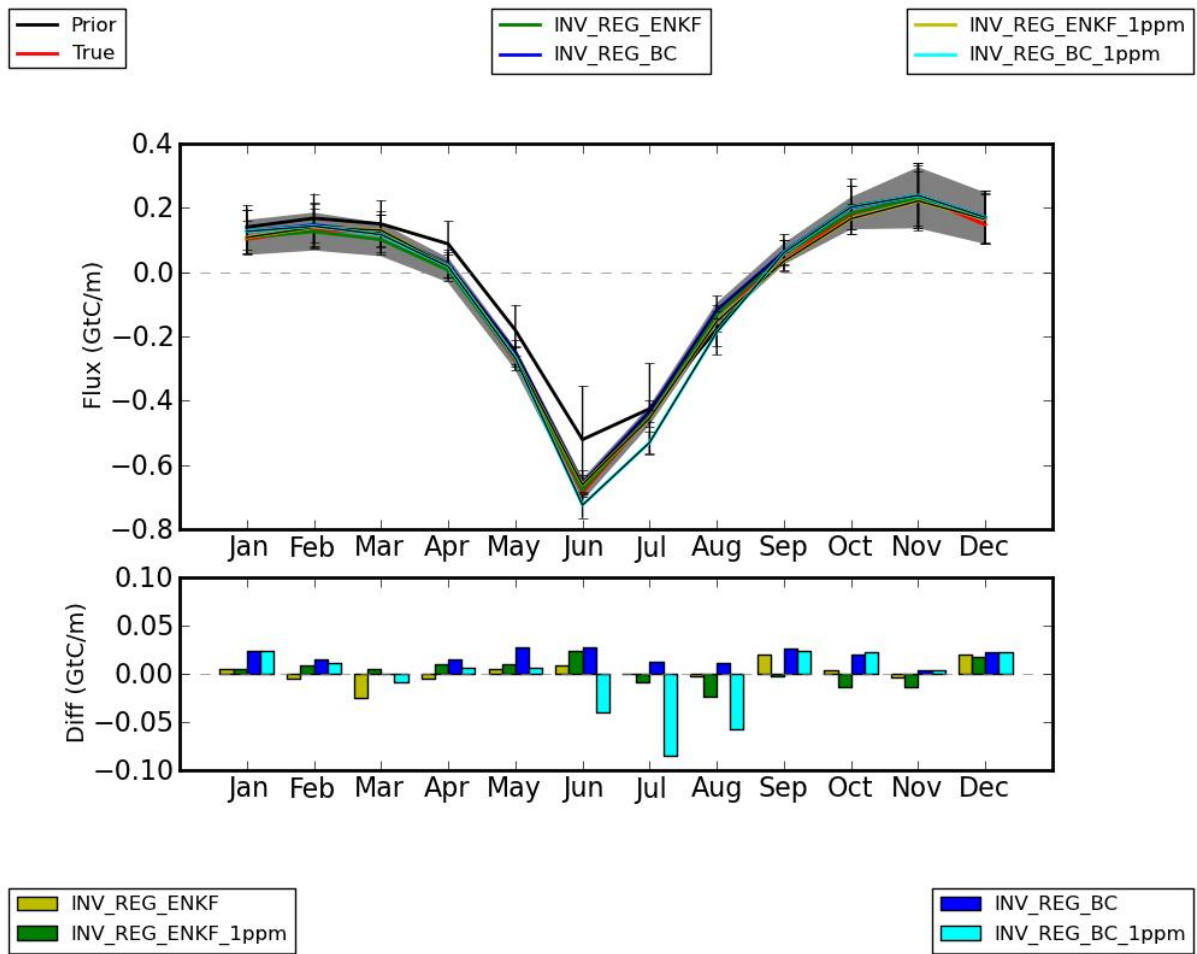
**Table A2:** The same as Table A1 but for Observation System Simulation Experiments, where we assimilate synthetic ACOS  $X_{CO_2}$  from model simulations forced by the assumed the 'True fluxes'.



**Figure A1:** As Figure 4, but for the comparisons between the quasi-regional inversions. All the inversion experiments assimilate the same ACOS data set over Europe, with the a priori for 12 European sub-regions taken from posterior estimates from INV\_TCCON. Fluxes outside Europe are fixed to the posterior estimates of INV\_TCCON (INV\_BD\_TCCON and INV\_BD\_TCCON\_BC) or to the estimates of INV\_ACOS (INV\_BD\_ACOS and INV\_BD\_ACOS\_BC). INV\_BD\_TCCON\_BC and INV\_BD\_ACOS\_BC also estimate the monthly bias across Europe as an additional parameter with an assumed a priori uncertainty of 100 pm estimated from ACOS data.

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**Figure A2:** As Figure 4, but for comparisons of the quasi-regional inversions for assimilation of synthetic ACOS retrievals against 'True' fluxes (INV\_TCCON). All the quasi-regional inversions have assumed the same a priori fluxes. But INV\_REG\_BC and INV\_REG\_BC\_1ppm also include the monthly observation bias across Europe, with a prior uncertainty of 100 pm, as additional parameters to be estimated from the synthetic observations. In INV\_REG\_ENKF\_1ppm and INV\_REG\_BC\_1ppm, 1ppm observation bias is added to the (synthetic) observations over a small south-west strip of Europe during the summer of 2010.



## Appendix B: Additivity of the increased European uptake estimates.

In the framework of Kalman Filter data assimilation [Feng et al., 2009], posterior flux estimates are determined by:

$$815 \quad \mathbf{F}^a = \mathbf{F}^f + \mathbf{K}(\mathbf{y}_{obs} - H(\mathbf{F}^f)), \quad (\text{B1})$$

where  $\mathbf{F}^a, \mathbf{F}^f$  are the prior and posterior estimates of monthly regional surface  $\text{CO}_2$  fluxes, respectively;  $\mathbf{y}_{obs}$  represents the GOSAT (real or simulated)  $X_{\text{CO}_2}$  retrievals.  $H$  is the observation operator for relating the surface fluxes to the observed GOSAT  $X_{\text{CO}_2}$ , which includes complicated atmospheric transporting as well as convolving of co-located model profiles with GOSAT averaging kernels [Feng et al., 2009; Chevallier et al. 2010]. Here, the Kalman gain matrix  $\mathbf{K}$  is given by

$$820 \quad \mathbf{K} = \mathbf{B}\mathbf{H}^T[\mathbf{H}\mathbf{B}\mathbf{H}^T + \mathbf{R}]^{-1}, \quad (\text{B2})$$

where  $\mathbf{B}$  is the a priori flux error covariance,  $\mathbf{R}$  is the observation error covariance, and  $\mathbf{H}$  is the Jacobian defined by

$$\mathbf{H} = \frac{\partial H(\mathbf{F}^f)}{\partial \mathbf{F}^f}. \quad (\text{B3})$$

825 Although the atmospheric transport is non-linear, the dependence of model concentrations (such as the column mixing ratios  $X_{\text{CO}_2}$ ) on the surface fluxes is nearly linear if we do not take into account any feedback of varying  $\text{CO}_2$  concentrations on atmospheric dynamics (for example, Chevallier et al., 2010; Baker et al., 2006). As a result, the gain matrix is eventually independent of actual observation values, but will still be affected by the location and uncertainty of observations.

830 As described in the main text, we split the actual (or simulated)  $X_{\text{CO}_2}$  observations into two parts: Part A for observations within Europe; and Part B for observations outside Europe. For the GOSAT inversions (such as INV\_ACOS), we denote the observation vector as:

$$\mathbf{y}_{obs} = \begin{bmatrix} \mathbf{G}^A \\ \mathbf{G}^B \end{bmatrix}, \quad (\text{B4})$$

The corresponding posterior flux estimate is given as:

$$835 \quad \mathbf{F}_G^a = \mathbf{F}^f + \mathbf{K} \left( \begin{bmatrix} \mathbf{G}^A \\ \mathbf{G}^B \end{bmatrix} - H(\mathbf{F}^f) \right). \quad (\text{B5})$$

In experiment INV\_MOD\_ALL, we replace the retrieved  $X_{\text{CO}_2}$  values by the reference model simulation (from INV\_TCCON), so that the observation vector becomes

$$\mathbf{y}_{obs} = \begin{bmatrix} \mathbf{M}^A \\ \mathbf{M}^B \end{bmatrix}, \quad (\text{B6})$$

840 so that the resulting flux estimates are:

$$\mathbf{F}_M^a = \mathbf{F}^f + \mathbf{K} \left( \begin{bmatrix} \mathbf{M}^A \\ \mathbf{M}^B \end{bmatrix} - H(\mathbf{F}^f) \right). \quad (\text{B7})$$

The gain matrix in Eq. B7 is the same as Eq. B5. Similarly, for INV\_MOD\_ONLYEU where GOSAT X<sub>CO2</sub> retrievals over Europe are replaced by model simulations, we have

$$\mathbf{F}_{MG}^a = \mathbf{F}^f + \mathbf{K} \left( \begin{bmatrix} \mathbf{M}^A \\ \mathbf{G}^B \end{bmatrix} - H(\mathbf{F}^f) \right). \quad (\text{B8})$$

845 And for INV\_MOD\_NOEU where GOSAT XCO2 retrievals outside Europe are replaced by model simulations, we have

$$\mathbf{F}_{GM}^a = \mathbf{F}^f + \mathbf{K} \left( \begin{bmatrix} \mathbf{G}^A \\ \mathbf{M}^B \end{bmatrix} - H(\mathbf{F}^f) \right). \quad (\text{B9})$$

From equations B5, B7, B8, and B9, we can directly obtain:

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$$\mathbf{F}_G^a - \mathbf{F}_M^a = (\mathbf{F}_{MG}^a - \mathbf{F}_M^a) + (\mathbf{F}_{GM}^a - \mathbf{F}_M^a). \quad (\text{B10})$$

Equation B10 demonstrates that elevated European uptake is the sum of the individual contributions from INV\_MOD\_NOEU and INV\_MOD\_ONLYEU. As discussed in Section 3, such additivity has also  
855 been found in our inversion results (Table 1), despite approximations in numerically solving posterior fluxes [Feng et al., 2009].

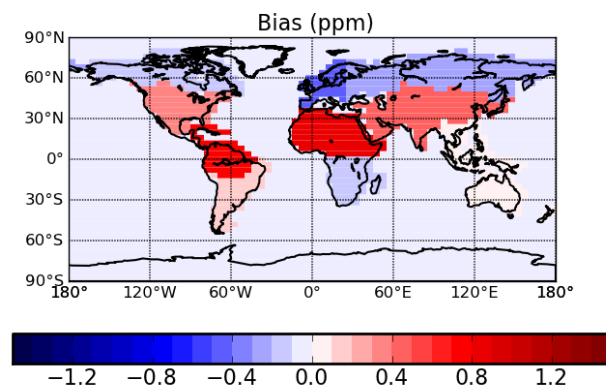
## Appendix C: Regional and Sub-regional systematic errors inferred in joint data assimilation.

860 In the joint data assimilation, we attempt to estimate and remove systematic errors at the regional and sub-regional scales from GOSAT XCO<sub>2</sub> retrievals. The assimilated XCO<sub>2</sub> retrieval can be described as :

$$y^c = y - bias(m, i), (C1)$$

865 where  $y$  represents GOSAT retrievals before the (extra) bias correction, and  $y^c$  is the bias-corrected XCO<sub>2</sub> data that we assimilate in our joint data assimilation experiments. For simplicity, we have assumed the regional (sub-regional) bias,  $bias(m, i)$ , is a function only of month ( $m$ ) and geographical region ( $i$ ).

870 In the joint data assimilation experiments, we consider  $bias(m, i)$  as part of the state vector that we infer from assimilating in situ and satellite observations. Figure C1 shows the resulting bias (in ppm) for March 2010. Like other model and GOSAT inter-comparisons (see for example, Lindqvist et al., 2015), our results demonstrate a strong spatial dependence of the derived systematic errors. As discussed in Section 4, our results reflect the mean differences between the inversion system and XCO<sub>2</sub> retrievals at (sub) regional scales, which does not necessarily suggest that the GOSAT XCO<sub>2</sub> bias (as well as the coverage) within these (sub) regions is homogeneous.



875 **Figure C1:** Inferred regional bias (in ppm) for March 2010 over TransCom regions and 2 European (West and North) sub-regions.