

Answer to Anonymous referee #2 comments

We would like to thank anonymous referee #2 for his comments that helped in improving the quality of our paper. The changes proposed by the referee are marked in the revised version of our manuscript with a blue color.

Major or specific comments:

1) “It would be very useful if you could draw a major conclusion as to which variable out of CFC, COT, and AOD were most important to the RegCM4-CM SAF SSR deviations over the whole of Europe and on an annual basis. This should be included in the paper and in the abstract.”

Answer: We address this issue in the revised version of our manuscript by adding a few lines in the abstract, in Section 3.6 and in the Conclusions section.

2) “Overall the data used in the analyses are presented clearly in section 2, but two of the subsections could be written more concisely dealing with the equations of Rel and Rei in section 2.1 and the CM SAF satellite data in 2.2. Please see the minor comments below for more details.”

Answer: Please find answers below (minor comments).

3) “Regarding the datasets used in the study, I recommend making a table to show all the variables and its source from the datasets used and their corresponding periods and original resolutions. The reader can simply refer to this table and see at once all the variables and datasets used for the analysis. Please see the minor comments for more details.”

Answer: We inserted a new table in the revised version of the manuscript (Table 1) with the parameters being analyzed in this work, their sources, the original resolution at which the data were acquired and the corresponding time periods.

4) “Detailed information was given on the biases found in the variables from the literature including the cloud property variables from CM SAF satellite but none on the other data (AOD, ASY, and SSA, ALB, and WV). Please add this in your section 2.3 Other data.”

Answer: We address this issue in the revised version of our manuscript by adding a short paragraph in the end of Section 2.3.

5) “As completed for the cloud fractional cover and the cloud optical thickness, additional comments should be made dealing with the spatial patterns seen in the cloud effective radius, aerosol optical properties, and other parameters (WV and ALB) compared to that of the SSR of RegCM. From a qualitative perspective, do

these parameters explain the SSR patterns seen in Figure 1? Such comments should be made respectively at the end of their sections, i.e. sections 3.3.2, 3.4, and 3.5.”

Answer: We addressed this issue by adding a few lines in the end of Section 3.3.2 and 3.4. However, we would like to comment here that a direct connection of the observed SSR bias patterns with atmospheric parameters is not a straightforward procedure. An effort to qualitatively assess the RegCM4-CM SAF differences is mostly reasonable in the case of CFC, COT and AOD, since, these are the main determinants of surface solar radiation. On the other hand, in some cases (e.g. WV and ALB) the radiative effect of the examined parameters is either negligible or the bias they cause in solar radiation is monotonous (overestimation or underestimation for the whole region). In these cases it is obvious that we cannot reach safe conclusions (e.g. Section 3.5) and this was the reason why we decided to introduce the quantitative approach with the use of a radiative transfer model in this paper.

6) “The conclusion section seems to be a repeat of the results. If you do this, I recommend to make a summary of the results by writing these paragraphs more concisely. Also, a few comments as a separate paragraph should be written on comparing and/or contrasting these SSR results to the ones in the references you cited in the Introduction, i.e. Jaeger et al. (2008) and Kothe et al. (2011). The new title of this section should then reflect these changes and called the Summary and Conclusions section.

Answer: We followed the referee’s suggestion and shortened the conclusions section by more than 30%. Our conclusions are now presented in a more condensed and precise way. The studies mentioned by the referee are focusing on the net surface solar radiation where albedo plays a major role. Therefore, we selected not to mention these studies in the conclusion section.

Minor comments:

1) “Regarding section 2.1 on description of the model, where it is mentioned that the emissions are monthly historical, are they also time independent or not changing in time? If so, this would affect your results of simulated SSR. Please account for this in your conclusions.”

Answer: It is mentioned in the revised version of the manuscript that the emissions of the anthropogenic aerosols are based on monthly, timed-dependent, historical emissions from CMIP5. There is only a marginal change in the emissions that RegCM4 takes into account for the years 2000-2009 and the sub-periods 2000-2005 (MFG) and 2006-2009 (MSG) that we examine in this paper. Therefore our results would not be affected by the use of changing emissions.

2) “Lines 170-173: You mention the influence of CFC, Re, and cloud water path (CWP), but is there any particular reason you analysed the cloud optical thickness (COT) instead of the CWP or not analyzing both?”

Answer: COT along with CFC is one of the basic optical properties describing clouds and there are numerous studies in the literature using this parameter. So, COT is considered an ideal parameter to describe the vertical development of clouds. Since we use COT and Re, the use of CWP would be meaningless, as these three parameters are connected with the following relationship: $CWP = \frac{2}{3} \rho Re COT$ (where ρ is the density of water).

3) “The equations that follow line 180 through line 194 can be all taken out and referred to from the studies of Giorgi et al. (2012), Slingo (1989), and Briegleg et al. (1992) if the reader is interested.”

Answer: Since our main target is this paper to serve as a textbook study for the evaluation of the ability of climate models to reproduce the SSR levels, we would prefer to keep these equations in the manuscript. This paper could serve as a bridge between the modelling and satellite community. Hence, we believe that details about the calculations done by the model and details about the satellite retrievals would be very helpful for members from both the communities to fully understand this research.

4) “Lines 223-229: This paragraph should be taken out and used instead in the introduction as you started in lines 82-85. Add this paragraph i.e, lines 223-229 to lines 82-85. As stated in the major revisions above, I recommend making a table at this point showing all the variables used, their data sources, periods, and original resolutions. It should also be made clear here in the text of this paragraph or somewhere in the introduction what period you will use for your main investigation. Following this in the introduction, you should also state here why you chose these data, such as its used as input for the radiative transfer model which is also used in the CM SAF SSR estimation as you pointed out in lines 377-380. It would be clearer to the reader if you pointed this out sooner as in the introduction..”

Answer: We addressed all these issues in the revised manuscript following the referee’s suggestions.

5) “Lines 251-298: The descriptions of the MagicSol-Heliosat algorithm and the MSG satellites should be written more concisely or condensed.”

Answer: The same answer as in minor comment 3.

6) “Lines 317-320: Is this homogeneity considered for Europe or globally?”

Answer: It is for Europe, we clarify this in the revised version of the paper.

7) “Lines 344-347: Does this bias refer to a global bias?”

Answer: It refers to SEVIRI’s disk; we also clarify this in the revised version of the paper.

8) “It is interesting that the results in Figure 10 for eastern Europe show that AOD contributes to the SSR positively for all months of the year, but why this is not reflected in the negative change in SSR in eastern Europe in Figure 1?”

Answer: We thank the reviewer for giving us the opportunity to clarify this. One should keep in mind that Figures 1 and 10 refer to % biases. For Eastern Europe the qualitative method predicts perfectly the relative seasonal variability of SSR bias, however, the surface radiation levels in winter and autumn are low, $\sim 38 \text{ W/m}^2$ and $\sim 85 \text{ W/m}^2$, respectively (among the sub-regions appearing in Figure 10 Eastern Europe exhibits the lowest SSR levels). This means that a 10% bias would be $\sim 4 \text{ W/m}^2$ and $\sim 8 \text{ W/m}^2$, which is below the combined CM SAF and radiative transfer model uncertainty. So, in this case one should not be very strict with the method and focus on the relative month by month seasonal variability of SSR bias.

Technical Comments:

1) “Line 8: Change the sentence to: “The SSR bias. . .”

Answer: Corrected in the revised version of the manuscript.

2) “The fonts of the figures in the main text of the paper should still be addressed as the fonts are still hard to read at that size.”

Answer: We will collaborate closely with the production team of the journal so as to make sure that the size of the figures and the size of the fonts will be optimal for reading.

3) “All figures: Larger fonts should be used for all parts of the figure. The same corrections should be made for all remaining figures.”

Answer: We will collaborate closely with the production team of the journal so as to make sure that the size of the figures and the size of the fonts will be optimal for reading.

4) “A black font or one that would be clearer to read should be used in figures 1,4, and 7. This refers to the text of different seasons on the upper left-hand corner of each panel in the map.

Answer: Unfortunately, when using black fonts this part of the maps appears very blurred. Prior to the submission of the paper we did several efforts with various colors and we concluded that white fonts with a black border were the optimal solution. However, once again we assure the referee that we will collaborate closely with the production team of the journal so as to make sure that the size of the figures and the size of the fonts will be optimal for reading.

Answer to Dr J. Trentmann comments

We would like to thank Dr J. Trentmann for his comments that helped us to improve the quality of our paper. The changes proposed by the referee are marked in the revised version of our manuscript with a red color.

Major or specific comments:

1) “The manuscript itself already provides a lot of tables with detailed information on the regional differences; in addition a 34-page supplement is accompanying the manuscript. Overall, by the huge amount of numbers, tables, and figures in the manuscript the main message of the manuscript sometimes is not clearly highlighted. Some of the tables and figures, in particular in the supplement, are not referred to in the manuscript. I suggest that the authors consider to remove some of the tables, in particular those without references in the text, and to focus the attention of the reader on the main results of the analysis, which are highly relevant. Please find more specific comments for the streamlining of the manuscript below.”

Answer: Since our main target is this paper to serve as a textbook study for the evaluation of the ability of climate models to reproduce the SSR levels, we would prefer to keep the electronic supplement in its current form. The reader will be able to find all the details about the various parameters utilized in this paper which would be very helpful for future follow-up studies with the same or other climate models.

2) “The differences between the model and the observations are provided with two digits. This accuracy does not seem to be appropriate considering the high spatial variability and the overall uncertainty. It would be sufficient, from my point of view, to provide most values in the text and in the table with one digit, sometimes even integer values would be appropriate.”

Answer: We agree with the referee, the values appearing in the text are given with one digit.

3) “Recently, the CM SAF released a new surface solar radiation data set: SARAH (http://dx.doi.org/10.5676/EUM_SAF_CM/SARAH/V001), which provides consistent data from 1983 to 2013. Likely, this data set has not been available during the research documented in this manuscript. However, the results of this manuscript will be much more robust and the manuscript will be much easier to follow if this new data set would be used for the assessment, since no differentiation would be required for the time periods prior and after 2005. If time and resources allow I recommend to redo the analysis using the SARAH data set and to replace the current results. The supplement could be substantially shortened or even removed.”

Answer: We agree with the referee that a future repetition of this work using the new CM SAF SARAH product and possibly various model set-ups would be very important. It is indeed in our plans to proceed to such a research in the near future. Taking into account the time and resources especially as far the radiative transfer

calculations are concerned we prefer to keep the original products in this manuscript. However, we mention in the conclusions section of the revised paper that an update of this work using the new CM SAF SARA product would be very interesting.

Specific comments:

1) “Page 18493, lines 12 ff: Please add a brief statement of the treatment of cloud ice and convective cloud coverage in the radiation scheme in RegCM4. Also add a brief statement on the aerosol scheme and their radiative treatment.”

Answer: We addressed this by adding a few lines in the two paragraphs prior to the one referred in the comment.

2) “Formulas (1) to (7): The diagnostic calculations of the different cloud parameters might not need to be explicitly stated here, a reference to the model describing paper would be sufficient.”

Answer: As discussed above, our main target is this paper to serve as a textbook study constituting a bridge between the modelling and satellite community. Hence, we believe that details about the calculations done by the model and details about the satellite retrievals would be very helpful for members from both the communities to fully understand this research.

3) “Section 2.2: The section on the CM SAF satellite data could be substantially shortened; details of the retrieval algorithm could be left out here with references to the corresponding articles.”

Answer: The same as in specific comment 2.

4) “Section 2.2: Please state clearly, which data set of the surface solar radiation has been used for the assessment. Two different data sets have been used, one for the time period prior to 2006 and one for the years 2006 to 2009. If possible, please provide the digital object identifiers for those data sets. I suspect that the MVIRI data set (DOI:10.5676/EUM_SAF_CM/RAD_MVIRI/V001) has been used for the years 2000 to 2005, and the surface radiation data set from the CM SAF CLAAS data set (DOI:10.5676/EUM_SAF_CM/CLAAS/V001) has been used for the years 2006 to 2009.”

Answer: We address this in the revised version of the manuscript, the digital object identifiers of the two datasets utilized in this paper mentioning the periods of MFG (2000-2005) and MSG (2006-2009) in the introduction and in the main body of the manuscript.

5) “Section 2.4: ...Please carefully check formula (9) and make sure that the sums are correctly calculated. Based on the right side of the formula the middle part should read:...”

Answer: We made sure that the formula is correct.

6) “Page 18500, line9 ff: The other statistical metrics are only mentioned here, but there is no clear definition; the values are listed in several Tables in the Appendix, but they are referred to at all in the text; I suggest to remove these tables.”

Answer: We agree with the referee that the other metrics should be defined prior to their use. So, in the revised manuscript we write “...other statistical metrics (correlation coefficient R , normalized standard deviation NSD , modified normalized mean bias $MNMB$, root mean square error $RMSE$) are also defined, calculated and presented in the electronic supplement of this manuscript...”

7) “Page 18503, lines 7ff, Figure 1: The strong positive bias observed in the Northern Europe during winter is likely due to the satellite data set; no such bias is observed for the period 2000 to 2005 (Fig. S3) when the other satellite data is used as reference.”

Answer: We thank the referee for giving us the opportunity to comment on this. Indeed, the strong positive bias appearing over Northern Europe is likely due to the MSG satellite data since no such bias is observed for the MFG data. We have to highlight that the SSR levels over the region are very low in winter (less than 20 W/m^2). Therefore, retrieval or model uncertainties of few Watts/m^2 would appear as a very large percent bias in the maps. This is the reason why we do not further comment on the strong winter bias over Northern Europe and we also did not proceed to radiative transfer calculations over this region.

8) “Tables 1 and 2: Please order the regions according to Figure 3: start with EU, LA, OC, and then go North – South: NE, CE, EE, IP, CM, EM, NA. Please check the significance of the bias; to me it appears that small NMBs like -1.16 might not be significant considering the high variability of the original data (134 ± 89 and 136 ± 83).”

Answer: The regions have been rearranged in the tables of the revised manuscript. Also, we checked again the significance results using equation (10) and we found that our results are correct. It has to be mentioned here that the values and the standard deviations appearing in the tables come from the whole timeseries of all the grid-cells that fall within each region and not from monthly spatial averages.

Main document changes and comments

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SSR

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on a monthly basis

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Overall, for the European domain, the underestimation of CFC by RegCM4 is the most important cause of the SSR overestimation on an annual basis.

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, SSA

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For the scopes of this study, the same parameters are extracted from satellite-based observational data (CM SAF, CERES), data from an aerosol climatology (MACv1) and data from the ERA-Interim reanalysis (see Table 1).

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(2000-2009)

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satellite-based

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SSR data from the Meteosat First Generation (MFG) satellites are available for the period 2000-2005 while data from the Meteosat Second Generation (MSG) satellites are available for the period 2006-2009.

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against data from MFG and MSG

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against data from MSG

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The CFC, COT, Re, AOD, ASY, SSA, ALB and WV datasets were chosen so as to be consistent with the CM SAF SSR dataset.

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for the MSG SSR period (2006-2009)

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, time-dependent,

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convection

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Convection is triggered when the buoyancy level is higher than the cloud base level. The cloud mixing is considered episodic and inhomogeneous, the convective fluxes being based on a model of sub-cloud-scale updrafts and downdrafts (Giorgi et al., 2012).

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per layer

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per layer

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The fraction (f_{ice}) of cloud water that consists of ice particles is given as a function of temperature (T), the fraction (f_{liq}) of the liquid water droplets being calculated as $f_{liq} = 1 - f_{ice}$.

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Overall, as discussed in the introduction, for the scopes of this research we used the following parameters from a RegCM4 simulation: 1) SSR, 2) CFC, 3) COT, 4) Re, 5) AOD, 6) ASY, 7) SSA, 8) WV and 9) ALB. The same parameters were extracted from satellite-based observational data (CM SAF, CERES), data from an aerosol climatology (MACv1) and data from the ERA-Interim reanalysis.

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Meteosat First and Second Generation (

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(DOI:10.5676/EUM_SAF_CM/RAD_MVIRI/V001)

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(DOI:10.5676/EUM_SAF_CM/
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respectively)

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over Europe

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for SEVIRI's disk

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To our knowledge, the uncertainty of the MACv1 aerosol parameters used here has not been reported somewhere in detail. However, due to the methodology followed for the production of the MACv1 climatology, the MACv1 data are consistent with the AERONET ground network. The CERES broadband surface albedo over land exhibits a relative bias of -2.4% compared to MODIS. Specifically, over deserts, the relative bias drops to -2.1% (Rutan et al., 2009). A detailed evaluation of the ERA-Interim WV total column product does not exist. Only recently, the upper troposphere – lower stratosphere WV data were evaluated against airborne campaign measurements showing a good agreement (30% of the observations were almost perfectly represented by the model) (Kunz et al., 2014).

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for 2000-2005

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for 2006-2009

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A constant underestimation of Rel and Rei is observed for the whole Europe (see Figs. S6 and S8).

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This is apparent in Figs. S13 and S15 where the RegCM4-CM SAF NMB maps are presented along with the latitudinal variability of the two products.

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Concluding, for the total of the six sub-regions, CFC and AOD are the most important factors that determine the SSR overestimation by RegCM4 on an annual basis. The underestimation of CFC and AOD by the model causes an annual overestimation of SSR by 4.8% and 2.6%, respectively.

4 Conclusions

In the present study, a decadal simulation (2000-2009) with the regional climate model RegCM4 is implemented in order to assess the model's ability to represent the SSR patterns over Europe. The RegCM4 SSR fields are evaluated against satellite-based observations from CM SAF. The annual bias patterns of RegCM4-CM SAF are similar for both MFG (2000-2005) and MSG (2006-2009) observations. The model slightly overestimates SSR compared to CM SAF over Europe, the bias being +1.5% for MFG and +3.3% for MSG observations. Moreover, the bias is much lower over land than over ocean while some differences appear locally between the seasonal and annual bias patterns.

In order to understand the RegCM4-CM SAF SSR deviations, CFC, COT and Re data from RegCM4 are compared against observations from CM SAF (MSG period). For the same reason, AOD, ASY, SSA, WV and ALB from RegCM4 are compared against data from MACv1, ERA-Interim reanalysis and CERES since these data are similar to the ones used as input in the retrieval of CM SAF SSR.

CFC is significantly underestimated by RegCM4 compared to CM SAF over Europe by 24.3% on annual basis. Part of the bias between REGCM4 and CM SAF SSR can be explained through CFC with the underestimation of CFC leading to a clear overestimation of SSR. It was also found that RegCM4 overestimates COT compared to CM SAF on an annual basis suggesting that COT may explain part of the RegCM4-CM SAF SSR deviations that could not be explained through CFC over specific regions. In addition, RegCM4 underestimates significantly Rel and Rei compared to CM SAF over the whole European domain on an annual basis. A comparison of the RegCM4 AOD seasonal patterns with AOD values from the MACv1 aerosol climatology reveals that RegCM4 overestimates AOD over the region of NA and underestimates it for the rest of the European domain. ASY and SSA are slightly underestimated by the model. The comparison of RegCM4 WV against data from ERA-Interim reanalysis, reveals a clear overestimation over Europe. In line with previous studies, RegCM4 underestimates ALB significantly over CE, EE and NA compared to climatological data from CERES with a striking difference between land and ocean.

The combined use of SBDART radiative transfer model with RegCM4, CM SAF, MACv1, CERES and ERA-Interim data for the common period 2006-2009 shows that the difference between RegCM4 and CM SAF SSR is mostly explained through CFC, COT and AOD deviations. In the majority of the regions, CFC leads to an overestimation of SSR by RegCM4. In some cases, COT leads to an underestimation of SSR by RegCM4, while for the majority of the regions leads to an overestimation. Apart from NA, where AOD leads to a significant underestimation of RegCM4 SSR, AOD is generally responsible for the overestimation of SSR. The other parameters (Re, ASY, SSA, WV and ALB) play a less significant role, except for NA where they have a significant impact on the RegCM4-CM SAF SSR deviations. Overall, CFC and AOD are the major determinants of the SSR overestimation by RegCM4 on an annual basis. The underestimation of CFC and AOD by the model causes an annual overestimation of SSR by 4.8% and 2.6%, respectively.

Overall, it is shown in this study that RegCM4 simulates adequately the SSR patterns over Europe. However, it is also shown that the model overestimates or underestimates significantly several parameters that determine the transmission of solar radiation in the

atmosphere. The good agreement between RegCM4 and satellite-based SSR observations from CM SAF is actually a result of the contradicting effect of these parameters. Our results suggest that there should be a reassessment of the way these parameters are represented within the model so that SSR is not only well simulated but also for the right reasons. This would also allow for a safer investigation of the dimming/brightening effect since the SSR deviations would be safely dedicated to the one or the other parameter. It is suggested here that a similar approach should be implemented in the future to the same or other regional climate models with various setups also utilizing new satellite products (e.g. CM SAF SARA).

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Kunz, A., Spelten, N., Konopka, P., Müller, R., Forbes, R. M., and Wernli, H.: Comparison of Fast In situ Stratospheric Hygrometer (FISH) measurements of water vapor in the upper troposphere and lower stratosphere (UTLS) with ECMWF (re)analysis data, *Atmos. Chem. Phys.*, 14, 10803-10822, doi:10.5194/acp-14-10803-2014, 2014.

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Rutan, D., Rose, F., Roman, M., Manalo-Smith, N., Schaaf, C., and Charlock, T.: Development and assessment of broadband surface albedo from Clouds and the Earth's Radiant Energy System Clouds and Radiation Swath data product, *J. Geophys. Res.*, 114, D08125, doi:10.1029/2008JD010669, 2009.

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Table 1. List of the parameters being analyzed in this work, their sources, the original resolution at which the data were acquired and the corresponding time periods.

Parameter	Source	Resolution	Period
SSR	CM SAF MFG	0.03° x 0.03°	2000-2005
SSR	CM SAF MSG	0.05° x 0.05°	2006-2009
CFC	CM SAF MSG	0.05° x 0.05°	2004-2009
COT	CM SAF MSG	0.05° x 0.05°	2004-2009
Re	CM SAF MSG	0.05° x 0.05°	2004-2009
AOD	MACv1	1° x 1°	Climatology
ASY	MACv1	1° x 1°	Climatology
SSA	MACv1	1° x 1°	Climatology
ALB	CERES	1° x 1°	Climatology
WV	ERA-Interim	1° x 1°	2006-2009
All above	RegCM4	50km x 50km	2000-2009

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while the regions are listed in alphabetical order.

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	ANN			DJF			MAM			JJA			SON		
	MOD	SAT	bias	MOD	SAT	bias	MOD	SAT	bias	MOD	SAT	bias	MOD	SAT	bias
EU	175.0±106.5	169.3±96.7	3.3	77.1±57.1	74.2±57.2	3.9	206.8±83.0	206.7±67.0	0.0*	281.6±70.6	265.2±55.2	6.2	126.3±77.4	123.3±71.3	2.4
LA	173.1±106.9	171.9±97.2	0.7	78.1±61.0	78.0±60.8	0.1*	202.7±85.7	208.7±68.6	-2.9	278.6±71.7	267.0±55.0	4.4	124.9±79.0	126.1±72.8	-0.9
OC	178.2±105.6	164.9±95.7	8.1	75.3±49.7	67.7±49.8	11.3	213.8±77.8	203.2±64.2	5.2	286.7±68.2	262.1±55.3	9.4	128.7±74.5	118.6±68.4	8.4
NE	104.0±81.2	113.7±93.4	-8.5	19.3±12.0	12.7±16.8	52.4	137.6±53.4	160.4±60.8	-14.2	198.7±45.5	219.4±43.3	-9.4	52.9±38.2	53.4±44.3	-1.0*
CE	134.5±89.2	136.1±83.1	-1.2	42.3±20.8	42.8±24.4	-1.1*	158.1±55.6	174.0±51.3	-9.1	245.6±47.9	228.9±38.2	7.3	84.4±46.8	90.9±48.2	-7.2
EE	132.3±92.0	139.5±89.8	-5.2	37.5±17.5	38.8±22.1	-3.4	155.2±61.2	179.4±57.7	-13.5	248.4±44.9	242.8±36.5	2.3	80.1±46.0	88.8±48.8	-9.8
IP	197.9±95.1	194.7±84.4	1.7	91.7±26.9	98.6±27.5	-7.0	224.8±56.5	224.0±46.3	0.4*	317.5±29.1	296.3±32.3	7.2	148.6±53.9	151.8±50.4	-2.1
CM	209.8±98.6	195.1±85.1	7.5	97.3±29.1	96.7±27.1	0.6*	243.7±59.2	225.9±46.2	7.9	331.3±27.3	299.9±25.1	10.4	157.7±53.5	149.8±45.4	5.3
EM	219.3±101.6	205.6±90.3	6.7	105.1±36.8	101.8±33.7	3.3	251.4±68.8	235.6±54.4	6.7	339.3±29.1	312.8±28.1	8.5	171.8±63.0	163.7±55.9	5.0
NA	261.8±82.3	243.8±69.5	7.4	164.7±35.2	161.8±31.9	1.8	303.8±41.3	280.2±33.7	8.4	353.5±20.5	320.5±21.6	10.3	217.2±49.5	205.8±39.7	5.5

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	ANN			DJF			MAM			JJA			SON		
	REG	CMS	bias	REG	CMS	bias	REG	CMS	bias	REG	CMS	bias	REG	CMS	bias
CE	134,52±89,22	136,1±83,09	-1,16	42,30±20,78	42,75±24,38	-1,05*	158,12±55,58	174,03±51,26	-9,14	245,57±47,93	228,91±38,24	7,27	84,38±46,83	90,92±48,19	-7,19
CM	209,83±98,64	195,12±85,11	7,53	97,29±29,07	96,68±27,13	0,63*	243,68±59,2	225,86±46,23	7,88	331,26±27,32	299,93±25,09	10,44	157,71±53,54	149,8±45,42	5,27
EE	132,29±91,96	139,54±89,84	-5,19	37,53±17,47	38,84±22,10	-3,38	155,22±61,16	179,35±57,73	-13,45	248,43±44,9	242,79±36,54	2,31	80,07±45,95	88,77±48,76	-9,79
EM	219,29±101,61	205,61±90,31	6,65	105,08±36,84	101,76±33,74	3,25	251,44±68,77	235,6±54,43	6,72	339,32±29,07	312,76±28,12	8,49	171,81±62,97	163,66±55,85	4,98
EU	174,99±106,45	169,33±96,69	3,34	77,08±57,06	74,17±57,20	3,92	206,76±83,03	206,67±67,03	0,04*	281,62±70,56	265,18±55,16	6,19	126,27±77,38	123,31±71,26	2,40
IP	197,88±95,07	194,66±84,41	1,65	91,71±26,94	98,58±27,47	-6,97	224,84±56,48	223,98±46,33	0,38*	317,51±29,05	296,28±32,29	7,16	148,61±53,88	151,77±50,36	-2,08
LA	173,07±106,88	171,92±97,18	0,66	78,12±60,95	78,01±60,82	0,14*	202,65±85,68	208,71±68,55	-2,90	278,63±71,74	266,99±55,01	4,36	124,87±79,02	126,07±72,78	-0,94
NA	261,83±82,27	243,79±69,53	7,40	164,74±35,19	161,80±31,87	1,81	303,81±41,30	280,2±33,73	8,42	353,48±20,51	320,52±21,65	10,28	217,2±49,53	205,79±39,71	5,54
NE	103,98±81,16	113,69±93,36	-8,54	19,31±12,00	12,67±16,78	52,44	137,62±53,40	160,43±60,82	-14,21	198,68±45,53	219,39±43,31	-9,43	52,88±38,18	53,42±44,28	-0,99*
OC	178,24±105,64	164,93±95,68	8,07	75,32±49,72	67,66±49,81	11,32	213,75±77,82	203,22±64,23	5,18	286,69±68,21	262,11±55,29	9,37	128,65±74,47	118,63±68,35	8,44

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	CFC	COT	Rel	Rei	AOD	ASY	SSA	ALB	WV
EU	-24.3	4.3	-36.1	-28.3	-35.3	-1.1	-4.2	1.6	12.0
LA	-13.7	7.3	-47.7	-26.4	-32.1	-1.8	-4.3	-28.3	11.4
OC	-38.4	-2.5	-18.3	-31.1	-42.0	0.1	-4.1	131.1	12.8
NE	-20.3	54.3	-32.8	-31.3	-75.9	1.0	-5.6	5.2	13.1
CE	-19.7	24.1	-45.1	-24.0	-63.6	0.0*	-5.9	-22.7	14.0
EE	-16.0	30.8	-44.6	-24.2	-64.6	2.1	-3.5	-40.7	10.8
IP	-13.7	-13.9	-46.1	-27.3	-7.4	-1.5	-4.8	-3.8	14.4
CM	-31.2	-30.7	-26.7	-27.6	-19.3	-0.7	-3.5	85.9	10.4
EM	-28.8	-22.0	-29.3	-28.4	-34.2	-0.0	-2.3	35.4	10.9
NA	0.4*	-39.8	-47.3	-30.0	25.0	-7.9	-3.5	-26.4	8.7

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On the ability of RegCM4 regional climate model to simulate surface solar radiation patterns over Europe: An assessment using satellite-based observations

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Abstract

In this work, we assess the ability of RegCM4 regional climate model to simulate surface solar radiation (SSR) patterns over Europe. A decadal RegCM4 run was implemented and evaluated against satellite-based observations from the Satellite Application Facility on Climate Monitoring (CM SAF) showing that the model simulates adequately the SSR patterns over the region. The [SSR](#) bias between RegCM4 and CM SAF is +1.54% for MFG (Meteosat First Generation) and +3.34% for MSG (Meteosat Second Generation) observations. The relative contribution of parameters that determine the transmission of solar radiation within

1 the atmosphere to the deviation appearing between RegCM4 and CM SAF SSR is also
2 examined. Cloud macrophysical and microphysical properties such as cloud fractional cover
3 (CFC), cloud optical thickness (COT) and cloud effective radius (Re) from RegCM4 are
4 evaluated against data from CM SAF. The same procedure is repeated for aerosol optical
5 properties such as aerosol optical depth (AOD) asymmetry factor (ASY) and single scattering
6 albedo (SSA), as well as other parameters including surface broadband albedo (ALB) and
7 water vapor amount (WV) using data from MACv1 aerosol climatology, from CERES
8 satellite sensors and from ERA-Interim reanalysis. It is shown here that the good agreement
9 between RegCM4 and satellite-based SSR observations can be partially attributed to
10 counteracting effects among the above mentioned parameters. The contribution of each
11 parameter to the RegCM4-CM SAF SSR deviations is estimated with the combined use of the
12 aforementioned data and a radiative transfer model (SBDART). CFC, COT and AOD are the
13 major determinants of these deviations [on a monthly basis](#); however, the other parameters also
14 play an important role for specific regions and seasons. [Overall, for the European domain, the
15 underestimation of CFC by RegCM4 is the most important cause of the SSR overestimation
16 on an annual basis.](#)

17

18 **1 Introduction**

19 Modeling climate on a regional scale is essential for assessing the impact of climate change
20 on society, economy and natural resources. Regional climate models are limited-area models
21 that simulate climate processes being often used to downscale dynamically global model
22 simulations or global reanalysis data for specific regions in order to provide more detailed
23 results (Laprise, 2008; Rummukainen, 2010). Several studies suggest that we can benefit from
24 the use of regional climate models, especially due to the higher resolution of stationary
25 features like topography, coastlines and from the improved representation of small-scale
26 processes such as convective precipitation (see Flato et al., 2013 and references therein).
27 Usually, regional climate models are evaluated and “tuned” according to their ability to
28 simulate temperature and precipitation (e.g. Giorgi et al., 2012; Vautard et al., 2013; Kotlarski
29 et al., 2014). However, as discussed in Katragkou et al. (2015), the role of other
30 climatological parameters should be included in the evaluation procedure of regional climate
31 models.

1 For example, the ability of regional climate models to assess surface solar radiation (SSR)
2 patterns has not received so much attention despite the fact that SSR plays a core role in
3 various climatic processes and parameters such as: 1) evapotranspiration (e.g. Teuling et al.,
4 2009), 2) hydrological cycle (e.g. Allen & Ingram, 2002; Ramanathan et al., 2001; Wang et
5 al., 2010; Wild and Liepert, 2010), 3) photosynthesis (e.g. Gu et al., 2002; Mercado et al.,
6 2009), 4) oceanic heat budget (e.g. Lewis et al., 1990; Webster et al., 1996; Bodas-Salcedo et
7 al., 2014), 5) global energy balance (e.g. Kim and Ramanathan, 2008; Stephens et al., 2012;
8 Trenberth et al., 2009; Wild et al., 2013) and solar energy production (Hammer et al., 2003)
9 and largely affects temperature and precipitation. The same stands for the parameters that
10 drive SSR levels, such as cloud macrophysical and microphysical properties (cloud fractional
11 cover CFC, cloud optical thickness COT and cloud effective radius R_e), aerosol optical
12 properties (aerosol optical depth AOD, asymmetry factor ASY and single scattering albedo
13 SSA), surface broadband albedo (ALB) and atmospheric water vapor amount (WV).
14 However, during the last years, there were a few regional climate model studies focusing on
15 the SSR levels or the net surface shortwave radiation, either to examine the
16 dimming/brightening effect (e.g. Zubler et al., 2011; Chiacchio et al., 2015) or to evaluate the
17 models (e.g. Jaeger et al., 2008; Markovic et al., 2008; Kothe and Ahrens, 2010; Kothe et al.,
18 2011; 2014; Güttler et al., 2014). These studies highlight the dominating effect of cloud cover
19 and surface albedo.

20 In this work, we go a step further, proceeding to a detailed evaluation of the ability of
21 RegCM4 regional climate model to simulate SSR patterns over Europe taking into account
22 not only CFC and ALB but also COT, R_e , AOD, ASY, [SSA](#) and WV. [For the scopes of this](#)
23 [study, the same parameters are extracted from satellite-based observational data \(CM SAF,](#)
24 [CERES\), data from an aerosol climatology \(MACv1\) and data from the ERA-Interim](#)
25 [reanalysis \(see Table 1\).](#) First a decadal simulation ([2000-2009](#)) is implemented with the
26 model and the output is evaluated against ~~satellite-based~~ observations from the [EUMETSAT](#)
27 [geostationary satellites of CM SAF. SSR data from the Meteosat First Generation \(MFG\)](#)
28 [satellites are available for the period 2000-2005 while data from the Meteosat Second](#)
29 [Generation \(MSG\) satellites are available for the period 2006-2009.](#) These data are
30 characterized by a high spatial (~3-5 km) and temporal resolution (15-30 min) and have been
31 validated in the past, constituting a well-established product. In Sect. 2.1., the basic features
32 of the model are described along with the simulation setup and the way various parameters
33 are calculated by the model. In Sects. 2.2. and 2.3., a description of the satellite data from CM

1 SAF and the other data which are used for the evaluation of RegCM4 is given, while, in Sect.
2 2.4., we discuss the methodology followed in this manuscript. Sect. 3.1. includes the
3 evaluation of RegCM4 SSR [against data from MFG and MSG](#), Sect. 3.2. and 3.3. the
4 evaluation of CFC, COT and Re [against data from MSG](#), Sect 3.4. the comparison of
5 RegCM4 AOD, ASY and SSA with data from MACv1 aerosol climatology and Sect 3.5. the
6 comparison of RegCM4 WV and ALB with data from ERA-Interim reanalysis and CERES
7 satellite sensors, respectively. [The CFC, COT, Re, AOD, ASY, SSA, ALB and WV datasets](#)
8 [where chosen so as to be consistent with the CM SAF SSR dataset.](#) –The potential
9 contribution of various parameters to the RegCM4-CM SAF SSR differences is estimated
10 with the combined use of the data mentioned above and a radiative transfer model [for the](#)
11 [MSG SSR period \(2006-2009\)](#). The results are presented in Sect. 3.6., while the main findings
12 of this manuscript are summarized in Sect.4.

13

14 **2 Model description, data and methods**

15 **2.1 RegCM4 description and simulation setup**

16 In this work, a decadal (2000-2009) simulation was implemented with RegCM4.4 (hereafter
17 denoted as RegCM4 or RegCM) for the greater European region with an horizontal resolution
18 of 50 km. The model's domain extends from 65° W to 65° E and 15° N to 75° N including the
19 largest part of the Sahara Desert and part of Middle East (see Fig. S1 in the Supplement of
20 this manuscript). RegCM is a hydrostatic, sigma-p regional climate model with a dynamical
21 core based on the hydrostatic version of NCAR-PSU's Mesoscale Model version 5 (MM5)
22 (Grell et al., 1994). Specifically, RegCM4 is a substantially improved version of the model
23 compared to its predecessor RegCM3 (Pal et al., 2007) by means of software code and
24 physics (e.g. radiative transfer, planetary boundary layer, convection schemes over land and
25 ocean, land types and surface processes, ocean-air exchanges). Details on the historical
26 evolution of RegCM from the late 1980s until today and a full description of RegCM4's basic
27 features are given in Giorgi et al. (2012).

28 Data from ECMWF's ERA-Interim reanalysis were used as lateral boundary conditions.
29 RegCM4 through a simplified aerosol scheme accounts for anthropogenic SO₂, sulfates,
30 organic and black carbon (Solmon et al., 2006). The emissions of these anthropogenic
31 aerosols are based on monthly [timed-dependent](#), historical emissions from the Coupled

1 Model Intercomparison Project Phase 5 (CMIP5) (Lamarque et al., 2010) with one year spin
2 up time (1999). This inventory is used by a number of climate models in support of the most
3 recent report of the Intergovernmental Panel on Climate Change (IPCC, 2013). The model
4 also accounts for maritime particles through a 2-bin sea salt scheme (Zakey et al., 2008) and
5 for dust through a 4-bin approach (Zakey et al., 2006). For our simulation, the MIT-Emanuel
6 convection scheme (Emanuel, 1991; Emanuel and Zivkovic-Rothman, 1999) was used.
7 Convection is triggered when the buoyancy level is higher than the cloud base level. The
8 cloud mixing is considered episodic and inhomogenous, the convective fluxes being based on
9 a model of sub-cloud-scale updrafts and downdrafts (Giorgi et al., 2012). Zanis et al. (2009)
10 reported for RegCM3 that the low stratiform clouds are systematically denser and more
11 persistent with the use of the Grell (Grell, 1993) convective scheme than with the Emanuel
12 scheme, a result with major importance for the cloud- radiation feedback. The boundary layer
13 scheme of Holtslag et al. (1990) was utilized while the Subgrid Explicit Moisture Scheme
14 (SUBEX) handles large-scale cloud and precipitation computations. The ocean flux scheme
15 was taken from Zeng et al. (1998) with the Biosphere-Atmosphere Transfer Scheme (BATS)
16 (Dickinson et al., 1993) accounting for land surface processes.

17 The Community Climate Model version 3 (CCM3) (Kiehl et al., 1996) radiative package
18 handles radiative transfer within RegCM4. The CCM3 scheme employs the δ -Eddington
19 approximation following its predecessor (CCM2) (Briegleb, 1992). Especially for the
20 shortwave radiation, the radiative transfer model takes into account the effect of atmospheric
21 water vapor and greenhouse gasses, aerosol amount and optical properties per layer (e.g.
22 aerosol optical thickness, asymmetry factor, single scattering albedo) as well as cloud
23 macrophysical (e.g. cloud fractional cover) and microphysical properties per layer (e.g.
24 effective droplet radius, liquid water path, cloud optical thickness) and land surface properties
25 (surface albedo). The radiative transfer equation is solved for 18 discrete spectral intervals
26 from 0.2 to 5 μm for the 18 RegCM vertical sigma layers from 50 hPa to the surface.

27 The effect of clouds on shortwave radiation is manifested by CFC, cloud droplet size and
28 cloud water path (CWP) which is based on the prognostically calculated parameter of cloud
29 water amount (Giorgi et al., 2012). Within the model, the effective droplet radius for liquid
30 clouds (Rel) is considered constant (10 μm) over the ocean while over land it is given as a
31 function of temperature (Kiehl et al., 1998; Collins et al., 2004). On the other hand, the ice

1 particle effective radius (Re_i) is given as a function of normalized pressure, starting from 10
 2 μm . The equations used for the calculation of Re_l and Re_i are given below.

3

$$4 \quad Re_l = \begin{cases} 5 \mu m & T > -10^\circ C \\ 5 - 5 \left(\frac{T+10}{20} \right) \mu m & -30^\circ C \leq T \leq -10^\circ C \\ Re_i & T < -30^\circ C \end{cases} \quad (1)$$

5

$$6 \quad Re_i = \begin{cases} Re_{i_{min}} & p / p_s > p_I^{high} \\ Re_{i_{min}} - (Re_{i_{max}} - Re_{i_{min}}) \left[\frac{(p / p_s) - p_I^{high}}{p_I^{high} - p_I^{low}} \right] \mu m & p / p_s \leq p_I^{high} \end{cases} \quad (2)$$

7

8 where $Re_{i_{max}}=30 \mu m$, $Re_{i_{min}}=10 \mu m$, $p_I^{high}=0.4$ and $p_I^{low}=0.0$.

9 The fraction (f_{ice}) of cloud water that consists of ice particles is given as a function of
 10 temperature (T), the fraction (f_{liq}) of the liquid water droplets being calculated as $f_{liq}=1-f_{ice}$.

11

$$12 \quad f_{ice} = \begin{cases} 0 & T > -10^\circ C \\ -0.05(T+10) & -30^\circ C \leq T \leq -10^\circ C \\ 1 & T < -30^\circ C \end{cases} \quad (3)$$

13

14 Then, the radiative properties of liquid and ice clouds in the shortwave spectral region are
 15 given by the following parameterizations, originally found in Slingo (1989) and revisited by
 16 Briegleb et al. (1992).

17

$$18 \quad COI_{ph}^\lambda = CWP \left[a_{ph}^\lambda + \frac{b_{ph}^\lambda}{Re_{ph}} \right] f_{ph} \quad (4)$$

$$19 \quad \omega_{ph}^\lambda = 1 - c_{ph}^\lambda - d_{ph}^\lambda Re_{ph} \quad (5)$$

$$1 \quad g_{ph}^{\lambda} = e_{ph}^{\lambda} + f_{ph}^{\lambda} \text{Re}_{ph} \quad (6)$$

$$2 \quad \varphi_{ph}^{\lambda} = (g_{ph}^{\lambda})^2 \quad (7)$$

3

4 where superscript λ denotes the spectral interval and subscript ph denotes the phase
 5 (liquid/ice). Also, ω is the single scattering albedo, g is the asymmetry factor and φ is the
 6 phase function of clouds. It has to be highlighted here that all the equations presented above
 7 are given in Kiehl et al. (1998) and Collins et al. (2004) with a slightly different annotation.
 8 The coefficients a-f for liquid clouds are given in Slingo (1989), while for ice clouds in Ebert
 9 and Curry (1992) for the four pseudo-spectral intervals (0.25-0.69, 0.69-1.19, 1.19-2.38 and
 10 2.38-4.00 μm) employed in the radiative scheme of RegCM. Especially for COT, in this paper
 11 we calculated it for the spectral interval 0.25-0.69 μm for both liquid and ice clouds so as to
 12 be comparable to the CM SAF satellite retrieved COT at 0.6 μm (see Sect. 2.2.). Following
 13 the approach of Cess (1985), to derive the bulk COT for the whole atmospheric column, the
 14 COTs calculated for each layer are simply added. The total COT for each layer is calculated
 15 by merging the COT values for liquid and ice clouds.

16 Within RegCM, CFC at each layer is calculated from relative humidity and cloud droplet
 17 radius. The surface radiation flux in RegCM4 is calculated separately for the clear and cloud
 18 covered part of the sky. The total CFC for each model grid-cell is an intermediate value
 19 between the one calculated using the random overlap approach, which leads to a maximum
 20 cloud cover, and the one found by assuming a full overlap of the clouds appearing in different
 21 layers, which minimizes cloud cover. As discussed in Giorgi et al. (2012), this approach
 22 allows for a more realistic representation of surface radiative fluxes.

23 ~~Overall, as discussed in the introduction, for the scopes of this research we used the following~~
 24 ~~parameters from a RegCM4 simulation: 1) SSR, 2) CFC, 3) COT, 4) Re, 5) AOD, 6) ASY, 7)~~
 25 ~~SSA, 8) WV and 9) ALB. The same parameters were extracted from satellite-based~~
 26 ~~observational data (CM SAF, CERES), data from an aerosol climatology (MACv1) and data~~
 27 ~~from the ERA-Interim reanalysis.~~

28 **2.2 CM SAF satellite data**

29 To evaluate the RegCM4 SSR simulations described previously, we use high resolution satellite
 30 data from the SIS (Surface Incoming Shortwave radiation) product of CM SAF. The datasets

1 were obtained from EUMETSAT's ~~Meteosat First and Second Generation~~ (MFG
2 ([DOI:10.5676/EUM_SAF_CM/RAD_MVIRI/V001](https://doi.org/10.5676/EUM_SAF_CM/RAD_MVIRI/V001)) and MSG ([DOI:10.5676/EUM_SAF_CM/](https://doi.org/10.5676/EUM_SAF_CM/CLAAS/V001)
3 [CLAAS/V001](https://doi.org/10.5676/EUM_SAF_CM/CLAAS/V001)) ~~respectively~~) geostationary satellites. SSR data are available from 1983 to 2005
4 from six Meteosat First Generation satellites (Meteosat 2-7) and from 2005 onwards from
5 Meteosat Second Generation satellites (Meteosat 8-10). These satellites fly at an altitude of
6 ~36000 km, being located at longitudes around 0° above the equator and covering an area
7 extending from 80° W to 80° E and from 80° S to 80° N. In the case of MFG satellites, the SSR
8 data are retrieved from measurements with the Meteosat Visible and Infrared Instrument
9 (MVIRI) sensor. MVIRI is a radiometer that takes measurements at 3 spectral bands (visible,
10 water vapor, infrared) every 30 minutes. SSR is retrieved using MVIRI's broadband visible
11 channel (0.45-1 μm) only, at a spatial resolution of ~2.5 km (at the sub-satellite point). The data
12 are afterwards re-gridded at a 0.03° x 0.03° regular grid.

13 The MagicSol-Heliosat algorithm, used for the derivation of the SSR data analyzed in this work,
14 has been extensively described in several papers (see Posselt et al., 2011a,b; Mueller et al.,
15 2011; Posselt et al., 2012; Sanchez-Lorenzo et al., 2013; Posselt et al., 2014). The algorithm
16 includes a modified version of the original Heliosat method (Beyer et al., 1996; Cano et al.,
17 1986). Heliosat utilizes the digital counts obtained from the visible channel to calculate the so-
18 called effective cloud albedo. The modified version incorporates the determination of the
19 monthly maximum normalized digital count (for each MVIRI sensor) that serves as a self-
20 calibration parameter. To derive the clear-sky background reflection, a 7-day running average of
21 the minimum normalized digital counts is used instead of fixed monthly mean values. This
22 method minimizes changes appearing in the radiance data recorded by different MVIRI sensors
23 due to the transition from the one Meteosat satellite to the other, ensuring an as much as
24 possible homogeneous dataset. Then, the clear-sky irradiances are derived using the look-up-
25 table based clear-sky model MAGIC (Mueller et al., 2009) and finally SSR is retrieved by
26 combining them with the effective cloud albedo.

27 On the other hand, MSG satellites carry the Spinning Enhanced Visible and Infrared Imager
28 (SEVIRI), a radiometer taking measurements at 12 spectral bands (from visible to infrared)
29 every 15 minutes with a spatial resolution of ~3 km (at the sub-satellite point). The data used
30 here are available at a 0.05° x 0.05° regular grid. The SEVIRI broadband high-resolution visible
31 channel (HRV) which is very close to MVIRI's broadband visible channel cannot be used for
32 the continuation of the SSR dataset, since, unlike MVIRI, it does not cover the full earth's disk.

1 On the other hand, the use of one of the SEVIRI's narrow band visible channels directly in the
2 same algorithm as MVIRI (MagicSol) is not feasible, first of all, because of the spectral
3 differences with MVIRI's broadband visible channel, and second, because of the sensitivity of
4 cloud albedo to spectral differences of the land surfaces below the clouds (especially for
5 vegetated areas) (see Posselt et al., 2011a; 2014). In this case, an artificial SEVIRI broadband
6 visible channel that corresponds to MVIRI's broadband visible channel is simulated following
7 the approach of Cros et al. (2006). SEVIRI's two narrow band visible channel (0.6 μm and 0.8
8 μm) and MVIRI's broadband channel spectral characteristics are used to establish a simple
9 linear model. This model is afterwards applied to SEVIRI's 0.6 μm and 0.8 μm radiance
10 measurements to calculate the broadband visible channel radiance (see Posselt et al., 2014 for
11 more details).

12 The CM SAF SSR satellite-based product is characterized by a threshold accuracy of 15 W/m^2
13 for monthly mean data and 25 W/m^2 for daily data (Mueller et al., 2011; Posselt et al., 2012;
14 Sanchez-Lorenzo et al., 2013; Posselt et al., 2014). Posselt et al. (2012) evaluated CM SAF SSR
15 data on a daily and monthly basis against ground-based observations from 12 BSRN (Baseline
16 Surface Radiation Network) stations around the world, showing that both daily and monthly
17 CM SAF data are below the target accuracy for $\sim 90\%$ of the stations. Specifically for Europe,
18 Sanchez-Lorenzo et al. (2013) using monthly SSR data from 47 GEBA (Global Energy Balance
19 Archive) ground stations proceeded to a detailed validation of the CM SAF SSR dataset for the
20 period 1983-2005. They found that CM SAF slightly overestimates SSR by 5.2 W/m^2 (4.4% in
21 relative values). Also, the mean absolute bias was found to be 8.2 W/m^2 which is below the
22 accuracy threshold of 15 W/m^2 (10 W/m^2 for the CM SAF retrieval accuracy and 5 W/m^2 for the
23 surface measurements uncertainties). Applying the Standard Normal Homogeneity Test (SNHT)
24 Sanchez-Lorenzo et al. (2013) revealed that the MFG SSR data [over Europe](#) can be considered
25 homogeneous for the period 1994-2005. Recently, Posselt et al. (2014) verified the results of the
26 previous two studies by using a combined MFG-MSG SSR dataset spanning from 1983 to 2010.
27 They found that the monthly mean dataset exhibits a mean absolute bias of 8.15 W/m^2
28 compared to BSRN which is again below the accuracy threshold of CM SAF. Also, the dataset
29 was found to be homogeneous for the period 1994-2010 in most of the investigated regions
30 except for Africa.

31 To investigate the differences appearing between the RegCM4 and CM SAF SSR fields we also
32 use CFC, COT and Re CM SAF observations from MSG satellites for the period 2004-2009. A

1 description of this cloud optical properties product, also known as CLAAS (CLoud property
 2 dAtAset using SEVIRI), can be found in Stengel et al. (2014). The MSG NWC software
 3 package v2010 is used for the detection of cloudy pixels, the determination of their type
 4 (liquid/ice) and their vertical placement (Derrien and Le Gléau, 2005; NWCSAF, 2010). The
 5 detection of cloudy pixels is based on a multispectral threshold method incorporating
 6 parameters such as illumination (e.g. daytime, twilight, night-time, sunglint) and type of
 7 surface. According to Kniffka et al. (2014), the CM SAF Cloud Mask accuracy is ~90%
 8 (successful detection of cloudy pixels for ~90% of the cases) when evaluated against satellite
 9 data from CALIOP/CALIPSO and CPR/CloudSat. The bias of the CFC product was found to be
 10 2% and 3% [for SEVIRI's disk](#) when compared to ground-based data from SYNOP (lidar-radar
 11 measurements) and satellite-based data from MODIS, respectively (Stengel et al., 2014). The
 12 Cloud Physical Properties (CPP) algorithm (Roebeling et al., 2006; Meirink et al., 2013) is used
 13 to retrieve COT at 0.6 μm , Re and CWP. The algorithm is based on the use of SEVIRI's
 14 spectral measurements at the visible (0.64 μm) and near infrared (1.63 μm) (Nakajima and
 15 King, 1990). First, COT and Re are retrieved for the cloudy pixels and then CWP is given by
 16 the following equation:

$$17$$

$$18 \quad \text{CWP}_{\text{ph}} = \frac{2}{3} \rho_{\text{ph}} \text{Re}_{\text{ph}} \text{COT}_{\text{ph}} \quad (8)$$

$$19$$

20 where ph stands for the clouds' phase (liquid/ice) and ρ is the density of water. According to
 21 Stengel et al. (2014), the CM SAF COT bias was estimated at -9.9% compared to MODIS
 22 observations. The corresponding bias for CWP is -0.3% for liquid phase clouds and -6.2% for
 23 ice phase clouds. COT and CWP data are available from CM SAF at a spatial resolution of
 24 $0.05^\circ \times 0.05^\circ$ on a daily basis. In this work, Re values were calculated from the COT and CWP
 25 CM SAF available data using Eq. (8).

26 **2.3 Other data**

27 In addition to the CM SAF SSR and cloud optical properties data used for the evaluation of
 28 RegCM4, we also use ancillary data from other sources, namely, AOD, ASY and SSA at 550
 29 nm monthly climatological values from the MACv1 climatology (Kinne et al., 2013), monthly
 30 climatological broadband surface shortwave fluxes retrieved from CERES sensors aboard EOS
 31 TERRA and AQUA satellites for a 14-year period starting from 3/2000 (Kato et al., 2013) and

1 finally monthly mean total column WV data from ECMWF's ERA-Interim reanalysis (Dee et
 2 al., 2011) for the period 2006-2009. All the data were obtained at a spatial resolution of $1^\circ \times 1^\circ$.
 3 It has to be highlighted that these data are similar to the ones used as input within the MAGIC
 4 clear sky radiative transfer code (Mueller et al., 2009) which is used for the calculation of CM
 5 SAF SSR. Therefore, they can be used in order to examine the reasons for possible deviations
 6 appearing between RegCM4 and CM SAF SSR (see Sect. 2.4.). To our knowledge, the
 7 uncertainty of the MACv1 aerosol parameters used here has not been reported somewhere in
 8 detail. However, due to the methodology followed for the production of the MACv1
 9 climatology, the MACv1 data are consistent with the AERONET ground network. The CERES
 10 broadband surface albedo over land exhibits a relative bias of -2.4% compared to MODIS.
 11 Specifically, over deserts, the relative bias drops to -2.1% (Rutan et al., 2009). A detailed
 12 evaluation of the ERA-Interim WV total column product does not exist. Only recently, the
 13 upper troposphere – lower stratosphere WV data were evaluated against airborne campaign
 14 measurements showing a good agreement (30% of the observations were almost perfectly
 15 represented by the model) (Kunz et al., 2014).

16

17 **2.4 Methodology**

18 In this study, first, the RegCM4 SSR fields are evaluated against SSR fields from CM SAF
 19 (MFG for 2000-2005 and MSG for 2006-2009) for the European region (box region in Fig. S1).
 20 Prior to the evaluation, the model and satellite data are averaged on a monthly basis and brought
 21 to a common $0.5^\circ \times 0.5^\circ$ spatial resolution. It has to be mentioned that the same temporal and
 22 spatial resolution was used for all the data utilized in this study. Maps with the normalized mean
 23 bias (NMB) (hereafter denoted as bias) are produced on an annual and seasonal basis. NMB is
 24 given by the following equation:

25

$$26 \quad NMB = \frac{\sum_{i=1}^N (\text{RegCM}_i - \text{CMSAF}_i)}{\sum_{i=1}^N \text{CMSAF}_i} 100\% = \left(\frac{\overline{\text{RegCM}}}{\overline{\text{CMSAF}}} - 1 \right) 100\% \quad (9)$$

27

1 where $RegCM_i$ and $CMSAF_i$ represent the RegCM4 and CM SAF mean values for each month
2 i , N is the number of months and \overline{RegCM} , \overline{CMSAF} are the RegCM4 and CM SAF mean
3 values. The statistical significance of the results at the 95% confidence level is checked by
4 means of a two independent sample t-test:

$$6 \quad t = (\overline{RegCM} - \overline{CMSAF}) / \sqrt{(\sigma_{RegCM}^2 + \sigma_{CMSAF}^2) / N} \quad (10)$$

7
8 where σ_{RegCM} and σ_{CMSAF} are the standard deviations of RegCM4 and CM SAF total means.
9 When $|t|$ is greater than a critical value that depends on the degrees of freedom (here $2n-1$) the
10 bias is considered statistically significant. In addition to the whole European region (EU), the
11 land covered (LA) and ocean covered (OC) part of Europe, seven other sub-regions are defined
12 for the generalization of our results: Northern Europe (NE), Central Europe (CE), Eastern
13 Europe (EE), Iberian Peninsula (IP), Central Mediterranean (CM), Eastern Mediterranean (EM)
14 and Northern Africa (NA) (see Figs. 1a and S1). The bias on an annual and seasonal basis is
15 calculated per region. Apart from bias, other statistical metrics (correlation coefficient R ,
16 normalized standard deviation NSD, modified normalized mean bias MNMB, root mean square
17 error RMSE) are also defined, calculated and presented in the electronic supplement of this
18 manuscript. The latitudinal variability of model and satellite-based SSR is examined by means
19 of seasonal plots. Finally, the seasonal variability of SSR from RegCM4 and CM SAF and their
20 differences is investigated for each of the 10 regions mentioned above. The same procedure is
21 done separately for MFG data (2000-2005) and MSG data (2006-2009) to see if the two datasets
22 lead to similar results. Our results are mostly focused on MSG satellite-based observations,
23 since CFC and cloud optical properties data are only available from MSG SEVIRI.

24 In order to interpret the observed differences between RegCM4 and CM SAF SSR, the same
25 detailed procedure is repeated for CFC and COT for the period 2004-2009. CFC and COT are
26 the two major determinants of the transmission of shortwave radiation through clouds (Gupta et
27 al., 1993) and along with AOD constitute the major controllers of SSR (Kawamoto and
28 Hayasaka, 2008). Therefore, we also proceed to a detailed comparison of RegCM4 AOD at 550
29 nm (AOD_{550}) against MACv1 climatological data. However, other cloud (Re) and aerosol
30 (ASY, SSA) related parameters also play a significant role. Here, RegCM4 Re is evaluated
31 against observational data from CM SAF while RegCM4 ASY and SSA are compared against

1 climatological data from MACv1 (see Supplement). Specifically, the comparison of RegCM4
2 data with MACv1 does not constitute an evaluation of the RegCM4 aerosol-related parameters,
3 like in the case of the cloud-related parameters above, since, MACv1 data (Kinne et al., 2013)
4 are climatological (based on a combination of models and observations) and not pure
5 observational data. However, a similar climatology (Kinne et al., 2006) is used for the
6 production of CM SAF SSR (Trentmann et al., 2013). In addition, Mueller et al. (2014) showed
7 that the use of MACv1 aerosol climatology instead of the Kinne et al. (2006) climatology does
8 not affect significantly the CM SAF SSR product. Hence, this comparison allows us to reach
9 useful conclusions about the effect of aerosol representation within RegCM4 on the simulated
10 SSR fields by the model. The same stands for the comparison of RegCM4 ALB data with
11 climatological data from CERES satellite sensors and RegCM4 WV data with WV data from
12 ERA-Interim reanalysis (see Supplement). The CERES ALB 14-year climatology is temporally
13 constant, similar to the CERES climatology used for the production of CM SAF SSR
14 (Trentmann et al., 2013). Finally, the ERA-Interim WV data used here are the same with the
15 WV data incorporated by the radiative scheme of CM SAF. Unlike the RegCM4 evaluation
16 results, the comparison results discussed in this paragraph are presented in the Supplement.

17 Apart from a qualitative approach, we also proceed to a quantitative study of the reasons that
18 lead to deviations between the RegCM4 and CM SAF SSR. Using data from RegCM4 and CM
19 SAF and the Santa Barbara DISORT Atmospheric Radiative Transfer (SBDART) model
20 (Ricchiazzi et al., 1998), we estimate the potential relative contribution of the parameters CFC,
21 COT, Re, AOD, ASY, SSA, ALB and WV to the percent RegCM4-CM SAF SSR difference
22 (Δ SSR), over the 7 sub-regions mentioned above. Δ SSR is given by Eq. (11), expressing the
23 percentage of SSR deviation caused by the observed difference between RegCM4 and CM SAF
24 for each parameter (p). First, a SBDART simulation is implemented with a 3-hour timestep for
25 the 15th day of each month (Ming et al., 2005) using monthly mean RegCM4 data as input
26 (control run) for each region. The average of all the timesteps per month expresses the monthly
27 SSR flux (SSR_{control}). The SSR fields simulated with SBDART are almost identical to the
28 RegCM4 SSR fields. This indicates that SBDART indeed can be used to study the sensitivity of
29 RegCM4's radiative scheme to various parameters. Then, several SBDART simulations are
30 implemented in the same way, replacing each time only one of the aforementioned input
31 parameters with corresponding values from CM SAF, MACv1 or ERA-Interim ($SSR(p)$).
32 SSR_{control} and $SSR(p)$ are then used in Eq. (11) to calculate Δ SSR for each month (i) and
33 parameter (p).

1

$$\Delta SSR^i(p) = 100(SSR_{control}^i - SSR^i(p)) / SSR_{control}^i \quad (11)$$

2

3

4 The results of this analysis are presented by means of bar plots for each sub-region. In addition,
5 a method like the one introduced by Kawamoto and Hayasaka (2008, 2010, 2011), which is
6 based on the calculation of the sensitivities of SSR on CFC, COT, AOD and WV, was also
7 implemented with similar results (not shown here).

8

9 **3 Results and Discussion**

10 **3.1 Surface Solar Radiation**

11 As discussed above, first, we examine the CM SAF and RegCM4 bias patterns for the MFG
12 (2000-2005) and MSG (2006-2009) periods, separately. This work focuses on the MSG
13 dataset, since, cloud properties data which are used in order to investigate the reasons of the
14 observed bias between CM SAF and RegCM4 at a later stage, are only available from MSG.
15 However, we investigate both the periods to examine if the observed biases are valid for the
16 whole simulation period and ensure that there are no differences when using the one or the
17 other dataset. As shown in Fig. S2a and b, the annual bias patterns are similar for both MFG-
18 RegCM4 and MSG-RegCM4. The main feature is a low negative bias over land and a low
19 positive bias over ocean. Overall, the RegCM4 simulations slightly overestimate SSR
20 compared to CM SAF over Europe with a bias of +1.54% in the case of MFG and +3.34% in
21 the case of MSG, while SSR from RegCM4 is much closer to SSR from CM SAF over land
22 (bias of -1.659% for MFG and +0.766% for MSG) than over ocean (bias of +7.29% for MFG
23 and +8.107% for MSG). These values can be found in Table 2+ for the RegCM4-MSG period
24 along with the corresponding values for the 7 sub-regions of interest appearing in Fig. 1a
25 while the same values for the RegCM4-MFG period can be found in Table S1 of the
26 Supplement. It has to be highlighted, that hereafter, only results for the MSG CM SAF SSR
27 dataset are presented within the paper while the results for the MFG dataset are included in
28 the Supplement.

29 As presented in Fig. 1, some differences appear in the seasonal bias patterns. A strong
30 positive bias is observed during winter over Northern Europe. For the rest of the regions the

1 winter patterns are very close to the spring and the annual patterns. Contrary to the annual
2 patterns, in summer, the positive bias extends over Europe until the latitudinal zone of 50°N,
3 while in autumn the bias patterns are pretty similar with the annual ones. In winter, the
4 RegCM4 simulations overestimate SSR compared to CM SAF for the whole European
5 domain, the bias being +3.92%. Over land the bias is nearly zero (+0.14%) while over ocean
6 there is a significant bias of +11.32%. As shown in Fig. 1a, NE is by far the sub-region with
7 the strongest bias (+52.44%). The seasonal and annual model and satellite-derived values with
8 the corresponding biases and their statistical significance at the 95% confidence level
9 according to a two independent sample t-test appear in Table 42. The latitudinal variability of
10 RegCM4 SSR, CM SAF SSR and their difference is presented in Fig. 2a. Overall, RegCM4
11 slightly overestimates SSR at latitudes lower than ~40°N, then a negligible difference between
12 RegCM4 and CM SAF is observed until the latitudinal zone of ~52°N, while, a significant
13 difference is observed for higher latitudes. In spring, a zero bias is observed between the
14 model and CM SAF for Europe. When discriminating between land and ocean covered
15 regions a negative bias is observed over land (-2.90%) and a positive over ocean (+5.218%).
16 The regions with the highest negative bias are NE (-14.21%), EE (-13.545%) and CE (-
17 9.14%), while the regions with the highest positive bias are NA (+8.42%), CM (+7.988%) and
18 EM (+6.72%) (see Table 42). This is also reflected in Fig. 2b where RegCM4 clearly
19 overestimates SSR for latitudes less than ~44°N, significantly underestimating SSR thereafter.
20 In summer, a positive bias of +6.219% is calculated for the whole European domain, the bias
21 being +4.436% over land and +9.437% over ocean. As seen in Table 24, the bias is positive
22 for all the sub-regions ranging from +2.31% (EE) to +10.44% (CM) except for NE (-9.43%).
23 RegCM4 clearly overestimates SSR for latitudes less than ~55°N and underestimates SSR for
24 higher latitudes (Fig. 2c). A positive bias of +2.40% is found for Europe in autumn with the
25 corresponding values being -0.94% over land and +8.44% over ocean covered regions. EE (-
26 9.879%) and CE (-7.219%) are the regions with the strongest negative bias while the regions
27 with the strongest positive bias are the ones at the south, namely, NA (+5.54%), CM
28 (+5.327%) and EM (+5.0498) (see also Table 24). This is also seen in Fig. 2d where
29 RegCM4 overestimates SSR for latitudes less than ~42°N.

30 The seasonal variability of RegCM4 SSR, CM SAF SSR and their difference for the whole
31 European domain, for the land and ocean covered part of Europe as well as for the 7 sub-
32 regions of interest are presented in Figs. 3a-j. For Europe as a whole, the largest difference
33 between RegCM4 and CM SAF SSR is observed in summer, July being the month with the

1 highest RegCM4-CM SAF difference (20.30 W/m²). Over land, the difference between
2 RegCM4 and CM SAF SSR is nearly zero for winter and autumn months. During spring, in
3 March and April, RegCM4 underestimates SSR while in summer SSR is overestimated,
4 especially in July. On the contrary, over ocean, SSR is overestimated by RegCM4 for the total
5 of the months. The highest RegCM4-CM SAF differences are observed during the warm
6 period (May-September). Over NE, RegCM4 underestimates SSR for the months from March
7 to September and overestimates SSR during the winter months. The seasonal variability of the
8 difference between RegCM4 and CM SAF is pretty similar over CE and EE. The simulations
9 underestimate SSR in spring (especially during April) and autumn and overestimate SSR in
10 summer. Over IP, SSR is overestimated again in May and during the summer and
11 underestimated in February, March, November and December. For CM and EM, the seasonal
12 variability of the difference between RegCM4 and CM SAF is almost identical. RegCM4
13 significantly overestimates SSR from April to October while for the rest of the months the
14 difference is nearly zero. Finally, over NA, the seasonal variability of the difference is close
15 to the one appearing over CM and EM, but here, SSR is overestimated by RegCM4 also in
16 March.

17 **3.2 Cloud Fractional Cover**

18 CFC plays a determinant role as far as SSR levels are concerned. Therefore, we compare the
19 CFC patterns simulated with RegCM4 against CFC patterns from MSG CM SAF for the
20 common period 2004-2009. Overall, CFC is underestimated by RegCM4 over Europe by
21 24.34% on annual basis (13.74% over land and 38.41% over ocean) despite the fact that over
22 specific regions (e.g. within IP and NA) CFC is overestimated (see Table 32).
23 Underestimation is observed for the total of the four seasons, NA being the only region with a
24 bias of +8.12% in winter and a bias of +13.14% in autumn (see Table S3). As shown in Figs.
25 4a-d, the underestimation of CFC from RegCM4 is stronger over ocean especially in summer,
26 while strong overestimation is observed over regions in western NA in winter and spring,
27 eastern NA in summer and the whole NA during autumn. The latitudinal variability of
28 RegCM4 CFC, CM SAF CFC and their difference is presented in Fig. 5. A clear, strong
29 underestimation of CFC from RegCM4 is observed for all the latitudinal bands and seasons
30 apart from latitudes around 30° N where CFC is slightly overestimated in autumn. The
31 seasonal variability of RegCM4 CFC, CM SAF CFC and their difference for the whole
32 European domain, for the land and ocean covered part of Europe and for the 7 sub-regions of

1 interest are presented in Figs. 6a-j. CFC is underestimated steadily by RegCM4 throughout a
2 year, the underestimation being much stronger over the ocean than over land (see Figs. 6b and
3 c). This underestimation is observed for all the sub-regions except for NA where CFC is
4 underestimated from April to September and overestimated for the rest of the months.

5 Generally, lower CFCs would lead to higher SSR levels. However, a comparison of the SSR
6 bias patterns appearing in Figs. 1a-d with the CFC bias patterns appearing in Figs. 4a-d and
7 also of the biases appearing in Table 1 and Table S3 reveals that for some areas and seasons
8 the RegCM4-CM SAF SSR deviations cannot be explained through the corresponding CFC
9 deviations (e.g. land covered regions during spring and autumn). This is in line with the
10 findings of Katragkou et al. (2015) where the WRF-ISCCP SSR deviations could not always
11 be attributed to CFC deviations. As discussed there the role of microphysical cloud properties
12 should also be taken into account. Following this, in the next paragraph we go a step further,
13 taking into account the effect of COT.

14 **3.3 Cloud Microphysical Properties**

15 **3.3.1 Cloud Optical Thickness**

16 COT is a measure of the transparency of clouds and along with CFC determines the
17 transmission of shortwave radiation through clouds (Gupta et al., 1993). In this paragraph, the
18 RegCM4 COT patterns are compared against COT patterns from MSG CM SAF for the
19 common period 2004-2009. Overall, COT is overestimated by RegCM4 over Europe by
20 4.329% on annual basis, the bias being positive over land (+7.329%) but negative over ocean
21 (-2.546%) (see Table 23). In addition, COT bias varies with seasons, being positive in spring
22 and autumn and negative in winter and summer (see Table S5). As shown in Figs. 7a-d,
23 positive biases are mostly observed over land covered regions of CE, EE and NE and negative
24 biases over NA and the regions around the Mediterranean Sea. In fact, there is a strong
25 latitudinal variability of the RegCM4-CM SAF COT difference for all the seasons as
26 presented in Figs. 8a-d. RegCM4 underestimates COT for latitudes below $\sim 45^\circ$ N in winter,
27 spring and autumn and for latitudes below $\sim 50^\circ$ N in summer. The seasonal variability of
28 RegCM4 COT, CM SAF COT and their difference for the whole European domain, for the
29 land and ocean covered part of Europe and for the 7 sub-regions of interest are presented in
30 Figs. 9a-j. In general, the RegCM4-CM SAF COT difference is not steadily positive or
31 negative but varies from month to month over both land and ocean. RegCM4 steadily

1 overestimates COT throughout a year only over NE and underestimates COT over CM and
2 NA. It has to be highlighted that there are no COT retrievals over NE for December and
3 January due to a limited illumination at that latitudes during this period of the year. This is
4 also the reason for the missing grid cells appearing in the top-right corner of Figs. 7a-d.

5 A comparison of the SSR bias patterns appearing in Figs. 1a-d with the CFC (Figs. 4a-d) and
6 the COT (Figs. 7a-d) bias patterns reveals that COT could explain part of the RegCM4-CM
7 SAF SSR deviations that could not be explained through CFC (e.g. NE, CE, EE). The same
8 conclusions can be reached by comparing the seasonal variability of SSR, CFC and COT over
9 the region of interest (see Figs. 3, 6 and 9). However, other parameters are expected to be
10 responsible for the remaining unexplained RegCM4-CM SAF SSR deviation.

11 **3.3.2 Cloud Effective Radius**

12 R_e is a microphysical optical property expressing the size of cloud droplets in the case of
13 liquid clouds and the size of ice crystals in the case of ice clouds. R_e of liquid (R_{el}) and ice
14 (R_{ei}) clouds plays a critical role in the calculation of the optical thickness of clouds as well as
15 their albedo (see Eqs. 4-7 in Sect. 2.1.). The evaluation of RegCM4 R_{el} and R_{ei} against
16 observational data from CM SAF reveals a significant underestimation over the whole
17 European domain (bias of -36.106% for R_{el} and -28.325% for R_{ei}). In the case of ice clouds,
18 the biases over land and ocean do not differ significantly. On the contrary, for liquid clouds,
19 the bias over land is more than double the bias over ocean (see Table 23). This is due to the
20 very low RegCM4 R_{el} values appearing over land while the CM SAF dataset does not exhibit
21 such a land-ocean difference. A possible explanation for this could be the fact that for liquid
22 clouds a different approach is used over land (constant R_{el} of 10 μm) and ocean (Eq. 1) while
23 for ice clouds the parameterization is the same for land and ocean (Eq. 2). The fact that the
24 average R_{el} value over land ($5.65 \pm 1.06 \mu\text{m}$) is very close to the lowest R_{el} boundary (5 μm)
25 according to Eq. (1), possibly points towards an underestimation of the liquid cloud height
26 and vertical development. Also, this R_{el} land-ocean difference is in charge of the COT land-
27 ocean difference (see Table 23) according to Eq. (4). In general, the underestimation of R_e
28 would result into more reflective clouds and hence into underestimated SSR levels. It has to
29 be mentioned here that the monthly variability of RegCM4 R_{el} and R_{ei} , CM SAF R_{el} and R_{ei}
30 and their difference for the whole European domain, for the land and ocean covered part of
31 Europe and for the 7 sub-regions are presented in the Supplement of this manuscript. [A](#)

1 [constant underestimation of Rel and Rei is observed for the whole Europe \(see Figs. S6 and](#)
2 [S8\).](#)

3 **3.4 Aerosol Optical Properties**

4 As discussed in Sect. 2.4., AOD along with CFC and COT constitute the major controllers of
5 SSR. A comparison of the RegCM4 AOD₅₅₀ seasonal patterns with climatological AOD₅₅₀
6 values from MACv1 is presented in Figs. S10a-d. On an annual basis, RegCM4 overestimates
7 AOD over the region of NA (bias of +25.0±%) (see Table 23). The overestimation is very
8 strong during winter being much weaker in spring and autumn (see Table S9). This
9 overestimation over regions affected by dust emission has been discussed comprehensively in
10 Nabat et al. (2012) and has to do with the dust particle size distribution schemes utilized by
11 RegCM4 (Alfaro and Gomes, 2001; Kok, 2011). Nabat et al. (2012) showed that the
12 implementation of Kok (2011) scheme generally returns AODs closer to that of MODIS
13 within the Mediterranean Basin. However, a first climatological comparison of RegCM4 dust
14 AODs with data from CALIOP/CALIPSO (A. Tsikerdekis, personal communication, 2015)
15 has shown that both schemes overestimate dust AOD over Europe and therefore the selection
16 of a specific dust scheme is not expected to change drastically our results. On the contrary,
17 AOD is significantly underestimated over the rest of the domain. This should be expected as
18 RegCM does not account for several types of aerosols, anthropogenic (e.g. nitrates,
19 ammonium and secondary organic aerosols, industrial dust) and natural (e.g. biogenic
20 aerosols) which potentially play an important role (Kanakidou et al., 2005; Zanis et al., 2012).
21 This overestimation/underestimation dipole in winter, spring and autumn is also reflected in
22 Fig. S11. RegCM4 overestimates AOD for latitudes below ~40° N in winter, for latitudes
23 below ~35° N in spring and for a narrow latitudinal band (~30-33° N) in autumn. In summer,
24 RegCM4 steadily underestimates AOD compared to MACv1. The seasonal variability of
25 RegCM4 AOD₅₅₀, MACv1 AOD₅₅₀ and their difference for the whole European domain, for
26 the land and ocean covered part of Europe and for the 7 sub-regions of interest are presented
27 in Figs. S12a-j. In general, RegCM4 clearly underestimates AOD throughout a year over
28 regions that are not affected heavily by Sahara dust transport. This underestimation would
29 cause an overestimation of SSR if all the other parameters were kept constant. The opposite
30 stands for the region of NA where AOD, except for summer, is significantly overestimated.

31 As in the case of COT and Re, in order to fully assess the contribution of aerosols to the
32 observed RegCM4-CM SAF SSR deviations, one has to take into account ASY and SSA

1 apart from AOD. A comparison of RegCM4 ASY and SSA with climatological values from
2 MACv1 reveals a small underestimation from RegCM4 over Europe (bias of -1.107% and -
3 4.23% respectively). While SSA is underestimated for the total of the investigated sub-
4 regions, in some cases ASY is slightly overestimated (see Table 23). [This is apparent in Figs.](#)
5 [S13 and S15 where the RegCM4-CM SAF NMB maps are presented along with the](#)
6 [latitudinal variability of the two products.](#)

7

8 **3.5 Other parameters**

9 Apart from the major (CFC, COT, AOD) and minor (Re, ASY, SSA) SSR determinants
10 which are discussed above in detail, there are also a number of other parameters that could
11 impact the simulation skills of RegCM4 compared to CM SAF, since these parameters are
12 used as input within the radiative scheme of the model.

13 As it was previously discussed, WV is another parameter that affects the transmission of solar
14 radiation within the atmosphere. RegCM4 is found here to overestimate WV compared to
15 ERA-Interim reanalysis all over Europe with a bias of ~12%. This becomes more than
16 obvious when looking into the seasonal and latitudinal variability of the two datasets (see
17 Figs. S17 and S18).

18 In line with the study of Güttler et al. (2014), RegCM4 exhibits a significant underestimation
19 of ALB over CE, EE and NA (see Table 23) compared to climatological data from CERES
20 (see Sect. 2.3.). In general, there is a striking difference between land and ocean covered
21 regions (see Fig. S19 and S20). Over land RegCM4 underestimates ALB by 28.327% while
22 over ocean ALB is strongly overestimated by 131%. As it was previously highlighted, the
23 comparisons of RegCM4 with non-observational data presented in this paragraph do not
24 constitute an evaluation of RegCM4. However, these comparisons give us an insight into how
25 several parameters affect the ability of RegCM4 to simulate SSR.

26 **3.6 Assessing the effect of various parameters on RegCM's SSR**

27 As discussed in detail in Sect. 2.4., the contribution of each one of the aforementioned
28 parameters in the deviation between RegCM4 and CM SAF SSR is assessed with the use of
29 SBDART radiative transfer model. The results of this analysis are presented in Fig. 10. The
30 percent contribution of each parameter to the RegCM4-CM SAF SSR difference is calculated

1 on a monthly basis. Results for NE are not included in this manuscript, since COT and Re are
2 not available from CM SAF during winter (December, January) and also due to the low
3 insolation levels for several months at high latitudes.

4 As seen in Fig. 10a, the percent RegCM4-CM SAF SSR difference (Δ SSR) over CE is mostly
5 determined by CFC, COT and AOD. However, for specific months, Re and the other
6 parameters also play an important role leading to an underestimation of SSR. The effect of
7 CFC ranges from a significant SSR underestimation (Δ SSR of -23.56% for April) to a
8 significant SSR overestimation (Δ SSR of +10.04% for June). Apart from July, COT leads to
9 an underestimation of SSR, April being the month with the highest underestimation (Δ SSR of
10 -13.325%). AOD on the other hand, leads to an overestimation of SSR over CE ranging from
11 +4.60% (June) to +9.545% (January).

12 In line with CE, Δ SSR over EE is mostly determined by CFC, COT and AOD (Fig. 10b).
13 Apart from April, CFC leads to an overestimation of SSR, December being the month with
14 the highest overestimation (+22.987%). Apart from June and July, COT causes an
15 underestimation of SSR, March/August being the month with the highest/lowest
16 underestimation (-15.878%/-0.249%). On the other hand, AOD leads to an overestimation of
17 SSR the whole year, December/May being the month with the highest/lowest overestimation
18 (+12.30%/+4.22%). Re also plays a role leading to an underestimation of SSR, that ranges
19 from -1.06% (July) to -2.54% (February). All the other parameters play a minor role,
20 generally leading to an underestimation of SSR.

21 Over IP, despite the fact that the dominant parameters are CFC and COT, for some months
22 AOD, SSA and Re contribute substantially in Δ SSR (Fig. 10c). CFC leads to an
23 overestimation of SSR, January/September being the month with the highest/lowest
24 overestimation of SSR (+9.106%/+1.106%). COT causes an important overestimation of SSR
25 from April to October (e.g. +3.765% in June) and a significant underestimation during March
26 (-2.876%). On the other hand, Re leads to an underestimation of SSR that ranges from -
27 1.327% in April to -0.30% in August. The same stands for SSA with an average annual SSR
28 underestimation of -1.20%, while AOD exhibits a mixed behavior leading to either
29 underestimation (a maximum of -6.14% in December) or overestimation (a maximum of
30 +4.93% in March).

31 As seen in Fig. 10d, Δ SSR over CM is mostly determined by CFC, COT, AOD and SSA.
32 CFC causes a significant overestimation of SSR ranging from +3.23% (July) to +11.94%

1 (December). COT leads to an overestimation of SSR on an annual basis, October being the
2 month with the highest overestimation (+4.60%). AOD causes an overestimation of SSR over
3 CM for the period from March to October (average Δ SSR of +2.21%) and an underestimation
4 during winter (average Δ SSR of -2.31%). SSA on the other hand, causes an underestimation of
5 SSR on an annual basis ranging from -0.54% (July) to -1.988% (December).

6 Δ SSR over EM is dominated by the relative contribution of CFC, AOD and COT (see Fig.,
7 10e). CFC causes an overestimation of SSR on an annual basis ranging from +1.70%
8 (August) to +12.216% (December). Apart from February, AOD causes a significant
9 overestimation ranging from +0.53% (March) to +6.05.99% (September). Apart from March,
10 COT leads to an overestimation of SSR, February being the month with the highest
11 overestimation (+4.327%). SSA also plays a role, in some cases comparable in magnitude to
12 that of COT or AOD (e.g. January, March).

13 Over NA Δ SSR is largely determined by AOD, SSA and COT (Fig. 10f). AOD causes a
14 significant underestimation of SSR during the period from November to April (a maximum of
15 -15.329% for February) and an overestimation from June to September (a maximum of
16 +3.987% for July). COT leads to a significant SSR overestimation on an annual basis ranging
17 from +1.31% (June) to +4.81% (September). SSA leads to a significant underestimation of
18 SSR, January being the month with the highest underestimation (Δ SSR of -3.70%). Important
19 is the contribution of ALB which also causes an SSR underestimation on annual basis
20 (average Δ SSR of -1.04%). It has to be highlighted here, that due to the high insolation levels
21 over the region of NA, the Δ SSR values correspond to higher absolute RegCM4-CM SAF
22 SSR deviations than in regions at higher latitudes. Also, the low cloud coverage in the region
23 leads to an update of the role of aerosol related parameters as shown in Fig. 10f.

24 Concluding, for the total of the six sub-regions, CFC and AOD are the most important factors
25 that determine the SSR overestimation by RegCM4 on an annual basis. The underestimation
26 of CFC and AOD by the model causes an annual overestimation of SSR by 4.8% and 2.6%,
27 respectively.

28

29 **4 Conclusions**

30 In the present study, a decadal simulation (2000-2009) with the regional climate model
31 RegCM4 is implemented in order to assess the model's ability to represent the SSR patterns

1 over Europe. The RegCM4 SSR fields are evaluated against satellite-based observations from
2 CM SAF. The annual bias patterns of RegCM4-CM SAF are similar for both MFG (2000-
3 2005) and MSG (2006-2009) observations. The model slightly overestimates SSR compared
4 to CM SAF over Europe, the bias being +1.5% for MFG and +3.3% for MSG observations.
5 Moreover, the bias is much lower over land than over ocean while some differences appear
6 locally between the seasonal and annual bias patterns.

7 In order to understand the RegCM4-CM SAF SSR deviations, CFC, COT and Re data from
8 RegCM4 are compared against observations from CM SAF (MSG period). For the same
9 reason, AOD, ASY, SSA, WV and ALB from RegCM4 are compared against data from
10 MACv1, ERA-Interim reanalysis and CERES since these data are similar to the ones used as
11 input in the retrieval of CM SAF SSR.

12 CFC is significantly underestimated by RegCM4 compared to CM SAF over Europe by
13 24.3% on annual basis. Part of the bias between REGCM4 and CM SAF SSR can be
14 explained through CFC with the underestimation of CFC leading to a clear overestimation of
15 SSR. It was also found that RegCM4 overestimates COT compared to CM SAF on an annual
16 basis suggesting that COT may explain part of the RegCM4-CM SAF SSR deviations that
17 could not be explained through CFC over specific regions. In addition, RegCM4
18 underestimates significantly Rel and Rei compared to CM SAF over the whole European
19 domain on an annual basis. A comparison of the RegCM4 AOD seasonal patterns with AOD
20 values from the MACv1 aerosol climatology reveals that RegCM4 overestimates AOD over
21 the region of NA and underestimates it for the rest of the European domain. ASY and SSA are
22 slightly underestimated by the model. The comparison of RegCM4 WV against data from
23 ERA-Interim reanalysis, reveals a clear overestimation over Europe. In line with previous
24 studies, RegCM4 underestimates ALB significantly over CE, EE and NA compared to
25 climatological data from CERES with a striking difference between land and ocean.

26 The combined use of SBDART radiative transfer model with RegCM4, CM SAF, MACv1,
27 CERES and ERA-Interim data for the common period 2006-2009 shows that the difference
28 between RegCM4 and CM SAF SSR is mostly explained through CFC, COT and AOD
29 deviations. In the majority of the regions, CFC leads to an overestimation of SSR by
30 RegCM4. In some cases, COT leads to an underestimation of SSR by RegCM4, while for the
31 majority of the regions leads to an overestimation. Apart from NA, where AOD leads to a
32 significant underestimation of RegCM4 SSR, AOD is generally responsible for the

1 [overestimation of SSR. The other parameters \(Re, ASY, SSA, WV and ALB\) play a less](#)
2 [significant role, except for NA where they have a significant impact on the RegCM4-CM](#)
3 [SAF SSR deviations. Overall, CFC and AOD are the major determinants of the SSR](#)
4 [overestimation by RegCM4 on an annual basis. The underestimation of CFC and AOD by the](#)
5 [model causes an annual overestimation of SSR by 4.8% and 2.6%, respectively.](#)
6 [Overall, it is shown in this study that RegCM4 simulates adequately the SSR patterns over](#)
7 [Europe. However, it is also shown that the model overestimates or underestimates](#)
8 [significantly several parameters that determine the transmission of solar radiation in the](#)
9 [atmosphere. The good agreement between RegCM4 and satellite-based SSR observations](#)
10 [from CM SAF is actually a result of the contradicting effect of these parameters. Our results](#)
11 [suggest that there should be a reassessment of the way these parameters are represented](#)
12 [within the model so that SSR is not only well simulated but also for the right reasons. This](#)
13 [would also allow for a safer investigation of the dimming/brightening effect since the SSR](#)
14 [deviations would be safely dedicated to the one or the other parameter. It is suggested here](#)
15 [that a similar approach should be implemented in the future to the same or other regional](#)
16 [climate models with various setups also utilizing new satellite products \(e.g. CM SAF](#)
17 [SARAH\).](#)
18

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32

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1 Table 1. List of the parameters being analyzed in this work, their sources, the original
 2 resolution at which the data were acquired and the corresponding time periods.

Parameter	Source	Resolution	Period
SSR	CM SAF MFG	0.03° x 0.03°	2000-2005
SSR	CM SAF MSG	0.05° x 0.05°	2006-2009
CFC	CM SAF MSG	0.05° x 0.05°	2004-2009
COT	CM SAF MSG	0.05° x 0.05°	2004-2009
Re	CM SAF MSG	0.05° x 0.05°	2004-2009
AOD	MACv1	1° x 1°	Climatology
ASY	MACv1	1° x 1°	Climatology
SSA	MACv1	1° x 1°	Climatology
ALB	CERES	1° x 1°	Climatology
WV	ERA-Interim	1° x 1°	2006-2009
All above	RegCM4	50km x 50km	2000-2009

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1 Table 42. Average RegCM4 SSR and CM SAF SSR (MSG SEVIRI) with their standard
 2 deviations ($\pm 1\sigma$) and the corresponding Normalized Mean Bias (NMB) per season and region.
 3 When the difference between RegCM4 and CM SAF SSR is statistically significant at the
 4 95% confidence level due to a two independent sample t-test, the NMB values are marked
 5 with bold letters while in the opposite case they are marked with an asterisk. Positive NMBs
 6 are marked with red color while negative NMBs with blue. ANN corresponds to annual, DJF
 7 to winter, MAM to spring, JJA to summer and SON to autumn results.

	ANN			DJF			MAM			JJA			SON		
	MOD	SAT	bias	MOD	SAT	bias	MOD	SAT	bias	MOD	SAT	bias	MOD	SAT	bias
EU	175.0±106.5	169.3±96.7	3.3	77.1±57.1	74.2±57.2	3.9	206.8±83.0	206.7±67.0	0.0*	281.6±70.6	265.2±55.2	6.2	126.3±77.4	123.3±71.3	2.4
LA	173.1±106.9	171.9±97.2	0.7	78.1±61.0	78.0±60.8	0.1*	202.7±85.7	208.7±68.6	-2.9	278.6±71.7	267.0±55.0	4.4	124.9±79.0	126.1±72.8	-0.9
OC	178.2±105.6	164.9±95.7	8.1	75.3±49.7	67.7±49.8	11.3	213.8±77.8	203.2±64.2	5.2	286.7±68.2	262.1±55.3	9.4	128.7±74.5	118.6±68.4	8.4
NE	104.0±81.2	113.7±93.4	-8.5	19.3±12.0	12.7±16.8	52.4	137.6±53.4	160.4±60.8	-14.2	198.7±45.5	219.4±43.3	-9.4	52.9±38.2	53.4±44.3	-1.0*
CE	134.5±89.2	136.1±83.1	-1.2	42.3±20.8	42.8±24.4	-1.1*	158.1±55.6	174.0±51.3	-9.1	245.6±47.9	228.9±38.2	7.3	84.4±46.8	90.9±48.2	-7.2
EE	132.3±92.0	139.5±89.8	-5.2	37.5±17.5	38.8±22.1	-3.4	155.2±61.2	179.4±57.7	-13.5	248.4±44.9	242.8±36.5	2.3	80.1±46.0	88.8±48.8	-9.8
IP	197.9±95.1	194.7±84.4	1.7	91.7±26.9	98.6±27.5	-7.0	224.8±56.5	224.0±46.3	0.4*	317.5±29.1	296.3±32.3	7.2	148.6±53.9	151.8±50.4	-2.1
CM	209.8±98.6	195.1±85.1	7.5	97.3±29.1	96.7±27.1	0.6*	243.7±59.2	225.9±46.2	7.9	331.3±27.3	299.9±25.1	10.4	157.7±53.5	149.8±45.4	5.3
EM	219.3±101.6	205.6±90.3	6.7	105.1±36.8	101.8±33.7	3.3	251.4±68.8	235.6±54.4	6.7	339.3±29.1	312.8±28.1	8.5	171.8±63.0	163.7±55.9	5.0
NA	261.8±82.3	243.8±69.5	7.4	164.7±35.2	161.8±31.9	1.8	303.8±41.3	280.2±33.7	8.4	353.5±20.5	320.5±21.6	10.3	217.2±49.5	205.8±39.7	5.5

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	ANN			DJF			MAM			JJA			SON		
	REG	CMS	bias	REG	CMS	bias	REG	CMS	bias	REG	CMS	bias	REG	CMS	bias
CE	134,52±89,22	136,1±83,09	-1,16	42,30±20,78	42,75±24,38	-1,05*	158,12±55,58	174,03±51,26	-9,14	245,57±47,93	228,91±38,24	7,27	84,38±46,83	90,92±48,19	-7,19
CM	209,83±98,64	195,12±85,11	7,53	97,29±29,07	96,68±27,13	0,63*	243,68±59,2	225,86±46,23	7,88	331,26±27,32	299,93±25,09	10,44	157,71±53,54	149,8±45,42	5,27
EE	132,29±91,96	139,54±89,84	-5,19	37,53±17,47	38,84±22,10	-3,38	155,22±61,16	179,35±57,73	-13,45	248,43±44,9	242,79±36,54	2,31	80,07±45,95	88,77±48,76	-9,79
EM	219,29±101,61	205,61±90,31	6,65	105,08±36,84	101,76±33,74	3,25	251,44±68,77	235,6±54,43	6,72	339,32±29,07	312,76±28,12	8,49	171,81±62,97	163,66±55,85	4,98
EU	174,99±106,45	169,33±96,69	3,34	77,08±57,06	74,17±57,20	3,92	206,76±83,03	206,67±67,03	0,04*	281,62±70,56	265,18±55,16	6,19	126,27±77,38	123,31±71,26	2,40
IP	197,88±95,07	194,66±84,41	1,65	91,71±26,94	98,58±27,47	-6,97	224,84±56,48	223,98±46,33	0,38*	317,51±29,05	296,28±32,29	7,16	148,61±53,88	151,77±50,36	-2,08
LA	173,07±106,88	171,92±97,18	0,66	78,12±60,95	78,01±60,82	0,14*	202,65±85,68	208,71±68,55	-2,90	278,63±71,74	266,99±55,01	4,36	124,87±79,02	126,07±72,78	-0,94
NA	261,83±82,27	243,79±69,53	7,40	164,74±35,19	161,80±31,87	1,81	303,81±41,30	280,2±33,73	8,42	353,48±20,51	320,52±21,65	10,28	217,2±49,53	205,79±39,71	5,54
NE	103,98±81,16	113,69±93,36	-8,54	19,31±12,00	12,67±16,78	52,44	137,62±53,40	160,43±60,82	-14,21	198,68±45,53	219,39±43,31	-9,43	52,88±38,18	53,42±44,28	-0,99*
OC	178,24±105,64	164,93±95,68	8,07	75,32±49,72	67,66±49,81	11,32	213,75±77,82	203,22±64,23	5,18	286,69±68,21	262,11±55,29	9,37	128,65±74,47	118,63±68,35	8,44

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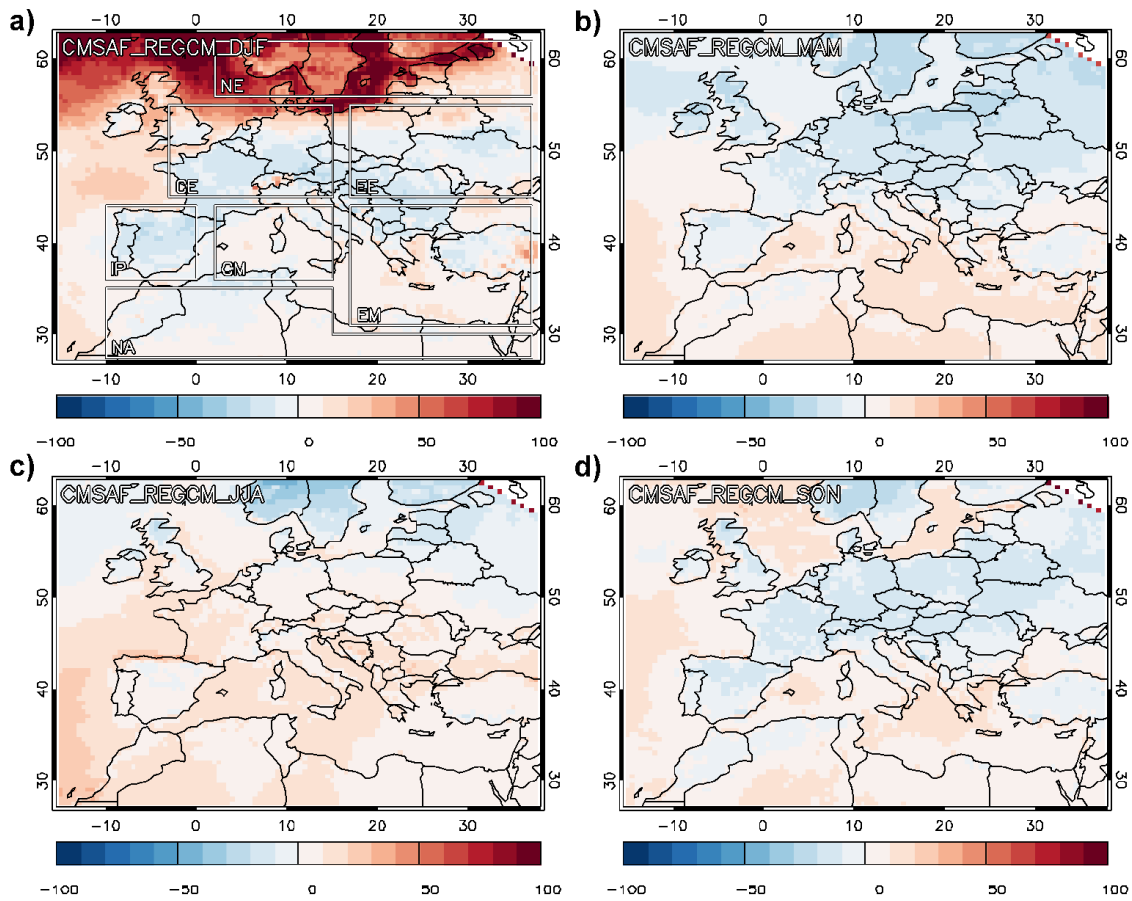
1 Table 23. Normalized Mean Bias (NMB) of RegCM4-CM SAF Rel and Rei, RegCM4-
 2 MACv1 ASY and SSA, RegCM4-CERES ALB and RegCM4-ERA-Interim WV. When the
 3 difference between RegCM4 and CM SAF or CERES or ERA-Interim is statistically
 4 significant at the 95% confidence level due to a two independent sample t-test, the NMB
 5 values are marked with bold letters while in the opposite case they are marked with an
 6 asterisk. Positive NMBs are marked with red color while negative NMBs with blue. ~~The~~
 7 ~~regions are listed in alphabetical order.~~

	CFC	COT	Rel	Rei	AOD	ASY	SSA	ALB	WV
EU	-24.3	4.3	-36.1	-28.3	-35.3	-1.1	-4.2	1.6	12.0
LA	-13.7	7.3	-47.7	-26.4	-32.1	-1.8	-4.3	-28.3	11.4
OC	-38.4	-2.5	-18.3	-31.1	-42.0	0.1	-4.1	131.1	12.8
NE	-20.3	54.3	-32.8	-31.3	-75.9	1.0	-5.6	5.2	13.1
CE	-19.7	24.1	-45.1	-24.0	-63.6	0.0*	-5.9	-22.7	14.0
EE	-16.0	30.8	-44.6	-24.2	-64.6	2.1	-3.5	-40.7	10.8
IP	-13.7	-13.9	-46.1	-27.3	-7.4	-1.5	-4.8	-3.8	14.4
CM	-31.2	-30.7	-26.7	-27.6	-19.3	-0.7	-3.5	85.9	10.4
EM	-28.8	-22.0	-29.3	-28.4	-34.2	-0.0	-2.3	35.4	10.9
NA	0.4*	-39.8	-47.3	-30.0	25.0	-7.9	-3.5	-26.4	8.7

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3 Figure 1. Seasonal NMB patterns of RegCM4-CM SAF SSR over Europe for (a) winter
 4 (DJF), (b) spring (MAM), (c) summer (JJA) and (d) autumn (SON) from MSG SEVIRI
 5 observations. The 7 sub-regions used for the generalization of the results are marked in Fig.
 6 1a: Northern Europe (NE), Central Europe (CE), Eastern Europe (EE), Iberian Peninsula (IP),
 7 Central Mediterranean (CM), Eastern Mediterranean (EM) and Northern Africa (NA).

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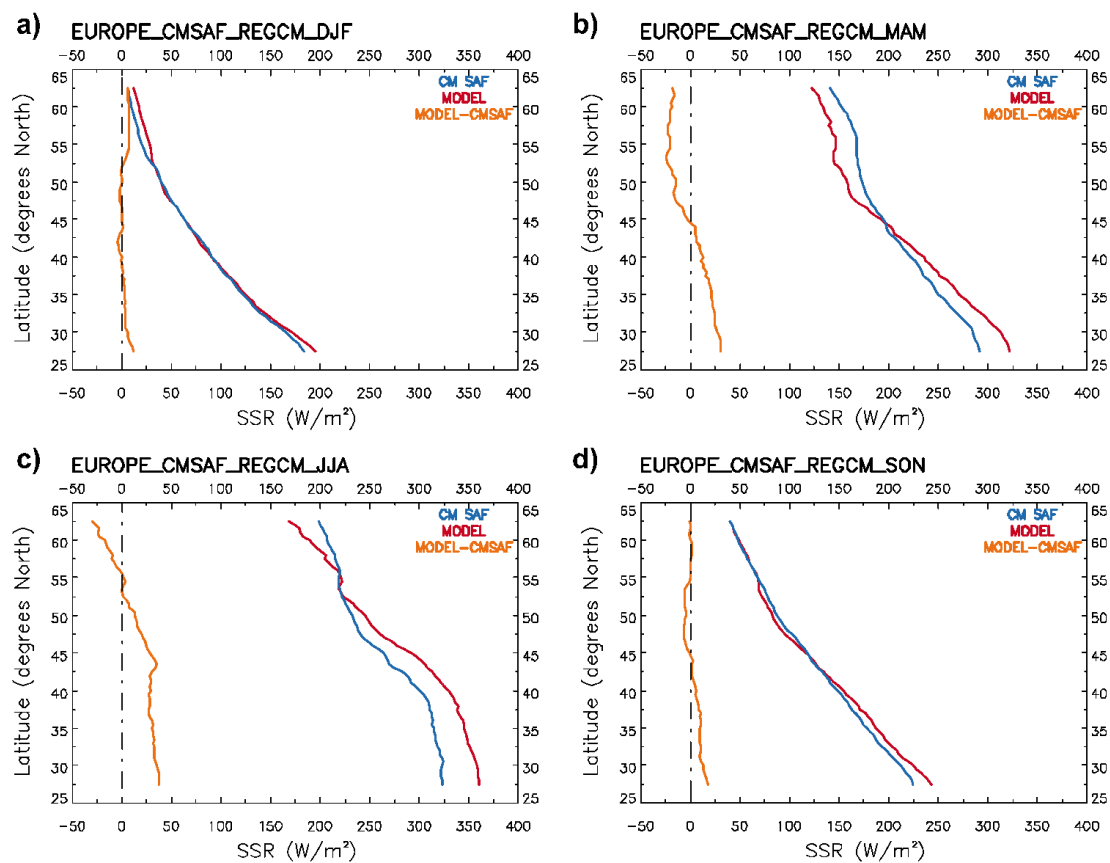
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3 Figure 2. Latitudinal variability of RegCM4 SSR (red), CM SAF SSR (blue) and their
 4 difference (orange) over Europe for (a) winter (DJF), (b) spring (MAM), (c) summer (JJA)
 5 and (d) autumn (SON) from MSG SEVIRI observations.

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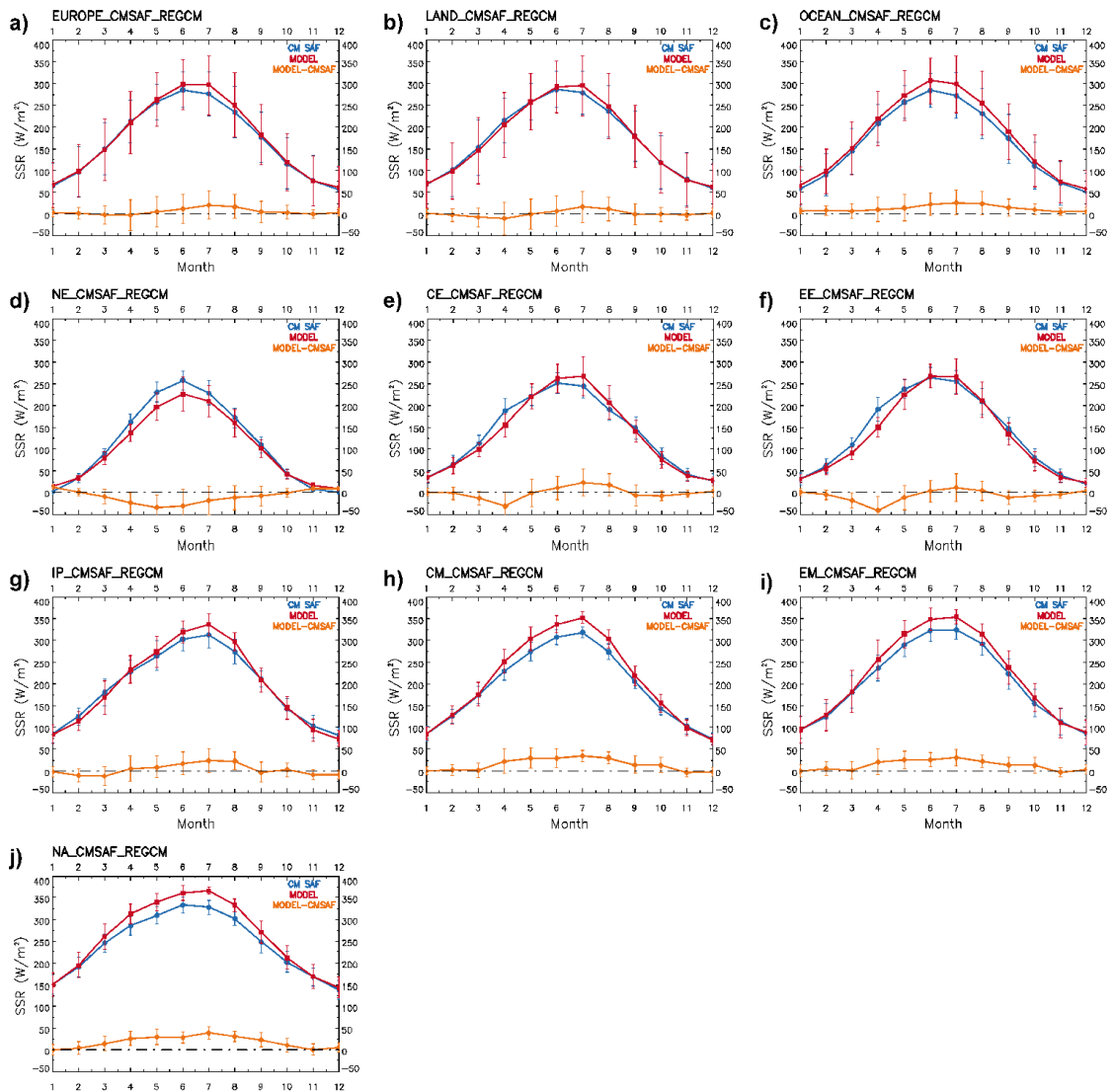
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2 Figure 3. Seasonal variability of RegCM4 SSR (red), CM SAF SSR (blue) and their
 3 difference (orange) over (a) the whole Europe, (b) Land, (c) Ocean, (d) NE, (e) CE, (f) EE, (g)
 4 IP, (h) CM, (i) EM, (j) NA from MSG SEVIRI observations.

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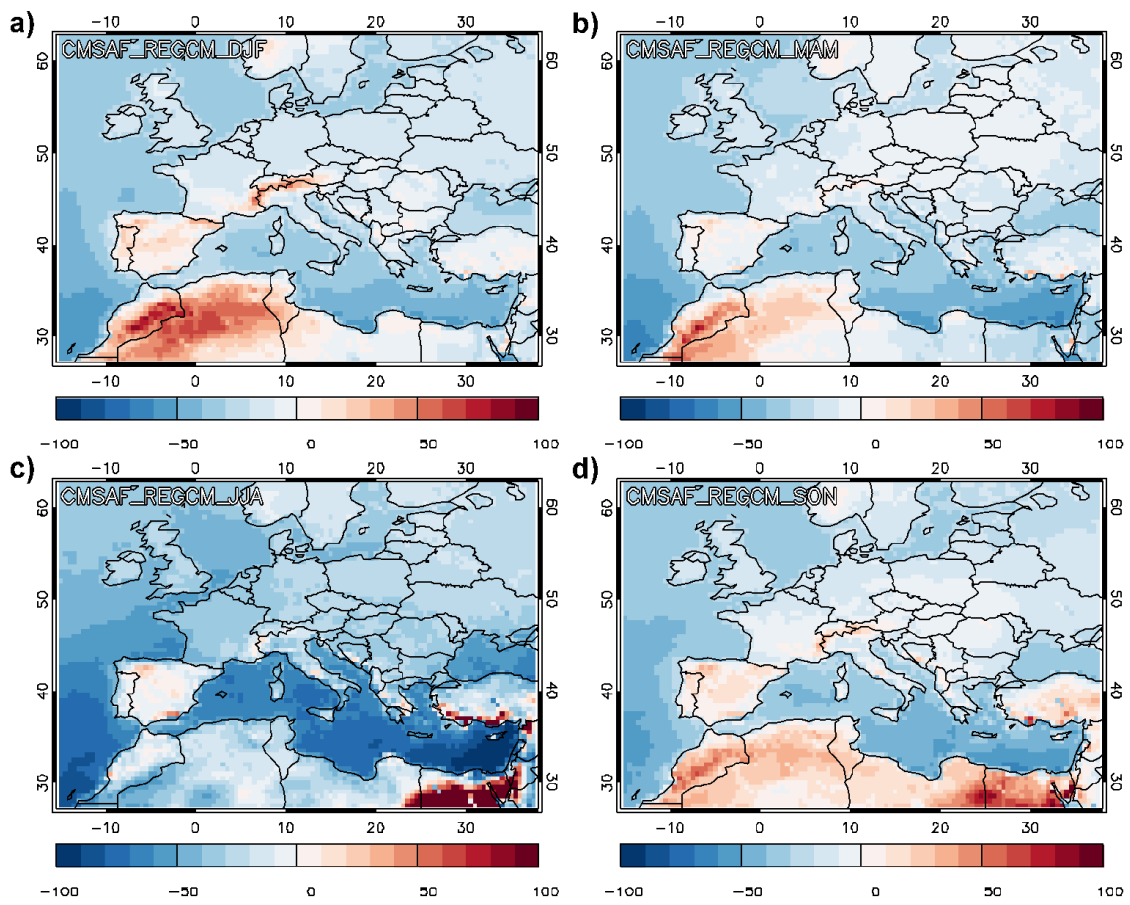
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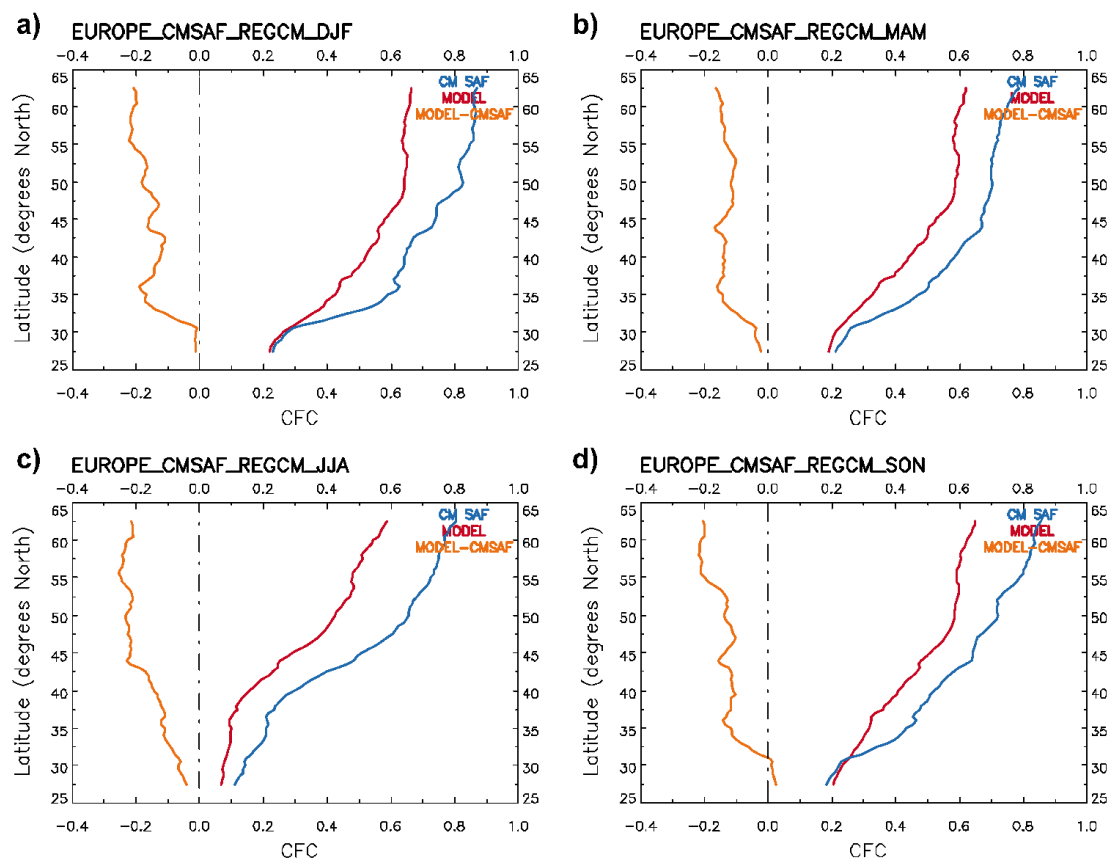
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Figure 4. The same as Fig. 3 but for RegCM4 and CM SAF CFC.



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3 Figure 5. The same as Fig. 4 but for RegCM4 and CM SAF CFC.

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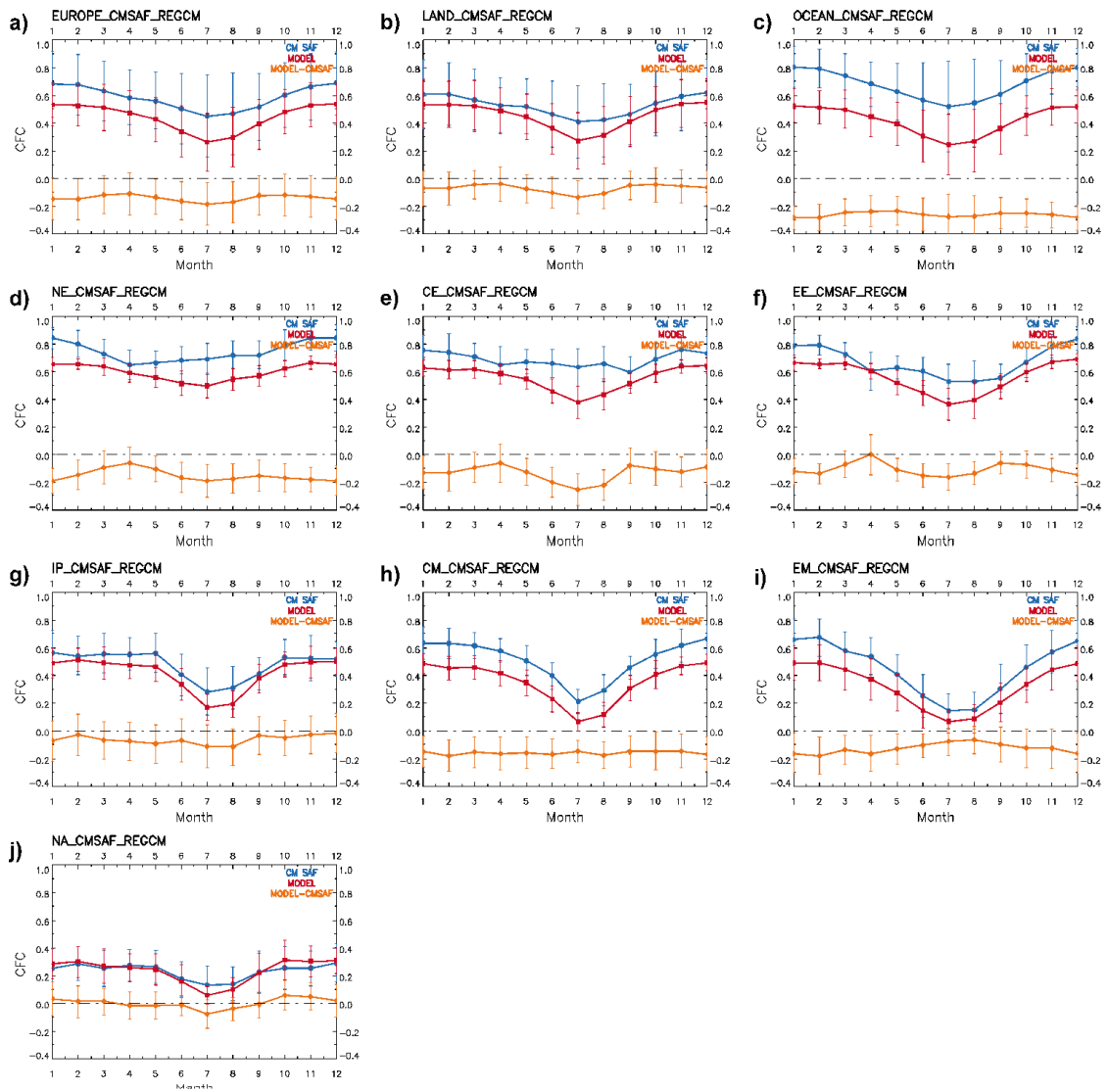
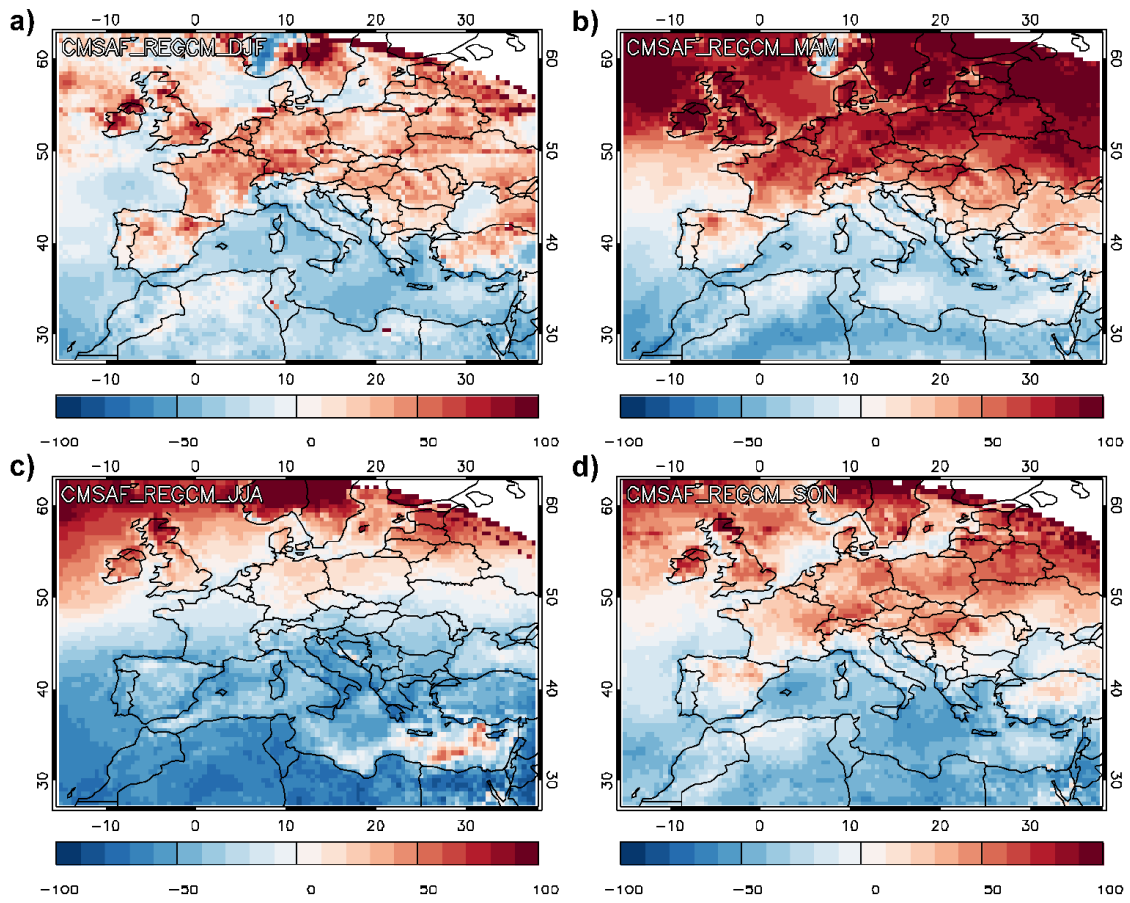


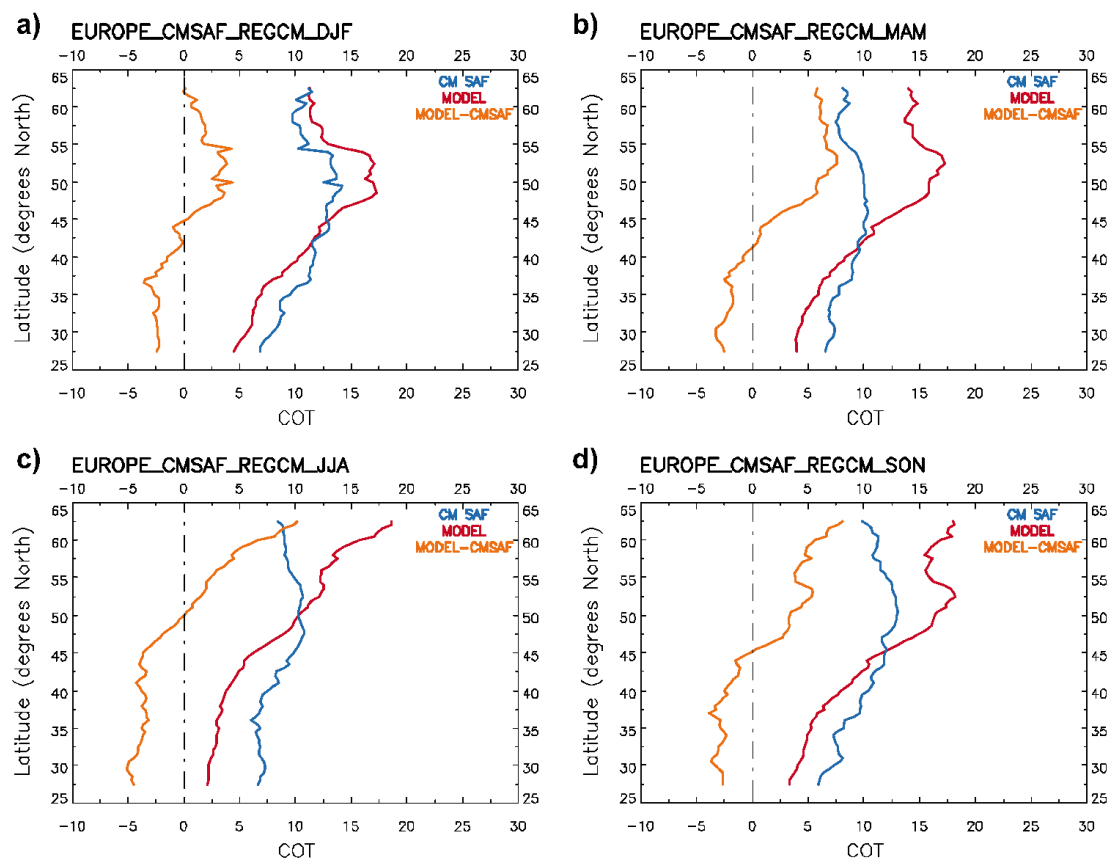
Figure 6. The same as Fig. 5 but for RegCM4 and CM SAF CFC.

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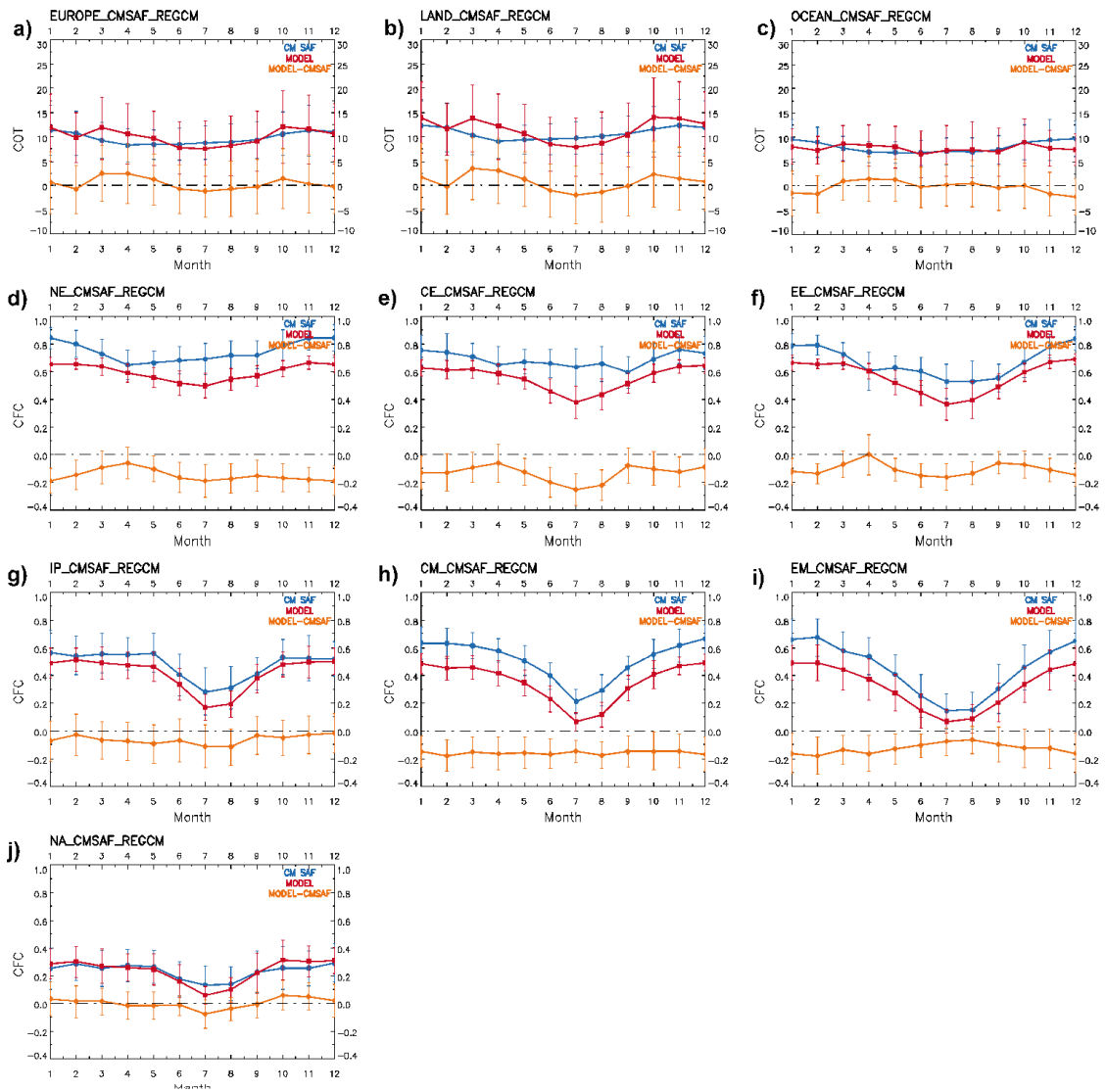
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Figure 7. The same as Fig. 3 but for RegCM4 and CM SAF COT.



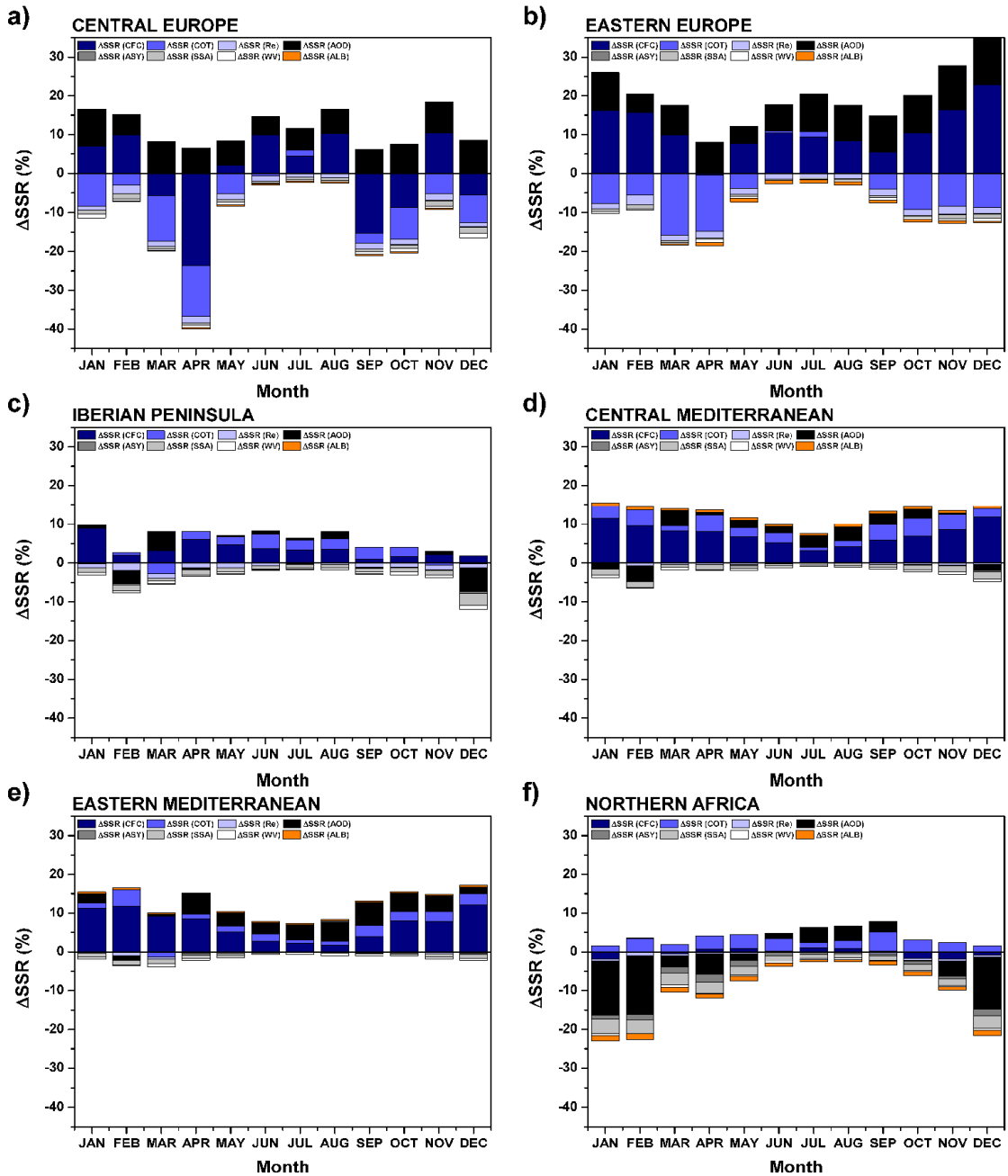
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Figure 8. The same as Fig. 4 but for RegCM4 and CM SAF COT.



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Figure 9. The same as Fig. 5 but for RegCM4 and CM SAF COT.



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2 Figure 10. Δ SSR (%) caused by CFC, COT, Re, AOD, ASY, SSA, WV and ALB for (a) CE,
 3 (b) EE, (c) IP, (d) CM, (e) EM and (f) NA.