

Laboratory Evidence of Organic Peroxide and Peroxyhemiacetal Formation in the Aqueous Phase and Implications for Aqueous OH

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Abstract

Aqueous chemistry in atmospheric waters (e.g., cloud droplets or wet aerosols) is considered a potentially important atmospheric pathway to produce secondary organic aerosol (SOA_{aq}). Water-soluble organic compounds with small carbon numbers (C₂-C₃) are precursors for SOA_{aq} and products include organic acids, organic sulfates, and high molecular weight compounds/oligomers. Fenton reactions and the uptake of gas-phase OH radicals are considered to be the major oxidant sources for aqueous organic chemistry. However, the sources and availability of oxidants in atmospheric waters are not well understood. The degree to which OH is produced in the aqueous phase affects the balance of radical and non-radical aqueous chemistry, the properties of the resulting aerosol, and likely its atmospheric behavior.

This paper demonstrates organic peroxide formation during aqueous photooxidation of methylglyoxal using ultra high resolution Fourier Transform Ion Cyclotron Resonance electrospray ionization mass spectrometry (FTICR-MS). Organic peroxides are known to form through *gas-phase* oxidation of volatile organic compounds. They contribute secondary organic aerosol (SOA) formation directly by forming peroxyhemiacetals, and epoxides (i.e., IEPOX), and indirectly by enhancing gas-phase oxidation through OH recycling. We provide

1 simulation results of organic peroxide/peroxyhemiacetal formation in clouds and wet aerosols
2 and discuss organic peroxides as a source of condensed-phase OH radicals and as a
3 contributor to aqueous SOA.

4

5 **1 Introduction**

6 Secondary organic aerosol (SOA) is a major component of atmospheric fine particulate matter
7 [PM_{2.5}] (Zhang et al., 2007), contributes to adverse health, and affects climate by scattering
8 (Seinfeld and Pandis, 1998) and sometimes by absorbing solar radiation (e.g., “brown
9 carbon”) (Andreae and Gelencser, 2006; Bones et al., 2010; Zhang et al., 2011). Although the
10 chemical and physical properties of aerosols are needed to predict effects, the properties of
11 SOA are poorly understood because SOA formation itself is poorly understood. Aqueous
12 chemistry in atmospheric waters (e.g., cloud droplets or wet aerosols) is a potentially
13 important pathway to produce SOA (SOA_{aq}; Blando and Turpin, 2000), and could be
14 comparable in magnitude to “traditional” SOA, formed via partitioning of semivolatile
15 organic products of gas-phase oxidation (SOA_{gas}) globally (Liu et al., 2012; Lin et al., 2012;
16 Henze et al., 2008) and in locations where relative humidity and aerosol hygroscopicity are
17 high (Carlton and Turpin, 2013; Carlton et al., 2008; Fu et al., 2008; Chen et al., 2007).
18 Because SOA_{aq} is formed from small water-soluble precursors with high O/C ratios, it forms
19 SOA (e.g., oligomers, organic salts) with high O/C ratios and may explain the highly
20 oxygenated nature of atmospheric organic aerosols, while SOA_{gas} is less oxygenated (Aiken et
21 al., 2008; Lim et al., 2010 and 2013).

22 OH radicals are important oxidants in clouds. In the high solute concentrations in wet
23 aerosols, however, besides OH radical reactions a more complex system of organic radical
24 and non-radical reactions occurs (Lim et al., 2010; McNeill et al., 2012; Ervens et al., 2014).
25 Thus, an understanding of the availability of OH radicals is important to assessing the relative
26 importance of radical and non-radical chemistry in aerosols. The uptake of gas-phase OH
27 radicals into atmospheric waters (Faust and Allen, 1993) and Fenton reactions in the
28 condensed/aqueous media (Arakaki and Faust, 1998) are considered the major oxidant
29 sources for aqueous organic chemistry. Oxidant sources in organic-containing cloud, fog and
30 aerosol waters and oxidant reactions with dissolved organic compounds have been
31 documented (Arakaki et al., 2013; Weller et al., 2014; Long et al., 2013). Depending on

1 sources of OH radicals, aqueous oxidation reactions could exhibit a surface area dependence
2 (e.g., controlled by OH uptake), or a volume dependence (e.g., controlled by OH production
3 through aqueous chemistry; Ervens et al., 2014). Herein, we explore the hypothesis that
4 organic peroxides produce OH radicals within the atmospheric aqueous phase; we also
5 demonstrate the formation of organic peroxides in the aqueous phase and their contribution to
6 *condensed* phase chemistry.

7 Organic peroxides (herein particularly, organic hydroperoxides = ROOH) are known to play
8 an important role in gas phase chemistry. They are commonly found in the atmosphere with
9 mixing ratios of 0.1-1 ppb (Lee et al., 1993; De Serves et al., 1994; Sauer et al., 2001;
10 Grossmann et al., 2003; Lee et al., 2000; Guo et al., 2014). They are known to form through
11 gas-phase reactions of volatile organic compounds (VOCs) with OH radical, NO₃ radical and
12 O₃ (Atkinson and Arey, 2003). While their chemistry is not fully understood, these
13 atmospheric organic species are “key” to peroxy radical/NO_x chemistry (Dibble, 2007;
14 Glowacki et al., 2012), lead to photochemical smog formation, important to the HO_x-NO_x-O₃
15 balance (Wennberg et al., 1998; Singh et al., 1995), contribute to O₃ formation or depletion in
16 the upper troposphere, and form SOA (Tobias and Ziemann, 2000; Ehn et al., 2014). Organic
17 peroxides (formed from gas-phase ozonolysis of monoterpenes, e.g., α- and β-pinenes) are
18 major constituents of SOA (Docherty et al., 2005). Monoterpenes have a global emission
19 second only to isoprene among non-methane VOCs and maybe the most efficient SOA_{gas}
20 precursor class (Kanakidou et al., 2005). Organic peroxides contribute to organic aerosol by
21 forming peroxyhemiacetal oligomers with atmospherically abundant organic carbonyls (e.g.,
22 aldehydes and ketones) via acid catalysis in aerosols (Tobias and Ziemann, 2000; Ziemann,
23 2002). In general, due to the characteristically weak O-O bonds of organic peroxides, the gas-
24 phase decomposition of organic peroxides through photolysis or intermolecular radical
25 reactions recycles OH radicals and can enhance gas-phase photooxidation of atmospheric
26 organic compounds. Recent field studies demonstrate that gas-phase OH recycling enhances
27 isoprene photooxidation (Paulot et al., 2009; Taraborrelli et al., 2012). And a recent lab study
28 (Badali et al., 2015) demonstrates that OH radicals are photolytically formed from the
29 solutions of SOA from terpene ozonolysis and OH formation is likely due to photolysis of
30 organic peroxides.

31 Organic peroxides are known to be moderately water soluble (Henry’s law constant up to
32 1000 M/atm). They are present in rainwater with concentrations of 0.1 - 10 μM (Lind et al.,

1 1986; Hellpointer and Gab, 1989; Liang et al., 2013), presumably by uptake from the gas
2 phase. Badali et al. (2015) measured OH radical formation from photolysis of terpene-O₃
3 SOA solutions and organic peroxide standard solutions (t-butyl hydroperoxide and cumene
4 hydroperoxide). However, photolysis of the terpene SOA generates twice as much OH as is
5 generated from a comparable amount of organic peroxide alone (i.e., standards). Since there
6 should exist plenty of aldehydes formed from ozone reactions, we argue that organic
7 peroxides could also be formed in the SOA solution during the photolytic experiments. In
8 this work, we show that organic peroxides are also produced from aqueous-phase OH
9 oxidation. We identify organic peroxide products from methylglyoxal and acid catalyzed
10 oligomers (i.e., peroxyhemiacetals formed with methylglyoxal) by ultra-high resolution mass
11 spectrometry. We simulate organic peroxide and peroxyhemiacetal formation under
12 atmospheric conditions and explore organic peroxide contributions to aqueous-phase OH
13 production and to SOA_{aq} formation.

14

15 **2 Experimental Section**

16 **2.1 Cuvette Chamber Reactions**

17 Reactions of methylglyoxal with OH radicals in the aqueous phase were conducted in a
18 cuvette chamber, which holds 10 cuvettes (3 mL each; Spectrocell) equidistant from a 254 nm
19 Hg UV lamp (Strahler). Cuvettes were immersed in a water bath to maintain the temperature
20 at 25 °C. In each cuvette, 10 mM of methylglyoxal (Sigma-Aldrich) was dissolved in 18 MΩ
21 Mili-Q water. OH radicals (10⁻¹³—10⁻¹² M) were generated in each cuvette by photolysis of
22 20 mM of H₂O₂ (Sigma-Aldrich) with the rate constant of 5.58e-5 M s⁻¹. The H₂O₂ photolysis
23 rate constant was determined from H₂O₂ + UV control experiments conducted in the same
24 cuvette chamber as described previously (Tan et al., 2010) and corrected for light absorption
25 by H₂O₂ (Lim et al., 2013). Liu et al. (2012) and Zhao et al. (2013) found that α-
26 hydroperoxides can form when methylglyoxal reacts with hydrogen peroxide in the dark.
27 However, this reaction cannot explain the formation of the identified peroxy hemiacetals in
28 this work (PHA₁ and PHA₂) since the molecular weight of the α-hydroperoxide is different
29 from those of R₁OOH and R₂OOH. Moreover, according to control experiments by Tan et al.
30 (2010) methylglyoxal degradation is much slower with hydrogen peroxide in the dark than it
31 is with hydrogen peroxide in the UV light. Therefore, we do not expect the formation of the
32 α-hydroperoxide in our photooxidation experiment.

1 The OH radical concentrations were estimated via modeling (Lim et al., 2013). It should be
2 noted that using 20 mM of H₂O₂ and the 254 nm UV lamp was not intended to simulate
3 tropospheric photolysis, rather to provide a source of OH radicals. According to our previous
4 control experiments (i.e., methylglyoxal + UV; methylglyoxal + H₂O₂), small amounts of
5 pyruvic, acetic and formic acids form slowly in control experiments. However, dicarboxylic
6 acids, the major products, did not form in the absence of OH radicals (i.e., in control
7 experiments; Tan et al., 2010). Photooxidation of methylglyoxal was allowed to proceed for
8 1 hr. After being removed from the chamber, the cuvettes were kept frozen until analysis. No
9 catalase was added in order to preserve organic peroxide products.

10 **2.2 Organic Peroxide and Peroxyhemiacetal Analysis**

11 Ultra high resolution Fourier Transform Ion Cyclotron Resonance Electrospray Ionization
12 Mass Spectrometer (FTICR-MS; Thermo-Finnigan LTQ-XL, Woods Hole Oceanographic
13 Mass Spectrometer Facility) was used to determine the elemental composition of organic
14 products as described previously (Altieri et al., 2008; Tan et al., 2012). The capillary voltage
15 and a capillary temperature were -30.00V and 300 °C, respectively for negative mode
16 analyses. Positive mode analyses were conducted with a capillary voltage of 20.00 V and a
17 capillary temperature of 260 °C. Both FTICR-MS and FTICR-MS/MS were used to analyze
18 organic peroxide products from aqueous photooxidation of methylglyoxal and a standard
19 solution, which was prepared by adding 10 mM of tert-butyl hydroperoxide (Sigma-Aldrich)
20 and 10 mM of methylglyoxal (Sigma-Aldrich). These samples were diluted 100 fold with
21 water (by volume), and diluted again with methanol (MeOH) by 2 fold (by volume). Thus,
22 the mobile phase consisted of 50% water and 50% MeOH; 0.1% of formic acid (by volume)
23 was also added. These diluted samples were immediately introduced into the electrospray
24 ionization source by direct infusion at 5 µL/min. Photooxidation products of methylglyoxal
25 were expected in both negative and positive modes due to a carboxylic group (negative mode)
26 and a hydroxyl group (positive mode) in their structure (Table 1), whereas tert-butyl
27 hydroperoxide is found only in the positive mode.

28

29 **3 Organic Peroxide Chemistry**

30 We hypothesize that aqueous-phase OH radical reactions of methylglyoxal lead to organic
31 peroxide formation as shown in Figure 1. OH radical reactions are initiated by H-atom

1 abstraction. Subsequent O₂ addition and HO₂ decomposition mainly lead to the formation of
2 pyruvic acid and acetic acid (Lim et al., 2013). Both pyruvic and acetic acid react further
3 with OH radicals and O₂, forming peroxy radicals (RO₂), which undergo bimolecular RO₂-
4 RO₂ reactions (Lim et al., 2013). However, substantial amounts of peroxy radicals could also
5 react with HO₂ forming organic peroxides (as indicated by a bold arrow) since HO₂ is a
6 common byproduct of aqueous photooxidation (Lim et al., 2010 and 2013) and is also water
7 soluble (Henry's law constant = 4e3 M/atm; this is ~ 100 times higher than that of OH
8 radicals).

9 We further expect organic peroxides to form peroxyhemiacetals with aldehydes via acid
10 catalysis in the aqueous phase (Figure 2A), as they do in dry aerosols (Tobias and Ziemann,
11 2000; Docherty et al., 2005). Below we document the formation of peroxyhemiacetals from a
12 commercially available organic peroxide, tert-butyl hydroperoxide and methylglyoxal in
13 aqueous solution (Figure 2B). Then we argue that organic peroxide products (R₁OOH and
14 R₂OOH in Figure 1) from the aqueous OH oxidation of methylglyoxal react further with
15 methylglyoxal in water to produce peroxyhemiacetals. Briefly, a carbonyl group (aldehyde)
16 in methylglyoxal is protonated by H⁺, then a hydroperoxyl group (-OOH) attacks a protonated
17 carbonyl group forming peroxyhemiacetal. This peroxyhemiacetal chemistry is a well
18 established oligomerization mechanism for SOA from gas-phase ozone reactions of alkenes in
19 smog chamber studies (Tobias and Ziemann, 2000). In Tobias and Ziemann study, organic
20 peroxides are first formed in the gas phase and become particles through gas-particle
21 partitioning. Then organic peroxides form peroxyhemiacetals with by-product aldehydes
22 through acid-catalyzed heterogeneous reactions on the particle surface. In current study, the
23 detection of peroxyhemiacetals in our aqueous chemistry experiments (see below) provides
24 evidence for organic peroxide formation through aqueous photochemistry.

25

26 **4 Results and Discussion**

27 **4.1 Standard Solution (Mixture of Methylglyoxal and t-Butyl Hydroperoxide)**

28 FTICR-MS and FTICR-MS/MS analyses of the aqueous mixture of methylglyoxal and t-butyl
29 hydroperoxide show methylglyoxal (m/z⁺ 127.03666, 145.04714 and 159.06278) and a
30 peroxyhemiacetal (PHA_{std}; m/z⁺ 185.07797) in the positive mode (Figure 3). In a mobile
31 phase of 50% MeOH and 50% water, methylglyoxal undergoes hydration with water and
32 hemiacetal formation with MeOH, and is detected as a sodium adduct (i.e., m/z⁺ 127.03666 =

1 [methylglyoxal + MeOH + Na]⁺, m/z⁺ 145.04714 = [methylglyoxal + H₂O + MeOH + Na]⁺,
2 and m/z⁺ 159.06278 = [methylglyoxal + 2MeOH + Na]⁺. Methylglyoxal solvation with
3 MeOH was verified by FTICR-MS/MS (Figure 4). The fragments of m/z⁺ 159.06278 are
4 m/z⁺ 141.05203 (H₂O loss), m/z⁺ 127.03668 (MeOH loss), and m/z⁺ 95.01041 (another
5 MeOH loss). The ion m/z⁺ 95.01041 is the sodium adduct to methylglyoxal (a theoretical
6 reading for [methylglyoxal + Na]⁺ is m/z⁺ 95.01090). We are confident that m/z⁺ 159.06278 is
7 a double hemiacetal of methanol with methylglyoxal, not a cluster of methylglyoxal with two
8 methanol molecules by the water loss in Fig. 4 and examination of ESI-MS standard runs for
9 glyoxal and methylglyoxal in the water mobile phase with and without methanol (See
10 Supplementary Material Fig. S5). A sodium adduct is also expected for a peroxyhemiacetals,
11 and seen in Figure 3 as m/z⁺ 185.07797 (= [PHA_{std} + Na]⁺). Details of FTICR-MS readings
12 and theoretical readings based on actual molecular/atomic masses are shown in Table 1.

13 The PHA_{std} peak, m/z⁺ 185.07802 (theoretical reading of [PHA_{std} + Na]⁺ = 185.07899) was
14 fragmented by infrared multiphoton dissociation (IRMPD). Major fragments (Figure. 5) are
15 m/z⁺ 153.05228 (MeOH loss) and m/z⁺ 95.01041 ([methylglyoxal + Na]⁺). In electron
16 impact (hard ionization), fragmentation of organic peroxides results in the loss of HO₂
17 (Tobias and Ziemann, 2000; Docherty et al., 2005). However, O₂ loss is expected for soft
18 ionization, IRMPD fragmentation in FTICR-MS/MS (M. Soule and E. Kujawinski, personal
19 communication, 2013; Detailed discussion in Supplementary Material). In Figure 5, the ion
20 m/z⁺ 81.06971 indicates the loss of O₂ from t-butyl hydroperoxide. We also observed the O₂
21 loss (m/z⁺ 153.07158) from PHA_{std}. FTICR-MS/MS and theoretical readings are provided in
22 Table 2.

23 Note that no organic peroxide peak was observed in the standard solution (nor in
24 methylglyoxal + OH samples). This is not surprising because 1) high temperature of the
25 capillary in an electrospray chamber (~ 250 °C) is likely to decompose organic peroxides
26 (Kharasch et al., 1950; M. Soule and E. Kujawinski, personal communication, 2013); 2) in the
27 ESI method, it is difficult to ionize organic peroxides (Witkowski and Gierczak, 2013) and
28 organic peroxides react with methylglyoxal to form peroxyhemiacetals. These
29 peroxyhemiacetals are much more stable and lesser volatile than organic peroxides (Tobias
30 and Ziemann, 2000). These peroxyhemiacetal peaks (and fragments) appear in FTICR-MS
31 (and FTICR-MS/MS) analysis of standard solutions and samples (see below), providing
32 evidence for the presence (and formation) of organic peroxides from methylglyoxal + OH.

1 FTICR-MS/MS of peroxyhemiacetal peaks show corresponding organic peroxide fragments,
2 methylglyoxal and other fragments as expected (Tobias and Ziemann, 2000; Docherty et al.,
3 2005).

4 **4.2 Aqueous Photooxidation Products of Methylglyoxal**

5 A FTICR mass spectrum of an aqueous methylglyoxal solution exposed to OH radicals for 60
6 minutes is shown in Figure 6 (negative mode). Main photooxidation products (Tan et al.,
7 2012; Lim et al., 2013) are seen at m/z^- 87.00862 (pyruvic acid) and m/z^- 177.04036 (2, 3-
8 dimethyltartaric acid; structure shown in Figure 6). This spectrum also provides evidence for
9 peroxyhemiacetal formation at m/z^- 163.02392 (= [PHA₁ - H]⁻) and m/z^- 191.01998 (= [PHA₂
10 - H]⁻) since these readings are very close to the theoretical readings, m/z^- 163.02405 for PHA₁
11 and m/z^- 191.01918 for PHA₂ (Table 1). Fragmentation of these peaks by FTICR-MS/MS
12 supports their identification as peroxyhemiacetals. In Figure 7A, m/z^- 131.01339 and m/z^-
13 131.01585 result from the losses of MeOH and O₂, respectively, from PHA₁ at m/z^-
14 163.02405 and m/z^- 59.01377 results from the loss of O₂ from R₁OOH, which is the organic
15 peroxide constituent of PHA₁. Similarly, in Figure 7B, m/z^- 159.02946 (C₆H₇O₅) results
16 from the loss of O₂ from PHA₂ at m/z^- 191.02000. m/z^- 87.00832 results from the loss of O₂
17 from R₂OOH, which is the organic peroxide constituent of PHA₂. Note that m/z^- 191.02000 is
18 PHA₂ while m/z^- 191.05540 is prominent as a parent ion. Due to the small intensity we were
19 unable to isolate m/z^- 191.02000 from m/z^- 191.05540 for MS/MS analyses. Therefore, for
20 the PHA₂ analysis, we cannot rule out the possibility that m/z^- 59.01377 [R₂OOH - O₂] could
21 be the fragment from m/z^- 191.05540 [C₇H₁₁O₆]⁻. As was the case with the mixed peroxide-
22 aldehyde standard, the organic peroxides themselves were not observed (see previous section).
23 Detected and theoretical readings are provided in Table 2.

24 The FTICR-MS/MS was also conducted in the positive mode (Figure 8) for PHA₁ (m/z^+
25 187.02069 in Figure 8A) and PHA₂ (m/z^+ 215.01565 in Figure 8B). The theoretical readings
26 are 187.02186 for PHA₁ and 215.01678 for PHA₂ (Table 1). The methylglyoxal fragments
27 (m/z^+ 95.1041 in Figure 8A and m/z^+ 95.0140 in Figure 8B) appear. This confirms that both
28 PHA₁ and PHA₂ are indeed acid-catalyzed products of methylglyoxal. Note that for the PHA₂
29 analysis in the positive mode, again, we cannot rule out the possibility that methylglyoxal
30 [m/z^+ 95.01040] could be the fragment of m/z 215.05151 [C₇H₁₂O₆Na]⁺.

31

1 **5 Atmospheric Implications**

2 Using ultra-high resolution FTICR-MS and FTICR-MS/MS, we observed the presence of
3 peroxyhemiacetals, after aqueous photooxidation of methylglyoxal and in aqueous
4 methylglyoxal-organic peroxide standard solutions. The presence of stable
5 peroxyhemiacetals is an indicator of the existence of the less stable organic peroxides. Thus,
6 this work provides evidence for the formation of organic peroxides through aqueous phase
7 OH radical oxidation of methylglyoxal.

8 **5.1 Organic Peroxide Production in Clouds and Wet Aerosols**

9 Below we demonstrate through chemical modeling that organic peroxides photochemically
10 form from organics present both in clouds and wet aerosols. We used the full kinetic model
11 for glyoxal and methylglyoxal (Lim et al., 2013) to simulate the formation of organic
12 peroxides and peroxyhemiacetals. The following updates were made to the model: 1) The
13 rate constant for the bimolecular reactions of RO_2 and HO_2 was given as $3\text{e}6 \text{ M}^{-1}\text{s}^{-1}$ (Reaction
14 213-219 in Table S1) based on the rate constant for $[\text{HO}_2 + \text{HO}_2] \sim 1\text{e}6 \text{ M}^{-1}\text{s}^{-1}$; 2) The
15 concentration of OH in the aqueous phase was set to $\sim 10^{-14}$ (previously $\sim 10^{-12}$) according to
16 recent estimations (Arakaki et al., 2013) (Figure S1A); 3) The concentration of HO_2
17 photochemically formed in the aqueous phase was estimated to be $\sim 10^{-8} \text{ M}$ maintained by the
18 Henry's law equilibrium; therefore, the excess HO_2 produced by photooxidation in the
19 aqueous phase was transported to the gas phase (Figure S1B). All the reactions included in
20 the model are listed in Table S1.

21 For wet aerosol simulations, 1 M (the initial concentration) of methylglyoxal was used in the
22 aqueous phase. Note that we do not expect that methylglyoxal is present at 1 M in aerosols.
23 However, water-soluble organic matter is present at 1-10 M. So this analysis treats all water-
24 soluble organic matter as if it behaves like methylglyoxal. Under wet aerosol conditions
25 ($[\text{methylglyoxal}]_{\text{initial}} = 1 \text{ M}$, $[\text{H}_2\text{O}_2]_{\text{initial}} = 0 \text{ M}$, $[\text{OH}] \sim 10^{-14} \text{ M}$, and $[\text{HO}_2] \sim 10^{-8} \text{ M}$), ~ 400
26 μM of organic peroxides during the 12-hr daytime were formed through aqueous OH radical
27 reactions (Figure 10A). The model also includes the sinks of aqueous-phase organic
28 peroxides: OH radical reactions (R220-225), photolysis (R230), and the evaporation to the gas
29 phase (R234) in Table S1. Note that organic peroxide (ROOH) formation in Figure 10A and
30 B does not change within the Henry's law constant, 100 to 1000 M/atm, and the evaporation
31 rate is assumed to be a diffusion-controlled transfer coefficient (Lelieveld and Crutzen, 1991;

1 Lim et al., 2005), which is the upper limit based on the equation provided by Lelieveld and
2 Crutzen (1991). Here, the gas-phase [ROOH] is assumed to be 1 ppb (R234 in Table S1).

3 In atmospheric cloud conditions ($[\text{methylglyoxal}]_{\text{initial}} = 10 \mu\text{M}$, $[\text{H}_2\text{O}_2]_{\text{initial}} = 0 \text{ M}$, $[\text{OH}] \sim 10^{-14}$
4 M , and $[\text{HO}_2] \sim 10^{-8} \text{ M}$), $\sim 0.4 \mu\text{M}$ of organic peroxide formation during the 12-hr daytime
5 is expected (Figure 10B) while all the sinks of organic peroxides listed above are included.
6 This concentration of aqueous-phase photochemically produced organic peroxides is within
7 the range of measured rainwater concentrations ($0.1 - 10 \mu\text{M}$) and similar to the concentration
8 expected by Henry's law equilibrium from gas-phase organic peroxides ($0.1 - 1 \text{ ppb}$).

9 **5.2 Peroxyhemiacetal Formation in Wet Aerosol**

10 The formation of peroxyhemiacetals competes with 1) hydration of methylglyoxal and 2)
11 photolysis and OH reactions of organic peroxides (Figure 9). Competing with methylglyoxal
12 hydration means that only a dehydrated methylglyoxal (DeMGLY), not hydrated
13 methylglyoxal (MGLY), forms a peroxyhemiacetal (PHA) with an organic peroxide (ROOH),
14 since the aldehyde reacts with peroxides. The dehydration equilibrium for methylglyoxal is
15 included in the model (R226 in Table S1). In wet aerosols, $\sim 0.4 \text{ mM}$ of DeMGLY out of 1 M
16 MGLY will undergo peroxyhemiacetal formation and react with OH radicals (R227 and R228
17 in Table S1) at the same time (Figure S2A). The main sink for peroxyhemiacetals is expected
18 to be OH reaction (no evaporation is expected). The peroxyhemiacetal formation equilibrium
19 (R229) and the OH reaction of peroxyhemiacetals (R231) are listed in Table S1. The
20 modified model simulates $\sim 0.4 \mu\text{M}$ of peroxyhemiacetal formation during the 12-hr daytime
21 and the minor increase during the nighttime (Figure 10A). Under cloud conditions,
22 peroxyhemiacetal formation is negligible (Note that the model simulates $\sim 4\text{e-}15 \text{ M}$
23 peroxyhemiacetal formation during the daytime from $10 \mu\text{M}$ of methylglyoxal photooxidation
24 in Figure 10B).

25 **5.3 OH Recycling Due to the Photolysis of Organic Peroxides in Atmospheric** 26 **Waters**

27 In both cloud and wet aerosol conditions, $7.5\text{e-}15 \text{ M}$ of aqueous-phase OH production is
28 expected from the photolysis of organic peroxides ($[\text{ROOH}]_{\text{initial}} \sim 400 \mu\text{M}$ in wet aerosols
29 and $[\text{ROOH}]_{\text{initial}} \sim 0.4 \mu\text{M}$ in cloud droplets) formed by aqueous photooxidation during the
30 12-hr daytime (Figure S3) while the sink of ROOH is OH reaction. Note that Badali et al.
31 (2015) confirmed OH formation from photolysis of solutions of organic peroxide SOA and

1 measured OH formation rates are comparable to an estimation by Arakaki et al. (2013), which
2 is $\sim 10^{-14}$ M OH in atmospheric waters, and $\sim 10^{-14} - 10^{-15}$ M of OH was previously estimated
3 in the core of the bulk phase (Jacob 1986). Thus, the aqueous production of organic
4 peroxides in atmospheric waters could be an important source of aqueous OH through organic
5 peroxide photolysis.

6

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16

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7

1 TABLE 1. FTICR-MS and theoretical readings for methylglyoxal, organic peroxides and
 2 peroxy hemiacetals.

Methylglyoxal		Organic Peroxide		Peroxyhemiacetal (= Methylglyoxal + Organic Peroxide)	
MW 72.02113		MW 90.06809		MW 162.08922	
Theoretical Reading ^a	FTICR-MS Reading	Theoretical Reading ^a	FTICR-MS Reading	Theoretical Reading ^a	FTICR-MS Reading
 [M + MeOH ^b + Na] ⁺ = 127.03715	m/z ⁺ 127.03666	[M + Na] ⁺ = 113.05786	m/z ⁺ (Not Detected)	[M + Na] ⁺ = 185.07899	m/z ⁺ 185.07797
		MW 92.01096		MW 164.03209	
 [M + H ₂ O + MeOH ^b + Na] ⁺ = 145.04768	m/z ⁺ 145.04714	[M - H] ⁻ = 91.00314	m/z ⁻ (Not Detected)	[M - H] ⁻ = 163.02427 [M + Na] ⁺ = 187.02186	m/z ⁻ 163.02405 m/z ⁺ 187.02069
		MW 120.00588		MW 192.02701	
 [M + 2MeOH ^b + Na] ⁺ = 159.06330	m/z ⁺ 159.06278	[M - H] ⁻ = 118.99805	m/z ⁻ (Not Detected)	[M - H] ⁻ = 191.01918 [M + Na] ⁺ = 215.01678	m/z ⁻ 191.01998 m/z ⁺ 215.01565

3 ^aTheoretical reading is based on actual atomic/molecular weights obtained by online software,

4 ^b“Molecular Isotopic Distribution Analysis (MIDAS)”

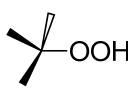
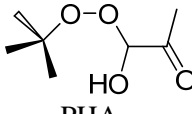
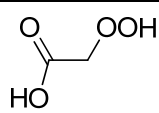
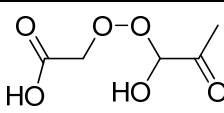
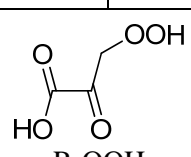
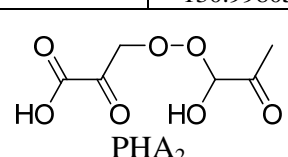
5 (<http://www.ncbi.nlm.nih.gov/CBBresearch/Yu/midas/index.html>).

6 ^bMeOH = Methanol. Note that the mobile phase contains 50% water (with 0.05% formic acid) and
 7 50% MeOH.

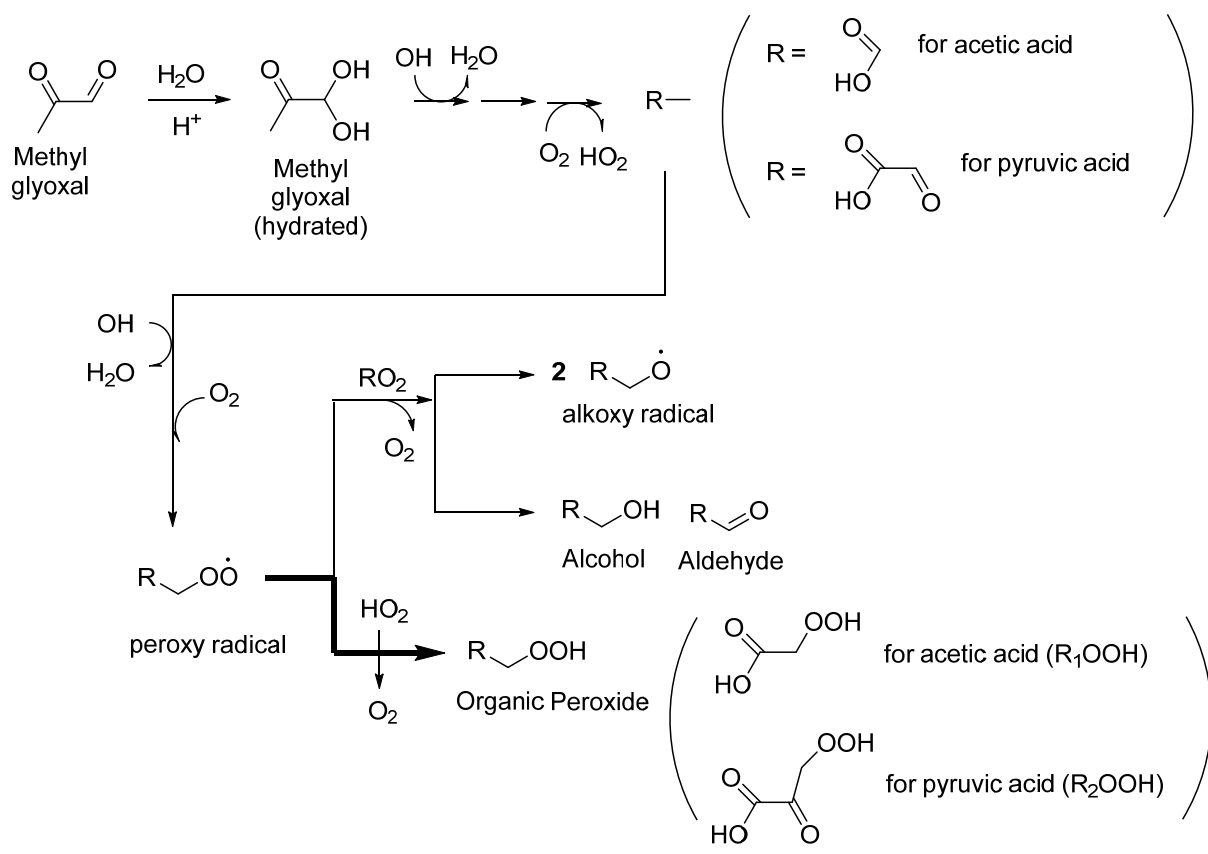
8 ^cMeO is a deprotonated MeOH.

9

1 TABLE 2. FTICR-MS/MS and theoretical readings for fragments of peroxyhemiacetals and
 2 organic peroxides

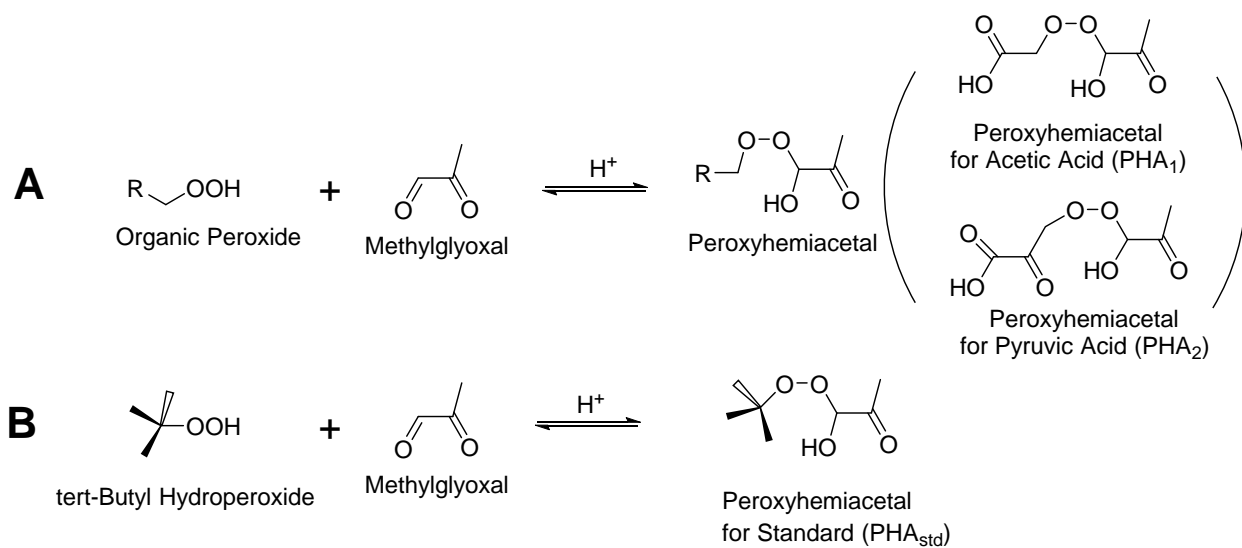
Organic Peroxide		Peroxyhemiacetal			
 t-butyl hydroperoxide		 PHA _{std}			
Fragment (loss of O ₂)		Fragment (loss of O ₂)		Fragment (loss of MeOH)	
Theoretical Reading	FTICR-MS/MS Reading	Theoretical Reading	FTICR-MS/MS Reading	Theoretical Reading	FTICR-MS/MS Reading
$[M + Na - O_2]^+ = 81.06803$	$m/z^+ = 81.06971$	$[M + Na - O_2]^+ = 153.08916$	$m/z^+ = 153.07158$	$[M + Na - MeOH]^+ = 153.05277$	$m/z^+ = 153.05228$
 R ₁ OOH		 PHA ₁			
Fragment (loss of O ₂)		Fragment (loss of O ₂)		Fragment (loss of MeOH)	
Theoretical Reading	FTICR-MS/MS Reading	Theoretical Reading	FTICR-MS/MS Reading	Theoretical Reading	FTICR-MS/MS Reading
$[M - H - O_2]^- = 59.01331$	$m/z^- = 59.01377$	$[M - H - O_2]^- = 131.03444$	$m/z^- = 131.01589$	$[M - H - MeOH]^- = 130.99805$	$m/z^- = 131.01339$
 R ₂ OOH		 PHA ₂			
Fragment (loss of O ₂)		Fragment (loss of O ₂)		Fragment (loss of MeOH)	
Theoretical Reading	FTICR-MS/MS Reading	Theoretical Reading	FTICR-MS/MS Reading	Theoretical Reading	FTICR-MS/MS Reading
$[M - H - O_2]^- = 87.00822$	$m/z^- = 87.00832$	$[M - H - O_2]^- = 159.02935^e$	$m/z^- = 159.02946$	$[M - H - MeOH]^- = 159.00079$	m/z^- (Not Detected)

3



1

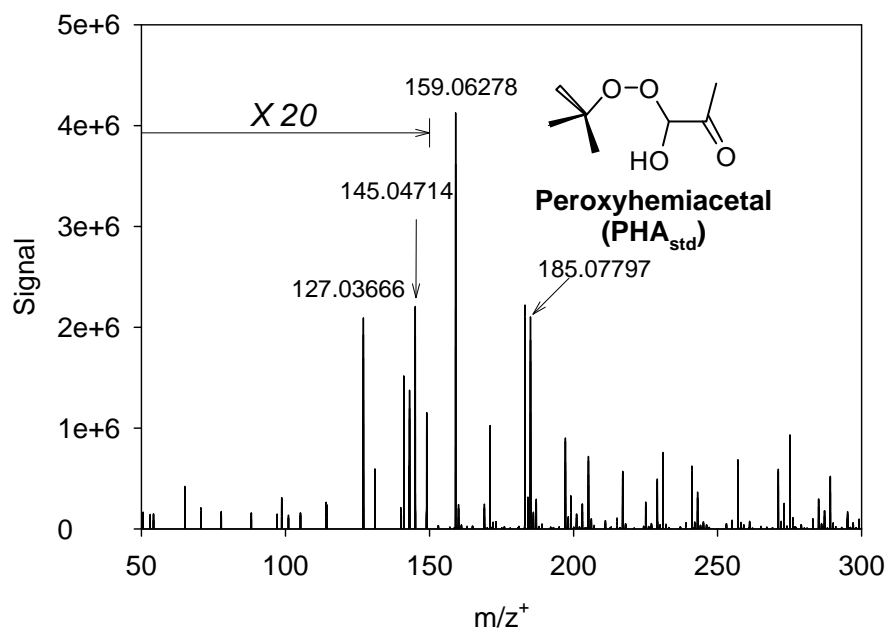
2 Figure 1. Organic peroxide formation from aqueous-phase OH radical reactions of
 3 methylglyoxal.



1
2

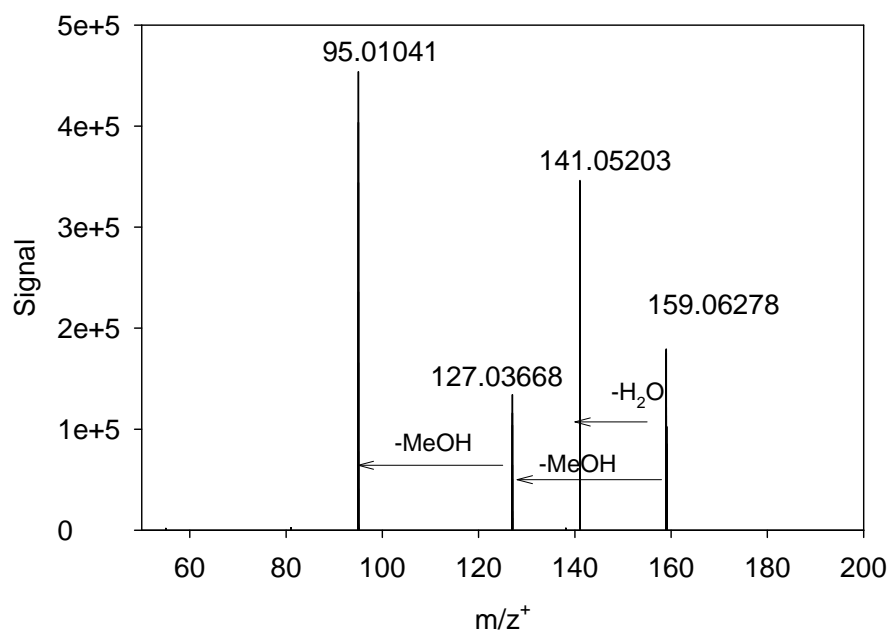
3 Figure 2. Acid-catalyzed peroxyhemiacetal formation from a precursor, methylglyoxal with
 4 organic peroxide products of aqueous photooxidation (A) and from methylglyoxal with tert-
 5 butyl hydroperoxide in standard solutions (B).

6



1
 2 FIGURE 3. A full FTICR-MS spectrum for the standard solution of methylglyoxal (10 mM)
 3 and tert-butyl hydroperoxide (10 mM) in the positive mode.

4
 5

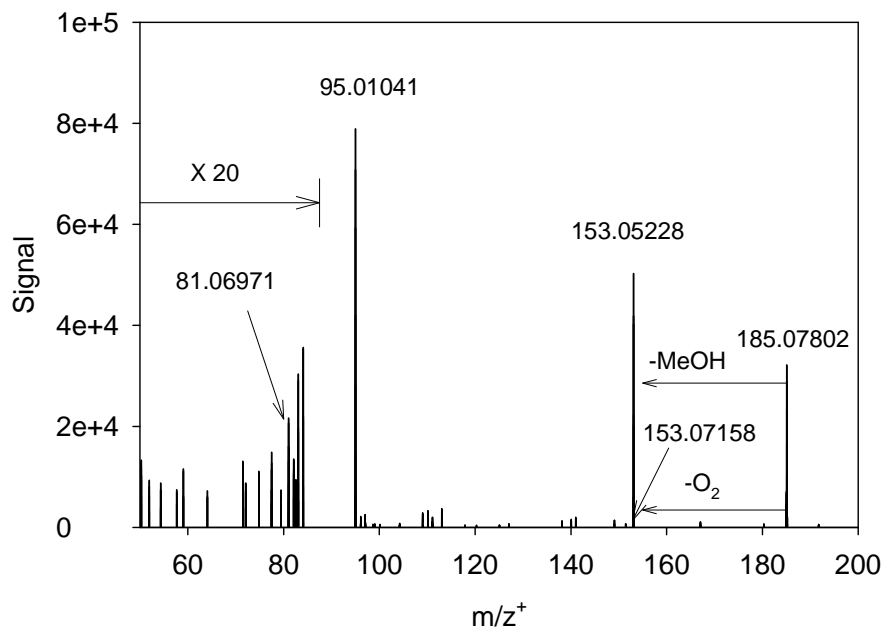


1

2 FIGURE 4. FTICR-MS/MS for m/z+ 159 (methylglyoxal).

3

1

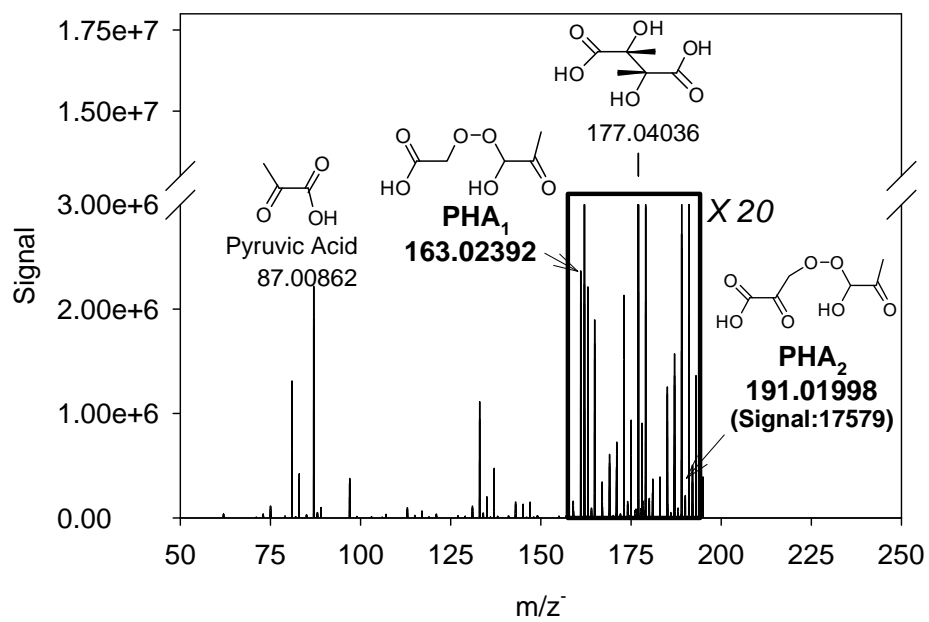


2

3 FIGURE 5. FTICR-MS/MS for m/z+ 185 (PHA_{std}).

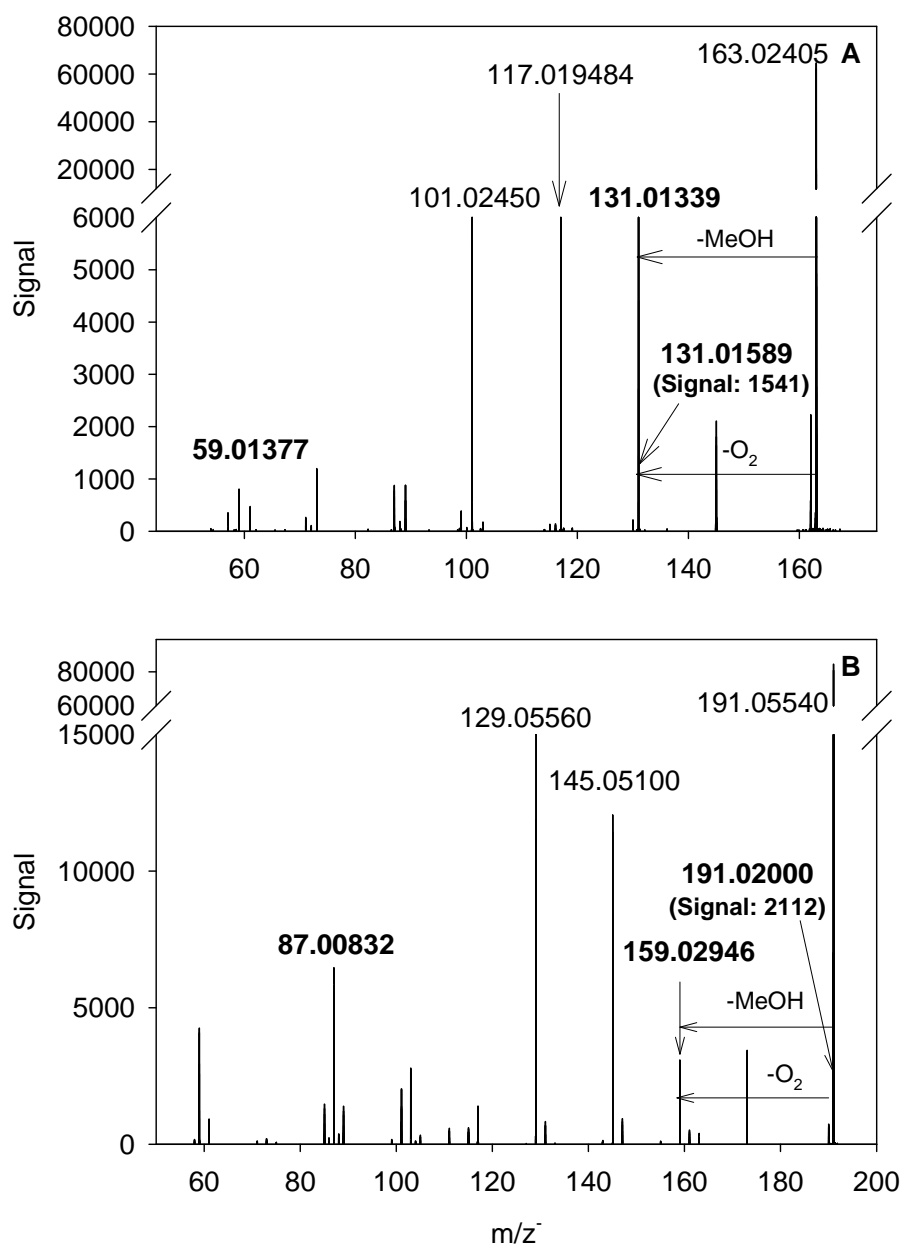
4

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1
 2 FIGURE 6. A full FTICR-MS spectrum for products of 1-hr aqueous photooxidation of
 3 methylglyoxal in the negative mode.
 4

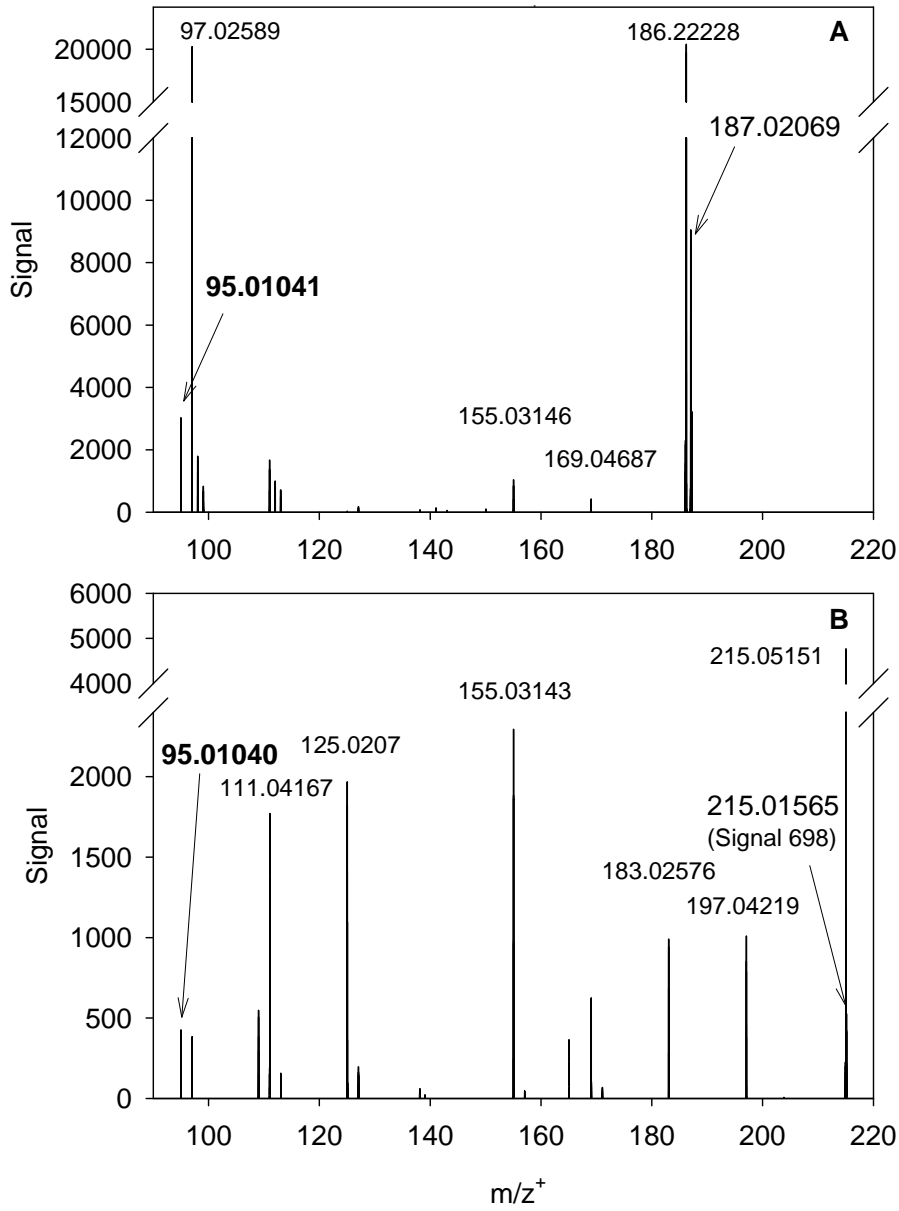
1



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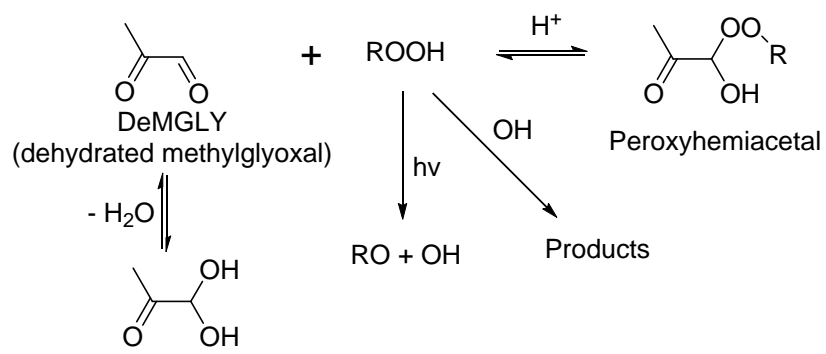
3 FIGURE 7. FTICR-MS/MS for m/z 163 for PHA₁ (A) and m/z 191 PHA₂ (B).

4



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4

FIGURE 8. FTICR-MS/MS for m/z^+ 187 for PHA₁ (**A**) and m/z^+ 215 for PHA₂ (**B**).



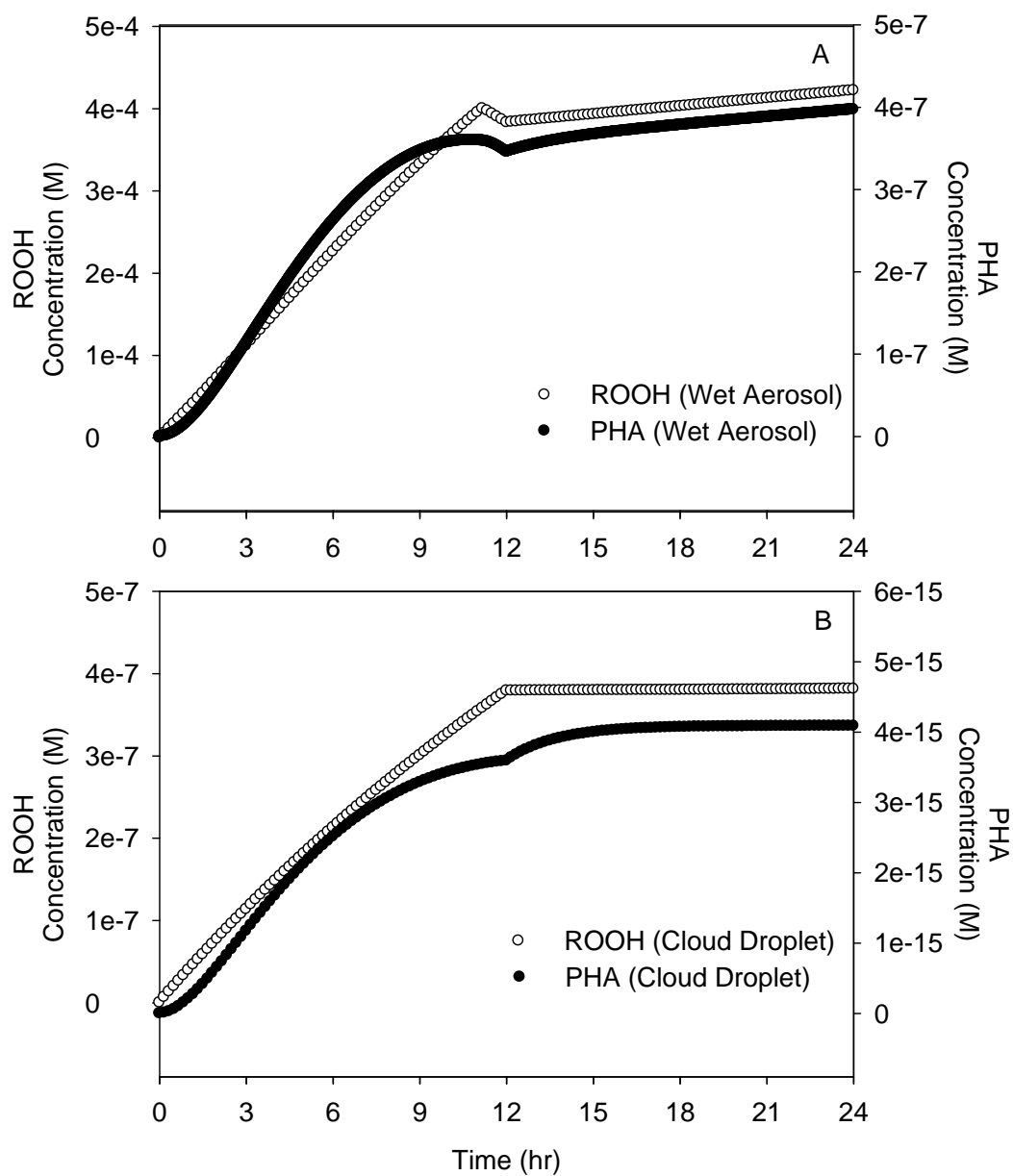
1 Methylglyoxal (hydrated)

2 FIGURE 9. Peroxyhemiacetal formation

3

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1
 2 FIGURE 10. The atmospheric simulated concentrations of ROOH (organic peroxides) and
 3 PHA (peroxyhemiacetals) in wet aerosols (A) and cloud droplets (B) during 24 hrs. The first
 4 12 hrs are daytime.

5