



Does GOSAT capture the true seasonal cycle of XCO₂?

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Does GOSAT capture the true seasonal cycle of XCO₂?

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Abstract

The seasonal cycle accounts for a dominant mode of total column CO₂ (XCO₂) annual variability and is connected to CO₂ uptake and release; it thus represents an important variable to accurately measure from space. We quantitatively evaluate the XCO₂ seasonal cycle of the Greenhouse Gases Observing Satellite (GOSAT) observations from the Atmospheric CO₂ Observations from Space (ACOS) retrieval system, and compare average regional seasonal cycle features to those directly measured by the Total Carbon Column Observing Network (TCCON). We analyze the mean seasonal cycle amplitude, dates of maximum and minimum XCO₂, as well as the regional growth rates in XCO₂ through the fitted trend over several years. We find that GOSAT generally captures the seasonal cycle amplitude within 1.0 ppm accuracy compared to TCCON, except in Europe, where the difference exceeds 1.0 ppm at two sites, and the amplitude captured by GOSAT is generally shallower compared to TCCON. This bias over Europe is not as large for the other GOSAT retrieval algorithms (NIES v02.21, RemoTeC v2.35, UoL v5.1, and NIES PPDF-S v.02.11), although they have significant biases at other sites. The ACOS bias correction was found to partially explain the shallow amplitude over Europe. The impact of the TCCON retrieval version, co-location method, and aerosol changes in the ACOS algorithm were also tested, and found to be few tenths-of-a-ppm and mostly non-systematic. We find generally good agreement in the date of minimum XCO₂ between ACOS and TCCON, but ACOS generally infers a date of maximum XCO₂ 2–3 weeks later than TCCON. We further analyze the latitudinal dependence of the seasonal cycle amplitude throughout the Northern Hemisphere, and compare the dependence to that predicted by current optimized models that assimilate in-situ measurements of CO₂. In the zonal averages, GOSAT agrees with the models to within 1.4 ppm, depending on the model and latitude. We also show that the seasonal cycle of XCO₂ depends on longitude especially at the mid-latitudes: the amplitude of GOSAT XCO₂ doubles from West US to East Asia at 45–50° N, which is only partially shown by the models. In general, we find that model-to-model differences can be larger

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than GOSAT-to-model differences. These results suggest that GOSAT retrievals of the XCO₂ seasonal cycle may be sufficiently accurate to evaluate land surface models in regions with significant discrepancies between the models.

1 Introduction

5 Satellites provide unprecedented spatial coverage of the variability of atmospheric carbon dioxide (CO₂) through retrievals of column mean dry mole fractions of CO₂ (XCO₂). XCO₂ shows temporal variability on different timescales: diurnal, synoptic, seasonal, inter-annual, and long-term (Olsen and Randerson, 2004; Keppel-Aleks et al., 2011). Variability is determined by the collective impact of CO₂ fluxes resulting from fossil
10 fuel emissions, biosphere–atmosphere exchange, and ocean–atmosphere exchange, and the imprint of these on regional XCO₂ can be strongly influenced by atmospheric dynamics, in addition to the regional origin of the fluxes. While the secular trend and multi-year interhemispheric CO₂ gradient are driven by the global build-up of CO₂ from fossil fuel combustion mainly in the Northern Hemisphere, the seasonal variability is
15 mainly controlled by variations in the terrestrial biospheric fluxes (Palmer et al., 2008; Keppel-Aleks et al., 2011). The ocean–atmosphere and fossil fuel CO₂ fluxes are, although seasonally varying, only minor contributors to the XCO₂ seasonal variability in the Northern Hemisphere. Therefore, the seasonal cycle of XCO₂ bears the signature of large-scale biospheric flux patterns, especially their north-south distribution.

20 Regional biospheric CO₂ fluxes are a critical part of land surface models that describe the biosphere–atmosphere carbon exchange in larger modeling systems, such as carbon cycle and climate models (Pitman, 2003). Inverse model systems use these land surface models in conjunction with atmospheric transport models, and optimize their CO₂ flux estimates by assimilating CO₂ measurements, but especially in regions
25 where the in-situ measurement network has sparse coverage, the inverse models can strongly disagree about the seasonality and magnitude of the fluxes (Lindqvist et al., 2015). Recently, this disagreement has been found to lead to large regional discrepan-

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cies of several ppm in the seasonal cycle amplitudes of modeled XCO₂ (Keppel-Aleks et al., 2012; Peng et al., 2015; Lindqvist et al., 2015). This finding not only suggests that regional XCO₂ can be indicative of local fluxes and that satellite-measured XCO₂ may be useful in constraining the models even without inversions, but also is another reminder that there is potentially much to be gained by assimilating space-based XCO₂ retrievals that vastly expand the current in-situ measurement network; a lesson shown previously by a number of studies (e.g., Rayner and O'Brien, 2001; Chevallier et al., 2007; Takagi et al., 2011, 2014; Maksuytov et al., 2013). In particular, the strength of the seasonal cycle drawdown is fundamentally connected to the magnitude of the carbon sink during the growing season. By studying the GOSAT XCO₂ seasonal cycle and its inter-annual variability, Wunch et al. (2013) showed that the variability in the drawdown correlates with surface temperature in the boreal regions, and Guerlet et al. (2013b) found a reduced carbon uptake during the 2010 Northern Hemisphere summer.

The Greenhouse Gases Observing Satellite (GOSAT; Yokota et al., 2009) and the Orbiting Carbon Observatory-2 (OCO-2; Crisp et al., 2004) are indeed designed to make near-global XCO₂ retrievals that will constrain the inverse model systems enough to provide a picture of the global carbon cycle with respect to regional sources and sinks. However, a crucial question still lingers: are the satellite observations accurate enough to reliably capture the seasonal variability of XCO₂? The question is fair because satellite-retrieved XCO₂ is subject to biases in the retrieval system (e.g., Wunch et al., 2011b), and also sampling biases due to the seasonally-dependent amount of solar radiation (e.g., Liu et al., 2014). Both of these may have an impact on the measured seasonal cycle. For the Atmospheric CO₂ Observations from Space (ACOS) retrieval system (O'Dell et al., 2012; Crisp et al., 2012), known biases in GOSAT retrievals are corrected using a global bias correction (Wunch et al., 2011b) but some parameters of the bias correction vary seasonally, for example surface albedo. Potential remaining biases, their seasonality, and impact on the seasonal cycles of XCO₂ are best identified through evaluation of the GOSAT seasonal cycle against the best

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available independent data – those from the Total Carbon Column Observing Network, TCCON (Wunch et al., 2011a). There have been several studies that compare GOSAT retrievals against the TCCON, some of them introducing novel methods for comparisons (Wunch et al., 2011b; Nguyen et al., 2014), some concentrating on quantifying biases in a specific retrieval algorithm (Butz et al., 2011; Cogan et al., 2012; Yoshida et al., 2013), and some focusing more on the intercomparison of different retrieval algorithms (Buchwitz et al., 2015; Oshchepkov et al., 2013a; Reuter et al., 2013; Dils et al., 2014). Overall, the collective message from the validation studies is that the agreement of GOSAT and TCCON has improved (i.e., the satellite biases have decreased) substantially from the earliest validation efforts (Morino et al., 2011), owing to major improvements and updates in the retrieval algorithms and the development of more sophisticated comparison methods. However, less attention has been paid to the evaluation of the seasonal cycle. Reuter et al. (2013, p. 1776) touched on this by showing averages of the seasonal cycle amplitude differences between all GOSAT retrievals and TCCON (and also a model, CarbonTracker CT2011_oi). More recently, Kulawik et al. (2015) studied the seasonality of GOSAT-TCCON biases (using the ACOS B3.4 retrieval algorithm for GOSAT data) and found notable station-to-station variability in the biases, but also persisting seasonal biases in latitudinally averaged results. These seasonal biases were reflected in the seasonal cycle amplitudes.

In this paper, we continue the evaluation of the GOSAT seasonal cycle from Kulawik et al. (2015). Five years of GOSAT observations and the updated TCCON GGG2014 retrievals lengthen the co-located time series sufficiently to evaluate the seasonal cycles regionally at 12 TCCON sites in the Northern Hemisphere. We extend the seasonal cycle analysis to four other retrieval algorithms to identify potential biases characteristic to the ACOS retrievals. Although the emphasis of the study is on these TCCON comparisons, we also compare the GOSAT seasonal cycle against models that assimilate in-situ data; because of their connection to measurements, models may be a reasonable representation of the truth in areas with high assimilated data density, such as North America or Western Europe. This seasonal cycle evaluation study lays

important ground work to the analysis of OCO-2 observations that also use the ACOS retrieval system and are, therefore, likely to be affected by any seasonal biases present in the GOSAT/ACOS retrievals that are due to the ACOS system itself.

2 GOSAT

5 The Greenhouse Gases Observing Satellite (GOSAT), developed by Japan Aerospace Exploration Agency (JAXA), was launched in January 2009 to make near-global greenhouse gas measurements from a polar orbit (Yokota et al., 2009). GOSAT measures scattered solar near-infrared radiation with a Fourier transform spectrometer (TANSO-FTS; Kuze et al., 2009). The diameter of a GOSAT sounding footprint is approximately
10 10 km, and the soundings repeat in a three-day cycle. We used GOSAT data taken in two primary modes: glint over oceans, and nadir view over land. Nadir data over land has two gain states: high gain (H) for most of the data, and medium (M) over bright surfaces, such as deserts.

Several retrieval algorithms have been developed for retrieving the column-averaged
15 CO₂ from the GOSAT near-infrared measurements; these algorithms have been recently reviewed and compared by Oshchepkov et al. (2013a) and Reuter et al. (2013). In this paper, we concentrate on the evaluation of the Atmospheric CO₂ Observations from Space build 3.5 (ACOS B3.5) retrieval algorithm (Crisp et al., 2012). The ACOS retrieval algorithm is described in detail by O'Dell et al. (2012). The most significant
20 subsequent updates and improvements to the operational algorithm include updated spectroscopy for the 1.6 and 2.1 μm CO₂ absorption bands, moving from static to dynamic vertical pressure levels, an improved prior profile of CO₂, and a complete change in the treatment of aerosol and cloud scattering. Instead of a globally constant aerosol model that was incorporated in ACOS B3.4 and earlier versions, B3.5 uses Modern-Era
25 Retrospective Analysis for Research and Applications (MERRA) reanalysis data of five aerosol types (mineral dust, sea salt, black carbon, sulphates, and organic carbon) to

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determine two most common types at a given GOSAT sounding location, and applies their respective optical properties in the retrieval.

3 Validation data

3.1 TCCON

5 The Total Carbon Column Observing Network (TCCON) is currently composed of 21 operating Fourier transform spectrometers that make ground-based measurements of atmospheric XCO₂ and other gases (Wunch et al., 2011a). Their validated and calibrated higher precision and accuracy compared to satellite observations, coupled with the fact that they measure the same quantity in essentially the same way as the satel-
10 lites, make them an ideal, independent validation source for GOSAT (Wunch et al., 2010; Messerschmidt et al., 2011). Though the seasonal cycle of TCCON has itself never been explicitly validated by comparison with aircraft, we implicitly assume that our inferred TCCON seasonal cycles for XCO₂ can be taken as truth, similar to the assumption in several previous studies (Messerschmidt et al., 2011; Keppel-Aleks et al.,
15 2012; Wunch et al., 2013), though in principle sub-ppm seasonal biases could remain. For instance, the TCCON retrieval performs a post-hoc airmass bias correction (Wunch et al., 2011a), errors in which could lead to small but nontrivial differences in the TCCON seasonal cycle. However, it is beyond the scope of this work to validate the accuracy of the TCCON seasonal cycle.

20 In this study, we used data from all Northern Hemisphere TCCON sites that had (1) at least two years of coincidental measurements with GOSAT; and (2) enough co-located data (see Sect. 4.1) to evaluate a seasonal cycle; i.e., both ACOS and TCCON observations available at the proximity of the site through most seasons. The first criterion eliminated the Four Corners and Caltech/Pasadena sites, while the second eliminated
25 the northernmost sites of Ny Ålesund and Eureka which have very little co-located data due to the high latitude. We did not include the Southern Hemisphere sites be-

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cause the seasonal changes in XCO₂ at those sites are minor, making the definition of an average seasonal cycle more ambiguous and sensitive to inter-annual variability. The sites that were used in this study are shown in Fig. 1. For these sites, we analyzed all co-located data between 23 April 2009, and 31 December 2013. We used the newest available TCCON retrievals for every site. GGG2014 retrievals were available for the following sites in our study: Bialystok (Messerschmidt et al., 2012; Deutscher et al., 2014), Garmisch (Sussmann and Rettinger, 2014), Izaña (Blumenstock et al., 2014), JPL (Wennberg et al., 2014a), Karlsruhe (Hase et al., 2014), Lamont (Wennberg et al., 2014c), Orleans (Warneke et al., 2014), Park Falls (Washenwelder et al., 2006; Wennberg et al., 2014b), Saga (Kawakami et al., 2014), Sodankylä (Kivi et al., 2014), and Tsukuba (Ohyama et al., 2009; Morino et al., 2014), whereas for Bremen, we used the GGG2012 retrievals. TCCON data were obtained from the TCCON Data Archive website at <http://tcon.ornl.gov/>.

3.2 Model CO₂ data

Because evaluation against TCCON is limited to 12 sites in the Northern Hemisphere, another validation source is necessary for obtaining a more thorough view of the accuracy of the GOSAT seasonal cycle. Therefore, we also analyzed XCO₂ from three models that assimilate in-situ CO₂ measurements to optimize their fluxes. The models were CarbonTracker (CT2013B; Peters et al., 2007, with updates documented at <http://carbontracker.noaa.gov>), MACC 13.1 (Chevallier et al., 2010, documentation and data available at <http://www.copernicus-atmosphere.eu/catalogue>), and the University of Edinburgh model (UoE; Feng et al., 2009, 2011, <http://www.palmergroup.org>). Relevant model properties are listed in Table 1. The models were resampled at GOSAT/ACOS observations in latitude, longitude and time, and integrated over all atmospheric layers to form the column-averaged CO₂. The ACOS averaging kernel correction was first considered for CT2013B, but as it had only a very minor effect on the total column (generally < 0.1 ppm difference in monthly averages), it was subsequently neglected for all models. However, seasonal effects of the averaging kernel correction

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are briefly assessed in Sect. 5.3. All model results were available from the beginning of GOSAT data (23 April 2009) but have different end dates: UoE and CT2013B run until the end of December 2012, and MACC 13.1 is available until the end of December 2013.

4 Methods

In this section, we describe the co-location of ground-based and satellite remote sensing measurements, filtering and bias correction for GOSAT/ACOS, and the averaging kernel correction, and define the average seasonal cycle. We demonstrate these steps with an example TCCON site at Park Falls, Wisconsin, US.

4.1 Co-locating GOSAT and TCCON

GOSAT/ACOS B3.5 XCO₂ observations were first co-located with the TCCON soundings, which were interpolated to local noon to exclude any effects from the diurnal cycle of XCO₂. The co-location can be done in several ways that were described and compared by Nguyen et al. (2014). In this study, we used the NOAA/Basu co-location technique that considers atmospheric transport of CO₂ in addition to spatiotemporal proximity of the TCCON and GOSAT observations (Guerlet et al., 2013a). This co-location technique has several advantages: high co-location data volume, good accuracy, and good sampling of parameter space, such as surface albedo.

The NOAA/Basu co-location technique works as follows. Temporally, any co-located observations need to be acquired on the same day, within 2 h of each other. The spatial region of matching TCCON and GOSAT changes dynamically based on how the inversion-derived estimates of local CO₂ surface fluxes are transported with the TM5 transport model: the region around a TCCON site over which modeled XCO₂ does not differ by more than 0.5 ppm from its value at the TCCON site sets the boundaries for co-location (as an upper spatial limit, GOSAT soundings need to be within $\pm 22.5^\circ$

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are presented in Fig. 3, in the lower left panel. For model results, the averaging kernel corrections were not applied.

Finally, we calculated daily averages of both GOSAT/ACOS and TCCON retrievals. This way, days with multiple soundings are not more dominant in the seasonal cycle fit than the days with fewer soundings. Time series of daily averages are shown in Fig. 3, in the lower right panel.

4.3 Seasonal cycle

In what follows, we parameterize the seasonal cycle of XCO_2 as a skewed sine wave with an upward trend, and find that it is generally a good model for the time series of XCO_2 in the Northern Hemisphere. We fitted an average seasonal cycle to the daily XCO_2 averages using the following six-parameter function

$$f(t) = a_0 + a_1 t + a_2 \sin(\omega[t - a_3] + \cos^{-1}[a_4 \cos(\omega[t - a_5])]), \quad (1)$$

where t is the time in days and ω is the annual period of 365 days. The first two terms with the parameters a_0 and a_1 (denoting the average growth rate) fit for a linear trend, and the third term, a sine wave with a time-dependent phase, fits for the seasonal cycle parameters a_2 – a_5 . In particular, $2|a_2|$ denotes the peak-to-peak amplitude of the sine wave and is, from here forwards, used to define the seasonal cycle amplitude. The nonlinear least squares fit was solved using a standard gradient-expansion algorithm. For Park Falls, the seasonal cycle fits for TCCON and ACOS are shown in Fig. 3, lower right panel, and the resulting seasonal cycle amplitude is 8.4 ± 0.1 ppm for TCCON, and 8.6 ± 0.2 ppm for ACOS. The errors of the fitted parameters are driven by the standard deviations σ of each daily XCO_2 , initially requiring $\sigma_{ACOS} \geq 1.5$ ppm and $\sigma_{TCCON} \geq 0.3$ ppm. Because the true errors in daily-averaged XCO_2 are not well known, we scaled the daily errors by multiplying them with the minimized quantity χ to yield $\chi^2 = 1$ from the least squares fit. For TCCON data fits, the original χ^2 values varied between $2 < \chi^2 < 10$, while for ACOS, the values were typically $\chi^2 < 1$, which implies that the initial errors σ_{TCCON} may have been underestimated and σ_{ACOS} overestimated.

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The fitting errors are purely statistical, and do not take into account systematic errors in the data. A more traditional Fourier series fit with an annual and semi-annual cycle (Wunch et al., 2013) was also tried, and the fitted seasonal cycle amplitudes were virtually identical (well within the fitting errors), but because some strange behavior during unobserved times of year could result, we opted for the fit in Eq. (1).

We recognize that there could be inter-annual variability in some or all of the fitted parameters, and that our results can be affected by that variability; especially we can expect sites with shorter co-located time series to be more sensitive. However, we do not fit for inter-annual variability because we are interested in identifying potential systematic errors in the average seasonal cycle captured by GOSAT and, in particular, the ACOS retrieval system. For the purposes of evaluating the average seasonal cycle of XCO₂, it is important to compare observations from the same time interval, which we take into account by co-locating the observations from TCCON and GOSAT.

5 Results and discussion

5.1 Evaluation against TCCON

Seasonal cycles for co-located TCCON and GOSAT/ACOS B3.5 XCO₂ soundings were studied at 12 TCCON sites in the Northern Hemisphere. Detrended average seasonal cycles for both retrievals at each site are shown in Fig. 4. Detrending removed a linear trend, i.e. XCO₂ average growth rate, that varied between 1.90–2.39 ppm year⁻¹ for ACOS and 2.02–2.58 ppm year⁻¹ for TCCON retrievals, depending on the site. We estimated the sensitivity of the average seasonal cycle parameters of Eq. (1) to the fitted trend from the error covariance matrix associated to the best-fit parameters. The error in the trend was generally weakly negatively correlated with the error in the seasonal cycle amplitude, for both TCCON and ACOS. The phase-related parameters a_3 – a_5 were not correlated with the trend. Therefore, the error from removing the trend should statistically have little effect on the parameters of the average seasonal cycle. Descrip-

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with the biased amplitude inferred by ACOS discussed below. While the maximum occurs within two spring months depending on location, the minimum is more seasonally restricted, varying from 15 August to 27 September for TCCON, and from 14 August to 25 September for GOSAT. During the minimum, the Northern Hemisphere receives solar light abundantly and is not snow-covered, so the number of co-located soundings is larger and the minimum is well captured by the satellite, within 6 days from TCCON, except for Tsukuba and Bremen. These values are generally in good agreement (within a few days) with Wunch et al. (2013, p. 9451), except for the TCCON seasonal cycle maximum date at the European sites Bialystok and Bremen. We evaluated the statistical errors of the dates of the maximum and minimum XCO₂ with a Monte Carlo approach, using the error covariance matrices associated with the fitted function parameters. On average, the TCCON maximum date had an error of 3.5 days, while the error for ACOS maximum date was 6.1 days. The corresponding average statistical errors for the date of the minimum were 2.2 days (TCCON) and 3.6 days (ACOS).

The seasonal cycle amplitudes are presented in Fig. 5a, in addition to Table 2. The amplitude is captured within the error bars of the regression at four sites: Izaña, Lamont, Saga, and Park Falls. The largest absolute differences are 1.6 ppm at Tsukuba and 1.4 ppm at Bremen, which are also the largest relative differences (28 and 18 %). Within 1.0 ppm difference, the amplitude is captured at most sites, excluding Orleans, Bremen, and Tsukuba. It should be noted that the latter only has data for two years and therefore substantial uncertainty in both the trend and amplitude, whereas the former two sites have sufficient data for evaluating an average seasonal cycle. A closer inspection of Figs. 4 and 5a reveals that the amplitude seen by GOSAT/ACOS is systematically shallower than TCCON at all five TCCON sites in continental Europe. This bias appears to be regionally very concentrated, because at the Northern European site Sodankylä, GOSAT captures the seasonal cycle reasonably well (within 0.8 ppm), considering the site suffers from data (and sunlight) deficiency in winter. Kulawik et al. (2015) noted the low bias as well, although they grouped all TCCON sites within latitudes 46–53°N together and found that, at this latitude range, the seasonal cycle of

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ACOS was biased low by 0.7 ± 0.7 ppm. Intuitively, a shallow-biased GOSAT seasonal cycle over Europe contradicts with the message from several recent flux inversion studies (Basu et al., 2013; Chevallier et al., 2014; Reuter et al., 2014), where the inversions using GOSAT XCO₂ observations inferred a stronger carbon sink over Europe compared to the inversions that assimilated in-situ measurements only, and compared with bottom-up inventories. However, according to the results by Reuter et al. (2014), these two results are not in a conflict. They showed in their regional flux inversion experiment that the sink enhancement is due to a North-West to South-East gradient in XCO₂ over Europe, and that most of the additional uptake takes place in Eastern Europe.

We explored several possible explanations for the low-biased seasonal cycle amplitude over continental Europe. First, we repeated the analysis using GOSAT/ACOS B3.4 retrievals (instead of B3.5), which have two constant aerosol types in the retrieval, different filtering, and bias correction. This did not have a systematic effect: the seasonal cycle amplitude of GOSAT increased at Bremen (+0.3 ppm) and Orleans (+0.5 ppm), and decreased at Bialystok (−0.2 ppm), Garmisch (−0.2 ppm), and Karlsruhe (−0.4 ppm).

We also studied the differences between TCCON GGG2012 and GGG2014 retrievals, and found that in the latter, the seasonal cycle amplitudes in Europe were shallower by up to 0.4 ppm (Orleans). The difference comes likely from the extended time series and the additional measurements present in the GGG2014 version. It is therefore possible that some of the discrepancy between GOSAT and Bremen TCCON is due to the use of the GGG2012 retrieval.

Next, we introduced variations to the co-location method to quantify its impact to the seasonal cycle amplitude. Our default co-location technique was the NOAA/Basu method with 0.50 ppm CO₂ gradient, maximum latitude difference 7.5°, and longitude 22.5°. We experimented with four modifications to it: (1) latitude 5.0°, longitude 15°, (2) latitude 2.5°, longitude 7.5°, (3) 0.25 ppm CO₂ gradient, and (4) 1.0 ppm CO₂ gradient. The latter increased the number of co-located points while the three former reduced it by making the co-location requirement stricter. We found that a smaller longitude-

latitude box and a tighter CO₂ gradient led to a better match-up in terms of the seasonal cycle amplitude at Bialystok (difference only 0.1 ppm), but not in other European sites where the difference either did not change or increased. The ACOS seasonal cycle amplitude at Garmisch site turned out to be highly dependent on the co-location details, varying from 5.0 to 5.9 ppm in these tests. The TCCON amplitudes changed typically only 0.1 ppm, but the fitting errors increased as the number of co-located soundings decreased. We also found that the co-location box dimensions had an impact on the seasonal cycle at JPL, which is located in the Los Angeles basin where large CO₂ gradients could be expected. With the default technique, the amplitude for ACOS was 0.5 ppm shallower than TCCON (10 % difference), but when decreasing the box size, the difference was reduced to 0.1 ppm (2 %).

In our last experiment, we tested the impact of the ACOS B3.5 bias correction for H gain over land; as Table 2 shows, all co-located soundings at the continental European sites were land gain H. We found that the bias correction increased the seasonal cycle amplitude at Park Falls by 1.4 ppm, mostly due to a correction for dust aerosol optical depth and surface albedo in the 2.1 μm band, but the bias correction had only a 0.1 ppm total impact on the amplitude at the European sites. It turned out that two of the bias correction parameters (related to the retrieved surface pressure and vertical CO₂ gradient) made the seasonal cycle over Europe consistently shallower by 0.3–0.4 ppm, depending on the site (see Fig. 5b). However, these parameters did not affect the seasonal cycle amplitude at Park Falls or Lamont, which are the two main sites used when optimizing the ACOS bias correction. An interesting finding is that removing these two terms from the bias correction made the ACOS seasonal cycle amplitude (Fig. 5b) and trend (not shown) agree better with TCCON at 10 of the 12 sites, even though it made the scatter worse in single-sounding statistics. This implies that the bias correction might be improved by designing it based on aggregated soundings in addition to single observation statistics.

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the algorithms is around 1 ppm. These results can be interpreted to support the ensemble median algorithm EMMA introduced by Reuter et al. (2013), which combines all individual retrievals into one data set that globally has the best agreement with TCCON.

The maximum and minimum days of the seasonal cycle reflect the drawdown season and are dependent on latitude and climate region. Both TCCON and GOSAT capture an earlier start of drawdown at the continental European sites compared to the other sites, the latest start being at the southernmost site, Izaña. The ACOS and NIES PPDF-S algorithms appear to be generally best in phase with TCCON regarding the date of maximum XCO₂. At the continental European sites, GOSAT and TCCON fits for the maximum day differ by several weeks, TCCON being systematically earlier. The minimum is better captured by all retrievals, with the spread varying from a few days to about 20 days; the performance of the individual algorithms is very site-specific.

5.3 Evaluation against models

The seasonal cycle amplitude of GOSAT/ACOS B3.5 was also compared to the inverse model systems MACC 13.1, CT2013B, and UoE in the Northern Hemisphere. As described in Sect. 3.2, these models have been optimized against assimilated flask and in-situ CO₂ measurements, though not exactly same data sets nor using the exact same weighting. For the comparison, latitudes from 0 to 70° were divided into 5° latitude bins (see Fig. 1 for the map), and the GOSAT/ACOS soundings within one latitude bin were collected into a single time series. The seasonal cycle was fitted on the daily averages of GOSAT/ACOS XCO₂ and the resampled models. The resulting seasonal cycle amplitudes are shown in Fig. 7. The amplitude increases significantly from the tropics towards high latitudes for both GOSAT and the models. Although the results are qualitatively similar, the models can show close to 2 ppm differences within latitude bands. ACOS is in excellent agreement to MACC from 0 to 50° N, whereas CT2013B and UoE have a shallower seasonal cycle from the tropics up to 35° N. Differences in the model seasonal cycle can be caused by a number of error sources, including their

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prior, transport, and inversion. Tropical and subtropical latitudes include large regions where the data constraint is weaker; therefore, the land surface prior (and its particular implementation) may impact the inversion results more than at those regions where the measurement network is dense. Both UoE and CT2013B use a variant of CASA as their prior biospheric flux model, as presented in Table 1 (in fact, CT2013B uses a unique combination of two flavors of CASA (Andy Jacobson, personal communication, 17 April 2015)). Even though different versions of CASA can differ in their seasonal cycle magnitude, our results may imply that the seasonal cycle of CASA fluxes is too shallow in some tropical regions or biomes. We first did the comparison using earlier versions of CarbonTracker (CT2011 and CT2013), and found that CarbonTracker and UoE results were nearly identical in these regions (see CT2013 and UoE in Figs. 7 and 8), which was surprising because the two models were different in every aspect (transport, in-situ data selection, inversion) except for their prior biospheric fluxes. However, a significant correction to the transport model's vertical mixing was introduced in CT2013B. This led to an increase of about 0.5 ppm in the CarbonTracker's seasonal cycle amplitude at all latitudes.

At 50–60° N in Fig. 7, ACOS agrees better with UoE and CT2013B. From 60 to 70°, ACOS has a higher seasonal cycle amplitude than most models. A similar result was also obtained by Belikov et al. (2014) using GOSAT/NIES v02.00 retrievals, NIES transport model, and LMDZ model. However, at high boreal latitudes, the satellite observations are associated with larger errors that are not reflected in the purely statistical fitting errors. ACOS results at these latitudes should therefore be interpreted with caution.

We tested how the ACOS bias correction and model averaging kernel correction affected the latitudinally averaged seasonal cycle amplitudes. The ACOS bias correction decreased the amplitude about 0.5 ppm at latitudes 10–40° N, but increased the amplitude at 40–70° N. The maximum increase was 1.0 ppm at latitudes 50–60° N, implying that before the bias correction, ACOS was in better agreement with MACC at these latitudes, but that after the bias correction, ACOS agreed better with UoE and CT2013B.

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Even though validation against models is part of the ACOS bias correction, the TCCON sites are likely to dominate the bias correction at mid-latitudes. We studied the potential seasonal impact of the averaging kernel correction for CT2013B. We found that the averaging kernel correction systematically decreased the model seasonal cycle amplitude in the Northern Hemisphere by 0.15 ppm on average. Overall, these changes are minor and do not affect our general conclusions about the model comparisons.

The latitudinal dependence of the CO₂ seasonal cycle amplitude has been previously shown in e.g. “the flying carpet” plot presented by Conway et al. (1994, Fig. 4), but we would like to emphasize that the amplitude can also depend on longitude. Especially in the mid-latitudes, its increase from west to east is notable; this is demonstrated in Fig. 8 for latitude band 45–50° N, where the seasonal cycle amplitude of GOSAT/ACOS is 6.4 ppm over the longitudes 180–120° W, and is doubled at 120–180° E. These GOSAT observations considered were taken over land, so in practice, this means that the seasonal cycle amplitude is dampened from the Eastern Asia over the North Pacific Ocean to the North-West US. In the lower troposphere, this dampening above 30° N latitude was shown by Nakazawa et al. (1992) who analyzed a three-year time series (1984–1986) of CO₂ measurements onboard container ships. The model results in Fig. 8 show a similar pattern of amplitude enhancement towards east, albeit the seasonal cycle amplitude of MACC is 2–3 ppm shallower compared to those of the other models and ACOS in the Eastern Asia. Despite this large discrepancy in the east where the data volume is small (see Fig. 8, right vertical axis), the zonally-averaged seasonal cycle amplitudes of MACC and ACOS agree within 0.1 ppm at the same latitude band (45–50° N). The CT2013B amplitudes are consistently higher than ACOS at all longitudes in Fig. 8, but they agree within 0.1 ppm in the Eastern Asia. Of the three models, UoE is most consistent with ACOS, agreeing about the seasonal cycle amplitude to within 1 ppm at these specific regions. The northern and mid-latitudinal regions of Asia are again regions where the in-situ measurement coverage is very limited, which explains the large spread between the individual model results.

6 Conclusions

The seasonal cycle of XCO₂ is profoundly connected to the biospheric fluxes that determine the global terrestrial net CO₂ sink. Satellite measurements of XCO₂ by the Greenhouse Gases Observing Satellite (GOSAT) and the Orbiting Carbon Observatory (OCO-2) expand the current in-situ measurement network tremendously and therefore have the potential to improve flux inversions. However, the satellite-measured seasonal cycle of XCO₂ can be affected by different retrieval biases, such as biases related to seasonally-varying parameters (e.g., surface albedo) and a sampling bias due to the seasonal variation in solar radiation. Mischaracterization of the seasonal cycle could lead to errors in the inverse model systems that assimilate satellite CO₂ data. Motivated by this, we evaluated the seasonal cycle of GOSAT observations using ACOS B3.5 retrievals from years 2009–2013.

Three independent approaches were used for the evaluation of the XCO₂ seasonal cycle: comparisons against the Total Carbon Column Observing Network (TCCON), other GOSAT retrievals (UoL v5.1, NIES v02.21, NIES PPDF-S v.02.11, and RemoTeC v2.35), and comparisons to optimized inversion models that assimilate in-situ measurements of CO₂. We found that ACOS captures the seasonal cycle amplitude of TCCON with an accuracy of better than 1.0 ppm at most of the 12 TCCON sites in the Northern Hemisphere considered in this study. As we also inferred the mean annual growth rate at each TCCON site in order to remove it, we found agreement of generally better than 0.2 ppm year⁻¹ in this quantity, with the ACOS-inferred growth rate most often being lower than TCCON. Over continental Europe, the seasonal cycle amplitude as measured by ACOS was biased low at all five sites, the largest difference being 18 % at Bremen. We also found that ACOS generally captured the seasonal cycle phase within a few days, except over Europe where the differences were 2–3 weeks, with ACOS measuring the date of maximum XCO₂ later than TCCON. Several other algorithms also had minor low biases in their seasonal cycle amplitudes over Europe. We explored the cause of the low bias for ACOS, and found that the bias correction param-

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eters related to the retrieved surface pressure and vertical CO₂ gradient were partially responsible, explaining 16–48 % of the difference. This suggests that the bias correction might benefit from considering aggregated soundings in addition to deviations at single-sounding level. Also, the selection of the co-located soundings was found to affect the seasonal cycle amplitude at few sites. Especially at JPL, which is in the Los Angeles basin, the agreement with TCCON improved notably when the co-location criteria were made sufficiently tight to not include soundings taken too far from the basin.

Model comparisons at latitudes 0–70° N revealed that qualitatively the models and satellite observations agreed well, but also that the model-to-model differences were (at most latitude bands studied) larger than model-to-ACOS differences. From the tropics up to 50° N, the zonally-averaged seasonal cycle amplitude of ACOS was in very good agreement with MACC 13.1, while between 50–60° N, ACOS agreed better with the University of Edinburgh model and CarbonTracker CT2013B. Both of the latter models had seasonal cycle amplitudes shallower than ACOS or MACC at tropical and subtropical latitudes, where the models lack direct constraints from measurements over land and are thus more affected by their prior fluxes (or by extra-tropical or ocean measurements through long-range transport). Therefore, the shallower seasonal cycle amplitude might be connected to their prior land surface models that are different variants of CASA. However, to verify this, one should investigate also the impact of transport, data assimilation, and inversion system differences. We also found that the longitudinal changes in the seasonal cycle amplitude at mid-latitudes can be notable. In particular, we showed that at 45–50° N latitudes, the amplitude of the GOSAT XCO₂ seasonal cycle doubles from the North-West US to Eastern Asia. The model results showed a gradient as well, although it was 1–3 ppm shallower, depending on the model.

As model-to-model differences in XCO₂ can be several ppm at regions poorly sampled by in-situ measurements, GOSAT observations that measure seasonal cycle amplitude to within 1.0 ppm, based on this study, could potentially be used directly (without elaborate inversions) to evaluate model differences at these regions. This idea is explored in more detail in a work under preparation (Lindqvist et al., 2015).

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Table 1. Models used in the evaluation of the GOSAT seasonal cycle.

Model	Biosphere	Transport	Resolution of the model run (lon × lat × time × layers)
CT2013B	CASA/GFED2 and CASA/GFED3.1	TM5/ERA-interim, ECMWF	3° × 2° × 3h × 25
UoE	CASA/GFED	GEOS-Chem/GEOS5	5° × 4° × 3h × 47
MACC 13.1	ORCHIDEE	LMDZ/ECMWF	3.75° × 1.9° × 3h × 39

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Table 2. Parameters describing the XCO₂ seasonal cycle for TCCON and bias-corrected GOSAT/ACOS B3.5. The fraction of gain H soundings over land is also shown. The validation sites are sorted according to their latitude.

Site	Time series (month year ⁻¹)	Retrieval	Growth rate (ppm year ⁻¹)	Amplitude (ppm)	Date of max. XCO ₂	Date of min. XCO ₂	Fraction of land gain H
Izaña	May 2009–Oct 2013	TCCON	2.41 ± 0.02	5.3 ± 0.1	16 May	19 Sep	12.2 %
		GOSAT	2.22 ± 0.04	5.3 ± 0.2	18 May	17 Sep	
Saga	Aug 2011–Oct 2013	TCCON	2.39 ± 0.09	6.7 ± 0.2	7 May	13 Sep	77.7 %
		GOSAT	1.92 ± 0.26	6.7 ± 0.4	28 Apr	14 Sep	
JPL	May 2011–Jun 2013	TCCON	2.34 ± 0.07	5.1 ± 0.2	2 May	27 Sep	87.2 %
		GOSAT	2.39 ± 0.11	4.6 ± 0.3	21 May	25 Sep	
Tsukuba	Aug 2011–Dec 2013	TCCON	2.58 ± 0.10	5.7 ± 0.2	23 Apr	10 Sep	91.9 %
		GOSAT	2.20 ± 0.22	7.3 ± 0.5	23 Apr	26 Aug	
Lamont	Apr 2009–Dec 2013	TCCON	2.33 ± 0.02	5.3 ± 0.1	4 May	20 Sep	96.5 %
		GOSAT	2.14 ± 0.03	5.2 ± 0.1	6 May	15 Sep	
Park Falls	Apr 2009–Dec 2013	TCCON	2.21 ± 0.03	8.4 ± 0.1	22 Apr	15 Aug	100 %
		GOSAT	2.16 ± 0.04	8.6 ± 0.2	27 Apr	14 Aug	
Garmisch	May 2009–Oct 2013	TCCON	2.03 ± 0.04	6.6 ± 0.1	25 Mar	27 Aug	100 %
		GOSAT	1.90 ± 0.07	5.7 ± 0.2	17 Apr	24 Aug	
Orleans	Aug 2009–Nov 2013	TCCON	2.29 ± 0.04	7.3 ± 0.1	30 Mar	28 Aug	100 %
		GOSAT	2.04 ± 0.07	6.2 ± 0.3	13 Apr	22 Aug	
Karlsruhe	Apr 2010–Oct 2013	TCCON	2.21 ± 0.06	7.4 ± 0.1	19 Mar	23 Aug	100 %
		GOSAT	2.01 ± 0.08	6.4 ± 0.2	1 Apr	27 Aug	
Bremen	Apr 2009–Apr 2013	TCCON	2.02 ± 0.09	7.9 ± 0.3	20 Mar	5 Sep	100 %
		GOSAT	1.91 ± 0.14	6.5 ± 0.4	8 Apr	22 Aug	
Bialystok	Apr 2009–Oct 2013	TCCON	2.18 ± 0.03	8.1 ± 0.1	16 Mar	18 Aug	100 %
		GOSAT	1.99 ± 0.06	7.5 ± 0.2	5 Apr	17 Aug	
Sodankylä	May 2009–Oct 2013	TCCON	2.15 ± 0.04	8.7 ± 0.3	16 Apr	15 Aug	100 %
		GOSAT	2.05 ± 0.09	9.5 ± 0.5	24 Apr	17 Aug	

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Figure 1. Twelve Northern Hemisphere TCCON sites used for GOSAT validation in this study.

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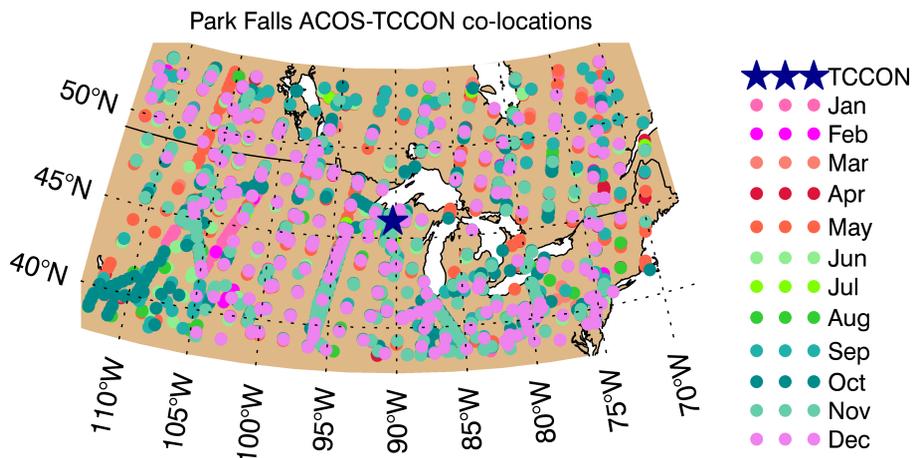


Figure 2. All co-located GOSAT/ACOS B3.5 soundings around the TCCON at Park Falls, Wisconsin, USA. The co-location technique was NOAA/Basu with 0.50 ppm CO₂ gradient, latitude limit of 7.5°, and longitude limit of 22.5°; see Guerlet et al. (2013a) for details of the method.

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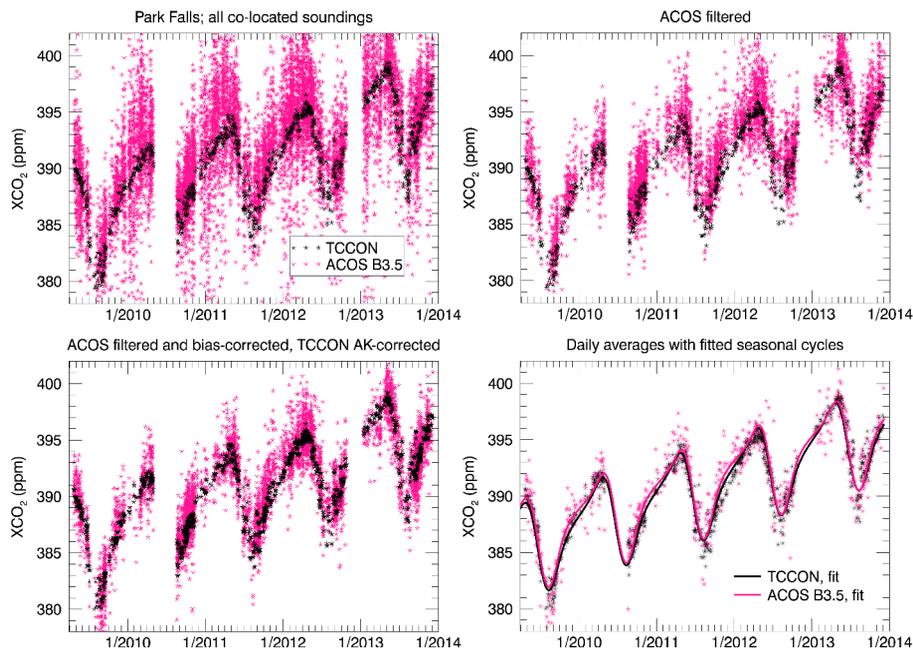


Figure 3. An example of data processing and the seasonal cycle fitting procedure at Park Falls. The upper left panel shows time series of the retrieved XCO₂ for all co-located TCCON (black) and GOSAT/ACOS (pink) soundings. The upper right figure shows only those ACOS L2 soundings that pass the post-processing filters. The lower left figure has bias correction applied for ACOS data and averaging kernel correction considered for TCCON soundings. The lower right panel shows the daily averages of XCO₂ and the respective seasonal cycle fits.

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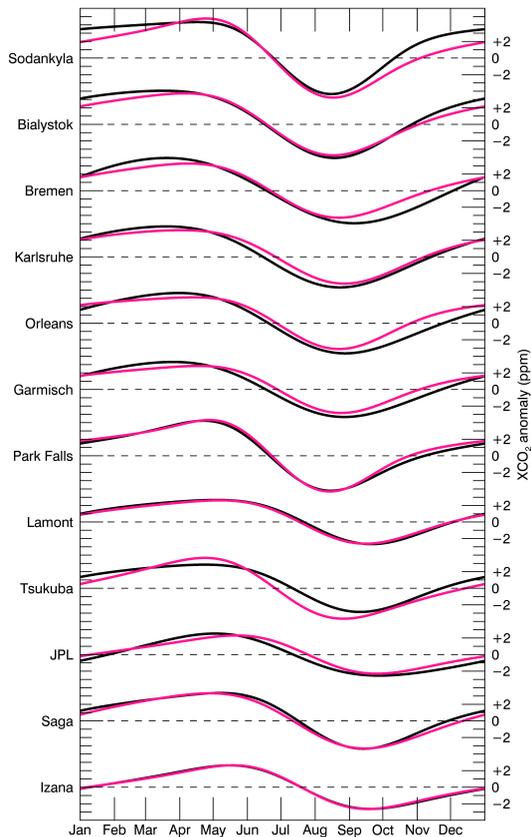


Figure 4. Detrended, best-fit seasonal cycles for GOSAT/ACOS (pink) and TCCON (black) at 12 validation sites in the Northern Hemisphere. The sites are organized according to their latitude (Izaña lowest, Sodankylä highest). On the vertical axis, one tick interval corresponds to 1.0 ppm XCO₂.

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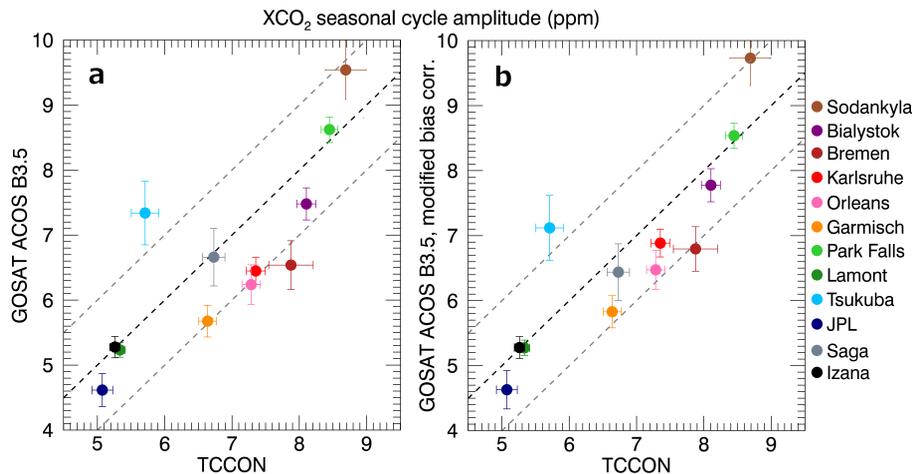


Figure 5. Seasonal cycle amplitude for ACOS (vertical axis) and TCCON (horizontal axis) for all the 12 NH sites used in the validation. The dashed black line corresponds to the one-to-one line, and the gray lines denote ± 1.0 ppm. Panel (a) shows the standard bias-corrected ACOS B3.5, and panel (b) shows ACOS B3.5 with a modified bias correction (see Sect. 5.1 for details).

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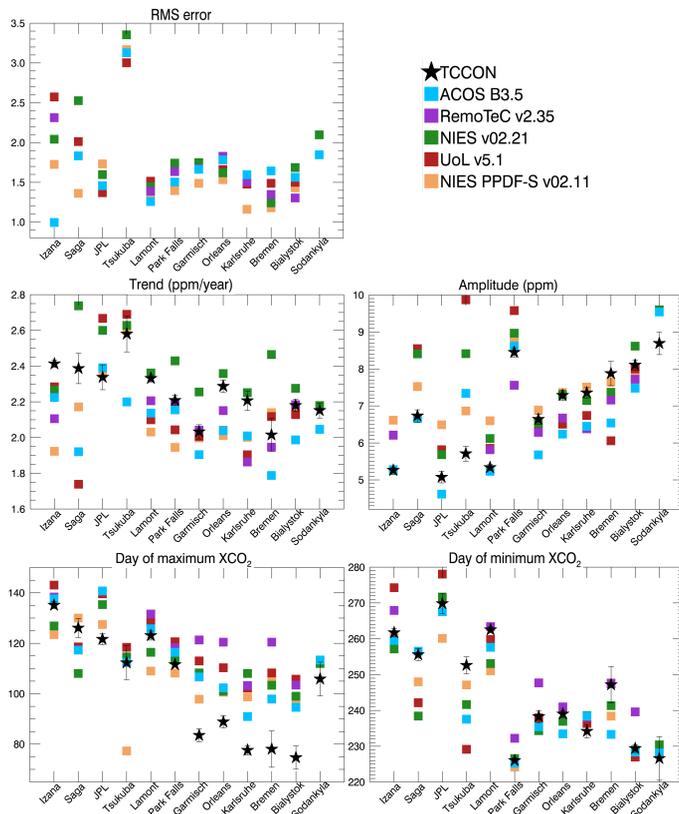


Figure 6. Comparison of the GOSAT and TCCON XCO₂ time series using the following parameters: root-mean-square (RMS) error (upper left panel), average trend (middle left panel), seasonal cycle amplitude (middle right panel), and the days of maximum and minimum XCO₂ (bottom row). Five retrieval algorithms were included to describe GOSAT observations. TCCON values were based on ACOS B3.5 co-located soundings. The 12 Northern Hemisphere validation sites are shown on the horizontal axis, their latitude increasing from left to right.

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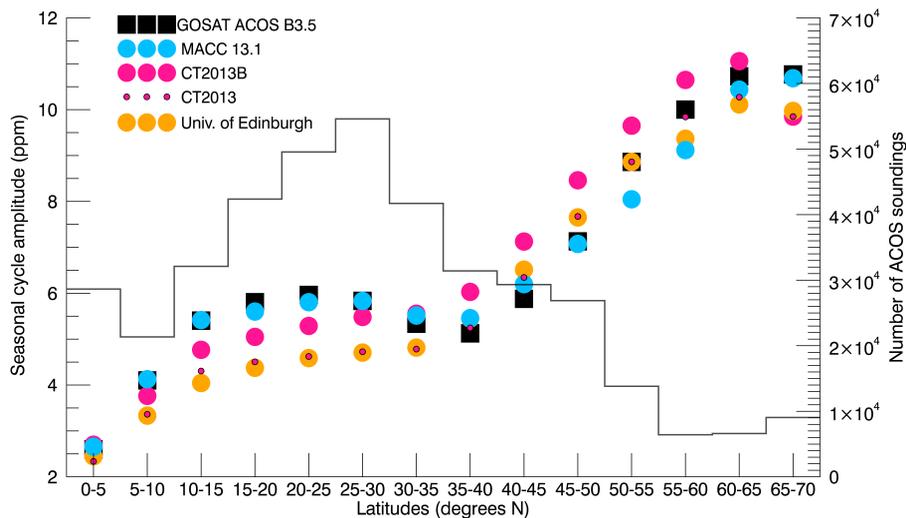


Figure 7. Latitudinal dependence of the seasonal cycle amplitude for bias-corrected ACOS B3.5 soundings and for three models resampled at the satellite soundings. For CarbonTracker, we show both CT2013 and CT2013B results, their difference being a major correction in the TM5 transport model. The left vertical axis shows the seasonal cycle amplitude in ppm, while the right vertical axis indicates the number of soundings that fall within each 5° latitude band.

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