# A point-by-point reply to the comments

#### **Reviewer 1#**

5 I suggest the authors check the data again. Because all the discussions are based on the good quality data.

The explanation of the data inconsistency between figure 1 and previously published paper are still confusing.

The explanation indicated that the fitted data (background or polluted cases) have higher AR than the observed data (all averaged together). Does that suggest you need to check your fitting criteria/methods?

In addition, if there is calibration and correction method applied to the raw data. It will affect both observed data and fitted data. Why only observed data will have a slight lower MAF?

- Re: Thanks for the suggestion. We do check the data carefully again. We are confident in the data
   quality. The reviewer is right, and if the correction is applied to the raw data, both the observed and fitted data will be affected. Thus, we also got a slight lower MAF for the fitted data (not shown in Figure 1). In the previously published paper, we didn't ignore and process some data points when the MAF value >1, which resulted in larger mean MAF. But in this paper, the data points with MAF > 1.0 were forced to 1 when D<sub>n</sub>>300 nm, which we thought could be activated
- 20 completely at even lower supersaturations but the MAF would never be >1.0. This has been stated in the revised paper (please see Page 11, lines 199-204).

#### **Reviewer 2#**

25 P16147 L10 These "relative deviations" are because of particle loss in the nafion dryer? Also, it is not clear what the "kinetic limitations" are.

Re: the relative deviations are due to the influence of dehydration-related particle mobility changes inside DMA. According to Mikhailov et al., 2009, void fractions as well as residual water in dried aerosol particles that are not water-free (due to kinetic limitations of drying or stable

30 hydrate formation) should be taken into account in Kohler model calculations of hygroscopic growth and CCN activation.

First, if the dehydration processes (efflorescence, restructuring, or desorption) inside DMA are completed within 0.1 s, then the resulting changes in particle mobility diameter should be fully captured with deviations <1%. And also, kinetically limited dehydration processes that lead to

35 progressive changes of particle mobility on a time scale of 0.1–10 s should significantly influence the particle sizing (deviations >1%) and lead to a broadening of the measured size distributions. Dehydration processes progressing on time scales >10 s should have no effect on particle sizing (no change of mobility diameter and no broadening of size distribution).

Thanks for your explanation. I agree that void fractions and residual water should be taken intoaccount; however, how can you do this with the measurements you have made?

Here is how I would proceed: You are making the assumption that drying imposed by the nafion, and by the dry sheath air inside the DMA, are sufficient to remove water associated with the ambient particles. Why not state this assumption, and move on without reference to "kinetic limitations" or Mikhailov et al. (2009)?

Re: Yes, you are correct, and your statement is exactly our assumption. Thank you! It has been revised (see Page 7, lines 107-108).

P16147L25

5 It is the inner diameter, not the outer diameter, that is relevant.
Re: it is just the outer diameter.
It's the inner diameter that is important. Please report that value.
Re: The inner diameter is 0.38 inch. (Revised)
P16148L19

- 10 "..temperature stability was zero." I don't understand what you are referring to here. Re: here it means the data is invalid if the "temperature stability" was flagged as "0". And for the valid data, the "temperature stability" was flagged as "1". The sentence has been revised. Thanks for your explanation. Why can't you explain what is meant by "stability"? Isn't this the average difference between what is preset and what is measured?
- **15** Re: sorry for not explaining it clearly. But you've got the point. The "temperature stability" refers to the  $T_1$ ,  $T_2$  and  $T_3$  in cloud chamber of the CCNc, which is set to obtain the target supersaturations. If the average differences between preset  $T_1$ ,  $T_2$  and  $T_3$  and the measured

values are larger than 0.4  $\,\,^\circ\text{C}$  , the "temperature stability" is flagged as "0". Thus, the data is

invalid and will be removed. This has been revised in the paper( see Page 9, lines 150-153)

#### 20

#### P16148L21

Here you define the "aerosol number (CN) size distribution spectrum." How is this different from the PSD mentioned on P16147L1 and on P16148L24?

Re: that's exactly the same thing. We have corrected all of them to particle number size distribution in the revised paper.

PSDs are shown in Figure 2 (revision). It is not clear how the quantity on the Y axis, once integrated over all sizes, becomes the CN concentration referred to in the text. For example, on L241 we are told that the CN concentration is  $\sim$ 1.7x104 cm-3. Typically in aerosol science the PSD is dN/dlogD and the latter has dimension cm-3. How are we to interpret the "N" on the Y

# 30 axis of Figures 2a-c? For an example see here in their Figure 12a (Atmos. Chem. Phys., 13, 7263–7278, 2013).

Re: Thanks for pointing this. It should be dN/dlogD (cm-3) in Figure 2. The figures have been revised. (see Figure 2)

#### 35 P16152L6

It is the "CCN activity", not the "aerosol activity", that is the focus here.

Re: corrected.

Please see L227 in the revision. This was not "corrected." If you do not feel the correction is needed, just say so in your response.

40 Re: sorry for my carelessness, I agree your suggestion. It has been corrected in the revision. (see Page 13, line 241)

P16153L26

Here is a relevant reference.

Snider, J.R., and, S.Guibert, J.-L. Brenguierand J.-P.Putaud, Aerosol activation in marine stratocumulus clouds: Part – II Köhler and parcel theory closure studies, J. Geophy. Res., 108, doi:10.1029/2002JD002692, 2003

Re: the reference has been added.
 I see this paper in the References, but I do not see it cited in the text.
 Re: Added it this time. (Page 4, line 27)

# P15154L7

10 What are "bulk ARs"?

Re: bulk ARs means the ratio of NCCN to NCN, which is calculated from the total CN and CCN number concentrations. Thus, we called it bulk ARs. For the size-resolved CCN measurements, we can get size-resolved ARs from size-resolved CCN and CN number concentrations.

Thanks for your explanation; however, what you wrote (revision) needs to be removed from the
 middle of the paragraph that discusses the size-resolved AR. I recommend that the removed text beput somewhere else (e.g., at the end of the paragraph). For example, on L52, you could state: "The bulk activation ratio (bulk AR) is defined as the CCN-measured concentration divided by the CN concentration. These values were measured in ambient air, every ?? minutes, and without particle size selection in the DMA."

20 Re: Thanks a lot for your comments. In this revision, this has been revised as follows (also see Page 10, lines 161-166),

"The size-resolved CCN activation ratio (size-resolved AR) is defined as the  $dN_{CCN}/dlogD_p$  divided by the  $dN_{CN}/dlogD_p$ . These values were measured by SMPS-DMT-CCNc with particle size selection in the DMA. The bulk activation ratio (bulk AR) is defined as the total CCN account of the state of

25 concentration divided by the total CN concentration. The total CCN and CN number concentrations are integrated by the measured CCN and CN size distribution respectively over the whole size range."

P16155L2What you are calling the "estimate" is the summed product of AR(D) and PSD(D)? Right?
By "Observation" you are talking about the direct measurement of the ambient Nccn(S) made \_\_without\_ the DMA in front of the CCN instrument. I did not see mention of the ambient Nccn(S) measurement (without the DMA in front) in Section 2.1.

 Re: We used a Scanning Mobility Particle Sizer (SMPS), combined with a Droplet Measurement Technologies-Cloud Condensation Nuclei Counter (DMT-CCNc) (Lance et al., 2006), for
 size-resolved CCN measurements as well as particle number size distribution (PSD)

measurements. The SMPS is just the DMA. To estimate NCCN, estimated CCN size distributions at the five supersaturations were calculated by multiplying the campaign-averaged CCN efficiency spectrum with the actually measured PSD. The estimated NCCN at the five supersaturations was then calculated by integrating the

40 estimated CCN size distribution over the whole size range. The measured CCN size distributions are integrated to produce the observed NCCN. Thanks for the explanation. In the revision, do you explain what you mean by the "observed NCCN"?

Re: Yes. This has also been explained in the text. (see Page 16, lines 321-328; Page 18, lines

#### 376-379)

P16155L19

5

20

I would reword this because the Figure 7 shows how the difference (estimated minus observed) varies with chi-org. The latter is the independent variable. There are other places in the manuscript where "sensitivity" is used. I would change the word order in some of these instances too. E.g., P16145L26. There are other places too.

Re: revised. Because in the revised version, the sensitivity of oxidation level (using f44, the fraction of m/z 44 in total organics, as an indicator) of organics to estimatimation of NCCN is also

examined. The section is thus rewritten (see Section 4.4, and Figure 5 and Figure 6).In my opinion this argument is presented backwards, both in the original manuscript, and in the revision.

Here is an example from the abstract (revised):

"The sensitivity of volume fraction of organic aerosols (chi-org) as well as oxidation level (using f44, the fraction m/z 44 in total organics, as an indicator) of organics on estimating NCCN is

15 f44, the fraction m/z 44 in total organics, as an indicator) of organics on estimating NCCN is examined."

I feel that the authors should reword these statements. Here is a suggested replacement sentence for the sentence above:

"The sensitivity of the estimated CCN number concentration (NCCN) to both volume fraction of organic material (chi-org) and aerosol oxidation level (using f44, the fraction m/z 44 in aerosol organic material) are examined."

Re: thank you very much for correcting the presentation regarding to this part. We totally agree to your comments and correction. That has been revised carefully (see Page 2, lines 2-5)

Another example is seen on L65 (revision). Going back to what I said in my first review, I feel
that the authors should be focused on how NCCN is sensitive to aerosol chemical composition, not the other way around.

Re: revised as "...The aim of this paper is to examine the sensitivity of CCN activity to aerosol physicochemical properties (especially aerosols containing large amounts of organics, as well as the oxidation level)..." (see pages 6, lines 66-68)

30 Another example is L294 (revision).

Re: revised as "...we examine the sensitivity of  $N_{CCN}$  to both volume fraction of organics ( $x_{org}$ ) and oxidation or aging of organics based on measurement at Xinzhou site...." (see Page 16, lines 307-309)

Another example is the caption of Figure 5 (revision).

**35** Re: revised as "...The sensitivity of  $N_{CCN}$  to both organics volume fraction ( $x_{org}$ ) and oxidation level (using  $f_{44}$ , the fraction of m/z 44 in aerosol organic material) of organics at supersaturation levels of" (see the caption of Figure 5)

Another example is L396 (revision)

Re: revised as "...we examine the sensitivity of N<sub>CCN</sub> to both volume fraction of organics (x<sub>org</sub>) and
 40 oxidation or aging of organics based on measurement at Xinzhou site." (see Page 20, lines 409-412)

#### Additional comments:

According to Petters et al. (ACP, 2007), the kappa for ammonium sulfate is 0.61 and the kappa for

ammonium nitrate is 0.67 (see their Table 1; CCN derived kappa). Hence, the sentence on L181 (revision) is backwards.

Re: revised. Thanks a lot for your careful check.

L117 (revision) - here you are referring to the "overall relative error"

5 Re: corrected. Thank you very much.

L275 (revision) – The sentence needs work (...."it should be caution")

Re: revised. (Page 15, lines 289-290)

L345 (revision) – misspelled word "usully")

Re: corrected.

10 L388 (revision) – This mentions "local primary biomass burning" but I do not see discussion of the topic in the analysis section of the manuscript.

Re: revised as "...CCN efficiency was largely reduced by local air masses..." (see Page 19, line 401) General – if the subscript is "org" for chi, it should also be "org" for epsilon (i.e., not "Org"). Re: corrected.

15

20

# Impacts of organic aerosols and its oxidation level on CCN activity from measurement at a suburban site in China

Fang Zhang<sup>1,2</sup>, Zhanqing Li\*<sup>1,2,3</sup>, Yanan Li<sup>1,2</sup>, Yele Sun<sup>4</sup>, Zhenzhu Wang<sup>5</sup>, Ping Li<sup>1,2</sup>,

Li Sun<sup>6</sup>, Maureen Cribb<sup>3</sup>, Chuanfeng Zhao<sup>1,2</sup>, Qingqing Wang<sup>4</sup>

<sup>1</sup>State Key Laboratory of Earth Surface Processes and Resource Ecology, College of Global Change and Earth System Science, Beijing Normal University, Beijing 100875, China

25 *Ch* 

<sup>2</sup>Joint Center for Global Change Studies, Beijing 100875, China
 <sup>3</sup>Earth System Science Interdisciplinary Center and Department of Atmospheric and Oceanic Science, University of Maryland, College Park, Maryland, USA.
 <sup>4</sup>State Key Laboratory of Atmospheric Boundary Layer Physics and Atmospheric

30 *Chemistry, Institute of Atmospheric Physics, Chinese Academy of Sciences, Beijing* 100029, *China* 

<sup>5</sup>Key Laboratory of Atmospheric Composition and Optical Radiation, Anhui Institute of Optics and Fine Mechanics, Chinese Academy of Sciences, Hefei 230031, China <sup>6</sup>Liaoning Weather Modification Office, Shenyang, 112000, China

35

\*correspondence to: Z. Li (zli@atmos.umd.edu)

# Abstract

This study is concerned with the impacts of organic aerosols on CCN activity based on field measurements made at a suburban site in north China. The sensitivity of the estimated CCN number concentration (N<sub>CCN</sub>) to both volume fraction of 5 organic aerosols material  $(x_{org})$  as well as and aerosol oxidation level (using  $f_{44}$ , the fraction of m/z 44 in total organics, as an indicator) of organics on estimating N<sub>CCN</sub> isaerosol organic material) are examined. A strong dependence of CCN number concentration (N<sub>CCN</sub>) on the  $x_{org}$  and  $f_{44}$  was noted. The sensitivity of volume fraction of organics to  $N_{CCN}$  increased with increasing  $x_{org}$ . The impacts of the aerosol particles 10 oxidization or aging level on estimating N<sub>CCN</sub> were also very significant. When the particles were mostly composed of organics ( $x_{org}$ >60%), the N<sub>CCN</sub> at the supersaturation of 0.075% and 0.13% was underestimated by 46% and 44% respectively if aerosol particles were freshly emitted with primary organics ( $f_{44} < 11\%$ ); while the underestimation decreased to 32% and 23% at the corresponding 15 supersaturations if the particles were with more hygroscopic secondary organics( $f_{44}$ >15%). The N<sub>CCN</sub> at the supersaturation of 0.76% was underestimated by 11% and 4% respectively at  $f_{44}$ <11% and  $f_{44}$ >15%. But for the particles composed of low organics (e.g.  $x_{org} < 40\%$ ), the effect caused by the  $f_{44}$  was quite insignificant both at high and low supersaturations. This is due to that the overall hygroscopicity of the particles is dominated by inorganics such as sulfate and nitrate, which are more 20 hygroscopic than organic compounds. Our results indicated that it would decrease the uncertainties in estimating N<sub>CCN</sub> and lead to a more accurate estimation of N<sub>CCN</sub> to

1	带格式的
---	------

1	带格式的
-(	带格式的
1	带格式的
-{	带格式的
-	带格式的



increase the proportion of secondary organics, especially when the composition of the aerosols is dominated by organics.

5

The applicability of the CCN activation spectrum obtained at Xinzhou to the Xianghe site, about 400 km to the northeast of Xinzhou, was investigated, with the aim of further examining the sensitivity of  $\underline{\text{CCNN}}_{\underline{\text{CCN}}}$  to aerosol type. Overall, the mean CCN efficiency spectrum derived from Xinzhou performs well at Xianghe when the supersaturation levels are >0.2% (overestimation of 2-4%). However,  $N_{CCN}$  was overestimated by ~20% at supersaturation levels of < 0.1%. This suggests that the overestimation is mainly due to the smaller proportion of aged and oxidized organic aerosols present at Xianghe compared with Xinzhou.

10

#### 11 **1. Introduction**

To reduce the uncertainty of aerosol indirect effects on the radiative balance of 12 the atmosphere, it is important to gain a good knowledge of the ability of aerosol 13 particles to form cloud condensation nuclei (CCN) at the typical supersaturations 14 found in the atmosphere. The CCN activity of aerosol particles is governed by the 15 Köhler theory (Köhler, 1936). This theory determines CCN from aerosol particle size 16 17 and physicochemical properties, which include the molar volume, activity coefficient, and effect on surface tension (McFiggans et al., 2006). These properties, however, are 18 19 difficult to measure.

20 Researchers have proposed single-parameter models to parameterize the CCN activation and hygroscopicity of multi-component aerosols (Hudson and Da, 1996; 21 Rissler et al., 2006; Petters and Kreidenweis, 2007; Wex et al., 2007). Field 22 experiments have been conducted with the aim of better characterizing particle 23 physicochemical parameters influencing cloud CCN activation. Due to the large 24 25 spatial variability of aerosol types and compositions, the CCN activation efficiency varies greatly over different regions. CCN number concentrations (N<sub>CCN</sub>) can often be 26 better predicted in the background atmosphere (Chuang et al., 2000; Dusek et al., 27 28 2003; VanReken et al., 2003; Snider et al., 2003; Rissler et al., 2004; Gasparini et al., 2006; Stroud et al., 2007; Bougiatioti et al., 2009). 29

The largest errors are associated with urban emissions (Sotiropoulou et al., 2007). This is likely due to the organics component of aerosol particles, which have the largest uncertainty and are not fully understood. Biomass burning aerosols and secondary organics formed from the oxidation of common biogenic emissions are often more difficult to activate (Mircea et al., 2005; VanReken et al., 2005; Lee et al., 2006; Varutbangkul et al., 2006; Clarke et al., 2007; Rose et al., 2010; Engelhart et al.,

2012; Paramonov et al., 2013; Lathem et al., 2013; Mei et al., 2013b; Zhang et al., 36 37 2014). Particles with aged/oxidized secondary organic components (e.g., organic acids) have been shown to be more hygroscopic (Raymond and Pandis, 2002; Hartz et al., 38 2006; Bougiatioti et al., 2011), but still much less hygroscopic than inorganic species. 39 The sensitivity of estimated N<sub>CCN</sub> to organics have been examined in a number of 40 recent studies (Wang et al., 2008; Reutter et al., 2009; Ervens et al., 2010; 41 Kammermann et al., 2010; Ward et al., 2010; Zhang et al., 2012; Mei et al., 2013a). It 42 43 is widely known that the estimated N<sub>CCN</sub> is sensitive to changes in organics due to the latter's complex components. The amounts and hygroscopicity parameter of organics 44  $(\kappa_{org})$  vary substantially and lead to significant biases in estimating CCN 45 concentrations and aerosol indirect forcing (Sotiropoulou et al., 2007; Hings et al., 46 2008; Liu and Wang, 2010). Therefore, field investigations regarding CCN activity 47 48 and organics impacts, especially in heavily polluted regions, are pivotal to better parameterize CCN in climate models. 49

Northern China is a fast developing and densely populated region of China, 50 where aerosol loading is high (Li et al., 2007, 2011), the particle composition is 51 complex, and severe haze pollution episodes are common (Guo et al., 2014). In recent 52 years, CCN measurements have been collected during field campaigns carried out in 53 the region (Wiedensohler et al., 2009; Gunthe et al., 2011; Yue et al., 2011; Deng et 54 al., 2011, 2013; Zhang et al., 2014). These studies have presented different 55 56 perspectives on the influence of particle size and composition on CCN activity. For example, Deng et al. (2013) evaluated various schemes for CCN parameterization and 57 recommended that the particle number size distribution (PSD) together with inferred 58 mean size-resolved activation ratios can be used to estimate CCN number 59 concentrations without considering the impact of particle composition. However, 60

associated with changes in particle composition. None of the above-mentioned studies have investigated the impact of organics on estimating  $N_{CCN}$  in Northern China. Zhang et al. (2012) noted a more significant influence of organics on CCN activity but without concerning the influences of particles oxidation or aging on CCN activity; in addition, the campaign average mass fraction of organics in their study was < 20%.

The aim of this paper is to examine the sensitivity of changes in CCN activity to 67 68 aerosol physicochemical properties (especially aerosols containing large amounts of organics, as well as the oxidation level) to CCN activity.), and also to see how much 69 uncertainty is incurred by applying the CCN efficiency spectra measured at one site to 70 another site in a heavily polluted region. The instrumentation and data used in the 71 study are described in section 2. The method for calculating the hygroscopicity 72 73 parameter ( $\kappa_{chem}$ ) is introduced in section 3. The sensitivity of  $x_{org}$  as well as oxidation 74 level of organics on estimating N<sub>CCN</sub> in section 4, and the ability of the CCN efficiency spectrum observed at the Xinzhou site to represent CCN at the Xianghe site, 75 76 are also presented and discussed at the last part of this section. Conclusions from the study are given in section 5. 77

## 78 2. Measurements and data

61

79 intensive observation An period field campaign similar to the Aerosol-CCN-Cloud Closure Experiment (Zhang et al., 2014), called the Atmosphere, 80 Aerosol, Cloud, and CCN (A<sup>2</sup>C<sup>2</sup>) experiment, was conducted from 22 July to 26 81 August of 2014 at Xinzhou (38.24°N, 112.43°E; 1500 m above sea level), a city with a 82 83 population of 0.51 million in Northern China. The site is located about 360 km southwest of the metropolitan Beijing area and about 10 km south of the local town 84 85 center. The site is surrounded by agricultural land (e.g., corn) with little local pollution

plums from motor vehicles and industrial activities. Sitting between two mountains 86 87 (Taihang Mountain to the east and Lüliang Mountain to the west), the site also experiences air masses from Xinzhou City to the north and from Taiyuan City to the 88 south, the capital of Shanxi Province. Air masses from the northeast and southwest 89 dominate over the site during summer. Depending on the wind direction, 90 measurements at the Xinzhou site can detect air parcels of urban, rural, or mixed 91 origins, including both fresh biogenic emissions around the site and aged aerosols 92 93 from advection.

#### 94 **2.1 Instruments and measurements**

During the field campaign, a Scanning Mobility Particle Sizer (SMPS),
combined with a Droplet Measurement Technologies-Cloud Condensation Nuclei
Counter (DMT-CCNc) (Lance et al., 2006), was used for size-resolved CCN
measurements as well as particle number size distribution (PSD) measurements. The
measured aerosol PSD is within the size range of 14-600 nm. Aerosol chemical
composition was measured simultaneously by an Aerodyne Aerosol Chemical
Speciation Monitor (ACSM) (Sun et al., 2012).

The aerosol inlet for the size distribution measurements was equipped with a TSI 102 Environmental Sampling System (Model 3031200), which consists of a sharp-cut 103 104 PM<sub>1</sub> cyclone and a bundled nafion dryer. The size-resolved CCN efficiency spectra 105 were measured by coupling the DMT-CCN<sub>C</sub> used with the SMPS (Rose et al., 2008). In this step, the particles are rapidly dried with RH < 30% upon entering the 106 Differential Mobility Analyzer (DMA). Thus, size selection is effectively performed 107 under dry conditions. The nation dryer and the sheath air inside the DMA are 108 sufficient to remove residual water associated with the ambient particles. Relative 109

deviations in particle diameter should be < 1% except for potential kinetic limitations 110 111 (Mikhailov et al., 2009).%. The sample flow exiting the DMA was split into two parts: 0.3 lpm for the CPC and 0.5 lpm for the CCN counter ( $CCN_c$ ). The DMA, controlled 112 113 by TSI-AIM software, scanned one size distribution every five minutes. The CCN<sub>C</sub> was operated at a total flow rate of 0.5 lpm with a sheath-to-aerosol flow ratio of 10. 114 115 The inlet RH for CCNc was < 30%. During the field campaign, the mean sample temperature and pressure measured by  $CCN_C$  sensors was  $(24.3\pm1.4)$  °C and 116 117 (898.4±11.7) hPa. The supersaturations levels of CCN<sub>C</sub> were calibrated with ammonium sulfate before and after the field campaign, following the procedures 118 outlined in Rose et al. (2008). During each CCN measurement cycle, calibrated 119 effective supersaturations were set at 0.075%, 0.13%, 0.17%, 0.39%, and 0.75%. The 120 overall <u>relative</u> error  $(1\sigma)$  for the supersaturation levels was estimated to be < 3.5%. 121 122 The completion of a full measurement cycle took 50 minutes (10 minutes for each 123 supersaturation level).

The measurement of non-refractory submicron aerosol species including 124 125 organics, sulfate, nitrate, ammonium, and chloride were made with an ACSM. During the field campaign, ambient aerosols were drawn inside through a 1/2 inch (outer 126 diameter, the inner diameter is 0.38 inch) stainless steel tube at a flow rate of ~3 L 127 min<sup>-1</sup>, of which ~84 cc min<sup>-1</sup> was sub-sampled into the ACSM. An URG cyclone 128 (Model: URG-2000-30ED) was also positioned in front of the sampling inlet to 129 remove coarse particles with a cut-off size of 2.5 µm. Before sampling into the ACSM, 130 aerosol particles were dried using a silica gel desiccant. The residence time in the 131 sampling tube was ~5 s. The ACSM was operated at a time resolution of ~15 min with 132 a scan rate of mass spectrometer at 500 ms  $amu^{-1}$  from m/z 10 to 150. Regarding the 133 calibration of the ACSM, mono-dispersed, size-selected 300-nm ammonium nitrate 134

particles within a range of concentrations were sampled into both the ACSM and a condensation particle counter (CPC). The ionization efficiency (IE) was then determined by comparing the response factors of the ACSM to the mass calculated with known particle size and number concentrations from the CPC. More detailed descriptions of the operation and calibration of the ACSM are given in Sun et al. (2012) and Ng et al. (2011). The campaign averaged mass concentration of PM<sub>1</sub> is  $31.6 \ \mu g \ m^{-3}$ .

In addition to the ACSM, the black carbon (BC) in  $PM_{2.5}$  was simultaneously measured at a time resolution of 5 min by a seven-wavelength aethalometer (Model AE31, Magee Scientific Corporation). The campaign averaged mass concentration of BC is ~2.5 µg m<sup>-3</sup>. During the experiment, the campaign area was generally hot and dry, with an average temperature of 21.6 °C and an average ambient RH of 69.5%.

# 147 2.2 Data

The raw CCN data were first filtered according to instrument recorded 148 parameters (e.g., temperature and flow). For example, if the relative difference 149 150 between the actual and preset sample flows was larger than 4%, the data are flagged as invalid. The data is also excluded if the "temperature stability" was flagged as "0". 151 These flagged data are not used for further analysis. Here, the "temperature stability" 152 153 refers to the  $T_1$ ,  $T_2$  and  $T_3$  in cloud chamber of the CCNc, which is set to obtain the target supersaturations. If the average differences between preset  $T_1$ ,  $T_2$  and  $T_3$  and the 154 measured values are larger than 0.4 °C, the "temperature stability" is flagged as "0". 155 Thus, the data is invalid and will be removed. A multiple charge correction and 156 transfer function (Deng et al., 2011) is applied to each PSD as well as to the CCN 157 efficiency spectrum. The CCN activation ratio (AR) is the ratio of N<sub>CCN</sub> to CN 158

|带格式的

159	concentration (N <sub>CN</sub> ). Bulk AR is calculated from the total N <sub>CN</sub> and N <sub>CCN</sub> . For the								
160	size-resolved CCN measurements, we get size-resolved AR from size-resolved CCN								
161	and CN number concentrations.								
162	——Size-resolved CCN and PSD data, measured with a DMT-CCNc and a-								
163	SMPS (with a particle size range of 10-700 nm) on 7-21 July 2013 at Xianghe (Zhang								
164	et al., 2014), are used in this study for comparisons with CCN activity at the Xinzhou								
165	site. Aerosol mass concentrations were processed using the ACSM standard data								
166	analysis software (version 1.5.3.0). Detailed procedures for the data analysis have								
167	been described by Ng et al. (2011) and Sun et al. (2012). The size-resolved CCN								
168	activation ratio (size-resolved AR) is defined as the dN <sub>CCN</sub> /dlogD <sub>p</sub> divided by the								
169	$dN_{CN}/dlogD_p$ . These values were measured by SMPS-DMT-CCNc with particle size								
170	selection in the DMA. The bulk activation ratio (bulk AR) is defined as the total CCN								
171	concentration divided by the total CN concentration. The total CCN and CN number								
172	concentrations are integrated by the measured CCN and CN size distribution								
173	respectively over the whole size range.								

174 **3. Derivation** of *κ*<sub>chem</sub>

In this study, we calculate  $\kappa_{chem}$  based on bulk chemical composition observations made during the field campaign. The method is very similar to that used by Zhang et al., (2014). As proposed by Petters and Kreidenweis (2007),  $\kappa_{chem}$  can be predicted using a simple mixing rule based on chemical volume fractions for a given internal mixture:

180 
$$\kappa_{chem} = \sum_{i} \varepsilon_i \kappa_i \tag{1}$$

181 where  $\kappa_i$  and  $\varepsilon_i$  are the hygroscopicity parameter and volume fraction, respectively, for 182 the individual (dry) components in the mixture and *i* is the number of components in 183 the mixture. **带格式的:**缩进:首行缩进: 2字 符

Measurements from the ACSM in Xinzhou show that the composition of 184 185 submicron particles was dominated by organics, followed by sulfate, ammonium, and nitrate. The contribution of chloride was negligible (volume fraction of about < 2%). 186 The analysis of the anion and cation balance suggests that anionic species (NO<sub>3</sub><sup>-</sup>, 187  $SO_4^{2^-}$ ) were essentially neutralized by  $NH_4^+$  over the relevant size range. For 188 189 refractory species, BC represented a negligible fraction of the total submicron aerosol volume (< 3%). Sea salt and dust are usually coarse mode particles with particle sizes > 190 191 1 µm (Whitby, 1978). The contribution of such types of aerosols is thus expected to be negligible for sizes  $< 1 \,\mu m$ . Therefore, the submicron particles measured by the ACSM 192 mainly consisted of organics, (NH<sub>4</sub>)<sub>2</sub>SO<sub>4</sub>, and NH<sub>4</sub>NO<sub>3</sub>. The particle hygroscopicity is 193 thus the volume average of the three participating species: 194

195 
$$\kappa_{chem} = \kappa_{org} \varepsilon_{org} = \kappa_{org} \varepsilon_{org} + \kappa_{(NH_4)_2 SO_4} \varepsilon_{(NH_4)_2 SO_4} + \kappa_{NH_4 NO_3} \varepsilon_{NH_4 NO_3}$$
(2)

Here, the values of  $\kappa$  for  $(NH_4)_2SO_4$  and  $NH_4NO_3$  are  $0.67\underline{61}$  and  $0.64\underline{67}$ , respectively. The following linear function derived by Mei et al. (2013) was used to estimate  $\kappa_{Orgorg}$  in this study:  $\kappa_{Orgorg} = 2.10 \times f_{44} - 0.11$ , where  $f_{44}$  is the fraction of m/z 44 in total organics. The mean value of  $\kappa_{Orgorg}$  during the field campaign is 0.115 \pm 0.019.

### 201 4. Results and discussion

### 202 4.1 CCN efficiency spectra

During the field campaign at the Xinzhou site, ~790 size-resolved CCN efficiency spectra at five supersaturation levels ranging from 0.075% to 0.76% were measured. Figure 1 shows campaign averaged spectra of the measured CCN efficiency at Xinzhou for supersaturation levels of 0.075%, 0.13%, 0.17%, 0.39%, and 0.76%. The observed averaged CCN efficiency spectra during Xianghe campaign in

208	summer 2013 are also shown. The Note that the maximum activation fraction (MAF)
209	for Xianghe site showed in Figure 1 is slight lower than that we plotted in Zhang et
210	al.(2014). Because some data points when the MAF value >1 were not processed
211	previously, as resulted in larger mean MAF. But in this paper, the data points with
212	<u>MAF &gt; 1.0 were forced to 1 when <math>D_p</math>&gt;300 nm, which could be completely activated</u>
213	at even lower supersaturations but the MAF would never be larger than 1.0. In Figure
214	1, the right panels show the mass concentration fraction of particle chemical
215	compositions at Xinzhou (top panel) and Xianghe (bottom panel) during their
216	respective observation periods. Significant differences in size-resolved CCN
217	efficiency spectra at the two sites are seen. Aerosol particles at Xinzhou activate more
218	efficiently (higher values of AR) at a given particle diameter $(D_p)$ for the same
219	supersaturation level. In the other words, a larger $D_p$ was required to reach the same
220	activation efficiency at Xianghe. This suggests that aerosol properties at each site
221	differ.

Zhang et al.: CCN measurement at Xinzhou 28 Jan 2015

Manuscript for .

222 The slope of AR with respect to diameters near  $D_p$  when AR=50% (defined here 223 as the cut-off diameter,  $D_{cut}$ ) provides information about the heterogeneity of the 224 composition for size-resolved particles. For an ideal case when all CCN-active particles have the same composition and size, a steep change in AR from 0 to 1 would 225 226 be observed as  $D_p$  reached  $D_{cut}$ . A gradual increase in size-resolved AR with  $D_p$ 227 suggests that aerosol particles have different hygroscopicities. The steeper slopes of AR around  $D_{cut}$  observed at Xinzhou suggest that the particle composition was less 228 229 heterogeneous with more hygroscopicity than particles at the Xianghe site. This can 230 be partially explained by the magnitudes of the mean  $\kappa_{chem}$  at the two sites (0.42 at 231 Xinzhou and 0.38 at Xianghe). Also, the  $f_{44}$  is greater at Xinzhou than at Xianghe. The m/z 44 signal is mostly due to acids (Takegawa et al., 2007; Duplissy et al., 2011) or 232

acid-derived species, such as esters.  $f_{44}$  is closely related to the organic oxidation level (Aiken et al., 2008). Oxidized/aged acids are generally more hygroscopic and easily activated. Moreover, the primary inorganic particles at the Xinzhou site are sulfates, with a mass fraction that is two times greater than that measured at Xianghe. Therefore, particles at the Xinzhou site consist of more hygroscopic sulfate-dominant inorganics and aged/oxidized secondary organics and can thus be more efficiently activated at a given  $D_p$ , as shown in Fig. 1.

# 240 4.2 Air mass influences on CCN activity: a case study

Because air mass back trajectories combined with ambient air measurements can 241 242 be used for analyzing large-scale air pollutant transport and source identification at a 243 receptor site (Stohl, 1996; Rousseau et al., 2004), in this study, we calculated five-day (120 hr) back trajectories using the Hybrid Single-Particle Lagrangian Integrated 244 245 Trajectory (HYSPLIT) model (Draxler and Hess, 1998) with National Centers for Environmental Prediction (NCEP) reanalysis data. TrajStat software (Wang et al., 246 2009) has been used to calculate trajectories. The arrival height of the trajectories at 247 248 the Xinzhou site was at the surface.

Three cases were selected to study air mass influences on aerosol<u>CCN</u> activity: (1) Case 1, 19 August 2014, 19:00-21:00 local time (LT); (2) Case 2, 9 August 2014, 03:00-10:00 LT; and (3) Case 3, 29 July 2014, 00:00-12:00 LT. Each case is associated with a different CCN efficiency spectrum, i.e., top, middle, and bottom panels of Fig. 2 are for Cases 1, 2, and 3, respectively. Their respective back trajectories are shown in Fig. 3.

In Case 1, air trajectories (red line in Fig. 3) originated from the southwest and passed through northern Shaanxi Province and northwestern Shanxi Province, then

rounded back to the site from the north/northeast. So, aerosols in this case are closely 257 258 associated with air parcels north/northeast of the site. The trajectories were very short, suggesting that the air flow was slow during the observational period. Under these 259 260 circumstances, aerosol loading would be largely impacted by local sources around the 261 site. A high mass fraction of organics (> 60%) with low  $f_{44}$  (~10%) and  $\kappa_{chem}$  (< 0.3) 262 values was measured during the observational period. Furthermore, the PSD showed one peak mode with  $D_p = 56$  nm and a high N<sub>CN</sub> (~1.7×10<sup>4</sup> cm<sup>-3</sup>), but low mass 263 concentration of  $PM_1$  (28.36 µg m<sup>-3</sup>). This suggests that particles may be composed of 264 freshly emitted primary aerosols (the biogenic emissions from the plants and trees 265 around the site). This type of aerosol is usually less hygroscopic with a single peak 266 mode primarily composed of fine particles (Whitby, 1978; Hussein et al., 2005). 267 These aerosols cannot activate efficiently. The maximum activation fraction (MAF) 268 269 shown in the top right panel of Fig. 2 is less than 0.6 at all supersaturation levels for particles with  $D_p > 300$  nm, indicating that the particles should be largely externally 270 271 mixed aerosols.

In Case 2 (blue line in Fig. 3), air parcels moved rapidly from the west to the site. 272 The site should then be influenced by the large-scale transport of air masses. For this 273 case, aerosols contain a small amount of organics (< 30%), but have high  $f_{44}$  (~14%) 274 and  $\kappa_{chem}$  values (~0.5). The PSD showed a double peak mode with an N<sub>CN</sub> of 275 ~ $1.3 \times 10^4$  cm<sup>-3</sup> and a relatively high mass concentration of PM<sub>1</sub> (81.45 µg m<sup>-3</sup>). The 276 double peak mode suggests that aerosols in this case are a mixture of aerosols from 277 local sources and from other regions (Whitby, 1978; Dal Maso et al., 2007). Because 278 aerosols are aged and oxidized during long-distance transport, these particles are 279 usually composed of secondary organic and inorganic components with more 280 hygroscopicity (Weber et al., 1999; Verver et al., 2000). These aerosols can activate 281

efficiently. The MAF is close to 1 and the slopes of AR around  $D_{cut}$  are steep at all supersaturation levels (middle right panel of Fig. 2). This CCN efficiency spectrum is similar to the ideal spectrum of pure ammonium sulfate.

In Case 3 (green line in Fig. 3), air parcels travelled from the northwest to the site. Air masses arriving at the site in this case had passed over densely populated regions with more heavy pollution. A gradual increase in size-resolved AR with  $D_p$  is seen (bottom right panel of Fig. 2). This is attributed to the diversity in aerosol hygroscopicity because of the complex nature of the chemical composition of aerosol particles.

# $291 \qquad \textbf{4.3 Correlation of } N_{CN} \text{ and } N_{CCN}$

Figure 4 shows N<sub>CN</sub> as a function of N<sub>CCN</sub> for different supersaturation levels at 292 the Xinzhou and Xianghe sites. They showed high or moderate correlations at high 293 supersaturation levels (e.g.,  $R^2 = 0.57$  at Xinzhou and  $R^2 = 0.85$  at Xianghe at a 294 supersaturation level of  $\sim 0.8\%$ ), but quite poor correlations at low supersaturation 295 levels. Although Andreae (2009) proposed using the relationship of CCN and CN, or 296 297 even aerosol optical depth (AOD), to parameterize CCN in models, it should be caution if one uses the correlationwould lead to large uncertainties especially for those 298 eases at low when the supersaturations because of the spatial variation in CCN 299 300 activity for maritime and continental aerosols are low. It was noticed that there was an 301 apparent higher degree of correlation at Xianghe site for each supersaturation than that derived at Xinzhou site. In view of the similar regimes from which the data are taken 302 303 and the same instruments by which they have been collected, the discrepancy between 304 Xianghe and Xinzhou should be caused largely by the spatial variations of aerosols types. 305 These variations are primarily attributed to variations in aerosol particle size, i.e., the 306 shape of the PSD as well as particle composition. As presented by Zhang et al. (2014), 19/44

the relationship between bulk activation ratios and  $N_{CN}$  was complex under polluted conditions and was heavily dependent on the physicochemical properties of atmospheric aerosols.

310 4.4 Impact of *x*<sub>org</sub> on N<sub>CCN</sub>

311 Precise quantification of N<sub>CCN</sub> is crucial for understanding aerosol indirect 312 effects and characterizing these effects in models. A CCN closure study is useful to examine the controlling physical and chemical factors and to help verify experimental 313 314 results. N<sub>CCN</sub> is usually derived from measured aerosol properties, such as PSD and composition or hygroscopicity based on the Köhler theory. Achieving such closure 315 316 under heavily polluted conditions is more challenging, especially due to the complex 317 effects of organics on CCN activity. In this section, we examine the sensitivity of <u>N<sub>CCN</sub> to both</u> volume fraction of organics  $(x_{org})$  as well as and oxidation or aging of 318 organics to N<sub>CCN</sub> estimation based on measurement at Xinzhou site. During the 319 observed period, aerosols at the Xinzhou site were dominated by organics, with 12%, 23%, 320 39%, and 25% of the data points corresponding to xorg >60 %, 50% < xorg <60 %, 40% < xorg 321 322 <50 % and  $x_{org} <40$  %, respectively. For the purpose of examining the sensitivity of 323 estimated N<sub>CCN</sub> to x<sub>org</sub> and oxidation/aging level, we sorted the size-resolved CCN data when the xorg >60 %, 50% < xorg <60 %, 40% < xorg <50 % and xorg <40 %. Furthermore, 324 325 for each level of  $x_{\text{org}}$ , we tested the impacts on N<sub>CCN</sub> estimation both from the most 326 oxidized (with  $f_{44}$  of higher than 15%) and least oxidized (those primary organic 327 aerosols with  $f_{44}$  of lower than 11%) organic particles. For example, the size-resolved CCN data points during the period when  $x_{org} > 60$  % and also  $f_{44} > 15$ % was averaged to 328 329 generate the averaged CCN efficiency spectra at the five supersaturations respectively. Then we used the produced averaged CCN efficiency spectra to estimate N<sub>CCN</sub>. 330

331 Estimated CCN size distributions at the five supersaturations were firstly 332 calculated by multiplying the averaged CCN efficiency spectrum (by using the averaged CCN efficiency spectra, the aerosol particles were assumed with uniform 333 334 chemical composition without considering the effects of the temporal variations of the activation curves on CCN activity) with the actually measured PSD. Then, we 335 336 integrated the estimated CCN size distribution over the whole size range to generate estimated N<sub>CCN</sub>. While the measured CCN size distributions are integrated to produce 337 338 the observed N<sub>CCN</sub>.

Observed and estimated N<sub>CCN</sub> at four supersaturation levels (0.075%, 0.13, 0.17 339 and 0.76%) were showed in Fig 5. The data points presented more disperse and 340 weaker correlations at lower supersaturations. The sensitivity of volume fraction of 341 organics to  $N_{CCN}$  increased with increasing  $x_{org}$ . This is especially for the case of these 342 343 primary organic particles with  $f_{44} < 11\%$ : the slopes obtained from a linear fit of estimated and measured N<sub>CCN</sub> in Fig 6 decreased rapidly (almost with a decrease of 344 ~50%) when the  $x_{\text{org}}$  varied from <40% to >60% at supersaturations of 0.075%, while 345 346 it didn't exhibit a lot of reduction (merely ~10%) along with the increasing of  $x_{org}$  for the supersaturation of 0.76%. N<sub>CCN</sub> was estimated most accurately at higher 347 supersaturation levels. This is likely because a large fraction of particles was already 348 349 CCN-active. Also, particle composition has relatively less influence on CCN activation at high supersaturations (Twohy and Anderson, 2008). For the oxidized or 350 351 aged particles with  $f_{44}>15\%$ , the slopes still follow the similar tendency with the variations of  $x_{org}$  at low and high supersaturations but changed more smoothly to the 352 *x*<sub>org</sub> attributing to the oxidized/aged organic particles being more hygroscopic. 353

However, the impacts of the aerosol particles oxidization level on estimating N<sub>CCN</sub> were also very significant. For example, when the particles were composed by 356 large amounts of organics ( $x_{org}$ >60%), the N<sub>CCN</sub> at the supersaturation of 0.075% and 357 0.13% was underestimated by 46% and 44% respectively at  $f_{44} < 11\%$ , while the underestimation decreased to 32% and 23% at the corresponding supersaturation level 358 359 at  $f_{44}$ >15%. The N<sub>CCN</sub> at ss=0.76 was underestimated by 11% and 4% respectively at  $f_{44}$ <11% and  $f_{44}$ >15%. One thus could conclude that the estimation of N<sub>CCN</sub> would be 360 361 largely improved if the aerosol particles were aged with high oxidation level, especially when the chemical composition of the particles is dominated by organics. 362 363 But for the particles with relative low organics ( $x_{org} < 40\%$ ), the effect caused by the  $f_{44}$ was quite insignificant both for high and low supersaturations. In Fig 6, the slopes 364 were all around 1.0 at the two cases of  $f_{44} < 11\%$  and  $f_{44} > 15\%$ . This can be easily 365 explained. When  $x_{org}$  is less than 40%, the overall hygroscopicity of the particles is 366 dominated by inorganic species such as sulfate and nitrate, which are more 367 368 hygroscopic ( $\kappa_{inorg}$  usually usually larger than 0.6) than organic compounds ( $\kappa_{org}$ usully usually smaller than 0.2). As a result, a larger fraction of particles can be 369 370 activated. According to the simple mixing rule based on chemical volume fractions 371 proposed by Petters and Kreidenweis (2007), the contribution from organics is quite small. If  $x_{org}$  is greater than 60%, organics will dominate the overall particle 372 373 hygroscopicity. Particles with a large  $f_{44}$  are much more hygroscopic and thus strongly 374 influence the estimated N<sub>CCN</sub>. Our results indicated that increasing the proportion of 375 secondary organics would decrease the uncertainties in estimating N<sub>CCN</sub> and lead to a 376 more accurate estimation of N<sub>CCN</sub>.

377 4.5 Applicability of CCN efficiency spectra

378 As a means of testing the applicability of the CCN activation spectra, campaign 379 mean CCN efficiency spectra at different supersaturations observed at the Xinzhou

site is used to estimate N<sub>CCN</sub> at the Xinzhou and Xianghe sites respectively, which 380 381 helps to further examine the sensitivity of N<sub>CCN</sub> to aerosol type. Data from the two sites were measured during the warm season so that the effect of temporal variations 382 in aerosols on CCN levels is reduced. Fitted campaign mean CCN efficiency spectrum 383 at the five supersaturations at Xinzhou (corresponding to spectra in Fig. 1) is 384 385 multiplied by dry PSDs actually measured at Xinzhou and at the Xianghe site respectively. This generates estimated CCN size distributions at the two sites. They 386 387 are then integrated over the whole size range (14-600 nm and 10-700 nm at the Xinzhou and Xianghe sites, respectively) to obtain the estimated N<sub>CCN</sub>. The measured 388 CCN size distributions at each site are integrated to produce the observed N<sub>CCN</sub>. 389

Figure 7 shows estimated  $N_{CCN}$  as a function of measured  $N_{CCN}$  for different 390 391 supersaturation levels at the two sites.  $N_{CCN}$  at Xinzhou was underestimated by 4-5% 392 at supersaturation levels of 0.39% and 0.76%, and was slightly overestimated (~2%) 393 at Xianghe for the same supersaturation levels. Good agreement is seen at the 0.39% and 0.76% supersaturation levels for data from both sites ( $R^2 > 0.92$ ). N<sub>CCN</sub> at 394 Xinzhou was underestimated by ~7% at supersaturation levels < 0.1% (R<sup>2</sup> = 0.87). At 395 Xianghe, however, N<sub>CCN</sub> was overestimated by 19-23% at supersaturation levels < 0.1% 396 397 although the correlation between calculated and measured N<sub>CCN</sub> was good. Because 398 size-resolved CCN efficiency spectra were applied here, excluding the impact of particle size, the influence of chemical composition on CCN activation can be 399 400 investigated. The poor estimates of CCN at low supersaturation levels could be attributed to the high sensitivity of N<sub>CCN</sub> to chemical composition. Because the mass 401 fractions of inorganics and organics measured at the two sites are similar (Fig. 1) and 402 the hygroscopicity for inorganic components is fixed, this overestimation is attributed 403 to the smaller proportion of aged and oxidized organic aerosols at Xianghe compared 404

405 with aerosols at Xinzhou ( $f_{44} = 17\%$  and 11% at Xinzhou and Xianghe, respectively).

### 406 **5. Summary and conclusions**

407 In this study, we have investigated the impacts of particle physicochemical 408 properties on CCN activity based on field measurements obtained from 22 July to 26 409 August of 2014 in the suburb of Xinzhou, China. Five-day back trajectories combined 410 with measurements were analyzed to examine air mass influences on CCN activity. CCN efficiency was largely reduced by local primary biomass burning eventsair 411 412 masses, and the MAF was low to <60%, suggesting externally-mixed and the heterogeneity of particle composition for local emitted aerosols. The CCN activation 413 414 efficiency was enhanced significantly when the site was under the influence of air 415 transported from far away, during which aerosols could be mixed well with more 416 hygroscopic secondary organic and inorganic components. The relationship between N<sub>CN</sub> and N<sub>CCN</sub> was generally poor. Large errors would arise if using the former to 417 418 estimate the latter, especially under low supersaturation conditions.

419 The sensitivity of  $N_{CCN}$  estimation to both  $x_{org}$  as well as and  $f_{44}$  on estimating  $\mathbb{N}_{\mathbb{CCN}}$  has also been examined. A strong dependence of  $N_{\mathbb{CCN}}$  on the both two 420 421 parameters was noted. The sensitivity of N<sub>CCN</sub> to volume fraction and particles oxidization or aging level of organics to  $N_{CCN}$  increased with increasing increase of  $x_{org}$ . 422 423 And also this dependence weakens as the supersaturation level increases. When the particles were mostly composed of organics ( $x_{org}$ >60%), the N<sub>CCN</sub> at the 424 supersaturation of 0.075% and 0.13% was underestimated by 46% and 44% 425 426 respectively if aerosol particles were freshly emitted with primary organics ( $f_{44} < 11\%$ ); 427 while the underestimation decreased to 32% and 23% at the corresponding supersaturations if the particles were with more hygroscopic secondary 428 429 organics( $f_{44}$ >15%). The N<sub>CCN</sub> at the supersaturation of 0.76% was underestimated by

带格式的

11% and 4% respectively at  $f_{44} < 11\%$  and  $f_{44} > 15\%$ . But for the particles composed of 430 low organics (e.g.  $x_{org} < 40\%$ ), the effect caused by the  $f_{44}$  was quite insignificant both 431 at high and low supersaturations. This is due to that the overall hygroscopicity of the 432 433 particles is dominated by inorganics such as sulfate and nitrate, which are more hygroscopic than organic compounds. Our results indicated that it would-decrease the 434 435 uncertainties in estimating  $N_{CCN}$  and lead to a more accurate estimation of  $N_{CCN}$  to increase the proportion of secondary organics, especially when the composition of the 436 437 aerosols is dominated by organics.

The applicability of the CCN efficiency spectrum measured at the Xinzhou site to the Xianghe site was examined and a good agreement was found when the supersaturation level was > 0.2%. However, N<sub>CCN</sub> at the Xianghe site was overestimated by 19-23% when the supersaturation level was < 0.1%. Because of the similar mass fractions of inorganics and organics measured at the two sites, we conclude that this overestimation was mainly caused by the smaller proportion of aged and oxidized organic aerosols at Xianghe compared with aerosols at Xinzhou.

# 445 Acknowledgements

This work was funded by the National Basic Research Program of China '973' (Grant 446 447 No. 2013CB955801, 2013CB955804), the Fundamental Research Funds for the Central Universities (Grant No. 2013YB35) and the NSCF-TAMU Collaborative 448 Research Grant Program (Grant No. 4141101031). We also acknowledge the members 449 of the  $A^2C^2$  team for their hard work during the campaign, including Mr. Du Wei 450 451 (from the State Key Laboratory of Atmospheric Boundary Layer Physics and 452 Atmospheric Chemistry of the Institute of Atmospheric Physics/Chinese Academy of Sciences for carrying out the chemical composition measurements) and Mr. Yuan 453 454 Cheng (from Nanjing University who helped make the size-resolved CCNc ( **带格式的:** 字体: Times New Roman

455 measurements).

456 **References** 

- 457 Aiken, A. C., DeCarlo, P. F., Kroll, J. H., Worsnop, D. R., Huffman, J. A., Docherty,
- 458 K. S., Ulbrich, I. M., Mohr, C., Kimmel, J. R., Sueper, D., Sun, Y., Zhang, Q.,
- 459 Trimborn, A., Northway, M., Ziemann, P. J., Canagaratna, M. R., Onasch, T. B.,
- 460 Alfarra, M. R., Prevot, A. S. H., Dommen, J., Duplissy, J., Metzger, A.,
- Baltensperger, U., and Jimenez, J. L.: O/C and OM/OC ratios of primary,
  secondary, and ambient organic aerosols with high-resolution time-of-flight aerosol
- 463 mass spectrometry, Environ. Sci. Technol., 42, 4478–4485, 2008.
- Andreae, M.O.: Correlation between cloud condensation nuclei concentration and
  aerosol optical thickness in remote and polluted regions, Atmos. Chem. Phys., 9,
  543-556, 2009.
- Bougiatioti, A., Fountoukis, C., Kalivitis, N., Pandis, S. N., Nenes, A., and
  Mihalopoulos, N.: Cloud condensation nuclei measurements in the marine
  boundary layer of the eastern Mediterranean: CCN closure and droplet growth
  kinetics. Atmos. Chem. Phys., 9, 7053–7066, 2009.
- Bougiatioti, A., Nenes, A., Fountoukis, C., Kalivitis, N., Pandis, S. N., and
  Mihalopoulos, N.: Size-resolved CCN distributions and activation kinetics of aged
  continental and marine aerosol, Atmos. Chem. Phys., 11, 8791-8808,
  doi:10.5194/acp-11-8791-2011, 2011.
- 475 Chuang, P. Y., Collins, D. R., Pawlowska, H., Snider, J. R., Jonsson, H. H., Brenguier,
- 476 J. L., Flagan, R. C., and Seinfeld, J. H.: CCN measurements during ACE-2 and
- their relationship to cloud microphysical properties, Tellus B, 52, 843–867, 2000.
- 478 Clarke, A., McNaughton, C., Kasputin, V. N., Shinozuka, Y., Howell, S., Dibb, J.,
- 279 Zhou, J., Anderson, B., Brekhovskikh, V., Turner, H., and Pinkerton, M.: Biomass

- burning and pollution aerosol over North America: Organic components and their
  influence on spectral optical properties and humidification response, J. Geophys.
  Res., 112, D12S18, doi:10.1029/2006JD007777, 2007.
- Dal Maso, M., L. Sogacheva, P. P. Aalto, I. Riipinen, M. Komppula, P. Tunved, L.
  Korhonen, V.SUUR©\USKI, A. Hirsikko, and T. KurtEN, Aerosol size
  distribution measurements at four Nordic field stations: identification, analysis
  and trajectory analysis of new particle formation bursts, Tellus B,2007,59(3),
  350-361.
- Deng, Z., Zhao, C., Ma, N., Liu, F., Ran, L., Xu, W., Liang, Z., Liang, S., Huang, M.,
  Ma, X., Zhang, Q., Quan, J., and Yan, P.: Size- resolved and bulk activation
  properties of aerosols in the North China Plain. Atmos. Chem. Phys., 11,
  3835-3846, 2011.
- Deng, Z., Zhao, C., Ma, N., Ran, L., Zhou, G., Lu, D., and Zhou, X.: An examination
  of parameterizations for the CCN number concentration based on in situ
  measurements of aerosol activation properties in the North China Plain, Atmos.
  Chem. Phys., 13, 6227–6237, doi:10.5194/acp-13-6227-2013, 2013.
- Draxler, R., R. and Hess, G., D. 1998. An overview of the HYSPLIT 4 modeling
  system for trajectories, dispersion, and deposition, *Aust. Meteorol. Mag.* 47, 295–
  308.
- Duplissy, J., DeCarlo, P. F., Dommen, J., Alfarra, M. R., Metzger, A., Barmpadimos,
  I., Prevot, A. S. H., Weingartner, E., Tritscher, T., Gysel, M., Aiken, A. C., Jimenez,
  J. L., Canagaratna, M. R., Worsnop, D. R., Collins, D. R., Tomlinson, J., and
  Baltensperger, U.: Relating hygroscopicity and composition of organic aerosol
- particulate matter, Atmos. Chem. Phys., 11, 1155–1165,
  doi:10.5194/acp-11-1155-2011, 2011.

- Dusek, U., Covert, D. S., Wiedensohler, A., Neususs, C., Weise, D., and Cantrell, W.:
  Cloud condensation nuclei spectra derived from size distributions and hygroscopic
  properties of the aerosol in coastal south-west Portugal during ACE-2, Tellus B, 55,
  35–53, 2003.
- Engelhart, G. J., Hennigan, C. J., Miracolo, M. A., Robinson, A. L., and Pandis, S. N.:
  Cloud condensation nuclei activity of fresh primary and aged biomass burning
  aerosol, Atmos. Chem. Phys., 12, 7285–7293, doi:10.5194/acp-12-7285-2012,
  2012.
- Ervens, B., Cubison, M. J., Andrews, E., Feingold, G., Ogren, J.A., Jimenez, J. L.,
  Quinn, P. K., Bates, T. S., Wang, J., Zhang, Q., Coe, H., Flynn, M., and Allan, J. D.:
  CCN predictions using simplified assumptions of organic aerosol composition and
  mixing state: a synthesis from six different locations, Atmos. Chem.Phys., 10,
  4795–4807, doi:10.5194/acp-10-4795-2010, 2010.
- Gasparini, R., Collins, D. R., Andrews, E., Sheridan, P. J., Ogren, J. A., and Hudson, J.
  G.: Coupling aerosol size distributions and size-resolved hygroscopicity to predict
  humidity-dependent optical properties and cloud condensation nuclei spectra., J.
  Geophys. Res., 111, D05S13, doi:10.1029/2005JD006092, 2006.
- Gunthe SS; Rose D; Su H; Garland RM; Achtert P; Nowak A; Wiedensohler A;
  Kuwata M; Takegawa N; Kondo Y; Hu M; Shao M; Zhu T; Andreae MO; Pöschl
  U (2011) Cloud condensation nuclei (CCN) from fresh and aged air pollution in
  the megacity region of Beijing, Atmospheric Chemistry and Physics, 11,
  pp.11023-11039. doi: 10.5194/acp-11-11023-2011
- Guo, S., Hu, M., Zamora, M. L., Peng, J., Shang, D., Zheng, J., Zhuofei Du, Zhijun
  Wu, Min Shao, Limin Zeng, Mario J. Molina,1, and Zhang, R. (2014).
- 529 Elucidating severe urban haze formation in China. Proceedings of the National

- Academy of Sciences of the United States of America, 111(49), 17373–17378.
  doi:10.1073/pnas.1419604111
- 532 Hartz, K. E. H., Tischuk, J. E., Chan, M. N., Chan, C. K., Donahue, N. M., and Pandis,
- S. N.: Cloud condensation nuclei activation of limited solubility organic aerosol,
  Atmos. Environ., 40, 605–617, 2006.
- 535 Hings, S. S., Wrobel, W. C., Cross, E. S., Worsnop, D. R., Davidovits, P., and Onasch,
- T. B.: CCN activation experiments with adipic acid: effect of particle phase andadipic acid coatings on soluble and insoluble particles, Atmos. Chem. Phys., 8,
- 538 3735–3748, doi:10.5194/acp-8-3735-2008, 2008.
- Hudson, J. G. and Da, X. Y.: Volatility and size of cloud condensation nuclei, J.
  Geophys. Res., 101, 4435–4442, 1996.
- 541 Hussein, T., M. Dal Maso, T. PETÄJÄ, I. K. KOPONEN, P. PAATERO, P. P.
- AALTO, K. HÄMERI, and M. KULMALA, Evaluation of an automatic
  algorithm for fitting the particle number size distributions, BOREAL
  ENVIRONMENT RESEARCH, 2005,10(5), 337-355.
- Junge, C. and McLaren, E.: Relationship of cloud nuclei spectra to aerosol size
  distribution and composition, J. Atmos. Sci., 28, 382–390, 1971.
- 547 Kammermann, L., Gysel, M., Weingartner, E., Herich, H., Cziczo, D. J., Holst, T.,
- 548 Svenningsson, B., Arneth, A., and Baltensperger, U.: Subarctic atmospheric aerosol
- 549 composition: 3. Measured and modeled properties of cloud condensation nuclei, J.
- 550 Geophys. Res., 115, D04202, doi:10.1029/2009JD012447, 2010.
- 551 Köhler, H.: The nucleus in and growth of hygroscopic droplets, Trans. Faraday Soc.,
- 552 32, 1152–1161, doi:10.1039/TF9363201152, 1936.
- 553 Lance, S., Medina, J., Smith, J., and Nenes, A.: Mapping the operation of the DMT
- continuous flow CCN counter, Aerosol Sci. Technol., 40, 242–254, 2006.

- Lathem, T. L., Beyersdorf, A. J., Thornhill, K. L., Winstead, E. L., Cubison, M. J., 555 Hecobian, A., Jimenez, J. L., Weber, R. J., Anderson, B. E., and Nenes, A.: Analysis 556 of CCN activity of Arctic aerosol and Canadian biomass burning during summer 557 2008, Atmos. Chem. Phys., 13, 2735-2756, doi:10.5194/acp-13-2735-2013, 2013. 558 Lee, Y. S., Collins, D. R., Li, R. J., Bowman, K. P., and Feingold, G.: Expected 559 560 impact of an aged biomass burning aerosol on cloud condensation nuclei and cloud droplet concentrations, J. Geophys. Res., 111, D22204, doi:10.1029/2005JD006464, 561 562 2006. Li, Z., Chen, H., Cribb, M., Dickerson, R. E., Holben, B., Li, C., Lu, D., Luo, Y., 563 Maring, H., Shi, G., Tsay, S.-C., Wang, P., Wang, Y., Xia, X., Zheng, Y., Yuan, T., 564 and Zhao, F.: Preface to special section on East Asian Studies of Tropospheric 565 Aerosols: An International Regional Experiment (EASTAIRE), J. Geophys. Res., 566 567 112, D22S00, doi:10.1029/2007JD008853, 2007. Li, Z., Li, C., Chen, H., Tsay, S.-C., Holben, B., Huang, J., Li, B., Maring, H., Qian, 568 Y., Shi, G., Xia, X., Yin, Y., Zheng, Y., and Zhuang, G.: East Asian Studies of 569 Tropospheric Aerosols and Impact on Regional Climate (EAST - AIRC): An 570 overview, J. Geophys. Res., 116, D00K34, doi:10.1029/2010JD015257, 2011. 571
- Liu, X. and Wang, J.: How important is organic aerosol hygroscopicity to aerosol
  indirect forcing? Environ. Res. Lett., 5, 044010,
  doi:10.1088/1748-9326/5/4/044010, 2010.
- 575 McFiggans, G., Artaxo, P., Baltensperger, U., Coe, H., Facchini, M. C., Feingold, G.,
- 576 Fuzzi, S., Gysel, M., Laaksonen, A., Lohmann, U., Mentel, T. F., Murphy, D. M.,
- 577 O'Dowd, C. D., Snider, J. R., and Weingartner, E.: The effect of physical and
- 578 chemical aerosol properties on warm cloud droplet activation, Atmos. Chem. Phys.,
- 579 6, 2593–2649, doi:10.5194/acp-6-2593-2006, 2006.

W. C., de Gouw, J. A., Jimenez, J. L., and Wang, J.: Droplet activation properties

580 Mei, F., Hayes, P. L., Ortega, A. M., Taylor, J. W., Allan, J. D., Gilman, J. B., Kuster,

of organic aerosols observed at an urban site during CalNex-LA, J. Geophys. Res.,

583 118(7), 2903-2917, doi: 10.1002/jgrd.50285, 2013b.

581

- Mei, F., Setyan, A., Zhang, Q., and Wang, J.: CCN activity of organic aerosols
  observed downwind of urban emissions during CARES, Atmos. Chem. Phys., 13,
  12155–12169,doi:10.5194/acp-13-12155-2013, 2013a.
- 587 Mikhailov, E., Vlasenko, S., Martin, S. T., Koop, T., and P<sup>...</sup> oschl, U.: Amorphous
  588 and crystalline aerosol particles interacting with water vapor: conceptual
  589 framework and experimental evidence for restructuring, phase transitions and
  590 kinetic limitations, Atmos. Chem. Phys., 9, 9491-9522,
  591 doi:10.5194/acp 9 9491-2009, 2009.
- 592 Mircea, M., Facchini, M. C., Decesari, S., Cavalli, F., Emblico, L., Fuzzi, S., Vestin, A., Rissler, J., Swietlicki, E., Frank, G., Andreae, M. O., Maenhaut, W., Rudich, Y., 593 and Artaxo, P.: Importance of the organic aerosol fraction for modeling aerosol 594 595 hygroscopic growth and activation: a case study in the Amazon Basin, Atmos. Chem. Phys., 5, 3111–3126, 2005, http://www.atmos-chem-phys.net/5/3111/2005/. 596 Ng, N. L., Herndon, S. C., Trimborn, A., Canagaratna, M. R., Croteau, P. L., Onasch, 597 598 T. B., Sueper, D., Worsnop, D. R., Zhang, Q., Sun, Y. L., and Jayne, J. T.: An Aerosol Chemical Speciation Monitor (ACSM) for Routine Monitoring of the 599 600 Composition and Mass Concentrations of Ambient Aerosol, Aerosol Sci. Tech., 45,
- 601 770–784, 2011.
- Paramonov, M., Aalto, P. P., Asmi, A., Prisle, N., Kerminen, V.-M., Kulmala, M., and
  Pet äj ä, T.: The analysis of size-segregated cloud condensation nuclei counter
  (CCNC) data and its implications for aerosol-cloud interactions, Atmos. Chem.

- 605 Phys. Discuss., 13, 9681-9731, doi:10.5194/acpd-13-9681-2013, 2013.
- Petters, M. D., and Kreidenweis, S. M.: A single parameter representation of
  hygroscopic growth and cloud condensation nucleus activity, Atmos. Chem. Phys.,
  7, 1961–1971, doi:10.5194/acp-7-1961-2007, 2007.
- Petters, M. D. and Kreidenweis, S. M.: A single parameter representation of
  hygroscopic growth and cloud condensation nucleus activity Part 2: Including
  solubility, Atmos. Chem. Phys., 8, 6273–6279, doi:10.5194/acp-8-6273-2008,
  2008.
- Raymond, T. M., and Pandis, S. N.: Cloud activation of single component organic
  aerosol particles, J. Geophys. Res., 107, 4787, doi:10.1029/2002JD002159, 2002.
- Reutter, P., Su, H., Trentmann, J., Simmel, M., Rose, D., Gunthe, S. S., Wernli, H.,
- Andreae, M. O., and Pöschl, U.: Aerosol- and updraft-limited regimes of cloud
  droplet formation: influence of particle number, size and hygroscopicity on the
  activation of cloud condensation nuclei (CCN), Atmos. Chem. Phys., 9, 7067–7080,
  doi:10.5194/acp-9-7067-2009, 2009.
- Rissler, J., Swietlicki, E., Zhou, J., Roberts, G., Andreae, M. O., Gatti, L. V., and
  Artaxo, P.: Physical properties of the submicrometer aerosol over the Amazon rain
  forest during the wet to dry season transition comparison of modeled and
  measured CCN concentrations, Atmos. Chem. Phys., 4, 2119–2143,
  http://www.atmos-chem-phys.net/4/2119/2004/, 2004.
- Rissler, J., Vestin, A., Swietlicki, E., Fisch, G., Zhou, J., Artaxo, P., and Andreae, M.
  O.: Size distribution and hygroscopic properties of aerosol particles from
  dry-season biomass burning in Amazonia, Atmos. Chem. Phys., 6, 471–491,
- 628 doi:10.5194/acp-6-471-2006, 2006.
- 629 Rose, D., Gunthe, S. S., Mikhailov, E., Frank, G. P., Dusek, U., Andreae, M. O., and

- Poschl, U.: Calibration and measurement uncertainties of a continuous-flow cloud
  condensation nuclei counter (DMT-CCNC): CCN activation of ammonium sulfate
  and sodium chloride aerosol particles in theory and experiment, Atmos. Chem.
  Phys., 8, 1153–1179, 2008, http://www.atmos-chem-phys.net/8/1153/2008/.
- Rose, D., Nowak, A., Achtert, P., Wiedensohler, A., Hu, M., Shao, M., Zhang, Y.,
  Andreae, M. O., and P<sup>\*</sup> oschl, U.: Cloud condensation nuclei in polluted air and
  biomass burning smoke near the mega-city Guangzhou, China Part 1:
  Size-resolved measurements and implications for the modeling of aerosol particle
  hygroscopicity and CCN activity, Atmos. Chem. Phys., 10, 3365–3383,
  doi:10.5194/acp-10-3365-2010, 2010.
- Rousseau, D., D., Duzer, D., Etienne, J., L., Cambon, G., Jolly, D. And coauthors.
  2004. Pollen record of rapidly changing air trajectories to the North Pole, *J. Geophys. Res.* 109, D06116, doi:10.1029/2003JD003985.
- Snider, J.R., Guibert, S., Brenguierand, J. L. and Putaud, J. P.: Aerosol activation in
  marine stratocumulus clouds: Part II Köhler and parcel theory closure studies, J.
  Geophy. Res., 108, doi:10.1029/2002JD002692, 2003
- Sotiropoulou, R. E. P., Nenes, A., Adams, P. J., and Seinfeld, J. H.: Cloud
  condensation nuclei prediction error from application of K öhler theory:
  Importance for the aerosol indirect effect, J. Geophys. Res., 112, D12202,
  doi:10.1029/2006JD007834, 2007.
- Stohl, A. 1996, Trajectory statistics a new method to establish source-receptor
  relationships of air pollutants and its application to the transport of particulate
  sulfate in Europe, *Atmos. Environ.* 30, 579–587.
- 653 Stroud, C. A., Nenes, A., Jimenez, J. L., DeCarlo, P., Huffman, J. A., Bruintjes, R.,
- 654 Nemitz, E., Delia, A. E., Toohey, D. W., Guenther, A. B., and Nandi, S.: Cloud

- Activating Properties of Aerosol Observed during CELTIC, J. Atmos. Sci., 64, 441–
  459, 2007.
- Sun, Y., Wang, Z., Dong, H., Yang, T., Li, J., Pan, X., Chen, P., and Jayne, J. T.:
  Characterization of summer organic and inorganic aerosols in Beijing, China with
  an Aerosol Chemical Speciation Monitor, Atmos. Environ., 51, 250–259,
  doi:10.1016/j.atmosenv.2012.01.013, 2012.
- Takegawa, N., Miyakawa, T., Kawamura, K., and Kondo, Y.: Contribution of selected
  di-carboxylic and omega-oxocarboxylic acids in ambient aerosol to the m/z 44
  signal of an aerodyne aerosol mass spectrometer, Aerosol Sci. Technol., 41, 418–
  437, doi:10.1080/02786820701203215, 2007.
- Twohy, C. H. and Anderson, J. R.: Droplet nuclei in non-precipitating clouds:
  composition and size matter, Environ. Res. Lett., 3, 045002,
  doi:10.1088/1748-9326/3/4/045002, 2008.
- 668 VanReken, T. M., Rissman, T. A., Roberts, G. C., Varutbangkul, V., Jonsson, H. H.,
- Flagan, R. C., and Seinfeld, J. H.: Toward aerosol/cloud condensation nuclei (CCN)
  closure during CRYSTAL-FACE, J. Geophys. Res., 108, 4633,
  doi:10.1029/2003JD003582, 2003.
- 672 VanReken, T. M., Ng, N. L., Flagan, R. C., and Seinfeld, J. H.: Cloud condensation
- nucleus activation properties of biogenic secondary organic aerosol, J. Geophys.
  Res., 110, D07206, doi:10.1029/2004JD005465, 2005.
- 675 Varutbangkul, V., Brechtel, F. J., Bahreini, R., Ng, N. L., Keywood, M. D., Kroll, J.
- H., Flagan, R. C., Seinfeld, J. H., Lee, A., and Goldstein, A. H.: Hygroscopicity of
- secondary organic aerosols formed by oxidation of cycloalkenes, monoterpenes,
- sesquiterpenes, and related compounds, Atmos. Chem. Phys., 6, 2367–2388, 2006.
- 679 Verver, G., F. Raes, D. Vogelezang, and D. Johnson, The 2nd Aerosol characterization

- Experiment (ACE-2): meteorological and chemical context, Tellus B, 2000,52(2),
  126-140.
- Wang, J., Lee, Y.-N., Daum, P. H., Jayne, J., and Alexander, M. L.: Effects of aerosol
  organics on cloud condensation nucleus (CCN) concentration and first indirect
  aerosol effect, Atmos. Chem. Phys., 8, 6325–6339, doi:10.5194/acp-8-6325-2008,
  2008.
- Wang, Y.Q., Zhang, X.Y. and Draxler, R., 2009. TrajStat: GIS-based software that
  uses various trajectory statistical analysis methods to identify potential sources
  from long-term air pollution measurement data. Environmental Modelling &
  Software, 24: 938-939
- Ward, D. S., Eidhammer, T., Cotton, W. R., and Kreidenweis, S. M.: The role of the
  particle size distribution in assessing aerosol composition effects on simulated
  droplet activation, Atmos. Chem.Phys., 10, 5435–5447,
  doi:10.5194/acp-10-5435-2010, 2010.
- Weber, R., P. H. McMurry, R. Mauldin, D. Tanner, F. Eisele, A. Clarke, and V.
  Kapustin,New particle formation in the remote troposphere: A comparison of
  observations at various sites, Geophysical Research Letters, 1999,26(3),
  307-310.
- Wex, H., Hennig, T., Salma, I., Ocskay, R., Kiselev, A., Henning, S., Massling, A.,
  Wiedensohler, A., and Stratmann, F.: Hygroscopic growth and measured and
  modeled critical super-saturations of an atmospheric HULIS sample, Geophys. Res.
  Lett., 34, L02818, doi:10.1029/2006GL028260, 2007.
- 702 Whitby, K., T.: The physical characteristics of sulfur aerosols. Atmos. Environ., 12,
- 703 135-159, 1967, Online publication date: 1-Jan-1978, 1978.
- 704 Wiedensohler A; Cheng YF; Nowak A; Wehner B; Achtelt P; Berghof M; Birmili W;

705

Wu ZJ; Hu M; Zhu T; Takegawa N; Kita K; Kondo Y; Lou SR; Hofeumahaus A;

706	Holland F; Wahner A; Gunthe SS; Rose D; Su H; Pöschl U (2009) Rapid aerosol
707	particle growth and increase of cloud condensation nucleus activity by secondary
708	aerosol formation and condensation: A case study for regional air pollution in
709	northeastern China, Journal of Geophysical Research: Atmospheres, 114, . doi:
710	10.1029/2008JD010884
711	Yue, D. L., Hu, M., Zhang, R. J., Wu, Z. J., Su, H., Wang, Z. B., Peng, J. F., He, L.Y.,
712	Huang, X. F., Gong, Y. G., and Wiedensohler, A.: Potential contribution of new
713	particle formation to cloud condensation nuclei in Beijing, Atmos. Environ., 45,
714	6070-6077, 2011.
715	Zhang, Q., Meng, J., Quan, J., Gao, Y., Zhao, D., Chen, P., and He, H.: Impact of
716	aerosol composition on cloud condensation nuclei activity, Atmos. Chem. Phys., 12,
717	3783-3790, doi:10.5194/acp-12-3783-2012, 2012.
718	Zhang, F., Z. Li, R. J. Li, L. Sun, C. Zhao, P. C. Wang, Y. L. Sun, Y. N. Li, X. G. Liu,
719	J. X. Li, P. R. Li, G. Ren, and T. Y. Fan., Aerosol hygroscopicity and cloud
720	condensation nuclei activity during the AC3Exp campaign: implications for cloud
721	condensation nuclei parameterization. Atmos. Chem. Phys., 14, 13423-13437,
722	2014
723	
724	
725	
726	
727	
728	
729	
730	
731	



Fig. 1. Mean CCN efficiency spectra at the Xinzhou site (black lines with asterisks) 736 measured from 22 July-26 August 2014 and at the Xianghe site (red lines with circles) 737 site measured from 7-21 July 2013 for different supersaturation levels. Error bars 738 739 representing one standard deviation are shown. Right panels show particle chemical composition in terms of mass concentration fractions at Xinzhou (top panel) and 740 Xianghe (bottom panel) during their respective observation periods. The campaign 741 average mass concentration of PM<sub>1</sub> is 31.6  $\mu$ g m<sup>-3</sup> and 72.4  $\mu$ g m<sup>-3</sup> at Xinzhou and 742 Xianghe respectively. Note that the preset supersaturation levels were 0.07%, 0.1%, 743 0.2%, 0.4% and 0.8% at both sites, but effective supersaturation levels showed 744 slightly different after calibration. 745

746

747

748

749

750



Zhang et al.: CCN measurement at Xinzhou 28 Jan 2015

Manuscript for .....

Fig. 2. Particle number size distribution (PSD) and CCN size distributions (left panels)
and CCN efficiency spectra (right panels) at different supersaturation levels for Case 1
(upper panels, 19 August 2014, 19:00-21:00 LT), Case 2 (middle panels, 9 August 38/44



Fig. 3. Five-day back trajectories for Case 1 (in red), Case 2 (in blue), and Case 3 (in
green) calculated using the Hybrid Single-Particle Lagrangian Integrated Trajectory
model with National Centers for Environmental Prediction reanalysis data. The arrival
height of the trajectories at the Xinzhou site was at the surface.

- 775
- 776
- 777
- 778
- 779





Fig. 4. Measured  $N_{\text{CN}}$  as a function of measured  $N_{\text{CCN}}$  for different supersaturation levels at the Xinzhou (left panel) and Xianghe (right panel) sites. The scatterplot between CCN\_Obs and CN\_Obs. were fitted with a linear function (in colored lines) and  $R^2$ refer to the correlations of them. .





**Fig. 5.** The sensitivity of  $N_{CCN}$  to both organics volume fraction  $(x_{org})$  as well as and oxidation level (using  $f_{44}$ , the fraction of m/z 44 in total organics, as an indicatoraerosol organic material) of organics to estimation of  $N_{CCN}$ -at supersaturation

825	levels of (a) 0.075%, (b) 0.13%, (c) 0.17% and (d) $0.76\frac{1}{200}$ for cases when $x_{\text{org}} = 35\%$									
826	(blue circles), 52% (green circles), and 66% (red circles). The size-resolved CCN									
827	data were sorted when the $x_{org} > 60 \%$ , $50\% < x_{org} < 60 \%$ , $40\% < x_{org} < 50\%$ and $x_{org} < 40\%$									
828	respectively to do the sensitivity examination. The results of 40% < xorg < 50 % was not									
829	plotted here. Mean values of the hygroscopic parameter $\kappa_{ehem}$ at $f_{44} < 11\%$ when									
830	xorg > 60 %, 50% < xorg < 60 %, 40% < xorg < 50 % and xorg < 40 % are 0.27, 0.34, 0.40 and									
831	0.46, respectively; while at $f_{44}$ >15% the value increased to 0.36, 0.42, 0.46 and 0.50									
832	respectivelyLinear best-fit lines through each group of points are shown. Slopes and									
833	$R^2$ values are given in parentheses.									
834										
835										
836										
837										
838										
839										
840										
841										
842										
843										
844										
845										
	1.1									
	Z 0.9									
	SS 0.8									
	Β Β Β Β Β Β Β Β Β Β Β Β Β Β									
	Ŏ     0.7       →     ss=0.39%,f44<11%									
	$ \begin{array}{c} \bullet & \bullet & \bullet \\ \bullet & \bullet & \bullet & \bullet \\ \bullet & \bullet & \bullet &$									
	$- = - s_{5} = 0.39\%, (14 \times 15\%)$ $- = - s_{5} = 0.39\%, (14 \times 15\%)$									
	0.5									
	30% 40% 50% 60% 70%									



42 / 44



Zhang et al.: CCN measurement at Xinzhou 28 Jan 2015

Manuscript for .....



Fig. 7. Estimated N<sub>CCN</sub> as a function of observed N<sub>CCN</sub> for different supersaturation 870 levels at (a) Xinzhou and (b) Xianghe. Note that the campaign mean CCN efficiency 871 spectra at Xinzhou are used for estimating N<sub>CCN</sub> at Xianghe. Linear best-fit lines 872

873	through	each	group	of	points	are	shown.	Slopes	and	$\mathbb{R}^2$	values	are	given	in
874	parenthe	ses.												
875														
876														
877														
878														
879														
880														

881

\_