

Impacts of organic aerosols and its oxidation level on CCN activity from measurement at a suburban site in China

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Abstract

This study is concerned with the impacts of organic aerosols on CCN activity based on field measurements made at a suburban site in north China. The sensitivity of the estimated CCN number concentration (N_{CCN}) to both volume fraction of organic material (x_{org}) and aerosol oxidation level (using f_{44} , the fraction of m/z 44 in aerosol organic material) are examined. A strong dependence of CCN number concentration (N_{CCN}) on the x_{org} and f_{44} was noted. The sensitivity of N_{CCN} to volume fraction of organics increased with increasing x_{org} . The impacts of the aerosol particles oxidation or aging level on estimating N_{CCN} were also very significant. When the particles were mostly composed of organics ($x_{\text{org}} > 60\%$), the N_{CCN} at the supersaturation of 0.075% and 0.13% was underestimated by 46% and 44% respectively if aerosol particles were freshly emitted with primary organics ($f_{44} < 11\%$); while the underestimation decreased to 32% and 23% at the corresponding supersaturations if the particles were with more hygroscopic secondary organics ($f_{44} > 15\%$). The N_{CCN} at the supersaturation of 0.76% was underestimated by 11% and 4% respectively at $f_{44} < 11\%$ and $f_{44} > 15\%$. But for the particles composed of low organics (e.g. $x_{\text{org}} < 40\%$), the effect caused by the f_{44} was quite insignificant both at high and low supersaturations. This is due to that the overall hygroscopicity of the particles is dominated by inorganics such as sulfate and nitrate, which are more hygroscopic than organic compounds. Our results indicated that it would decrease the uncertainties in estimating N_{CCN} and lead to a more accurate estimation of N_{CCN} to increase the proportion of secondary organics, especially when the composition of the

aerosols is dominated by organics. The applicability of the CCN activation spectrum obtained at Xinzhou to the Xianghe site, about 400 km to the northeast of Xinzhou, was investigated, with the aim of further examining the sensitivity of N_{CCN} to aerosol type. Overall, the mean CCN efficiency spectrum derived from Xinzhou performs well at Xianghe when the supersaturation levels are $> 0.2\%$ (overestimation of 2-4%). However, N_{CCN} was overestimated by $\sim 20\%$ at supersaturation levels of $< 0.1\%$. This suggests that the overestimation is mainly due to the smaller proportion of aged and oxidized organic aerosols present at Xianghe compared with Xinzhou.

9 1. Introduction

10 To reduce the uncertainty of aerosol indirect effects on the radiative balance of
11 the atmosphere, it is important to gain a good knowledge of the ability of aerosol
12 particles to form cloud condensation nuclei (CCN) at the typical supersaturations
13 found in the atmosphere. The CCN activity of aerosol particles is governed by the
14 Köhler theory (Köhler, 1936). This theory determines CCN from aerosol particle size
15 and physicochemical properties, which include the molar volume, activity coefficient,
16 and effect on surface tension (McFiggans et al., 2006). These properties, however, are
17 difficult to measure.

18 Researchers have proposed single-parameter models to parameterize the CCN
19 activation and hygroscopicity of multi-component aerosols (Hudson and Da, 1996;
20 Rissler et al., 2006; Petters and Kreidenweis, 2007; Wex et al., 2007). Field
21 experiments have been conducted with the aim of better characterizing particle
22 physicochemical parameters influencing cloud CCN activation. Due to the large
23 spatial variability of aerosol types and compositions, the CCN activation efficiency
24 varies greatly over different regions. CCN number concentrations (N_{CCN}) can often be
25 better predicted in the background atmosphere (Chuang et al., 2000; Dusek et al.,
26 2003; VanReken et al., 2003; Snider et al., 2003; Rissler et al., 2004; Gasparini et al.,
27 2006; Stroud et al., 2007; Bougiatioti et al., 2009).

28 The largest errors are associated with urban emissions (Sotiropoulou et al., 2007).
29 This is likely due to the organics component of aerosol particles, which have the
30 largest uncertainty and are not fully understood. Biomass burning aerosols and
31 secondary organics formed from the oxidation of common biogenic emissions are
32 often more difficult to activate (Mircea et al., 2005; VanReken et al., 2005; Lee et al.,
33 2006; Varutbangkul et al., 2006; Clarke et al., 2007; Rose et al., 2010; Engelhart et al.,

2012; Paramonov et al., 2013; Lathem et al., 2013; Mei et al., 2013b; Zhang et al., 2014). Particles with aged/oxidized secondary organic components (e.g., organic acids) have been shown to be more hygroscopic (Raymond and Pandis, 2002; Hartz et al., 2006; Bougiatioti et al., 2011), but still much less hygroscopic than inorganic species. The sensitivity of estimated N_{CCN} to organics have been examined in a number of recent studies (Wang et al., 2008; Reutter et al., 2009; Ervens et al., 2010; Kammermann et al., 2010; Ward et al., 2010; Zhang et al., 2012; Mei et al., 2013a). It is widely known that the estimated N_{CCN} is sensitive to changes in organics due to the latter's complex components. The amounts and hygroscopicity parameter of organics (κ_{org}) vary substantially and lead to significant biases in estimating CCN concentrations and aerosol indirect forcing (Sotiropoulou et al., 2007; Hings et al., 2008; Liu and Wang, 2010). Therefore, field investigations regarding CCN activity and organics impacts, especially in heavily polluted regions, are pivotal to better parameterize CCN in climate models.

Northern China is a fast developing and densely populated region of China, where aerosol loading is high (Li et al., 2007, 2011), the particle composition is complex, and severe haze pollution episodes are common (Guo et al., 2014). In recent years, CCN measurements have been collected during field campaigns carried out in the region (Wiedensohler et al., 2009; Gunthe et al., 2011; Yue et al., 2011; Deng et al., 2011, 2013; Zhang et al., 2014). These studies have presented different perspectives on the influence of particle size and composition on CCN activity. For example, Deng et al. (2013) evaluated various schemes for CCN parameterization and recommended that the particle number size distribution (PSD) together with inferred mean size-resolved activation ratios can be used to estimate CCN number concentrations without considering the impact of particle composition. However,

59 Zhang et al. (2014) demonstrated that the 30–40% uncertainties in N_{CCN} are mainly
60 associated with changes in particle composition. None of the above-mentioned studies
61 have investigated the impact of organics on estimating N_{CCN} in Northern China.
62 Zhang et al. (2012) noted a more significant influence of organics on CCN activity but
63 without concerning the influences of particles oxidation or aging on CCN activity; in
64 addition, the campaign average mass fraction of organics in their study was $< 20\%$.

65 The aim of this paper is to examine the sensitivity of CCN activity to aerosol
66 physicochemical properties (especially aerosols containing large amounts of organics,
67 as well as the oxidation level), and also to see how much uncertainty is incurred by
68 applying the CCN efficiency spectra measured at one site to another site in a heavily
69 polluted region. The instrumentation and data used in the study are described in
70 section 2. The method for calculating the hygroscopicity parameter (κ_{chem}) is
71 introduced in section 3. The sensitivity of estimating N_{CCN} to x_{org} as well as oxidation
72 level of organics is investigated in section 4, and the ability of the CCN efficiency
73 spectrum observed at the Xinzhou site to represent CCN at the Xianghe site, are also
74 presented and discussed at the last part of this section. Conclusions from the study are
75 given in section 5.

76 2. Measurements and data

77 An intensive observation period field campaign similar to the
78 Aerosol-CCN-Cloud Closure Experiment (Zhang et al., 2014), called the Atmosphere,
79 Aerosol, Cloud, and CCN (A^2C^2) experiment, was conducted from 22 July to 26
80 August of 2014 at Xinzhou ($38.24^\circ N$, $112.43^\circ E$; 1500 m above sea level), a city with a
81 population of 0.51 million in Northern China. The site is located about 360 km
82 southwest of the metropolitan Beijing area and about 10 km south of the local town
83 center. The site is surrounded by agricultural land (e.g., corn) with little local pollution

84 plums from motor vehicles and industrial activities. Sitting between two mountains
85 (Taihang Mountain to the east and Lüliang Mountain to the west), the site also
86 experiences air masses from Xinzhou City to the north and from Taiyuan City to the
87 south, the capital of Shanxi Province. Air masses from the northeast and southwest
88 dominate over the site during summer. Depending on the wind direction,
89 measurements at the Xinzhou site can detect air parcels of urban, rural, or mixed
90 origins, including both fresh biogenic emissions around the site and aged aerosols
91 from advection.

92 **2.1 Instruments and measurements**

93 During the field campaign, a Scanning Mobility Particle Sizer (SMPS),
94 combined with a Droplet Measurement Technologies-Cloud Condensation Nuclei
95 Counter (DMT-CCN_C) (Lance et al., 2006), was used for size-resolved CCN
96 measurements as well as particle number size distribution (PSD) measurements. The
97 measured aerosol PSD is within the size range of 14-600 nm. Aerosol chemical
98 composition was measured simultaneously by an Aerodyne Aerosol Chemical
99 Speciation Monitor (ACSM) (Sun et al., 2012).

100 The aerosol inlet for the size distribution measurements was equipped with a TSI
101 Environmental Sampling System (Model 3031200), which consists of a sharp-cut
102 PM₁ cyclone and a bundled nafion dryer. The size-resolved CCN efficiency spectra
103 were measured by coupling the DMT-CCN_C used with the SMPS (Rose et al., 2008).
104 In this step, the particles are rapidly dried with RH < 30% upon entering the
105 Differential Mobility Analyzer (DMA). Thus, size selection is effectively performed
106 under dry conditions. The nafion dryer and the sheath air inside the DMA are
107 sufficient to remove residual water associated with the ambient particles. Relative

108 deviations in particle diameter should be $< 1\%$. The sample flow exiting the DMA
109 was split into two parts: 0.3 lpm for the CPC and 0.5 lpm for the CCN counter
110 (CCN_C). The DMA, controlled by TSI-AIM software, scanned one size distribution
111 every five minutes. The CCN_C was operated at a total flow rate of 0.5 lpm with a
112 sheath-to-aerosol flow ratio of 10. The inlet RH for CCN_C was $< 30\%$. During the
113 field campaign, the mean sample temperature and pressure measured by CCN_C
114 sensors was $(24.3 \pm 1.4) ^\circ\text{C}$ and (898.4 ± 11.7) hPa. The supersaturations levels of CCN_C
115 were calibrated with ammonium sulfate before and after the field campaign, following
116 the procedures outlined in [Rose et al. \(2008\)](#). During each CCN measurement cycle,
117 calibrated effective supersaturations were set at 0.075%, 0.13%, 0.17%, 0.39%, and
118 0.75%. The overall relative error (1σ) for the supersaturation levels was estimated to
119 be $< 3.5\%$. The completion of a full measurement cycle took 50 minutes (10 minutes
120 for each supersaturation level).

121 The measurement of non-refractory submicron aerosol species including
122 organics, sulfate, nitrate, ammonium, and chloride were made with an ACSM. During
123 the field campaign, ambient aerosols were drawn inside through a $\frac{1}{2}$ inch (outer
124 diameter, the inner diameter is 0.38 inch) stainless steel tube at a flow rate of ~ 3 L
125 min^{-1} , of which ~ 84 cc min^{-1} was sub-sampled into the ACSM. An URG cyclone
126 (Model: URG-2000-30ED) was also positioned in front of the sampling inlet to
127 remove coarse particles with a cut-off size of $2.5 \mu\text{m}$. Before sampling into the ACSM,
128 aerosol particles were dried using a silica gel desiccant. The residence time in the
129 sampling tube was ~ 5 s. The ACSM was operated at a time resolution of ~ 15 min with
130 a scan rate of mass spectrometer at 500 ms amu^{-1} from m/z 10 to 150. Regarding the
131 calibration of the ACSM, mono-dispersed, size-selected 300-nm ammonium nitrate
132 particles within a range of concentrations were sampled into both the ACSM and a

133 condensation particle counter (CPC). The ionization efficiency (IE) was then
134 determined by comparing the response factors of the ACSM to the mass calculated
135 with known particle size and number concentrations from the CPC. More detailed
136 descriptions of the operation and calibration of the ACSM are given in Sun et al.
137 (2012) and Ng et al. (2011). The campaign averaged mass concentration of PM_{10} is
138 $31.6 \mu\text{g m}^{-3}$.

139 In addition to the ACSM, the black carbon (BC) in $PM_{2.5}$ was simultaneously
140 measured at a time resolution of 5 min by a seven-wavelength aethalometer (Model
141 AE31, Magee Scientific Corporation). The campaign averaged mass concentration of
142 BC is $\sim 2.5 \mu\text{g m}^{-3}$. During the experiment, the campaign area was generally hot and
143 dry, with an average temperature of 21.6°C and an average ambient RH of 69.5%.

144 2.2 Data

145 The raw CCN data were first filtered according to instrument recorded
146 parameters (e.g., temperature and flow). For example, if the relative difference
147 between the actual and preset sample flows was larger than 4%, the data are flagged
148 as invalid. The data is also excluded if the “temperature stability” was flagged as “0”.
149 Here, the “temperature stability” refers to the T_1 , T_2 and T_3 in cloud chamber of the
150 CCNc, which is set to obtain the target supersaturations. If the average differences
151 between preset T_1 , T_2 and T_3 and the measured values are larger than 0.4°C , the
152 “temperature stability” is flagged as “0”. Thus, the data is invalid and will be removed.
153 A multiple charge correction and transfer function (Deng et al., 2011) is applied to
154 each PSD as well as to the CCN efficiency spectrum. Size-resolved CCN and PSD
155 data, measured with a DMT-CCNc and a SMPS (with a particle size range of 10-700
156 nm) on 7-21 July 2013 at Xianghe (Zhang et al., 2014), are used in this study for

157 comparisons with CCN activity at the Xinzhou site. Aerosol mass concentrations were
158 processed using the ACSM standard data analysis software (version 1.5.3.0). Detailed
159 procedures for the data analysis have been described by Ng et al. (2011) and Sun et al.
160 (2012). The size-resolved CCN activation ratio (size-resolved AR) is defined as the
161 $dN_{CCN}/d\log D_p$ divided by the $dN_{CN}/d\log D_p$. These values were measured by
162 SMPS-DMT-CCNc with particle size selection in the DMA. The bulk activation ratio
163 (bulk AR) is defined as the total CCN concentration divided by the total CN
164 concentration. The total CCN and CN number concentrations are integrated by the
165 measured CCN and CN size distribution respectively over the whole size range.

166 3. Derivation of κ_{chem}

167 In this study, we calculate κ_{chem} based on bulk chemical composition observations
168 made during the field campaign. The method is very similar to that used by Zhang et
169 al., (2014). As proposed by Petters and Kreidenweis (2007), κ_{chem} can be predicted
170 using a simple mixing rule based on chemical volume fractions for a given internal
171 mixture:

$$172 \quad \kappa_{chem} = \sum_i \varepsilon_i \kappa_i \quad (1)$$

173 where κ_i and ε_i are the hygroscopicity parameter and volume fraction, respectively, for
174 the individual (dry) components in the mixture and i is the number of components in
175 the mixture.

176 Measurements from the ACSM in Xinzhou show that the composition of
177 submicron particles was dominated by organics, followed by sulfate, ammonium, and
178 nitrate. The contribution of chloride was negligible (volume fraction of about < 2%).
179 The analysis of the anion and cation balance suggests that anionic species (NO_3^- ,
180 SO_4^{2-}) were essentially neutralized by NH_4^+ over the relevant size range. For
181 refractory species, BC represented a negligible fraction of the total submicron aerosol

182 volume (< 3%). Sea salt and dust are usually coarse mode particles with particle sizes >
183 1 μm (Whitby, 1978). The contribution of such types of aerosols is thus expected to be
184 negligible for sizes < 1 μm . Therefore, the submicron particles measured by the ACSM
185 mainly consisted of organics, $(\text{NH}_4)_2\text{SO}_4$, and NH_4NO_3 . The particle hygroscopicity is
186 thus the volume average of the three participating species:

$$187 \quad \kappa_{chem} = \kappa_{org}\varepsilon_{org} + \kappa_{(\text{NH}_4)_2\text{SO}_4}\varepsilon_{(\text{NH}_4)_2\text{SO}_4} + \kappa_{\text{NH}_4\text{NO}_3}\varepsilon_{\text{NH}_4\text{NO}_3} \quad (2)$$

188 Here, the values of κ for $(\text{NH}_4)_2\text{SO}_4$ and NH_4NO_3 are 0.61 and 0.67, respectively.

189 The following linear function derived by Mei et al. (2013) was used to estimate κ_{org} in
190 this study: $\kappa_{org} = 2.10 \times f_{44} - 0.11$, where f_{44} is the fraction of m/z 44 in total organics.

191 The mean value of κ_{org} during the field campaign is 0.115 ± 0.019 .

192 4. Results and discussion

193 4.1 CCN efficiency spectra

194 During the field campaign at the Xinzhou site, ~790 size-resolved CCN
195 efficiency spectra at five supersaturation levels ranging from 0.075% to 0.76% were
196 measured. Figure 1 shows campaign averaged spectra of the measured CCN
197 efficiency at Xinzhou for supersaturation levels of 0.075%, 0.13%, 0.17%, 0.39%, and
198 0.76%. The observed averaged CCN efficiency spectra during Xianghe campaign in
199 summer 2013 are also shown. Note that the maximum activation fraction (MAF) for
200 Xianghe site showed in Figure 1 is slight lower than that we plotted in Zhang et
201 al.(2014). Because some data points when the MAF value >1 were not processed
202 previously, as resulted in larger mean MAF. But in this paper, the data points with
203 MAF > 1.0 were forced to 1 when $D_p > 300$ nm, which could be completely activated
204 at even lower supersaturations but the MAF would never be larger than 1.0. In Figure
205 1, the right panels show the mass concentration fraction of particle chemical

206 compositions at Xinzhou (top panel) and Xianghe (bottom panel) during their
207 respective observation periods. Significant differences in size-resolved CCN
208 efficiency spectra at the two sites are seen. Aerosol particles at Xinzhou activate more
209 efficiently (higher values of AR) at a given particle diameter (D_p) for the same
210 supersaturation level. In the other words, a larger D_p was required to reach the same
211 activation efficiency at Xianghe. This suggests that aerosol properties at each site
212 differ.

213 The slope of AR with respect to diameters near D_p when AR=50% (defined here
214 as the cut-off diameter, D_{cut}) provides information about the heterogeneity of the
215 composition for size-resolved particles. For an ideal case when all CCN-active
216 particles have the same composition and size, a steep change in AR from 0 to 1 would
217 be observed as D_p reached D_{cut} . A gradual increase in size-resolved AR with D_p
218 suggests that aerosol particles have different hygroscopicities. The steeper slopes of
219 AR around D_{cut} observed at Xinzhou suggest that the particle composition was less
220 heterogeneous with more hygroscopicity than particles at the Xianghe site. This can
221 be partially explained by the magnitudes of the mean κ_{chem} at the two sites (0.42 at
222 Xinzhou and 0.38 at Xianghe). Also, the f_{44} is greater at Xinzhou than at Xianghe. The
223 m/z 44 signal is mostly due to acids (Takegawa et al., 2007; Duplissy et al., 2011) or
224 acid-derived species, such as esters. f_{44} is closely related to the organic oxidation level
225 (Aiken et al., 2008). Oxidized/aged acids are generally more hygroscopic and easily
226 activated. Moreover, the primary inorganic particles at the Xinzhou site are sulfates,
227 with a mass fraction that is twice greater than that measured at Xianghe. Therefore,
228 particles at the Xinzhou site consist of more hygroscopic sulfate-dominant inorganics
229 and aged/oxidized secondary organics and can thus be more efficiently activated at a
230 given D_p , as shown in Fig. 1.

231 4.2 Air mass influences on CCN activity: a case study

232 Because air mass back trajectories combined with ambient air measurements can
233 be used for analyzing large-scale air pollutant transport and source identification at a
234 receptor site (Stohl, 1996; Rousseau et al., 2004), in this study, we calculated five-day
235 (120 hr) back trajectories using the Hybrid Single-Particle Lagrangian Integrated
236 Trajectory (HYSPLIT) model (Draxler and Hess, 1998) with National Centers for
237 Environmental Prediction (NCEP) reanalysis data. TrajStat software (Wang et al.,
238 2009) has been used to calculate trajectories. The arrival height of the trajectories at
239 the Xinzhou site was at the surface.

240 Three cases were selected to study air mass influences on CCN activity: (1) Case
241 1, 19 August 2014, 19:00-21:00 local time (LT); (2) Case 2, 9 August 2014,
242 03:00-10:00 LT; and (3) Case 3, 29 July 2014, 00:00-12:00 LT. Each case is
243 associated with a different CCN efficiency spectrum, i.e., top, middle, and bottom
244 panels of Fig. 2 are for Cases 1, 2, and 3, respectively. Their respective back
245 trajectories are shown in Fig. 3.

246 In Case 1, air trajectories (red line in Fig. 3) originated from the southwest and
247 passed through northern Shaanxi Province and northwestern Shanxi Province, then
248 rounded back to the site from the north/northeast. So, aerosols in this case are closely
249 associated with air parcels north/northeast of the site. The trajectories were very short,
250 suggesting that the air flow was slow during the observational period. Under these
251 circumstances, aerosol loading would be largely impacted by local sources around the
252 site. A high mass fraction of organics ($> 60\%$) with low f_{44} ($\sim 10\%$) and κ_{chem} (< 0.3)
253 values was measured during the observational period. Furthermore, the PSD showed
254 one peak mode with $D_p = 56$ nm and a high N_{CN} ($\sim 1.7 \times 10^4$ cm $^{-3}$), but low mass
255 concentration of PM_{10} (28.36 $\mu\text{g m}^{-3}$). This suggests that particles may be composed of

256 freshly emitted primary aerosols (the biogenic emissions from the plants and trees
257 around the site). This type of aerosol is usually less hygroscopic with a single peak
258 mode primarily composed of fine particles (Whitby, 1978; Hussein et al., 2005).
259 These aerosols cannot activate efficiently. The maximum activation fraction (MAF)
260 shown in the top right panel of Fig. 2 is less than 0.6 at all supersaturation levels for
261 particles with $D_p > 300$ nm, indicating that the particles should be largely externally
262 mixed aerosols.

263 In Case 2 (blue line in Fig. 3), air parcels moved rapidly from the west to the site.
264 The site should then be influenced by the large-scale transport of air masses. For this
265 case, aerosols contain a small amount of organics ($< 30\%$), but have high f_{44} ($\sim 14\%$)
266 and κ_{chem} values (~ 0.5). The PSD showed a double peak mode with an N_{CN} of
267 $\sim 1.3 \times 10^4$ cm^{-3} and a relatively high mass concentration of PM_{10} (81.45 $\mu\text{g m}^{-3}$). The
268 double peak mode suggests that aerosols in this case are a mixture of aerosols from
269 local sources and from other regions (Whitby, 1978; Dal Maso et al., 2007). Because
270 aerosols are aged and oxidized during long-distance transport, these particles are
271 usually composed of secondary organic and inorganic components with more
272 hygroscopicity (Weber et al., 1999; Verver et al., 2000). These aerosols can activate
273 efficiently. The MAF is close to 1 and the slopes of AR around D_{cut} are steep at all
274 supersaturation levels (middle right panel of Fig. 2). This CCN efficiency spectrum is
275 similar to the ideal spectrum of pure ammonium sulfate.

276 In Case 3 (green line in Fig. 3), air parcels travelled from the northwest to the
277 site. Air masses arriving at the site in this case had passed over densely populated
278 regions with more heavy pollution. A gradual increase in size-resolved AR with D_p is
279 seen (bottom right panel of Fig. 2). This is attributed to the diversity in aerosol
280 hygroscopicity because of the complex nature of the chemical composition of aerosol

281 particles.

282 **4.3 Correlation of N_{CN} and N_{CCN}**

283 Figure 4 shows N_{CN} as a function of N_{CCN} for different supersaturation levels at
284 the Xinzhou and Xianghe sites. They showed high or moderate correlations at high
285 supersaturation levels (e.g., $R^2 = 0.57$ at Xinzhou and $R^2 = 0.85$ at Xianghe at a
286 supersaturation level of $\sim 0.8\%$), but quite poor correlations at low supersaturation
287 levels. Although [Andreae \(2009\)](#) proposed using the relationship of CCN and CN, or
288 even aerosol optical depth (AOD), to parameterize CCN in models, it would lead to
289 large uncertainties especially when the supersaturations are low. It was noticed that
290 there was an apparent higher degree of correlation at Xianghe site for each
291 supersaturation than that derived at Xinzhou site. In view of the similar regimes from
292 which the data are taken and the same instruments by which they have been collected, the
293 discrepancy between Xianghe and Xinzhou should be caused largely by the spatial variations
294 of aerosols types. These variations are primarily attributed to variations in aerosol
295 particle size, i.e., the shape of the PSD as well as particle composition. As presented
296 by [Zhang et al. \(2014\)](#), the relationship between bulk activation ratios and N_{CN} was
297 complex under polluted conditions and was heavily dependent on the
298 physicochemical properties of atmospheric aerosols.

299 **4.4 Impact of x_{org} on N_{CCN}**

300 Precise quantification of N_{CCN} is crucial for understanding aerosol indirect
301 effects and characterizing these effects in models. A CCN closure study is useful to
302 examine the controlling physical and chemical factors and to help verify experimental
303 results. N_{CCN} is usually derived from measured aerosol properties, such as PSD and
304 composition or hygroscopicity based on the Köhler theory. Achieving such closure

305 under heavily polluted conditions is more challenging, especially due to the complex
306 effects of organics on CCN activity. In this section, we examine the sensitivity of
307 N_{CCN} to both volume fraction of organics (x_{org}) and oxidation or aging of organics
308 based on measurement at Xinzhou site. During the observed period, aerosols at the
309 Xinzhou site were dominated by organics, with 12%, 23%, 39%, and 25% of the data points
310 corresponding to $x_{org} > 60\%$, $50\% < x_{org} < 60\%$, $40\% < x_{org} < 50\%$ and $x_{org} < 40\%$, respectively.
311 For the purpose of examining the sensitivity of estimated N_{CCN} to x_{org} and
312 oxidation/aging level, we sorted the size-resolved CCN data when the $x_{org} > 60\%$,
313 $50\% < x_{org} < 60\%$, $40\% < x_{org} < 50\%$ and $x_{org} < 40\%$. Furthermore, for each level of x_{org} , we
314 tested the impacts on N_{CCN} estimation both from the most oxidized (with f_{44} of higher
315 than 15%) and least oxidized (those primary organic aerosols with f_{44} of lower than
316 11%) organic particles. For example, the size-resolved CCN data points during the
317 period when $x_{org} > 60\%$ and also $f_{44} > 15\%$ was averaged to generate the averaged CCN
318 efficiency spectra at the five supersaturations respectively. Then we used the produced
319 averaged CCN efficiency spectra to estimate N_{CCN} .

320 Estimated CCN size distributions at the five supersaturations were firstly
321 calculated by multiplying the averaged CCN efficiency spectrum (by using the
322 averaged CCN efficiency spectra, the aerosol particles were assumed with uniform
323 chemical composition without considering the effects of the temporal variations of the
324 activation curves on CCN activity) with the actually measured PSD. Then, we
325 integrated the estimated CCN size distribution over the whole size range to generate
326 estimated N_{CCN} . While the measured CCN size distributions are integrated to produce
327 the observed N_{CCN} .

328 Observed and estimated N_{CCN} at four supersaturation levels (0.075%, 0.13, 0.17
329 and 0.76%) were showed in Fig 5. The data points presented more disperse and

330 weaker correlations at lower supersaturations. The sensitivity of N_{CCN} to volume
331 fraction of organics increased with increasing x_{org} . This is especially for the case of
332 these primary organic particles with $f_{44}<11\%$: the slopes obtained from a linear fit of
333 estimated and measured N_{CCN} in Fig 6 decreased rapidly (almost with a decrease of
334 $\sim 50\%$) when the x_{org} varied from $<40\%$ to $>60\%$ at supersaturations of 0.075% , while
335 it didn't exhibit a lot of reduction (merely $\sim 10\%$) along with the increasing of x_{org} at
336 the supersaturation of 0.76% . N_{CCN} was estimated most accurately at higher
337 supersaturation levels. This is likely because a large fraction of particles was already
338 CCN-active. Also, particle composition has relatively less influence on CCN
339 activation at high supersaturations (Twohy and Anderson, 2008). For the oxidized or
340 aged particles with $f_{44}>15\%$, the slopes still follow the similar tendency with the
341 variations of x_{org} at low and high supersaturations but changed more smoothly to the
342 x_{org} attributing to the oxidized/aged organic particles being more hygroscopic.

343 However, the impacts of the aerosol particles oxidization level on estimating
344 N_{CCN} were also very significant. For example, when the particles were composed by
345 large amounts of organics ($x_{org}>60\%$), the N_{CCN} at the supersaturation of 0.075% and
346 0.13% was underestimated by 46% and 44% respectively at $f_{44}<11\%$, while the
347 underestimation decreased to 32% and 23% at the corresponding supersaturation level
348 when $f_{44}>15\%$. The N_{CCN} at $ss=0.76$ was underestimated by 11% and 4% respectively
349 at $f_{44}<11\%$ and $f_{44}>15\%$. The results showed that the estimation of N_{CCN} would be
350 largely improved if the aerosol particles were aged with high oxidation level,
351 especially when the chemical composition of the particles is dominated by organics.
352 But for the particles with relative low organics ($x_{org}<40\%$), the effect caused by the f_{44}
353 was quite insignificant both for high and low supersaturations. In Fig 6, the slopes
354 were all around 1.0 at the two cases of $f_{44}<11\%$ and $f_{44}>15\%$. This can be easily

355 explained. When x_{org} is less than 40%, the overall hygroscopicity of the particles is
356 dominated by inorganic species such as sulfate and nitrate, which are more
357 hygroscopic (κ_{inorg} usually larger than 0.6) than organic compounds (κ_{org} usually
358 smaller than 0.2). As a result, a larger fraction of particles can be activated. According
359 to the simple mixing rule based on chemical volume fractions proposed by [Petters and](#)
360 [Kreidenweis \(2007\)](#), the contribution from organics is quite small. If x_{org} is greater
361 than 60%, organics will dominate the overall particle hygroscopicity. Particles with a
362 large f_{44} are much more hygroscopic and thus more CCN-active. Our results indicated
363 that increasing the proportion of secondary organics would decrease the uncertainties
364 in estimating N_{CCN} and lead to a more accurate estimation of N_{CCN} .

365 **4.5 Applicability of CCN efficiency spectra**

366 As a means of testing the applicability of the CCN activation spectra, campaign
367 mean CCN efficiency spectra at different supersaturations observed at the Xinzhou
368 site is used to estimate N_{CCN} at the Xinzhou and Xianghe sites respectively, which
369 helps to further examine the sensitivity of N_{CCN} to aerosol type. Data from the two
370 sites were measured during the warm season so that the effect of temporal variations
371 in aerosols on CCN levels is reduced. Fitted campaign mean CCN efficiency spectrum
372 at the five supersaturations at Xinzhou (corresponding to spectra in Fig. 1) is
373 multiplied by dry PSDs actually measured at Xinzhou and at the Xianghe site
374 respectively. This generates estimated CCN size distributions at the two sites. They
375 are then integrated over the whole size range (14-600 nm and 10-700 nm at the
376 Xinzhou and Xianghe sites, respectively) to obtain the estimated N_{CCN} . The measured
377 CCN size distributions at each site are integrated to produce the observed N_{CCN} .

378 Figure 7 shows estimated N_{CCN} as a function of measured N_{CCN} for different

379 supersaturation levels at the two sites. N_{CCN} at Xinzhou was underestimated by 4-5%
380 at supersaturation levels of 0.39% and 0.76%, and was slightly overestimated (~2%)
381 at Xianghe for the same supersaturation levels. Good agreement is seen at the 0.39%
382 and 0.76% supersaturation levels for data from both sites ($R^2 > 0.92$). N_{CCN} at
383 Xinzhou was underestimated by ~7% at supersaturation levels $< 0.1\%$ ($R^2 = 0.87$). At
384 Xianghe, however, N_{CCN} was overestimated by 19-23% at supersaturation levels $< 0.1\%$
385 although the correlation between calculated and measured N_{CCN} was good. Because
386 size-resolved CCN efficiency spectra were applied here, excluding the impact of
387 particle size, the influence of chemical composition on CCN activation can be
388 investigated. The poor estimates of CCN at low supersaturation levels could be
389 attributed to the high sensitivity of N_{CCN} to chemical composition. Because the mass
390 fractions of inorganics and organics measured at the two sites are similar (Fig. 1) and
391 the hygroscopicity for inorganic components is fixed, this overestimation is attributed
392 to the smaller proportion of aged and oxidized organic aerosols at Xianghe compared
393 with aerosols at Xinzhou ($f_{44} = 17\%$ and 11% at Xinzhou and Xianghe, respectively).

394 5. Summary and conclusions

395 In this study, we have investigated the impacts of particle physicochemical
396 properties on CCN activity based on field measurements obtained from 22 July to 26
397 August of 2014 in the suburb of Xinzhou, China. Five-day back trajectories combined
398 with measurements were analyzed to examine air mass influences on CCN activity.
399 CCN efficiency was largely reduced by local air masses, and the MAF was low to
400 $< 60\%$, suggesting externally-mixed and the heterogeneity of particle composition for
401 local emitted aerosols. The CCN activation efficiency was enhanced significantly
402 when the site was under the influence of air transported from far away, during which
403 aerosols could be mixed well with more hygroscopic secondary organic and inorganic

404 components. The relationship between N_{CN} and N_{CCN} was generally poor. Large
405 errors would arise if using the former to estimate the latter, especially under low
406 supersaturation conditions.

407 The sensitivity of N_{CCN} estimation to both x_{org} and f_{44} has also been examined. A
408 strong dependence of N_{CCN} on the both two parameters was noted. The sensitivity of
409 N_{CCN} to volume fraction and particles oxidization or aging level of organics increased
410 with increase of x_{org} . And also this dependence weakens as the supersaturation level
411 increases. When the particles were mostly composed of organics ($x_{\text{org}} > 60\%$), the N_{CCN}
412 at the supersaturation of 0.075% and 0.13% was underestimated by 46% and 44%
413 respectively if aerosol particles were freshly emitted with primary organics ($f_{44} < 11\%$);
414 while the underestimation decreased to 32% and 23% at the corresponding
415 supersaturations if the particles were with more hygroscopic secondary
416 organics ($f_{44} > 15\%$). The N_{CCN} at the supersaturation of 0.76% was underestimated by
417 11% and 4% respectively at $f_{44} < 11\%$ and $f_{44} > 15\%$. But for the particles composed of
418 low organics (e.g. $x_{\text{org}} < 40\%$), the effect caused by the f_{44} was quite insignificant both
419 at high and low supersaturations. This is due to that the overall hygroscopicity of the
420 particles is dominated by inorganics such as sulfate and nitrate, which are more
421 hygroscopic than organic compounds. Our results indicated that it would lead to a
422 more accurate estimation of N_{CCN} to increase the proportion of secondary organics,
423 especially when the composition of the aerosols is dominated by organics.

424 The applicability of the CCN efficiency spectrum measured at the Xinzhou site
425 to the Xianghe site was examined and a good agreement was found when the
426 supersaturation level was $> 0.2\%$. However, N_{CCN} at the Xianghe site was
427 overestimated by 19-23% when the supersaturation level was $< 0.1\%$. Because of the
428 similar mass fractions of inorganics and organics measured at the two sites, we

429 conclude that this overestimation was mainly caused by the smaller proportion of
430 aged and oxidized organic aerosols at Xianghe compared with aerosols at Xinzhou.

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440 **References**

- 441 Aiken, A. C., DeCarlo, P. F., Kroll, J. H., Worsnop, D. R., Huffman, J. A., Docherty,
442 K. S., Ulbrich, I. M., Mohr, C., Kimmel, J. R., Sueper, D., Sun, Y., Zhang, Q.,
443 Trimborn, A., Northway, M., Ziemann, P. J., Canagaratna, M. R., Onasch, T. B.,
444 Alfarra, M. R., Prevot, A. S. H., Dommen, J., Duplissy, J., Metzger, A.,
445 Baltensperger, U., and Jimenez, J. L.: O/C and OM/OC ratios of primary,
446 secondary, and ambient organic aerosols with high-resolution time-of-flight aerosol
447 mass spectrometry, *Environ. Sci. Technol.*, 42, 4478–4485, 2008.
- 448 Andreae, M.O.: Correlation between cloud condensation nuclei concentration and
449 aerosol optical thickness in remote and polluted regions, *Atmos. Chem. Phys.*, 9,
450 543-556, 2009.
- 451 Bougiatioti, A., Fountoukis, C., Kalivitis, N., Pandis, S. N., Nenes, A., and
452 Mihalopoulos, N.: Cloud condensation nuclei measurements in the marine

- 453 boundary layer of the eastern Mediterranean: CCN closure and droplet growth
454 kinetics. *Atmos. Chem. Phys.*, 9, 7053–7066, 2009.
- 455 Bougiatioti, A., Nenes, A., Fountoukis, C., Kalivitis, N., Pandis, S. N., and
456 Mihalopoulos, N.: Size-resolved CCN distributions and activation kinetics of aged
457 continental and marine aerosol, *Atmos. Chem. Phys.*, 11, 8791–8808,
458 doi:10.5194/acp-11-8791-2011, 2011.
- 459 Chuang, P. Y., Collins, D. R., Pawlowska, H., Snider, J. R., Jonsson, H. H., Brenguier,
460 J. L., Flagan, R. C., and Seinfeld, J. H.: CCN measurements during ACE-2 and
461 their relationship to cloud microphysical properties, *Tellus B*, 52, 843–867, 2000.
- 462 Clarke, A., McNaughton, C., Kasputin, V. N., Shinozuka, Y., Howell, S., Dibb, J.,
463 Zhou, J., Anderson, B., Brekhovskikh, V., Turner, H., and Pinkerton, M.: Biomass
464 burning and pollution aerosol over North America: Organic components and their
465 influence on spectral optical properties and humidification response, *J. Geophys.*
466 *Res.*, 112, D12S18, doi:10.1029/2006JD007777, 2007.
- 467 Dal Maso, M., L. Sogacheva, P. P. Aalto, I. Riipinen, M. Komppula, P. Tunved, L.
468 Korhonen, V. SUUR@USKI, A. Hirsikko, and T. KurtEN, Aerosol size
469 distribution measurements at four Nordic field stations: identification, analysis
470 and trajectory analysis of new particle formation bursts, *Tellus B*, 2007, 59(3),
471 350–361.
- 472 Deng, Z., Zhao, C., Ma, N., Liu, F., Ran, L., Xu, W., Liang, Z., Liang, S., Huang, M.,
473 Ma, X., Zhang, Q., Quan, J., and Yan, P.: Size- resolved and bulk activation
474 properties of aerosols in the North China Plain. *Atmos. Chem. Phys.*, 11,
475 3835–3846, 2011.
- 476 Deng, Z., Zhao, C., Ma, N., Ran, L., Zhou, G., Lu, D., and Zhou, X.: An examination
477 of parameterizations for the CCN number concentration based on in situ

- 478 measurements of aerosol activation properties in the North China Plain, *Atmos.*
479 *Chem. Phys.*, 13, 6227–6237, doi:10.5194/acp-13-6227-2013, 2013.
- 480 Draxler, R., R. and Hess, G., D. 1998. An overview of the HYSPLIT 4 modeling
481 system for trajectories, dispersion, and deposition, *Aust. Meteorol. Mag.* 47, 295–
482 308.
- 483 Duplissy, J., DeCarlo, P. F., Dommen, J., Alfarra, M. R., Metzger, A., Barmpadimos,
484 I., Prevot, A. S. H., Weingartner, E., Tritscher, T., Gysel, M., Aiken, A. C., Jimenez,
485 J. L., Canagaratna, M. R., Worsnop, D. R., Collins, D. R., Tomlinson, J., and
486 Baltensperger, U.: Relating hygroscopicity and composition of organic aerosol
487 particulate matter, *Atmos. Chem. Phys.*, 11, 1155–1165,
488 doi:10.5194/acp-11-1155-2011, 2011.
- 489 Dusek, U., Covert, D. S., Wiedensohler, A., Neususs, C., Weise, D., and Cantrell, W.:
490 Cloud condensation nuclei spectra derived from size distributions and hygroscopic
491 properties of the aerosol in coastal south-west Portugal during ACE-2, *Tellus B*, 55,
492 35–53, 2003.
- 493 Engelhart, G. J., Hennigan, C. J., Miracolo, M. A., Robinson, A. L., and Pandis, S. N.:
494 Cloud condensation nuclei activity of fresh primary and aged biomass burning
495 aerosol, *Atmos. Chem. Phys.*, 12, 7285–7293, doi:10.5194/acp-12-7285-2012,
496 2012.
- 497 Ervens, B., Cubison, M. J., Andrews, E., Feingold, G., Ogren, J.A., Jimenez, J. L.,
498 Quinn, P. K., Bates, T. S., Wang, J., Zhang, Q., Coe, H., Flynn, M., and Allan, J. D.:
499 CCN predictions using simplified assumptions of organic aerosol composition and
500 mixing state: a synthesis from six different locations, *Atmos. Chem. Phys.*, 10,
501 4795–4807, doi:10.5194/acp-10-4795-2010, 2010.
- 502 Gasparini, R., Collins, D. R., Andrews, E., Sheridan, P. J., Ogren, J. A., and Hudson, J.

503 G.: Coupling aerosol size distributions and size-resolved hygroscopicity to predict
504 humidity-dependent optical properties and cloud condensation nuclei spectra., *J.*
505 *Geophys. Res.*, 111, D05S13, doi:10.1029/2005JD006092, 2006.

506 Gunthe SS; Rose D; Su H; Garland RM; Achtert P; Nowak A; Wiedensohler A;
507 Kuwata M; Takegawa N; Kondo Y; Hu M; Shao M; Zhu T; Andreae MO; Pöschl
508 U (2011) Cloud condensation nuclei (CCN) from fresh and aged air pollution in
509 the megacity region of Beijing, *Atmospheric Chemistry and Physics*, 11,
510 pp.11023-11039. doi: [10.5194/acp-11-11023-2011](https://doi.org/10.5194/acp-11-11023-2011)

511 Guo, S., Hu, M., Zamora, M. L., Peng, J., Shang, D., Zheng, J., Zhuofei Du, Zhijun
512 Wu, Min Shao, Limin Zeng, Mario J. Molina,¹ and Zhang, R. (2014).
513 Elucidating severe urban haze formation in China. *Proceedings of the National*
514 *Academy of Sciences of the United States of America*, 111(49), 17373–17378.
515 doi:10.1073/pnas.1419604111

516 Hartz, K. E. H., Tischuk, J. E., Chan, M. N., Chan, C. K., Donahue, N. M., and Pandis,
517 S. N.: Cloud condensation nuclei activation of limited solubility organic aerosol,
518 *Atmos. Environ.*, 40, 605–617, 2006.

519 Hings, S. S., Wrobel, W. C., Cross, E. S., Worsnop, D. R., Davidovits, P., and Onasch,
520 T. B.: CCN activation experiments with adipic acid: effect of particle phase and
521 adipic acid coatings on soluble and insoluble particles, *Atmos. Chem. Phys.*, 8,
522 3735–3748, doi:10.5194/acp-8-3735-2008, 2008.

523 Hudson, J. G. and Da, X. Y.: Volatility and size of cloud condensation nuclei, *J.*
524 *Geophys. Res.*, 101, 4435–4442, 1996.

525 Hussein, T., M. Dal Maso, T. PETÄJÄ, I. K. KOPONEN, P. PAATERO, P. P.
526 AALTO, K. HÄMERI, and M. KULMALA, Evaluation of an automatic
527 algorithm for fitting the particle number size distributions, *BOREAL*

- 528 ENVIRONMENT RESEARCH, 2005,10(5), 337-355.
- 529 Junge, C. and McLaren, E.: Relationship of cloud nuclei spectra to aerosol size
530 distribution and composition, *J. Atmos. Sci.*, 28, 382–390, 1971.
- 531 Kammermann, L., Gysel, M., Weingartner, E., Herich, H., Cziczo, D. J., Holst, T.,
532 Svenningsson, B., Arneth, A., and Baltensperger, U.: Subarctic atmospheric aerosol
533 composition: 3. Measured and modeled properties of cloud condensation nuclei, *J.*
534 *Geophys. Res.*, 115, D04202, doi:10.1029/2009JD012447, 2010.
- 535 Köhler, H.: The nucleus in and growth of hygroscopic droplets, *Trans. Faraday Soc.*,
536 32, 1152–1161, doi:10.1039/TF9363201152, 1936.
- 537 Lance, S., Medina, J., Smith, J., and Nenes, A.: Mapping the operation of the DMT
538 continuous flow CCN counter, *Aerosol Sci. Technol.*, 40 , 242–254, 2006.
- 539 Latham, T. L., Beyersdorf, A. J., Thornhill, K. L., Winstead, E. L., Cubison, M. J.,
540 Hecobian, A., Jimenez, J. L., Weber, R. J., Anderson, B. E., and Nenes, A.: Analysis
541 of CCN activity of Arctic aerosol and Canadian biomass burning during summer
542 2008, *Atmos. Chem. Phys.*, 13, 2735-2756, doi:10.5194/acp-13-2735-2013, 2013.
- 543 Lee, Y. S., Collins, D. R., Li, R. J., Bowman, K. P., and Feingold, G.: Expected
544 impact of an aged biomass burning aerosol on cloud condensation nuclei and cloud
545 droplet concentrations, *J. Geophys. Res.*, 111, D22204, doi:10.1029/2005JD006464,
546 2006.
- 547 Li, Z., Chen, H., Cribb, M., Dickerson, R. E., Holben, B., Li, C., Lu, D., Luo, Y.,
548 Maring, H., Shi, G., Tsay, S.-C., Wang, P., Wang, Y., Xia, X., Zheng, Y., Yuan, T.,
549 and Zhao, F.: Preface to special section on East Asian Studies of Tropospheric
550 Aerosols: An International Regional Experiment (EASTAIRE), *J. Geophys. Res.*,
551 112, D22S00, doi:10.1029/2007JD008853, 2007.
- 552 Li, Z., Li, C., Chen, H., Tsay, S.-C., Holben, B., Huang, J., Li, B., Maring, H., Qian,

- 553 Y., Shi, G., Xia, X., Yin, Y., Zheng, Y., and Zhuang, G.: East Asian Studies of
 554 Tropospheric Aerosols and Impact on Regional Climate (EAST - AIRC): An
 555 overview, *J. Geophys. Res.*, 116, D00K34, doi:10.1029/2010JD015257, 2011.
- 556 Liu, X. and Wang, J.: How important is organic aerosol hygroscopicity to aerosol
 557 indirect forcing? *Environ. Res. Lett.*, 5, 044010,
 558 doi:10.1088/1748-9326/5/4/044010, 2010.
- 559 McFiggans, G., Artaxo, P., Baltensperger, U., Coe, H., Facchini, M. C., Feingold, G.,
 560 Fuzzi, S., Gysel, M., Laaksonen, A., Lohmann, U., Mentel, T. F., Murphy, D. M.,
 561 O'Dowd, C. D., Snider, J. R., and Weingartner, E.: The effect of physical and
 562 chemical aerosol properties on warm cloud droplet activation, *Atmos. Chem. Phys.*,
 563 6, 2593–2649, doi:10.5194/acp-6-2593-2006, 2006.
- 564 Mei, F., Hayes, P. L., Ortega, A. M., Taylor, J. W., Allan, J. D., Gilman, J. B., Kuster,
 565 W. C., de Gouw, J. A., Jimenez, J. L., and Wang, J.: Droplet activation properties
 566 of organic aerosols observed at an urban site during CalNex-LA, *J. Geophys. Res.*,
 567 118(7), 2903-2917, doi: 10.1002/jgrd.50285, 2013b.
- 568 Mei, F., Setyan, A., Zhang, Q., and Wang, J.: CCN activity of organic aerosols
 569 observed downwind of urban emissions during CARES, *Atmos. Chem. Phys.*, 13,
 570 12155–12169, doi:10.5194/acp-13-12155-2013, 2013a.
- 571 Mircea, M., Facchini, M. C., Decesari, S., Cavalli, F., Emblico, L., Fuzzi, S., Vestin,
 572 A., Rissler, J., Swietlicki, E., Frank, G., Andreae, M. O., Maenhaut, W., Rudich, Y.,
 573 and Artaxo, P.: Importance of the organic aerosol fraction for modeling aerosol
 574 hygroscopic growth and activation: a case study in the Amazon Basin, *Atmos.*
 575 *Chem. Phys.*, 5, 3111–3126, 2005, <http://www.atmos-chem-phys.net/5/3111/2005/>.
- 576 Ng, N. L., Herndon, S. C., Trimborn, A., Canagaratna, M. R., Croteau, P. L., Onasch,
 577 T. B., Sueper, D., Worsnop, D. R., Zhang, Q., Sun, Y. L., and Jayne, J. T.: An

- 578 Aerosol Chemical Speciation Monitor (ACSM) for Routine Monitoring of the
579 Composition and Mass Concentrations of Ambient Aerosol, *Aerosol Sci. Tech.*, 45,
580 770–784, 2011.
- 581 Paramonov, M., Aalto, P. P., Asmi, A., Prisle, N., Kerminen, V.-M., Kulmala, M., and
582 Petäjä T.: The analysis of size-segregated cloud condensation nuclei counter
583 (CCNC) data and its implications for aerosol-cloud interactions, *Atmos. Chem.*
584 *Phys. Discuss.*, 13, 9681-9731, doi:10.5194/acpd-13-9681-2013, 2013.
- 585 Petters, M. D., and Kreidenweis, S. M.: A single parameter representation of
586 hygroscopic growth and cloud condensation nucleus activity, *Atmos. Chem. Phys.*,
587 7, 1961–1971, doi:10.5194/acp-7-1961-2007, 2007.
- 588 Petters, M. D. and Kreidenweis, S. M.: A single parameter representation of
589 hygroscopic growth and cloud condensation nucleus activity – Part 2: Including
590 solubility, *Atmos. Chem. Phys.*, 8, 6273–6279, doi:10.5194/acp-8-6273-2008,
591 2008.
- 592 Raymond, T. M., and Pandis, S. N.: Cloud activation of single component organic
593 aerosol particles, *J. Geophys. Res.*, 107, 4787, doi:10.1029/2002JD002159, 2002.
- 594 Reutter, P., Su, H., Trentmann, J., Simmel, M., Rose, D., Gunthe, S. S., Wernli, H.,
595 Andreae, M. O., and Pöschl, U.: Aerosol- and updraft-limited regimes of cloud
596 droplet formation: influence of particle number, size and hygroscopicity on the
597 activation of cloud condensation nuclei (CCN), *Atmos. Chem. Phys.*, 9, 7067–7080,
598 doi:10.5194/acp-9-7067-2009, 2009.
- 599 Rissler, J., Swietlicki, E., Zhou, J., Roberts, G., Andreae, M. O., Gatti, L. V., and
600 Artaxo, P.: Physical properties of the submicrometer aerosol over the Amazon rain
601 forest during the wet to dry season transition – comparison of modeled and
602 measured CCN concentrations, *Atmos. Chem. Phys.*, 4, 2119–2143,

- 603 <http://www.atmos-chem-phys.net/4/2119/2004/>, 2004.
- 604 Rissler, J., Vestin, A., Swietlicki, E., Fisch, G., Zhou, J., Artaxo, P., and Andreae, M.
605 O.: Size distribution and hygroscopic properties of aerosol particles from
606 dry-season biomass burning in Amazonia, *Atmos. Chem. Phys.*, 6, 471–491,
607 doi:10.5194/acp-6-471-2006, 2006.
- 608 Rose, D., Gunthe, S. S., Mikhailov, E., Frank, G. P., Dusek, U., Andreae, M. O., and
609 Poschl, U.: Calibration and measurement uncertainties of a continuous-flow cloud
610 condensation nuclei counter (DMT-CCNC): CCN activation of ammonium sulfate
611 and sodium chloride aerosol particles in theory and experiment, *Atmos. Chem.*
612 *Phys.*, 8, 1153–1179, 2008, <http://www.atmos-chem-phys.net/8/1153/2008/>.
- 613 Rose, D., Nowak, A., Achtert, P., Wiedensohler, A., Hu, M., Shao, M., Zhang, Y.,
614 Andreae, M. O., and Poschl, U.: Cloud condensation nuclei in polluted air and
615 biomass burning smoke near the mega-city Guangzhou, China – Part 1:
616 Size-resolved measurements and implications for the modeling of aerosol particle
617 hygroscopicity and CCN activity, *Atmos. Chem. Phys.*, 10, 3365–3383,
618 doi:10.5194/acp-10-3365-2010, 2010.
- 619 Rousseau, D., D., Duzer, D., Etienne, J., L., Cambon, G., Jolly, D. And coauthors.
620 2004. Pollen record of rapidly changing air trajectories to the North Pole, *J.*
621 *Geophys. Res.* 109, D06116, doi:10.1029/2003JD003985.
- 622 Snider, J.R., Guibert, S., Brenguierand, J. L. and Putaud, J. P.: Aerosol activation in
623 marine stratocumulus clouds: Part – II Köhler and parcel theory closure studies, *J.*
624 *Geophys. Res.*, 108, doi:10.1029/2002JD002692, 2003
- 625 Sotiropoulou, R. E. P., Nenes, A., Adams, P. J., and Seinfeld, J. H.: Cloud
626 condensation nuclei prediction error from application of Köhler theory:
627 Importance for the aerosol indirect effect, *J. Geophys. Res.*, 112, D12202,

- 628 doi:10.1029/2006JD007834, 2007.
- 629 Stohl, A. 1996, Trajectory statistics - a new method to establish source-receptor
630 relationships of air pollutants and its application to the transport of particulate
631 sulfate in Europe, *Atmos. Environ.* 30, 579–587.
- 632 Stroud, C. A., Nenes, A., Jimenez, J. L., DeCarlo, P., Huffman, J. A., Bruintjes, R.,
633 Nemitz, E., Delia, A. E., Toohey, D. W., Guenther, A. B., and Nandi, S.: Cloud
634 Activating Properties of Aerosol Observed during CELTIC, *J. Atmos. Sci.*, 64, 441–
635 459, 2007.
- 636 Sun, Y., Wang, Z., Dong, H., Yang, T., Li, J., Pan, X., Chen, P., and Jayne, J. T.:
637 Characterization of summer organic and inorganic aerosols in Beijing, China with
638 an Aerosol Chemical Speciation Monitor, *Atmos. Environ.*, 51, 250–259,
639 doi:10.1016/j.atmosenv.2012.01.013, 2012.
- 640 Takegawa, N., Miyakawa, T., Kawamura, K., and Kondo, Y.: Contribution of selected
641 di-carboxylic and omega-oxocarboxylic acids in ambient aerosol to the m/z 44
642 signal of an aerodyne aerosol mass spectrometer, *Aerosol Sci. Technol.*, 41, 418–
643 437, doi:10.1080/02786820701203215, 2007.
- 644 Twohy, C. H. and Anderson, J. R.: Droplet nuclei in non-precipitating clouds:
645 composition and size matter, *Environ. Res. Lett.*, 3, 045002,
646 doi:10.1088/1748-9326/3/4/045002, 2008.
- 647 VanReken, T. M., Rissman, T. A., Roberts, G. C., Varutbangkul, V., Jonsson, H. H.,
648 Flagan, R. C., and Seinfeld, J. H.: Toward aerosol/cloud condensation nuclei (CCN)
649 closure during CRYSTAL-FACE, *J. Geophys. Res.*, 108, 4633,
650 doi:10.1029/2003JD003582, 2003.
- 651 VanReken, T. M., Ng, N. L., Flagan, R. C., and Seinfeld, J. H.: Cloud condensation
652 nucleus activation properties of biogenic secondary organic aerosol, *J. Geophys.*

- 653 Res., 110, D07206, doi:10.1029/2004JD005465, 2005.
- 654 Varutbangkul, V., Brechtel, F. J., Bahreini, R., Ng, N. L., Keywood, M. D., Kroll, J.
655 H., Flagan, R. C., Seinfeld, J. H., Lee, A., and Goldstein, A. H.: Hygroscopicity of
656 secondary organic aerosols formed by oxidation of cycloalkenes, monoterpenes,
657 sesquiterpenes, and related compounds, *Atmos. Chem. Phys.*, 6, 2367–2388, 2006.
- 658 Verver, G., F. Raes, D. Vogelezang, and D. Johnson, The 2nd Aerosol characterization
659 Experiment (ACE-2): meteorological and chemical context, *Tellus B*, 2000, 52(2),
660 126-140.
- 661 Wang, J., Lee, Y.-N., Daum, P. H., Jayne, J., and Alexander, M. L.: Effects of aerosol
662 organics on cloud condensation nucleus (CCN) concentration and first indirect
663 aerosol effect, *Atmos. Chem. Phys.*, 8, 6325–6339, doi:10.5194/acp-8-6325-2008,
664 2008.
- 665 Wang, Y.Q., Zhang, X.Y. and Draxler, R., 2009. TrajStat: GIS-based software that
666 uses various trajectory statistical analysis methods to identify potential sources
667 from long-term air pollution measurement data. *Environmental Modelling &*
668 *Software*, 24: 938-939
- 669 Ward, D. S., Eidhammer, T., Cotton, W. R., and Kreidenweis, S. M.: The role of the
670 particle size distribution in assessing aerosol composition effects on simulated
671 droplet activation, *Atmos. Chem. Phys.*, 10, 5435–5447,
672 doi:10.5194/acp-10-5435-2010, 2010.
- 673 Weber, R., P. H. McMurry, R. Mauldin, D. Tanner, F. Eisele, A. Clarke, and V.
674 Kapustin, New particle formation in the remote troposphere: A comparison of
675 observations at various sites, *Geophysical Research Letters*, 1999, 26(3),
676 307-310.
- 677 Wex, H., Hennig, T., Salma, I., Ocskay, R., Kiselev, A., Henning, S., Massling, A.,

678 Wiedensohler, A., and Stratmann, F.: Hygroscopic growth and measured and
679 modeled critical super-saturations of an atmospheric HULIS sample, *Geophys. Res.*
680 *Lett.*, 34, L02818, doi:10.1029/2006GL028260, 2007.

681 Whitby, K., T.: The physical characteristics of sulfur aerosols. *Atmos. Environ.*, 12,
682 135-159, 1967, Online publication date: 1-Jan-1978, 1978.

683 Wiedensohler A; Cheng YF; Nowak A; Wehner B; Achtert P; Berghof M; Birmili W;
684 Wu ZJ; Hu M; Zhu T; Takegawa N; Kita K; Kondo Y; Lou SR; Hofeumahaas A;
685 Holland F; Wahner A; Gunthe SS; Rose D; Su H; Pöschl U (2009) Rapid aerosol
686 particle growth and increase of cloud condensation nucleus activity by secondary
687 aerosol formation and condensation: A case study for regional air pollution in
688 northeastern China, *Journal of Geophysical Research: Atmospheres*, 114, . doi:
689 [10.1029/2008JD010884](https://doi.org/10.1029/2008JD010884)

690 Yue, D. L., Hu, M., Zhang, R. J., Wu, Z. J., Su, H., Wang, Z. B., Peng, J. F., He, L. Y.,
691 Huang, X. F., Gong, Y. G., and Wiedensohler, A.: Potential contribution of new
692 particle formation to cloud condensation nuclei in Beijing, *Atmos. Environ.*, 45,
693 6070-6077, 2011.

694 Zhang, Q., Meng, J., Quan, J., Gao, Y., Zhao, D., Chen, P., and He, H.: Impact of
695 aerosol composition on cloud condensation nuclei activity, *Atmos. Chem. Phys.*, 12,
696 3783-3790, doi:10.5194/acp-12-3783-2012, 2012.

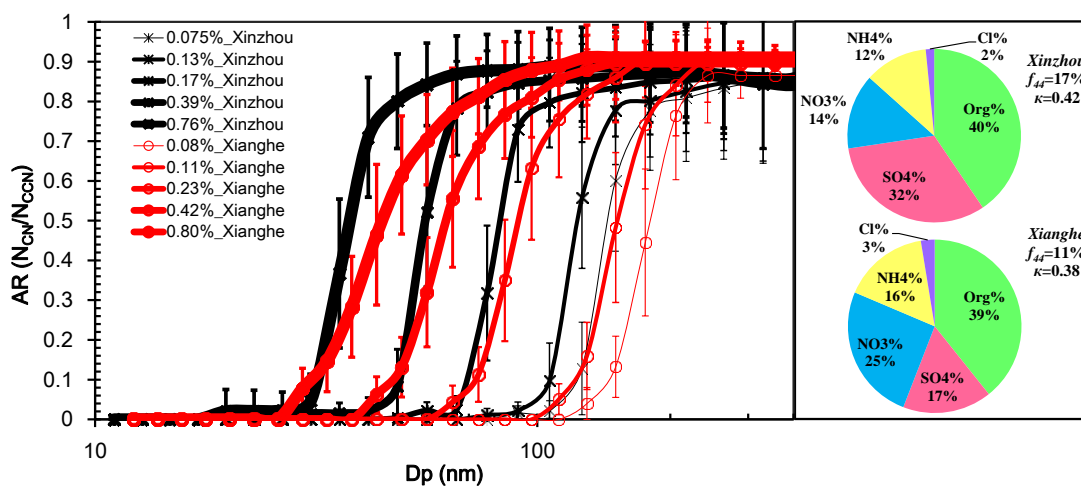
697 Zhang, F., Z. Li, R. J. Li, L. Sun, C. Zhao, P. C. Wang, Y. L. Sun, Y. N. Li, X. G. Liu,
698 J. X. Li, P. R. Li, G. Ren, and T. Y. Fan., Aerosol hygroscopicity and cloud
699 condensation nuclei activity during the AC3Exp campaign: implications for cloud
700 condensation nuclei parameterization. *Atmos. Chem. Phys.*, 14, 13423–13437,
701 2014

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Figures

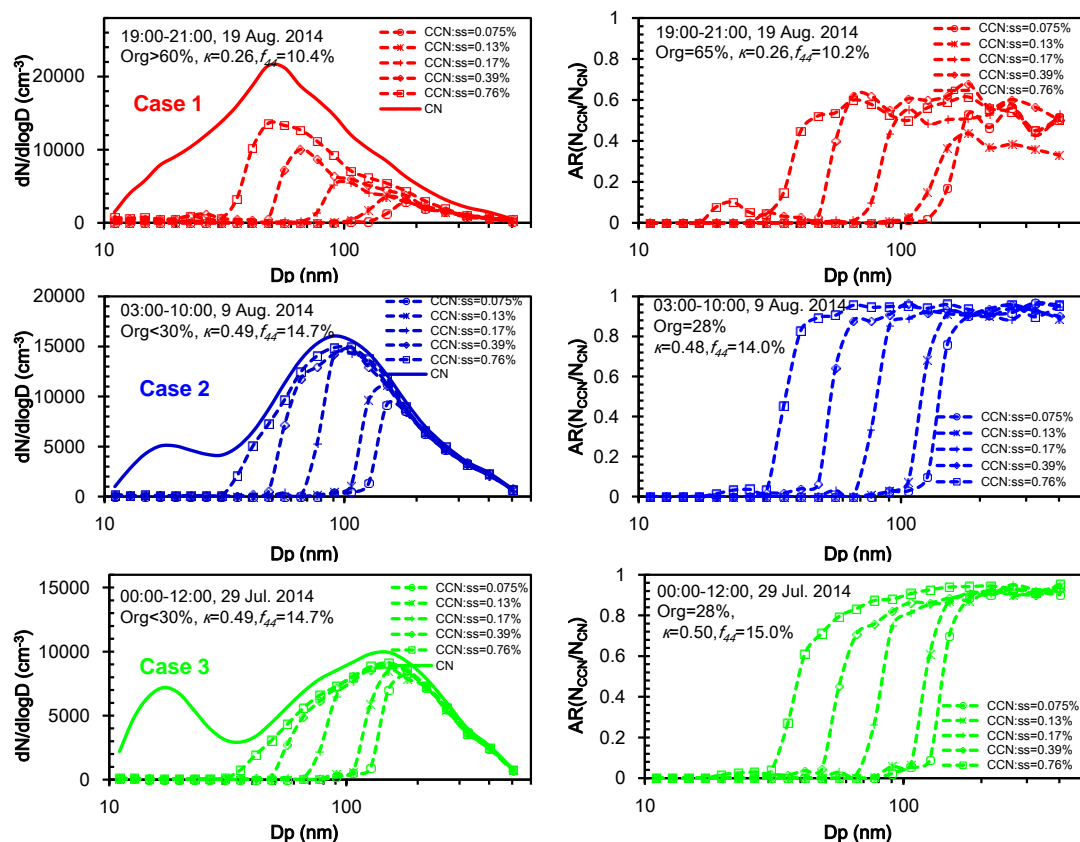


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Fig. 1. Mean CCN efficiency spectra at the Xinzhou site (black lines with asterisks) measured from 22 July-26 August 2014 and at the Xianghe site (red lines with circles) site measured from 7-21 July 2013 for different supersaturation levels. Error bars representing one standard deviation are shown. Right panels show particle chemical composition in terms of mass concentration fractions at Xinzhou (top panel) and Xianghe (bottom panel) during their respective observation periods. The campaign average mass concentration of PM₁ is 31.6 μg m⁻³ and 72.4 μg m⁻³ at Xinzhou and Xianghe respectively. Note that the preset supersaturation levels were 0.07%, 0.1%, 0.2%, 0.4% and 0.8% at both sites, but effective supersaturation levels showed slightly different after calibration.

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 731 **Fig. 2.** Particle number size distribution (PSD) and CCN size distributions (left panels)

732 and CCN efficiency spectra (right panels) at different supersaturation levels for Case 1

733 (upper panels, 19 August 2014, 19:00-21:00 LT), Case 2 (middle panels, 9 August

734 2014, 03:00-10:00 LT), and Case 3 (bottom panels, 29 July 2014, 00:00-12:00 LT).

 735 Total CN number concentrations are 16671 cm^{-3} , 12869 cm^{-3} , and 10134 cm^{-3} for

 736 Case 1, Case 2, and Case 3, respectively. Mass concentrations of PM_{10} are $28.36 \mu\text{g m}^{-3}$

 737 m^{-3} , $81.45 \mu\text{g m}^{-3}$, and $78.73 \mu\text{g m}^{-3}$ for Case 1, Case 2 and Case 3, respectively.

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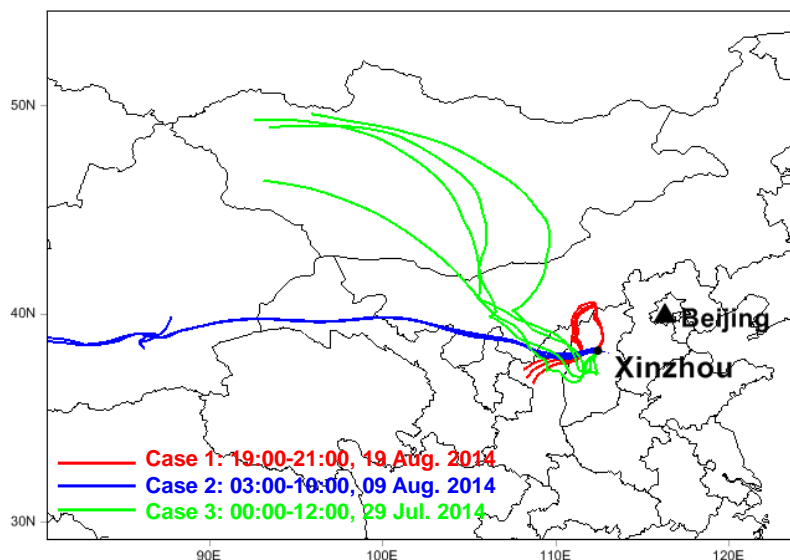
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749 **Fig. 3.** Five-day back trajectories for Case 1 (in red), Case 2 (in blue), and Case 3 (in
 750 green) calculated using the Hybrid Single-Particle Lagrangian Integrated Trajectory
 751 model with National Centers for Environmental Prediction reanalysis data. The arrival
 752 height of the trajectories at the Xinzhou site was at the surface.

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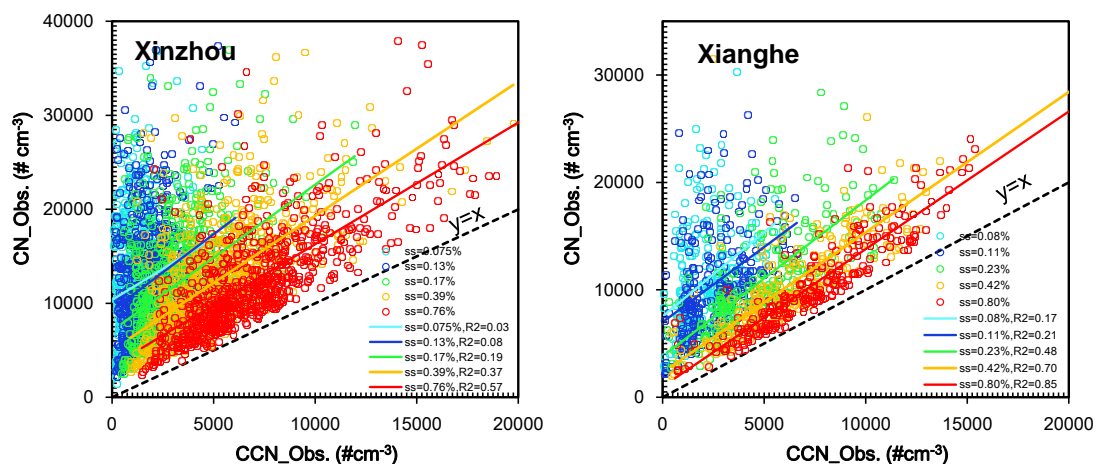
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774 **Fig. 4.** Measured N_{CN} as a function of measured N_{CCN} for different supersaturation
 775 levels at the Xinzhou (left panel) and Xianghe (right panel) sites. The scatterplot
 776 between CCN_{Obs} and CN_{Obs} were fitted with a linear function (in colored lines) and R^2
 777 refer to the correlations of them.

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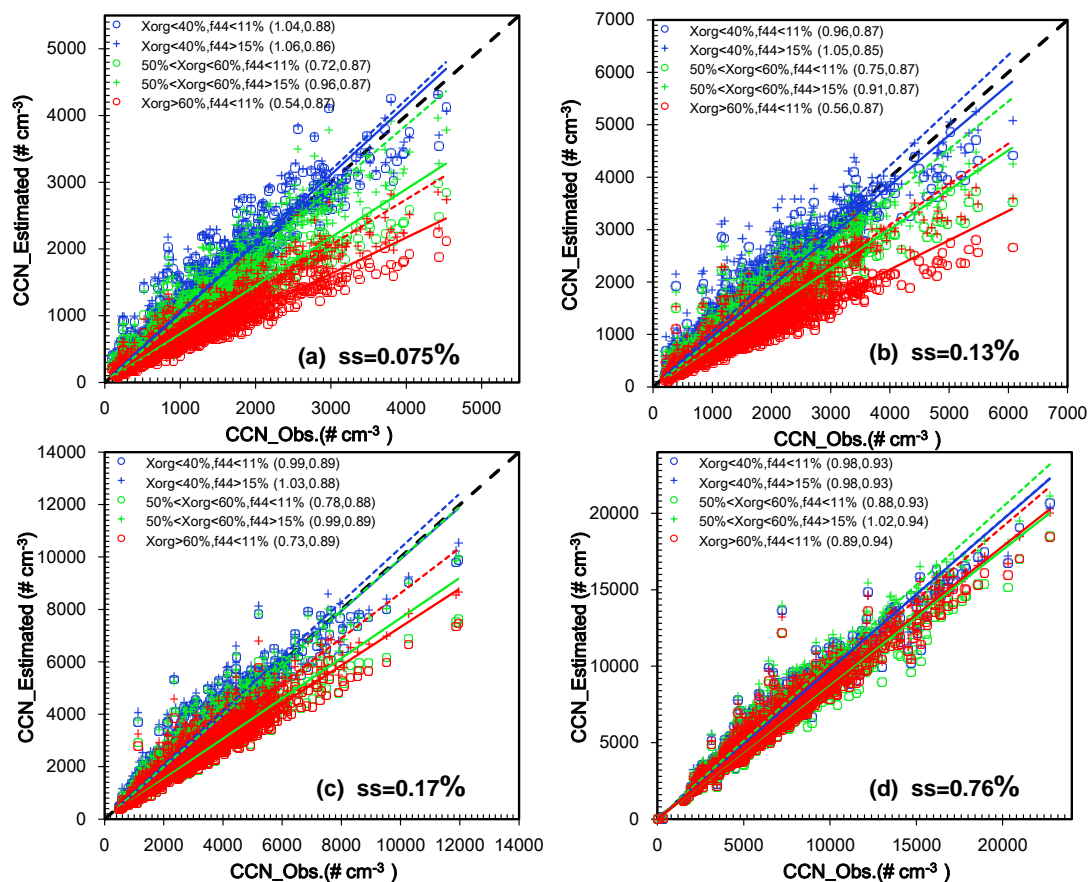
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800 **Fig. 5.** The sensitivity of N_{CCN} to both organics volume fraction (x_{org}) and oxidation
 801 level (using f_{44} , the fraction of m/z 44 in aerosol organic material) of organics at
 802 supersaturation levels of (a) 0.075%, (b) 0.13%, (c) 0.17% and (d) 0.76% for cases
 803 when $x_{org} = 35\%$ (blue circles), 52% (green circles), and 66% (red circles). The
 804 size-resolved CCN data were sorted when the $x_{org} > 60\%$, $50\% < x_{org} < 60\%$ and $x_{org} < 40\%$
 805 respectively to do the sensitivity examination. Linear best-fit lines through each group of
 806 points are shown. Slopes and R^2 values are given in parentheses.

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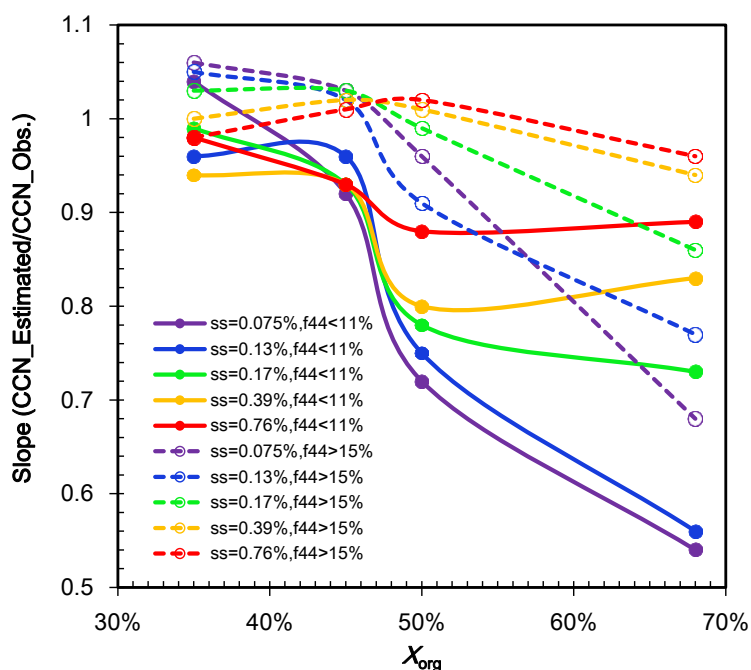
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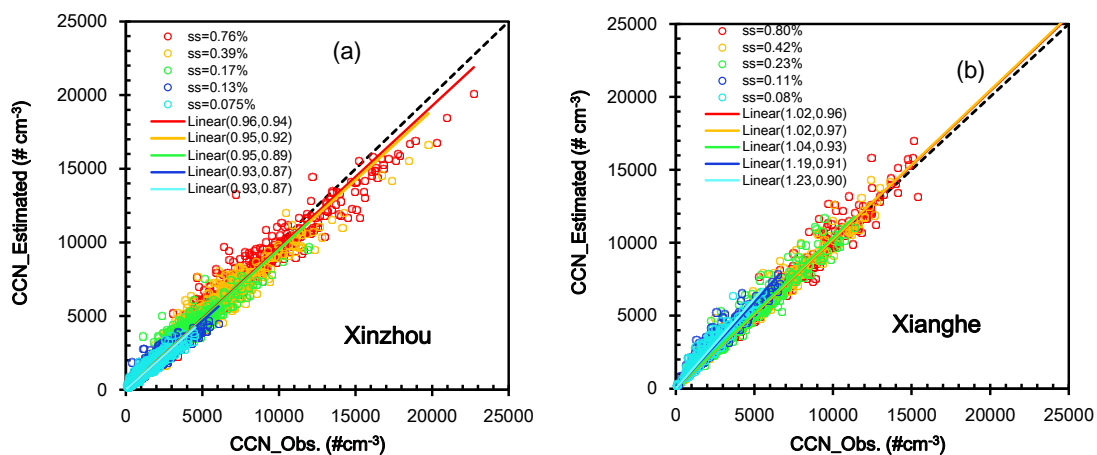
820 **Fig. 6.** Slopes of the linear fit of estimated and observed N_{CCN} dependence on volume
821 fraction of organics (x_{org}) at $f_{44}<11\%$ and $f_{44}>15\%$ for different supersaturation levels.
822 Mean values of the hygroscopic parameter κ_{chem} at $f_{44}<11\%$ when $x_{org} >60\%$, $50\% < x_{org}$
823 $<60\%$, $40\% < x_{org} < 50\%$ and $x_{org} < 40\%$ are 0.27, 0.34, 0.40 and 0.46, respectively;
824 while at $f_{44}>15\%$ the value increased to 0.36, 0.42, 0.46 and 0.50 respectively.

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843 **Fig. 7.** Estimated N_{CCN} as a function of observed N_{CCN} for different supersaturation
 844 levels at (a) Xinzhou and (b) Xianghe. Note that the campaign mean CCN efficiency
 845 spectra at Xinzhou are used for estimating N_{CCN} at Xianghe. Linear best-fit lines
 846 through each group of points are shown. Slopes and R^2 values are given in
 847 parentheses.

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