

1     **Response to comments of reviewer 1 on**

2     **”Ice phase in Altocumulus Clouds over Leipzig: Remote sensing observa-**  
3     **tions and detailed modelling” by Simmel et al.**

4     <http://www.atmos-chem-phys-discuss.net/15/1573/2015/>

5     We thank the reviewer for his/her constructive suggestions and for generally accepting  
6     the paper when the proposed revisions are realised.

7     Comments of the reviewers are cited in *italic*.

8     *General coments*

9     *The authors study two altocumulus cloud case studies that were observed over Germa-*  
10    *ny via ground-based remote sensing. The cases were selected to represent the warmest*  
11    *possible ice formation and more typically cold ice formation within altocumulus. The aut-*  
12    *hors apply an axisymmetric 1D model with spectral microphysics. The model is initialized*  
13    *with observed or model-derived thermodynamic profiles. Varying assumptions are made*  
14    *regarding prescribed vertical motions, aerosol and ice nucleus properties, and ice habit.*  
15    *Generally little work has been done on altocumulus microphysics, but that which has been*  
16    *done requires more review in the introduction and conclusions to motivate this work and*  
17    *to place the results into context. The approach is generally sound, but not enough details*  
18    *are provided to allow the work to be reproduced. The observations should be shown and*  
19    *described more completely. Overall, this work merits publication after revisions to the*  
20    *manuscript that can readily address specific comments below.*

21    *Specific comments (page/line number if relevant)*

22    1. *The scientific questions to be addressed are not adequately stated. Ice nucleation is*  
23    *discussed in the very short introduction, but no questions are targeted for this study. This*  
24    *is perhaps related to the problem that the authors provide no background on altocumulus.*  
25    *Has any study simulated such clouds before? Why did the authors choose to use a model?*  
26    *Why this model with elaborate microphysics but simple dynamics? Has any literature*  
27    *drawn conclusions about altocumulus relevant to this study? Does this study produce*  
28    *conclusions that are consistent with past literature? References should include Fleishauer*  
29    *et al. (JGR 59:1779, 2002), for instance.*

30    Altocumulus clouds are a good example for shallow mixed-phase clouds with compara-  
31    bly simple vertical structure — at least for the single-layered cases as they are considered  
32    here. Therefore, a dynamically simple model setup with prescribed vertical velocity was  
33    chosen to remain close to the observations. Feedback of microphysics on dynamics is

34 not considered to concentrate on primary microphysical effects and to avoid misleading  
35 conclusions about secondary effects due to changed dynamics. A model intercomparison  
36 study by Ovchinnikov et al. (2013, doi:10.1002/2013MS000282) has shown that bulk  
37 microphysical models tend to underestimate ice growth by vapor deposition due to the  
38 underlying ice distribution assumptions. In contrast to this, bin models directly simulate  
39 the shape of the distributions and, therefore, no assumptions concerning the shape have  
40 to be made.

41 The underlying topic is mixed-phase microphysics and the interaction between the  
42 three phases of water. It is well-known that due to the different saturation pressure  
43 of water vapor with respect to liquid water and ice, a mixed-phase cloud is in a non-  
44 equilibrium state which, nevertheless, may lead to a quasi-steady existence (e.g., Korolev  
45 and Field, 2008, JAS). To study those interactions, a bin model is suited well, since  
46 condensational/depositional growth is not only described by saturation adjustment but  
47 by a detailed description of sub-/supersaturation of each size bin resulting in different  
48 growth rates. This automatically results in a very detailed description of the Wegener-  
49 Bergeron-Findeisen (WBF) process which drives the phase interaction.

50 The main drivers for this phase transfer are vertical velocity (leading to supersatu-  
51 ration and subsequent droplet formation) and ice particle formation and growth (WBF  
52 starts) leading to sedimentation of the typically fast growing ice particles (WBF ends  
53 due to removal of ice). The motivation of this work is to shed more light on the relative  
54 contributions of the different processes involved in these complex interactions (see also  
55 response to review 3).

56 The manuscript was changed accordingly.

57 *2. The observations that motivated the selection of these cases, and which are relied*  
58 *upon, are not adequately shown and their uncertainty properties are not described. Figure*  
59 *1b makes a good start at showing case 1 cloud conditions, but other case 1 figures are*  
60 *truncated in time. Please show all five of the following fields between 23:45 and 0:40*  
61 *for case 1 (providing important context for the narrow 20-minute window used for the*  
62 *study) and for case 2: lidar backscatter, radar reflectivity, retrieved IWC, retrieved LWC,*  
63 *retrieved vertical wind. Only the first is shown for the full time range for case 1. LWC*  
64 *is never shown now. Also please report the stated or estimated uncertainty properties of*  
65 *IWC, LWC and vertical wind speed. Are there no clear-air vertical wind retrievals from*  
66 *the Doppler lidar? Please explain why the vertical wind speeds shown in Figure 2 appear*  
67 *as they do for lidar. Finally, please show plots of the initial soundings used, including*  
68 *RHI and RH.*

69 Additional pictures are shown for both cases in the revised version. However, for case  
70 1 full time range is shown only for 2 parameters (RC signal, radar reflectivity, new Fig.  
71 1) because at 0:22 h, a new cloud appears to form at a lower level (compare also humid  
72 layer in profile, Fig. 7) around 3000 m.

73 Accuracy of the IWC is +/-50 %. For the LWC calculated by the scaled adiabatic  
74 approach the same order of magnitude applies. Vertical wind speeds are measured di-

75 rectly by evaluation from the recorded cloud radar and Doppler lidar spectra. Errors are  
76  $\pm 0.15$  m/s for the cloud radar and  $\pm 0.05$  m/s for the Doppler lidar. These errors are  
77 mainly due to the pointing accuracy of the two systems.

78 The Doppler lidar (right panel) shows the motion of small cloud droplets at the predo-  
79 minantly liquid cloud top. Hence, in this plot the cloud-top turbulence becomes visible.  
80 The cloud radar (left) mainly shows particles falling from the top layer, therefore, par-  
81 ticles are mainly moving downwards (green color). Only at the very top at about 4300  
82 m particles are small enough to still be lifted upwards (yellow colors).

83 There are no possibilities to derive clear air velocity with a coherent doppler wind  
84 lidar, because this instrument depends on tracer targets like aerosol particles or cloud  
85 droplets. However, clear air motions around a cloud is a very interesting quantity which  
86 can, e.g., be derived with radar wind profilers. See for example: [http://www.atmos-meas-](http://www.atmos-meas-tech-discuss.net/8/353/2015/amtd-8-353-2015.pdf)  
87 [tech-discuss.net/8/353/2015/amtd-8-353-2015.pdf](http://www.atmos-meas-tech-discuss.net/8/353/2015/amtd-8-353-2015.pdf)

88 Initial soundings (T, rh, rhi) are shown as new Fig. 7.

89 The manuscript was changed accordingly.

90 *3. The authors acknowledge that the specification of vertical winds is a controlling*  
91 *parameter, but they do not discuss the general nature of these winds, which seems to be*  
92 *important to understanding the relationship of the model setup to the large-scale con-*  
93 *ditions. Is the mean vertical wind described large-scale in nature whereas the stochastic*  
94 *components are turbulence? I would expect the updrafts and downdrafts within altocu-*  
95 *mulus to be driven by cloud-top cooling rather than large-scale winds. In the case that*  
96 *cloud-top cooling-driven turbulence is driving mixing between downdrafts and updrafts, I*  
97 *would expect it to drive the supply of IN. However, the authors state that the mean wind*  
98 *is driving the supply of IN. Does that mean that large-scale convergence is driving the*  
99 *supply of IN to updrafts and downdrafts whereas turbulence does not play a role in the*  
100 *supply of IN?*

101 We share your statement that cloud-top cooling is an important driver for altocumulus  
102 clouds. We consider this effect to be included in the observations as well as in the  
103 prescribed vertical velocity.

104 In the paper, we state that the supply of IN is driven by the horizontal exchange with  
105 the outer cylinder. The horizontal exchange is driven by the change of vertical wind  
106 speed with height (see Eq. (4)). This means that turbulence (which is responsible for  
107 the direction of vertical wind speed) plays a major role in IN supply.

108 The mean vertical wind can be considered as large-scale driving force, however, due  
109 to the model configuration, the strength of the mean updraft has to be chosen larger  
110 than observed.

111 *4. The model vertical resolution is 25 m, but what is the size of the inner and outer*  
112 *cylinder? How was it decided how large to make the inner and outer cylindrical coor-*  
113 *dinates? Are results sensitive to the specification of cylinder relative size? Is the inner*  
114 *cylinder considered to be the whole 20-min cloud observed (both updrafts and downdrafts)*

115 *whereas the outer cylinder is the air surrounding the cloud? If so, how much air sur-*  
116 *rounding the cylinder? Or is the cylinder specified to be an updraft element size, similar*  
117 *to deep convection studies?*

118 The radius was chosen to be 100 m for the inner cylinder and 1000 m for the outer  
119 cylinder. In an Asai-Kawahara model the ratio of the radius of the inner cylinder to  
120 the radius of the outer cylinder is the dominating parameter and the chosen value of  
121 1:10 is a typical value for an Asai-Kawahara model setup. The results are sensitive to  
122 the radius ratio when the outer cylinder is chosen too small. Then the influence of the  
123 inner on the outer cylinder increases and the outer cylinder cannot serve as a proper  
124 background any more. However, the geometric configuration of the model is not intended  
125 to describe or to match the geometry of the clouds (and cloud-free spaces in between)  
126 as observed. It should rather be understood as a possibility to describe a vertically  
127 resolved cloud evolution and to provide the possibility of horizontal exchange with a  
128 cloud-free background (see also response to reviews 2 and 3). Neither is it intended to  
129 directly model the cases presented. They rather should serve as frame to judge whether  
130 the model simulations lead to results close enough to reality to apply the model to  
131 sensitivity studies. Therefore, the 60 minute model runs are not compared directly to a  
132 20 minute period of observations. However, the inner cylinder gives the relevant results  
133 for both, updrafts and downdrafts.

134 The manuscript was changed accordingly.

135 *5. Aggregation and riming are neglected? Please provide some literature support for*  
136 *why that would be appropriate or otherwise explain.*

137 The observed clouds are rather shallow and a large fraction of the ice is formed at/near  
138 cloud base which means that there is not that much possibility of ice particles to rime.  
139 Aggregation can be neglected due to the rather low ice particle number concentrations  
140 for case 1 (relative little probability of collision between particles) and the relative low  
141 temperatures for case 2 (reducing sticking efficiency). This assumption is corroborated  
142 by the findings of Smith et al. (2009, doi:10.1029/2008JD011531) stating that water  
143 vapor deposition (and sublimation), balanced by sedimentation are more important than  
144 accretional growth.

145 The manuscript was changed accordingly.

146 *6. (1581/21) M1 and M6 both have lower free troposphere aerosol. Why did you choose*  
147 *M6 for case 1 and M1 for case 2? How did you apportion  $1e5/kg$  aerosol among the three*  
148 *modes?*

149 For case 2 the upper free troposphere aerosol distribution of M1 was used, whereas  
150 for case 1 the lower free troposphere aerosol of M6 was used. The choice of the aerosol  
151 distributions is quite arbitrary, however, one intention was to use M6 LFT measurements  
152 with and without a polluted layer for case 1. Nevertheless, the polluted layer run was

153 not reported since Lidar observations showed no polluted layers for case 1. For the UFT,  
154 no polluted layers were observed in Petzold et al., therefore, we decided to use M1 from  
155 the beginning.

156 We assume that  $1e5/kg$  particles are larger (in radius) than 250 nm according to the  
157 parameterization of DeMott et al. to calculate and initialize the temperature-dependent  
158 INP field. This has to be considered separately from the AP distributions used for the  
159 initialization of the combined AP/drop spectrum. Those are taken as described in the  
160 paper cited.

161 *7. (1586/6, 1578/4) DeMott et al. (2010) did not analyze measurements colder than*  
162 *-9 C, to my knowledge. Did you extrapolate their relationship to colder temperatures? If*  
163 *so, how did you decide at what temperature to stop extrapolating when approaching 0 C?*

164 Yes. DeMott et al. (2010) only shows observations for temperatures below -9 C. We  
165 extrapolated the relationship to higher temperatures (-5 C). We did not have to decide  
166 where to stop the extrapolation in these case studies since in the model used ice formation  
167 by immersion freezing could only take place in the vicinity of drops which were only  
168 present at temperatures below -5/-6 C.

169 The manuscript was changed accordingly.

170 *8. (1579/2) Because the relationships in Mitchell et al. (1996) and past literature*  
171 *have been derived from observations over limited size ranges, it is not uncommon to*  
172 *use more than one relationship to represent columns of various sizes (e.g., Sölch et*  
173 *al. QJRMS 136:2074, 2010, table AII). Please provide sufficient information re exactly*  
174 *which relationships you used and over what size ranges for this work to be reproduced.*

175 We used the relationships in Mitchell et al. (1996) in their Tab. 1 for hexagonal plates  
176 and hexagonal columns. The mass-dimension power laws were transformed to aspect  
177 ratios for the given shapes. For columns, three size ranges (30 to 100  $\mu\text{m}$ , 100 to 300  $\mu\text{m}$ ,  
178 and above 300  $\mu\text{m}$  in diameter) with different coefficients are given whereas for plates  
179 the coefficients are valid for diameters from 15 to 3000  $\mu\text{m}$ .

180 The manuscript was changed accordingly.

181 *9. (1575/5) Could preconditioned ice nuclei be nucleated as warm as -1 degrees C*  
182 *or some other temperature limit? Please explain mechanistically how preconditioning*  
183 *could introduce ice nuclei relevant in this study, with reference to literature and relevant*  
184 *temperature range.*

185 The statement was removed from the text since it was too speculative.

186 *10. (1576/30) IWC is shown to 2000 m in Figure 1, which apparently is warmer than*  
187 *0 C according to the text, which states that IWC extends to only 3000 m. Please clarify.*

188 Indeed there is no ice detected below 0 C. The IWC is derived by the parameterization  
189 of Hogan 2006 which computes IWC as a simple function of radar reflectivity and tem-  
190 perature. The equation is mathematically valid for  $T > 0$  C, so the usage of this equation  
191 has to be restricted to temperatures below 0 C. That restriction was, however, not done  
192 properly done in this case. The figure was therefore corrected and now shows IWC only  
193 up to 0 C.

194 *11. What is the model time step used?*

195 For the dynamics as well as for the microphysics a time step of 1 s is used.  
196 The manuscript was changed accordingly.

197 *12. (1582/22) It is stated that "ice forms primarily at cloud base". Does this mean*  
198 *that ice is primarily nucleated at cloud base? Cloud base is warmer than cloud top, so I*  
199 *would expect more rapid nucleation at cloud top. Please explain.*

200 When drops form at cloud base all available INP active at cloud base temperature  
201 can contribute to primary ice formation in the immersion mode. The unfrozen droplets  
202 are transported further upwards which results in cooling. Nevertheless, if the cloud is  
203 relatively shallow (which is the case here) the temperature difference between cloud  
204 base and top is rather small. Therefore, the additional number of active INP causing ice  
205 nucleation in the upper parts of the cloud remains also relatively small. Therefore, in  
206 summary, more ice particles are nucleated near cloud base than near cloud top in the  
207 cases presented here.

208 *Technical corrections*

209 *1 (1576/3). Please define TROPOS.*

210 TROPOS is the Leibniz-Institute for Tropospheric Research.  
211 The manuscript was changed accordingly.

212 **Response to comments of reviewer 2 on**

213 **”Ice phase in Altocumulus Clouds over Leipzig: Remote sensing observa-**  
214 **tions and detailed modelling” by Simmel et al.**

215 **<http://www.atmos-chem-phys-discuss.net/15/1573/2015/>**

216 We thank the reviewer for his/her constructive suggestions and for generally accepting  
217 the paper when the proposed revisions are realised.

218 Comments of the reviewers are cited in *italic*. Revised manuscript with highlighted  
219 changes is given as supplement.

220 *This paper has great potential as a comparison between modeled and observed mixed*  
221 *phase clouds. Its strengths are the high quality remote observations and the relatively*  
222 *direct modeling approach that allows for straightforward implementation and comparison*  
223 *of different ice nucleus (IN) concentrations and ice crystal shapes. The paper stumbles*  
224 *before reaching the finishing line, so I encourage the authors to improve the paper to its*  
225 *potential. There are a number of problems with the analysis and presentation, as detailed*  
226 *below, but these are relatively minor aspects that can be improved with modest effort. The*  
227 *major shortcoming of the paper is the complete absence of comparison between modeled*  
228 *and observed properties in section 5. This is the section where the most interesting science*  
229 *is finally addressed, through variation in IN concentration and ice crystal shape and fall*  
230 *speed. Is it possible to vary these parameters and obtain results that compare with the*  
231 *lidar and radar observations with higher fidelity? And as such, can the suitability of IN*  
232 *parameterization or crystal habit representation be evaluated? As a single example, there*  
233 *is discussion of the “stronger tilting of the virgae” for nonspherical ice (page 1589). This*  
234 *seems like a perfect aspect to compare to observations. It is only one example, and in*  
235 *general, there needs to be a much more thorough and, to the extent possible, quantitative*  
236 *comparison between modeled and observed mixed-phase properties in this section.*

237 The presentation of observations is extended by showing IWC and LWC for both ca-  
238 ses. Comparison between model and observation seems to be difficult for case 1 where  
239 the observations are close to the detection limit. Additionally, the INP parameterization  
240 of DeMott is rather insensitive to the number of aerosol particles at rather high tempe-  
241 ratures of -5 C. For case 2 it seems to be clear that either too many INP or non-spherical  
242 particles could easily lead to an overestimation of the ice-phase and even the complete  
243 depletion of the liquid phase which is in contradiction to the observations.

244 Conclusions about possible ice shapes being consistent with (a) laboratory studies and  
245 (b) our observations are drawn in the final section.

246 *The following points should also be addressed:*

247 - Abstract: “warm temperatures” should be “high temperatures” (air is warm, tempe-  
248 ratures are high).

249 - Pg 1574 line 22: “attributed the aerosol” should be “attributed to the aerosol”.

250 The changes were done according to the suggestion of the reviewer.

251 - Pg 1574 line 25: I do not understand the statement that only biological particles form  
252 ice above -15 C. In the parameterization employed, which is mostly describing dust, IN  
253 exist at much higher temperatures.

254 There is an ongoing discussion of this topic. In laboratory studies, it was shown, that  
255 biological material is able to initiate ice at those high temperatures. However, there are  
256 at least two possibilities for dust to form ice above -15 C: (a) Pure dust is also able to  
257 form ice above -15 C if only enough material (surface) is available. This is a question of  
258 detection limits in lab studies (frozen drop fractions). Experiments with large drop on  
259 freezing arrays at least hint to this possibility. (b) Dust is mixed with biological material  
260 (forming soil dust). Ice formation in this case is triggered by the biological material at  
261 least at higher temperatures.

262 - Pg 1575 line 1: should be “to what extent”.

263 The changes were done according to the suggestion of the reviewer.

264 - Pg 1575 line 5-6: reference needed for this statement.

265 Statement was removed from the text since it was too speculative.

266 - Pg 1576 line 16: define GDAS.

267 - Pg 1576 line 24: “could be observed” should be “was observed”.

268 - Pg 1577 line 1: “an LWP” should be “a LWP”.

269 The changes were done according to the suggestion of the reviewer.

270 - Note: I stopped correcting minor grammatical errors after section 2. Authors, please  
271 proofread the paper carefully.

272 Careful proofreading was done.

273 - Pg 1577, sec 3: Asai-Kasahara type model should be described more thoroughly, e.g.,  
274 be clearer on cylindrical geometry, boundary conditions, etc.



275 For initialization of the Asai-Kasahara model, only a vertical profile of temperature  
276 and humidity is needed (now shown in new Fig. 7). No additional boundary conditions  
277 are needed. The radius was chosen to be 100 m for the inner cylinder and 1000 m for the  
278 outer cylinder. In an Asai-Kasahara model the ratio of the radius of the inner cylinder  
279 to the radius of the outer cylinder is the dominating parameter and the chosen value of  
280 1:10 is a typical value for an Asai-Kasahara model setup. The results are sensitive to  
281 the radius ratio when the outer cylinder is chosen too small. Then the influence of the  
282 inner on the outer cylinder increases and the outer cylinder cannot serve as a proper  
283 background any more. However, the geometric configuration of the model is not intended  
284 to describe or to match the geometry of the clouds (and cloud-free spaces in between) as  
285 observed. It should rather be understood as a possibility to describe a vertically resolved  
286 cloud evolution and to provide the possibility of horizontal exchange with a cloud-free  
287 background (see also response to reviews 1 and 3).

288 The manuscript was changed accordingly.

289 - *Sec 3.1.1: Regarding “Immersion freezing occurs as soon as liquid drops above a*  
290 *certain size limit are present”, why is there a drop size dependence? Freezing probability*  
291 *should be related to IN properties, not to volume of drop.*

292 One possible way of the drop volume to influence freezing is the concentration of solved  
293 chemical species which may lead to a freezing point depression in the case of relative  
294 large aerosol particles with relatively little water mass. The size limit is intended to make  
295 sure that there are supercooled drops available for freezing and to avoid the freezing of  
296 aerosol particles which are present in the same joint spectral liquid-phase field.

297 - *Sec. 3.2.1 and Fig 5: What is the advantage of using a stochastic forcing for vertical*  
298 *velocity? It seems to only add complexity, with no obvious illumination of new physics.*  
299 *Why not force with a deterministic, e.g., sinusoidal, vertical velocity, for example?*

300 A simpler profile was tested (constant up- and downdrafts for given times with short  
301 linearly interpolated transitions between both), however, it appears that this more com-  
302 plex stochastic forcing gives more realistic results and better shows the variety of cloud’s  
303 LWC and IWC since it better matches the temporal patterns of the updraft. Addition-  
304 ally, the more often changes between up- and downdrafts on smaller time scales provide  
305 a certain horizontal exchange between inner and outer cylinder which is important for  
306 the supply of fresh INP.

307 - *Pg 1583 line 27: I think “presence time” is clearer as “residence time”.*

308 The change was done according to the suggestion of the reviewer.

309 - *Sec 4: Comparison with figures and reported results is not straightforward: for ex-*  
310 *ample, figures are in g/kg, observations are in kg/m<sup>3</sup>. Please be consistent.*

311 To be consistent, all model results were changed to  $\text{g/m}^3$ .

312 - *Fig 1: Left panel should be labeled as  $\log_{10}$  of IWC.*

313 The change was done according to the suggestion of the reviewer.

314 - *Fig 1: What is the purpose of the dashed box? Maybe I missed it in the text, but it*  
315 *should also be specified in the caption.*

316 Fig. 1a and 1b show different height and time ranges. Since Fig. 1b shows a larger  
317 part of the data, the dashed box indicates the region shown in Fig. 1a. This was clarified  
318 in the revised manuscript (new Fig. 1).

319 - *Figs 6, 8, 10, 12: How useful are these comparisons? The differences between the*  
320 *panels are so small that it is not clear to me that they need to be presented graphically.*  
321 *The numerical results such as max LWMR and max IWMR may be adequate, unless*  
322 *details of the plots are specifically discussed in the text.*

323 In our opinion, there are quite significant differences between the cases shown which  
324 were already mentioned in the manuscript. In Fig. 6, the main differences can be seen  
325 in the liquid phase (contours), but also in the ice phase (increase of the ice phase cloud  
326 base). Fig. 8 shows substantial differences in both phases, Fig. 10 again shows differences  
327 mainly in the liquid phase. Fig. 12 (upper panels) show large differences in both phases,  
328 whereas Fig. 12 (lower panels) illustrates the different timing due to the changed forcing.

329 - *Figs 6, 8, 10, 12, 14, 15, 17, 18: Maximum ice and liquid water values are reported*  
330 *with 7 significant digits. It cannot be that such accuracy is valid. (Also note that there*  
331 *is some inconsistency in using max LWMR versus max. drop water, etc.).*

332 There is no valid 7 digit accuracy. This was changed to 3 significant digits. In the  
333 revised paper, LWC and IWC are used consistently.

334 - *Figs 15 and 18: The captions state that liquid is denoted by color and ice water mass*  
335 *is denoted by contours. That seems to be backwards.*

336 Yes. This was changed.

337 - *Fig 16: In the bottom of the left panel, please confirm that all lines are plotted (i.e.,*  
338 *are they identical and cannot be distinguished?)*

339 All lines are identical. Compared to the liquid fraction, the ice fraction is so small  
340 that changes as modelled in the sensitivity runs are too small to affect the liquid phase  
341 considerably (compare LWMR (now LWC) maxima in Table 3 of case 1).

342 **Response to comments of reviewer 3 on**

343 **"Ice phase in Altocumulus Clouds over Leipzig: Remote sensing observa-**  
344 **tions and detailed modelling" by Simmel et al.**

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346 We thank the reviewer for his/her constructive suggestions and for generally accepting  
347 the paper when the proposed revisions are realised.

348 Comments of the reviewers are cited in *italic*. Revised manuscript with highlighted  
349 changes is given as supplement.

350 *General comments: The authors simulate two mixed-phase cloud layers that were ob-*  
351 *erved by remote sensing. However, the dynamical model used is unsophisticated, and*  
352 *key observations needed to initialize, force, and validate the simulations are unavailable.*  
353 *For instance, the study concludes that IWP is sensitive to IN number, but there are no*  
354 *IN measurements to assess how much IN number should be varied in the sensitivity stu-*  
355 *dy. Hence the conclusions must necessarily be regarded as tentative. Furthermore, only*  
356 *two cloud cases are examined, limiting the conclusions' generality. If the authors wish*  
357 *to simulate a case study, then I recommend that they choose a more complete dataset to*  
358 *simulate, one that uses more accurate (e.g. in situ) measurements. Also, I recommend*  
359 *that they use a more sophisticated model (e.g. LES). If the authors wish to do an ob-*  
360 *servational study, then I recommend that they exploit the instruments they have. Given*  
361 *the facts that the set of instruments is incomplete, that none are in situ, but that they*  
362 *can be run continuously, the instruments seem better suited to assessing climatological*  
363 *relationships between variables. If, instead, the authors wish to invest the time to maxi-*  
364 *mize the usefulness of the present study, I would attempt to better quantify the statement*  
365 *"the liquid phase is mainly determined by the model dynamics (location and strength*  
366 *of vertical velocity) whereas the ice phase is much more sensitive to the microphysical*  
367 *parameters (ice nuclei (IN) number, ice particle shape)." In particular, instead of vary-*  
368 *ing w\_ave from 0.1 to 0.4 m/s, I would vary it by "observed" values taken from obs or*  
369 *reanalyses or the literature. Instead of varying N\_AP by a factor of 10, vary it by the*  
370 *suitable range given by values in the literature. That would provide a better sense of the*  
371 *practical sensitivity of LWP and IWP to w\_ave versus N\_AP. Consider doing likewise*  
372 *for the other sensitivity experiments.*

373 The general idea of the paper is to use a two step approach: In the first step, the model  
374 is used to simulate a cloud which is close to the one observed (in terms of model input –  
375 e.g., temperature/humidity profile, vertical velocity – and output – e.g., cloud evolution,  
376 liquid and ice phase). If this is done successfully, the model can be used for a second  
377 step which is a sensitivity study with respect to certain parameters. This sensitivity

378 study is done within the 'model world' by varying the respective input data (e.g., INP  
379 number) and comparing the results.

380 The variation of N<sub>AP</sub> is in the range of the observations which are the basis of  
381 the parameterization used. Only the very high concentrations were omitted since no  
382 polluted layers were observed by the lidar.

383 It must be stated that the variation of the w<sub>ave</sub> is caused by the model configuration:  
384 If w<sub>ave</sub> is chosen to be smaller than about 0.1 m/s the model will not be able to reach  
385 supersaturation and to form a cloud due to the horizontal exchange with the background.  
386 On the other hand, if w<sub>ave</sub> is chosen much larger than 0.4 m/s, the downdrafts will be  
387 too weak and too short to lead to cloud-free spaces in between the clouds.

388 *Specific comments:*

389 *The abstract is well written, but the introductory section could more clearly introduce*  
390 *the main issues that will be addressed in the paper. What is the gap in knowledge, and*  
391 *how will it be addressed in the subsequent sections?*

392 The underlying topic is mixed-phase microphysics and the interaction between the  
393 three phases of water. It is well-known that due to the different saturation pressure  
394 of water vapor with respect to liquid water and ice, a mixed-phase cloud is in a non-  
395 equilibrium state which, nevertheless, may lead to a quasi-steady existence (e.g., Korolev  
396 and Field, 2008, JAS). For this purpose, a bin model is suited well, since condensatio-  
397 nal/depositional growth is not only described by saturation adjustment but by a de-  
398 tailed description of sub-/supersaturation of each size bin resulting in different growth  
399 rates. This automatically results in a very detailed description of the Wegener-Bergeron-  
400 Findeisen (WBF) process which drives the phase interaction.

401 The main drivers for this phase transfer are vertical velocity (leading to supersatu-  
402 ration and subsequent droplet formation) and ice particle formation and growth (WBF  
403 starts) leading to sedimentation of the typically fast growing ice particles (WBF ends  
404 due to removal of ice). The motivation of this work is to shed more light on the relative  
405 contributions of the different processes involved in these complex interactions (see also  
406 response to review 1).

407 *p. 1576: "The liquid part of the cloud extends from about 4250 to 4450 m height at*  
408 *temperatures of about -6 C according to the GDAS reanalysis data for Leipzig." Some*  
409 *of the discussion relates to the temperature at which various IN are active. Therefore,*  
410 *it is of relevance to know: What are the error bars on the temperature measurement? I*  
411 *wouldn't expect a reanalysis to be terribly accurate.*

412 Temperature errors of the GDAS data compared to radiosonde profiles over Leipzig  
413 has been determined to be +/-1K during the DRIFT-project by Patric Seifert (see  
414 <http://onlinelibrary.wiley.com/doi/10.1029/2009JD013222/pdf>). These errors seem to

415 be sufficiently small to allow for a strong connection between temperature (as deduced  
416 from GDAS reanalysis) and potential ice formation processes.

417 *p. 1577: “For the model studies an Asai–Kasahara type model is used (Asai and Kasa-*  
418 *hara, 1967). The model geometry is axisymmetric and consists of an inner and an outer*  
419 *cylinder.” By today’s standards, the Asai-Kasahara model is crude. Instead, I recommend*  
420 *using a large-eddy simulation (LES) model. These days, LES are affordable and easy to*  
421 *configure. If not LES, then I recommend trying a prescribed dynamics model like the*  
422 *Kinematic Driver (KiD) model, because it will provide flexibility and control.*

423 Maybe the term “Asai-Kasahara” is misunderstood. As it is explained, vertical dyna-  
424 mics is prescribed which to our understanding is rather similar to the KiD model. Only  
425 the model geometry assumption (cylinder-symmetric with an inner and an outer cylin-  
426 der) and the exchange between the cylinders (see Eq. (4)) relates to the Asai-Kasahara  
427 model.

428 However, the geometric configuration of the model is not intended to describe or to  
429 match the geometry of the clouds (and cloud-free spaces in between) as observed. It  
430 should rather be understood as a possibility to describe a vertically resolved cloud evo-  
431 lution and to provide the possibility of horizontal exchange with a cloud-free background  
432 (see also response to reviews 1 and 2).

433 *p. 1577: “Since during the above mentioned observations no measurements of the IN*  
434 *are available, the parameterization of DeMott et al. (2010) is used assuming that all*  
435 *IN are active in the immersion freezing mode.” The observations needed to address the*  
436 *scientific questions are lacking. Consider focusing your efforts on addressing a question*  
437 *that your instruments are better positioned to answer.*

438 In general, ambient INP measurements are sparse and typically not available for long-  
439 term observations. We do not think that this fact should deter us from investigating  
440 those cases. It is a common approach to use certain assumptions (here about INP) and  
441 to check how the model results based on those assumptions compare to observations.  
442 Additionally, sensitivity studies are carried out to check how important the respective  
443 parameter is for the whole situation.

444 *p. 1579: “For case 1, profiles from both methods show a similar general behaviour but*  
445 *the radiosonde profile of Meiningen measured at 00:00 UTC is used since it provides a*  
446 *finer vertical resolution than the GDAS reanalysis data. However, for case 2 the Meinin-*  
447 *gen RS profile misses the humidity layer at the level where the clouds were observed and,*  
448 *therefore, GDAS reanalysis data for Leipzig at 21:00 UTC were chosen.” Apparently, the*  
449 *observations are too inaccurate to initialize the simulations.*

450 The Meiningen profile was not representative for Leipzig for case 2. Therefore, the  
451 GDAS profile was chosen as a substitute. Despite the coarser height resolution, cloud

452 formation was triggered in the model when vertical updrafts similar to the observed  
453 ones were prescribed. Again, we have to emphasize that the aim of the study was not  
454 to model the observed cases in detail but more to obtain reasonable model results that  
455 allow for sensitivity studies which are in turn transferable to the "real world".

456 *p. 1581: "Since no in situ aerosol measurements are available, literature data is used."*  
457 *The dataset is inadequate for the purpose of studying sensitivity to IN.*

458 The Lidar shows that no dust layers or similar pronounced features concerning aerosol  
459 could be observed. Therefore, we consider it reasonable to use those literature data. The  
460 aim of the study is not to study sensitivity of the clouds with respect to INP on the  
461 basis of observations. If this was the case, we would have to have measurements of both,  
462 cloud ice phase as well as INP, to obtain e.g., statistical correlations between both data  
463 sets. However, we use a two step approach mentioned above which allows us to perform  
464 the sensitivity study in the 'model world'.

465 **Relevant changes made in the manuscript**

466 (also compare the version of the manuscript with the highlighted changes compared to  
467 the original version)

468 **Section 1: Introduction** In the introduction, material concerning altocumulus in gene-  
469 ral and the Wegener-Bergeron-Findeisen process was added.

470 **Section 2: Observations** In section 2, more observational data is presented in the  
471 revised manuscript (two additional figures and the corresponding text) and the error  
472 range of the observations is discussed shortly.

473 **Section 3: Model description and initialization**

- 474 • The model description was extended with respect to model geometry, time step  
475 and driving velocity.
- 476 • The reasons for neglecting collision processes of ice particles was explained and  
477 supporting literature was cited.
- 478 • Figure 7 showing the vertical profiles of temperature as well as relative humidities  
479 with respect to liquid water and ice for both cases was added.

480 **Section 4: Model results: Dynamics** Reasons for preferred ice formation near liquid  
481 cloud base were explained.

482 **Section 5: Sensitivity studies** Drying of the layer below liquid cloud base due to ice  
483 particle growth with subsequent sedimentation and consequences for LWC were explai-  
484 ned.

# Ice phase in Altopcumulus Clouds over Leipzig: Remote sensing observations and detailed modelling

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**Abstract.** The present work combines remote sensing observations and detailed cloud modeling to investigate two altopcumulus cloud cases observed over Leipzig, Germany. A suite of remote sensing instruments was able to detect primary ice at rather warm-high temperatures of  $-6^{\circ}\text{C}$ . For comparison, a second mixed phase case at about  $-25^{\circ}\text{C}$  is introduced. To further look into the details of cloud microphysical processes a simple dynamics model of the Asai-Kasahara type is combined with detailed spectral microphysics forming the model system AK-SPECS. Vertical velocities are prescribed to force the dynamics as well as main cloud features to be close to the observations. Subsequently, sensitivity studies with respect to ice microphysical parameters are carried out with the aim to quantify the most important sensitivities for the cases investigated. For the cases selected, the liquid phase is mainly determined by the model dynamics (location and strength of vertical velocity) whereas the ice phase is much more sensitive to the microphysical parameters (ice nuclei (IN) nucleating particle (INP) number, ice particle shape). The choice of ice particle shape may induce large uncertainties which are in the same order as those for the temperature-dependent IN-INP number distribution.

## 1 Introduction

Altopcumulus clouds. According to Warren et al. (1998a,b) altopcumulus and altostratus clouds together cover 22 % of the earth's surface. For single-layered altopcumulus clouds, observations by Buehl et al. (2013) show the typical feature with a maximum of liquid water in the upper part of the cloud (increasing with height) and an ice maximum in the lower part of the cloud, mostly below liquid cloud base down in the virgae. This was previously reported from Fleishauer et al. (2002) and Carey et al. (2008).

Fleishauer et al. (2002) also emphasized a lack of significant temperature inversions or wind shears as a major feature of these clouds. Kanitz et al. (2011) show that the ratio of ice-containing clouds increases with decreasing temperature. However, the numbers are different for different locations with similar dynamics but with different aerosol burden, e.g., at northern and southern midlatitudes, underlining the question for the influence of ice-nucleating particles (INP). The observations with the highest temperatures are close to the limit at which the best atmospheric ice nuclei are known to nucleate ice in the immersion mode. This can only be attributed to the aerosol particles which are formed out of or at least contain biological material such as bacteria (Hartmann et al., 2013), fungi, or pollen. This is corroborated by the review of Murray et al. (2012) stating that only biological particles are known to form ice above  $-15^{\circ}\text{C}$ . However, these observations are from laboratory studies and it is still unclear whether or to which-what extent these extremely efficient ice nuclei are abundant in atmosphere, especially above the boundary layer. One idea is that freezing is caused by soil dust with biological particles dominating the freezing behaviour (O'Sullivan et al., 2014) which could explain on the one hand the atmospheric abundance of biological material and on the other hand the relatively high freezing temperatures above  $-15^{\circ}\text{C}$  of ambient measurements. Seeding from ice clouds above can be excluded for the cases presented ; however, preconditioning of ice nuclei (IN) could be a reason for these high ice activation temperatures - which means that ice has formed at the cloud temperatures observed.

Ice nucleation still is a large source of uncertainty in cloud modeling. Recently, several studies use combinations of vertically fine resolved models with rather detailed representation of the ice nucleation processes. Often, wave clouds are used as comparison since they represent rather ideal condi-



tions when they are not influenced by ice seeding from layers above. Field et al. (2012) apply a 1D kinematic model with bulk microphysics but prognostic ~~IN-INP~~. Eidhammer et al. (2010) use a Lagrangian parcel model for the comparison of the ice nucleation schemes of Phillips et al. (2008) and DeMott et al. (2010) under certain constraints. A 1D column model with a very detailed 2D spectral description of liquid and ice phase is employed by Dearden et al. (2012). Sun et al. (2012) use a 1.5D model with spectral microphysics for shallow convective clouds for a sensitivity study of immersion freezing due to bacteria and its influence on precipitation formation.

Most ice microphysics descriptions in models lack from the fact that ice nuclei are not represented as a prognostic variable. These models diagnose the number of ice particles based on thermodynamical parameters such as temperature and humidity (MEYERS et al., 1992) and are, therefore, not able to consider whether ~~IN-INP~~ were already activated at previous time steps in the model.

However, despite its important contribution, ice nucleation does not determine the entire microphysics of mixed-phase clouds alone. It is rather the complex transfer between the three phases of water: water vapor, liquid water and ice described by the Wegener-Bergeron-Findeisen (WBF) mechanism (Wegener, 1911; Bergeron, 1935; Findeisen, 1938).

It is well-known that due to the different saturation pressures of water vapor with respect to liquid water and ice, a mixed-phase is in a non-equilibrium state which, nevertheless, may lead to a quasi-steady existence (Korolev and Field, 2008). The main drivers for this phase transfer are vertical velocity (leading to supersaturation and subsequent droplet formation) and ice particle formation and growth (WBF starts) leading to sedimentation of the typically fast growing ice particles (WBF ends due to removal of ice). The motivation of this work is to shed more light on the relative contributions of the different processes involved in these complex interactions.

The paper is structured as follows. Section 2 describes the remote sensing observations of two mixed-phase altocumulus cloud cases above Leipzig. The dynamical model as well as the process descriptions and initial data used for this study are specified in section 3. Section 4 refers to changes in the dynamic parameters of the model to identify base cases ~~–The results for the which describe the observations sufficiently well to perform~~ sensitivity studies with respect to ~~the microphysical parameters~~ microphysical parameters. The results for those sensitivity studies are presented in section 5 and section 6 closes with a discussion of the results.

## 2 Remote sensing observations

Altocumulus and altostratus clouds are regularly observed with the Leipzig Aerosol and Cloud Remote Observations

System (LACROS) at the [Leibniz Institute for Tropospheric Research TROPOS](#). LACROS combines the capabilities of Raman/depolarization lidar (Althausen et al., 2009), a MIRA-35 cloud radar (Bauer-Pfundstein and Görsdorf, 2007), a Doppler lidar (Bühl et al., 2012), a microwave radiometer, a sun-photometer and a disdrometer to measure height-resolved properties of aerosols and clouds. The Cloudnet framework (Illingworth et al., 2007) is used to derive microphysical parameters like liquid-water content (Pospichal et al., 2012) or ice-water content (Hogan et al., 2006). The following two cases have been selected to illustrate this variety and to serve as examples to be compared to model results.

### 2.1 Case 1: Warm mixed-phase cloud

One of the warmest mixed-phase clouds within the data set was observed on 17 September 2011 between 00:00 and 00:22 UTC (see Fig. ?? 1). The liquid part of the cloud extends from about 4250 m to 4450 m height at temperatures of about  $-6^{\circ}\text{C}$  according to the GDAS ([Global Data Assimilation System](#)) reanalysis data for Leipzig. ~~Liquid water~~–Liquid water content (LWC) is between  $0.1\text{ g/m}^{-3}$  to  $1\text{ g/m}^{-3}$  whereas ice water content (IWC) is about 3–4 orders of magnitude smaller and reaches its maximum value within the virgae (see Fig. 2). Liquid water path (LWP) measured by a microwave radiometer varies between 20 and  $50\text{ g/m}^2$  (mostly about  $25\text{ g/m}^2$ ) whereas ice water path (IWP) is only slightly above the detection limit of about  $0.01\text{ g/m}^2$  implying a rather large uncertainty with correspondingly large error bars. Virgae (falling ice) are observed down to about 3000 m, which is close to the  $0^{\circ}\text{C}$  level. This is supported by Fig. 3 where the cloud radar (right panel) mainly shows particles falling from the top layer. Therefore, particles are mainly moving downwards (green color) and can be identified as ice particles by their size. Only at the very top at about 4300 m particles are small enough to still be lifted upwards (yellow colors). The Doppler lidar (left panel), however, shows the motion of small cloud droplets at the predominantly liquid cloud top. Hence, in this plot the cloud-top turbulence becomes visible. Vertical windspeeds range from about  $-1.5\text{ m/s}$  to  $1.0\text{ m/s}$  with pdf maxima at  $-0.5\text{ m/s}$  and  $0.5\text{ m/s}$ , respectively (Fig. ?? 3).

### 2.2 Case 2: Colder mixed-phase cloud

A much colder case ~~could be~~ was observed on 2 August 2012 between 21:00 UTC and 21:40 UTC (see Fig. ?? 4). Liquid water was measured around 7500 m at about  $-25^{\circ}\text{C}$  with ~~an~~–a LWP between 10 and  $30\text{ g/m}^2$  ~~–and a LWC of up to~~  $0.1\text{ g/m}^{-3}$  which is much smaller than for case 1. As can be expected due to the lower temperature, the ice phase was much more massive than in case 1 and reached down to about 5500 m with an IWP of about  $1\text{--}10\text{ g/m}^2$  ~~–and an IWC of up to~~  $0.01\text{ g/m}^{-3}$  which means that in some parts of the cloud,

ice and liquid water reach the same order of magnitude (see Fig. 5). Vertical wind speeds were in the same range as for the warmer case described above (Fig. 6).

Accuracy of the IWC is  $\pm 50\%$ . For the LWC calculated by the scaled adiabatic approach the same order of magnitude applies. Vertical wind speeds are measured directly by evaluation from the recorded cloud radar and Doppler lidar spectra. Errors are  $\pm 0.15$  m/s for the cloud radar and  $\pm 0.05$  m/s for the Doppler lidar. These errors are mainly due to the pointing accuracy of the two systems.

### 3 Model Description and initialization

For the model studies an Asai-Kasahara type model is used (Asai and Kasahara, 1967). The model geometry is axisymmetric and consists of an inner and an outer cylinder with radii of 100 m and 1000 m, respectively, resulting in a radius ratio of 1:10 which is typical for this setup. However, the geometric configuration of the model is not intended to match the geometry of the clouds (and the cloud-free spaces between the clouds) but is rather meant to provide the possibility of horizontal exchange between clouds and a cloud-free background.

The vertical resolution is constant with height and is chosen to be  $\Delta z = 25$  m to give a sufficient resolution of the cloud layer and to roughly match the vertical resolution of the observations. In contrast to a parcel model, the vertically resolved model grid allows for a description of hydrometeor sedimentation. This is important especially for the fast growing ice crystals to realistically describe their interaction with the vapor and liquid phase (Bergeron-Findeisen Wegener-Bergeron-Findeisen process). A time step of 1 s was used for the dynamics as well as for the microphysics.

However, in contrast to other Asai-Kasahara model studies, updrafts are not initialized by a heat and/or humidity pulse in certain layers for a given period of time. Instead, vertical velocity (updrafts and downdrafts) in the inner cylinder is prescribed, which is more similar to a kinematic model like the Kinematic Driver Model KiD (Shipway and Hill, 2012). In that way dynamics can be controlled to make sure that it is close to the observations.

The cloud microphysics is described by the mixed-phase spectral microphysics module SPECS (Simmel and Wurzler, 2006; Diehl et al., 2006). SPECS provides a joint spectrum for the liquid phase (soluble wetted aerosol particles as well as cloud and rain drops) and one spectrum for the ice phase.

For this case study, collision processes between ice particles and drops (riming) and between ice particles and ice particles (accretion) are not taken into account. One the one hand, this is to exclude further uncertainties which would be introduced by the collision/collection kernel for those interactions, one the other hand, only small or neglectable effects are expected. Clouds are shallow which means that

there is not much time for the ice particles to interact with droplets (especially when the ice is preferentially formed near cloud base and sediments out soon). In addition, for case 1 ice particle concentrations are low which highly limits the probability of collisions. At the low temperatures of case 2 sticking efficiency is expected to be low. This assumption is corroborated by the findings of Smith et al. (2009) stating that water vapor deposition (and sublimation), balanced by sedimentation are more important than accretional growth.

### 3.1 Description of ice microphysics

In the following, the differences in the description of the microphysics compared to Diehl et al. (2006) are described.

#### 3.1.1 Immersion freezing

For this study, immersion freezing is assumed to be the only primary ice formation process. Since during the above mentioned observations no *in situ* measurements of the  $\text{IN-INP}$  are available, the parameterization of DeMott et al. (2010) is used assuming that all  $\text{IN-INP}$  are active in the immersion freezing mode. The parameterization of DeMott et al. (2010) is based on an empirical relation of  $\text{IN-INP}$  and the number of aerosol particles with radii  $> 250$  nm ( $N_{AP,r>250nm}$ ). To cover case 1, the parameterization is extrapolated to  $-5^\circ\text{C}$  despite the fact that the underlying measurements were only taken at  $-9^\circ\text{C}$  and below. As base case  $N_{AP,r>250nm} = 10^5 \text{ kg}^{-1}$  air is used as input data for the parameterization resulting in about 0.01 active  $\text{IN-INP}$  per liter for  $-6^\circ\text{C}$  (case 1) and about 0.5  $\text{IN-INP}$  per liter for  $-25^\circ\text{C}$  (case 2), respectively, at standard conditions. This corresponds to a relatively low number of larger aerosol particles but is well within the range observed by DeMott et al. (2010).

For the potential  $\text{IN-INP}$  a prognostic temperature resolved field with 20 temperature bins with a resolution of 1 K is introduced into SPECS. It ranges from  $-5^\circ\text{C}$  to  $-25^\circ\text{C}$  to cover the temperature range for the selected cases and can easily be changed for other case studies. This is a simplified version of the method used by Fridlind et al. (2007). The potential  $\text{IN-INP}$  field is initially defined in every grid cell (layer) and is transported vertically with the given up/downdrafts and horizontally exchanged between inner and outer cylinder in the same way as the other hydrometeor fields (drops and ice crystals). Immersion freezing occurs as soon as liquid drops above a certain size limit are present and the temperature of a certain potential  $\text{IN-INP}$  bin is reached. Then the respective amount of drops freezes (if available) instantaneously and is transferred from the liquid to the frozen spectrum. If more drops larger than the size threshold of  $10 \mu\text{m}$  than active  $\text{IN-INP}$  are present, the  $\text{IN-INP}$  are distributed evenly over all drop size bins above the threshold value. The drop size threshold was chosen to restrict freezing to droplets and to prevent (large) non-activated aerosol particles at high relative humidity (but subsaturated

275 wrt water) outside the cloud from freezing. If ice crystals 325  
melt below the freezing level, they contribute to the poten-  
tial ~~IN-INP~~ field at that level.

### 3.1.2 Ice particle shape

280 It is well known that ice particle shape highly influences wa-  
ter vapor deposition (described by changing the capacitance  
of the particle) as well as terminal fall velocity of the ice  
particle. Therefore, instead of the previously chosen spher-  
ical ice particle shape, ice particles now can be prescribed  
285 as hexagonal columns or plates. The aspect ratio can be ei-  
ther constant for all size bins or be changed with size fol-  
lowing the approach of Mitchell (1996). Typically, with in-  
creasing particle size, the deviation from an uniform aspect 335  
ratio increases. In our simulations, a constant uniform aspect  
ratio ( $ar=1$ ) is used as base case. From Mitchell (1996) the  
size-varying aspect ratios for plates ~~and columns are (ranging~~  
290 ~~from 15 to 3000  $\mu\text{m}$  with a single description) and columns~~  
~~(for size ranges of 30 to 100  $\mu\text{m}$ , 100 to 300  $\mu\text{m}$ , and above~~  
~~300  $\mu\text{m}$  in diameter) are calculated from the mass-dimension 340~~  
~~power laws and used for sensitivity studies.~~

295 The (relative) capacitance needed for the calculation of  
deposition growth of the ice crystals is modeled using the  
method of Westbrook et al. (2008) for the aspect ratios given  
above. Ice crystal terminal fall velocities are calculated ac- 345  
cording to Heymsfield and Westbrook (2010) using the same  
300 aspect ratios.

## 3.2 Model initialization

### 3.2.1 Thermodynamics

The Asai-Kasahara model has to be initialized with vertical  
profiles of temperature and dewpoint temperature either from  
305 reanalysis data (here GDAS) or radiosonde profiles from  
nearby stations (here Meiningen, Thuringia). Fig. 7 shows  
profiles of temperature and relative humidities with respect  
to liquid water and to ice, respectively, for both cases. For  
case 1, profiles from both methods show a similar general  
behaviour but the radiosonde profile of Meiningen measured  
310 at 00 UTC is used since it provides a finer vertical resolution  
than the GDAS reanalysis data ~~(cp. Fig. 7).~~ However, for 360  
case 2 the Meiningen RS profile misses the humidity layer  
at the level where the clouds were observed ~~and, therefore,~~  
315 ~~This means that the profile is not representative for the given~~  
~~meteorological situation. Therefore,~~ GDAS reanalysis data  
for Leipzig at 21 UTC were chosen. Finally, both profiles 365  
used show a sufficiently humid layer where the clouds were  
observed, so that lifting of these layers lead to supersatura-  
320 tion and subsequent cloud formation.

~~In contrast to other Asai-Kasahara model studies, updrafts~~  
~~are not initialized by a heat and/or humidity pulse in certain~~  
~~layers for a given period of time. Instead, As mentioned 370~~  
~~above,~~ vertical velocity (updrafts and downdrafts) in the in-

ner cylinder is prescribed at cloud level ranging from  $h_{bot}$  to  
 $h_{top}$ . The center of this interval is given by  $h_{mid} = (h_{top} +$   
 $h_{bot})/2$  and its half-depth by  $h_{depth} = (h_{top} - h_{bot})/2$ .  $h_{bot}$   
ranges from 3800 m to 4100 m for case 1 and from 7000 m to  
7300 m for case 2. The respective values for  $h_{top}$  are 4500 m  
and 7700 m. The vertical dependency (compare Fig. 8, left)  
is given by

$$f_h(h) = \frac{h_{depth}^2 - (h - h_{mid})^2}{h_{depth}^2} \quad \text{for } h_{bot} \leq h \leq h_{top} \quad (1)$$

resulting in the time- and height-dependent function

$$w(h, t) = w_{mid}(t) f_h(h) \quad \text{for } h_{bot} \leq h \leq h_{top} \quad (2)$$

and  $w(h, t) = 0$  otherwise, defining  $w_{mid}(t)$  as the updraft  
velocity at  $h_{mid}$ . In order to match the observed wind field  
distributions rather closely,  $w_{mid}(t)$  is chosen as a stochastic  
function

$$w_{mid}(t) = w_{ave} + f_{scal} \frac{\delta(t)^3}{|\delta(t)|} \quad (3)$$

where  $w_{ave}$  is the average ('large-scale') updraft velocity at  
 $h_{mid}$  varying between 0.1 m/s and 0.4 m/s,  $f_{scal}$  is the scal-  
ing factor determining the range of updraft velocities (chosen  
as 4 m/s to obtain a difference of minimum and maximum ve-  
locity of 2 m/s), and  $\delta(t)$  is a random number ranging from  
-0.5 to +0.5 obtained from a linear stochastic process pro-  
vided by FORTRAN. After 30 s model time a new  $\delta(t)$  is  
created. Different realizations of the stochastic process are  
tested (see below). E.g.,  $w_{mid}(t)$  ranges from -0.7 m/s to  
1.3 m/s if  $w_{ave} = 0.3$  m/s and  $f_{scal} = 4$  m/s as it is shown in  
350 the temporal evolution and the histogram in ~~Figure Fig. 8.~~

Due to the height dependent vertical velocity  $w$ , a horizon-  
tal transport velocity  $u_k$  (exchange between inner and outer  
cylinder) is induced in the Asai-Kasahara formulation for a  
given model layer  $k$ .

$$u_k = - \frac{w_{k+\frac{1}{2}} \rho_{k+\frac{1}{2}} - w_{k-\frac{1}{2}} \rho_{k-\frac{1}{2}}}{f_r \Delta z \rho_k} \quad (4)$$

Full indices  $k$  indicate values at level centers whereas  
half indices ( $k + \frac{1}{2}$ ,  $k - \frac{1}{2}$ ) describe values at level inter-  
faces.  $f_r = 2/r_i$  is a geometry parameter with the radius  
 ~~$r_i = 1000$~~   $r_i = 100$  m of the inner cylinder.

The prescribed velocity field leads to the following effects  
(all descriptions are related to the inner cylinder if not stated  
otherwise explicitly):

- In the updraft phase: In the upper part (between  $h_{mid}$   
and  $h_{top}$ ) of the updraft, mixing occurs from the inner  
to the outer cylinder whereas in the lower part (between  
 $h_{low}$  and  $h_{mid}$ ), horizontal transport is from the outer  
cylinder into the inner one
- For downdrafts it is the other way: This means that be-  
low  $h_{mid}$  drops and ice particles are transported from  
the inner cylinder to the outer one and are therefore re-  
moved from the inner cylinder

- below  $h_{low}$  or above  $h_{top}$ , no horizontal exchange takes place.

The question arises to which extent this dynamical behaviour reflects the real features of the observed clouds and whether this is critical for the topics aimed at in this study.

Prescribing vertical velocity in any way also means that a feedback of microphysics on dynamics due to phase changes (e.g., release of latent heat for condensing water vapor or freezing/melting processes) is not considered by the model.

### 3.2.2 Aerosol distribution

Since no in situ aerosol measurements are available, literature data is used. The Raman lidar observations do not show any polluted layers for both cases; therefore data from LACE98 (Petzold et al., 2002) are used which should be representative for the free troposphere over Leipzig. For case 1 values for the lower free troposphere (M6), for case 2 those from the upper free troposphere (M1) are used (see Petzold et al., 2002, Tab. 6) (see Petzold et al., 2002, Tab. 6).

## 4 Model results: Dynamics

In a first step, the aim is to achieve a sufficient agreement concerning macroscopic cloud features as well as (liquid phase) microphysics as far as they were observed. The following parameters describing model dynamics (updraft velocity) are varied to identify a ‘best case’ which in the second step can be used to perform sensitivity studies with respect to (ice) microphysics (see also Tables 1 and 2).

- $h_{low}$ : ranging from 3800 m to 4100 m for the warmer and from 7000 m to 7300 m for the colder case. This parameter influences the vertical cloud extent and, therefore, liquid water content and liquid water path.
- $w_{ave}$ : ranging from 0.1 m/s to 0.4 m/s. Higher average updraft also leads to higher LWC. Due to the lateral mixing processes the model setup requires a positive updraft velocity in average to form and maintain clouds.
- $\delta$ : Four different realizations of the stochastic process are used. This influences the timing of the cloud occurrence as well as LWC and LWP but not systematically.

All model results shown refer to the inner cylinder.

### 4.1 Case 1: Warm mixed-phase cloud

Figs. 9 and 11 show time-height plots of the liquid (contours, linear scale) and ice (colours, logarithmic scale) water mixing-ratio content for case 1 illustrating the cloud sensitivity with respect to variation of cloud base ( $h_{bot}$ ), average vertical updraft ( $w_{ave}$ ), and the realization of the stochastic

process, respectively. Liquid clouds form in the updraft regions (cp. Fig. 8) whereas in the downdrafts the liquid phase vanishes at least partly. ~~Ice forms mainly at cloudbase and immediately starts.~~ If active INP are available ice formation can take place within the liquid part of the cloud. The INP are partly already active near liquid cloud base which means that they trigger freezing as soon as the droplets are formed. Less efficient INP become active after further cooling above cloud base. After ice formation rapid depositional growth takes place and the ice particles almost immediately start to sediment. Due to the supersaturation with respect to ice even below liquid cloud base, ice particles still grow while sedimenting, reaching their maximum size before, finally, subsaturated regions are reached and ~~evaporation~~ sublimation sets in. Figs. 10 and 12 show the time evolution of liquid (lower panel) and ice water path (upper panel) for the same parameters varied, reflecting the same temporal patterns. ~~Table Tab. 1~~ summarizes the maximum values for liquid and ice water ~~mixing-ratio~~ (LWMCcontent (LWC/IWMC)), liquid and ice water path (LWP/IWP) as well as cloud droplet and ice particle number concentration (CDN/IPN) for all dynamics sensitivity runs for case 1.

One can clearly observe, that a lower  $h_{bot}$  (Fig. 9) results in a lower cloud base, larger vertical cloud extent as well as more liquid water. The ~~same trend with similar intensity~~ LWC maxima are within a factor of 2 for varying  $h_{bot}$ . A similar trend is observed for the ice phase (see also Fig. 10) 10, but IWC maxima differ only by about 25 %. However, the values of the two maxima of the condensed phase after about 15–20 min and about 40 min model time are quite different. The first maximum is more pronounced for the ice phase whereas the second one is larger for the liquid phase. While the liquid phase is dominated by the updraft velocity (see Fig. 8) the ice phase additionally depends on ~~IN~~ INP supply. In the first ice formation event at 15 min, all ~~IN~~ INP active at the current temperature actually form ice leading to an ~~IN-IPN~~ INP depletion. Due to the horizontal exchange with the outer cylinder the ~~IN-IPN~~ INP reservoir is refilled, but only to a certain extent when the second cloud event after 40 min sets in. Due to the limited ~~IN-IPN~~ INP supply the second ice maximum is weaker than the first one. The stochastic velocity fluctuations cause fluctuations in relative humidity, which are directly reflected by the liquid phase parameters whereas the ice phase generally reacts much slower. Sensitivity of CDN and IPN with respect to change of  $h_{bot}$  does not seem to be systematic.

Increasing the average updraft velocity  $w_{ave}$  leads to a similar increase of liquid water and ice as lowering  $h_{bot}$  (see Figs. 11, upper panel and 12, left). This can be expected since more water vapor flows through the cloud and is able to condense. However, a certain limit seems to be reached for W\_w04, since the increase of LWP slows down (see maximum value at 40 min in Fig. 12, left). This is due to the enhanced horizontal exchanged following eq. (4). Additionally, the stronger updrafts allow the ice particles a longer presence



time in the vicinity of the cloud and, therefore, an enhanced growth at comparably high supersaturation with respect to ice before sedimentation sets in at larger sizes. This also leads to an accumulation of ice particles and, therefore, to a higher IPN. Surprisingly, CDN ~~only depends~~ depends only weakly and not systematically on  $w_{ave}$  which is in contrast to the typical enhancement of CDN with increasing updraft velocities.

Figs. 11 (lower panel) and 12 (right) show that different realizations of the stochastic process (as explained above in section 3.2.1) lead to different temporal cloud evolutions. However, differences in maximum LWP and ~~LWLR-LWC~~ LWLR-LWC are much smaller than those discussed above. Variations in maximum IWP and ~~IWLR-IWC~~ IWLR-IWC as well as CDN and IPN are in the range of about 30%. This is also true for average LWP ranging from 18 g/m<sup>2</sup> for W\_r1 to 26 g/m<sup>2</sup> for W\_r3. However, despite the different maxima and temporal evolutions of IWP, average IWP is almost identical for the different stochastic realizations (0.023 g/m<sup>2</sup>). This shows that changing the stochastic realization influences cloud evolution in detail (timing) but does not change the overall picture.

With maximum values between 17 and 57 g/m<sup>2</sup> the modeled liquid water path is in the same range as the observed values (20–50 g/m<sup>2</sup>), especially for the 'wetter' runs (smaller  $h_{bot}$ , larger  $w_{ave}$ ). Average LWP typically is about half (40–60%) of the maximum value for most of the runs which also fits well into the observations. Ice forms within the liquid layer and sediments to about 3800 m for most runs which is less than for the observations. The (maximum) modeled ice mixing ratio is in the same order of magnitude as the observed one (about 10<sup>-7</sup> kg/m<sup>3</sup>). The same holds for the ice water path with values of about 0.01 g/m<sup>2</sup> for both, model and observation. For the other values, no observational data is available for comparison.

## 4.2 Case 2: Colder mixed-phase cloud

Due to the ~~colder-lower~~ colder-lower temperatures of case 2 much more ~~INP are active and much more~~ ice is produced than in case 1 (see Figs. 13–16 as well as Tab. 2). ~~Therefore This~~ also means that near cloud base much more active INP are available and that a further cooling within the clouds only slightly increases the number of active INP leading again to a preferential ice nucleation near liquid cloud base. Due to the lower temperatures and the more massive ice formation, the virgae reach down to more than 1500 m below liquid cloud base which is in concordance with the observations. The principal behaviour with respect to the sensitivity parameters is similar as in case 1: The liquid phase is enhanced by either decreasing  $h_{bot}$  or increasing  $w_{ave}$ , showing the 'saturation' effect slightly more pronounced as in case 1. Different stochastic realizations only weakly influence the maximum and average values of the liquid phase but change the timing of occurrence. Generally, the variability of the ice phase is weaker than in case 1. The different stochastic realizations

show the highest variability in ~~IWLR-IWC~~ IWLR-IWC and IWP. Different variations of  $h_{bot}$  show almost identical IWPs, whereas changing  $w_{ave}$  at least slightly influences maximum ~~IWLR-IWC~~ IWLR-IWC and IWP, which again can be attributed to the ice particle accumulation in the updraft. Liquid water path is smaller than in case 1 and reaches maximum values between 10 and 43 g/m<sup>2</sup> which well covers the observed maximum value of about 20 g/m<sup>2</sup>. Cloudnet observations show an IWC of 10<sup>-7</sup> – 10<sup>-5</sup> kg/m<sup>3</sup> which is an increase by a factor of 10–100 compared to case 1. Similar values are obtained by the model results underlining the strong temperature dependency of the ice nucleation process.

## 5 Sensitivity studies

In the previous section it could be shown that dynamical parameters can be chosen in a way that the model results (in terms of LWP, IWP as well as cloud geometry) are in good agreement with the observations. This allows to perform sensitivity studies with respect to cloud microphysics. To cover the proper sensitivities we have to answer the question which microphysical parameters are expected to have a large influence on mixed phase microphysics and are rather uncertain to be estimated. This leads to (temperature-dependent) ~~IN~~ number ( $N_{INP}$ ) ~~number ( $N_{INP}$ )~~ number ( $N_{INP}$ ) which directly influences ice particle number but mostly is poorly known. To be consistent with the freezing parameterization of the model,  ~~$N_{INP}$~~   $N_{INP}$  is varied by changing  $N_{AP,r>250nm}$  which additionally is easier to observe in most cases. A second parameter is the shape of the ice particles which does not influence the primary freezing process but the subsequent growth by water vapor deposition onto existing ice particles and, therefore, the total ice mass produced. Their relative importance shall be quantified and also be compared to the influence of dynamics discussed above.

### 5.1 ~~IN-INP~~ number

Changing  $N_{AP,r>250nm}$  leads to a temperature-dependent change of ~~IN-INP~~ number which is relatively small for warmer conditions. However, the effect increases with decreasing temperature. This is illustrated by the following numbers. The parameterization of DeMott et al. (2010) gives about 0.009 active ~~IN-INP~~ number per liter at standard conditions ( ~~$N_{INP}$~~   $N_{INP}$ ) when  $N_{AP,r>250nm} = 10^5$  kg<sup>-1</sup> at  $T = -5^\circ\text{C}$ . A tenfold increase to  $N_{AP,r>250nm} = 10^6$  kg<sup>-1</sup> results in about 0.012 active ~~IN-INP~~ number per liter which is a rise of only about 35%. For  $T = -7^\circ\text{C}$ , ~~IN-INP~~ number rises by about 65% for a tenfold increase of  $N_{AP,r>250nm}$ . This shows that for those rather ~~warm-high~~ temperatures considered for case 1, a massive change in  $N_{AP,r>250nm}$  leads to relatively small changes in  ~~$N_{INP}$~~   $N_{INP}$  and only a small effect on the ice phase can be expected. This is confirmed by Fig. 17 (left) showing liquid and ice water ~~mixing ratios~~ mixing ratios

575 [contents](#) for W\_in6. Ice-mass-IWC is enhanced by less than 60 % for W\_in6 and by about 160 % for W\_in7 which is consistent for the given temperature range (see Tab. 3). Similar values are obtained for the change in IPN. This directly leads to the conclusion that the individual ice particles grow independently from each other. Their individual growth history is (in contrast to drop growth) only influenced by thermodynamics as long as their number is low enough which seems to be the case here.

580 This is confirmed by Fig. 18 showing drop and ice particle size distributions at the time when the maximum IWP is reached (16 min for case 1, 17 min for case 2). For case 1 (upper panel), the liquid phase (contours) is unaffected by the IN-INP enhancement. Despite the increase of ice particle number and mass the shape of the ice particle size distribution (colors) is not changed. The smallest ice particles can be observed at three discrete height (and temperature) levels caused by the temperature resolved parameterization of the potential IN-INP described in section 3.1.1. In reality this part of the spectrum showing rather freshly nucleated and fast growing ice particles should be continuous over the height range from about 4100 m to 4400 m. Nevertheless, the total number of ice particles formed is described correctly.

585 One can conclude that increasing IN-INP number therefore increases ice particle number as well as ice mass proportionally. Generally, the ice mass remains small and the liquid phase is not affected by the ice mass increase. Those results are supported by Fig. 19 (left) showing an unchanged LWP and a proportionally growing IWP for increased IN-INP numbers.

590 For the colder case 2 the parameters are varied in the same way. However, one big difference is that a tenfold increase of  $N_{AP,r>250nm}$  at  $T = -25^\circ\text{C}$  results in a much larger change in active IN-INP. Their number rises by 300 % from about 0.5 per liter to about 2 per liter following the parameterization. This is reflected by the IPN values in [Table Tab. 4](#). Fig. 17 (right) and [Table Tab. 4](#) show that ice mass increases in such a way that liquid water is depleted partially (C\_in6 by about 50 %) or almost totally (C\_in7) due to the Bergeron-Findeisen-Wegener-Bergeron-Findeisen process. Compared to C\_base, ice is enhanced by a factor of 3–4 for C\_in6 and about 10 for C\_in7 whereas IPN increases by a factor of 12. This can also be seen in the IWP (Fig. 19, right, red lines) showing a limited increase for C\_in7, especially for the first maximum after 17 min. This means that the results for C\_in6 are still consistent with an independent growth of the individual ice particles (as described above) despite the relatively high ice occurrence.

605 This is verified by the size distributions in Fig. 18 (lower panel). As in case 1 the ice particle size distributions only differ by the number/mass, but not by shape. Additionally, the decrease in the liquid phase is reflected also in the drop spectrum showing a more shallow liquid part of cloud as well as droplet distribution shifted to smaller sizes.

However, for C\_in7 the ice particles compete for water vapor which becomes clear from (i) the depletion of liquid water (resulting in a lower supersaturation with respect to ice) and (ii) the ice mass enhancement factor being below the value expected from the ice nucleation parameterization and below that of IPN. This means that despite the higher number of IN-INP and, therefore, ice particles, the amount of ice is limited by the thermodynamic conditions which results in the production of more but smaller ice particles, similar to the Twomey effect for drop activation.

As mentioned earlier, ice particle growth is not restricted to the liquid part of the cloud but also occurs in the layer below liquid cloud base, which is still supersaturated with respect to ice. This leads to a decrease in relative humidity in this part of the cloud, which in turn weakens or suppresses droplet formation by shifting liquid cloud base to higher altitudes. The lower LWC for the runs with higher IWC therefore cannot only be attributed to the WBF processes but also to this indirect effect.

## 5.2 Ice particle shape

As discussed previously, for most of the cases (except for C\_in7) changing the parameters in the section above does neither influence the ice particles themselves nor their individual growth. Additionally, due to their low number, there is almost no competition of the ice particles for water vapor, and, therefore, ice water content scales linearly with ice particle number. In contrast to this, changing the ice particle shape from quasi-spherical ( $ar=1$ ) to columns or plates with size-dependent axis ratios deviating from unity results in an increase of water vapor deposition on the individual ice particles leading to enhanced ice water content due to larger individual particles when ice particle numbers remain unchanged. This is due to (i) enhanced relative capacitance resulting in faster water vapor deposition and (ii) lower terminal velocities of the ice particles leading to longer residence times in vicinity of conditions with supersaturation with respect to ice.

610 Fig. 20 (left) shows the results for the runs using hexagonal columns (W\_col) as prescribed ice particle type. Compared to the previous results (W\_base, W\_in6, W\_in7) more ice mass is produced (see [Table Tab. 3](#)) but still the liquid part of the cloud remains unaffected (compare also LWP and IWP in Fig. 19, left). Similar results are obtained for the assumption of plate-like ice particles (W\_pla). The mass increase results from the larger ice particle size due to the reasons discussed above which can be seen from Fig. 21 showing the size distributions for W\_col at different times. On the upper left panel W\_col is shown after 16 min corresponding to Fig. 18. Compared to W\_base, larger ice particles are produced leading to more ice mass (equivalent radius up to  $300\ \mu\text{m}$  compared to  $189\text{--}238\ \mu\text{m}$  for the base case). Additionally, due to the lower fall speed of the columns ( $1.03\ \text{m/s}$  vs.  $1.75\text{--}2.24\ \text{m/s}$ ), the maximum of the ice is at about 4200 m compared to 4100 m

for the base case. On the upper right panel, size distributions after 21 min are shown corresponding to the IWP maximum of W\_col. Ice particles have grown larger (equivalent radius up to  $378 \mu\text{m}$ , length of the columns increases from about 3 mm to 4.5 mm) and sedimentation has developed further with increasing terminal velocity (1.13 m/s). Similar results are obtained for plates (W\_pla) with terminal velocities of 0.89–1.21 m/s, equivalent radii of 300–476  $\mu\text{m}$  and maximum dimension of 1.8–3.2 mm.

The lower terminal velocity of columns and plates despite their larger size is leading to the stronger tilting of the virgae. Additionally, ice particle number IPN is enhanced by about 30% although ice nucleation is identical to W\_base. This can be attributed to the lower fall velocities, too, leading to an accumulation of ice particles. The differences between W\_col and W\_pla are caused by both, the higher relative capacitances of and lower terminal fall velocities of plates compared to columns (at least when their axis ratios are chosen following Mitchell (1996)).

For case 2 (C\_col and C\_pla), the liquid water reduction due to the Bergeron-Findeisen process is similar to C\_in6 (see Fig. 20, right, and Table Tab. 4). In contrast to the respective case 1 runs, less ice is produced than for C\_in7. The tilting of the virgae is not as strong as in W\_col which is due to the larger ice particle sizes leading to higher terminal fall velocities (1.43–1.60 m/s). Additionally, the lower air density leads to an increase of terminal velocity of more than 10% independently from shape. Fig. 21 (lower panels) show panel shows the size distributions for C\_col at different times. Due to the longer growth time larger individual ice particles than in case 1 are produced (equivalent radius up to 600  $\mu\text{m}$  compared to 300  $\mu\text{m}$  for the base case).

To decide whether independent ice particle growth or competition occurs, further runs with less IN-INP (C\_col\_in4 and C\_pla\_in4) are discussed (see Fig. 19, right). IWMR-IWC and IWP of these runs (in4) are about one third of the values of the respective runs with more IN-INP (in5). For ice particle number, a factor of slightly more than three occurs which means that a weak competition for water vapor occurs for C\_col and C\_pla resulting in slightly smaller individual ice particles compared to C\_col\_in4 and C\_pla\_in4.

## 6 Conclusions

The model system AK-SPECS was applied to simulate dynamical and microphysical processes within altocumulus clouds. Sensitivity studies on relative contributions on cloud evolution as well as comparisons to observations were made.

Variation of the dynamic parameters as it was done in section 4 leads to systematic differences mainly in the liquid phase (LWMRLWC, LWP) which can easily be explained. More liquid water is produced when either cloud base is lowered (corresponding to a larger vertical cloud extent) or vertical wind velocity is increased. However, the effects of

the dynamics on the ice phase are surprisingly small, at least smaller than those on the liquid phase. Increasing vertical velocity leads to an accumulation of the smaller ice particles in the enhanced updraft.

On the other hand, much larger differences in terms of IWMR-IWC and IWP were found when microphysical parameters like IN-INP number or ice particle shape were varied under identical dynamic conditions. This is valid for both cases studied. However, at least for the ice nucleation parameterization used, sensitivity of IN-INP number strongly increased with decreasing temperature.

This means that relatively large differences concerning the ice phase can only be reached when either IN-INP number differs considerably or ice particle shape is different (which should not be the case for relatively similar thermodynamical conditions). After Fukuta and Takahashi (1999) for case 1 with temperatures of about  $-6^\circ\text{C}$  column-like ice particles with  $ar = 0.1$  could be expected (corresponding to W\_col) whereas for case 2 ( $T < -24^\circ\text{C}$ ) hexagonal particles with  $ar = 1$  are most likely (e.g., C\_base). Those ice shapes were observed in laboratory studies at water saturation which was also valid for the observed cases when ice formed by immersion freezing within the liquid layer of the cloud. However, below liquid cloud base supersaturation with respect to ice decreases. These ice shapes can also explain why a depletion of the liquid phase was not observed in case 2 as it was predicted by the sensitivity studies using either columns or plates as prescribed shape. Generally, the liquid phase is affected considerably only when enough ice particles are present which typically is the case for cold conditions with a sufficient amount of IN-INP and fast growing ice particle shapes (most effective for large deviations from spherical shapes).

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**Table 1.** Overview of the model results for the dynamic sensitivity runs for the warmer case 1 (maximum values of L/FWMR/IWC: liquid/ice water mixing-ratio/content, L/IWP: liquid/ice water path, CDN: cloud drop number, IPN: Ice particle number).

run	parameter value differing from base case	<u>LW</u> <u>MR</u> - <u>LWC</u> g/kg·m <sup>-3</sup>	<u>IW</u> <u>MR</u> - <u>IWC</u> 10 <sup>-3</sup> g/kg·m <sup>-3</sup>	LWP g/m <sup>2</sup>	IWP 10 <sup>-3</sup> g/m <sup>2</sup>	CDN cm <sup>-3</sup>	IPN l <sup>-1</sup>
W_base	—	<u>0.277</u> - <u>0.355</u>	<u>0.304</u> - <u>0.379</u>	41.33	62.27	46.89	0.0197
W_h38	$h_{bot} = 3800$ m	<u>0.332</u> - <u>0.426</u>	<u>0.329</u> - <u>0.408</u>	57.05	73.11	48.63	0.0235
W_h40	$h_{bot} = 4000$ m	<u>0.225</u> - <u>0.289</u>	<u>0.286</u> - <u>0.357</u>	28.58	58.12	61.48	0.0240
W_h41	$h_{bot} = 4100$ m	<u>0.170</u> - <u>0.219</u>	<u>0.259</u> - <u>0.324</u>	18.23	45.81	59.53	0.0208
W_w01	$w_{ave} = 0.1$ m/s	<u>0.147</u> - <u>0.187</u>	<u>0.160</u> - <u>0.200</u>	17.41	31.73	43.36	0.0138
W_w02	$w_{ave} = 0.2$ m/s	<u>0.232</u> - <u>0.297</u>	<u>0.241</u> - <u>0.300</u>	32.86	47.18	54.57	0.0175
W_w04	$w_{ave} = 0.4$ m/s	<u>0.297</u> - <u>0.382</u>	<u>0.359</u> - <u>0.448</u>	44.48	78.25	52.66	0.0219
W_r1	stoch. realiz. r1	<u>0.261</u> - <u>0.336</u>	<u>0.254</u> - <u>0.316</u>	40.32	54.85	64.26	0.0163
W_r3	stoch. realiz. r3	<u>0.296</u> - <u>0.381</u>	<u>0.252</u> - <u>0.314</u>	42.88	54.48	43.03	0.0167
W_r4	stoch. realiz. r4	<u>0.269</u> - <u>0.346</u>	<u>0.197</u> - <u>0.245</u>	40.91	46.93	47.42	0.0151

**Table 2.** Overview of the model results for the dynamic sensitivity runs for the colder case 2 (maximum values of L/FWMR/IWC: liquid/ice water mixing-ratio/content, L/IWP: liquid/ice water path, CDN: cloud drop number, IPN: Ice particle number).

run	parameter value differing from base case	<u>LW</u> <u>MR</u> - <u>LWC</u> g/kg·m <sup>-3</sup>	<u>IW</u> <u>MR</u> - <u>IWC</u> g/kg·m <sup>-3</sup>	LWP g/m <sup>2</sup>	IWP g/m <sup>2</sup>	CDN cm <sup>-3</sup>	IPN l <sup>-1</sup>
C_base	—	<u>0.215</u> - <u>0.377</u>	<u>0.026</u> - <u>0.041</u>	29.35	10.71	70.56	0.462
C_h70	$h_{bot} = 7000$ m	<u>0.258</u> - <u>0.452</u>	<u>0.030</u> - <u>0.048</u>	43.06	11.34	71.33	0.432
C_h72	$h_{bot} = 7200$ m	<u>0.169</u> - <u>0.296</u>	<u>0.022</u> - <u>0.035</u>	18.71	10.11	90.51	0.396
C_h73	$h_{bot} = 7300$ m	<u>0.122</u> - <u>0.215</u>	<u>0.018</u> - <u>0.028</u>	10.54	9.27	77.61	0.337
C_w01	$w_{ave} = 0.1$ m/s	<u>0.126</u> - <u>0.219</u>	<u>0.025</u> - <u>0.040</u>	17.19	8.01	76.98	0.292
C_w02	$w_{ave} = 0.2$ m/s	<u>0.181</u> - <u>0.316</u>	<u>0.027</u> - <u>0.044</u>	25.89	9.42	74.40	0.415
C_w04	$w_{ave} = 0.4$ m/s	<u>0.229</u> - <u>0.402</u>	<u>0.028</u> - <u>0.045</u>	30.58	11.85	98.37	0.439
C_r1	stoch. realiz. r1	<u>0.209</u> - <u>0.366</u>	<u>0.014</u> - <u>0.023</u>	29.37	6.57	86.64	0.257
C_r3	stoch. realiz. r3	<u>0.228</u> - <u>0.399</u>	<u>0.029</u> - <u>0.046</u>	30.22	9.95	79.65	0.341
C_r4	stoch. realiz. r4	<u>0.213</u> - <u>0.373</u>	<u>0.031</u> - <u>0.049</u>	29.53	8.33	95.89	0.419

**Table 3.** Overview of the model results for the microphysical sensitivity runs for the warmer case 1 (maximum values of L/FWMR/IWC: liquid/ice water mixing-ratio/content, L/IWP: liquid/ice water path, CDN: cloud drop number, IPN: Ice particle number).

run	parameter value differing from base case	<u>LW</u> <u>MR</u> - <u>LWC</u> g/kg·m <sup>-3</sup>	<u>IW</u> <u>MR</u> - <u>IWC</u> 10 <sup>-3</sup> g/kg·m <sup>-3</sup>	LWP g/m <sup>2</sup>	IWP g/m <sup>2</sup>	CDN cm <sup>-3</sup>	IPN l <sup>-1</sup>
W_in6	$N_{AP,r>250nm} = 10^6$ kg <sup>-1</sup>	<u>0.276</u> - <u>0.354</u>	<u>0.496</u> - <u>0.619</u>	41.31	0.10	46.69	0.0296
W_in7	$N_{AP,r>250nm} = 10^7$ kg <sup>-1</sup>	<u>0.276</u> - <u>0.354</u>	<u>0.801</u> - <u>1.000</u>	41.24	0.17	41.61	0.0450
W_col	ice shape: columns	<u>0.275</u> - <u>0.353</u>	<u>1.467</u> - <u>1.830</u>	41.20	0.27	42.90	0.0257
W_pla	ice shape: plates	<u>0.275</u> - <u>0.353</u>	<u>2.285</u> - <u>2.850</u>	41.13	0.45	43.41	0.0267

**Fig. 1.** Lidar and radar observations on 17 September 2011 (case 1). Left: Ice-water content-IWC, right: Lidar range-corrected 1064 nm signal (in logarithmic scale, arbitrary units a. u.), right: radar reflectivity. The dashed box denotes the region for which case 1 observations are shown in the following figures.

**Table 4.** Overview of the model results for the microphysical sensitivity runs for the colder case 2 (maximum values of L/IWMP/IWC: liquid/ice water ~~mixing-ratio~~ content, L/IWP: liquid/ice water path, CDN: cloud drop number, IPN: Ice particle number).

run	parameter value differing from base case	<del>LWMP-LWC</del> g/kg-m <sup>3</sup>	<del>IWMP-IWC</del> g/kg-m <sup>3</sup>	LWP g/m <sup>2</sup>	IWP g/m <sup>2</sup>	CDN cm <sup>-3</sup>	IPN l <sup>-1</sup>
C.in6	$N_{AP,r>250nm} = 10^6 \text{ kg}^{-1}$	<del>0.128</del> <u>0.224</u>	<del>0.089</del> <u>0.140</u>	13.09	34.75	80.29	1.380
C.in7	$N_{AP,r>250nm} = 10^7 \text{ kg}^{-1}$	<del>0.021</del> <u>0.036</u>	<del>0.265</del> <u>0.446</u>	2.58	57.98	46.67	5.208
C.col	ice shape: columns	<del>0.135</del> <u>0.237</u>	<del>0.139</del> <u>0.223</u>	14.33	46.78	78.40	0.462
C.col.in4	ice shape: columns, $N_{AP,r>250nm} = 10^4 \text{ kg}^{-1}$	<del>0.216</del> <u>0.378</u>	<del>0.048</del> <u>0.076</u>	30.01	14.93	74.75	0.139
C.pla	ice shape: plates	<del>0.104</del> <u>0.182</u>	<del>0.183</del> <u>0.294</u>	9.94	57.11	39.41	0.472
C.pla.in4	ice shape: plates, $N_{AP,r>250nm} = 10^4 \text{ kg}^{-1}$	<del>0.207</del> <u>0.362</u>	<del>0.064</del> <u>0.102</u>	27.80	19.21	74.44	0.129

**Fig. 2.** ~~Vertical velocity-Cloudnet derived water contents~~ for case 1. Left: ~~derived from radar observations (valid for large particles; virgae)~~ Liquid water content, right: ice water content (both in logarithmic scale).

**Fig. 3.** Vertical velocity for case 1. Left: derived from lidar (valid for more numerous smaller droplets at cloud base), right: derived from radar observations (valid for large particles; virgae).

**Fig. 4.** Lidar and radar observations on 2 August 2012 (case 2). Left: Ice-water content IWC, right: 1064-532 nm attenuated backscatter coefficient, right: radar reflectivity.

~~Vertical velocity for case 2, derived from radar observations.~~

**Fig. 5.** Cloudnet derived water contents for case 2. Left: Liquid water content, right: ice water content (both in logarithmic scale).

**Fig. 6.** Vertical velocity for case 2. Left: derived from lidar (valid for more numerous smaller droplets at cloud base), right: derived from radar observations (valid for large particles; virgae).

**Fig. 7.** Vertical profiles of temperature (left) and relative humidity (right) with respect to liquid water (full lines) and ice (dashed lines) based on a radiosonde observation (Meiningen) for case 1 (black) and from GDAS (grid point Leipzig) for case 2 (red).

**Fig. 8.** Vertical velocity field of the inner cylinder for case 1. Left: Height dependence (red line) and temporal evolution of one realization of the stochastic vertical velocity field (black line) for  $w_{ave} = 0.3 \text{ m/s}$  at  $h_{mid}$ . Right: Histogram of velocity field. Vertical velocity for case 2 is identical but for heights between 7100 m and 7700 m.

**Fig. 9.** Liquid LWC (contours) and ice-water-mixing-ratio-LWC (colours, logarithmic scale) for case 1. Comparison of different values for  $h_{bot}$  (Upper left: W\_base,  $h_{bot} = 3900$  m, upper right: W\_h38,  $h_{bot} = 3800$  m, lower left: W\_h40,  $h_{bot} = 4000$  m, lower right: W\_h41,  $h_{bot} = 4100$  m.)

**Fig. 10.** Liquid (lower panel) and ice water paths (upper panel) for case 1. Comparison of the different values for  $h_{bot}$ .

**Fig. 11.** Liquid-LWC (contours) and ice-water-mixing-ratio-IWC (colours, logarithmic scale) for case 1. Comparison of different average updraft velocities  $w_{ave}$  (Upper panel: Left: W\_w01,  $w_{ave} = 0.1$  m/s, right: W\_w04,  $w_{ave} = 0.4$  m/s.) and different stochastic realizations (Lower: Left: W\_r1, r1, right: W\_r4, r4).

**Fig. 12.** Liquid (lower panels) and ice water paths (upper panels) for case 1. Comparison of the different values for  $w_{ave}$  (left) and the different stochastic realizations (right).

**Fig. 13.** Liquid-LWC (contours) and ice-water-mixing-ratio-IWC (colours, logarithmic scale) for case 2. Comparison of different values for  $h_{bot}$  (Upper left: C\_base,  $h_{bot} = 7100$  m, upper right: C\_h70,  $h_{bot} = 7000$  m, lower left: C\_h72,  $h_{bot} = 7200$  m, lower right: C\_h73,  $h_{bot} = 7300$  m.)

**Fig. 14.** Liquid (lower panel) and ice water paths (upper panel) for case 2. Comparison of the different values for  $h_{bot}$ .

**Fig. 15.** Liquid-LWC (contours) and ice-water-mixing-ratio-IWC (colours, logarithmic scale) for case 2. Comparison of different average updraft velocities  $w_{ave}$  (Upper: Left: C\_w01,  $w_{ave} = 0.1$  m/s, right: C\_w04,  $w_{ave} = 0.4$  m/s) and the different stochastic realizations (Lower: Left: C\_r1, r1, right: C\_r4, r4).

**Fig. 16.** Liquid (lower panels) and ice water paths (upper panels) for case 2. Comparison of the different values for  $w_{ave}$  (left) and the different stochastic realizations (right).

**Fig. 17.** Liquid-LWC (colours) and ice-water-mixing-ratio-IWC (contours, logarithmic scale) for case 1 (W\_in6, left) and case 2 (C\_in6, right). Enhancing IN by increasing  $N_{AP,r>250nm}$  by a factor of 10.

**Fig. 18.** Liquid-LWC (colourscontours) and ice-water-mass-IWC per bin (contourscolours, both logarithmic scale) for case 1 (upper panel) and case 2 (lower panel) for the respective base case (left) and the case with enhanced IN number (right; in6) after 16 and 17 minutes model time, respectively, corresponding to the IWP maximum of the base case runs.

**Fig. 19.** Liquid (lower panel) and ice water paths (upper panel) for case 1 (left) and case 2 (right). Comparison of the sensitivities with respect to IN number and ice particle shape.

**Fig. 20.** Liquid-LWC (colourscontours) and ice-water-mixing-ratio-IWC (contourscolours, logarithmic scale). Results for changing ice particle shape to hexagonal columns for case 1 (W\_col, left) and case 2 (C\_col, right).

**Fig. 21.** Liquid-LWC (colourscontours) and ice water mass per bin (contourscolours, both logarithmic scale) for case 1 (upper panel) and case 2 (lower panel) assuming columns as ice particle shape at IWP maximum of the respective base case (left) and at IWP of the run (right).