

1 **Impact of vehicular emissions on the formation of fine particles in the Sao Paulo**  
2 **Metropolitan Area: A numerical study with the WRF-Chem model**

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18

19 **Abstract**

20 The objective of this work is to evaluate the impact of vehicular emissions on the  
21 formation of fine particles (PM<sub>2.5</sub>;  $\leq 2.5$   $\mu\text{m}$  in diameter) in the Sao Paulo Metropolitan  
22 Area (SPMA) in Brazil, where ethanol is used intensively as a fuel in road vehicles. The  
23 Weather Research and Forecasting with Chemistry (WRF-Chem) model, which  
24 simulates feedbacks between meteorological variables and chemical species, is used as  
25 photochemical modelling tool to describe the physico-chemical processes leading to

26 evolution of number and mass size distribution of particles through gas-to-particle  
27 conversion. A vehicular emission model based on statistical information of vehicular  
28 activity is applied to simulate vehicular emissions over the studied area. The simulation  
29 has been performed for a one month period (7 August - 6 September 2012) to cover the  
30 availability of experimental data from the NUANCE-SPS (Narrowing the Uncertainties  
31 on Aerosol and Climate Changes in Sao Paulo State) project that aims to characterize  
32 emissions of atmospheric aerosols in the SPMA. The availability of experimental  
33 measurements of atmospheric aerosols and the application of the WRF-Chem model  
34 made it possible to represent some of the most important properties of fine particles in  
35 the SPMA such as the mass size distribution and chemical composition, besides  
36 allowing us to evaluate its formation potential through the gas-to-particle conversion  
37 processes. Results show that the emission of primary gases, mostly from vehicles, led to  
38 a production of secondary particles between 20 and 30 % in relation to the total mass  
39 concentration of  $PM_{2.5}$  in the downtown SPMA. Each of  $PM_{2.5}$  and primary natural  
40 aerosol (dust and sea salt) contributed with 40-50% of the total  $PM_{10}$  (i.e. those  $\leq 10 \mu m$   
41 in diameter) concentration. Over 40% of the formation of fine particles, by mass, was  
42 due to the emission of hydrocarbons, mainly aromatics. Furthermore, an increase in the  
43 number of small particles impaired the ultraviolet radiation and induced a decrease in  
44 ozone formation. The ground level  $O_3$  concentration decreased by about 2% when the  
45 aerosol-radiation feedback is taken into account.

46

## 47 **1. Introduction**

48 The Sao Paulo Metropolitan Area (SPMA), in the southeast region of Brazil, is  
49 considered a megalopolis comprised of Sao Paulo city and more 38 municipalities. One  
50 of the main concern in the SPMA is the occurrence of violations of air quality standards

51 for ozone and fine particles at different air quality stations from the Sao Paulo  
52 Environmental Agency (CETESB). The air pollutant emissions in the SPMA are related  
53 to the burning of the fuels: ethanol, gasohol (gasoline with 25% ethanol) and diesel.  
54 Recent work of Carvalho et al. (2014) reported a substantial increase in number of road  
55 vehicles from 1 million in 2000 to almost 7 million in 2014, together with an overview  
56 of the pollutants concentration, fuel use in the SPMA and the relationship between the  
57 emissions and the improvement in the air quality in past years.

58         They constitute the main cause of impairment to air quality in the SPMA, but  
59 the number of air quality standard violations has decreased for almost all pollutants with  
60 the exception of PM<sub>2.5</sub> and O<sub>3</sub>. Both these pollutants are impacted by the vehicular  
61 emissions and have experienced an increase in the number of violations of local air  
62 quality standards as discussed in detail by Carvalho et al. (2014). Pérez-Martínez et al.  
63 (2015) have analyzed the monthly mean values for the regulated pollutants from 2000 to  
64 2013 for the air quality stations in the SPMA. They found a decrease in the average  
65 concentration of NO<sub>x</sub>, CO and PM<sub>10</sub> by 0.65, 0.37 and 0.71 % month<sup>-1</sup>, respectively,  
66 although the sales of the fuels (gasoline, ethanol, and diesel) had increased by 0.26, 1.96  
67 and 0.38 % month<sup>-1</sup>, respectively.

68         A recent report from CETESB (CETESB, 2013) highlighted that, in 2012, the  
69 vehicles contributed with about 40% of the total PM<sub>10</sub> mass concentrations through  
70 direct emissions. If we consider the secondary aerosols, which were about 25% of PM<sub>10</sub>  
71 as estimated by CETESB (2013), these were mainly found to be formed by chemical  
72 reactions between gases released from exhaust of vehicles.

73         The implementation of the Program for the Control of Vehicular Emission  
74 (PROCONVE) established by the Brazilian Government in the 80's, enforcing measures  
75 such as use of catalytic converters and ethanol as additive to gasoline in substitution of

76 tetraethyllead, led to decrease in emissions of CO and VOCs and hence their ambient  
77 concentration. Although the emissions have been controlled by regulations, the number  
78 of vehicles has increased substantially and faster than the replacement of the old  
79 vehicles by the new ones (Pérez-Martínez et al., 2014). According to CETESB (2013),  
80 the road vehicles contributed up to about 97, 87 and 80% of CO, VOCs and NO<sub>x</sub>  
81 emissions in 2012, respectively.

82         To date, many studies assessing the impact of biofuels on the air quality have  
83 been performed in Brazil. For example, Anderson (2009) conducted a review  
84 concerning the use of ethanol fuel in Brazil. His work highlighted that the atmospheric  
85 concentrations of acetaldehyde and ethanol are much higher in Brazil in comparison  
86 with the other areas of the world. Costa and Sodr  (2010) showed that exhaust  
87 emissions of hydrous ethanol reduced CO and Hydrocarbons (HC), but increased CO<sub>2</sub>  
88 and NO<sub>x</sub> levels.

89         A number of past studies has shown the significant participation of the  
90 carbonaceous compounds in the concentration of fine particles in the SPMA  
91 (Albuquerque et al., 2011; Miranda and Andrade, 2005; Ynoue and Andrade, 2004;  
92 Castanho and Artaxo, 2001). Studies conducted on ambient air pollution in the SPMA  
93 have also shown that black carbon (BC) explains 21% of mass concentrations of fine  
94 particles (PM<sub>2.5</sub>;  $\leq 2.5$   $\mu\text{m}$  in diameter) compared with 40% of organic carbon (OC),  
95 20% of sulfates, and 12% of soil dust (Andrade et al., 2012). Most of the observed  
96 ambient PM<sub>2.5</sub> mass concentration usually originates from precursors gases such as  
97 sulphur dioxide (SO<sub>2</sub>), ammonia (NH<sub>3</sub>), nitrogen oxides (NO<sub>x</sub>) and volatile organic  
98 compounds (VOCs) as well as through the physico-chemical processes such as the  
99 oxidation of low volatile hydrocarbons noted above transferring to the condensed phase  
100 (McMurry et al., 2004; Heal et al., 2012). Since these processes are often photo-

101 chemically driven, the resultant aerosol usually falls into the category of secondary  
102 photochemical pollutant (Jenkin and Clemitshaw, 2000). Oxidation of VOCs can  
103 produce species of sufficiently low vapor pressure to be condensable, leading to the  
104 formation of secondary organic aerosol (SOA) (Kroll and Seinfeld, 2008). Fine particles  
105 in SPMA have a great participation on its composition of SOA, formed from the  
106 emissions of VOCs, which have the same origin of the primary compounds involved in  
107 the formation of ozone, from the burning of fuels. The participation of the biogenic  
108 emission is considered to be small in the formation of particles in the metropolitan area  
109 of the city according to previous studies of Martins et al. (2006).

110         The impact of the fine particles has been discussed in previous works, with  
111 evaluation of the scattering and absorbing effects of the aerosol (e.g. Li et al., 2005;  
112 Real et al., 2011). Vehicular emissions of particulate matter (PM) in the SPMA have a  
113 high percentage of BC (Brito et al., 2013), which after emitted to the atmosphere can  
114 enhance the absorption coefficient and thus the attenuation rates.

115         One of the most important aspects of this work is the quantitative analysis of the  
116 formation of  $PM_{2.5}$  and ozone ( $O_3$ ) in the SPMA. Photolysis of  $O_3$  by ultraviolet light in  
117 the presence of water vapor is the main source of hydroxyl radical (OH), the most  
118 important radical in the atmosphere in terms of reactivity (Monks, 2004). At the same  
119 time, OH levels in the atmosphere directly determine the oxidation rate of the precursors  
120 of secondary aerosols. Oxidation products of VOCs and semi-VOCs by OH are the  
121 most important precursors of SOA (Li et al., 2011a). Although VOCs and  $NO_x$  are  
122 precursors of both  $O_3$  and a fraction of atmospheric PM ( $NO_3^-$  and secondary organics)  
123 while they influence indirectly the formation of the rest of the secondary PM  
124 components like  $SO_4^{2-}$ , their control strategies that are optimal for  $O_3$  controls may even  
125 increase  $PM_{2.5}$  concentrations (McMurry et al., 2004). Such an analysis is important to

126 evaluate the contribution of the vehicular fleet using different kind of fuels to the  
127 concentration of fine particles. In this sense, a numerical study with an adequate  
128 physical approach, representing particles in the modelling system, is important to  
129 understand the formation of secondary aerosols from primary emission of gases in a  
130 metropolitan area where the composition of fuel in vehicular fleet has changed  
131 significantly over the past years. Therefore, the goal of the present study is to evaluate  
132 the impact of vehicular emissions on the formation of fine particles in the SPMA,  
133 focusing especially on the potential formation of secondary particles from the primary  
134 emission of gases coming from on-road vehicles. The impact of aerosol particles on the  
135 ozone photochemistry is also examined by means of numerical simulations.  
136 Measurements were performed to provide input data to evaluate the modelling  
137 performance and estimate the vehicular emission factors. Aerosol measurements were  
138 taken from field campaigns that were carried out as part of the Narrowing the  
139 Uncertainties on Aerosol and Climate Changes in Sao Paulo State (NUANCE-SPS)  
140 project (<http://nuance-lapat.iag.usp.br/>). These campaigns took place between July and  
141 September 2012. An online-coupled meteorology and chemistry model, i.e., the  
142 Weather Research and Forecasting with Chemistry (WRF-Chem) model, has been used  
143 to characterize and describe the physico-chemical processes involved in both the  
144 formation and growth of new particles over the SPMA in southern Brazil. The details of  
145 the experimental campaigns, WRF-Chem model and emissions are described in Section  
146 2. Results from modelling experiments and comparison with measurements are  
147 presented in Section 3. Finally, the summary and conclusions are given in Section 4.

148

## 149 **2. Methodology**

### 150 **2.1. Observational datasets**

151           The study period starting from 7 August until 6 September 2012 was selected for  
152 comparison with the modelled results (Section 2.2) due to the availability of  
153 experimental data from the NUANCE-SPS project. The aim of NUANCE-SPS was to  
154 evaluate the impact of emissions in the SPMA on the air quality and changing climatic  
155 conditions, and feedback mechanisms between climatic perturbations produced by both  
156 primary and secondary emissions and urban atmospheric processes. Aerosol observation  
157 datasets used in this work were collected using a Dichotomous sampler (Wedding et al.,  
158 1980) and a Micro-Orifice Uniform Deposit Impactor (MOUDI, model 100; MSP  
159 Corporation - Marple et al., 1986). The MOUDI impactor collected particles in 10 size  
160 classes with nominal 50% cut-off diameters: 10, 5.6, 3.2, 1.8, 1.0, 0.56, 0.32, 0.18, 0.1  
161 and 0.06  $\mu\text{m}$ . Particles smaller than 0.06  $\mu\text{m}$  were collected in a subsequent stage or  
162 after-filter. The samples collected with the MOUDI impactor were deposited on a  
163 polycarbonate membrane filter with 0.4  $\mu\text{m}$  porous and for the Dichotomous sampler  
164 the substrate was a teflon membrane filter with 2  $\mu\text{m}$  porous. The after-filter in the  
165 MOUDI impactor is a 33 mm teflon membrane filter, which was not submitted to the  
166 reflectance analysis. The collected membrane filters sampled with the Dichotomous and  
167 MOUDI samplers were analyzed to the identification of trace elements of mass through  
168 X-ray diffraction analysis, mass concentration through gravimetric analysis, and black  
169 and organic carbon through reflectance and thermo analysis using a thermal-optical  
170 transmittance (TOT) (Sunset Laboratory Inc. – Birch and Cary, 1996). Ion  
171 concentrations were evaluated through the ion chromatography analysis of the soluble  
172 material collected on the membrane filters (sulphate, nitrate, ammonium, sodium, and  
173 chloride). All these samplings were performed on the roof of the main building of the  
174 Institute of Astronomy, Geophysics and Atmospheric Sciences of the University of Sao  
175 Paulo (IAG-USP) (hereafter also referred as IAG-USP measurement site or simply

176 IAG-USP), which is inside a small green-park (approximately 7.4 km<sup>2</sup>), with local  
177 traffic during the day and surrounded by major roads with intense traffic by light and  
178 heavy-duty vehicles (Nogueira et al., 2014). Table 1 lists the aerosol instrumentation  
179 deployed roughly at the IAG-USP measurement site. In addition, ambient data from the  
180 CETESB's air quality monitoring network and the IAG-USP's climatological station  
181 (hereafter also referred as AF-IAG) were also considered for evaluation of numerical  
182 simulations. The locations of measurement sites are depicted in Fig. 1 whereas  
183 geographic coordinates and the list of pollutants and meteorological parameters  
184 monitored at each site is available in Table 2.

185

## 186 **2.2. WRF-Chem model**

187 The WRF-Chem model is a fully coupled online meteorological and chemical  
188 transport model (Grell et al., 2005), supported by National Center for Atmospheric  
189 Research (NCAR) of the USA and several other research institutions around the world.  
190 This model is a system of two key components. The WRF-Chem meteorological  
191 component, the Weather Research and Forecasting (WRF), is a system configured for  
192 both research and operational applications. The dynamical core used in this study is the  
193 Advanced Research WRF (ARW). Model's equations into ARW are solved to non-  
194 hydrostatic conditions on a fully compressible atmosphere. Further details on the  
195 modelling system can be found on the WRF model website ([http://www.wrf-](http://www.wrf-model.org)  
196 [model.org](http://www.wrf-model.org)). On the other hand, the WRF-Chem chemical component treats chemical  
197 processes such as dry deposition, gas-phase chemistry, photolysis rates, and aerosols  
198 chemistry. A detailed description of the WRF-Chem model can be found on its website  
199 (<http://ruc.noaa.gov/wrf/WG11>). Since both meteorological and chemical components  
200 are fully coupled, the transport of all chemical species is on-line. The gas-phase



201 chemistry and aerosol modules employed in this study are the Regional Acid Deposition  
202 Model, version 2 (RADM2) (Chang et al., 1989) and the Modal Aerosol Dynamics  
203 Model for Europe - Secondary Organic Aerosol Model (MADE - SORGAM)  
204 (Ackermann et al., 1998; Schell et al., 2001), respectively. The inorganic species  
205 included in the RADM2 mechanism are 14 stable species, 4 reactive intermediates, and  
206 3 abundant stable species (oxygen, nitrogen and water). Atmospheric organic chemistry  
207 is represented by 26 stable species and 16 peroxy radicals. The RADM2 mechanism  
208 represents organic chemistry through a reactivity aggregated molecular approach  
209 (Middleton et al., 1990). Similar organic compounds are grouped together in a limited  
210 number of model groups through the use of reactivity weighting. The aggregation  
211 factors for the most emitted VOCs are given in Middleton et al. (1990).

212         On the other hand, the most important process for the formation of secondary  
213 aerosol particles is the homogeneous nucleation in the sulfuric acid-water system. It is  
214 parameterized in MADE, following the method of Kulmala et al. (1998). Aerosol  
215 growth by condensation occurs in two steps: the production of condensable material  
216 (vapor) by the reaction of chemical precursors, and the condensation and evaporation of  
217 ambient volatile species on aerosols. The inorganic chemistry system, based on the  
218 Model for an Aerosol Reacting System (MARS) (Saxena et al., 1986) and its  
219 modifications by Binkowski and Shankar (1995), calculates the chemical composition  
220 of a sulphate-nitrate-ammonium-water aerosol according to equilibrium  
221 thermodynamics. The organic aerosol chemistry is based on the SORGAM, which  
222 assumes that SOA compounds interact and form a quasi-ideal solution (Grell et al.,  
223 2005). The SOA formation in SORGAM follows the two-product approach (Odum et  
224 al., 1996) where the oxidation of hydrocarbons produces two types of modelled  
225 semivolatile compounds that are partitioned between the gas and particle phases after

226 considering the absorptive partitioning theory (Pankow, 1994a; b). The primary organic  
227 aerosol (POA) in MADE is calculated from the primary anthropogenic emission of OC.  
228 Then, one may calculate the predicted OC concentration from the sum of both SOA and  
229 POA. The concurrent organic matter (OM) can be obtained from the OC concentration  
230 by application of a conversion factor. Brown et al. (2013) showed that the average  
231 OM:OC ratio was 1.54 (with a standard deviation of 0.2) for sites with low amount of  
232 secondary aerosol formation. It is important to note that this ratio can change from one  
233 place to another. In areas impacted by biomass burning the ratio can be higher. Gorin et  
234 al. (2006) assumed a ratio of 1.6 for the conversion from OC to OM over an area that  
235 experiences a significant wood smoke influence.

236

### 237 **2.2.1. Model configuration**

238 WRF-Chem version 3.5 was configured with three nested grid cells: coarse (75  
239 km), intermediate (15 km) and fine (3 km). The coarse grid cell covered a big region of  
240 Brazil and also of the Atlantic Ocean. The intermediate grid covered the southeast  
241 Brazil while the fine grid cell covered barely the SPMA and metropolitan areas nearest  
242 to it. Fig. 1 shows the arrangement of measurement sites and topography in the  
243 downtown area of the 3-km modelling domain. The initial and boundary meteorological  
244 conditions are from the National Center for Environmental Prediction's Final  
245 Operational Global Analysis with 1° of grid spacing, 26 vertical levels and are available  
246 every six hours: 00, 06, 12 and 18 UTC (<http://rda.ucar.edu/datasets/ds083.2/>). The  
247 initial and boundary chemical conditions for representing gases and aerosols  
248 background concentration were obtained from the Model for Ozone and Related  
249 chemical Tracers, version 4 (MOZART-4; Emmons et al., 2010). This model was  
250 driven by meteorological inputs from the Goddard Earth Observing System Model,

251 version 5 at a horizontal resolution of  $1.9^{\circ} \times 2.5^{\circ}$ , 56 vertical levels that are also available  
252 every six hours. Table 3 lists the WRF-Chem configuration options employed by this  
253 study.

254 WRF-Chem simulation with coupled primary aerosol (dust, sea salt and  
255 anthropogenic) and gas (biogenic and anthropogenic) emission modules, together with  
256 the direct effect of aerosol particles turned on, is performed as the control simulation in  
257 order to evaluate the model performance (hereafter referred to as Case\_0). For  
258 secondary aerosols, a simulation scenario (Case\_1) with biogenic and anthropogenic  
259 gases emission is performed to evaluate its formation potential. An additional  
260 simulation (Case\_2) is also performed to evaluate the impact of aerosols on ozone  
261 photochemistry. Notation and description of simulations are listed in Table 4. The first  
262 seven days of each simulation were not analyzed and used for model spin-up.

263

## 264 **2.3. Emissions**

### 265 **2.3.1. Anthropogenic emissions**

266 Anthropogenic emissions of trace gases and particles in both 3 and 15 km  
267 modelling domains were considered to include emissions only coming from on-road  
268 vehicles through the use of a vehicular emission model developed by the IAG-USP's  
269 Laboratory of Atmospheric Processes (LAPAt). Basically, this model considers the  
270 number of vehicles, vehicular emission factors, and average driving kilometers for  
271 vehicle per day as basic parameters for the calculation of exhaust emissions considering  
272 different vehicle types (light-duty vehicles, heavy-duty vehicles, and motorcycles) and  
273 different fuel types (ethanol, gasohol, combination of any proportion of gasohol and  
274 ethanol, and diesel) according to CETESB (2012). The details of this model are  
275 available in Andrade et al. (2015). In the case of VOCs, there are other two relevant

276 emissions (fuel transfer and evaporative processes) associated with the vehicles, besides  
277 the exhaust emissions. Because of the complexities in the spatial representation due to a  
278 numerous factors such as emissions at service stations, such emission sources are  
279 assumed to be emitted by exhaust of vehicles for the sake of simplicity. The vehicular  
280 fleet and intensity of use datasets are provided by the National Department of Traffic  
281 (DENATRAN) and the Sao Paulo Transporte (SPTrans), respectively. Emission factors  
282 for road vehicles for most pollutants were considered from previous studies performed  
283 inside the road tunnels (i.e. Janio Quadros, referred as JQ tunnel, and the tunnel 3 of the  
284 Rodoanel Mario Covas that is referred hereafter as RA tunnel) located within the SPMA  
285 (Pérez-Martínez et al., 2014; Nogueira et al., 2014). However, emission factors for  
286 VOCs are considered from dynamometer protocols (CETESB, 2010). VOCs and  
287 particulate matter speciation profiles used by pas-phase and aerosol chemical modules  
288 were also obtained from NUANCE-SPS experimental campaigns performed in 2011  
289 (tunnel measurements) and 2012 (ambient data). It is important to note that due to the  
290 lack of information on vehicular emission factors and intensity of use for most of the  
291 other metropolitan areas inside both modelling domains (e.g. the Campinas  
292 Metropolitan Area, which is shown by the second largest grey stain in Fig. 2), the  
293 calculation of vehicular emissions for these urban areas was carried out on the basis of  
294 the parameters found for the SPMA. The number of vehicles in any modelling domain  
295 is calculated from the sum of the number of vehicles in each one of the main urban  
296 areas inside the modelling domain in question.

297         Spatial distribution of emissions for the 3 km modelling domain resolution was  
298 based on road density products compiled by the OpenStreetMap project and extracted  
299 from the Geofabrik's free download server (<http://download.geofabrik.de>). Urban areas  
300 were assumed to allocate high emissions since these concentrate a road density greater

301 than other areas. In the case of the 15 km modelling domain, emissions are based on  
302 night-time lights data from the Defense Meteorological Satellite Program  
303 (<http://ngdc.noaa.gov/eog/dmsp/downloadV4composites.html>). These images are 30 arc  
304 second grids, spanning from  $-180^{\circ}$  to  $+180^{\circ}$  longitude and  $-65^{\circ}$  to  $+75^{\circ}$  latitude and  
305 contain the lights from cities, towns and other sites with persistent lighting, including  
306 gas flares. Cleaned up night-time light points with no ephemeral events such as forest  
307 fires are used to allocate emissions. To estimate the number of vehicles in each grid  
308 point of both domains, the sum of individual intensities at each point (i.e. total road  
309 length for the 3 km modelling domain and night-time light for the 15 km modelling  
310 domain) is firstly normalized by the total fleet, and then distributed uniformly using the  
311 total fleet distribution so that emissions in urban areas are mainly represented by  
312 emissions coming from their vehicles. Furthermore, due to the complexity involved in  
313 describing the temporal variation of emissions at each grid point, median values for  
314 vehicular traffic obtained from measurements inside the JQ and RA tunnels (Pérez-  
315 Martínez et al., 2014) were used for distributing the emissions during the day in both  
316 domains. This approximation followed the approach used by Fast et al. (2006) where  
317 emission profiles were calculated from median diurnal variations on weekdays and  
318 weekends. We have applied the same constant diurnal cycle at all grid points where  
319 emissions have values greater than zero. VOC and PM emission profiles were assumed  
320 to be the same as for CO and NO<sub>x</sub> emission profiles since these pollutants are also  
321 characteristic tracers of emissions of light-duty and heavy-duty vehicles, respectively.  
322 Fig. 2 shows the maximum hourly emission rates for aromatic VOCs in the 3 km  
323 modelling domain. Anthropogenic emissions were not considered in the 75 km  
324 modelling domain.

325           The Another Assimilation System for WRF-Chem (AAS4WRF) chemical  
326 emissions pre-processor developed by the Latin American Observatory (OLE2; Muñoz  
327 et al., 2010; 2012) was used to scale emission rates on WRF curvilinear coordinates.  
328 AAS4WRF is appropriate to write chemical emission rates from both surface and  
329 elevated sources in the proper WRF data file format, providing an alternative tailored  
330 way to assimilate emissions to WRF-Chem. The method is explained in the OLE2 Wiki  
331 pages in detail ([http://www.cmc.org.ve/mediawiki/index.php?title=Calidad\\_de\\_Aire](http://www.cmc.org.ve/mediawiki/index.php?title=Calidad_de_Aire)).

332

### 333 **2.3.2. Other emissions**

334           Biogenic emissions are calculated online based on the Guenther scheme  
335 (Guenther et al., 1993; 1994). The Guenther biogenic emissions model calculates the  
336 emission rates using temperature, photo-synthetically active radiation flux and land-use  
337 data as the U.S. Geological Survey (USGS) land-use cover system classification if  
338 coupled with the WRF model. However, as indicated in the WRF-Chem emissions  
339 guide ([http://ruc.noaa.gov/wrf/WG11/Emission\\_guide.pdf](http://ruc.noaa.gov/wrf/WG11/Emission_guide.pdf)), several key chemical  
340 species would have been representing relatively low emission rates because of the  
341 limited vegetation types in the simulation, and thus their impacts are anticipated to be  
342 much lower than those from vehicular emissions.

343           Dust and sea salt emissions are calculated online following the works of Ginoux  
344 et al. (2001) and Gong (2003), respectively. The calculation of Ginoux et al. (2001) for  
345 the uplifting of dust particles is based on the surface wind speed, wetness and  
346 information on soil characteristics. The model then solves the continuity equation  
347 including the emission, chemistry, advection, convection, diffusion, dry deposition, and  
348 wet deposition of each species. The parameterization of sea salt aerosol source function  
349 of Gong (2003) is an extended parameterization of Monahan et al. (1986), which scales

350 the generation of marine aerosols from mechanical disruption of wave crests by the  
351 wind and sea surface covered by whitecaps.

352

### 353 **3. Results and discussion**

#### 354 **3.1. Characterization of meteorological conditions**

355 In order to study and understand the spatial and temporal variability of  
356 atmospheric aerosols, O<sub>3</sub>, and other pollutants (i.e. CO, NO<sub>x</sub>) during the study period, it  
357 was first necessary to analyze the behavior of main meteorological systems acting on  
358 the atmospheric environment of the SPMA and surrounding areas.

359 According to the monthly climate reports from the IAG-USP's Climate Research  
360 Group (GrEC), the observed precipitation rates were lower than the climatological value  
361 in SPMA (anomaly of -38.6 mm) and larger part of the Sao Paulo State during August  
362 2012. Negative anomalies on the precipitation were caused by the intensification of the  
363 South Atlantic Subtropical High (SASH). These conditions established an easterly wind  
364 anomaly pattern at the 850 hPa level. Conditions were unfavorable for relative humidity  
365 coming from the Amazon due to the Low Level Jet (LLJ) and less intense Alisian winds  
366 in the Tropical Atlantic (GrEC, 2012a). However, the action of frontal systems favored  
367 the rain accumulation in September 2012, mainly in western Sao Paulo State where the  
368 greater positive amount of anomalies was observed. Precipitation events were  
369 predominantly observed during the second half of the month. In this case, the wind  
370 pattern showed an opposite configuration to that observed in August 2012 as a result of  
371 the weakening of the SASH (GrEC, 2012b). The IAG-USP's climatological station  
372 recorded an accumulated precipitation of about 1.3 mm on three days of occurrence (28  
373 August, 30 August and 4 September 2012) and an easterly wind pattern with a median  
374 intensity of 2 m s<sup>-1</sup> during the period between 07 August and 06 September 2012. Fig. 3

375 shows the hourly accumulated precipitation and relative humidity observed at the IAG-  
376 USP's climatological station.

377

### 378 **3.2. Analysis of aerosol species**

379 Aerosol analysis included species such as organic carbon (OC), elemental  
380 carbon (EC), sulphate, nitrate, ammonium, sodium and chloride in addition to other  
381 elemental constituent of PM. All the sampling for these species were performed at IAG-  
382 USP. Results showed that the major contributors to the concentration of fine particles  
383 are OM (55.7%; OM:OC ratio of 1.5 found by Brito et al. (2013)) and EC (15%),  
384 followed by sulphate (2.9%), ammonium (2.1%), sodium (1.9%), nitrate (0.5%) and  
385 chloride (0.3%). The remaining mass (21.6%) is calculated by determining of the  
386 difference between the total mass of PM<sub>2.5</sub> (from the gravimetric analysis) and the sum  
387 of the masses of 7 individual compounds, as noted above. Part of this remaining mass is  
388 related to the water content of aerosols (Andrade et al., 2012).

389 On the other hand, PM<sub>2.5</sub>, PM<sub>10</sub> and size distribution of particles measured at  
390 IAG-USP show that the study period was characterized by a reduction in the  
391 concentrations up to the end of August 2012 when their minimum values were  
392 achieved. This reduction was related to the action of a semi-stationary front between the  
393 coasts of Sao Paulo and Parana States. After the passage of this system, aerosol  
394 concentrations have significantly increased what could be related to an increase in  
395 relative humidity once the SASH system is moved away from the continent, as well as  
396 the transport of aerosol particles produced by forest fires in the central-west region of  
397 Brazil and the Sao Paulo State. Several studies have shown the contribution of forest



398 fires on the atmospheric aerosol concentrations in SPMA (Vieira-Filho et al., 2013;  
399 Vasconcellos et al., 2010). One way to qualitatively evaluate the contribution of forest  
400 fires on aerosol concentrations is by using the air mass trajectories. The Hybrid Single-  
401 Particle Lagrangian Integrated Trajectory (HYSPLIT) model (Draxler and Hess, 1998)  
402 was used to calculate backward trajectories of air masses in order to identify  
403 atmospheric transport of air mass from forest fire areas. Fig. 4 shows the three-day  
404 backward trajectories of air masses starting at IAG-USP for the days 9 and 31 August  
405 and 5 September, when increases in the OC and EC concentrations were observed at  
406 IAG-USP. The pink markers on the map represent the observed fire locations during the  
407 study period considering different satellite products (GOES, AQUA, TERRA, NOAA).

408 Fig. 5 shows the concentration of OC, EC and some species of PM<sub>2.5</sub> during the  
409 study period at IAG-USP. We can observe eleven exceedances of PM<sub>2.5</sub> concentration  
410 with respect to the air quality standard of 25  $\mu\text{g m}^{-3}$  (see grey line in Fig. 5a) established  
411 by the World Health Organization (WHO). These exceedances have mainly occurred at  
412 the beginning and at the end of the study period when an increase in the concentrations  
413 of OC and EC were observed. The increasing organic matter could be associated to  
414 traffic incidents which may raise the emissions, which in case of less favorable  
415 meteorological conditions (e.g. lower height of lower planetary boundary layer, PBL, or  
416 slow transport of air pollutants) may have led to a more efficient formation of secondary  
417 particles. Castanho and Artaxo (2001) analyzed the behavior of the aerosol composition  
418 in SPMA and showed the increase in the concentration of inorganic and organic  
419 material in the winter season compared to the summer season, explaining this behavior  
420 with the meteorological characteristics: dry conditions with low height inversion layer  
421 in the wintertime and a rainy summer.

422 Size distributions of aerosol mass indicate that the majority of sulphate,  
423 ammonium and PM<sub>10</sub> mass concentration is distributed in the size range with diameters  
424 between 0.1 and 1 µm, commonly known as accumulation mode particles (Kumar et al.,  
425 2010). In the cases of nitrate, sodium, and chloride, most part of mass was concentrated  
426 in particles with diameters greater than 1 µm.

427

### 428 **3.3. Comparison of baseline simulation with observations**

429 All the numerical results presented in this section, for the purpose of comparison  
430 with the measurements, were obtained from the baseline simulation (Case\_0). The  
431 predicted temperature, humidity, and wind distribution have been compared to  
432 measurements from the AF-IAG and INT measurement sites. Overall, the model  
433 captured the diurnal variation of temperature, humidity, and wind directions reasonably  
434 well. However, the predicted wind speeds were slightly lower than the observed values.  
435 To evaluate the model performance in solving the meteorology and chemical species,  
436 we computed the statistics correlation coefficient (R), Bias (B), and root mean square  
437 error (RMSE<sub>UB</sub>). The definitions of these statistics are given in the Appendix. Table 5  
438 presents the summary of these statistics, showing comparisons between WRF-Chem  
439 predictions and observations. Fig. 6 shows the predicted average of wind vectors at 10  
440 m and temperature at 2 m for the whole study period in the 3 km modelling domain.  
441 Blue dots represent the locations of AF-IAG and INT sites, while the numbers in cyan  
442 indicate the observed average temperatures (i.e. 17.7 °C at AF-IAG and 17.8 °C at  
443 INT). On an average, the predicted wind direction was easterly in SPMA, which has  
444 somewhat affected the spatial distribution of aerosol particles as examined later in this  
445 section. Likewise, a good agreement is found between the predicted PM<sub>2.5</sub>, PM<sub>10</sub> and O<sub>3</sub>  
446 concentration and measurements at most of the sites. Figs. 7, 8 and 9 show the observed

447 and predicted temporal variations of PM<sub>2.5</sub>, PM<sub>10</sub> and O<sub>3</sub> concentrations at 3, 10 and 6  
448 sites in the SPMA, respectively, with some measurement sites sharing the same grid  
449 point for comparisons due to the geographical proximity. These figures suggest that  
450 predicted concentrations did not present any significant spatial variation in the  
451 downtown SPMA and were generally underestimated when compared to measurements.  
452 This under prediction could be associated with an underestimation on the vehicular  
453 emissions as well as other emission sources (e.g. emissions coming from industry) that  
454 are disregarded in this study, in addition to predicted surface winds more intense than  
455 those observed, leading to a dilution of aerosol particles in SPMA. The high  
456 concentrations of PM<sub>2.5</sub> and PM<sub>10</sub> observed at the beginning and at the end of the study  
457 period, whose variability and trends were reasonably well captured by the model, could  
458 be related with the emission of high aerosol loadings due to traffic incidents as well as  
459 the establishment of lower PBL heights, commonly observed under post-frontal  
460 situations. The results for this simulation (Case\_0) show that overall the predicted PBL  
461 heights (not shown here) have a regular diurnal variation in the downtown SPMA with  
462 averaged daily values around 500 m at both the beginning and the end, and of up to 700  
463 m in the middle of the study period, when lower concentrations of aerosols were  
464 observed. Statistics to quantify the model performance in the representation of PM<sub>2.5</sub>,  
465 PM<sub>10</sub> and O<sub>3</sub> concentrations can be visualized along with the Taylor Diagram (Taylor,  
466 2001) shown in Fig. 10. In general, most of evaluated parameters present good  
467 correlation coefficients, mainly those for PM<sub>10</sub>, but with negative biases and standard  
468 deviations lower than those for observations. The mean biases for PM<sub>2.5</sub>, PM<sub>10</sub> and O<sub>3</sub>  
469 concentrations were -8.84 µg m<sup>-3</sup>, -14.13 µg m<sup>-3</sup> and -0.85 ppb, respectively (see Table  
470 5).

471           Figures 11-13 show the predicted average surface distribution of  $PM_{2.5}$ ,  $PM_{10}$   
472 and  $PM_{2.5}:PM_{10}$  ratio for the 3 km modelling domain, respectively. Red dots and cyan  
473 numbers represent the locations and the observed mean PM concentrations (or mean  
474 PM concentration ratios) at the measurement sites, respectively. Major contributions of  
475  $PM_{2.5}$  on the total  $PM_{10}$  concentration were observed mainly over offshore continental  
476 areas (see Fig. 13). High  $PM_{2.5}:PM_{10}$  concentration ratios would be firstly associated  
477 with the transportation of fine particles and gases from upwind regions (see Fig. 6),  
478 followed by a production of fine particles from biogenic emissions. Additional  
479 comparisons between the observed and predicted concentrations of OC and EC at IAG-  
480 USP (the only site with measurements of OC and EC) are shown in Fig. 14. In addition  
481 to an underestimation of emissions, under predicted OC concentrations could also be  
482 associated with an underestimation of SOA probably due to the absence of oxidation of  
483 monoterpenes and a limited treatment of anthropogenic VOCs oxidation in the RADM2  
484 mechanism, as discussed by Tuccella et al. (2012). The SORGAM aerosol module  
485 considers the formation of anthropogenic SOAs from the oxidation of alkane, alkene  
486 and aromatic VOCs as well as the biogenic SOA formation from the oxidation of alpha-  
487 pinene, limonene and isoprene VOCs. Recent studies coupling non-traditional SOA  
488 models (volatility basis set approaches) in WRF-Chem show improvements in the  
489 predicted SOA concentrations, although these are still lower than those observed (e.g.  
490 Li et al., 2011b; Ahmadov et al., 2012; Shrivastava et al., 2013).

491           On the other hand, measurements of mass size distribution were also made with  
492 a MOUDI impactor at IAG-USP, following the protocol describe in Miranda and  
493 Andrade (2005). Constituents of aerosol were subsequently determined by X-Ray  
494 fluorescence analysis and ion chromatography analysis. As previously indicated in this  
495 section, the main identified species are  $SO_4$ ,  $NO_3$ ,  $NH_4$ , Na and Cl. The observed

496 average aerosol composition is derived using measurements from both MOUDI  
497 impactor and SUNSET analyzer. To perform the comparisons of mass size distribution,  
498 we adequately joined the MOUDI bin sizes according to the three modes used by the  
499 MADE aerosol module: Aitken ( $<0.1 \mu\text{m}$ ), accumulation ( $0.1\text{-}1 \mu\text{m}$ ) and coarse ( $>1$   
500  $\mu\text{m}$ ). The observed and predicted aerosol mass size distributions averaged over the same  
501 sampling time period (16 days along the study period) are shown in Fig. 15. Over the  
502 downtown SPMA, both the observed and predicted fine particles from accumulation  
503 mode account for majority of the total  $\text{PM}_{2.5}$  mass. Since the formation-growth  
504 processes of aerosols in question are explicitly treated in the Aitken and accumulation  
505 modes, the predicted concentrations for particles larger than  $1 \mu\text{m}$  are assumed to be  
506 zero. In this case, the mass of particles larger than  $1 \mu\text{m}$  is allocated to the  $\text{PM}_{10}$  aerosol  
507 variable (see Fig. 15). The comparison between the observed and predicted average  
508 contributions for the main identified aerosol constituents at IAG-USP is shown in Fig.  
509 16. Both the observed and predicted OC and EC make up the largest fraction of  $\text{PM}_{2.5}$   
510 mass with contributions of 55 and 40%, respectively. In addition, it was found that the  
511 predicted SOA concentrations contribute 17% of the predicted total OC concentration at  
512 this measurement site. Various global and regional scale SOA simulations have been  
513 conducted using mass-based yield and partitioning coefficients, but they have  
514 underestimated the SOA concentrations by roughly an order of magnitude, especially  
515 over urban regions (Matsui et al., 2014). Using the same SOA formation approach  
516 employed by this study and a conversion factor of 1.6 to convert the emissions of OC to  
517 OM, Tuccella et al. (2012) found simulated SOA:OM ratios in the 5-40% range against  
518 the observed range of 50-80%. Although the predicted average  $\text{PM}_{2.5}$  concentration  
519 ( $14.48 \mu\text{g m}^{-3}$ ) was lower than observed ( $22.32 \mu\text{g m}^{-3}$ ), the mean aerosol chemical  
520 composition was reasonably well represented by the model (see Fig. 16).

521 **3.4. Contribution of dust-sea salt and coarse anthropogenic aerosols to PM**  
522 **concentration**

523 The evaluation of the contribution of dust and sea salt aerosols on PM<sub>10</sub>  
524 concentration is performed from the sum of their concentrations divided by the PM<sub>10</sub>  
525 concentration. The simulated average ratio between dust–sea salt aerosols and the total  
526 PM<sub>10</sub> mass concentration is shown in Fig. 17b. High concentration ratios have been  
527 observed over the ocean where sea salt emissions are by far the most important aerosols  
528 source. Unlike high concentration ratios over the ocean, lower concentration ratios are  
529 observed over the continent far away from the coast. In this region, the main sources of  
530 atmospheric aerosols would be the emission of primary biological aerosol, SOA formed  
531 from the emission of biogenic volatile organic compounds (BVOCs), and forest fires.  
532 However, particles could also be transported from remote areas. In addition, we can also  
533 observe that dust and sea salt aerosols have a contribution between 40 and 50% of the  
534 total PM<sub>10</sub> concentration in the downtown SPMA. Furthermore, it is possible to estimate  
535 the contribution of all the other PM<sub>10</sub> (i.e., the coarse anthropogenic aerosol) to the total  
536 PM<sub>10</sub> mass concentration. It may be directly calculated from the model or estimated  
537 from the Figs. 13 and 17b once the sum of concentrations of PM<sub>2.5</sub>, dust and sea salt,  
538 and coarse anthropogenic aerosol represents 100% of the total PM<sub>10</sub> mass concentration.  
539 For example, we found that the coarse anthropogenic aerosol represents around 10% of  
540 PM<sub>10</sub> in the downtown SPMA.

541 **3.5. Evaluation of secondary aerosol formation**

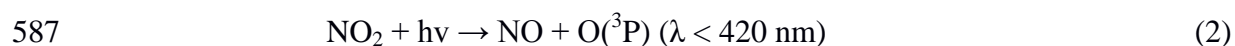
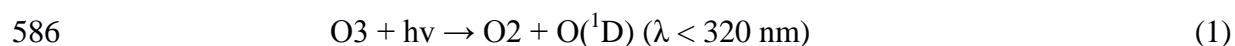
542 As described in Section 2.1, aerosol module employed by this study  
543 (MADE/SORGAM) includes the homogeneous nucleation in the sulphuric acid-water  
544 system. The sulphuric acid is the most significant condensable molecule formed in the  
545 atmosphere, which has also been long recognised as the most important molecule from

546 the point of view of the nucleation of new particles (Jenkin and Clemitshaw, 2000;  
547 Seinfeld and Pandis, 2006). However, for the SPMA, the importance of SOA formed  
548 from the anthropogenic emission of fuel used by the transport sector was noted (Salvo  
549 and Geiger, 2014). According to the official emission inventory developed by the Sao  
550 Paulo Environmental Protection Agency (CETESB, 2013), the SOA explains 51% of  
551 the fine particle mass concentration, with the vehicular emission being its main source.  
552 The subsequent growth processes involve aerosol growth by condensation of  
553 condensable material onto existing particles, and by coagulation of particles to form  
554 larger particles (Kumar et al., 2011; 2014). For example, particles in the accumulation  
555 mode emerge through coagulation of particles from the Aitken mode (Kumar et al.,  
556 2011). It is important to emphasize that the boundaries were updated with gas and  
557 aerosol background concentrations coming from the 15 km modelling domain during  
558 the whole simulation period. Thereafter, the impact of vehicular emissions on the  
559 formation of fine particles was calculated from the predicted  $PM_{2.5}$  concentration  
560 considering an emission scenario (Case\_1) in which only emission of gases from  
561 vehicles and vegetation are taken into account to be emitted to the atmosphere from the  
562 surface. The predicted average  $PM_{2.5}$  (Case\_1): $PM_{2.5}$  (Case\_0) ratio is shown in Fig.  
563 17a. A contribution between 20 and 30% in the predicted baseline  $PM_{2.5}$  concentration  
564 in downtown SPMA is found to correspond to the fine particles formation and  
565 transportation processes. Higher concentration ratios over the SPMA surroundings (30-  
566 50%) could be associated with more efficient biogenic emissions. Overall, it is observed  
567 that the formation efficiency increases towards the northwest from the ocean. Deep red  
568 areas in Fig. 17a could also be associated with the transportation of fine particles and  
569 gases from other regions, in addition to having a more efficient production of fine  
570 particles from biogenic emissions. For example, given the distribution of winds in Fig.

571 6, the northern boundary could represent the main source of particles and gases over this  
572 part of the simulation domain. Additionally, the comparison between the predicted and  
573 observed OC and EC concentrations at IAG-USP shown in Fig. 14 includes the Case\_1  
574 simulation in which only emission of primary gases is taken into account in the  
575 assessment of fine particles formation. The concentration peaks observed at the  
576 beginning and at the end of the study period may be associated with the transport of  
577 aerosol particles from both biomass and fossil fuel burning areas (see Fig. 4).  
578 Considering the Case\_1 simulation, we can observe very low concentrations for EC  
579 (mean concentration of  $0.01 \mu\text{g m}^{-3}$ ), as expected. This is because these particles are not  
580 produced by photochemical processes in the atmosphere, but associated mainly with the  
581 diesel exhaust.

### 582 **3.6. Aerosol impact on O<sub>3</sub> photochemistry**

583 Ozone photochemistry production mainly depends on the two key photolysis  
584 rates, as shown in Eqs. (1) and (2), i.e., shortwave radiation able to reach the surface to  
585 break molecules of O<sub>3</sub> and NO<sub>2</sub>.



588 Therefore, the impact of aerosols on O<sub>3</sub> photochemistry has been evaluated from  
589 the impact of aerosols on downward shortwave radiation. Attenuation (scattering and  
590 absorption) of downward shortwave radiation by aerosols may substantially modify the  
591 photolysis rates, and thereby affecting the ozone photochemistry production.

592 The average percentage change in surface O<sub>3</sub> concentrations at 16:00 h (local  
593 time) with and without the aerosol-radiation feedback module turned on are shown Fig.



594 17c. Overall O<sub>3</sub> is destroyed or formed (incoming transport from other regions) in small  
595 quantities between -1 and +1% in relation to its total concentration. In addition, it was  
596 observed that the surface O<sub>3</sub> concentration decreased by around 2% in the downtown  
597 SPMA. Li et al. (2011a) found that the impact of aerosols on O<sub>3</sub> formation in Mexico  
598 City was most pronounced in the morning with the O<sub>3</sub> reduction of 5-20%, but the  
599 reduction is less than 5% in the afternoon. Low reductions in the O<sub>3</sub> concentration in the  
600 downtown SPMA compared to results from other studies may be explained by the lower  
601 predicted PM<sub>10</sub> concentrations, which can lead to a minor attenuation of the incoming  
602 solar radiation. Simulated mean downward shortwave fluxes at ground surface (not  
603 shown) were up to 5% higher for the Case\_2 than for the Case\_0 during the afternoon.  
604 The inclusion of the direct effect of aerosol particles was found to have small reductions  
605 in the surface temperature (changes by around 2%), presumably due to an increase in  
606 the number of atmospheric processes involving downward longwave fluxes over this  
607 area. Forkel et al. (2012) found an underestimation of predicted downward longwave  
608 radiation over the southern Baltic Sea when the direct effect of aerosol particles was  
609 neglected. Despite the highly non-linear behavior of tropospheric O<sub>3</sub>, the reduction in  
610 the predicted O<sub>3</sub> concentrations indicates a high efficiency of aerosols to attenuate the  
611 downward shortwave radiation, what is plausible once it was found that low PM<sub>10</sub>  
612 concentrations have a capability to reduce ground level O<sub>3</sub> concentrations in a few ppb.

#### 613 **4. Summary and conclusions**

614 The WRF-Chem community model has been used to evaluate the impact of  
615 vehicular emissions on the fine particles formation in the SPMA. Three thirty-one day  
616 simulations, covering a period from 7 August to 6 September 2012, have been  
617 performed. The aims were to evaluate the impact of fine particles formation (both  
618 inorganic and SOA) from gases emitted by road vehicles as well as the aerosol impacts

619 on the ozone formation photochemistry. The results were compared with the  
620 measurements available from the NUANCE-SPS project.

621 The predicted temporal variations of meteorology,  $PM_{2.5}$ ,  $PM_{10}$  and  $O_3$  were  
622 found to agree well with the measurements at most of the sites during the entire  
623 simulation period. However, the predicted concentrations of  $PM_{2.5}$ ,  $PM_{10}$  and  $O_3$  (but in  
624 minor intensity) were lower than the observed values. This difference could be  
625 associated with an underestimation of the vehicular emissions and other emission  
626 sources such as industry, heating and cooking, which are not considered in this study.  
627 Wind speed and direction played an important role in the distribution of fine particles  
628 over the simulation domain. Backward trajectories analysis suggested that aerosol  
629 particles from biomass burning were transported to SPMA, impacting on the PM  
630 concentration over this region.

631 The baseline simulation (Case\_0) showed that dust and sea salt aerosols made a  
632 contribution between 40 and 50% of the total  $PM_{10}$  concentration in the downtown  
633 SPMA. On the other hand, the Case\_1, which represents simulations with gaseous  
634 emissions only, indicates that the emissions of primary gases coming mainly from  
635 vehicles have a potential to form new particles between 20 and 30% in relation to the  
636 baseline  $PM_{2.5}$  concentration found in the downtown SPMA. Finally, the Case\_2, which  
637 represents simulations with aerosol-radiation feedback turned on, reveals a reduction in  
638 the surface  $O_3$  concentration by around 2% in the afternoon (16:00 h; local time) when  
639 the aerosol-radiation feedback is taken into account.

640 This study provides a first step to understand the impact of vehicular emissions  
641 on the secondary particles formation in the SPMA. Nevertheless, more experimental  
642 campaigns are recommended for future work in order to characterize aerosols in

643 ambient air and to improve their emission estimates so that a better understanding of  
 644 physical and chemical properties and their formation can be established. This study also  
 645 evaluates the importance of the VOCs in the formation of not only O<sub>3</sub> but also of fine  
 646 particles. These compounds play an important role concerning health impacts and  
 647 climate change, and the control of their concentrations requires the description of their  
 648 formation mechanisms.

## 649 **Appendix A**

650 The statistics used in this study are defined as follows:

651 1. Correlation coefficient (R)

$$R = \frac{\sum_{i=1}^N (M_i - \bar{M})(O_i - \bar{O})}{\sqrt{\sum_{i=1}^N (M_i - \bar{M})^2} \sqrt{\sum_{i=1}^N (O_i - \bar{O})^2}}$$

652

653 2. Root mean square error UB (RMSE<sub>UB</sub>)

$$RMSE_{UB} = \sqrt{\frac{1}{N} \sum_{i=1}^N [(M_i - \bar{M}) - (O_i - \bar{O})]^2}$$

654

655 3. Bias (B)

$$B = \frac{1}{N} \sum_{i=1}^N (M_i - O_i)$$

656

657 where

658  $\bar{O} = \frac{1}{N} \sum_{i=1}^N O_i$  and  $\bar{M} = \frac{1}{N} \sum_{i=1}^N M_i$  are the average values of the individual observed

659 and predicted values,  $O_i$  and  $M_i$ , respectively. N is the number of observations.

660

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679

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889

890 Table 1. Description of aerosol sampling campaigns performed at IAG-USP.

Parameter	Sampling frequency	Period of sampling	Sampling device
Aerosol mass size distribution	24 hours	July-September	MOUDI impactor
PM <sub>2.5</sub> and PM <sub>10</sub> concentration	12 hours	July-September	Dichotomous sampler
OC and EC concentration	12 hours	August-September	Sunset OC-EC analyser

891

892

893 Table 2. Description of measurement sites.

Initials	Name	Latitude	Longitude	Measured species
NSO	Nossa S. do O	-23.4796	-46.6916	PM <sub>10</sub> , O <sub>3</sub>
SAN	Santana	-23.5055	-46.6285	PM <sub>10</sub>
PDP	Parque D. Pedro	-23.5448	-46.6276	PM <sub>10</sub> , O <sub>3</sub>
MOO	Mooca	-23.5497	-46.5984	PM <sub>10</sub> , O <sub>3</sub>
CCE	Cerqueira Cesar	-23.5531	-46.6723	PM <sub>10</sub>
IAG-USP	IAG-USP	-23.5590	-46.7330	PM <sub>10</sub> , PM <sub>2.5</sub> , OC, EC Aerosol mass size distrib. <sup>a</sup>
IPEN-USP	IPEN-USP	-23.5662	-46.7374	PM <sub>2.5</sub> , O <sub>3</sub> , NO <sub>x</sub> , CO
IBI	Ibirapuera	-23.5914	-46.6602	PM <sub>10</sub> , O <sub>3</sub> , NO <sub>x</sub> , CO
CON	Congonhas	-23.6159	-46.6630	PM <sub>10</sub> , PM <sub>2.5</sub>
AF-IAG	AF-IAG	-23.6500	-46.6167	T, RH, WS, WD <sup>b</sup>
SAM	Santo Amaro	-23.6545	-46.7095	PM <sub>10</sub>
INT	Interlagos	-23.6805	-46.6750	PM <sub>10</sub> , O <sub>3</sub> , T, RH, WS, WD

894 <sup>a</sup>includes SO<sub>4</sub><sup>2-</sup>, NO<sub>3</sub><sup>-</sup>, NH<sub>4</sub><sup>+</sup>, Na<sup>+</sup>, Cl<sup>-</sup> and PM<sub>10</sub>.

895 <sup>b</sup>T, RH, WS, and WD denote temperature, relative humidity, wind speed and wind  
 896 direction, respectively.

897

898 Table 3. Selected WRF-Chem configuration options.

Atmospheric Process	WRF-Chem option
Longwave radiation	RRTM
Shortwave radiation	Goddard
Surface layer	Monin-Obukhov
Land surface	Noah
Boundary layer	YSU
Cumulus clouds <sup>a</sup>	Grell 3D
Cloud microphysics	Lin
Gas-phase chemistry	RADM2
Aerosol chemistry	MADE/SORGAM
Photolysis	Fast-J

899 <sup>a</sup>Outer domains only

900



901 Table 4. Description of WRF-Chem simulations.

Label	Description
Case_0 (Baseline simulation)	Emission of gases Emission of aerosols Aerosol-radiation feedback turned on
Case_1	Emission of gases No emission of aerosols Aerosol-radiation feedback turned on
Case_2	Emission of gases Emission of aerosols Aerosol-radiation feedback turned off

902

903

904 Table 5. Performance statistics for WRF-Chem predictions at all sites<sup>a</sup>

Index	PM <sub>2.5</sub>	PM <sub>10</sub>	O <sub>3</sub>	NO <sub>x</sub>	CO	T	RH	U <sup>b</sup>	V <sup>c</sup>
R	0.73	0.72	0.63	0.42	0.54	0.71	0.62	0.48	0.44
B	-8.84	-14.1	-0.85	-8.75	-0.27	0.65	-5.74	-0.96	0.75
RMSE <sub>UB</sub>	6.83	10.59	27.45	30.35	0.57	3.21	20.06	1.04	1.02

905 <sup>a</sup>Values are averaged from all the individual indexes found at the measurement sites.

906 Individual indexes are calculated from both hourly observed and predicted values.

907 <sup>b</sup>Zonal wind component

908 <sup>c</sup>Meridional wind component

909

910 Figure 1. Downtown area of the 3 km modelling domain (d03) showing the locations of  
911 measurement sites and WRF topography in the vicinity of SPMA. Red dots  
912 represent sites with information on O<sub>3</sub> and PM. Yellow dots represent only  
913 sites with information on PM. Blue dot represents the location of the IAG-  
914 USP's climatological station.

915 Figure 2. Emission rates for Aromatic VOCs at 19 UTC in the 3 km modelling  
916 domain.

917 Figure 3. Hourly accumulated precipitation and relative humidity observed at the  
918 IAG-USP's climatological station during the study period.

919 Figure 4. HYSPLIT three-day backward trajectories and locations of fires in Sao Paulo  
920 State and part of central-west region of Brazil. Pink markers represent the  
921 observed fire locations during the study period considering different satellite  
922 products (GOES, AQUA, TERRA, NOAA). The backward trajectories starting  
923 at IAG-USP were calculated for the days 9 and 31 August and 5 September 2012  
924 at three different altitudes: 500 m (red lines), 1000 m (blue lines), and 2000 m  
925 (green lines).

926 Figure 5. Daily (top), diurnal (bottom), and nocturnal (middle) mean concentrations for  
927 EC, OC, PM<sub>10</sub>, PM<sub>2.5-10</sub>, PM<sub>2.5</sub> (left panels), and Na, Fe<sub>2</sub>SO<sub>3</sub>, SiO<sub>2</sub>, K<sub>2</sub>O, and S  
928 (right panels). The PM<sub>2.5-10</sub> aerosol variable is defined as particulate matter with  
929 aerodynamic diameter between 2.5 and 10 μm. The grey line indicates the WHO  
930 air quality standard for PM<sub>2.5</sub> (25 μg m<sup>-3</sup>).

931 Figure 6. The predicted average of wind vectors at 10 m and temperature at 2 m from  
932 the baseline simulation (Case\_0) for the whole study period in the 3 km  
933 modelling domain. Blue dots represent the locations of the measurement sites,

934            whereas cyan numbers represent the observed average temperature in those  
935            sites: 17.7 °C in AF-IAG and 17.8 °C in INT.

936   Figure 7. The observed and predicted daily variations of PM<sub>2.5</sub> concentrations at three  
937            sites in SPMA for the 3 km modelling domain.

938   Figure 8. The observed and predicted daily variations of PM<sub>10</sub> concentrations at ten sites  
939            in SPMA for the 3 km modelling domain.

940   Figure 9. The observed and predicted hourly variations of O<sub>3</sub> concentrations at six sites  
941            in SPMA for the 3 km modelling domain.

942   Figure 10. Taylor diagram showing the individual correlation coefficients, biases, and  
943            normalized standard deviations for the PM<sub>2.5</sub>, PM<sub>10</sub>, and O<sub>3</sub> concentrations.

944   Figure 11. The predicted average surface distribution of PM<sub>2.5</sub> concentrations for the  
945            whole study period in the 3 km modelling domain. Red dots represent the  
946            locations of the measurement sites with information on PM<sub>2.5</sub>, whereas cyan  
947            numbers represent the observed average PM<sub>2.5</sub> concentration in those sites: 23.4  
948            μg m<sup>-3</sup> in IPEN-USP, 21.3 μg m<sup>-3</sup> in IAG-USP, and 22.2 μg m<sup>-3</sup> in CON.

949   Figure 12. The predicted average surface distribution of PM<sub>10</sub> concentrations for the  
950            whole study period in the 3 km modelling domain. Red dots represent the  
951            locations of the measurement sites with information on PM<sub>10</sub>, whereas cyan  
952            numbers represent the observed average PM<sub>10</sub> concentration in those sites: 49.5  
953            μg m<sup>-3</sup> in IAG-USP and 38.7 μg m<sup>-3</sup> in CON.

954   Figure 13. The predicted average surface distribution of the PM<sub>2.5</sub>:PM<sub>10</sub> ratio for the  
955            whole study period in the 3 km modelling domain. Red dots represent the  
956            locations of the measurement sites with information on both PM<sub>2.5</sub> and PM<sub>10</sub>,  
957            whereas cyan numbers represent the observed average PM<sub>2.5</sub>:PM<sub>10</sub> ratio in those  
958            sites: 0.43 in IAG-USP and 0.57 in CON.

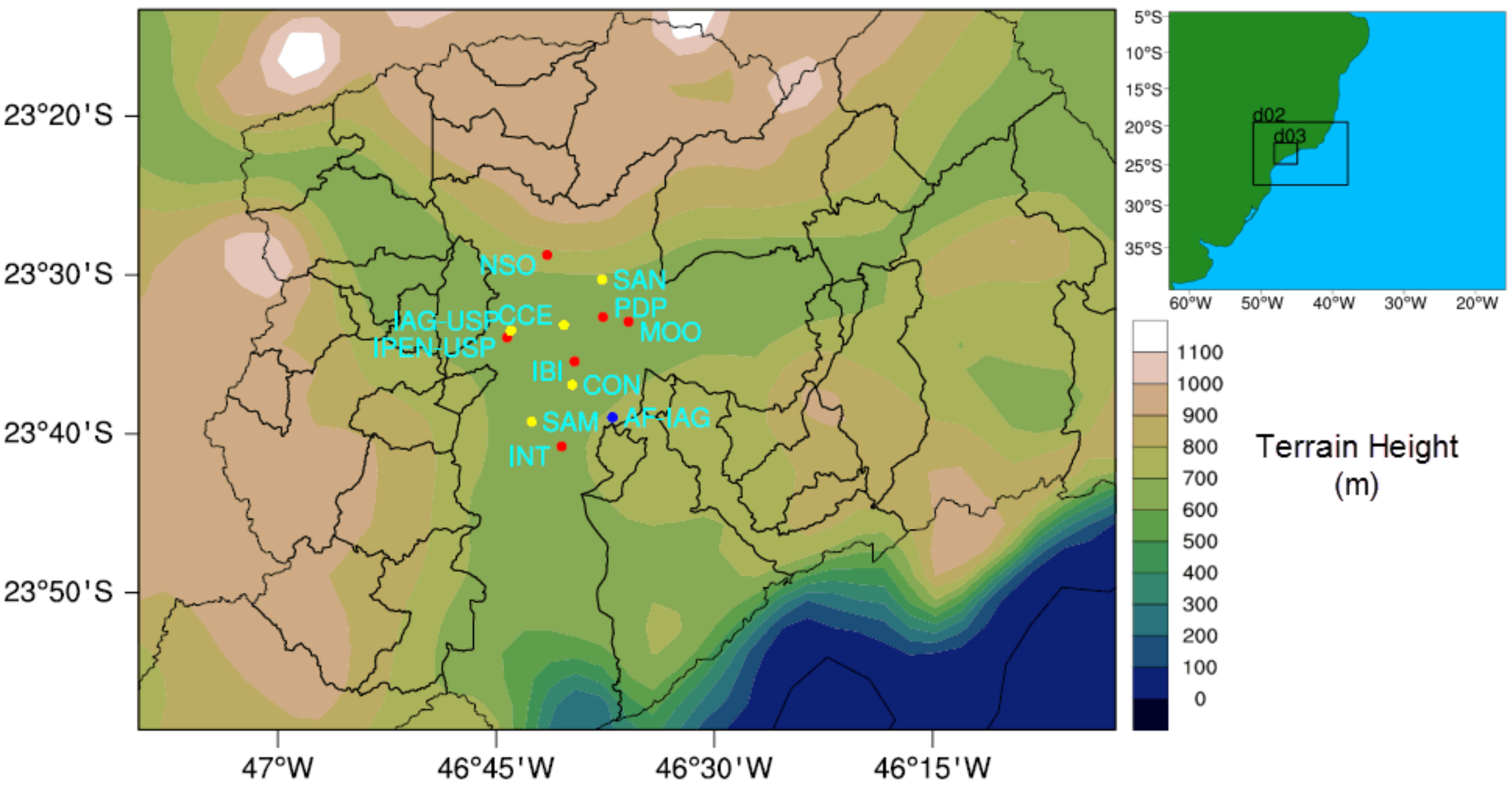
959 Figure 14. The observed and predicted daily variations of OC and EC concentrations  
960 at IAG-USP.

961 Figure 15. The observed and predicted average aerosol mass size distribution for SO<sub>4</sub>,  
962 NO<sub>3</sub>, NH<sub>4</sub>, Na, Cl, and other PM<sub>10</sub> constituents at IAG-USP. The observed  
963 aerosol distributions were collected in ten size classes using a rotated impactor  
964 (MOUDI) and joined adequately according to the three modes used by the  
965 MADE aerosol scheme: Aitken (<0.1 μm), accumulation (0.1-1 μm) and coarse  
966 (>1 μm). The five inorganic ions carried in MADE are only calculated for the  
967 Aitken and accumulation modes. The WRF's PM<sub>10</sub> aerosol variable does not  
968 include neither OC nor EC for this comparison.

969 Figure 16. The observed and predicted average contributions for the main identified  
970 constituents of PM<sub>2.5</sub> at IAG-USP.

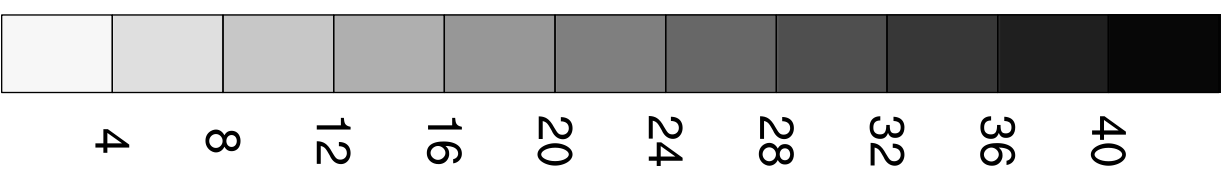
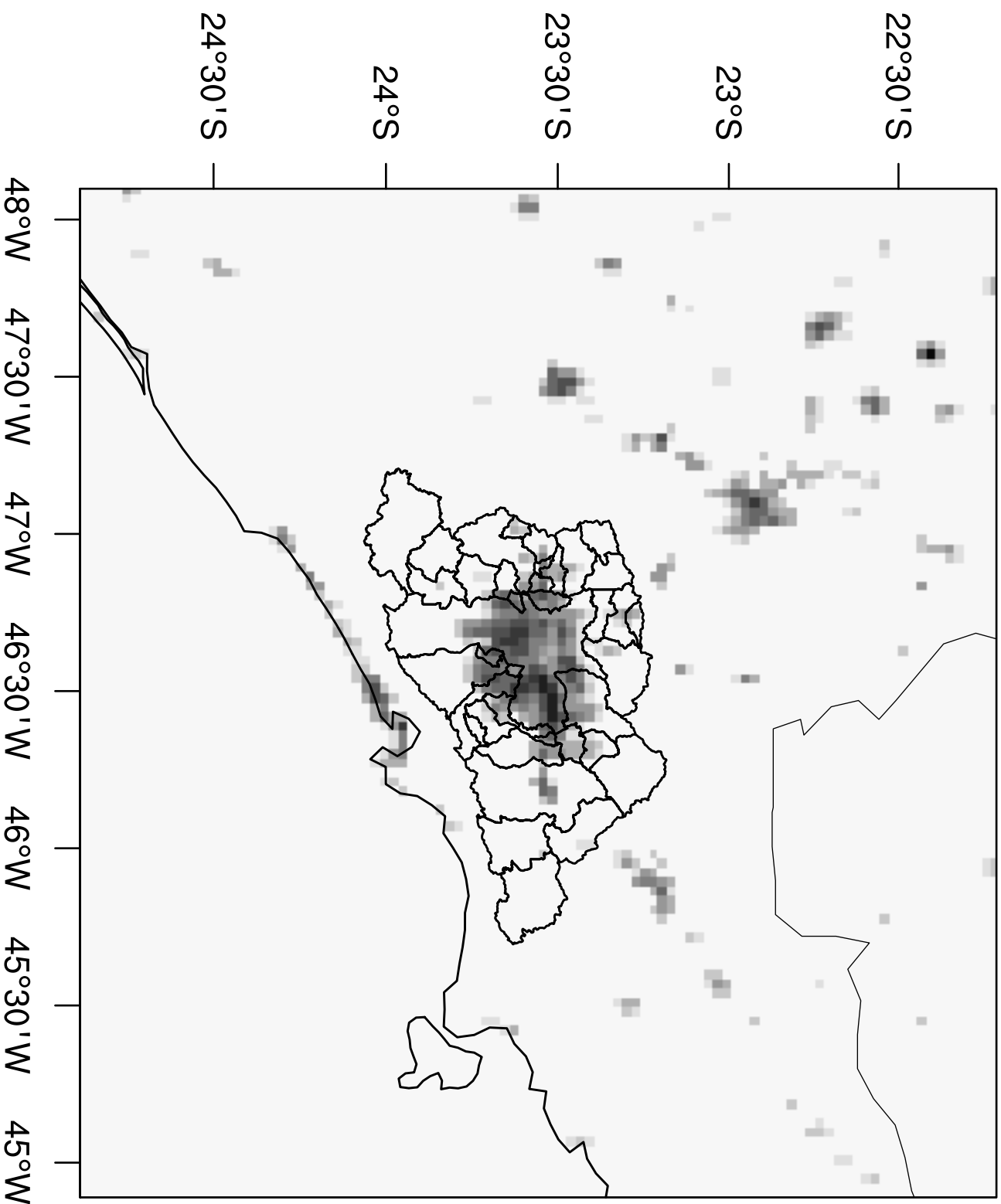
971 Figure 17. The impact of (a) emissions of primary gases on the fine particles formation,  
972 (b) emissions of dust-sea salt aerosols on the PM<sub>10</sub> concentration, and (c) aerosol  
973 direct effect on the ground level O<sub>3</sub> concentrations at 16:00 h (local time).

974

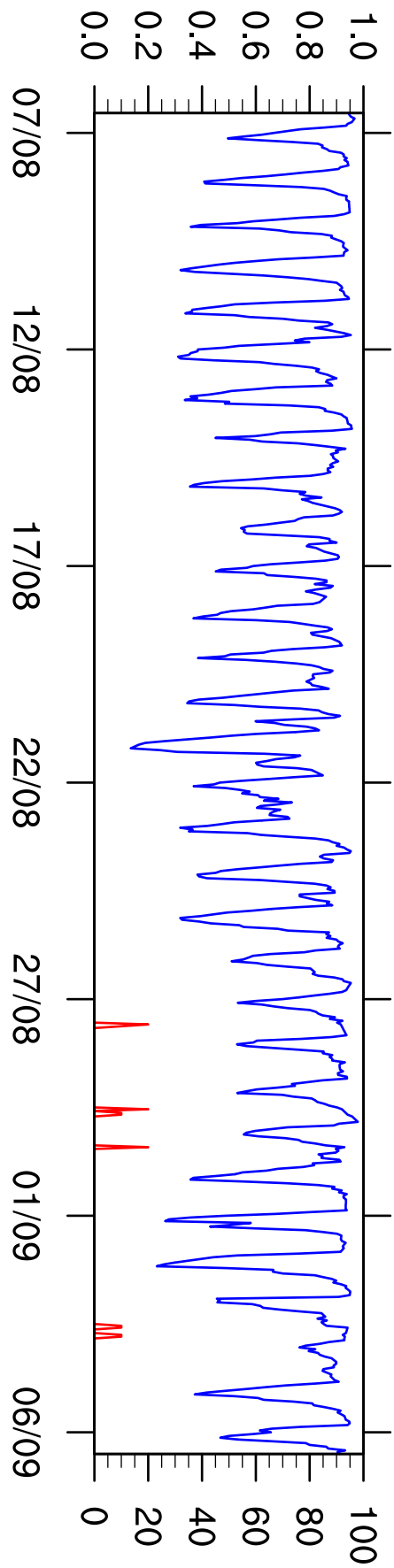


19 UTC

E\_TOL (mol/km<sup>2</sup>/hr)



Precipitation (mm) [red]

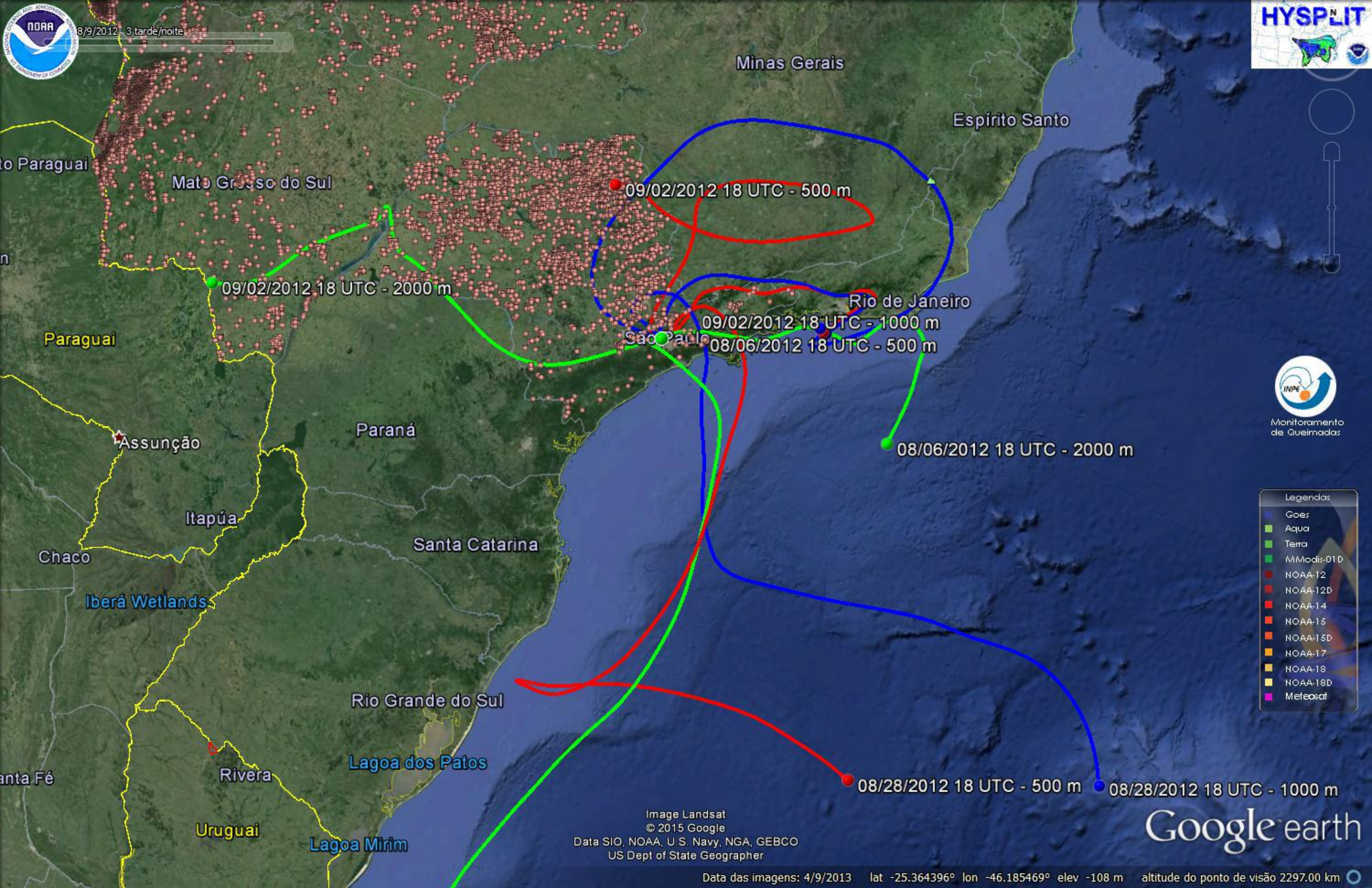


Humidity (%) [blue]





8/9/2012 3 tarde/noite



09/02/2012 18 UTC - 500 m

09/02/2012 18 UTC - 2000 m

09/02/2012 18 UTC - 1000 m

08/06/2012 18 UTC - 500 m

08/06/2012 18 UTC - 2000 m

08/28/2012 18 UTC - 500 m

08/28/2012 18 UTC - 1000 m

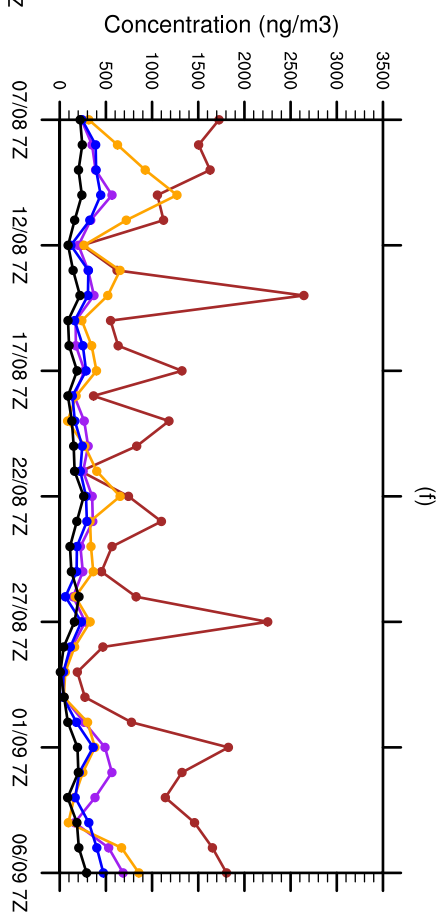
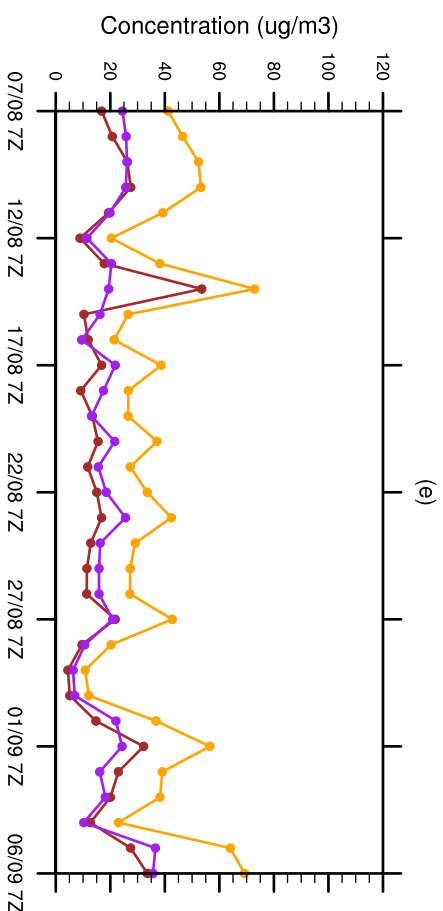
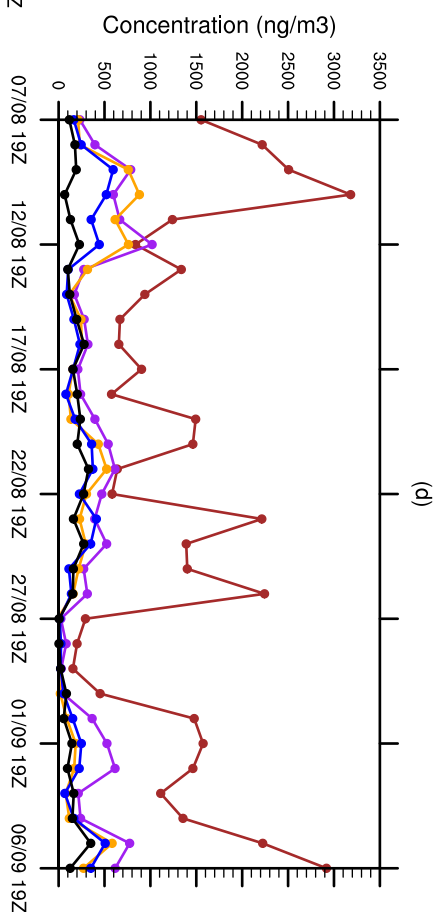
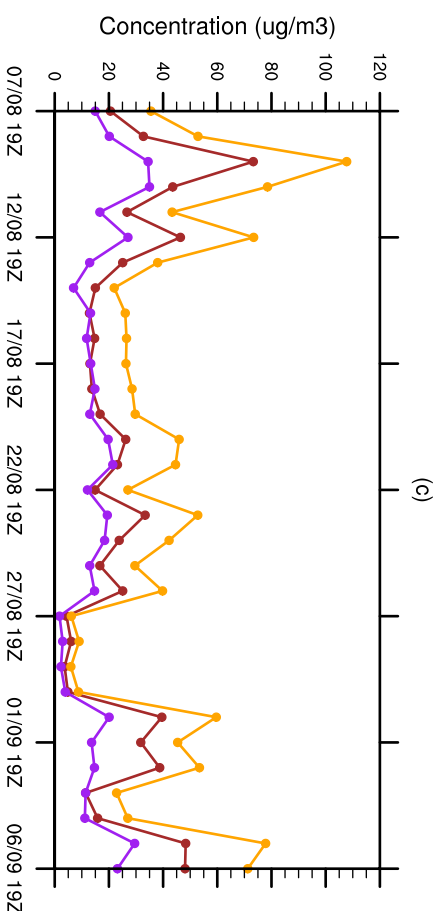
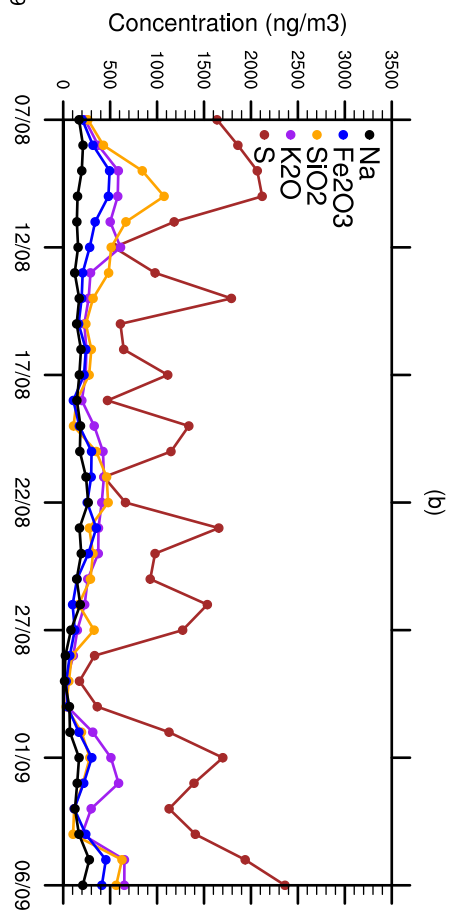
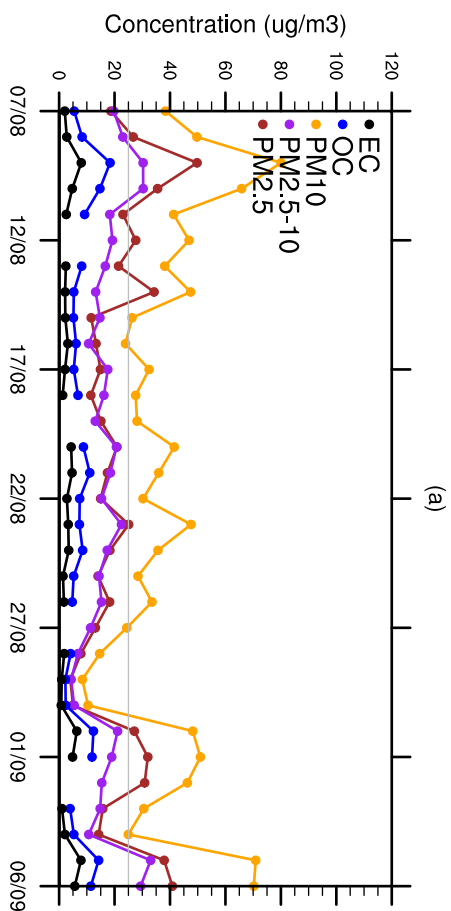
- Legendas
- Goes
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  - NOAA-12D
  - NOAA-14
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  - NOAA-15D
  - NOAA-17
  - NOAA-18
  - NOAA-18D
  - Meteosat



Monitoramento de Queimadas

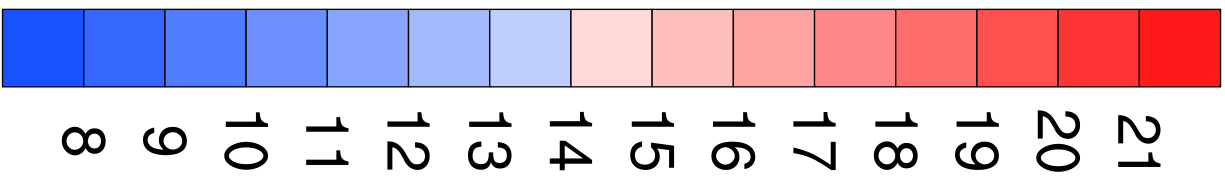
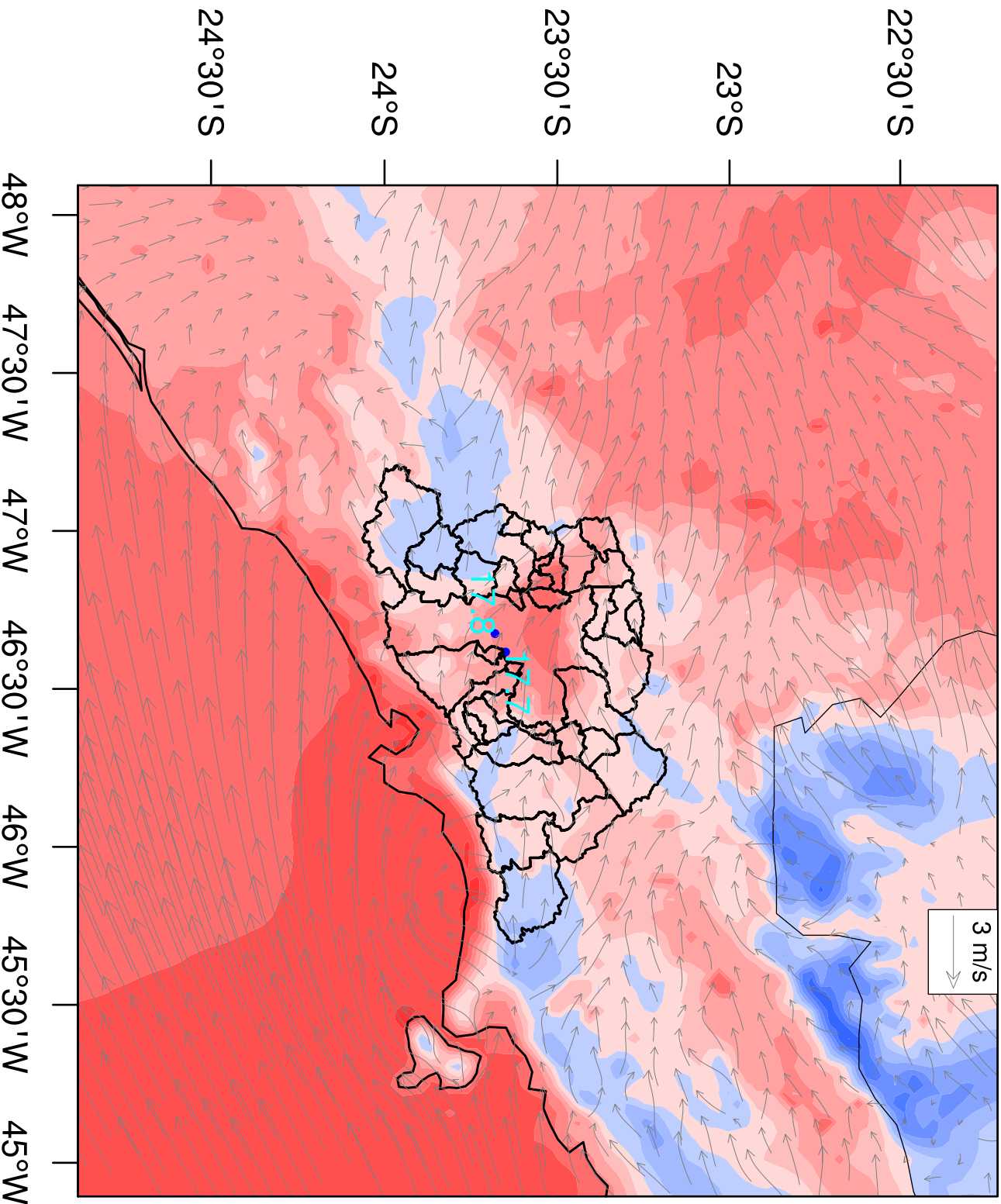
Image Landsat  
© 2015 Google  
Data SIO, NOAA, U.S. Navy, NGA, GEBCO  
US Dept of State Geographer

Google earth

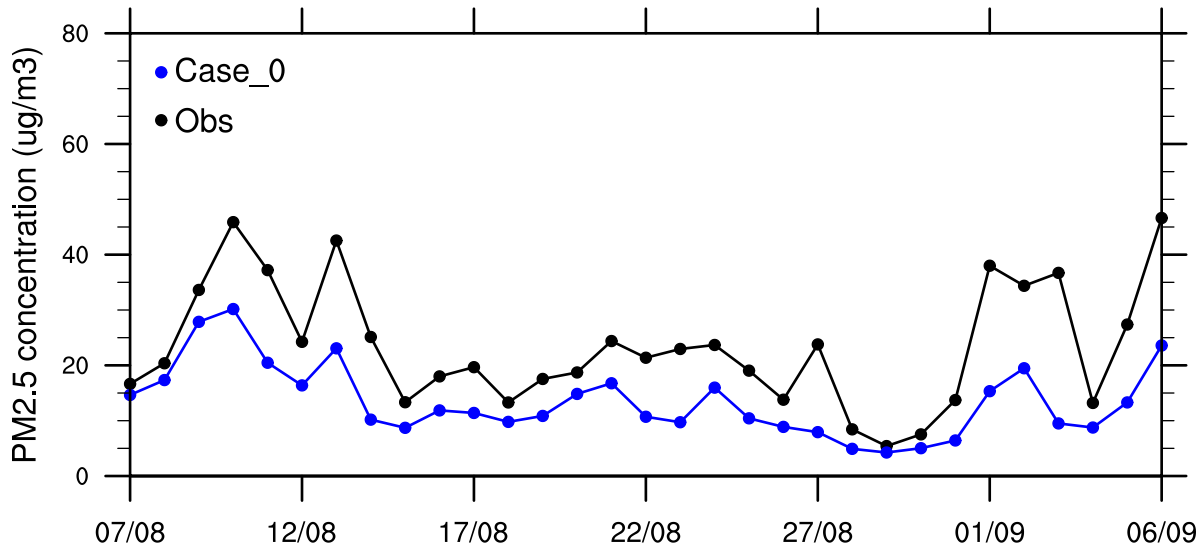


Temperature at 2m (Case\_0)

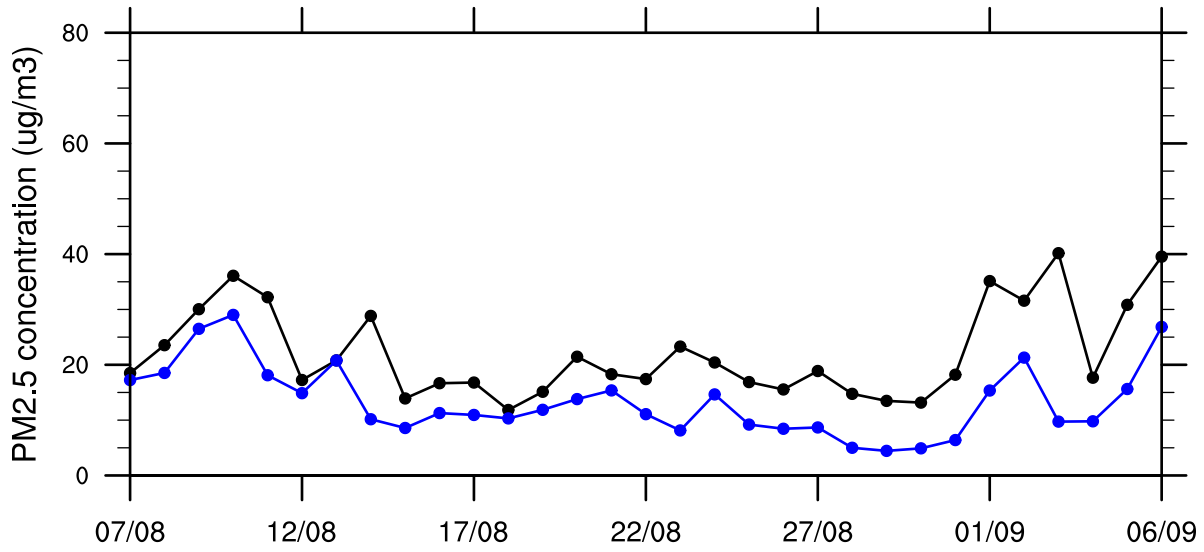
degrees C



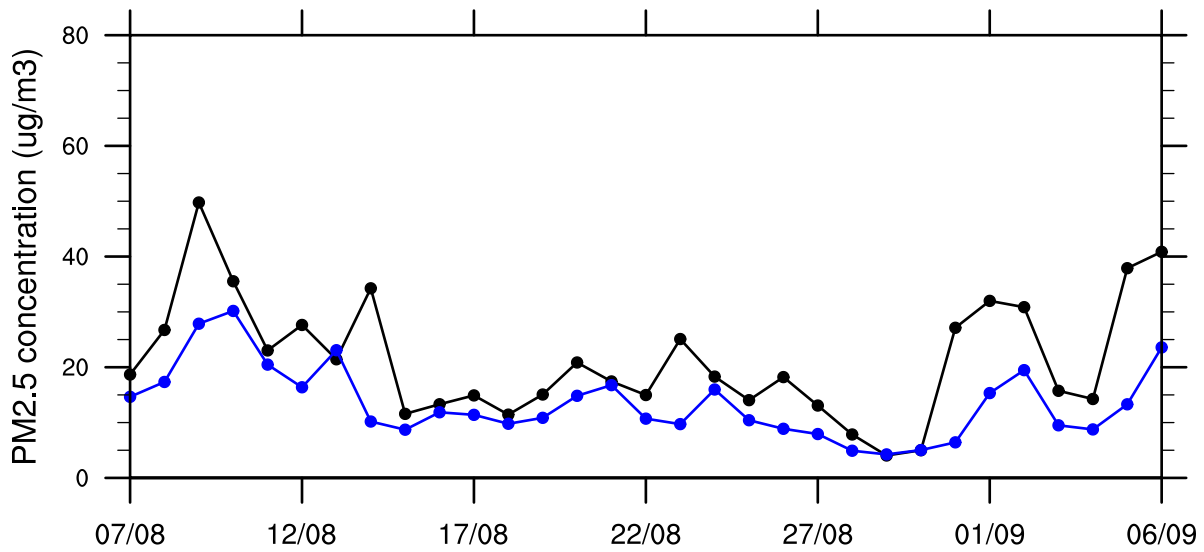
### IPEN-USP

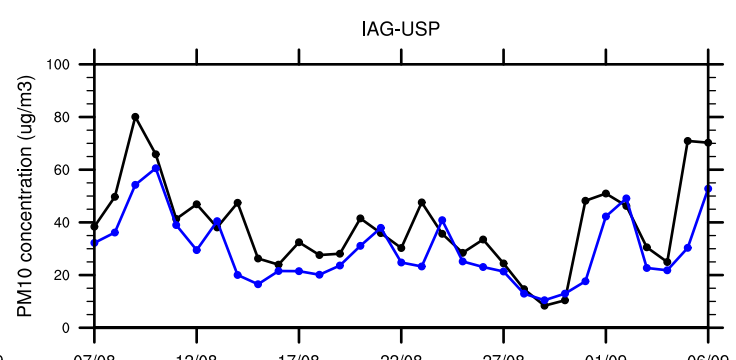
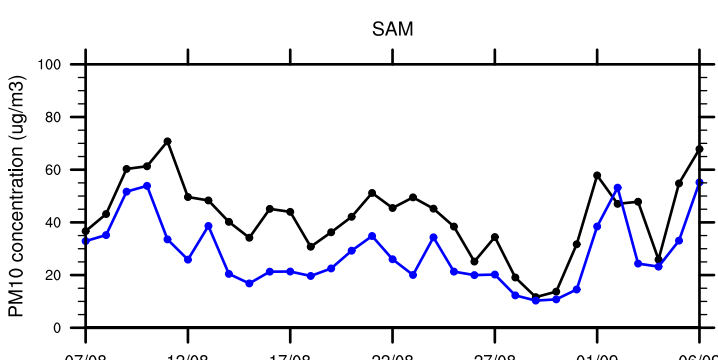
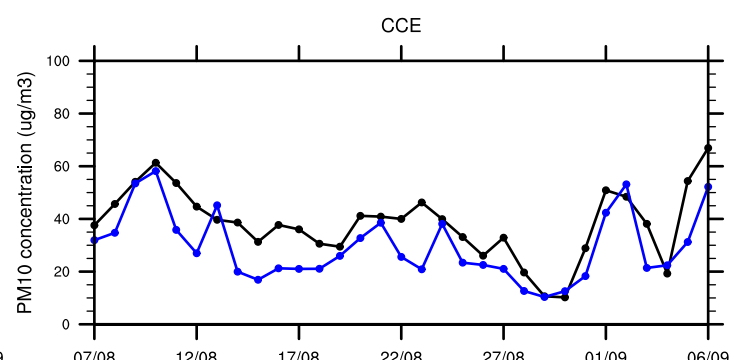
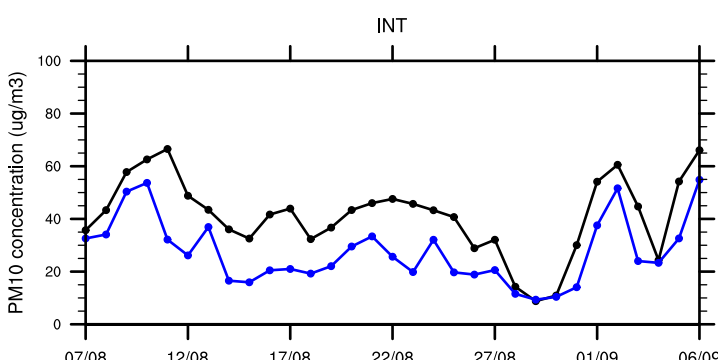
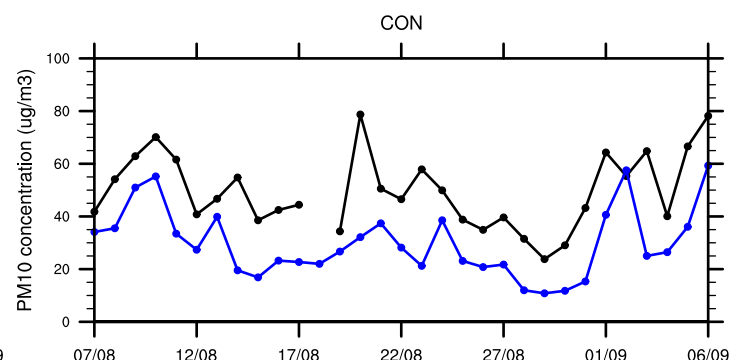
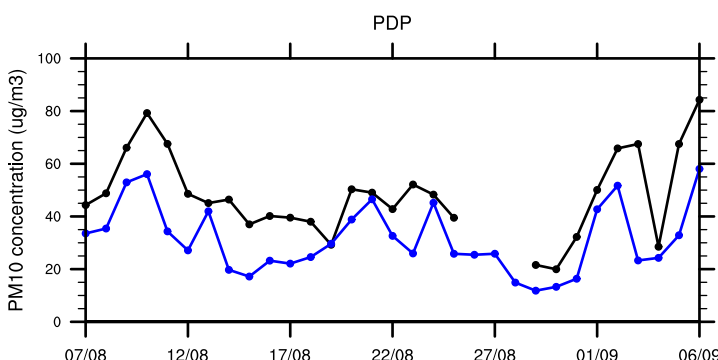
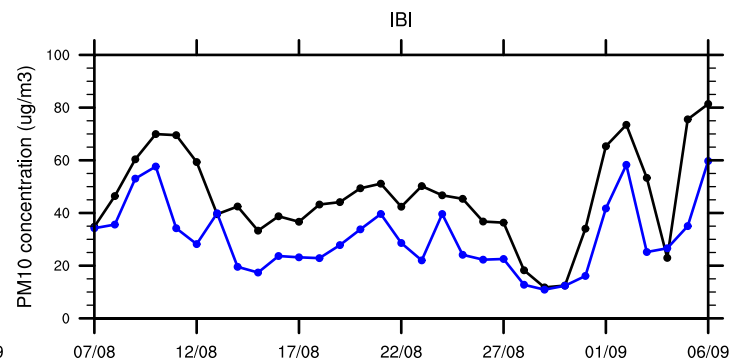
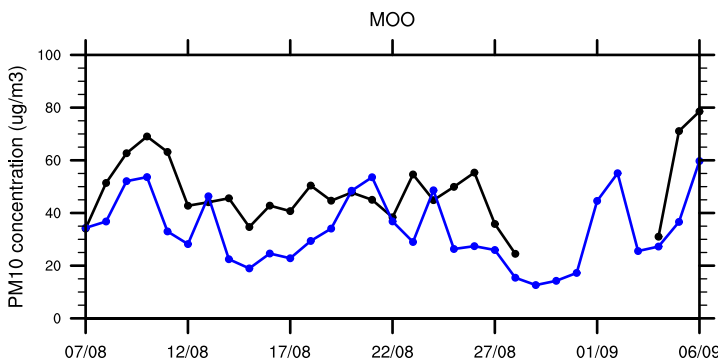
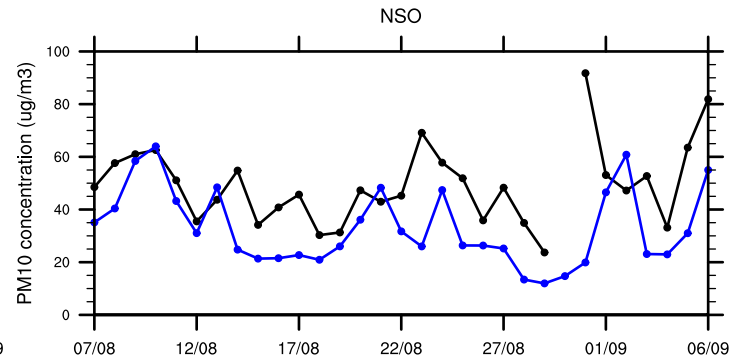
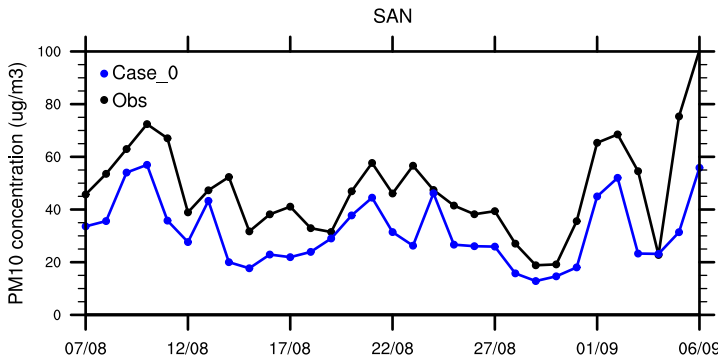


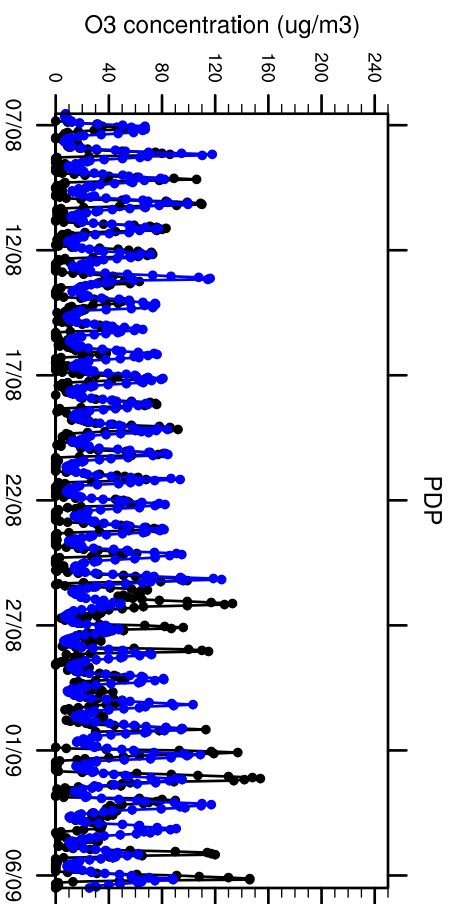
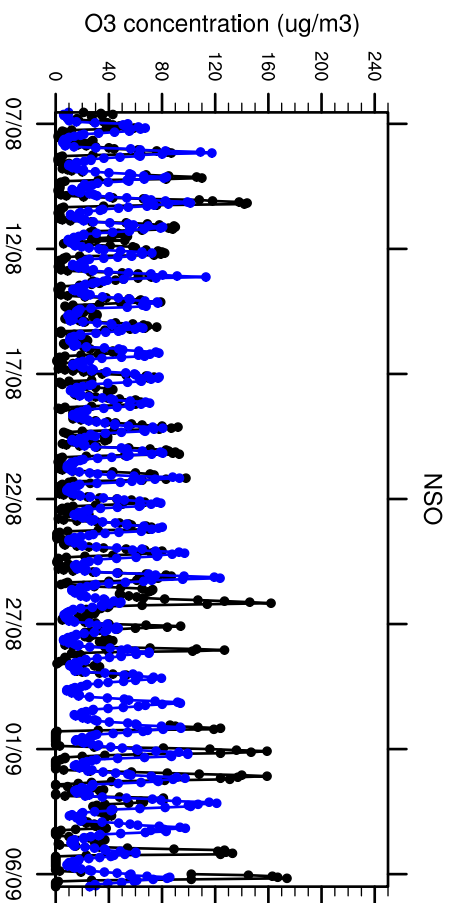
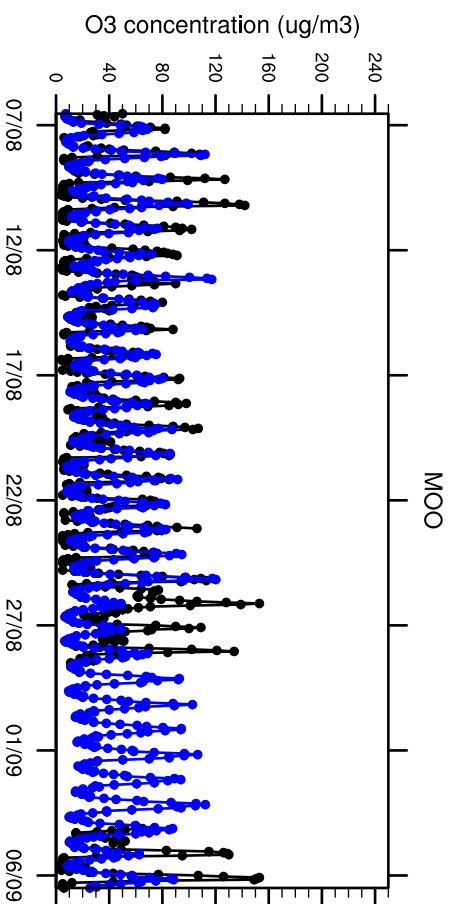
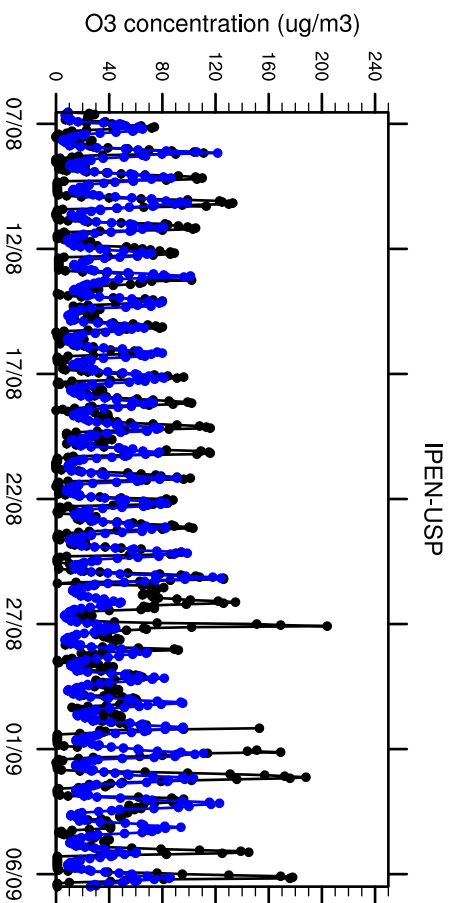
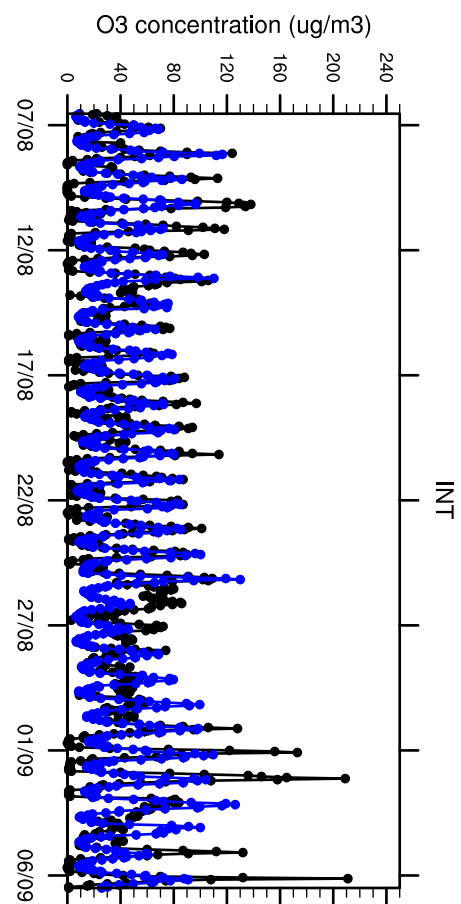
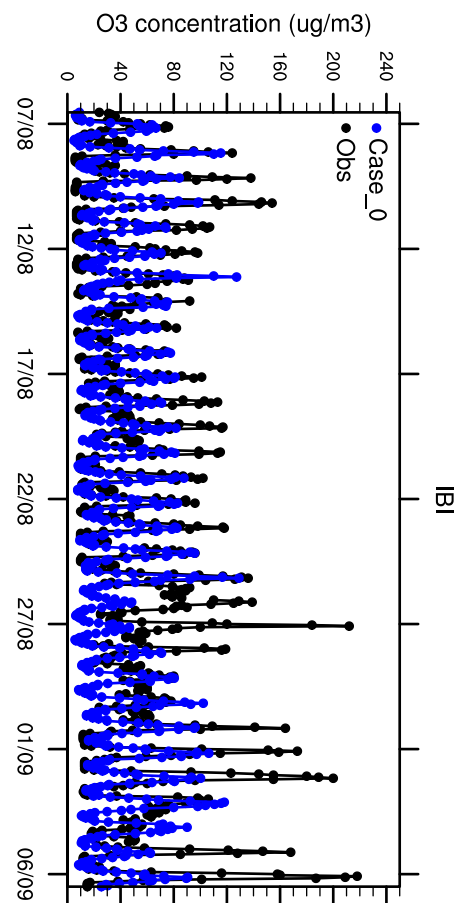
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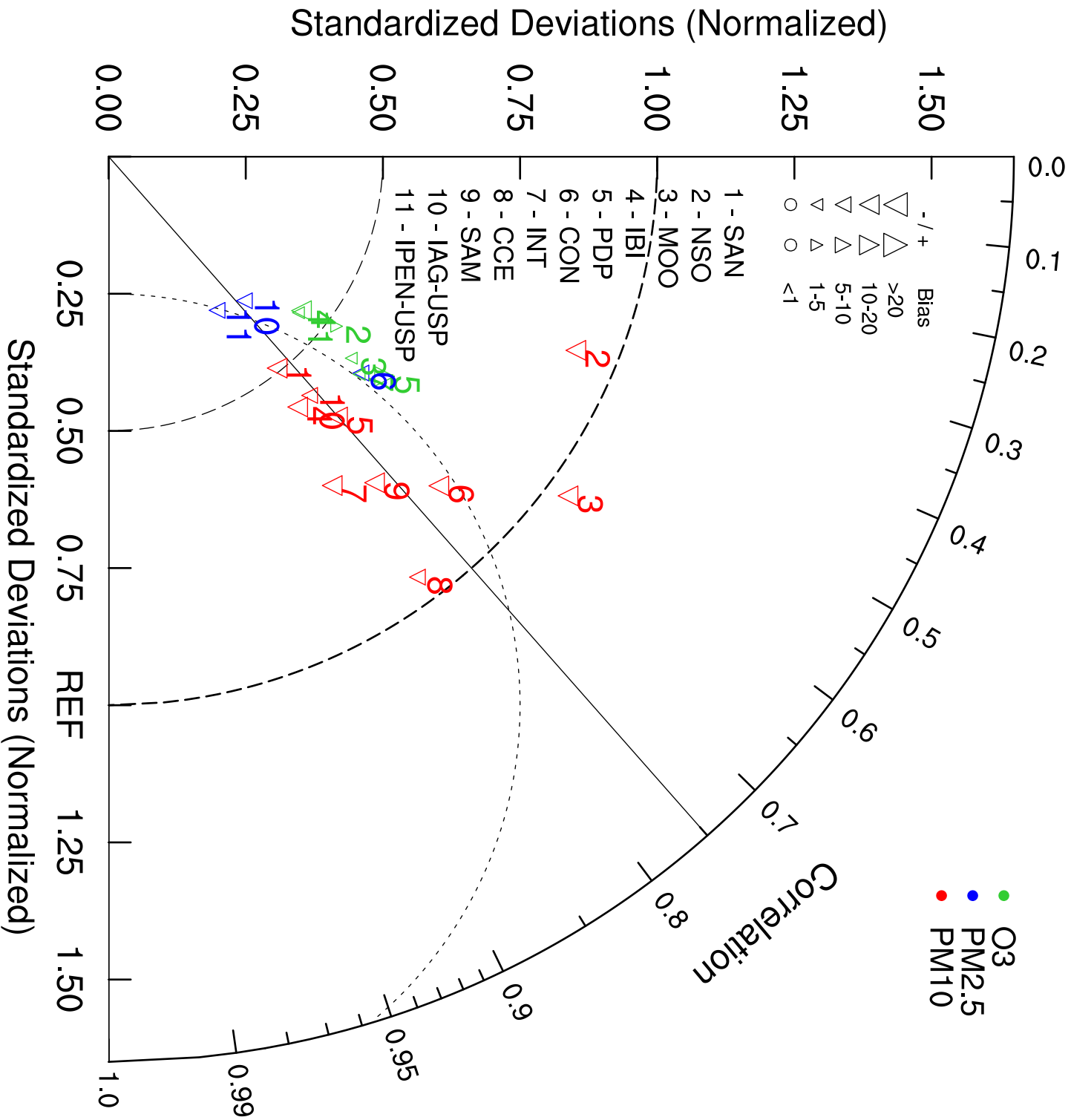


### IAG-USP



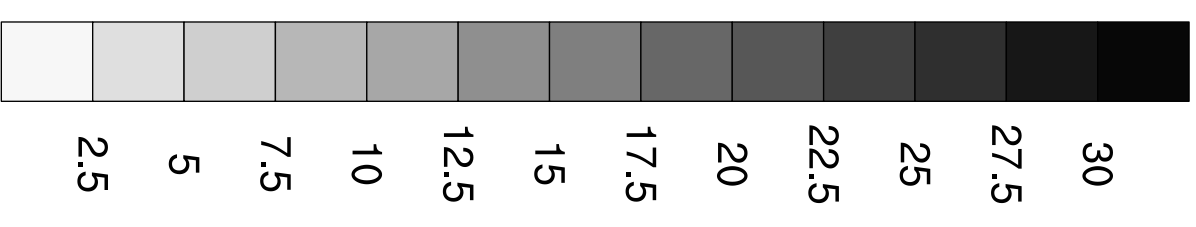
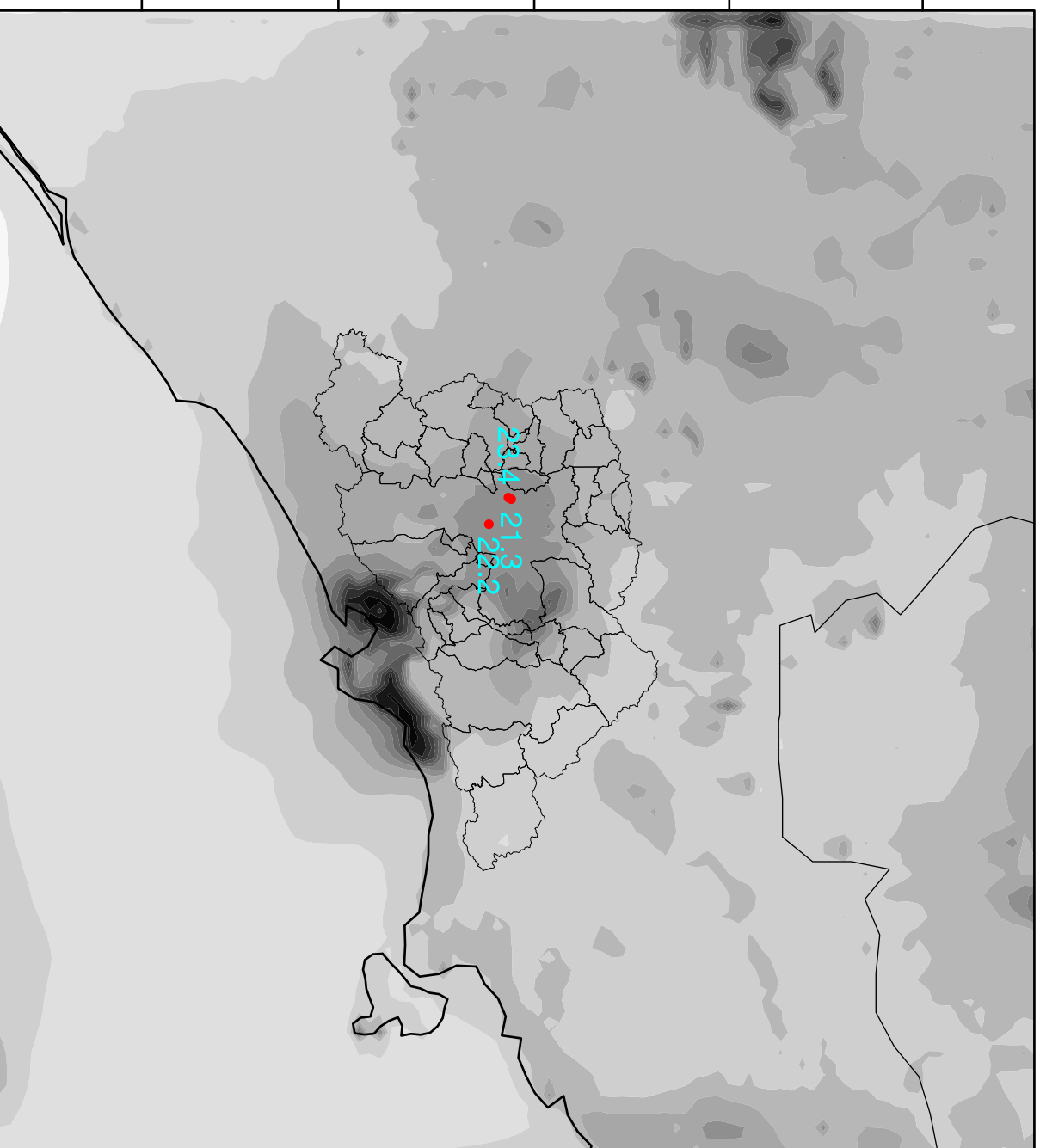






PM2.5 (Case\_0)

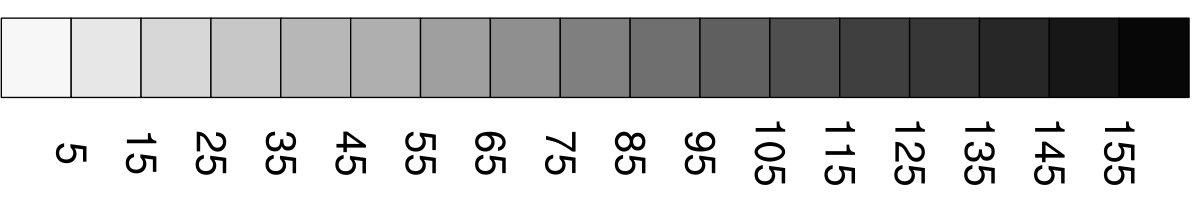
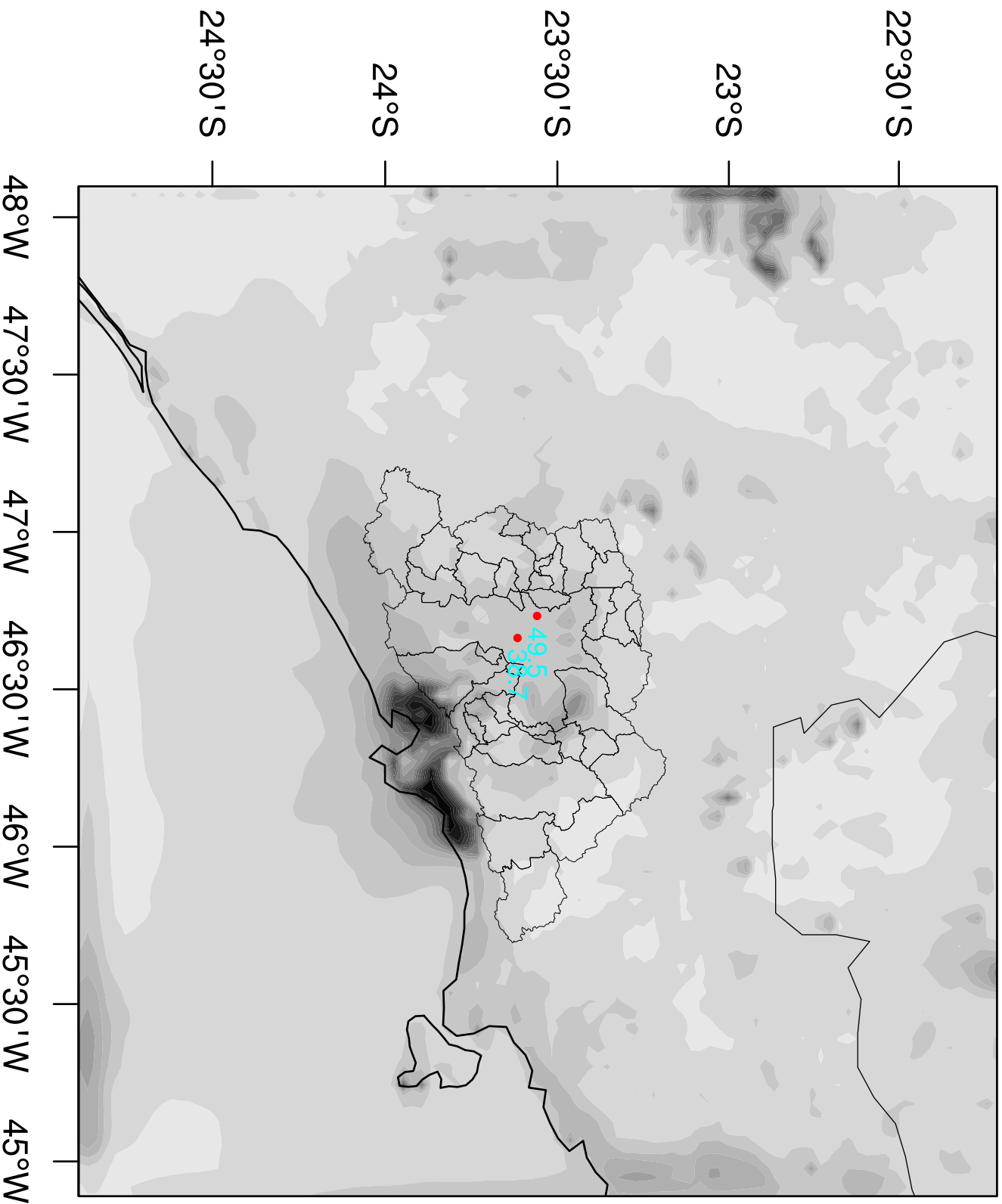
ug/m3





# PM10 (Case\_0)

ug/m3



# PM2.5 / PM10 (Case\_0)

