

Interactive comment on “An estimation of the $^{18}\text{O} / ^{16}\text{O}$ ratio of UT/LMS ozone based on artefact CO in air sampled during CARIBIC flights” by S. Gromov and C. A. M. Brenninkmeijer

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The authors have observed CO increase in the samples taken in UT/LMS and stored in stainless steel canisters. Elevated CO correlates with elevated ozone and also demonstrates an increase in $\text{d}^{18}\text{O}(\text{CO})$. Based on that oxidation of hydrocarbons (possibly methane) by ozone in canisters, being catalysed by stainless steel is proposed as the most likely mechanism. Finally the authors discuss that elevated CO has (maybe indirectly) recorded d^{18}O signals of ozone in UT/LMS. Notably, the authors write: “...the molecular lifetime (the period through which the species’ isotope reservoir becomes entirely renewed, as opposed to the “bulk” lifetime) of O_3 encountered along the C1

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flight routes is estimated on the order of minutes to hours at daylight (H. Riede, MPI-C, 2010), thus the isotope composition of the photochemically regenerated O_3 resets quickly according to the local conditions.” Thus the d^{18}O values observed reflect not the original signature of ozone from the stratosphere but the ozone being photochemically recycled in UT/LMS. This is very interesting scientific finding.

Given that, I would suggest giving some supporting references and/or more detailed info (or discussion) on the reactions involved. Second, can the authors speculate on the dependence of d^{18}O signals of the recycled ozone on the local conditions in UT/LMS? How large variations could be expected? Third, is there any information on ozone reactions catalysed by stainless steel?

As a technical comment – I suggest to address the method of $\text{d}^{18}\text{O}(\text{CO})$ measurements and how the d^{18}O signal of CO is preserved. One can give a few lines supported by references.

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