

variations in atmospheric aerosols because of meteorological and climatic factors (Reid et al., 2012). The trans-boundary and long-range transport of aerosols interact with their local counterparts (e.g. cloud droplets), enhance the microphysical properties of aerosols, and affect their radiative properties and precipitation processes (Ichoku et al., 2004; Lin et al., 2013; Rosenfeld, 2007; Andreae and Rosenfeld, 2008). The global effects of aerosols on the Earth's climate are hardly quantifiable because of the lack of extensive and reliable measurements in most world regions (Tripathi et al., 2005; Russell et al., 2010; Hansen et al., 1997; Kaskaoutis and Kambezidis, 2008; Kaskaoutis et al., 2007).

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Aerosol optical depth (AOD) conveniently analyzes air quality/pollution, radiation budget and radiation forcing, climate change, atmospheric corrections in remote sensing from space, and aerosol characteristics. The spatial and temporal variations in AOD are large because of production sources, transport and removal processes, and prevalent meteorological conditions. Given the large uncertainty in aerosol characterization, local analyses essentially verify the satellite imagines because the extraction of aerosol optical properties from remote sensing data exhibits limited accuracy despite its capability to provide global-scale coverage (Levy et al., 2005; Tripathi et al., 2005; Yoram et al., 2002; Gupta et al., 2013; Zhong et al., 2007). Local studies on the optical properties of aerosols have been conducted using sun photometers and sky radiometers (Salinas et al., 2009; Holben et al., 1998; Remer et al., 2008). However, these methods are limited to space coverage in contrast to satellite imagery. Therefore, ground- and space-based measurements complementarily perform reliable and comprehensive studies on atmospheric aerosols.

The accuracy of satellite-derived daily AOD is often assessed by comparing satellite-based AOD with the AEROSOL ROBOTIC NETWORK (AERONET), a network of ground-based sun photometers. AERONET is widely used to monitor, investigate, and characterize the optical properties of aerosols (Holben et al., 1998). This network provides a database to atmospherically correct and validate satellite-based aerosol retrievals. However, cloud-contaminated data should be removed from the AERONET database

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(Smirnov et al., 2000); the process is termed as cloud screening. Hence, only a limited dataset of level 2 AOD (data have been cloud screened and quality assured) can be obtained. Meanwhile, AODs obtained from satellites, such as those from MODIS (Retalis et al., 2010), are limited because these satellites are orbiting. Continuous retrieval of AOD data is difficult. Thus, several models have been proposed to efficiently predict and retrieve AOD. REFERENCES

are available
 in sun-synchronous orbit

Previous studies have used single parameters from ground measurements to estimate the atmospheric columnar AOD, such as in situ horizontal visibility (Vis) or particulate matter (PM) with diameters less than 10 or 2.5 μm (PM_{10} or $\text{PM}_{2.5}$). The high concentrations of atmospheric aerosols increase the AOD to effectively scatter light and reduce Vis. PM_{10} and $\text{PM}_{2.5}$ are used to physically quantify the concentration of PM at ground level. High-quantity PM records imply high aerosol concentrations at the ground surface. AOD is proportional to air quality (Müller et al., 2012; Cordero et al., 2012; Mogo et al., 2012; Mielonen et al., 2012; Wang and Christopher, 2003) but inversely proportional to Vis (Horvath, 1995; Bäumer et al., 2008; Li and Lu, 1997; Peppler et al., 2000; Singh and Dey, 2012). Vis and air quality interact with columnar AOD; hence, these parameters should be considered into the algorithm to predict AOD through multiple regression analysis. The complementary combination increases the relative accuracy of prediction.

REFS?

Three types of measurement data were used in this study, namely (i) AOD, (ii) Vis and (iii) air pollution index (API). The AOD measurements were obtained through the AERONET site located in Universiti Sains Malaysia (USM). The Vis and API data were taken from the meteorological stations at the Penang international airport and USM. All data were taken between 2012 and 2013. The aerosol characteristics in Penang were comprehensively analyzed based on changes in seasonal monsoons. A near real-time AOD model was established based on multiple regression analysis. The accuracy and efficiency of the model were validated and evaluated to assess the atmospheric pollution in Penang.

Toth et al. 2014
 State clearly what you are doing and why

lat/lon coordinates

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June–September, and (iv) October–November. The fifth or “overall” set comprised the annual data. The number of data points for December–March, April–May, June–September, and October–November were 257, 132, 235, and 166, respectively. The data for each seasonal monsoon were further divided into two subsets that were sourced from alternatively selected data (in temporal sequence) for cross-validation. For example, consider that data with a particular seasonal monsoon period takes a sequential form (D1, D2, D3, D4, D5, . . .). Thus, the subsets are in the form of (D1, D3, D5, . . .) and (D2, D4, D6, . . .). The first data subset was used to determine the correlation between the parameters and AOD at 500 nm (Eq. 1), which was the original model of Tan et al. (2014a), and given as follows:

$$\begin{aligned} \text{AOD} = & a_0 + a_1(\text{RH}) + a_2(\text{RH})^2 + a_3(\text{RH})^3 + a_4(\text{Vis}) + a_5(\text{Vis})^2 + a_6(\text{Vis})^3 \\ & + a_7(\text{API}) + a_8(\text{API})^2 + a_9(\text{API})^3 \end{aligned} \quad (1)$$

where RH is the relative humidity, the second data subset was used to predict AOD in each seasonal monsoon and validate the accuracy of the prediction based on the parameters (e.g., a_0 and a_1) obtained from the correlation procedure. The algorithm of Tan et al. (2014a) was tested to determine the correlation at 95% confidence level for each seasonal monsoon. The root mean square error (RMSE), coefficient of determination (R^2), and percent mean relative error (%MRE) between the measured and predicted AOD for each period were calculated. The %MRE parameter was used to quantify the systematic differences between the concentration levels. This parameter is given as follows: %MRE = [(mean predicted AOD – mean measured AOD)/mean measured AOD] × 100. The ability of the proposed model to produce reliable AOD estimates for temporal air monitoring can be quantitatively justified or falsified based on the quality of the resultant %MRE.

Aerosols could be hydrophilic or hydrophobic, and these properties could give rise to non-trivial contribution to AOD retrieval (Ramachandran and Srivastava, 2013; Singh and Dey, 2012; de Meij et al., 2012; Tang, 1996; Song et al., 2007; van Beelen et al., 2014; Wang et al., 2013). However, to discriminate whether the aerosols are hydrophilic

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or hydrophobic requires addition resources beyond the reach of the present study. On the other hand, our pre-analysis showed that RH does not contribute significantly to AOD prediction in the proposed model. If RH was considered as a predictor, its related factors (e.g., aerosol stratification (dust or smoke aloft), convection, and hysteresis in particles) should be taken into account. The contribution of RH to the aerosol properties was integrated in the aerosol model (Srivastava et al., 2012) because the net effect of RH on aerosol and related factors were hardly quantifiable. The RH contribution can be disregarded in the present model, yielding Eq. (2). The results were obtained from the correlation analysis based on Eq. (2) given as follows:

$$\text{AOD} = a_0 + a_1(\text{Vis}) + a_2(\text{Vis})^2 + a_3(\text{Vis})^3 + a_4(\text{API}) + a_5(\text{API})^2 + a_6(\text{API})^3 \quad (2)$$

Lee et al. (2012) excluded the days when the deviation between the measured and predicted values was greater than RMSE, or when the estimated AOD slope was negative because of measurement errors and cloud-contaminated AOD. Given the previous findings, the outliers in our model were removed using the approach of (Lee et al., 2012). The predicted AOD was compared with the measured counterpart from AERONET to determine the accuracy of the generated model. Equation (2) was applied to retrieve the AOD for specific days when no AOD values were available. The features of predicted AOD were compared against those of the measured counterpart. The under- and overpredicted AOD were examined by RAYMETRICS LIDAR system. However, examination can only be performed when LIDAR data were available. When LIDAR data were available for examination, only the data that can clearly elucidate the under- and over-predicted AOD were selected. The backscatter coefficients of the aerosol were determined using the method of Fernald (1984). The LIDAR signals were pre-analyzed based on the published works of Tan et al. (2013, 2014c).

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* See Omar et al. 2005 (JGR)

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observed during the northeast monsoon period. Almost no frequency data were obtained for PW < 3.5, except the northeast monsoon period with about 14% less than this value. The most humid period took place in April–May, with PW ranging from 5.0 to 5.5 (approximately 74% of the total occurrence).

5.3 Seasonal discrimination of aerosol types based on the relationship between AOD and Angstrom exponent

Aerosols have been widely classified by the scatter plots of AOD and Angstrom exponent. AOD provides information of aerosol loading in the atmosphere column through the extinction of radiation rate for a specific wavelength. The Angstrom exponent determines the aerosol size in coarse and fine modes from the slope with wavelengths that depend on AOD in logarithmic coordinates. Therefore, the AOD–Angstrom exponent scatter plots indicate the amount and dimension of the observed aerosols. The corresponding distribution pattern was grouped into a few clusters to determine the aerosol species. Related studies have been analyzed using AERONET data; these datasets have been applied at different locations, such as the Persian Gulf (Smirnov et al., 2002a); Brazil, Italy, Nauru, and Saudi Arabia (Kaskaoutis et al., 2007); Spain (Toledano et al., 2007); Singapore (Salinas et al., 2009); several oceanic regions (Smirnov et al., 2011); Kuching (Jalal et al., 2012); and the Multi-filter Rotating Shadowband Radiometer in Central Mediterranean (Pace et al., 2006). The scatter plot of AOD₅₀₀ or AOD₄₄₀ against Angstrom_{440–870} was used to identify the aerosol type. The wavelength range of Angstrom_{440–870} was used because of its nearness to the typical size range of aerosol based on spectral AOD (Eck et al., 1999). The relation between AOD values at 500 nm and Angstrom 440–870 is usually used for aerosol classification in scatter plot diagram. The AOD values at 500 nm are normally used to indicate the turbidity conditions (Cachorro et al., 2001; Smirnov et al., 2002b, 2003; Kaskaoutis et al., 2007; Pace et al., 2006; Salinas et al., 2009). Optically, 500 nm is an effective visible wavelength suitable for aerosol study (Stone, 2002). In this study,

you've said all of this?

AOD₄₄₀–Angstrom_{440–870} and AOD₅₀₀–Angstrom_{440–870} plots were used to classify the aerosols.

The aerosols were classified into five types, namely dust, maritime, continental/urban/industrial, biomass burning, and mixed aerosols (Ichoku et al., 2004); mixed aerosols in practice represents indistinguishable aerosol type that cannot be categorized into any of the previous types. To effectively identify the aerosol distribution types in our study sites, the results were compared using different threshold criteria (Table 2). The results of aerosol classification using different threshold criteria are presented in Fig. 2. The thresholds proposed by Kaskaoutis et al. (2007) and Pace et al. (2006) failed to determine the maritime aerosol (MA) and dust aerosol (DA) for each season. Instead, they showed that mixed-type aerosols (MIXA) were dominant in Penang (50–72%). Urban and industrial (UIA) and biomass burning (BMA) aerosols were grouped into a single class (28–50% of the total occurrence). Meanwhile, the threshold suggested by Smirnov et al. (2002, 2003, 2011) failed to identify DA, UIA, and BMA, but efficiently identified MA. As a result, a large amount of MIXA was obtained (> 80% of the total occurrence). These results reveal the extent of uncertainty; the indistinguishable aerosol types in the study sites were large. Thus, other options should be considered.

order of references

Salinas et al. (2009) suggested that the determination of DA and BMA did not correspond entirely to the range of threshold used in our study, in which the amount of MIXA (approximately 43% of the total occurrence) was large. Jalal et al. (2012) efficiently identified the aerosol types using an alternative threshold criterion. Using their threshold, we yielded a low amount of MIXA, approximately 21%. However, the determination of DA was unsatisfactory. The threshold criteria of Toledano et al. (2007) provided the least MIXA (< 5%; Fig. 2). All thresholds consistently increased from June to September (Fig. 2c) and coincided with the occurrence of haze. UIA was constantly and highly distributed over Penang. Overall, the thresholds provided by Toledano et al. (2007) were properly suited for our study site to determine the aerosol types.

best

The thresholds of AOD–Angstrom_{440–870} scatter plots by Toledano et al. (2007) used to classify the aerosol types revealed that higher amount of pollutants in UIA class were identified, and directly affected the air quality in Malaysia (Fig. 3). The MA observed in Penang was high because of its geolocation (i.e., surrounded by the sea). The study site was minimally affected by coarse particles and DA, which were less than 5% in each seasonal monsoon. BMA was one of the major pollutants in Penang because of the active burning activities. Furthermore, haze occurred during the southwest monsoon because of the trans-boundary aerosols from Indonesia. These results were in accordance with the records from DOE (2010). BMA, UIA, and MA obtained in our study during the southwest monsoon were about 45, 24, and 19%, respectively. During the northeast monsoon period, UIA (approximately 38%) was the major aerosol in Penang, followed by MA (30%), BMA (20%), dust (4%), and unidentified substances (8%). However, MIXA reached 17% from April to May, which was the highest among the seasonal monsoons. MA and UIA were 38%; the MA level was significant from October to November (51%), followed by UIA (40%) and BMA (< 1%). The aerosol distribution in Penang was highly season dependent.

where?

haze is STILL around!

3.4 Seasonal flow patterns of air parcel from the HYSPLIT_4 model for identification of aerosol origins

Given the seven-day seasonal plot of the back-trajectory frequency by the HYSPLIT_4 model, the flow patterns of the air parcel in Penang site were obtained (Fig. 4) for each monsoon season in terms of percentage averaged between the ground surface up to an altitude of 5000 m. Residence time analysis was performed to generate the frequency plot and determine the time percentage of a specific air parcel in a horizontal grid cell across the domain.

During the northeast monsoon period, the air parcel flowed southwestward from the northern part of southeast Asia (Fig. 4a), which illustrated that the aerosol sources to Penang were from the former (open burning season, Lin et al., 2013), including Indochina, and transported through South China Sea to reach Penang. The aerosols

during the northeast monsoon period were also locally produced, whereas those obtained during southwest monsoon period were from Andaman Sea, Malacca Strait, Sumatra (site of open active burning), and other local areas.

The patterns in seasonal relative frequency of air parcel movement were significantly different (Fig. 4a). Comparison with Fig. 1b indicated the differences in the patterns of the seasonal relative frequency of occurrence for Angstrom_{440–870} during the northeast monsoon. These differences were attributed to the mixing of various aerosol sources from the northern (e.g., Indochina, Philippines, Taiwan, and eastern China) and southern (e.g., Malaysia and Indonesia) parts of Southeast Asia. As a result, the bimodal pattern was only obtained during the northeast monsoon period (December–March) because the local aerosol sources were mixed with several sources from Indochina that contained different sizes compared with those from the southern counterpart.

Figure 1b reveals that the distribution patterns of Angstrom exponent between the post-monsoon and northeast monsoon are similar. Figure 4a and d also indicate the similarities of the air flow patterns for these monsoon seasons. Hence, a clear correspondence was observed between Fig. 1b with Fig. 4a and d. The similarity in the patterns of Angstrom exponents for post-monsoon and northeast monsoon was attributed to the mixture of aerosols from northern and southern parts of Southeast Asia. Given the classification results (Fig. 3), MA was the major aerosol during the post-monsoon and northeast monsoon. The large amount of MA originated from South China Sea and Andaman Sea.

For the pre-monsoon period, the aerosols observed at Penang originated from the Malacca Strait, Andaman Sea, the northern and some eastern areas of Sumatra, and the western part of peninsular Malaysia, especially the local regions marked in yellow (Fig. 4b). During this season, the air flow patterns were similar to those during the southwest monsoon (Fig. 4c). However, a small percentage of aerosol was transported from the northern part of Southeast Asia to Penang. A clear correspondence was observed between Fig. 1b with Fig. 4b and c during pre-monsoon and southwest monsoon.

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The dominant aerosol types were UIA and MA (Fig. 3). The yellow portions in Fig. 4a–e indicate that Penang, the second largest city in Malaysia and one of the most industrially concentrated cities, was a major aerosol trap because of local and industrial emissions. MA contribution to the overall aerosol distribution could be significant because the study site was surrounded by the sea.

likely significantly influenced by proximity of the surrounding sea

3.5 Examination of predicted AOD values

Various Monthly AOD and Angstrom exponents from climatological data implied that each period exhibit different aerosol distributions in Penang. Seasonal analysis of the relative frequency occurrence of AOD₅₀₀, Angstrom_{440–870}, and PW clearly distinguished the dominant optical properties of aerosol for each monsoonal season. We hypothesized that the proposed model should exhibit different accuracies each season because the sensitivity for AOD prediction depended on the distribution patterns of the measured AOD; these values were used as inputs to derive the correlation parameters of the model. The sensitivity of AOD prediction was affected when the major occurrence frequency was clustered around small AOD values. The insensitivity of the aerosol models in clearing atmospheric conditions was also previously observed (Zhong et al., 2007). Conversely, the model appropriately predicted the AOD data when the corresponding input data were clustered around large values.

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The model performance for each monsoonal season was tested (Table 3). The pre-monsoon and southwest periods exhibited R^2 of 0.65 (RMSE = 0.114) and 0.77 (RMSE = 0.172). However, for the transition period between post-monsoon to northeast monsoon, $R^2 < 0.45$ and RMSE ranged from 0.06 to 0.11. The increased amount of atmospheric aerosol enhanced the predicted AOD and vice versa. This result was in agreement with the previous hypothesis. Meanwhile, the "overall" 22 month data was satisfactory with $R^2 = 0.72$ and RMSE = 0.133. The low value of %MRE (< 1) indicated that the model yielded accurate results for all seasons. Given the criteria that a low %MRE corresponded to a good prediction, the "overall" dataset yielded the least biased prediction.

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The aerosol types and distribution patterns could be elucidated from the model results. These parameters strongly depended on the changes in wind flow from one seasonal monsoon to the next. Aerosols were transported by monsoonal flows which were combined into the atmosphere from different sources. Haze could become noticeable (i.e., AOD value is high) when higher amount of aerosols was injected into the atmosphere from different sources, especially during large-scale open burning activities. High correlations were observed between the measured and predicted AOD for pre-monsoon and southwest monsoon, in which similar air flow patterns occurred (Fig. 4b and c). Figure 1b displays the relative frequencies of the occurrence of Angstrom_{440–870}. The frequency spectra for pre-monsoon and southwest monsoon also indicated the same patterns for AOD (Fig. 4b and c). The spectrum of the Angstrom frequency exhibited narrow peaks at 1.6 and 1.7 Å for pre-monsoon and southwest monsoon, respectively.

The accuracy of the prediction of the AOD model was moderate, the aerosols in Penang were locally mixed with those from foreign sources, because of the winds during post-monsoon and northeast monsoon characterized by similar air flow patterns (Fig. 4a and d). Correlations between Fig. 1b with Fig. 4a and d were observed for these monsoonal periods. The spectrum of the Angstrom frequency exhibited a broad region from 1.3 Å to 1.7 Å for post-monsoon and northeast monsoon. The broadened region implied that the particle size was largely distributed. The relationship between AOD to the air quality at ground surface depended on environmental factors, such as RH, aerosol size distribution, and chemical composition. These factors were disregarded in the AOD model, yielding deviations in the predicted values (Gupta et al., 2013; Lee et al., 2012).

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3.6 Validation of the predicted AOD

The optimized coefficients, a_i (Eq. 2), were obtained from the first subset in the "overall" dataset. To validate the model accuracy, a_i was used to predict the AOD from the second subset of the "overall" dataset (Fig. 5). The predicted AOD exhibited high

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than those at the ground level. This result was in accordance with the lower value in the predicted AOD of approximately 0.044, which was consistent with the expected result. Therefore, the predicted AOD values were acceptable because they exhibited small deviations against the measured AOD; this result was valid as long as the aerosols did not considerably differ at altitude levels beneath the planetary boundary layer. The LIDAR data should be considered as an independent validation method for ground-based prediction models. Comparing the consistency between the predicted results against LIDAR data could falsify or verify the correctness of the prediction model with high confidence. In reality, aerosols are not frequently well mixed in the atmosphere; several environmental factors can cause ambiguity in the predictions (Gupta et al., 2013; Lee et al., 2012). The small group of highly underpredicted results (Fig. 5) was attributed to the significant heterogeneity of aerosols in the atmosphere (e.g., aerosol residual layers) and the large amount of high-level transported aerosol (Tan et al., 2014b, c).

3.7 Applications of the proposed model in the absence of measured AOD data

Our proposed model generates AOD data when those from AERONET are unavailable. We described the procedure to predict AOD data. Only the API data for 7.00 a.m., 11.00 a.m., and 5.00 p.m. (local time) were available from the web site (<http://apims.doe.gov.my>) before 24 June 2013. The API data were provided hourly beyond this date. Any in situ visibility data with a value of -9999 and those recorded as fog, rain, or thunderstorms were removed. In this study, approximately 5% of the data were discarded, and only 4493 data points were retained. Figure 7 shows the predicted results from 2012 to 2013, which overlapped with the measured AOD data to simplify the comparison. The average AOD was 0.31 based on 4493 predicted data for the entire study period, which was near that of AERONET (about 0.29). The good agreement between the predicted and measured average AOD suggested that the model was sufficiently feasible to perform predictions.

As an illustration, we selectively zoom into three separate data windows (28 September, 17 October, and 30–31 October 2013; Fig. 8a–c) to analyze the variations in the

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predicted and measured AOD values at the scale of days. The predicted AOD and CIMEL sun photometer data were shown as blue and red dotted lines, respectively. The availability of the measured data points are often limited because of the unavailability of AERONET data caused by the presence of clouds and robotic errors. The predicted graphs exhibited temporal variation trends that tally with those measured at the same time scale (days).

AOD variations were continuously generated by the proposed model based on the hourly data from ground-based measurements. The unrecorded information by the sun photometer could be reproduced by the proposed method (Fig. 8). The model coefficients were trained under cloud-free conditions. Hence, the hourly AOD data could be generated anytime to compensate for the absence of measured AOD data during cloudy periods. In addition, the proposed model can generate daytime and nighttime temporal data in contrast to AERONET. Our model can be highly beneficial in monitoring the air concentration cycle because it generates continuous hourly data, hence, complementary information are provided.

The proposed model was independently verified using four selective sets of LIDAR data. We generated these data and compared them against the temporal plots of the aerosol backscattering coefficient signal (Fig. 9). The rectangles in Fig. 9a corresponded to the window periods for the LIDAR signal (Fig. 9b). The variation patterns in the retrieved AOD for the given window periods (Fig. 9a) corresponded well to the intensity variations in the aerosol backscattering coefficient signal (Fig. 9b). The LIDAR signals revealed the correctness of our predicted AOD, because the low (high) intensities of aerosol backscattering coefficient signal corresponded to low (high) AOD. The high intensities at 1–1.5 km altitudes (low cloud distributions) are represented by green ovals. Although clouds were present within the selected time windows, the retrieved AOD remained invariant. Therefore, this result strengthened the robustness of the proposed model to perform reliable and accurate prediction and retrieval of AOD. Our model could provide complementary retrieval of AOD data when AERONET data are unavailable because of the presence of clouds.

