A Global Model Simulation for 3-D Radiative Transfer Impact on Surface Hydrology over Sierra Nevada and Rocky Mountains

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Abstract. We investigate 3-D mountain effects on solar flux distributions and their impact on surface hydrology over the Western United States, specifically the Rocky Mountains and Sierra Nevada using CCSM4 (CAM4/CLM4) global model with a $0.23° \times 0.31°$ resolution for simulations over 6 years. In 3-D radiative transfer parameterization, we have updated surface topography data from a resolution of 1 km to 90 meters to improve parameterization accuracy. In addition, we have also modified the upward-flux deviation [3D – PP (plane-parallel)] adjustment to ensure that energy balance at the surface is conserved in global climate simulations based on 3-D radiation parameterization. We show that deviations of the net surface fluxes are not only affected by 3-D mountains, but also influenced by feedbacks of cloud and snow in association with the long-term simulations. Deviations in sensible heat and surface temperature generally follow the patterns of net surface solar flux. The monthly snow water equivalent (SWE) deviations show an increase in lower elevations due to reduced snowmelt, leading to a reduction in cumulative runoff. Over higher elevation areas, negative SWE deviations are found because of increased solar radiation available at the surface. Simulated precipitation increases for lower elevations, while decreases for higher elevations with a minimum in April. Liquid runoff significantly decreases in higher elevations after April due to reduced SWE and precipitation.

1. Introduction

Orographic forcing is an efficient and dominant mechanism for harnessing water vapor into consumable fresh water in the form of precipitation, snowpack, and runoff. It has been estimated that about 60 – 90% of water resources originate from mountains worldwide. Mountain water resources not only support human activities, but are also vital to diverse terrestrial and aquatic ecosystems. There is strong observational evidence that mountain water resources have been and continue to be threatened by global warming trends, which lead to snowpack reduction (Mote et al. 2007; Kapnick and Hall, 2012) and alter the timing and amount of runoff (McCabe and Clark, 2005). Observations and modeling studies have suggested that warming trends are amplified in mountains compared to lowlands because of the moist adiabatic structure of the atmosphere - the lapse-rate effect and snow-albedo feedback (Leung et al., 2004). Also, mountains are an integral part of global monsoon systems in which elevated warming may have important influence on monsoon circulation and the associated water cycle. However, accurate predictions of mountain snowpack have been limited by uncertainty in projecting future changes in temperature and precipitation due to model limitations in representing snow processes and their interactions with radiative transfer and other terrestrial processes in mountain environments.

The spatial and temporal distributions of surface solar radiation are the primary energy sources that contribute to the energy and water balance at 3-D and inhomogeneous mountain surfaces, with particularly strong influence on snowmelt processes (Geiger, 1965; Bonan, 2002; Gu et al., 2002; Müller and Scherer, 2005). The spatial orientation and inhomogeneous features of mountains/snow that interact with direct and diffuse solar beams are intricate and complex. Quantifying the interactions of direct and diffuse solar beams with mountain topography and reliably determining total surface solar fluxes for incorporation in a land surface model has been a challenging task that has yet to be accomplished in regional and high-resolution global climate modeling. Essentially all modern climate models have used a plane-parallel (PP) radiative transfer program in performing radiation parameterization; however, the potential errors have never been quantified.

In conjunction with radiative transfer in mountains/snow regions, we have developed a Monte Carlo photon tracing program specifically applicable to intense and intricate inhomogeneous mountains and demonstrated that the effect of mountains on surface radiative balance is substantial in terms of subgrid variability as well as domain average conditions (Liou et al., 2007; Lee et al., 2011; 2013). Because of the computational burden required by the 3-D Monte Carlo photon tracing program, an innovative parameterization approach has been developed in terms of deviations from PP radiative transfer results readily available in climate models for the five component of surface solar flux: direct and diffuse fluxes, direct- and diffuse-reflected fluxes, and coupled mountain-mountain flux (Lee et al., 2011). We have derived five regression equations for flux deviations which are linear and have a general 5 by 5 matrix form and successfully incorporated this efficient parameterization into the Weather Research Forecasting (WRF) model, which was used as the testbed in connection with the Fu-Liou-Gu PP radiation scheme (Fu and Liou 1992, 1993; Gu et al. 2010, 2011) that has been included in the WRF physics package. We have investigated 3-D mountain/snow effect on solar flux distribution and their impact on surface hydrology over the Western United States, specifically the Rocky Mountains and Sierra Nevada using the WRF applied at a 30 km grid resolution (Gu et al. 2012; Liou et al. 2013)

More recently, the 3-D radiative transfer parameterization has been incorporated into Community Climate System Model version 4 (CCSM4) global model with a 0.23°×0.31°

resolution to investigate the long-term 3-D effect on the simulated surface solar insolation patterns and associated sensible and latent heat fluxes, surface temperature, and surface hydrology over mountains/snow in the Western United States covering both the narrow coastal Sierra-Nevada Range and the broad continental Rocky mountains. Marked by complex terrain and with surface hydrology dominated by seasonal precipitation and snow accumulation and melt (e.g., Leung et al., 2003 a, b), the surface hydrology of the Western United States has been shown to be extremely sensitive to climate change (Leung et al., 2004; Kapnick and Hall, 2010). Thus, understanding factors leading to uncertainties in modeling snowpack and runoff is important for improving hydrologic predictions from seasonal to century time scales from the perspective of a global model

The organization of the present study is as follows. In Section 2 we describe CCSM4 with a brief discussion on the incorporation of the improved 3-D parameterization for surface solar radiation over mountain surfaces, followed by a discussion in Section 3 on the significance of 3-D radiation effect on the seasonal and elevation-dependent variations in solar flux, sensible and latent heat fluxes, surface temperature, and surface hydrology, including precipitation, snow water equivalent (SWE), and runoff, as well as a discussion on the potential impact of 3-D parameterization of surface solar radiation on vegetation. Concluding remarks are given in Section 4.

2. 3-D Radiation Parameterization in CCSM4

To study the long-term effect of 3-D mountain radiation effect over mountains/snow on the surface energy and hydrology, simulations using CCSM4 have been performed. CCSM is a general circulation model developed by the National Center for Atmospheric Research (NCAR).

The fourth version CCSM4 (Gent et al. 2011) is composed of atmosphere (Community Atmosphere Model, CAM4), land (Community Land Model, CLM4), sea ice (Community Ice Code, CICE4), and ocean (Parallel Ocean Program, POP2). The detail description of CCSM4 has already been given in Gent et al. (2011); thus only a brief outline of the components relevant to our study is presented here. Compared to the previous version, CAM4 used the finite-volume dynamical core (Lin, 2004) with the revised deep convection parameterization developed by Neale et al. (2008) that includes convective momentum transport. CLM4 was substantially modified (Lawrence et al., 2011) to include a carbon-nitrogen cycle (CLM-CN), a Snow and Ice Aerosol Radiation model (SNICAR, Flanner and Zender, 2006), and a dynamic vegetation model.

To investigate the impact of complex topography on surface solar radiation, the parameterization developed by Lee et al. (2011, 2013) has been incorporated in CCSM4. We have carried out 6-year simulations at a horizontal resolution of 0.23°×0.31° with prescribed sea surface temperatures and sea ice, greenhouse gases, and aerosols corresponding to Year 2000. The carbon-nitrogen cycle in CLM4 has also been activated. Although our goal is not to investigate 3-D mountain effects on vegetation, which would require long-term simulations to simulate vegetation response to different climate forcing, we included the carbon-nitrogen cycle in our simulations to provide preliminary indications of how vegetation processes may respond to changes in solar radiation due to mountain topography. Since a global high-resolution initial condition for CLM-CN is not available, our simulations were initialized using arbitrary initial conditions of land surface and vegetation states. Hence we note the caveat that slow processes such as groundwater table and carbon and nitrogen pools in our 6-years long simulations are far from reaching an equilibrium state and will have some influence on our results even with our focus on comparing simulations with and without 3-D mountain effects.

We have designed two experiments as follows: the PP experiment is the control run with default plane-parallel radiative transfer scheme, while the 3D experiment is identical to the PP experiment, except that the parameterization for 3-D solar flux is implemented. In this study, we focus on a domain covering the Rocky Mountains and Sierra Nevada from 120-105°W and 35 -45°N. Figure 1 displays the elevation map of the Western United States at a 0.23°×0.31° resolution, and the box is the area where the spatial average is calculated (see Liou et al., 2013).

In the previous WRF studies of 3-D radiative transfer, surface topography with a 1 km resolution was used, which was taken from the HYDRO1k geographic database available from the USGS' National Center for Earth Resources Observation and Science Data Center. We have since updated the surface topography data using the Shuttle Radar Topography Mission (SRTM) global dataset at a resolution of 90 meter (Jarvis et al., 2008) to perform 3-D Monte Carlo photon tracing simulations to improve parameterization accuracy (Lee et al., 2013). Because SRTM data cover the land surface between 56 S and 60 N, the parameterization is applied to all area within this range. Moreover, Lee et al. (2013) have shown that the parameterization can be applied to any grid box with a size larger than 10×10 km. Therefore, it is suitable for CCSM4 at a quarter-degree resolution.

In addition, we have also accounted for the adjustment involving upward flux deviations in the parameterization for application to climate models. It should be noted that the parameterization in our previous studies only adjusts downward solar fluxes calculated by the conventional radiative transfer scheme in a weather or climate model, while the upward fluxes remain unchanged. The impact of upward flux adjustment is normally insignificant and can be

Deleted: The magnitudes of upward flux adjustment are normally insignificant and can be neglected in surface energy analysis associated with a regional model neglected in regional model simulations since the contribution from the upward solar flux, which is only a fraction of the downward flux associated with surface albedo, to the atmospheric heating rate is much smaller than the downward flux. This slight adjustment for upward fluxes will ensure the total energy balance at the surface for simulations involving 3-D radiative transfer parameterization in a global model. Specifically, in the structure of a global climate model, land-surface model computes the surface albedo taking into account land types, snow cover, soil moisture, and other factors. This albedo is then employed as a boundary condition in the global climate model for radiative transfer calculations. We can use the parameterization for 3-D radiative transfer to adjust the land surface albedo, the ratio of the upward flux to the downward flux such that the downward flux adjustment remains unchanged. In this manner, a balance of the total energy flux at the surface would be ensured, which is critical for long-term climate simulations.

Following Lee et al. (2011), the downward surface solar flux can be categorized into: (1) The direct flux (F_{dir}) is composed of photons travelling from the Sun to the surface without encountering reflection or scattering. (2) The direct-reflected flux (F_{rdir}) is the reflection of F_{dir}. (3) The diffuse flux (F_{dif}) is associated with photons experiencing single and/or multiple scattering. (4) The diffuse-reflected flux (F_{rdif}) is the reflection of F_{dif}. The components related to downward direct solar radiation received by the real topography, F_{dir} and F_{rdir}, can be expressed as:

$$F_{\text{dir}} = (1 + f_{\text{dir}})\hat{F}_{\text{dir}} \text{ and } F_{\text{rdir}} = f_{\text{rdir}}\hat{F}_{\text{dir}} \tag{1}$$

where $\hat{F}_{\text{dir}}$ is the direct downward solar flux calculated by a plane-parallel radiative transfer scheme. f_{dir} and f_{rdir} are the relative deviations evaluated by parameterization and are functions of solar incident angle, standard deviation of elevation within a model grid box, sky view factor (the fraction of sky visible to the target), and terrain configuration factor (the area of surrounding mountains seen by the target). F_{rdir} is assumed to be proportional to the direct downward surface solar flux because conventional plane-parallel radiative transfer schemes do not explicitly calculate reflected fluxes. With the surface albedo for direct fluxes, α_{dir}, calculated by the land model, the direct radiation absorbed by the surface is equal to $(F_{\mathrm{dir}} + F_{\mathrm{rdir}}) \times (1 - \alpha_{\mathrm{dir}})$. We can now introduce the adjusted albedo for direct radiation in mountains, denoted as α'_{dir}. To keep the solar radiation absorbed by the surface unchanged, we must have

$$\hat{F}_{\mathrm{dir}}(1 - \alpha'_{\mathrm{dir}}) = (F_{\mathrm{dir}} + F_{\mathrm{rdir}})(1 - \alpha_{\mathrm{dir}}) . \quad (2)$$

Substituting Eq. (1) into Eq. (2) leads to

$$\alpha'_{\mathrm{dir}} = 1 - (1 + f_{\mathrm{dir}} + f_{\mathrm{rdir}})(1 - \alpha_{\mathrm{dir}}) . \quad (3)$$

Therefore, given the surface albedo provided by the land model and f_{dir} and f_{rdir} defined by the original parameterization, the adjusted albedo for direct flux can be obtained. Note that the adjusted albedo is independent from the value of incoming solar radiation, indicating that it can be calculated first and then used in the plane-parallel radiative transfer scheme to account for the topography effect. Correspondingly, the same procedure can be applied to the diffuse and diffuse-reflected fluxes, since CLM4 calculates albedos for direct and diffuse fluxes separately.

3. Model Simulation Results

3.1 3-D mountain effects on the geographic distribution of energy and hydrology

As mentioned above, we have conducted two 6-year CCSM4 simulations, PP and 3D. In the following presentation we have used the results determined from the last 5 years in the analysis. The 5-year mean net surface solar flux (FSNS), clear-sky surface solar flux (FSNSC), and total cloud fraction for April simulated with the incorporation of 3-D parameterization as a function of latitude and longitude are shown in Figs. 2a, 2b, and 2c, respectively, where the contour lines represent terrain height (km). FSNS generally follows the FSNSC and also depicts a pattern reflecting the negative modulation by the cloud fraction computed from the model. More clouds are generally found over the top of the mountains, where FSNS is relatively smaller because of reflection by snow over high elevation areas. The corresponding deviations (3D - PP) are displayed in Figs. 2d, 2e, and 2f. It reveals that the difference in FSNS is generally dominated by the difference in FSNSC. In this study, FSNSC is controlled by the adjusted albedo, which is related to snow cover and 3-D topography effect. Differences in FSNSC in Fig. 2e are mostly due to changes in the snow field, which will be discussed later. The 3-D topography effect can be found over the Sierra Nevada, where negative/positive deviation appears in the northern/southern slope.

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Changes in the surface downward solar flux distribution can affect cloud formation, which in turn will impact the transfer of solar flux reaching the surface. Figure 2f displays deviations (3D - PP) of total cloud fraction, which increases over mountain summits in the vicinity of northern Rockies around 45°N and 110°W (Fig. 2f) where the downward solar radiation decreases (Fig. 2d). In high-elevation areas, because of more reflection and less shading, the surface generally receive more solar radiation in the morning when the sky is clear. The additional insolation due to the topography effect can trigger convection earlier than 1D simulation, and then the larger cloud fraction produced by including the 3-D parameterization can reduce total daily insolation. For the broad south facing side of the mountains south of 38°N, increases in surface solar radiation correspond to decreases in cloud fraction.

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Figure 3a depicts the monthly mean SWE map for April simulated from CCSM4 with the inclusion of 3-D radiation parameterization for mountains. Significant SWE is mostly seen over the vast Rocky Mountain region and the narrow Sierra Nevada region. Generally, the SWE pattern shows relatively larger values on the west side of the mountains in response to enhanced precipitation on the windward slopes associated with orographic forcing. However, SWE displays smaller values at the highest elevation and on the east side of mountains in response to the reduced precipitation and the largest solar flux available at mountain tops. Contours of differences (3D – PP) in the simulated SWE are shown in Fig. 3b. Due to 3-D mountain effect, SWE generally decreases over mountain tops, especially in the area south of 42°N. In the Rocky Mountains (~37°N and 107°W), for example, reduction in SWE is as high as 100 mm or 40%. Decreased/increased SWE patterns correspond closely to increased/decreased net surface solar radiation patterns, as shown in Fig. 2d.

3.2 3-D mountain effects on seasonal variation

Figure 4 shows the 5-year mean deviations (3D – PP) in the domain-averaged monthly net surface solar flux, sensible heat fluxes, total cloud cover, and surface temperature as a function of month for different elevations over Sierra Nevada and Rocky Mountain areas. For long-term simulations during which cloud fields are modified through interactions with radiation, cloud feedback can play an important role in radiation field variation. As a matter of fact, the pattern of change in net solar flux is generally opposite to that of the total cloud fraction, where increases/decreases in the net solar flux correspond to decreases/increases in cloud cover (Figs. 4a and 4c). For higher elevations above 2.5 km, the net solar flux shows positive deviations largely throughout the year, indicating that mountain tops tend to receive enhanced solar radiation due to the 3-D effects. For valley areas with elevations lower than 2 km, while solar fluxes reaching the surface are also generally larger in the 3-D case, the magnitude of the increase in smaller than higher altitude regions due to the shading effect, as shown in our short-

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term WRF simulations for the same region (Liou et al. 2013). However, negative deviations mainly occur during December-January and in June due to increases in total cloud fraction (Figs. 4a and 4c). 3-D mountain effects lead to the reduction in total cloud fraction most of the year, except for January and June. Mountain clouds normally develop in response to surface solar heating, which gradually build up at the onset of morning hours. Furthermore, upslope flows contribute to convection and cloud formation as the elevated surface in mountains heats up relative to the surrounding air. A reduction in surface insolation can therefore reduce upslope flow and convection, leading to reduced clouds. Therefore, the reduced solar insolation in lower elevations due to the 3-D mountain effect tends to cool the surface and weaken the convection over mountain regions, resulting in less cloud water. Since cloud formation is primarily dominated by dynamical processes, enhanced surface heating over mountains tops due to the 3-D effect may not be sufficiently large to initiate cloud formation (Gu et al., 2012). However, during summer (June) when the surface is heated up, or during winter (January), which is the rainy season over the Sierra Nevada and Rocky Mountains in association with frontal systems, additional surface heating from the 3-D mountain effect could enhance cloud formation. Changes in sensible heat flux and surface temperature generally follow the patterns of net solar flux (Figs. 4b and 4d).

Figure 5 depicts the SWE, precipitation, and liquid runoff for the 3D experiment and differences between 3D and PP experiments. It is shown that SWE reaches its maximum in February in lower elevations and in March for higher elevations (Fig. 5a). Due to the 3-D mountain effect, decreases in SWE are found for the higher elevation zone (> 2.5 km) because more solar radiation is intercepted at mountain tops, while increases are found in lower elevations because of topographic shading (Fig. 5d). Positive deviations become smaller after

January because the sun is moving northward and getting closer to the overhead position during spring, leading to a reduced shading effect. The monthly mean precipitation (mm) as a function of elevation over the simulation domain is shown in Fig. 5b. Generally, precipitation increases with elevation due to orographic forcing. Precipitation shows maximum values around July for higher elevation zones and in January for all elevations in the rainy season (Fig. 5b). Differences in precipitation (Fig. 5e) are mostly negative values except for January and follow the pattern of total cloud fraction (Fig. 4c). The liquid runoff reveals a significant increase during April - June for the higher elevation range associated with the sun's position (Fig. 5c). Differences in liquid runoff are the combined results from snowmelt and precipitation. For higher elevations, due to more solar radiation, runoff first increases during February-March and then decreases after March related to less available snow and reduced precipitation (Fig. 5f). For valley areas, liquid runoff shows positive deviations beginning in January associated with more available snow amount and precipitation. Thus, the impact of 3-D mountain effect is to speed up snowmelt at mountain tops, and at the same time extend snowmelt and snowmelt-driven runoff into the warm season for lower elevations.

3-D mountain effects could have an important impact on surface vegetation. Many plant ecological studies, particularly those performed in mountainous terrain, have revealed that relationships exist between vegetation and the aspect and inclination of slopes (e.g. Killick, 1963; Edwards, 1967; Kruger, 1974; Granger and Schulze, 1977), which results largely from differences in the amounts of light, i.e. solar radiation, intercepted by different slopes. Solar radiation variation has been known to affect not only surface energy budgets (Garnier, 1968) and temperatures, but also soil moisture balances and photosynthesis processes. Such topographically induced incoming radiation differences may be regarded as one of the most fundamental variables of plant environment. Over a long-term period, plant would likely respond to differences in light amount (Granger and Schulze, 1977).

Figure 6 illustrates deviations of the domain-averaged monthly net vegetation absorbed solar radiation, sensible heat from vegetation, vegetation temperature, and total leaf area index (LAI) as a function of elevation. It is shown that the 3-D mountain induced changes in these vegetation related parameters, which will affect photosynthesis process and vegetation phenology, follow deviation patterns in the surface solar flux produced in part by elevation dependence. For example, for the vegetation absorbed solar radiation, positive deviations are seen for higher elevations (>2.5 km) with a maximum value in April, whereas negative deviations are found for valley areas (<1.5 km) with the largest reduction occurring in January (Fig. 6a), which largely follows the net surface solar flux patterns as shown in Fig. 4a. While the global radiation budget at the top of the atmosphere and surface, precipitation, and surface temperature do not have significant interannual variation, large fluctuations are seen in the temporal evolution of LAI over the Western United States. Clearly the vegetation results obtained from a 5-year simulation have not reached equilibrium as biomass continues to build up after model initialization. Still it is interesting to see how the difference in LAI between 3D and PP varies over the seasonal cycle with larger differences developing in early summer (Fig. 6d), following larger changes in the solar flux absorbed by the vegetation (Fig. 6a). However, much longer simulations with spun up carbon and nitrogen pools will be needed to obtain meaningful results for vegetation response to mountain-radiation interactions, a subject requiring further investigations in regards to the 3-D mountain effects on radiation and vegetation interaction and feedback.

4. Concluding remarks

The 3-D radiative transfer parameterization developed for the computation of surface solar fluxes has been incorporated into CCSM4 and applied at a resolution of 0.23°×0.31° over the Rocky Mountains and Sierra Nevada in the Western United States. We have carried out 6-year simulations with prescribed SST to understand the long-term effect of 3-D mountains on the monthly variation of surface radiative and heat fluxes and the consequence of snowmelt and precipitation on different elevations.

3-D mountain effects play an important role in the distribution of energy and water sources. Significant increases of net surface solar radiation are mainly found over mountain tops, while reductions, on the other hand, are mostly observed over valley areas. Changes in the surface downward solar flux distribution can affect the clouds and snow fields, which in turn will impact the transfer of solar flux reaching the surface. As a result, increases/decreases in surface solar radiation generally correspond to decreases/increases in cloud fraction and snow amount. Changes in clouds are mostly negative throughout the year due to the reduced solar radiation reaching the surface of lower elevations. The enhanced surface insolation at mountain tops appears to assist cloud formation during summer (June) related to surface heating or in January associated with frontal systems. Deviations in the surface solar radiation field can significantly alter the distribution of mountain snow. Decreases/increases in SWE correspond closely to increases/decreases in net surface solar radiation.

3-D mountain features also affect the seasonal variation of surface fluxes and hydrology. Deviations of the monthly mean surface solar flux produced by 3-D mountain effects, as compared to PP results, over the Rocky Mountain and Sierra Nevada regions are a function of elevation and at the same time, modulated by cloud feedback. Deviations in the net solar flux show opposite patterns to changes in the total cloud fraction. Deviations in the surface solar radiation field can affect heat fluxes, while changes in the surface energy balance are reflected in surface temperature variation. Changes in heat flux and surface temperature generally follow the deviation patterns in the net surface solar flux. Due to the 3-D mountain effect, decreases in SWE are found at higher elevation zones as a result of more solar radiation intercepted at mountain tops, while increases are found in lower elevations.

Differences in precipitation are mostly negative throughout the year, except for January, which follow the patterns of total cloud fraction. Differences in liquid runoff are produced by the combined results from snowmelt and precipitation. For higher elevations, due to increased solar radiation, runoff first increases during February and March but then decreases after March associated with reduced snow and precipitation. For valley areas, liquid runoff shows positive deviations after January associated with more available snow amount. Therefore, one of the important impacts of 3-D mountain effect is to speed up the snowmelt at mountain tops, while extend snowmelt and snowmelt-driven runoff into the warm season for lower elevations.

Finally, we wish to note that compared to our previous WRF studies of 3-D radiative transfer over mountains (Liou et al. 2013), similar 3-D mountain effects have been manifested in CCSM4 global simulations. Additionally, long-term simulations show that cloud feedback through cloud-radiation interactions exerts an important impact on surface fluxes and hydrology.

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Figure Captions

Fig. 1. The elevation map over a 0.23°×0.31°° resolution grid for the Rocky-Sierra areas in the Western United States. The box on the map displays major mountainous areas where simulation results are analyzed and presented in the paper.

Fig. 2. The April mean (a) net surface solar flux (W m^{-2}), (b) clear-sky net surface solar flux (W m^{-2}), and (c) total cloud fraction simulated for the 3-D case, and differences (3D – PP) in (d) net surface solar flux, (e) clear-sky net surface solar flux, and (f) total cloud fraction.

Fig. 3. The April mean (a) SWE (mm) and (b) corresponding differences (3D – PP).

Fig. 4. Deviations (3D - PP) of the domain-averaged monthly (a) net solar flux, (b) sensible heat flux, (c) total cloud fraction, and (d) surface temperature for a 12-month period as a function of elevation, lower that 1.5 km (red), 1.5-2 km (orange), 2-2.5 km (green), above 2.5 km (blue), and the whole domain (black).

Fig. 5. The monthly mean (a) Snow Water Equivalent (SWE, mm), (b) cumulative precipitation (mm), (c) cumulative runoff and the corresponding deviations (3D – PP) in (d) SWE, (e) precipitation, and (f) runoff, averaged over the simulation domain for a 12-month period as a function of elevation, lower that 1.5 km (red), 1.5-2 km (orange), 2-2.5 km (green), above 2.5 km (blue), and the whole domain (black).

Fig. 6 Deviations (3D - PP) of the domain-averaged monthly (a) vegetation absorbed solar flux, (b) sensible heat flux from vegetation, (c) vegetation temperature, and (d) total leaf-area index for a 12-month period as a function of elevation, lower that 1.5 km (red), 1.5-2 km (orange), 2-2.5 km (green), above 2.5 km (blue), and the whole domain (black).