- 1 We thank the Reviewer #1 for evaluating our paper, both here, and at the technical level.
- 2 The reviewer's comments, and our replies/revisions, are in red and black, respectively.
- 3 Page 26594
- 4 line 15: change 'in' to 'on'
- 5 Author's Change to Manuscript: Second, we use our measurement to derive a temperature-
- 6 and aerosol-dependent fit of N_{IC} based on Eq. (1).
- 7 line 24: 'heterogeneous ice generation can be distinguished ...': Please be more specific
 8 in how they can be distinguished
- 9
- 10 Author's Change to Manuscript: The foundations of our investigation are the cold-season
- 11 middle-tropospheric wave cloud studies of Cooper and Vali (1981), Cotton and Field
- 12 (2002), Eidhammer et al. (2010) and Field et al. (2012). The prior research demonstrated
- 13 that an assessment of wave cloud kinematics can be used to distinguish heterogeneous
- 14 nucleation from homogeneous ice nucleation and that crystal production occurs primarily
- 15 via the previously-mentioned freezing nucleation pathways. Further, no compelling
- 16 evidence for secondary ice production was reported in those prior studies.

17

19 Page 26596

line 25 ff.: to readers who are not familiar with FSSP, PCASP and 2DC measurements it
might not be clear from the description that evaporation is intended or may be a problem
of the instrument. Please be more explicit here.

- 23
- Author's Response: We feel that the manuscript makes it clear that evaporation, due to
- 25 heating of the sample stream, is an issue for measurements made with the PCASP (haze
- 26 particles and cloud droplets), but not for measurements made with the FSSP (cloud
- droplets) or with the 2DC (ice crystals). Further, we used the words "evaporate" and
- ²⁸ "evaporation" in our description of the PCASP measurement (bottom of p. 26596).
- Finally, we also stated that the PCASP measurements used to evaluate $n_{0.5}$ were
- acquired outside of cloud (Sections 3.3 and 3.4) where neither our work (Snider and
- Petters, 2006), or the work of Strapp et al. (1992), reveals a "problem of the instrument."
- 32 To emphasize this point, we modified the sentence on Page 26602, Line 11.
- 33
- 34 Author's Change to Manuscript: These were averaged outside of cloud during the five-
- second time windows used for thermodynamic-property averaging (Sect. 3.1).

- 37 Page 26600
- 38 line 7: 'This is shown, for the example,..': delete 'the'
- Author's Change to Manuscript: An example of this is shown in Fig. 1d.
- 40 Page 26602
- 41 line 6: Please describe more explicit the effect you observe of ice nucleation on cloud
- 42 properties which is evident in Figure 1.
- 43
- 44 Author's Change to Manuscript: The effect of ice development on cloud properties is
- 45 evident at the downwind track-streamline intersection in Figs. 1 and 2. Most noticeable
- 46 are the enhanced lidar depolarization ratios seen at $x \ge 15$ km in Fig. 1c and the enhanced
- 47 diameter-integrated crystal concentrations seen at $x \ge 15$ km in Fig. 2d.

- 49 Page 26604
- 50 line 7: A short description of D10's three-step procedure would help the reader to follow
- 51 the method described in this manuscript.
- 52 L5-L6, P26604 was modified to this:
- 53 Author's Change to Manuscript: ...using the three-step procedure described in D10. We
- refer to the latter as method #2 and describe our implementation of that method in
- 55 Appendix B (attached below).
- 56 Page 26605
- 57 line 4 ff.: Is the fraction of the measured crystal concentrations that plot within a factor of
- two of the fit still significantly different if the error of the measured ice crystalconcentrations is taken into account?
- 60 Author's Response: In the table shown below, we evaluate the effect of the Poisson
- 61 sampling error on the fractions. Regardless of how the comparison is made the
- 62 conclusion is the same: More points plot within a factor-of-two when using Eq. (1) with
- 63 our Method #1 coefficients.

64 Fraction of N_{IC} Measurements within a Factor of two of the Fit

	Eq. (1) with	Eq. (1) with
	Method #1 Fit Coefficients	D10 Fit Coefficients
Manuscript	0.69	0.66
L. 4 / P. 26605		
Measurements		
Increased	0.74	0.72
by Poisson Sampling		
Error		
Measurements	0.65	0.52
Decreased		
by Poisson Sampling		
Error		

65

- 67 Page 26606
- 68 line 11 ff.: Is should be stressed more that already the original D10 equation fits well to
- the measured data. This is of high value because of the very different measurementmethods.
- 71 Author's Change to Manuscript: The result we present in Tab. 2, with fit coefficients
- 72 generally consistent, in a statistical sense, with those reported by D10, is important
- 73 because it validates D10's approach using different methodology.

74 Page 26606

- 75 line 20: insert 'ice' before 'nuclei'
- 76 Author's Change to Manuscript: We also probed the conjecture that the duration of ice
- nuclei exposure to water-saturated conditions is a determinant of IC concentration.

78 **References**

- Cooper, W.A., and G. Vali, The origin of ice in mountain cap clouds, J. Atmos. Sci., 38,
 1244-1259, 1981
- 81 Cotton, R. and P. Field, Ice nucleation characteristics of an isolated wave cloud, Q. J.
- 82 Roy. Meteor. Soc., 128, 2417-2437, 2002
- 83 DeMott, P.J., A.J. Prenni, X. Liu, S.M. Kreidenweis, M.D. Petters, C.H. Twohy, M.S.
- 84 Richardson, T. Eidhammer and D.C. Rogers, Predicting global atmospheric ice nuclei
- distributions and their impacts on climate, P. Natl. Acad. Sci., 107, 11217-11222, 2010
- Eidhammer, T., P. J. DeMott, A. J. Prenni, M. D. Petters, C. H. Twohy, D. C. Rogers, J.
- 87 Stith, A. Heymsfield, Z. Wang, K. A. Pratt, K. A. Prather, S. M. Murphy, J. H. Seinfeld,
- 88 R. Subramanian, and S. M. Kreidenweis, Ice initiation by aerosol particles: Measured and
- 89 predicted ice nuclei concentrations versus measured ice crystal concentrations in an
- 90 orographic wave cloud, J. Atmos. Sci., 67, 2417-2436, 2010
- 91 Field, P.R., A.J. Heymsfield, B.J. Shipway, P.J. DeMott, K.A. Pratt, D.C. Rogers, J. Stith
- 92 and K.A. Prather, Ice in clouds experiment-layer clouds. Part II: Testing characteristics of
- heterogeneous ice formation in lee wave clouds, J. Atmos. Sci., 69, 1066-1079, 2012
- Snider, J.R., and M.D. Petters, Optical particle counter measurement of marine aerosol
 hygroscopic growth, Atmos. Chem. Phys., 8, 1949-1962, 2008
- 96 Strapp, J.W., W.R. Leaitch and P.S.K. Liu, Hydrated and dried aerosol-size-distribution
- 97 measurements from the particle measuring systems FSSP-300 probe and the deiced
- 98 PCASP-100x Probe, J. Atmos. Oceanic Tech., 9, 548-555, 1992
- 99

100 Appendix B

Here we describe how we fitted our 80 measurements of the set $\{N_{IC}, n_{0.5}, T_{low}\}$ using the three step procedure developed by D10 (herein method #2). In the first step, the data were binned into four $(273.16 - T_{low})$ subsets; the number of samples in the four subsets is provided in Table 3. In the second step, values of $ln(p_i)$ and q_i were derived for each subset by regression. Here "*i*" indicates the temperature subset and the form of the fitted equation is

107
$$ln(N_{IN,i}) = ln(p_i) + q_i \cdot ln(n_{0.5,i})$$
 (B1)

In the third step, the values of $ln(p_i)$ were regressed vs. $ln(273.16 - T_{low,i})$, and also, the values of q_i were regressed vs. $T_{low,i}$. In these regressions the $T_{low,i}$ is the average of the subset. The slopes and intercepts of these regressions define the coefficients ln(a), b, c and d for method #2.

112
$$ln(a) = intercept(ln(p_i) vs. ln(273.16 - T_{low,i}))$$
 (B2)

113
$$b = \text{slope}(ln(p_i) \text{ vs. } ln(273.16 - T_{low,i}))$$
 (B3)

114
$$c = \text{slope}(q_i \ vs. (273.16 - T_{low,i}))$$
 (B4)

115
$$d = \text{intercept}(q_i \ vs. (273.16 - T_{low,i}))$$
 (B5)