1	Insights into characteristics, sources and evolution of submicron aerosols
2	during harvest seasons in Yangtze River Delta (YRD) region, China
3	
4	Y. J. Zhang ^{1,2} , L. L. Tang ^{2,1} , Z. Wang ¹ , H. X. Yu ³ , Y. L. Sun ⁴ , D. Liu ⁵ , W. Qin ² ,
5	F. Canonaco ⁶ , A. S. H. Pr évôt ⁶ , H. L. Zhang ⁷ , and HC. Zhou ¹
6	
7	¹ Jiangsu Key Laboratory of Atmospheric Environment Monitoring and Pollution Control,
8	School of Environmental Science and Engineering, Nanjing University of Information
9	Science and Technology, Nanjing 210044, China
10	² Jiangsu Environmental Monitoring Center, Nanjing 210036, China
11	³ State Key Laboratory of Pollution Control and Resource Reuse, School of the Environment,
12	Nanjing University, Nanjing 210093, China
13	⁴ State Key Laboratory of Atmospheric Boundary Layer Physics and Atmospheric Chemistry,
14	Institute of Atmospheric Physics, Chinese Academy of Sciences, Beijing 100029, China
15	⁵ Centre for Atmospheric Science, School of Earth, Atmospheric and Environmental Sciences,
16	University of Manchester, Manchester M13 9PL, UK
17	⁶ Laboratory of Atmospheric Chemistry, Paul Scherrer Institute, Villigen PSI 5232,
18	Switzerland
19	⁷ Handix LLC, Boulder, CO 8031, USA
20	Correspondence to: L. L. Tang (lily3258@163.com)

22 Abstract

Atmospheric submicron particulate matter (PM₁) is one of the most significant pollution 23 components in China. Despite its current popularity in the studies of aerosol chemistry, the 24 characteristics, sources and evolution of atmospheric PM₁ species are still poorly understood 25 in China, particularly for the two harvest seasons, namely the summer wheat harvest and 26 autumn rice harvest. An Aerodyne Aerosol Chemical Speciation Monitor (ACSM) was 27 deployed for online monitoring of PM₁ components during summer and autumn harvest 28 seasons in urban Nanjing, in the Yangtze River Delta (YRD) region of China. PM₁ 29 30 components were shown to be dominated by organic fraction (OA, 39% and 41%) and nitrate (23% and 20%) during the harvest seasons (the summer and autumn harvest). Positive matrix 31 factorization (PMF) analysis of the ACSM OA mass spectra resolved four OA factors: 32 33 hydrocarbon-like mixed with cooking related OA (HOA + COA), fresh biomass burning OA (BBOA), oxidized biomass burning-influenced OA (OOA-BB), and highly oxidized OA 34 (OOA); in particular the oxidized BBOA contributes ~80% of the total BBOA loadings. Both 35 fresh and oxidized BBOA exhibited apparent diurnal cycles with peak concentration at night, 36 when the high ambient relative humidity and low temperature facilitated the partitioning of 37 semi-volatile organic species into the particle phase. The fresh BBOA concentrations for the 38 harvests are estimated as BBOA = $15.1 \times (m/z \ 60 - 0.26\% \times OA)$, $m/z \ 60$ as a marker for 39 40 levoglucosan-like species. The (BBOA + OOA-BB)/ Δ CO, (Δ CO is the CO minus 41 background CO), decreases as a function of f_{44} (fraction of m/z 44 in OA signal), which might indicate that BBOA was oxidized to less volatile OOA, e.g., more aged and less volatile OOA 42 (LV-OOA) during the aging process. Analysis of air mass back-trajectories indicates that the 43 44 high BB pollutant concentrations are linked to the air masses from the western (summer harvest) and southern (autumn harvest) areas. 45

47 **1 Introduction**

Particulate matter (PM) that is suspended in the atmosphere as atmospheric aerosol plays a 48 crucial role in regional and global climate system (Ramanathan et al., 2001; Kaufman et al., 49 50 2002), air pollution (Sun et al., 2013), ambient visibility reduction (Watson, 2002) and human health (Ge et al., 2011). Significant amounts of PM can be generated from human activities. 51 In particular, biomass burning (BB) activities, e.g., forest fires, wildfire, and agricultural fires, 52 can become the main sources of fine particulate matter (PM_{2.5}, particulates $\leq 2.5 \mu m$ in 53 aerodynamic diameter) and/or submicron particulate matter (PM₁, particulates $\leq 1 \mu m$ in 54 55 aerodynamic diameter) (Andreae and Merlet, 2001, Aiken et al., 2010; DeCarlo et al., 2010; Lee et al., 2010; Cubison et al., 2011; Reche et al., 2012; Bougiatioti et al., 2014). 56 Agricultural residues burning is one of the most serious sources leading to severe air quality 57 58 problems during harvest seasons in China (Li et al., 2007; Wang et al., 2009a; Du et al., 2011; Cheng et al., 2013; Ding et al., 2013). Moreover, China is an agricultural country which has 59 1.8 billion cultivated fields with a large amount of agricultural crop residue (Zhang et al., 60 2008). Recently, the use of agricultural residues as fuel in China declined. During harvest 61 seasons, farmers usually harvest the crop in the daytime and burn agricultural residues in 62 their fields directly, which results in BB emissions. The investigation of the compositions, 63 sources and processes of atmospheric aerosol particles during harvest seasons is urgently 64 needed to better understand the impact of aerosol particles from BB sources on air quality. 65

Organic aerosol (OA) composes a large fraction of atmospheric aerosol particles (Zhang et al., 2007). Combination of positive matrix factorization (PMF, Paatero, 1997) and a PMF Evaluation Toolkit (PET, Ulbrich et al., 2009) has been well used to identify and apportion the sources of OA in recent studies (e.g. Lanz et al., 2007; Ulbrich et al., 2009; Allan et al., 2010; Zhang et al., 2005a, 2011; Crippa et al., 2013, 2014; Sun et al., 2013). In addition, an IGOR-based Source Finder (SoFi, Canonaco et al., 2013) with a multilinear engine algorithm 72 (ME-2, Paatero, 1999) can also resolve the emission sources of OA. The current PMF and ME-2 method can only be employed to analyze OA datasets a posteriori (Sun et al., 2012; 73 Zhang et al., 2011; Canonaco et al., 2013), but cannot be easily utilized in the real-time 74 75 online estimation of atmospheric OA sources. To identify the sources of atmospheric OA online, an algorithm based solely on organic mass fragments, namely m/z 57 (mostly C₄H₉⁺) 76 and m/z 44 (mostly CO₂⁺), was developed to estimate hydrocarbon-like OA (HOA) and 77 oxygenated OA (OOA), respectively (Zhang et al., 2005a, 2005b; Ng et al., 2011c). Mohr et 78 al. (2012) also identified cooking OA (COA) in ambient datasets based on the fractions of 79 80 COA tracers at m/z 55 (mostly C₄H₇⁺) and m/z 57 organic mass fragments. Biomass burning organic aerosol (BBOA) is one of the major atmospheric OA species during BB periods 81 (Aiken et al., 2010; Allan et al., 2010). However, limited information on developing the 82 83 tracer-based method a posteriori is available for estimating the source apportionment of 84 BBOA.

The evolution processes of atmospheric OA, e.g., aging and/or oxidation, can 85 significantly influence the physicochemical properties of OA (Aiken et al., 2008; Jimenez et 86 al., 2009; Sun et al., 2011b). In the presence of BB source, various volatile and semi-volatile 87 organic precursors can be emitted from the field burning of agricultural wastes, and SOA can 88 be formed from these precursors rapidly (Jimenez et al. 2009; Grieshop et al., 2009; Heringa 89 90 et al., 2011; Kawamura et al., 2013). What is more, BB plumes can be mixed with urban and 91 regional pollutants during aging processes (DeCarlo et al., 2010; Cubison et al., 2011). In addition, the secondary formation, atmospheric transport and diffusion, as well as the mass 92 loadings and oxidation state of ambient OA can be also affected by the aging processes of 93 94 OA (Jimenez et al., 2009; Cubison et al., 2011; Sun et al., 2011b). Thus, it is important for understanding the nature of atmospheric OA to investigate the evolution of OA and the 95 evolution process effects. 96

97 This study investigates the characteristics of PM₁ species using an Aerodyne Aerosol 98 Chemical Speciation Monitor (ACSM), and OA mass spectra are analyzed with PMF model 99 during summer and autumn harvests in the YRD region; the evolution of OA and the effects 100 of the evolution process on PM burden were also investigated. Combination of 101 back-trajectory analysis and local wind meteorology was used to investigate the origins.

102

103 2 Experimental methods

104 **2.1 Sampling site description**

With a population of more than 8 million and an area of ~ 6597 km², Nanjing, is a 105 representative Chinese city in terms of the pollution characteristics of the YRD region. Local 106 107 and regional air pollution events frequently occur in Nanjing, mainly caused by emissions of 108 mixed aerosols from fossil fuel burning, residential activities, and agricultural residues 109 burning (Wang et al., 2009a; Ding et al., 2013). In this study, all kinds of data was collected in urban Nanjing (118 46'N, 32 05'E) from June 1 to 15, and October 15 to 30, 2013, 110 111 corresponding to two harvest seasons in a year, namely the summer wheat harvest and autumn rice harvest. The sampling site was located on the roof of a six-story building 112 approximately 18 m above ground level, ~15 m from the nearest heavy-traffic road, and ~50 113 m from the nearest restaurants and residents. As a matter of fact, there is no agricultural field 114 in the urban Nanjing areas. However, there are some agricultural fields in the rural areas 115 116 around Nanjing (Figure S1). This means that the urban Nanjing site is significantly influenced by the BB plumes originated from the rural areas. In addition, the local cooking 117 and traffic emissions can also significantly affect the PM pollution in this sampling site. 118 119 Therefore, in the presence of BB plumes, the mixed/complicated air pollution will occur in urban Nanjing during the harvest seasons. 120

122 **2.2 Instrumentation and data analysis**

123 2.2.1 Measurements

The ambient non-refractory submicron aerosol (NR-PM₁) species, i.e., OA, nitrate, sulfate, 124 ammonium, chloride, were continuously measured using ACSM from June 1 to 15, and 125 October 15 to 30, 2013. In addition, the measurement without significant agricultural burning 126 impacts (little/negligible BB-influence) was also performed at the same site from July 1 to 8. 127 Detailed descriptions of ACSM can be found in previous studies (Ng et al., 2011a; Sun et al., 128 2012). Briefly, the ambient aerosols were drawn into the room using a ¹/₂ inch (outer diameter) 129 stainless steel tube at a flow rate of $\sim 3 \, \mathrm{l} \, \mathrm{min}^{-1}$, of which $\sim 84 \, \mathrm{cm}^3 \, \mathrm{min}^{-1}$ was sub-sampled into 130 the ACSM. Moreover, ACSM was operated at a time resolution of about 15 min with a scan 131 from m/z 10 to 150 amu at 500 ms amu⁻¹ rate, which corresponds to the settings of Sun et al. 132 133 (2012).

An online analyzer, Monitoring of Aerosols and Gases (MARGA, model ADI 2080 134 Applikon Analytical B. V. Corp., the Netherlands), was deployed to measure the mass 135 concentrations of a major water-soluble inorganic ion (potassium ion, K⁺) in the aerosols. A 136 PM_{2.5} cyclone inlet was used to remove coarse particles. Ambient air was sampled into a 137 liquid with a flow rate of 16.7 l min⁻¹. The detection limit of K^+ is 0.09 µg m⁻³. The MET 138 ONE BAM-1020 and the 7-wavelength aethalometer (Magee AE31) were also employed to 139 measure PM1 and ambient atmospheric BC in PM2.5, respectively. CO was measured using a 140 141 gas analyzer (Thermo Scientific, Model 48i). Ambient meteorological parameters including ambient temperature (T), relative humidity (RH), precipitation, wind speed (WS), and wind 142 direction (WD) were obtained from a ground level meteorology station located on the same 143 144 six-story building as the sampling site.

Daily fire locations used in this study were available from MODIS (Moderate-resolution
 Imaging Spectroradiometer) mounted on NASA's Terra and Aqua satellites, NASA's Earth

Observing System (EOS) (https://earthdata.nasa.gov/data/near-real-time-data/firms). MODIS can present fire distributions in details at 1 km resolution through Fire Information for Resource Management System (FIRMS) on global scale (Justice et al., 2002; Kaufman et al., 2003). As shown in Figure S1 – S2, all agricultural fire locations (red dots) in the YRD region were detected by the remote sensing retrieval of MODIS from June 1 to 15, and October 15 to 30, 2013 (https://firms.modaps.eosdis.nasa.gov/firemap/).

153

154 2.2.2 ACSM data analysis

155 An ACSM Data Analysis Software package, ACSM Local (Ver. 1.5.2.0.0, released April 25, 2012) written in Wavemetrics IgorTM, was used to analyze the ACSM dataset. More details of 156 procedures have been described in the studies of Ng et al. (2011a) and Sun et al. (2012). The 157 158 ACSM calibration is based on a combination of a Differential Mobility Analyzer (DMA, TSI model 3080) and Condensation Particle Counter (CPC, TSI model 3775) for the ionization 159 efficiency (IE) and relative ionization efficiencies (RIEs). Pure ammonium nitrate (NH₄NO₃) 160 particles (size selected by 300 nm) are used for the calibration of the instrument, because 161 NH₄NO₃ vaporizes with 100% efficiency (Ng et al., 2011a). The RIEs values usually used in 162 Aerosol Mass Spectrometer (AMS) ambient concentration calculations (Canagaratna et al., 163 2007) are the default values of organics (1.4), nitrate (1.1), sulfate (1.2), and chloride (1.3) in 164 this study. Moreover, the RIE value of ammonium is 7.04, and the response factor (RF) value 165 of nitrate is 3.96×10^{-11} in this study. In addition, the mass concentrations of ambient aerosol 166 need to be corrected for particle collection efficiency (CE) (Middlebrook et al., 2011). CE = 167 0.5 is found to be representative with data uncertainties generally within 20% (Canagaratna et 168 169 al., 2007; Middlebrook et al., 2011). The CE values observed in previous studies range from 0.43 to 1, due to (a) shape-related collection losses at the vaporizer from inefficient focusing 170 of non-spherical particles, (b) particle losses at the vaporizer because of bouncing of solid 171

particles before they are completely vaporized, and (c) particle losses in the aerodynamic inlet as a function of particle diameter (Allan et al., 2004; Zhang et al., 2005b; Canagaratna et al. 2007; Middlebrook et al., 2011). In this study, we selected the CE value for OA, nitrate, sulfate, ammonium, and chloride, respectively, according to the equation CE = max (0.45, 0.0833 + 0.9167 × ANMF) (Middlebrook et al., 2011), in which ANMF is the mass fraction of NH₄NO₃ measured by the ACSM.

The PMF method, combining PMF2 executables with the PMF Evaluation Tool (PET) 178 (Ulbrich et al., 2009), was applied to analyze OA datasets from the ACSM. More details of 179 180 procedure for the PMF model can be found in previous studies (Ulbrich et al., 2009; Zhang et al., 2005a, 2011). Due to large interferences of internal standard of naphthalene at m/z's 127 181 - 129, only m/z < 120 was used for PMF analysis (Sun et al., 2012, 2013). Based on the OA 182 183 dataset from the ACSM, the PMF analysis was performed for 1 to 7 factors. A summary of the PMF results is presented in Figure S6 - S15. For the chosen number of factors, fpeaks 184 were varied in steps of 0.1 from -1 to 1 for the data of the summer and autumn harvest. Four 185 186 OA factors, i.e., hydrocarbon-like mixed with cooking OA (HOA + COA), fresh biomass-burning OA (BBOA), oxidized BB-influenced OA (OOA-BB), and highly 187 oxygenated OA (OOA) were resolved in this study. The HOA + COA was considered as a 188 factor mixing with COA and traffic HOA in the 4-factor solution, while the "HOA + COA" 189 factor splits into factors with very similar time series in the five- to a seven-factor solution. 190 191 This means that the PMF analysis of the ACSM OA mass spectra has difficulties distinguishing the COA from the traffic HOA in this study. Sun et al. (2010) and Sun et al. 192 (2012) also found a similar phenomenon for distinguishing COA from traffic HOA in a 193 194 Quadrupole AMS (Q-AMS) and an ACSM OA mass spectra using PMF analysis in Beijing. For the 3-factor solution, BBOA might be mixed with OOA-BB, while the 4-factor solution 195 (which contained two BB-related BBOA factors, i.e., BBOA and OOA-BB) seemed to be 196

better (more details can be found in section 3.2) than the other solutions. The detailed lists of
explanation on the reasons for the selection of the 4-factor solution can be found in Table S4
- S5. In addition, the OA source apportionment for the two harvests will be further discussed
in section 3.2.

201

202 2.2.3 Back-trajectory analysis

Impacts of various source regions on the PM pollution during the harvest seasons have been investigated using the HYbrid Single Particle Lagrangian Integrated Trajectory (HYSPLIT-4) model developed by NOAA/ARL (Draxler and Rolph, 2003). Accordingly, 48 h back-trajectories (BTs) at 500 m arrival height above ground level were calculated every 2 h starting at China Standard Time (CST) using a Trajectory Statistics (TrajStat) software developed by Wang et al. (2009b). In this study, the 48 h back-trajectories of air masses were used for further analysis.

210

211 **3 Results and discussion**

212 **3.1 Meteorological factors and PM1 components**

213 **3.1.1 Time series of meteorological factors and PM1 components**

Figure 1 shows the time series of NR-PM₁ species and BC in the presence of different 214 meteorological conditions during the harvest seasons in urban Nanjing, i.e., WS, WD, RH, T, 215 216 and precipitation. During the summer harvest, the average values were 70.7 \pm 15.3%, 3.7 \pm 1.7 m s⁻¹, and 23.4 \pm 4.1 °C for the ambient RH, WS, and T, respectively. In the autumn 217 harvest, the average values were 54.3 \pm 13.7 %, 2.6 \pm 1.4 m s⁻¹, and 18.1 \pm 3.6 °C for the 218 ambient RH, WS, and T, respectively. The frequency distribution of hourly averaged wind 219 direction and speed throughout the summer and autumn harvests are shown in Figure S2 (a -220 221 b).

As shown in Figure S3, there is a strong correlation between the MET ONE PM₁ 222 measured by MET ONE BAM-1020 and the PM₁ (= NR-PM₁ + BC) mass concentrations (r^2 223 = 0.88, *slope* = 1.11), indicating that the ambient submicron aerosols consisted mainly of the 224 225 NR-PM₁ and BC. Note that the mass concentration of BC in the PM₁ may be overestimated due to the fact that the mass concentration of BC was measured by the 7-wavelength 226 aethalometer for PM_{2.5}. An overestimation was previously suggested by Huang et al. (2011). 227 The average PM₁ mass for the summer harvest is 38.5 μ g m⁻³ with hourly average ranging 228 from 3.6 to 270.6 μ g m⁻³, which is similar to that observed in the autumn harvest (42.3 μ g 229 m^{-3}) with hourly average ranging 8.1 to 191.5 µg m^{-3} . Indeed, PM₁ consisted of OA (39%), 230 nitrate (23%), ammonium (16%), sulfate (12%), BC (8%), and chloride (1%) during the 231 summer harvest. During the autumn harvest, PM₁ was composed of OA (41%), nitrate (20%), 232 233 ammonium (14%), sulfate (11%), BC (13%), and chloride (1%). Table 1 presents a comparison of the average composition of PM₁ between the summer harvest and autumn 234 harvest periods. The average bulk composition of PM₁ during the summer harvest shows 235 similar dominance of OA to the PM pollution during the autumn harvest, but lower mass 236 fractions for other species except nitrate. Overall, those species with the exception of BC also 237 show a similar contribution between the summer and autumn harvest to the PM₁ mass, 238 suggesting that the PM pollution could be affected by similar pollution sources for the two 239 harvests. 240

As shown in Figure 1, all aerosol species exhibited very dynamic variations in mass concentrations due to the changes of source emissions, meteorology factors (such as WD, RH, *T*, and planetary boundary layer height), photochemical reactions and regional transport (e.g. the BB plumes). For example, the aerosol species dramatically reduced because of the quick removal processes associated with heavy wet scavenging (e.g. 6 - 8 June) during the summer harvest. However, the wet scavenging plays a minor role in changing aerosol loadings with

little precipitation during the autumn harvest. OA shows a significant dynamic variation in 247 mass concentrations during the harvest seasons (Fig. 1c), likely due to the changes of source 248 emissions (such as cooking, traffic and/or BB emissions). There are three sharp peaks during 249 the summer harvest (case 1) and autumn harvest (case 2 and case 3). The relationships 250 between the PM pollution, meteorology and chemical composition are presented in three case 251 events (Table S1). The case 1, at 21:00 - 22:00 on 10 June, with the highest PM₁ mass (253.1) 252 μ g m⁻³) during the summer harvest is characterized by high loadings of K⁺, BBOA, OOA-BB, 253 chloride, and BC, indicating the significant impacts of agricultural burning from the 254 255 northwest of Nanjing (Fig. 1). Apart from the high loadings of BB-related components, such as BBOA, OOA-BB, and K⁺ seen in the case 2 and case 3 periods, local source-related 256 components, e.g., HOA + COA and BC, also present high concentrations. This suggests that 257 258 both local primary sources emission and regional BB plumes dominate the PM pollution during the case 2 and case 3 periods. Therefore, those findings indicate that indeed BB 259 contributes significantly in the area during the specific time period. 260

261

262 **3.1.2 Diurnal variations of meteorological factors and PM1 components**

Figure 2 depicts the diurnal variations of the meteorological factors, i.e., RH, T, and WS, and 263 PM₁ species (including OA, nitrate, sulfate, ammonium, chloride, and BC). Generally, the 264 diurnal variations of the meteorological parameters and PM₁ species are similar during the 265 266 summer and autumn harvest. However, the ambient RH and T during summer harvest were higher than those during autumn harvest. OA obviously exhibits three peaks occurring 267 between 6:00 - 8:00, 11:00 - 14:00, and 19:00 - 22:00, which is in agreement with the 268 269 contributions of pollution sources, e.g., traffic, cooking and/or BB (Allan et al., 2010; Huang et al., 2012; Sun et al., 2012; Crippa et al., 2013). More details of the diurnal variations of the 270 271 OA components will be presented in section 3.2.

Sulfate does not show any significant diurnal trend during both summer and autumn 272 harvest, and shows a similar concentration during the two harvests. This means the 273 non-volatile character and regional pollution of sulfate in the YRD region during the summer 274 275 and autumn harvest. Similar diurnal trend of sulfate was also found by Huang et al. (2012) in the eastern YRD region. Nitrate presents in a higher fraction of the total PM₁ compared with 276 sulfate, yet with lower concentrations in the afternoon and higher concentrations in the 277 evening during the harvests. Similarly, nitrate also shows a similar concentration for the two 278 harvests during the whole day. In addition, chloride shows a similar diurnal cycle with nitrate 279 280 during the two harvest seasons. This is in accordance with the volatile and gas-particle partitioning properties of ammonium nitrate and ammonium chloride dependent on ambient T281 and RH (Lanz et al., 2007; Sun et al., 2011b, 2012). This also reflects that the photochemical 282 production of HNO₃ cannot compensate for the evaporative loss at the relatively high T283 conditions during the two harvests, which is similar to previous results observed by Huang et 284 al. (2012) in the eastern YRD region and Sun et al. (2012) in Beijing. Furthermore, the higher 285 boundary layer may dilute their loadings during the daytime, and then influence their diurnal 286 cycles (Sun et al., 2012). Chloride is mainly ammonium chloride (NH₄Cl) and/or organic 287 chlorine-containing species (Huffman et al., 2009; Huang et al., 2012; Sun et al., 2012). 288 During the harvest seasons, the evening high values of nitrate and chloride might be affected 289 290 by the BB emissions and/or formed via gas-phase and aqueous-phase oxidations.

BC shows a classic diurnal variation with higher loadings appearing in early morning and during nighttime, which is consistent with traffic rush hours in early morning (07:00 - 08:00)and during nighttime (20:00 - 21:00). As in previous studies, atmospheric BC is strongly associated with combustion emissions (including traffic and BB sources emissions), particular for BB periods (Sandradewi et al., 2008; Liu et al., 2011, 2014; Crippa et al., 2013). Therefore, the reason for the peak values of BC during the nighttime may be also caused by the BB emissions during the harvest seasons, apart from the effect of traffic source on the BC
loadings. The lower concentrations of BC in the afternoon can be associated with the dilution
effects of higher planetary boundary layer and reduced traffic emissions.

300

301 3.2 Organic source apportionment

Four OA factors (i.e. HOA + COA, BBOA, OOA-BB, and OOA) were identified, as illustrated in Figure 3 and Figure 4. The mean mass concentrations of HOA + COA, BBOA, OOA-BB and OOA during the harvest seasons are presented in Table 1. HOA + COA, BBOA, OOA-BB, and OOA accounted on average for 15% (28%), 7% (7%), 29% (33%) and 49% (32%) of the total OA mass concentrations during the summer (autumn) harvest, respectively.

308 **3.2.1 Hydrocarbon-like and cooking-emission related OA (HOA + COA)**

The prominent hydrocarbon ion series of $C_nH_{2n+1}^+$ and $C_nH_{2n-1}^+$ (e.g. 27, 29, 41, 43, 55, 57...) 309 obtained from the mass spectrum were characterized as components of HOA (Zhang et al., 310 311 2005a, 2011; Mohr et al. 2009; Allan et al., 2010). As reported in previous studies, m/z 57 $(C_3H_5O^+ \text{ and/or } C_4H_9^+)$ and m/z 55 $(C_3H_3O^+ \text{ and/or } C_4H_7^+)$ are commonly considered as 312 tracers for the primary organic emissions of combustion sources in urban areas, including 313 COA and HOA (Zhang et al., 2005a; Ng et al., 2010, 2011b; He et al., 2010; Sun et al., 2012, 314 315 2013; Hu et al., 2013). It is found that there is no significant difference in the mass spectrum 316 between the summer harvest and the autumn harvest (Fig. 3a). Compared with traffic-like OA (Liu et al., 2011; Crippa et al., 2013), the mass spectrum obtained in the present study shows 317 a higher m/z 55/57 ratio. Previous studies indicated that high m/z 55/57 together with a unique 318 319 diurnal variation can be used as a diagnostics for the presence of COA (Mohr et al., 2009; Allan et al., 2010; Sun et al., 2012). The mass spectrum of HOA in this study is characterized 320 by more abundant ions, i.e., m/z 41 (mainly C₃H₅⁺), m/z 55 (mainly C₄H₇⁺) and m/z 57 (Fig. 321

322 3a), which is similar to the characteristics of COA mass spectrum measured by He et al. 323 (2010). As shown in Figure 3, the diurnal variation of HOA + COA shows two pronounced 324 peaks corresponding to noon (a weak peak) and evening traffic/cooking activities (a strong 325 peak). Hence, HOA + COA in this study refers to the sum of traffic-related HOA and COA. 326 Similarly, Sun et al. (2010) and Sun et al. (2012) also found that HOA species in urban 327 ambient were influenced by both traffic and cooking-like emissions.

328

329 **3.2.2 Fresh biomass burning OA (BBOA)**

330 As shown in Figure 3b, the mass spectrum of BBOA extracted in this study shows a prominent peak of m/z 60 which is a well-known tracer ion for BB emissions (Alfarra et al., 331 2007; Aiken et al., 2009; Cubison et al., 2011; Huang et al., 2011; Liu et al., 2011). 332 333 Levoglucosan was shown to contribute to m/z 60 and was found in large amounts in urban, suburban, and rural background atmosphere during BB periods (Maenhaut et al., 2012). In 334 addition, the BBOA is also characterized by higher peaks at masses m/z 27, 29, 41, 43, 55, 57, 335 77 and 91 that are indicative of freshly emitted organic aerosol, because fresh m/z 43 ... m/z336 57 can be also from BB-related emissions (Aiken et al., 2009; Heringa et al., 2011; 337 Bougiatioti et al., 2014). For example, primary BBOA (P-BBOA) has a significant 338 contribution from a non-oxygenated ion $C_3H_7^+$ at m/z 43, but not from an oxygenated ion 339 $C_2H_3O^+$ (m/z, 43) in smog chamber experiments by Heringa et al. (2011). The BBOA 340 spectrum profiles with the lack of m/z 44 signal (CO₂⁺) during the summer and autumn 341 harvest show high correlation ($r^2 = 0.82$ and $r^2 = 0.87$) with a result in Paris (Crippa et al., 342 2013). Moreover, the spectrum of BBOA in this study is qualitatively similar to published BB 343 spectra from the fresh BB smoke in a smog chamber (Grieshop et al., 2009). These findings 344 suggest that this factor can be related to BBOA with low atmospheric oxidants, and thus this 345 factor might this factor might be associated with fresh/primary BBOA during the harvests. 346

347 Using soluble K⁺ as a tracer for BB has also been reported by previous analyses of BB campaign data (Gilardoni et al., 2009; Aiken et al., 2010; Du et al., 2011; Crippa et al., 2013). 348 The time series of BBOA along with K⁺ measured by MARGA is shown in Figure 4b. BBOA 349 is strongly correlated with K^+ ($r^2 = 0.95$ and $r^2 = 0.78$) during the summer and autumn 350 harvest respectively (Fig. 5a), suggesting that BBOA and K⁺ were from the same source. In 351 addition, the diurnal variation of BBOA shows a pronounced peak at the nighttime (Fig. 3), 352 which is consistent with the effects of the BB emissions (Fig. 3). This means that BBOA 353 contributes to primary organic aerosol (POA) during the nighttime mainly. This finding is 354 also consistent with the habit of the farmers in the YRD region, namely that they usually 355 harvest wheat or rice in the daytime and burn off straw in the nighttime during the harvest 356 seasons each year. In addition, chloride correlates well with BBOA ($r^2 = 0.61$ and $r^2 = 0.66$) 357 and K^+ ($r^2 = 0.60$ and $r^2 = 0.64$) during the harvest seasons (Figure S5). This suggests that 358 chloride was mainly from the BB emissions and might be in the form of KCl during the BB 359 periods. 360

361

362 **3.2.3 Oxygenated OA (OOA) and oxidized BB-influenced OA (OOA-BB)**

The mass spectrum of both OOA components (Fig. 3c and d) was characterized by the 363 prominent $C_xH_yO_z^+$ fragments, which has been denoted previously found in many AMS 364 studies (Zhang et al., 2005a; Lanz et al., 2007; Sun et al., 2010; Crippa et al., 2013). The 365 mass spectra of OOA by the prominent peak of m/z 44 (mainly CO₂⁺) (22.9% and 25.5% of 366 the total OOA signal respectively) during the summer and autumn harvest are strongly 367 consistent with more oxidized OOA component determined ($r^2 = 0.91$ and $r^2 = 0.89$, Fig. 3d) 368 in BB-period in Paris (Crippa et al., 2013) and OOA components resolved at other urban sites 369 (Lanz et al., 2007; Ulbrich et al., 2009). 370

371 In Figure 4d the time series of OOA is compared with the sulfate mass loadings. A

correlation was observed between time series of OOA and sulfate mass loadings ($r^2 = 0.60$ 372 and $r^2 = 0.46$, Fig. 6) during the summer and autumn harvests, respectively. Previous studies 373 performed at various sites also showed that these two species were secondary with 374 375 low-volatility property in the atmosphere (Zhang et al., 2005a; Lanz et al., 2007; Ulbrich et al., 2009; Sun et al., 2011a; Huang et al., 2012). Overall, the diurnal pattern of OOA shows 376 relatively stable trend throughout the whole day (Fig. 3). OOA often remains at a high 377 concentration across several days until a change of air mass occurs, which shows a regional 378 production (Sun et al., 2012, 2013). This may be a main reason causing the relatively stable 379 380 trend through the whole day in this study. Nevertheless, OOA shows a slight increase at around 12:00 - 15:00, suggesting that more oxidized OOA might be formed by 381 photochemical processing. OOA also exhibits higher loadings during the nighttime, probably 382 383 caused by the aging of BB plumes, in which BB emissions will be further oxidized and begin to transition into OOA (Jimenez et al., 2009, DeCarlo et al., 2010). The uniform distribution 384 of its concentrations is almost in association with all kinds of WD during the summer and 385 386 autumn harvest respectively (Figure S4). This is a good evidence for explaining the regional pollution of OOA in the YRD region during the harvest seasons. 387

Additionally, an oxygenated factor with the high degree of oxygenation during the 388 summer and autumn harvest (m/z 44, 18.2% and 14.5% of the total factor signal respectively) 389 in its mass spectrum has been resolved and identified as oxidized BB-influenced OA 390 391 (OOA-BB, Fig. 3c). The mass spectra of OOA-BB are characterized by both the oxidized ions (m/z 18, 29, 43 and 44) and the typical marker of BB (m/z 60) during the summer and 392 autumn harvest, which correlates well with those of BB-emission related OOA 393 (OOA₂-BBOA) ($r^2 = 0.85$ and $r^2 = 0.86$) during BB periods at an urban site in Paris (Crippa et 394 al., 2013). It is also highly similar to the mass spectrum of the aged BBOA identified by 395 DeCarlo et al. (2010) for airborne measurements during the MILAGRO campaign, and very 396

397 in agreement with the aged BBOA from a BB experiment in a chamber study by Heringa et al. (2011). In addition, the mass spectrum of OOA-BB shows more oxygenated degree, 398 compared to mass spectrum of fresh/primary BBOA from PMF analysis in the atmosphere 399 400 and from laboratory open wood burning (Aiken et al., 2009) and from BBOA in this study. The OOA-BB spectrum in this study is also very similar to the spectrum of the aged OA 401 produced from aged biomass smoke in a smog chamber (Grieshop et al., 2009). OOA-BB 402 presents a pronounced diurnal cycle with the highest concentration in the evening and early 403 morning during the harvests (Fig. 3), which is very consistent with the diurnal variations of 404 405 BBOA. This means that the OOA-BB production from open BB is rapid in short timescales with the high RH and low T conditions in the nighttime. OOA-BB also shows relatively low 406 loadings in the daytime, due to the dilution effects by enhanced mixing in the planetary 407 408 boundary layer and the evaporative loss of semi-volatile components.

As shown in Figure 4c, the OOA-BB time series strongly correlates with K⁺ and $\Delta m/z$ 409 60 (= m/z 60 – 0.26% × OA, in which applied metric of background f_{60} = 0.26% of OA will 410 411 be discussed in section 3.4) during the summer and autumn harvest, supporting the BB influence. In addition, the sum of BBOA and OOA-BB also shows high correlation with K⁺ 412 and $\Delta m/z$ 60 for the two harvests (Fig. 5a – b). This suggests that OOA-BB represents an 413 atmospheric mixture of BBOA and OOA, which is similar to a recent HR-ToF-AMS study by 414 Crippa et al. (2013). It is interesting that OOA-BB correlates well with nitrate ($r^2 = 0.30$ and 415 $r^2 = 0.54$), yet shows lower correlation with sulfate ($r^2 = 0.16$ and $r^2 = 0.30$) for the summer 416 and autumn harvest respectively (Fig. 6). Also, the time series of OOA-BB shows a similar 417 trend as chloride during the two harvest seasons (Fig. 4c). This implies an indication of the 418 419 semi-volatile character of OOA-BB, which is consistent with the results from a recent field study in the eastern Mediterranean (Bougiatioti et al., 2014) and some laboratory chamber 420 studies (Lipsky and Robinson, 2006; Robinson et al., 2007; Yee et al., 2013). Particularly, this 421

also means that aged biomass burning OA (OOA-BB) may be significantly mixed with nitrate
in the BB plumes. Healy et al. (2013) also found a similar result in Paris using single-particle
mass spectrometer (SP-AMS) and HR-ToF-AMS measurements.

425

426 **3.3 Effects of Chemical components on PM pollution**

Figure 7 presents the average contributions of PM₁ species and OA components during the 427 summer and autumn harvest, respectively. It is also compared with other sites, including 428 megacities (Mexico City, Paris, Beijing, and Shanghai), and suburban/remote areas (Crete, 429 430 Jiaxing, and Pearl River Delta) (Aiken et al., 2009; Crippa et al., 2013; Huang et al., 2012, 2013; Sun et al., 2012; Bougiatioti et al., 2014). Using the relative contribution of the sum of 431 BBOA and OOA-BB to OA, the harvest season was separated into 3 time periods, i.e., low 432 433 BB (L-BB, 28% and 29%) period, medium BB (M-BB, 49% and 38%) period, and high BB (H-BB, 93% and 50%) period, during the summer and autumn harvest respectively. We also 434 include averages of some meteorological parameters (i.e. RH, T, WS, and WD) for the 435 reference, and these averages are shown in Table S2. Compared with other sites including 436 Mexico City, Paris, Crete, Jiaxing, and Pearl River Delta (Fig. 7), the BB source shows the 437 largest contribution to aerosol pollution during the H-BB both in the summer and autumn 438 harvest in urban Nanjing. In addition, in the absence of BB source the sum of OOA and 439 440 OOA-BB shows a higher fraction of total OA mass during the harvests in urban Nanjing in 441 comparison with Beijing and Shanghai in China (Fig. 7). This means that the BB source significantly contributes to the SOA pollution in urban Nanjing during harvest seasons. 442

As shown in Figure 7, OA is important in PM pollution in the summer and autumn harvest (39% and 41%). Furthermore, the average fraction of BBOA to OA during the summer harvest (7%) is highly consistent with that in the autumn (7%), while BC shows a higher fraction during the autumn harvest (12%) than that in the summer harvest (8%). This 447 is also corresponding to the fraction of HOA + COA, which shows a higher contribution during the autumn harvest (28%) than that in the summer harvest (15%). The different 448 boundary layer height and primary sources emission influences on primary pollutants 449 450 (including BC, HOA and COA) may be all potential causes of such seasonal differences. On average, the total oxidized fraction of OA (including OOA and OOA-BB) accounts for more 451 than 60% (78% for summer and 65% for autumn harvest), which indicates that regional OOA 452 453 plays an important role in PM pollution in urban Nanjing during the harvest seasons. As a comparison, OOA-BB shows a higher fraction to OA in H-BB period than in L-BB period. 454 455 The fraction of OOA-BB to OA is higher than the fraction of BBOA during the harvest seasons, even in the H-BB period. These findings indicate that "aged" BBOA plays a more 456 significant role in PM pollution than BBOA in the BB plumes, particularly in the H-BB 457 458 period. This is consistent with recent studies (Grieshop et al., 2009; Heringa et al., 2011; Lathem et al., 2013; Yee et al., 2013; Bougiatioti et al., 2014) indicating that the fresh BB 459 emission OA can be rapidly surpassed by SOA formation within a few hours after its 460 emission. 461

The secondary inorganic aerosols (including sulfate, nitrate, and ammonium) can be 462 seen in lower fraction in the H-BB period than in the L-BB period. However, the mass 463 concentrations of sulfate, nitrate and ammonium are higher in the H-BB period than in the 464 L-BB period (Table S2) respectively. Therefore, these findings indicate that the BB source 465 466 contributes more to OA than secondary inorganic components. It is interesting that the contribution of nitrate to PM₁ is higher than the contribution of sulfate in the H-BB periods 467 during the two harvest seasons. For example, the average contribution of nitrate to PM₁ is 468 469 ~18% in the H-BB periods, which is almost twice higher than that of sulfate. However, the contribution of nitrate to PM₁ is very similar to the sulfate contribution in the L-BB periods. 470 471 All of those indicate the BB is a much more important source of nitrate, compared to sulfate. 472 Similar results have been observed by Crippa et al. (2013), Healy et al. (2013) and
473 Bougiatioti et al. (2014) during open BB periods.

Figure 8 presents the mass fractions of PM₁ species and OA components as a function of 474 total PM₁ mass loadings, as well as the probability density of total PM₁ mass loadings during 475 the summer and autumn harvest respectively. Overall, the total OA fraction increases from 476 about ~15% to 40% and from ~30% to 45% as a function of the PM_1 loadings during summer 477 and autumn, respectively. Indeed, OOA-BB and BBOA show a significant increase as a 478 function of the PM_1 loadings respectively during the harvest seasons. The contribution of 479 480 OOA-BB to PM₁ increases from $\sim 3\%$ ($\sim 5\%$) to 33% (26%) during the summer (autumn) harvest. And the contribution of BBOA increases from ~2% (~4%) to 8% (8%). The results 481 highlight the contribution of OOA-BB arising from BB emissions to PM pollution in the 482 483 harvests. During the summer harvest, the HOA + COA and BC mass fractions display a slight decrease, suggesting that local primary sources play an important in the low PM pollution 484 period. In addition, the nitrate and sulfate contributions show a slight increase and decrease 485 respectively, indicating additional production of nitrate mass during high PM episodes. Note 486 that the mass fraction of OOA shows a slight decrease with the increasing of total PM₁ 487 loadings during the autumn harvest. This suggests that OOA mainly contributes to the low 488 PM pollution, and OOA-BB mainly contributes to the high PM pollution. However, the 489 490 contribution of HOA + COA, BC, and the secondary inorganic species to the total PM_1 491 loadings did not show clear PM-mass loading dependency, which indicates that the high PM pollution during the autumn harvest may be caused by the synergistic effects of all pollutants. 492 493

494 **3.4 Estimation of BBOA directly from a tracer** ($\Delta m/z$ 60)

495 The BBOA mass loadings during the harvest seasons were estimated a posteriori using a 496 simple method. As described in previous studies, the parameter f_{60} , fraction of m/z 60 in total 497 OA, is considered as a marker of fresh/primary BBOA (Alfarra et al., 2007; DeCarlo et al., 2008; Aiken et al., 2009; Cubison et al., 2011). To estimate the real value of the BBOA 498 loadings, the background fraction of f_{60} (0.26 \pm 0.1%) during little/negligible BB-influence 499 periods (non-BB periods) was determined (Fig. 9). Aiken et al. (2009) and Cubison et al. 500 (2011) also obtained a similar background level of f_{60} (0.3 ± 0.06%) for an urban city in 501 Mexico Therefore, the levoglucosan-like species in ambient BB plumes was estimated by Δ 502 m/z 60 ($\Delta m/z$ 60 = m/z 60 - background value of $f_{60} \times OA$). As shown in Figure 5b, the 503 strong correlations ($r^2 = 0.95$, $r^2 = 0.98$, and $r^2 = 0.97$) between the BBOA and $\Delta m/z$ 60 with 504 505 the similar slopes, i.e., 16.3 for summer, 14.6 for autumn, and 15.1 for the total harvest seasons, were observed. The OOA-BB mass loadings also show the high correlations with Δ 506 m/z 60 ($r^2 = 0.95$ and $r^2 = 0.97$), but with very different slopes (74.8 and 64.4) during the 507 summer and autumn harvest respectively (Fig. 5b). Aiken et al. (2009) also found that BBOA 508 strongly correlated with $\Delta m/z$ 60 mass loadings ($r^2 = 0.91$, Slope = 34) during the 509 BB/wood-smoke periods in Mexico City. Furthermore, Lee et al. (2010) obtained a strong 510 relationship between BBOA and m/z 60 mass loadings ($r^2 = 0.92$, Slope = 34.5) through a 511 wildland fuels fire experiment in the lab. Thus, we reconstructed the time series of BBOA to 512 compare the relationship between the extracted BBOA by PMF model (PMF BBOA) and the 513 estimated BBOA. As shown in Figure S16, an excellent agreement is observed between the 514 identified and reconstructed BBOA concentrations during the total harvest seasons ($r^2 = 0.97$). 515 516 Therefore, the BBOA component during the BB periods in urban Nanjing of the YRD region can be estimated with the equation of BBOA = $15.1 \times (m/z \ 60 - 0.26\% \times OA)$ for the harvest 517 518 seasons.

519

520 **3.5 Evaluation of OA**

521 To further investigate the probable importance of the aging and/or mixing processes of

BB plumes, the total BB-related OA (BBOA + OOA-BB) to Δ CO ratio as a function of the 522 f_{44} during the summer and autumn harvest respectively is shown in Figure 10. The CO 523 background is determined as 14.9 μ g m⁻³ for summer harvest and 17.9 μ g m⁻³ for autumn 524 harvest, respectively, based on an average of the lowest 5% CO during two plumes 525 (Takegawa et al., 2006). The ratio of BBOA + OOA-BB to Δ CO can remove the effect of 526 dilution in the atmosphere (de Gouw 2005; Dunlea et al., 2009; DeCarlo et al., 2008, 2010). 527 As discussed in de Gouw et al. (2005), Aiken et al. (2008), Jimenez et al. (2009), and Ng et al. 528 (2010), the f_{44} can be considered as indicator of atmospheric aging due to photochemical 529 530 aging processes leading to the increasing of f_{44} in the atmosphere. Overall, the (BBOA + OOA-BB) / Δ CO ratio shows an obvious reduction with increasing of f₄₄ values during the 531 summer and autumn harvest respectively, in the absence of traffic and cooking-like plumes. 532 533 This is likely due to a synergistic effect of the rapid formation of OOA from BB plumes and the mixing of BBOA with regional OOA and/or CO. Similar results have been found by 534 DeCarlo et al. (2010), from aircraft measurements during MILAGRO in Mexico City and the 535 Central Mexican Plateau. 536

Figure 11a depicts the evolution process of OA with the f_{44} vs. f_{43} space during two harvest seasons. The BBOA and HOA + COA show similar low oxidative properties with varying f_{43} , which are located at the left-bottom of the triangular region during the summer and autumn harvest, respectively. With the aging process in the atmosphere, OA clusters within a well-defined triangular region and shows more similar oxidative properties to OOA-BB and/or OOA (Fig. 11a). This implies that OOA-BB and/or BBOA might be further oxidized, and might be transformed into highly oxidized OOA.

Furthermore, the formation and transformation of primary and secondary BBOA during BB periods can be described by f_{44} vs. f_{60} plot (Cubison et al., 2011). In the f_{44} vs. f_{60} space of Figure 11b, OA shows a trend toward higher f_{44} and lower f_{60} values with the aging of BB plumes, appearing into the low-volatility OOA (LV-OOA) range. This is very consistent with previous reports in aircraft and laboratory studies (Cubison et al., 2011) with a similar trend. In a smog chamber experiment, Grieshop et al. (2009) also found that the relative contribution at m/z 44 and m/z 60 rapidly increases and decreases, respectively during aging process, which presents the characteristics of fresh and aged BBOA.

552

553 **3.6 Impacts of various source regions on the PM pollution**

Figure 12 presents the BTs clusters of air masses at 500 m arrival height above ground 554 level at intervals of two hours (i.e. 00:00, 02:00, 04:00, ..., 22:00) starting at CST using the 555 HYSPLIT model (Draxler and Rolph, 2003) in Nanjing (118 46'N, 32 05'E). The 556 corresponding BTs can be broadly classified into four principal clusters of air masses based 557 558 on the spatial distributions during the summer and autumn harvests, respectively, i.e., northeasterly (NE) back-trajectories (BTs), easterly marine (EM) BTs, southeasterly marine 559 (SEM) BTs, and southwesterly continental (SWC) for the summer harvest; northerly 560 continental (NC) BTs, northeasterly marine (NEM) BTs, easterly marine (EM) BTs and 561 southerly continental (SC) for the autumn harvest. The air masses in Nanjing in this study 562 563 were mainly from the SEM BTs (accounting for 57.4% of all the BTs) during the summer harvest, while predominantly from the NC and EM BTs (at frequencies of 43.8% and 24%, 564 respectively) during the autumn harvest (Fig. 12 and Table S3). 565

The average PM₁ loadings are the highest (71.3 μ g m⁻³) for the WC BTs, which is almost twice higher than that of the lowest (24.4 μ g m⁻³) for the EM BTs during the summer harvest. This suggests that the long-range transported pollutants from southwestern areas can cause the high PM pollution in the YRD region during the summer harvest. Similarly, the highest average concentration of PM₁ (80.9 μ g m⁻³) is associated with the continental-related air masses during the autumn harvest. Therefore, source regions related to the fire locations (Figure S1) are of utmost importance to the high air pollution in the YRD region during theharvest seasons.

The PM₁ chemical compositions show also significantly different fraction among the 574 four clusters during the summer and autumn harvest respectively, which might be associated 575 with the different source regions of air pollution. The lowest PM1 loadings are associated 576 with the EM BTs, but with the high contribution of HOA + COA during the summer harvest 577 578 (Fig. 12a). This suggests that the local sources play a key controller of aerosol pollution during relatively clean periods in the summer harvest. For the NE BTs, the OOA, nitrate, and 579 580 sulfate provide high fractions of the total PM₁ mass, suggesting that regional pollution plays a key role in controlling the PM pollution. With regards to the marine-related air masses, PM₁ 581 loadings associated with the NEM and NE BTs are higher in the summer harvest than in the 582 583 autumn harvest (Fig. 12b), which may be due to the fact that high local POA contributes to the PM pollution. This suggests that local sources play a more important role in the aerosol 584 pollution in the summer harvest than in the autumn harvest. Compared with other clusters 585 during the autumn harvest, the BB-related components (e.g. BBOA, OOA-BB and chloride) 586 contribute the highest fractions to the PM₁ mass in air masses originated from the SC BTs, 587 indicating that BB plumes can contribute to the highest heaviest PM pollution during the 588 autumn harvest. Apart from the high contributions of nitrate and OOA, HOA + COA also 589 590 accounted a higher fraction to PM₁ mass in the NEM and EM BTs than in other clusters 591 during the autumn harvest. In addition, the PM₁ components show the lowest concentrations for the NC BTs, compared to the other clusters during the autumn harvest. When removing 592 the mass concentrations of BB related OA (BBOA and OOA-BB), the mean concentration of 593 PM₁ (31.6 μ g m⁻³) for the NC BTs is corresponding to a result (28.7 μ g m⁻³) for a similar 594 cluster during a non-BB period (Huang et al., 2012). 595

596 **4 Conclusions**

The characteristics, sources and evolution of atmospheric PM_1 species in urban Nanjing, the 597 YRD region of China were investigated using an Aerodyne ACSM during the two harvest 598 599 seasons, namely the summer wheat harvest (June 1 to 15, 2013) and the autumn rice harvest (October 15 to 30, 2013). The PM₁ species show a similar contribution, which on average 600 account for 39% (41%) OA, 23% (20%) nitrate, 16% (14%) ammonium, 12% (11%) sulfate, 601 8% (13%) BC, and 1% (1%) chloride during the summer (autumn) harvest. Secondary 602 inorganic species, i.e., nitrate, sulfate and ammonium, show highly similar diurnal patterns 603 604 between the summer and autumn harvest, meaning its similar chemical processing and physical processes (e.g. gas-particle partitioning). In particular, OA, chloride and BC present 605 higher concentrations in the diurnal cycles during the autumn harvest than during the summer 606 607 harvest, due to larger impacts of BB and/or local primary emissions during the autumn harvest. 608

PMF analysis was performed on the ACSM OA mass spectra to investigate organic 609 610 source apportionment during the two harvests. Four OA components were resolved including two POA factors associated with traffic and cooking (HOA + COA) and biomass burning OA 611 (BBOA) emissions and two secondary factors associated with regional and highly oxidized 612 OOA and less oxidized BB-like OA (OOA-BB). Apart from HOA + COA, BBOA and 613 614 OOA-BB also present pronounced diurnal cycles during the harvests, with the highest 615 concentrations occurring at night due to the nighttime BB plumes over urban Nanjing. This suggests that BBOA components may be quickly oxidized to OOA-BB during the nighttime 616 with the high RH and low T conditions. The diurnal profiles of OOA are similar to that of 617 618 sulfate with relatively flat variations, reflecting their regional origin. OA was dominated by secondary organics (OOA and OOA-BB) with the fraction more than 60% to total OA mass. 619 POA shows a lower contribution to OA during the summer (autumn) harvest, traffic and 620

621 cooking 15% (28%) and BB 7% (7%) emissions. The background level of f_{60} (0.26 ±0.1%) was determined using the f_{44} vs. f_{60} space during the non-BB periods (in July). Thus, we 622 suggest a simpler method for estimating the fresh BBOA loadings based on the equation of 623 BBOA = $15.1 \times (m/z 60 - 0.26\% \times OA)$ during the harvests. The (BBOA + OOA-BB) / ΔCO 624 ratios decrease with the increasing of the f_{44} , suggesting that BBOA components may be 625 oxidized to more aged and less volatile OOA, e.g., LV-OOA during the aging process. Air 626 627 mass trajectory analysis indicates that the high PM pollution is mainly contributed by nitrate, BBOA, and OOA-BB, which is associated with air masses originated from the western 628 629 (summer harvest) and southern (autumn harvest) areas.

630

631 Acknowledgements

This work was funded by the Natural Science Key Research of Jiangsu Province High 632 Education (11KJA170002), the Foundation Research Project of Jiangsu Province 633 (BK2012884, BK20140987), the Project Funded by the Jiangsu Province Science & 634 Technology Support Program (BE2012771), the Environmental Monitoring Scientific 635 Research Foundation of Jiangsu Province (1016), and the National Natural Science 636 Foundation of China (21407080). We are very grateful for the help and support from Dr. 637 Douglas R. Worsnop and Dr. John T. Jayne (Aerodyne Research Inc.) in the ACSM 638 measurements. We also would like to thank Dr. P. Chen (Handix LLC), and Dr. W. Li (South 639 Coast Air Quality Management District) for their constructive suggestions in improving the 640 contents. We thank the reviewers for their valuable comments on the manuscript. 641

642

643 **References**

Aiken, A. C., Decarlo, P. F., Kroll, J. H., Worsnop, D. R., Huffman, J. A., Docherty, K. S., 26

Ulbrich, I. M., Mohr, C., Kimmel, J. R., Sueper, D., Sun, Y., Zhang, Q., Trimborn, A.,
Northway, M., Ziemann, P. J., Canagaratna, M. R., Onasch, T. B., Alfarra, M. R., Prevot,
A. S. H., Dommen, J., Duplissy, J., Metzger, A., Baltensperger, U., and Jimenez, J. L.:
O/C and OM/OC ratios of primary, secondary, and ambient organic aerosols with
high-resolution time-of-flight aerosol mass spectrometry, Environ. Sci. Technol., 42,
4478–4485, 2008.

Aiken, A. C., Salcedo, D., Cubison, M. J., Huffman, J. A., DeCarlo, P. F., Ulbrich, I. M., 651 Docherty, K. S., Sueper, D., Kimmel, J. R., Worsnop, D. R., Trimborn, A., Northway, M., 652 653 Stone, E. A., Schauer, J. J., Volkamer, R. M., Fortner, E., de Foy, B., Wang, J., Laskin, A., Shutthanandan, V., Zheng, J., Zhang, R., Gaffney, J., Marley, N. A., Paredes-Miranda, G., 654 Arnott, W. P., Molina, L. T., Sosa, G., and Jimenez, J. L.: Mexico City aerosol analysis 655 during MILAGRO using high resolution aerosol mass spectrometry at the urban 656 supersite (T0) – Part 1: Fine particle composition and organic source apportionment, 657 Atmos. Chem. Phys., 9, 6633-6653, doi:10.5194/acp-9-6633-2009, 2009. 658 Aiken, A. C., de Foy, B., Wiedinmyer, C., DeCarlo, P. F., Ulbrich, I. M., Wehrli, M. N., 659 Szidat, S., Prévôt, A. S. H., Noda, J., Wacker, L., Volkamer, R., Fortner, E., Wang, J., 660

Laskin, A., Shutthanandan, V., Zheng, J., Zhang, R., Paredes-Miranda, G., Arnott, W. P.,
Molina, L. T., Sosa, G., Querol, X., and Jimenez, J. L.: Mexico city aerosol analysis
during MILAGRO using high resolution aerosol mass spectrometry at the urban
supersite (T0) - Part 2: Analysis of the biomass burning contribution and the non-fossil
carbon fraction, Atmos. Chem. Phys., 10, 5315-5341, doi:10.5194/acp-10-5315-2010,
2010.

Alfarra, M. R., Pr év ât, A. S. H., Szidat, S., Sandradewi, J., Weimer, S., Schreiber, D., Mohr,
M., and Baltensperger, U.: Identification of the mass spectral signature of organic
aerosols from wood burning emissions, Environ. Sci Technol., 41, 5770-5777, 2007.

670	Allan, J. D., Bower, K. N., Coe, H., Boudries, H., Jayne, J. T., Canagaratna, M. R., Millet, D.
671	B., Goldstein, A. H., Quinn, P. K., Weber, R. J., W. D. R.: Submicron aerosol
672	composition at Trinidad Head, California, during ITCT 2K2: Its relationship with gas
673	phase volatile organic carbon and assessment of instrument performance, J. Geophys.
674	Res., 109: D23S24, 2004.
675	Allan, J. D., Williams, P. I., Morgan, W. T., Martin, C. L., Flynn, M. J., Lee, J., Nemitz, E.,
676	Phillips, G. J., Gallagher, M. W., and Coe, H.: Contributions from transport, solid fuel
677	burning and cooking to primary organic aerosols in two UK cities, Atmos. Chem. Phys.,

678 10, 647-668, doi:10.5194/acp-10-647-2010, 2010.

- Andreae, M. O., and Merlet, P.: Emission of trace gases and aerosols from biomass burning,
 Global Biogeochem. Cycles, 15, 955-966, 2001.
- Bougiatioti, A., Stavroulas, I., Kostenidou, E., Zarmpas, P., Theodosi, C., Kouvarakis, G.,
 Canonaco, F., Pr év ât, A. S. H., Nenes, A., Pandis, S. N., and Mihalopoulos, N.:
 Processing of biomass-burning aerosol in the eastern Mediterranean during summertime,

684 Atmos. Chem. Phys., 14, 4793-4807, doi:10.5194/acp-14-4793-2014, 2014.

- 685 Canagaratna, M., Jayne, J., Jimenez, J. L., Allan, J. A., Alfarra, R., Zhang, Q., Onasch, T.,
- 686 Drewnick, F., Coe, H., Middlebrook, A., Delia, A., Williams, L., Trimborn, A., Northway,
- M., Kolb, C., Davidovits, P., and Worsnop, D.: Chemical and microphysical
 characterization of aerosols via Aerosol Mass Spectrometry, Mass Spectrom. Rev., 26,
 185-222, 2007.
- Canonaco, F., Crippa, M., Slowik, J. G., Baltensperger, U., and Prévôt, A. S. H.: SoFi, an
 IGOR-based interface for the efficient use of the generalized multilinear engine (ME-2)
 for the source apportionment: ME-2 application to aerosol mass spectrometer data,
 Atmos. Meas. Tech., 6, 3649-3661, doi:10.5194/amt-6-3649-2013, 2013.

- Cheng, Y., Engling, G., He, K.-B., Duan, F.-K., Ma, Y.-L., Du, Z.-Y., Liu, J.-M., Zheng, M.,
 and Weber, R. J.: Biomass burning contribution to Beijing aerosol, Atmos. Chem. Phys.,
 13, 7765-7781, doi:10.5194/acp-13-7765-2013, 2013.
- Crippa, M., DeCarlo, P. F., Slowik, J. G., Mohr, C., Heringa, M. F., Chirico, R., Poulain, L., 697 Freutel, F., Sciare, J., Cozic, J., Di Marco, C. F., Elsasser, M., Nicolas, J. B., 698 Marchand, N., Abidi, E., Wiedensohler, A., Drewnick, F., Schneider, J., Borrmann, S., 699 Nemitz, E., Zimmermann, R., Jaffrezo, J.-L., Pr év ât, A. S. H., and Baltensperger, U.: 700 Wintertime aerosol chemical composition and source apportionment of the organic 701 fraction in the metropolitan area of Paris, Atmos. Chem. Phys., 13, 961-981, 702 doi:10.5194/acp-13-961-2013, 2013. 703
- Crippa, M., Canonaco, F., Lanz, V. A., Äijälä, M., Allan, J. D., Carbone, S., Capes, G.,
 Ceburnis, D., Dall'Osto, M., Day, D. A., DeCarlo, P. F., Ehn, M., Eriksson, A., Freney, E.,
- Hildebrandt Ruiz, L., Hillamo, R., Jimenez, J. L., Junninen, H., Kiendler-Scharr, A.,
- 707 Kortelainen, A.-M., Kulmala, M., Laaksonen, A., Mensah, A. A., Mohr, C., Nemitz, E.,
- O'Dowd, C., Ovadnevaite, J., Pandis, S. N., Petäjä, T., Poulain, L., Saarikoski, S.,
- 709 Sellegri, K., Swietlicki, E., Tiitta, P., Worsnop, D. R., Baltensperger, U., and Pr év ôt, A. S.
- 710 H.: Organic aerosol components derived from 25 AMS data sets across Europe using a
- consistent ME-2 based source apportionment approach, Atmos. Chem. Phys., 14,
 6159-6176, doi:10.5194/acp-14-6159-2014, 2014.
- Cubison, M. J., Ortega, A. M., Hayes, P. L., Farmer, D. K., Day, D., Lechner, M. J.,
 Brune, W. H., Apel, E., Diskin, G. S., Fisher, J. A., Fuelberg, H. E., Hecobian, A.,
 Knapp, D. J., Mikoviny, T., Riemer, D., Sachse, G. W., Sessions, W., Weber, R. J.,
 Weinheimer, A. J., Wisthaler, A., and Jimenez, J. L.: Effects of aging on organic aerosol
 from open biomass burning smoke in aircraft and laboratory studies, Atmos. Chem.
 Phys., 11, 12049-12064, doi:10.5194/acp-11-12049-2011, 2011.

719 DeCarlo, P. F., Dunlea, E. J., Kimmel, J. R., Aiken, A. C., Sueper, D., Crounse, J., Wennberg, P. O., Emmons, L., Shinozuka, Y., Clarke, A., Zhou, J., Tomlinson, J., 720 Collins, D. R., Knapp, D., Weinheimer, A. J., Montzka, D. D., Campos, T., 721 and 722 Jimenez, J. L.: Fast airborne aerosol size and chemistry measurements above Mexico City and Central Mexico during the MILAGRO campaign, Atmos. Chem. Phys., 8, 723 4027-4048, doi:10.5194/acp-8-4027-2008, 2008. 724

- DeCarlo, P. F., Ulbrich, I. M., Crounse, J., de Foy, B., Dunlea, E. J., Aiken, A. C., Knapp, D.,
 Weinheimer, A. J., Campos, T., Wennberg, P. O., and Jimenez, J. L.: Investigation of the
 sources and processing of organic aerosol over the Central Mexican Plateau from
 aircraft measurements during MILAGRO, Atmos. Chem. Phys., 10, 5257-5280,
 doi:10.5194/acp-10-5257-2010, 2010.
- de Gouw, J. A., Middlebrook, A. M., Warneke, C., Goldan, P. D., Kuster, W. C., Roberts, J.

M., Fehsenfeld, F. C., Worsnop, D. R., Canagaratna, M. R., Pszenny, A. A. P., Keene, W.
C., Marchewka, M., Bertman, S. B., and Bates, T. S.: Budget of organic carbon in a
polluted atmosphere: Results from the New England Air Quality Study in 2002, J.
Geophys. Res. Atmos., 110, D16305, 2005.

- 735 Ding, A. J., Fu, C. B., Yang, X. Q., Sun, J. N., Pet äj ä, T., Kerminen, V.-M., Wang, T., Xie, Y.,
- Herrmann, E., Zheng, L. F., Nie, W., Liu, Q., Wei, X. L., and Kulmala, M.: Intense
 atmospheric pollution modifies weather: a case of mixed biomass burning with fossil
 fuel combustion pollution in eastern China, Atmos. Chem. Phys., 13, 10545-10554,
 doi:10.5194/acp-13-10545-2013, 2013.
- Draxler, R. R., Rolph, G. D.: HYSPLIT (HYbrid Single-Particle Lagrangian Integrated
 Trajectory) Model Access via NOAA ARL READY Website. NOAA Air Resources
- Laboratory, Silver Spring, MD. <u>http://www.arl.noaa.gov/ready/hysplit4.html</u>, 2003.

- 743 Du, H. H., Kong, L.D., Cheng, T.T., Chen, J.M., Du, J.F., Li, L., Xia, X.G., Leng, C.P., Huang,
- G. H.: Insights into summertime haze pollution events over Shanghai based on online
 water-soluble ionic composition of aerosols, Atmos. Environ., 45, 5131-5137, 2011.
- Dunlea, E. J., DeCarlo, P. F., Aiken, A. C., Kimmel, J. R., Peltier, R. E., Weber, R. J.,
- Tomlinson, J., Collins, D. R., Shinozuka, Y., McNaughton, C. S., Howell, S. G., Clarke,
- A. D., Emmons, L. K., Apel, E. C., Pfister, G. G., van Donkelaar, A., Martin, R. V.,
- Millet, D. B., Heald, C. L., and Jimenez, J. L.: Evolution of Asian aerosols during
 transpacific transport in INTEX-B, Atmos. Chem. Phys., 9, 7257-7287,
 doi:10.5194/acp-9-7257-2009, 2009.
- Ge, W. Z., Chen, R. J., Song, W. M., and Kan, H. D.: Daily visibility and hospital admission
 in Shanghai, China, Biomed Environ. Sci., 24(2): 117-121, 2011.
- 754 Gilardoni, S., Liu, S., Takahama, S., Russell, L. M., Allan, J. D., Steinbrecher, R., Jimenez, J.
- L., De Carlo, P. F., Dunlea, E. J., and Baumgardner, D.: Characterization of organic ambient aerosol during MIRAGE 2006 on three platforms, Atmos. Chem. Phys., 9, 5417-5432, doi:10.5194/acp-9-5417-2009, 2009.
- Grieshop, A. P., Donahue, N. M., and Robinson, A. L.: Laboratory investigation of
 photochemical oxidation of organic aerosol from wood fires 2: analysis of aerosol mass
 spectrometer data, Atmos. Chem. Phys., 9, 2227-2240, doi:10.5194/acp-9-2227-2009,
 2009.
- He, L.-Y., Lin, Y., Huang, X.-F., Guo, S., Xue, L., Su, Q., Hu, M., Luan, S.-J., and
 Zhang, Y.-H.: Characterization of high-resolution aerosol mass spectra of primary
 organic aerosol emissions from Chinese cooking and biomass burning, Atmos. Chem.
 Phys., 10, 11535-11543, doi:10.5194/acp-10-11535-2010, 2010.
- Healy, R. M., Sciare, J., Poulain, L., Crippa, M., Wiedensohler, A., Pr év α̂, A. S. H.,
 Baltensperger, U., Sarda-Est ève, R., McGuire, M. L., Jeong, C.-H., McGillicuddy, E.,

768	O'Connor, I. P., Sodeau, J. R., Evans, G. J., and Wenger, J. C.: Quantitative
769	determination of carbonaceous particle mixing state in Paris using single-particle mass
770	spectrometer and aerosol mass spectrometer measurements, Atmos. Chem. Phys., 13,
771	9479-9496, doi:10.5194/acp-13-9479-2013, 2013.

- Heringa, M. F., DeCarlo, P. F., Chirico, R., Tritscher, T., Dommen, J., Weingartner, E., Richter,
 R., Wehrle, G., Pr év α̂, A. S. H., and Baltensperger, U.: Investigations of primary and
 secondary particulate matter of different wood combustion appliances with a
 high-resolution time-of-flight aerosol mass spectrometer, Atmos. Chem. Phys., 11,
 5945-5957, doi:10.5194/acp-11-5945-2011, 2011.
- Hu, W. W., Hu, M., Yuan, B., Jimenez, J. L., Tang, Q., Peng, J. F., Hu, W., Shao, M., 777 Wang, M., Zeng, L. M., Wu, Y. S., Gong, Z. H., Huang, X. F., and He, L. Y.: Insights on 778 779 organic aerosol aging and the influence of coal combustion at a regional receptor site of central eastern China, Atmos. Chem. Phys., 13, 10095-10112, 780 doi:10.5194/acp-13-10095-2013, 2013. 781

Huang, X.-F., He, L.-Y., Hu, M., Canagaratna, M. R., Sun, Y., Zhang, Q., Zhu, T., Xue, L.,
Zeng, L.-W., Liu, X.-G., Zhang, Y.-H., Jayne, J. T., Ng, N. L., and Worsnop, D. R.:
Highly time-resolved chemical characterization of atmospheric submicron particles
during 2008 Beijing Olympic Games using an Aerodyne High-Resolution Aerosol Mass
Spectrometer, Atmos. Chem. Phys., 10, 8933-8945, doi:10.5194/acp-10-8933-2010,
2010.

Huang, X.-F., He, L.-Y., Hu, M., Canagaratna, M. R., Kroll, J. H., Ng, N. L., Zhang, Y.-H.,
Lin, Y., Xue, L., Sun, T.-L., Liu, X.-G., Shao, M., Jayne, J. T., and Worsnop, D. R.:
Characterization of submicron aerosols at a rural site in Pearl River Delta of China using
an Aerodyne High-Resolution Aerosol Mass Spectrometer, Atmos. Chem. Phys., 11,
1865-1877, doi:10.5194/acp-11-1865-2011, 2011.

- 793 Huang, X.-F., He, L.-Y., Xue, L., Sun, T.-L., Zeng, L.-W., Gong, Z.-H., Hu, M., and Zhu, T.: Highly time-resolved chemical characterization of atmospheric fine particles during 794 12, 4897-4907, 795 2010 Shanghai World Expo, Atmos. Chem. Phys., 796 doi:10.5194/acp-12-4897-2012, 2012.
- 797 Huang, X.-F., Xue, L., Tian, D.-X., Shao, W.-W., Sun, T.-L., Gong, Z.-H., Ju, W.-W., Jiang,
- B., Hu, M., and He, L.-Y.: Highly time-resolved carbonaceous aerosol characterization
- in Yangtze River Delta of China: Composition, mixing state and secondary formation,
 Atmos. Environ., 64, 200-207, 2013.
- Huffman, J. A., Docherty, K. S., Aiken, A. C., Cubison, M. J., Ulbrich, I. M., DeCarlo, P. F.,
 Sueper, D., Jayne, J. T., Worsnop, D. R., Ziemann, P. J., and Jimenez, J. L.:
 Chemically-resolved aerosol volatility measurements from two megacity field studies,
 Atmos. Chem. Phys., 9, 7161-7182, doi:10.5194/acp-9-7161-2009, 2009.
- Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Pr év ôt, A. S. H., Zhang, Q., Kroll, J. H.,
- BOG DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I.
- 807 M., Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V.
- A., Hueglin, C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J.,
- 809 Vaattovaara, P., Ehn, M., Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J.,
- E, Dunlea, J., Huffman, J. A., Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K.,
- 811 Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K.,
- 812 Salcedo, D., Cottrell, L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimono,
- A., Sun, J. Y., Zhang, Y. M., Dzepina, K., Kimmel, J. R., Sueper, D., Jayne, J. T.,
- Herndon, S. C., Trimborn, A. M., Williams, L. R., Wood, E. C., Middlebrook, A. M.,
- Kolb, C. E., Baltensperger, U., and Worsnop, D. R.: Evolution of organic aerosols in the
- atmosphere, Science, 326, 1525-1529, 2009.

- Justice, C. O., Giglio, L., Korontzi, S., Owens, J., Morisette, J.T., Roy, D., Descloitres, J.,
 Alleaume, S., Petitcolin, F., Kaufman, Y.: The MODIS fire products, Remote Sens.
 Environ., 83, 244-262, 2002.
- Kaufman, Y. J., Tanre, D., and Boucher, O.: A satellite view of aerosols in the climate system,
 Nature, 419: 215-23, 2002.
- Kaufman, Y. J., Ichoku, C., Giglio, L., Korontzi, S., Chu, D. A., Hao, W. M., Li R. R., Justice,
 C. O.: Fire and smoke observed from the Earth Observing System MODIS
 instrument-products, validation, and operational use, Int. J. Remote Sensing, 24,
 1765-781, 2003.
- Kawamura, K., Tachibana, E., Okuzawa, K., Aggarwal, S. G., Kanaya, Y., and Wang, Z. F.:
 High abundances of water-soluble dicarboxylic acids, ketocarboxylic acids and
 α-dicarbonyls in the mountaintop aerosols over the North China Plain during wheat
 burning season, Atmos. Chem. Phys., 13, 8285-8302, doi:10.5194/acp-13-8285-2013,
 2013.
- Lanz, V. A., Alfarra, M. R., Baltensperger, U., Buchmann, B., Hueglin, C., and
 Pr év ât, A. S. H.: Source apportionment of submicron organic aerosols at an urban site by
 factor analytical modelling of aerosol mass spectra, Atmos. Chem. Phys., 7, 1503-1522,
 doi:10.5194/acp-7-1503-2007, 2007.
- Lathem, T. L., Beyersdorf, A. J., Thornhill, K. L., Winstead, E. L., Cubison, M. J., Hecobian,
- A., Jimenez, J. L., Weber, R. J., Anderson, B. E., and Nenes, A.: Analysis of CCN
- activity of Arctic aerosol and Canadian biomass burning during summer 2008, Atmos.
 Chem. Phys., 13, 2735-2756, doi:10.5194/acp-13-2735-2013, 2013.
- Lee, T., Sullivan, A. P., Mack, L., Jimenez, J. L., Kreidenweis, S. M., Onasch, T. B., Worsnop,
 D. R., Malm, W., Wold, C. E., Hao, W. M., and Collett, J. L.: Chemical smoke marker

- emissions during flaming and smoldering phases of laboratory open burning of wildland
 fuels, Aerosol Sci. Tech., 44, 1-5, 2010.
- Li, X. H., Wang, S. X., Duan, L., Hao, J. M., Li, Y. S., and Yang, L.: Particulate and trace gas
 emissions from open burning of wheat straw and corn stover in China, Environ. Sci.
 Technol., 41 (17): 6052-6058, 2007.
- Lipsky, E. M. and Robinson, A. L.: Effects of dilution on fine particle mass and partitioning
 of semivolatile organics in diesel exhaust and wood smoke, Environ. Sci. Technol., 40,
 155–162, 2006.
- Liu, D., Allan, J., Corris, B., Flynn, M., Andrews, E., Ogren, J., Beswick, K., Bower, K., 849 Burgess, R., Choularton, T., Dorsey, J., Morgan, W., Williams, P. I., and Coe, H.: 850 Carbonaceous aerosols contributed by traffic and solid fuel burning at a polluted rural 851 site in Northwestern England, Atmos. Chem. Phys., 11, 1603-1619. 852 doi:10.5194/acp-11-1603-2011, 2011. 853
- Liu, D., Allan, J. D., Young, D. E., Coe, H., Beddows, D., Fleming, Z. L., Flynn, M. J.,
 Gallagher, M. W., Harrison, R. M., Lee, J., Prevot, A. S. H., Taylor, J. W., Yin, J.,
 Williams, P. I., and Zotter, P.: Size distribution, mixing state and source apportionments
 of black carbon aerosols in London during winter time, Atmos. Chem. Phys. Discuss.,
 14, 16291-16349, doi:10.5194/acpd-14-16291-2014, 2014.
- Maenhaut, W., Vermeylen, R., Claeys, M., Vercauteren, J., Matheeussen, C., Roekens, E.:
 Assessment of the contribution from wood burning to the PM₁₀ aerosol in Flanders,
 Belgium, Sci. Total Environ., 437, 226-236, 2012.
- Middlebrook, A. M., Bahreini, R., Jimenez, J. L., and Canagaratna, M. R.: Evaluation of
 composition-dependent collection efficiencies for the aerodyne aerosol mass
 spectrometer using field data, Aerosol Sci. Tech., 46, 258-271, 2011.

Mohr, C., Huffman, J. A., Cubison, M. J., Aiken, A. C., Docherty, K. S., Kimmel, J. R.,
Ulbrich, I. M., Hannigan, M., and Jimenez, J. L.: Characterization of primary organic
aerosol emissions from meat cooking, trash burning, and motor vehicles with
High-Resolution Aerosol Mass Spectrometry and comparison with ambient and chamber
observations, Environ. Sci. Technol., 43, 2443-2449, 2009.

- Mohr, C., DeCarlo, P. F., Heringa, M. F., Chirico, R., Slowik, J. G., Richter, R., Reche, C., 870 Peñuelas, J., 871 Alastuey, A., Ouerol, X., Seco, R., Jim énez, J. L., Crippa, M., Zimmermann, R., Baltensperger, U., and Pr év ât, A. S. H.: Identification 872 and 873 quantification of organic aerosol from cooking and other sources in Barcelona using mass aerosol spectrometer data, Chem. Phys., 12, 1649-1665, 874 Atmos. doi:10.5194/acp-12-1649-2012, 2012. 875
- 876 Ng, N. L., Canagaratna, M. R., Zhang, Q., Jimenez, J. L., Tian, J., Ulbrich, I. M., Kroll, J. H., Docherty, K. S., Chhabra, P. S., Bahreini, R., Murphy, S. M., Seinfeld, J. H., 877 Hildebrandt, L., Donahue, N. M., DeCarlo, P. F., Lanz, V. A., Pr év ôt, A. S. H., Dinar, E., 878 879 Rudich, Y., and Worsnop, D. R.: Organic aerosol components observed in Northern Hemispheric datasets from aerosol mass spectrometry, Atmos. Chem. Phys., 10, 880 4625-4641, doi:10.5194/acp-10-4625-2010, 2010. 881
- Ng, N. L., Herndon, S. C., Trimborn, A., Canagaratna, M. R., Croteau, P. L., Onasch, T. B.,
 Sueper, D., Worsnop, D. R., Zhang, Q., Sun, Y. L., and Jayne, J. T.: An aerosol chemical
 speciation monitor (ACSM) for routine monitoring of the composition and mass
 concentrations of ambient aerosol, Aerosol Sci. Technol., 45: 7, 770 -784, 2011a.
- Ng, N. L., Canagaratna, M. R., Jimenez, J. L., Chhabra, P. S., Seinfeld, J. H., 886 and Worsnop, D. R.: Changes in organic aerosol composition with aging inferred from 887 aerosol Atmos. Chem. Phys., 11, 6465-6474, 888 mass spectra, doi:10.5194/acp-11-6465-2011, 2011b. 889

- Ng, N. L., Canagaratna, M. R., Jimenez, J. L., Zhang, Q., Ulbrich, I. M., and Worsnop, D. R.:
- Real-time methods for estimating organic component mass concentrations from aerosol
 mass spectrometer data, Environ. Sci. Technol., 45, 910-916, 2011c.
- Paatero, P.: Least squares formulation of robust non-negative factor analysis, Chemom. Intell.
 Lab. Syst., 37, 23-35, 1997.
- Paatero, P.: The multilinear engine A table-driven, least squares program for solving
 multilinear problems, including the n-way parallel factor analysis model, J. Comput.
 Graph. Stat., 8, 854–888, 1999.
- Ramanathan, V., Crutzen, P. J, Kiehl, J. T., and Rosenfeld, D.: Atmosphere, aerosols, climate,
 and the hydrological cycle, Science, 294, 2119–2124, 2001.
- Reche, C., Viana, M., Amato, F., Alastuey, A., Moreno, T., Hillamo, R., Teinila, K., Saarnio,
 K., Seco, R., Penuelas, J., Mohr, C., Pr év ôt, A. S. H., and Querol, X,.: Biomass burning
 contributions to urban aerosols in a coastal Mediterranean City, Sci. Total Environ., 427,
 175-190, 2012.
- Robinson, A. L., Donahue, N. M., Shrivastava, M. K., Weitkamp, E. A., Sage, A. M.,
 Grieshop, A. P., Lane, T. E., Pierce, J. R., and Pandis, S. N.: Rethinking organic aerosols:
 Semivolatile emissions and photochemical aging, Science, 315, 1259–1262,
 doi:10.1126/science.1133061, 2007.
- Sandradewi, J., Prevot, A. S. H., Szidat, S., Perron, N., Alfarra, M. R., Lanz, V. A.,
 Weingartner, E., and Baltensperger, U.: Using aerosol light absorption measurements for
 the quantitative determination of wood burning and traffic emission contributions to
 particulate matter, Environ. Sci. Technol., 42, 3316–3323, 2008.
- 912 Sun, J. Y., Zhang, Q., Canagaratna, M. R., Zhang, Y. M., Ng, N. L., Sun, Y. L., Jayne, J. T.,
- 213 Zhang, X. C., Zhang, X. Y., and Worsnop, D. R.: Highly time-and size-resolved

914 characterization of submicron aerosol particles in Beijing using an Aerodyne Aerosol
915 Mass Spectrometer, Atmos. Environ., 44, 131-140, 2010.

916 Sun, Y.-L., Zhang, Q., Schwab, J. J., Demerjian, K. L., Chen, W.-N., Bae, M.-S., Hung, H.-M.,

- Hogrefe, O., Frank, B., Rattigan, O. V., and Lin, Y.-C.: Characterization of the sources
 and processes of organic and inorganic aerosols in New York city with a high-resolution
 time-of-flight aerosol mass spectrometer, Atmos. Chem. Phys., 11, 1581-1602,
 doi:10.5194/acp-11-1581-2011, 2011a.
- 921 Sun, Y. L., Zhang, Q., Schwab, J. J., Chen, W. N., Bae, M. S., Lin, Y. C., Hung, H. M., and
- Demerjian, K. L.: A case study of aerosol processing and evolution in summer in New
 York City, Atmos. Chem. Phys., 11, 12737-12750, doi:10.5194/acp-11-12737-2011,
 2011b.
- Sun, Y. L., Wang, Z. F., Dong, H. B., Yang, T., Li, J., Pan, X. L., Chen, P., and Jayne, J. T.:
 Characterization of summer organic and inorganic aerosols in Beijing, China with an
 Aerosol Chemical Speciation Monitor, Atmos. Environ., 51, 250-259, 2012.
- 928 Sun, Y. L., Wang, Z. F., Fu, P. Q., Yang, T., Jiang, Q., Dong, H. B., Li, J., and Jia, J. J.:
- Aerosol composition, sources and processes during wintertime in Beijing, China, Atmos.
 Chem. Phys., 13, 4577-4592, doi:10.5194/acp-13-4577-2013, 2013.
- Takegawa, N., Miyakawa, T., Kondo, Y., Jimenez, J. L., Zhang, Q., Worsnop, D. R., and
 Fukuda, M.: Seasonal and diurnal variations of submicron organic aerosol in Tokyo
 observed using the Aerodyne aerosol mass spectrometer, J. Geophys. Res., 111, D11206,
 2006.
- Ulbrich, I. M., Canagaratna, M. R., Zhang, Q., Worsnop, D. R., and Jimenez, J. L.:
 Interpretation of organic components from Positive Matrix Factorization of aerosol mass
 spectrometric data, Atmos. Chem. Phys., 9, 2891–2918, doi:10.5194/acp-9-2891-2009,
 2009.

- 939 Wang, G. H., Kawamura, K., Xie, M. J., Hu, S. Y., Cao, J. J., An, Z. H., Waston J. G., and
- Chow, J. C.: Organic molecular compositions and size distributions of Chinese summer
 and autumn aerosols from Nanjing: characteristic haze event caused by wheat straw
 burning, Environ. Sci. Technol., 43 (17): 6493-6499, 2009a.
- Wang, Y. Q., Zhang, X. Y., and Draxler, R. R.: TrajStat: GIS-based software that uses various
 trajectory statistical analysis methods to identify potential sources from long-term air
 pollution measurement data, Environ. Modell. Softw., 24: 938-939, 2009b.
- Watson, J. G.: Visibility: Science and regulation, J. Air Waste Manage. Assoc., 52, 628-713,
 2002.
- 948 Yee, L. D., Kautzman, K. E., Loza, C. L., Schilling, K. A., Coggon, M. M., Chhabra, P. S.,
- Chan, M. N., Chan, A. W. H., Hersey, S. P., Crounse, J. D., Wennberg, P. O., Flagan, R.
 C., and Seinfeld, J. H.: Secondary organic aerosol formation from biomass burning
 intermediates: phenol and methoxyphenols, Atmos. Chem. Phys., 13, 8019-8043,
 doi:10.5194/acp-13-8019-2013, 2013.
- Zhang, H., Ye, X., Cheng, T., Chen, J., Yang, X., Wang, L., Zhang, R.: A laboratory study of
 agricultural crop residue combustion in China: Emission factors and emission inventory,
 Atmos. Environ., 42, 8432-8441, 2008.
- Zhang, Q., Worsnop, D. R., Canagaratna, M. R., and Jimenez, J. L.: Hydrocarbon-like and
 oxygenated organic aerosols in Pittsburgh: insights into sources and processes of organic
 aerosols, Atmos. Chem. Phys., 5, 3289-3311, doi:10.5194/acp-5-3289-2005, 2005a.
- 259 Zhang, Q., Alfarra, M.R. Worsnop, D., Allan, J. D., Coe, H., Cangaratna, M. R., Jimenez, J.
- L.: Deconvolution and quantification of hydrocarbon-like and oxygenated organic
 aerosols based on aerosol mass spectrometry, Environ. Sci. Technol., 39, 4938-4952,
 2005b.
- 263 Zhang, Q., Jimenez, J. L., Canagaratna, M. R., Allan, J. D, Coe, H., Ulbrich, I. M., Alfarra, M.

964	R., Takami, A., Middlebrook, A. M., Sun, Y. L., Dzepina, K., Dunlea, E., Docherty, K.,
965	DeCarlo, P, F., Salcedo, D., Onasch, T., Jayne, J. T., Miyoshi, T., Shimono, A.,
966	Hatakeyama, S., Takegawa, N., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S.,
967	Weimer, S., Demerjian, K., Williams, P., Bower, K., Bahreini, R., Cottrell, L., Griffin, R.
968	J., Rautiainen, J., Sun, J. Y., Zhang, Y. M., Worsnop, D. R.: Ubiquity and dominance of
969	oxygenated species in organic aerosols in anthropogenically-influenced Northern
970	Hemisphere midlatitudes, Geophys. Res. Lett., 34(13): L13801, 2007.
971	Zhang, Q., Jimenez, J. L., Canagaratna, M. R., Ulbrich, I. M., Ng, N. L., Worsnop, D. R., Sun,
972	Y. L.: Understanding atmospheric organic aerosols via factor analysis of aerosol mass
973	spectrometry: a review, Anal. Bioanal. Chem., 401:3045-3067, 2011.

975 Table 1

976 Mean mass concentrations (µg m⁻³) and standard deviation (S.D.) of PM₁ (NR-PM₁ + BC) species, OA

977	components and	meteorological t	factors (i.e.	WS, RH, and T)	during the harvest seasons.
	1	U	· · ·		0

	Summer harvest		Autumn harvest	
-	Mean	S.D.	Mean	S.D.
Aerosol species				
NO ₃	9.0	7.1	9.2	6.2
SO_4	5.0	2.4	4.7	2.5
NH_4	7.0	3.5	6.4	3.5
Chl	0.4	0.9	0.7	0.8
OA	15.4	12.8	22.3	17.5
BC	3.2	2.2	6.0	3.8
PM_1	38.5	24.3	46.4	27.0
OA component	S			
HOA + COA	2.2	2.4	5.7	7.6
BBOA	1.1	1.0	1.5	1.6
OOA-BB	4.1	4.6	6.5	7.3
OOA	7.1	3.6	6.6	3.2
Meteorological	factors			
WS (m s^{-1})	3.5	1.7	2.6	1.4
RH (%)	70.7	15.3	54.3	13.7
T (°C)	24.1	4.1	18.1	3.5



Fig. 1. Time series of (a) wind speed (WS) and wind direction (WD); (b) relative humidity (RH), temperature (*T*) and precipitation (Precip); (c) submicron aerosol species, i.e., organic aerosol (OA), ammonium (NH₄), nitrate (NO₃), sulfate (SO₄), chloride (Chl) and black carbon (BC); and (d) mass fraction during the harvest seasons. Three case events are marked and discussed in the text.



Fig. 2. Diurnal variation patterns of meteorological factors (i.e. RH, *T*, and WS), PM₁ species including
organic aerosol (OA), nitrate (NO₃), sulfate (SO₄), ammonium (NH₄), chloride (Chl), and black carbon
(BC) during the harvest seasons.



990

Fig. 3. Mass spectra profiles (left) and diurnal variations (right) of four OA factors, i.e., hydrocarbon-like and cooking-emission related OA (HOA + COA), fresh biomass burning (BB) OA (BBOA), oxygenated BB-influenced OA (OOA-BB), and highly oxygenated OA (OOA). Note that: reference mass spectra (MS) are obtained from the results by Crippa et al. (2013), and oxygenated BBOA components have been resolved (OOA₂-BBOA).



997

Fig. 4. Time series of OA factors (left) and relevant tracer species (right): (**a**) HOA + COA; (**b**) BBOA and a surrogate of levoglucosan (m/z 60) and potassium ion (K⁺); (**c**) OOA-BB, nitrate and chloride; (**d**) OOA and SO₄. Note that different BBOA mass concentrations for low biomass burning period (L-BB), medium biomass burning period (M-BB), and high biomass burning period (H-BB).



Fig. 5. Comparison of biomass burning-related PMF factors (BBOA and OOA-BB) and biomass related species: (a) K⁺; and (b) the ACSM m/z 60 minus 0.26% × OA (applied metric of background $f_{60} = 0.26$ %

- 1006 of OA is discussed in the section 3.4 of the text) during the summer and autumn harvest.
- 1007



1008 $NO_3 (\mu g m^2)$ $SO_4 (\mu g m^2)$ $SO_4 (\mu g m^2)$ 1009 Fig. 6. Comparison of two kinds of oxygenated OA (OOA-BB and OOA) and two kinds of secondary 1010 inorganic species, i.e., nitrate (NO₃) and sulfate (SO₄), during the harvest seasons. Colored by the f_{60} as a 1011 biomass-burning marker. The three dashed lines in the plot refer to 2 : 1, 1 : 1, and 1 : 2 lines, respectively. 1012





Fig. 7. Average relative contributions of PM₁ species and OA components for low biomass burning period
(L-BB), medium biomass burning period (M-BB), and high biomass burning period (H-BB), as well as
entire period during the harvest seasons and other sites including mega-cities (Mexico city, Paris, Beijing
and Shanghai), suburban area (Jiaxing), remote background site (Crete), and PRD (Pearl River Delta,
China). Note that OOA in this plot includes OOA₂-BBOA in Paris.



1021 Fig. 8. The mass fractions of PM1 species and OA components as a function of PM1 mass loadings (left), and probability density of PM_1 mass loadings (right, with the white lines in the plots) during the summer and autumn harvest respectively.



1026 **Fig. 9.** Summary plots showing f_{44} vs. f_{60} for measurements with little or negligible biomass burning 1027 influence. Colored by the summer harvest (blue), autumn harvest (orange), and little or negligible biomass 1028 burning influence period (gray, July 1 to 8, 2013), respectively. Also shown is the average background 1029 level of f_{60} (~0.3%, red dashed line) in other studies from Aiken et al. (2009) and Cubison et al. (2011) for 1030 references.

1031



Fig. 10. The (BBOA + OOA-BB) / Δ CO ratio as a function of f_{44} during the summer and autumn harvest.







1038Fig. 11. Summary plots showing (a) triangle plot (f_{44} vs. f_{43}), SV-OOA and LV-OOA indicate semi-volatile1039OOA and low-volatility OOA respectively. The dots are colored by f_{60} as a biomass-burning marker; (b) f_{44}

1040 as a function of f_{60} (f_{44} vs. f_{60}), colored by the PM₁ mass concentration and sized by the OA loadings.





Fig. 12. Average composition of PM₁ (pie charts) and OA factors (bar charts) for each cluster. The four clusters are: (**a**) northeasterly (NE) back-trajectories (BTs), easterly marine (EM) BTs, southeasterly marine (SEM) BTs and westerly continental (WC) during the summer harvest; and (**b**) northerly continental (NC) BTs, northeasterly marine (NEM) BTs, easterly marine (EM) BTs and southerly continental (SC) during the autumn harvest. The markers on the trajectories indicate 6 h interval.