

1 **Organic aerosol concentration and composition over Europe:**
2 **Insights from comparison of regional model predictions with aerosol**
3 **mass spectrometer factor analysis**

4
5 **C. Fountoukis¹, A.G. Megaritis², K. Skyllakou², P.E. Charalampidis³, C. Pilinis³,**
6 **H.A.C. Denier van der Gon⁴, M. Crippa^{5,18}, F. Canonaco⁵, C. Mohr^{5,19}, A.S.H.**
7 **Prévôt⁵, J. D. Allan^{6,7}, L. Poulain⁸, T. Petäjä⁹, P. Tiitta^{10,11}, S. Carbone¹², A.**
8 **Kiendler-Scharr¹³, E. Nemitz¹⁴, C. O'Dowd¹⁵, E. Swietlicki¹⁶ and S.N. Pandis^{2,17,*}**

- 9
10 [1] Institute of Chemical Engineering Sciences, Foundation for Research and
11 Technology Hellas (FORTH), Patras, Greece
12 [2] Department of Chemical Engineering, University of Patras, Patras, Greece
13 [3] Department of Environment, University of the Aegean, Mytilene, Greece
14 [4] Netherlands Organization for Applied Scientific Research TNO, Utrecht, The
15 Netherlands
16 [5] Laboratory of Atmospheric Chemistry, Paul Scherrer Institute, PSI Villigen,
17 Switzerland
18 [6] National Centre for Atmospheric Science, University of Manchester, Manchester,
19 UK
20 [7] School of Earth, Atmospheric and Environmental Sciences, University of
21 Manchester, Manchester, UK
22 [8] Leibniz Institute for Tropospheric Research, Leipzig, Germany
23 [9] Department of Physics, University of Helsinki, Finland
24 [10] Department of Environmental Science, Univ. of Eastern Finland, Kuopio,
25 Finland
26 [11] Department of Applied Physics, Univ. of Eastern Finland, Kuopio, Finland
27 [12] Atmospheric Composition Research, Finnish Meteorological Institute, Helsinki,
28 Finland
29 [13] Institut für Energie- und Klimaforschung: Troposphäre (IEK 8),
30 Forschungszentrum Jülich GmbH, Jülich, Germany
31 [14] Centre for Ecology and Hydrology, Bush Estate, Penicuik, Midlothian, EH26
32 0QB, UK
33 [15] School of Physics & Centre for Climate & Air Pollution Studies, National
34 University of Ireland Galway, Ireland
35 [16] Division of Nuclear Physics, University of Lund, 221 00 Lund, Sweden
36 [17] Department of Chemical Engineering, Carnegie Mellon University, Pittsburgh,
37 USA
38 [18] Now at: EC Joint Research Centre (JRC), Inst. Environment & Sustainability,
39 Via Fermi, Ispra, Italy
40 [19] Now at: Department of Atmospheric Sciences, University of Washington, Seattle
41 WA 98195, USA
42
43

44 *Correspondence to: S. N. Pandis (spyros@andrew.cmu.edu)

45 **Abstract**

46 A detailed three-dimensional regional chemical transport model (PMCAMx)
47 was applied over Europe focusing on the formation and chemical transformation of
48 organic matter. Three periods representative of different seasons were simulated,
49 corresponding to intensive field campaigns. An extensive set of AMS measurements
50 was used to evaluate the model and, using factor analysis results, gain more insight
51 into the sources and transformations of organic aerosol (OA). Overall, the agreement
52 between predictions and measurements for OA concentration is encouraging with the
53 model reproducing two thirds of the data (daily average mass concentrations) within a
54 factor of two. Oxygenated OA (OOA) is predicted to contribute 93% to total OA
55 during May, 87% during winter and 96% during autumn with the rest consisting of
56 fresh primary OA (POA). Predicted OOA concentrations compare well with the
57 observed OOA values for all periods with an average fractional error of 0.53 and a
58 bias equal to -0.07 (mean error = $0.9 \mu\text{g m}^{-3}$, mean bias = $-0.2 \mu\text{g m}^{-3}$). The model
59 systematically underpredicts fresh POA in most sites during late spring and autumn
60 (mean bias up to $-0.8 \mu\text{g m}^{-3}$). Based on results from a source apportionment
61 algorithm running in parallel with PMCAMx, most of the POA originates from
62 biomass burning (fires and residential wood combustion) and therefore biomass
63 burning OA is most likely underestimated in the emission inventory. The sensitivity
64 of POA predictions to the corresponding emissions' volatility distribution is
65 discussed. The model performs well at all sites when the PMF-estimated low volatility
66 OOA is compared against the OA with saturation concentrations of the OA surrogate
67 species $C^* \leq 0.1 \mu\text{g m}^{-3}$ and semivolatile OOA against the OA with $C^* > 0.1 \mu\text{g m}^{-3}$
68 respectively.

69 **1. Introduction**

70 Organic aerosol (OA) is a significant component (20 – 90%) of atmospheric
71 fine particulate matter (Zhang et al., 2007) and thus strongly affects the
72 physicochemical properties of aerosols. Despite its importance, OA remains today the
73 least understood component of the atmospheric aerosol system. OA has hundreds of
74 sources, both anthropogenic and natural, while it can undergo complex atmospheric
75 chemical and physical processing (Hallquist et al., 2009). The description of these
76 emissions and processes in Chemical Transport Models (CTMs) is not a trivial task.

77 OA has been traditionally characterized as either primary (POA) or secondary
78 OA (SOA). POA is introduced in the atmosphere in the particulate phase while SOA
79 is formed from oxidation products of volatile organic compounds (VOCs). Murphy
80 and Pandis (2009) classified OA into “fresh” POA, oxidized POA (OPOA),
81 anthropogenic SOA (aSOA) and biogenic SOA (bSOA). Fresh POA is primary OA
82 that is emitted in the particulate phase and has not undergone chemical processing,
83 while OPOA refers to POA compounds that evaporate and undergo oxidation in the
84 gas phase, which allows them to reduce their volatility and re-condense back to the
85 particulate phase. SOA produced from the oxidation of intermediate volatility
86 compounds (IVOCs) was also included in OPOA mainly because the IVOC emissions
87 were calculated based on the POA emissions. SOA (either aSOA or bSOA) consists
88 of compounds of low volatility that are formed when VOCs are oxidized in the gas
89 phase (Kroll et al., 2011; Hallquist et al., 2009; Robinson et al., 2007). Additional
90 oxidation of the semivolatile SOA, POA and OPOA components in the gas phase is
91 known as “aging” of OA which, assisted by transport, can increase OA concentrations
92 in areas far away from sources and is responsible for the regional distribution of OA.

93 Quantification, characterization and speciation of organic aerosol are hindered
94 by analytical difficulties (Kanakidou et al., 2005; Turpin et al., 2000). For instance,
95 conventional techniques (e.g., GC-MS) can only speciate a small fraction of the OA
96 mass. Recently, several new measurement techniques have emerged that can quantify
97 and characterize to some degree all the OA mass present in fine aerosol. The Aerosol
98 Mass Spectrometer (AMS) is used to measure the size-resolved mass concentration
99 and total mass spectrum of organic aerosols with high time resolution (Canagaratna et
100 al., 2007). Information about processes or sources contributing to the OA levels can
101 be provided by the Positive Matrix Factorization (PMF) method (Paatero and Tapper,
102 1994; Lanz et al., 2007; Ulbrich et al., 2009; Ng et al., 2010), the multi-linear engine
103 (ME-2) (Lanz et al., 2008, Canonaco et al., 2013) or custom principal component
104 analysis (Zhang et al., 2005) of the AMS measurements. These methods allow a
105 classification of the OA into different types based on their different temporal and
106 mass spectral signatures. Two major components often resolved by the analysis of the
107 AMS measurements (Zhang et al., 2005) are hydrocarbon-like organic aerosol (HOA)
108 and oxygenated organic aerosol (OOA). HOA represents fresh POA from fossil fuel
109 combustion while OOA is OA of secondary nature. Often, factor analysis can further
110 classify OOA into a more oxygenated low-volatility OOA component (LV-OOA) and
111 a less oxygenated semi-volatile OOA part (SV-OOA) (Lanz et al., 2007; Ulbrich et
112 al., 2009; Jimenez et al., 2009; Ng et al., 2010; Crippa et al., 2013). Biomass burning
113 OA (BBOA), marine-related OA (MOA) and cooking OA (COA) are other classes
114 that the factor analysis may identify (Zhang et al., 2007; Crippa et al., 2013).

115 Earlier measurement campaigns in Europe have shown that biomass burning
116 and fossil fuel combustion are the main sources of OA in winter, while secondary OA
117 from non-fossil sources is dominant during summer and in European boreal regions

118 (Szidat et al., 2006; Tunved et al., 2006; Gelencser et al., 2007). More recently Lanz
119 et al. (2010) analyzed AMS data (using PMF and ME-2) in Central Europe and found
120 a strong impact of BBOA in the Alpine region. Morgan et al. (2010) performed
121 aircraft AMS measurements in North-Western Europe and found significant chemical
122 processing of OA downwind of major pollution sources, with the LV-OOA
123 component becoming increasingly dominant as the distance from source and
124 photochemical age increased. AMS measurements at Finokalia, a remote site in the
125 Mediterranean (influenced by air masses from different source regions), revealed two
126 OOA components which, however, did not appear to correspond to different OA
127 sources but instead to different limits of the extent of OA oxidation observed during
128 the campaign (Hildebrandt et al., 2010).

129 Air quality and climate models have until recently treated POA emissions as
130 non-volatile and non-reactive while SOA formation from VOCs is usually simulated
131 using a semi-empirical 2-product oxidation parameterization. Recent advancements in
132 OA modeling have introduced the volatility basis set framework (VBS) (Donahue et
133 al., 2006) in which POA is considered semi-volatile and photochemically reactive and
134 the saturation concentrations (C^*) of OA surrogate species are fixed and usually range
135 from $0.01 \mu\text{g m}^{-3}$ up to $10^6 \mu\text{g m}^{-3}$ with logarithmically spaced bins. Recently, CTMs
136 have successfully adopted the new developments in OA modeling improving
137 predictions of OA when compared to measurements (Murphy and Pandis, 2009;
138 Tsimpidi et al 2010; Fountoukis et al., 2011; Bergstrom et al., 2012; Zhang et al.,
139 2013). However, uncertainties still exist in these models regarding the volatility
140 distribution of the primary OA emissions, the simplistic parameterization of the
141 chemical aging of the OA or errors in either the anthropogenic or biogenic emissions
142 (Fountoukis et al., 2011; Bergstrom et al., 2012). Bergstrom et al. (2012) applied the

143 EMEP model with the VBS framework over Europe during 2002-2007. A comparison
144 with AMS data from one campaign in Switzerland (during June 2006) was conducted
145 while other long-term datasets were also used for the model evaluation. They
146 concluded that the volatility distribution of primary emissions and the emissions of
147 biogenic VOCs are two main sources of uncertainty in the model. Zhang et al. (2013)
148 applied the CHIMERE CTM over the greater Paris area and conducted a thorough
149 comparison with PMF analyzed AMS data for a summer period. They found that
150 adopting the new paradigm in OA modeling significantly improves model predictions
151 of SOA, while depending on the emission inventory used SOA levels tend to be
152 overestimated. This was attributed to emissions errors, the choice of (potentially high)
153 yields or uncertainty in the chemical aging of biogenic species. Interestingly, it was
154 found that if some OPOA had been included in the HOA estimated by the PMF, the
155 model bias would be reduced.

156 The factor analysis of AMS measurements can allow more in-depth evaluation
157 of CTMs and further constrain the corresponding uncertain parameters. However, to
158 date such model-measurement comparison studies on a regional scale are rare. In this
159 work we apply PMCAMx (Fountoukis et al., 2011; 2013) over Europe during 3
160 periods, representative of different seasons, and use an extensive set of AMS
161 measurements to evaluate the model. Using factor analysis data, we attempt to gain
162 more insight into the formation and evolution of OA, as well as to identify strengths
163 and limitations of the current OA modeling framework.

164

165 **2. OA simulation in PMCAMx**

166 A regional chemical transport model, PMCAMx (Karydis et al., 2010;
167 Fountoukis et al., 2011; 2013), is used in this study. The model describes the

168 processes of horizontal and vertical dispersion, horizontal and vertical advection, gas-
169 and aqueous-phase chemistry, wet and dry deposition. The gas-phase chemistry
170 mechanism is based on the SAPRC99 mechanism (Environ, 2003; Carter, 2010). For
171 the aerosol processes, bulk equilibrium is assumed. More details about this version of
172 the model can be found in Fountoukis et al. (2011).

173 The OA scheme in PMCAMx treats all organic species (primary and
174 secondary) as chemically reactive using the Volatility Basis Set (VBS) approach
175 (Donahue et al., 2006). Primary OA in PMCAMx is assumed to be semivolatile
176 (Shrivastava et al., 2008) with nine surrogate POA species used, corresponding to
177 nine effective saturation concentrations ranging from 10^{-2} to 10^6 $\mu\text{g m}^{-3}$ (at 298 K) in
178 logarithmically spaced bins. POA is simulated in the model as “fresh” (unoxidized)
179 POA (fPOA) and oxidized POA (OPOA). The IVOC emissions are assumed to be
180 proportional to the emitted primary OA mass (1.5 times POA emissions) (Tsimpidi et
181 al., 2010; Shrivastava et al., 2008), since the existing emission inventories do not
182 include these compounds. The products of oxidation of IVOCs are allowed to
183 partition between gas and particle phase according to their volatilities, forming OPOA
184 in the aerosol phase. BBOA is not simulated separately but is included in the fresh
185 POA, while processed (oxidized) BBOA is included in OPOA. The POA emissions
186 are assumed to have the volatility distribution used by Tsimpidi et al. (2010). Recently
187 May et al. (2013a; 2013b; 2013c) have estimated volatility distributions for the major
188 OA sources. The sensitivity of our results to assumed POA volatility distribution is
189 explored in Section 4.

190 The SOA is described using four volatility bins ($1, 10, 10^2, 10^3$ $\mu\text{g m}^{-3}$)
191 following Lane et al. (2008a). The SOA module incorporates NO_x -dependent SOA
192 yields (Lane et al., 2008b) and includes anthropogenic aerosol yields based on the

193 studies of Ng et al. (2006) and Hildebrandt et al. (2009). The modeled oxygenated OA
194 is defined as the sum of aSOA, bSOA and OPOA.

195 Chemical aging is modeled through gas-phase oxidation of organic
196 compounds assuming a gas-phase OH reaction with a rate constant of $k=1\times 10^{-11}$ cm³
197 molec⁻¹ s⁻¹ for anthropogenic SOA and $k=4\times 10^{-11}$ cm³ molec⁻¹ s⁻¹ for the primary OA
198 and the IVOCs (Murphy and Pandis, 2009). Each reaction is assumed to effectively
199 decrease the volatility of the compound by one order of magnitude. The base-case
200 simulation assumes that the chemical aging reactions of biogenic SOA (including both
201 functionalization and fragmentation reactions) do not result in a net change of the
202 volatility distribution (and hence the bSOA concentration) and thus it effectively
203 neglects the chemical aging of biogenic SOA. The sensitivity of the model predictions
204 to this assumption will be investigated in a subsequent section.

205 The VBS approach was used by Skyllakou et al. (2013) to extend the Particle
206 Source Apportionment Algorithm (PSAT) of Wagstrom et al. (2008) so that the
207 semivolatile POA and the continuous aging of OA can be considered in source-
208 receptor analyses. In this study we applied the extended PSAT in PMCAMx in the
209 European domain. PSAT keeps track of the sources of all OA components (fresh
210 primary, oxidized primary, anthropogenic and biogenic secondary) which are
211 distributed into different volatility bins and are tracked by PSAT as separate species.

212

213 **3. Model application**

214 We simulate 3 periods (1 – 29 May 2008, 25 February – 24 March 2009, and
215 15 September – 17 October 2008) during which intensive campaigns were performed
216 as part of EUCAARI (Kulmala et al., 2009; 2011) and EMEP (Tørseth et al., 2012).
217 During all three campaigns, AMS measurements were performed at several sites

218 across Europe while further analysis of the OA sources was also performed using
219 factor analysis techniques (Crippa et al., 2013). Crippa et al. (2013) proposed a
220 standardized methodology, tailored to the datasets under investigation, to perform
221 source apportionment using PMF with the multi-linear engine (ME-2) (Paatero, 1999)
222 on AMS data (Canonaco et al., 2013). This differs from the standard, unconstrained
223 PMF analysis described by Ulbrich et al. (2009) in that target factors and relaxation
224 parameters can be provided in a manner similar to Lanz et al. (2008). The target
225 factors are added to the analysis following a step-by-step process described in section
226 3.2.1 of Crippa et al. (2013). This means that additional factors can be quantified that
227 might not otherwise have been found through unconstrained PMF analysis, while the
228 use of a common set of rules maintains consistency and objectivity of the analysis. In
229 this work all factor analysis data (except for the analysis in Section 4.6) are taken
230 from the analysis of Crippa et al. (2013). For most of the sites, they retrieved four
231 factors (HOA, BBOA, SV-OOA and LV-OOA). In addition, factor analysis of AMS
232 data measured at Cabauw during May 2008 identified a second LV-OOA component
233 with characteristic spectral profile representing approximately 15% of the OA.
234 Paglione et al. (2014) showed that this OOA factor was associated with polluted
235 continental air masses and exhibited a good correlation with off-line measurements of
236 HULIS (humic-like substances). Therefore, the Cabauw HULIS component is
237 included in the OOA concentration for the purposes of our study.

238 The modeling domain covers a $5400 \times 5832 \text{ km}^2$ region in Europe with 36×36
239 km grid resolution and 14 vertical layers extending up to 6 km in height with a surface
240 layer depth of 55 m. PMCAMx was set to perform simulations on a rotated polar
241 stereographic map projection. The dimensions of the modeling domain are the same in
242 the simulations of all 3 periods. Figure 1 shows a map of the modeling domain of

243 PMCAMx with the location of the stations with available AMS measurements for
244 each period. We have used the same boundary conditions as in Fountoukis et al.
245 (2011).

246 The necessary meteorological inputs to the model were generated from the
247 WRF (Weather Research and Forecasting) model (Skamarock et al., 2008) and
248 include horizontal wind components, vertical diffusivity, temperature, pressure, water
249 vapor, clouds and rainfall. Anthropogenic and biogenic hourly emissions for gases
250 and primary particulate matter were developed for all three periods. Volatile organic
251 compounds are split based on the speciation proposed by Visschedijk et al. (2007).
252 Anthropogenic particulate matter mass emissions of organic and elemental carbon are
253 based on the EUCAARI (Kulmala et al., 2011, 2009) Pan-European Carbonaceous
254 Aerosol Inventory. The biogenic emissions were produced by MEGAN (Model of
255 Emissions of Gases and Aerosols from Nature) (Guenther et al., 2006). A marine
256 aerosol emission model (O'Dowd et al., 2008) was also used for the estimation of
257 mass fluxes for both accumulation and coarse mode including an organic fine mode
258 aerosol fraction. Wildfire emissions are also included (Sofiev et al., 2009). The OA
259 emissions in PMCAMx were distributed by volatility using the volatility distributions
260 of Tsimpidi et al. (2010). The enthalpies of vaporization that are used for primary and
261 secondary OA are the same as in Murphy and Pandis (2009).

262

263 **4. Results and discussion**

264 **4.1 Total fine organic aerosol concentrations in Europe**

265 Figure 2 shows the PMCAMx predicted average ground level concentration of
266 PM₁ OA during each simulation period. Overall, the domain-average contribution of
267 OA concentration to total PM₁ mass is similar (ranges from 31 to 33%) during the

268 three simulation periods. However, the absolute concentration levels and spatial
 269 distribution are quite different. During May the model predicts elevated
 270 concentrations (up to $6 \mu\text{g m}^{-3}$) in a large area covering the UK, northern France,
 271 Belgium, the Netherlands and northwestern Germany while in central and southern
 272 Europe the model predicts lower concentrations ($\sim 2 \mu\text{g m}^{-3}$). During winter the
 273 situation is different with the model predicting high OA values at urban and heavily
 274 industrialized areas (up to $15 \mu\text{g m}^{-3}$), a result of the strong influence of primary
 275 emissions. The largest (on average) OA concentrations are predicted for the autumn
 276 period with a peak monthly average of $5.7 \mu\text{g m}^{-3}$ at the Po Valley in Italy. Contrary
 277 to the late spring period, OA predictions during September/October are relatively low
 278 in central and Northern Europe ($1 - 2 \mu\text{g m}^{-3}$), a result of northwesterly winds
 279 prevailing in the last two weeks of the simulation period.

280 Figure 3 shows a comparison of predicted vs. observed daily average PM_{10} OA
 281 concentrations from all three measurement periods (18 measurement datasets in total).
 282 The prediction skill of PMCAMx is quantified in terms of the mean bias (MB), the
 283 mean absolute gross error (MAGE), the fractional bias (FBIAS), and the fractional
 284 error (FERROR),

$$\begin{aligned}
 285 \quad MB &= \frac{1}{n} \sum_{i=1}^n (P_i - O_i) & MAGE &= \frac{1}{n} \sum_{i=1}^n |P_i - O_i| \\
 286 \quad FBIAS &= \frac{2}{n} \sum_{i=1}^n \frac{(P_i - O_i)}{(P_i + O_i)} & FERROR &= \frac{2}{n} \sum_{i=1}^n \frac{|P_i - O_i|}{(P_i + O_i)}
 \end{aligned}$$

287 where P_i represents the model predicted value for data point i , O_i is the corresponding
 288 observed value and n is the total number of data points. The overall agreement
 289 between AMS measurements and model predictions is encouraging. The majority
 290 (63%) of the data points for PM_{10} OA lay within the 1:2 and 2:1 error lines. The error
 291 is mostly scatter (FERROR=0.57) rather than systematic bias (FBIAS=-0.09). There

292 is a notable variation of model performance among the three periods. The model
293 performs best during May reproducing 84% of the data within a factor of 2 with a
294 fractional error of 0.35 and a fractional bias of 0.01. The winter 2009 simulation
295 period shows the largest discrepancies between model and observations
296 (FERROR=0.68, FBIAS=0.02, 50% of the data predicted within a factor of 2) while
297 during September/October 2008 the model performs better than during winter,
298 reproducing 66% of the data within a factor of 2 (FERROR=0.58, FBIAS=-0.3). The
299 OA formation during the late spring and autumn period is governed by photochemical
300 reactions while during winter primary OA emissions play a major role. The remaining
301 one third of the data that lay out of the 1:2 and 2:1 error lines of Fig. 3 is mostly
302 scattered over many sites and appears as either an over- or an under-prediction during
303 mostly the winter and autumn periods. The overall comparison using hourly resolved
304 data shows similar model performance compared to the daily data, with the model
305 reproducing 57% of the data within a factor of 2 (FERROR=0.68, FBIAS=-0.06).

306 The comparison for fine particulate sulfate concentrations is similar to that for
307 OA in all three periods and sites (55% of the data are predicted within a factor of 2
308 with FERROR=0.6 and FBIAS=0.03). The best performance is seen in May when the
309 model reproduces 70% of the data within a factor of 2 (FERROR=0.4, FBIAS=0.1).
310 For EC the model tends to under predict concentrations (FERROR=0.8, FBIAS=
311 -0.35) during all periods especially at suburban sites influenced by local pollution
312 (e.g. Melpitz) although the number of available data were limited.

313

314 **4.2 Predictions of OA composition in Europe**

315 The average ground level concentration predictions of all the PM₁ OA
316 components during each simulation period are shown in Figure 4. Fresh POA is

317 higher during the winter period ranging from 0.5 to 1.5 $\mu\text{g m}^{-3}$ on average, with peak
318 values up to 10 $\mu\text{g m}^{-3}$ at urban centers. The model predicts that POA contributes 7%
319 to total OA during May, 13% during winter and 4% during autumn with the rest
320 comprising OOA. This is in close agreement with Bergstrom et al. (2012) who
321 predicted (using the EMEP model) less than 10% contribution of fresh POA to total
322 fine OA, over a 6-yr (2002 – 2007) period in Europe. bSOA is predicted to be the
323 dominant oxygenated OA component with a contribution of ~40% to total OOA
324 during May and 60% during both winter and autumn. Oxidized POA is predicted to be
325 the second largest contributor to OOA. OPOA comprises around one third of OOA
326 during May and around 25% during the other two periods. The model suggests that
327 what was traditionally thought as primary (non-volatile and non-reactive) organic
328 matter can actually undergo evaporation, gas phase oxidation and transformation to
329 lower volatility products which can condense back to the particulate phase. This OOA
330 is transported and shows regional rather than local characteristics in its spatial
331 distribution. Anthropogenic SOA levels range from 0.2 to 1.0 $\mu\text{g m}^{-3}$. Interestingly,
332 aSOA shows a maximum over the UK and Benelux area during May and a minimum
333 over the same area during winter. The model predicts the highest concentrations of
334 OPOA and aSOA during the May period and of bSOA during autumn. In each one of
335 the simulation periods, the spatial patterns of concentrations for OPOA, aSOA and
336 bSOA are quite similar to each other, while for fresh POA they are different
337 indicating differences in sources and production mechanisms.

338

339 **4.3 Comparison with factor analysis (ME-2) AMS data**

340 The prediction skill metrics of PMCAMx against AMS HOA and OOA
341 measurements from all stations are summarized in Tables 1 and 2 while Fig. 5 shows

342 an overall comparison of modeled versus measured values for the two components.
343 The AMS HOA component is typically associated with primary fossil fuel
344 combustion organic matter and thus we compare it with the POA in the model. The
345 oxygenated OA AMS component is compared against the sum of aSOA, bSOA and
346 OPOA.

347 Overall, the average AMS HOA is higher than the POA concentrations (by
348 roughly $0.4 \mu\text{g m}^{-3}$). However, the discrepancies vary considerably depending on the
349 site and period (Fig. 5). During winter PMCAMx underpredicts HOA at Barcelona
350 and Chilbolton and overpredicts at Cabauw, Hyytiälä and Helsinki. During the fall
351 there is systematic HOA underprediction at almost all sites. During the May period,
352 the model correctly predicts very low concentrations of POA (less than $0.3 \mu\text{g m}^{-3}$) at
353 Melpitz and Finokalia throughout the month while at Mace Head it underpredicts
354 HOA with a mean error of $-0.25 \mu\text{g m}^{-3}$. At Cabauw the model predicts an average
355 concentration of $0.6 \mu\text{g m}^{-3}$ compared to a $1 \mu\text{g m}^{-3}$ of HOA estimated during the May
356 period. Overall, the model performance is better during the late spring period with a
357 mean error of $0.26 \mu\text{g m}^{-3}$ while the mean bias and error is always less than $1 \mu\text{g m}^{-3}$
358 during all three periods (with the exception of Barcelona).

359 The agreement between predictions and observations is better for OOA with
360 the model reproducing 83% of the data within a factor of 2 during May, 55% during
361 winter and 68% during the autumn period while the average fractional error and bias
362 are 0.53 and -0.07, respectively ($\text{MAGE} = 0.9 \mu\text{g m}^{-3}$, $\text{MB} = -0.2 \mu\text{g m}^{-3}$).

363 The error for HOA concentrations is most likely an indication of errors in the
364 emissions rates of OA and/or errors in their assumed volatility distribution. Although
365 most previous studies have considered OPOA as OOA (Robinson et al., 2007;
366 Shrivastava et al., 2008; Murphy and Pandis, 2009; Hodzic et al., 2010, Fountoukis et

367 al., 2011, etc.) it has been argued (Cappa and Wilson, 2012; Aumont et al., 2012) that
368 not all OPOA is oxidized enough to be assigned to the OOA mass fraction by factor
369 analysis. Zhang et al. (2013) suggested that if some OPOA was measured as part of
370 HOA, the CTM bias would be reduced. In our case, if 50% of simulated OPOA is
371 considered as fPOA, the model bias for HOA is reduced (MAGE = $0.6 \mu\text{g m}^{-3}$, MB =
372 $-0.08 \mu\text{g m}^{-3}$ compared to MAGE = $0.7 \mu\text{g m}^{-3}$ and MB = $-0.4 \mu\text{g m}^{-3}$), but the average
373 model performance for OOA deteriorates (fractional error and bias are 0.6 and -0.24
374 compared to 0.54 and -0.08 in the basecase). Errors in the emissions inventory are
375 likely to be the source of bias for HOA as well as the uncertain distribution of OA
376 emissions in the low volatility bins which can strongly influence the initial
377 partitioning between the gas and the aerosol phase and thus the predicted POA
378 concentrations (Tsimpidi et al., 2011).

379 Figure 6 shows average diurnal profiles of PM_{10} HOA and OOA components at
380 Cabauw during May 2008. The observed morning HOA peak is reproduced by the
381 model but there is a tendency towards underprediction throughout the day. The model
382 compares well with the measured OOA when HULIS is included in the OOA (Fig.
383 6b). On average the model predicts $3.6 \mu\text{g m}^{-3}$ OOA compared to $3.4 \mu\text{g m}^{-3}$ measured
384 by the AMS. If HULIS was assigned to HOA the fractional error increases from 0.55
385 to 0.97 for HOA and from 0.25 to 0.46 for OOA.

386 Biomass burning OA was detected by the factor analysis at two sites during
387 May (Cabauw and Mace Head), five sites during autumn (Hyytiälä, K-Pusztta,
388 Vavihill, Harwell and Melpitz) and seven during winter (Cabauw, Melpitz, Hyytiälä,
389 Helsinki, Mace Head, Chilbolton, and Barcelona). For these datasets, the predicted
390 POA has been compared so far against the sum of the AMS BBOA and HOA.
391 However the model systematically underpredicts fresh POA at most of these sites,

392 especially during autumn with a mean bias up to $-3.4 \mu\text{g m}^{-3}$ (fractional bias up to -1.3
393 $\mu\text{g m}^{-3}$). Figure 7 shows average diurnal profiles of PM_{10} HOA and OOA components
394 at K-Pusztá during the autumn period. If measured BBOA is excluded from the model
395 vs. measurement comparison, the agreement for fPOA is much better and the
396 fractional bias is reduced substantially from -0.95 to -0.1 . The same applies in other
397 sites such as at Melpitz where the FBIAS drops from -1 to -0.47 or at Cabauw where
398 FBIAS is reduced from -0.51 to -0.02 . This indicates that the quantity of fresh primary
399 OA emissions from biomass burning sources may be underestimated in the inventory.
400 This is consistent with recent studies (e.g., Bergstrom et al., 2012; Denier van der Gon
401 et al., 2013; Kostenidou et al., 2013) who pointed towards large uncertainties in the
402 biomass burning emission estimates in many European areas. To further explore this,
403 the PSAT algorithm was used in parallel with the main CTM to calculate the
404 apportionment for each source and for each species. Source types tracked by PSAT
405 include all the major anthropogenic emission sectors (e.g., energy sector, non-
406 industrial combustion, industry, fossil fuel production, solvent use, road transport,
407 waste disposal, non-road transport, agriculture, and shipping) as well as emission
408 from wildfires, ecosystems and sea-salt emission sources. The fraction of fresh
409 primary OA concentration that was attributed to the three most important biomass
410 burning related sources (wildfires, residential combustion and agricultural waste
411 processing) during the spring period is shown in Fig. 8. A large fraction of POA
412 concentrations in Europe during the late spring period is attributed to wildfires. In
413 certain areas such as Russia and south Europe wildfire emissions are predicted to
414 contribute up to 95% to the primary OA concentrations. The second largest
415 contributor over land is the residential (wood and coal) combustion ranging on
416 average from 30 to 60%. Fresh POA concentrations originating from this source are

417 more regionally distributed with peak contributions (up to 60%) in east Turkey,
418 Norway, France and Portugal. Lower fractional contributions to POA are predicted
419 from waste processing and agricultural sector emissions ranging from 10 to 30%
420 mostly in central and Eastern Europe. By directly comparing the AMS BBOA
421 concentrations with the fresh POA from the sum of the biomass burning sources
422 predicted by PSAT during the spring period, a clear model underprediction is seen in
423 both sites where BBOA data are available. At Mace Head the factor analysis
424 estimated average BBOA concentration is three times higher than the model predicted
425 fresh POA from the sum of the three major sources of Fig. 8. At Cabauw the model
426 underpredicts the observed BBOA by approximately 40%.

427 The factor analysis in 13 out of the 18 datasets resulted in two oxygenated OA
428 components, LV-OOA and SV-OOA. The model does not simulate these two
429 components explicitly. Therefore, an attempt to identify certain parts of the modeled
430 OOA (based on their volatility distribution) that correlate with these two observed
431 components was made. The LV-OOA is considered highly oxygenated and processed
432 OA and thus it can be initially attributed to the long-range transported OA in the
433 model. The SV-OOA is considered less processed and of higher volatility and thus it
434 is expected to correlate well with the modeled OOA covering volatility bins that
435 correspond to higher saturation concentrations. However, as these two components
436 are thought to represent the two end states of the OOA oxidation in a given dataset,
437 the respective LV-OOA and SV-OOA components may not be chemically identical
438 between sites. Figure 9 shows a comparison of estimated LV-OOA and SV-OOA by
439 factor analysis against predicted concentrations at Finokalia during the late spring
440 period. Predictions shown include OA transported from outside the domain plus the
441 first 3 volatility bins (with C^* of 10^{-2} , 10^{-1} and $1 \mu\text{g m}^{-3}$) for the LV-OOA comparison,

442 and OOA with C^* from 1 to $10^3 \mu\text{g m}^{-3}$ for the SV-OOA comparison. A number of
443 different combinations were tested for all datasets. Statistically the model performs
444 the best in all sites for both LV-OOA and SV-OOA when LV-OOA is compared
445 against the OA with $C^* \leq 10^{-1} \mu\text{g m}^{-3}$ and SV-OOA against the OA with $C^* > 10^{-1} \mu\text{g}$
446 m^{-3} respectively. However, this model performance should not be overinterpreted, as
447 these results are sensitive to the combination of boundary conditions used and the
448 assumed volatility distribution. Furthermore, the 2-D VBS scheme (Donahue et al.,
449 2011) tracking both the volatility and oxidation state could be more helpful for such a
450 comparison, as in some sites the two OOA components analyzed by factor analysis
451 may differ in the extent of oxidation but show similarities in volatility (Hildebrandt et
452 al., 2010).

453

454 **4.4 Sensitivity to the volatility distribution**

455 The volatility distribution of primary OA emissions affects the gas-particle
456 partitioning of POA and may be another source of uncertainty in the predicted POA
457 concentrations. Based on the PSAT results for the late spring period, the largest part
458 of POA concentrations and emissions in continental Europe originates from biomass
459 burning (fires and residential wood combustion). Here we tried a different volatility
460 distribution than that of Shrivastava et al. (2008) used in the base case, which was
461 recently proposed by May et al. (2013c) for biomass burning POA emissions. The
462 simulation with the new volatility distribution results in higher on average POA
463 concentrations ($0.37 \mu\text{g m}^{-3}$ compared to $0.28 \mu\text{g m}^{-3}$ in the basecase run) due to a
464 larger fraction of emissions assigned to the lower volatility bins. As a result, the
465 underprediction for POA is lower (mean bias is reduced from $-0.23 \mu\text{g m}^{-3}$ to $-0.15 \mu\text{g}$
466 m^{-3}) and the mean error is slightly decreased (by $0.03 \mu\text{g m}^{-3}$). However, the new

467 volatility distribution produces significantly lower OOA values, with PMCAMx
468 predicting an average concentration of $1.6 \mu\text{g m}^{-3}$ at the 4 sites during May, compared
469 to a concentration of $2.0 \mu\text{g m}^{-3}$ with the base case run. This resulted in an increase of
470 the average fractional error from 0.37 to 0.56 and a systematic underprediction of
471 OOA (FBIAS decreased from 0.09 to -0.43). However, this sensitivity simulation is
472 assigning the fresh BBOA volatility distribution to all other sources. This should be
473 viewed as a sensitivity test, not as a test of the actual volatility distributions. Future
474 work should include utilizing multiple volatility distributions depending on the
475 various sources of primary OA emissions. The estimation of the absolute emission
476 rate from the measured organic emission factors is an additional source of uncertainty.
477 The corresponding scaling depends on the organic concentration levels and
478 temperature of the measurements and these are often not well documented or are not
479 used appropriately in the development of the emission inventories.

480

481 **4.5 Sensitivity to SOA aging rate**

482 A sensitivity test was conducted to explore the role of the assumed aging in
483 biogenic SOA. In a similar sensitivity test, Fountoukis et al. (2011) assumed that
484 biogenic SOA ages with the same aging rate constant as the anthropogenic SOA ($k= 1$
485 $\times 10^{-11} \text{ cm}^3 \text{ molec}^{-1} \text{ s}^{-1}$) and found a significant increase of OA concentrations
486 resulting in model overprediction of OA at all studied sites. Murphy and Pandis
487 (2010) tried different combinations of aging for aSOA and bSOA and concluded that
488 aging biogenic and anthropogenic SOA together with a reduced aging reaction rate
489 ($2.5 \times 10^{-12} \text{ cm}^3 \text{ molec}^{-1} \text{ s}^{-1}$) resulted in reasonable model performance and a slightly
490 increased summertime OA formation in the eastern US. Following Murphy and
491 Pandis (2010) we tried the same reduced aging rate constant ($2.5 \times 10^{-12} \text{ cm}^3 \text{ molec}^{-1}$

492 s^{-1}) for both aSOA and bSOA. The model predicts a small increase of OA
493 concentrations during May (between 4 and 14%) and the autumn period (0.2 – 3%)
494 and a decrease during winter (0.2 – 2%) mainly due to a stronger effect of bSOA
495 aging during periods with enhanced photochemistry. The sensitivity run predicts an
496 increase of the contribution of bSOA concentration to total OA from 48 % to 57 % in
497 May, from 68 % to 73 % in autumn and from 64 % to 66 % in winter, on a domain-
498 average basis. Overall the model performance for total PM_{10} OA did not improve (the
499 average fractional error increased from 0.57 to 0.58). Moreover there was a notable
500 increase of the bias for OOA concentrations during the late spring period (the
501 fractional bias increased from 0.01 to 0.1).

502

503 **4.6 Sensitivity to factor analysis methodology**

504 Crippa et al. (2013) performed source apportionment using factor analysis
505 with the multilinear engine (ME-2) on AMS data from all three campaigns. They
506 suggested a standardized methodology that could be applied to any future
507 measurement site or campaign instead of relying on performing unconstrained PMF
508 analysis on each dataset separately by each research group, which is subject to
509 differences in each researcher's approach. Here we explore the sensitivity of our
510 model's performance to the use of the ME-2 method. In 6 datasets HOA
511 concentrations were not available in the initial PMF analysis performed by the
512 individual research groups. The ME-2 data, however, included HOA concentrations in
513 these sites improving the model's performance. Overall the average fractional bias for
514 HOA concentrations was reduced from -0.4 to -0.12 with the use of the ME-2 data.
515 OOA concentrations were only slightly changed using the ME-2 method (by $-0.1 \mu g$
516 m^{-3} or -5% on average). In all cases the use of ME-2 results contributed to improved

517 model performance in all studied sites reducing the average fractional error for OOA
518 from 0.56 to 0.53. Furthermore, the number of datasets where the factor analysis
519 identified two OOA components (LV-OOA and SV-OOA) instead of just one,
520 increased (from 9 to 13) with the use of the ME-2 method. This facilitated the analysis
521 of section 4.3. The systematic application of the ME-2 method should probably be
522 preferred in future modeling studies that use extensive factorization for model
523 evaluation.

524

525 **5. Conclusions**

526 In this work, a three-dimensional regional chemical transport model
527 (PMCAMx) was applied over Europe focusing on the formation and chemical
528 transformation of organic matter. The Volatility Basis Set, which considers both
529 primary and secondary organics as semi-volatile and photochemically reactive, was
530 used in both the CTM and the source apportionment algorithm PSAT. Three periods
531 were simulated corresponding to intensive measurement campaigns at various sites in
532 Europe, and the model predictions are compared against factor analysis AMS data.
533 The overall agreement between predictions and measurements for total fine OA mass
534 is encouraging with the model reproducing the majority (more than 63%) of the data
535 points within a factor of two. Interestingly, the model performance fluctuates
536 substantially among the three periods, showing the lowest error (FERROR=0.35)
537 during spring and the highest during the winter period (FERROR=0.68).

538 On average, the predicted oxygenated OA contributes 93% to total OA during
539 May, 87% during winter and 96% during autumn with the rest comprising fresh
540 primary OA. The model predicts the highest concentrations of biogenic secondary OA
541 during autumn and of oxidized primary OA and anthropogenic secondary OA during

542 May. Biogenic secondary OA is predicted to be the dominant oxygenated OA
543 component with a contribution of ~40-60% to total oxygenated OA in all three
544 periods. Predicted oxygenated OA concentrations compare well with the AMS
545 measured values for all periods.

546 The error for hydrocarbon-like OA concentrations (MAGE = $0.7 \mu\text{g m}^{-3}$, mean
547 bias = $-0.4 \mu\text{g m}^{-3}$) is most likely an indication of errors in the emissions rates of
548 primary OA and/or errors in their assumed volatility distribution. Including a portion
549 of simulated oxidized primary OA in the fresh primary OA rather than oxygenated
550 OA does not generally improve model performance. The model systematically
551 underpredicts fresh primary OA in most sites during late spring and autumn with a
552 mean bias up to $-0.8 \mu\text{g m}^{-3}$ (fractional bias up to -0.95). Based on the PSAT results
553 the biomass burning OA is most likely underestimated in the emission inventory. The
554 largest part of primary OA concentrations and emissions in continental Europe
555 originates from biomass burning (fires and residential wood combustion). A different
556 volatility distribution (representative of biomass burning primary OA emissions)
557 applied to primary OA emissions from all sectors results in higher on average primary
558 OA concentrations with an increased error for oxygenated OA concentrations (FBIAS
559 increased from 0.09 to -0.43).

560 The model performs well in all sites when the measured low-volatility
561 oxygenated OA is compared against the OA with saturation concentration $C^* \leq 10^{-1} \mu\text{g}$
562 m^{-3} and the semi-volatile oxygenated OA against the OA with $C^* > 10^{-1} \mu\text{g m}^{-3}$
563 respectively. Assuming that both biogenic and anthropogenic secondary OA age
564 together with a reduced aging reaction rate ($2.5 \times 10^{-12} \text{ cm}^3 \text{ molec}^{-1} \text{ s}^{-1}$) does not
565 improve model performance for total OA while it increases the bias for oxygenated
566 OA concentrations.

567 The comparisons are sensitive to the factor analysis methodology and the
568 observations were found to compare better with the model when the ME-2 based
569 approach proposed by Crippa et al. (2013) is applied to all datasets.

570 Other possible sources of uncertainty that are not explored here but have been
571 investigated in past applications of PMCAMx include uncertainties in the aging
572 scheme, the magnitude of IVOC emissions, and others. For example Murphy et al.
573 (2011) explored a 2-bin reduction in volatility upon one oxidation step with a
574 simultaneous decrease (by a factor of 2) in the aging rate constants. They found a
575 slight underprediction of the OA mass observed at Finokalia during May 2008
576 compared to the base case 1-bin shift. Hodzic et al. (2010) and Grieshop et al. (2009)
577 investigated a 2-bin reduction (in addition to the 1-bin base case saturation
578 concentration reduction) with a reduced OH reaction rate constant and found both to
579 perform adequately. Murphy et al. (2012) added a detailed functionalization scheme
580 to approximate the effect on volatility of adding relevant functional groups to the
581 carbon backbone (Donahue et al., 2011). This approach alone resulted in a significant
582 increase of the OA mass. Adding fragmentation to the detailed functionalization
583 scenario decreased OA mass concentrations to the approximate magnitude predicted
584 by the base case (which employs a simplified scheme that is currently used in
585 PMCAMx) and brought the model into reasonable agreement with the OA mass
586 concentration measurements. In our base case aging scheme we use this simplified
587 scenario that tries to describe the net effect of the chemical aging reactions (both
588 functionalization and fragmentation) without treating any of the two types explicitly.
589 The sensitivity of the model's predictions to the uncertain IVOC emissions was
590 stressed by Tsimpidi et al. (2010). An additional SOA formation pathway that is not
591 simulated here is the in-cloud SOA formation from glyoxal and methylglyoxal.

592 Murphy et al. (2012) explored the contribution of this pathway to OOA concentrations
593 in several European sites that are also studied in our work (Mace Head, Cabauw,
594 Finokalia) during both May 2008 and February/March 2009. They reported small
595 enhancements to both average OA mass loadings ($< 3\%$) and O:C ($< 10\%$) at the
596 surface. Their estimated OA production from this pathway was found lower (in
597 absolute magnitude) than that seen by Carlton et al. (2008) or measured during the
598 CalNEX campaign but the contribution to total SOA formed was similar (0-4%).
599 Aqueous-phase SOA formation from glyoxal and methylglyoxal was also investigated
600 by Myriokefalitakis et al. (2011) using the global 3-D chemistry/transport model
601 TM4-ECPL. They reported significant contributions of oxalate to SOA mass mainly
602 over oceans, possibly due to long range transport of oxidation products of terrestrial
603 biogenic VOC and subsequent cloud processing, as well as to the multiphase
604 processing of the marine VOC emissions.

605 Although the base case OA scheme seems to represent the average
606 atmospheric chemistry of OA reasonably well there is clearly the possibility of
607 compensating errors, since several required parameters for describing all relevant
608 processes in a framework like the VBS are highly uncertain.

609

610 *Acknowledgements*

611 We thank M. Dall'Osto and J. Ovadnevaite for provision of the Mace Head data, and
612 L. Hildebrandt for the Finokalia measurements. This work was funded by the
613 European Community's 7th Framework Programme EU projects PEGASOS (contract
614 265307) and the ERC project ATMOPACS (contract 267099).

615

616

617

618 **References**

- 619 Aumont, B., Valorso, R., Mouchel-Vallon, C., Camredon, M., Lee-Taylor, J., and
620 Madronich, S.: Modeling SOA formation from the oxidation of intermediate
621 volatility n-alkanes, *Atmos. Chem. Phys.*, 12, 7577–7589, 2012.
- 622 Bergstrom, R., Denier van der Gon, H. A. C., Prevot, A. S. H., Yttri, K. E., and
623 Simpson, D.: Modelling of organic aerosols over Europe (2002–2007) using a
624 volatility basis set (VBS) framework: application of different assumptions
625 regarding the formation of secondary organic aerosol, *Atmos. Chem. Phys.*, 12,
626 8499–8527, 2012.
- 627 Canagaratna, M. R., Jayne, J. T., Jimenez, J. L., Allan, J. D., Alfarra, M. R., Zhang,
628 Q., Onasch, T. B., Drewnick, F., Coe, H., Middlebrook, A., Delia, A., Williams,
629 L. R., Trimborn, A. M., Northway, M. J., DeCarlo, P. F., Kolb, C. E., Davidovits,
630 P., and Worsnop, D. R.: Chemical and microphysical characterization of ambient
631 aerosols with the aerodyne aerosol mass spectrometer, *Mass Spectrometry
632 Reviews*, 26, 185-222, 2007.
- 633 Canonaco, F., Crippa, M., Slowik, J. G., Baltensperger, U., and Prévôt, A. S. H.: SoFi,
634 an Igor based interface for the efficient use of the generalized multilinear engine
635 (ME-2) for source apportionment: application to aerosol mass spectrometer data,
636 *Atmos. Meas. Tech. Discuss.*, 6, 6409-6443, 2013.
- 637 Cappa, C. D. and Wilson, K. R.: Multi-generation gas-phase oxidation, equilibrium
638 partitioning, and the formation and evolution of secondary organic aerosol,
639 *Atmos. Chem. Phys.*, 12, 5, 9505–9528, 2012.
- 640 Carlton, A. G., Turpin, B. J., Altieri, K. E., Seitzinger, S. P., Mathur, R., Roselle, S.
641 J., and Weber, R. J.: CMAQ model performance enhanced when in-cloud
642 secondary organic aerosol is included: comparisons of organic carbon predictions
643 with measurements, *Environ. Sci. Technol.*, 42, 8798–8802,
644 doi:10.1021/Es801192n, 2008.
- 645 Carter, W. P. L.: Programs and Files Implementing the SAPRC-99 Mechanism and its
646 Associates Emissions Processing Procedures for Models-3 and Other Regional
647 Models: <http://www.engr.ucr.edu/~carter/SAPRC99/> (last access: 30 March
648 2010), 2010.

649 Crippa, M., F. Canonaco, V.A. Lanz, M. Äijälä, J.D. Allan, S. Carbone, G. Capes, M.
650 Dall'Osto, D.A. Day, P.F. DeCarlo, C.F. Di Marco, M. Ehn, A. Eriksson, E.
651 Freney, L. Hildebrandt Ruiz, R. Hillamo, J.-L. Jimenez, H. Junninen, A.
652 Kiendler-Scharr, A.-M. Kortelainen, M. Kulmala, A.A. Mensah, C. Mohr, E.
653 Nemitz, C. O'Dowd, J. Ovadnevaite, S.N. Pandis, T. Petäjä, L. Poulain, S.
654 Saarikoski, K. Sellegri, E. Swietlicki, P. Tiitta, D.R. Worsnop, U. Baltensperger,
655 and A.S.H. Prévôt: Organic aerosol components derived from 25 AMS datasets
656 across Europe using a newly developed ME-2 based source apportionment
657 strategy, *Atmos. Chem. Phys. Discuss.*, 13, 23325-23371, 2013.

658 Denier van der Gon, H., Visschedijk, A., Fountoukis, C., Pandis, S. N., Bergstrom, R.,
659 Simpson, D., and Johansson, C.: Particulate emissions from residential wood
660 combustion in Europe – revised estimates and an evaluation, in preparation, 2014.

661 Donahue, N. M., Robinson, A. L., Stanier, C. O., and Pandis, S. N.: Coupled
662 partitioning, dilution, and chemical aging of semivolatile organics, *Environ. Sci.*
663 *Technol.*, 40, 2635–2643, 2006.

664 Donahue, N. M., Epstein, S. A., Pandis, S. N., and Robinson, A. L.: A two-
665 dimensional volatility basis set: 1. organic-aerosol mixing thermodynamics,
666 *Atmos. Chem. Phys.*, 11, 3303–3318, doi:10.5194/acp-11-3303-2011, 2011.

667 ENVIRON: User's Guide to the Comprehensive Air Quality Model with Extensions
668 (CAMx), Version 4.02, ENVIRON Int. Corp., Novato, Calif., 2003.

669 Fountoukis, C., Racherla, P.N., Denier van der Gon, H.A.C., Polymeneas, P.,
670 Charalampidis, P.E., Pilinis, C., Wiedensohler, A., Dall'Osto, M., O'Dowd, C.,
671 and Pandis, S.N.: Evaluation of a three-dimensional chemical transport model
672 (PMCAMx) in the European domain during the EUCAARI May 2008 campaign,
673 *Atmos. Chem. Phys.*, 11, 10331-10347, 2011.

674 Fountoukis, C., Koraj, Dh., Denier van der Gon, H.A.C., Charalampidis, P.E., Pilinis,
675 C., and Pandis, S. N.: Impact of grid resolution on the predicted fine PM by a
676 regional 3-D chemical transport model, *Atmos. Environ.*, 68, 24-32, 2013.

677 Gelencser, A., May, B., Simpson, D., Sanchez-Ochoa, A., Kasper-Giebl, A.,
678 Puxbaum, H., Caseiro, A., Pio, C., and Legrand, M.: Source apportionment of
679 PM_{2.5} organic aerosol over Europe: Primary/secondary, natural/anthropogenic,

680 and fossil/biogenic origin, *J. Geophys. Res.*, 112, D23S04,
681 doi:10.1029/2006JD008094, 2007.

682 Grieshop, A. P., Logue, J. M., Donahue, N. M., and Robinson, A. L.: Laboratory
683 investigation of photochemical oxidation of organic aerosol from wood fires 1:
684 measurement and simulation of organic aerosol evolution, *Atmos. Chem. Phys.*,
685 9, 1263–1277, 2009.

686 Guenther, A., Karl, T., Harley, P., Wiedinmyer, C., Palmer, P. I., and Geron, C.:
687 Estimates of global terrestrial isoprene emissions using MEGAN (Model of
688 Emissions of Gases and Aerosols from Nature), *Atmos. Chem. Phys.*, 6, 3181–
689 3210, doi:10.5194/acp-6-3181-2006, 2006.

690 Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M.,
691 Dommen, J., Donahue, N. M., George, C., Goldstein, A. H., Hamilton, J. F.,
692 Herrmann, H., Hoffmann, T., Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L.,
693 Kiendler-Scharr, A., Maenhaut, W., McFiggans, G., Mentel, Th. F., Monod, A.,
694 Prevot, A. S. H., Seinfeld, J. H., Surratt, J. D., Szmigielski, R., and Wildt, J.: The
695 formation, properties and impact of secondary organic aerosol: current and
696 emerging issues, *Atmos. Chem. Phys.*, 9, 5155–5236, 2009.

697 Hildebrandt, L., Engelhart, G. J., Mohr, C., Kostenidou, E., Lanz, V. A., Bougiatioti,
698 A., DeCarlo, P. F., Prevot, A. S. H., Baltensperger, U., Mihalopoulos, N.,
699 Donahue, N. M., and Pandis, S. N.: Aged organic aerosol in the Eastern
700 Mediterranean: the Finokalia Aerosol Measurement Experiment – 2008, *Atmos.*
701 *Chem. Phys.*, 10, 4167–4186, doi:10.5194/acp-10-4167-2010, 2010.

702 Hildebrandt, L., Donahue, N. M., and Pandis, S. N.: High formation of secondary
703 organic aerosol from the photo-oxidation of toluene, *Atmos. Chem. Phys.*, 9,
704 2973–2986, 2009.

705 Hodzic, A., Jimenez, J. L., Madronich, S., Canagaratna, M. R., DeCarlo, P. F.,
706 Kleinman, L., and Fast, J.: Modeling organic aerosols in a megacity: potential
707 contribution of semi-volatile and intermediate volatility primary organic
708 compounds to secondary organic aerosol formation, *Atmos. Chem. Phys.*, 10,
709 5491–5514, doi:10.5194/acp-10-5491-2010, 2010.

710 Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q.,
711 Kroll, J. H., DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C.,
712 Docherty, K. D., Ulbrich, I. M., Grieshop, A. P., Robinson, A. L., Duplissy, J.,
713 Smith, J. D., Wilson, K. R., Lanz, V. A., Hueglin, C., Sun, Y. L., Tian, J.,
714 Laaksonen, A., T., R., Rautiainen, J., Vaattovaara, P., Ehn, M., Kulmala, M.,
715 Tomlinson, J. M., Collins, D. R., Cubison, M. J., Dunlea, E. J., Huffman, J. A.,
716 Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo, Y., Schneider,
717 J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D., Cottrell,
718 L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimojo, A., Sun, J. Y.,
719 Zhang, Y. M., Dzepina, K., Kimmel, J. R., Sueper, D., Jayne, J. T., Herndon, S.
720 C., Trimborn, A. M., Williams, L. R., Wood, E. C., Kolb, C. E., Baltensperger,
721 U., and Worsnop, D. R.: Evolution of organic aerosol in the atmosphere, *Science*,
722 326, 1525–1529, 2009.

723 Kanakidou, M., Seinfeld, J. H., Pandis, S. N., Barnes, I., Dentener, F. J., Facchini, M.
724 C., Van Dingenen, R., Ervens, B., Nenes, A., Nielsen, C. J., Swietlicki, E.,
725 Putaud, J. P., Balkanski, Y., Fuzzi, S., Horth, J., Moortgat, G. K., Winterhalter,
726 R., Myhre, C. E. L., Tsigaridis, K., Vignati, E., Stephanou, E. G., and Wilson, J.:
727 Organic aerosol and global climate modeling: a review, *Atmos. Chem. Phys.*, 5,
728 1053–1123, 2005.

729 Karydis, V. A., Tsimpidi, A. P., Fountoukis, C., Nenes, A., Zavala, M., Lei, W.,
730 Molina, L. T., and Pandis, S. N.: Simulating the fine and coarse inorganic
731 particulate matter concentrations in a polluted megacity, *Atmos. Environ.*, 44,
732 608–620, 2010.

733 Kostenidou, E., Kaltsonoudis, C., Tsiglikiotou, M., Louvaris, E., Russell, L. M., and
734 S. N. Pandis, Burning of olive tree branches: a major organic aerosol source in
735 the Mediterranean, *Atmos. Chem. Phys. Discuss.*, 13, 7223-7266, 2013.

736 Kroll, J. H., Donahue, N. M., Jimenez, J. L., Kessler, S. H., Canagaratna, M. R.,
737 Manjula, R. Wilson, K. R., Altieri, K. E., Mazzoleni, L. R., Wozniak, A. S.,
738 Bluhm, H., Mysak, E. R., Smith, J. D., Kolb, C. E., and Worsnop, D. R.: Carbon
739 oxidation state as a metric for describing the chemistry of atmospheric organic
740 aerosol, *Nature Chemistry*, 3, 133-139, 2011.

741 Kulmala, M., Asmi, A., Lappalainen, H. K., Carslaw, K. S., Poschl, U., Baltensperger,
742 U., Hov, Ø., Brenguier, J.-L., Pandis, S. N., Facchini, M. C., Hansson, H.-C.,
743 Wiedensohler, A., and O'Dowd, C. D.: Introduction: European Integrated Project
744 on Aerosol Cloud Climate and Air Quality interactions (EUCAARI) – integrating
745 aerosol research from nano to global scales, *Atmos. Chem. Phys.*, 9, 2825–2841,
746 doi:10.5194/acp-9-2825-2009, 2009.

747 Kulmala, M., Asmi, A., Lappalainen, H. K., Baltensperger, U., Brenguier, J.-L.,
748 Facchini, M. C., Hansson, H.-C., Hov, Ø., O'Dowd, C. D., Pöschl, U.,
749 Wiedensohler, A., Boers, R., Boucher, O., de Leeuw, G., Denier van der Gon, H.
750 A. C., Feichter, J., Krejci, R., Laj, P., Lihavainen, H., Lohmann, U., McFiggans,
751 G., Mentel, T., Pilinis, C., Riipinen, I., Schulz, M., Stohl, A., Swietlicki, E.,
752 Vignati, E., Alves, C., Amann, M., Ammann, M., Arabas, S., Artaxo, P., Baars,
753 H., Beddows, D. C. S., Bergstrom, R., Beukes, J. P., Bilde, M., Burkhardt, J. F.,
754 Canonaco, F., Clegg, S. L., Coe, H., Crumeyrolle, S., D'Anna, B., Decesari, S.,
755 Gilardoni, S., Fischer, M., Fjaeraa, A. M., Fountoukis, C., George, C., Gomes, L.,
756 Halloran, P., Hamburger, T., Harrison, R. M., Herrmann, H., Hoffmann, T.,
757 Hoose, C., Hu, M., Hyvarinen, A., Horrak, U., Iinuma, Y., Iversen, T., Josipovic,
758 M., Kanakidou, M., Kiendler-Scharr, A., Kirkevåg, A., Kiss, G., Klimont, Z.,
759 Kolmonen, P., Komppula, M., Kristjansson, J.-E., Laakso, L., Laaksonen, A.,
760 Labonnote, L., Lanz, V. A., Lehtinen, K. E. J., Rizzo, L. V., Makkonen, R.,
761 Manninen, H. E., McMeeking, G., Merikanto, J., Minikin, A., Mirme, S.,
762 Morgan, W. T., Nemitz, E., O'Donnell, D., Panwar, T. S., Pawlowska, H.,
763 Petzold, A., Pienaar, J. J., Pio, C., Plass-Duelmer, C., Prevot, A. S. H., Pryor, S.,
764 Reddington, C. L., Roberts, G., Rosenfeld, D., Schwarz, J., Seland, Ø., Sellegri,
765 K., Shen, X. J., Shiraiwa, M., Siebert, H., Sierau, B., Simpson, D., Sun, J. Y.,
766 Topping, D., Tunved, P., Vaattovaara, P., Vakkari, V., Veefkind, J. P.,
767 Visschedijk, A., Vuollekoski, H., Vuolo, R., Wehner, B., Wildt, J., Woodward,
768 S., Worsnop, D. R., van Zadelhoff, G.-J., Zardini, A. A., Zhang, K., van Zyl, P.
769 G., Kerminen, V.-M., S Carslaw, K., and Pandis, S. N.: General overview:
770 European Integrated project on Aerosol Cloud Climate and Air Quality
771 interactions (EUCAARI) – integrating aerosol research from nano to global
772 scales, *Atmos. Chem. Phys.*, 11, 13061–13143, doi:10.5194/acp-11-13061-2011,
773 2011.

- 774 Lane, T. E., Donahue, N. M., and Pandis, S. N.: Simulating secondary organic aerosol
775 formation using the volatility basis-set approach in a chemical transport model,
776 *Atmos. Environ.*, 42, 7439–7451, 2008a.
- 777 Lane, T. E., Donahue, N. M., and Pandis, S. N.: Effect of NO_x on secondary organic
778 aerosol concentrations, *Environ. Sci. Technol.*, 42, 6022–6027, 2008b.
- 779 Lanz, V. A., Prevot, A. S. H., Alfarra, M. R., Weimer, S., Mohr, C., DeCarlo, P. F.,
780 Gianini, M. F. D., Hueglin, C., Schneider, J., Favez, O., D’Anna, B., George, C.,
781 and Baltensperger, U.: Characterization of aerosol chemical composition by
782 aerosol mass spectrometry in Central Europe: an overview, *Atmos. Chem. Phys.*,
783 10, 10453–10471, 2010.
- 784 Lanz, V. A., Alfarra, M. R., Baltensperger, U., Buchmann, B., Hueglin, C., Szidat, S.,
785 Wehrli, M. N., Wacker, L., Weimer, S., Caseiro, A., Puxbaum, H., and Prevot, A.
786 S. H.: Source attribution of submicron organic aerosols during wintertime
787 inversions by advanced factor analysis of aerosol mass spectra, *Environ. Sci.*
788 *Tech.*, 42, 214-220, 2008.
- 789 Lanz, V. A., Alfarra, M. R., Baltensperger, U., Buchmann, B., Hueglin, C., and
790 Prevot, A. S. H.: Source apportionment of submicron organic aerosols at an urban
791 site by factor analytical modeling of aerosol mass spectra, *Atmos. Chem. Phys.*,
792 7, 1503–1522, 2007.
- 793 May, A. A., Presto, A. A., Hennigan, C. J., Nguyen, N. T., Gordon, T. D., and
794 Robinson, A. L.: Gas-particle partitioning of primary organic aerosol emissions:
795 (1) Gasoline vehicle exhaust, *Atmos. Environ.* 77, 128-139, 2013a.
- 796 May, A. A., Presto, A. A., Hennigan, C. J., Nguyen, N. T., Gordon, T. D., and
797 Robinson, A. L.: Gas-particle partitioning of primary organic aerosol emissions:
798 (2) Diesel vehicles, *Environ. Sci. Technol.*, 47, 8288 – 8296, doi:
799 10.1021/es400782j, 2013b.
- 800 May, A. A., Levin, E. J. T., Hennigan, C. J., Riipinen, I., Lee, T., Collett, J. L.,
801 Jimenez, J. L., Kreidenweis, S. M., and Robinson, A. L.: Gas-particle partitioning
802 of primary organic aerosol emissions: 3. Biomass burning, *J. Geophys. Res.*, 118,
803 11327 – 11338, 2013c.

804 Morgan, W. T., Allan, J. D., Bower, K. N., Highwood, E. J., Liu, D., McMeeking, G.
805 R., Northway, M. J., Williams, P. I., Krejci, R., and Coe, H.: Airborne
806 measurements of the spatial distribution of aerosol chemical composition across
807 Europe and evolution of the organic fraction, *Atmos. Chem. Phys.*, 10, 4065–
808 4083, doi:10.5194/acp-10-4065-2010, 2010.

809 Murphy, B. N. and Pandis, S. N.: Simulating the formation of semivolatile primary
810 and secondary organic aerosol in a regional chemical transport model, *Environ.*
811 *Sci. Technol.*, 43, 4722–4728, 2009.

812 Murphy, B. N. and Pandis, S. N.: Exploring summertime organic aerosol formation in
813 the eastern United States using a regional-scale budget approach and ambient
814 measurements, *J. Geophys. Res.*, 115, D24216, doi:10.1029/2010JD014418,
815 2010.

816 Murphy, B. N., Donahue, N. M., Fountoukis, C., and Pandis, S. N.: Simulating the
817 oxygen content of ambient organic aerosol with the 2D volatility basis set,
818 *Atmos. Chem. Phys.*, 11, 7859–7873, 2011.

819 Murphy, B. N., Donahue, N. M., Fountoukis, C., Dall’Osto, M., O’Dowd, C.,
820 Kiendler-Scharr, A., and Pandis, S. N.: Functionalization and fragmentation
821 during ambient organic aerosol aging: application of the 2-D volatility basis set to
822 field studies, *Atmos. Chem. Phys.*, 12, 10797–10816, 2012.

823 Myriokefalitakis, S., Tsigaridis, K., Mihalopoulos, N., Sciare, J., Nenes, A.,
824 Kawamura, K., Segers, A., and Kanakidou, M.: In-cloud oxalate formation in the
825 global troposphere: a 3-D modeling study, *Atmos. Chem. Phys.*, 11, 5761–5782,
826 2011.

827 Ng, N. L., Canagaratna, M. R., Zhang, Q., Jimenez, J. L., Tian, J., Ulbrich, I. M.,
828 Kroll, J. H., Docherty, K. S., Chhabra, P. S., Bahreini, R., Murphy, S. M.,
829 Seinfeld, J. H., Hildebrandt, L., Donahue, N. M., DeCarlo, P. F., Lanz, V. A.,
830 Prevot, A. S. H., Dinar, E., Rudich, Y., and Worsnop, D. R.: Organic aerosol
831 components observed in Northern Hemispheric datasets from Aerosol Mass
832 Spectrometry, *Atmos. Chem. Phys.*, 10, 4625–4641, doi:10.5194/acp-10-4625-
833 2010, 2010.

834 Ng, N. L., Kroll, J. H., Keywood, M. D., Bahreini, R., Varutbangkul, V., Flagan, R.
835 C., and Seinfeld, J. H.: Contribution of first- versus second-generation products to
836 secondary organic aerosols formed in the oxidation of biogenic hydrocarbons,
837 *Environ. Sci. Technol.*, 40, 2283–2297, 2006.

838 O’Dowd, C. D., Langmann, B., Varghese, S., Scannell, C., Ceburnis, D., and
839 Facchini, M. C.: A combined organic-inorganic sea-spray source function,
840 *Geophys. Res. Lett.*, 35, L01801, doi:10.1029/2007GL030331, 2008.

841 Paatero, P., and Tapper, U.: Positive matrix factorization - a nonnegative factor model
842 with optimal utilization of error-estimates of data values, *Environmetrics*, 5, 111-
843 126, 1994.

844 Paatero, P.: Least squares formulation of robust non-negative factor analysis,
845 *Chemometr. Intell. Lab.*, 37, 23–35, 1997.

846 Paatero, P.: The multilinear engine - A table-driven, least squares program for solving
847 multilinear problems, including the n-way parallel factor analysis model, *J.*
848 *Comput. Graph. Stat.*, 8, 854-888, 1999.

849 Paglione, M., Kiendler-Scharr, A., Mensah, A. A., Finessi, E., Giulianelli, L.,
850 Sandrini, S., Facchini, M. C., Fuzzi, S., Schlag, P., Piazzalunga, A., Tagliavini,
851 E., Henzing, J. S., and Decesari, S.: Identification of humic-like substances
852 (HULIS) in oxygenated organic aerosols using NMR and AMS factor analyses
853 and liquid chromatographic techniques, *Atmos. Chem. Phys.*, 14, 25–45,
854 doi:10.5194/acp-14-25-2014, 2014.

855 Presto, A. A., Miracolo, M. A., Kroll, J. H., Worsnop, D. R., Robinson, A. L., and
856 Donahue, N. M.: Intermediate-volatility organic compounds: A potential source
857 of ambient oxidized organic aerosol, *Environ. Sci. Technol.*, 43, 4744–4749,
858 2009.

859 Robinson, A. L., Donahue, N. M., Shrivastava, M. K., Weitkamp, E. A., Sage, A. M.,
860 Grieshop, A. P., Lane, T. E., Pierce, J. R., Pandis, S. N.: Rethinking organic
861 aerosol: semivolatile emissions and photochemical aging, *Science*, 315, 1259–
862 1262, 2007.

863 Seinfeld, J. H. and Pandis, S. N.: Atmospheric chemistry and physics: From air
864 pollution to climate change. 2nd ed.; John Wiley and Sons, Hoboken, NJ, 2006.

865 Shrivastava, M. K., Lane, T. E., Donahue, N. M., Pandis, S. N., and Robinson, A. L.:
866 Effects of gas-particle partitioning and aging of primary emissions on urban and
867 regional organic aerosol concentrations, *J. Geophys. Res.*, 113, D18301,
868 doi:10.1029/2007JD009735, 2008.

869 Skamarock, W. C., Klemp, J. B., Dudhia, J., Gill, D. O., Barker, D. M., Duda, M. G.,
870 Huang, X. Wang, W., and Powers, J. G.: A Description of the Advanced
871 Research WRF Version 3, NCAR Technical Note, available at:
872 http://www.mmm.ucar.edu/wrf/users/docs/arw_v3.pdf, last access: 30 June 2008,
873 2008.

874 Skyllakou, K., Murphy, B. N., Megaritis, A. G., Fountoukis, C., and Pandis, S. N.:
875 Contributions of local and regional sources to fine PM in the Megacity of Paris,
876 *Atmos. Chem. Phys. Discuss.*, 13, 25769-25799, 2013.

877 Sofiev, M., Vankevich, R., Lotjonen, M., Prank, M., Petukhov, V., Ermakova, T.,
878 Koskinen, J., Kukkonen, J.: An operational system for the assimilation of the
879 satellite information on wild-land fires for the needs of air quality modeling and
880 forecasting, *Atmos. Chem. Phys.*, 9, 6833 – 6847, 2009.

881 Szidat, S., Jenk, T. M., Synal, H. A., Kalberer, M., Wacker, L., Hajdas, I., Kasper-
882 Giebl, A., and Baltensperger, U.: Contributions of fossil fuel, biomass-burning,
883 and biogenic emissions to carbonaceous aerosols in Zurich as traced by (14)C, *J.*
884 *Geophys. Res.*, 111, D07206, doi:10.1029/2005JD006590, 2006.

885 Tørseth, K., Aas, W., Breivik, K., Fjæraa, A. M., Fiebig, M., Hjellbrekke, A. G., Lund
886 Myhre, C., Solberg, S., and Yttri, K. E.: Introduction to the European Monitoring
887 and Evaluation Programme (EMEP) and observed atmospheric composition
888 change during 1972–2009, *Atmos. Chem. Phys.*, 12, 5447-5481, 2012.

889 Tsimpidi, A. P., Karydis, V. A., Zavala, M., Lei, W., Bei, N., Molina, L., and Pandis,
890 S. N.: Sources and production of organic aerosol in Mexico City: insights from
891 the combination of a chemical transport model (PMCAMx-2008) and
892 measurements during MILAGRO, *Atmos. Chem. Phys.*, 11, 5153–5168, 2011.

893 Tsimpidi, A. P., Karydis, V. A., Zavala, M., Lei, W., Molina, L., Ulbrich, I. M.,
894 Jimenez, J. L., and Pandis, S. N.: Evaluation of the volatility basis-set approach

895 for the simulation of organic aerosol formation in the Mexico City metropolitan
896 area, *Atmos. Chem. Phys.*, 10, 525–546, 2010.

897 Tunved, P., Hansson, H. C., Kerminen, V. M., Strom, J., Dal Maso, M., Lihavainen,
898 H., Viisanen, Y., Aalto, P. P., Komppula, M., and Kulmala, M.: High natural
899 aerosol loading over boreal forests, *Science*, 312, 5771, 261–263,
900 doi:10.1126/science.1123052, 2006.

901 Turpin, B. J., Saxena, P., and Andrews, E.: Measuring and simulating particulate
902 organics in the atmosphere: problems and prospects, *Atmos. Environ.*, 34, 2983–
903 3013, 2000.

904 Ulbrich, I. M., Canagaratna, M. R., Zhang, Q., Worsnop, D. R., and Jimenez, J. L.:
905 Interpretation of organic components from Positive Matrix Factorization of
906 aerosol mass spectrometric data, *Atmos. Chem. Phys.*, 9, 2891–2918, 2009.

907 Visschedijk, A. J. H., Zandveld, P., and Denier van der Gon, H. A. C.: A high
908 resolution gridded European emission database for the EU integrated project
909 GEMS, TNO Report 2007 A-R0233/B: Organization for Applied Scientific
910 Research, Netherlands, available at: [http://lap.phys.auth.gr/gems/docu/
911 TNO%20Short%20Emissions%20Report.pdf](http://lap.phys.auth.gr/gems/docu/TNO%20Short%20Emissions%20Report.pdf), last access: 31 March 2007, 2007.

912 Wagstrom, K.M., Pandis, S.N., Yarwood, G., Wilson, G.M., and Morris, R.E.:
913 Development and application of a computationally efficient particulate matter
914 apportionment algorithm in a three-dimensional chemical transport model,
915 *Atmos. Environ.*, 42, 5650–5659, 2008.

916 Zhang, Q., Worsnop, D. R., Canagaratna, M. R., and Jimenez, J. L.: Hydrocarbon-like
917 and oxygenated organic aerosols in Pittsburgh: insights into sources and
918 processes of organic aerosols, *Atmos. Chem. Phys.*, 5, 3289–3311,
919 doi:10.5194/acp-5-3289-2005, 2005.

920 Zhang, Q., Jimenez, J. L., Canagaratna, M. R., Allan, J. D., Coe, H., Ulbrich, I.,
921 Alfarra, M. R., Takami, A., Middlebrook, A. M., Sun, Y. L., Dzepina, K.,
922 Dunlea, E., Docherty, K., De-Carlo, P., Salcedo, D., Onasch, T. B., Jayne, J. T.,
923 Miyoshi, T., Shimon, A., Hatakeyama, N., Takegawa, N., Kondo, Y., Schneider,
924 J., Drewnick, F., Weimer, S., Demerjian, K. L., Williams, P. I., Bower, K. N.,
925 Bahreini, R., Cottrell, L., Griffin, R. J., Rautianen, J., and Worsnop, D. R.:

926 Ubiquity and dominance of oxygenated species in organic aerosols in
927 anthropogenically-influenced Northern Hemisphere midlatitudes, *Geophys. Res.*
928 *Lett.*, 34, L13801, doi:10.1029/2007GL029979, 2007.

929 Zhang, Q. J., Beekmann, M., Drewnick, F., Freutel, F., Schneider, J., Crippa, M.,
930 Prevot, A. S. H., Baltensperger, U., Poulain, L., Wiedensohler, A., Sciare, J.,
931 Gros, V., Borbon, A., Colomb, A., Michoud, V., Doussin, J.-F., Denier van der
932 Gon, H. A. C., Haeffelin, M., Dupont, J.-C., Siour, G., Petetin, H., Bessagnet, B.,
933 Pandis, S. N., Hodzic, A., Sanchez, O., Honore, C., and Perrussel, O.: Formation
934 of organic aerosol in the Paris region during the MEGAPOLI summer campaign:
935 evaluation of the Volatility-Basis-Set approach within the CHIMERE model,
936 *Atmos. Chem. Phys.*, 13, 5767-5790, 2013.

937

938 **Table 1.** Prediction skill metrics of PMCAMx against factor analysis AMS data for
 939 OOA.

940

OOA	Mean predicted ($\mu\text{g m}^{-3}$)	Mean observed ($\mu\text{g m}^{-3}$)	FERROR	FBIAS	MAGE ($\mu\text{g m}^{-3}$)	MB ($\mu\text{g m}^{-3}$)
<i>May 2008</i>						
Melpitz	3.4	4.6	0.33	-0.32	1.3	-1.3
Cabauw	3.6	3.4	0.25	0.15	0.7	0.3
Finokalia	2.7	2.3	0.28	0.18	0.65	0.43
Mace						
Head	1.9	0.8	0.68	0.68	1.0	1.0
<i>February/March 2009</i>						
Melpitz	0.8	1.2	0.73	-0.5	0.65	-0.44
Cabauw	0.9	0.9	0.58	0.07	0.45	-0.2
Finokalia	2.3	1.5	0.52	0.36	1.09	0.8
Hyytiälä	1.6	1.3	0.51	0.24	0.7	0.3
Barcelona	2.4	4.5	0.68	-0.53	2.3	-2.1
Helsinki	2.3	2.1	0.3	0.1	0.5	0.2
Mace						
Head	0.8	0.9	1.0	0.58	0.9	-0.1
Chilbolton	0.7	1.6	0.75	-0.63	0.95	-0.9
<i>September/October 2008</i>						
Melpitz	2.3	2.6	0.46	-0.18	0.98	-0.3
Hyytiälä	0.9	0.7	0.58	0.2	0.45	0.2
K-Pusztá	3.7	3.8	0.43	-0.18	1.2	-0.14
Puijo	0.5	0.8	0.7	-0.49	0.44	-0.34
Vavihill	0.6	1.8	1.1	-1.1	1.2	-1.2
Harwell	2.4	2.9	0.4	-0.2	1.0	-0.6

941

942

943

944 **Table 2.** Prediction skill metrics of PMCAMx against factor analysis AMS data for
 945 HOA.

HOA	Mean predicted ($\mu\text{g m}^{-3}$)	Mean observed ($\mu\text{g m}^{-3}$)	FERROR	FBIAS	MAGE ($\mu\text{g m}^{-3}$)	MB ($\mu\text{g m}^{-3}$)
<i>May 2008</i>						
Melpitz	0.2	0.25	0.38	-0.19	0.09	-0.04
Cabauw	0.6	0.99	0.55	-0.51	0.45	-0.41
Finokalia	0.04	0.09	0.83	-0.83	0.05	-0.05
Mace Head	0.1	0.35	1.1	-1.1	0.25	-0.25
<i>February/March 2009</i>						
Melpitz	0.3	0.3	0.48	-0.12	0.12	-0.04
Cabauw	1.1	0.4	0.96	0.91	0.8	0.8
Finokalia	-	-	-	-	-	-
Hyytiälä	0.8	0.13	1.4	1.4	0.68	0.68
Barcelona	0.7	4.1	1.3	-1.3	3.4	-3.4
Helsinki	1.6	0.85	0.68	0.62	0.84	0.76
Mace Head	0.1	0.2	0.86	-0.74	0.13	-0.13
Chilbolton	0.4	1.0	1.1	-1.1	0.7	-0.7
<i>September/October 2008</i>						
Melpitz	0.2	0.6	1	-1	0.45	-0.45
Hyytiälä	0.2	0.1	0.8	0.56	0.09	0.07
K-Pusztá	0.4	1.2	0.95	-0.95	0.8	-0.8
Puijo	0.1	0.2	0.85	-0.72	0.14	-0.1
Vavihill	0.2	0.8	1.3	-1.2	0.6	-0.6
Harwell	0.3	0.7	0.8	-0.8	0.4	-0.4

946

947

948

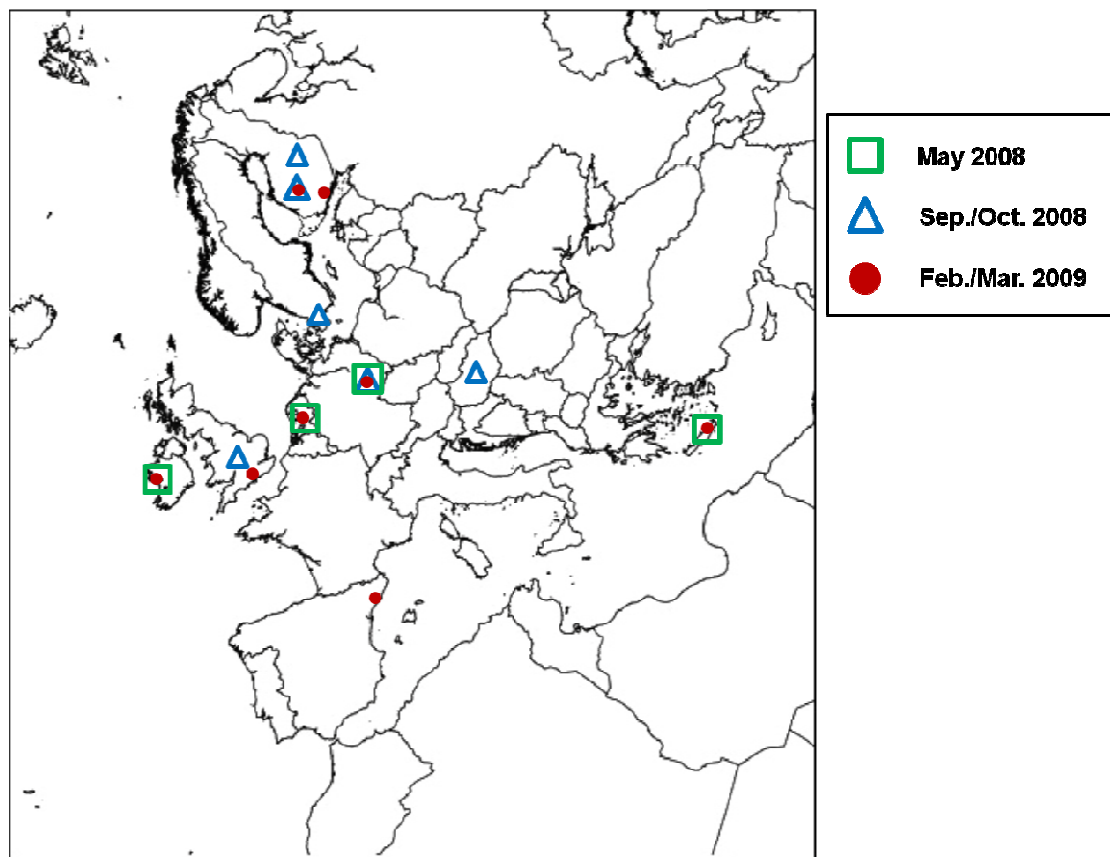
949

950

951

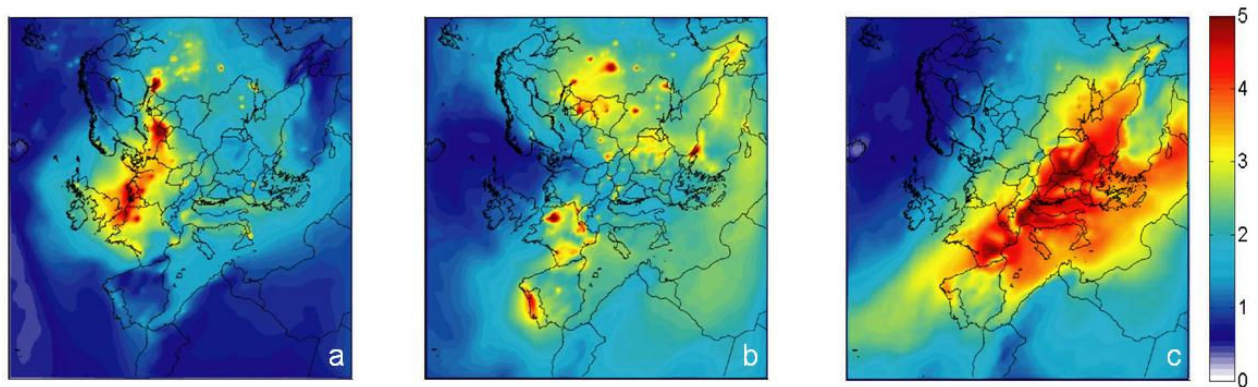
952

953



955

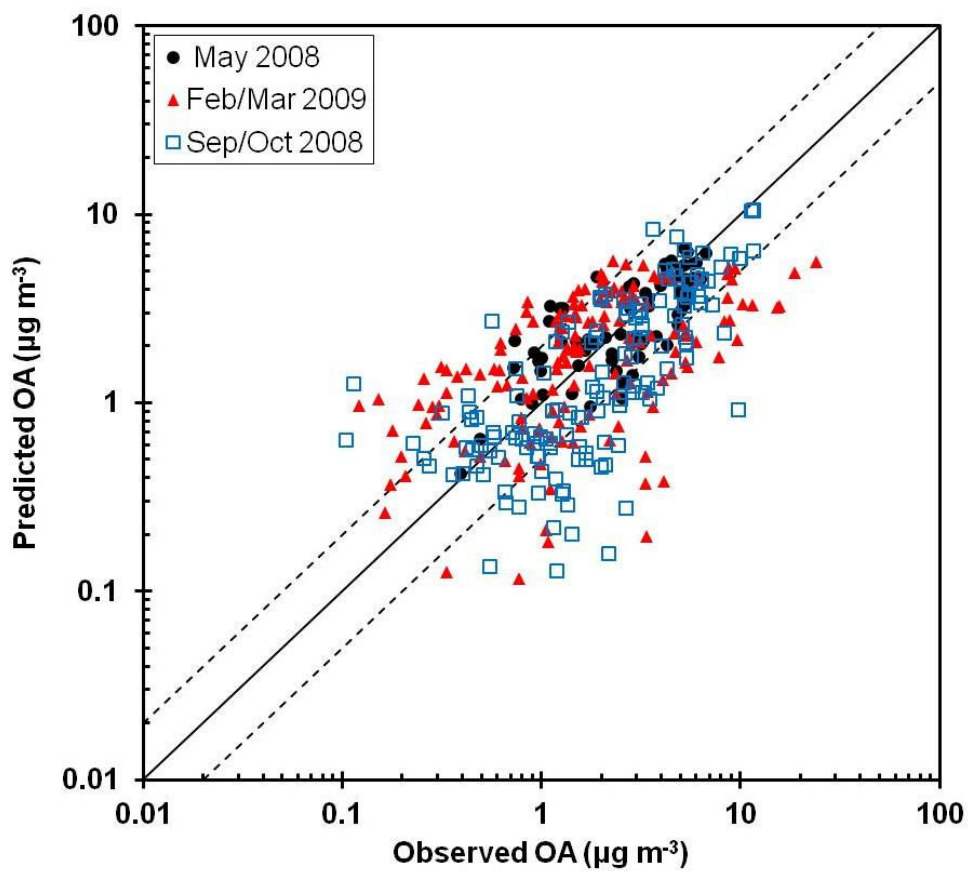
956 **Figure 1.** Modeling domain of PMCAMx for Europe. Symbols show the location of
957 measurement sites for each period. Green squares represent the May 2008 period, blue
958 triangles the September/October 2008 period and red circles the February/March 2009
959 period.



962 **Figure 2.** Ground-level concentration predictions of PM₁ OA (µg m⁻³) averaged over
963 the entire simulation period for (a) 1 – 29 May 2008, (b) 25 February – 24 March
964 2009, and (c) 15 September – 17 October 2008.
965

966

967

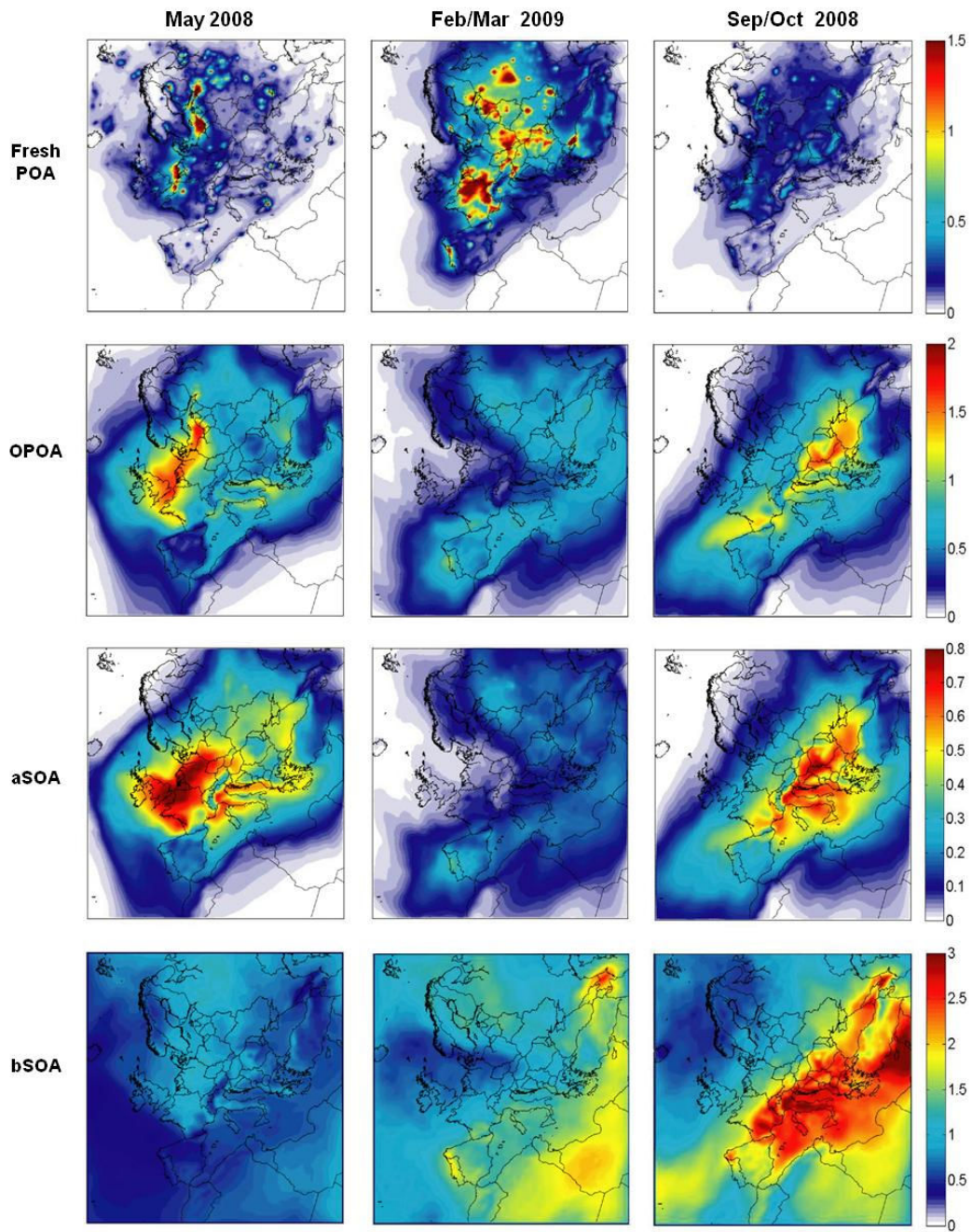


968

969

970 **Figure 3.** Comparison of predicted vs. observed (AMS) PM₁ OA ($\mu\text{g m}^{-3}$) for the 3
971 measurement periods (18 measurement sites in total). Each point is a daily average
972 value. Also shown the 1:1, 2:1 and 1:2 lines.

973

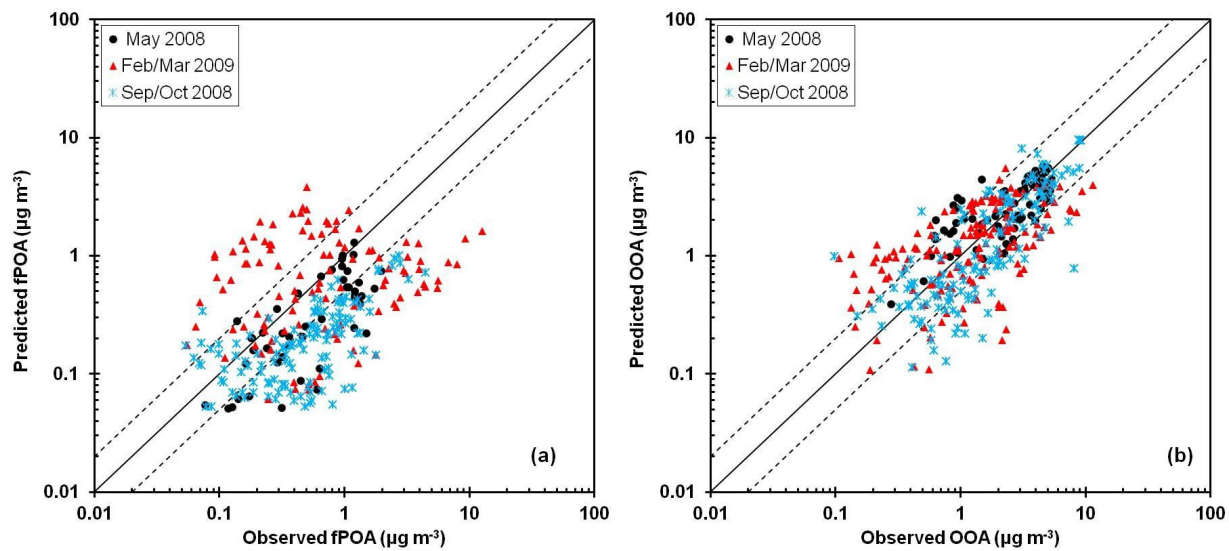


975
976
977
978

Figure 4. Average ground-level concentration predictions for fresh POA, OPOA, aSOA and bSOA for each period. Different scales are used.

979

980



981

982

983

984

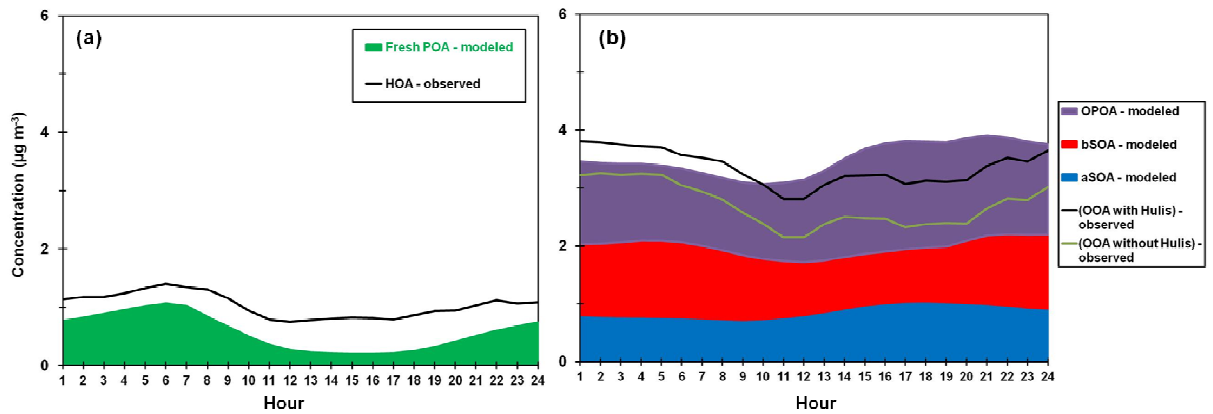
985

986

Figure 5. Comparison of predicted vs. observed PM₁ a) HOA and b) OOA ($\mu\text{g m}^{-3}$) from 3 measurement periods (18 measurement sites in total). Each point corresponds to a daily average value. Also shown the 1:1, 2:1 and 1:2 lines. Observed data represent AMS/factor-analysis measurements.

987

988



989

990

991

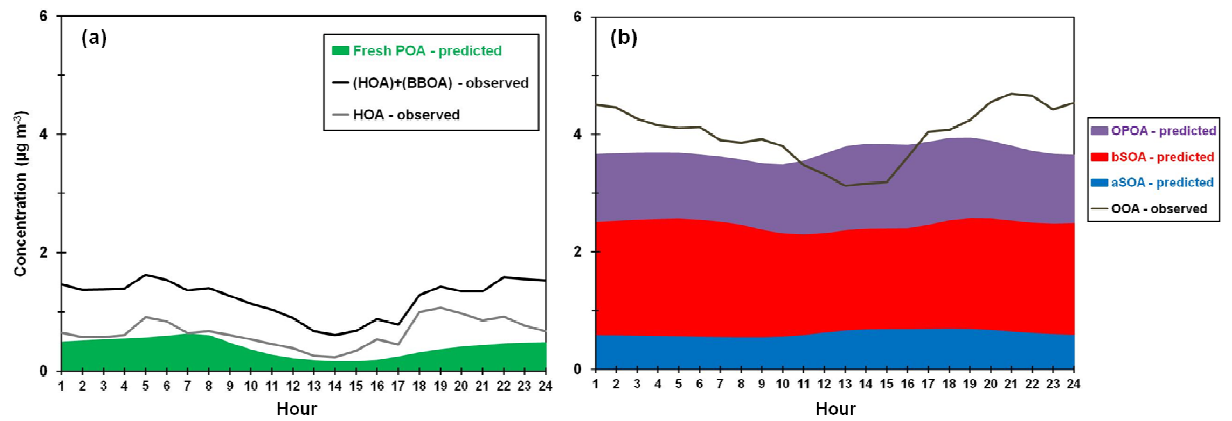
992

993

Figure 6. Average diurnal profiles of PM₁ a) HOA and b) OOA components at Cabauw during the EUCAARI May 2008 campaign.

994

995



996

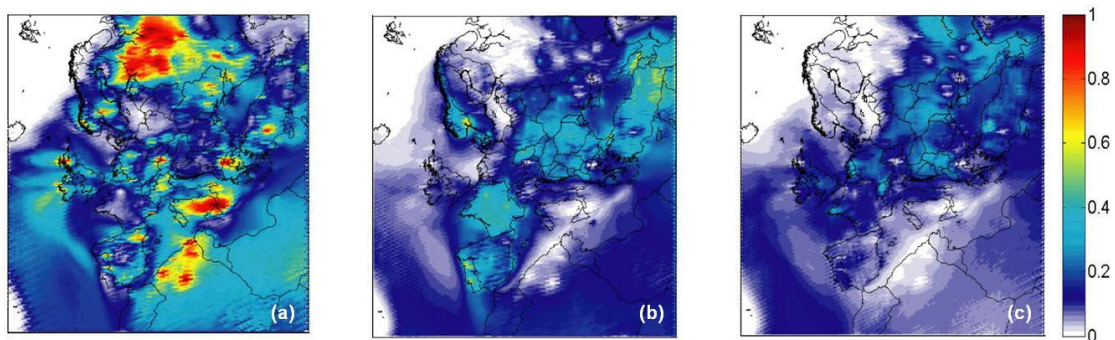
997

998

999

Figure 7. Average diurnal profiles of PM₁ a) HOA and b) OOA components at K-Pusztá during the September/October 2008 period.

1000
1001
1002

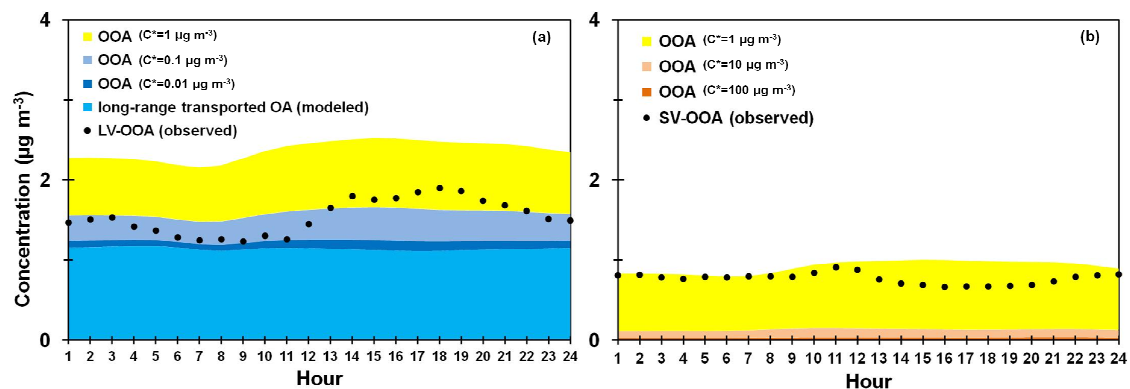


1003
1004
1005
1006
1007

Figure 8. Average fraction of fine fresh primary OA concentrations during May 2008 attributed to a) wildfire emissions, b) residential combustion, and c) waste processing and agricultural activity emissions.

1008

1009



1010

1011

1012

1013

Figure 9. Average diurnal profiles of PM₁ a) LV-OOA and b) SV-OOA components at Finokalia during the May 2008 period.