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Long-term trends in aerosol and precipitation composition over the western North Atlantic Ocean at Bermuda

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Abstract

Since the 1980s, emissions of SO_2 and NO_x (NO + NO₂) from anthropogenic sources in the United States (US) and Europe have decreased significantly suggesting that the export of oxidized S and N compounds from surrounding continents to the atmosphere

- overlying North Atlantic Ocean (NAO) has also decreased. The chemical compositions of aerosols and precipitation sampled daily on Bermuda (32.27° N, 64.87° W) from 1989 to 1997 and from 2006 to 2009 were evaluated to quantify the magnitudes, significance, and implications of associated tends in atmospheric composition. The chemical data were stratified based on FLEXPART retroplumes into four discrete transport regimes:
- ¹⁰ westerly flow from the eastern North America (NEUS/SEUS); easterly trade-wind flow from northern Africa and the subtropical NAO (Africa); long, open-ocean, anticyclonic flow around the Bermuda High (Oceanic); and transitional flow from the relatively clean open ocean to the polluted northeastern US (North). Based on all data, annual average concentrations of non-sea-salt (nss) SO₄²⁻ associated with aerosols and annual VWA
- ¹⁵ concentrations in precipitation decreased significantly (by 22 and 49 %, respectively) whereas annual VWA concentrations of NH_4^+ in precipitation increased significantly (by 70%). Corresponding trends in aerosol and precipitation NO_3^- and of aerosol NH_4^+ were insignificant. Nss SO_4^{2-} in precipitation under NEUS/SEUS and Oceanic flow decreased significantly (61% each) whereas corresponding trends in particulate nss SO_4^{2-} under both flow regimes were insignificant.
- ²⁰ SO_4^{2-} under both flow regimes were insignificant. Trends for precipitation were driven in part by decreasing emissions of SO_2 over upwind continents and associated decreases in anthropogenic contributions to nss SO_4^{2-} concentrations. Under NEUS/SEUS and Oceanic flow, the ratio of anthropogenic to biogenic contributions to to nss SO_4^{2-} in the column scavenged by precipitation were relatively greater than those in near surface ²⁵ aerosol, which implies that, for these flow regimes, precipitation is a better indicator of
- overall anthropogenic impacts on the lower troposphere. Particulate nss SO_4^{2-} under African flow also decreased significantly (34 %) whereas the corresponding decrease in nss SO_4^{2-} associated with precipitation was marginally insignificant. We infer that these





trends were driven in part by reductions in the emissions and transport of oxidized S compounds from Europe. The lack of significant trends in NO₃ associated with aerosols and precipitation under NEUS/SEUS flow is notable in light of the large decrease (39%) in NO, emissions in the US over the period of record. Rapid chemical processing of oxidized N in marine air contributed to this lack of correspondence. Decreasing ratios of nss SO_4^{2-} to NH_4^+ and the significant decreasing trend in precipitation acidity (37 %) indicate that the total amount of acidity in the multiphase gas-aerosol system in the western NAO troposphere decreased over the period of record. Decreasing aerosol acidities would have shifted the phase partitioning of total NH_3 (NH_3 + particulate NH_4^+) towards the gas phase thereby decreasing the atmospheric lifetime of total NH₃ against wet plus dry deposition. The trend of increasing NH_4^+ in precipitation at Bermuda over the period of record suggests that NH₃ emissions from surrounding continents also increased. Decreasing particulate nss SO_4^{2-} in near-surface air under NEUS/SEUS flow over the period of record suggests a lower limit for net warming in the range of $0.1-0.3 \,\mathrm{Wm^{-2}}$ resulting from the decreased shortwave scattering and absorption by 15

 5 0.1–0.3 With resulting from the decreased since SO_{4}^{2-} and associated aerosol constituents.

1 Introduction

It has been recognized for many years that material emitted to the atmosphere from continental sources can be transported long distances over adjacent oceans. For
 example, reports of crustal aerosol transported from northern Africa over the open North Atlantic Ocean (NAO) have been published in the scientific literature since the late 1700s (e.g., Dobson, 1781; Darwin, 1846). More recently, trace elements and reaction products from S and N species emitted by anthropogenic combustion and industrial sources were detected in atmospheric aerosols and precipitation over the NAO (e.g., Zoller et al., 1973; Jickells et al., 1982; Chen and Duce, 1983). Subsequent, studies have attempted to quantitatively differentiate relative contributions from major source types (continental vs. marine; anthropogenic vs. biogenic vs. crustal) and major



source regions based on analysis of elemental and molecular tracers (e.g., Arimoto et al., 1995; Savioe et al., 2002), radionuclides (e.g., Arimoto et al., 1999), isotopic composition (e.g, Turekian et al., 2001, 2003; Hastings et al., 2003), air-mass transport history (e.g., Moody and Galloway, 1988; Galloway et al., 1989; Moody et al., 1995,

- 2014), and satellite remote sensing (e.g., Zhao et al., 2008; Zhang and Reid, 2010; Hsu et al., 2012; Prospero et al., 2012). These and other studies indicate that the transport of anthropogenic emissions and reaction products from North America is the dominant source of pollutants associated with aerosols and precipitation over the western NAO at Bermuda. However, anthropogenic emissions from surrounding continents have
 changed significantly over the past guarter century suggesting that relative impacts
- of anthropogenic sources to air quality over the western NAO have also changed. Implementation of the Clean Air Act and associated amendments in the United

States (US) and similar legislation in Europe in the late 1980s and early 1990s led to substantial reductions in emissions of SO_2 and NO_x to the atmosphere surrounding the North Atlantic basin (e.g., Vestreng, 2007; Hand et al., 2012). For example, emissions

- ¹⁵ North Atlantic basin (e.g., Vestreng, 2007; Hand et al., 2012). For example, emissions inventories prepared by the US Environmental Protection Agency (EPA, 2013) indicate that, between 1980 and 2012, total emissions of SO₂ and NO_x in the US decreased by factors of 80 and 50 %, respectively. In contrast, reported emissions of NH₃ in the US between 1990 (the earliest year available) and 2012 differed by less than 1 % (EPA,
- ²⁰ 2013). Thus, the relative mixture of acids (from SO₂ and NO_x oxidation) and bases in the planetary boundary layer (PBL) over the US and, presumably, associated pHdependent chemical processes, also evolved temporally over this period. Reductions in SO₂ and NO_x emissions led to corresponding decreases in annual average SO₂ and NO_x mixing ratios in ambient air over the US during the same period (factors of
- ²⁵ 78 and 56 %, respectively) (EPA, 2013). Concentrations of particulate-phase reaction products from the oxidation of SO₂ and NO_x as well as aerosol optical depth (AOD) also decreased and visibility improved. For example, between 2000 (the earliest year available) and 2012, average annual concentrations of PM_{2.5} mass in near-surface air over the US decreased by 33 % (EPA, 2013), between 2000 and 2010, particulate SO₄²⁻





concentrations in rural areas of the US decreased by 27 % (Hand et al., 2012), and, between 1980 and 2006, simulated AOD over the US decreased by 38 % (Streets et al., 2009). Reductions in emissions have also led to large decreases in the atmospheric deposition of SO_4^{2-} , NO_3^- and H⁺ and the associated recovery of some degraded watersheds in the eastern US (Kahl et al., 2004; Webb et al., 2004; Leibensperger et al., 2012).

The above trends for the US suggest that the physicochemical properties of the western NAO atmosphere have also changed significantly over the past quarter century analogous to the sharp decreases in lead over the NAO that followed phase out of leaded gasoline in the US (Shen and Boyle, 1987). Although not significant, Zhang and Reid (2010) reported a decreasing trend in AOD over the western NAO between 2000 and 2009. During the same period SO₂ and NO_x emissions in the US decreased by 49 and 31 %, respectively, and simulated AOD over the continental US also decreased (Streets et al., 2009). Based on satellite observations of AOD and the chemical composition of near-surface aerosols, Moody et al. (2014) estimate that under transport from the portheastern US during the period 2006–2009 pollutant.

- that, under transport from the northeastern US during the period 2006–2009, pollutant aerosol accounted for direct radiative cooling at Bermuda of 1.0–2.2 Wm⁻². However, potential changes in AOD resulting from emissions reductions in the US were not evaluated.
- In this paper, we investigate long-term trends in the chemical composition of aerosols and precipitation over the western NAO at Bermuda with a focus on evaluating impacts of emissions reductions and influences of different source regions. This work is based on measurements between 1989 and 1997 under the auspices of the Atmosphere Ocean Chemistry Experiment (AEROCE) and comparable measurements between 25 2006 and 2009 at the same location.



2 Methods

2.1 Sampling site, periods, and protocols

From June 1988 through July 1998, aerosols in sectored (on-shore) flow at the Tudor Hill Atmospheric Observatory (THAO) on Bermuda (32.27° N, 64.87° W) were sampled
daily for chemical characterization from the top of a 23 m scaffolding tower situated on a steep slope about 25 m a.s.l. Samples were collected in bulk at a nominal flow rate of 1.0 m³ min⁻¹ on Whatman 41 filters (20.3 cm × 25.4 cm) as part of AEROCE (Galloway et al., 1993; Savoie et al., 2002). Data were blank corrected based on analysis of paired bulk-filter cassettes deployed periodically in parallel with samples but through which no air was pulled.

 From July 2006 through June 2009, dichotomous aerosol (nominal super- and subµm-diameter size fractions) was sampled daily at a rate of 0.1 m³ min⁻¹ with a customdesigned and fabricated MSP model 130 high-flow cascade impactor configured with Liu-Pui type omnidirectional inlet (Liu et al., 1983). Relative to MSP's Micro-Orifice
 ¹⁵ Uniform Deposit Impactor (MOUDI) (Marple et al., 1991) that is in more widespread use by the research community, these hi-flow impactors yield greater signal per unit deployment time while segregating aerosol size fractions using similar nozzle technology. Impactors were deployed at the top of the THAO tower (Moody et al., 2014). The calculated inlet passing efficiency for 20-µm-diameter particles was 95 % and the

²⁰ 50 % aerodynamic cut between the two size factions was 0.8 µm ambient diameter. Impactors were configured with quartz-fiber (Pallflex 2500 QAT-UP) substrates (75 mm diameter) and back filters (90 mm diameter). Dichotomous data were blank corrected based on analysis of impactor substrates and back filters through which air was briefly pulled (~ 15 s). Between 2006 and 2009, a blank impactor was exposed and processed approximately once every two weeks.

With the exception of the differences noted above, aerosol sampling and handling procedures prior analysis during both periods were virtually identical. Sampling was controlled by sensors that activated pumps only during periods of no precipitation





when surface winds were off the ocean sector at speeds greater than 1 ms^{-1} . Air volumes were measured with sharp-edged flow tubes and normalized to standard temperature and pressure (0 °C and 1 atm). Bulk cassettes and impactors were cleaned with $18 \text{ M}\Omega \text{ cm}^{-1}$ deionized water (DIW), and dried, loaded, and unloaded in a Class

- ⁵ 100 clean bench mounted in a laboratory container at the base of the tower. After recovery, exposed sample and blank filters and substrates were folded in half, sealed in clean polyethylene bags, and stored (and shipped) frozen prior to analysis. Samples and corresponding blanks during each period were processed and analyzed using identical analytical procedures.
- ¹⁰ From July 1988 through April 1997 and from July 2006 through June 2009, wetonly precipitation in unsectored air was sampled daily from the top of the THAO tower in precleaned 13.2 L polyethylene buckets mounted in an automated collector (Galloway et al., 1993). Precipitation amount was measured in parallel with both bulk and recording rain gauges; precipitation amounts reported herein correspond to those
- ¹⁵ measured with the bulk gauge. After recovery, sample aliquots (250 mL or less for low volume events) were transferred to precleaned polyethylene bottles, sterilized on site via addition of 500 µL CHCl₃ to prevent microbial activity (Keene et al., 1983; Herlihy et al., 1987), and stored refrigerated prior to analysis. During both periods, precipitation was sampled and processed prior to chemical analysis using identical procedures.

20 2.2 Sample analysis

Aerosol sampled from 1988 to 1998 was analyzed at the University of Miami (UM) for SO₄²⁻, NO₃⁻, and NH₄⁺ by suppressed ion chromatography (IC), and for Na⁺ by flame atomic absorption spectroscopy (AA) (Galloway et al., 1993; Savoie et al., 2002). From 1988 to 1996, particulate CH₃SO₃⁻ was also measured at UM by IC. Aerosols sampled from 2006 to 2009 were analyzed at the University of Virginia (UVA) by IC for CH₃COO⁻, HCOO⁻, (COO)₂²⁻, CH₃SO₃⁻, SO₄²⁻, CI⁻, Br⁻, NO₃⁻, NH₄⁺, Na⁺, K⁺, Mg²⁺, and Ca²⁺ (Moody et al., 2014).





Precipitation sampled during both periods was analyzed for major ionic constituents at UVA. From 1988 to 1997, H⁺ was measured by electrode and meter, anions (CH₃COO⁻, HCOO⁻, CH₃SO₃⁻, SO₄²⁻, Cl⁻, Br⁻, and NO₃⁻) were measured by IC, NH₄⁺ was measured by the automated indophenol blue technique, and base cations (Na⁺, K⁺, Mg²⁺, and Ca²⁺) were measured by AA (Galloway et al., 1993). From 2006 to 2009, H⁺ was measured by electrode and meter and the other major ions (the same suite as indicated above for aerosols during this period) were measured by IC using procedures described by Moody et al. (2014).

2.3 Data quality and comparability

- ¹⁰ The quality of data for aerosols and precipitation sampled from 1988 to 1998 is described in detail by Galloway et al. (1993) and Savoie et al. (2002). During both the earlier and later sampling periods, data for aerosol samples that corresponded to in-sector times of less than 2.4 h (10 % of a day) exhibit low signal-to-noise and may not be representative; consequently, these results were excluded from the final quality-
- assured data set. In addition, five samples contained unusually high concentration of sea-salt constituents (30–380% higher than the sample with the next highest concentrations) and/or operator notes that suggest direct contamination by rainwater or splash; data for these samples were also excluded from the final quality-assured data set.
- Detection limits (DLs) for all analytes measured from 2006 to 2009 were estimated following Keene et al. (1989). Because in-sector sampling times for aerosols included in the data set varied from 10 to 100% of the corresponding deployment times, DLs for particulate-phase species varied among samples and, thus, were calculated individually for each analyte in each sample (Moody et al., 2014). Analytical performance was verified by intercomparison of the UVA laboratory with those at UM and the University of New Hampshire, among others; routine analysis of audit solutions from the National Institute of Standards and Technology, the World Meteorological Association, and the US Environmental Protection Agency; periodic analysis of



standard additions to samples, and evaluation of ion balances and constituent ratios. These comparisons indicate that the ionic data generated by both UVA and UM are unbiased.

Based on their thermodynamic properties (Henry's Law and dissociation constants),
the equilibrium phase partitioning of HNO₃ and NH₃ varies as a function of aerosol solution pH; HNO₃ partitions preferentially with the less acidic super-um size fractions of marine aerosol whereas NH₃ partitions preferentially with the more acidic sub-um size fractions (e.g., Moody et al., 2014). In addition, larger aerosol size fractions in ambient air may be undersaturated with respect to the gas phase because (1) their
surface-to-volume ratios are relatively low and thus equilibration times are relatively slow and (2) they exhibit relatively short atmospheric lifetimes against deposition (e.g., Keene et al., 2004). When chemically distinct marine aerosols are sampled in bulk (or in relatively coarse size fractions) the pH of the mixed sample may diverge from those for the size fractions with which most NO₃⁻ and NH₄⁺ was associated in ambient air and

- thereby drive artifact phase changes (e.g., Keene et al., 1990). Explicit evaluation of the magnitude of such artifacts is beyond the scope of this study but, because different sampling methodologies were employed during the earlier and later periods, artifacts of this nature could influence temporal trends based on measured particulate-phase concentrations of these analytes. We return to this point below. Because H₂SO₄ is
- highly soluble over the reported range in aerosol pH, particulate SO_4^{2-} is not subject to sampling artifacts of this nature.

2.4 Calculations

Contributions to measured SO_4^{2-} from sea-salt and non-sea-salt sources were differentiated following (Keene et al., 1986) for all samples for which the measured concentrations of total SO_4^{2-} and the sea-salt reference species were above detection limits (DL). Between 1988 and 1998, nss SO_4^{2-} associated with aerosols was calculated using Na⁺ as the sea-salt references species and the mass ratio of SO_4^{2-} to Na⁺



in surface seawater reported by Millero and Sohn (1992) (0.2516). As indicated above, between 2006 and 2009, super- and sub- μ m aerosol size fractions were sampled on 75 mm and 90 mm diameter quartz-fiber substrates and filters, respectively. Background concentrations of the two mostly commonly used sea-salt reference species (Na⁺ and Mg²⁺) in extracts of blanks for the two filter sizes differed such that Na⁺ offered greater resolution in calculating nss SO₄²⁻ associated with the super- μ m-diameter size fraction. Super- and sub- μ m-diameter nss SO₄²⁻ was calculated accordingly based on the mass ratios of SO₄²⁻ to Na⁺ and SO₄²⁻ to

Mg²⁺ in surface seawater reported by Wilson (1975), (0.2518 and 2.102, respectively). Minor differences between the seawater compositions reported by Millero and Sohn (1992) and Wilson (1975) (less than 0.1%) used in the above calculations were an insignificant source of bias in resulting nss concentrations. Over the entire period of record, nss SO₄²⁻ in precipitation was calculated based on Na⁺ as the reference
 species and the seawater composition reported by Wilson (1975).

Concentrations of ionic species associated with super- and sub-µm-diameter aerosol size fractions measured during 2006–2009 were summed (Moody et al., 2014) for comparison with the chemical composition of aerosol sampled in bulk during the earlier period of record. To evaluate temporal trends over the entire period of record, data were

- ²⁰ binned into sample years based on the midpoint of the each daily sampling interval. Aerosol constituents are reported as annual average concentrations and precipitation constituents are reported as annual volume-weighted-average (VWA) concentrations. FLEXPART retroplumes (see below) were not available to characterize transport for the samples collected prior to 1 January 1989. Consequently, from 1989 through 1997,
- sample years were defined as 1 January through 31 December and temporal trends were interpreted based on the corresponding mid-year date of 1 July. Because the later 3 year period of record started in mid 2006 and ended in mid 2009, sample years were defined as 1 July through 30 June and interpreted based on the mid-year date of 1 January. This approach yielded three full (12 month) sample years over the period of





record. If based on calendar year, the last six months of 2006 and the first six months of 2009 would have been excluded yielding only two years for the analysis.

2.5 Atmospheric transport

The particle dispersion model FLEXPART (Stohl et al., 1998, 2005) was run in backward mode (Stohl et al., 2003; Seibert and Frank, 2004) to generate the column residence times and footprint residence plots for source regions associated with the sampled air parcels. FLEXPART was driven with ECMWF analyses of 0.36° resolution and accounts for turbulence and deep convection in addition to the transport by gridresolved winds. At 15:00 UTC (the approximate midpoint of the deployment times for aerosol and precipitation samplers), 40 000 particles were released from the location of the measurement site and followed backward in time for 10 days. Primary flow regimes were characterized based on proportional footprint residence times within four prescribed source regions from which air was transported to Bermuda. See Moody et al. (2014) for detailed explanation of the procedure employed to identify primary air-

mass source regions for aerosols sampled at Bermuda during 2006 through 2009. An identical approach was applied to characterize source regions for aerosols sampled during 1989–1997 and for precipitation sampled during both periods. Wind fields were not available for the first six months of 1990. Because retroplumes could not be calculated for that period, annual statistics for 1990 were not included in the evaluation of trends associated with air transported from different source regions.

Moody et al. (2014) identified two distinct flow regimes that transported aerosols from eastern North America to Bermuda. One was associated with predominant flow from the northeastern United States (NEUS) and southern Canada and the other with predominant flow from southeastern US (SEUS) and the Gulf of Mexico. To evaluate

trends associated with the combined flow off northern and southern regions of eastern North America, the NEUS and SEUS flow regimes were merged for this analysis into a single group (referred to as "NEUS/SEUS"). Transport from northern Africa and the tropical NAO region (referred to as "Africa") was associated with trade-wind





flow along the southern portion Bermuda High (BH). The "Oceanic" flow regime was generally associated with long trajectories over the open ocean under anticyclonic flow around the BH. The "North" transport regime represented transitional flow from the relatively clean open ocean to the polluted northeastern US. Data evaluated herein were classified accordingly; the total numbers of samples and corresponding mean 5 precipitation amounts in the final data base partitioned by flow regime and year are summarized in Table 1. The relatively fewer numbers of aerosol samples for the period 2006–2009 reflects the fact that (1) dichotomous rather than bulk aerosol was sampled during the later period and thus the super- and sub-um-diameter size fractions of each sample were segregated and analyzed separately and (2) sampling rates through 10 the high-flow impactor were a factor of 10 lower that those through the hi-volume bulk sampler. Although analytical resolution per unit analyte concentration in extract solutions was higher during the latter period, the absolute amount of analyte present in each size fraction was lower relative to bulk aerosol during the earlier part of the record. Consequently, relatively fewer aerosol constituents were present in samples at 15 concentrations above DLs during 2006-2009.

2.6 Statistical evaluations

The statistical package SPSS was used to evaluate temporal trends in the data using two complimentary approaches. (1) Trends were characterized based on slopes of standard linear regressions (SLRs) for annual average concentrations of aerosol constituents vs. time and for annual VWA concentrations of precipitation constituents vs. time. (2) Temporal trends in annual mean (for aerosols) and annual VWA (for precipitation) constituent concentrations were also characterized based on slopes calculated using the bootstrapping method (Freedman, 1981). The significance of each bootstrap slope was evaluated using a two-tailed test based on 1000 random samples of the corresponding data subset. Relative to SLRs, this method provides a more robust evaluation of significance in trends. For smaller data sets, the approach resulted in redundant subsampling. In these cases the actual significance thresholds may



have deviated somewhat from the calculated values. However, sensitivity evaluations employing smaller numbers of random samples yielded virtually identical results suggesting that any such deviations were small to negligible. Interannual variability in mean or VWA constituent concentrations is the only source of variance in the SLRs whereas seasonal variability contributes significantly to overall variance based on the bootstrap method. Consequently, substantially less overall variance in long-term trends is explained by bootstrapping. Because this analysis focuses on long-term trends, reported variances correspond to those for SLRs.

In addition to the calculations summarized above, trends in annual median aerosol and precipitation concentrations and in per-event wet-deposition fluxes were also quantified for selected data subsets to evaluate the sensitivity of results to the underlying sample statistics used in the analysis. Results were reasonably consistent with those based on annual mean concentrations for aerosols and annual VWA concentrations for precipitation. We interpret results based on these latter more widely used conventions for reporting and evaluating concentrations of aerosol and precipitation constituents thereby facilitating direct comparison with previously published results.

3 Results and discussion

5

3.1 Temporal trends in the chemical composition of aerosols and precipitation

Temporal trends in annual average concentrations of particulate nss SO₄²⁻, NO₃⁻, and NH₄⁺ and corresponding trends in annual VWA concentrations in precipitation based on all quality-assured data over the period of record at Bermuda are depicted in Fig. 1. Nss SO₄²⁻ concentrations for both aerosols and precipitation decreased significantly but the percentage decrease in VWA concentrations based on the SLR for precipitation (49%) was greater than the corresponding decrease in annual mean concentrations for aerosols (24%, Table 2). In contrast, the VWA concentration of NH₄⁺ in precipitation



increased significantly (70%) over the period of record whereas the trend in particulate NH_4^+ was insignificant (Table 2). Concentrations of NO_3^- associated with aerosols and precipitation also did not vary significantly (Table 2). These results suggest long-term changes in atmospheric composition and also indicate that relative trends in the composition of precipitation and near-surface aerosols diverge. However, substantial interannual variability is evident for all species in both phases (Fig. 1a and b) and the corresponding linear regressions explain less than 50% of the variance over the period

3.2 Temporal trends based on FLEXPART flow patterns

of record.

- ¹⁰ Year-to-year variability in frequencies of transport from source regions is one of the major factors that drives interannual variability in atmospheric composition at a given location. Such influences can be minimized by evaluating trends in concentrations as a function of transport regime. As was the case for the data subset corresponding to aerosols sampled from 2006 to 2009 (Moody et al., 2014), segregation by source
- ¹⁵ region differentiated analyte concentrations into chemically distinct subgroups. For most years, the upper limits for mean concentrations of aerosol and precipitation constituents were associated with polluted NEUS/SEUS flow from North America whereas the lower limits were associated with background Oceanic flow around the BH (Fig. 2). Analytes associated with the Africa and North flow regimes generally fell
- within the range bounded by the other two regimes. However, numbers of samples per year varied substantially among flow regimes (Table 1). The relatively small numbers of samples collected under Oceanic and North flow in particular constrain the power of statistical approaches for evaluating associated temporal trends and thus caution is warranted in their interpretation.
- Because regional transport varies seasonally, segregating data by source region also segregated to some extent by season. For example, efficient transport of North American emissions over the western NAO is typically associated with frontal passages, which are most frequent during winter and spring, whereas transport from





Northern Africa to the western NAO in association with the easterly trade wind regime is most frequent during summer. Factors other than precursor emissions within and the frequencies of flow from source regions also contribute to seasonal variability in aerosol and precipitation composition among flow regimes. Relative to colder months,

- ⁵ higher ultraviolet radiation and temperatures during warmer months sustain faster rates of precursor oxidation and secondary aerosol production. The temperature-dependent emissions of NH₃ contribute to seasonal variability in the nucleation and growth of new particles. Seasonal variability in precipitation fields and in the temperature-dependent partitioning of semi-volatile species between the gas and particulate phases also drive
- seasonal variability in atmospheric lifetimes and associated transport and deposition fields. Although these and other processes contribute to overall variability in aerosol and precipitation composition as a function of source region, our data lack adequate resolution to quantitatively differentiate relative influences. See Moody et al. (2014) for a detailed evaluation of season variability in the sources and composition of aerosols at Bermuda between 2006 and 2009.

3.2.1 nss SO₄²⁻

Based on slopes for SLRs fit to annual average concentrations, VWA nss SO₄²⁻ in precipitation associated with the NEUS/SEUS flow regime decreased significantly (61%) over the period of record whereas the corresponding trend in particulate nss SO₄²⁻ was marginally insignificant (Fig. 2a and b, Table 2). The absence of a significant trend in particulate nss SO₄²⁻ for the NEUS/SEUS flow was driven in part by the missing data for 1990. Based on all data (Fig. 1a), 1990 exhibited the highest mean concentration over the period and most samples during each year were associated with the NEUS/SEUS regime (Table 1). Between 1989 and 2009, SO₂ emissions over the US decreased by 64% (Fig. 3; EPA, 2013). Although the overall decrease in VWA

nss SO_4^{2-} in precipitation based on the SLR was similar to that for SO_2 emissions over the US during the period of record, the temporal patterns within the period of record





were distinct. Most of the decrease in VWA nss SO_4^2 was associated with a significant negative trend between 1991 and 1996 (SLR slope = -0.62, $r^2 = 0.84$, Fig. 2b). Based on the SLR for 1991 through 1996, annual VWA nss SO_4^{2-} during the period decreased by 51 % whereas, over the same period, SO_2 emissions in the US decreased by only

- ⁵ 18% (Fig. 3, EPA, 2013). It is evident from the above that (1) the declines in nss $SO_4^{2^-}$ associated with both near-surface aerosols and precipitation under NEUS/SEUS flow were not directly proportional to that for SO_2 emissions over the US and (2) factors other than SO_2 emissions in the US influenced trends in nss $SO_4^{2^-}$ associated with near-surface aerosols and precipitation under NEUS/SEUS flow differentially.
- ¹⁰ Significant amounts of nss SO_4^{2-} associated with aerosols and precipitation at Bermuda originate from the atmospheric oxidation of dimethylsulfide ((CH₃)₂S) produced in the surface ocean by marine biota and subsequently emitted to the atmosphere (Galloway et al., 1989; Savoie et al., 2002; Moody et al., 2014); potential variability in biogenic contributions may have contributed to the differences in trends
- ¹⁵ noted above. To evaluate potential influences of biogenic S, we employed the mass ratio of particulate nss SO_4^{2-} to $CH_3SO_3^{-}$ yields from $(CH_3)_2S$ oxidation estimated from measurements at Bermuda (18.8±2.2, Savoie et al., 2002) to differentiate relative contributions from marine biogenic and anthropogenic sources to annual average concentrations of particulate nss SO_4^{2-} associated with NEUS/SEUS flow (Fig. 4a).
- ²⁰ Corresponding trends in annual VWA nss SO_4^{2-} in precipitation were calculated (Fig. 4b) based on the assumption that the same product yields apply to precipitation. Trends in $CH_3SO_3^-$ (and thus biogenic contributions) for both aerosols and precipitation were insignificant, which indicates that long-term temporal variability in nss SO_4^{2-} associated with both aerosols and precipitation was driven primarily by differential variability in anthropogenic contributions. These results also indicate that relative contributions of anthropogenic sources to VWA nss SO_4^{2-} in precipitation at Bermuda have decreased more rapidly that those to mean nss SO_4^{2-} associated with aerosols (Fig. 4). For example, based on the corresponding SLRs, percentage contributions of





anthropogenic sources to nss SO_4^{2-} in precipitation and aerosols under NEUS/SEUS flow were roughly similar at the beginning of the record (73 and 72%, respectively) whereas they diverge to a greater degree by the end of the record (41 and 68%, respectively). We return to this issue below.

VWA nss SO_4^{2-} associated with the Oceanic flow regime decreased significantly whereas the corresponding trend in particulate nss SO_4^{2-} was insignificant (Fig. 2a and b, Table 2). Available evidence suggests that virtually all particulate nss SO_4^{2-} in near-surface air sampled under Oceanic flow at Bermuda between 2006 and 2009 originated from the oxidation of (CH₃)₂S (Moody et al., 2014). Corresponding trends in CH₃SO₃ associated with both aerosols and precipitation sampled under 10 Oceanic flow over the entire period of record were insignificant. However, the relatively small numbers of observations per year (Table 1) coupled with seasonal variability in CH₃SO₃ production (Moody et al., 2014) constrains resolution in explicitly evaluating biogenic vs. anthropogenic contributions to total nss SO_4^{2-} for this regime based on product yields as was done for the NEUS/SEUS regime. However, the similarities in annual mean particulate-phase nss SO_4^{2-} under Oceanic flow during both the earlier and later periods of record (as reflected in the corresponding lack of a significant trend) indicates that the background particulate nss SO_4^{2-} from biogenic sources in near-surface aerosol remained fairly constant over this 20 year period. In contrast, VWA nss SO_4^{2-} in precipitation sampled under the same flow regime exhibited 20 significant decreasing concentrations. Most nss SO_4^{2-} in precipitation originates from the aqueous-phase oxidation of SO₂ in cloud droplets and scavenging of particulate nss SO_{A}^{2-} from the air column (e.g., Mari et al., 2000; von Glasow et al., 2002) whereas aerosols evaluated herein were sampled near the surface. Our results suggest the possibility that the transport of pollutant S emitted from surrounding continents 25 sustained relatively higher concentrations of oxidized S in aged marine air aloft, which





with precipitation relative to aerosols in NEUS/SEUS flow discussed above would also be consistent with this hypothesis. Air masses transported off eastern North America over the western NAO may rise up over and become decoupled from the marine boundary layer (MBL) (Neuman et al., 2006) whereas most biogenic $(CH_3)_2S$ emitted

- ⁵ from the ocean surface is oxidized within the MBL (e.g., Savoie et al., 2002). SO₂ in the MBL is subject to oxidization in super- μ m marine aerosols that have short lifetimes against deposition (e.g., Keene et al., 1998; von Glasow et al., 2002). In addition, both SO₂ and sub- μ m particulate nss SO₄²⁻ in the MBL dry deposit directly to the surface ocean. Consequently, oxidized S above the MBL, which is primarily
- ¹⁰ anthropogenic, exhibits longer atmospheric lifetimes against removal relative to that within the MBL, which includes most of the biogenic component. The differential vertical distributions and associated atmospheric lifetimes of oxidized S from anthropogenic vs. biogenic sources probably account for relatively greater anthropogenic contributions to nss SO_4^{2-} in precipitation vs. near-surface aerosols. As anthropogenic emissions from
- ¹⁵ surrounding continents decreased over the period of record, contributions of pollutant S to background concentrations in aged marine air aloft over the NAO would have also decreased thereby contributing to the observed decline (61 %) in VWA nss SO₄²⁻ in precipitation sampled under Oceanic flow (Fig. 2b, Table 2).

Particulate nss SO_4^{2-} associated with the Africa flow regime decreased significantly

- ²⁰ (34%) over the period of record (Fig. 2a, Table 2). Corresponding VWA nss SO_4^{2-} concentration in precipitation under Africa flow generally decreased, but unlike aerosols, this trend was marginally insignificant (Fig. 2b, Table 2). We hypothesize that large reductions in SO₂ emissions over Europe (Vestreng et al., 2007) coupled with transport over the Atlantic Ocean in association with the easterly trade winds (e.g.,
- ²⁵ Savoie et al., 1989) contributed to the trend of decreasing particulate nss SO_4^{2-} in the Africa regime. To the extent that interactions with mineral aerosol influence that atmospheric lifetime of oxidized S against deposition, the long-term temporal variability in mineral aerosol over the western NAO (Evan et al., 2006) may have influenced the trend in particulate nss SO_4^{2-} . Trends in nss SO_4^{2-} associated with aerosols and



precipitation sampled under North flow were both not significant (Fig. 2a and b, Table 2).

3.2.2 NO₃⁻

Under NEUS/SEUS flow, trends based on SLRs fit to NO₃ concentrations associated with both aerosols and precipitation over the entire period of record were insignificant (Fig. 2c and d; Table 2). In contrast, emissions of NO_x over the US decreased by 39% during this period (Fig. 3, EPA, 2013). The reported increase in NO_x emissions between 2001 and 2002 reflects an artifact that resulted from a change in the method used to estimate road emissions (R. L. Dennis, US EPA, personal communication, 2014). Like VWA nss SO_4^{2-} in precipitation, VWA NO_3^{-} also decreased significantly between 1991 and 1996 (SLR slope = -0.50, $r^2 = 0.83$) but, unlike nss SO₄²⁻, NO₃⁻ concentrations in precipitation recovered during the latter period (2006 to 2009) to levels similar to those during the late 1980s and early 1990s (Fig. 2d). In contrast, NO_x emissions over the US decreased by only 2% between 1991 and 1996 (Fig. 3, EPA, 2013). Comparable decreases in particulate NO₃⁻ between 1991 and 1996 were not evident (Fig. 2c). The annual average NO_3^- concentrations associated with aerosols under the NEUS/SEUS regime were generally higher than those for other regimes whereas VWA concentrations for precipitation overlapped to a greater degree (Fig. 2c and d). However, temporal trends in NO₃⁻ associated with both near-surface aerosols and precipitation under NEUS/SEUS flow appear to be largely decoupled from those of 20 upwind NO_x emissions in the US. In addition, temporal variability in particulate NO₃⁻ in near-surface aerosols was distinct from that for VWA NO₃⁻ in precipitation. We speculate that the efficient oxidation of NO_x to HNO₃ in marine air coupled with the subsequent partitioning of most HNO₃ with large, short-lived marine-aerosol (e.g., Russell et al., 2003; Keene et al., 2009) tended to dampen temporal variability in near-25

surface particulate NO_3^- under a given flow regime relative to both upwind emissions of precursor NO_x and particulate species such as nss SO_4^{2-} and NH_4^+ that are associated





primarily with longer-lived sub-um aerosol in marine air. The decreasing concentrations of primary marine aerosol with altitude and corresponding influences on HNO_3 phase partitioning may have also contributed to divergence in the temporal variability in near-surface particulate NO_3^- relative to NO_3^- scavenged from the column by precipitation. However, corresponding trends in mean particulate Na^+ and VWA Na^+ in precipitation at Bermuda (not shown) over the study period were not significant, which implies that interannual variability in production of marine aerosol did not directly influence tends in NO_3^- .

VWA NO₃⁻ in precipitation associated with flow from the North revealed a marginally
 significant trend of increasing concentrations (Fig. 2d, Table 2). All other trends in NO₃⁻ concentrations for both aerosols and precipitation were insignificant.

3.2.3 NH₄⁺

Temporal trends in NH_4^+ associated with both aerosols and precipitation under NEUS/SEUS flow were similar to those for NO_3^- ; none were significant (Fig. 2e and f and Table 2). The lack of significant trends in NH_4^+ concentrations is consistent with expectations based on trends in reported NH_3 emissions over the US, which increased by less than 1 % between 1990 and 2009 (Fig. 3, EPA 2013). The reported decrease in NH_3 emissions between 2000 and 2001 is an artifact associated with a change in the method used to estimate emissions (R. L. Dennis, US EPA, personal communication, 2014). Like both nss SO_4^{2-} and NO_3^- , VWA NH_4^+ concentrations in precipitation decreased significantly during the early period of record (1991–1996) (SLR slope = -0.19, $r^2 = 0.40$) and then recovered (Fig. 2f). In contrast, between 1991 and 1996, NH_3 emissions in the US increased by 8 % (EPA, 2013).

Like nss SO₄²⁻, particulate NH₄⁺ associated with the Africa regime decreased significantly over the study period. In contrast, VWA NH₄⁺ in precipitation sampled under Africa flow increased significantly (Fig. 2e and f, Table 2). As discussed in more detail below, these opposite trends may be driven in part by decreases in the acidity of the





multiphase system and associate shifts in gas-aerosol phase partition of NH_3 . All other trends in NH_4^+ were insignificant.

3.3 Implications

It is evident from the above that, in many cases, interannual variability and long-term trends in mean nss SO_4^{2-} , NO_3^{-} , and NH_4^{+} in near-surface aerosol associated with the major flow regimes that transport air to Bermuda do not mirror corresponding patterns in VWA SO_4^{2-} , NO_3^{-} , and NH_4^{+} in precipitation (Fig. 2). Precipitation scavenges both soluble gases and particles from the column whereas the measured aerosol composition reflects near-surface conditions. In addition, relative to the column, nearsurface air is chemically processed to a greater degree via transformations involving 10 freshly produced, large, and short-lived marine aerosol; direct interactions with the ocean surface via the dry deposition of both precursor gases and particles; and, emissions of (CH₃)₂S and, under clean conditions, NH₃ (e.g., Bouwman et al., 1997) from the ocean surface. These near-surface processes contribute to differential variability in the composition of near-surface aerosols and precipitation reported herein. These results imply that, relative to near-surface aerosols, the composition of precipitation may be a better indicator of column-integrated trends in soluble atmospheric constituents. We also note that aerosol sampling was sector controlled whereas precipitation sampling was not. However, available evidence suggests that local emissions have minor to negligible influences on the composition of wet-only 20 precipitation at Bermuda (Galloway et al., 1988, 1989, 1993).

Trends in aerosol and precipitation composition associated with the NEUS/SEUS flow regime do not track corresponding trends in emissions over the US. Indeed, the significant and roughly similar proportionate decreases in VWA nss SO_4^{2-} (factor

 $_{25}$ of 2.2), NO_3^- (factor of 1.9), and NH_4^+ (factor of 1.8) in precipitation associated with NEUS/SEUS flow between 1991 and 1996 (Fig. 2b, d and f) occurred in conjunction with reductions in SO₂ and NO_x emissions over the US of only 18 and 2 %, respectively, and an increase on NH₃ emissions of 8 % (Fig. 3, EPA, 2013). Corresponding patterns





of decreasing particulate phase concentrations were not evident (Fig. 2a, c and e). These results imply that physical processes such as trends in the efficiency of upwind removal via wet and dry deposition as opposed to trends in upwind emissions drove the similar variability evident in the VWA concentrations of these species during this period.

- In this regard, we note that variability in mean precipitation amount per sample (Table 1) did not correlate with variability in VWA precipitation composition, which indicates that dilution effects associated with storm size (e.g., Galloway et al., 1989) was not a major factor driving interannual variability or trends in precipitation composition.
- In contrast to the differences in concentrations associated with NEUS/SEUS flow and emissions in the US evident in our data, between 1980 and 2010, wetdeposition fluxes of SO_4^{2-} via precipitation and particulate SO_4^{2-} concentrations in near-surface air over the eastern US decreased by 58 and 40%, respectively and the corresponding SO_2 emissions over the US decreased by 56% (Leibensperger et al., 2012). Between 1980 and 2009, wet deposition fluxes of NO_3^- over the eastern
- ¹⁵ US decreased 33 % while corresponding NO_x emissions decreased 36 %. Particulate NO₃⁻ concentrations over the eastern US were relatively uniform from 1990 to 2000 but decreased by 23 % between 2000 and 2009 (Leibensperger et al., 2012). Wet deposition fluxes of NH₄⁺ over the eastern US during this period exhibited no significant trend whereas the corresponding particulate NH₄⁺ concentrations decrease by 30 %.
- ²⁰ Comparisons with measurements suggest that model calculations of these trends significantly underestimate the decreases in particulate NO₃⁻ and NH₄⁺ over this period (Leibensperger et al., 2012). Measured and simulated trends over the downwind western NAO were not reported.

It is evident from the above that temporal variability in (1) spatial distributions of emission fields over the NEUS/SEUS source region during the period of record, (2) the spatial distribution of transport within the NEUS/SEUS flow regime, and/or (3) spatial distributions of deposition fields under NEUS/SEUS flow contributed to temporal variability in aerosol and precipitation composition at Bermuda. However, available information, and particularly the lack of spatially resolved measurements of aerosol





and precipitation composition over the western NAO precludes quantitative evaluation of these factors. Our results highlight the difficulty in characterizing influences of emissions reductions based on measurements at individual sites located downwind.

Despite these limitations, results do provide unique insight regarding long-term trends in regional atmospheric composition. For example, although only the trend in particulate NO₃⁻⁻ was marginally significant, all trends in NO₃⁻⁻ and NH₄⁺⁻ associated with aerosols and precipitation sampled under Oceanic flow over the period of record were positive and within a factor of about 2 in magnitude (31–67 %. Figure 2c–f). The consistency of these patterns suggests that background concentrations of inorganic N species in aged marine air within the NAO basin may have increased over the past two decades. In contrast, the corresponding concentrations of VWA nss SO₄²⁻

in precipitation associated with Oceanic flow decreased significantly (Fig. 2b, Table 2). The changing mixture of acids and bases in the NAO troposphere impacts pHdependent chemical processes including the phase partitioning and associated atmospheric lifetimes against deposition for compounds with pH-dependent solubilities. For example, trends of both aerosols and precipitation based on all data (Fig. 1) and the subset of data associated with NEUS/SEUS flow (Fig. 2) suggest that, on average, concentrations of nss SO₄²⁻ decreased, NH₄⁺ increased, and NO₃⁻ remained essentially unchanged over the period of record. As noted above, negative bias in particulate NH₄⁺

²⁰ resulting from artifact volatilization of NH₃ from aerosol sampled in bulk during the early part of the record may have contributed to the positive trend in particulate NH₄⁺ but the positive trend for precipitation is not subject to such artifacts. These data imply that, on average, acidities of sub-µm-diameter aerosol solutions, with which most particulate nss SO₄²⁻ and NH₄⁺ at Bermuda was associated (Moody et al., 2014), decreased over the period, particularly during the earlier part of the record. Decreasing aerosol acidities would have shifted the phase partitioning of total NH₃ (NH₃ + NH₄⁺) towards the gas

phase (e.g., Smith et al., 2007) as is evident in the changing molar ratios of nss $SO_4^{2^-}$ to NH_4^+ over the period (Fig. 5). For example, based on the trend lines (Fig. 2, Table 2), between 1989 and 2009, ratios of mean nss $SO_4^{2^-}$ to NH_4^+ in near-surface aerosols





associated with NEUS/SEUS flow decreased from 1.2 to 1.1 (factor of 1.1) whereas the corresponding ratios for VWA concentrations in precipitation decreased from 2.0 to 0.48 (factor of 4.3). The greater decrease in ratios for precipitation suggests that relative contributions from the scavenging of gaseous NH_3 vs. particulate NH_4^+ to NH_4^+

- ⁵ in precipitation increased over the period. This increase in relative contributions from NH_3 would be consistent with expectations based on a pH-dependent shift in the phase partitioning of total NH_3 towards the gas phase coupled with efficient scavenging of both NH_3 and particulate NH_4^+ by precipitation. In addition, because dry deposition velocities to the ocean surface for NH_3 are greater than those for sub-µm aerosol size
- ¹⁰ fractions with which most NH_4^+ is associated in ambient air (Smith et al., 2007), this shift would have resulted in an increase in the average dry-deposition flux of total NH_3 to the ocean surface, a corresponding decrease in the fraction of total NH_3 removed via wet deposition, and a decrease in the atmospheric lifetime of total NH_3 against wet plus dry deposition. The long-term change in atmospheric acidity is also evident in the
- ¹⁵ significant decrease (37 %) in VWA H⁺ associated with precipitation sampled under NEUS/SEUS flow over the period of record (SLR slope = -0.29, $r^2 = 0.49$, Fig. 6). We recognize that trends reported herein are associated with large uncertainties, the corresponding changes in phase partition and deposition are nonlinear, and explicit evaluation of trends in aerosol solution pH and associated implications are not possible
- without paired measurements of particulate and gas phase species. However, these relationships provide relevant quantitative context for aspects of S and N cycling over the North Atlantic that are driven by changing emissions over surrounding continents. Increasing concentrations and deposition fluxes of total NH₃ over the NAO based on

all data and under the NEUS/SEUS flow regime imply that NH_3 emissions in source regions have increased over the period of record. However, EPA (2013) indicates that NH_3 emissions from the US did not vary significantly between 1990 and 2009. Our results suggest the possibility that EPA (2013) may underestimate increasing trends in NH_3 emissions over the US or, alternatively, that increasing transport of NH_3 and particulate NH_4^+ from more distant sources in Asia are contributing to rising



atmospheric concentrations over the NAO. In this regard, Moody et al. (2014) detected significant mineral aerosol associated with NEUS flow that available evidence suggests originated in Asia. Other studies have reported that the transport of emission products from Asia significantly impact air quality over the western US although corresponding impacts in the eastern US are relatively small to negligible (e.g., Koch et al., 2007; Lin et al., 2014).

Because HNO₃ partitions primarily with less-acidic super-µm-diameter aerosols, its cycling in marine air is largely decoupled from that of NH₃. In addition, most acidity added to larger aerosol size fractions via accumulation of HNO3 and H2SO4 is displaced to the gas phase in association with HCl volatilization, which acts to regulate the pH of super-um marine aerosols to a fairly narrow range (low 3s to high 4s) over highly variable chemical regimes (e.g., Keene et al., 2009). Consequently, relative to NH₃, HNO₃ cycling over the ocean is less sensitive to changes in atmospheric acidity. Similarly, particulate NO_3^- is also less sensitive than NH_4^+ to artifact volatilization from marine aerosol sampled in bulk.

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Our results also provide insight regarding the potential influence of reductions in SO₂ emissions from the US on direct radiative forcing by pollutant aerosols over the western NAO at Bermuda. Assuming that the marginally insignificant trend in particulate nss SO_4^{2-} associated with NEUS/SEUS flow is reasonably representative of the actual long-term decline, the SLR indicates that between 1989 and 2009, annual average 20 concentrations for this flow regime decreased by roughly 2.4 nmol m⁻³ or about 10% (Fig. 2a, Table 2). A regression fit to the scatter plot of satellite AOD at 550 nm vs. nss SO_4^{2-} associated with near-surface aerosol under NEUS flow regime during 2006– 2009 reported by Moody et al. (2014) yields a slope of 0.002 nmol⁻¹ m³. Assuming that trends in other aerosol constituents that scatter and/or absorb radiation under this 25 flow regime co-varied with particulate nss SO_4^{2-} , the corresponding decrease in AOD (0.005) coupled with satellite-derived estimates for the radiative efficiency of aerosols

per unit AOD in North American outflow over the ocean (-27 to $-60 \text{ Wm}^{-2} \delta^{-1}$) (Anderson et al., 2005) yield an estimated net warming of 0.1–0.3 W m⁻² that resulted



from the decreased shortwave scattering and absorption by nss SO_4^{2-} and associated aerosol constituents under NWUS/SEUS flow at Bermuda between 1989 and 2009. As discussed above, the greater decreasing trend for nss SO_4^{2-} in precipitation relative to near-surface aerosols suggests that the decline in particulate nss SO_4^{2-} in near surface air was less than that for oxidized S in the column scavenged by precipitation. Consequently, this estimated range in warming should be considered a lower limit. We appreciate that this approach is associated with large and poorly constrained uncertainties but it is reasonably consistent with expectations based on model estimates of the trend in net radiative warming of 0.8 W m⁻² over the US resulting from the reduction in scattering and absorption by pollutant aerosols between 1990 and 2010 (Leibensperger et al., 2012).

4 Summary and conclusions

Since the 1980s, mandated reductions in emissions of SO₂ and NO_x from anthropogenic sources in the US and Europe have resulted in decreased export
of oxidized S and N compounds from surrounding continents to the atmosphere overlying NAO. The chemical compositions of aerosols and precipitation sampled daily at Bermuda from 1989 to 1997 and from 2006 to 2009 were evaluated to quantify the magnitudes, significance, and environmental implications of associated tends in atmospheric composition. Based on all data, annual average concentrations of nss SO₄²⁻ associated with aerosols and corresponding annual VWA concentrations in precipitation decreased significantly (by 22 and 49%, respectively) whereas annual VWA concentrations of NH₄⁺ in precipitation increased significantly (by 70%) (Fig. 1, Table 2). Corresponding trends in NO₃⁻ associated with aerosols and precipitation and of particulate NH₄⁺ were insignificant.

To assess influences of emissions in upwind source regions, the chemical data were stratified based on FLEXPART retroplumes into four discrete transport regimes:





westerly flow from the eastern North America (NEUS/SEUS); easterly trade-wind flow from northern Africa and the subtropical NAO (Africa); long, open-ocean, anticyclonic flow around the BH (Oceanic); and transitional flow from the relatively clean open ocean to the polluted northeastern US (North).

- ⁵ Nss SO₄²⁻ in precipitation sampled under NEUS/SEUS and Oceanic flow decreased significantly (61 % each) over the period of record whereas corresponding trends in nss SO₄²⁻ associated with near-surface aerosols sampled under both flow regimes were insignificant (Fig. 2a and b, Table 2). Available evidence suggests that the trends for precipitation were driven primarily by decreasing emissions of SO₂ in upwind
- ¹⁰ continents and associated decreases in anthropogenic contributions to total nss SO_4^{2-} . Results support the hypothesis that, for these flow regimes, ratios of anthropogenic to biogenic contributions to nss SO_4^{2-} in the column scavenged by precipitation were relatively greater than those in near surface aerosol and, thus, precipitation provides a better indicator of column-integrated trends.
- ¹⁵ Nss SO₄²⁻ associated with aerosols sampled under African flow also decreased significantly (34%) whereas the corresponding decrease of roughly similar magnitude in nss SO₄²⁻ associated with precipitation was marginally insignificant (Fig. 2a and b, Table 2). We infer that these trends were driven in part by reductions in the emissions and transport of oxidized S compounds from Europe.
- Although partitioning the data based on flow regime segregated the results into chemically distinct subsets, the lack of significant trends in NO_3^- associated with aerosols and precipitation sampled under NEUS/SEUS flow (Fig. 2c and d, Table 2) did not reflect the large decrease in NO_x emissions in the US over the period of record (Fig. 3). The rapid oxidation of NO_x in marine air coupled with partitioning of most HNO₃ with large, short-lived marine aerosol may have dampened trends in $NO_3^$ concentrations relative to upwind NO_x emissions.

Trends of both aerosols and precipitation based on all data (Fig. 1) and the subset of data associated with NEUS/SEUS flow (Fig. 2) suggest that, on average, concentrations of nss SO_4^{2-} decreased, NH_4^+ increased, and NO_3^- remained essentially



unchanged, which implies that the total amount of acidity in the multiphase gasaerosol system in the lower NAO troposphere decreased over the period of record. This interpretation is consistent with the decreasing trend in VWA H⁺ in precipitation sampled under NEUS/SEUS flow (Fig. 6). These data suggest that, on average, acidities of sub- μ m-diameter aerosol solutions, with which most particulate nss SO²₄ and NH_4^+ at Bermuda was associated (Moody et al., 2014), decreased over the period. Decreasing aerosol acidities would have shifted the phase partitioning of total NH₃ towards the gas phase, increased the dry deposition velocities for total NH₃ to the ocean surface, increased relative contributions from the scavenging of gaseous NH₃ vs. particulate NH_4^+ to NH_4^+ in precipitation, and decreased the atmospheric lifetime 10 of total NH₃ against wet plus dry deposition. The trend of increasing NH₄⁺ associated with NEUS flow suggests that either the actual NH₃ emissions over the US may have increased faster than reported by EPA (2013) or increasing NH₃ emissions from Asia may have contributed to total NH₃ transported from eastern North America to the western NAO. 15

Assuming that the marginally insignificant trend in particulate nss $SO_4^{2^-}$ in nearsurface air associated with NEUS/SEUS flow is reasonably representative of long-term temporal variability (Fig. 2a, Table 2), decreasing concentrations over the period of record suggests a lower limit for net warming in the range of 0.1–0.3 W m⁻² that resulted from the decreased electrony and absorption by page $SO_2^{2^-}$ and appaging the

²⁰ from the decreased shortwave scattering and absorption by nss SO_4^{2-} and associated aerosol constituents transported from North America over the western NAO.

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Table 1. Numbers (Ns) of daily aerosol and precipitation samples and mean precipitation amounts (cm) per sample for all data evaluated in this study partitioned by flow regime and year. Annual data for 1989 through 1997 are binned by calendar year (1 January through 31 December) and those for 2006 through 2009 are binned from 1 July through 30 June.

Mid Year Aerosols, N	1989	1990	1991	1992	1993	1994	1995	1996	1997	2007	2008	2009	All
All Data ¹ NEUS/SEUS Africa Oceanic North	168 119 30 9 10	184 N/A N/A N/A N/A	199 136 41 11 11	186 107 45 17 17	130 95 24 3 8	152 101 28 11 12	174 104 40 9 21	182 132 32 8 10	215 133 49 10 23	113 64 26 6 17	111 68 30 8 5	84 54 22 5 3	1898 1113 367 97 137
Precipitation,	Ν												
All Data ² NEUS/SEUS Africa Oceanic North	101 59 34 5 3	72 N/A N/A N/A N/A	71 43 15 5 8	70 37 17 9 7	76 51 17 3 5	82 46 19 7 10	76 42 17 7 10	84 53 20 6 5	N/A N/A N/A N/A N/A	99 61 19 10 9	100 55 21 11 13	86 40 34 6 6	917 487 213 69 76
Precipitation	amount	t, cm											
All Data NEUS/SEUS Africa Oceanic North	0.87 0.81 1.02 0.90 0.35	1.08 N/A N/A N/A N/A	1.09 0.89 1.07 0.56 2.53	1.15 1.08 1.35 1.28 0.90	1.66 1.71 1.36 1.00 2.49	1.39 1.36 1.32 1.93 1.31	1.06 1.21 0.97 1.06 0.60	1.08 1.06 1.43 0.51 0.57	N/A N/A N/A N/A N/A	0.94 0.95 1.07 1.11 0.36	1.02 1.05 1.02 0.76 1.14	0.83 0.79 0.77 0.99 1.31	1.09 1.09 1.10 1.03 1.16

N/A indicates that either the chemical data or the corresponding FLEXPART retroplumes were not available for the indicated period (see text). ¹ No correspond to particulate page SO^{2-} NO⁻ and NH⁺, for particulate CH SO⁻ table N was 1966 due to the shorter pagied of pager and

¹ Ns correspond to particulate nss SO_4^{2-} , NO_3^{-} , and NH_4^+ ; for particulate $CH_3SO_3^{-}$, total N was 1266 due to the shorter period of record and relatively fewer samples with concentrations greater than DLs.

² Ns corresponds to nss SO₄²⁻, NO₃⁻, and NH₄⁺ in precipitation; for CH₃SO₃⁻ in precipitation, total N was 845 due to the relatively fewer samples with concentrations greater than DLs.



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Species/ Flow Regime	Bootstrap Slope (Std.Error)	SLR Slope	SLR	SLR Change ¹	Bootstrap Slope (Std. Error)	SLR Slope	SLR	SLR Change ¹
0	(nmolm ⁻³ yr ⁻¹)	$(nmol m^{-3} yr^{-1})$	r^2	(%)	$(\mu mol L^{-1} yr^{-1})$	$(\mu mol L^{-1} yr^{-1})$	r ²	(%)
		Aerosol				Precipitatior	ı	
nss SO ₄ ²⁻								
All Data	-0.26(0.08)*	-0.24	0.17	-22 %	-0.10(0.03)*	-0.11	0.67	-49 %
NEUS/SEUS	-0.16(0.11)	-0.12	-	-	-0.16(0.05)*	-0.17	0.61	-61 %
Africa	-0.28(0.12)*	-0.31	0.29	-34 %	-0.06(0.06)	-0.06	-	-
Oceanic	0.14(0.12)	0.14	-	-	-0.09(0.03)*	-0.08	0.31	-61 %
North	0.36(0.34)	0.22	-	-	0.10(0.06)	0.04	-	-
NO ₃								
All Data	-0.01(0.05)	-0.01	-	-	0.05(0.03)	0.04	-	-
NEUS/SEUS	0.03(0.07)	0.03	-	-	0.02(0.05)	0.01	-	-
Africa	-0.12(0.06)**	-0.13	0.26	-20 %	0.08(0.05)	0.07	-	-
Oceanic	0.17(.09)**	0.18	0.53	46 %	0.06(0.05)	0.06	-	-
North	0.23(0.17)	0.31	-	-	0.17(0.06)*	0.13	0.32	76%
NH ₄ ⁺								
All Data	-0.05(0.08)	-0.06	-	-	0.08(0.03)*	0.08	0.68	70%
NEUS/SEUS	0.05(0.10)	0.04	-	-	0.08(0.05)	0.09	-	-
Africa	-0.33(0.11)*	-0.33	0.35	-44 %	0.09(0.05)**	0.09	0.48	102 %
Oceanic	0.05(0.16)	0.09	-	-	0.03(0.03)	0.03	-	-
North	0.15(0.26)	0.05	-	-	0.09(0.06)	0.05	-	-

Table 2. Regression statistics and, for significant trends, percent changes in concentrations over the period of record.

* p = 0.05; significant at 95 % confidence. ** p = 0.10; significant at 90 % confidence.

¹ Percent change over the period of record relative to 1989 based on the SLR slope.

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Fig. 1. Temporal trends in **(a)** annual average concentrations of bulk particulate nss SO_4^{2-} (red), NO_3^- (blue), and NH_4^+ (green) in near-surface air (white backgrounds) and **(b)** corresponding annual VWA concentrations in precipitation (light blue backgrounds) at Bermuda based on all quality-assured data generated over the period of record. Color-coded dashed lines correspond to SLRs for annual averages and VWAs vs. time.











Fig. 2. Temporal trends in annual average concentrations of bulk particulate (a) nss SO_4^{2-} , (c) NO_3^- , and (e) NH_4^+ in near-surface air (white backgrounds) and annual VWA concentrations of (b) nss SO_4^{2-} , (d) NO_3^- , and (f) NH_4^+ in precipitation (light blue backgrounds) associated with NEUS/SEUS (red), African (yellow), Oceanic (blue), and North (green) flow regimes. Color-coded dashed lines correspond to SLRs for annual mean and VWA concentrations vs. time.



Fig. 3. Temporal trends in SO_2 , NO_x , and NH_3 emitted to the atmosphere over the continental US (EPA, 2013).





Fig. 4. Temporal trends in anthropogenic and biogenic contributions to nss SO_4^{2-} associated with **(a)** aerosols (white backgrounds) and **(b)** precipitation (light blue backgrounds) sampled under NEUS/SEUS flow.











Fig. 6. Temporal trend in annual VWA H⁺ in precipitation associated with NEUS/SEUS flow.



