Enhancing non-refractory aerosol apportionment from an urban industrial site through receptor modelling of complete high time-resolution aerosol mass spectra

4

M. L. McGuire¹, R. Y. –W. Chang^{1,2*}, J. G. Slowik^{1,2*}, C. –H. Jeong¹, R. M. Healy¹,
G. Lu³, C. Mihele³, J. P. D. Abbatt^{1,2}, J. R. Brook³, G. J. Evans¹

7 [1](Southern Ontario Centre for Atmospheric Aerosol Research, University of Toronto, 200
8 College St., Toronto, Ontario, Canada)

9 [2](Department of Chemistry, University of Toronto, St. George St., Toronto, Ontario,10 Canada)

[3](Air Quality and Research Division, Science and Technology Branch, Environment
Canada, 4905 Dufferin St., Toronto, Ontario, Canada)

*Now at (School of Engineering and Applied Sciences and Department of Earth and
Planetary Sciences, Harvard University, Cambridge, MA, United States)

15 *Now at (Paul Scherrer Institute, Laboratory of Atmospheric Chemistry, General Energy

16 Research Department, Paul Scherrer Institute, Villigen, Switzerland, Villigen, Switzerland)

- 17 Correspondence to: G. J. Evans (greg.evans@utoronto.ca)
- 18

19 Abstract

20 Receptor modelling was performed on quadrupole unit mass resolution aerosol mass 21 spectrometer (Q-AMS) sub-micron particulate matter (PM) chemical speciation 22 measurements from Windsor, Ontario, an industrial city situated across the Detroit River from 23 Detroit, Michigan. Aerosol and trace gas measurements were collected on board 24 Environment Canada's CRUISER mobile laboratory. Positive matrix factorization (PMF) was performed on the AMS full particle-phase mass spectrum (PMF_{Full MS}) encompassing 25 26 both organic and inorganic components. This approach was compared to the more common method of analysing only the organic mass spectra (PMF_{Org MS}). PMF of the full mass 27 28 spectrum revealed that variability in the non-refractory sub-micron aerosol concentration and 1 composition was best explained by six factors: an amine-containing factor (Amine); an 2 ammonium sulphate and oxygenated organic aerosol containing factor (Sulphate-OA); an 3 ammonium nitrate and oxygenated organic aerosol containing factor (Nitrate-OA); an 4 ammonium chloride containing factor (Chloride); a hydrocarbon-like organic aerosol (HOA) 5 factor; and a moderately oxygenated organic aerosol factor (OOA). PMF of the organic mass 6 spectrum revealed three factors of similar composition to some of those revealed through 7 PMF_{Full MS}: Amine, HOA and OOA.

8 Including both the inorganic and organic mass proved to be a beneficial approach to analysing 9 the unit mass resolution AMS data for several reasons. First, it provided a method for 10 potentially calculating more accurate sub-micron PM mass concentrations, particularly when 11 unusual factors are present, in this case, an Amine factor. As this method does not rely on a priori knowledge of chemical species, it circumvents the need for any adjustments to the 12 13 traditional AMS species fragmentation patterns to account for atypical species, and can thus lead to more complete factor profiles. It is expected that this method would be even more 14 15 useful for HR-ToF-AMS data, due to the ability to better understand the chemical nature of 16 atypical factors from high resolution mass spectra. Second, utilizing PMF to extract factors 17 containing inorganic species allowed for the determination of extent of neutralization, which could have implications for aerosol parameterization. Third, subtler differences in organic 18 19 aerosol components were resolved through the incorporation of inorganic mass into the PMF 20 The additional temporal features provided by the inorganic aerosol components matrix. 21 allowed for the resolution of more types of oxygenated organic aerosol than could be reliably resolved from PMF of organics alone. Comparison of findings from the PMF_{Full MS} and 22 23 PMF_{Org MS} methods showed that for the Windsor airshed, the PMF_{Full MS} method enabled 24 additional conclusions to be drawn in terms of aerosol sources and chemical processes. While 25 performing PMF_{Org MS} can provide important distinctions between types of organic aerosol, it 26 is shown that including inorganic species in the PMF analysis can permit further 27 apportionment of organics for unit mass resolution AMS mass spectra.

28 **1** Introduction

Atmospheric aerosol or particulate matter (PM) is known to have important implications on atmospheric visibility (Watson, 2002), climate change (IPCC, 2013), and human health (Pope and Dockery, 2006; Anderson et al., 2012; Brook et al., 2010). Understanding the sources and processes responsible for PM composition and concentration is critical to enacting

effective PM reduction strategies. Receptor modeling of PM chemical speciation data is one 1 2 method towards achieving this. Historically, receptor modeling studies have focused on 3 understanding integrated filter measurements, which have been particularly useful for 4 providing an overview of the main chemical components of major source categories, and their 5 longer-term temporal trends (Gordon, 1980; Hopke, 2003; Watson et al., 2008). More recently, receptor modeling analyses have been focused on online high time-resolution 6 7 chemical analysis techniques, as they can provide additional insight on sources and processes 8 not captured by the chemical or temporal resolution of daily filter measurements.

9 Aerosol mass spectrometry is among the most widely used high-time resolution PM chemical 10 speciation methods that can be used to quantify the impacts of non-refractory source 11 components, including both organic and inorganic components. Receptor modeling using 12 Positive Matrix Factorization (PMF), has become a useful analytical technique to further 13 understand the origins of AMS measured aerosol. Among these studies, most have focused 14 on the organic fraction of the AMS mass spectrum (e.g., Lanz et al., 2007; Ulbrich et al., 15 2009) in an effort to resolve uncertainty regarding the sources and processes contributing to secondary organic aerosol (SOA), an aerosol component with potential implications on 16 17 climate (IPCC, 2013). Many of these studies have focused on the application of factor 18 analysis to the organic mass fraction in an attempt to deconvolve it into descriptive sub-19 components, namely a hydrocarbon-like organic aerosol (HOA) factor, and an oxygenated 20 organic aerosol (OOA) factor. Examination of datasets from numerous, diverse environments 21 has shown that the OOA fraction often splits into two sub-components, OOAI and OOAII (Zhang et al., 2011). Observations of their temporal behavior have shown that these two 22 23 factors typically exhibit different volatility regimes, whereby OOAII exhibits significant diurnal variability associated with condensation and volatilization from temperature cycling. 24 25 By contrast, OOAI has been mainly associated with synoptic flow regimes, with no 26 significant association with temperature cycling. The semi-volatile OOAII type factor was 27 first reported in a study by (Lanz et al., 2007), although its volatile nature was first 28 substantiated with external measurements in a study by Huffman et al. (2009); in the latter 29 study it was shown that decreased volatility of OA factors was associated with increasing 30 oxygenation, or oxygen to carbon (O/C) ratio (Huffman et al., 2009). As a result, the OOAI and OOAII factors are often referred to in the literature as low-volatility OOA (LV-OOA), 31 32 and semi-volatile OOA (SV-OOA) respectively (Jimenez et al., 2009). While the HOA and 33 OOA components have been the most widely observed organic components deconvolved in

the multitude of AMS studies performed to date (Zhang et al., 2007), other factors have been 1 identified, such as biomass burning organic aerosol (BBOA) (e.g., Aiken et al., 2009), amine-2 3 containing organic aerosol (e.g., Aiken et al., 2009; Sun et al., 2012; Docherty et al., 2011; 4 Hildebrandt et al., 2011) and even cooking organic aerosol (COA) (e.g., Lanz et al., 2007; 5 Allan et al., 2010; Sun et al., 2011; Mohr et al., 2012; Crippa et al., 2013a). In many studies, 6 correlation analysis of organics with inorganic species has been used to further ascertain the 7 sources and processes contributing to a particular factor. For instance, significant correlations have been found between OOAI and SO_4^{2-} , and between OOAII and NO_3^{-} (Lanz et al., 2007; 8 9 Ulbrich et al., 2008).

10 This study presents a receptor modeling analysis of high time-resolution unit mass resolution 11 quadrupole AMS measurements. Aerosol and trace gas speciation measurements were made onboard Environment Canada's CRUISER (Canadian Regional and Urban Investigation 12 13 System for Mobile Research) mobile laboratory at MicMac Park in Windsor, Ontario in the 14 winter of 2005. A different approach was taken in this study with respect to the majority of 15 previous receptor modeling analyses of the non-refractory sub-micron chemical composition, as PMF was applied to the full mass spectrum, comprising both the organic and inorganic 16 17 components. To the authors' knowledge, combined PMF analysis of organic and inorganic 18 AMS mass spectra has been performed in only three other studies (Chang et al., 2011; Sun et 19 al., 2012; Crippa et al., 2013b). Of these, the study by Sun et al. was the only one to include all inorganic and organic species together. While PMF had previously been applied to data 20 21 including AMS derived bulk concentrations of inorganic and organic species (e.g., Buset et 22 al., 2006; Crippa et al., 2013b), Chang et al. were the first to apply the PMF multivariate 23 deconvolution algorithm to combined organic and inorganic mass spectra, in that case to 24 ambient arctic aerosol (Chang et al., 2011). However, low ambient aerosol concentrations, 25 and low associated signal to noise ratios, prevented the inclusion of all species in the analysis. As a result, NH₄⁺ was excluded, which precluded certain conclusions regarding aerosol 26 27 neutralization from being drawn. Nonetheless, four factors were extracted in that study, 28 including factors representing marine biogenic emissions (containing methanesulphonic acid 29 or MSA), continental emissions, ship emissions, and OOA. Each factor was characterized by 30 differing degrees of cross-apportionment between organic and inorganic species. Eight factors were identified in the study by Sun et al., many more than found in the Arctic study, 31 32 mainly due to the urban sampling location in New York City (Sun et al., 2012). Similar to 33 Chang et al., significant organic and inorganic cross-apportionment was noted for most of the 1 factors. In the study by Crippa et al., only SO_4^{2-} related ions were included in the PMF 2 analysis in addition to organics, which allowed for the apportionment of SO_4^{2-} ions to marine 3 and terrestrial aerosol factors (Crippa et al., 2013b).

This study focuses on the physical interpretation of cross-apportionment between organic and inorganic non-refractory sub-micron PM species between factors. Drawing upon cold condition measurements from a complex, industrialized site, this analysis illustrates how this methodology can help to better understand underlying aerosol sources, and processes, and identify new scientific and methodological conclusions.

9 2 Experimental Methods

10 **2.1** Aerosol mass spectrometer measurements

11 Aerosol and trace gas measurements were collected on board Environment Canada's 12 CRUISER mobile laboratory, which was stationed at MicMac Park in Windsor (42°17'5.38"N, 83° 4'31.42"W) in January and early February 2005. Located next to Detroit 13 14 on the Canada-US border in south-western Ontario, Windsor has historically been known experienced frequent episodes of poor air quality. Local sources of PM are numerous and 15 diverse due to a large manufacturing base, including sources such as steelmaking, salt and 16 17 gypsum mining, petrochemical refining, and coal-fired power generation. Another significant 18 local source is traffic, given that the Windsor-Detroit border crossing is the largest 19 international border crossing between Canada and the US. Regionally, Windsor is impacted 20 by many sources, but perhaps most significantly by coal-fired power plants to the south in the 21 Ohio River valley. A map of the Windsor/Detroit area is shown in Figure 1.

22 Chemical speciation measurements of submicron PM were made on board CRUISER using a 23 unit mass resolution quadrupole aerosol mass spectrometer (Q-AMS) (Aerodyne Research 24 Inc., Billerica, MA, USA). The AMS sampled from 12 January – 2 February 2005, except for 25 a period of three days between 15 - 18 January when it was not operating. Sampling occurred at a 15min time resolution, except for a short period at the beginning of the campaign (12 to 26 27 13 January), when it sampled at a 5min time resolution. CRUISER samples air at 4m above 28 ground level through a PM_{2.5} sharp cut cyclone (Rupprecht and Patashnick, East Greenbush, NY, USA) at a flow rate of 16.7 litres per min (lpm), to supply a primary sampling line from 29 30 which several on board PM and gas instruments are connected. The temperature in the 31 sampling enclosure in wintertime is approximately 20°C, although sheath air surrounding the

primary 1.25" diameter stainless steel sampling line maintains it at near ambient temperature. 1 2 The AMS is connected to CRUISER's primary sampling line by 0.8m of conductive tubing, and samples at a flow rate of 1 lpm. Although the aerosol was not dried prior to AMS 3 4 sampling, the relative humidity (RH) remained close to that of the ambient air due to the use 5 of sheath air, and relatively short sampling line. Filtered AMS measurements were performed at several times during the campaign, and were removed from the dataset, resulting in 1745 6 7 observations. The operating principles of the AMS instrument have been documented 8 elsewhere (Jayne et al., 2000; Canagaratna et al., 2007b), although those of the Q-AMS are briefly outlined here. Particles are sampled through the particle inlet and focused into a 9 10 collimated beam using an aerodynamic lens. The stream of particles impact a porous tungsten surface heated to ~600°C, whereupon the non-refractory components of the particles flash 11 vaporize. The resulting gases are ionized by electron impact (EI, 70 eV), and the resulting 12 13 ions are measured using a quadrupole mass spectrometer.

The AMS was calibrated for ionization efficiency by atomizing a NH_4NO_3 solution and then size-selecting 300nm particles using a TSI 3071 electrostatic classifier. A relative ionization efficiency (vs. NO_3^+) of 4.5 for NH_4^+ was required to obtain ion balance for the bulk NH_4NO_3 calibration particles, and this value was applied to the ambient data. Default relative ionization efficiencies were assumed for organics (1.4), chloride (1.3), nitrate (1.1), and sulphate (1.2).

20 The collection efficiency of the AMS is often estimated by comparison of the measured mass 21 with that of a collocated instrument. If collocated external PM₁ mass measurements are 22 unavailable, CE is often assumed by comparison of the combined AMS sub-micron PM mass and BC with an external measure of PM_{2.5}. Middlebrook et al., have also shown that a 23 24 composition dependent CE can be estimated from the bulk aerosol composition (Middlebrook 25 et al., 2012). These two options were investigated to determine whether a CE other than a default of 1 could be applied to the data, and the results of this investigation are presented in 26 27 the Supplement. Ultimately, this analysis did not yield a reliable estimate of CE; as such no 28 correction was applied to these data, and a default, simple integer collection efficiency of 29 unity was assumed for this campaign. This value has been used in other studies (Lanz et al., 2007; Richard et al., 2011; Chirico et al., 2011), and reflects a lower bound for the non-30 31 refractory mass concentration. While an accurate estimate of collection efficiency is required for overall mass determination, it remains a multiplier (either constant, or composition 32

1 dependent) applied to the total mass concentration, and ultimately does not affect the primary

2 study conclusions with respect to identifying and characterizing factors.

The time series of collected AMS mass spectra were separated into chemically resolved mass spectra $(NH_4^+, NO_3^-, SO_4^{2-}, Cl^-, and organics)$ using pre-defined fragmentation patterns (Allan et al., 2004), as implemented in the Q-AMS analysis software Deluxe v1.43 for the IGOR Pro

6 software package (Wavemetrics, Inc.).

7 2.2 Positive matrix factorization

22

8 Positive matrix factorization or PMF, (Paatero and Tapper, 1993, 1994; Paatero, 1997) is a 9 non-negative factor analysis model that can be used to represent a time series of 10 measurements as a linear combination of static factor profiles (ideally corresponding to 11 specific sources and/or processes) and their time-dependent intensities. It is applied to an n x 12 m matrix of data, X, and solves the general receptor equation:

$$x_{ij} = \sum_{k=1}^{r} g_{ik} f_{kj} + e_{ij}, \qquad (1)$$

where n is the number of samples and m is the number of species; x_{ij} is the concentration of the jth species in the ith sample; g_{ik} is the contribution of the kth factor to the ith sample; f_{kj} is the mass fraction of the jth species contributing to the kth factor; e_{ij} is the residual concentration of the jth species in the ith sample; and p is the number of independent factors as chosen by the user. The general receptor equation is solved iteratively, using a Gauss-Newton, weighted least-squares algorithm, until the object function Q, is minimized:

$$Q = \sum_{i=1}^{n} \sum_{j=1}^{m} \left(\frac{e_{ij}}{s_{ij}}\right)^2$$
(2)

20

23

where s_{ij} is an element in the n x m matrix, S, of uncertainties corresponding to x_{ij} . The expected Q value is defined as:

$$\boldsymbol{Q}_{expected} = \boldsymbol{n}\boldsymbol{m} - \boldsymbol{p}(\boldsymbol{n} + \boldsymbol{m}) \tag{3}$$

PMF analysis was performed using EPA PMF 4.1 (Norris et al., 2010). Two different approaches were taken in this study. One approach involved application of PMF to the organic mass spectra only (PMF_{Org MS}), and the other involved application of PMF to the full mass spectra (PMF_{Full MS}). As such, two different methodologies were required for data pre treatment, and formulation of the data and error matrices.

The data and error matrices for the PMF_{Org MS} analysis were prepared following principles 3 4 outlined in Ulbrich et al. (2009). A total of 167 m/z's were included in the PMF_{OrgMS} analysis 5 $(m/z \ 12 - 200)$, with m/zs excluded due to either low signal (e.g., 19-23), signal dominated by 6 inorganic (e.g., 14) or gaseous (e.g., 28) species, or high background levels (e.g., 149) (Ulbrich et al., 2009; Zhang et al., 2005). Uncertainties for the PMF_{Org MS} analysis were 7 8 calculated according to the method of Allan et al. (2003), and a minimum error corresponding 9 to the counting of a single ion was enforced throughout the dataset (Ulbrich et al., 2009). 10 Within the AMS organic fraction extraction process, certain m/z (16, 17, 18, and 44) are 11 assumed to be a constant fraction of m/z 44. The uncertainties of these ions were accordingly 12 multiplied by sqrt(4) to prevent them from being over weighted by the PMF algorithm. The signal to noise ratio (S/N) for each m/z was calculated to identify weak (0.2 < S/N < 2), and 13 bad (S/N < 0.2) variables; a downweighting policy was applied such that weak variables are 14 15 downweighted by a factor of two, and bad variables excluded. No variables were designated 16 as bad for this analysis.

17 The data and error matrices for the PMF_{Full MS} analysis were prepared following principles 18 outlined by Chang et al. (2011). The data matrix was calculated in nitrate equivalent (NO_{3eq}) 19 mass (refer to section 2.3) and calculated by taking the entire raw MS matrix ("All"), and 20 from it, subtracting the mass spectral matrices of species not of interest to the analysis (i.e., air and water), leaving the contributions from fragments associated with NH_4^+ , NO_3^- , SO_4^{2-} , CI^- , 21 and organics. Downweighting of selective organic or inorganic peaks as is required when 22 23 conducting the PMF_{Org MS} analysis was not required in the PMF_{Full MS} analysis, as the PMF_{Full} MS matrix resulted from the subtraction of the "Air" and "Water" components from the 24 original "All" matrix. In other words, the PMF_{Full MS} matrix was not constructed from 25 application of the fragmentation scheme to create "Org", "SO₄²⁻", "NO₃⁻", "Chl", and "NH₄⁺" 26 matrices which could be added together to generate a "Full MS" matrix, but was a result of 27 the subtraction of the "Air" and "Water" components subtracted from the original "All" 28 29 matrix. The corresponding error matrix (in NO_{3eq} mass) was then constructed by adding in quadrature the "All err", "Water err", and "Air err" matrices. Similar to the organic matrix 30 31 preparation, a minimum uncertainty corresponding to a single ion was enforced, and the same 32 S/N policy was applied, although no variables were designated as bad (Ulbrich et al., 2009).

1 A total of 173 m/zs were included in the PMF_{Full MS} analysis, with m/zs excluded due to either 2 low signal (e.g., 19-23), known interferences (e.g., 18), signal dominated by gaseous species (e.g., 28), high background levels (e.g., 149), or non-linear contributions (m/z, 39). 3 While m/z 39 (potassium) could be useful in a PMF analysis of the full mass spectrum for the 4 5 potential identification of certain factors (e.g., biomass burning), it was excluded due to nonlinearities in the signal. Potassium can ionize by two different ionization pathways, namely 6 7 electron impact and surface ionization, each bearing a different relative ionization efficiency 8 (RIE) (Drewnick et al., 2006). Thus, the amount of signal measured from potassium depends 9 not only on the actual initial potassium mass, but also the particle's history within the AMS. 10 Quantifying the relative degree of vaporization via electron impaction vs. surface ionization is 11 difficult, as this ratio is not entirely stable over time (e.g., minor drifts in tuning, fluctuations 12 of the vaporizer temperature), and may depend on the particle composition. Initial PMF tests 13 indicated that potassium inclusion did not aid in the extraction of a biomass burning organic 14 aerosol factor. Although some potassium is also found at m/z 41, this fragment was dominated by organics (potassium contribution <7%). This signal could be removed, 15 16 although doing so requires referring to non-linear m/z 39. Thus, due to the low contribution 17 of potassium at m/z 41, this fragment was left unaltered to avoid introducing additional noise 18 to the matrix.

19 In addition to the uncertainties as described above, a global uncertainty of 5% of the data (the 20 C3 parameter) was added to the uncertainty matrix, in a similar fashion to (Brown et al., 21 2012). Solutions were interpreted based on the resulting factor profiles and temporal trends. 22 The factor profile mass spectra were compared with those extracted from other PMF studies 23 of AMS data, however it should be noted that these comparisons were interpreted with care 24 due to methodological differences between $PMF_{Full\ MS}$ and $PMF_{Org\ MS}$ analyses. Factor 25 temporal trends were examined for particular behaviors such as diurnal trends, and 26 correlations with meteorological conditions and external species (e.g., gases and PM mass).

Two methods were employed to test the robustness of the factor analysis of the AMS data: FPeak rotational analysis and bootstrapping. The effect of global matrix rotations through the FPeak parameter was mainly evaluated in terms of the mass spectra: changes in the mass spectra could be more objectively evaluated than changes in temporal trends due to the availability of comparison mass spectra from other studies, and lack of a priori knowledge on source temporal trends. However, the effect of FPeak rotations on correlations between some

1 factors and key external measurements was also evaluated. It should be noted that since EPA 2 PMF 4.1 utilizes the multilinear engine (ME-2), FPeak values approximately five times 3 greater than those typically used for PMF2 in order to achieve a similar degree of rotation 4 (i.e., in PMF2, FPeaks explored typically range from -2 to 2) (Norris et al., 2010). Similar to 5 the approach used by Brown et al. (2012), FPeak rotations were calculated from -10 to 10 in 6 increments of 2 (Brown et al., 2012). This range led to increases in Q/Q_{exp} of ~1%, indicating 7 that the base solution appeared rotationally robust. Furthermore, this rotational range 8 appeared sufficient to provide indication of the relative robustness of factors, by comparing 9 the relative degree of rotational ambiguity between factors: the robustness of each factor was 10 examined by applying the AMS fragmentation species extraction algorithm (Allan et al., 11 2003) to the resulting FPeak factor profiles, and the species mass fractions across FPeak 12 values was examined. In terms of the bootstrap analysis, 100 bootstrap iterations were 13 performed. Bootstrap results were mainly interpreted according to the number of unmapped 14 factors (factors which could not be "mapped" to the base case using a threshold uncentered 15 correlation coefficient of 0.6). The results of these tests are described in the supplementary 16 material.

17 **2.3** Aerosol mass spectrometer mass quantification

The use of NO_{3eq} mass proved to be a useful method for obtaining better mass estimates, as 18 the various relative ionization efficiencies (RIEs) of component species can be considered in 19 20 the mass quantification of resolved PMF factors. In the case of PMF of organic MS, only a 21 single multiplicative factor of is applied to the dataset as a whole to account for the RIE of organics (RIE of 1.4). However, another approach is required for mass estimates of multi-22 23 component, combined inorganic and organic mass spectra. Firstly, PMF analysis of the full mass spectrum was performed using nitrate equivalent mass (NO_{3ea}), whereby instrument 24 25 signal was converted to mass using a single RIE (in this case, that for nitrate). Following PMF, the factor species composition was determined through application of the fragmentation 26 27 scheme to factor mass spectra (Allan et al., 2003; Allan et al., 2004), and an effective factor 28 RIE calculated through weighted average of RIEs according to the factor composition (Chang 29 et al., 2011). Default RIE values were assumed for the main AMS species, and used to convert the NO_{3eq} factor mass to "real" mass. It should be noted that this method works well 30 when the defining species chemical nature is well understood, and fragmentation patterns and 31 RIE values are available (i.e., as for NO₃⁻, SO₄²⁻, NH₄⁺, chloride, organics). However, the 32

1 AMS has been known to detect other species, such as methanesulphonic acid (Zorn et al., 2 2008; Langley et al., 2010; Chang et al., 2011), and amines (Silva et al., 2008; Docherty et al., 2011; Hildebrandt, et al., 2011), for which less information is available. In particular, it has 3 been shown that depending on their chemical nature, amines may display a wide range of 4 5 fragmentation patterns and RIE values (i.e., from 1.3 to 10) (Silva et al., 2008). Thus, an indication of the chemical nature of the factor species may be integral to the factor mass 6 7 quantification calculation. Further discussion of the implication of these assumptions is 8 provided in section 3.2.

9 2.4 Supporting measurements

10 Trace gases were measured using a variety of techniques, namely by quadrupole Proton 11 Transfer Reaction Mass Spectrometry (PTR-MS) (Ionicon, Innsbruck, Austria) to measure 12 VOC's, as well as other gas analyzers to measure NO_x , SO_2 , O_3 and CO. Particle number concentration measurements were provided by a condensation particle counter (model 3010, 13 14 TSI Inc., Shoreview, MN, USA). Black carbon measurements were also available from an aethalometer (Magee Scientific), and measurements derived from absorption at 880nm were 15 16 used. As reliable collocated sub-micron PM mass concentration measurements were 17 unavailable, PM mass comparisons were made to 5min PM_{2.5} measurements obtained by a 18 TEOM (TEOM model 1400ab, Rupprecht and Patashnick, East Greenbush, NY, USA) 19 onboard CRUISER. Hourly meteorological measurements were supplied courtesy of 20 Environment Canada, from a meteorological station located 10km to the east of MicMac Park 21 in an open field at Windsor Airport (42°16'48"N, 82°57'36"W). Measurements of wind direction and speed, RH, and visibility were used in this analysis. 22

23 **2.5** Assessment of geographic origins

24 The geographic origins of the AMS PMF factors were assessed using the conditional 25 probability function (CPF), and the potential source contribution function (PSCF), which have been described elsewhere (Ashbaugh et al., 1985; McGuire et al., 2011). In this study, the 26 CPF threshold was set to the top 25th percentile, and probabilities associated with infrequently 27 observed wind directions (winds < 5% of the time), were downweighted by 3. For the PSCF 28 analysis, each cell was chosen to be 0.5° in both latitude and longitude, and the threshold for 29 the PSCF was set to the top 50th percentile. For the purposes of the Sulphate-OA factor PSCF 30 analysis, three short events known to be associated with local sources were removed from the 31

analysis (see section 3.1.2 for further description). The result is a probability distribution map
where higher probabilities indicate more probable regional source regions.

3 **3 Results and discussion**

An overview of the meteorological conditions observed during the campaign is presented in 4 5 Figure 2. Unusually warm January temperatures for southwestern Ontario were observed at the beginning of the campaign. Higher wind speeds were associated with southerly air flow. 6 7 Wind speeds dropped dramatically towards the end of the campaign resulting in a stagnation 8 period, resulting in a significant increase in PM mass concentration. The time series of the 9 AMS non-refractory species, as calculated from the AMS data analysis package, is shown in Figure 3, and descriptive statistics for these species are listed in Table 1. On average, prior to 10 PMF analysis, organic aerosol (37%), and NO_3^- (31%) were found to contribute most to the 11 12 non-refractory sub-micron PM mass. The following sections first outline results from PMF analysis of the full mass spectrum, followed by PMF analysis of the organics. Finally, results 13 14 from both analyses are compared.

15 **3.1 PMF of AMS full mass spectra**

16 PMF_{Full MS} analysis showed that six factors best captured the variability in the data. The 17 following factors were retrieved: Amine; Sulphate-OA, containing mostly ammonium 18 sulphate; Nitrate-OA, containing mostly ammonium nitrate; Chloride, composed mostly of ammonium chloride; HOA, a hydrocarbon-like organic factor, which represented primary 19 20 organic aerosol; and OOA, an oxygenated organic aerosol factor. Figures 4 and 5 show the 21 time series (in local time) and factor profiles respectively. Figure 6 shows the mass spectra of 22 each factor's organic components, and Figure 7 details each factor's chemical composition by 23 species and factor component.

Solutions containing two through ten factors were analyzed. In brief, as with the six factor 24 25 solution, the five, seven, and eight factor solutions contained almost the same five factors (Sulphate-OA, Nitrate-OA, Chloride, HOA, Amine). While the five factor solution did not 26 27 extract an OOA factor, the seven factor solution split the OOA resolved from the six factor 28 solution into a similar OOA factor, as well as an Other OA factor which did not sufficiently resemble any known mass spectra. The eight factor solution added a Local Sulphate factor. 29 30 More detailed solution descriptions and justification of the six to eight factor solutions are 31 presented in the supplement. The six factor solution was chosen because: among 100 random runs, all runs converged and displayed a constant, global minimum Q value; higher order solutions did not explain significantly more variance in the data; and factors from this solution were the most physically meaningful. The following sections detail findings for each factor with a focus on new insights into aerosol sources and processes due to the incorporation of both the organic and inorganic aerosol fractions into the PMF analysis.

6 **3.1.1 Amine factor**

7 The Amine factor's time series, shown in Figure 4, was characterized by several short 8 duration events. The Amine factor MS (Figure 5) was distinctly different from the other 9 factors due to the presence of fragments such as m/z 30, 58, 86, and 100. This factor also 10 contained significant signal at m/z 30, with the m/z 30/46 ratio much higher than that for 11 nitrate, suggesting the presence of other ions (e.g., CH_4N^+). The organic functionality of this 12 factor was examined through the delta (Δ) pattern displayed by its mass spectral profile, 13 whereby $\Delta = m/z - 14n + 1$ (where n is the number of CH₂ groups left on the functional group) (McLafferty and Tureček, 1993). Given its characteristic fragments, and strong $\Delta = 3$ pattern 14 15 (i.e., 30, 44, 58, 72, 86, 100, etc.) representative of alkyl amines ($C_nH_{2n+2}N$), this factor was assigned as Amine. The Amine factor was robust in the solution, emerging in each solution 16 17 involving at least three factors. In terms of assessing rotational ambiguity from FPeak analysis, the Amine factor appeared robust, and rotationally fixed (Figure 5). 18

19 Gas and particle phase amines have been recorded in the troposphere for some time, and can 20 be emitted from a variety of sources. The largest global sources of amines are animal 21 husbandry, industrial operations, and waste-water treatment (Ge et al., 2011). Gaseous aliphatic amines at high concentrations can have serious consequences for human health, with 22 effects ranging from irritation of mucosal membranes, to blood clots, and potentially cancer 23 24 (Greim et al., 1998). Particle-phase amines have been measured in widely different settings, 25 ranging from rural areas in Utah (Silva et al., 2008) and Ontario (McGuire et al., 2011; Rehbein et al., 2011), to heavily urbanized areas, such as Mexico city (Aiken et al., 2009), 26 27 Riverside, California (Pratt et al., 2009), and Toronto (Tan et al., 2002; Rehbein et al., 2011). In this study, the measurement site was located in an urban industrial setting, with known 28 29 amine sources located nearby: a waste-water reclamation plant, and a major amine chemical manufacturer were located 1 and 13km to the southwest, respectively. According to the TRI 30 31 and NPRI inventories, the amine manufacturer was the largest monitored emitter of TEA in 2005 in the Windsor/Detroit region (Environment Canada, 2013; US EPA, 2013). The strong 32

southwest directionality observed in the CPF (Figure 8a) indicated that both of these sources
 may have contributed. The sharp bursts in temporality were consistent with local sources,
 such as fugitive emissions from industrial operations.

4 In order to quantify the Amine factor's mass concentration, it was necessary to obtain an 5 estimate of the factor's effective RIE. Unfortunately, an effective RIE could not be 6 determined through application of the traditional AMS fragmentation table, as the factor It is possible to alter the 7 contained amines that were not represented in the table. 8 fragmentation table to include additional species, provided the nature of the measured species 9 is known, and a species fragmentation pattern is available. This has previously been 10 accomplished, for instance, with methanesulphonic acid (MSA) (Zorn et al., 2008; Langley et 11 al., 2010; Chang et al., 2011). In a study by Chang et al., MSA could be positively identified due to unique marker fragments and a lack of interfering species, and its mass could be 12 13 calculated through application of a laboratory determined fragmentation pattern and RIE (Chang et al., 2011). Taking this approach was not obvious for the present study, as the 14 15 particular amine compound(s) could not be easily identified, and there was a possibility that 16 the factor represented a linear combination of different amine species with different RIEs and 17 fragmentation patterns.

18 Nonetheless, the nature of the Amine factor was investigated to determine a potential factor 19 RIE for mass estimation. Amines have been shown to exhibit a wide range of RIEs, 20 depending on their chemical nature. AMS measurements of amines present in salt form, such as methylammonium chloride, dimethylammonium chloride, and trimethylammonium 21 22 chloride, have shown that the amine fraction in these compounds can display RIEs from 5 to 23 10 (Silva et al., 2008). However, oxidized alkyl amines such as trimethylamine-n-oxide 24 (TMAO) have been shown to ionize with an RIE of 1.3, a value closer to organics (RIE = 1.4) 25 (Silva et al., 2008). It has been hypothesized that aminium salts exhibit a relatively high RIE due to surface ionization on the AMS vaporizer, similar to that observed for potassium salts 26 (Silva et al., 2008). Thus, depending on the type of amine compound or mixture of 27 compounds the Amine factor represents, its RIE may lie within a wide range (i.e., 1.3 - 10). 28

A reasonable estimate for an effective RIE for the Amine factor was sought by examining the data for a dominant particle phase amine formation pathway, namely for signs of aqueous dissolution, acid-base reaction, or oxidation (Ge et al., 2011). First, the Amine factor time series was examined relative to external measurements. Dissolution into water droplets was

investigated by comparing the time series with periods of rain, fog, or high RH: no 1 2 association could be identified as the Amine factor often appeared on clear days with lower RH. Reaction with acidic species was also considered through time series analysis of the 3 extent of neutralization, a useful metric for determining periods of particle acidity. However, 4 5 this metric cannot provide reliable information, as NO_3^- cannot be properly quantified prior to PMF analysis due to amine interferences. The temporality of the Amine factor was also 6 7 investigated because the daytime maximum identified for a similar factor identified by Sun et 8 al. suggested that photo-oxidation and condensation can also be a likely formation pathway 9 (Sun et al., 2012). However, no consistent diurnal trend was noted. Docherty et al. reported similar difficulty in determining the origins of an amine related factor through time series 10 11 analysis of results from a PMF_{Org MS} analysis (Docherty et al., 2011).

12 Mass spectral comparison to laboratory generated amine MS provided better indication as to 13 the chemical nature of the Amine factor. Among comparisons with mass spectra from 14 suspected amine compounds reported in the NIST library, the Amine factor's profile was most similar to that of triethylamine (TEA: C₆H₁₅N, 101 g mol⁻¹), as demonstrated in Figure S 15 2.1 ($r^2 = 0.23$) (Stein, 2013). However, direct comparisons between AMS and NIST mass 16 spectra are interpreted with caution, as they use different ionization techniques which can lead 17 to mass spectral differences (Canagaratna et al., 2007a). Nonetheless, the amine spectra 18 19 showed the same characteristic peaks (i.e., m/zs 56, 58, and 86).

20 Examination of AMS mass spectra of amines provided further perspective. Amines have been studied by AMS in chamber experiments to examine potential reaction pathways, for 21 22 example oxidation, such as by nitrate radical (Murphy et al., 2007; Silva et al., 2008; Malloy 23 et al., 2009), and reaction with acid gases such as HNO₃ (Murphy et al., 2007; Silva et al., 24 2008). These different mechanisms can actually lead to similar mass spectra (Malloy et al., 25 2009). The MS of the Amine factor in this study was determined to be very similar to that of TEAN reported by Murphy et al. (2007), resultant from reaction between TEA and HNO₃, 26 27 with signal at m/zs 30, 46, 58, 86 and 100 (Murphy et al., 2007). One sign of reaction formation of amine salts from reaction with HNO₃, as reported by Malloy et al., is the 28 presence of significant signal from NO⁺ and NO₂⁺ (at m/z 30 and 46) (Malloy et al., 2009). 29 Examination of the MS from the Amine factor showed that signal was very high at m/z 30, 30 and some signal was also present at m/z 46, although as mentioned previously, m/z 30 can also 31 32 represent CH_4N^+ , and CH_2O^+ . There were no mass spectral signs to suggest an oxidation 1 mechanism over salt formation. Ultimately, the factor was interpreted as being dominated by

2 TEAN.

3 With this interpretation, an effective RIE for the factor was calculated. This was achieved by 4 assuming a neutralized factor, and that TEAN was the only component. Though there did appear to be other contributions (e.g., SO_4^{2-}), these appeared very low. The nitrate fraction 5 was calculated using the nitrate fragmentation pattern, with m/z 30 altered to reflect the 6 7 isotopic ratio between m/z 30 and 46, obtained from calibration. The RIE of the amine 8 fraction was determined by assuming factor neutrality, and that the remaining mass following 9 subtraction of nitrate was triethylammonium. An RIE for the amine fraction of 6.0 was 10 determined from this calculation, which fell within the range of RIEs previously measured for 11 various amine salts. With the nitrate fraction taken into consideration, an effective RIE of 4.3 12 This amine salt interpretation appeared to provide reasonable mass was established. 13 concentrations, as the calculated RIE resulted in spikes (<2hrs) reaching a magnitude of 4.8µg m⁻³, while an RIE of 1.3 reflective of oxidized amines, resulted in concentrations exceeding 14 15µg m⁻³. While the factor was assumed to be dominated by TEAN, its exact nature could 15 16 not be validated; no amine fragmentation pattern proved an exact match, and external high 17 time resolution collocated sub-micron PM mass measurements were not available to validate the RIE through PM mass comparison. In considering acid-base chemistry, calculations by 18 19 Murphy et al. have showed that the formation of aminium salts from the reaction of HNO_3 20 and TEA is only thermodynamically favorable in the presence of very low NH₃ (Murphy et 21 al., 2007). These conditions may have been provided by plumes from a nearby source.

22 Four studies to date have identified an amine related factor through PMF of AMS mass spectra to the authors' knowledge (Aiken et al., 2009; Docherty et al., 2011; Hildebrandt et 23 24 al., 2011; Sun et al., 2012). The analyses by Aiken et al., Docherty et al., and Hildebrandt et 25 al. all extracted the amine factors from PMF of the organic MS, while Sun et al. extracted it from the full MS. Since the former three applied PMF to only the organic MS, significant 26 27 mass that may have been associated with this factor (i.e., m/z 30) was potentially not captured. 28 Furthermore, the latter study did not take into account the potential for a wide range of RIEs 29 for the amine-related species, as highlighted by Silva et al. (2008), and discussed in this study. Regardless of the exact methodology, it can be seen that PMF can be effective in resolving 30 31 atypical factors, such as amines.

1 **3.1.2 Sulphate-OA factor**

2 The time series and mass spectral profile of the Sulphate-OA factor are shown in Figures 4 3 and 5 respectively, and the chemical composition breakdown for this factor is shown in 4 Figure 7. A mass spectral comparison between the Sulphate-OA factor and the published MS of atomized (NH₄)₂SO₄ (Hogrefe et al., 2004), shows that they exhibit the same major peaks 5 at m/z 16, 17, 48, 64, 80, and 81, and compare well with an r^2 of 0.74. The Sulphate-OA 6 factor on average contributed 1.81µg m⁻³, or 25% to the sub-micron PM mass, and showed the 7 highest mass concentrations towards the beginning of the campaign when air masses 8 9 originated from the south. Overall, this factor showed a slight correlation with PM_{2.5} mass concentration ($r^2 = 0.21$). A moderate correlation with SO₂ ($r^2 = 0.31$) implied that this factor 10 11 was likely not only influenced by long range transport, but also by local sources. A more 12 local influence was determined through examination of these temporal trends, which showed 13 that several spikes in the Sulphate-OA factor coincided with large spikes (up to 58ppb) of 14 SO_2 . Local (within the metropolitan area) and local-to-regional (within ~ 100km) geographic 15 origins for the Sulphate-OA factor were investigated by CPF. This highlighted a strong association with emissions from the southwest (Figure 8b), which was consistent with some 16 local and local-to-regional coal fired power plants (Figure 1). Three large SO_4^{2-} spikes were 17 observed on January 19th, 24th, and 27th: the first two were associated with the west-18 19 southwest, and the last with southeast, and all of which were likely associated with local coal-20 fired power plants. A smaller, yet still significant influence from the northeast was also 21 observed, which may have been associated with emissions from coal plants located to the 22 northeast. The aforementioned spikes were all associated with large excursions in SO_2 (24, 34, and 58ppb respectively), and each lasted about 2-6 hours. While the CPF on the whole 23 demonstrated strong directionality to the southwest, this method cannot resolve how far away 24 25 the responsible source(s) actually are located. Since southwesterly winds in Windsor are also 26 consistent with typical synoptic flows for this region even in winter, and there are known large SO₂ emissions sources located further away in the Ohio River Valley, regional 27 influences for the Sulphate-OA factor were also investigated using the PSCF (Figure 9a). The 28 PSCF calculations for all factors were performed using data greater than the 50th percentile. 29 but for the Sulphate-OA factor, the three spikes associated with more local emissions were 30 31 removed. The PSCF highlighted high probability source regions around the Ohio River valley, an area known as a major SO₂ source due to the presence of many large coal fired 32 33 power plants. A dominant regional influence was also demonstrated by a reasonably constant diurnal profile (Figure 11a). The geographic origins of the factor were also consistent with
 those of a Sulphate factor derived from a long-term receptor modeling study of Windsor
 (Jeong et al., 2011), and the nearby rural location of Harrow, Ontario (McGuire et al., 2011).

As the NH₄⁺ and SO₄²⁻ contained within the Sulphate-OA factor likely formed an (NH₄)₂SO₄ salt, the extent of neutralization was calculated (Table 2). It was assumed that the only species capable of participating in the neutralization reaction, in this factor and in others, were NH₄⁺, NO₃⁻, SO₄²⁻, and Cl⁻. Aminium species were assumed to have been effectively separated into the Amine factor through PMF and thus it was assumed they did not contribute to the neutralization of other factors. The extent of neutralization (Neut_{Ext}), reported as the ratio of cations/anions in units of molar equivalents, was defined by:

$$Neut_{Ext} = \left(\frac{\frac{NH_4^{+}}{18}}{\frac{2SO_4^{2-}}{96} + \frac{NO_3^{-}}{62} + \frac{Chl}{35.5}}\right),$$
(4)

11

12 where a neutral factor shows a Neut_{Ext} of 1, and acidic factors show values less than 1.

The extent of neutralization of the Sulphate-OA factor was 0.99, indicating that the factor was 13 14 reasonably neutral. It should be noted that a source of uncertainty in this value lies in the use 15 of default RIE values for most species. Despite this potential uncertainty, the factor appeared 16 neutral, similar to several other factors. Rotational analysis showed that the composition of this factor, and thus the degree of neutralization, did not change with FPeak rotations (Figure 17 S-2.5). This suggested that regional rather than local sources of SO_4^{2-} had a greater influence 18 on the chemical composition of this factor. While the Sulphate-OA factor appeared 19 20 neutralized, simultaneous SO_2 and SO_4^{2-} spikes suggested contributions from more local, possibly primary SO_4^{2-} emissions sources, which may have been more acidic. As highlighted 21 in Figure 10a, these SO_4^{2-} events, while high in mass, did not account for a significant fraction 22 of the total sulphur due to the magnitude of the coincident SO₂ spikes (up to 65ppb, 1 min 23 24 average). This could be attributed to the observed winter conditions that do not favor rapid oxidation of SO₂. Interestingly, a smaller SO_4^{2-} size distribution was observed during these 25 spikes (from AMS p-ToF data), substantiating local SO_4^{2-} contributions. Figure 10b shows 26 that SO422- measured over the entire campaign showed an average modal size of 500nm 27 (accumulation mode consistent with regional transport), while the size modes observed during 28

1 the spikes were much smaller (150 – 250nm). These results indicate that proximate sources 2 of SO_4^{2-} contributed to the total SO_4^{2-} , particularly during the largest SO_4^{2-} spikes.

Examination of the residuals from the six factor solution shows that some AMS signal cannot 3 be accounted for during the Sulphate-OA factor/ SO_4^{2-} spikes, indicating that another factor 4 may be required to more fully explain the mass during these spikes. As a result, higher order 5 6 solutions were investigated, and are reported in the supplementary material. Figure S-2.12 7 shows that at eight factors, the Sulphate-OA factor split into two factors: a Regional Sulphate 8 factor that was characterized by synoptic-scale temporal rises, and a Local Sulphate factor that captured the SO_4^{2-} spikes. While the Local Sulphate factor appeared meaningful in that it 9 captured residual SO_4^{2-} , the eight factor solution could not be justified for several reasons, 10 11 which are further detailed in the supplement. First, while the Local Sulphate factor appeared acidic (Neut_{Ext} = 0.25), consistent with more local origins, the Regional Sulphate factor in the 12 13 eight factor solution now appeared over neutralized (Neut_{Ext} = 1.27), relative to a reasonably 14 unchanged, and neutral Nitrate-OA factor. Second, the OOA factor appeared split, resulting 15 in an OA factor which did not show a strong enough resemblance to any known, ambient deconvolved OA. Finally, the correlation between the HOA factor MS from the eight factor 16 17 solution and the reference HOA was worse than in lower order solutions. Mass spectral examination showed that more signal was apportioned to m/z 16 and 17 in higher order 18 solutions; this effect was mainly attributed to solution uncertainty, as FPeak analysis of the 19 20 six factor solution showed some variability in these fragments upon rotation. Although the 21 Sulphate-OA factor from the six factor solution did not fully capture the variability and types of SO_4^{2-} observed, it appeared stable, with a low degree of rotational uncertainty (Figure S-22 23 2.5). It is possible that the Local Sulphate factor could be extracted more definitively if it were more prominent in the dataset. However, resolving acidic factors may not always be 24 25 realistic, as data quality, receptor site complexity and atmospheric dynamics can all influence 26 factor resolution. Nonetheless, resolving acidic factors could be useful from a 27 parameterization perspective for resolving the effects of acidic aerosols on health, impacts on 28 materials, or acidic seed particle chemistry, such as the competition between NH₃ and organic gas uptake to acidic SO_4^{2-} containing particles (Liggio et al., 2011; Liggio and Li, 2013). 29

The organic fraction of the Sulphate-OA factor was extracted for comparison with the organic fraction of other factors, as well as published organic factor MS (Figure 6). Uncertainty in the organic fraction was assessed through rotational analysis, and was found to be low suggesting

that these organics were reasonably rotationally fixed. The organic composition amongst 1 2 factors extracted using the PMF_{Full MS} analysis was assessed against that from other studies. Only one of the past three similar studies, that of Sun et al., (2012), was sufficiently similar 3 for comparison; Chang et al., involved Arctic aerosol, and the analysis by Crippa et al., only 4 involved the organics and SO_4^{2-} . The Sulphate-OA factor contained 21% organics by mass, 5 6 accounting for 16% of the total measured organics; interestingly this value was the same as 7 that produced in the analysis by Sun et al., (2012). Although both of these studies were 8 conducted in highly urban environments, such an agreement is interesting, considering that 9 different organic factors were found. Compared to other factors from this study with 10 significant organic content (i.e., > 15% by mass), the Sulphate-OA factor organics were most similar to typical OOA mass spectra, and were the most highly oxidized according to the 11 associated F44 and O/C ratio. The F44 metric (fractional contribution of m/z 44 within the 12 13 organic MS) has been used in PMF_{Org MS} studies to assess the degree of oxygenation of OOA factors, as it primarily represents CO_2^+ , the most prevalent oxygenated fragment within the 14 organic MS. An empirical relation between F44 and the atomic oxygen to carbon ratio (O/C) 15 16 within a PMF OOA factor has been developed based on a collection of laboratory and field 17 study data (based on PMF_{Org MS} analyses) (Aiken et al., 2008), and has been used to estimate the O/C ratio for similar factors in other studies, including previous PMF studies of combined 18 19 organic and inorganic mass spectra (Chang et al., 2011;Sun et al., 2012;Crippa et al., 2013b). 20 An F44 of 0.15 (estimated O/C = 0.65) was calculated for this factor, the highest value from 21 this study and consistent with the average values obtained for OOA for two-component OA 22 datasets (i.e. HOA and OOA) across the Northern Hemisphere (F44 = 0.14 ± 0.04 , O/C = 0.6223 ± 0.15 , mean $\pm 1\sigma$) (Ng et al., 2010).

24 Many previous AMS PMF studies have shown that a high degree of correlation exists between SO₄²⁻ and OOA (Lanz et al., 2008; Ulbrich et al., 2009; Slowik et al., 2010; Richard 25 26 et al., 2011). However, this analysis has been useful for sub-apportionment of the oxygenated 27 organic fraction of OA to different factors, as the Sulphate-OA factor was found to contain a 28 notable fraction (31%) of the total OOA (defined here as the proportion of m/z 44 signal 29 apportioned to this factor, excluding the Amine factor), and in this case, the most oxidized 30 fraction of the OOA. The higher degree of oxygenation exhibited by this factor is consistent with its regional origins from the south, and a longer atmospheric lifetime. Although the 31 32 mixing state of particles associated with this factor cannot be positively deduced from these 33 data, two extremes exist. Either this factor mainly represents externally mixed OOA and

20

(NH₄)₂SO₄ particles that exhibit similar temporality due to regional transport, or it represents
 internally mixed (NH₄)₂SO₄ particles coated by lower volatility SOA during regional transit.

In summary, the Sulphate-OA factor was neutralized and associated with the most highly oxygenated organics. Cross-apportionment of the most oxygenated organics measured during this campaign further reinforced this factor's aged, regional nature. While the Sulphate-OA factor appears to be dominated by regional transport of neutralized ammonium sulphate, minor contributions from more local, primary SO₄²⁻ sources were present as well.

8 3.1.3 Nitrate-OA factor

9 The Nitrate-OA factor time series and mass spectrum are shown in Figures 4 and 5 10 respectively, and its contribution to total mass is shown in Figure 7. Of all the factors, it 11 contributed most to the sub-micron PM mass, with an average mass concentration of $3.19\mu g$ 12 m⁻³ (45%). Significant accumulation was observed towards the end of the campaign, when a 13 severe nitrate episode occurred, and this factor's mass alone exceeded $20\mu g$ m⁻³.

Deconvolution of the Nitrate-OA factor into its component species (Table 2) shows that it was 14 15 mostly composed of NH_4^+ and NO_3^- (78% by mass). However, other species such as organics and SO_4^{2-} also comprise a significant mass fraction (16% and 6% respectively). As with the 16 17 Sulphate-OA factor, the Nitrate-OA factor demonstrated identical organic composition to the NO₃-OA factor found in the study by Sun et al., (2012). That both of these largely inorganic 18 19 factors were highly similar to another study conducted in New York City is a possible 20 indication of consistent internal and/or external mixing between these combined inorganic 21 and organic species in these types of inorganic factors. The dominance of ammonium nitrate in this factor's mass spectrum was confirmed by a high correlation with published NH₄NO₃ 22 standard spectra ($r^2 = 0.96$) with peaks at m/z 17, 18, 30, and 46 (Hogrefe et al., 2004). 23 24 Comparison with NH₄NO₃ calibration mass spectra showed that it also compared well, 25 although the factor's m/z 30 to 46 ratio was 19% higher than the calibration mass spectrum, 26 which was likely due to the presence of organics. The Nitrate-OA factor contained 20% of 27 the total measured organics during the campaign, or 31% of the oxidized organics (as defined 28 by the fraction of m/z 44 apportioned to this factor). The nitrate-bound organics displayed the second highest F44 (0.12) of all the factors from the PMF_{Full MS} analysis. The F44 of 0.12 29 30 corresponded with an estimated O/C = 0.54, which when compared to OA components from PMF_{Org MS} studies, falls between typical values for OOAI and OOAII indicating moderately 31 oxidized organics (Ng et al., 2010). Several previous PMFOrg MS studies have found an 32

association between NO_3^- and OOAII, suggestive of temperature dependent volatility (Lanz et 1 2 al., 2007; Ulbrich et al., 2009). The diurnal trend for the Nitrate-OA factor partly suggests semi-volatile behavior, due to a slight inverse relationship with temperature (Figures 11b and 3 e), however slightly higher nighttime values may also be associated with decreased mixing. 4 5 As a stronger inverse relationship was noted for the OOA factor than Nitrate-OA, it was likely that the Nitrate-OA factor was likely more regionally influenced. This was reinforced by an 6 7 m/z 44 to 43 ratio of 1.5, which was lower than that for the Sulphate-OA factor, yet still 8 higher than that expected for OOAII. The extent of neutralization shows that the ratio of 9 cations to anions was near unity (1.04), suggestive of a reasonably neutral factor. FPeak rotational analysis demonstrates that the Nitrate-OA factor appears rotationally fixed (Figure 10 11 S-2.5).

12 Sources of NH_4NO_3 precursor gases (i.e., NH_3 and NO_3) in the Windsor region are abundant: 13 constant vehicle traffic, in large part from the nearby border crossing, provides a constant supply of NO_x (38.0 \pm 35.0 ppb), as well as minor contributions of NH₃ (Li et al., 2006; Godri 14 15 et al., 2009). While the Windsor/Detroit area may not comprise as significant a source region for NH₃ as rural agricultural areas, many point industrial (e.g., wastewater reclamation plant) 16 17 and diffuse (e.g., traffic, population) sources do emit large quantities of NH₃. According to Environment Canada's NAPS monitoring network, NH₃ levels are lowest in wintertime in 18 19 Windsor, yet are still detectable: using the multi-year average NH₃ wintertime concentration 20 measured by ion chromatography, the average wintertime NH_3 concentration in downtown 21 Windsor between 2004 and 2007 was 0.5 ± 0.6 ppb, which exceeds the 0.1 ppb detection limit for this analytical method. Figure 11b presents the diurnal trend for the Nitrate-OA factor. 22 23 Noted along with a slight diurnal profile that was likely influenced by both temperature (Figure 11b), as well as reduced mixing overnight and into the early morning, were two small 24 25 peaks consistent with morning and afternoon rush-hour traffic, at 09:00 and 17:00. These 26 peaks were likely due to NH₄NO₃ formation from traffic emissions. A more general traffic contribution to the Nitrate-OA factor was observed by moderate and slight correlations with 27 black carbon (BC) ($r^2 = 0.31$), and NO ($r^2 = 0.17$) respectively, both traffic tracers. 28

The geographic origins of the Nitrate-OA factor were explored. The CPF, presented in Figure 8c, did not reveal any significant directionality; this indicated that this factor was not significantly influenced by any local point sources, but did not discount potential contributions from local diffuse sources, such as traffic. The regional nature of the factor was

explored through the PSCF (Figure 9b), which highlighted areas of high source probability to 1 2 the south and southwest. In comparison with the PSCF from the Sulphate-OA factor, the Nitrate-OA factor also appeared more regionally influenced than the Sulphate-OA factor, yet 3 4 from less specific source regions. Regional influences are more likely in winter compared to 5 summer in this location, as nitrate can persist and be transported greater distances provided cold enough temperatures. In summary, the Nitrate-OA factor contributed most to sub-6 7 micron PM mass, was associated with moderately oxidized organics, and was likely attributed 8 to both local and regional sources, with regional contributions likely dominating.

9 **3.1.4 Chloride factor**

10 Presented in Figures 4 and 5 respectively are the time series and mass spectral profile of the 11 Chloride factor. The chemical composition breakdown for this factor is shown in Figure 7. Strong peaks observed at m/zs 16 and 17 indicate significant NH₄⁺ content, while peaks at 12 m/zs 35, 36, 37 and 38 indicate chloride. The factor appeared to be dominated by NH₄Cl, as 13 it compared well with the NIST mass spectrum for this species ($r^2 = 0.71$). The isotopic ratio 14 of m/z 35 to 37 matched that from the AMS chloride fragmentation pattern (3.09). 15 Application of the AMS fragmentation table to the MS demonstrates that NH_4^+ and Cl^- 16 17 contributed 28% and 27% respectively to the mass of this factor, however, other species were also evident, namely NO₃⁻ (13%), organics (13%) and SO₄²⁻ (18%). This study represents the 18 first PMF analysis of AMS mass spectral data to apportion NH_4^+ and Cl^- to a unique factor 19 20 suggestive of mostly ammonium chloride. The extent of neutralization (1.09) shows that this 21 factor was fairly close to neutral, taking into account all identifiable ions (i.e., NH₄⁺, Cl⁻, NO₃⁻ , and SO_4^{2-}), and considering potential PMF error and uncertainty in RIEs. 22

23 Interestingly, the factor was composed 13% by mass of organics (Figure 6). The most 24 significant organic peaks were m/z 29, 41, and 55, and mass spectral delta pattern, with a $\Delta =$ 25 0 (i.e., m/z 27, 41, 55, 69, etc.) was noted (McLafferty and Tureček, 1993). However, despite a distinct delta pattern, the factor showed rotational uncertainty that precluded over-26 interpretation of the factor's organics. With negative rotations, the Cl^{-} and NH_4^{+} content 27 28 varied only slightly. However, rotation in both the positive and negative direction caused 29 significant changes to the organic fraction delta pattern: negative rotations caused the $\Delta = 0$ pattern to decrease and the $\Delta = 2$ (m/z 29, 43, 57, 71, etc.) to increase, while there was 30 31 insignificant change in the organic mass spectral profile with positive rotations. As the 32 organic profile and content varied most significantly among all of the Chloride factor

1 components (from 18% to 40% by mass), the delta pattern could not be used with confidence 2 for factor identification. It should be noted that little uncertainty was observed for the $SO_4^{2^-}$ 3 mass fraction of the Chloride factor, indicating that the $SO_4^{2^-}$ cross-apportionment was 4 relatively robust. Although the Chloride factor showed the most uncertainty among all PMF 5 factors (Figure S-2.5), its appearance from five to eight factors suggested it was robust.

6 The presence of NH_4Cl aerosol in the atmosphere has previously been noted. Pio et al. 7 performed early experiments to assess thermodynamic behavior of NH₄Cl under tropospheric 8 conditions and found that like NH₄NO₃, NH₄Cl formation is favored by lower temperatures 9 and higher RH (Pio and Harrison, 1987a, b). The conditions associated with the Chloride factor were examined (i.e., temperature, RH, cloudiness, fog, rain), to determine if rapid 10 11 meteorological changes triggered gas to particle transition. No association could be discerned. However, it should be noted that from January 14th onward, on the whole, the conditions for 12 NH₄Cl were favorable with a mean temperature of $-8 \pm 5^{\circ}$ C, and RH of $73 \pm 13^{\circ}$. 13

14 Although not directly measured, particulate NH₄Cl has been speculated to be present in 15 several settings due to high correlation between its component species. Perron et al. have 16 reported NH₄Cl-containing particles in a Swiss alpine valley setting due to significant correlation in NH₄⁺ and Cl⁻ spikes, with Cl⁻ likely emitted as KCl during biomass combustion 17 (Perron et al., 2010). In a study in Mexico City, Salcedo et al., reported large, yet short-lived 18 plumes of chloride, which were coincident with spikes in both NH_4^+ and organics; the 19 simultaneous appearance of both NH₄⁺ and Cl⁻ indicated NH₄Cl formation (Salcedo et al., 20 21 2006). Precursor gases (e.g., NH₃, HCl, and HNO₃) and their sources are plentiful in this 22 region. Although NH₃ is typically the limiting agent in these reactions, in this region, point 23 industrial (e.g., wastewater reclamation plant), and diffuse sources (e.g., traffic, population) of 24 NH₃ are numerous. There are thus two plausible mechanisms for NH₄Cl formation in the Windsor/Detroit region. In the first case, NH₃ could react with primary HCl which has been 25 26 emitted directly by anthropogenic sources. According to NPRI and TRI emissions 27 inventories, the largest primary emitter of HCl in Windsor/Detroit in 2005 was the River Rouge Power Plant, located 3km southwest of the receptor site (Environment Canada, 2013; 28 US EPA, 2013). However, secondary reactions may lead to HCl production, and eventually 29 NH₄Cl formation as well. It is possible that acidic, ammonia poor, industrial plumes with 30 31 sufficient HNO₃ content may liberate HCl from certain sources, which then subsequently repartitions as NH₄Cl to the particle phase under NH₃ rich conditions. Some of these sources 32

include NaCl particles from road salt and KCl from biomass burning emissions (including 1 2 coal combustion and domestic wood burning). Waste incineration has also been identified as a potential source of short duration, strong chloride events; Moffet et al. reported Cl⁻ spikes 3 alongside with spikes in Zn and Pb-containing particles measured by ATOFMS in Mexico 4 5 City (Moffet et al., 2008a; Moffet et al., 2008b). These events were attributed to rapid secondary HCl formation from the reaction of HNO₃ with PbCl and ZnCl in a waste 6 7 incineration plume. All of these sources were likely in this study: as mentioned previously, 8 there are several coal plants near the site, and a waste incinerator is located to the northwest. 9 Furthermore, given the winter-time conditions, domestic wood burning was prevalent in the 10 region.

11 Examination of the CPF (Figure 8d) and PSCF (not shown) did not provide any additional information with respect to local or regional source origins. However, concurrent appearance 12 of sharp spikes in Cl^{-} and SO_4^{2-} (Figure 3), and moreover the apportionment of significant 13 SO_4^{2-} to this factor suggested an association with industrial emissions. Both of these spikes 14 corresponded with a 230-250° wind direction, consistent with coal plants to the west-15 southwest. Despite these spikes having been partially captured by their respective, reasonably 16 neutralized PMF factors, the Cl⁻ and SO_4^{2-} observed in these spikes was likely not well 17 18 neutralized, as concurrent spikes in NH_4^+ at these times were not as strong as those observed 19 for its counter ions. Any available NH₃ at this time was likely partitioned as (NH₄)₂SO₄ first, 20 because $(NH_4)_2SO_4$ formation should be thermodynamically more favorable than NH_4Cl 21 formation under the given meteorological conditions. As such, while the Chloride factor 22 appeared reasonably close to neutral on the whole, there was likely some variation in acidity 23 that was not captured by this factor, as more acidic particles appeared to be present during the spikes. Furthermore, the rotational uncertainty determined through FPeak analysis may have 24 25 been an indication of the more dynamic nature of this factor's components. In summary, the 26 Chloride factor was most likely associated with local atmospheric processing, whereby 27 ammonium chloride secondary aerosol formed rapidly in-plume from the reaction of NH₃ and 28 HCl emissions. The source for HCl emissions may have been primary emissions from an 29 industrial facility, and/or it may have formed from secondary reactions.

30 3.1.5 HOA factor

A hydrocarbon-like organic factor (HOA) was extracted by PMF and its time series and mass
 spectral profile are presented in Figures 4 and 5 respectively. The factor's chemical

1 composition is shown in Figure 7. Accounting for 0.94 μ g m⁻³ or 13% of the sub-micron PM 2 mass during the campaign, the HOA factor represented the third largest contributor to sub-3 micron PM mass. Delta patterns of Δ =0 and Δ =2 were observed in the mass spectral profile, 4 most likely corresponding to alkanes and H₂-neutral losses from alkyl fragments and/or 5 alkenes, respectively (McLafferty and Tureček, 1993;Zhang et al., 2005). The extracted HOA 6 factor was similar to those from the Pittsburgh fall 2002 (Ulbrich et al., 2009), and the Zurich 7 winter 2006 (Lanz et al., 2008) studies, with r^2 values of 0.56 and 0.66 respectively.

8 Deconvolution of the HOA MS into organic and inorganic species was performed using the 9 AMS fragmentation table. Contributions from NH_4^+ (5%) and NO_3^- (9%) were noted, indicating that this factor was likely mixed with NH₄NO₃. FPeak rotational analysis shows 10 11 that although contributions of these species were moderate, their variation with rotation precluded their over-interpretation (Figure S-2.5). A less than plausible extent of 12 13 neutralization of 1.62 also served to highlight the rotational ambiguity associated with the 14 inorganics, although this uncertainty could be a result of low overall inorganic contributions. 15 With an organic composition of 85%, it was not as similar to the HOA factor found in Sun et al. (97%), as it was for other factors. 16

17 HOA is often referred to as a surrogate for fossil fuel combustion emissions, with traffic cited 18 as the most significant source contributor (Aiken et al., 2008). The same can be said for this 19 study; moderate r values of 0.48, and 0.64 were found between the HOA factor with BC and NO respectively, pollutants which are typically associated with relatively fresh traffic 20 emissions. The HOA factor also displayed moderate and slight correlations with CO (r^2 = 21 0.41), particle number concentration ($r^2 = 0.21$), and aromatic hydrocarbons measured by 22 PTR-MS, such as benzene (m/z, 79, $r^2 = 0.43$), toluene (m/z 93, $r^2 = 0.32$), C₂ - benzene (m/z23 107, $r^2 = 0.58$), and C₃ - benzene (m/z 121, $r^2 = 0.54$), which are also typically associated with 24 primary anthropogenic emissions, such as traffic (Vlasenko et al., 2009). Distinct rush hour 25 26 peaks in the morning (05:00 - 09:00) and evening (18:00 - 22:00) further supported source attribution to traffic emissions (Figure 11c). The CPF for the HOA factor provides modest 27 support for a traffic source, as excursions in this factor were moderately associated with the 28 north-east wind sector (0 - 90°), consistent with higher traffic density from the international 29 border crossing (Figure 8e). While general circulation traffic is also likely to be a significant 30 31 contributor to the HOA factor, many industrial operations in the region combust fossil fuels in their processes, for instance heavy machinery used for onsite material transport. As such, 32

1 other sources may also contribute to this factor. Thus, the HOA factor represented primary

2 organic aerosol emissions, with its most significant contributions likely emitted from traffic.

3 **3.1.6 OOA factor**

A factor consistent with oxygenated organic aerosol (OOA) was extracted using PMF_{Full MS}, 4 5 and its time series and MS are presented in Figures 4 and 5 respectively. On average, the OOA factor comprised 0.76 μ g m⁻³ or 11% of the sub-micron PM mass. The fragmentation 6 table was applied to this factor to decouple the inorganic and organic species for species-7 8 based mass spectral and temporal examination. The only other species associated with this 9 factor was NO₃⁻ (7% by mass). While the OOA factor represented the lowest F44 compared 10 to Sulphate-OA and Nitrate-OA (F44 = 0.09), it contained the highest fraction of oxygenated 11 organic aerosol, as 35% of m/z 44 was apportioned to this factor. The robustness of the factor mass spectrum was examined using FPeak analysis. It was found that rotations caused the 12 13 minor NO₃⁻ concentration to vary (Figure S-2.5). With a lack of neutralizing species (i.e., NH_4^+) present, the nitrate contribution may have been attributed to model error. In examining 14 15 the robustness of the F44 content through FPeak analysis (Table S2.1), it varied only slightly, 16 suggesting minor rotational ambiguity with respect to the oxygenated organic fragments. In 17 terms of organic composition, this factor's organic composition (93%) was most comparable to the LV-OOA factor from Sun et al., (2012), which comprised 94% organics. 18

19 Comparing this F44 value with previous studies, it is apparent that the F44 observed in the 20 OOA factor in this study is lower than others from two component organic results (Ng et al., This observation highlights a distinct difference in PMF_{Org MS} vs. PMF_{Full MS}. 21 2010). 22 Typically, the contribution of m/z 44 to the total organic spectrum is highest in the OOA 23 factor when PMF is applied to only the organic MS. However, in this study it represents the 24 third most oxygenated organic component among the PMF factors, yet the highest contributor 25 to oxygenated organic aerosol. Comparing with mass spectral profiles from other studies, this OOA factor more closely resembled semi-volatile OOAII factors, instead of highly 26 27 oxygenated OOAI factors, given the comparatively low m/z 44/43 ratio. The diurnal trend provided further insight into the nature of the OOA factor (Figure 11d). It can be seen that 28 29 this factor was significantly higher overnight than during daylight hours, suggesting either buildup in concentration in the nighttime boundary layer, or semi-volatile nighttime gas to 30 31 particle partitioning. The latter was suspected, as lower concentrations were observed on 32 warmer nights, suggestive of a temperature/volatility concentration dependence. In addition, the OOA factor exhibited a stronger anti-correlation with temperature than the Nitrate-OA factor (r = -0.26), suggesting that temperature played a greater role in explaining the variability in the OOA factor than in the Nitrate-OA factor. This was evident in increased contributions in the OOA factor mid-way through to the end of the campaign, when the air was colder. These results show that less oxygenated, and thus likely more local and semivolatile OOA, appeared to be a significant fraction of the total oxygenated organic aerosol.

The moderately oxidized and semi-volatile behavior of the OOA factor was further demonstrated through the highest correlation amongst all factors with acetaldehyde ($r^2 =$ 0.31), a VOC with an atmospheric lifetime of less than a day. Slowik et al., also report observing high correlations between a semi-volatile, OOAII factor and acetaldehyde in Toronto, ON (Slowik et al., 2009), and a similar conclusion can be drawn here that this factor's variability is influenced by both oxidation timescale as well as gas-particle partitioning. Ultimately, this factor more closely resembles OOAII rather than OOAI.

14 The geographic origins of the OOA factor were investigated by both CPF and PSCF. The 15 CPF, shown in Figure 8f, shows a similar result as was found for the HOA factor, with a 16 moderate association with emissions to the north-east, suggestive of contributions from 17 traffic. However, on the whole, no strong directionality was observed, supporting local 18 oxidation of organic aerosol. The PSCF for this factor was also generated (not shown), 19 although no distinct source regions could be identified, reinforcing the notion that transport 20 was not responsible for the temporal behavior of OOA in Windsor's winter-time air. Hence, 21 the OOA factor mainly represented local atmospheric processing, and appeared to be 22 characterized by less oxygenated, semi-volatile organic aerosol.

23 3.2 PMF_{Full MS} vs. PMF_{Org MS}

24 As only two other studies to date have performed PMF on the complete or near complete MS, 25 the following section investigates the effect of including both inorganics and organics in the 26 analysis. The two to six factor solutions from PMF_{Org MS} were examined and compared to the 27 six factor PMF_{Full MS} solution. Based on the mass spectral profiles and PMF diagnostics, the 28 three factor solution provided the most physically meaningful results. A two factor solution generated Amine and OOA factors, and moving to a three factor solution produced Amine, 29 OOA and HOA factors similar in profile to those from the PMF_{Full MS} analysis. Adding a 30 31 fourth factor caused the OOA factor to split, leading to another with an MS which did not

1 sufficiently resemble any known factors (nor any of the organic fractions contained within the 2 factors from the PMF_{Full MS} solution). The same applied to the five and six factor solutions. A 3 comparison between higher order solutions from the PMF_{Org MS} analysis and the organic fraction of the PMF_{Full MS} six factor solution showed that they did not contain the same 4 5 information, indicating that organics were resolved differently between the analyses. As the three factor solution led to the most physically meaningful results, it was compared with the 6 7 six factor PMF_{Full MS} solution, and is discussed below. The three factor solution justification, 8 along with descriptions of the four to six factor solutions are presented in the supplement.

9 Overall, the organic fraction was mainly split into the HOA and OOA factors, whereby these 10 factors each contributed 53 and 47% relatively to the total OA as assessed by the $PMF_{Org MS}$ 11 analysis. This nearly 50:50 division between these two organic components was somewhat 12 different than that found in other two component studies in urban areas, with HOA making a 13 greater contribution (Zhang et al., 2007). Such a difference could possibly be attributed to the 14 significant industrial activity and traffic contributions from the international border crossing.

15 **3.2.1 Amine factor comparison**

The mass spectra of the Amine factors extracted from the two PMF analyses are compared in Figure 12a. The most notable difference between the two profiles was the presence of a significant contribution from m/z 30 in PMF_{Full MS} analysis. The majority of this signal was removed during extraction of the organic MS as it is mostly attributed to NO⁺ in the fragmentation table. Excellent agreement was found between the factor time series ($r^2 =$ 0.99), indicating that this factor was similarly extracted in the two analyses (Figure 13a).

The extraction of the Amine factor in this study highlights advantages of applying PMF to the 22 full AMS mass spectrum for unit mass resolution data. Firstly, it can be seen that PMF can 23 provide better resolution of AMS species (i.e., SO₄²⁻, NO₃⁻, NH₄⁺, Cl⁻, organics) than using 24 the traditional AMS fragmentation scheme, as this method does not pre-assume the nature of 25 26 chemical species. In this study, the presence of the Amine factor had consequences for the 27 estimation of organics mass concentrations, and ultimately sub-micron PM mass. А comparison between the AMS resolved species calculated pre-PMF with those calculated 28 from the factors post-PMF_{Full MS} analysis (excluding any mass associated with the Amine 29 30 factor) is presented in Table 1. It can be seen that organics may have been overestimated by up 11%. Although the PMF_{Org MS} analysis could be improved by stoichiometrically fixing the 31

1 contribution of the subtracted inorganic component of m/z 30 (NO⁺) to m/z 46 to include 2 amine ions in the analysis, this was not performed for this analysis in order to highlight the 3 potential benefits of using the PMF_{Full MS} method (namely that no a priori assumptions are 4 required regarding the chemical nature of the aerosol). Ultimately, when comparing the pre-5 PMF sub-micron PM mass to the post-PMF reconstructed mass, it can be seen that the mass 6 may have been overestimated by up to 5% using the traditional fragmentation scheme.

7 Second, PMF of the full MS also provided better factor resolution and mass estimates than 8 PMF of the organic MS; with PMF_{Org MS}, the organic MS was extracted through application of 9 the organic fragmentation pattern, while in the PMF_{Full MS} analysis, the full MS remained as 10 measured, except for subtraction of air and water. Comparison between the PMF_{Full MS} and PMF_{Org MS} Amine factor time series showed that the mass of the Amine factor from the 11 12 PMF_{Org MS} was 2.6 times greater than that from the PMF_{Full MS} analysis. This difference 13 accounted for both the differences in RIE (an effective RIE of 4.3 was assumed in the PMF_{Full} 14 _{MS} analysis), as well as mass contributions from other fragments (namely m/z 30, which is 15 typically associated with nitrate but may also represent amines). A more direct comparison 16 between the Amine factors resolved in the PMF_{Full MS} and PMF_{Org MS} analyses using NO₃⁻eq mass showed that the factor resolved in the PMF_{Full MS} analysis contained 25% more mass, 17 18 given inorganic contributions. This suggests that some factors may not be fully resolved 19 using the PMF_{Org MS} method. As amine species were assigned as organics in the PMF_{Org MS} 20 analysis (with an RIE of 1.4 as opposed to 4.3), it can be seen that errors of up to two and a 21 half times the mass may arise from not considering alternate possibilities for the chemical 22 nature of certain species, which may become clearer through the extraction of the factor's full 23 MS. It is possible that in some cases (e.g., amine rich areas with amines of higher RIE), this 24 type of error could outweigh other AMS errors such as collection efficiency.

25 **3.2.2 HOA factor comparison**

Figure 12b presents the mass spectral comparison between the HOA factors resolved from PMF analysis of the full MS vs. the organic MS. The HOA factor mass spectra from the two PMF analyses compared very well ($r^2 = 0.98$), and good agreement was also found between their time series (r^2 of 0.98); a slope of 0.91 for the time series comparison indicated that a portion of this factor had been cross-apportioned to other factors in the full MS analysis (Figure 13b). Due to minimal differences between MS, it was difficult to distinguish across which factors this HOA was further distributed in the PMF_{Full MS} analysis.

1 **3.2.3 OOA factor comparison**

2 Figure 12c provides a comparison between the OOA factor profiles extracted in the PMF_{Full} 3 _{MS} and PMF_{Org MS} analyses. Enhanced cross-apportionment of m/z 44 to other factors in the PMF_{Full MS} analysis is evident due to the significant difference in the relative contributions of 4 5 F44 in the OOA factors (i.e., $PMF_{Full MS} F44 = 0.09$, $PMF_{Org MS} F44 = 0.15$). An F44 = 0.15 6 found in the PMF_{Org MS} in this factor is comparable to that found in other studies, also 7 indicative that this factor included the highly oxidized organics that were separated into the 8 Sulphate-OA factor in PMF_{Full MS}. Comparing the time-series of the OOA factors between the PMF_{Full MS} and PMF_{Org MS} analyses, a comparatively low $r^2 = 0.72$ highlighted that some OOA 9 10 signal was cross-apportioned to inorganic factors in the PMF_{Full MS} analysis (Figure 13c). A 11 mass difference was also observed between the factors, with the PMF_{Full MS} weighing on average 0.76µgm⁻³, and the PMF_{Org MS} measuring on average 1.01µgm⁻³. A higher mass 12 13 attributed to the PMF_{Org MS} OOA factor is likely linked with the higher contribution from m/z14 44. A comparison of just the organic contributions from the factors containing oxidized organic aerosol (Nitrate-OA, Sulphate-OA and OOA) with the OOA factor from the PMF_{Org} 15 _{MS} three factor solution showed that the r^2 was much improved from 0.72 to 0.94 with the 16 additional variability from the organic component of these factors (Figure 13d). A slope 17 18 slightly higher than unity suggested that the organic aerosol extracted from these three factors 19 was likely not entirely oxidized organic aerosol as defined from the OOA factor from the 20 PMF_{Org MS} analysis, and may have contained some less oxidized aerosol as well. This was corroborated by a slope of slightly less than unity for the HOA factor comparison. 21

22 Interestingly, the OOA factor could not be split into more OOA factors (i.e., OOAI and 23 OOAII) using PMF analysis of the organics alone. Moving to a four factor solution resulted 24 in an OOA factor, and an Other OA factor, which could not be fully justified, as its mass 25 spectrum did not sufficiently resemble any known mass spectra to accept the solution (see supplement for further details). It is likely that the inorganics in the PMF_{Full MS} analysis are 26 27 providing additional correlational structure to more effectively apportion organics. If moderately oxygenated organics are calculated as those apportioned to the OOA and Nitrate-28 29 OA factors, and highly oxygenated organics are those apportioned to the Sulphate-OA factor, 30 it can be seen that highly oxygenated organics comprise only about a quarter of the total 31 oxygenated organics. With the low prevalence of highly oxygenated organics observed in this 32 campaign, these species may not have carried sufficient weight in the model to result in the splitting of the OOA factor in the PMF_{Org MS} analysis into OOAI and OOAII types. With the
 OOA dominated by moderately oxidized organics, the addition of the inorganics introduced
 correlational structure for these more minor, more oxidized OAA types could be isolated.

The effect of having performed PMF on the full MS is evident when comparing the time 4 series between ${SO_4}^{2\text{-}}$ and the $PMF_{Full\ MS}$ and $PMF_{Org\ MS}$ OOA factors. In many previous 5 studies, particularly in urban locations, the OOA factor from PMF of the organic MS has been 6 highly correlated with SO_4^{2-} . Here an r^2 of 0.38 was found between SO_4^{2-} and OOA from 7 PMF of the organic fraction. However, following PMF of the full MS, the correlation 8 between the PMF_{Full MS} OOA factor and SO₄²⁻ was markedly lower ($r^2 = 0.08$). Thus, the 9 PMF_{Full MS} OOA factor shows a different temporal character, due to removal of the highly 10 11 oxidized contribution. The remaining OOA that characterizes this factor thus more resembles the OOAII typically found when two OOA factors are extracted during PMF_{Org MS}. Thus, 12 13 apportionment of organics of varying degrees of oxidation was more effectively accomplished through the PMF_{Full MS} over the PMF_{Org MS} method, most likely due to correlations with 14 inorganic species of differing atmospheric behavior (e.g., volatility), and origins. 15

16 **4 Conclusions**

17 Six factors were identified through PMF analysis of the full AMS mass spectrum: an Amine 18 factor suspected to be triethylamine nitrate; a Sulphate-OA factor composed of mostly 19 $(NH_4)_2SO_4$ and highly oxidized organics; a Nitrate-OA factor consisting of mostly NH_4NO_3 20 along with moderately oxidized organics; a Chloride factor hypothesized to be mostly NH_4Cl 21 along with organics and sulphate; an HOA factor classified as mostly primary organics from 22 fossil fuel combustion; and an OOA factor, consisting of moderately oxidized organics.

23 Factor variability was governed by chemical processing, primary source emissions, and The OOA factor appeared to be mostly influenced by local processing as it 24 transport. 25 exhibited a strong nighttime diurnal trend consistent with gas to particle partitioning. While both the Amine and Chloride factors showed temporal trends with brief excursions suggestive 26 27 of primary source emissions, rapid secondary aerosol formation could not be ruled out in 28 either case. Rapid in-plume secondary formation of the Chloride factor was supported by its sporadic appearance, and concurrent appearance with short-lived spikes in SO_4^{2-} , consistent 29 with industrial emissions plumes. While the precise mechanism of formation for the Amine 30 31 factor could not be validated, it was hypothesized to be a result of acid-base reaction of primary triethylamine gaseous emissions with HNO₃. Two significant industrial sources of 32

amines located to the southwest of the measurement site were likely the main contributing sources. By contrast, the main determinant for the HOA factor was primary source emissions, given its strong daytime diurnal trend, and high correlations with short-lived traffic related gases, such as NO. Regional transport was most likely the dominant determinant for variability for the Sulphate-OA and Nitrate-OA factors, although local gas to particle partitioning also appeared to contribute to the Nitrate-OA factor.

7 Allowing for PMF to cross-apportion inorganic and organic species between factors led to 8 richer conclusions regarding the potential sources, chemical nature, and behavior of certain 9 factors than would otherwise have been obtained through application of PMF of organics 10 only. First, cross-apportionment of inorganic species between factors led to some enhanced 11 aerosol chemistry conclusions, such as the relative degree of neutralization of co-varying inorganic species. It was found that NH_4^+ was cross-apportioned to the Sulphate-OA, Nitrate-12 13 OA and Chloride factors, and on the whole, each of these factors appeared reasonably neutral, 14 at least relative to one another. In an eight factor solution, a more acidic Local Sulphate 15 factor appeared. While this solution on the whole could not be justified, it suggests that acidic factors could be extracted in other studies. Second, inclusion of inorganics led to enhanced 16 17 apportionment of organic aerosol, which in turn led to a better understanding of each factor's degree of oxygenation and chemical nature. For instance, less oxygenated organics were 18 19 associated with the Chloride and HOA factors, which supported local, primary emissions 20 Inclusion of the inorganics also provided additional correlational structure to sources. 21 apportion the oxygenated organic fraction, with the most highly oxygenated OA apportioned to the regional Sulphate-OA factor, and the moderately oxygenated OA apportioned to the 22 23 Nitrate-OA and OOA factors. Using this technique, the resulting OOA factor displayed behavior typical of temperature-dependent gas to particle partitioning, while the more 24 25 oxidized organics contained in the Nitrate-OA and Sulphate-OA factors appeared longer 26 lived, and more associated with regional transport. Without the inclusion of inorganics, the 27 highly oxygenated and moderately oxygenated organics could not be effectively separated.

The methodology used in this study also proved advantageous for extracting atypical factors, and for obtaining more accurate mass quantification for data obtained from a unit mass resolution AMS. In this study, PMF of the full MS resulted in the extraction of the Amine factor, which was only partially extracted using PMF of only the organic MS. A more complete mass spectral profile including inorganic fragments (e.g., m/z 30) was extracted,

which permitted further exploration into the factor's chemical nature. Although its formation 1 2 mechanism could not be positively validated, the Amine factor was ascertained to have resulted from an acid-base neutralization reaction, and a factor specific relative ionization 3 efficiency was calculated to provide improved mass estimates. This method was useful for 4 5 preventing overestimation of certain species, such as nitrate, when the aerosol is influenced by species unaccounted for by the typical AMS analysis scheme. It is likely that this method 6 7 would be even more useful for HR-ToF-AMS data, due to the ability to better understand the 8 chemical nature of atypical factors from high resolution mass spectra. In summary, PMF of 9 the full AMS unit mass resolution MS has been shown to be a useful method to obtain 10 additional insights into the sources and processes governing fluctuations in non-refractory 11 sub-micron PM chemical composition, in comparison to PMF of the organics alone.

12 Acknowledgements

13 These measurements were supported by Environment Canada. Funding for SOCAAR was 14 provided by the Canadian Foundation for Innovation, the Ontario Innovation Trust, and the 15 Ontario Research Fund. MLM is grateful to Environment Canada for funding this analysis 16 through the Research Affiliate Program (RAP). The authors gratefully acknowledge the 17 NOAA Air Resources Laboratory (ARL) for the provision of the HYSPLIT transport and 18 dispersion model and/or READY website (http://www.arl.noaa.gov/ready.php) (Draxler and 19 Rolph, 2010; Rolph, 2010) used in this publication. The authors also thank anonymous referees for helpful comments. 20

21

1 References

Aiken, A. C., Decarlo, P. F., Kroll, J. H., Worsnop, D. R., Huffman, J. A., Docherty, K. S.,
Ulbrich, I. M., Mohr, C., Kimmel, J. R., Sueper, D., Sun, Y., Zhang, Q., Trimborn, A.,
Northway, M., Ziemann, P. J., Canagaratna, M. R., Onasch, T. B., Alfarra, M. R., Prevot, A.
S. H., Dommen, J., Duplissy, J., Metzger, A., Baltensperger, U., and Jimenez, J. L.: O/C and
OM/OC ratios of primary, secondary, and ambient organic aerosols with high-resolution timeof-flight aerosol mass spectrometry, Environmental Science and Technology, 42, 4478-4485,
2008.

9 Aiken, A. C., Salcedo, D., Cubison, M. J., Huffman, J. A., DeCarlo, P. F., Ulbrich, I. M., Docherty, K. S., Sueper, D., Kimmel, J. R., Worsnop, D. R., Trimborn, A., Northway, M., 10 Stone, E. A., Schauer, J. J., Volkamer, R., Fortner, E., De Foy, B., Wang, J., Laskin, A., 11 12 Shutthanandan, V., Zheng, J., Zhang, R., Gaffney, J., Marley, N. A., Paredes-Miranda, G., 13 Arnott, W. P., Molina, L. T., Sosa, G., and Jimenez, J. L.: Mexico City aerosol analysis 14 during MILAGRO using high resolution aerosol mass spectrometry at the urban supersite 15 (T0) - Part 1: Fine particle composition and organic source apportionment, Atmospheric Chemistry and Physics, 9, 6633-6653, 2009. 16

- 17 Allan, D., Williams, P. I., Morgan, W. T., Martin, C. L., Flynn, M. J., Lee, J., Nemitz, E.,
- 18 Phillips, G. J., Gallagher, M. W., and Coe, H.: Contributions from transport, solid fuel
- 19 burning and cooking to primary organic aerosols in two UK cities, Atmospheric Chemistry
- 20 and Physics, 10, 647-668, 2010.
- 21 Allan, J. D., Jimenez, J. L., Williams, P. I., Alfarra, M. R., Bower, K. N., Jayne, J. T., Coe,
- 22 H., and Worsnop, D. R.: Quantitative sampling using an Aerodyne aerosol mass spectrometer:
- 1. Techniques of data interpretation and error analysis (vol 108, art no 4090, 2003), Journal of
- 24 Geophysical Research-Atmospheres, 108, 1, 4283, 2003, doi:10.1029/2002JD002358.
- Allan, J. D., Delia, A. E., Coe, H., Bower, K. N., Alfarra, M. R., Jimenez, J. L., Middlebrook,
 A. M., Drewnick, F., Onasch, T. B., Canagaratna, M. R., Jayne, J. T., and Worsnopf, D. R.: A
 generalised method for the extraction of chemically resolved mass spectra from aerodyne
 aerosol mass spectrometer data, Journal of Aerosol Science, 35, 909-922, 2004.
- Anderson, J. O., Thundiyil, J. G., and Stolbach, A.: Clearing the Air: A Review of the Effects
 of Particulate Matter Air Pollution on Human Health, Journal of Medical Toxicology, 8, 166175, 2012.
- Ashbaugh, L. L., Malm, W. C., and Sadeh, W. Z.: A residence time probability analysis of
 sulfur concentrations at Grand-Canyon National Park, Atmospheric Environment, 19, 12631270, 1985.
- 35 Brook, R. D., Rajagopalan, S., Pope, C. A., Brook, J. R., Bhatnagar, A., Diez-Roux, A. V.,
- 36 Holguin, F., Hong, Y., Luepker, R. V., Mittleman, M. A., Peters, A., Siscovick, D., Smith, S.
- C., Whitsel, L., and Kaufman, J. D.: Particulate matter air pollution and cardiovascular
 disease: An update to the scientific statement from the american heart association,
 Circulation 121 2221 2278 2010
- 39 Circulation, 121, 2331-2378, 2010.
- 40 Brown, S. G., Lee, T., Norris, G. A., Roberts, P. T., Collett, J. L., Paatero, P., and Worsnop,
- 41 D. R.: Receptor modeling of near-roadway aerosol mass spectrometer data in Las Vegas,
- 42 Nevada, with EPA PMF, Atmospheric Chemistry and Physics, 12, 309-325, 10.5194/acp-12-
- 43 309-2012, 2012.

- 1 Buset, K. C., Evans, G. J., Leaitch, W. R., Brook, J. R., and Toom-Sauntry, D.: Use of
- 2 advanced receptor modelling for analysis of an intensive 5-week aerosol sampling campaign,
- 3 Atmospheric Environment, 40, S482-S499, 2006.
- 4 Canagaratna, M. R., Jayne, J. T., Jimenez, J. L., Allan, J. D., Alfarra, M. R., Zhang, Q.,
- 5 Onasch, T. B., Drewnick, F., Coe, H., Middlebrook, A., Delia, A., Williams, L. R., Trimborn,
- 6 A. M., Northway, M. J., DeCarlo, P. F., Kolb, C. E., Davidovits, P., and Worsnop, D. R.:
- 7 Chemical and microphysical characterization of ambient aerosols with the aerodyne aerosol
- 8 mass spectrometer, Mass Spectrometry Reviews, 26, 185-222, 2007a.
- 9 Canagaratna, M. R., Jayne, J. T., Jimenez, J. L., Allan, J. D., Alfarra, M. R., Zhang, Q.,
- 10 Onasch, T. B., Drewnick, F., Coe, H., Middlebrook, A., Delia, A., Williams, L. R., Trimborn,
- 11 A. M., Northway, M. J., DeCarlo, P. F., Kolb, C. E., Davidovits, P., and Worsnop, D. R.:
- 12 Chemical and microphysical characterization of ambient aerosols with the aerodyne aerosol
- 13 mass spectrometer, Mass Spectrom. Rev., 26, 185–222, 2007b.
- 14 Chang, R. Y. W., Leck, C., Graus, M., Müller, M., Paatero, J., Burkhart, J. F., Stohl, A., Orr,
- L. H., Hayden, K., Li, S. M., Hansel, A., Tjernström, M., Leaitch, W. R., and Abbatt, J. P. D.:
 Aerosol composition and sources in the central Arctic Ocean during ASCOS, Atmospheric
- 17 Chemistry and Physics, 11, 10619-10636, 2011.
- 18 Chirico, R., Prevot, A. S. H., DeCarlo, P. F., Heringa, M. F., Richter, R., Weingartner, E., and
- 19 Baltensperger, U.: Aerosol and trace gas vehicle emission factors measured in a tunnel using
- 20 an Aerosol Mass Spectrometer and other on-line instrumentation, Atmospheric Environment,
- 21 45, 2182-2192, 10.1016/j.atmosenv.2011.01.069, 2011.
- 22 Crippa, M., DeCarlo, P. F., Slowik, J. G., Mohr, C., Heringa, M. F., Chirico, R., Poulain, L.,
- 23 Freutel, F., Sciare, J., Cozic, J., Di Marco, C. F., Elsasser, M., José, N., Marchand, N., Abidi,
- 24 E., Wiedensohler, A., Drewnick, F., Schneider, J., Borrmann, S., Nemitz, E., Zimmermann,
- 25 R., Jaffrezo, J.-L., Prévôt, A. S. H., and Baltensperger, U.: Wintertime aerosol chemical
- 26 composition and source apportionment of the organic fraction in the metropolitan area of
- 27 Paris, Atmos. Chem. Phys., 13, 961-981, 2013a.
- 28 Crippa, M., El Haddad, I., Slowik, J. G., Decarlo, P. F., Mohr, C., Heringa, M. F., Chirico, R.,
- Marchand, N., Sciare, J., Baltensperger, U., and Prévôt, A. S. H.: Identification of marine and continental aerosol sources in Paris using high resolution aerosol mass spectrometry, Journal
- 21 of Coophysical Bassarch D. Atmospheres, 118, 1050, 1062, 2012h
- 31 of Geophysical Research D: Atmospheres, 118, 1950-1963, 2013b.
- 32 Docherty, K. S., Aiken, A. C., Huffman, J. A., Ulbrich, I. M., DeCarlo, P. F., Sueper, D.,
- 33 Worsnop, D. R., Snyder, D. C., Peltier, R. E., Weber, R. J., Grover, B. D., Eatough, D. J.,
- 34 Williams, B. J., Goldstein, A. H., Ziemann, P. J., and Jimenez, J. L.: The 2005 Study of
- 35 Organic Aerosols at Riverside (SOAR-1): instrumental intercomparisons and fine particle
- 36 composition, Atmospheric Chemistry and Physics, 11, 12387-12420, 10.5194/acp-11-12387 37 2011, 2011.
- 38 Drewnick, F., Hings, S. S., Curtius, J., Eerdekens, G., and Williams, J.: Measurement of fine
- particulate and gas-phase species during the New Year's fireworks 2005 in Mainz, Germany,
 Atmospheric Environment, 40, 4316-4327, 2006.
- 41 Environment Canada: National Pollutant Release Inventory, 2013, downloadable at 42 http://www.ec.gc.ca/inrp-npri/, accessed December 2013.
- 43 Ge, X., Wexler, A. S., and Clegg, S. L.: Atmospheric amines Part I. A review, Atmospheric
- 44 Environment, 45, 524-546, 2011.

- 1 Godri, K. J., Evans, G. J., Slowik, J., Knox, A., Abbatt, J., Brook, J., Dann, T., and Dabek-
- 2 Zlotorzynska, E.: Evaluation and application of a semi-continuous chemical characterization
- 3 system for water soluble inorganic PM2.5 and associated precursor gases, Atmospheric
- 4 Measurement Techniques, 2, 65-80, 2009.
- 5 Gordon, G. E.: Receptor Models, Environmental Science & Technology, 14, 792-800, 1980.
- Greim, H., Bury, D., Klimisch, H. J., Oeben-Negele, M., and Ziegler-Skylakakis, K.: Toxicity
 of aliphatic amines: Structure-activity relationship, Chemosphere, 36, 271-295, 1998.
- 8 Hildebrandt, L., Kostenidou, E., Lanz, V. A., Prevot, A. S. H., Baltensperger, U.,
- 9 Mihalopoulos, N., Laaksonen, A., Donahue, N. M., and Pandis, S. N.: Sources and
- 10 atmospheric processing of organic aerosol in the Mediterranean: Insights from aerosol mass
- spectrometer factor analysis, Atmospheric Chemistry and Physics, 11, 12499-12515, 2011.
- 12 Hogrefe, O., Drewnick, F., Lala, G. G., Schwab, J. J., and Demerjian, K. L.: Development,
- 13 operation and applications of an aerosol generation, calibration and research facility, Aerosol
- 14 Science and Technology, 38, 196-214, 2004.
- Hopke, P. K.: Recent developments in receptor modeling, Journal of Chemometrics, 17, 255-265, 2003.
- 17 Huffman, J. A., Docherty, K. S., Aiken, A. C., Cubison, M. J., Ulbrich, I. M., Decarlo, P. F.,
- 18 Sueper, D., Jayne, J. T., Worsnop, D. R., Ziemann, P. J., and Jimenez, J. L.: Chemically-
- resolved aerosol volatility measurements from two megacity field studies, AtmosphericChemistry and Physics, 9, 7161-7182, 2009.
- IPCC: Intergovernmental Panel on Climate Change, Fifth Assessment Report, 2013,
 downloadable at http://www.ipcc.ch/, accessed December 2013.
- 23 Jayne, J. T., Leard, D. C., Zhang, X. F., Davidovits, P., Smith, K. A., Kolb, C. E., and
- 24 Worsnop, D. R., Development of an aerosol mass spectrometer for size and composition
- analysis of submicron particles, Aerosol Sci. Tech., 33, 49–70, 2000.
- 26 Jeong, C.-H., McGuire, M. L., Herod, D., Dann, T., Dabek-Zlotorzynska, E., Wang, D., Ding,
- 27 L., Celo, V., Mathieu, D., and Evans, G. J.: Receptor Model Based Identification of the
- 28 Sources of PM2.5 in Canadian Cities, Atmospheric Pollution Research, 2, 158-171, 2011.
- 29 Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H.,
- 30 DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. M.,
- 31 Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. A.,
- Hueglin, C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara, B., Ehn, M., Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J., Dunlea, E. J.,
- Huffman, J. A., Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo, Y.,
- 35 Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D., Cottrell,
- 36 L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J. Y., Zhang, Y.
- 37 M., Dzepina, K., Kimmel, J. R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M.,
- 38 Williams, L. R., Wood, E. C., Middlebrook, A. M., Kolb, C. E., Baltensperger, U., and
- 39 Worsnop, D. R.: Evolution of organic aerosols in the atmosphere, Science, 326, 1525-1529,
- 40 2009.
- 41 Langley, L., Leaitch, W. R., Lohmann, U., Shantz, N. C., and Worsnop, D. R.: Contributions
- 42 from DMS and ship emissions to CCN observed over the summertime North Pacific,
- 43 Atmospheric Chemistry and Physics, 10, 1287-1314, 2010.

- 1 Lanz, V. A., Alfarra, M. R., Baltensperger, U., Buchmann, B., Hueglin, C., and Prevot, A. S.
- 2 H.: Source apportionment of submicron organic aerosols at an urban site by factor analytical
- 3 modelling of aerosol mass spectra, Atmospheric Chemistry and Physics, 7, 1503-1522, 2007.
- 4 Lanz, V. A., Alfarra, M. R., Baltensperger, U., Buchmann, B., Hueglin, C., Szidat, S., Wehrli,
- 5 M. N., Wacker, L., Weimer, S., Caseiro, A., Puxbaum, H., and Prevot, A. S. H.: Source
- 6 attribution of submicron organic aerosols during wintertime inversions by advanced factor
- 7 analysis of aerosol mass spectra, Environmental Science and Technology, 42, 214-220, 2008.
- 8 Li, Y. Q., Schwab, J. J., and Demerjian, K. L.: Measurements of ambient ammonia using a
- 9 tunable diode laser absorption spectrometer: Characteristics of ambient ammonia emissions in
- 10 an urban area of New York City, Journal of Geophysical Research-Atmospheres, 111, 11,
- 11 D10s02, 2006, doi:10.1029/2005JD006275.
- Liggio, J., Li, S. M., Vlasenko, A., Stroud, C., and Makar, P.: Depression of ammonia uptake
 to sulfuric acid aerosols by competing uptake of ambient organic gases, Environmental
 Science and Technology, 45, 2790-2796, 2011.
- Liggio, J., and Li, S. M.: A new source of oxygenated organic aerosol and oligomers,
 Atmospheric Chemistry and Physics, 13, 2989-3002, 2013.
- 17 Malloy, Q. G. J., Qi, L., Warren, B., Cocker, D. R., Erupe, M. E., and Silva, P. J.: Secondary
- 18 organic aerosol formation from primary aliphatic amines with NO3 radical, Atmospheric
- 19 Chemistry and Physics, 9, 2051-2060, 2009.
- 20 McGuire, M. L., Jeong, C. H., Slowik, J. G., Chang, R. Y. W., Corbin, J. C., Lu, G., Mihele,
- C., Rehbein, P. J. G., Sills, D. M. L., Abbatt, J. P. D., Brook, J. R., and Evans, G. J.:
 Elucidating determinants of aerosol composition through particle-type-based receptor
 modeling, Atmospheric Chemistry and Physics, 11, 8133-8155, 10.5194/acp-11-8133-2011,
- 24 2011.
- McLafferty, F. W., and Tureček, F.: Interpretation of mass spectra, 4th ed., University
 Science Books, Sausalito, Calif., xviii, 371 p. pp., 1993.
- 27 Middlebrook, A. M., Bahreini, R., Jimenez, J. L., and Canagaratna, M. R.: Evaluation of
- 28 composition-dependent collection efficiencies for the Aerodyne aerosol mass spectrometer
- using field data, Aerosol Science and Technology, 46, 258-271, 2012.
- 30 Moffet, R. C., de Foy, B., Molina, L. T., Molina, M. J., and Prather, K. A.: Measurement of 31 ambient aerosols in northern Mexico City by single particle mass spectrometry, Atmospheric
- 32 Chemistry and Physics, 8, 4499-4516, 2008a.
- 33 Moffet, R. C., Desyaterik, Y., Hopkins, R. J., Tivanski, A. V., Gilles, M. K., Wang, Y.,
- 34 Shutthanandan, V., Molina, L. T., Abraham, R. G., Johnson, K. S., Mugica, V., Molina, M. J.,
- 35 Laskin, A., and Prather, K. A.: Characterization of aerosols containing Zn, Pb, and Cl from an
- 36 industrial region of Mexico City, Environmental Science & Technology, 42, 7091-7097, doi:
- 37 10.1021/es7030483, 2008b.
- 38 Mohr, C., DeCarlo, P. F., Heringa, M. F., Chirico, R., Slowik, J. G., Richter, R., Reche, C.,
- Alastuey, A., Querol, X., Seco, R., Peñuelas, J., Jiménez, J. L., Crippa, M., Zimmermann, R.,
- 40 Baltensperger, U., and Prévôt, A. S. H.: Identification and quantification of organic aerosol 41 from cooking and other sources in Barcelona using aerosol mass spectrometer data,
- 42 Atmospheric Chemistry and Physics, 12, 1649-1665, 2012.
- 43 Murphy, S. M., Sorooshian, A., Kroll, J. H., Ng, N. L., Chhabra, P., Tong, C., Surratt, J. D., 44 Knipping, E., Flagan, R. C., and Seinfeld, J. H.: Secondary aerosol formation from

- atmospheric reactions of aliphatic amines, Atmospheric Chemistry and Physics, 7, 2313-2337,
 2007.
- 3 Ng, N. L., Canagaratna, M. R., Zhang, Q., Jimenez, J. L., Tian, J., Ulbrich, I. M., Kroll, J. H.,
- 4 Docherty, K. S., Chhabra, P. S., Bahreini, R., Murphy, S. M., Seinfeld, J. H., Hildebrandt, L.,
- 5 Donahue, N. M., Decarlo, P. F., Lanz, V. A., Prévôt, A. S. H., Dinar, E., Rudich, Y., and
- 6 Worsnop, D. R.: Organic aerosol components observed in Northern Hemispheric datasets
- 7 from Aerosol Mass Spectrometry, Atmospheric Chemistry and Physics, 10, 4625-4641, 2010.
- 8 Norris, G., Vedantham, R., Duvall, R., Wade, K., Brown, S., Prouty, J., Bai, S., and
- 9 DeWinter, J.: EPA Positive Matrix Factorization 4.1 Fundamentals & User Guide, EPA,
- 10 Research Triangle Park, North Carolina, 2010.
- Paatero, P., and Tapper, U.: Analysis of Different Modes of Factor-Analysis as Least-Squares
 Fit Problems, Chemometrics and Intelligent Laboratory Systems, 18, 183-194, 1993.
- Paatero, P., and Tapper, U.: Positive Matrix Factorization A Nonnegative Factor Model
 With Optimal Utilization of Error-Estimates of Data Values, Environmetrics, 5, 111-126,
- 15 1994.
- Paatero, P.: Least squares formulation of robust non-negative factor analysis, Chemometricsand Intelligent Laboratory Systems, 37, 23-35, 1997.
- 18 Perron, N., Sandradewi, J., Alfarra, M. R., Lienemann, P., Gehrig, R., Kasper-Giebl, A.,
- 19 Lanz, V. A., Szidat, S., Ruff, M., Fahrni, S., Wacker, L., Baltensperger, U., and Prévôt, A. S.
- 20 H.: Composition and sources of particulate matter in an industrialised Alpine valley,
- 21 Atmospheric Chemistry and Physics Discussions, 10, 9391-9430, 2010.
- Pio, C. A., and Harrison, R. M.: The equilibrium of ammonium chloride aerosol with gaseous
 hydrochloric acid and ammonia under tropospheric conditions, Atmospheric Environment, 21,
 1243-1246, 1987a.
- Pio, C. A., and Harrison, R. M.: Vapour pressure of ammonium chloride aerosol: Effect of
 temperature and humidity, Atmospheric Environment, 21, 2711-2715, 1987b.
- Pope, C. A., and Dockery, D. W.: Health effects of fine particulate air pollution: Lines that
 connect, Journal of the Air & Waste Management Association, 56, 709-742, 2006.
- Pratt, K. A., Hatch, L. E., and Prather, K. A.: Seasonal volatility dependence of ambient
 particle phase amines, Environmental Science and Technology, 43, 5276-5281, 2009.
- 31 Rehbein, P. J. G., Jeong, C.-H., McGuire, M. L., Yao, X., Corbin, J., and Evans, G. J.: Cloud
- and Fog Processing Enhanced Gas-to-Particle Partitioning of Trimethylamine, Environmental
 Science and Technology, 45, 4346-4352, 2011.
- 34 Richard, A., Gianini, M. F. D., Mohr, C., Furger, M., Bukowiecki, N., Minguillón, M. C.,
- 35 Lienemann, P., Flechsig, U., Appel, K., Decarlo, P. F., Heringa, M. F., Chirico, R.,
- 36 Baltensperger, U., and Prévôt, A. S. H.: Source apportionment of size and time resolved trace
- 37 elements and organic aerosols from an urban courtyard site in Switzerland, Atmospheric
- 38 Chemistry and Physics, 11, 8945-8963, 2011.
- 39 Salcedo, D., Onasch, T. B., Dzepina, K., Canagaratna, M. R., Zhang, Q., Huffman, J. A.,
- 40 DeCarlo, P. F., Jayne, J. T., Mortimer, P., Worsnop, D. R., Kolb, C. E., Johnson, K. S.,
- 41 Zuberi, B., Marr, L. C., Volkamer, R., Molina, L. T., Molina, M. J., Cardenas, B., Bernabe, R.
- 42 M., Marquez, C., Gaffney, J. S., Marley, N. A., Laskin, A., Shutthanandan, V., Xie, Y.,
- 43 Brune, W., Lesher, R., Shirley, T., and Jimenez, J. L.: Characterization of ambient aerosols in

- 1 Mexico City during the MCMA-2003 campaign with Aerosol Mass Spectrometry: results
- 2 from the CENICA Supersite, Atmospheric Chemistry and Physics, 6, 925-946, 2006.
- 3 Silva, P. J., Erupe, M. E., Price, D., Elias, J., Malloy, Q. G. J., Li, Q., Warren, B., and Cocker
- 4 Iii, D. R.: Trimethylamine as precursor to secondary organic aerosol formation via nitrate
- 5 radical reaction in the atmosphere, Environmental Science and Technology, 42, 4689-4696,
- 6 2008.
- Slowik, J. G., Vlasenko, A., McGuire, M., Evans, G. J., and Abbatt, J. P. D.: Simultaneous
 factor analysis of organic particle and gas mass spectra: AMS and PTR-MS measurements at
- 9 an urban site, Atmospheric Chemistry and Physics Discussions, Submitted, 2009.
- 10 Slowik, J. G., Vlasenko, A., McGuire, M., Evans, G. J., and Abbatt, J. P. D.: Simultaneous
- 11 factor analysis of organic particle and gas mass spectra: AMS and PTR-MS measurements at
- 12 an urban site, Atmospheric Chemistry and Physics, 10, 1969-1988, 2010.
- Stein, S. E.: "Mass Spectra", NIST Chemistry WebBook, NIST Standard Reference Database
 Number 69, National Institute of Standards and Technology, Gaithersburg MD, 20899, 2013.
- 15 Sun, Y. L., Zhang, Q., Schwab, J. J., Demerjian, K. L., Chen, W. N., Bae, M. S., Hung, H. M.,
- 16 Hogrefe, O., Frank, B., Rattigan, O. V., and Lin, Y. C.: Characterization of the sources and
- 17 processes of organic and inorganic aerosols in New York city with a high resolution time-of-
- 18 flight aerosol mass spectrometer, Atmospheric Chemistry and Physics, 11, 1581-1602, 2011.
- 19 Sun, Y. L., Zhang, Q., Schwab, J. J., Yang, T., Ng, N. L., and Demerjian, K. L.: Factor
- analysis of combined organic and inorganic aerosol mass spectra from high resolution aerosol
- 21 mass spectrometer measurements, Atmospheric Chemistry and Physics, 12, 8537-8551, 2012.
- Tan, P. V., Evans, G. J., Tsai, J., Owega, S., Fila, S., Malpica, O., and Brook, J. R.: On-line
 analysis of urban particulate matter focusing on elevated wintertime aerosol concentrations,
 Environmental Science and Technology, 36, 3512-3518, 2002.
- Ulbrich, I. M., Canagaratna, M. R., Zhang, Q., Worsnop, D. R., and Jimenez, J. L.:
 Interpretation of organic components from positive matrix factorization of aerosol mass
 spectrometric data, Atmospheric Chemistry and Physics Discussions, 8, 6729-6791, 2008.
- Ulbrich, I. M., Canagaratna, M. R., Zhang, Q., Worsnop, D. R., and Jimenez, J. L.:
 Interpretation of Organic Components from Positive Matrix Factorization of Aerosol Mass
 Spectrometric Data, Atmospheric Chemistry and Physics, 9, 2891-2918, 2009.
- US EPA: National Emission Inventory (NEI) 2013: Inventory Data: Point Sector Data ALLNEI HAP Annual, Research Triangle Park, North Carolina, 2013.
- 33 US EPA: Toxics Release Inventory, Washington, DC, 2013.
- 34 Vlasenko, A., Slowik, J. G., Bottenheim, J. W., Brickell, P. C., Chang, R. Y. W., Maedonald,
- 35 A. M., Shantz, N. C., Sjostedt, S. J., Wiebe, H. A., Leaitch, W. R., and Abbatt, J. P. D.:
- 36 Measurements of VOCs by proton transfer reaction mass spectrometry at a rural Ontario site:
- 37 Sources and correlation to aerosol composition, J. Geophys. Res., 114, D21305,
- 38 doi:10.1029/2009JD012025.
- Watson, J. G.: Visibility: Science and Regulation, Journal of the Air and Waste ManagementAssociation, 52, 628-713, 2002.
- 41 Watson, J. G., Chen, L. W. A., Chow, J. C., Doraiswamy, P., and Lowenthal, D. H.: Source
- 42 apportionment: Findings from the U.S. supersites program, Journal of the Air and Waste
- 43 Management Association, 58, 265-288, 2008.

- 1 Zhang, Q., Alfarra, M. R., Worsnop, D. R., Allan, J. D., Coe, H., Canagaratna, M. R., and
- 2 Jimenez, J. L.: Deconvolution and quantification of hydrocarbon-like and oxygenated organic
- 3 aerosols based on aerosol mass spectrometry, Environmental Science & Technology, 39,
- 4 4938-4952, 2005.

5 Zhang, Q., Jimenez, J. L., Canagaratna, M. R., Allan, J. D., Coe, H., Ulbrich, I., Alfarra, M. 6 R., Takami, A., Middlebrook, A. M., Sun, Y. L., Dzepina, K., Dunlea, E., Docherty, K.,

7 DeCarlo, P. F., Salcedo, D., Onasch, T., Jayne, J. T., Miyoshi, T., Shimono, A., Hatakeyama,

8 S., Takegawa, N., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S.,

- 9 Demerjian, K., Williams, P., Bower, K., Bahreini, R., Cottrell, L., Griffin, R. J., Rautiainen,
- 10 J., Sun, J. Y., Zhang, Y. M., and Worsnop, D. R.: Ubiquity and dominance of oxygenated
- 11 species in organic aerosols in anthropogenically-influenced Northern Hemisphere
- 12 midlatitudes, Geophysical Research Letters, 34, 6, L13801, doi:10.1029/2007GL029979,
- 13 2007.

14 Zhang, Q., Jimenez, J. L., Canagaratna, M. R., Ulbrich, I. M., Ng, N. L., Worsnop, D. R., and

Sun, Y.: Understanding atmospheric organic aerosols via factor analysis of aerosol mass
 spectrometry: A review, Analytical and Bioanalytical Chemistry, 401, 3045-3067, 2011.

17 Zorn, S. R., Drewnick, F., Schott, M., Hoffmann, T., and Borrmann, S.: Characterization of

18 the South Atlantic marine boundary layer aerosol using an aerodyne aerosol mass

19 spectrometer, Atmospheric Chemistry and Physics, 8, 4711-4728, 2008.

1 Table Captions

- 2 Table 1: Descriptive statistics for AMS measured non-refractory sub-micron PM species, pre- and post-PMF analysis ($\mu g m^{-3}$).
- 4 Table 2: Chemical composition of the six factors from the PMF_{Full MS} analysis, extent of neutralization, and F44 for each factor's organic fraction.
- 6

7 Figure Captions

- Figure 1: The location of the sampling site (MicMac Park) in Windsor, Ontario relative to major industrial sources, namely coal fired power plants, steel mills, and potential large amine sources. It can be seen that the measurement site was located close to the largest international border crossing between the US and Canada (Huron Church road and the Ambassador Bridge).
- Figure 2: Meteorological conditions and PM_{2.5} mass concentration from the CRUISER
 TEOM for the MicMac Park winter campaign.
- 15 Figure 3: Time series of AMS-measured non-refractory sub-micron PM species.
- 16Figure 4:Time series of the factors from the six factor solution from $PMF_{Full MS}$ analysis.17The solid line represents FPeak = 0, and the range of uncertainties through18FPeak analysis (-10 and +10) is shown in the shaded regions.
- 19Figure 5:Mass spectra of factors from the six factor solution from PMF_{Full MS} analysis.2020Bars represent the central, chosen rotation, and dots show the range in mass21spectral variation from FPeak rotations (-10 and 10).
- 22Figure 6:Mass spectra of the organic fraction of the factors from the six factor solution23from $PMF_{Full MS}$ analysis. The Amine factor is not shown due to insufficient24information regarding the chemical nature of the Amine factor. Bars represent25the central, chosen rotation, and dots show the range in mass spectral variation26from FPeak rotations (-10 and 10).
- Figure 7: Chemical composition by factor and species components of the six PMF factors from the PMF_{Full MS} analysis.
- 29Figure 8:Conditional probability function (CPF) plots for the six factors from the six30factor solution of the $PMF_{Full MS}$ analysis, along with a wind rose plot (wind31speed in m s⁻¹). The strongest wind dependence is observed for the Amine and32Sulphate-OA factors, which show strong, yet slightly different directional33associations to the southwest.
- 34Figure 9:Potential source contribution function (PSCF) plot for the Sulphate-OA (a) and35the Nitrate-OA factor (b) from the $PMF_{Full MS}$ analysis. Both factors showed36regional source influences, with the Sulphate-OA factor showing more37prominent, distant influences to the south and southwest. The Nitrate-OA38factor showed the most prominent influences to the southwest, over Indiana39and Illinois.
- 40 Figure 10: Particulate sulphate concentration $(SO_4^{2^-})$ vs. $SO_{2(g)}$ concentration, and fraction 41 of $SO_4^{2^-}$ / Total S (Total S = $SO_4^{2^-}$ + SO_2) (a). Due to the winter conditions,

- 1 the fractional contribution of pSO_4^{2-} to total S was typically very low. Shown 2 in (b) are the averaged particle size distributions of AMS SO_4^{2-} over the entire 3 campaign, as well as during the extreme SO_4^{2-} spikes. Particulate SO_4^{2-} 4 measured during the spikes appeared highly local due to their significantly 5 smaller modal diameter.
- 6 Figure 11: Boxplots of diurnal trends for the Sulphate-OA (a), Nitrate-OA (b), HOA (c), 7 and OOA factors (d), from the PMF_{Full MS} analysis, along with temperature (e). 8 Boxes indicate interquartile ranges, horizontal lines indicate median hourly 9 values, cross markers indicate hourly means, and whiskers represent the 5th and 95th percentiles. The HOA factor demonstrated a strong diurnal trend 10 consistent with traffic patterns, while the OOA factor demonstrated a trend 11 12 more consistent with daytime lows and overnight highs. The Nitrate-OA factor 13 showed a minor diurnal trend indicative of more regional contributions as 14 compared to OOA, while the Sulphate-OA factor showed minimal diurnal 15 trend.
- 16 Figure 12: Comparison of factor profiles from the PMF_{Org MS} analysis to comparable 17 factors found by PMF of the full mass spectra (a, b, and c). The factors are 18 normalized to the total organic fraction signal, except for the Amine factor 19 which is normalized to the total factor signal. Only minor differences in mass 20 spectra are noted for the HOA factor. The most significant difference for the 21 Amine factor is represented by the addition of the m/z 30 peak, and for the 22 OOA factor with the difference in magnitude of the m/z 44 peak. Dots show 23 the range in mass spectral variation from FPeak rotations (-10 and 10). Also 24 shown in (d) is the average factor composition of organics for both the PMF of 25 the full mass spectrum and of the organics only.
- Figure 13: Scatter plots Amine (a), HOA (b), and OOA (c and d) factors comparing their mass derived from PMF_{Full MS} vs. PMF_{Org MS} analyses. One observation (01/31/2005 21:00) was removed from the HOA comparison to enhance the fit of the trendline.
- 30
- 31

Table 1: Descriptive statistics for AMS measured non-refractory sub-micron PM species, pre-and post-PMF analysis ($\mu g m^{-3}$).

	Org		$\mathrm{NH_4}^+$		NO ₃ ⁻		SO4 ²⁻		Cl		Total	
	Pre	Post	Pre	Post	Pre	Post	Pre	Post	Pre	Post	Pre	Post
Mean	2.73	2.45	1.13	1.15	2.28	2.20	1.18	1.13	0.12	0.12	7.45	7.10
1σ	2.41	2.17	1.01	1.04	2.58	2.53	0.97	0.91	0.18	0.17	6.30	6.12
Min	0.29	0.00	0.00	0.00	0.01	0.00	0.03	0.00	0.00	0.00	0.58	0.00
Max	28.08	24.61	5.80	6.22	12.61	12.56	9.09	7.24	3.14	2.79	43.79	41.15

Table 2: Chemical composition of the six factors from the $PMF_{Full\ MS}$ analysis, extent of neutralization, and F44 for each factor's organic fraction.

PMF Factor	Sub-micron PM Mass		AMS S	Neut _{Ext}	F44				
	(µg m ⁻³)	(%)	SO4 ²⁻	Org	NO ₃ ⁻	$\mathrm{NH_4}^+$	Cl		
Sulphate-OA	1.81	25	0.49	0.21	0.09	0.21	0.01	0.99	0.15
Nitrate-OA	3.19	45	0.06	0.16	0.58	0.20	0.00	1.04	0.12
Chloride	0.34	5	0.18	0.13	0.13	0.27	0.28	1.09	0.03
HOA	0.94	13	0.01	0.85	0.09	0.05	0.00	1.62	0.01
OOA	0.78	11	0.00	0.93	0.07	0.00	0.00	-	0.09
Amine	0.07	1	-	-	0.38	-	-	-	-





2 Figure 1: The location of the sampling site (MicMac Park) in Windsor, Ontario relative to

3 major industrial sources, namely coal fired power plants, steel mills, and potential large amine

4 sources. It can be seen that the measurement site was located close to the largest international

- 5 border crossing between the US and Canada (Huron Church road and the Ambassador6 Bridge).
- 7

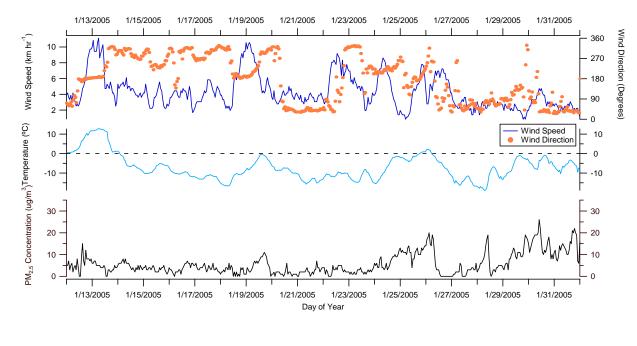


Figure 2: Meteorological conditions and $PM_{2.5}$ mass concentration from the CRUISER TEOM for the MicMac Park winter campaign. 6

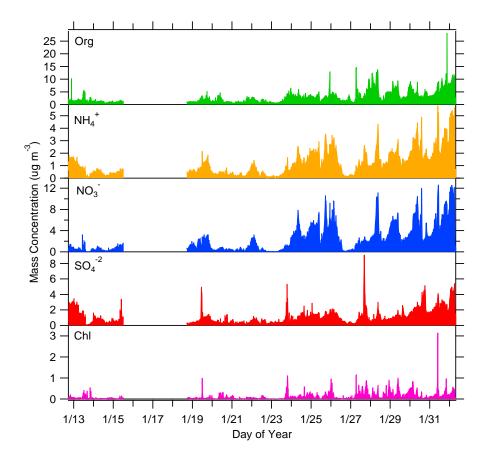


Figure 3: Time series of AMS-measured non-refractory sub-micron PM species.

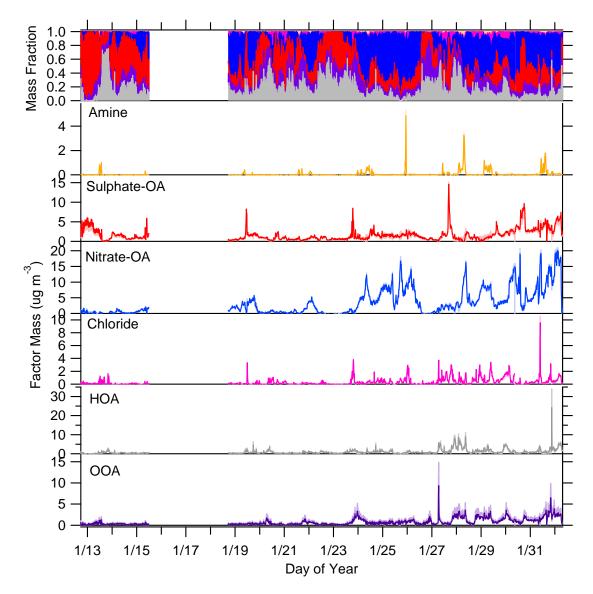




Figure 4: Time series of the factors from the six factor solution from $PMF_{Full MS}$ analysis. The solid line represents FPeak = 0, and the range of uncertainties through FPeak analysis (-10 6 and +10) is shown in the shaded regions.



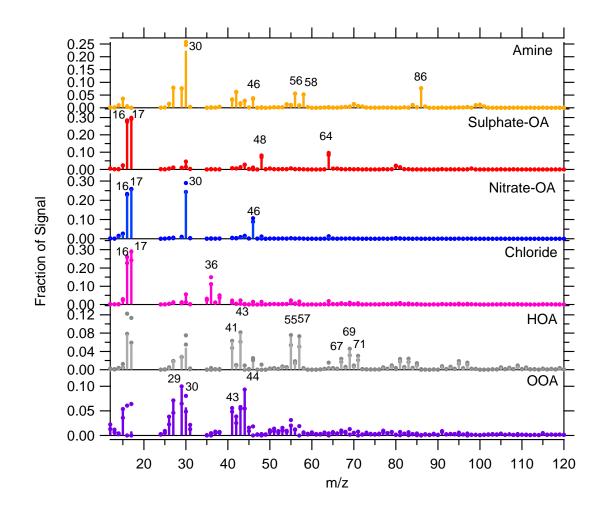
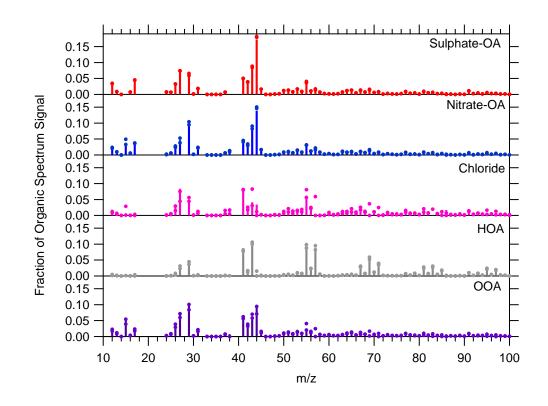


Figure 5: Mass spectra of factors from the six factor solution from PMF_{Full MS} analysis. Bars
represent the central, chosen rotation, and dots show the range in mass spectral variation from
FPeak rotations (-10 and 10).



1

2

4

5 Figure 6: Mass spectra of the organic fraction of the factors from the six factor solution from 6 $PMF_{Full MS}$ analysis. The Amine factor is not shown due to insufficient information regarding 7 the chemical nature of the Amine factor. Bars represent the central, chosen rotation, and dots 8 show the range in mass spectral variation from FPeak rotations (-10 and 10).

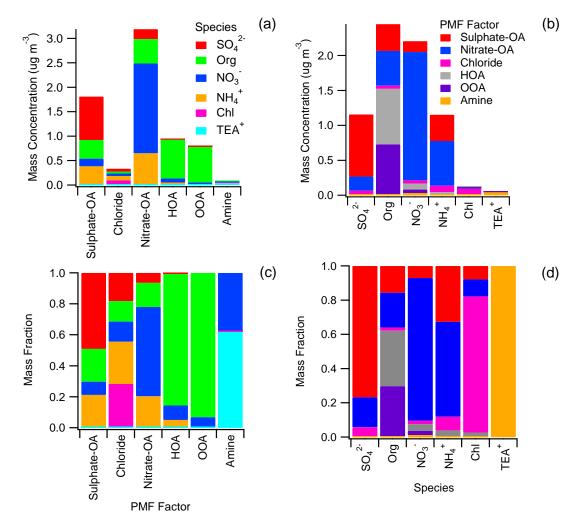
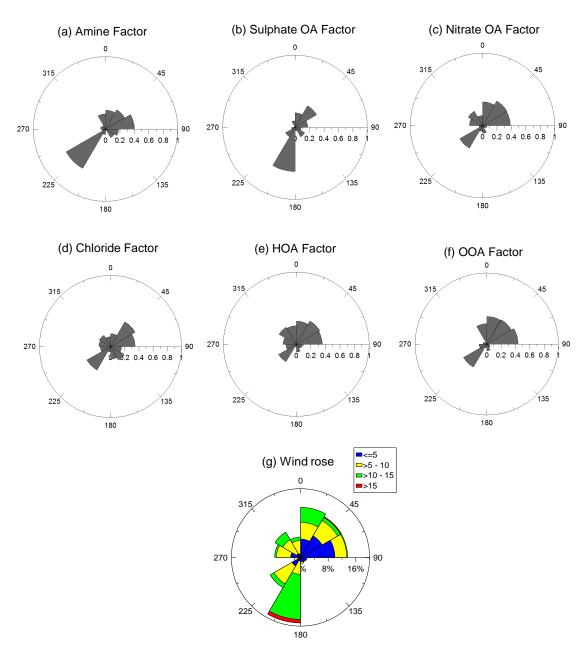




Figure 7: Chemical composition by factor and species components of the six PMF factors

3 from the PMF_{Full MS} analysis.



2

Figure 8: Conditional probability function (CPF) plots for the six factors from the six factor solution of the PMF_{Full MS} analysis, along with a wind rose plot (wind speed in m s⁻¹). The strongest wind dependence is observed for the Amine and Sulphate-OA factors, which show strong, yet slightly different directional associations to the southwest.

7

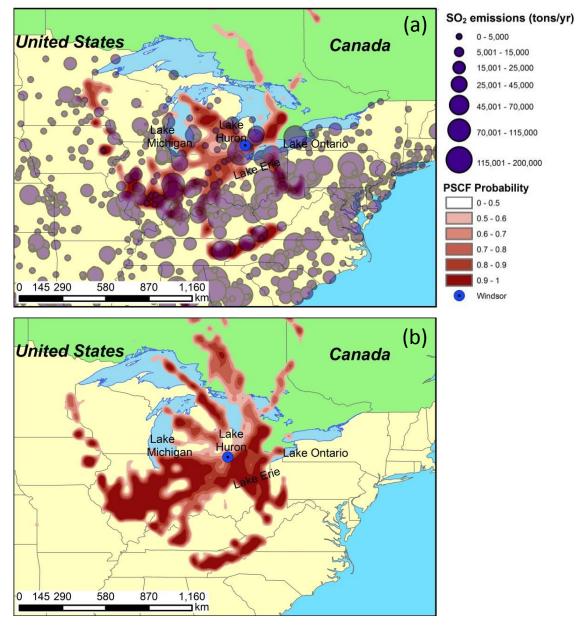


Figure 9: Potential source contribution function (PSCF) plot for the Sulphate-OA (a) and the Nitrate-OA factor (b) from the PMF_{Full MS} analysis. Both factors showed regional source influences, with the Sulphate-OA factor showing more prominent, distant influences to the south and southwest. The Nitrate-OA factor showed the most prominent influences to the southwest, over Indiana and Illinois. SO_2 emissions represent emissions from major coal powered energy generating facilities.

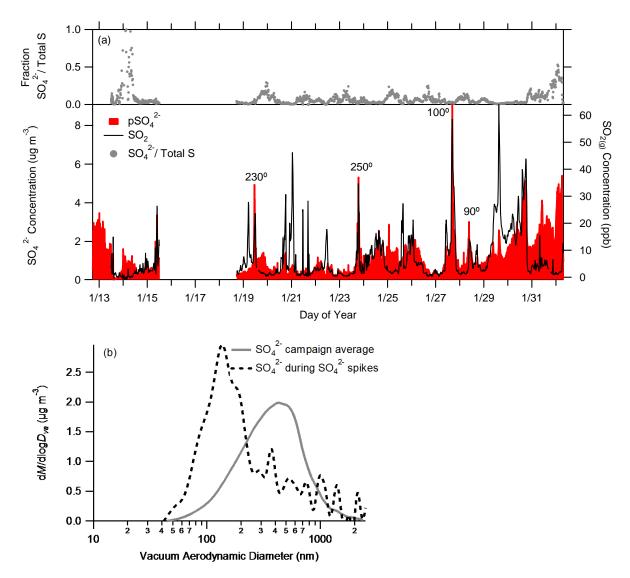
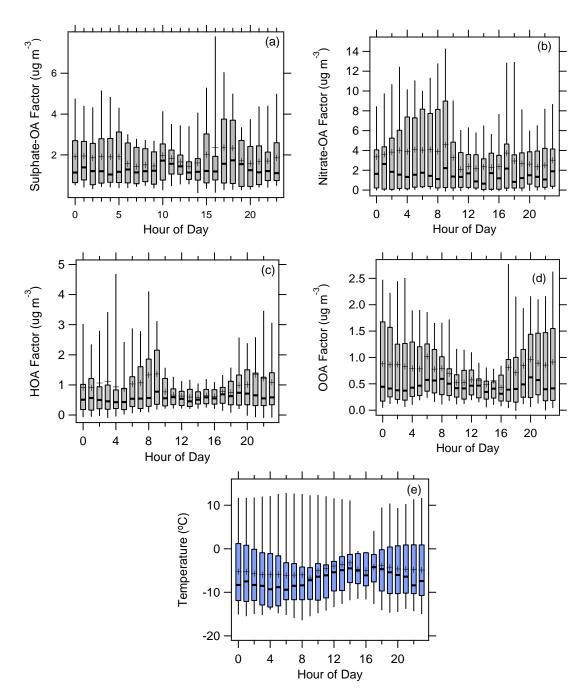


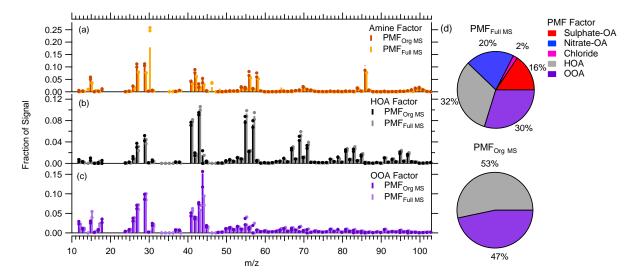


Figure 10: Particulate sulphate concentration $(SO_4^{2^-})$ vs. $SO_{2(g)}$ concentration, and fraction of SO_4^{2^-} / Total S (Total S = SO_4^{2^-} + SO_2) (a). Due to the winter conditions, the fractional contribution of $pSO_4^{2^-}$ to total S was typically very low. Shown in (b) are the averaged particle size distributions of AMS SO_4^{2^-} over the entire campaign, as well as during the extreme SO_4^{2^-} spikes. Particulate SO_4^{2^-} measured during the spikes appeared highly local due to their significantly smaller modal diameter.

- 8
- 9
- 10



2 Figure 11: Boxplots of diurnal trends for the Sulphate-OA (a), Nitrate-OA (b), HOA (c), and 3 OOA factors (d), from the PMF_{Full MS} analysis, along with temperature (e). Boxes indicate 4 interquartile ranges, horizontal lines indicate median hourly values, cross markers indicate hourly means, and whiskers represent the 5th and 95th percentiles. 5 The HOA factor 6 demonstrated a strong diurnal trend consistent with traffic patterns, while the OOA factor 7 demonstrated a trend more consistent with daytime lows and overnight highs. The Nitrate-8 OA factor showed a minor diurnal trend indicative of more regional contributions as 9 compared to OOA, while the Sulphate-OA factor showed minimal diurnal trend.



2

3 Figure 12: Comparison of factor profiles from the PMF_{Org MS} analysis to comparable factors 4 found by PMF of the full mass spectra (a, b, and c). The factors are normalized to the total 5 organic fraction signal, except for the Amine factor which is normalized to the total factor 6 signal. Only minor differences in mass spectra are noted for the HOA factor. The most 7 significant difference for the Amine factor is represented by the addition of the m/z 30 peak, 8 and for the OOA factor with the difference in magnitude of the m/z 44 peak. Dots show the 9 range in mass spectral variation from FPeak rotations (-10 and 10). Also shown in (d) is the 10 average factor composition of organics for both the PMF of the full mass spectrum and of the 11 organics only.

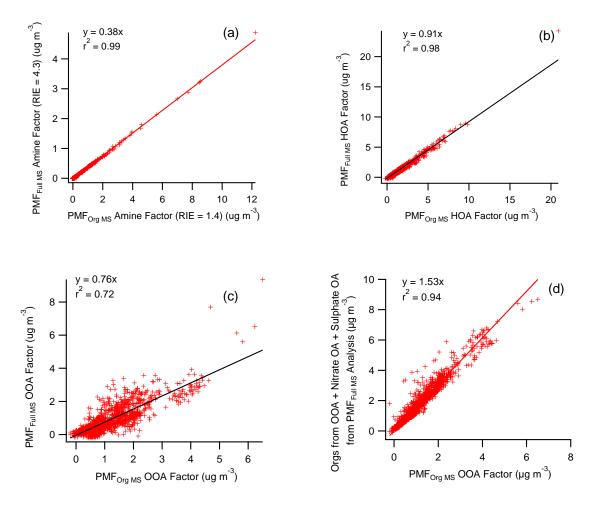


Figure 13: Scatter plots Amine (a), HOA (b), and OOA (c and d) factors comparing their mass
derived from PMF_{Full MS} vs. PMF_{Org MS} analyses. One observation (01/31/2005 21:00) was
removed from the HOA comparison to enhance the fit of the trendline.