1	Influence of Particle Phase State on the Hygroscopic
2	Behavior of Mixed Organic-Inorganic Aerosols
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15 Abstract

Recent work has demonstrated that organic and mixed organic-inorganic particles can 16 exhibit multiple phase states depending on their chemical composition and on ambient conditions 17 such as relative humidity (RH). To explore the extent to which water uptake varies with particle 18 phase behavior, hygroscopic growth factors (HGFs) of nine laboratory-generated, organic and 19 20 organic-inorganic aerosol systems with physical states ranging from well-mixed liquids, to phase-separated particles, to viscous liquids or semi-solids were measured with the Differential 21 Aerosol Sizing and Hygroscopicity Spectrometer Probe at RH values ranging from 40 - 90%. 22 Water-uptake measurements were accompanied by HGF and RH-dependent thermodynamic 23 24 equilibrium calculations using the Aerosol Inorganic-Organic Mixtures Functional groups Activity Coefficients (AIOMFAC) model. In addition, AIOMFAC-predicted growth curves are 25 compared to several simplified HGF modeling approaches: (1) representing particles as ideal, 26 well-mixed liquids, (2) forcing a single phase, but accounting for non-ideal interactions through 27 28 activity coefficient calculations, and (3) a Zdanovskii-Stokes-Robinson-like calculation in which

complete separation between the inorganic and organic components is assumed at all RH values, 1 2 with water-uptake treated separately in each of the individual phases. We observed variability in 3 the characteristics of measured hygroscopic growth curves across aerosol systems with differing phase behaviors, with growth curves approaching smoother, more continuous water uptake with 4 decreasing prevalence of liquid-liquid phase separation and increasing oxygen:carbon ratios of 5 the organic aerosol components. We also observed indirect evidence for the dehydration-induced 6 formation of highly viscous semi-solid phases and for kinetic limitations to the crystallization of 7 ammonium sulfate at low RH for sucrose-containing particles. AIOMFAC-predicted growth 8 curves are generally in good agreement with the HGF measurements. The performances of the 9 simplified modeling approaches, however, differ for particles with differing phase states. This 10 suggests that no single simplified modeling approach can be used to capture the water-uptake 11 behavior for the diversity of particle phase behavior expected in the atmosphere. Errors in HGFs 12 calculated with the simplified models are of sufficient magnitude to produce substantial error in 13 estimates of particle optical and radiative properties, particularly for the assumption that water 14 uptake is driven by absorptive equilibrium partitioning with ideal particle-phase mixing. 15

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17 **1 Introduction**

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19 Atmospheric aerosols alter the Earth's radiation budget, reduce visibility, and are associated with adverse health effects (Finlayson-Pitts and Pitts, 2000; Pöschl, 2005; Seinfeld 20 and Pandis, 2006). The magnitude of these impacts is influenced by aerosol water content, as this 21 is a major determinant of aerosol particle size. Further, aerosol water can impact gas-phase 22 23 photochemistry and secondary organic aerosol (SOA) concentrations by serving as a sink for reactive gases and as a medium for aqueous-phase and heterogeneous reactions (Ervens et al., 24 25 2011, 2013). As a result, clear understanding of the hygroscopicity of atmospheric aerosols is key to representing aerosol properties and behavior in atmospheric models and to improving our 26 27 understanding of their impacts on climate and air quality.

Organic aerosol (OA) comprises a substantial fraction of total aerosol mass (20 - 90%; Seinfeld and Pandis, 2006). Moreover, particle formation and transformation processes commonly lead to the formation of internally mixed organic-inorganic particles (Seinfeld and Pankow, 2003; Marcolli et al., 2004; Zhang et al., 2005; Murphy et al., 2006; Goldstein and

Galbally, 2007; Hallquist et al., 2009). Multiple studies have sought to elucidate the hygroscopic 1 2 properties of OA, as well as the influence of organic aerosol components on the hygroscopic 3 behavior and phase transitions of inorganic salts. Much of this work has focused on single- and multi-component aerosols comprised of carboxylic, dicarboxylic, and humic acids (e.g., Prenni et 4 al., 2001; Choi and Chan, 2002a; Brooks et al., 2004; Chan et al., 2006; Moore and Raymond, 5 2008; Hatch et al., 2009; Pope et al., 2010; Lei et al., 2014), as well as mixtures of organic acids 6 with inorganic salts (e.g., Cruz and Pandis, 2000; Choi and Chan, 2002b; Prenni et al., 2003; 7 Wise et al., 2003; Brooks et al., 2004; Svenningsson et al., 2006; Sjogren et al., 2007; Gao et al., 8 9 2008). Recent studies have explored water uptake by sugars, higher molecular weight organics, and polymers (Gysel et al., 2004; Mochida and Kawamura, 2004; Tong et al., 2011; Zobrist et 10 al., 2011; Lei et al., 2014; Xu et al., 2014). Such studies have aimed to characterize the 11 12 hygroscopicity of biomass burning aerosols, highly oxygenated, aged SOA, and oligomers. This body of research has demonstrated that the water-uptake behavior of OA components and their 13 influence on the phase transitions of inorganics depend on multiple factors, including the 14 composition and relative amounts of the organic and inorganic aerosol fractions, the 15 16 physiochemical properties of the organic components, and ambient conditions. Controlled laboratory studies have also served as a basis for the development and evaluation of 17 18 thermodynamic models (Clegg et al., 2001; Chan et al., 2005; Raatikainen and Laaksonen, 2005; Clegg and Seinfeld, 2006; Marcolli and Krieger, 2006; Svenningsson et al., 2006; Moore and 19 Raymond, 2008; Zardini et al., 2008; Zuend et al., 2011; Lei et al., 2014), with the aim of 20 representing water uptake by organic and mixed organic-inorganic aerosols. 21

Current regional and global chemical transport models include a simplified treatment of 22 aerosol hygroscopicity. In CMAQ (Community Multi-scale Air Quality model), for example, 23 only water uptake by the inorganic fraction is considered (Binkowski and Roselle, 2003; 24 25 http://www.epa.gov/AMD/Research/Air/aerosolModule.html). In WRF-Chem (Weather Research and Forecasting model coupled with Chemistry), aerosol hygroscopicity depends on 26 the aerosol module implemented. In MADE (Modal Aerosol Dynamics model for Europe) and 27 MOSAIC (Model for Simulating Aerosol Interactions and Chemistry), only the inorganic aerosol 28 components are considered in estimates of aerosol liquid water (Zaveri et al., 2008). In 29 GOCART (Goddard Chemistry Aerosol Radiation and Transport), however, the influence of 30 water uptake on the optical properties of OA is prescribed in tabulated, RH-dependent growth 31

factors (Koepke et al., 1997; http://acmg.seas.harvard.edu/geos). The Modal Aerosol Module 1 2 (MAM) assumes the hygroscopicity of primary and secondary organic aerosol can each be 3 described with a single parameter based on Köhler theory (κ ; Petters and Kreidenweis, 2007). Similarly, aerosol water associated with organic aerosol components is considered in GEOS-4 Chem (Goddard Earth Observing System coupled with chemistry) using a set of RH-dependent 5 HGFs for total particulate organic carbon (http://wiki.seas.harvard.edu/geos-6 chem/index.php/Aerosol optical properties). In the Advanced Particle Microphysics (APM) 7 module of GEOS-Chem, hygroscopicity is again parameterized based on Köhler theory (Yu et 8 al., 2012). When water uptake by both inorganic and organic components is taken into account in 9 these large-scale chemical transport models, total aerosol liquid water concentrations are 10 commonly calculated either assuming particles are externally mixed or using a volume- or mass-11 12 weighted average of the κ values or water contents of all simulated components (Drury et al., 2008; Fu et al., 2012). This simplified treatment of water uptake leads to substantial uncertainties 13 in predicted direct and indirect radiative forcing of atmospheric aerosols (Chung and Seinfeld, 14 2002; Kanakidou et al., 2005; Liu and Wang, 2010). Further, the assumption of external or 15 16 additive mixing of the aerosol components fails to account for the influence that particle mixing state and particle-phase morphology can have on aerosol hygroscopicity and optical properties. 17

Recent work has demonstrated that organic and mixed organic-inorganic particles can 18 19 exist in multiple phase states depending on their chemical composition and on ambient conditions such as RH and temperature (Cappa et al., 2008; Zobrist et al., 2008; Ciobanu et al., 20 2009; Virtanen et al., 2010; Bertram et al., 2011; Koop et al., 2011; Krieger et al., 2012). For 21 example, non-ideal interactions between aerosol components can result in a liquid-liquid phase 22 separation (LLPS) in which an inorganic-electrolyte-rich phase and an organic-rich phase co-23 exist within a single particle (Erdakos and Pankow, 2004; Ciobanu et al., 2009; Zuend et al., 24 2010; Bertram et al., 2011; Pöhlker et al., 2012; Song et al., 2012a; Zuend and Seinfeld, 2012; 25 You et al., 2012, 2013, 2014). Laboratory studies have demonstrated that ambient OA can exist 26 as a highly viscous liquid, semi-solid, or glass under atmospherically relevant conditions (Zobrist 27 et al., 2008, 2011; Mikhailov et al., 2009; Koop et al., 2011; Tong et al., 2011). Amorphous solid 28 (glassy) SOA has been observed both in a laboratory chamber and in the field (Virtanen et al., 29 2010; Saukko et al., 2012). Such complex phase behavior has major implications for the 30 partitioning of water and semi-volatile organic species to the particle phase (Ciobanu et al., 2009; 31

Koop et al., 2011; Krieger et al., 2012; Mikhailov et al., 2009; Bones et al., 2012; Song et al., 1 2 2012a, 2014; Zaveri et al., 2014). Diffusion coefficients in solid or semi-solid particles have been 3 estimated to be up to seven orders of magnitude smaller than in liquids (Vaden et al., 2011; Abbatt et al., 2012), resulting in the inhibition of mass transfer through the aerosol bulk and 4 delayed uptake and evaporation of water (Koop et al., 2011; Bones et al., 2012; Shiraiwa et al, 5 2011, 2013; Lienhard et al., 2014). Assuming that multi-component particles exist as well-6 mixed, single phase liquids when two separate phases are actually present can result in errors as 7 large as 200% in predicted particle mass formed through the partitioning of organic vapors to the 8 condensed phase (Zuend et al., 2010; Zuend and Seinfeld, 2012). The present study explores the 9 extent to which phase separations and other complex phase behavior influence the partitioning of 10 water vapor to the particle phase. 11

A variety of methods have been used to characterize the factors that influence the 12 prevalence of LLPS and amorphous solid OA. Coupling single-particle techniques with 13 microscopy has enabled the observation of particle phase transitions with changing ambient 14 15 conditions (Krieger et al., 2012). Song et al. (2012a), for example, evaluated the prevalence of LLPS as a function of RH and characterized the chemical composition of phases present within 16 mixed organic-inorganic aerosols using a high-speed video camera and Raman microscopy. 17 Moisture-induced glass transitions have been observed for sucrose solutions using single particle 18 19 techniques and differential scanning calorimetry (Zobrist et al., 2008, 2011). Similarly, phase states (solid/semi-solid versus liquid) of SOA as a function of RH have been inferred based on 20 the fraction of particles that bounced when impacted on a steel substrate (Sauuko et al., 2012). A 21 combination of bounce-fraction measurements and electron microscopy of newly formed OA in 22 a boreal forest provided the first evidence that SOA formed in the atmosphere can behave as 23 amorphous solids (Virtanen et al., 2010). These analyses have also shown that phase separation 24 and particle viscosity vary with chemical composition (e.g., the organic:inorganic mass ratio), 25 the molecular properties of the organic fraction of the aerosols (e.g., oxygen:carbon [O:C] ratio, 26 molar mass, hydrophilicity), and RH (Bertram et al., 2011; Song et al., 2012a, b; You et al., 27 2013). While these studies have provided valuable information regarding the influence of RH 28 and particle water content on particle phase state, investigations of the influence of complex 29 phase behavior on water uptake are limited. Further, the single-particle techniques commonly 30 used to study particle-phase morphology generally require particle sizes on the order of 1 - 1031

μm (Krieger et al., 2012). There is a need for the study of the properties and behavior of submicrometer particles with complex phase morphologies. Herein, we explore variability in the hygroscopic behavior of accumulation-mode particles with phase morphologies ranging from well-mixed liquids, to phase-separated particles, to amorphous solids or semi-solids and evaluate the importance of accounting for such complex phase states and related phase transitions when modeling water uptake.

7

8 2 Methods

9 2.1. Hygroscopic Growth Factor Measurements

Diameter hygroscopic growth factors of nine laboratory-generated aerosol systems that 10 serve as atmospheric aerosol surrogates were measured at RHs ranging from 40 to 90% with the 11 Differential Aerosol Sizing and Hygroscopicity Spectrometer Probe (DASH-SP). The DASH-SP 12 has been described in detail previously (Sorooshian et al., 2008). Briefly, in the DASH-SP, 13 aerosols are dried in a Nafion dryer, pass through a ²¹⁰Po neutralizer, and are then size-selected 14 with a differential mobility analyzer, resulting in a monodisperse aerosol population. The size-15 selected aerosols are split into four flows in which they are exposed to humidified air in parallel 16 Nafion humidifiers at four different, controlled RH values. Residence times in the Nafion dryer 17 and in each of the humidifiers are 1 s and 4 s, respectively. The aerosols were also sent through a 18 silica gel diffusion dryer with a residence time of $\sim 3 - 5$ s prior to entering the DASH-SP. The 19 size distributions of the particles following humidification are measured with four dedicated 20 optical particle counters (OPCs). A data processing algorithm that accounts for the dependence 21 of the OPC signal on particle diameter and refractive index has been developed to calculate wet 22 particle diameter (i.e., particle diameter after exposure to humidified air) from the OPC signal 23 24 (Sorooshian et al., 2008). This calculation requires knowledge of the refractive index of the dry particle. One of the flow regions of the DASH-SP is kept dry (RH < 8%) to enable calculation of 25 the effective refractive index using an empirical relationship between OPC signal height, 26 refractive index, and particle size. This empirical relationship was developed using calibration 27 28 salts with known refractive indices (Sorooshian et al., 2008). It is assumed that the particles are spherical and that they scatter, but not do absorb, light. Hygroscopic diameter growth factors are 29 calculated as the ratio of the mode diameter of the wet particle size distribution (i.e., at a 40% – 30 90% RH set point) to the dry mode particle diameter (HGF = D_{wet}/D_{dry}). In this work, HGFs were 31

measured for particles with dry mobility diameters of 250 nm. HGF measurements were performed at room temperature (~298 K). A minimum of 1500 particles were counted and sized at each RH value to produce the dry and wet size distributions. Experiments were repeated ten times for each aerosol system and reported growth factors were calculated as the average across those ten runs. The overall average uncertainty in DASH-SP-measured HGFs is approximately 5%, but can approach 8% at lower values of RH (Sorooshian et al., 2008).

7 With the aim of exploring the extent to which water uptake varies with particle phase 8 state, hygroscopic growth curves were measured for chemical systems for which particle phase 9 behavior as a function of RH has previously been characterized. For all systems studied here, RH-dependent phase states had previously been observed at room temperature. Most of those 10 studies were based on single-particle techniques, in which a single-particle (typically 11 supermicron-sized) is isolated in a controlled environment and probed with various optical and 12 13 microscopy techniques (Tong et al., 2011; Zobrist et al., 2011; Song et al., 2012a; You et al., 2013). Briefly, You et al. (2013) explored the prevalence of liquid-liquid phase separations as a 14 15 function of RH for internally mixed organic-inorganic aerosol systems comprising one organic compound and one inorganic salt. Single particles were deposited on a glass slide coated with a 16 hydrophobic substrate that was then mounted in a RH- and temperature-controlled flow cell. 17 Phase transitions during RH cycling were observed with an optical reflectance microscope. Song 18 19 et al. (2012a) studied the phase behavior of more complex particle mixtures, each consisting of three dicarboxylic acids with 5, 6, or 7 carbon atoms (C_5 , C_6 , C_7) and ammonium sulfate, 20 deposited on a glass slide using a Raman microscope equipped with a high-speed camera. 21 Zobrist et al. (2011) and Tong et al. (2011) studied the phase-state and water-uptake behavior of 22 sucrose particles using an electrodynamic balance coupled with various optical techniques and 23 optical tweezers coupled with brightfield imaging, respectively. 24

Seven aerosol systems with RH-dependent phase morphologies ranging from well-mixed liquids, to phase-separated particles, to amorphous solids or semi-solids were chosen for study in the present work from the above-described studies. Table 1 summarizes the compositions and phase behaviors of these chemical systems as determined in these prior studies. Two additional chemical systems consisting of sucrose and ammonium sulfate with varying organic:inorganic ratios were also studied here. A concurrent study (Robinson et al., 2014) also explored the hygroscopic behavior of submicron sucrose-ammonium sulfate particles, with the aim of characterizing the influence of glassy aerosol components on the optical properties of organicinorganic particles. Sucrose was selected as a model compound in that work and in the present study due to its high glass-transition temperature ($T_g = 331 - 335.7$ K), as characterized by Zobrist et al., (2008) and Dette et al. (2014). The phase behavior of these systems as a function of RH, however, has only been characterized for a sucrose:ammonium sulfate dry mass ratio of 2:1 (You and Bertram, 2015, Table1). The phase behavior of the 1:1 sucrose-ammonium sulfate system has not been characterized with single-particle techniques.

8 In addition to differences in phase behavior, the aerosol systems represent variations in 9 their complexity (in terms of dry composition). Particle compositions range from singlecomponent organic systems, to two-component systems consisting of one organic and 10 ammonium sulfate, to more complex systems consisting of dicarboxylic acid mixtures and 11 ammonium sulfate (Table 1). To evaluate the performance of the DASH-SP and the HGF-12 13 calculation algorithm, control runs were also performed for pure ammonium sulfate aerosols. For all chemical systems, aerosols were generated by atomizing aqueous solutions prepared by 14 15 dissolving the organic and inorganic components with the mass ratios given in Table 1 in Milli-Q water (resistivity $\geq 18.2 \text{ M}\Omega$ cm). Ammonium sulfate (purity $\geq 99\%$) and ACS reagent grade 16 sucrose were purchased from Macron Fine Chemicals; all other organic components were 17 purchased from Sigma-Aldrich (purity \geq 98%). 18

19 2.2 Modeling Hygroscopic Growth

Hygroscopic growth curves were modeled for the nine aerosol systems with the Aerosol 20 Inorganic-Organic Mixtures Functional groups Activity Coefficients (AIOMFAC) model (Zuend 21 et al., 2008, 2010, 2011; Appendix A). AIOMFAC is a thermodynamic model that explicitly 22 accounts for molecular interactions between all components in an aqueous solution through the 23 calculation of activity coefficients. A group-contribution concept based on UNIFAC (UNIversal 24 quasi-chemical Functional group Activity Coefficients), in which the thermodynamic properties 25 26 of organic compounds are determined on the basis of their molecular structures, is employed to account for interactions between organic functional groups, inorganic ions, and water (Zuend et 27 al., 2008, 2011). In these AIOMFAC-based equilibrium calculations, the potential existence of a 28 LLPS is determined and corresponding liquid-phase compositions are computed by the method 29 30 of Zuend and Seinfeld (2013). Diameter growth factors are calculated for RH values ranging

from near 0 to 99% for both dehydration (from high RH to low RH) and hydration (low RH to 1 2 high RH) cycles for all systems. For dehydration-branch calculations, growth factors at RH 3 values lower than the particle/inorganic salt deliquescence point are representative of metastable conditions, where a supersaturated solution is present with respect to the dissolved inorganic salt. 4 Hydration calculations include the presence of a solid inorganic phase at RH values before the 5 full deliquescence of ammonium sulfate and the partial dissolution of ammonium sulfate into the 6 aqueous organic solution (solid-liquid equilibrium; SLE). The hydration calculations are the 7 8 most relevant to the hygroscopic growth experiments, as the atomized aerosols were dried before being exposed to elevated humidities in the DASH-SP. AIOMFAC-calculated diameter growth 9 factors are compared to measured values to provide a detailed evaluation of the interactions 10 likely to be occurring between aerosol chemical components. 11

Following Zuend and Seinfeld (2012), in addition to the full AIOMFAC-based 12 13 equilibrium calculations, several simplified calculations were performed to explore the influence of phase separation and the effects of other non-ideal interactions on hygroscopic growth. These 14 15 comparisons also evaluate the need for accounting for such interactions in modeling the wateruptake behavior of atmospheric aerosols. Two hygroscopic growth calculations in which no 16 LLPS was allowed to occur were performed: (1) one in which non-ideal interactions are taken 17 into account through AIOMFAC-calculated activity coefficients and (2) one in which it is 18 19 assumed that the condensed phase behaves as an ideal mixture. Water-uptake calculations were also performed in a mode in which complete separation between an aqueous inorganic 20 electrolyte phase and an organic phase is assumed and, thus, no organic-ion interactions are 21 accounted for. This latter case is similar to a Zdanovskii-Stokes-Robinson (ZSR) relation 22 assumption, since the water uptake of individual components (here two separate phases) are 23 added up to estimate the HGF. In our ZSR-like calculation case, the influence of non-ideality on 24 25 water activity and, therefore water content, is accounted for within the individual phases with AIOMFAC-based activity coefficients. A calculation in which a solid organic phase is assumed 26 at all RH values was also performed for comparison against the DASH-SP measurements in 27 order to evaluate whether the presence of a solid organic phase was likely in any of the chemical 28 systems studied. All growth-curve calculations were performed at 298 K. 29

30 Note that AIOMFAC predicts phase compositions, which can be used to derive mass 31 growth factors, but not the densities of the phases necessary to calculate diameter growth factors

(assuming spherical particles). Solid and liquid-state densities of ammonium sulfate were taken 1 2 from Clegg and Wexler (2011). The density of citric acid was calculated based on the data and 3 parameterizations of Lienhard et al. (2012). The densities of all other organic compounds were estimated using the structure-based method of Girolami (1994) with the online density-4 calculation tools available on the E-AIM website (http://www.aim.env.uea.ac.uk/aim/aim.php). 5 Densities of mixtures within particles and those of individual phases were calculated assuming 6 additive molar volumes of each component (i.e., ideal mixing in terms of density). Organic 7 8 component densities are given in Table 1.

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10 **3 Results and Discussion**

3.1 Experimental Hygroscopic Growth Curves and AIOMFAC Modeling

3.1.1 Two-Component Carboxylic Acid-Ammonium Sulfate Systems

Measured and AIOMFAC-predicted HGFs for the two-component carboxylic acid-13 14 ammonium sulfate systems (organic:inorganic dry mass ratios = 2:1) are shown in Fig. 1. Panels below the hygroscopic growth curves illustrate AIOMFAC predictions of the occurrence of 15 16 phase separation and the composition of each phase. All of the two-component carboxylic acidammonium sulfate systems demonstrate suppressed water uptake at high RH values as compared 17 18 to pure ammonium sulfate (Fig. 1a - 1c; Fig. B1). This is expected because, based on approximately linear additivity in water uptake by the different particle components, a lesser 19 20 degree of water uptake by the organic fraction in comparison to inorganic salts (per unit mass) is typically the case. At RH = 80%, ratios of the HGFs for the mixed organic-inorganic aerosols to 21 22 those for pure ammonium sulfate were 0.88, 0.86, and 0.89 for the systems containing diethylmalonic acid, 2-methylglutaric acid, and citric acid, respectively. Similarly, these values 23 were 0.90, 0.88, and 0.88 at RH = 90%. At low to moderate RH values (40 - 70%), however, the 24 shapes and characteristics of the growth curves vary across the systems due, at least in part, to 25 differences in phase behavior, as is discussed below. 26

Mixed diethylmalonic acid-ammonium sulfate particles demonstrated little to no water uptake (HGF \approx 1.0) before the particle deliquesced at an RH of approximately 80% (Fig. 1a). This suggests the presence of a viscous semi-solid, an amorphous solid, or a crystalline solid

organic phase (with no water uptake on the timescale of the 4 s residence time in the humidifier). 1 2 The presence of a solid organic at low to moderate RH is supported by a comparison between 3 measured and AIOMFAC-calculated growth factors. For RH = 40 - 70%, model-measurement agreement is closest for the case in which diethylmalonic acid is assumed to be solid. Measured 4 values then approach the full equilibrium AIOMFAC calculation in which the organic is 5 assumed to be in a liquid state. Note that all AIOMFAC-based calculations shown by the solid 6 curves assume that the organic compounds are in a liquid-like state at all RH levels. For this 7 system, AIOMFAC predicts a SLE state, with diethylmalonic acid dominating the liquid phase 8 and with the solid phase fully comprised of ammonium sulfate. Following the deliquescence of 9 ammonium sulfate at RH = 79.75%, the particle is predicted to undergo LLPS. The organic-10 dominated and the ammonium sulfate-dominated phases are then predicted to merge to a single 11 phase at RH = 94%. The AIOMFAC-predicted RH at which merging of the two phases occurs is 12 5% higher than that observed by You et al. (2013). It is likely that this small discrepancy can be 13 attributed to uncertainty in modeled organic-ion interactions. However, it is also possible that 14 optical methods cannot discern the presence of two phases at $RH \ge 89\%$ if one of the two 15 16 predicted liquid phases is small in mass or volume compared to the other phase.

In addition, You et al. (2013) performed dehydration experiments to explore the onset of 17 phase separation (from high to low RH), while the present study focused on hydration 18 experiments (from low to high RH). AIOMFAC predictions are made under the assumption that 19 there is no hysteresis between LLPS and the merging of two liquid phases to a single phase in 20 terms of the onset RH. This is because experiments (e.g., Song et al., 2012a) show that there is 21 little to no hysteresis in such a phase transition, at least for systems with liquid-like viscosities. In 22 contrast to the typical hysteresis behavior of liquid-crystal/crystal-liquid phase transitions (i.e., 23 deliquescence vs. crystallization), liquid-liquid to single-liquid phase transitions involve only 24 disordered phase states (rather than crystalline solids with long-range order). Exceptions to this 25 may exist for some systems in a particular composition range involving the metastable region of 26 a liquid-liquid equilibrium phase diagram (e.g. Zuend et al., 2010). However, the energy barrier 27 for the nucleation and growth of a new liquid phase is small in comparison to the larger energy 28 barrier that needs to be overcome when a new crystalline phase is formed. Because the merging 29 of the phases is predicted to occur at an RH at which aqueous diethylmalonic acid is expected to 30

be of low viscosity (Figure 1a), no hysteresis behavior is expected. Thus, we do not expect that a
 hysteresis behavior influenced the disagreement in the RH of phase merging discussed above.

Like the diethylmalonic acid-ammonium sulfate system, the 2-methylglutaric-containing 3 aerosols demonstrate a marked increase in water uptake at RH = 80%; however, more gradual 4 and continuous water-uptake was observed prior to particle deliquescence (Fig. 1b). Good 5 agreement was achieved between measured and AIOMFAC-predicted growth curves at all RH 6 values. AIOMFAC again predicts a SLE state prior to the deliquescence of ammonium sulfate at 7 8 RH = 79.25%. This is followed by the presence of two liquid phases before full merging of the phases into a single liquid phase at RH = 82.25%. As was true for the diethylmalonic acid-9 ammonium sulfate system, this LLPS phase transition value is higher than the RH of 75% 10 observed by You et al. (2013). As noted above, this could be related to uncertainty associated 11 with the parameterization of the AIOMFAC group-contribution method. In addition, with the 12 13 gradual merging of the phases, it is possible that a remaining LLPS was not visible in the experiments at RH > 75%. Again, as is explained above, no hysteresis is expected for the phase 14 15 separation and merging for hydration and dehydration conditions for this system.

16 For the citric acid-ammonium sulfate particles, measured growth factors suggest gradual, 17 continuous water uptake at all experimental RH values (Fig. 1c). This is in agreement with measurements for this system at different dry organic:inorganic ratios reported by Zardini et al., 18 (2008). The AIOMFAC-predicted growth curve indicates a small increase in slope with the 19 deliquescence of ammonium sulfate at RH = 70.25%, lower than that for pure ammonium 20 21 sulfate. This is expected in the presence of a hygroscopic organic compound like citric acid (O:C = 1.17) due to the partial solubility of ammonium sulfate in the aqueous, citric acid-rich phase. 22 Measured and predicted growth factors are in very good agreement at all RH values. For this 23 system, a single, well-mixed liquid phase is predicted following the complete deliquescence of 24 ammonium sulfate. This is in agreement with the experimental results of You et al. (2013) and 25 Bertram et al. (2011), who observed no evidence of LLPS for citric acid-ammonium sulfate 26 particles. 27

A comparison of hygroscopic growth curves across the three two-component systems suggests that growth curve shape reflects phase behavior in mixed organic-ammonium sulfate systems. Growth curves approach smoother, more continuous water uptake with decreasing

prevalence of LLPS (i.e., looking from left to right for Fig. 1, panels a - c). In addition to higher 1 2 miscibility of the aerosol components for particles with no LLPS or those for which LLPS 3 persists for only a small range of RH values, this can likely be attributed to the fact that both the prevalence of LLPS and aerosol hygroscopicity vary with the O:C ratios of the organic 4 components of mixed organic-inorganic particles (Massoli et al., 2010; Bertram et al., 2011; 5 Duplissy et al., 2011; Song et al., 2012a; You et al., 2013). Note that the O:C ratios of 6 diethylmalonic, 2-methylglutaric, and citric acid are 0.57, 0.67, and 1.17, respectively. Thus, the 7 smoothing of the growth curves with increasing O:C (and, incidentally, lower prevalence of 8 LLPS) is consistent with the higher propensity of the more polar compounds to take up water and 9 dissolve some ammonium sulfate (in a SLE) at lower RH. 10

3.1.2 Complex Mixed Dicarboxylic Acid-Ammonium Sulfate Systems

12 In general, water-uptake behavior was similar across the more complex dicarboxylic acid-ammonium sulfate mixtures, regardless of the carbon number of the acids included in the 13 system and the differences in the phase morphologies of the particles observed previously in 14 experiments by Song et al., (2012a). All systems demonstrated gradual water uptake with growth 15 factors increasing from 1.0 to \sim 1.05 prior to a rapid increase in particle diameter of \sim 30 - 45% at 16 17 $RH \approx 80\%$ (Fig. 2). This is in agreement with the results of Song et al. (2012a), who observed nearly complete or complete particle deliquescence at RH values of 79 - 80% for all three 18 systems. At $RH \ge 80\%$, suppressed hygroscopic growth as compared to ammonium sulfate is 19 evident, but to a lesser degree than for the two-component systems discussed above, with 20 21 reductions in diameter growth factors ranging from 1% to 7%. This is likely attributable to the smaller organic: inorganic ratios of these systems as compared to the two-component systems 22 (dry mass ratios of 1:1 rather than 2:1). In addition, mixing effects due to the presence of 23 multiple organics, which can inhibit crystallization (Cruz and Pandis, 2000, Marcolli et al., 24 2004), likely influenced the water uptake behavior of these systems. One constituent of the C_7 25 acid mixture, diethylmalonic acid, demonstrated behavior indicative of the presence of a 26 crystalline solid or highly viscous organic phase in the two-component mixture discussed above. 27 In the more complex C_7 acid mixture, however, it is possible that there is limited, but more 28 29 gradual water uptake at low to moderate RH, suggesting a liquid (perhaps moderately viscous) 30 phase. Owing to their complexity, the multi-acid aerosol systems may provide a better analog for the majority of mixed organic-inorganic atmospheric aerosols than the two-component systems. 31

As was true for the two-component carboxylic acid-ammonium sulfate systems, we observed a smoothing of the growth curves with increasing O:C ratio of the C₅, C₆, and C₇ dicarboxylic acid mixtures (O:C = 0.80, 0.67, 0.57, respectively).

As determined both through previous experimental work and with the AIOMFAC 4 modeling, there is variability in phase behavior across the mixed dicarboxylic acid-ammonium 5 sulfate aerosol systems. For the C₅ dicarboxylic acid-ammonium sulfate system, AIOMFAC 6 predicts a SLE state of ammonium sulfate prior to complete deliquescence at RH = 77% (Fig. 7 8 2a). At this point, ammonium sulfate fully dissolves into the organic-dominated liquid phase. For the C₆ and C₇ systems, however, AIOMFAC predicts LLPS following the complete 9 deliquescence of ammonium sulfate at RH = 79.25 and 79.75%, respectively (Figs. 2b, 2c). For 10 the C₆ system, the inorganic- and organic-dominated phases are then predicted to merge to a 11 single phase at RH = 83%, while LLPS persists up to RH = 92.75% for the C_7 system. 12 AIOMFAC predictions of phase behavior are in qualitative agreement with the observations of 13 Song et al. (2012a). In agreement with AIOMFAC predictions of particle phase state during the 14 15 humidification process (i.e., starting from dry conditions), no evidence of LLPS was observed for the C₅ system (Song et al., 2012a). Contrary to the experiments reported by Song et al. 16 (2012a), AIOMFAC does predict LLPS at RH values below 80% for the dehydration process of 17 the C_5 system (i.e., when no solid ammonium sulfate phase is allowed to form). For both the C_6 18 and C₇ systems, AIOMFAC overpredicts the RH at which separated phases merge to a single 19 phase by 10% and 3%, respectively, in comparison to the observations by Song et al. (2012a). 20 Song et al. (2012a) also observed differences in the morphology of the LLPS states of the C₆ and 21 C₇ systems. For the C₆ system, an ammonium sulfate core was partially engulfed by an organic-22 23 dominated outer phase containing ammonium sulfate satellite inclusions. The C₇ system displayed a distinct, fully engulfed core-shell morphology. Despite the small degree of model-24 experiment disagreement in the RHs at which merging of two liquid phases occur in the C₆ and 25 C₇ systems, good agreement between measured and AIOMFAC-predicted growth curves is 26 achieved for the C₅, C₆, and C₇ systems (Figures 2a - c). This suggests that the presence or 27 absence of LLPS, as well as the morphology of LLPS (which is not taken into account in 28 AIOMFAC), might influence water-uptake to a lesser degree than other thermodynamic 29 properties. This is explored subsequently when we compare various modeling methods that vary 30 in the degree to which they consider thermodynamic non-ideality and LLPS. 31

1 3.1.3 Sucrose-Containing Aerosols

The pure sucrose aerosols demonstrate continuous, but limited, water uptake with growth 2 factors reaching only 1.25 at RH = 90% (Figure 3a). Consequently, as RH approaches 100%, the 3 growth factor is expected to show a steep increase. Such a steep increase is predicted by the 4 AIOMFAC-based equilibrium model at RH \approx 95%. For RH \leq 50%, average measured values of 5 HGF fall below the equilibrium calculations of AIOMFAC; however, this deviation is within 6 experimental uncertainty. Previous studies demonstrated that the glass transition RH of sucrose 7 8 droplets at room temperature (~ 298 K) ranges between 25 and 53%, depending on the rate of dehumidification (Zobrist et al., 2011; Tong et al., 2011). Thus, it is possible that a glass 9 transition occurred with the rapid drying of the atomized sucrose droplets (total residence time in 10 the silica gel diffusion and Nafion dryers ~ 5 s). However, no quantitative conclusions can be 11 drawn about the extent to which this affected HGFs in our measurements. Nevertheless, such a 12 13 glass transition may have influenced the growth factor measurement at RH < 53 % and could partially explain the observed discrepancy between the measured and modeled growth factors at 14 15 those RH levels. Measured HGFs are in agreement with equilibrium-condition measurements and modeling of Tong et al. (2011) and Zobrist et al (2011), as well as AIOMFAC-modeled 16 HGFs, particularly at $RH \ge 60\%$. 17

Like pure sucrose, the sucrose-ammonium sulfate systems demonstrate continuous water-18 uptake behavior, with smaller HGFs as compared to pure ammonium sulfate at RH = 80 and 19 90% (Fig. 3b,c). The magnitude of the deviations in HGF for the sucrose-ammonium sulfate 20 21 particles from those for pure ammonium sulfate varies with the organic:inorganic ratio, but is within the range observed for the other aerosol systems. At RH = 80%, ratios of the HGFs for the 22 sucrose-ammonium sulfate particles to those for pure ammonium sulfate were 0.86 and 0.91 for 23 the 2:1 and 1:1 sucrose-ammonium sulfate systems, respectively. These values were 0.85 and 24 0.89 at RH = 90%. 25

Evidence for the influence of a highly viscous phase state on hygroscopic behavior is stronger for the mixed sucrose-ammonium sulfate aerosols. While for stable thermodynamic equilibrium AIOMFAC predicts distinct deliquescence behavior at RH = 80% for both organic:inorganic ratios, we observed smooth, continuous water uptake for RH = 40 - 90% in the DASH-SP experiments (Figure 3b, c). Note that before entering the DASH-SP, particles were

dried after atomization from a solution, at which point ammonium sulfate crystallization may 1 2 occur. Since the viscosity of the solution is expected to increase considerably toward lower RH, 3 however, the high viscosity (potentially accompanied by a glass transition) may suppress the crystallization of ammonium sulfate in these systems. The corresponding AIOMFAC 4 calculations for the humidification process are shown by the solid curves in Fig. 3 (where 5 ammonium sulfate is in a crystalline state at low RH), while the dashed curves show the model 6 predictions of HGF for the case where both sucrose and ammonium sulfate are present in a liquid 7 (potentially viscous) solution. This latter scenario is often observed experimentally during a 8 dehydration process (starting at very high RH with a homogenous liquid solution). The measured 9 HGFs are in good agreement with the metastable conditions of the AIOMFAC-modeled 10 "dehydration branch." This is in agreement with the results of Robinson et al. (2014), who also 11 12 observed continuous water uptake and evidence for the inhibition of the crystallization of ammonium sulfate in the presence of viscous sucrose. Minor differences ($\leq 10\%$) in the HGFs 13 reported in their work (RH = 60 - 80%) and those described here can likely be attributed to 14 differences in technique, as they derived growth factor based on light extinction using Mie 15 16 theory, or variances in particle drying and/or humidification times (humidification residence time = 1 s versus 4 s in this work). Our results suggest that rapid particle drying within HGF 17 18 instrumentation can induce a steep increase in particle viscosity, potentially even leading to a glass transition. As a result, such measurements may not accurately reflect the equilibrium water-19 20 uptake behavior of viscous particles. This has implications for the measurement of HGFs for ambient aerosol such as the highly oxygenated SOA for which sucrose serves as a surrogate in 21 these experiments. 22

The formation of a highly viscous liquid or semi-solid phase may also lead to kinetic 23 limitations, affecting the loss of water by evaporation during the drying process prior to 24 25 humidification in the DASH-SP. There is increasing evidence from laboratory and field studies that viscous liquid or semi-solid SOA components may be present in atmospheric aerosol (e.g., 26 Virtanen et al., 2010; Vaden et al., 2011; Saukko et al. 2012; Renbaum-Wolff et al. 2013). Thus, 27 accounting for kinetic limitations to water uptake and release is crucial to accurately modeling 28 the dynamic hygroscopic behavior of SOA. However, the good agreement between measured 29 HGFs and the AIOMFAC-based dehydration-branch equilibrium calculations indicates that 30 water loss was not substantially inhibited during particle drying. If a glassy sucrose shell had 31

formed in these particles and this shell was of sufficient thickness to inhibit water evaporation 1 2 during the ~ 5 s residence time in the dryers, the measured effective "dry" reference diameter 3 would have been affected by the higher water content. This oversized shell would, in turn, affect all the experimental HGFs obtained at higher RH due to the normalization by the "dry" particle 4 diameter. If this were the case, HGFs from the experimental data would likely be lower and 5 discrepancies from the AIOMFAC-based calculations of the dehydration growth curve would be 6 larger. Another possible factor contributing to model-measurement disagreement for the sucrose-7 ammonium sulfate systems is the lower confidence in the treatment of ether groups by 8 AIOMFAC as compared to functional groups present in the other systems studied here. This is 9 due to limited experimental data for the development of parameterizations of ether group-ion 10 interactions (Zuend et al., 2011). This might also contribute to uncertainty in the predicted phase 11 12 behavior of the sucrose-ammonium sulfate aerosols, which indicates a SLE prior to the complete deliquescence of ammonium sulfate at RH = 80%. It is predicted that LLPS then persists up to 13 RH = 94.25% for both organic:inorganic ratios explored here (Fig. 3b, c). The predicted 14 occurrence of LLPS for sucrose (O:C = 0.92) contrasts with previous studies that suggest that 15 16 LLPS is unlikely for organic compounds with O:C > 0.7 - 0.8 (Bertram et al., 2011; Song et al., 2012b; You et al., 2013), as well as the experimental results of You and Bertram (2015), in 17 18 which no LLPS was observed for sucrose-ammonium sulfate particles with an organic to inorganic dry mass ratio of 2:1. 19

20 3.1.4 Measurement and Modeling Limitations and Uncertainty

21 As noted above, data availability limitations and variability in experimental conditions for the datasets used in developing parameterizations of functional group-ion interactions in 22 AIOMFAC contribute to uncertainties in ether-ion interactions and other functional-group-ion 23 interactions. In addition, other measurement and modeling limitations can contribute to 24 uncertainty in measured and predicted HGFs. Because experimental data regarding the densities 25 of organic compounds are limited, the densities of many of the organic compounds studied in 26 this work were estimated with a group-contribution method (Girolami et al., 1994). Further, for 27 all aerosol systems and all phases present within the particles, it was assumed that the molar 28 volumes of the aerosol components are additive (i.e., ideal mixing in terms of volume and 29 30 density contributions), regardless of the thermodynamic properties of the mixture under 31 consideration.

Complex particle-phase morphologies can also present unique sources of error and 1 2 uncertainty in HGF measurement methods that use optical methods, such as the DASH-SP. For 3 example, the algorithm that calculates diameter growth factor assumes that the refractive index of the non-water aerosol components is constant at the value measured in the dry DASH-SP 4 channel. For systems with non-uniform surfaces (e.g., the C₆ dicarboxylic acid-ammonium 5 sulfate system, which has been shown to have an organic outer shell with satellite ammonium 6 sulfate inclusions or partially engulfed morphology; Song et al., 2012a), this assumption might 7 not be accurate. In addition, the effective refractive index of the organic and inorganic particle 8 components might evolve upon humidification as previously separated phases merge. This 9 behavior has the potential to influence hygroscopicity measurements for systems that have core-10 shell morphologies at low RH to moderate RH that then transition to a single liquid-phase state 11 with increasing water content and might contribute to the greater model-measurement 12 disagreement at RH = 80% observed for the C_7 dicarboxylic acid-ammonium sulfate system. 13 However, the overall agreement achieved between measured HGFs and AIOMFAC-calculated 14 values in this work suggests that these factors, and their corresponding uncertainties, are not a 15 16 substantial issue for the measurement of HGFs for the aerosol compositions and particle size range studied here. 17

In addition, as noted above, the DASH-SP HGF experiments were performed with 18 19 particles much smaller than those used in the microscopy or electrodynamic balance experiments that had previously been used to directly characterize LLPS and glass transitions in the particle 20 systems studied. Thus, the phase behavior of the particles studied here was not characterized 21 directly. There is some limited evidence that the prevalence of LLPS can vary with particle size. 22 Veghte et al., (2013) observed that LLPS did occur in larger particles comprised of ammonium 23 sulfate and succinic acid or pimelic acid (diameters ≥ 170 and 270 nm, respectively), but that 24 LLPS was not evident in smaller particles with the same composition and organic:inorganic mass 25 ratios. While the DASH-SP experiments do not directly reveal whether a phase separation is 26 present in the particles, the observed hygroscopic growth in comparison to the model 27 calculations is in agreement with such phase behavior in the corresponding LLPS RH ranges. 28 Finally, the timescale of humidification in our experiments differs from that in the single-particle 29 studies. For example, in the mixed dicarboxylic acid-ammonium sulfate single-particle 30 experiments, RH was changed at a rate of 0.14 - 0.34% min⁻¹ (Song et al., 2012). The extent to 31

which the timescale of humidification influences the occurrence of phase separation is unknown;
however, again, agreement between the measured hygroscopic growth and those calculated
under equilibrium conditions suggests humidification timescale did not have a substantial impact
on phase behavior.

5 3.2 Evaluation of Simplified Thermodynamic Assumptions

6 For the multicomponent systems for which we expect that observed water-uptake 7 behavior is governed by thermodynamic equilibrium conditions (i.e., excluding the sucrose-8 containing systems, which display evidence of kinetic limitations to the crystallization of ammonium sulfate), we compared the rigorous thermodynamic modeling of the AIOMFAC-9 10 based equilibrium HGF predictions ("AIOMFAC, equilibrium" in Figure 4) to that based on several simplified thermodynamic assumptions: (1) representing particles as ideal, well-mixed 11 12 liquids ("Ideal - well-mixed liquid"), (2) forcing a single liquid phase following the deliquescence of ammonium sulfate, but accounting for non-ideal interactions through activity 13 14 coefficient calculations and allowing for a SLE of ammonium sulfate ("No LLPS – non-ideal"), and (3) a ZSR-like calculation in which complete separation between the inorganic and organic 15 16 components is assumed at all RH levels ("Complete Phase Separation (ZSR)"). Water is the only 17 component allowed to partition to both phases in this ZSR-like calculation case. In all of these simplified calculation cases, the formation of a solid ammonium sulfate phase is allowed to exist 18 (below its deliquescence point at the given temperature), except for the single-phase, ideal 19 mixture case. We evaluate the extent to which these simple, relatively computationally 20 21 inexpensive modeling approaches capture the hygroscopic behavior of particles with varied and complex phase states. Note that for the diethylmalonic acid-ammonium sulfate system, we first 22 focus on the calculation for which the presence of a solid organic is predicted prior to particle 23 deliquescence ("Solid Organic" in Figure 4). At $RH \ge 80\%$, we then consider the full AIOMFAC 24 equilibrium calculation in our discussion of model error, as this offered the best agreement 25 between measured and modeled values (Figure 1a). For all other systems, we focus on the full 26 AIOMFAC-based equilibrium calculations at all RH values. 27

Figure 4 shows a comparison between the AIOMFAC-predicted HGFs for the hydrationbranch of a humidity cycle and those calculated with the simplified modeling approaches. The performance of each of the simplified modeling approaches differs for particles depending on

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phase state. This suggests that a single simplified modeling approach cannot be used to capture 1 2 the water-uptake behavior for the diversity of particle-phase behaviors expected in the 3 atmosphere. For all systems, except the citric acid-ammonium sulfate particles, the assumption that the particles could be represented as thermodynamically ideal liquid mixtures leads to the 4 greatest deviation from AIOMFAC-predicted growth curves. We also note that such 5 discrepancies depend on the organic: inorganic dry state mass ratios. Generally, the smaller the 6 ammonium sulfate mass fraction, the lower the degree of hysteresis behavior of 7 8 hydration/dehydration processes. Overpredictions of diameter growth factors increase from those at low RH to a maximum just prior to the rapid increase in water uptake associated with 9 deliquescence of ammonium sulfate, then drop to within $\sim 5\%$ of AIOMFAC values (Fig. 4). 10 Maximum errors for the assumption of ideality ranged from 10 - 26% and increased with 11 12 increased prevalence of LLPS (i.e., looking from right to left from Figure 4 panels a - c and from left to right from panels d - f). For the citric acid-ammonium sulfate system, which does not 13 undergo LLPS, ideal and AIOMFAC calculated growth curves are in good agreement for RH \leq 14 80%. Error then increases slightly to \sim 5%. 15

Not surprisingly, the forced single-phase calculations, in which non-ideal interactions are 16 taken into account but LLPS is not, perform well for the C₅ dicarboxylic-ammonium sulfate and 17 citric acid-ammonium sulfate systems, which demonstrated no LLPS (Figure 4c, d). For these 18 systems, the ZSR-like calculation leads to largest disagreement, with maximum errors of 19 between 10 and 18% occurring between RH values of 60 and 80%. This can be attributed to 20 delayed particle deliquescence in the ZSR-like calculations, since the partial dissolution of 21 ammonium sulfate in the aqueous organic phase is not considered in this case. Similarly, for 22 systems for which LLPS persists for only a small range of RH values, the single-phase 23 assumption performs well, with errors never exceeding 5% (e.g., the 2-methylglutaric-24 ammonium sulfate and C₆ dicarboxylic acid-ammonium sulfate systems; Figure 4b, d). The 25 single-phase assumption also performs well for the systems for which LLPS does occur over a 26 wider RH range, but delayed prediction of particle deliquescence leads to a large spike in 27 underprediction of particle diameter at RH \approx 80%. The good performance of the single-phase 28 assumption at most RH values suggests that accounting for LLPS in calculations of hygroscopic 29 growth might be less important than accounting for other non-ideal interactions between 30 condensed-phase components. However, this result is likely not applicable to mixed organic-31

inorganic systems with less water-soluble (i.e., less hygroscopic) organic components, where
LLPS is more prevalent and water uptake by the organic fraction is limited (Zuend and Seinfeld,
2012).

4 For systems that do undergo LLPS, the ZSR-like calculation also performs relatively well across the range of RH values studied, but displays discrepancies to the AIOMFAC-based 5 equilibrium prediction of 12 - 18% at moderate RH values due to deviations in the predicted 6 SLE and deliquescence transition of ammonium sulfate. For these systems, the ZSR-like 7 8 calculation also underpredicts water uptake at RH values above the point at which separated 9 liquid phases merged to a single phase, with relative deviations approaching 10%. As expected, growth curves begin to converge towards the AIOMFAC equilibrium predictions as RH 10 approaches 100% for all systems, as the solutions become very dilute (Fig. 4). Our results 11 suggest that lack of accounting for non-ideal interactions and phase-separations leads to errors in 12 13 predicted sub-saturated hygroscopic growth. Note that while maximum deviations in HGFs for the simplified approaches (compared to the AIOMFAC-based equilibrium calculation) are 14 15 generally on the order of 10 - 25%, the corresponding errors in particle size and refractive index can substantially impact estimates of aerosol scattering and radiative forcing (Finlayson-Pitts and 16 Pitts, 2000). 17

18

19 4 Conclusions

Measurements and detailed thermodynamic modeling of the water uptake of model 20 21 organic-inorganic atmospheric aerosol systems demonstrate variability in hygroscopic behavior 22 across aerosol systems with differing RH-dependent phase behavior. Measured and modeled 23 growth curves approach smoother, more continuous water uptake with decreasing prevalence of LLPS and increasing O:C ratios of the OA components. AIOMFAC-predicted growth curves 24 25 reproduce the measured hygroscopic behavior reasonably well for all systems. A comparison 26 between measured and modeled HGFs for the sucrose-ammonium sulfate particles indicates the presence of a viscous semi-solid phase that inhibits the crystallization of ammonium sulfate. We 27 conclude that particle drying within HGF instrumentation may induce the formation of a highly 28 viscous, amorphous phase (potentially accompanied by a moisture-loss-related glass transition). 29 As a result, such measurements may not accurately reflect equilibrium water-uptake behavior. 30

This is an important consideration when applying similar instruments to measure the hygroscopic 1 2 behavior of ambient aerosols, particularly for the highly oxygenated SOA for which sucrose 3 serves as a surrogate in our experiments. Our results add support to the growing body of literature suggesting that accounting for the influence of viscous liquid or semi-solid phases to 4 water uptake and release can be important for accurately modeling the hygroscopic behavior of 5 atmospheric aerosols. The performance of simplified approaches for modeling water uptake 6 differs for particles with differing phase states/equilibria, suggesting that a single simplified 7 modeling approach cannot be used to capture the water-uptake behavior for the diversity of 8 particle phase behavior expected in the atmosphere. Errors in HGFs calculated using the 9 simplified models are of sufficient magnitude to contribute substantially to uncertainties in 10 estimates of particle optical and radiative properties. Parameterizations of LLPS and other 11 12 complex phase behavior based on commonly measured variables such as O:C (e.g., Bertram et al., 2011; Koop et al., 2011; Song et al., 2012a) may prove valuable in applying the simplified 13 14 HGF calculation approaches explored here in large-scale models. Average carbon oxidation state (\overline{OS}_{C}) has also been presented as an indicator of the degree of aerosol oxidation (Kroll et al., 15 2011) and may be a useful measure when the goal is to track changes in hygroscopicity with the 16 progression of oxidation and fragmentation of organic molecules in an atmospheric chemistry 17 model. Parameterizations of hygroscopicity based on \overline{OS}_{C} might also lead to advancements in the 18 modeling of aerosol water uptake for complex organic and mixed organic-inorganic systems and 19 should be a consideration in future work. Finally, while the majority of field-based 20 hygroscopicity studies focus on relatively high RH values, future measurements could also focus 21 22 on the growth factors of atmospheric aerosol at low to moderate RH values, as this is the region 23 where water-uptake behavior demonstrates the greatest variability with particle-phase behavior.

24

25 Appendix A: AIOMFAC Model and Phase Equilibrium Calculations

AIOMFAC is a group-contribution, thermodynamic model for the calculation of component activity coefficients in binary and multicomponent mixtures. It was developed to explicitly account for molecular interactions between organic functional groups and inorganic ions in aqueous solutions relevant to atmospheric aerosol chemistry. Descriptions of model details and parameterizations are available elsewhere (Zuend et al., 2008, 2010, 2011; Zuend and Seinfeld, 2012; http://www.aiomfac.caltech.edu). Herein, we provide a brief overview of the key
 aspects of the AIOMFAC model.

3 Within the model, organic molecules are represented as assemblies of functional groups. This treatment of organic molecules is based on the concept that the physiochemical properties 4 of organic compounds can be related to their chemical structure and characteristic structural 5 groups, which allows for treatment of the hundreds to thousands of organic compounds that 6 characterize atmosphere organic aerosol. The organic functional groups included in AIOMFAC 7 8 (alkyl (standard), alkyl (in alcohols), alkyl (in hydrophobic tails of alcohols), alkyl (bonded to 9 hydroxyl group), alkenyl, aromatic hydrocarbons, hydroxyl, aromatic carbon-alcohol, ketone, aldehyde, ester, ether, carboxyl, hydroperoxide, peroxyacid, peroxyacid, peroxyacyl 10 nitrate, organonitrate) allow for the representation of a large variety of compounds observed in 11 atmospheric aerosols. In addition AIOMFAC includes seven atmospherically-relevant cations 12 $(H^+, Li^+, Na^+, K^+, NH_4^+, Mg^{2+}, Ca^{2+})$, five anions $(Cl^-, Br^-, NO_3^-, HSO_4^-, SO_4^{2-})$, and water. 13

Non-ideality (i.e., deviations from Raoult's law) in organic-inorganic aqueous solutions is accounted for through the calculation of activity coefficients for all components in a given mixture. When considering the partitioning of water vapor to a multicomponent liquid mixture, the vapor pressure of water (p_w) over the mixture is related to water activity by:

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$$p_w = p_w^o a_w^{(x)} \tag{A1},$$

where $p_w^0 = p_w^0(T)$ is the saturation vapor pressure over pure liquid water at temperature *T* and $a_w^{(x)}$ is the water activity defined on a mole fraction basis. Under thermodynamic equilibrium conditions, relative humidity and water activity are related by

22
$$RH = a_w^{(x)} = \gamma_w^{(x)} x_w$$
 (A2)

where $\gamma_w^{(x)}$ is the activity coefficient of water on a mole fraction basis and x_w is the mole fraction of water in the particle-phase liquid mixture (here, for the case of a single liquid phase). Similarly, activity coefficients are key to accurately describing the partitioning of semi-volatile organic compounds between the gas and particle phases under thermodynamic equilibrium.

In AIOMFAC, activity coefficients are derived from expressions for the long-range, middle-range, and short-range molecular interactions that contribute to total Gibbs excess

energy, which is a descriptor of the overall non-ideality of a thermodynamic system. In addition 1 2 to their application in calculations of the gas-particle partitioning of water and other semi-3 volatile species (i.e., vapor-liquid equilibria), activity coefficients of all components in a multicomponent mixture are required for the calculation of solid-liquid (SLE) and liquid-liquid 4 equilibria (LLE). The prevalence of liquid-liquid phase separation and the composition of each 5 phase is calculated in this work by application of AIOMFAC to compute activity coefficients in 6 distinct liquid phases based on a reliable and efficient algorithm for finding the phase 7 compositions that correspond to an equilibrium state. A liquid-liquid equilibrium state of a 8 closed thermodynamic system is a state of minimum Gibbs energy of that system. The same 9 applies to SLE and, likewise, to coupled vapor-liquid-liquid-solid equilibrium calculations, such 10 as those performed in this work at given temperature and relative humidity to determine the 11 number and composition of the particle phases at equilibrium. Hence, while the AIOMFAC 12 model is at the heart of such equilibrium calculations, the distinct phases and their compositions 13 14 are determined using a more general thermodynamic equilibrium model, as described by Zuend et al. (2010) and Zuend and Seinfeld (2012). For the calculation of a potential liquid-liquid phase 15 16 separation, the equilibrium model essentially solves a system of nonlinear equations numerically to determine the phase state (i.e., one liquid phase versus two liquid phases) that achieves a 17 18 minimum in Gibbs energy for a given overall particle-phase composition at constant temperature and pressure. Full details regarding the algorithm used to diagnose the prevalence of LLPS and 19 20 to calculate the corresponding phase composition is available in Zuend and Seinfeld (2013).

21

22 Appendix B: Control Hygroscopic Growth Experiments

Following the same methods as described in the main text of the paper, control 23 hygroscopic growth experiments were conducted for pure ammonium sulfate aerosols with dry 24 mobility diameters of 250 nm. HGF experiments were conducted at room temperature (~298 K). 25 The control experiments were conducted for comparison against the mixed organic-ammonium 26 sulfate aerosol systems (i.e., to explore the influence of the organic species on the mixed organic-27 inorganic particle hygroscopicity when starting with the same dry size) and to evaluate the 28 performance of the DASH-SP and the growth-factor-calculation algorithm using this well-29 characterized aerosol system. DASH-SP measurements reproduced previous experimental 30

characterizations of ammonium sulfate aerosol hygroscopicity (e.g., Tang, 1980; Sorooshian et 1 al., 2008) and are in agreement with AIOMFAC-predicted diameter growth factors (Figure B1). 2

3

Author Contributions 4

N. H., J. H. S., and R. C. F. designed the experiments. N. H. carried out the experiments with 5 assistance from W. M. A. Z. developed the modeling tools and performed all simulations. N. H. 6 prepared the manuscript with contributions from all co-authors. 7

8

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1 References

2	Abbatt, J. P. D., Lee, A. K. Y., and Thornton J.A.: Quantifying trace gas uptake to tropospheric
3	aerosol: recent advances and remaining challenges, Chem. Soc. Rev., 41, 6555 - 6581,
4	2012.
5	Bertram, A. K., Martin, S. T., Hanna, S. J., Smith, M. L., Bodsworth, A., Chen, Q., Kuwata, M.,
6	Liu, A., You, Y., and Zorn, S. R.: Predicting the relative humidities of liquid-liquid phase
7	separation, efflorescence, and deliquescence of mixed particles of ammonium sulfate,
8	organic material, and water using the organic-to-sulfate mass ratio of the particle and the
9 10	oxygen-to-carbon elemental ratio of the organic component, Atmos. Chem. Phys., 11, 10995–1006, 2011.
11	Binkowski, F. S. and Roselle, S. J.: Models-3 Community Multiscale Air Quality (CMAQ)
12	model aerosol component 1. Model description, J. Geophys. Res., 108 4183, doi:
13	10.1029/2001JD001409, 2003
14	Bones, D. L., Reid, J. P., Lienhard, D. M., and Krieger, U. K.: Comparing the mechanism of
15	water condensation and evaporation in glassy aerosol, Proc. Natl. Acad. Sci., 109, 11613 -
16	11618, 2012.
17	Brooks, S. D., DeMott, P. J., and Kreidenweis, S. M.: Water uptake by particles containing
18	humic materials and mixtures of humic materials with ammonium sulfate, Atmos. Environ.,
19	38, 1859-1868, 2004.
20	Cappa, C. D., Lovejoy, E. R., and Ravishankara, A. R.: Evidence for liquid-like and nonideal
21	behavior of a mixture of organic aerosol components, Proc. Natl. Acad. Sci., 105, 18687-
22	18691, 2008.
23	Chan, M. N., Choi, M. Y., Ng, N. L., and Chan, C. K.: Hygroscopicity of water-soluble organic
24	compounds in atmospheric aerosols: amino acids and biomass burning derived organic
25	species, Environ. Sci. Technol., 39, 1555-1562, 2005.
26	Chan, M. N., Lee, A. K. Y., and Chan, C. K.: Responses of ammonium sulfate particles coated
27	with glutaric acid to cyclic changes in relative humidity: hygroscopicity and Raman
28	characterization, Environ. Sci. Technol., 40, 6983-6989, 2006.

1 2	Choi, M. Y. and Chan, C. K.: Continuous measurements of the water activities of aqueous droplets of water-soluble organic compounds. J. Phys., Chem. A., 106, 4566-4572, 2002a.
3	Choi, M. Y. and Chan, C. K.: The effects of organic species on the hygroscopic behaviors of inorganic species Environ. Sci. Technol. 36, 2422, 2428, 2002b
4	morganic aerosois, Environ. Sci. Technol., 30, 2422-2428, 20020.
5 6	Chung, S. H. and Seinfeld, J. J.: Global distribution and climate forcing of carbonaceous aerosols, J. Geophys. Res., 107, 4407, doi: 10.1029/2001JD001397, 2002.
7 8 9	Ciobanu, V. G., Marcolli, C., Krieger, U. K., Weers, U., and Peter, T.: Liquid-liquid phase separation in mixed organic/inorganic aerosol particles. J. Phys. Chem., 113 10966–10978, 2009.
10 11 12	Clegg, S. L., Seinfeld, J. H., and Brimblecombe, P.: Thermodynamic modeling of aqueous aerosols containing electrolytes and dissolved organic compounds, J. Aerosol Sci., 32, 713- 738, 2001.
13 14 15	Clegg, S. L. and Seinfeld, J. H.: Thermodynamic models of aqueous solutions containing inorganic electrolytes and dicarboxylic acids at 298.25 K. 1. The acids as nondissociating components, J. Phys. Chem. A, 110, 5692-5717, 2006.
16 17 18 19 20	 Clegg, S. L. and Wexler, A. S.: Densities and Apparent Molar Volumes of Atmospherically Important Electrolyte Solutions. 1. The Solutes H₂SO₄, HNO₃, HCl, Na₂SO₄, NaNO₃, NaCl,(NH₄) ₂SO₄, NH₄NO₃, and NH₄Cl from 0 to 50° C, including extrapolations to very low temperature and to the pure liquid state, and NaHSO₄, NaOH, and NH₃ at 25° C." J. Phys. Chem. A, 115, 3393-3460, 2011.
21 22	Cruz, C., and Pandis, S.N.: Deliquescence and hygroscopic growth of mixed inorganic-organic atmospheric aerosol. Environ. Sci. Technol., 32, 4313 – 4319, 2000.
23 24 25	Dette, H. P., Qi, M., Schröder, D. C., Godt, A., and Koop, T.: Glass-forming properties of 3- methylbutane-1,2,3-tricarboxylic acid and its mixtures with water and pinonic acid. J. Phys. Chem. A., 118, 7024-7033, 2014.
26 27 28	Drury, E., Jacob, D. J., Wang, J., Spurr, R. J. D., and Chance, K.: Improved algorithm for MODIS satellite retrievals for aerosol optical depth over western North America, J. Geophys. Res., 113, D16204, doi: 10.1029/2007JD009573, 2008.

1	Duplissy, J., DeCarlo, P. F., Dommen, J., Alfarra, M. R, Metzger, A., Barmpadimos, I., Prevot,
2	A. S. H., Weingartner, E., Tritscher, T., Gysel, M., Aiken, A. C., Jimenez, J. L.,
3	Canagaratna, M. R., Worsnop, D. R., Collins, D. R., Tomlinson, J., and Baltensperger, U.:
4	Relating hygroscopicity and composition of organic aerosol particulate matter, Atmos.
5	Chem. Phys., 11, 1155-1165, 2011.
6	Erdakos, G. B. and Pankow, J. F.: Gas/particle partitioning of neutral and ionizing compounds
7	to single- and multi-phase aerosol particles. 2. Phase separation in liquid particulate matter
8	containing both polar and low-polarity organic compounds, Atmos. Environ., 38, 1005 -
9	1013, 2004.
10	Ervens, B., Turpin, B. J., and Weber, R. J.: Secondary organic aerosol formation in cloud
11	droplets and aqueous particles (aqSOA): a review of laboratory, field, and model studies.
12	Atmos. Chem. Phys., 11, 11069-11102, 2011.
13	Ervens, B., Wang, Y., Eagar, J., Leaitch, W. R., Macdonald, A. M., Valsaraj, K. T., and
14	Herckes, P.: Dissolved organic carbon (DOC) and select aldehydes in cloud and fog water:
15	the role of the aqueous phase in impacting trace gas budgets. Atmos. Chem. Phys., 13,
16	5117-5135, 2013.
17	Finlayson-Pitts B. J. and Pitts Jr., J. N.: Chemistry of the Upper and Lower Atmosphere,
18	Academic Press, San Diego, CA, 2000.
19	Fu, Y., Luo, G., and Ma, X.: Regional and global modeling of aerosol optical properties with a
20	size, composition, and mixing state, resolved particle microphysics model, Atmos. Chem.
21	Phys., 12, 5719-5736, 2012.
22	Gao, Y., Yu, L. E., and Chen, S. B.: Effects of organics on the efflorescence relative humidity
23	of ammonium sulfate or sodium chloride particles, Atmos. Environ., 42, 4433-4445, 2008.
24	Girolami, G. S.: A simple "back of the envelope" method for estimating the densities and
25	molecular volumes of liquids and solids, J. Chem. Educ., 71, 962-964, 1994.
26	Goldstein, A. H. and Galbally, I. E.: Known and unexplored organic constituents in the Earth's
27	atmosphere, Environ. Sci. Technol., 41, 1514-1521, 2007.

- Gysel, M., Weingartner, E., Nyeki, S., Paulsen, D., Baltensperger, U., Galambos, I., and Kiss, 1 2 G.: Hygroscopic properties of water-soluble matter and humic-like organics in atmospheric 3 fine aerosol, Atmos. Chem. Phys., 4, 35-50, 2004. Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen, 4 5 J., Donahue N. M., George, C., Goldstein, A. H., Hamilton, J. F., Herrmann, H., Hoffmann, T., Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr, A., Maenhaut, W., 6 7 McFiggans, G., Mentel, T. F., Monod, A., Prevot, A. S. H., Seinfeld, J. H., Surratt, J. D., 8 Szmigielski, R., and Wildt, J.: The formation, properties and impact of secondary organic 9 aerosol: current and emerging issues, Atmos. Chem. Phys., 9, 5155–5236, 2009. 10 Hatch, C. D., Gierlus, K. M, Zahardis, J., Schuttlefield, J., and Grassian, V. H.: Water uptake of 11 humic and fulvic acid: measurements and modelling using single parameter Köhler theory, 12 Environ. Chem., 6, 380-388, 2009. 13 Kanakidou, M., Seinfeld, J. H., Pandis, S. N., Barnes, I., Dentener, F. J., Facchinni, M. C., Van Dingenen, R., Ervens, B., Nenes, A., Nielsen, C. J., Swietlicki, E., Putaud, J. P., Balkanski, 14 Y., Fuzzi, S., Horth, J., Moortgat, G. K., Winterhalter, R., Myhre, C. E. L., Tsigaridis, K., 15 Vgnati, E., Stephanou, E. G., and Wilson, J.: Organic aerosol and global climate modeling: 16 a review, Atmos. Chem. Phys., 5, 1053-1123, 2005. 17 Koepke, P., Hess, M., Schult, I., and Shettle, E. P: Global Aerosol Data Set, Rep. No. 243, Max-18 Planck-Institut fur Meteorol., Hamburg, Germany, 1997. 19
- Koop, T., Bookhold, J., Shiraiwa, M., and Pöschl, U.: Glass transition and phase state of organic
 compounds: dependency on molecular properties and implications for secondary organic
 aerosols in the atmosphere, Phys. Chem. Chem. Phys., 13, 1923–19255, 2011.

Krieger, U., Marcolli, C., and Reid, J. P.: Exploring the complexity of aerosol particle
properties and processes using single particle techniques, Chem. Soc. Rev., 41, 6631–6662,
2012.

- 26 Kroll, J. H., Donahue, N. M., Jimenez, J. L., Kessler, S. H., Canagaratna, M. R., Wilson, K. R.,
- 27 Altieri, K. E., Mazzoleni, L. R., Wozniak, A. S., Bluhm, H., Mysak, E. R., Smith, J. D.,
- 28 Kolb, C. E., and Worsnop, D. R.: Carbon oxidation state as a metric for describing the
- chemistry of atmospheric organic aerosol, Nature Chemistry, 3, 133-139, 2011.

1	Lei, T., Zuend, A., Wang, W. G., Zhang, Y. H., and Ge, M. F.: Hygroscopicity of organic
2	compounds from biomass burning and their influence on the water uptake of mixed
3	organic-ammonium sulfate aerosols, Atmos. Chem. Phys., 14, 111625-11183, 2014.
4	Lienhard, D. M., Bones, D. L., Zuend, A., Krieger, U. K., Reid, J. P., and Peter, T.:
5	Measurements of thermodynamic and optical properties of selected aqueous organic and
6	organic-inorganic mixtures of atmospheric relevance, J. Phys. Chem. A, 116, 9954-9968,
7	2012.
8	Lienhard, D. M., Huisman, A. J., Bones, D. L., Te, Y-F., Luo, B. P., Krieger, U. K., and Reid, J.
9	P.: Retrieving the translational diffusion coefficient of water from experiments on single
10	levitated aerosol droplets, Phys. Chem. Chem. Phys., 16, 16677-16683, 2014.
11	Liu, X. and Wang, J.: How important is organic aerosol hygroscopicity to aerosol indirect
12	forcing? Environ. Res. Lett., 5, 044010, doi:10.1088/1748-9326/5/4/044010, 2010.
13	Marcolli, C., Luo, B. P., Peter, Th., and Wienhold, F. G.: Internal mixing of the organic aerosol
14	by gas phase diffusion of semivolatile organic compounds, Atmos. Chem. Phys., 4, 2593-
15	2599, 2004.
16	Marcolli, C. and Krieger, U. K.: Phase changes during hygroscopicity cycles of mixed
17	organic/inorganic model systems of tropospheric aerosols, J. Phys. Chem. A, 110, 1881-
18	1893, 2006.
19	Marcolli, C., Luo, B. P., Peter, Th., and Wienhold, F. G.: Internal mixing of the organic aerosol
20	by gas diffusion of semivolatile organic compounds, Atmos. Chem. Phys., 4, 2593-2599,
21	2006.
22	Massoli, P., Lambe, A. T., Ahern, A. T., Williams, L. R., Ehn, M., Mikkilä, J., Canagaratna, M.
23	R., Brune, W. H., Onasch, T. B., Jayne, J. T., Petäjä, T., Kulmala, M., Laaksonen, A.,
24	Kolb, C. E., Davidovits, P., and Worsnop, D. R.: Relationships between aerosol oxidation
25	level and hygroscopic properties of laboratory generated secondary organic aerosol (SOA)
26	particles, Geophys. Res. Lett., 37, L24801, doi: 10.1029/2010GL045258, 2010.
27	Mikhailov, E., Vlasenko, S., Martin, S. T., Koop, T., and Poschl, U.: Amorphous and crystalline
28	aerosol particles interacting with water vapor: conceptual framework and experimental

- evidence for restructuring, phase transitions and kinetic limitations, Atmos. Chem. Phys., 9,
 9491–9522, 2009.
- Mochida, M. and Kawamura, K.: Hygroscopic properties of levoglucosan and related organic
 compounds characteristic to biomass burning aerosol particles, J. Geophys. Res., 109,
 D21202, doi: 10.1029/2004JD004962, 2004.
- Moore, R. H. and Raymond, T. M.: HTDMA analysis of multicomponent dicarboxylic acid
 aerosols with comparison to UNIFAC and ZSR, J. Geophys. Res., 113, doi:
 10.1029/2007JD008660, 2008.
- Murphy, D. M., Cziczo, D. J., Froyd, K. D., Hudson, P. K., Matthew, B. M., Middlebrook, A.
 M., Peltier, R. E., Sullivan, A., Thomson, D. S., and Weber, R. J.: Single-particle mass
 spectrometry of tropospheric aerosol particles, J. Geophys. Res., 111, D23S32, doi:

12 10.1029/2006JD007340, 2006.

- Petters, M. D. and Kreidenweis, S. M.: A single parameter representation of hygroscopic
 growth and cloud condensation nucleus activity, Atmos. Chem. Phys., 7, 1961-1971, 2007.
- Pope, F. D., Harper, L., Dennis-Smither, B. J., Griffiths, P. T., Clegg, S. L., and Cox, R. A.:
 Laboratory and modeling study of the hygroscopic properties of two model humic acid
 aerosol particles, J. Aerosol Sci., 41, 457-467, 2010.
- 18 Pöhlker, C., Wiedemann, K. T., Sinha, B., Shiraiwa, M., Gunthe, S. S., Smith, M., Su, H.,
- 19 Artaxo, P., Chen, Q., Cheng, Y. F., Elbert, W., Gilles, M. K., Kilcoyne, A. L. D., Moffet,
- R. C., Weigand, M., Martin, S. T., Pöschl, U., and Andreae, M. O.: Biogenic potassium salt
 particles as seeds for secondary organic aerosol in the Amazon, Science 337, 1075-1078,
- **22** 2012.
- Pöschl, U.: Atmospheric aerosols: composition, transformation, climate and health effects,
 Angew. Chem. Int. Edit., 44, 7520–7540, doi:10.1002/anie.200501122, 2005.
- Prenni, A. J., DeMott, P. J, Kreidenweis, S. M, Sherman, D. E., Russell, L. M., and Ming, Y.:
 The effects of low molecular weight dicarboxylic acids on cloud formation, J. Phys. Chem.,
 105, 11240-11248, 2001.

1 2	Prenni, A. J., DeMott, P. J., and Kreidenweis, S. M.: Water uptake of internally mixed particles containing ammonium sulfate and dicarboxylic acids, Atmos. Environ., 37, 4243-4251,
3	2003.
4	Raatikainen, T. and Laaksonen, A.: Application of several activity coefficient models to water-
5	organic-electrolyte aerosols of atmospheric interest, Atmos. Chem. Phys., 5, 2475-2495,
6	2005.
7	Renbaum-Wolff, L., Grayson, J. W., Bateman, A. P., Kuwata, M., Sellier, M., Murray, B. J.,
8	Shilling, J. E., Martin, S. T., and Bertram, A. K.: Viscosity of α-pinene secondary organic
9	material and implications for particle growth and reactivity, Proc. Natl. Acad. Sci., 110,
10	8014-8019.
11	Robinson, C. B., Schill, G. P., and Tolbert, M. A.: Optical growth of highly viscous
12	organic/sulfate particles, J. Atmos. Chem., 71, 145-156, 2014.
13	Saukko, E., Lambe A. T., Massoli, P., Koop, T., Wright, J. P., Croasdale, D. R., Pederna, D. A,
14	Onasch, T. B., Laaksonen, A., Davidovits, P., Worsnop, D. R., and Virtanen, A.: Humidity-
15	dependent phase state of SOA particles from biogenic and anthropogenic precursors,
16	Atmos. Chem. Phys., 12, 7515–7529, 2012.
17	Seinfeld, J. H. and Pankow, J. F.: Organic atmospheric particulate matter, Rev. Phys. Chem.,
18	54, 121-140, 2003.
19	Seinfeld, J. H. and Pandis S. N.: Atmospheric Chemistry and Physics, 2 nd Edition, John Wiley
20	& Sons, New York, NY, 2006.
21	Shiraiwa, M., Ammann, M., Koop, T., and Pöschl, U.: Gas uptake and chemical aging of
22	organic aerosol particles, Proc. Natl. Acad. Sci., 108, 11003–11008, 2011.
23	Shiraiwa, M., Zuend, A., Bertram, A. K., and Seinfeld, J. H.: Gas-particle partitioning of
24	atmospheric aerosols: interplay of physical state, non-ideal mixing and morphology, Phys.
25	Chem. Chem. Phys., 15, 1141-11453, 2013.
26	Sjogren, S., Gysel, M., Weingartner, E., Baltensperger, U., Cubison, M. J., Coe, H., Zardini, A.
27	A., Marcolli, C., Krieger, U. K., and Peter T.: Hygroscopic growth and water uptake
28	kinetics of two-phase aerosol particles consisting of ammonium sulfate, adipic and humic
29	acid mixtures, J. Aerosol Sci., 38, 157 – 171, 2007.

- Song, M., Marcolli, C., Krieger, U. K., Zuend, A., and Peter, T.: Liquid-liquid phase separation
 and morphology of internally mixed dicarboxylic acids/ammonium sulfate/water particles,
 Atmos. Chem. Phys., 12, 2691-2712, 2012a.
- Song, M., Marcolli, C., Krieger, U. K., Zuend, A., and Peter, T.: Liquid-liquid phase separation:
 Dependence on O:C, organic functionalities, and compositional complexity, Geophys. Res.
 Lett., 39, L19801, doi: 10.1029/2012GL052807, 2012b.
- Song, M., Marcolli, C., Krieger, U. K., Lienhard, D. M., and Peter, T.: Morphologies of mixed
 organic/inorganic/aqueous aerosol droplets, Faraday Discuss., 165, 289-316, 2014.
- Sorooshian, A., Hersey, S., Brechtel, F. J., Corless, A., Flagan, R. C., and Seinfeld, J. H.: Rapid,
 size-resolved aerosol hygroscopic growth measurements: Differential Aerosol Sizing and
 Hygroscopicity Spectrometer Probe (DASH-SP). Aerosol Sci. Technol., 42, 445-464, 2008.
- Svenningsoon, B., Rissler, J., Swietlicki, E., Mircea, M., Bilde, M., Facchini, M. C., Decesari,
 S., Fuzzi, S., Zhou, J., Monster, J., and Rosenorn, T.: Hygroscopic growth and critical
 supersaturations for mixed aerosol particles of inorganic and organic compounds of
 atmospheric relevance, Atmos. Chem. Phys., 6, 1937-1952, 2006.
- Tang, I. N.: Deliquescence properties and particle size change of hygroscopic aerosols, in:
 Generation of Aerosols and Facilities for Exposure Experiments, K. Willekie, Ed., Ann
 Arbor Science Publishers, Ann Arbor, MI, 153-167, 1980.
- Tong, H.-J., Reid, J. P., Bones, D. L., Luo, B. P., and Krieger, U. K.: Measurements of the
 timescales for the mass transfer of water in glassy aerosol at low relative humidity and
 ambient temperature, Atmos. Chem. Phys., 11, 4739-4754, 2011.
- Vaden, T. D., Imre, D., Beranek, J., Shrivastava, M., and Zelenyuk, A.: Evaporation kinetics
 and phase of laboratory and ambient secondary organic aerosol, Proc. Natl. Acad. Sci., 108,
 2190 2195, 2011.
- Veghte, D. P., Altaf, M. B., and Freedman, M. A.: Size dependence of the structure of organic
 aerosol, J. Am. Chem. Soc., 135, 16046-16049.
- Virtanen, A., Joutsensaari, J., Koop, T., Kannosto, J., Yli-Pirila, P., Leskinin, J., Makela, J. M.,
 Holopainen, J. K., Poschl, U., Kulmala, M., Worsnop, D. R., and Laaksonen A. An
 amorphous solid state of biogenic secondary organic particles. Nature, 467, 824–827, 2010.

1 2 3	Wise, M. W., Surratt, J. D., Curtis, D. B., Shilling, J. E., and Tolbert, M. A.: Hygroscopic growth of ammonium sulfate/dicarboxylic acids, J. Geophys. Res., 108, doi: 10.1029/2003JD003775, 2003.
4 5 6	 Xu, W., Guo, S., Gomez-Hernandez, M., Zambora, M. L., Secrest, J., Marrero-Ortiz, W., Zhang, A. L., Collins, D. R., and Zhang, R.: Cloud forming potential of oligomers relevant to secondary organic aerosols, Geo. Phys. Res. Lett., 41,6538–6545, 2014.
7 8 9 10	 You, Y., Renbaum-Wolff, L., Carreras-Sospedra, M., Hanna, S. J., Hiranuma, N., Kamal, S., Smith, M. L., Zhang, X. L., Weber, R. J., Shilling, J. E., Dabdub, D., Martin, S. T., and Bertram, A. K.: Images reveal that atmospheric particles can undergo liquid-liquid phase separations, Proc. Natl. Acad. Sci. 109, 13188-13193, 2012.
11 12 13	You, Y., Renbaum-Wolff, L., and Bertram, A. K.: Liquid-liquid phase separation in particles containing organics mixed with ammonium sulfate, ammonium bisulfate, ammonium nitrate, or sodium chloride, Atmos. Chem. Phys., 13, 11723-11734, 2013.
14 15 16	You, Y. and Bertram, A. K.: Effects of molecular weight and temperature on liquid-liquid separation in particles containing organic species and ammonium sulfate, Atmos. Chem. Phys., 15, 1351-1365, 2015.
17 18 19	You, Y., Smith, M. L., Song, M., Martin, S. T., and Bertram, A. K.: Liquid-liquid phase separation in atmospherically relevant particles consisting of organic species and inorganic salts, Int. Rev. Phys. Chem., 33, 43-77, 2014.
20 21 22	Yu, F., Luo, G., and Ma, X.: Regional and global modeling of aerosol optical properties with a size, composition, and mixing state resolved particle microphysics model, Atmos. Chem. Phys., 12, 5719-5736, 2012.
23 24 25	Zardini, A. A., Sjogren, S., Marcolli, C., Krieger, U. K., Gysel, M., Weingartner, E., Baltensperger, U., and Peter, T.: A combined particle trap/HTDMA hygroscopicity study of mixed organic/inorganic aerosol particles, Atmos. Chem. Phys., 5589-5601, 2008.
26 27 28	Zaveri, R. A., Easter, R. C., Fast, J. D., and Peters, L. K.: Model for Simulating Aerosol Interactions and Chemistry (MOSAIC), J. Geophys. Res., 113, D13204, doi: 10.1029/2007JD008782, 2008.

1 2 3	Zaveri, R. A., Easter, R. C., Shilling, J. E., and Seinfeld, J. H.: Modeling kinetic partitioning of secondary organic aerosol and size distribution dynamics: representing effects of volatility, phase state, and particle-phase reaction, Atmos. Chem. Phys., 14, 5153-5181, 2014.
4	Zhang, Q., Canagaratna, M. R., Jayne, J. T., Worsnop, D. R., and Jimenez, J. L.: Time- and
5	size-resolved chemical composition of submicron particles in Pittsburgh: Implications for
6	aerosol sources and processes. J. Geophys. Res., 110, D07S09, doi:
7	10.1029/2004JD004649, 2005.
8 9	Zobrist, B., Marcolli, C., Pedernera, D. A., and Koop, T. Do atmospheric aerosols form glasses? Atmos. Chem. Phys., 8, 5221–5244, 2008.
10	Zobrist, B., Soonsin, V., Luo, B. P., Kriegerm U. K., Marcolli, C., Peter, T, Koop, T.: Ultra-
11	slow water diffusion in aqueous sucrose glasses, Phys. Chem. Chem. Phys., 13, 3514-3526,
12	2011.
13	Zuend, A., Marcolli, C., Luo, B. P., and Peter T.: A thermodynamic model of mixed organic-
14	inorganic aerosols to predict activity coefficients, Atmos. Chem. Phys., 8, 4559-4593,
15	2008.
16	Zuend, A., Marcolli, C., Peter, T., and Seinfeld, J. H.: Computation of liquid-liquid equilibria
17	and phase stabilities: implications for RH-dependent gas/particle partitioning of organic-
18	inorganic aerosols, Atmos. Chem. Phys., 10, 7795-7820, 2010.
19	Zuend, A., Marcolli, C., Booth, A. M., Lienhard, D. M., Soonsin, V., Krieger, U. K., Topping,
20	D. O., McFiggans, G., Peter, T., and Seinfeld J. H.: New and extended parameterization of
21	the thermodynamic model AIOMFAC: calculation of activity coefficients for organic-
22	inorganic mixtures containing carboxyl, hydroxyl, carbonyl, ether, ester, alkenyl, alkyl, and
23	aromatic functional groups, Atmos. Chem. Phys., 11, 9155–9206, 2011.
24	Zuend, A. and Seinfeld, J. H.: Modeling the gas-particle partitioning of secondary organic
25	aerosol: the importance of liquid-liquid phase separation. Atmos. Chem. Phys., 12, 3857-
26	3882, 2012.
27	Zuend, A. and Seinfeld, J. H.: A practical method for the calculation of liquid-liquid equilibria
28	in multicomponent organic-water-electrolyte systems using physiochemical constraints,
29	Fluid Phase Equilibria, 337, 201-213, 2013.

Aerosol System	Organic-Component(s) Structures	Organic- Component(s) Density (kg/m ³)	Organic:Inorganic Dry Mass Ratio	Previous Experimentally- Determined Phase Behavior	References
Citric Acid + Ammonium Sulfate	HO OH OH	1.580	2:1	No LLPS observed	Bertram et al., (2011), You et al. (2013)
Diethylmalonic Acid + Ammonium Sulfate	HO OH	1.131	2:1	LLPS at RH ≤ 89%	You et al. (2013)
2-Methylglutaric Acid + Ammonium Sulfate	HO OH	1.169	2:1	LLPS at RH ≤ 75%	You et al. (2013)
C ₅ dicarboxylic acids mixture (glutaric, methylsuccinic, dimethlymalonic) + Ammonium Sulfate	H H H H H H H H H H H H H H H H H H H	1.219	1:1	No LLPS observed	Song et al. (2012a)

Table 1. Aerosol systems studied

C7 dicarboxylic acid mixture (3- methyladipic, 3,3- dimethylglutaric, diethylmalonic) + Ammonium Sulfate	C ₆ dicarboxylic acid mixture (2- methylglutaric, 3 - methylglutaric, 2,2- dimethylsuccinic) + Ammonium Sulfate
Here and the second sec	HO OH O OH O OH O OH O OH
1.131	1.169
1:1	1:1
Transition from LLPS with core- shell morphology to single liquid phase at RH = 89 – 90%	Transition from LLPS with partially- engulfed morphology to single liquid phase at RH = 74%
Song et al. (2012a)	Song et al. (2012a)

Sucrose	5	1.309	N/A	Glass transition at	Tong et al. (2011),
	O DE			RH = 24 - 53%, depending on timescale of RH	Zobrist et al. (2011)
	HO OH HO OH			change	
	HOMINI				
Sucrose +			1:1	Unknown	This study
Ammonium					
Sulfate					
Sucrose +			2:1	No LLPS observed	You and Bertram
Ammonium					(2015)
Sulfate					



phase 2, and a solid phase) as a function of relative humidity. Empty boxes indicate that no second liquid phase is predicted curves indicate the AIOMFAC-predicted composition of each of three possible phases present in the particles (liquid phase 1, liquid experiments and error bars indicate the standard deviation of the measured growth factors. Panels below the hygroscopic growth ammonium sulfate particles, (b) 2-methylglutaric acid-ammonium sulfate particles, and (c) citric acid-ammonium sulfate particles with dry organic: inorganic mass ratios of 2:1 for all systems. The black circles indicate the average growth factor measured across ten Figure 1. Top panels: Measured and AIOMFAC-predicted hygroscopic diameter growth factors for (a) diethylmalonic acid-



growth factor measured across ten experiments and error bars indicate the standard deviation of the measured growth factors. Panels ammonium sulfate particles, with dry organic: inorganic mass ratios of 1:1 for all systems. The black circles indicate the average ammonium sulfate particles, (b) C_6 dicarboxylic acid mixture-ammonium sulfate particles, and (c) C_7 dicarboxylic acid mixtureparticles (liquid phase 1, liquid phase 2, and a solid phase) as a function of relative humidity. Empty boxes indicate that no second below the hygroscopic growth curves indicate the AIOMFAC-predicted composition of each of three possible phases present in the Figure 2. Top panels: Measured and AIOMFAC-predicted hygroscopic diameter growth factors for (a) C₅ dicarboxylic acid mixtureliquid phase is predicted.

organic: inorganic ratio of 2:1. The black circles indicate the average growth factor measured across ten experiments and error bars ammonium sulfate particles with an organic: inorganic ratio of 1:1, and (c) mixed sucrose-ammonium sulfate particles with an AIOMFAC-based equilibrium model, while dashed lines illustrate the dehydration branch in which a supersaturated solution is present Figure 3. Top panels: Measured and AIOMFAC-predicted hygroscopic growth factors for (a) sucrose particles, (b) mixed sucroseindicate the standard deviation of the measured growth factors. Solid lines indicate the hydration curve calculated with the



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predicted. phase 2, and a solid phase) as a function of relative humidity. Empty boxes indicate that no second liquid phase or no solid phase is predicted composition of each of three possible phases present in the particles during the hydration process (liquid phase 1, liquid with respect to the dissolved ammonium sulfate. The three panels below the hygroscopic growth curves indicate the AIOMFAC-



given in parentheses. The performance of the simplified modeling approaches varies across the systems with variations in phase component systems for which we expect that observed water-uptake behavior is governed by thermodynamic equilibrium conditions. moderate RH (RH = 20 - 80%) behavior. Disagreement between the full AIOMFAC-based equilibrium calculations and the simplified models is greatest at low to Organic:inorganic dry mass ratios, which can substantially influence the extent to which non-ideal interactions affect water-uptake, are Figure 4. Comparison of simplified thermodynamic assumptions to the full AIOMFAC hygroscopic growth calculations for the multi-