

Manuscript prepared for Atmos. Chem. Phys. Discuss.  
with version 2014/09/16 7.15 Copernicus papers of the  $\LaTeX$  class copernicus.cls.  
Date: 20 March 2015

# **Analysis of actinic flux profiles measured from an ozone sonde balloon**

**P. Wang, M. Allaart, W. H. Knap, and P. Stammes**

Royal Netherlands Meteorological Institute (KNMI), De Bilt, the Netherlands

Correspondence to: P. Wang (ping.wang@knmi.nl)

## Abstract

A green light sensor has been developed at KNMI to measure actinic flux profiles using an ozone sonde balloon. In total, 63 launches with ascending and descending profiles were performed between 2006 and 2010. The measured uncalibrated actinic flux profiles are analyzed using the Doubling Adding KNMI (DAK) radiative transfer model. Values of the cloud optical thickness (COT) along the flight track were taken from the Spinning Enhanced Visible and Infrared Imager (SEVIRI) Cloud Physical Properties (CPP) product. The impact of clouds on the actinic flux profile is evaluated on the basis of the cloud modification factor (CMF) at the cloud top and cloud base, which is the ratio between the actinic fluxes for cloudy and clear-sky scenes. The impact of clouds on the actinic flux is clearly detected: the largest enhancement occurs at the cloud top due to multiple scattering. The actinic flux decreases almost linearly from cloud top to cloud base. Above the cloud top the actinic flux also increases compared to clear-sky scenes. We find that clouds can increase the actinic flux to 2.3 times of the clear-sky value at cloud top and decrease it to about 0.05 at cloud base. The relationship between CMF and COT agrees well with DAK simulations, except for a few outliers. Good agreement is found between the DAK simulated actinic flux profiles and the observations for single layer clouds in fully overcast scenes. The instrument is suitable for operational balloon measurements because of its simplicity and low cost. It is worth to further develop the instrument and launch it together with atmospheric chemistry composition sensors.

## 1 Introduction

Atmospheric trace gases such as ozone and nitrogen dioxide are involved in a series of chemical reactions driven by solar radiation at UV wavelengths (Crutzen and Zimmermann, 1991). Actinic flux – which is the integral of the radiance over all directions, i.e.  $4\pi$  solid angle – is relevant to the process of photodissociation. Clouds have a large impact on the actinic flux in the atmosphere and, consequently, on photodissociation rates (Calbó et al., 2005).

Therefore, the actinic flux profile is important for the study of the change in concentration of chemically reactive components in the atmosphere and is preferably measured together with the atmospheric chemical composition. Such actinic flux profiles have been measured  
30 by means of balloons and aircrafts during several campaigns.

During the Atlantic Stratocumulus Experiment (ASTEX), tethered balloon soundings were performed on Santa Maria, Azores. Vilà-Guerau de Arellano et al. (1994) compared measured actinic flux profiles with simulations using a delta-Eddington model. Excellent agreement was found for fully cloudy scenes. The authors reported that the actinic flux decreased  
35 from cloud top to cloud base. At cloud top the actinic flux was higher than the clear-sky actinic flux, while at cloud base the actinic flux was lower than the clear-sky values. In the First ISCCP Regional Experiment (FIRE III) Arctic Cloud Experiment actinic fluxes were measured above sea ice in May 1998. De Roode et al. (2001) found that the actinic flux profile within clouds is nearly constant with height, except in a shallow layer below cloud top where  
40 the actinic flux revealed a large increase. The authors attributed this feature to the bright surface of sea ice (high albedo). Actinic fluxes have been measured on the ground and on an aircraft during the INSPECTRO campaign to study the effect of clouds on the spectral actinic flux in East Anglia England in 2003. Kylling et al. (2005) showed that the spectral actinic flux can be reproduced with a 1-D radiative transfer model for clear-sky and fully cloudy  
45 cases. They reported that the actinic flux (in the UV wavelength range) could be enhanced by as much as 60–100% above clouds and reduced by 55–65% below clouds, as compared to the clear-sky situation. Junkermann (1994) measured  $J(O^1D)$  actinic flux within and above stratiform clouds and above snow surfaces from a hang glider with flight altitude of up to 1 km above ground. The actinic flux showed a very strong contribution of reflected or  
50 backscattered radiation within the planetary boundary layer. Palancar et al. (2011) reported extensive comparisons between aircraft-based measurements (between 0.1–11.9 km altitude) of actinic fluxes and Tropospheric Ultraviolet-Visible (TUV) model simulations. They found a good agreement between the measured and TUV clear-sky model actinic fluxes (integrated from 298 to 422 nm). Including both cloudy and clear-sky conditions, the ratio of  
55 observed actinic flux to the TUV clear-sky model value was  $1.1 \pm 0.3$ . Furthermore, accu-

rate spectral actinic flux measurements (280 - 420 nm) between ground and 12 km altitude were made by Hofzumahaus et al. (2002) from an aircraft. Actinic flux profiles have also been measured by balloons in the stratosphere (Schiller et al., 1994) and in the troposphere and stratosphere (Kylling et al., 2003).

60 On the ground, UV monitoring stations usually measure spectral irradiances. The relation between the actinic flux and the irradiance was studied in several papers (e.g. Madronich, 1987; van Weele et al., 1995; Kazadzis et al., 2000; McKenzie et al., 2001). In general, the actinic flux correlates well with the irradiance on the surface but the relationship depends on wavelength, surface albedo, solar zenith angle, and cloud conditions.

65 Most actinic flux profiles presented in the literature were measured in the lower troposphere in the UV wavelength range. Actinic flux observations at green wavelengths (about 510 nm) are more representative for photodissociation by visible light. In combination with ozone and NO<sub>2</sub> observations, the actinic flux profile observations are useful to investigate the photostationary state relationship between NO, NO<sub>2</sub> and O<sub>3</sub> in cloudy scenes in detail (Cantrell et al., 1993; Mannschreck et al., 2004). Knowledge of the chemical inter-relationship between O<sub>3</sub> and NO<sub>2</sub> is important to better constrain their vertical profile in air quality models and in satellite retrievals of O<sub>3</sub> and NO<sub>2</sub>. The actinic flux at UV wavelengths is more relevant to the photodissociation of NO<sub>2</sub> and O<sub>3</sub> than the actinic flux at visible wavelengths. However, good measurements and simulations of the actinic flux at visible wavelengths will give some confidence in our simulations at the UV wavelengths. It was intended to be a cheap, disposable instrument and no harm for the environment. Therefore, we measured actinic flux profiles using a green light sensor attached to an ozone sonde. Another advantage of using an operational ozone sonde is the large altitude range (from surface up to 35 km) and the regularity of launching. The aims of the actinic flux profile measurements are to evaluate the impact of clouds on the actinic flux profiles and to better constrain the O<sub>3</sub> and NO<sub>2</sub> chemical inter-relationship in atmospheric chemistry models.

80 The cloud modification factor (CMF) is often used in the analysis of cloud effects on UV radiation (e.g. Seckmeyer et al., 1996; Mayer et al., 1998; Schwander et al., 2002; Antón et al., 2012; Mateos et al., 2014). The cloud modification factor is the ratio between radia-

85 tion under cloudy and clear-sky conditions. The radiation for clear-sky scenes is calculated  
using the same atmospheric states as for cloudy scenes. It has been demonstrated and  
explained that the surface CMF in the UV has a wavelength dependence (Seckmeyer et al.,  
1996; Crawford et al., 2003). The transmission of clouds alone does not have a significant  
90 wavelength dependence in the UV but Rayleigh scattering does. According to Kylling et al.  
(1997), the observed wavelength dependence is due to photons reflected by clouds, then  
the photons are scattered downward by Rayleigh scattering above the clouds and transmit-  
ted through the clouds to the ground surface. CMF has been used to evaluate the cloud  
effects on irradiance, actinic flux and photolysis rate. We also use CMF in our analysis of  
the actinic flux profile at a green wavelength.

95 In this paper we will describe the instrument and measurements in Sect. 2. The simulation  
method for the actinic flux profiles is presented in Sect. 3. The results are shown in Sect. 4.  
Conclusions are drawn in Sect. 5.

## 2 Instruments and measurements

100 A light sensor has been developed at KNMI using a commercial green LED (Light Emitting  
Diode) with a diameter of 5 mm made of Gallium-Phosphide (GaP). It is mounted in the  
center of a 10 cm Styropor (Polystyrene) sphere. A picture of the light sensor and sphere is  
shown in Fig. 1. The sphere acts both as a thermal insulation and a light diffuser. The de-  
tector and amplifier are temperature stabilized at 25 °C. The instrument has been designed  
to have uniform response to sunlight. The point on the sphere where the LED sensor was  
105 inserted is called north pole. The sensor was glued at a position in the sphere where its  
sensitivity to sunlight was comparable when the sphere was lit from the north or the south  
pole. After the instrument was completed, it was held in a beam of sunlight that would enter  
a single window in an otherwise dark room. The response of the instrument was checked  
by rotating the sphere. If it was found to be more sensitive on one side, a black permanent  
110 marker pen was used to make marks on that side. Typically it was found that the sensitivity  
was higher (10 % or so) on the equator compared to the poles. This process was repeated

until the response of the instrument varied less than 2% when the sphere was rotated. So the spheres have a number of black lines around the equator to make the sensitivity uniform. The light sensor is not absolutely calibrated and the measured actinic profile is in an arbitrary unit. However, before launching a light sensor, inter-comparison measurements have been performed on the ground together with one reference light sensor which is kept and not used for launching. If the reference sensor does not change with time, the actinic flux profiles measured with different sensors should be intercomparable. The exact spectral response of the light sensor is not measured. The spectral response of the same type of sensor has been shown by Brooks and Mims (2001) in their Fig. 2 for the RadioShack green LED. The spectral response of the RadioShack green LED covers the wavelength range 450–580 nm with the peak value at about 560 nm and center at 525 nm. At present, the light sensor is mainly used to evaluate the cloud effects on actinic flux profiles. The green light sensor is chosen because it is not sensitive to ozone and water vapor absorptions, so that the impacts of clouds and aerosols on the actinic fluxes can be separated from the effects of gas absorptions.

The light sensor is launched together with the ozone sonde balloon and the data are transmitted during flight. The light sensor attached to a line and located at about 20 m below the balloon and 2 m below the ozonesonde. Because the payloads of the balloon are rotating during flight, the light sensor could rotate in the shadow of the balloon when the balloon is at high altitude and the sun is almost overhead ( $\text{SZA} < 30^\circ$ ). This can be identified as a sharp dip in the measured actinic flux profile from the 1-second data record. The measurements impacted by the shadow of the balloon have been removed before the data are averaged over 10 s and stored. The instrument weighs about 100 gram, is low cost and can be reused after recovery. The ozone sonde is launched at 11:30 UTC on every Thursday at De Bilt (5.18° E, 52.10° N), the Netherlands. The vertical velocity of the ozone sonde is typically  $5 \text{ m s}^{-1}$  and reaches an altitude of 30–35 km in about 100 min. After the burst of the balloon the sonde drops back to the ground with a parachute in about 60 min. The data stream includes ascending and descending values of time, altitude, pressure,

140 temperature, relative humidity, ozone partial pressure, and actinic flux. The trajectory of the balloon is obtained from GPS data.

### 3 Methods for the actinic flux profile simulations

#### 3.1 Radiative transfer modelling and input data

The Doubling-Adding KNMI (DAK) radiative transfer code (De Haan et al., 1987; Stammes et al., 1989; Stammes, 2001) is used to simulate the measured actinic flux profiles. DAK is a 1-D plane-parallel radiative transfer code with a pseudo-spherical correction for solar zenith angles (SZAs) greater than  $75^\circ$ . Because the measurements were made at noon, the SZAs are usually smaller than  $75^\circ$  except in December and January. The DAK model has been used by van Weele et al. (1995) in the UV-A actinic flux calculations. The DAK simulated surface shortwave broadband irradiances have good agreement with the ground-based BSRN (Baseline Surface Radiation Network) measurements for clear-sky and over-cast water clouds scenes (Wang et al., 2009, 2011). In this paper, the DAK simulations were performed for  $\lambda = 550$  nm to represent the green wavelength range. Absorption by ozone is taken into account by using the shape of a climatological ozone profile and scaling it to the assimilated total ozone column at 12:00 UTC using OMI total ozone data (Eskes et al., 2003). The ozone absorption cross section is taken from Johnston (1997). Water vapor absorption is not taken into account in the simulations because it is weak at 550 nm. The Rayleigh scattering cross section is calculated according to Bodhaine et al. (1999).

160 For clear-sky scenes, aerosol optical thickness (AOT) data are taken from the SPUV sunphotometer at Cabauw ( $4.93^\circ$  E,  $51.97^\circ$  N, 20 km SW of De Bilt) because AOT is not measured in De Bilt. For cloudy scenes, a default AOT of 0.18 at 550 nm is used, which is close to the climatological value for Cabauw. The LOWTRAN rural aerosol profile is used, with addition of a well-mixed aerosol layer from the surface to the top of the boundary layer for clear-sky scenes. The latter is determined from the lidar measurements in Cabauw.

165 For cloudy scenes, the sky is assumed to be fully overcast (cloud fraction of 1). The  
cloud optical thickness, cloud phase, effective radius, cloud top height and cloud mask are  
obtained from the SEVIRI/MSG Cloud Physical Property (CPP) product (Roebeling et al.,  
2006). This product is available every 15 min. The pixel size of SEVIRI measurement is  
170 about  $3 \text{ km} \times 6 \text{ km}$  (longitude x latitude) in the Netherlands. During one actinic flux profile  
measurement there are 10 SEVIRI images taken from 11:30 to 13:45 UTC. The cloud prop-  
erties are selected from every SEVIRI image according to the time and geolocation of the  
balloon along its trajectory from launching to landing. These times and locations are exactly  
matching between the SEVIRI images and the radiosonde. Cloud properties are also se-  
lected when the balloon reaches 5 special altitudes: surface (launching, landing), cloud top  
175 height (ascending, descending), and maximum flight altitude. At these 5 points, the cloud  
properties are selected at the exact geolocation of the balloon but on the SEVIRI images  
taken at the time most close to the balloon measurement. The resulting 15 COT values from  
SEVIRI, with possibly 2 COT values from one image at two geolocations, largely follow the  
COT variations during one profile measurement. The cloud top height is determined from  
180 the peak value in the actinic flux profile and is checked with the cloud top height in the  
SEVIRI data. Note that the location of the actinic flux peak in cloudy scenes depends pri-  
marily on the cloud optical thickness, surface albedo and solar zenith angle (assuming no  
aerosols). At small solar zenith angle, the actinic flux peak usually occurs slightly below the  
cloud top for optically thick clouds over dark surface (Madronich, 1987). The cloud base  
185 heights are taken from operational ceilometer measurements in De Bilt and Cabauw. The  
cloud scattering phase matrices were calculated using Mie theory for water droplets and  
using the effective radius from SEVIRI CPP data. The effective radius and cloud extinction  
coefficient are assumed to be constant inside the clouds. For example, if the cloud occupies  
5 100 m thick layers of the atmospheric model profile and has a COT of 10, then every layer  
190 has a COT of 2. The DAK simulations are performed for 15 COT values covering the SEVIRI  
data. COT of 1 and 100 are also used in the DAK simulations to improve the interpolations.

Although there are 15 collocated COT values from SEVIRI during every actinic flux profile  
measurement, the COT may not be representative of the actual cloud conditions during an



195 actinic flux profile measurement, because of retrieval uncertainty and subpixel cloud variability. The error of the COT from SEVIRI is about 15% ([http://msgcpp.knmi.nl/mediawiki/index.php/MSGCPP\\_product\\_description](http://msgcpp.knmi.nl/mediawiki/index.php/MSGCPP_product_description)). Including the mismatching of the SEVIRI measurement and actinic flux profile measurement, the uncertainty of COT estimation during actinic flux profile measurement is probably larger.

200 In DAK, the altitude levels are specified using the atmospheric profile (height, pressure, temperature) which is taken from the ozone sonde data. In order to get high vertical resolution simulations inside the clouds, the thickness of one atmospheric profile layer is about 100 m for the cloudy atmosphere and coarser ( $\sim 500$  m) for the cloud-free atmosphere. Above 30 km, the atmospheric profile is extended using the mid-latitude summer atmospheric profile (Anderson et al., 1986). The surface albedo is assumed to be 0.15 for grass-covered surface, because the surroundings of the De Bilt site are largely covered by green grass throughout the year. As the altitude of the instrument increases, the surface albedo might change because the instrument sees a larger area. The actinic flux profiles are simulated at different SZA and COT values. Then, the DAK simulated actinic fluxes are interpolated at the actual SZA and COT values along the flight track to get the best simulation for a single measured actinic flux profile.

205  
210

### 3.2 Wavelength dependence of actinic flux profile

Because we do not know the exact spectral response of our green LED, we have chosen in this paper the wavelength of 550 nm to represent the wavelength band of the LED. Here a sensitivity study is performed using DAK to check the wavelength dependence of the actinic flux profile. The actinic flux profile is simulated at 450, 500, 550, 600, and 650 nm for a cloudy scene. These wavelengths cover the green LED spectral response completely. The DAK inputs of aerosol, atmospheric profile, ozone and surface albedo data have been described in Sect. 3.1. The water cloud layer is located from 0.8 to 1.3 km altitude and the cloud optical thickness is 30. The particle size (effective radius) of the water clouds is 10  $\mu\text{m}$ . The cloud fraction is 1 (overcast). The extraterrestrial solar irradiance is assumed to be  $1 \text{ W m}^{-2} \text{ nm}^{-1}$  perpendicular to the solar beam for every wavelength in the simulation.

215  
220

The simulated actinic flux profiles for this cloudy case with and without normalization are shown in Fig. 2 at SZA of 30 and 75°, respectively. The normalized actinic flux profiles are the simulated actinic flux profiles divided by the actinic fluxes at 30 km.

225 As shown in Fig. 2, the simulated actinic fluxes (AF) have a strong wavelength dependence above the cloud due to Rayleigh scattering. Inside the cloud, there is hardly any wavelength dependence of AF because multiple Mie scattering by the cloud particles is dominant. Below the cloud, the wavelength dependence of actinic flux is also small. In this optically thick cloud case, photons scattered above the cloud top cannot easily transmit  
230 through the cloud layer; consequently there is almost no wavelength dependence of AF below the cloud. A wavelength dependence of the actinic flux below the cloud only occurs if the cloud is optically thin (say COT of 8).

The actinic flux profiles have similar shape for the 5 wavelengths at SZA = 30° (see Fig. 2a). After normalization at 30 km, the light profiles for different wavelengths almost overlap.  
235 Similar behavior occurs at SZA smaller than about 60°. However, when the SZA is getting larger than 60°, the actinic flux profile shape depends strongly on the wavelength. After being normalized at 30 km, the 5 actinic flux profiles have large differences at the cloud top. An example is shown in Fig. 2b for SZA of 75°. This indicates that we can use the actinic flux profile shape at 550 nm to represent the actinic flux profile shape of the wavelength  
240 range 450–650 nm at small SZA but not for large SZA.

## 4 Results

### 4.1 Actinic flux profile measurements

Actinic flux profiles were measured during 63 launches, of which 30 launches were made in 2006, 27 in 2007, 4 in 2008, and 2 in 2010. The flight trajectories were mainly between 4–  
245 9° E and 51–53° N. Most launches had one ascending profile and one descending profile; some launches had only ascending profiles while the descending profiles were not received

properly. In the data set there are 14 clear-sky profiles distributed over 9 days and the rest are cloudy profiles.

250 All the profiles are illustrated in Fig. 3a. The profiles are separated into two groups, because of a calibration issue which occurred in the year 2007. Figure 3b shows the ratio between the global irradiance measured at the ground by a pyranometer and the actinic flux at 4 m height (the lowest point of the profile) at 11:30 UTC as a function of SZA. The data points shown as filled circle are measured from 1 June to 21 December 2006. As shown in Fig. 3b, the ratio seems quite consistent in 2006 except for two outliers which are  
255 caused by very small actinic fluxes. The scatter of the points in 2006 could be due to partly cloudy scenes. This indicates that the instruments are comparable to each other and the sensors do not depend on the SZA and temperature. The calibration issue can be identified from the ratio in 2007 which appears to have a different magnitude from that in 2006 and is too variable. The ratio of irradiance to actinic flux at the surface could be used to rule out  
260 large instrument errors before launch, such as the calibration issue in 2007. The ratio is not suitable for the calibration of the light sensor because the relationship between irradiance and actinic flux is general non-linear (Madronich, 1987). The ratio is dependent on solar zenith angle, surface albedo, and presence of the direct solar beam, which requires cloud and aerosol information. In the comparison with DAK simulated profiles, the measured actinic flux profiles were normalized at 30 km and are thus in an arbitrary unit.  
265

As shown in Fig. 3a, the actinic flux profiles have large vertical variability below 10 km because of the presence of clouds and aerosols. Above 10 km, the profiles are less variable. The ascending and descending profiles often overlap between 15 and 35 km altitude, because the sonde falls down rapidly after the burst of the balloon. The descending profiles often do not extend down to the surface, because of loss of radio signal at long distance.  
270

## 4.2 Impact of aerosols and clouds on actinic flux profiles

Figure 4a shows four clear-sky actinic flux profiles measured on 11, 12 and 13 September 2006 and 17 April 2010. The clear-sky scenes are determined from MODIS images and Total Sky Imager data at Cabauw. Figure 4b shows the AOT measured at Cabauw for the

275 same days. The smallest SZA values during the actinic flux profile measurements are about  
42° on 17 April 2010 and 48° for the measurements in September 2006. According to the  
lidar extinction coefficient profiles at Cabauw, the boundary layer heights at 11:30 UTC for  
the 4 days are 1.0 km (17 April 2010), 1.2, 0.5 km, and 0.7 km (11–13 September 2006),  
280 respectively. Because the AOT varies during the day (Fig. 4b) we can expect that the AOT  
also varies along the flight track. On 17 April 2010 the AOT is about 0.1, which is the low-  
est value in the four days. On this day the actinic flux profile above the boundary layer is  
also lower than on the other days. The AOT is about 0.3 on 12 September (12:00 UTC) and  
about 0.4 on 11 and 13 September. This agrees with the actinic flux measurements, namely,  
285 the actinic fluxes at 1–5 km on 11 and 13 September are slightly larger than those on 12  
September. Due to the scattering by aerosols, the actinic flux is enhanced at the top of  
the aerosol layer and above. This further demonstrates the consistency of the instruments  
when the inter-comparison works well. The upper part of the actinic flux profiles ( $> 5$  km)  
has no correlation with the AOT at Cabauw, because of the spatial variation of AOT and  
surface albedo.

290 The actinic flux profiles (ascending profiles only) for 6 cloudy cases are shown in Fig. 5.  
The actinic flux profile peaks at the upper boundary of the high relative humidity (RH) layer.  
Inside clouds, RH values are close to 100%. RH values at temperatures below 0 °C are  
corrected to RH with respect to ice, which is larger than the measured RH with respect to  
water. When the balloon is close to the cloud top, the light sensor detects a sharp peak in  
295 the actinic flux profile (e.g. Fig. 5a–d). Inside the cloud, the actinic flux profile decreases  
until the base of the cloud, and becomes relatively stable below the cloud.

The altitude dependence of the directional radiation inside a scattering atmosphere or  
cloud layer depends mainly on the optical depth  $\tau$  (and its vertical distribution), solar zenith  
angle and surface albedo (Stammes et al., 1989). The fact that the peak in the diffuse radi-  
ative flux is not located at the top of the cloud layer but inside the layer, is due to enhanced  
300 multiple scattering in the top layer of clouds. Since direct sunlight is incident at the top of  
the clouds, the amount of scattered sunlight first increases going from the top downwards  
due to multiple scattering, reaches a maximum, and then decreases again. In first order,

the diffuse radiative flux is linear in  $\tau$  at the top of the clouds. It has been found by Stammes  
305 et al. (1989) that for an isotropically scattering cloud the downward flux increases descend-  
ing into the cloud layer for high sun ( $\mu_0 > 0.5$ ), but decreases from the cloud top for low  
sun ( $\mu_0 < 0.5$ ). Simulations for actinic fluxes have been done by Madronich (1987) using  
a simple isotropic model and a Delta-Eddington model. This finding was corroborated with  
multiple scattering calculations. In the field of actinic flux studies, the peak below the cloud  
310 top has been found in actinic flux calculations and in balloon measurements (Van Weele  
et al., 1995; Van Weele, 1996; Vilà-Guerau de Arellano et al., 1994).

As can be seen in Figs. 4–5, the actinic flux profiles have small high frequency varia-  
tions. The variations are relatively larger in cloud-free scenes and above clouds where the  
instrument is exposed to direct sunlight and smaller inside and below clouds, where mainly  
315 diffuse radiation is present. It is probably caused by the rotation of the balloon and the light  
sensor (sphere). The instrument has been adjusted before launch to have less than 2%  
variation during rotation (see Sect. 2). Therefore, it is interesting to check the relative varia-  
tion using real profile measurements. The relative variation is determined using the actinic  
fluxes between 15 and 30 km in every ascending profile. In 63 ascending profiles, 59 pro-  
320 files have relative variations smaller than 2%; 2 profiles have relative variations of 2.1% and  
3.2 %, respectively; 2 profiles were excluded because of not reaching high altitude. The  
mean variation is 0.7 % for the 61 profiles. If the data between 20 and 25 km are used for  
the test, the mean variation is also about 0.7 %. This demonstrates that the instrument has  
a quite uniform response to sunlight.

### 325 **4.3 Actinic flux profiles compared with surface radiation measurements**

Figure 6a shows the ratio between the actinic flux below the cloud base close to the surface  
and at the cloud top (the peak of the actinic flux profile) ( $R_{\text{base/top}}$ ) as a function of the sur-  
face solar irradiance measured in De Bilt for 63 ascending profiles. Although some profiles  
have a calibration issue, this has no impact on ratios calculated from one profile. The ratio  
330  $R_{\text{base/top}}$  is determined by cloud optical thickness, surface albedo and solar zenith angle.

Similar to the cloud modification factor (CMF) used in UV radiation studies (e.g. Antón et al., 2012), ratios of the actinic flux profiles in cloudy scenes and clear-sky scenes are calculated at cloud top ( $CMF_{top}$ ) and below the cloud base (close to the ground surface;  $CMF_{base}$ ). For clear-sky scenes, the CMF is 1 by definition. In principle, the clear-sky actinic flux profile should be calculated using a radiative transfer model using an atmospheric state which is identical to the cloudy scene. Here we used the atmospheric state (temperature, pressure, aerosol optical thickness) on the clear-sky day of 11 September 2006 as a reference for all the cloudy actinic flux profiles in 2006. However, the clear-sky profiles are simulated at the same SZA as occurred for the cloudy profiles. The AOT is not available for cloudy scenes. The AOT is 0.34 on 11 September. According to the scaling factor derived from the clear-sky actinic flux profiles for 11 September 2006, the DAK simulated clear-sky profiles are converted to the same scale as the measured profiles. Otherwise, the calculated CMF would not fulfill the definition of 1 for clear-sky scenes. The simulation of the clear-sky profile is discussed in Sect. 4.4.

In Fig. 6b the CMF as a function of COT is shown for cloud top and cloud base, which we denote by  $CMF_{top}$  and  $CMF_{base}$ . The lines are simulations of  $CMF_{top}$  and  $CMF_{base}$  for  $SZA = 60^\circ$  and  $30^\circ$  assuming single-layer water clouds (effective droplet radius  $8 \mu\text{m}$ ) with cloud top at 2.3 km and cloud base at about 1.5 km. As shown in Fig. 6b, most of the CMFs derived from the measurements are within the ranges of the simulated values. The CMF at cloud base decreases from 1 to about 0.05 with increasing COT. For the actinic flux at the surface, clouds have a shielding effect. The surface albedo is relatively small, so there is not much reflection between the cloud base and surface. Therefore, the optically thicker the clouds, the less light can penetrate the clouds. At COT close to 0, at cloud base there are some  $CMF_{base}$  values of 1.4–1.5. This indicates that the actinic fluxes in cloudy scenes are larger than for the clear-sky scenes. This could happen when broken clouds are present and 3-D effects become important. Enhancement of surface actinic flux in the presence of broken clouds has been measured and explained in earlier studies (e.g. Lantz et al., 1996; Crawford et al., 2003). This enhancement can happen if the sun is not obscured by

clouds and the diffuse actinic flux of clouds is larger than the diffuse actinic flux of clear-sky  
360 (Crawford et al., 2003).

$CMF_{top}$  increases with increasing COT. The enhancement of the actinic flux at cloud top  
is caused by multiple scattering at cloud top. As illustrated by the simulated CMFs in Fig.  
6b, the CMF at cloud top has a stronger SZA dependence than the CMF at cloud base. For  
example, at COT of 40, from SZA of 30 to 60°, the CMF changes 0.7 at cloud top and 0.2 at  
365 cloud base. This may partly explain the scatter of  $CMF_{top}$ .  $CMF_{top}$  depends also on the cloud  
top height due to the height dependence of the clear-sky actinic fluxes. For the measured  
actinic flux profiles, the cloud top heights are in the range of 1–10 km. At SZA = 60°, the  
clear-sky actinic fluxes can increase by 10 % from 1 to 10 km and, consequently, the  $CMF_{top}$   
can change by 10 %. The outliers of the  $CMF_{top}$  at COT < 5 could be due to the used COT  
370 values. For small local clouds, the COT from SEVIRI may not be presentative for the clouds  
'seen' by the light sensor. This can happen because of mismatching between the SEVIRI  
pixel and balloon location or due to the subpixel clouds in the SEVIRI retrievals.

If the assumed atmospheric state, especially AOT, is not the same as the actual atmo-  
spheric state, the CMF may not be 1 at COT = 0. According to our simulations, the ratio of  
375 clear-sky actinic flux for AOT of 0.25 and of 0.34 is 1.04 at the surface, 0.98 at 1 km, and  
0.99 at 10 km for SZA = 60°. We estimate that the uncertainty in the CMF due to the AOT  
uncertainty is up to 5%.

#### 4.4 Simulations of actinic flux profile for clear-sky scenes

As shown in Fig. 4b, the AOT on 11 September 2006 varies from 0.3 to 0.5 between 11:00  
380 and 14:00 UTC. This may indicate that the AOT is also spatially variable. This is confirmed  
by MODIS AOT image which reveals a typical spatial variation in AOT between 0.2 and  
0.4 for the Netherlands (see Fig. 7a and b). The missing AOT data are mostly caused by  
presence of clouds. We used the DAK model to simulate actinic flux profiles using AOT  
values varying from 0 to 0.5. Then the simulated actinic flux profiles are interpolated to  
385 the measured AOT and SZA values on 11 September 2006. The simulated and measured  
actinic flux profiles are shown in Fig. 7c. The simulated actinic flux profile and the measured

profile are normalized at 30 km with a scaling factor of 4.13 and 1.002, respectively. The simulated actinic flux profile closely follows the shape of the measured profile, although there are some small deviations. The simulated profile consists of several profiles having different AOT because the balloon moved both vertically and horizontally. In the simulated profile, when the balloon was below 1 km, the AOT was 0.5, after the balloon moved above 1 km, the AOT is 0.34. AOT values of 0.5 and 0.34 occur at different locations, not in one vertical profile. As shown in the trajectory in Fig. 7b, the measured actinic flux profile is in fact three-dimensional, because the sonde drifts away from the launch location. This feature is more significant during cloudy conditions due to the large spatial variation of cloud optical thickness.

#### 4.5 Simulations of actinic flux profiles for fully cloudy scenes

Global irradiances derived from SEVIRI, called surface solar irradiances (SSI), are calculated using the SEVIRI COT derived from the CPP algorithm and aerosols from a monthly climatology (Greuell et al., 2013). In order to check the quality of the SEVIRI COT in the SSI calculations, we compared the SEVIRI SSI with the ground-based shortwave global irradiance measurement at De Bilt. On the 63 actinic flux profile measurement days, we compared SEVIRI SSI at 11:30 UTC with ground-based 10 min mean (11:30–11:40 UTC) pyranometer measurements of global irradiance in De Bilt. As shown in Fig. 8, the SEVIRI SSI and ground-based measurements agree well with a correlation coefficient of 0.965 and a bias of  $22 \text{ W m}^{-2}$  (for the mean SSI). This suggests that the COT from SEVIRI is of good quality and the actinic flux profiles calculated using these COTs can be realistic. The mean SSI value of 4 SEVIRI pixels and the nearest single pixel SSI are often quite different for SSI values between 200 and  $800 \text{ W m}^{-2}$ . This is caused by the inhomogeneity of clouds. The cases where  $\text{SSI} < 200 \text{ W m}^{-2}$  or  $\text{SSI} > 800 \text{ W m}^{-2}$  usually correspond to fully cloudy or clear-sky scenes, respectively. Therefore, the COT of the single SEVIRI pixel that is most close to the balloon measurement time and location may not be representative of the clouds that actually determine the actinic flux profile.



415 The cloudy actinic flux profiles are simulated for 5 single-layer water cloud cases in 2006, of which the results are shown in Figs. 9–13. The original actinic flux profiles (without normalization) have already been shown in Fig. 5. The simulations are presented in the order of complexity of the cases. We start with the simplest case on 5 September (Fig. 9), which is followed by an optically very thick cloudy case on 15 June (Fig. 10). During these launches, the COT mainly decreased during the balloon flight. On 22 June (Fig. 11) and 10 August  
420 (Fig. 12), the COT increased after launch and decreased when the balloon was at its maximum height. On 22 December (Fig. 13) the cloud layer was low and stable but the SZA was large ( $> 75^\circ$ ).

On 5 September 2006, the prevailing winds were from the west, so the balloon drifted eastwards. The SEVIRI cloud optical thickness images of 11:30 UTC and 13:00 UTC are the closest to the start and end of the ascending profile. As shown in Fig. 9a it was fully  
425 cloudy in De Bilt at 11:30 UTC when the ozone sonde was launched. The cloud optical thickness at 11:30 UTC was about 25. The clouds were optically thicker in the north-east and optically thinner south of De Bilt. At 13:00 UTC the balloon drifted towards an area with COT of about 10 (see Fig. 9b). In the first 30 min, the balloon flew over clouds having similar  
430 COT, then the clouds became thinner.

The COT for the trajectory of the balloon is shown as a function of balloon altitude in Fig. 9c. The SEVIRI pixels are selected within  $0.1^\circ$  of the balloon lat/lon box which usually has up to 5 pixels. The number of SEVIRI pixels can be less than 5 because of missing  
435 COT data (clear-sky scene or no retrieval). Please note that the COT values in the maps are averaged over  $0.1^\circ \times 0.1^\circ$  grids, while the COT values in Fig. 9c are given for every SEVIRI pixel.

If the balloon is above the cloud layer, variations in cloud optical thickness and cloud height will not produce sharp peaks in the actinic flux profile. However, all cloud variations along the flight track will show up in the actinic flux profile, because it includes the radiation  
440 from all directions and heights. The measured actinic flux profile in Fig. 9d shows two peaks, one at about 0.5 km, another at about 2 km. This indicates that there are two cloud layers, with cloud top heights at about 0.5 and 2 km, respectively. This is confirmed by the radar

measurements at Cabauw, which show one cloud layer at 0.3–0.5 km and another cloud layer at 1–2 km.

445 In the simulations the clouds are assumed to be single-layer, with cloud top height at 2 km and cloud base height at 1 km. Therefore, the simulated actinic flux profile has a single peak at 2 km. The COT values used in the simulations are shown in Fig. 9c by the black line. The simulated actinic flux profile is in good agreement with the measured flux profile. Although the measured and simulated actinic flux profiles are both normalized at 30 km, this does  
450 not influence the shapes of the profiles. The consistency between the SEVIRI SSI and the ground-based global irradiance measurements suggest that the COT at De Bilt is quite representative. During the flight, the SZA changes from 29° to 34°, which has been taken into account in the simulation. The variation in the actinic flux profile shape is mainly due to variation in the COT.

455 The COT values and the measured and simulated actinic flux profiles for 15 June 2006 are shown in Fig. 10. The mean COT was very high, 95, at the launch in De Bilt and the actinic flux profile shows a very strong peak at 1.5 km. The balloon drifted to the north, then turned west, and reached an altitude of about 30 km over a region where the COT is small. The dips in the actinic flux profile at 2 and 4 km are probably related to optically thin clouds,  
460 cloud holes, cloud shadows or clear-sky regions (Los et al., 1997).

The cloud top and base heights in simulation were at 1.5 and 0.5 km, respectively. The dips at 2 and 4 km were ignored in the simulation. As shown by the SEVIRI images of COT (Figs. 10a and b), clouds were getting thinner during the flight. The change of the measured actinic flux profile at 10 km suggests the change of COT but unfortunately there is no SEVIRI  
465 image available at this time. Therefore, the COT is extrapolated at this altitude. Above 10 km, the SEVIRI COT had to be reduced from about 30 to 20 to get a better simulation. Since the COT used in the simulation allows for closure with the measured ground-based global irradiance, the COT values used for the actinic flux profile simulation are possibly correct. In view of the uncertainty in the SEVIRI COT retrieval, the mentioned adjustment is also  
470 reasonable.

On 22 June 2006, the spatial variability in the SEVIRI COT along the balloon flight track is considerable (Figs. 11a and b). As shown in Fig. 11c, the COT decreases along the flight track and shows a large scatter. Between 10 and 20 km altitude, there are often only 2 COT values available in the 0.1-degree box in the SEVIRI image instead of 4 COT values, which suggests some clear-sky pixels. At launch the COT is about 30 and at 2 km altitude the COT decreases slightly. In the simulation, the cloud top height and base height are 3 and 2 km, respectively. When the balloon passed through the cloud top, the flux profile showed a sharp peak at 3 km. When the balloon was flying at 4 km, the COT of underlying clouds increased, forming the broad peak at 3–6 km in the actinic flux profile. The simulated actinic flux profile follows the measured actinic flux profile shape very well. The actinic flux profile below cloud top decreases in two steps, which does not occur in the simulated profile. According to the measured actinic flux profile, cloud base could be at 1 km and the cloud optical thickness increases during the ascent from 1–3 km.

On 10 August 2006 (Fig. 12), the clouds are optically thinner than the clouds described in the previous cases. According to SEVIRI, the COT is only about 15 at the launch in De Bilt. The main cloud layer is at 3 km. As can be seen from Fig. 12d, the balloon went through several cloud layers between 2 and 5 km. The balloon passed over some optically thin clouds when it was at 5–10 km altitude, then the COT increased again. The simulated cloud base and top heights are 2 and 3 km, respectively. As shown in Fig. 12d the simulated actinic flux profile follows the general shape of the measured profile and the largest peak in the measurement is being reproduced in the simulation.

The previously described cases all have SZA smaller than  $60^\circ$ , so the application of a pseudo-spherical correction is not needed. However, during the actinic flux profile measurement on 21 December 2006 (Fig. 13), the SZA values ranged between  $75^\circ$  and  $78^\circ$ , therefore the pseudo-spherical correction has to be applied in the actinic flux profile simulations. Based on radar measurements in Cabauw, the cloud layer was very close to the ground surface and the cloud top height was 1 km. Although it was winter, the surface temperature was about  $7^\circ\text{C}$  at the launch in De Bilt and there was no snow on the surface. As presented in Fig. 13a the SEVIRI COT is about 30 in De Bilt and relatively stable until

500 the balloon reaches an altitude of 20 km, where the COT starts to decrease. The pixel-to-pixel variation of COT is small. Therefore, this should be a simple case to simulate. As shown in Fig. 13d, the simulated actinic profile agrees well with the measured profile shape above 10 km and below the cloud top. However, the simulated actinic flux profile from cloud top (1.5 km) to about 10 km is much smaller than the measured actinic flux profile. As discussed in Sect. 3, for this case the wavelength dependence of actinic flux is important. The agreement between simulated and observed actinic flux profile might be improved if the simulations were performed at a fine wavelength grid and convolved with the spectral response function of the LED. The simulations should also use the solar irradiance spectrum over the LED spectral range instead of a constant value ( $1 \text{ Wm}^{-2}\text{nm}^{-1}$ ) used in this paper.

510 A challenge for the DAK model appears to be the simulation of the actinic flux profile of 6 September 2006 (see Fig. 5e). The actinic flux profile of this day is similar to the clear-sky profiles, without sharp peaks. However, the value of the actinic flux profile above cloud top is larger than for the clear-sky scenes. According to the total sky imager data at Cabauw and MODIS Terra 1 km resolution imagery, there were many small broken clouds on this day. 515 Since SEVIRI cannot resolve the small broken clouds, the COT is quite low, about 1–3. This case probably requires a 3-D radiative transfer model for correctly simulating the actinic flux profile.

The simulated actinic flux ratio (cloudy/clear-sky) (CMF) for the cloudy cases of Figs. 9–13 are presented in Fig. 14. The simulated cloudy actinic flux profiles have already been 520 shown in Figs. 9–13 after normalization. The variation of CMF in a single profile is mainly caused by the cloud optical thickness. The differences of the CMFs in the 5 cases include the impact of SZA, COT, cloud vertical distribution and cloud scattering phase matrices, because the other DAK input data are the same for these cases. As shown in Fig. 14, the CMFs are larger than 1 above the cloud top and smaller than 1 below the cloud base. The maximum CMF happens just below the cloud top and decreases almost linearly towards 525 the cloud base. Following Madronich (1987) the features in Fig. 14 can be well understood. The actinic flux above clouds is larger than the clear-sky value because the effective albedo of cloud and surface together is greater than the surface albedo alone. Since the low cloud

530 layers in Figs. 9–13 are all optically thick clouds, the incident light at cloud top cannot easily  
transmit through the cloud layer. Therefore, the CMFs are smaller than 1 below cloud base.  
Inside clouds, the peak of CMF happens in a layer below cloud top where the direct incident  
light is strongly reduced but has given rise to a diffuse downward flux while the lower part of  
the cloud is still optically thick. In this case the lower part of the clouds creates a high albedo,  
535 the top layer of the cloud reflects downward the reflected light again. These components of  
the diffuse radiation field create a strong enhancement inside the cloud.

## 5 Conclusions

A green light sensor has been developed and launched from De Bilt on an ozone sonde  
balloon between June 2006 and April 2010. Several copies of this sensor were calibrated  
to one reference sensor to make them intercomparable, although not absolutely calibrated.  
540 The calibration failed in 2007, therefore mainly the 2006 data (ascending profiles) were  
used in the analysis. The instrument is not sensitive to ozone and water vapor and it is  
temperature-stabilized. In total there were 63 launches (one ascending profile and one de-  
scending profile per launch), of which 9 launches were made during clear-sky conditions.  
The other launches were made through cloudy atmospheres. The actinic flux profiles con-  
545 tain very little noise and the impact of clouds on the actinic flux profile could be clearly  
detected.

The actinic flux profile reveals an large enhancement at the cloud top. Above the cloud  
top, the actinic flux profile is still influenced by the clouds but the enhancement is smaller  
than at the cloud top. The actinic flux profile decreases from the cloud top to cloud base,  
550 almost linearly. Below the cloud base, the actinic flux profile is relatively constant, especially  
for the low clouds. The balloon often passed through several layers of clouds. Therefore,  
the actinic flux of different heights reflects different cloud conditions along the flight track.  
The impact of cloud is quantified by the cloud modification factor (CMF). The CMF is in  
the range of 1–2.3 at the cloud top and 1–0.1 at the cloud base. The CMF depends on  
555 cloud properties and SZA. The measured CMF values agree with the DAK simulated CMF,

especially for cloud base CMF. The CMF at the cloud top is more complicated because sometimes the location of the cloud top can be ambiguous due to the horizontal movement of the balloon in the clouds or due to holes in vertical direction (broken clouds). For the calculation of the clear-sky actinic flux, the same aerosol setting has been used for all

560

The actinic flux profiles for single-layer water clouds during fully cloudy conditions were simulated using DAK. The shapes of simulated actinic flux profiles are in good agreement with the actinic flux profile measurements, except for  $\text{SZA} > 75^\circ$ . The most important input for the simulation is the COT along the flight track, which was obtained from 15 min SEVIRI observations. However, when the optical properties of the clouds are variable at the pixel-to-pixel scale, the SEVIRI COT has to be modified to get a better simulation of the actinic flux profile. Because of the good agreement between the measured and simulated actinic flux profile shapes, it would be possible to convert the measured actinic flux profile to absolute values.

565

For clear-sky conditions, the most important factors determining the shape and magnitude of the actinic flux profiles are SZA, surface albedo, aerosol optical thickness and aerosol height. Using the aerosol optical thickness data for Cabauw, the measured actinic flux profiles could be simulated reasonably well. The simulations could be improved if we would have aerosol data along the flight track of the balloon.

570

The green light sensor that we used for the actinic flux profile measurements is cheap and stable. In combination with a ground-based irradiance measurement and retrieved COT from SEVIRI, actinic flux profiles can be calculated. The light sensor is useful for the evaluation of the impact of clouds on actinic flux profiles. Applications of measured actinic flux profiles and simultaneous ozone and  $\text{NO}_2$  profiles (Sluis et al., 2010) in atmospheric chemistry models will be studied in the future.

575

580

*Acknowledgements.* We would like to thank our colleague M. van Weele (KNMI) for his comments and suggestions about the manuscript. We appreciate J. F. Meirink (KNMI) for his help with the SEVIRI CPP product, available on <http://msgcpp.knmi.nl>. We acknowledge the MODIS mission scientists and associated NASA personnel for the production of the MODIS AOT data. This research

585 is supported by NWO (Netherlands Organization for Scientific Research) through the GO project  
“Cloud radiative forcing profile and vertical energy redistribution in the atmosphere (CRAFT)”, project  
number ALW-GO-AO/09-22.

## References

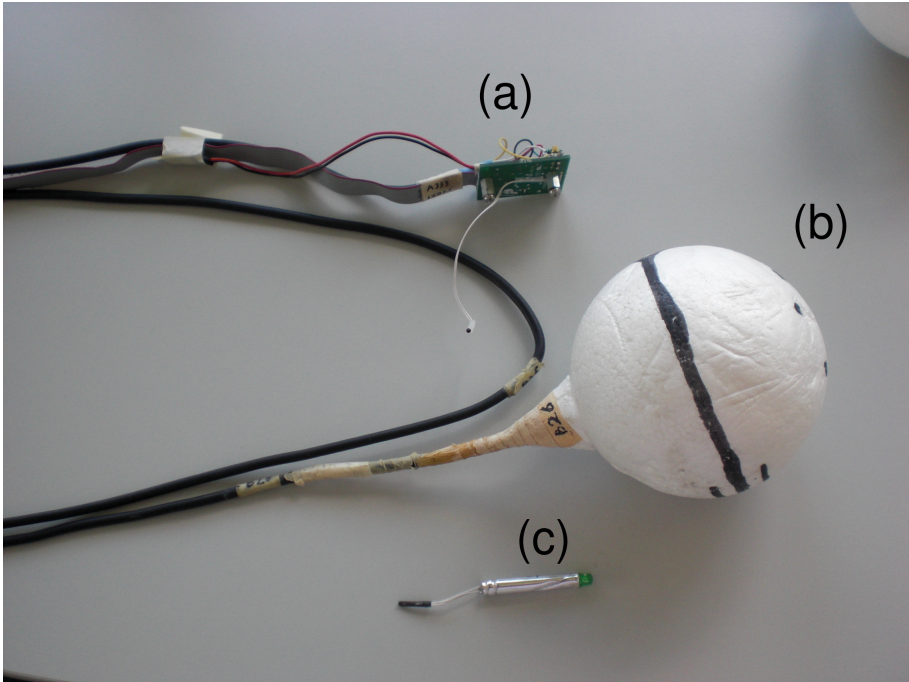
- 590 Anderson, G. P., Clough, S. A., Kneizys, F. X., Chetwynd, J. H., and Shettle, E. P.: AFGL atmo-  
spheric constituent profiles, Tech. Rep. AFGL-TR-86-0110, Air Force Geophys. Lab., Hanscom  
AFB, Mass, 1986.
- Antón, M., Alados-Arboledas, L., Guerrero-Rascado, J. L., Costa, M. J., C Chiu, J., and Olmo, F. J.:  
Experimental and modeled UV erythema irradiance under overcast conditions: the role of cloud  
optical depth, *Atmos. Chem. Phys.*, 12, 11723–11732, doi:10.5194/acp-12-11723-2012, 2012.
- 595 Bodhaine, B.A., Wood, N.B., Dutton, E.G., and Slusser, J.R.: On Rayleigh optical depth cal-  
culations, *J. Atmos. Oceanic Technol.*, 16, 1854–1861, doi: [http://dx.doi.org/10.1175/1520-0426\(1999\)016<1854:ORODC>2.0.CO;2](http://dx.doi.org/10.1175/1520-0426(1999)016<1854:ORODC>2.0.CO;2), 1999.
- Brooks, D. R. and Mims III, F. M.: Development of an inexpensive handheld LED-  
based Sun photometer for the GLOBE program, *J. Geophys. Res.*, 106(D5), 4733–4740,  
600 doi:10.1029/2000JD900545, 2001.
- Calbó, J., Pagès, D., and González, J.-A.: Empirical studies of cloud effects on UV radiation: a review,  
*Rev. Geophys.*, 43, RG2002, doi:10.1029/2004RG000155, 2005.
- Cantrell, C. A., Shetter, R. E., Calvert, J. G., Parrish, D. D., Fehsenfeld, F. C., Goldan, P. D.,  
Kuster, W., Williams, E. J., Westberg, H. H., Allwine, G., Martin, R.: Peroxy radicals as measured in  
605 ROSE and estimated from photostationary state deviations, *J. Geophys. Res.*, 98, 18355–18366,  
doi:10.1029/93JD01794, 1993.
- Crawford, J., Shetter, R. E., Lefer, B., Cantrell, C., Junkermann, W., Madronich, S., and  
Calvert, J.: Clouds and trace gas distributions during TRACE-P, *J. Geophys. Res.*, 108, 8818,  
doi:10.1029/2002JD003177, D21, 2003.
- 610 Crutzen, P. J. and Zimmermann, P. H.: The changing photochemistry of the troposphere, *Tellus B*,  
43, 136–151, doi:10.1034/j.1600-0889.1991.t01-1-00012.x, 1991.
- De Haan, J. F., Bosma, P. B., and Hovenier, J. W.: The adding method for multiple scattering com-  
putations of polarized light, *Astron. Astrophys.*, 183, 371–391, 1987.

- 615 De Roode, S. R., Duynkerke, P. G., Boot, W., and Van der Hage, J. C. H.: Surface and tethered-balloon observations of actinic flux: effects of arctic stratus, surface albedo, and solar zenith angle, *J. Geophys. Res.*, 106, 27497–27507, doi:10.1029/2001JD900236, 2001.
- Eskes, H., van Velthoven, P., Valks, P., and Kelder, H.: Assimilation of GOME total ozone satellite observations in a three-dimensional tracer transport model, *Q. J. Roy. Meteor. Soc.*, 129, 1663–1681, doi:10.1256/qj.02.14, 2003.
- 620 Greuell, W., Meirink, J. F., and Wang, P.: Retrieval and validation of global, direct, and diffuse irradiance derived from SEVIRI satellite observations, *J. Geophys. Res.-Atmos.*, 118, 2340–2361, doi:10.1002/jgrd.50194, 2013.
- Hofzumahaus, A., Kraus, A., Kylling, A. and Zerefos, C.: Solar actinic radiation (280–420 nm) in the cloud-free troposphere between ground and 12 km altitude: Measurements and model results, *J. Geophys. Res.*, 107, doi=10.1029/2001JD900142, 2002.
- 625 Johnston, H. S., personal communication (1997) to E.-P. Röth, R. Ruhnke, G. Moortgat, R. Meller, and W. Schneider, Berichte des Forschungszentrums Jülich, jül-3340, 1997. ([http://satellite.mpic.de/spectral\\_atlas/cross\\_sections/Ozone/O3\\_Johnston\(1997\)\\_298K\\_408-599nm.txt](http://satellite.mpic.de/spectral_atlas/cross_sections/Ozone/O3_Johnston(1997)_298K_408-599nm.txt))
- Junkermann, W.: Measurements of the  $J(O^1D)$  actinic flux within and above stratiform clouds and above snow surfaces, *Geophys. Res. Lett.*, 21, 793–796, doi:10.1029/93GL03498, 1994.
- 630 Kazadzis, S., Bais, A. F., Balis, D., Zerefos, C. S., and Blumthaler, M.: Retrieval of downwelling UV actinic flux density spectra from spectral measurements of global and direct solar UV irradiance, *J. Geophys. Res.*, 105, 4857–4864, 2000.
- Kylling, A., Albold, A., Seckmeyer, G.: Transmittance of a cloud is wavelength-dependent in the UV-range: Physical interpretation, *Geophys. Res. Lett.*, 24, 397–400, DOI: 10.1029/97GL00111, 1997.
- Kylling, A., Danielsen, T., Blumthaler, M., Schreder, J., and Johnsen, B.: Twilight tropospheric and stratospheric photodissociation rates derived from balloon borne radiation measurements, *Atmos. Chem. Phys.*, 3, 377–385, doi:10.5194/acp-3-377-2003, 2003.
- 640 Kylling, A., Webb, A. R., Kift, R., Gobbi, G. P., Ammannato, L., Barnaba, F., Bais, A., Kazadzis, S., Wendisch, M., Jäkel, E., Schmidt, S., Kniffka, A., Thiel, S., Junkermann, W., Blumthaler, M., Silbernagl, R., Schallhart, B., Schmitt, R., Kjeldstad, B., Thorseth, T. M., Scheirer, R., and Mayer, B.: Spectral actinic flux in the lower troposphere: measurement and 1-D simulations for cloudless, broken cloud and overcast situations, *Atmos. Chem. Phys.*, 5, 1975–1997, doi:10.5194/acp-5-1975-2005, 2005.
- 645

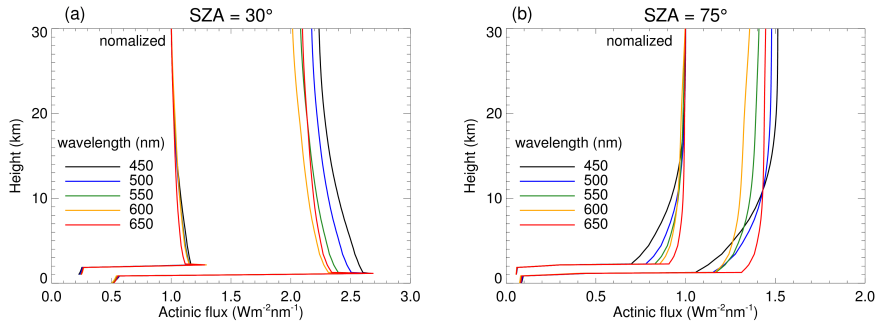


- Lantz, K. O., Shetter, R. E., Cantrell, C. A., Flocke, S. J., Calvert, J. G. and Madronich, S.: Theoretical, actinometric, and radiometric determinations of the photolysis rate coefficient of NO<sub>2</sub> during the Mauna Loa Observatory Photochemistry Experiment 2, *J. Geophys. Res.*, 101(D9), 14613–14630, doi:10.1029/96JD00215, 1996.
- 650 Los, A., van Weele, M., and Duynkerke, P. G.: Actinic fluxes in broken cloud fields, *J. Geophys. Res.*, 102, 4257–4266, doi:10.1029/96JD03123, 1997.
- Madronich, S.: Photodissociation in the atmosphere 1. Actinic flux and the effect of ground reflections and clouds, *J. Geophys. Res.*, 92, 9740–9752, doi:10.1029/JD092iD08p09740, 1987.
- 665 Mannschreck, K., Gilge, S., Plass-Duelmer, C., Fricke, W., and Berresheim, H.: Assessment of the applicability of NO-NO<sub>2</sub>-O<sub>3</sub> photostationary state to long-term measurements at the Hohenpeisenberg GAW Station, Germany, *Atmos. Chem. Phys.*, 4, 1265–1277, doi:10.5194/acp-4-1265-2004, 2004.
- Mateos, D., Pace, G., Meloni, D., Bilbao, J., di Sarra, A., de Miguel, A., Casasanta, G., and Min, Q.: Observed influence of liquid cloud microphysical properties on ultraviolet surface radiation, *J. Geophys. Res.-Atmos.*, 119, 2429–2440, doi:10.1002/2013JD020309, 2014.
- 660 Mayer, B., Kylling, A., Madronich, S., and Seckmeyer, G.: Enhanced absorption of UV radiation due to multiple scattering in clouds: experimental evidence and theoretical explanation, *J. Geophys. Res.*, 103, 31241–31254, doi:10.1029/98JD02676, 1998.
- McKenzie, R. L., Johnston, P. V., Smile, D., Bodhaine, B. A., and Madronich, S.: Altitude effects on UV spectral radiance deduced from measurements at Lauder, New Zealand, and Mauna Loa Observatory, Hawaii, *J. Geophys. Res.*, 106, 22845–22860, doi:10.1029/2001JD900135, 2001.
- Palancar, G. G., Shetter, R. E., Hall, S. R., Toselli, B. M., and Madronich, S.: Ultraviolet actinic flux in clear and cloudy atmospheres: model calculations and aircraft-based measurements, *Atmos. Chem. Phys.*, 11, 5457–5469, doi:10.5194/acp-11-5457-2011, 2011.
- 670 Roebeling, R. A., Feijt, A. J., and Stammes, P.: Cloud property retrievals for climate monitoring: implications of differences between Spinning Enhanced Visible and Infrared Imager (SEVIRI) on Meteosat-8 and Advanced Very High Resolution Radiometer (AVHRR) on NOAA-17, *J. Geophys. Res.*, 111, D20210, doi:10.1029/2005JD006990, 2006.
- Schwander, H., Koepke, P., Kaifel, A., and Seckmeyer, G.: Modification of spectral UV irradiance by clouds, *J. Geophys. Res.*, 107, 4296, doi:10.1029/2001JD001297, 2002.
- 675 Schiller, C., Hofzumahaus, A., Müller, M., Klein, E., Röth, E.-P., and Schmidt, U.: Ultraviolet actinic flux in the stratosphere: An overview of balloon-borne measurements during EASOE, 1991/92, *Geophys. Res. Lett.*, 21, 1239–1242, 1994.

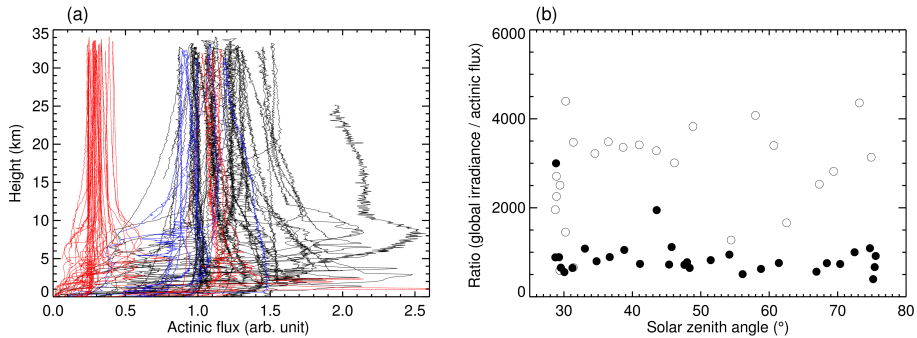
- 680 Seckmeyer, G., Erb, R., and Albold, A.: Transmittance of a cloud is wavelength-dependent in the  
UV-range, *Geophys. Res. Lett.*, 23, 2753–2755, 1996.
- Sluis, W. W., Allaart, M. A. F., Pitters, A. J. M., and Gast, L. F. L.: The development of a nitrogen  
dioxide sonde, *Atmos. Meas. Tech.*, 3, 1753–1762, doi:10.5194/amt-3-1753-2010, 2010.
- Stammes, P.: Spectral radiance modelling in the UV-Visible range, in: *IRS 2000: Current Problems in  
Atmospheric Radiation*, edited by: Smith, W. L. and Timofeyev, Y. M., A. Deepak Publ., Hampton,  
685 VA, 385–388, 2001.
- Stammes, P., de Haan, J. F., and Hovenier, J. W.: The polarized internal radiation field of a planetary  
atmosphere, *Astron. Astrophys.*, 225, 239–259, 1989.
- Wang, P., Knap, W. H., Kuipers-Munneke, P., and Stammes, P.: Clear-sky shortwave radiative closure  
for the Cabauw Baseline Surface Radiation Network site, the Netherlands, *J. Geophys. Res.*, 114,  
690 D14206, doi:10.1029/2009JD011978, 2009.
- Wang, P., Knap, W. H., and Stammes, P.: Cloudy-sky shortwave radiative closure for a Baseline  
Surface Radiation Network site, *J. Geophys. Res.*, 116, D08202, doi:10.1029/2010JD015141,  
2011.
- Van Weele, M.: Effect of clouds on ultraviolet radiation, photodissociation rates of chemical species  
in the troposphere, Ph.D. thesis, University of Utrecht, Utrecht, the Netherlands, 1996.
- Van Weele, M., Vilà-Guerau de Arellano, J., and Kuik, F.: Combined measurements of UV-A actinic  
flux, UV-A irradiance and global radiation in relation to photodissociation rates, *Tellus B*, 47, 353–  
364, doi:10.1034/j.1600-0889.47.issue3.6.x, 1995.
- Vilà-Guerau de Arellano, J., Duynderke, P. G., and van Weele, M.: Tethered-balloon measure-  
ments of actinic flux in a cloud-capped marine boundary layer, *J. Geophys. Res.*, 99, 3699–3705,  
700 doi:10.1029/93JD03090, 1994.



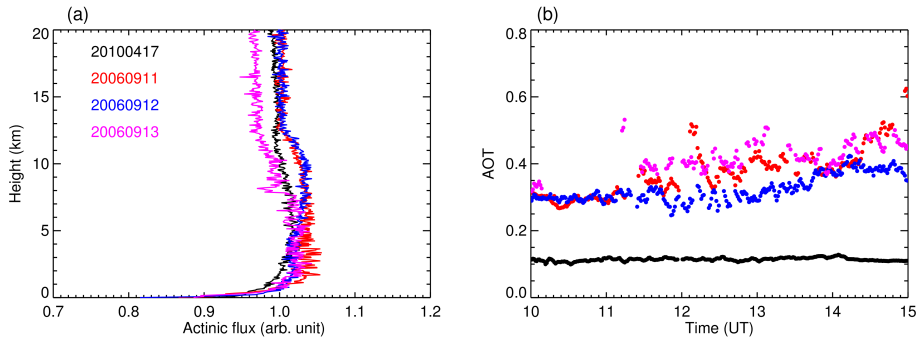
**Figure 1.** Photo of the light sensor: (a) data transmitter, (b) sphere, (c) green LED.



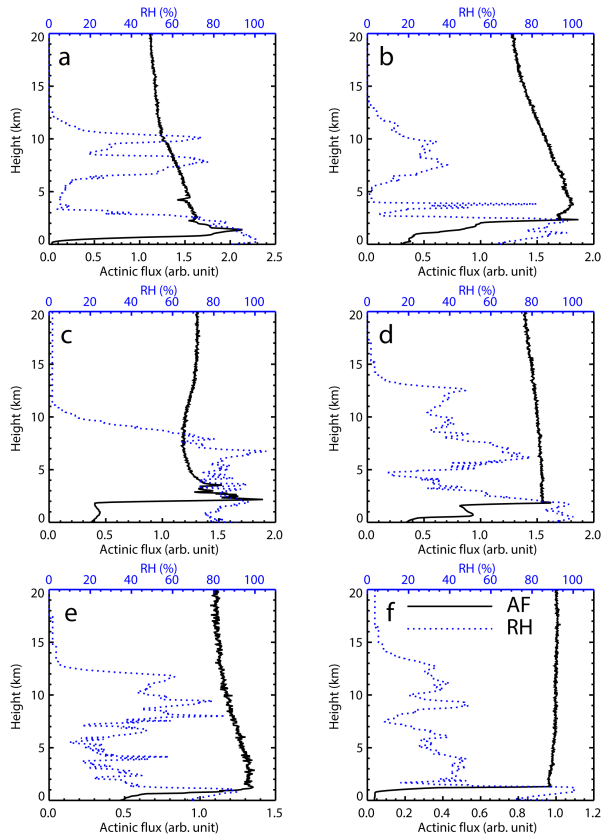
**Figure 2.** Simulated actinic flux profiles at wavelengths: 450, 500, 550, 600, 650 nm at SZA = (a) 30 and (b) 75 °. The cloud top is at 1.3 km, cloud base is at 0.8 km, cloud optical thickness is 30. Particle size of the water clouds is 10  $\mu\text{m}$ . The aerosol optical thickness is 0.18 at 550 nm and surface albedo is 0.15. At every wavelength the normalized actinic flux profile is the simulated actinic flux profile divided by the actinic flux at 30 km. The normalized actinic fluxes are shifted 1 km up in the y-axis.



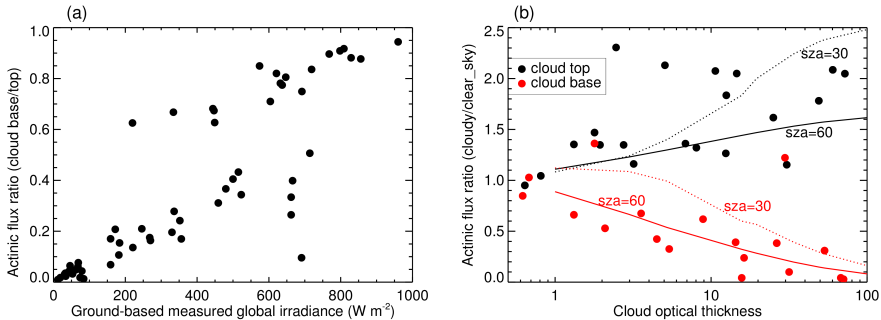
**Figure 3.** (a) Measured actinic flux profiles in 2006 (black lines), 2007 (red lines), 2008 and 2010 (blue lines). (b) Ratio between the global irradiance at the ground measured at 11:30 UTC at De Bilt and the actinic flux profile measurement at 4 m height at 11:30 UTC at De Bilt. Results for 2006 are marked as filled circles, results for 2007 are in open circles.



**Figure 4.** (a) Clear-sky actinic flux profiles measured on 11–13 September 2006 and 17 April 2010. (b) AOT at 501 nm measured in Cabauw for the same days. The profiles are the original measurements, not normalized.

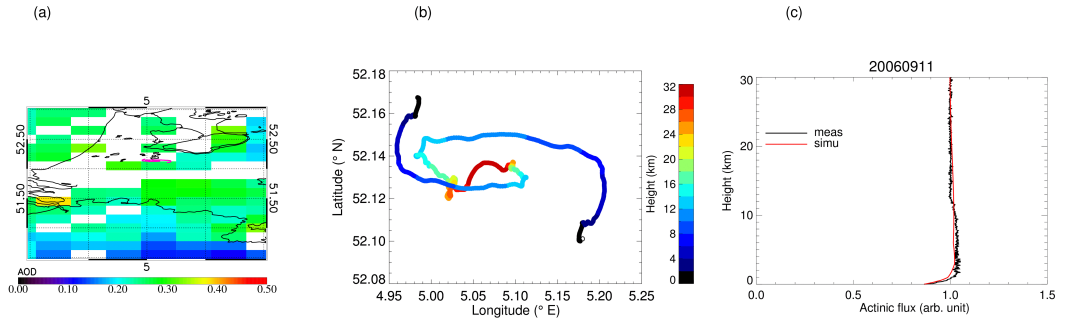


**Figure 5.** Actinic flux (AF) and relative humidity (RH) profiles for cloudy scenes measured on **(a)** 15 June 2006, SZA = 28.8°, **(b)** 22 June 2006, SZA = 28.7°, **(c)** 10 August 2006, SZA = 36.7°, **(d)** 5 September 2006, SZA = 45.4°, **(e)** 6 September 2006, SZA = 45.7°, and **(f)** 21 December 2006, SZA = 75.6°. The profiles are ascending measurements and are not normalized. The SZA value is at 11:30 UTC, the start of the profile.

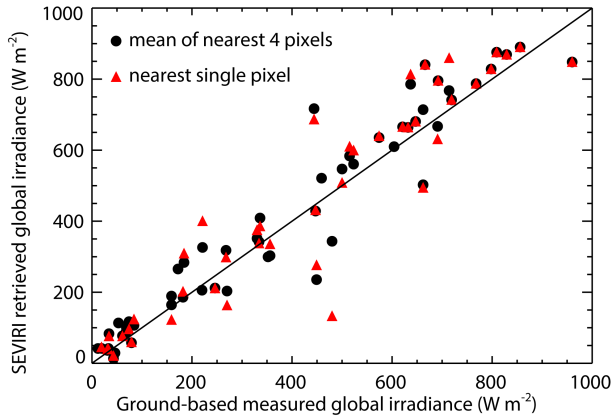


**Figure 6.** (a) Ratio between the measured actinic fluxes at cloud base and at cloud top vs. the measured global irradiance at the surface, for all data. (b) Ratio between measured actinic fluxes of cloudy and clear-sky scenes (CMF) at cloud top (black) and cloud base (red) vs. SEVIRI cloud optical thickness at 11:30 UTC for data in 2006. The dots are measurements, the lines are simulations for CMF at cloud top and cloud base for SZA = 30 and 60 °.

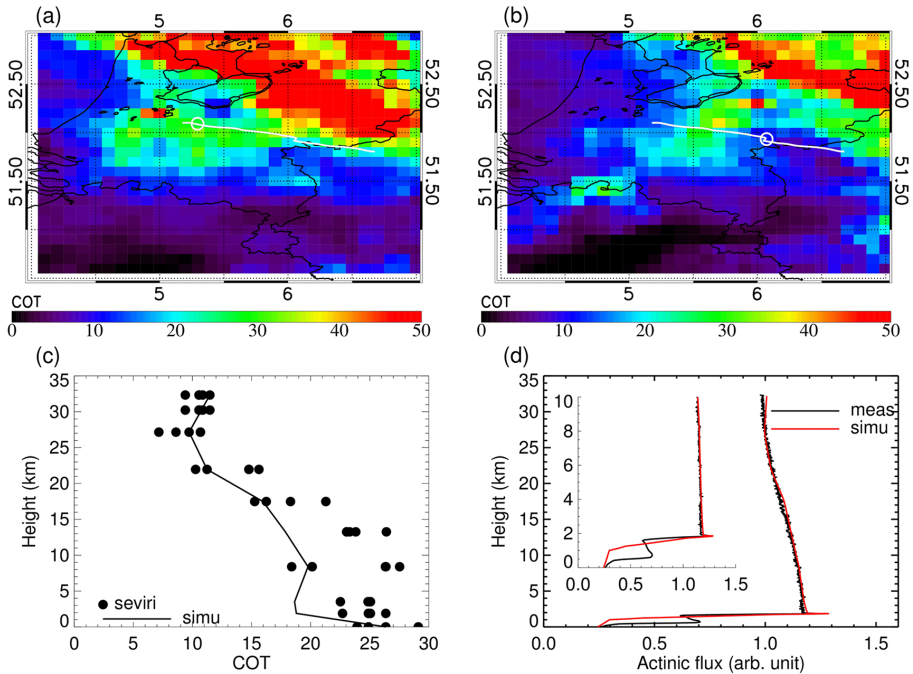




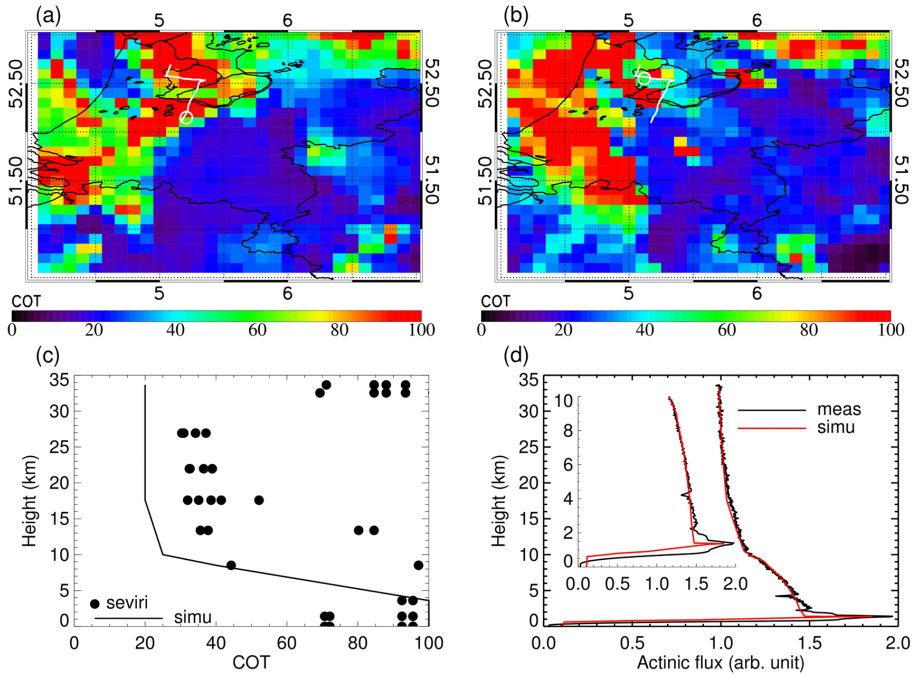
**Figure 7.** Clear-sky case on 11 September 2006. **(a)** MODIS aerosol optical thickness image at 11:45 UTC, the purple curve shows the trajectory of the balloon. **(b)** Trajectory of the balloon, with its height indicated by color. The location of De Bilt is marked with a circle. **(c)** Measured actinic flux profile and simulated actinic flux profile.



**Figure 8.** Scatter plot of SEVIRI retrieved global irradiance at the surface (surface solar irradiance, SSI) vs. ground-based measured global irradiance at 11:30 UTC at De Bilt on all 63 actinic flux profile measurement days. The black line is the one-to-one line. The black dots indicate the mean SSI of all SEVIRI pixels in a  $0.1^\circ \times 0.1^\circ$  (latitude  $\times$  longitude) grid box. The red triangles indicates the nearest single pixel SSI in the grid box.



**Figure 9.** Cloudy sky case on 5 September 2006. **(a)** SEVIRI cloud optical thickness image at 11:30 UTC (balloon launch) and **(b)** at 13:00 UTC (balloon maximum height). The flight track of balloon is indicated as a white line. The location of balloon at the time the SEVIRI image is taken is indicated as a white circle. **(c)** SEVIRI single pixel cloud optical thickness values in  $0.1^\circ \times 0.1^\circ$  (latitude  $\times$  longitude) boxes along the trajectory of the balloon (dots) as a function of the height of the balloon. The COT values used in the simulations are connected by the black line. **(d)** Measured actinic flux profile and simulated actinic flux profile (both normalized at 30 km altitude, respectively). The actinic flux profile is also shown zoomed-in at 0–10 km.



**Figure 10.** Same as Fig. 9 but for 15 June 2006.

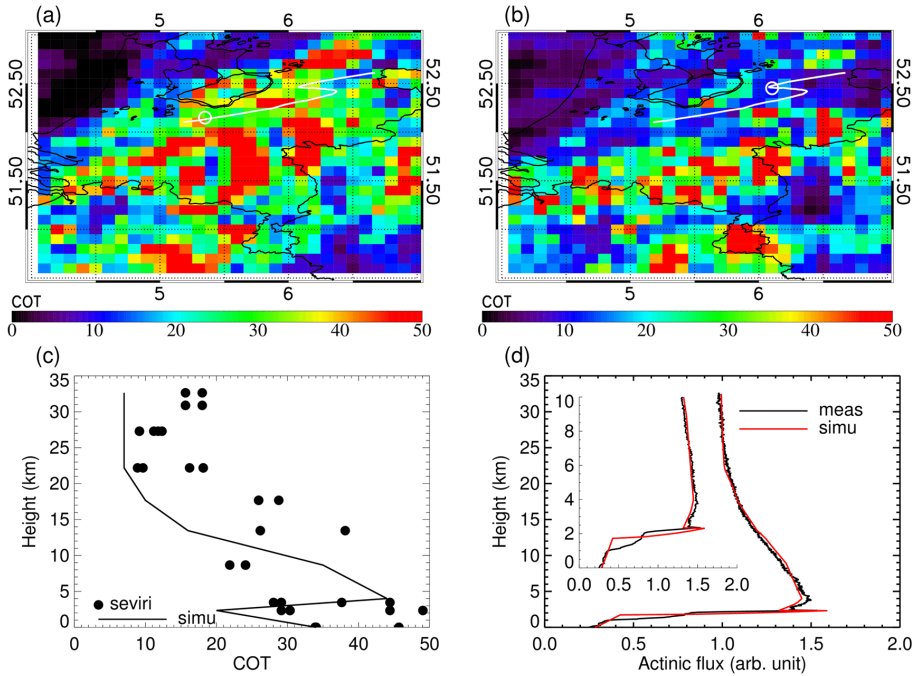


Figure 11. Same as Fig. 9 but for 22 June 2006.

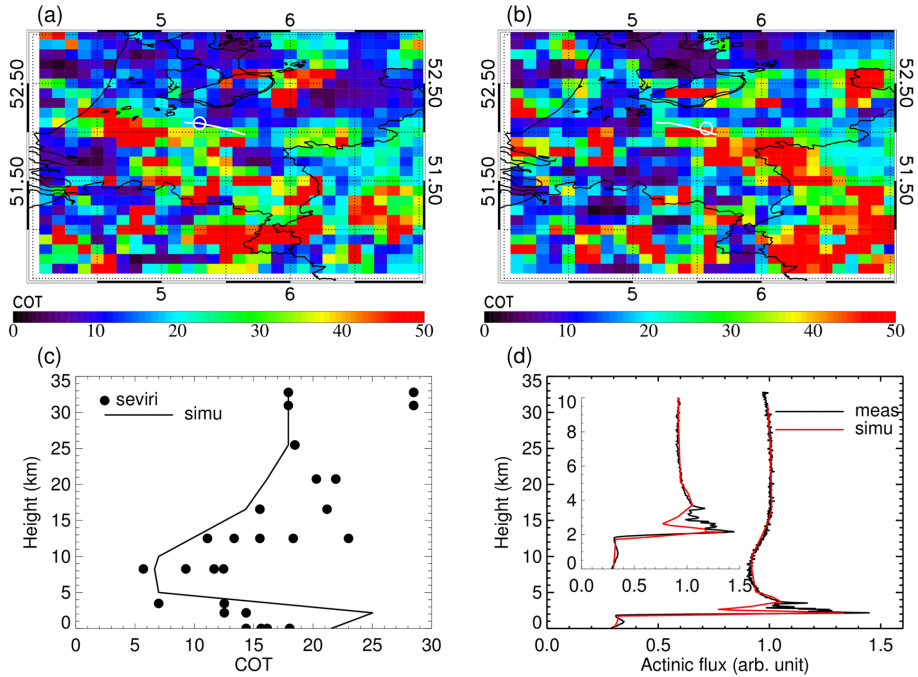
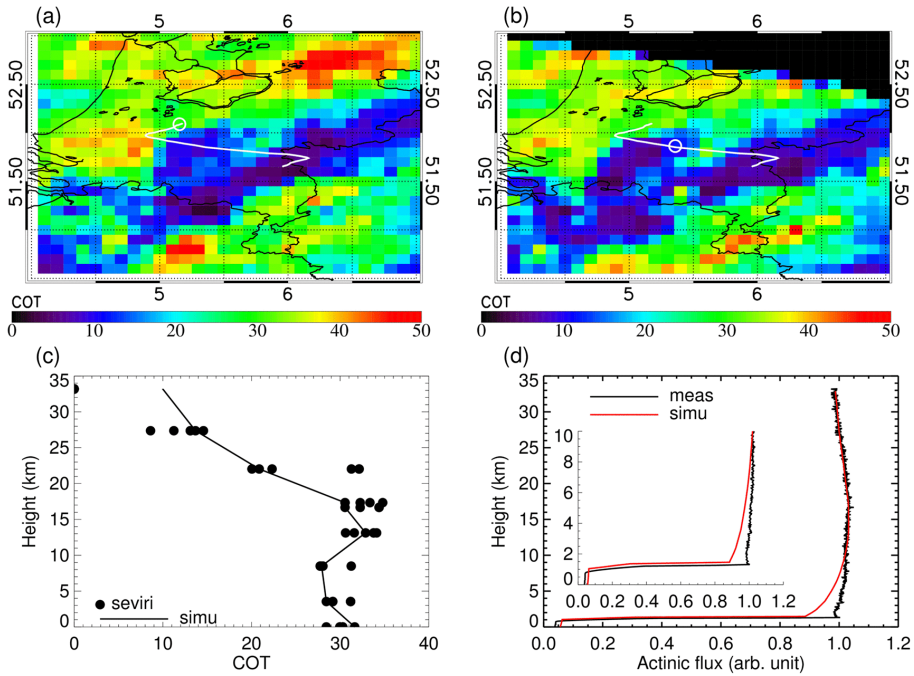
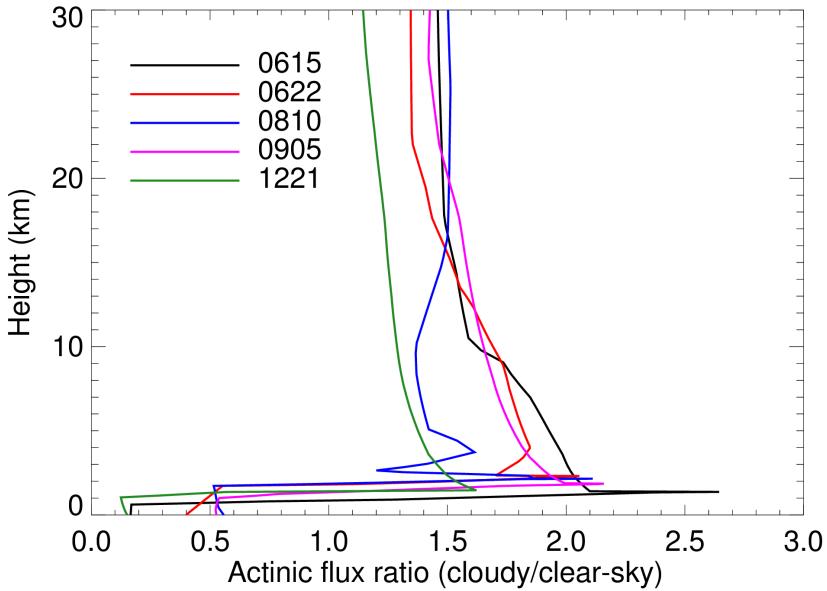


Figure 12. Same as Fig. 9 but for 10 August 2006.



**Figure 13.** Same as Fig. 9 but for 21 December 2006. **(b)** The SEVIRI image was acquired at 12:45 UTC instead of 13:00 UTC.



**Figure 14.** Simulated actinic flux ratio (cloudy/clear-sky) for the cloudy cases in Figs. 9-13.