

Response to interactive comment on “Atmospheric oxidation of isoprene and 1,3-butadiene: influence of aerosol acidity and relative humidity on secondary organic aerosol” by M. Lewandowski, M. Jaoui, J.H. Offenberg, J.D. Krug, and T.E. Kleindienst

Commentator A. Nenes

Received and published: 1 December 2014

The interpretation of the experimental results relies heavily on the ability of $[H^+]_{air}$ to capture the in-situ pH of atmospheric particles. The methodology of Surrat et al.(2007) involves extracting filter samples in water, measuring its pH and then back-calculating the amount of H^+ in the original aerosol (assuming that H^+ is conserved). This is largely equivalent to an “ion balance” method, because the H^+ measured corresponds to an aerosol sample diluted by orders of magnitude, so that the ions present are largely dissociated (this is affected by the degree of dilution used, but does not change the inherent nature of the methodology). However, two studies currently in discussion in ACPD (Guo et al., 2014; Hennigan et al., 2014) raise doubts on whether such methods can be used as a proxy for in-situ aerosol pH, as it is shown that H^+ from ion balance is not correlated with in-situ aerosol pH for a wide range of aerosol conditions. This is especially relevant for this study as chamber RH is varied; the latter affects the amount of in-situ aerosol liquid water (hence pH) in a way that is not reflected in $[H^+]_{air}$. Even if a constant RH were used as in Surrat et al.(2007), the variable chemical composition (and potentially phase state) of the particles across experiments can still have a nontrivial effect on aerosol liquid water content, hence in-situ pH.

With the above said, we are not implying that the study here is fundamentally flawed. However, it would be good to see the acidity interpretation accompanied by both $[H^+]_{air}$ and pH calculations from a thermodynamic model (e.g., AIM, ISORROPIA, MESA) applied to the chamber conditions. Guo et al.(2014) and Hennigan et al.(2014) have shown that models applied to complex SE US and Mexico City aerosol, even at relatively low RH, can provide a much better representation of aerosol pH than ion-balance methods, provided that they are applied in “forward mode”. This is what we recommend here as well to be done. In the absence of gas-phase inorganic component measurements, one can still apply thermodynamic models in a useful manner (e.g.,Guo et al.,2014).

We generally agree with the reviewer’s assessment of the limitations of the $[H^+]_{air}$ measurement. Although we believe that the limitations inherent in the measurement may not be as pronounced under the controlled conditions of a laboratory experiment as they are under ambient sampling conditions, it remains at best an indirect measurement of the aerosol properties of interest. We have retained the $[H^+]_{air}$ metric in the early part of the paper's results (isoprene/ SO_2 and 1,3-butadiene vs $[H^+]$) in part for consistency with previous studies, especially the isoprene/ SO_2 comparison to Surratt et al 2007. While these experiments could certainly benefit from additional information regarding the actual aerosol pH levels, we do not believe that this information is vital to the comparisons presented.

In the varied RH experiments, where $[H^+]_{air}$ measurements are of limited value (since the maximum dissociated H^+ in the extracts remains unchanged, but actual aerosol pH is expected to change with liquid water content), we agree with the reviewer that aerosol pH levels or aerosol liquid water concentrations would be of tremendous value to the interpretation of the results. Unfortunately, we do not have sufficient composition information to do the modeling with ISORROPIA or AIM appropriately. While chamber temperature, RH, and particle sulfate loading are known for each reaction step, particle phase ammonium and nitrate were not measured in these experiments. And, although not strictly necessary, no gas-phase ammonia or

nitric acid concentrations are available (and, as high-NO_x experiments, nitric acid concentrations should be non-trivial), further complicating model predictions. Additionally, as with the SOAS field measurements reported in Guo et al (2014), we have a significant aerosol fraction composed of isoprene-related organic aerosol to contend with, which can further contribute to aerosol phase water content, but is not accounted for in the models.

Given these limitations, we believe that any modeled pH levels or aerosol liquid water concentrations that we could generate from our existing data would be suspect and potentially counterproductive to the analysis (due more to our under-analysis of the aerosol composition than to any limitations inherent in the models). While we readily admit that humidity provides, at best, an indirect measure of the physically important aerosol parameters of interest in these comparisons, it is nevertheless the most reliable measurement surrogate that we have to work with in this data set. As such, we have retained it in the comparisons, but have added additional text to the Methods section further clarifying the limitations of both $[H^+]_{air}$ measurement in lieu of aerosol pH, and the use of humidity level as a surrogate for aerosol liquid water content (Line 165):

“While this method provides a simple, easily repeatable measure of bulk acidity, it does not fully capture the actual acidity of individual aerosol particles, which is more likely to be of physical significance in these chemical systems. It is also of limited value in experiments where the relative humidity is varied, as the extraction of the collected aerosol effectively masks the effects of changing particulate liquid water concentrations. Further limitations of the $[H^+]_{air}$ measurement techniques have been described in detail in Hennigan et al. (2014). While a number of methods have been developed to measure aerosol liquid water content directly or estimate it through the use of thermodynamic models such as ISORROPIA (Fountoukis and Nenes, 2007) or AIM (Wexler and Clegg, 2002), liquid water measurements were not available for this study, and insufficient aerosol compositional information was collected for accurate use of thermodynamic modeling.”

We have also added a Summary section which includes an additional review of the limitations of the experiments presented in the paper, including areas requiring additional investigation in future work, which includes discussion of the need for aerosol liquid water or pH assessments in future work (line 470):

“While these experiments are suggestive, they also include a number of shortcomings that need to be addressed in future work. Perhaps the most significant is the use of absolute humidity and $[H^+]_{air}$ as surrogate measures of aerosol liquid water content and aerosol pH. Determination of the effective pH in the aerosol particles through the application of thermodynamic models, such as ISORROPIA or AIM, should provide a more realistic assessment of actual acidity than the $[H^+]_{air}$ approach, provided adequate gas and particle composition data is obtained experimentally. However, even these models generally account for only the influence of inorganic species, while the presence of isoprene SOA products has been reported to also contribute significantly to water uptake on ambient aerosols (Guo et al., 2014). This suggests that for laboratory experiments with high organic aerosol concentrations, particularly from isoprene-related parent hydrocarbons, direct measurements of aerosol liquid water content may be required.”

Anonymous Referee #1

Received and published: 9 January 2015

Do the authors think that water vapour plays a role in the VOC oxidation chemistry leading to a lower SOC yield at an elevated humidity level or is it more related to particle phase chemistry such as a lower yield organosulfates, enhanced hydrolysis or reduced absorptive partitioning of organic compounds due to lower organic mass? Can the authors discuss this in the manuscript?

Given the relatively limited organic compositional analysis available in these experiments, it is difficult to assess the potential changes in VOC chemistry occurring as RH is varied. As no significant changes in NO_x or ozone concentrations were observed, and that the delta hydrocarbon remained essentially constant at all RH levels, it would appear that at least the early generations of VOC oxidation chemistry remained essentially unchanged (in contrast, the changes in the yield of the SO₂ to sulfate reactions with changing RH did result in observable changes in NO_x, O₃, and delta SO₂ concentrations, although this data was not presented in detail in the paper). Reductions in the formation of oligomeric species or organosulfates in a more water-rich environment seem plausible, but we currently lack the detailed organic analyses needed to back up these ideas. Further experiments are pending on these topics. Additional discussion added to the manuscript at Line 371:

“The reason for this reduction in SOC formation is not entirely clear. Gas-phase NO_x and O₃ concentrations do not appear to change significantly as a function of humidity level, as does the concentration of isoprene consumed in the reactions. This suggests that early-generation gas-phase oxidation reactions are probably not altered significantly by changing humidification. Changes in the aerosol liquid water content may affect the gas-particle partitioning of later-generation isoprene oxidation products, or increased water content may affect particle-phase organosulfate formation or the formation of oligomeric species (Pye et al., 2013). Further analysis of gas- and particle-phase organic constituents is required to further investigate this effect.”

Pp. 29441 Line 21 onwards: Is the sulfate acidity effect purely related to the reactivity of oxidation products to acidic sulfate? I imagine that the layer of SOA formed on the seed particle likely hinders acid catalysed accretion reactions and absorptive partitioning largely controls the subsequent SOA formation. In other words, VOCs with higher ΔM_{SOA} may not show a strong acidity effect. How do the placements of these VOC relate to ΔM_{SOA} from neutral seed experiments?

Although interface and/or transport effects may impose limits on acid catalyzed reactions at some concentration level, no clear indications were observed in these experiments. As noted in Section 3.1, variations in the particulate sulfate loading in the isoprene/SO₂ experiments (due to the concentration of SO₂ employed in different stages) produced linear OC vs [H⁺] behavior essentially identical to previous experiments conducted with constant sulfate concentrations generated with nebulized sulfate aerosol. In addition, Offenberget al (2009) previously considered different organic loading with α -pinene SOC with no obvious differences in the results. It is possible that the organic to sulfate ratios used in these experiments simply do not reach a critical threshold at which interface effects would become limiting. However, it is also possible that there are additional factors, such as aerosol phase and/or mixing state, which would negate the limiting interface conditions of the core/shell structure suggested by the reviewer. The need for further examination of the physical characteristics of the particle phase are now included in the Summary section (Line 486):

“Further examination of the mixing state and phase of the generated aerosols would also like be of value, given that the available interfacial area could affect interactions

between H^+ , sulfate, and SOA, which may in turn affect organosulfate formation or other pathways to increased SOA yield.”

Pp. 29442 Line 19 onwards: Temperature data should be given in Tables 1 and 2. How large was the temperature differences between these experiments? If the temperature had a significant impact on the RH, it is likely that the temperature also had an influence on the ΔM_{SOA} . Can the authors discuss about this in the manuscript?

The referenced text applies to the relative humidity variation experiments, which are detailed in Figure 4 (not Figures 1 and 2). Figure 4 has been modified to include temperature and relative humidity ranges for the described experiments, in addition to the absolute humidity ranges already provided.

Pp. 29442 Line 25 and Pp. 29443 Line 21 onwards: There are several thermodynamic models available to calculate aerosol water content and pH values. The authors may want to use one of these tools to estimate the values for these experiments and add them to Tables 1 and 2, and discuss if these parameters had an influence on acid catalysed accretion reactions or not.

See response to comments from A. Nenes, above.

Pp. 29444 Line 27: I would like the authors to discuss a bit deeper about the structural differences of oxidation products here. What sort of oxidation products do the authors expect when there is an additional methyl group in a precursor VOC? Does this lead to more reactive intermediate oxidation products towards acidic sulfate?

Given the current length of the paper, and the additional analytical techniques (GC-MS and/or LC-MS) that would need to be incorporated in order to provide a detailed discussion of the organic compositions of the isoprene and 1,3-butadiene systems, we have chosen to not include a more detailed discussion of the oxidation products here. We have instead qualified that this proposed pathway is speculative at this time, included a recent reference to 1,3-butadiene reaction mechanisms in general, and have suggested that this topic will need further discussion in a future manuscript. See Line 334 in Results and Discussion:

“This could represent a substituent effect which influences the sensitivity of the gas-phase precursors to reaction by the acidic sulfate nucleophile, but further organic analysis of the aerosol phase constituents would be required to examine this possibility in detail. Although recent studies have compared the reaction pathways and products formed for 1,3-butadiene oxidation versus isoprene oxidation (Jaoui et al., 2014), these studies did not focus on acid-influenced reactions or organosulfate formation.”

and line 493 in the Summary:

“A further examination of the similarities and differences between acid-influenced OC formation in the isoprene and 1,3-butadiene systems would also likely be valuable. A more rigorous organic analysis of product distributions may help reveal why isoprene appears to be significantly more sensitive to acidic conditions than 1,3-butadiene at low humidity levels, and whether this difference is due to structural effects related to the additional methyl group affecting gas-phase chemistry, particle-phase organosulfate formation, or due to other phenomena, such as differences in volatility and partitioning of oxidized intermediates. A more detailed comparison of isoprene and 1,3-butadiene

organic chemistry under acidic conditions may help with the development of more accurate mechanisms for inclusion in air quality models.”

As another reviewer commented, the authors should discuss their findings in atmospheric relevant context at the end of the manuscript.

A Summary section has been added to the paper, which incorporates a discussion of the relevance of these studies to atmospheric conditions, as well as a thorough assessment of the current limitations of the experiments presented (Lines 445-469):

“These experiments support previous studies suggesting that acidic aerosol can lead to increased SOA formation from the photooxidation of isoprene under laboratory conditions. Changing the source of the acidity from nebulized inorganic aerosol to a more atmospherically relevant photochemical conversion of SO₂ into acidic sulfate aerosol nearly identical results as previous nebulized sulfate aerosol experiments. In addition, 1,3-butadiene, a chemically similar compound released from primarily anthropogenic sources, was also demonstrated to produce higher concentrations of SOA under acidic conditions, albeit to a lesser extent than was seen with isoprene. The humidity experiments further suggest that aerosol liquid water content can have a substantial effect on SOA formation from isoprene and 1,3-butadiene. Increasing humidity produces a notable reduction in SOA formation in both the isoprene and 1,3-butadiene photochemical systems, which is more pronounced in both systems in the presence of acidic inorganic aerosols, and was most pronounced for the acidified isoprene/NO system.

In the isoprene/NO photochemical systems examined in this study, SOA enhancement due to the presence of acidic inorganic aerosol was observed to be negligible at absolute humidity levels above approximately 11 g H₂O m⁻³. This lower SOA enhancement at elevated humidities may explain, in part, the difficulties in detecting increased SOA formation under acidic conditions in field studies of ambient air masses, particularly in humid climates like the southeastern US. This work suggests that a more detailed understanding of the role of humidity and of aerosol liquid water content is likely required in order to accurately predict the impact of acidity-influenced oxidation chemistry on overall SOA yields. While the data presented here may suggest that enhanced SOA formation via acid-influenced pathways is more constrained than previous studies may have suggested, it does still appear to represent a viable pathway for additional SOA formation from a number of precursor hydrocarbons, which may need to be incorporated into air quality models in order to accurately estimate secondary PM concentrations in certain locations.

While these experiments are suggestive, they also include a number of shortcomings that need to be addressed in future work. Perhaps the most significant is the use of absolute humidity and [H⁺]_{air} as surrogate measures of aerosol liquid water content and aerosol pH. Determination of the effective pH in the aerosol particles through the application of thermodynamic models, such as ISORROPIA or AIM, should provide a more realistic assessment of actual acidity than the [H⁺]_{air} approach, provided adequate gas and particle composition data is obtained experimentally. However, even these models generally account for only the influence of inorganic species, while the presence of isoprene SOA products has been reported to also contribute significantly to water uptake on ambient aerosols (Guo et al., 2014). This suggests that for laboratory experiments with high organic aerosol concentrations, particularly from isoprene-

related parent hydrocarbons, direct measurements of aerosol liquid water content may be required.

Additionally, the relative humidity experiments presented here consider predominantly systematic increases in relative humidity. As some phenomena related to aerosol liquid water content are known to display hysteresis, such as deliquescence of sulfate aerosols, further testing. Experiments incorporating descending relative humidities and larger humidity steps are warranted, particularly along with the inclusion of liquid water content measurements. Further examination of the mixing state and phase of the generated aerosols would also like be of value, given that the available interfacial area could affect interactions between H^+ , sulfate, and SOA, which may in turn affect organosulfate formation or other pathways to increased SOA yield.

Finally, further research is needed to examine changes in organic composition triggered by the effects considered in this work. In particular, the mechanism through which humidity level affects OC production from isoprene, even under non-acidified conditions, could be important to the selection of appropriate SOC yields in air quality models. A further examination of the similarities and differences between acid-influenced OC formation in the isoprene and 1,3-butadiene systems would also likely be valuable. A more rigorous organic analysis of product distributions may help reveal why isoprene appears to be significantly more sensitive to acidic conditions than 1,3-butadiene at low humidity levels, and whether this difference is due to structural effects related to the additional methyl group affecting gas-phase chemistry, particle-phase organosulfate formation, or due to other phenomena, such as differences in volatility and partitioning of oxidized intermediates. A more detailed comparison of isoprene and 1,3-butadiene organic chemistry under acidic conditional may help with the development of more accurate mechanisms for inclusion in air quality models.”

Pp. 29432 Line 2 and elsewhere: I feel that $[H_3O^+]_{seed}$ or $[H_3O^+]_{aerosol}$ is more appropriate for expressing the aerosol acidity. It is very unlikely that free H^+ exists in the air under the experimental conditions used by the authors.

Although we agree that the $[H^+]_{air}$ terminology, generally defined as the dissociable H^+ in the aerosol per cubic meter of air sampled, can be misleading and subject to misinterpretation, it does have prior established use in the literature. In order to maintain consistency with these prior publications, the terminology has been retained here.

Anonymous Referee #2

Received and published: 24 December 2014

1) A major point of concern in this manuscript, as identified by the authors and further detailed in the short comment by A. Nenes, is the measurement of $[H^+]$ and its relationship with particle acidity and relative humidity. While the method used provides a valid bulk measurement, it is not clear whether bulk $[H^+]$ is particularly atmospherically relevant when other factors (e.g. particle water content) can influence acidity. In light of these concerns, I would suggest that the authors find addition methods, such as calculating particle acidity with a thermodynamic model, for approaching this intractable issue. As A. Nenes describes, such models have been shown to give reasonable estimates of aerosol pH, and when used in conjunction with measured bulk $[H^+]$, these estimates will provide a clearer interpretation of acidity in the particle phase. Of particular concern, in light of the difficulty of measuring particle acidity, are experiments in which multiple factors influencing acidity change together, such as the relative humidity experiments detailed in this manuscript. For reasons discussed in the manuscript, it can be hard to know how much the change in SOC yield reflects variations in humidity directly, or indirectly by changes in particle acidity, which may be changing despite a constant $[H^+]$ due to differences in particle water content. Employing thermodynamic models to estimate particle pH would be particularly useful for these experiments.

[See response to comments from A. Nenes, above.](#)

Additionally, considering the difficulty of separating the effects of humidity and acidity in these experiments, discussion of further controls would be beneficial in Section 3.3. For example, the authors mention temperature differences between experiments (L320); could these temperature differences also have an effect on SOC yields? Also, when the relative humidity was stepped up and down in these experiments, could hysteresis effects influence SOC yield, by which particles formed at one humidity and then brought to another humidity have different organic content than particles formed at the second humidity level (e.g. by irreversible particle-phase reactions)? A discussion of these temperature and hysteresis effects, and particularly any control experiments run to investigate these effects, would provide valuable insight into the many factors at play in these humidity experiments.

[Although it is difficult to say with certainty how strongly the SOC yields may be impacted by temperature variations, we believe that the effects should be relatively small in the directly compared data sets presented in the paper. In the 1,3-butadiene/humidity experiments, the temperature differences between the two runs are approximately 3°C, with the seed experiment having the higher temperature. In the isoprene/humidity experiments, the overall temperatures were higher than the 1,3-butadiene experiments \(by approximately 2 and 5 degrees C\), but consistent with one another to within about 1°C. This may account for some of the differences in the observed behavior of the two systems, such as the greater separation of the butadiene curves at elevated humidity levels, but other sources of experimental error are likely to have a greater overall effect \(as suggested by the scatter in the OC and yield plots\) than the temperature differences. However, since converting from relative to absolute humidity could correct for the roughly 10-15% difference in water concentrations produced by these temperature differences, we felt that incorporating this correction improve the comparability of the datasets. We have adjusted the text to more directly provide the temperatures at which the different experiments have been conducted \(in various locations\), and have included an explicit mention of the possibility of temperature differences affecting the 1,3-butadiene RH experiments \(Line 436\):](#)

“Additionally, the temperature difference between the two experiments, although relatively small (approximately 3 °C on average) may be sufficient to introduce differences in the gas-particle partitioning between the two experiments.”

With regards to hysteresis effect, in all described experiments the chamber system was allowed to equilibrate for at least 4 residence times following any change to the reaction conditions. This should allow for near-complete (>95%) turnover of all products formed under the previous conditions, which should minimize or eliminate any hysteresis effects from the causes suggested by the reviewer. However, other sources of hysteresis, such as deliquescence behavior of the inorganic sulfate, are more difficult to rule out. A limited number of additional data points were collected to attempt to test for hysteresis due to the direction or magnitude of the RH changes. However, due to the challenges of these long-running experiments, not enough data was obtained to produce a statistically significant assessment. Text has been added to the Summary section identifying this as an outstanding experimental issue warranting further study (Line 482):

“Additionally, the relative humidity experiments presented here consider predominantly systematic increases in relative humidity. As some phenomena related to aerosol liquid water content are known to display hysteresis, such as deliquescence of sulfate aerosols, further testing. Experiments incorporating descending relative humidities and larger humidity steps are warranted, particularly along with the inclusion of liquid water content measurements.”

2) Some discussion is needed at the end of the manuscript about the atmospheric relevance and implications of the results. Do these experiments provide any insight into field observations, and their persistent disparities from chamber experiments regarding the dependence of SOC yield on particle acidity? What effects might these humidity and acidity dependencies have on particle SOC yields in atmospheric conditions where isoprene and 1,3-butadiene are found?

A Summary section has been added to better summarize outstanding experimental issues related to this study and potential atmospheric implications for ambient SOA formation, particularly for isoprene (Lines 445-469). The text is provided above in the response to Reviewer #1.

3) Minor clarification questions about methods and instrumentation:

L151 – In the discussion of measuring $[H^+]$, some detail on the subject of uncertainty (e.g. the error bars shown on Figures 1, 2, 3, and 5) would be useful. How are these errors estimated, and what factors are (and are not) included?

L165 – A similar treatment of uncertainty in the discussion of particle organic carbon measurements would be useful as well.

Error bars for $[H^+]$ were based on the rated variability of the pH probe used, converted into $[H^+]$ units. Error bars on the OC concentrations were based on the variability in the replicate semi-continuous OC measurements conducted during each sampling period (typically, $n > 20$). Error bars on yield values (requested below) incorporate variability in replicated hydrocarbon measurements and OC measurements during each sampling period. Error bars on absolute humidity were based on variability in the measured temperature and RH data, converted to absolute humidity on a 5-minute basis throughout each sampling period ($n > 200$). Text has been added following the presentation of Figure 1 (Line 253), Figure 3 (Line 363), and Figure 4 (Line 387) clarifying the basis for the error bars presented.

L214 – What is the $[SO_2]$ background?

As given in Table 1, a background SO₂ concentration of 11 ppb was detected in the absence of SO₂ addition to the chamber. This reading was likely due to a small hydrocarbon interference on the SO₂ monitor, as this stage of the experiment was conducted last and the monitor's scrubbers had been subjected to significant hydrocarbon, SO₂, and aerosol concentrations in prior stages.

L579 – I believe figure 4 should have error bars similar to those in Figures 1,2,3 and 5.

Figure 4 has been revised to include error bars on both axes.

General – additional details that would provide useful insight include how SO₂ and particle sulfate were measured; the temperatures at which experiments were conducted; and whether any oxidant source was added (and how much) to initiate the oxidation of the hydrocarbons.

Additional clarification has been added to the methods section to identify the SO₂ monitor employed (Line 147):

“NO and NO_y were measured with a TECO model 42C (Franklin, MA) oxides of nitrogen chemiluminescent analyzer, SO₂ was monitored by pulsed fluorescence detection (TECO, Model 43A), and O₃ was measured with a chemiluminescent ozone monitor (Bendix Model 8002, Lewisburg, WV).”

and the measurement of particle sulfate (Line 157):

“Extracts were analyzed for sulfate (SO₄²⁻) ions using a Dionex DX500 Ion Chromatography system equipped with an electrical conductivity detector. Anion analysis was conducted using a Dionex IonPac AS14A column and an isocratic 4 mM sodium carbonate/0.5 mM sodium bicarbonate eluent.”

Clarification of experiment temperatures for the isoprene/SO₂ experiment and the 1,3-butadiene vs. [H⁺] experiment have been added to the text, while Table 4 has been updated to include temperature ranges for the RH experiments, as requested above by Reviewer #1. Only NO was used as an oxidant in the experiments.

Technical corrections: The manuscript is largely free of typographical and grammatical errors.

L90 – an extra space in “by anthropogenic”

Corrected

L109 – “of” should be “in”

Corrected

Edited Manuscript: All changes made to the text of the manuscript have been highlighted in yellow.

1 Atmospheric Oxidation of Isoprene and 1,3-Butadiene: Influence of Aerosol
2 Acidity and Relative Humidity on Secondary Organic Aerosol

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11 **Abstract**

12 The effects of acidic seed aerosols on the formation of secondary organic aerosol (SOA)
13 have been examined in a number of previous studies, several of which have observed strong
14 linear correlations between the aerosol acidity (measured as nmol H⁺ per m³ air sample volume)
15 and the percent change of secondary organic carbon (SOC). The measurements have used several
16 precursor compounds representative of different classes of biogenic hydrocarbons including
17 isoprene, monoterpenes, and sesquiterpenes. To date, isoprene has displayed the most
18 pronounced increase in SOC, although few measurements have been conducted with
19 anthropogenic hydrocarbons. In the present study, we examine several aspects of the effect of
20 aerosol acidity on the secondary organic carbon formation from the photooxidation of 1,3-
21 butadiene, as well as extending the previous analysis of isoprene.

22 The photooxidation products measured in the absence and presence of acidic sulfate
23 aerosols were generated either through photochemical oxidation of SO₂ or by nebulizing
24 mixtures of ammonium sulfate and sulfuric acid into a 14.5 m³ smog chamber system. The
25 results showed that, like isoprene and β-caryophyllene, 1,3-butadiene SOC yields linearly
26 correlate with increasing acidic sulfate aerosol. The observed acid sensitivity of 0.11 %SOC
27 increase per nmol m⁻³ increase in H⁺ was approximately a factor of three less than that measured
28 for isoprene. The results also showed that the aerosol yield decreased with increasing humidity

29 for both isoprene and 1,3-butadiene, although to different degrees. Increasing the absolute
30 humidity from 2 to 12 g m⁻³ reduced the 1,3-butadiene yield by 45% and the isoprene yield by
31 85%.

32

33 **1. Introduction**

34

35 The role of aerosol acidity to increase formation of secondary organic aerosol (SOA) in
36 the atmosphere continues to be a topic of considerable debate. Field studies at ground level have
37 indicated that increases in ambient secondary organic carbon (SOC) due to ambient acidity are
38 likely subtle. Zhang et al. (2007) examined increases of SOA species in the Pittsburgh area under
39 acidic conditions, found at most a 25% increase in ambient SOA from the Pittsburgh area that
40 could be attributed to acid catalyzed effects. In another study from the SEARCH network,
41 Tanner et al. (2009) report low apparent impacts to aerosol acidity at the rural sites at Yorkville,
42 GA and Centreville, AL, where biogenic hydrocarbons and anthropogenic oxidants from nearby
43 urban centers might be expected to produce relatively high levels of aerosol acidity in the
44 presence of the oxidation products of biogenic hydrocarbons.

45 Most laboratory studies aimed at addressing the impact of aerosol acidity on SOA
46 concentrations have focused on isoprene. Emissions of isoprene (C_5H_8) from vegetation
47 constitute the greatest worldwide source of nonmethane hydrocarbons (Guenther et al., 1995).
48 SOC formation from isoprene has been shown to increase in the presence of sulfate acidity in
49 smog chamber experiments (Edney et al., 2005; Surratt et al., 2007), with a variety of
50 organosulfate compounds detected in the aerosol phase (Surratt et al., 2008; Surratt et al., 2010).

51 The effect of acidity to produce organosulfates has been studied mainly for aerosols with
52 strong biogenic inputs. Surratt et al. (2007) initially showed that sulfate esters were formed in the
53 aerosol products from photooxidations of isoprene and α -pinene in the presence of acidic seed
54 aerosol. These products were then compared to those found in ambient aerosol collected at
55 ground sites in the Southeast U.S. (i.e., the SEARCH network) and found to be similar to the

56 laboratory aerosol (Jaoui et al., 2008). Additional studies (Froyd et al., 2010) showed that
57 products of isoprene oxidation could render a single organosulfate compound (IEPOX-sulfate),
58 which comprised up to 3% of the organic aerosol mass under some conditions in the free
59 troposphere.

60 On a broader basis, laboratory studies have readily shown that acidic sulfate aerosol
61 produces increased organic aerosol yields from the products of biogenic and anthropogenic
62 oxidation systems (e.g., Jang et al., 2002). Since the initial studies, efforts have been undertaken
63 to quantify the magnitude of the aerosol acidity effect. Surratt et al. (2007) investigated the effect
64 of sulfate acidity on photooxidation products from the isoprene/NO_x system. They found that
65 secondary organic carbon increases linearly with aerosol acidity, [H⁺]_{air}, an acidity measure that
66 gives its air concentration (nmol m⁻³) rather than an aerosol pH. Offenberg et al. (2009) extended
67 this same analysis to examine the acidity effects on monoterpenes (α-pinene) and sesquiterpenes
68 (β-caryophyllene). For α-pinene aerosol products, the effect of acidity was found to be
69 independent of organic carbon mass present and was a factor of eight lower than the effect for
70 isoprene. The β-caryophyllene aerosol products, by contrast, showed an effect similar to that for
71 isoprene and a factor of five higher than that for α-pinene. Analysis by Chan et al. (2011)
72 confirmed the presence of organosulfate compounds in β-caryophyllene SOA formed under these
73 conditions. Zhang et al. (2012) performed acidity experiments for 2-methyl-3-buten-2-ol
74 (MBO), a compound structurally related to isoprene. MBO was shown to be less influenced by
75 acidity than isoprene or β-caryophyllene, but more affected than α-pinene. However, this
76 comparison is complicated by the fact that the MBO experiments were conducted under dry
77 conditions using the photolysis of hydrogen peroxide to generate OH radicals; in contrast, Surratt

78 et al. (2007) and Offenberg et al. (2009) relied upon NO_x photochemistry conducted at 30%
79 relative humidity to generate their data.

80 Concentrations of the isoprene SOA tracer products, 2-methylthreitol and 2-
81 methylerythritol, have also been found to rise with increased aerosol sulfate acidity. These
82 results suggested that particle phase reactions could contribute to the increased isoprene aerosol
83 yields and compound concentrations. Mechanisms for C₅ and C₁₀ organosulfate formation in the
84 atmosphere have been proposed (Surratt et al., 2008). Subsequent studies by Paulot et al. (2009)
85 gave strong evidence that the atmospheric formation of isoprene sulfates under conditions of low
86 nitrogen oxides involved a stable gas-phase C₅-hydroperoxide epoxide. Once uptake of the
87 epoxide into acidified aerosol occurs, inorganic sulfate nucleophiles were able to convert the
88 epoxide to organosulfates, and hydrolysis led to the formation of the 2-methyl tetrols, depending
89 on the competitive rates of different nucleophiles in the aerosol. However, a recent study by Lin
90 et al. (2013) reports measurements made in Chapel Hill, NC, an area impacted by anthropogenic
91 oxidant emissions, that show epoxide formation also occurs through NO_x channel reactions. In
92 these reactions, methylacryloylperoxy nitrate (MPAN), an intermediate stable product from
93 isoprene oxidation, reacts with OH radicals leading to methyl acrylic epoxide (MAE).

94 While considerable effort has been expended studying acidic effects of biogenic
95 precursors, far less effort has been made to examine such effects on hydrocarbons having an
96 anthropogenic origin. An interesting anthropogenic compound for consideration is 1,3-butadiene
97 (C₄H₆). The main source for this compound is from automotive exhaust emissions, although
98 additional sources from cigarette smoke, evaporative emissions of gasoline, and from biomass
99 combustion have been reported (Anttinen-Klemetti et al., 2006; Dollard et al., 2001; Eatough et
100 al., 1990; Hurst, 2007; Pankow et al., 2004; Penn and Snyder, 1996; Sorsa et al., 1996; Thornton-

101 Manning et al., 1997; Ye et al., 1998). It has been classified as a hazardous compound in the
102 1990 Clean Air Act Amendments (US EPA, 1996), a carcinogenic and toxic pollutant, and a
103 genotoxic chemical in humans and other mammals (Acquavella, 1996; US EPA, 2002). With
104 respect to aerosol formation, 1,3-butadiene is also of interest as a structural analog for isoprene.
105 SOA formation from 1,3-butadiene has been examined in a number of recent studies (Angove at
106 al., 2006; Sato, 2008; Sato et al., 2011; Jaoui et al., 2014), although with only limited
107 consideration of the effects of aerosol acidity.

108 The main focus of the present study is to explore some additional aspects of the role of
109 acidic sulfate aerosol **in** the formation of SOA from isoprene and 1,3-butadiene. For isoprene, we
110 examine the increase of SOA using acidic sulfate derived from the photooxidation of SO₂ to see
111 if the results are consistent with those using nebulized acidic sulfate seed aerosol. In addition, we
112 have measured the extent to which the isoprene analog – 1,3-butadiene – also shows an increase
113 in SOA formation in the presence of acidic aerosol. The results are then compared to biogenic
114 compounds previously studied to determine the relative magnitudes of the effect. In addition,
115 this study attempts to extend the analysis over a broader range of humidities in an effort to assess
116 the impact of aerosol water content on acidic influenced SOA formation. In the previous studies
117 by Surratt et al. (2007) and Offenberg et al. (2009), all measurements were conducted at a single
118 humidity level (30% relative humidity), while Zhang et al. (2012) examined only dry conditions.
119 Extending these studies to a wider range of hydrocarbons and across a more realistic range of
120 humidities should provide data of greater atmospheric relevance and contribute to further
121 development of acidity-influenced SOA chemistry in air quality models.

122

123

124 2. Experimental

125

126 Secondary organic aerosol was generated in a 14.5 m³ fixed-volume, Teflon-coated
127 reaction chamber. The chamber used a combination of UV-fluorescent bulbs that provided
128 radiation from 300-400 nm with a distribution similar to that of solar radiation to the extent that
129 can be achieved with UV bulbs (Kleindienst et al., 2006). The reaction chamber was operated as
130 a continuous stirred tank reactor having a residence time of 4 h, to produce a constant, steady-
131 state aerosol distribution which could be repeatedly sampled at different seed aerosol acidities.

132 To supply isoprene and 1,3-butadiene, high concentration gas mixtures were produced in
133 high-pressure cylinders diluted with nitrogen (N₂). Tank concentrations were approximately
134 2000 ppm for isoprene and 4500 ppm for 1,3-butadiene. The hydrocarbons, NO, and SO₂ (when
135 used) were added through flow controllers into the inlet manifold, where they were diluted and
136 mixed prior to introduction into the chamber. Inorganic aerosol was added to the chamber by
137 nebulizing dilute aqueous solutions of ammonium sulfate and/or sulfuric acid (TSI, Model 9302,
138 Shoreville, MN), with total sulfate concentration of the combined solution held constant in order
139 to maintain stable inorganic concentrations in the chamber. The seed aerosol stream then passed
140 through a ⁸⁵Kr neutralizer (TSI, Model 3077, Shoreville, MN) and equilibrated to the computer-
141 controlled relative humidity designated for a particular experiment. To change the acidity of the
142 seed aerosol, the ratio of the ammonium sulfate and sulfuric acid solutions was changed to
143 produce a constant aerosol sulfate concentration (typically ~30 μg m⁻³) across the range of
144 acidities used.

145 Concentrations of isoprene and 1,3-butadiene in the inlet manifold and chamber were
146 measured using a gas chromatograph with flame ionization detection (Hewlett-Packard, Model

147 5890 GC). NO and NO_y were measured with a TECO model 42C (Franklin, MA) oxides of
148 nitrogen chemiluminescent analyzer, SO₂ was monitored by pulsed fluorescence detection
149 (TECO, Model 43A), and O₃ was measured with a chemiluminescent ozone monitor (Bendix
150 Model 8002, Lewisburg, WV). Temperature and relative humidity were measured with an
151 Omega Digital Thermo-Hydrometer (Model RH411, Omega Engineering, Inc., Stamford, CT).

152 Aerosol samples were collected on 47 mm Teflo membrane filters (Pall Corporation, Ann
153 Arbor, MI) for determination of the particulate sulfate concentration (in select experiments) and
154 the aerosol hydrogen ion concentration per unit volume of air sampled, or [H⁺]_{air}, expressed as
155 nmol H⁺ m⁻³. Aerosol produced in the chamber was collected at a rate 10 to 20 L min⁻¹ over a
156 period of approximately 4 h. Filters were extracted by sonication for 30 min using 10 mL of
157 distilled, deionized water in a 50 mL polypropylene vial. Extracts were analyzed for sulfate
158 (SO₄²⁻) ions using a Dionex DX500 Ion Chromatography system equipped with an electrical
159 conductivity detector. Anion analysis was conducted using a Dionex IonPac AS14A column and
160 an isocratic 4 mM sodium carbonate/0.5 mM sodium bicarbonate eluent.

161 For [H⁺]_{air} determination, the extract were allowed to cool to room temperature and the
162 pH of each extract was measured with a temperature-compensated Oakton 300 series
163 pH/conductivity meter (OAKTON Instruments, Vernon Hills, IL). The [H⁺]_{air} was calculated by
164 dividing the measured aqueous concentration of hydrogen ions by the volume of air collected, as
165 described by Surratt et al. (2007). While this method provides a simple, easily repeatable
166 measure of bulk acidity, it does not fully capture the actual acidity of individual aerosol particles,
167 which is more likely to be of physical significance in these chemical systems. It is also of
168 limited value in experiments where the relative humidity is varied, as the extraction of the
169 collected aerosol effectively masks the effects of changing particulate liquid water

170 concentrations. Further limitations of the $[H^+]_{air}$ measurement techniques have been described in
171 detail in Hennigan et al. (2014). While a number of methods have been developed to measure
172 aerosol liquid water content directly or estimate it through the use of thermodynamic models
173 such as ISORROPIA (Fountoukis and Nenes, 2007) or AIM (Wexler and Clegg, 2002), liquid
174 water measurements were not available for this study, and insufficient aerosol compositional
175 information was collected for accurate use of thermodynamic modeling. Nevertheless, in the
176 absence of a true aerosol pH measurement, the $[H^+]_{air}$ approach appears to provide a useful, if
177 limited, surrogate measure under sufficiently constrained experimental conditions.

178 Measurements of particulate organic carbon were performed with an on-line thermal
179 optical transmittance carbon analyzer using a parallel plate, carbon strip denuder (Sunset
180 Laboratories, Tigard, OR; Birch and Cary, 1997) prior to aerosol collection on the quartz filter
181 within the instrument. Other details of operation for the carbon analyzer on the photochemical
182 reaction chamber are described elsewhere (Offenberg et al., 2007). The duty cycle for this
183 measurement was 0.75 h (i.e., 0.5 h sampling and 0.25 h analysis times, respectively). All
184 particulate carbon concentrations measured during the interval of aerosol acidity filter collections
185 were averaged for comparison with the integrated measurements of aerosol acidity.

186 Four different sets of experiments were performed, each involving multiple stages: (1) an
187 isoprene/NO experiment in which different concentrations of SO_2 were used to generate varied
188 levels of aerosol acidity, (2) a 1,3-butadiene/NO experiment in which different nebulizer
189 solutions were used to generate varied levels of aerosol acidity, (3) a pair of isoprene/NO
190 experiments, one using a low concentration ammonium sulfate seed and the other using an acidic
191 inorganic component, in which the inorganic compositions were held constant while the

192 humidity levels were varied, and (4) a comparable pair of 1,3-butadiene/NO experiments in
193 which humidity levels were systematically varied.

194 In the isoprene/NO experiment (ER370), the initial mixture of isoprene, NO, and SO₂
195 was irradiated in the chamber until the reaction mixture reached a steady-state concentration. For
196 each of the three successive stages, the SO₂ concentration was progressively reduced and the
197 reaction mixture was allowed to equilibrate. In the final stage, SO₂ was turned off to generate a
198 “base case” aerosol from the isoprene/NO_x reaction alone. In all cases, filter measurements were
199 conducted only after the steady-state condition was achieved.

200 For the 1,3-butadiene/NO experiment (ER444), an ammonium sulfate solution was used
201 to generate approximately 35 μg m⁻³ of inorganic aerosol to provide a base case. In subsequent
202 stages, the seed aerosol was made progressively more acidic by reducing the proportion of
203 ammonium sulfate and adding increasing fractions of sulfuric acid to the solution. This approach
204 offers two main advantages over the SO₂ oxidation method described above. First, it provides a
205 consistent level of inorganic sulfate aerosol at all stages; in contrast, the SO₂ oxidation produces
206 variable inorganic concentrations, and effectively no inorganic content in the base case without
207 SO₂ addition. Second, the addition of the seed aerosol should have a negligible effect on the gas-
208 phase radical chemistry, which may otherwise be affected by the conversion of SO₂ to sulfuric
209 acid.

210 For the humidity studies, each hydrocarbon was examined with two different
211 experiments. First, each hydrocarbon/NO system was tested at multiple humidity levels using
212 only a low concentration (1 μg m⁻³) ammonium sulfate seed aerosol (ER666 for 1,3-butadiene;
213 ER667 for isoprene). This provided a base case for exploring the changes in SOC formation and
214 aerosol yield in the absence of significant aerosol acidity. Relative humidities were varied

215 between roughly 10% and 60%, which corresponded to absolute humidities of approximately 2
216 to 14 g m⁻³. For isoprene, this base case experiment was then repeated in the presence of a
217 moderately acidic sulfate aerosol, which was held constant across the full range of humidities
218 examined (ER662). For 1,3-butadiene, a more acidic inorganic aerosol, generated using a
219 solution incorporating a higher fraction of sulfuric acid solution to ammonium sulfate solution,
220 was employed (ER444).

221 3. Results and Discussion

222

223 The experiments presented here support previous studies suggesting that acidic aerosol
224 can lead to increased SOA formation from the photooxidation of isoprene under laboratory
225 conditions. Changing the source of the acidity from nebulized inorganic aerosol to a more
226 atmospherically relevant photochemical conversion of SO₂ into acidic sulfate aerosol produced
227 only a minor change in the resulting percent increase in SOC per unit increase in [H⁺]_{air}. In
228 addition, 1,3-butadiene, a chemically similar compound released from primarily anthropogenic
229 sources, was also demonstrated to produce higher concentrations of SOA under acidic
230 conditions. The humidity experiments further suggest that humidity level, and likely aerosol
231 liquid water content, can have a substantial effect on SOA formation from isoprene and 1,3-
232 butadiene. Increasing humidity produces a notable reduction in SOC formation in both the
233 isoprene and 1,3-butadiene photochemical systems. However, this reduction is more pronounced
234 in both systems in the presence of acidic inorganic aerosols, and was most pronounced for the
235 isoprene/NO system.

236

237 3.1 Isoprene Acidity Variation

238 Data for the isoprene/SO₂ acidity experiment are provided in Table 1. For this
239 experiment, the initial isoprene concentration was 8.4 ppmC, the initial NO was 0.37 ppm, and
240 the relative humidity averaged 30% at 25 °C (6.5 g m⁻³ absolute humidity, on average). SO₂
241 ranged from near background to 0.23 ppm. Residual SO₂ might have contributed to the
242 background [H⁺]_{air} of 54 nmol m⁻³, although this value is more likely due to aerosol-phase
243 organic products of isoprene oxidation, particularly organic acids. However, in terms of the

244 relative changes of percent OC increase, this background value is of little consequence.
245 Generating the acidity with SO₂ allows the [H⁺]_{air} to achieve values in excess of 1500 nmol m⁻³,
246 a value much greater than can be reliably maintained using nebulized solutions. However, unlike
247 nebulized aerosol, the concentrations of inorganic sulfate in the product aerosol vary at each
248 stage of the experiment, as shown in Table 1. Sulfate concentrations were measured by ion
249 chromatography at each stage of this experiment.

250 With no added SO₂ (stage ER370-9), the organic carbon from the isoprene reaction
251 resulted in 5.3 μgC m⁻³ of SOC formed (corrected for chamber losses). Percent increases over
252 this base case value ranged from 62% to 459% at the highest acidity level (1524 nmol m⁻³).
253 Figure 1 provides a plot of the percent change in OC against the aerosol acidity. Error bars for
254 [H⁺]_{air} are derived from the rated variability of the pH probe used, converted into nmol m⁻³ using
255 the sampling and extraction volumes employed. Error bars on the OC concentrations are based
256 on the variability in the replicate semi-continuous OC measurements conducted during each
257 sampling period (typically, n>20). As seen in the figure, the relative increase in organic carbon
258 correlated well with increasing acidity with an R² of 0.985. The negative intercept resulted from
259 the small amount of acidity measured under the condition without SO₂, and the slope indicates a
260 0.31 %SOC increase per nmol m⁻³ of increased [H⁺]_{air}.

261 Despite employing different mechanisms for generating the acidic aerosol, the agreement
262 in the data between this study and Surratt et al. (2007) is excellent. The %SOC increase appears
263 to be quite consistent (0.31 for SO₂ photooxidation vs. 0.32 via nebulization), suggesting both
264 pathways lead to comparable acid enhancements. The results also suggest that variations in the
265 inorganic aerosol loading do not strongly impact the observed %SOC increase, at least under the
266 range of conditions considered, which is consistent with the results previously reported by

267 Offenberg et al. (2009) for α -pinene/NO acidity experiments conducted at different SOC
268 concentrations.

269 Attempting to expand the SO₂ experiment to incorporate additional humidity conditions
270 revealed a further challenge for the use of SO₂ versus nebulization of sulfate aerosols in these
271 acidity experiments. Changes to the chamber humidification also resulted in changes in the
272 amount of SO₂ converted to aerosol-phase acidic sulfate, with higher humidity resulting in lower
273 aerosol sulfate concentrations. Nebulized sulfate aerosols, in contrast, appear to retain stable
274 aerosol sulfate concentrations and [H⁺]_{air} levels under variable humidity conditions. This
275 limitation could potentially be overcome through the use of a direct measure of acidity in aerosol
276 particles. However, given the inherent limitations of the [H⁺]_{air} measurement, the nebulization
277 approach provides a cleaner evaluation of the effects of humidity on SOC formation. For this
278 reason, the remainder of the experiments presented will focus on nebulized inorganic sulfate for
279 the generation of aerosol acidity.

280

281 3.2 1,3-Butadiene Acidity Variation

282 Data for the 1,3-butadiene acidity experiment are provided in Table 2. For these
283 experiments, the initial 1,3-butadiene and NO concentrations were 6.8 ppmC and 0.34 ppm,
284 respectively. The first acidity condition once the reaction started was the base case of pure
285 ammonium sulfate, which rendered a [H⁺]_{air} of 48 nmol m⁻³. The next condition used a nebulizer
286 solution of nominally one-third sulfuric acid and two-thirds ammonium sulfate to give an [H⁺]_{air}
287 of 259 nmol m⁻³; the third case was a nominal one-third ammonium sulfate and two-thirds
288 sulfuric acid giving an [H⁺]_{air} of 666 nmol m⁻³; and the last case used sulfuric acid solution only
289 for an [H⁺]_{air} of 963 nmol m⁻³. The aerosol sulfate concentration was measured as approximately

290 35 $\mu\text{g m}^{-3}$ for the ammonium sulfate nebulization prior to the start of photochemistry; previous
291 measurements have shown that sulfate concentrations remain stable as the ammonium
292 sulfate/sulfuric acid ratio of the nebulizer solution is varied. The 1,3-butadiene consumed by
293 reaction ranged from 4.9 to 5.2 ppmC and averaged 5.03 ppmC.

294 Organic carbon concentrations increased with increasing acidity at the fixed relative
295 humidity of 30% (at an average temperature of 22 °C) from the base case of 22.6 $\mu\text{gC m}^{-3}$ to 44.7
296 $\mu\text{gC m}^{-3}$ at the highest acidity condition. SOC concentrations and percent increases from the
297 base case (ammonium sulfate) for the four stages are given in Table 2. The %SOC increases
298 monotonically with sulfate acidity up to nearly a 100% increase at the highest acidity condition.
299 The yield determined as $[\text{SOC}] / \Delta[1,3\text{-butadiene}]_{\text{carbon}}$ was calculated for each condition and
300 found to increase from 0.009 at the lowest acidity condition to 0.019 at the highest. Since ΔHC
301 remained nearly constant over the entire experiment, the increase in yield was essentially
302 equivalent to the increase in SOC, that is, a factor of two.

303 Figure 2 provides a plot of the percent change in organic carbon versus the $[\text{H}^+]_{\text{air}}$ for 1,3-
304 butadiene SOA at 30% relative humidity. As seen in the figure, the relative increase in organic
305 carbon correlated well with increasing acidity with an R^2 of 0.967. The negative intercept
306 resulted from the small amount of acidity measured in the base case with the ammonium sulfate
307 nebulizer solution. The plot shows an increase of 0.112 %SOC for each nmol m^{-3} increase in
308 $[\text{H}^+]_{\text{air}}$.

309 Figure 2 also compares the results from the 1,3-butadiene system with similar acidity
310 measurements from this laboratory. Superimposed on the sulfate acidity effect from 1,3-
311 butadiene SOA products are measurements made for three biogenic hydrocarbons previously
312 studied: isoprene (Surratt et al., 2007), α -pinene, and β -caryophyllene (Offenberg et al., 2009).

313 In those studies, SOA formation from isoprene, β -caryophyllene, and α -pinene was found to
314 correlate with aerosol acidity as linear relationships with different slopes. From the present work,
315 the 1,3-butadiene case also follows the same trend with a slope larger than that of α -pinene and
316 smaller than that of β -caryophyllene. In all five of these studies, a relative humidity of 30% was
317 used.

318 Table 3 further summarizes all the data from these $[H^+]_{air}$ variation experiments. $[H^+]_{air}$
319 and absolute OC concentrations are given as ranges for the individual studies. For most of the
320 experiments, Figure 2 shows the relationship between the percent change in SOC concentration
321 compared to the “neutral” base case. All data from experiments with isoprene, α -pinene, and β -
322 caryophyllene are from prior studies in this laboratory (Surratt et al., 2007; Offenberg et al.
323 2009) and use a chamber relative humidity of 30%. Table 3 also includes data for the MBO
324 experiment described by Zhang et al. (2012) where SOA was produced under conditions of low
325 NO_x with the aerosol generated through $RO_2 + HO_2$ and $RO_2 + RO_2$ reactions. Unlike the other
326 experiments presented in Table 3, the MBO experiment was conducted under dry conditions
327 (less than 3% relative humidity).

328 Overall, the sulfate acidity effect follows the order (from greatest to least effect):
329 isoprene; β -caryophyllene; MBO; 1,3-butadiene; and α -pinene. However, the exact placement
330 of MBO in this range is somewhat questionable given the dramatic differences in experimental
331 conditions used in that study (low NO_x chemistry and dry conditions) compared to the others. In
332 comparing the relative sensitivity of isoprene and 1,3-butadiene to sulfate acidity, there is about a
333 factor of three difference in the %SOC response to increasing $[H^+]_{air}$ despite the general
334 structural similarity of the compounds. This could represent a substituent effect which influences
335 the sensitivity of the gas-phase precursors to reaction by the acidic sulfate nucleophile, but

336 further organic analysis of the aerosol phase constituents would be required to examine this
337 possibility in detail. Although recent studies have compared the reaction pathways and products
338 formed for 1,3-butadiene oxidation versus isoprene oxidation (Jaoui et al., 2014), these studies
339 did not focus on acid-influenced reactions or organosulfate formation.

340

341 3.3 Isoprene Humidity Variation

342 Table 4 provides the initial conditions for the two isoprene/NO experiments designed to
343 examine changes in SOC formation and yield resulting from changes in humidity. In the base
344 case experiment (ER667), the reaction was conducted in the presence of only a low
345 concentration ($\sim 1 \mu\text{g m}^{-3}$) of inorganic aerosol produced through the nebulization of a 10 mg L^{-1}
346 ammonium sulfate solution. The relative humidity was then changed in stages from 9% to 49%
347 in $\sim 10\%$ increments (at an overall average temperature of $28 \text{ }^\circ\text{C}$). At each stage, the chamber was
348 allowed to equilibrate before a complete set of $[\text{H}^+]_{\text{air}}$, SOC, and ΔHC measurements were made.
349 Measured $[\text{H}^+]_{\text{air}}$ values averaged 54 nmol m^{-3} over the course of the experiment, a level
350 consistent with previous non-acidified isoprene/NO systems (both Surratt et al., 2007, and
351 ER370 reported above). In addition, a comparable experiment (ER662) was conducted using a
352 moderately acidic inorganic aerosol generated via nebulization of a mixed ammonium sulfate
353 and sulfuric acid solution. In this experiment, duplicate measurements were made at steady-state
354 relative humidity levels of 8, 28, 44, and 18%. The overall average temperature over the course
355 of the experiment was $27 \text{ }^\circ\text{C}$. In this experiment, the measured $[\text{H}^+]_{\text{air}}$ values averaged 275 nmol
356 m^{-3} . Based upon previous isoprene acidity experiments, this modest level of sulfate acidity
357 would be expected to produce an increase in SOC of approximately 50-75% at a relative
358 humidity of 30%.

359 Figure 3 provides a plot of the measured SOC levels as a function of humidity for these
360 two isoprene systems. Due to temperature differences between these experiments (and, more
361 importantly, the 1,3-butadiene experiments described below), measures of chamber relative
362 humidity have been converted into absolute humidity ($\text{g H}_2\text{O m}^{-3}$) to provide a common basis for
363 all four experiments. Error bars on the humidity axis are determined from the variability in
364 absolute humidities calculated on a 5-minute basis throughout the sampling periods. It is unclear
365 whether relative humidity or absolute humidity is of greater physical significance in the systems
366 under consideration. A direct measure of aerosol liquid water content would likely be a more
367 appropriate metric than either relative or absolute humidity for this study. However, no method
368 for the analysis of aerosol liquid water content was available for these experiments.

369 For the base case experiment, the SOC values range from a high of $13.3 \mu\text{gC m}^{-3}$ at the
370 lowest humidity level ($2.6 \text{ g H}_2\text{O m}^{-3}$) to just over $3 \mu\text{gC m}^{-3}$ at the higher humidities (10.4 to
371 13.1 g m^{-3}). The reason for this reduction in SOC formation is not entirely clear. Gas-phase NO_x
372 and O_3 concentrations do not appear to change significantly as a function of humidity level, as
373 does the concentration of isoprene consumed in the reactions. This suggests that early-generation
374 gas-phase oxidation reactions are probably not altered significantly by changing humidification.
375 Changes in the aerosol liquid water content may affect the gas-particle partitioning of later-
376 generation isoprene oxidation products, or increased water content may affect particle-phase
377 organosulfate formation or the formation of oligomeric species (Pye et al., 2013). Further
378 analysis of gas- and particle-phase organic constituents is required to further investigate this
379 effect.

380 For the acidified experiment, SOC declined from above $30 \mu\text{gC m}^{-3}$ at the lowest
381 humidity level (2.2 g m^{-3}) to around $4 \mu\text{gC m}^{-3}$ under the highest humidity condition (11.3 g m^{-3}).

382 Although the absolute humidities considered in the two experiments do not correspond precisely,
383 the percent increase in SOC for the acidic experiment versus the base case ranges from
384 approximately 140% at the lowest humidity levels, to approximately 65-75% in the mid-range
385 (where these experiments best overlap with the previous SOC versus $[H^+]_{\text{air}}$ studies), to virtually
386 no statistical difference between SOC levels above approximately $11 \text{ g H}_2\text{O m}^{-3}$. Figure 4
387 provides SOC yield curves for these two isoprene/NO scenarios. Error bars on the SOC yields
388 incorporate variability in the replicate measurements of both the inlet and chamber hydrocarbon
389 concentration as well as in the semi-continuous OC measurements throughout each sampling
390 period. As in the experiments described in the previous sections, the humidity changes
391 performed here had a minimal impact on the measured ΔHC . As a result, the isoprene/NO yield
392 plots follow essentially the same pattern as that seen for SOC formation in Figure 3.

393 These results suggest that humidity can have a profound effect on the acid-derived
394 enhancement of SOC formation from isoprene. Although the range of conditions explored is
395 limited (only a single bulk acidity level; only a partial range of relative humidities; and only a
396 comparatively narrow temperature range, by atmospheric standards), the data imply that under
397 some circumstances, high humidity (or perhaps high aerosol water content) can essentially
398 suppress enhanced SOC formation from isoprene photochemistry. These results also reinforce
399 the fundamental weakness of the $[H^+]_{\text{air}}$ measurement as a surrogate for acidity levels in actual
400 aerosol particles. Although the bulk acidic potential of the systems, as measured by $[H^+]_{\text{air}}$, does
401 not change significantly over the range of humidities considered, the resulting changes in the
402 SOC concentrations suggest that the pH in aerosol particles may be changing significantly due to
403 variations in aerosol liquid water content, solution ionic strength, or other factors not effectively
404 captured by the $[H^+]_{\text{air}}$ measurement.

405

406 3.4 1,3-Butadiene Humidity Variation

407 Conditions for the two 1,3-butadiene/NO experiments for examining changes due to
408 humidity variations are presented in Table 4. As described above, the base case experiment
409 (ER666) was conducted in the presence of $\sim 1 \mu\text{g m}^{-3}$ of ammonium sulfate aerosol. The relative
410 humidity was then changed in stages from 10% to 60% in increments of roughly 10% each, at an
411 overall average temperature of 25 °C. This was compared with an additional experiment (ER444)
412 employing an acidic inorganic aerosol nebulized from solution, with measurements made at
413 steady-state relative humidity levels of 31, 50, 10, and 62%, at an overall average temperature of
414 25 °C. The nebulizer solutions used in ER444 used higher levels of sulfuric acid relative to
415 ammonium sulfate than the isoprene experiment described above (ER662). This produced a more
416 acidic inorganic aerosol, with measured $[\text{H}^+]_{\text{air}}$ values of 718 nmol m^{-3} on average observed for
417 the 1,3-butadiene acidic aerosol experiment.

418 Figure 5 provides a plot of the measured SOC levels as a function of humidity for these
419 two 1,3-butadiene systems. For the base case experiment, the SOC values range from a high of
420 $45.1 \mu\text{gC m}^{-3}$ at the lowest humidity level ($2.5 \text{ g H}_2\text{O m}^{-3}$) to $24.7 \mu\text{gC m}^{-3}$ at the higher humidity
421 (13.6 g m^{-3}). For the acidified experiment, SOC declined from $60.3 \mu\text{gC m}^{-3}$ at the lowest
422 humidity level (1.9 g m^{-3}) to $31.1 \mu\text{gC m}^{-3}$ under the highest humidity (12.3 g m^{-3}). The range in
423 SOC enhancement from the base case to the acidified case is far lower than that observed in the
424 isoprene system, ranging from approximately 35% at low humidity to 25% at high humidity.
425 These enhancements are somewhat lower than would be expected for this level of acidity based
426 on the data presented in Figure 2. SOC yield curves, provided in Figure 4, follow this same
427 trend, as the ΔHC shows only minimal variation with humidity.

428 These results are markedly different from those seen for isoprene/NO, both in terms of
429 the level of SOC enhancement under the acidic condition and the extent to which the SOC
430 enhancement declines with increasing humidification. It is not clear what factors are driving this
431 difference in behavior. Part of the difference likely derives from structural differences between
432 the two molecules, as was described above with respect to the SOC versus $[H^+]_{air}$ studies.
433 Additionally, the higher level of $[H^+]_{air}$ used for the 1,3-butadiene experiment may be partially
434 offsetting the impact of increasing humidity, as more aerosol liquid water would be needed to
435 reduce actual particle acidity under these conditions. Other factors, such as the relative
436 hygroscopicity of isoprene and 1,3-butadiene SOA, may also be contributing. Additionally, the
437 temperature difference between the two experiments, although relatively small (approximately 3
438 °C on average) may be sufficient to introduce differences in the gas-particle partitioning between
439 the two experiments. Further experimentation is needed to attempt to better understand which
440 aspects of these aerosol systems are physically significant for activation or deactivation of these
441 acid-influenced reaction pathways, in order to determine if these pathways are ultimately
442 important to SOA formation in the ambient atmosphere.

443 4. Summary

444

445 These experiments support previous studies suggesting that acidic aerosol can lead to
446 increased SOA formation from the photooxidation of isoprene under laboratory conditions.
447 Changing the source of the acidity from nebulized inorganic aerosol to a more atmospherically
448 relevant photochemical conversion of SO₂ into acidic sulfate aerosol nearly identical results as
449 previous nebulized sulfate aerosol experiments. In addition, 1,3-butadiene, a chemically similar
450 compound released from primarily anthropogenic sources, was also demonstrated to produce
451 higher concentrations of SOA under acidic conditions, albeit to a lesser extent than was seen
452 with isoprene. The humidity experiments further suggest that aerosol liquid water content can
453 have a substantial effect on SOA formation from isoprene and 1,3-butadiene. Increasing
454 humidity produces a notable reduction in SOC formation in both the isoprene and 1,3-butadiene
455 photochemical systems, which is more pronounced in both systems in the presence of acidic
456 inorganic aerosols, and was most pronounced for the acidified isoprene/NO system.

457 In the isoprene/NO photochemical systems examined in this study, SOC enhancement
458 due to the presence of acidic inorganic aerosol was observed to be negligible at absolute
459 humidity levels above approximately 11 g H₂O m⁻³. This lower SOC enhancement at elevated
460 humidities may explain, in part, the difficulties in detecting increased SOA formation under
461 acidic conditions in field studies of ambient air masses, particularly in humid climates like the
462 southeastern US. This work suggests that a more detailed understanding of the role of humidity
463 and of aerosol liquid water content is likely required in order to accurately predict the impact of
464 acidity-influenced oxidation chemistry on overall SOA yields. While the data presented here
465 may suggest that enhanced SOA formation via acid-influenced pathways is more constrained

466 than previous studies may have suggested, it does still appear to represent a viable pathway for
467 additional SOA formation from a number of precursor hydrocarbons, which may need to be
468 incorporated into air quality models in order to accurately estimate secondary PM concentrations
469 in certain locations.

470 While these experiments are suggestive, they also include a number of shortcomings that
471 need to be addressed in future work. Perhaps the most significant is the use of absolute humidity
472 and $[H^+]_{\text{air}}$ as surrogate measures of aerosol liquid water content and aerosol pH. Determination
473 of the effective pH in the aerosol particles through the application of thermodynamic models,
474 such as ISORROPIA or AIM, should provide a more realistic assessment of actual acidity than
475 the $[H^+]_{\text{air}}$ approach, provided adequate gas and particle composition data is obtained
476 experimentally. However, even these models generally account for only the influence of
477 inorganic species, while the presence of isoprene SOA products has been reported to also
478 contribute significantly to water uptake on ambient aerosols (Guo et al., 2014). This suggests
479 that for laboratory experiments with high organic aerosol concentrations, particularly from
480 isoprene-related parent hydrocarbons, direct measurements of aerosol liquid water content may
481 be required.

482 Additionally, the relative humidity experiments presented here consider predominantly
483 systematic increases in relative humidity. As some phenomena related to aerosol liquid water
484 content are known to display hysteresis, such as deliquescence of sulfate aerosols, further testing.
485 Experiments incorporating descending relative humidities and larger humidity steps are
486 warranted, particularly along with the inclusion of liquid water content measurements. Further
487 examination of the mixing state and phase of the generated aerosols would also like be of value,

488 given that the available interfacial area could affect interactions between H⁺, sulfate, and SOA,
489 which may in turn affect organosulfate formation or other pathways to increased SOA yield.

490 Finally, further research is needed to examine changes in organic composition triggered
491 by the effects considered in this work. In particular, the mechanism through which humidity
492 level affects OC production from isoprene, even under non-acidified conditions, could be
493 important to the selection of appropriate SOC yields in air quality models. A further
494 examination of the similarities and differences between acid-influenced OC formation in the
495 isoprene and 1,3-butadiene systems would also likely be valuable. A more rigorous organic
496 analysis of product distributions may help reveal why isoprene appears to be significantly more
497 sensitive to acidic conditions than 1,3-butadiene at low humidity levels, and whether this
498 difference is due to structural effects related to the additional methyl group affecting gas-phase
499 chemistry, particle-phase organosulfate formation, or due to other phenomena, such as
500 differences in volatility and partitioning of oxidized intermediates. A more detailed comparison
501 of isoprene and 1,3-butadiene organic chemistry under acidic conditions may help with the
502 development of more accurate mechanisms for inclusion in air quality models.

503

504 **Acknowledgments**

505 The U.S. Environmental Protection Agency through its Office of Research and Development
506 funded and collaborated in the research described here under Contract EP-D-10-070 to Alion
507 Science and Technology. It has been subject to Agency review and approved for publication.
508 Mention of trade names or commercial products does not constitute an endorsement or
509 recommendation for use.

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674 acidity and its influence on secondary organic aerosol, *Environ. Sci. Technol.*, 41, 3213 – 3219,
675 2007.

676 Table 1. Isoprene SOA as a function of sulfate acidity from the photooxidation of SO₂.

Stage	Initial SO ₂ (ppb)	SO ₄ ⁻² (μgC m ⁻³)	[H ⁺] _{air} (nmol m ⁻³)	OC (μgC m ⁻³)	OC Increase (%)	SOC Yield
ER370-9	11 ^a	0	54	5.3	0.0	0.002
ER370-8	55	8.7	324	8.6	62	0.003
ER370-7	88	15.3	457	10.9	105	0.004
ER370-4	136	31.1	912	16.8	214	0.006
ER370-1	231	59.2	1524	29.8	459	0.011

677 ^a Measurement subject to possible HC interference. No SO₂ was added in stage ER370-9.

678

679 Table 2. Conditions and OC data for 1,3-butadiene photooxidation with the nebulized inorganic
680 aerosol. For each stage, the initial 1,3-butadiene was 6.8 ppmC; initial NO was 0.34 ppm; and
681 relative humidity was 30% (6.1 g m⁻³ absolute humidity).

Stage	[H ⁺] _{air} (nmol m ⁻³)	OC (μgC m ⁻³)	OC % Increase	ΔHC (ppmC)	SOC Yield
ER444-1	48	22.6	0.0	5.0	0.009
ER444-2	259	28.3	25	5.0	0.012
ER444-3	666	41.6	84	5.2	0.016
ER444-4	963	44.7	98	4.9	0.019

682

683 Table 3. Summary of the normalized yields for sulfate acidity effect for precursor hydrocarbons
 684 studied to date. SOA formed in the presence of NO_x at 30% relative humidity, except where
 685 indicated.

SOA precursor	[H ⁺] _{air} (nmol m ⁻³)	[OC] (μgC m ⁻³)	Normalized OC Change ^a	Reference
1,3-Butadiene	48 – 963	22.6 – 44.7	0.11	(this work)
Isoprene ^b	54 - 1524	5.3 – 29.8	0.31	(this work)
Isoprene	32 – 517	12.2 – 31.1	0.32	Surratt et al. (2007)
α-pinene (low OC)	68 – 1229	8.0 – 11.6	0.044	Offenberg et al. (2009)
α-pinene (high OC)	153 – 1014	40.5 – 55.3	0.039	Offenberg et al. (2009)
β-caryophyllene	112 – 1147	10.0 – 34.0	0.22	Offenberg et al. (2009)
2-Methyl-3- butene-2-ol (MBO) ^c	125 – 1590	6.5 – 21.9	0.14	Zhang et al. (2012)

686 ^a %SOC change per [H⁺]_{air}; ^b acidity generated from SO₂ photooxidation; ^c experiment
 687 conducted in the absence of NO_x under dry conditions

688

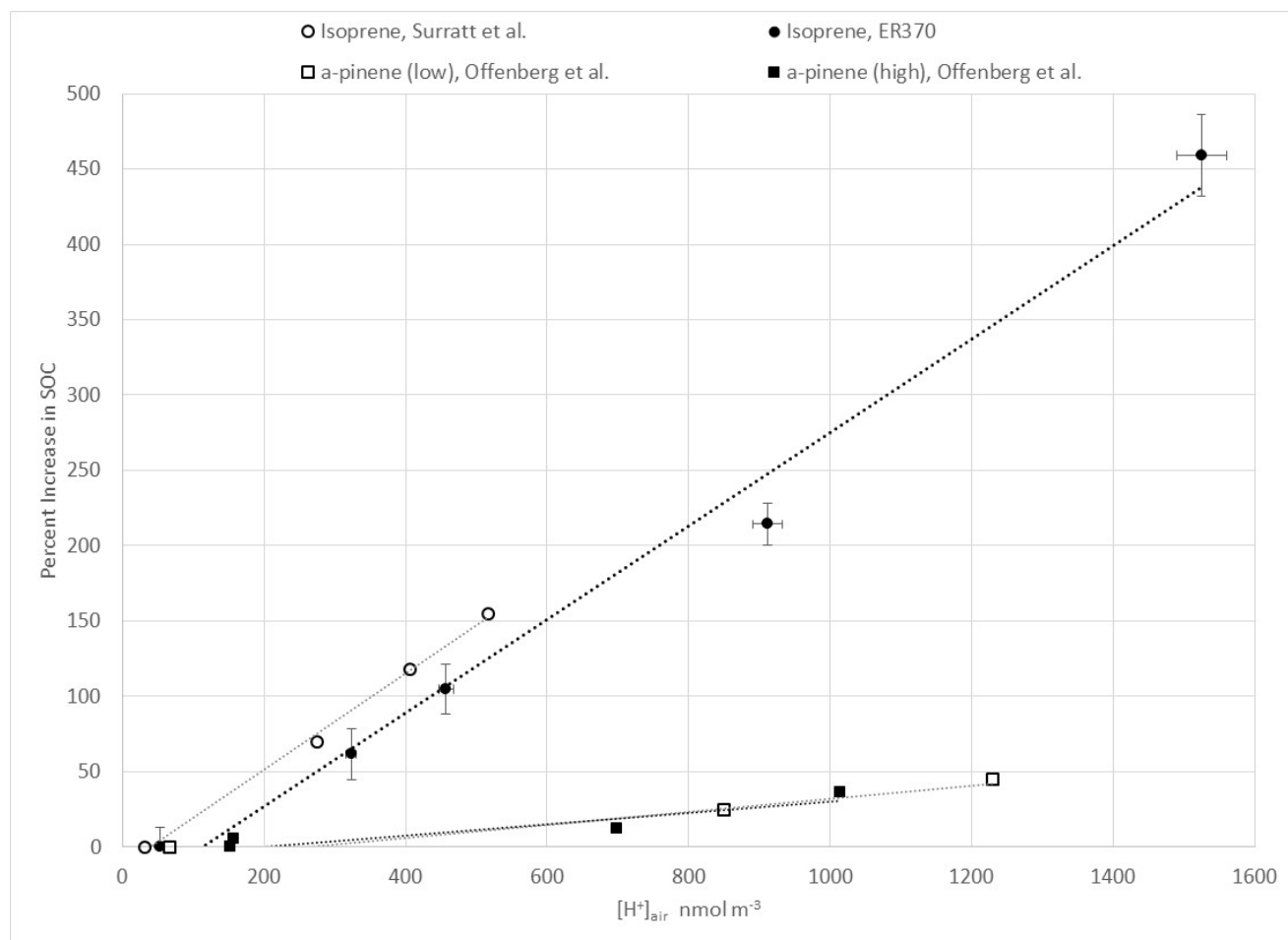
689 Table 4. Reaction conditions for humidity variation experiments.

Exp	Hydrocarbon	Inorganic	HC (ppmC)	NO _x (ppm)	Temp (°C)	Humidity (%)	Humidity (g m ⁻³)
ER667	Isoprene	Low Conc (NH ₄) ₂ SO ₄	8.2	0.35	28 – 29	9 – 49	2.6 – 13.1
ER662	Isoprene	1/2 (NH ₄) ₂ SO ₄ , 1/2 H ₂ SO ₄	7.0	0.29	27 – 28	8 – 44	2.2 – 11.3
ER666	1,3-Butadiene	Low Conc (NH ₄) ₂ SO ₄	7.1	0.42	25 – 26	11 - 60	2.5 – 13.6
ER444	1,3-Butadiene	1/3 (NH ₄) ₂ SO ₄ , 2/3 H ₂ SO ₄	6.9	0.34	22 – 23	10 - 62	1.9 – 12.3

690

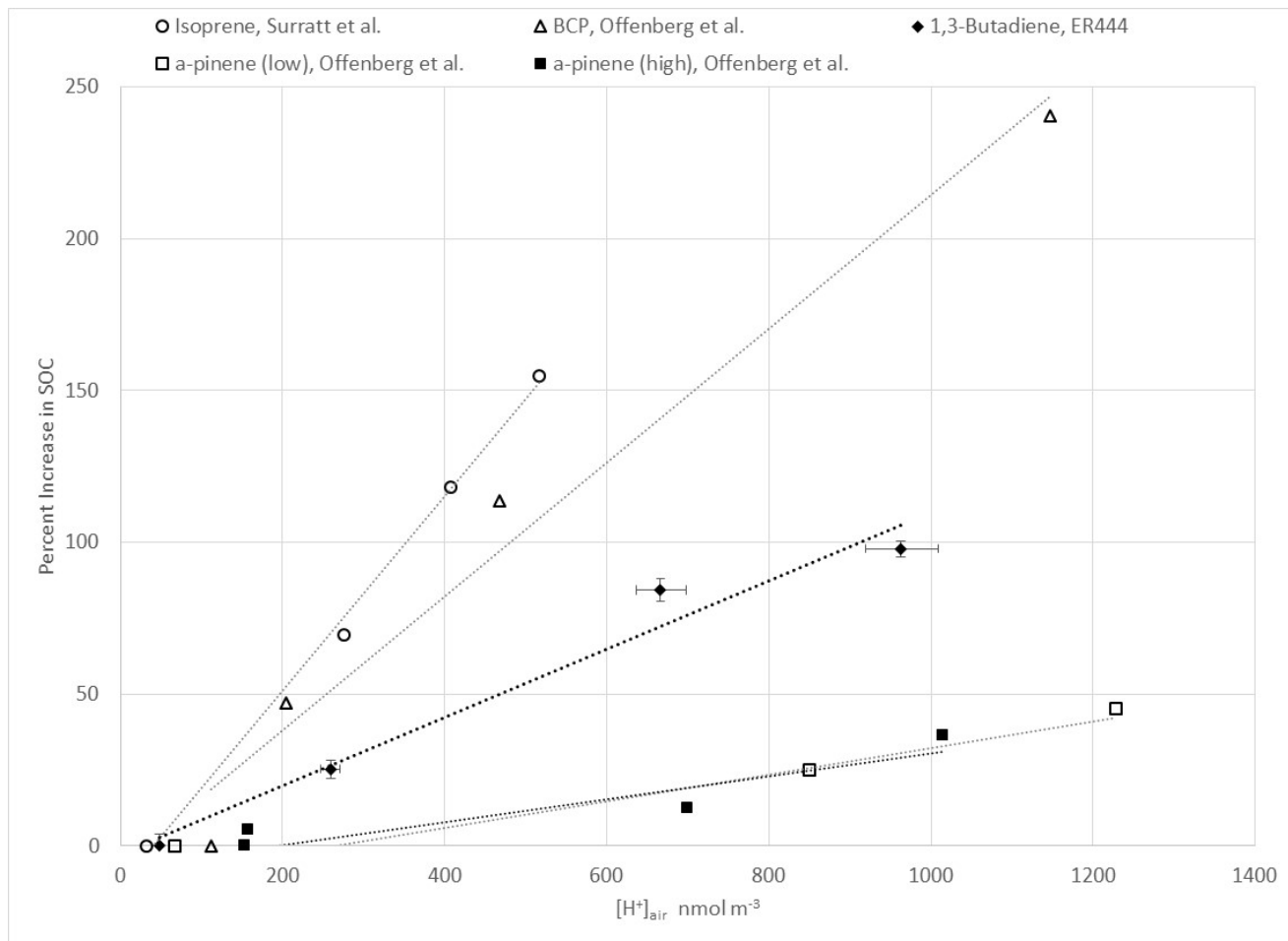
691

692 Figure 1. A comparison of the sulfate acidity effect for isoprene SOA. For Surratt et al., 2007
693 (open circles), the acidity was derived from nebulized sulfate aerosol. In the present study
694 (closed circles), the acidity was derived from the photooxidation of SO₂.



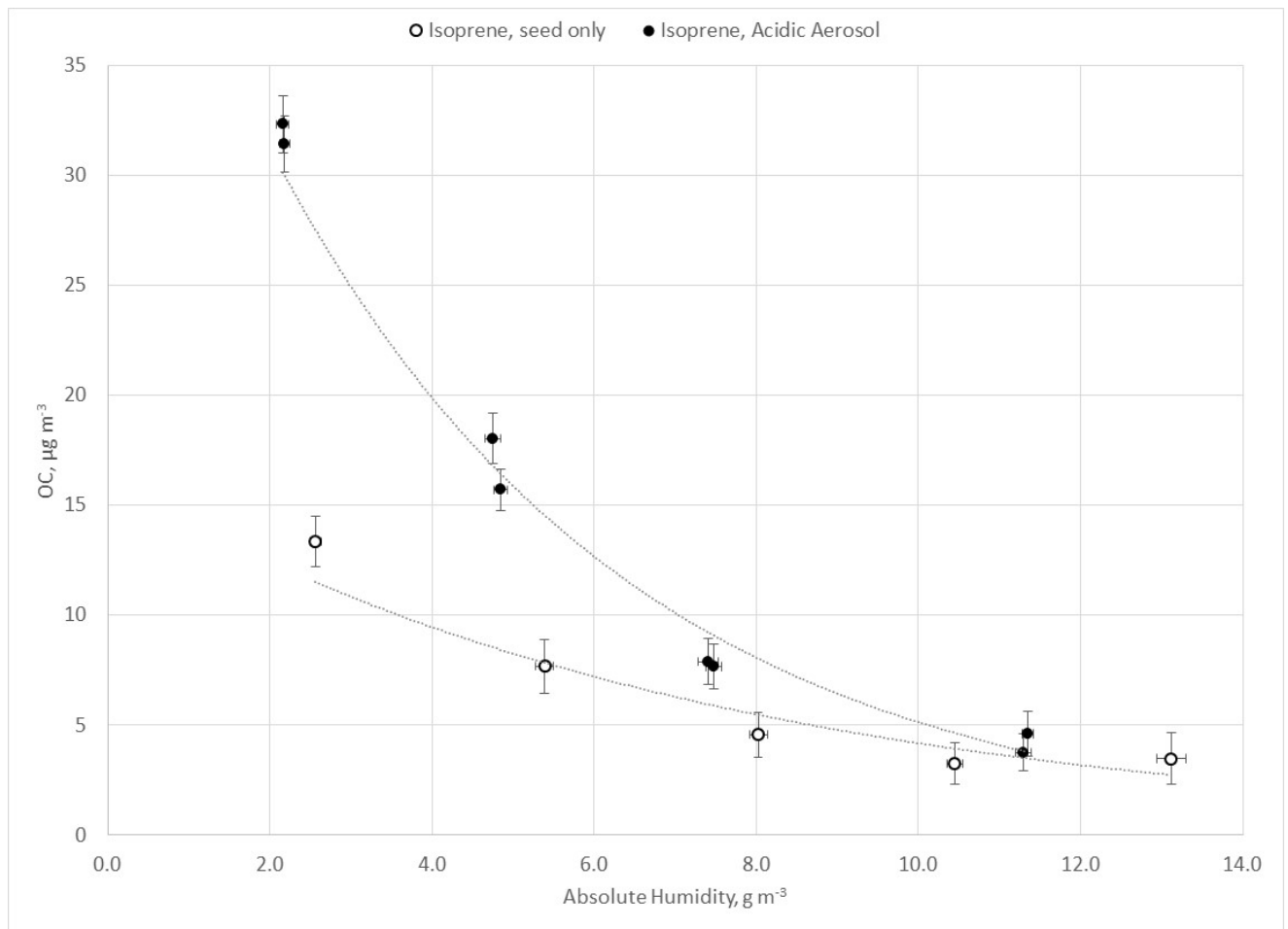
695

696 Figure 2. Aerosol acidity effect for 1,3-butadiene/NO, relative to previously published data
697 (Surratt et al., 2007; Offenberg et al., 2009). All experiments were conducted with nebulized
698 sulfate aerosol at 30% relative humidity.



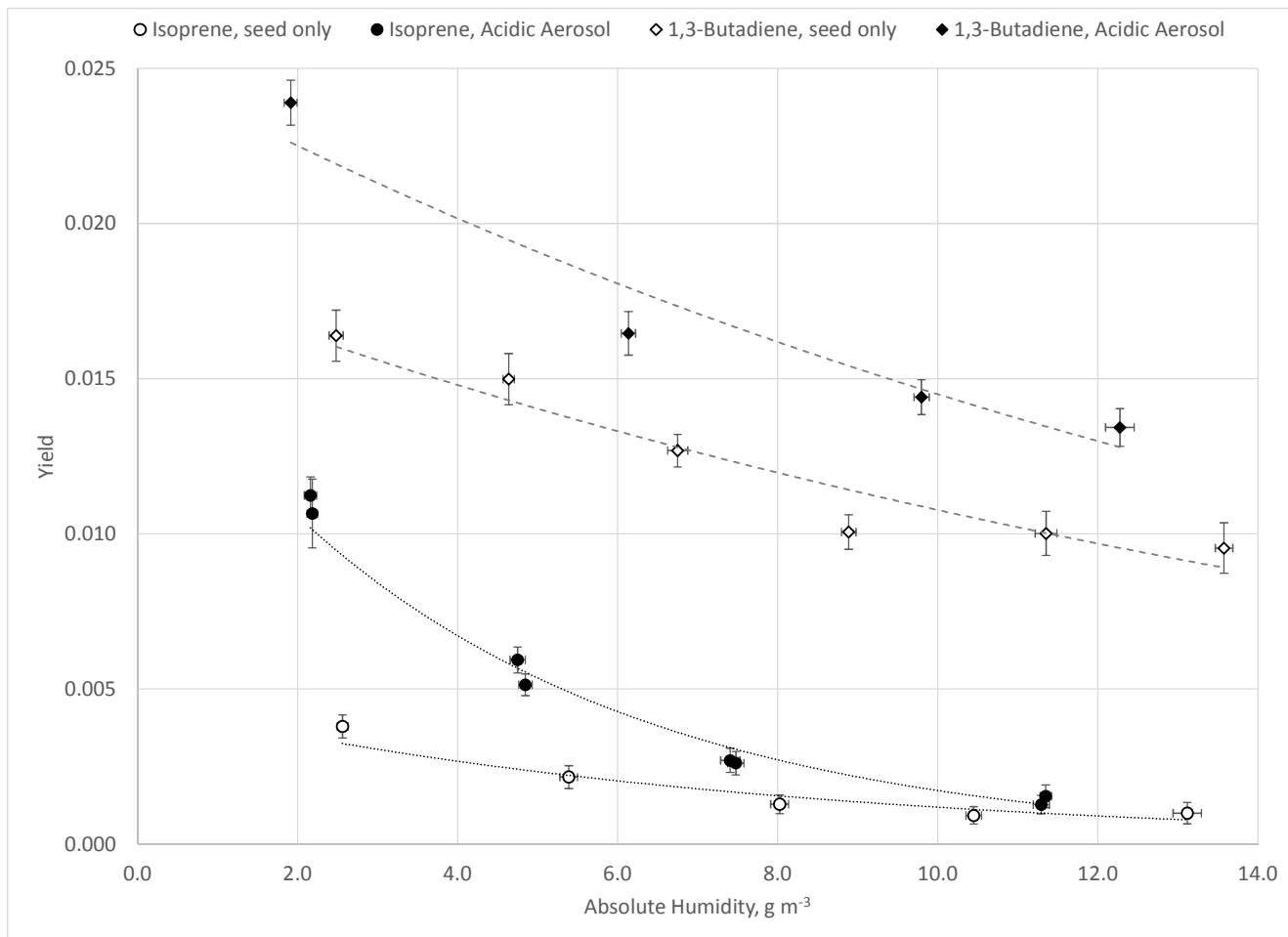
699

700 Figure 3. A comparison of the effects of humidity variation on isoprene/NO SOC formation. In
701 ER667 (open circles), only a low concentration ammonium sulfate seed aerosol was present. In
702 ER662 (closed circles), a moderately acidic sulfate aerosol was generated via nebulization.



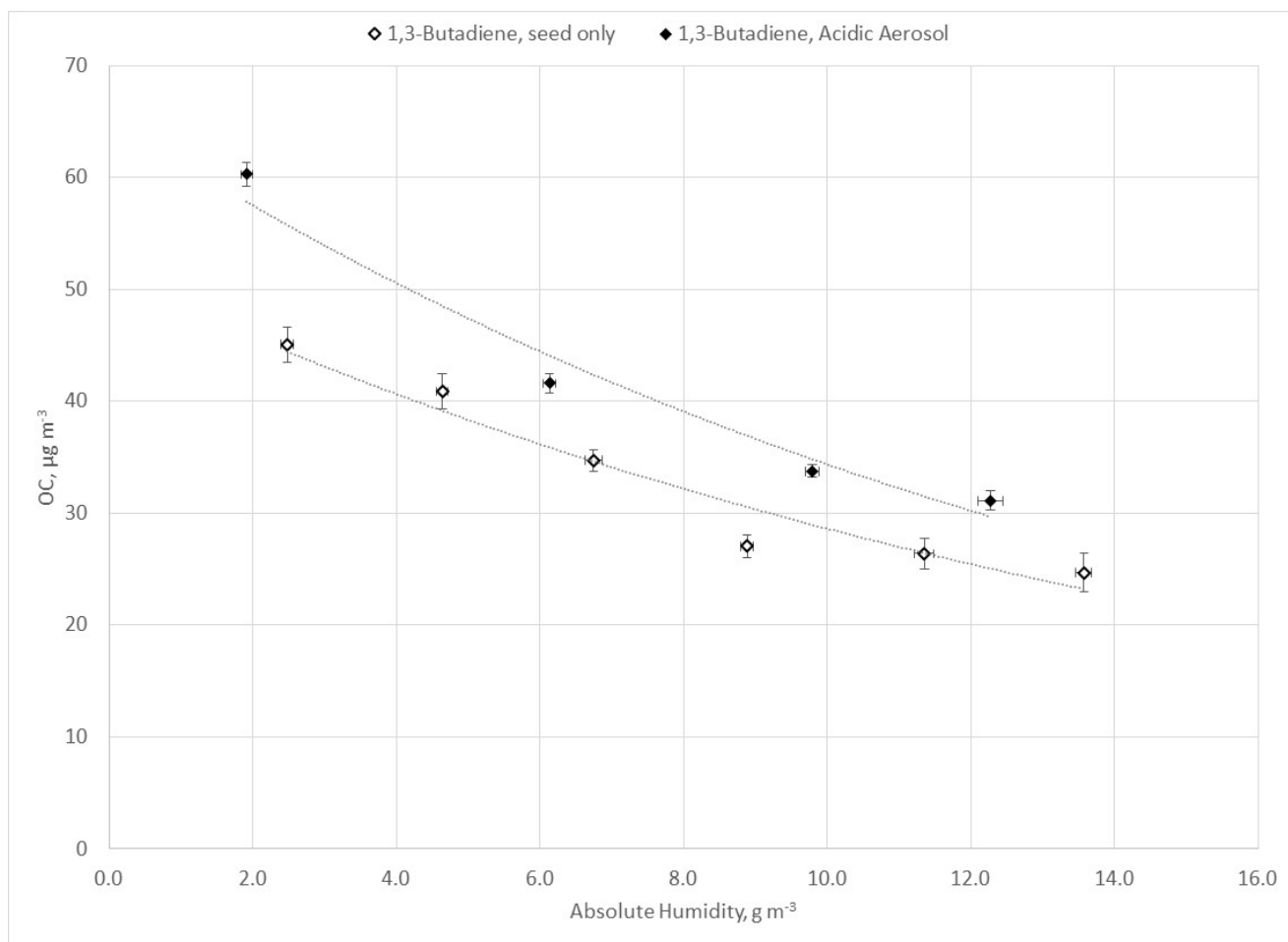
703

704 **Figure 4.** SOC yields for isoprene/NO and 1,3-butadiene/NO as a function of absolute humidity.
705 In ER667 (isoprene, open circles) and ER666 (1,3-butadiene, open diamonds), only a low
706 concentration ammonium sulfate seed aerosol was present. In ER662 (isoprene, closed circles)
707 and ER444 (1,3-butadiene, closed diamonds), an acidic sulfate aerosol was present.



708

709 Figure 5. A comparison of the effects of humidity variation on isoprene/NO SOC formation. In
710 ER666 (open diamonds), only a low concentration ammonium sulfate seed aerosol was present.
711 In ER444 (closed diamonds), an acidic sulfate aerosol was generated via nebulization.



712