Connecting the solubility and CCN activation of complex organic aerosols: A theoretical study using solubility distributions

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- 12
- 13 Abstract

14 We present a theoretical study investigating the cloud activation of multicomponent organic 15 particles. We modeled these complex mixtures using solubility distributions (analogous to 16 volatility distributions in the VBS, i.e. volatility basis set, approach), describing the mixture as 17 a set of surrogate compounds with varying water-solubilities in a given range. We conducted 18 Köhler theory calculations for 144 different mixtures with varying solubility range, number of 19 components, assumption about the organic mixture thermodynamics and the shape of the 20 solubility distribution, yielding approximately 6000 unique cloud condensation nucleus (CCN)-21 activation points. The results from these comprehensive calculations were compared to three 22 simplifying assumptions about organic aerosol solubility: 1) complete dissolution at the point 23 of activation; 2) combining the aerosol solubility with the molar mass and density into a single 24 effective hygroscopicity parameter κ ; 3) assuming a fixed water-soluble fraction ε_{eff} . The 25 complete dissolution was able to reproduce the activation points with a reasonable accuracy only when the majority (70-80%) of the material was dissolved at the point of activation. The 26 single parameter representations of complex mixture solubility were confirmed to be powerful 27 28 semi-empirical tools for representing the CCN activation of organic aerosol, predicting the 29 activation diameter within 10% in most of the studied supersaturations. Depending mostly on

the condensed-phase interactions between the organic molecules, material with solubilities 1 2 larger than about 0.1-100 g L⁻¹ could be treated as soluble in the CCN activation process over atmospherically relevant particle dry diameters and supersaturations. Our results indicate that 3 understanding the details of the solubility distribution in the range of 0.1 to 100 g L^{-1} is thus 4 5 critical for capturing the CCN activation, while resolution outside this solubility range will 6 probably not add much information except in some special cases. The connections of these 7 results to the previous observations of the CCN activation and the molecular properties of 8 complex organic mixture aerosols are discussed. The presented results help unravel the 9 mechanistic reasons behind observations of hygroscopic growth and CCN activation of 10 atmospheric secondary organic aerosol (SOA) particles. The proposed solubility distribution 11 framework is a promising tool for modeling the interlinkages between atmospheric aging, 12 volatility and water-uptake of atmospheric organic aerosol.

13

14 **1 Introduction**

15 Interactions of atmospheric aerosol particles with ambient water vapour determine to a large extent the influence that aerosols have on climate. On one hand the water content of aerosol 16 17 particles at atmospheric relative humidity (RH) below 100% contributes significantly to the direct effect they have on the global radiative balance (Seinfeld and Pandis, 2006; Petters and 18 19 Kreidenweis, 2007; Swietlicki et al., 2008; Zieger et al., 2011; Rastak et al., 2014). On the other hand the water-affinity of aerosol constituents, together with their dry size, defines the 20 21 efficiency with which these particles can activate as cloud condensation nuclei (CCN) at 22 supersaturated conditions (RH > 100%), form cloud droplets, and thus affect the properties of 23 clouds (Twomey 1974; Aalbrecht, 1989; McFiggans et al., 2006). To quantify the effects of 24 aerosol particles on clouds and climate it is thus necessary to understand the ways that aerosol 25 constituents interact with water.

Organic compounds contribute a large fraction (20-90%, depending on the environment) of atmospheric submicron particulate mass (Jimenez et al., 2009) – which is the part of the aerosol size distribution that typically dominates the CCN numbers. A significant fraction of this organic aerosol (OA) is secondary – i.e. produced in the atmosphere from the condensation of oxidation products of volatile, intermediate volatility and semi-volatile organic compounds (VOCs, IVOCs and SVOCs). Emissions of biogenic VOCs such as monoterpenes, isoprene and sesquiterpenes, followed by their subsequent oxidation and condensation in the atmosphere are
thought to be the dominant source of secondary organic aerosol (SOA) on a global scale
(Hallquist et al., 2009 and references therein) – although recent studies suggest also a notable
anthropogenic component to the global SOA (Volkamer et al., 2006; Hoyle et al., 2011;
Spracklen et al., 2011).

6 The solubility in water is one of the key properties governing the water-absorption (i.e. 7 hygroscopic growth) and CCN activation of aerosol particles. Together with aqueous phase 8 activity coefficients, surface tension, density and dry mass of the particle, water-solubility 9 affects the aerosol particle water content in thermodynamic equilibrium (Pruppacher and Klett, 10 1997; Seinfeld and Pandis, 2006; Topping et al., 2012). Atmospheric organic compounds have 11 a wide range of solubilities (Raymond and Pandis, 2003; Chan et al., 2008; Psichoudaki and 12 Pandis, 2013). OA is thus a complex mixture of molecules with different CCN-behaviour than 13 pure compounds. To accurately predict the water content and CCN activation of atmospheric 14 OA, information on the dissolution behaviour and aqueous phase interactions of these complex 15 mixtures is needed.

16 Representation of the complexity of OA is a major challenge for atmospheric chemical 17 transport models: OA consists of thousands of different compounds whose properties are poorly 18 known (Golstein and Galbally, 2007; Hallquist et al., 2009; Kroll et al., 2011). Approaches that simplify the complex nature of the OA mixture, yet reproducing its behaviour accurately 19 20 enough, are required to be able to assess the climate and air quality effects of atmospheric organics in large-scale modelling applications. One example of such approach is the 21 22 representation of the condensation and evaporation of SOA using a limited number of surrogate 23 compounds with a range of saturation concentrations, known as the volatility basis set (VBS, 24 Donahue et al., 2006; 2011; 2012). Similar simplifying approaches are needed to represent the 25 hygroscopic growth and CCN activation of OA as well.

When interpreting laboratory and field studies on hygroscopicity and CCN activation, a number of simplifying assumptions about the OA properties have been made, for instance: 1) assuming that organics completely dissolve in water at the point of activation (Huff-Hartz et al., 2006); 2) assuming a fraction (ε_{eff}) of organics to be completely soluble and the remaining fraction ($1 - \varepsilon_{eff}$) completely insoluble in water (e.g. Pruppacher and Klett, 1997; Engelhart et al., 2008); 3) lumping the phase-equilibrium thermodynamics, molar masses and densities of the OA constituents into a single semi-empirical parameter. One of the most commonly-used formulations is the hygroscopicity parameter κ , which relates the water activity in the aqueous solution with the water and dry particle volumes, and can be modified to account for limited solubility as well if the solubilities of the individual aerosol constituents are known (Petters and Kreidenweis, 2007; 2008; 2012, Petters et al., 2009a-c; Farmer et al., 2015; see also Rissler et al., 2004 and Wex et al., 2007 for alternative single-parameter formulations). These common simplifications of organic aerosol solubility and hygroscopicity are summarized in Table 1.

7 Laboratory studies on different types of organic aerosols have provided important 8 insights into the relationship between CCN-activation, hygroscopic growth and water-solubility 9 of the atmospheric OA constituents. Raymond and Pandis (2002, 2003) and Chan et al. (2008) investigated the CCN activation of single- and multi-component aerosol particles consisting of 10 11 organic compounds with known solubilities in water, and found that the particles activated at 12 lower supersaturations than would have been expected based on the bulk solubility of their 13 constituents. As an example, the laboratory studies by Chan et al. (2008) indicate that the CCN 14 activation of material with water solubility as low as 1 g L⁻¹ could be predicted assuming complete dissolution. For some model systems the surface properties (wettability) of the aerosol 15 16 particles, instead of the bulk water-solubility, seemed a more important factor defining their 17 CCN activation (Raymond and Pandis, 2002). Huff Hartz et al. (2006) assigned part of this 18 effect to residual water left in the particles upon their generation, causing the particles to exist 19 as metastable aqueous solutions and thus activate at lower supersaturations than the 20 corresponding dry material. The rest of the apparent increase in solubility was assigned to 21 potential impurities in the particles. In general the results reported by Huff Hartz et al. (2006) suggested that compounds with water-solubilities above 3 g L^{-1} behaved as if they were 22 23 completely soluble in water, in general agreement with the earlier results of Hori et al. (2003).

24 Secondary organic aerosol particles generated in the laboratory through oxidation 25 chemistry and condensation of the reaction products have also been found to activate as cloud droplets and thus contribute to the atmospheric CCN budgets (Cruz and Pandis, 1997; 1998; 26 27 Huff Hartz et al., 2005; VanReken et al., 2005; Prenni et al., 2007; King et al., 2007; 2009; Engelhart et al., 2008; 2011; Asa-Awuku, 2009; 2010). These particles probably resemble the 28 real atmospheric SOA more closely than individual organic species or their simple mixtures, 29 30 but the theoretical interpretation of their CCN-behaviour is complicated by the variety of their 31 constituents. Despite the fact that CCN-activity of SOA has been reported to vary with the 32 volatile precursor identity and loading (Varutbangkul et al., 2006; King et al., 2009; Good et

al., 2010), photochemical aging (Duplissy et al., 2008; Massoli et al., 2010, and temperature 1 2 (Asa-Awuku et al., 2009), the reported hygroscopicity parameter κ values determined for different SOA types are remarkably similar, being typically around 0.1 for the overall SOA and 3 4 0.3 for the dissolved fraction extracted from the aqueous sample (Asa-Awuku et al., 2010; King 5 et al., 2010). Similarly, Huff Hartz et al. (2005) reported effective solubilities of as high as 100 g L^{-1} for both mono- and sesquiterpene SOA – although both are known to consist of a range 6 7 of compounds with different solubilities. These results demonstrate the importance of knowing 8 the water-soluble fraction of SOA at varying conditions but suggest that its exact speciation is 9 probably not necessary for predictive understanding of the CCN activity of SOA particles (Asa-10 Awuku et al., 2010; Engelhart et al., 2011). The κ values inferred from sub- or supersaturated 11 conditions for the same SOA mixtures, on the other hand, are not always consistent, the sub-12 saturated κ values being typically lower than the super-saturated ones (Prenni et al., 2007; 13 Duplissy et al., 2008; Wex et al., 2009; Topping et al., 2012). Multiple possible reasons for this 14 have been presented in the literature, including incomplete dissolution of the aerosol constituents at sub-saturated conditions (Petters et al., 2009a), surface tension effects (Good et 15 al., 2010), RH-driven effects on the reaction chemistry and thus the composition (solubility and 16 activity) of the formed SOA (Poulain et al., 2010), evaporation and condensation of semi-17 18 volatile organic compounds (Topping et al., 2012), or the non-ideality of the mixtures being more pronounced at the sub-saturated conditions (Kreidenweis et al., 2006; Petters et al., 2009a; 19 20 Good et al., 2010).

21 While the basic theory of cloud droplet activation for pure water-soluble compounds is relatively well-established (Pruppacher and Klett, 1997; Asa-Awuku et al., 2007; Topping et 22 23 al., 2012; Farmer et al., 2015), and a number of theoretical and experimental studies on the different aspects controlling the CCN-activation of SOA have been presented (see above, 24 25 McFiggans et al., 2006; Dusek et al., 2006 and references therein), only few of these studies 26 have investigated the implications of the water-solubilities of complex organic mixtures on 27 CCN activation. Understanding the relationship between the dissolution behaviour and CCN-28 activation of complex organic mixtures is, however, needed to constrain the water-soluble 29 fraction of SOA in varying conditions as well as to systematically unravel the mechanisms 30 causing the apparent simplicity in the CCN-behaviour of complex organic mixtures.

In this work we introduce a framework for representing the mixture components with a continuous distribution of solubilities, similar to the VBS (Donahue et al., 2006; 2011; 2012).

Using this framework in a theoretical model, we investigate the dissolution behaviour of 1 2 complex organic mixtures and their CCN activity, focusing on the impact of mixture solubility on CCN-activation. In particular, we study the response of the CCN-activation to varying 3 4 solubility ranges, distribution shapes, and numbers of components in the mixture. Furthermore, 5 we compare the CCN-activation predictions using the simplified solubility representations 6 outlined above (complete dissolution, soluble fraction ε_{eff} , and hygroscopicity parameter κ 7 without including knowledge about the component solubilities) with the more detailed 8 description using the full solubility distributions, and study the relationship of the simplified 9 solubility parameters ε_{eff} and κ to the true mixture solubility distribution. Although the solubility 10 ranges and other thermodynamic properties of the mixture have been chosen to represent SOA, 11 many of the concepts and approaches introduced here can be applied to any particles consisting 12 of complex mixtures of organic compounds with varying water-solubilities. Finally, we discuss 13 the applicability of the introduced framework for describing the water-interactions of realistic 14 SOA mixtures and the relevant future directions.

15

16 2 Methods

17 2.1 Theoretical predictions of CCN activation of complex organic mixtures

18 Figure 1 schematically summarizes the model system considered in this study. We consider a 19 monodisperse population of spherical aerosol particles consisting of an internal mixture of 20 organic compounds. When exposed to water vapour, these particles grow reaching 21 thermodynamic equilibrium between the water vapour and the particle phase. The wet particle 22 is allowed to consist of maximum two phases: the insoluble organic phase and the aqueous 23 phase. The compositions of the organic and aqueous phases are determined on the one hand by 24 the equilibrium between the aqueous phase and the water vapour, and on the other hand by the equilibrium of the aqueous phase with the organic insoluble phase. To isolate the effects of 25 solubility from organic volatility effects, we do not allow the organics to evaporate from the 26 27 droplet -i.e. we assume that the equilibrium vapour pressures of the organics is zero above the droplet surface. Similarly, no condensation of organics from the gas phase to the particles is 28 29 allowed to take place. The validity of this assumption depends on the gas-phase concentrations 30 of the organic species as well as the atmospheric temperature during the cloud formation process. Testing it at different atmospherically relevant conditions deserves some future 31

attention, accounting for the dynamics of the atmospheric gas-phase as well (see also Topping et al., 2012). In this study, however, we focus strictly on the CCN-activation process. The organic composition and dry particle size were treated as an input to a model calculating the final equilibrium composition, wet size, and CCN activation behaviour of these particles. Note that while the solubility in the equations presented in the next Sections 2.1 and 2.2 is nondimensional (g g_{H2O}^{-1}), in the presentation of the results it is converted into g L⁻¹, assuming constant unit density of water.

8 2.1.1 Equilibrium between water vapour and an aqueous phase containing dissolved 9 material

10 The Köhler equation (Pruppacher and Klett, 1997) is used to link the ambient water vapour 11 saturation ratio *S* with the size, composition and water content of the aerosol particles in 12 thermodynamic equilibrium (lower panel of Fig. 1):

13
$$S = \frac{p_{w,eq}}{p_{w,sat}} = a_w \exp\left(\frac{4\sigma v_w}{RTD_{p,wet}}\right),$$
(1)

14 where $p_{w,eq}$ (Pa) is the equilibrium vapor pressure of water over the droplet surface, $p_{w,sat}$ (Pa) 15 the saturation vapor pressure over a pure flat water surface, σ (N m⁻¹) is the surface tension of 16 the droplet, v_w the molar volume of water in the aqueous phase, M_w (kg mol⁻¹) the molar mass 17 of water, ρ (kg m⁻³) the density of the aqueous phase, $D_{p,wet}$ (m) the diameter of the droplet, T18 (K) the temperature and R (J mol⁻¹ K⁻¹) the universal gas constant. a_w is the water activity, 19 defined as the product of the water mole fraction X_w and water activity coefficient in the aqueous 20 phase Γ_w :

$$21 a_w = X_w \Gamma_w. (2)$$

The activity coefficient describes the interactions between water molecules and the dissolved organic molecules in the mixture. The saturation ratio at which the particles of dry size $D_{p,dry}$ activate as cloud droplets (i.e. continue growing in size even if the saturation ratio decreases), is referred to as the critical saturation ratio S_c . Mathematically this corresponds to the highest local maximum in the $S(D_{p,wet})$ curve, usually referred to as the Köhler curve.

27 **2.1.2** Equilibrium between the aqueous and insoluble organic phases

The composition of the droplet and the distribution of material between the organic insoluble and the aqueous phases can be calculated applying the principles of mass conservation and the 1 thermodynamic equilibrium of the organic components in an aqueous mixture with the 2 insoluble organic phase. As the mass transfer of organics between the particles and the gas 3 phase is neglected, the total mass of the dry particle m_{dry} , being the sum over all components *i*, 4 is equal to the total organic mass in the wet droplet (see Fig. 1):

5
$$m_{dry} = \sum_{i}^{n} m_{i,inso\ lub\ le} + \sum_{i}^{n} m_{i,aqueous} , \qquad (3)$$

6 where *n* is the total number of organic compounds, $m_{i,insoluble}$ is the mass of compound *i* in the 7 insoluble organic phase and $m_{i,aqueous}$ the mass of compound *i* in the aqueous phase. The same 8 holds for each organic compound individually:

9
$$m_{i,dry} = y_{i,dry} m_{dry} = m_{i,inso\,lub\,le} + m_{i,aqueous}, \qquad (4)$$

10 where $y_{i,dry}$ is the mass fraction of *i* in the dry organic particle. On the other hand, the 11 concentration of each organic compound in the aqueous phase is determined by the 12 thermodynamics of the two-phase system consisting of the insoluble organic phase and the 13 aqueous solution phase. The mass of each organic compound *i* in the aqueous phase can be 14 expressed as (Prausnitz et al., 1998; Banerjee, 1984):

15
$$m_{i,aqueous} = \begin{cases} \gamma_i Y_{i,wet} c_{sat,pure,i} m_w, & Y_{i,wet} > 0\\ m_{i,dry}, & Y_{i,wet} = 0 \end{cases}$$
(5)

16 where γ_i is the activity coefficient of *i* in the insoluble organic phase (where the reference state is the pure component dissolution to water), $Y_{i,wet}$ and $c_{sat,pure,i}$ (here in g g_{H20}⁻¹) are the organic 17 18 phase mole fraction and pure component solubility (saturation concentration) of i, and m_w is the 19 total mass of water in the droplet. The former equation corresponds to the situation where the 20 particle contains an insoluble organic core in thermodynamic equilibrium, the latter to the case where only the aqueous phase exists, i.e. all the organic material has dissolved to the water. 21 22 Although the mole fraction and the corresponding molar activity coefficient have been used in 23 Eq. 5, a similar relationship can be defined using the mass fraction in the organic phase and a 24 corresponding mass-based activity coefficient. For a multicomponent system in which the 25 molar mass of the organic species varies, the mole and mass fractions of a given species are not 26 necessarily equal. In this study, however, we assume a constant molar mass throughout the 27 organic mixture for simplicity, leading to the mass and mole fractions in the organic phase to 28 be the same, i.e. $Y_i = y_i$ for all compounds. All the equations presented below can be re-derived in a relatively straightforward manner taking into account a potential difference between themole and the mass fractions in the organic phase.

Finding the organic and aqueous phase compositions that satisfy Eqs. 3-5 for given water and dry particles masses (m_w and m_{dry} , respectively) requires solving *n* coupled equations. These equations were expressed using the ratio χ_i of organic compound *i* in the insoluble core of the wet particle to the total mass of the compound (Raymond and Pandis, 2003; Petters and Kreidenweis 2008):

8
$$\chi_i = \frac{m_{i,inso\ lub\ le}}{m_{i,inso\ lub\ le} + m_{i,aqueous}} = \frac{m_{i,inso\ lub\ le}}{m_{i,dry}} = \frac{m_{i,inso\ lub\ le}}{Y_{i,dry}m_{dry}}.$$
(6)

9 The mole fraction (equal to the mass fraction for the mixtures considered here) of *i* in the 10 insoluble core is defined as

11
$$Y_{i,wet} = \frac{m_{i,inso\,lub\,le}}{\sum_{i} m_{i,inso\,lub\,le}} = \frac{\chi_i m_{i,dry}}{\sum_{i} \chi_i m_{i,dry}} = \frac{\chi_i Y_{i,dry}}{\sum_{i} \chi_i Y_{i,dry}}.$$
(7)

12 Finally, combining equations 3-7, we get *n* equations of the form

13
$$\chi_i = 1 - \frac{\chi_i Y_{i,dry} c_{i,sat,pure} m_w}{m_{i,dry} \sum_i \chi_i Y_{i,dry}} , \qquad (8)$$

14 which can be solved for χ_i with the constraint $0 \le \chi_i \le 1$ for given water and dry particle masses.

152.1.3 Representation of complex organic mixtures: Solubility distributions and16thermodynamic properties

17 A novel aspect of this study as compared with previous theoretical work is the representation 18 of complex mixtures using their aqueous solubility distribution of the individual species. In our 19 calculations we used mixtures of *n* compounds, whose water-solubilities ranged from $c_{sat,min}$ to *c*_{sat.max}, either on a linear or logarithmic basis. The shape of the distribution could vary as well. 20 21 In this work we studied essentially three types of mass fraction distributions in the dry particle: 22 a uniform distribution in which all solubilities are equally abundant, distribution increasing steadily (linearly or logarithmically), and a distribution decreasing steadily (linearly or 23 24 logarithmically). The 72 studied solubility distributions are specified in Table 2, and the 25 solubility distributions for n = 10, $c_{sat,min} = 0.1$ g L⁻¹ and $c_{sat,max} = 1000$ g L⁻¹ are presented in 26 Fig. 2 as examples.

1 For simplicity, we assumed that water forms an ideal solution with the dissolved 2 organics, i.e. $\Gamma_w = 1$, thus yielding an activity equal to the mole fraction of water, $a_w = X_w$ in Eq. 1. Since information about the activity coefficients of organic species in purely organic mixtures 3 4 is still scarce, we studied two alternative approaches to represent the dissolution 5 thermodynamics of the SOA mixture in Eqs. 9-12: 1) assuming an ideal organic mixture where $\gamma = 1$ for all compounds in the insoluble phase; 2) assuming a constant organic phase activity 6 7 $y_i Y_{i,wet}$ of unity for all compounds – in which case the dissolution behaviour of each *i* is similar 8 to their behaviour as pure components. These cases probably represent the limiting cases for 9 the dissolution of SOA components in CCN activation reasonably well, the former representing 10 a lower limit and the latter an upper limit for the overall solubility of the dry particle. Applying 11 the two limiting assumptions about the interactions of the compounds in the organic phase for 12 the 72 different solubility distributions (Table 2) thus results in total 144 unique representative 13 model mixtures.

14 The density, surface tension, and molar masses assumed for water and the organic compounds are summarized in Table 3. Although the density, surface tension and molar mass 15 of the organics are likely to vary with the solubility, we kept them constant throughout the 16 17 organic mixture to isolate the solubility effects on the CCN behaviour. The values were chosen 18 based on literature studies of the CCN behaviour of SOA (Engelhart et al., 2008; Asa-Awuku 19 et al., 2010). The surface tension σ was approximated by the surface tension of water, and the 20 molar volume of water in the aqueous phase was assumed the same as for pure water. 21 Furthermore, we assumed no dissociation of the organics in the aqueous phase.

22 2.1.4 Model calculations

We solved Eq. 8 for organic mixtures with the Matlab internal function *fsolve*, for varying water 23 24 and dry particle masses m_w and m_{dry} , covering 50 different dry particle diameters between 20 25 and 500 nm. The calculations yielded the composition of the insoluble organic and the aqueous 26 phase, and thus the mole fraction of water in the aqueous solution X_{w} . From these results the 27 Köhler curves $S(D_{p,wet})$ corresponding to each dry particle mass could be calculated using Eq. 28 1 (see Fig. 3a for an example of the Köhler curves). The critical supersaturations s_c (defined as 29 $S_c - 1$) corresponding to specific dry particle diameters $D_{p,dry}$ (termed also as activation 30 diameters $D_{p,act}$ at a given saturation ratio S or supersaturation s) were determined from the 31 maxima of the Köhler curves (see Fig. 3a). The temperature was assumed to be 298 K in all 32 calculations. These calculations for the 144 unique organic model mixtures corresponded to 1 7200 Köhler curves yielding 5957 ($D_{p,act}$, s_c) pairs (activation points, see Fig. 3b). For the 2 remaining 1143 curves no activation points were found with the given combinations of mixture 3 properties and dry diameters. For comparisons with the simple solubility representations, the 4 dissolved organic fraction defined as

5
$$\varepsilon = \frac{\sum m_{i,aqueous}}{m_{dry}}$$
 (9)

6 was extracted from the model output.

7 2.2 Comparison of the full model output to simple solubility representations

To investigate the performance of the simple solubility representations given in Table 1 in 8 9 reproducing the CCN activation of complex mixtures, we fitted the $(D_{p,act}, s_c)$ data created by the full model using these simpler models. No fitting is required for the complete solubility 10 11 approach. Using the obtained solubility parameters from the optimal fit and the corresponding 12 simplified forms of the Köhler equation we then recalculated new $(D_{p,act}, s_c)$ pairs and compared them to the predictions by the full model. Furthermore, we investigated the relationships 13 14 between the true mixture solubility distribution and the simplified solubility parameters. The 15 details of the approach used for each simple model are outlined below.

16 **2.2.1 Complete dissolution**

In the case where all of the organic material is assumed to completely dissolve at the point ofactivation, the calculation of the aqueous solution composition becomes trivial as

$$19 m_{i,dry} = m_{i,aqueous} (10)$$

for all the compounds, and the water mole fraction can simply be calculated based on the dryparticle mass as

$$22 X_w = \frac{\frac{m_w}{M_w}}{\frac{m_w}{M_w} + \frac{m_{dry}}{M_{org}}}, (11)$$

where m_w is the water mass in the droplet, m_{dry} the dry particle mass (related to $D_{p,dry}$ through the organic density ρ_{org}) and M_{org} the organic molar mass. The X_w calculated in this way was inserted into Eq. 1 to yield the corresponding ($D_{p,act}$, s_c) predictions and was also applied for calculating the solution density and surface tension as mass-weighted averages of the water and
 pure organic values.

3 2.2.2 Hygroscopicity parameter κ

In many practical applications the water activity and the difference in the densities and molar
masses of water and the dry material are expressed with a single hygroscopicity parameter κ,
introduced by Petters and Kreidenweis (2007), defined as:

$$7 \qquad \frac{1}{a_w} = 1 + \kappa \frac{V_s}{V_w},\tag{12}$$

8 where V_s and V_w are the volumes of the dry material and water, respectively. The following 9 formulation of the relationship between water saturation ratio, aerosol size and composition is 10 referred to as the κ -Köhler equation:

11
$$S = \frac{D_{p,wet}^{3} - D_{p,dry}^{3}}{D_{p,wet}^{3} - D_{p,dry}^{3}(1-\kappa)} \exp\left(\frac{4\sigma M_{w}}{RT\rho D_{p,wet}}\right)$$
(13)

12 yielding an approximate expression for the relationship between s_c and $D_{p,act}$ defined as

13
$$s_c = \frac{2}{3} \left(\frac{4M_w \sigma}{RT \rho} \right)^{\frac{3}{2}} \left(3\kappa D_{p,act}^3 \right)^{-\frac{1}{2}},$$
 (14)

14 Equation (14) was fitted to all $(D_{p,act}, s_c)$ data produced for a given organic mixture composition (see Table 2) by the full model, thus assuming a constant κ value for a given organic mixture. 15 16 To mimic the application of Eq. 14 to experimental data with no knowledge on the exact solute composition, in this case we assumed the surface tension and density to be those of water when 17 18 fitting the κ values to the full model data. The above formulation of κ , which is often used in 19 the interpretation of experimental data as well, thus contains information about solubility, 20 potential aqueous-phase non-ideality, as well as molar mass and density of the solutes (see 21 Farmer et al., 2015).

22 **2.2.3** Soluble fraction ε_{eff}

For an ideal solution of water and an organic solute the κ is directly proportional to the dissolved fraction and the ratio of the molar volumes of water and the solute i.e.: $\kappa = \varepsilon \kappa_{max}$, where $\kappa_{max} = (M_w/M_{org})(\rho_{org}/\rho_w)$. Assuming that a single soluble fraction ε_{eff} can represent a given organic 1 mixture (see Table 2) at all considered supersaturations, and substituting these relationships

2 into Eq. 16 yields

$$3 \qquad s_c = \frac{2}{3} \left(\frac{4M_w \sigma}{RT\rho} \right)^{3/2} \left(3 \frac{M_w}{M_s} \frac{\rho_s}{\rho_w} \varepsilon_{eff} D_{p,act}^3 \right)^{-1/2}, \tag{15}$$

4 the corresponding form of the Köhler equation being (see Huff Hartz et al., 2005)

5
$$S = \frac{M_{org} \rho_w (D_{p,wet}^3 - D_{p,dry}^3)}{D_{p,wet}^3 (M_{org} \rho_w) + D_{p,dry}^3 (\varepsilon_{eff} M_w \rho_{org} - M_{org} \rho_w)} \exp\left(\frac{4\sigma M_w}{RT\rho D_{p,wet}}\right).$$
 (16)

6 Again, we fitted Eq. 17 to the data produced by the full model and assumed the aqueous solution 7 density and surface tension to be equal to those of water. When $\varepsilon_{eff} < 1$, the following 8 relationship has been used to estimate the effective saturation concentration of the mixture 9 (Raymond and Pandis, 2002; Huff Hartz et al., 2005)

$$10 \qquad c_{sat,eff} = \frac{\rho_w (D_{p,wet}^3 - D_{p,act}^3)}{\varepsilon_{eff} \rho_{org} D_{p,act}^3}.$$
(17)

11 **2.2.4** Connection between ε_{eff} and the solubility distribution of the mixture

12 Let us now assume that the dissolved fraction at the point of activation for each considered 13 mixture can be expressed as a sum of two terms, the contribution from the compounds below a 14 threshold solubility bin i_t and the contribution from the compounds over the threshold:

15
$$\varepsilon = \frac{\sum_{i=1}^{i_{t}} m_{i,aqueous} + \sum_{j=i_{t}+1}^{n} m_{j,aqueous}}{m_{dry}} = \frac{\sum_{i=1}^{i_{t}} m_{i,aqueous} + (\sum_{j=i_{t}+1}^{n} m_{j,dry} - \sum_{j=i_{t}+1}^{n} m_{j,inso\,lub\,le})}{m_{dry}}.$$
 (18)

16

We now hypothesize that assuming a single soluble fraction for a given aerosol mixture is in fact equivalent to assuming that everything above i_t is completely dissolved while all the material below this threshold remain undissolved, i.e.

$$20 \qquad \mathcal{E}_{eff} = \frac{\sum_{j=i_t+1}^{n} m_{j,dry}}{m_{dry}}.$$
(19)

21 On the other hand, $\varepsilon = \varepsilon_{eff}$ if the following condition is fulfilled (see Eq. 20):

$$1 \qquad \sum_{i=1}^{i_{t}} m_{i,aqueous} = \sum_{j=i_{t}+1}^{n} m_{j,inso\,lub\,le} = \sum_{j=i_{t}+1}^{n} m_{j,dry} - \sum_{j=i_{t}+1}^{n} m_{j,aqueous} .$$
(20)

2 Substituting Eq. 5 to Eq. 20 we now have

$$3 \qquad F_{w} \sum_{i=1}^{i_{t}} \gamma_{i} Y_{i,wet} c_{sat,pure,i} = \sum_{j=i_{t}+1}^{n} Y_{dry,j} - F_{w} \sum_{j=i_{t}+1}^{n} \gamma_{j} Y_{j,wet} c_{sat,pure,j} , \qquad (21)$$

4 where

5
$$F_w = \frac{m_w}{m_{dry}} \,. \tag{22}$$

6 At the limit of large *n* and in the case of a symmetric distribution of material between the 7 insoluble organic and aqueous phases, Eq. 21 is satisfied by setting the threshold solubility i_t so 8 that

9
$$\lim_{i \to i_t} F_w \gamma_i Y_{i,wet} c_{sat,pure,i} = Y_{dry,i_t} - F_w \gamma_i Y_{i,wet} c_{sat,pure,i}$$
(23)

10 In this case the threshold solubility c_t is found from the bin for which

11
$$c_{sat, pure, i_t} = c_t \approx \frac{Y_{dry, i_t}}{Y_{wet, i_t}} \cdot \frac{1}{\gamma_{i_t}} \cdot \frac{1}{2F_w}$$
(24)

12 This is also equal to the bin where 50% of the material is partitioned in the insoluble phase, i.e.

13
$$\chi_{i_t} = \frac{m_{i_t, inso \, \text{lub} \, le}}{m_{i_t, inso \, \text{lub} \, le} + m_{i_t, aqueous}} = \frac{1}{1 + \frac{m_{i_t, aqueous}}{m_{i_t, inso \, \text{lub} \, le}}} = \frac{1}{2}$$
 (25)

Finding the solubility threshold c_t requires knowledge on the ratio F_w (Eq. 22). F_w , on the other hand, depends on the ambient supersaturation and the total soluble mass – thus introducing a supersaturation-dependence to the ε given by Eq. 18 as well. The magnitude of F_w as a function of supersaturation can be estimated by substituting Eqs. 15 and 17 into the definition of F_w (Eq. 22) and after some rearranging, yields:

$$19 F_w = \frac{\rho_w}{\rho_{org}} \left(\frac{2\varepsilon}{s_c} \cdot \frac{\rho_{org}}{\rho_w} \cdot \frac{M_w}{M_{org}} - 1 \right) = \frac{\rho_w}{\rho_{org}} \left(\frac{2\varepsilon\kappa_{\max}}{s_c} - 1 \right). (26)$$

1 3 Results

2 Figure 3a displays examples of the Köhler curves obtained from solving Eqs. 1 and 6 for 3 distribution 1 (the flat logarithmic distribution) with varying solubilities and n = 5, assuming 4 that the organics form an ideal mixture with each other. Each curve corresponds to a different 5 dry size, and the dots indicate the activation point (S_{crit} corresponding to the activation dry 6 diameter $D_{p,act}$). Black dots indicate incomplete dissolution ($\varepsilon < 0.99$) at the point of activation 7 while red dots indicate that in practice all the organics are dissolved into the aqueous phase at 8 the point of activation ($\varepsilon > 0.99$). Qualitatively similar behaviour was observed for all the 9 considered distributions: as the overall solubility of the mixture increases, the dissolution of the compounds increases, leading eventually to complete dissolution at the point of activation. The 10 11 transition from a regime with two phases (aqueous + insoluble) to a single aqueous phase is 12 visible in the two maxima in Fig. 3a, in accordance with Shulman et al. (1996) and Petters and 13 Kreidenweis (2008).

14 Figure 3b illustrates the parameter space probed in this study, showing the 5957 (scrit, $D_{p,act}$ points corresponding to the Köhler curve maxima calculated for all the considered 15 organic mixtures. The relationships between the critical supersaturation, activation diameter, 16 17 and dissolved fraction ε at the point of activation are also schematically shown. The chosen dry 18 diameters and supersaturations represent a conservative range of typical atmospheric conditions 19 - as the total aerosol number concentrations are dominated by ultrafine (diameters smaller than 20 100 nm) particles at most locations. In most considered cases the dissolved fractions fall 21 between 0.1 and 1, but the lowest dissolved fractions at the point of activation are of the order 22 of only a few percent - thus mimicking nearly insoluble aerosols. Therefore the cases 23 considered here represent a reasonable sample of atmospherically relevant conditions and SOA 24 mixture compositions. The water-to-organic mass ratios F_w corresponding to the probed 25 conditions and mixtures (see Sect. 2.2.4) range from values below 1 up to 1000, with most 26 values around 10-100. In many of the following plots and considerations we have chosen four 27 specific supersaturations, 0.1%, 0.3%, 0.6% and 1%, as representative values for typical 28 laboratory experiments, which are also indicated in Fig. 3b.

An example of the dependence of the activation diameter $D_{p,act}$ on the solubility range for all the studied distributions (see Table 2) and n = 5 is presented in Fig. 4. As expected, the activation diameter decreases with increasing supersaturation and solubility range for a given solubility distribution. The solubility distribution is reflected in the overall magnitude of the 1 activation diameters: the distributions that have larger fractions of material in the higher end of 2 the solubility range (distributions 2, 3, and 4) have generally lower activation diameters for a 3 given supersaturation as compared with the other distributions. The case when unity activity in 4 the organic phase is assumed results in smaller activation diameters for the same supersaturation 5 as compared with the ideal organic mixture case (see Sect. 2.1.3 for the definitions of the cases).

6 Figure 5 presents the activation diameters predicted using the simplified solubility 7 descriptions (Table 1) based on best fits to all available data as compared with the full 8 description of the solubility distributions (Table 2). The results clearly show, not surprisingly, 9 that assuming complete dissolution for all the mixtures consistently under-predicts the 10 activation diameters (Fig. 5a). Representing the dissolution behaviour with only one additional parameter, i.e. the hygroscopicity parameter κ (Eqs. 15-16) or the effective soluble fraction ε_{eff} 11 12 (Eqs. 17-18) improves the agreement between the activation diameters considerably (Figs. 5b and 5c). Adding the knowledge about the molar mass and density of the organic mixture, which 13 14 is the only difference between using the ε_{eff} instead of the single κ , adds only marginal 15 improvements in predicting the activation diameters for a given supersaturation. The 16 disagreements between the simplified models and the full theoretical treatment are largest for 17 the smallest supersaturations. These are the cases with the widest range of possible ε values at 18 the point of activation (see Fig. 3b), and the effect is most obvious for the complete dissolution 19 model: the larger the deviation from complete dissolution at the point of activation (i.e. the $\varepsilon \sim$ 20 1 case in Fig. 3b) the more significant error we introduce. The activation diameters predicted 21 assuming complete dissolution are within 10% of the correct values if the real dissolved fraction 22 ε is larger than about 0.7-0.8 at the point of activation.

The performance of the simple solubility models for all the studied Köhler curves is summarized in Fig. 6: while the complete dissolution assumption results in systematic underprediction (up to 40%) of the activation diameter, the κ - and ε_{eff} -based models are generally within 10% (in most cases within 5%) of the activation diameter predicted for the full solubility distribution representation.

Figure 7 compares the fitted parameters representing the mixture dissolution to the corresponding values inferred from the full mixture data for the 144 different mixtures. In Fig. 7a, the effective soluble fractions ε_{eff} are compared to the actual dissolved fractions ε at the point of activation for all the studied mixtures. While in the fits a single constant ε_{eff} has been assumed to represent a given mixture (see Eq. 15-16), in reality ε varies with supersaturation (Eq. 9).

Thus, while the fitted ε_{eff} for a given mixture correlates very well with the average ε over all 1 2 activation points (the markers in Fig. 7a), the performance of the approach can vary considerably with supersaturation (the grey lines in Fig. 7a). In practice this means that 3 4 describing a given complex mixture with a fixed soluble fraction yields representative average dissolution behaviour, but does not guarantee correct solubility description for a specific s_c if 5 6 fitted over a range of supersaturations. The corresponding comparison for the hygroscopicity 7 parameter κ values describing the data are shown in Fig. 7b. A clear correlation between the 8 fitted κ and the average ε is observed as expected (see Sect. 2.2.3), but the variation of ε with 9 supersaturation again adds scatter to the data – suggesting a dependence of κ on s_c . The 10 maximum κ , on the other hand, is defined primarily by the molar masses and densities of the 11 organics. For our mixtures with constant M_{org} and ρ_{org} the value of κ_{max} is 0.15, which is 12 indicated in Fig. 7b. The points above this theoretical maximum are a result of using the pure 13 water density instead of the mixture value in the Kelvin term of the Köhler equation (Eq. 16). 14 These results thus suggest that the κ values of 0.1-0.2 typically observed for SOA particles 15 (Duplissy et al., 2011) are controlled by the molar masses and densities of the SOA mixtures to a large extent and can result from quite different SOA mixtures in terms of their solubilities. 16

17 To illustrate the relationship of the fitted ε_{eff} to the dissolution of a given mixture, the 18 partitioning between the aqueous and insoluble organic phase is presented in Fig. 8 for 19 distribution 1 with the "low" solubility range and n = 100 at the point of activation when $s_c =$ 20 0.1% (see Table 2). Figure 8a shows the partitioning for the case where ideal organic mixture 21 has been assumed and Fig. 8b shows the corresponding data for the unity activity case (see Eq. 22 5). The point of 50%-partitioning (c_t , Eqs. 20-21) is also shown. As described in Sect. 2.2.4 we 23 expect c_t to be a reasonable estimate for the limit for complete dissolution, if the complex 24 mixture is reduced into a two-component mixture of completely soluble and insoluble 25 components. It should be noted, however, that the water content and ε of the droplet at the point 26 of activation depend on supersaturation (see Eq. 25), causing also a dependence between c_t and 27 s_c . Furthermore, Fig. 8 illustrates a difference in the solubility-dependence of the partitioning 28 behavior for the two organic activity assumptions. The ideal mixture displays a symmetrical sigmoidal dependence around c_t . For the unity activity case, on the other hand, the undissolved 29 fraction is asymmetric around the 50%-value – dropping rapidly to zero above c_t but 30 approaching 1 asymptotically below c_t . 31

1 Figure 9 shows the distributions of the solubility bins containing the 50%-partitioning 2 points (c_t , Eqs. 20-21) on a decadal basis for all the activation points studied, illustrating also 3 the differences between the two assumptions about the organic phase activity. The c_t values for 4 the ideal organic mixture (Fig. 9a, based on 2465 points) display a symmetrical distribution around the median value of about 10 g L⁻¹. Also, a modest dependence of c_t on the number of 5 components is observed: the cases with 3 and 5 components display slightly higher c_t values as 6 7 compared with the cases with larger n. This apparent dependence is probably due to the discrete 8 nature of the solubility distributions in combination with the fact that for the different solubility 9 ranges (see Table 2) only the lower end of the distribution is changed while the upper end is always at 1000 g L⁻¹. The unity activity case displays a much stronger *n*-dependence (Fig. 9b, 10 based on 3492 points): if analysed separately, the median c_t shifts from about 0.1 to 10 g L⁻¹ 11 when *n* changes from 100 to 10 and 5, and up to 100 g L⁻¹ for n = 3. Unlike the ideal mixture 12 13 this behavior is explained by the actual dissolution thermodynamics: in a system where the 14 components do not affect each other's solubility directly (i.e. the dissolution of a compound *i* 15 is only limited by its own presence in the aqueous phase), the amount of dissolved material is 16 only dependent on the total water content and is larger the larger the number of dissolvable 17 components. If all the different mixtures are integrated together the median c_t for the unity activity assumption lies at about 1 g L^{-1} – a decade lower than for the ideal mixture case. Figure 18 10 provides a more detailed look on the sensitivity of the c_t to supersaturation, n and molar mas 19 20 for one of the distributions (Distribution 1, mid solubility range, see Table 2). As expected, c_t 21 depends considerably on supersaturation. It can also be seen that while c_t shows some sensitivity 22 to the number of components (in line with Fig. 9) and molar mass, in our case by far the most 23 critical assumption is related to the organic mixture thermodynamics.

24 The fitted ε_{eff} values are compared in Figure 11 to the fraction of mass with solubilities 25 above the median c_t for the 134 mixtures that activated at the probed conditions, individually 26 for both organic activity assumptions. The fitted dissolved fraction corresponds well to the 27 fraction of mass with solubilities above the 50%-partitioning point, as predicted by the 28 theoretical principles outlined in Sect. 2.2.4. Also, the solubilities of the different distributions 29 with varying shapes, numbers of components and solubility ranges can be represented reasonably well with a single median c_t (equal to 10 g L⁻¹ for the ideal mixture case and 1 g L⁻ 30 31 ¹ for the unity activity case, see Fig. 9) with median deviations between the fitted ε_{eff} and the 32 fraction above c_t of 9% and 8%, respectively. On the other hand, these results indicate that the 33 soluble fractions determined from experimental data on CCN activation provide information

about the fraction of material with solubility above c_t . There are 8 points that do not seem to 1 2 follow the general trend, however: a group of points with all the material above the threshold solubility can display a variety of fitted ε_{eff} values. These are all points that correspond to the 3 4 "high" solubility ranges. Distributions 5 and 6 (see Table 2) with n = 3 are among these points 5 for both organic activity assumptions. For the ideal mixture case also distribution 1 with 3 components diverges from the general trend, and for the unity activity assumption distribution 6 7 5 falls into the category regardless of the number of components. These points thus contribute 8 to the high ends of the c_t distributions in Fig. 9.

9 The deviations in the activation diameters as predicted by the three simplified solubility 10 representations (complete dissolution, ε and κ) are displayed in Figure 12 as a function of the 11 mixture properties for $s_c = 0.1\%$, 0.3%, 0.6% and 1%. Again, the two different assumptions 12 about the organic phase activity are treated separately due to their different limiting solubilities 13 c_t (Fig. 9) and different shapes of the partitioning distributions (Fig. 8). Also, the points close 14 to complete dissolution at the point of activation ($\varepsilon \ge 0.8$, see Fig. 5 and its explanation in the 15 text) are presented with a different color (gray symbols) than the points where the activation diameter differs significantly from the complete dissolution prediction ($\varepsilon < 0.8$, black symbols). 16 17 As expected, the complete dissolution assumption performs better for the more water soluble 18 organic mixtures. Figs. 12a and 12b illustrate this by showing the relationship between the norm 19 of the error in the predicted $D_{p,act}$ and the fraction of material below the median c_t (10 and 1 g L⁻¹) for the two organic activity assumptions. The larger the amount of material below the 20 21 solubility limit, the larger the deviation from the full model predictions. For the κ and ε models, 22 on the other hand, the variable best correlating with the error in $D_{p,act}$ induced by the 23 simplification is different for the different organic activity assumptions – although close to 24 complete dissolution these models also do well, nearly independent of the solubility of the distribution. For the ideal mixture case the fraction of mass between 1 and 100 g L⁻¹ correlates 25 26 better with the error (Fig. 12c, 12e) than the mass fraction below any solubility limit (not 27 shown), while for the unity activity case the material at the low end of the distribution (mass 28 fraction below 1 g L^{-1}) performs better (see Fig. 12d, f). The reason for this lies in the different 29 shapes of the partitioning distributions resulting from the two assumptions (Fig. 8). For the 30 symmetric partitioning curve of the ideal mixture case, the predicted ε and κ are most sensitive to differences in the partitioning behaviour between compounds within the range of c_t 31 corresponding to the supersaturation and particle diameter ranges studied here, i.e. 1-100 g L⁻¹ 32 (see Figs. 12c, 12e). Anything outside these boundaries will behave as completely soluble or 33

insoluble throughout the studied supersaturation space, thus not introducing a significant error 1 2 when constant ε is assumed to describe the mixture. However, the more the material that can behave as either insoluble or soluble depending on the conditions, the larger error we introduce 3 4 by assuming a constant ε for a given mixture at any conditions. The story is different for the 5 unity activity case (Figs. 12d, 12f): as the shape of the partitioning distribution (Fig. 8) does not depend on c_t , the compounds with solubilities below c_t will contribute relatively much more the 6 7 fitted ε_{ff} than for the previous case, and thus the more material there is in the "tail" of the 8 partitioning distribution, the worse the assumption about a single ε for the whole distribution.

9 To relate the theoretical work conducted here to realistic atmospheric organic aerosol 10 mixtures, Fig. 13 displays an example of a solubility distribution representing SOA formed from dark ozonolysis of α -pinene (Chen et al., 2008). The solubilities have been estimated with 11 12 SPARC (see Wania et al., 2014 and references therein). The average molar masses and O:C ratios in each solubility bin are also displayed, along with the c_t values corresponding to the 13 14 activation points with limited solubility – assuming that the organics form an ideal mixture with 15 each other. Most of the material is predicted to have solubilities between 1 and 100 g L^{-1} , indicating that this fresh α -pinene SOA is at the critical range of solubilities for limited 16 17 dissolution at the point of activation. This in turn suggests that the observed difference between 18 the κ values inferred from hygrosopicity and CCN-activation for this mixture might largely 19 result simply from the distribution of solubilities present.

20

21 **4** Discussion and conclusions

22 We have studied the relationship between CCN activation and solubility of 144 different theoretically constructed complex organic mixtures using Köhler theory, accounting for the 23 24 partial solubility of the compounds in water and assuming ideal interactions between the 25 dissolved molecules and water. The mixtures encompassed a wide variety of solubilities, and 26 were represented by solubility distributions with various solubility ranges and shapes 27 (analogously to the volatility basis set, VBS). Two limiting assumptions (ideal mixture vs. unity 28 activity) about the interactions between the organics in the insoluble organic phase were tested. 29 The results using this comprehensive solubility representation (termed as "the full model") were 30 compared to commonly-used simplified descriptions of solubility: 1) assuming complete 31 dissolution; 2) representing the mixture with single hygroscopicity parameter κ ; 3) representing the mixture with a single soluble fraction ε_{eff} . The calculations were carried out for particle dry sizes ranging from 20 to 500 nm and supersaturations between 0.03% and 5%, thus probing an atmospherically representative parameter space and resulting in total 5957 unique activation points.

5 Comparing the full model predictions to the simplified solubility descriptions, we find 6 that assuming complete dissolution under-predicts the activation diameter up to about a factor 7 of two for the studied mixtures. Our results indicate that about 70-80% of the material needs be 8 dissolved at the point of activation for the complete dissolution assumption to predict activation 9 diameters that are within 10% of that produced by the full solubility treatment. Adding a single parameter to describe the mixture solubility improved the situation considerably: the 10 11 predictions of activation diameters based on a single ε or κ for a given mixture were within 10% 12 of the full model predictions, the difference between these two approaches being only marginal.

The fitted soluble fractions, ε_{eff} , describing the solubility distribution (and thus the fitted 13 14 κ which is directly proportional to ε) were found to correspond well to the fraction of dry particle 15 material with solubilities larger than a given threshold solubility c_t . For the ideal organic mixture assumption the median c_t was 10 g L⁻¹, most of the values falling between 1 and 100 g 16 L⁻¹, depending somewhat on the supersaturation. Since the material with solubilities outside 17 18 this range can generally be treated as completely soluble or insoluble in CCN activation 19 calculations, the error made by using the single soluble fraction increased when a larger fraction 20 of material was present in this critical range. For the unity activity case the median c_t was 1 g 21 L^{-1} , but decreased with the number of components present in the mixture, n. For the range of n = 3-100 studied here, the typical c_t values were between 0.1 and 10 g L⁻¹. Due to the asymmetric 22 23 shape of the aqueous-organic phase partitioning of the organics in the unity activity case, the 24 simplified models performed better, the more material with solubilities larger than c_t was 25 present in the particles. In general, the median values for c_t represented the soluble fraction with 26 a reasonable accuracy in most of the studied mixtures, although the exact composition of the 27 mixtures varied considerably.

Our values for the limiting solubilities for complete dissolution are in agreement with the values of 3 g L^{-1} and 1 g L^{-1} previously reported by Huff Hartz et al. (2006) and Chan et al. (2008) based on experimental data on specific mixtures. Our results on the two different assumptions about the organic phase activities indicate that the mixtures investigated by these past studies were probably somewhat non-ideal, where the compounds hindered each other's dissolution less than would be expected for a fully ideal mixture. On the other hand, in light of our findings, the observations of the close to complete dissolution of SOA at activation (Huff Hartz et al., 2005; Engelhart et al., 2011) indicate that the majority of the material in the studied SOA mixtures had solubilities larger than 10 g L⁻¹. Our results suggest that even with vastly different solubility distributions one can yield very similar CCN-activation behavior (and consequently values of κ or ε), as the parameter that matters is the material above c_t .

7 The above results suggest that the solubility range corresponding to limited solubility in CCN activation is between 0.1-100 g L⁻¹, and resolving the solubility distributions of aerosol 8 9 mixtures outside this range provides little added value for understanding their CCN activation. In fact, this is probably a conservative estimate, as in most cases most material below 1 g L⁻¹ is 10 practically insoluble and most material above 10 g L^{-1} completely soluble – even considering 11 the uncertainty in the organic mixture activity. These results can be used to guide the 12 13 representation of the cloud activation properties of complex mixtures, and provide quantitative 14 support for the previous notion that knowing the water-soluble fraction of the aerosol mixture 15 in question is the key in most applications. We provide quantitative estimates on how this soluble fraction should be defined in the case of complex mixtures, and when such a simplified 16 17 model is not expected to perform well.

18 There are, however, some limitations to our approach to keep in mind when applying 19 the results to laboratory experiments or atmospheric data. Since the focus of this work was 20 strictly on the links between solubility and CCN activation, we did not explore in depth how variation of surface activity, molecular mass, pure-component density, the gas-droplet 21 22 partitioning of the organic compounds or non-ideality of water with respect to the aqueous 23 phase would affect the results (see also Suda et al. 2012, 2014; Topping et al., 2012). 24 Furthermore, temperature was assumed to stay constant at 298 K. Since many of the 25 thermodynamic properties relevant to CCN activation are temperature-dependent (see e.g. 26 Christensen et al., 2012), future work investigating the impact of temperature on the phenomena 27 studied here is needed. Furthermore, the solubility and organic phase activity should naturally be linked to the aqueous phase activity coefficients predicted and used in a number of previous 28 studies (see e.g. Topping et al., 2013 and references therein), although lack of well-defined 29 experimental data on organic phase activities and mixture solubilities currently hinders 30 31 quantitative evaluation of the current multi-phase mixture thermodynamic models (see e.g. 32 Cappa et al., 2008). Evaluation of the concepts and approaches presented here (e.g. the solubility limits *c_t*) with laboratory studies on well-defined complex mixtures over a wide range
 of solubilities, supersaturations and particle diameters would therefore be warranted.

3 Another important future step would be applying the introduced solubility distribution 4 framework in the atmospheric context. On one hand, the framework is likely to be useful in 5 modeling the evolution of the CCN-activity of secondary organics. We expect the solubility 6 distributions (and thus c_t) to depend on the SOA mixture properties such as the O:C ratio and 7 the molar masses of the mixture constituents, which in turn evolve due to atmospheric chemistry 8 (Kuwata et al., 2012; Suda et al., 2014), coupling the solubility distributions to the different 9 dimensions of the VBS (Donahue et al., 2006; Donahue et al., 2011; Kroll et al., 2011; Donahue 10 et al., 2012; Shiraiwa et al., 2014). A systematic study investigating the interlinkages between 11 these variables in light of the available experimental data from field and laboratory would thus be a valuable future contribution. Futhermore, as atmospheric aerosol particles are typically 12 mixtures of organic and inorganic constituents, the molecular interactions between atmospheric 13 14 organics and inorganics as well as their effect of the pure-component solubility should be 15 expanded. On the other hand, the solubility distribution can be used as a simplifying concept aiding in large-scale model simulations coupling atmospheric chemistry with the dynamics of 16 17 cloud formation. With the assumptions applied here the CCN-activation calculation itself is 18 computationally relatively light, slowing down considerably only if n is of the order 10000 or 19 larger with typical present-day computational resources. Therefore, using the solubility 20 distribution framework within an atmospheric model is a good option if accuracy beyond the 21 simple one-parameter approaches is required, or as an intermediate tool linking atmospheric 22 age to the effective ε or κ describing a given mixture.

23

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29 **References**

Albrecht, B. A.: Aerosols, cloud microphysics, and fractional cloudiness, Science, 245, 1227–
1230, 1989.

- 1 Asa-Awuku, A., and Nenes, A.: Effect of solute dissolution kinetics on cloud droplet formation:
- 2 Extended Köhler theory, Geophys. Res., 112, D22201, doi: 10.1029/2005JD006934, 2007.
- 3 Asa-Awuku, A., Engelhart, G. J., Lee, B. H., Pandis, S. N., and Nenes, A.: Relating CCN
- 4 activity, volatility, and droplet growth kinetics of β -caryophyllene secondary organic aerosol,
- 5 Atmos. Chem. Phys., 9, 795–812, 2009.
- 6 Asa-Awuku, A., Nenes, A., Gao, S., Flagan, R. C and Seinfeld, J. H.: Water-soluble SOA from
- 7 alkene ozonolysis: composition and droplet activation kinetics inferences from analysis of CCN
- 8 activity, Atmos. Chem. Phys., 10, 1585–1597, 2010.
- 9 Banerjee, S.: Solubility of organic mixtures in water, Environ. Sci. Technol., 18, 587–591,
 10 1984.
- 11 Cappa, C. D., Lovejoy, E. R., and Ravishankara, A. R.: Evidence for liquid-like and nonideal
- 12 behavior of a mixture of organic aerosol components, Proc. Natl. Acad. Sci. U. S. A. 105,
- 13 18687, 2008.
- 14 Chan, M. N., Kreidenweis, S. M., and Chan, C. K.: Measurements of the hygroscopic and
- deliquescence properties of organic compounds of different solubilities in water and their
 relationship with cloud condensation nuclei activities, Environ. Sci. Technol. 42 3602–3608,
 2008.
- 18 Chen, Q., Liu, Y., Donahue, N. M., Shilling, J. E., and Martin, S. T.: Particle-phase chemistry
- 19 of secondary organic material: Modeled compared to measured O:C and H:C elemental ratios
- 20 provide constraints, Environ. Sci. Technol., 45, 4763, doi:10.1021/es104398s, 2011.
- Christensen, S. I. and Petters, M. D.: The role of temperature in cloud droplet activation., J.
 Phys. Chem. A, 116, 9706–17, doi:10.1021/jp3064454, 2012.
- Cruz, C. N. and Pandis, S. N.: A study of the ability of pure secondary organic aerosol to act as
 cloud condensation nuclei, Atmos. Environ., 31, 2205–2214, doi:10.1016/S13522310(97)00054-X, 1997.
- Cruz, C. N. and Pandis, S. N.: The effect of organic coatings on the cloud condensation nuclei
 activation of inorganic atmospheric aerosol, J. Geophys. Res., 103, 13 111–13 123, 1998.
- 28 Donahue, N. M., Epstein, S. a., Pandis, S. N. and Robinson, a. L.: A two-dimensional volatility
- 29 basis set: 1. organic-aerosol mixing thermodynamics, Atmos. Chem. Phys., 11, 3303–3318,
- 30 doi:10.5194/acp-11-3303-2011, 2011.

- 1 Donahue, N. M., Kroll, J. H., Pandis, S. N. and Robinson, a. L.: A two-dimensional volatility
- 2 basis set Part 2: Diagnostics of organic-aerosol evolution, Atmos. Chem. Phys., 12(2), 615–
- 3 634, doi:10.5194/acp-12-615-2012, 2012.
- 4 Donahue, N. M., Robinson, a L., Stanier, C. O. and Pandis, S. N.: Coupled partitioning, dilution,
- 5 and chemical aging of semivolatile organics., Environ. Sci. Technol., 40, 2635–43 [online]
- 6 Available from: http://www.ncbi.nlm.nih.gov/pubmed/16683603, 2006.
- 7 Duplissy, J., DeCarlo, P. F., Dommen, J., Alfarra, M. R., Metzger, A., Barmpadimos, I., Prevot,
- 8 a. S. H., Weingartner, E., Tritscher, T., Gysel, M., Aiken, a. C., Jimenez, J. L., Canagaratna, M.
- 9 R., Worsnop, D. R., Collins, D. R., Tomlinson, J. and Baltensperger, U.: Relating
- 10 hygroscopicity and composition of organic aerosol particulate matter, Atmos. Chem. Phys., 11,
- 11 1155–1165, doi:10.5194/acp-11-1155-2011, 2011.
- 12 Duplissy, J., Gysel, M., Alfarra, M. R., Dommen, J., Metzger, A., Prevot, S. H., Weingartner,
- 13 E., Laaksonen, A., Raatikainen, T., Good, N., Turner, S. F., McFiggans, G., and Baltensperger,
- 14 U.: Cloud forming potential of secondary organic aerosol under near atmospheric conditions,
- 15 Geophys. Res. Lett., 35, L03818, doi:10.1029/2007GL031075, 2008.
- 16 Dusek, U., Frank, G. P., Hildebrandt, L., Curtius, J., Schneider, J., Walter, S., Chand, D.,
- 17 Drewnick, F., Hings, S., Jung, D., Borrmann, S., and Andreae, M. O.: Size matters more than
- 18 chemistry for cloud-nucleating ability of aerosol particles, Science, 312, 1375–1378, 2006.
- Engelhart, G. J., Hildebrandt, L., Kostenidou, E., Mihalopoulos, N., Donahue, N. M., and
 Pandis, S. N.: Water content of aged aerosol, Atmos. Chem. Phys., 11, 911–920, 2011.
- 21 Engelhart, G. J., Moore, R. H., Nenes, A., and Pandis, S. N.: Cloud condensation nuclei activity
- of isoprene secondary organic aerosol, J. Geophys. Res., 116, D02207,
 doi:10.1029/2010JD014706, 2011.
- Farmer, D. K., Cappa, C. D. and Kreidenmweis, S. M.: Atmospheric processes and their
 controlling influence on cloud condensation nuclei activity, Chem. Rev., doi:
 10.1021/cr5006292, 2015.
- Goldstein, A. H. and Galbally, I. E.: Known and unexplored organic constituents in the earth's
 atmosphere, Environ. Sci. Technol., 41, 1514–1521, 2007.
- 29 Good, N., Topping, D. O., Duplissy, J., Gysel, M., Meyer, N. K., Metzger, a., Turner, S. F.,
- 30 Baltensperger, U., Ristovski, Z., Weingartner, E., Coe, H. and McFiggans, G.: Widening the

- 1 gap between measurement and modelling of secondary organic aerosol properties, Atmos.
- 2 Chem. Phys., 10, 2577–2593, doi:10.5194/acp-10-2577-2010, 2010.
- 3 Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M. and
- Dommen, J.: The formation, properties and impact of secondary organic aerosol: current and
 emerging issues, Atmos. Chem. Phys., 5155–5236, 2009.
- 6 Hori, M., Ohta, S., Murao, N., and Yamagata, S.: Activation capability of water soluble organic
- 7 substances as CCN, Aerosol Sci., 34, 219, 2003.
- 8 Hoyle, C. R., Boy, M., Donahue, N. M., Fry, J. L., Glasius, M., Guenther, A., Hallar, A. G.,
- 9 Huff Hartz, K., Petters, M. D., Petäjä, T., Rosenoern, T. and Sullivan, A. P.: A review of the
- 10 anthropogenic influence on biogenic secondary organic aerosol, Atmos. Chem. Phys., 11, 321–
- 11 343, doi:10.5194/acp-11-321-2011, 2011.
- 12 Huff Hartz, K. E., Rosenoern, T., Ferchak, S. R., Raymond, T. M., Bilde, M., Donahue, N. M.,
- 13 and Pandis, S. N.: Cloud condensation nuclei activation of monoterpene and sesquiterpene
- 14 secondary organic aerosol, J. Geophys. Res., 110, D14208, doi:10.1029/2004JD005754, 2005.
- 15 Huff-Hartz, K. E. H., Tischuk, J. E., Chan, M. N., Chan, C. K., Donahue, N. M., and Pandis, S.
- N.: Cloud condensation nuclei activation of limited solubility organic aerosol, Atmos. Environ.,
 40, 605–617, 2006.
- 18 Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, a S. H., Zhang, Q., Kroll, J. H.,
- 19 DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, a C., Docherty, K. S., Ulbrich, I. M.,
- 20 Grieshop, a P., Robinson, a L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. a, Hueglin,
- 21 C., Sun, Y. L., Tian, J., Laaksonen, a, Raatikainen, T., Rautiainen, J., Vaattovaara, P., Ehn, M.,
- 22 Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J., Dunlea, E. J., Huffman, J. a,
- 23 Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo, Y., Schneider, J., Drewnick,
- 24 F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D., Cottrell, L., Griffin, R., Takami, a,
- 25 Miyoshi, T., Hatakeyama, S., Shimono, a, Sun, J. Y., Zhang, Y. M., Dzepina, K., Kimmel, J.
- 26 R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, a M., Williams, L. R., Wood, E. C.,
- 27 Middlebrook, a M., Kolb, C. E., Baltensperger, U. and Worsnop, D. R.: Evolution of organic
- 28 aerosols in the atmosphere., Science, 326, 1525–1529, doi:10.1126/science.1180353, 2009.
- 29 King, S. M., Rosenoern, T., Shilling, J. E., Chen, Q. and Martin, S. T.: Cloud condensation
- 30 nucleus activity of secondary organic aerosol particles mixed with sulfate, Geophys. Res. Lett.,
- 31 34, L24806, doi:10.1029/2007GL030390, 2007.

- 1 King, S. M., Rosenoern, T., Shilling, J. E., Chen, Q., Wang, Z., Biskos, G., McKinney, K. A.,
- 2 Pöschl, U and Martin, S. T.: Cloud droplet activation of mixed organic-sulfate particles
- 3 produced by the photooxidation of isoprene, Atmos. Chem. Phys., 10, 3953–3964, 2010.
- 4 King, S. M., Rosenoern, T., Shilling, J. E., Chen, Q., and Martin, S. T.: Increased cloud
- 5 activation potential of secondary organic aerosol for atmospheric mass loadings, Atmos. Chem.
- 6 Phys., 9, 2959–2972, 2009.
- 7 Kreidenweis, S. M., Petters, M. D., and DeMott, P. J.: Deliquescence-controlled activation of
- 8 organic aerosols, Geophys. Res. Lett., 33, L06801, doi: 10.1029/2005GL024863, 2006.
- 9 Kroll, J. H., Donahue, N. M., Jimenez, J. L., Kessler, S. H., Canagaratna, M. R., Wilson, K.
- 10 R., Altieri, K. E., Mazzoleni, L. R., Wozniak, A. S., Bluhm, H., Mysak, E. R., Smith, J. D.,
- 11 Kolb, C. E., Worsnop, D. R. Carbon oxidation state as a metric for describing the chemistry of
- 12 atmospheric organic aerosol, Nature Chem., 3, 133-139, 2011.
- 13 Kuwata, M., Shao, W., Lebouteiller, R., Martin, S. T.: Classifying organic materials by oxygen-
- 14 to-carbon elemental ratio to predict the activation regime of cloud condensation nuclei (CCN),
- 15 Atmos. Chem. Phys., 13, 5309-5324, 2013.
- 16 Massoli, P., Lambe, A. T., Ahern, A. T., Williams, L. R., Ehn, M., Mikkil a, J., Canagaratna,
- 17 M. R., Brune, W. H., Onasch, T. B., Jayne, J. T., Petäjä, T., Kulmala, M., Laaksonen, A., Kolb,
- 18 C. E., Davidovits, P., and Worsnop, D. R.: Relationship between aerosol oxidation level and
- 19 hygroscopic properties of laboratory generated secondary organic aerosol (SOA) particles,
- 20 Geophys. Res. Lett., 37, L24801, doi: 10.1029/2010GL045258, 2010.
- 21 McFiggans, G., Artaxo, P., Baltensperger, U., Coe, H., Facchini, M., Feingold, G., Fuzzi, S.,
- 22 Gysel, M., Laaksonen, A., Lohmann, U., Mentel, T., Murphy, D., O'Dowd, C., Snider, J., and
- 23 Wein- gartner, E.: The effect of physical and chemical aerosol properties on warm cloud droplet
- 24 activation, Atmos. Chem. Phys., 6, 2593–2649, 2006.
- 25 Petters, M. D., and Kreidenweis, S. M.: A single parameter representation of hygroscopicity
- 26 growth and cloud condensation nucleus activity, Atmos. Chem. Phys., 7, 1961–1971, 2007.
- 27 Petters, M. D., and Kreidenweis, S. M.: A single parameter representation of hygroscopicity
- 28 growth and cloud condensation nucleus activity Part 2: Including solubility, Atmos. Chem.
- 29 Phys., 8, 6273–6279, 2008.

- 1 Petters, M. D., and Kreidenweis, S. M.: A single parameter representation of hygroscopic
- 2 growth and cloud condensation nucleus activity Part 3: Including surfactant partitioning,
- 3 Atmos. Chem. Phys. Discuss., 12, 22687–22712, 2012.
- 4 Petters, M. D., H. Wex, E. Hallberger, L. Poulain, C. M. Carrico, G. McMeeking, S. M.
- 5 Kreidenweis, F. Stratmann, and Massling, A.: Towards closing the gap between hygroscopic
- 6 growth and activation for secondary organic aerosol Part 2: Theoretical approaches, Atmos.
- 7 Chem. Phys., 9, 3999–4009, 2009a.
- 8 Petters, M. D., Kreidenweis, S. M., Prenni, A.J., Sullivan, R. C., Carrico, C. M., Koehler, K.
- 9 A., and Ziemann, P.J.: Role of molecular size in cloud droplet activation, J. Geophys. Res., 114,
- 10 D07209, doi: 10.1029/2008JD011532, 2009b.
- 11 Petters, M. D., Carrico, C. M., Kreidenweis, S. M., Prenni, A. J., DeMott, P.J., Collett Jr, J. L.,
- 12 and Hans Moosmüller, H.: Cloud condensation nucleation activity of biomass burning aerosol,
- 13 J. Geophys. Res., 114, D22205, doi: 10.1029/2009JD012353, 2009c.
- 14 Poulain, L., Wu, Z., Petters, M. D., Wex, H., Hallbauer, E., Wehner, B., Massling, a.,
- 15 Kreidenweis, S. M. and Stratmann, F.: Towards closing the gap between hygroscopic growth
- 16 and CCN activation for secondary organic aerosols Part 3: Influence of the chemical
- 17 composition on the hygroscopic properties and volatile fractions of aerosols, Atmos. Chem.

18 Phys., 10(8), 3775–3785, doi:10.5194/acp-10-3775-2010, 2010.

- Prausnitz, J. M., Lichtenhaler, R. M., Gomez de Azevedo, E.: Molecular thermodynamics offluid-phase equilibria, Prentice Hall; 3 edition, 1998.
- Prenni, A. J., M. D. Petters, S. M. Kreidenweis, P. J. DeMott, and Ziemann, P.J.: Cloud droplet
 activation of secondary organic aerosol, J. Geophys. Res., 112, D10223, doi:
 10.1029/2006JD007963, 2007.
- Pruppacher, H.R., and J.D. Klett, Microphysics of Clouds and Precipitation, 954 pp., Kluwer
 Acad., Norwell, Mass., 1997.
- Psichoudaki, M. and Pandis, S. N.: Atmospheric aerosol water-soluble organic carbon
 measurement: a theoretical analysis, Environ. Sci. Technol., 47, 9791–8,
 doi:10.1021/es402270y, 2013.
- 29 Rastak, N., Silvergren, S., Zieger, P., Wideqvist, U., Ström, J., Svenningsson, B., Maturilli, M.,
- 30 Tesche, M., Ekman, a. M. L., Tunved, P. and Riipinen, I.: Seasonal variation of aerosol water

- 1 uptake and its impact on the direct radiative effect at Ny-Ålesund, Svalbard, Atmos. Chem.
- 2 Phys., 14, 7445–7460, doi:10.5194/acp-14-7445-2014, 2014.
- Raymond, T. M. and Pandis, S. N.: Cloud activation of single-component organic aerosol
 particles, J. Geophys. Res., 107, 4787, doi:10.1029/2002JD002159, 2002.
- Raymond, T. M. and Pandis, S. N.: Formation of cloud droplets by multicomponent organic
 particles, J. Geophys. Res., 108, 4469, doi:10.1029/2003JD003503, 2003.
- 7 Seinfeld, J. H. and Pandis, S. N.: Atmospheric chemistry and physics: From air pollution to
 8 climate change, John Wiley, New York, 2006.
- 9 Shiraiwa, M., Berkemeier, T., Schilling-Fahnestock, K. A., Seinfeld, J. H., and Pöschl, U.:
- 10 Molecular corridors and kinetic regimes in the multiphase chemical evolution of secondary
- 11 organic aerosol, Atmos. Chem. Phys., 14, 8323-8341, doi:10.5194/acp-14-8323-2014, 2014.
- 12 Shulman, M. L., Jacobson, M. C., Charlson, R. J., Synovec, R. E., and Young, T. E.: Dissolution
- 13 behavior and surface tension effects of organic compounds in nucleating cloud droplets. J.
- 14 Geophys. Res., 23, 277-280, 1996.
- Sotiropoulou, R. E. P., Medina, J., and Nenes, A.: CCN predictions: Is theory sufficient for
 assessments of the indirect effect?, Geophys. Res. Lett., 33, L05816,
 doi:10.1029/2005GL025148, 2006.
- 18 Spracklen, D. V., Jimenez, J. L., Carslaw, K. S., Worsnop, D. R., Evans, M. J., Mann, G. W.,
- 19 Zhang, Q., Canagaratna, M. R., Allan, J., Coe, H., McFiggans, G., Rap, A., and Forster, P.:
- 20 Aerosol mass spectrometer constraint on the global secondary organic aerosol budget, Atmos.
- 21 Chem. Phys., 11, 12109-12136, doi:10.5194/acp-11-12109-2011, 2011.
- 22 Suda, S. R., Petters, M. D., Matsunaga, A., Sullivan, R. C., Ziemann, P. J., and Kreidenweis,
- 23 S. M.: Hygroscopicity frequency distributions of secondary organic aerosols, J. Geophys. Res.
- 24 VOL. 117, D04207, doi:10.1029/2011JD016823, 2012.
- 25 Suda, S. R., Petters, M. D., Yeh, K., Strollo, C., Matsunaga, A., Faulhaber, A., Ziemann, P. J.,
- Prenni, A. J., Carrico, C. M., Sullivan, R. C. and Kreidenweis, S. M.: Influence of functional
 groups on organic aerosol cloud condensation nucleus activity, 2014.
- 28 Swietlicki, E., Hansson, H.-C., Hämeri, K., Svenningsson, B., Massling, A., Mcfiggans, G.,
- 29 Mcmurry, P. H., Petäjä, T., Tunved, P., Gysel, M., Topping, D., Weingartner, E., Baltensperger,
- 30 U., Rissler, J., Wiedensohler, A. and Kulmala, M.: Hygroscopic properties of submicrometer

- atmospheric aerosol particles measured with H-TDMA instruments in various environments a 1
- 2 review, Tellus B, doi:10.1111/j.1600-0889.2008.00350.x, 2008.
- 3 Topping, D. O., McFiggans, G. B.: Tight coupling of particle size, number and composition in 4 atmospheric cloud droplet activation, Atmos. Chem. Phys., 12, 3253–3260, 2012.
- 5 Topping, D.O., Barley, M., and McFiggans, G.: Including phase separation in a unified model
- 6 to calculate partitioning of vapours to mixed inorganic-organic aerosol particles, Faraday
- 7 Discuss., 165, 273, 2013.

- 8 Twomey, S.: Pollution and the planetary albedo, Atmos. Environ., 8, 1251–1256, 1974.
- 9 VanReken, T. M., Ng, N. L., Flagan, R. C., and Seinfeld, J. H.: Cloud condensation nucleus
- 10 activation properties of biogenic secondary organic aerosol, J. Geophys. Res., 110, D07206, 11 doi:10.1029/2004JD005465, 2005.
- 12 Varutbangkul, V., Brechtel, F. J., Bahreini, R., Ng, N. L., Keywood, M. D., Kroll, J. H., Flagan,
- R. C., Seinfeld, J. H., Lee, A., and Goldstein, A. H.: Hygroscopicity of secondary organic 13 14 aerosols formed by oxidation of cycloalkenes, monoterpenes, sesquiterpenes, and related
- compounds, Atmos. Chem. Phys. ,6, 2367-2388, 2006. 15
- 16 Volkamer, R., Jimenez, J. L., San Martini, F., Dzepina, K., Zhang, Q., Salcedo, D., Molina, L.
- 17 T., Worsnop, D. R. and Molina, M. J.: Secondary organic aerosol formation from anthropogenic
- air pollution: Rapid and higher than expected, Geophys. Res. Lett., 33, L17811, 18 19 doi:10.1029/2006GL026899, 2006.
- 20 Wania, F., Lei, Y. D., Wang, C., Abbatt, J. P. D., and Goss, K.-U.: Novel methods for predicting gas-particle partititioning during the formation of secondary organic aerosol, Atmos. Chem. 21 22 Phys. 14, 13189, 2014.
- 23 Wex, H., Petters, M. D., Carrico, C. M., Hallbauer, E., Massling, A., McMeeking, G. R.,
- Poulain, L., Wu, Z., Kreidenweis, S. M., and Stratmann, F.: Towards closing the gap between 25 hygroscopic growth and activation for secondary organic aerosol: Part I - evidence from
- measurements, Atmos. Chem. Phys., 9, 3987-3997, 2009. 26
- 27 Zieger, P., Weingartner, E., Henzing, J., Moerman, M., de Leeuw, G., Mikkilä, J., Ehn, M.,
- 28 Petäjä, T., Clémer, K., van Roozendael, M., Yilmaz, S., Frieß, U., Irie, H., Wagner, T.,
- Shaiganfar, R., Beirle, S., Apituley, a., Wilson, K. and Baltensperger, U.: Comparison of 29
- 30 ambient aerosol extinction coefficients obtained from in-situ, MAX-DOAS and LIDAR

- measurements at Cabauw, Atmos. Chem. Phys., 11, 2603-2624, doi:10.5194/acp-11-2603-
- 2011, 2011.

- 6 7 8

Mixture model	Number of	Solubility	Other input parameters
	components	Presentation	
Complete dissolution	1	$c_{sat} \rightarrow \infty^{-1}$	M_{org}, ho_{org}
κ	1	$\kappa(c_{sat}, M_{org}, \rho_{org})$	-
Soluble fraction ε_{eff}	2	$c_{sat,1} \rightarrow \infty \ c_{sat,2} = 0$	M_{org}, ho_{org}

Table 1. Simplified descriptions of organic mixture solubilities.

 $\overline{}^{I} c_{sat}$ the solubility (saturation concentration) in aqueous solution.

Distribution ¹	Shape	Number of	[Csat,min, Csat,max,]
		components	(g L ⁻¹)
1	Flat, log <i>c</i> -axis		
2	Flat, linear <i>c</i> -axis		Low: [10 ⁻⁵ , 10
3	Log. increasing	3, 5, 10, 100	Mid: $[0.1, 10^3]$
4	Linear increasing		High: [10, 10 ³
5	Log. decreasing		
6	Linear decreasing		

Table 2. Solubility distributions of the organic mixtures considered in this study.

¹For all solubility distributions two assumptions about the organic phase activity coefficients: 18

19 (1) ideal mixture and (2) unity activity (see text for details).

 c_{sat} the solubility (pure component saturation concentration) in aqueous solution. 20

21

Table 3. Properties of water and organic compounds used in Köhler curve calculations (see Eq.

2 1).

Property (unit) ¹	Water	Organic <i>i</i>
ρ (kg m ⁻³)	1000	1500
σ (N m ⁻¹)	0.073	-
M (kg mol ⁻¹)	0.018	0.18

¹ These properties were chosen based on literature on the effective molar masses and densities
determined for laboratory SOA (Engelhart et al., 2008; Asa-Awuku et al., 2010), and assumed
to be same for every organic compound *i*.

- _



Figure 1. Schematic of the conceptual model used in the equilibrium composition calculations. The dry particle is assumed to consist of *n* organic compounds, each denoted with a subscript *i*. The wet particle is assumed to consist of a dry organic (insoluble) phase and an aqueous phase with water and dissolved organics. The aqueous phase is assumed to be in equilibrium with the ambient water vapour. Y refers to mole fractions in the organic phase, X to mole fractions in the aqueous phase, and m to the masses of the organic constituents and water. $c_{i,eq}$ refers to the equilibrium concentration of each organic compound in the aqueous solution and $p_{w,eq}$ to the equilibrium vapour pressure of water above the aqueous solution.



Figure 2. Examples of solubility distributions used in the calculations for saturation
concentrations ranging from 0.1 to 1000 g L⁻¹. (a) Linear and logarithmic flat distributions; (b)
Linear and logarithmic increasing distributions; (c) Linear and logarithmic decreasing
distributions. The numbers of the distributions refer to the numbering in Table 2 (see Sect.
2.1.4).





Figure 3. a) Examples of Köhler curves for the flat logarithmically spaced solubility distribution (Distribution 1 in Table 2) with n = 5 and the high solubility range (Table 2). The dots indicate the point of activation, black indicating incomplete dissolution ($\varepsilon < 1$) and red complete dissolution ($\varepsilon = 1$). b) The activation points determined from the model calculations (in total 7200 Köhler curves, see Table 2), corresponding to in total 5957 points in the activation dry diameter vs. supersaturation space. Also the dependence of the dissolved fraction at the point of activation is illustrated.



Figure 4. The dependence of the activation diameter for four different supersaturations (s = 1%, 0.6%, 0.3% and 0.1%, see also Fig. 3b) on the solubility range for the solubility distributions outlined in Table 2 for n = 5, and the two assumptions about the organic phase activity.



Figure 5. The activation diameter calculated using the solubility distributions (Table 2, referred to as the full model) and the simplified dissolution descriptions (Table 1) a) complete dissolution assumption; b) the hygroscopicity parameter κ ; c) the soluble fraction ε_{eff} to describe the solubility of the organic mixture. The symbols correspond to the best fits to the full model data. The black line shows the 1:1 correspondence between the two data sets.



Figure 6. The performance of the simplified solubility representations (see Table 1) in
predicting the activation diameter for a given supersaturation as compared with the full model.
The black bars depict the 25- and 75-percentiles and the gray bars the 10- and 90-percentiles.





Figure 7. a) The fitted soluble fraction ε_{eff} as function of the true dissolved fraction ε for each considered mixture (see Table 2). Symbols: mean ε over all activation points. Grey lines: the range of ε values at the different activation points for a given mixture; b) The fitted κ values as a function as function of the true dissolved fraction ε for each considered mixture. The red dashed line denotes the limit of $\kappa_{max} = 0.15$ which applies to all the studied mixtures. 1:1 lines are also indicated.



2 Figure 8. The dissolution behavior of the organic mixture corresponding to Distribution 1 with 3 n = 100 and the low solubility range (see Table 2) at the activation point for $s_c = 0.1\%$ (see Fig. 4 4). The figures depict the distribution of material in each solubility bin between the aqueous 5 and the insoluble organic phases for the two different assumptions about the organic phase 6 activity. c_t refers to the 50%-point of the partitioning (Eqs. 24-25). a) The ideal organic mixture; 7 b) the unity activity assumption.





Figure 9. The distributions of the c_t values (i.e. the 50% partitioning point, see Eqs. 24-25) at the point of activation for all the considered mixtures (Table 2) and activation points, and the two assumptions about the organic phase activity. a) The ideal organic mixture; b) the unity activity assumption.





Figure 10. Sensitivity of the c_t values to supersaturation, molar mass and number of components for Distribution 1 with the mid solubility range (see Table 2). Only points with limited solubility ($0 < \varepsilon < 1$) at the point of activation are included.





Figure 11. The relationship between the fitted dissolved fraction ε_{eff} at the point of activation and the mass fraction over the median c_t for all the considered mixtures and s = 1%, 0.6%, 0.3% and 0.1%. Closed symbols: the ideal organic mixture assumption (median $c_t = 10$ g L⁻¹). Open symbols: the unity activity assumption for the organics (median $c_t = 1$ g L⁻¹).



1 Figure 12. The solubility distribution properties best explaining the performance of the three 2 simplified solubility models, illustrated with the norm of the relative deviation of $D_{p,act}$ as 3 compared with the full model predictions for s = 1%, 0.6%, 0.3% and 0.1%. The performance 4 of the complete dissolution assumption as a function of the mass fraction with solubilities below 5 the median c_t for a) the ideal organic mixture (median $c_t = 10 \text{ g L}^{-1}$) and b) the unity organic 6 activity (median $c_t = 1 \text{ g L}^{-1}$) assumptions. The performance of the κ model as a function of the mass fraction with solubilities c) between $1-100 \text{ g L}^{-1}$ for the ideal organic mixture assumption; 7 8 d) below 1 g L⁻¹ for the unity organic activity assumption. The performance of the ε model as a function of the mass fraction with solubilities e) between 1-100 g L⁻¹ for the ideal organic 9 mixture assumption; f) below 1 g L⁻¹ for the unity organic activity assumption. The points close 10 to to complete dissolution ($\varepsilon \ge 0.8$) are shown with lighter grey than the rest of the points. The 11 12 error bars represent the variability with supersaturation and particle size. 13





Figure 13. A) An example of a solubility distribution for SOA generated in dark-ozonolysis of a-pinene SOA. The expected composition has been taken from Chen et al. (2008), and the purecomponent solubilities have been calculated with the SPARC prediction system (e.g. Wania et al., 2014 and references therein). "Misc." refers to completely miscible components. B) The dependence of the c_t value at the point of activation on supersaturation assuming an ideal organic mixture, with only points corresponding to limited solubility ($0 < \varepsilon < 1$) displayed.