Modelling the contribution of biogenic volatile organic compounds to new particle formation in the Jülich plant atmosphere chamber

4

P. Roldin^{1,2}, L. Liao¹, D. Mogensen¹, M. Dal Maso³, A. Rusanen¹, V.-M.
Kerminen¹, T. F. Mentel⁴, J. Wildt⁵, E. Kleist⁵, A. Kiendler-Scharr⁴, R. Tillmann⁴,
M. Ehn¹, M. Kulmala¹, and M. Boy¹

- 8 [1]{Department of Physics, University of Helsinki, P.O. Box 64, 00014 Helsinki, Finland}
- 9 [2] {Division of Nuclear Physics, Lund University, P.O. Box 118, 221 00 Lund, Sweden}

10 [3]{Department of Physics, Tampere University of Technology, P.O. Box 692, 33101
11 Tampere, Finland}

- [4] {Institute for Energy- and Climate Research (IEK-8), Forschungszentrum Jülich, 52425
 Jülich, Germany}
- 14 [5] {Institute of Biogeosciences (IBG-2), Forschungszentrum Jülich, 52425 Jülich, Germany}

15 Correspondence to: P. Roldin (pontus.roldin@nuclear.lu.se)

16

17 Abstract

We used the Aerosol Dynamics gas- and particle-phase chemistry model for laboratory 18 19 CHAMber studies (ADCHAM) to simulate the contribution of BVOC plant emissions to the observed new particle formation during photooxidation experiments performed in the Jülich 20 21 Plant-Atmosphere Chamber and to evaluate how well smog chamber experiments can mimic 22 the atmospheric conditions during new particle formation events. ADCHAM couples the 23 detailed gas-phase chemistry from Master Chemical Mechanism with a novel aerosol 24 dynamics and particle phase chemistry module. Our model simulations reveal that the 25 observed particle growth either may have been controlled by the formation rate of semi- and low-volatility organic compounds in the gas-phase or by acid catalyzed heterogeneous 26 27 reactions between semi-volatility organic compounds in the particle surface layer (e.g. peroxyhemiacetal dimer formation). The contribution of extremely low-volatility organic gas-28

phase compounds to the particle formation and growth was suppressed because of their rapid and irreversible wall losses, which decreased their contribution to the nano-CN formation and growth compared to the atmospheric situation. The best agreement between the modelled and measured total particle number concentration ($R^2 > 0.95$) was achieved if the nano-CN was formed by kinetic nucleation involving both sulphuric acid and organic compounds formed from OH oxidation of BVOCs.

7

8 1 Introduction

9

New particle formation, including formation of nano condensation nucleii (nano-CN) 10 11 (McMurry et al., 2011) and their growth to larger sizes, has been observed world-widely in 12 continental boundary layers and free troposphere (Kulmala et al., 2004; Mirme et al., 2010). 13 Field observations, laboratory experiments and model simulations indicate that gaseous 14 sulphuric acid (H₂SO₄) plays an important role in atmospheric nano-CN formation, yet H₂SO₄ 15 alone appears not to be able to explain all the steps of this process (Kulmala et al., 2000; Boy 16 et al., 2003; Sipilä et al., 2010; Riipinen et al., 2007; Sihto et al., 2006; Kerminen et al., 2010; 17 Kulmala et al., 2013, 2014). Basic compounds like ammonia and certain amines have been 18 proposed to act as stabilizing compounds in nano-CN clusters (Berndt et al., 2010; Almeida et al., 2013; Kurtén et al., 2008), while subsequent steps of atmospheric new particle formation 19 20 seem to rely on the presence of low-volatility organic compounds (LVOCs) (e.g., Metzger et 21 al., 2010; Paasonen et al., 2010; Riipinen et al., 2012; Ehn et al., 2014; Schobesberger et al., 2013). 22

23 Oxidation products of biogenic volatile organic compounds (BVOCs) constitute the largest 24 source of secondary organic aerosol (SOA) in the global atmosphere (Tsigaridis and 25 Kanakidou, 2003; Hallquist et al., 2009; Spracklen et al., 2011), accounting for the main 26 composition of SOA condensational growth (VanReken et al., 2006; Hao et al., 2009; Riipinen et al., 2012). BVOC oxidation also produces extremely low-volatility organic 27 28 compounds (ELVOCs) essential to the new particle formation process in the atmosphere (Ehn 29 et al., 2014). The most abundant group of BVOCs, accounting for more than half of their 30 global emissions, are terpenoids (Guenther et al., 1995). Terpenoids include compounds consisting of one to several isoprene units, e.g. isoprene (C_5H_8), monoterpenes ($C_{10}H_{16}$), and 31 sesquiterpenes ($C_{15}H_{24}$). Oxidation products of monoterpenes have substantial contribution to 32

2

SOA formation (Hoffmann et al., 1997, 1998; Laaksonen et al., 2008), and low-volatility substances produced by sesquiterpene-ozone reactions may also initiate SOA formation (Bonn and Moortgat, 2003). Oxidation of isoprene leads to the formation of SOA (Surratt et al., 2006; Claeys et al., 2004), yet isoprene may also suppress the new particle formation process due to its high reactivity with OH (Kiendler-Scharr et al., 2009). Overall, the exact contribution of BVOCs to SOA still remains uncertain, especially with respect to the initial steps of atmospheric new particle formation.

8 In this study, we used the Aerosol Dynamics, gas- and particle-phase chemistry model for 9 laboratory CHAMber studies (ADCHAM) (Roldin et al., 2014), to investigate the nano-CN 10 formation and growth during a measurement campaign conducted in the Jülich Plant 11 Atmosphere Chamber (JPAC). The in-depth analysis of the chamber measurements is 12 discussed in another paper (Dal Maso et al., 2014). Here, we use the full chamber dataset 13 including gas and particle phase measurements, either as model input or for evaluation of the 14 model results.

The main objectives of this work were to evaluate how well the JPAC experiments could mimic the real atmospheric conditions during new particle formation events over the boreal forest and to constrain the dominating mechanisms responsible for the nano-CN formation and growth.

19

20 2 Measurement set up

21

The experiments were conducted in the JPAC located at Forschungszentrum Jülich, Germany. 22 23 Detailed description regarding the chamber facility was given in previous articles (e.g., 24 Mentel et al., 2009; Schimang et al., 2006). In addition, more details about this measurement 25 campaign can be found from Dal Maso et al. (2014). In brief, the system consisted of two 26 borosilicate glass chambers with PTFE Teflon floors. The chambers were operated as 27 continuously stirred tank reactors (CSTR) with Teflon fans ensuring homogeneous air mixing. 28 Each chamber housed adjustable temperature between 10 and 50 °C with a stability of ± 0.5 °C. The small chamber (1150 L) served as plant chamber and was connected to the larger 29 chamber that worked as reaction chamber (1450 L, surface-area-to-volume ratio 4.87 m⁻¹). 30

31 Three small trees aging from three to four years were brought from Hyytiälä, Finland, and

1 included Scots pine (Pinus sylvestris), Norway spruce (Picea abies), and Silver birch (Betula 2 *pendula*). These trees, representing the main boreal forest species in Finland, were placed in 3 the JPAC plant chamber for almost two weeks before the intensive experiment campaign 4 started. This allowed them to adjust to the chamber environment. Starting from the day when 5 the trees were installed in the plant chamber, tree emissions were transferred into the reaction 6 chamber, where O₃ was added together with water vapour directly and OH was generated 7 periodically by turning on and off the UV light, in the same fashion as applied in the intensive 8 phase. New particles were formed every day during the UV-light on periods and gases and 9 particles deposited on the chamber walls. Discharge lamps (Osram HQI 400 W/D) were used 10 for illumination to simulate the solar light spectrum in both chambers. Filters (OptoChem, 11 type IR3) that reflect wavelengths between 750 and 1050 nm were used as heat shields to avoid infrared radiation inflicted plant overheating. For the model simulations we used the 12 measured spectrum for the discharge lamps in the wavelength range of 280-650 nm and a 13 14 single UV-light peak at a wavelength of 254 nm which represents the UV-spectrum from UVlight source, a Philips, TUV 40W lamp (Fig. S1). The UV-light source intensity corresponds 15 to a O₃ to O(¹D) photolysis rate of 2.9×10^{-3} s⁻¹ (Mentel et al., 2009). 16

17 The ambient air was purified by an adsorption dryer (Zander, KEA 70) and a palladium 18 catalyst (450 °C). O_3 , NO, NO₂ and VOC levels decreased significantly after passing the 19 purification system. The flow through the plant chamber was 115 L min⁻¹ from which a 20 fraction of 20 L min⁻¹ was transferred to the reaction chamber. This flow was kept nearly 21 constant by keeping the pressure drop constant. In addition, the 10 L min⁻¹ flow containing 22 ozone was added, controlled by a second flow controller.

23 The concentrations of O₃, CO₂ and H₂O were measured by commercial analytical instruments. Two Gas Chromatography Mass Spectrometer (GC-MS) systems were used, one to measure 24 the VOC concentrations from C_5 to C_{20} in the outflow air from the plant chamber (Heiden et 25 26 al., 2003), and another to identify the OH concentration by determining the decrease in the 27 concentration of 2-butanol in the reaction chamber (Kiendler-Scharr et al., 2009). Meanwhile, 28 the VOC concentration was continuously measured by an on-line Proton Transfer Reaction Mass Spectrometer (PTR-MS) in the plant and reaction chamber. The gas phase H₂SO₄ 29 30 concentration in the reaction chamber was measured by a Chemical Ionization Mass Spectrometer (CIMS) (Petäjä et al., 2009; Mauldin et al., 1998). A prototype Airmodus 31 Particle Size Magnifier (PSM) coupled with a TSI condensation particle counter (CPC) was 32

used to count the total number concentration of particles larger than ca. 1.6 nm in diameter
(Vanhanen et al., 2011) and a TSI CPC (TSI3022A) was used to measure the total
concentration of particles larger than ca. 7 nm in diameter. A Scanning Mobility Particle Sizer
(SMPS TSI3071 + TSI3025A) was used to measure the particle size distribution in the size
range of 14 - 600 nm.

The real plant emissions and the simulated day and night conditions make these experiments
suitable for evaluation of methods used to describe the atmospheric transformation (ageing) of
BVOCs and SOA beyond the first gas-phase VOC oxidation stage (e.g. the 2-Dimensional
Volatility Basis Set (2D-VBS) Donahue et al., 2011).

10

11

3 Description of the ADCHAM model and its application

12

ADCHAM is a model primarily developed for simulations of laboratory chamber experiments on SOA formation and ageing. The model includes modules for reversible partitioning of organic compounds to and from the chamber walls, all fundamental aerosol dynamics processes, detailed gas- and particle-phase chemistry and a kinetic multilayer model which can be used to simulate mass transfer limited mixing of compounds in the particle phase (Fig. 1). Below we describe how ADCHAM was set up in this work. For a more detailed description of ADCHAM we refer to Roldin et al. (2014).

20 **3.1 Gas-phase chemistry**

The gas-phase chemistry reactions were selected from the Master Chemical Mechanism (MCM) v3.2 (Jenkin et al., 1997, 2012; Saunders et al., 2003) via website: http://mcm.leeds.ac.uk/MCM/. The MCM is a near-explicit chemical degradation mechanism that simplifies the chemical path of compounds by lumping products beyond the second oxidation step. We used the Kinetic Pre-Processor (KPP) version 2.1 (Damian et al., 2002) to simulate the gas-phase chemistry.

Among all the compounds measured by GC-MS in the JPAC plant chamber, 28 organic compounds were included in the gas-phase chemistry mechanism (see Table 1). In the table, the "other MTs" equals to the concentration difference between the summation of the concentrations of the 13 selected monoterpenes by GC-MS and the total monoterpene concentration measured by PTR-MS, while the "other SQTs" equals to the summation of 1 other sesquiterpene isomers besides the listed four sesquiterpenes measured by GC-MS.

2 In total, the gas-phase chemistry mechanism has 2294 species and 6487 chemical and 3 photochemical reactions for the chemical calculations. These include relevant inorganic reactions and the full MCM chemistry path for isoprene, α -pinene, β -pinene, limonene, β -4 5 caryophyllene, toluene, 2-butanol and hexanal. We also included the initial reactions of OH, O₃, and NO₃ with the following organic compounds: myrcene, sabinene, camphene, ocimene, 6 Δ^3 -carene, "other MTs" (which we assumed to have the same rate coefficients as α -pinene), 7 8 cineole (also known as eucalyptol), farnesene, "other SQTs" (which assumed the same rate 9 coefficients as β -carvophyllene), α -terpinene, Δ -terpinene, α -phellandrene, β -phellandrene, 10 and terpinolene. Furthermore, we included the initial reactions between OH and the following 11 organic compounds: tricyclene, nonanal, bornyl acetate and methyl salicylate. Finally, the initial reactions between O_3 and the following organic compounds were included: α -12 13 humulene, α -longipinene, and Δ -cardinene. Table S1 lists the reaction rate coefficients that were used for all 28 VOCs, together with references to the publications where the rate 14 15 coefficients can be found. For the BVOCs with very small contribution to the total BVOC concentration in the chamber (myrcene, sabinene, camphene, cineole, α -terpinene, Δ -16 17 terpinene, α -phellandrene, β -phellandrene, terpinolene, tricyclene, nonanal, bornyl acetate, methyl salicylate, α -longipinene and α -humulene), the initial reaction was only considered as 18 19 a sink of OH, O₃ and NO₃, without any other influence on the MCM gas-phase chemistry scheme. The oxidation of the sesquiterpenes farnesene, Δ -cardinene and "other SQTs" were 20 21 treated in a similar manner, except that we assumed that on average 1/2 OH molecule was formed for each sesquiterpene molecule that reacted with O₃. To only consider the first 22 23 oxidation reactions of some of the BVOCs influence the predicted radical and ozone 24 concentrations. However, because the concentrations of these BVOCs were relatively low this 25 effect was most likely small.

Ocimene constituted a major fraction of MT on the first experiment day. The chemistry path for ocimene is not available in the MCM. However, we included the ocimene chemistry by approximating its chemistry beyond the first oxidation step with that of limonene, for which the full MCM chemistry path is available. We also approximated the chemical path for Δ^3 carene after the first oxidation step with that of α -pinene, since both are bicyclic monoterpenes with an endocyclic double bond, as was done by Boy et al. (2013). For the "other MTs", we approximated the chemistry beyond the first oxidation step assuming that 50 1 % of the oxidation products end up as the corresponding MCM α -pinene first generation 2 oxidation products and 50 % end up as the corresponding MCM β -pinene first generation 3 oxidation products."

4

Based on the recent finding of rapid formation of extremely low-volatility organic compounds 5 6 (ELVOCs) from ozonolysis of monoterpenes containing endocyclic double bonds (Ehn et al., 2014) we also included a simplified ELVOC formation mechanism in the MCM gas-phase 7 chemistry code (R1), assuming that 7 mole % of the α -pinene + O₃ and Δ^3 -carene + O₃ 8 oxidation products were ELVOCs, with a molar mass of 325 g/mol and a vapour pressure of 9 10^{-10} Pa, which approximately corresponds to the VBS bin of $\log_{10}(C^*/\mu g m^{-3}) = -5$. We also 10 performed simulations where we considered that ELVOCs were formed from ozonolvsis of 11 12 any monoterpene or sesquiterpene, with the same ELVOC molar yield as for α -pinene. MCM_{ox,prod} represents the MCMv3.2 oxidation products that were formed from the same 13 14 reactions as the ELVOCs but via a different reaction pathway.

15
$$VOC+O_3 \rightarrow 0.07 ELVOC+0.93 MCM_{ox.prod}$$
 (R1)

16 Ehn et al. (2014) also observed ELVOC formed from OH-oxidation of α -pinene with an 17 estimated maximum molar yield of 1 %. In this work we evaluated the potential contribution 18 of ELVOCs formed from OH-oxidation by using an ELVOC molar yield of 1 % for any 19 monoterpene and sesquiterpene that reacts with OH (R2).

$$20 \quad \text{VOC+OH} \rightarrow 0.01 \text{ELVOC+} 0.99 \text{MCM}_{\text{ox,prod}} \tag{R2}$$

In order to evaluate the potential influence of specific ELVOCs (e.g. dimers) which may be involved in the nano-CN formation (here denoted ELVOC_{nucl}) we also included the possibility to add separate reactions where the monoterpene and sesquiterpene (VOC_{MT/SQT}) that react with O₃ or OH form trace amounts of ELVOC_{nucl} (R3-R4). In order to have little influence on the gas-phase chemistry or SOA formation, we used a very low ELVOC_{nucl} yield (γ_{ELVOC}) of 0.001 mole % and scaled the nucleation rate coefficient accordingly.

27
$$VOC_{MT/SQT} + O_3 \rightarrow \gamma_{ELVOC} ELVOC_{nucl} + (1 - \gamma_{ELVOC}) MCM_{ox,prod}$$
 (R3)

28
$$VOC_{MT/SQT} + OH \rightarrow \gamma_{ELVOC} ELVOC_{nucl} + (1 - \gamma_{ELVOC})MCM_{ox.prod}$$
 (R4)

The reactions R3 and R4 were only used in those simulations where $ELVOC_{nucl}$ were involved in the nano-CN formation (Eq. 3-6).

1 3.2 Aerosol dynamics

2 The aerosol dynamics module in ADCHAM is based on the aerosol dynamics code from the 2-D Lagrangian model for Aerosol Dynamics, gas-phase CHEMistry and radiative transfer 3 4 (ADCHEM) (Roldin et al., 2011). It includes subroutines for nano-CN formation, 5 condensation/evaporation, Brownian coagulation and particle deposition onto the chamber walls. ADCHAM simulates the condensation, dissolution and evaporation of sulphuric acid, 6 7 ammonia, nitric acid, hydrochloric acid and an unlimited number of organic compounds using 8 the analytic prediction of condensation scheme (Jacobson, 1997) and prediction of non-9 equilibrium growth scheme (Jacobson, 2005). The aerosol particle water content is calculated 10 with a thermodynamics model (Sect. 3.4).

11 ADCHAM considers the deposition of particles onto the chamber walls and keep track of the 12 amount of deposited material on the walls. In Roldin et al. (2014) we kept track of each 13 compound in each particle size bin that deposited on the chamber walls, and explicitly 14 simulated the mass-transfer-limited gas-particle partitioning between the gas-phase and the 15 wall deposited particles, assuming that the particles deposited on the walls remain as spherical particles on the walls. In this work, we instead assumed that the deposited SOA particles lose 16 17 their individual particle identity and merge into the VOC wall matrix together with the gasphase VOCs that deposit directly to the chamber walls (Sect. 3.5). 18

During the JPAC experiments the particles were formed by nucleation and, as a result of 19 20 coagulation, wall losses and dilution, their average lifetime in the chamber was relatively 21 short (less than 45 minutes, wherein dilution generally is the dominant loss process). Therefore, we expect that most of the formed particles were under charged with respect to the 22 Bolzmann charge equilibrium distribution so that we do not have to consider the enhanced 23 deposition rates induced by the particle charge (McMurry and Rader, 1985, Pierce et al., 2008 24 and Roldin et al., 2014). Thus, we calculated the first order wall deposition loss rate 25 26 assuming that all particles were non-changed using the indoor deposition loss rate model from 27 Lai and Nazaroff (2000), which accounts for different deposition loss rates on upward-, 28 downward- and vertically facing surfaces. In the model by Lai and Nazaroff (2000) the fundamental (but unknown) parameter for the particle loss rate is the friction velocity, u*. In 29 this work we found that with a relatively small value of u^* (0.02 m s⁻¹), the model best 30 31 captures the observed particle number and particle volume concentration loss rates after the UV-light is turned off. With this value of u* the deposition loss rate is equal to the observed 32

1 ELVOC molecule loss rate measured by Ehn et al. (2014) when the particle (molecule) 2 diameter is equal to 0.8 nm. We therefore used the value of $u^* = 0.02 \text{ m s}^{-1}$ for all the model 3 results presented in this work.

4 We evaluated six nano-CN formation parameterizations (Eq. 1-6) and compared them against 5 base-case simulations with a fixed nano-CN formation rate (J). The first mechanism (Eq. 1) is sulphuric acid activation nucleation (Kulmala et al., 2006), Eq. 2 is used for kinetic nucleation 6 of two H₂SO₄ molecules (McMurry and Friedlander, 1979), Eq. 3 is used for kinetic 7 nucleation of one H₂SO₄ molecule and one ELVOC_{nucl} Eq. 4 is the nano-CN parameterization 8 9 proposed by Riccobono et al., (2014) based on experiments in the CLOUD chamber, Eq. 5 10 represents a mechanism were single ELVOCs serve as nano-CN (Ehn et al., 2014), and Eq. 6 11 represent a kinetic type of nucleation mechanism with ELVOC_{nucl}.

$$12 \quad J = A_1[\mathrm{H}_2\mathrm{SO}_4] \tag{1}$$

13
$$J = K_2 [H_2 S O_4]^2$$
 (2)

$$14 \quad J = K_3[H_2SO_4][ELVOC_{nucl}] \tag{3}$$

15
$$J = K_4 [H_2 SO_4]^2 [ELVOC_{nucl}]$$
(4)

$$16 \quad J = A_5[\text{ELVOC}_{\text{nucl}}] \tag{5}$$

$$17 J = K_6 [\text{ELVOC}_{\text{nucl}}]^2 (6)$$

18 A (s⁻¹) and K (cm³ s⁻¹) in Eq. 1-3 and Eq. 5-6 are formation rate coefficients for activation type 19 and kinetic type of nucleation, respectively. For Eq. 4 the formation rate coefficient has the 20 unit cm⁶ s⁻¹.

The composition of the nucleation clusters was chosen in order to match the respective new particle formation mechanism. The dry nano-CN volume was composed of equal mole fractions of H_2SO_4 and $ELVOC_{nucl}$ when we used Eq. 1, 3 or 5, equal mole fractions of H_2SO_4 and NH_3 when we used Eq. 2, and pure $ELVOC_{nucl}$ when we used Eq. 6. The molar mass of $ELVOC_{nucl}$ was assumed to be 500 g/mol. The equilibrium water content of the nano-CN clusters was calculated with the thermodynamics model. The dry particle size of the nano-CN was assumed to be 1.5 nm.

Both $ELVOC_{nucl}$ and H_2SO_4 were assumed to be non-volatile. The $ELVOC_{nucl}$ first order wall loss rate was 0.011 s⁻¹ according to Ehn et al. (2014). The $ELVOC_{nucl}$ condensation sink was modelled explicitly with the condensation algorithm in ADCHAM. Chemical degradation of 1 ELVOC_{nucl} was not considered.

The ELVOC_{nucl} involved in the nano-CN formation were assumed to be formed 2 3 instantaneously after the first oxidation stage of monoterpenes and sesquiterpenes (R3-R4). With Eq. 3 and 4 we investigated six different sources of $ELVOC_{nucl}$; (i) as a product formed 4 exclusively from the ozonolysis of endocyclic monoterpenes (α -pinene and Δ^3 -carene), (ii) 5 from the ozonolysis of all monoterpene and sesquiterpene, (iii) from the ozonolysis of 6 7 sesquiterpenes, (iv) from the OH and O_3 oxidation of monoterpenes and sesquiterpenes, (v) 8 when monoterpenes and sesquiterpenes react with OH, or (vi) when sesquiterpenes react with 9 O_3 or OH. When using Eq. 5 and 6 we only considered the ELVOC_{nucl} that was formed from 10 monoterpenes and sesquiterpenes reacting with OH.

11 **3.3 Size distribution structures**

12 ADCHAM can be operated with the full-moving, fixed-sections or moving-centre particle size distribution approach (Roldin et al., 2011). In this work, we have tested both the fixed-13 14 sections approach and full-moving method using different numbers of size bins. The largest 15 advantage of the full-moving approach compared to the fixed-sections approach is that it does 16 not introduce any numerical diffusion problems during condensation/evaporation because it allows the particles to growth/shrink to their exact size. The main disadvantage is that new 17 18 particle size bins need to be introduced when new particles are formed. In this work, we 19 handled this by only introducing new particles formed by nucleation once every minute and at 20 the same time as we added a new size bin for the freshly nucleated particles, we also removed 21 the size bin containing the largest (oldest) particles. In order to not remove any particles 22 within a time frame substantially longer than their average residence time in the reaction 23 chamber, which was about 45 minutes, we used 400 size bins for the full-moving approach. 24 Thus, with this method we kept track of the nucleated particles for 400 minutes (almost 9) 25 times the average residence time in the chamber), before they were removed from the 26 modelled size distribution. This method was compared with the fixed-section approach using a different number of size bins in the diameter (D_p) range 1.5 to 800 nm. Based on these tests 27 28 we could conclude that the fixed-section approach was not suitable due to numerical diffusion 29 if the number of size bins were less than 400 (Fig. S2 in the supplementary material). With 30 the full-moving approach the particle volume (PV) and particle number concentrations (PN) were well conserved although the last size bin was removed once every minute (Fig. S3). 31 Therefore, we decided to use the full-moving method to represent the particle number size 32

1 distribution in this work.

One difficulty with the full-moving method is that the particle number size distribution needs to be mapped back onto a fixed diameter grid in order to illustrate it as a $dPN/dlogD_p$ distribution. This grid needs to be relatively coarse in order to not get a "jumpy" graphical representation of the particle number size distribution. For this we used 50 size bins between 1.5 and 360 nm in diameter.

7 **3.4** Particle phase chemistry and phase-state

8 ADCHAM includes a detailed particle-phase chemistry module, which is used to calculate the 9 particle equilibrium water content, the particle acidity, nitric acid and hydrochloric acid equilibrium vapour pressures for each particle size bin, and the non-ideal interactions between 10 11 organic compounds, water and inorganic ions using the activity coefficient model AIOMFAC (Zuend et al., 2008, 2011). In this work, we did not model the specific interactions between 12 the organic and inorganic compounds but assumed a complete phase-separation of the 13 inorganic- and organic particle phase. We used AIOMFAC to calculate the equilibrium water 14 15 content in both the inorganic and organic particle phase and the individual compound activity 16 coefficients. The organic compound activity coefficients were used when deriving the organic compounds equilibrium vapour pressures above each particle size (Sect. 3.6). 17

18 The particle phase chemistry module also contains subroutines that can be used to calculate 19 organic salt formation, oligomerization and heterogeneous oxidation (Roldin et al., 2014). 20 Recently, Shiraiwa et al. (2013) illustrated what peroxyhemiacetal (PHA) formation between 21 organic compounds containing aldehydes and hydroperoxide functional groups may proceed fast and contribute to a large and rapid increase of the formed SOA mass during 22 23 photooxidation experiments. In this work we evaluate if this type of heterogeneous dimer formation mechanism may explain the observed nano-particle growth during the JPAC 24 experiment. For this we tested to use a constant value of the PHA formation rate (k_{PHA}) of 12 25 M^{-1} s⁻¹ adopted from Shiraiwa et al. (2013). We also tested to model k_{PHA} as a 26 27 parameterization of the sulphate particle mole concentration $(x_{S(VI)})$, assuming that the PHA formation is acid catalysed by the co-condensing H_2SO_4 (Eq. 7). 28

29
$$k_{PHA} = B \cdot x_{S(VI)} (M^{-1}s^{-1})$$
 (7)

30 Here *B* is a constant, the value of which we varied in the range 1-500 M^{-1} s⁻¹.

31 In this work we used the kinetic-multi-layer module in ADCHAM for investigating weather

the phase-state of the SOA particles might have influenced the evolution of the particle 1 2 number size distribution. In order to do this, we divided each particle into three layers (an 3 approximately monolayer thick surface layer of 0.7 nm, and two bulk-layers). We considered the two extreme conditions where the SOA particles either were considered to be completely 4 5 liquid-like (no concentration difference between the surface and bulk layers) or solid-like (no 6 molecule transport between the surface and the particle bulk layers). Still, this had only a 7 minor effect on the modelled particle growth (SOA formation), but by treating the SOA as 8 solid-like improved the agreement between the modelled and measured SOA particle volume 9 decay when the UV-light was turned off (Fig. S4). Thus, in the simulations used to produce 10 the results presented in Sect. 4, we treated the SOA particles as solid-like with the assumption 11 that the molecule transport between the particle bulk and the particle surface-layer is relatively slow compared to the time scale it takes for the condensation to form a new 12 13 monolayer thick surface layer. However, it is important to mention that the effect of the SOA 14 phase-state cannot be unambiguously distinguished from the effect of the SOA volatility 15 when only looking at total particle volume loss rates.

16 **3.5 Reversible VOC wall loss**

17 The JPAC reaction chamber was mixed with a Teflon fan with mixing times <2 min. The first 18 order VOC wall loss rate to the chamber walls (k_w) was therefore governed by the molecular 19 diffusion across the boundary layer near the chamber walls and by the uptake rate at the wall 20 surface. According to McMurry and Grosjean (1985), the first order VOC wall loss rate for 21 FEP Teflon films ($k_{w,\text{FEP}}$) can be modelled with Eq. 8, which has two key parameters; the VOC wall mass accommodation coefficient (α_w) and the coefficient of eddy diffusion (k_e). 22 23 Unfortunately neither α_w or k_e can be derived easily. Based on the observed wall losses of particles, McMurry and Radar (1985) estimated k_e to be 0.12 s⁻¹ in a 60 m³ FEP Teflon film 24 chamber. Zhang et al. (2015) estimated k_e to be 0.075 s⁻¹ and 0.015 s⁻¹ in two not actively 25 mixed FEP Teflon film chambers with volumes of 24 and 28 m³. 26

27
$$k_{w,\text{FEP}} = \frac{A_w}{V_{chamber}} \left(\frac{\alpha_w \overline{\nu}/4}{1 + (\frac{\pi}{2})\alpha_w \overline{\nu}/(4\sqrt{(k_e D)})} \right)$$
(8)

- Here A_w is the chamber wall surface area, $V_{chamber}$ is the chamber volume, \bar{v} is the mean thermal speed of the gas molecules and *D* is the molecular diffusion coefficient.
- 30 In the JPAC reaction chamber Ehn et al. (2014) observed ELVOC first order wall loss rates in
- 31 the range 0.013-0.011 s⁻¹. By inserting a value of 0.011 s⁻¹ for $k_{w,\text{FEP}}$ in Eq. 8 and assuming

1 that the surface wall uptake rate is not limiting the ELVOC wall loss rate ($\alpha_w > 10^{-3}$) we get a 2 k_e of 4.2 s⁻¹ for the JPAC reaction chamber. This value is substantially larger than what was 3 estimated by Zhang et al. (2014) and McMurry and Radar (1985) and is probably because the 4 JPAC reaction chamber was actively mixed and has a smaller volume.

From measurements in a 4 m³ FEP Teflon chamber Kokkola et al. (2014) observed that for 5 nopinone which has a pure-liquid equilibrium saturation vapour pressure (p_0) of 53 Pa, the 6 gas-wall equilibrium was reached within a few minutes and $k_{w,\text{FEP}}$ was $\geq 0.03 \text{ s}^{-1}$. The 7 8 observations by Ehn et al. (2014) and Kokkola et al. (2014) indicate that gas-wall 9 equilibration can be reached rapidly for both volatile and low volatile VOCs and that their 10 uptake on the chamber walls is primarily limited by the transport to the chamber walls and not by surface uptake (α_w) . In contrast, Zhang et al. (2015) observed that the gas-phase 11 concentration of 25 different oxidized VOCs ($p_0 = 6x10^{-6} - 20$ Pa) slowly decreased over more 12 13 than 18 hours without reaching gas-wall equilibrium. These experiments were performed in a 24 m³ FEP Teflon film chamber that was not actively mixed. Based on Eq. 8, Zhang et al. 14 (2015) concluded that the VOC wall loss rate onto the FEB Teflon film walls was primarily 15 16 limited by the surface uptake (α_w) and not by the molecule diffusion to the chamber walls. 17 Zhang et al., (2015) also derived a parameterization of α_w as a function of the compounds pure liquid equilibrium saturation vapour pressure. 18

Based on the rapid ELVOC wall losses observed by Ehn et al. (2014), we assumed that the VOC losses to the JPAC reaction chamber walls was primarily governed by the molecule diffusion to the chamber walls and used a constant k_w equal to 0.011 s⁻¹ for all condensable organic compounds. However, we also performed test simulations using Eq. 8 and the α_w parameterization from Zhang et al. (2015).

According to Matsunaga and Ziemann (2010) the loss rate of VOCs from FEP Teflon chamber walls back to the gas-phase (k_g (s⁻¹)) can be represented by Eq. 9,

26
$$k_{g,i} = \frac{k_w}{(RT/p_{0,i}c_w/\gamma_{w,i})}$$
 (9)

where C_w in Eq. 9 is an effective mole concentration of organic compounds on the chamber walls (mol m⁻³) and $\gamma_{w,i}$ is the activity coefficient of compounds *i* in the organic film on the chamber walls, *T* is the temperature in Kelvin and *R* is the universal gas constant (J mol⁻¹ K⁻¹). According to Matsunaga and Ziemann (2010), the FEP Teflon film walls serve as a large 1 organic pool where VOCs can absorb $(C_w/\gamma_{w,i} \text{ on the order of } 100 \,\mu\text{mol m}^{-3})$. Thus, the actual 2 VOCs deposited on the chamber walls is considered to have a negligible contribution to C_w .

3 We assumed that there was practically no C_w on the glass for the VOCs to dissolve into. We also estimated that even if the deposited VOCs and SOA were distributed as a molecule 4 5 monolayer on the walls, the VOCs formed during the simulated experiment could only cover maximum 2 % of the total glass wall surface area in the end of the measurement campaign 6 7 (see supplementary material). Thus, we assumed that the VOC losses onto the glass wall 8 surfaces could be treated as a condensation process but without dissolution (absorption) into 9 an organic matrix on the walls. With this approach the uptake of compound *i* is governed by 10 the difference between the concentration in the gas-phase $(c_{g,i})$ and the pure liquid equilibrium saturation concentration $(c_{0,i})$ (Eq. 10-11). Thus, as long as $c_{g,i} < c_{0,i}$ the VOC will not 11 12 condense on the glass walls. For many of the semi-volatility organic compounds (SVOCs), the limit at which $c_{g,i}$ become larger than $c_{o,i}$ was never reached during the experiments. 13

14
$$\frac{dc_{g,i}}{dt} = -k_w (c_{g,i} - c_{0,i}) \qquad (if \ c_{g,i} > c_{0,i} \ or \ if \ c_{w,i} > 0)$$
(10)

15
$$\frac{dc_{w,i}}{dt} = k_w (c_{g,i} - c_{0,i})$$
 (11)

For the PTFE Teflon floor, we used the theory developed for FEP Teflon films (Eq. 9, 12 and 13), assuming that the PTFE Teflon surface (in total 15 % of the total chamber surface area) has a $C_w/\gamma_{w,i} = 100 \ \mu\text{mol m}^{-3}$ and $k_{w,i} = 1.7 \times 10^{-3} \text{ s}^{-1} (0.15 \times 0.011 \text{ s}^{-1})$.

19
$$\frac{ac_{g,i}}{dt} = -k_{w,i}c_{g,i} + k_{g,i}c_{w,i}$$
(12)

20
$$\frac{dc_{w,i}}{dt} = k_{w,i}c_{g,i} - k_{g,i}c_{w,i}$$
(13)

In order to mimic the actual experiment procedure (Sect. 2), we started the model simulations by running ADCHAM 14 days prior to the actual experiments started, with conditions similar to Day-1 of the experiment campaign (see Sect. 3.7). This allowed the VOC concentration to build up on the chamber walls.

25 **3.6 Condensable organic compound properties**

In this work, we used three different approaches to simulate the SOA formation. In the first approach, the SOA formation was modelled by considering the gas-particle partitioning of all non-radical organic compounds from the MCMv3.2 gas-phase chemistry code with estimated pure-liquid equilibrium saturation vapour pressure lower than 10^{-2} Pa at T=289 K (in total 488 1 compounds). The pure-liquid equilibrium saturation vapour pressures of these compounds 2 were estimated using the boiling point and vapour pressure extrapolation method from 3 Nannoolal et al. (2004; 2008), hereafter referred to as the Nannooal method. MCMv3.2 only 4 includes one sesquiterpene: β -caryophyllene. As an attempt to take into account the SOA 5 formation from the other sesquiterpens, exclusively in these simulations all the sesquiterpenes 6 were assumed to be emitted as β -caryophyllene. In this work we will refer to this SOA 7 formation representation as the MCM compound SOA formation mechanism.

8 The second method is the 2D-VBS approach (Jimenez et al., 2009; Donahue et al., 2011), 9 which is a simplified approach to describe how the volatility distribution of the VOCs (as a 10 function of their Oxygen to Carbon ratio (O:C)) evolves beyond the first oxidation step. A 11 detailed description of the structure and general assumptions of the 2D-VBS method used in 12 ADCHAM is given by Hermansson et al. (2014). Here we primarily describe how the 2D-13 VBS was modified in order to better capture the SOA formation observed during the JPAC 14 experiments.

15 The 2D-VBS compounds were distributed across 11 volatility classes separated by powers of 10 in equilibrium saturation concentration (at the standard temperature 298 K (C_{298}^*)), ranging 16 from 10⁻⁷ to 10³ µg m⁻³, and 12 discrete O:C from 0.1 to 1.2, in steps of 0.1. VOCs were 17 introduced into the 2D-VBS by scaling the amount of reacted monoterpenes and 18 19 sesquiterpenes with stoichiometric VBS yield parameterizations that represent the volatility 20 distribution of the first generation oxidation products. We used the parameterizations from 21 Table 3 in Henry et al. (2012). These parameterizations represent the first generation product 22 volatility distributions formed during experiments where α -pinene was primarily oxidized by ozone (no OH-scavenger) or where α -pinene was oxidized by OH. We used the former 23 parameterization to represent the first generation terpene (monoterpenes and sesquiterpenes) 24 25 oxidation products during the UV-off periods and the later for the UV-light on periods. ELVOCs formed as first generation oxidation products from R1 were introduced into the 2D-26 VBS assuming that they had a O:C of 1 and $C_{298}^* = 10^{-7} \,\mu\text{g m}^{-3}$. 27

Hermansson et al. (2014) illustrated that the modelled SOA formation is sensitive to the volatility distribution and the assumed O:C of the first generation oxidation products. Moreover it depends on how the functionalization and fragmentation pattern is represented in the VBS and the reaction rate constant between the VOCs and OH. Equation 14 from Jimenez et al. (2009) was used to represent how the fraction of VOCs which fragmentize upon 1 oxidation varies as a function of their O:C.

$$2 \qquad f_{\text{frag}} = (\text{O:C})^{\text{y}} \tag{14}$$

Jimenez et al. (2009) assumed that y was equal to 1/6, that the 2D-VBS compounds were 3 oxidized by OH with a reaction rate coefficient (k_{OH}) of 3×10^{-11} cm³ s⁻¹, and that the 4 compounds that fragmentize had equal probabilities to split at any of the carbon bonds. 5 6 However, with this 2D-VBS setup, ADCHAM substantially overestimated the SOA formation 7 (Fig. S5). This was mainly because a substantial fraction of the 2D-VBS fragmentation 8 products were still less volatile than the reacting VOCs. If we instead assumed that all 9 fragmenting compounds (on a mole fraction basis with equal proportions) end up into the 10 VBS bins where C_{298}^* is at least 3 order of magnitude larger than the corresponding functionalization products, and decreased the probability of fragmentation somewhat (y=1/3), 11 12 the modelled particle volume concentration agreed better with the measurements (Fig. S6). Thus, in the proceeding sections we used this fragmentation parameterization. We also 13 performed sensitivity tests where we varied the value of k_{OH} in the range of 3×10^{-11} to 5×10^{-11} 14 $cm^3 s^{-1}$ and the O:C of the first generation oxidation products in the range of 0.3 to 0.5 (Fig. 15 S6). Based on these tests, we decided to use the values of $k_{OH} = 5 \times 10^{-11} \text{ cm}^3 \text{ s}^{-1}$ and O:C = 0.4 16 for the first generation oxidation products in the simulations used to produce the results in 17 18 Sect. 4.

19 The temperature dependence of the 2D-VBS equilibrium saturation concentrations were 20 derived using the Clausius Clapeyron equation and assuming the following form for the heat 21 of vaporization (ΔH ; Epstein et al., 2010):

22
$$\Delta H = -11 \times \log_{10} C_{300}^* + 129 \text{ kJ mol}^{-1}$$
 (15)

As the third approach we also tested to represent the SOA formation with a one-product 23 model. With this method, only a single non-volatile condensable organic compound was used 24 25 to represent the SOA formation. This compound was formed as a first generation oxidation product when any of the monoterpenes or sesquiterpenes reacted with O₃ or OH, assuming a 26 molar yield of 25 %. The non-volatile condensable organic compound was assumed to have a 27 28 molar mass of 325 g/mol. This simulation represents the limiting condition where the SOA 29 formation is purely kinetically limited and not absorption partition limited, and it was also 30 used for estimating how large the ELVOC yield would need to be if the particles were only growing by condensation of ELVOCs. 31

Both with the 2D-VBS and when the MCM compounds were used to represent the condensable organic compounds, the Raoult's law and the Kelvin equation were used to calculate the equilibrium vapour pressure $(p_{s,i,j})$ for each compound (i) above each particle size (j) (Eq. 16). With the 2D-VBS we assumed that the organic compounds were mixed ideally (unity activity coefficients (γ_i)). However, for the MCM compounds the molecule structure is known, so we calculated the organic compound activity coefficients in the SOA particle mixtures using AIOMFAC.

8
$$p_{s,i,j} = p_{0,i} x_{i,j} \gamma_{i,j} e^{\left(\frac{4M_i \sigma_i}{RT \rho_j D_{p,j}}\right)}$$
(16)

9 where $x_{i,j}$ is the mole fraction of compound *i* in the particle surface layer of particles in size 10 bin *j*, M_i is the molar mass of compound *i* and ρ is the density of the absorbing phase. In this 11 work the surface tension of the organic compounds (σ_i) was estimated to be equal to 0.05 N 12 m⁻¹ following Riipinen et al. (2010), even though we also tested the values of σ_i in the range 13 0.02 to 0.07 N m⁻¹. These sensitivity tests revealed that the modelled total particle number 14 concentration decreased by approximately 10 % and the total particle volume concentration 15 by approximately 30 % when the value of σ_i was increased from 0.02 to 0.07 N m⁻¹ (Fig. S7).

16 **3.7** Optimizing model input parameters

We used the ADCHAM model to simulate the nano-CN formation and growth during one measurement campaign conducted in the JPAC chamber. We chose to study the first four continuous days from the measurement campaign because these days had continuous measurements, coverage of the dataset was complete and UV light-on period was the same.

The measured temperature and relative humidity (RH) in the JPAC reaction chamber were 21 22 used directly in the model as input. The values from the first four days are illustrated in Fig. 2. 23 The purple bars in the figures illustrate the UV-on periods. The temperature stabilized around 24 noon on Day-1, after that the temperature was ca. 16 °C for all days. The RH had minor 25 fluctuations during the experiment. RH was kept to ca. 60% during UV-on periods for the 26 first three days. On Day-4, the RH was around 55 %. In addition, 8 discharge lamps were 27 used to simulate solar illumination on Day-1 in the plant chamber, but only 4 lamps were used 28 on the remaining days to generate lower levels of VOC emissions from the trees.

Because the inflow of ambient air into the JPAC chamber was purified by an adsorption dryer, by default, the concentrations of NO, NO₂, SO₂, and CO in the inflow to the reaction

1 chamber were assumed to be 0.03, 0.2, 0.015 and 15 ppby, respectively. However, in order to 2 evaluate how sensitive the model results was, in particular the modelled VOC composition 3 and SOA formation, we also performed sensitivity tests where we varied the inflow 4 concentrations of NO_x (NO+NO₂) within the range of 0.05-1 ppby. Concentrations of VOCs 5 in the inflow to the reaction chamber were based on the GC-MS and PTR-MS data. The initial 6 concentrations of all VOCs were set to zero. The measured concentrations of isoprene, 7 monoterpenes and sesquiterpenes (in the plant chamber) are illustrated in Fig. 3. The 8 monoterpene and sesquiterpene concentrations displayed a certain diurnal distribution pattern, 9 because the discharge lamps mimicked sunlight in the plant chamber. The total monoterpene 10 concentration exceeded 4 ppbv on Day-1, and then decreased to a value below 2 ppbv on the 11 following three days. In particular, the ocimene concentration was the highest on Day-1, 12 which can be explained by the fact that ocimene emission is light dependent (Owen et al., 13 2002), and the simulated solar light intensity in the plant chamber from the discharge lamps 14 was two times higher on Day-1 than during Day-2 – Day-4.

The isoprene concentration and the total sesquiterpene concentration were about an order of magnitude lower than the monoterpene concentration. Overall, the measured concentrations of terpenes were at similar levels as ambient air concentrations in boreal forest environments (Ruuskanen et al., 2009).

19 At the first attempt, we tried to simulate the O_3 and OH concentrations with the gas-phase 20 chemistry mechanism. An OH sink has been determined in the empty chamber by direct OH-21 LIF measurements (Diss. Sebastian Broch) and has been always found to be within 2 - 4 s⁻¹. 22 However, even with a OH sink of 4 s⁻¹, the model could not capture the observed gradual 23 decrease in the O_3 concentration and increase in OH concentration during the UV-light on 24 periods without either varying the UV-light intensity or the OH sink (Fig. 4).

25 Therefore, we decided to use the measured O₃ and H₂SO₄ concentrations, and estimated OH 26 concentration, as input to the model for the simulations presented in Sect. 4. Figure 4a shows the measured O₃ concentration during the experiment, Fig. 4b shows the estimated OH 27 concentration based on the observed loss rate of 2-butanol and Fig. 4c shows the measured 28 29 H₂SO₄ concentration. When the UV-light was turned on the H₂SO₄ concentration increased 30 very rapidly and had a distinct peak. This is most likely due to an initial peak in the OH concentration at the moment when the UV-lamps was turned on (indicated by the modelled 31 OH concentration). In the model we represented this by setting the OH concentration to 5×10^7 32

1 molecules cm⁻³ for the 12 first minutes after the UV-light were turned on.

2 The photolysis rates were simulated by using the quantum yields and absorption cross 3 sections reported at the MCMv3.2 web site. In order to evaluate the direct influence of the 254 nm wavelength UV-light on the VOC composition and SOA formation, we performed 4 5 model simulations where the 254 nm UV-light was not considered when calculating the photolysis rates (J_x) for all compounds except O₃. These test showed that the modelled 6 7 condensable organic compound composition, O₃ concentration (Fig. 4a), OH concentration 8 (Fig. 4b) and the SOA formation were not significantly influenced directly by the 254 nm 9 UV-light source, but strongly by the OH generated from the photolysis of O_3 . Apart from O_3 , 10 the compound that was affected most by the 254 nm wavelength UV-light was H_2O_2 $(J_{\rm H_2O_2}=1.26\times10^{-4} \text{ s}^{-1} \text{ with the UV-light source and } J_{\rm H_2O_2}=3.72\times10^{-8} \text{ s}^{-1} \text{ without the UV-light}$ 11 source). Some of the VOCs containing carbonyl groups and nitrate groups were also 12 influenced by the UV-light source. The most prominent change in the modelled photolysis 13 rate in these two groups of VOCs occurred for glyoxal (a dialdehyde) $(J_{alyoxal}=2.93\times10^{-5} \text{ s}^{-1})$ 14 with the UV-light source and $J_{glyoxal}$ =5.04×10⁻⁷ s⁻¹ without the UV-light source) and methyl 15 nitrate (CH₃NO₃) ($J_{CH_3NO_3}$ =5.89×10⁻⁵ s⁻¹ with the UV-light source and $J_{CH_3NO_3}$ =2.51×10⁻⁸ s⁻¹ 16 17 without the UV-light source). These values can be compared with the rates at which glyoxal and CH₃NO₃ were oxidized by OH. At the OH concentration of 5×10^7 molecules cm⁻³ (typical 18 for the UV-light on periods) this rate was 5.00 $\times 10^{-4}$ s⁻¹ for glyoxal and 1.08 $\times 10^{-6}$ s⁻¹ for 19 CH₃NO₃. Thus, for glyoxal the reaction with OH was still ~10 times faster than the photolysis 20 21 reaction rate, while for CH₃NO₃ the photolysis reaction rate was ~2 times larger than the rate 22 at which CH₃NO₃ was oxidized by OH. This indicates that short wavelength UV-light sources 23 (as the one used in JPAC) may influence the VOC composition, especially if a large fraction of the VOCs contain nitrate functional groups (i.e. at high NO_x concentrations). 24

25

26 4 Results and discussion

27

Table 2 summarizes the different model tests that we performed in this work in order to constrain the VOC wall losses, the aerosol dynamics-, gas- and particle-phase chemistry mechanisms that can explain the nano-CN formation and growth observed during the JPAC experiments presented in Sect. 2.

1 4.1 Time series of BVOC concentrations

2 To investigate the potential contribution of BVOCs to the nano-CN formation and growth, it 3 is essential to properly predict the time series of the BVOC concentrations. The modelled 4 isoprene, monoterpene, and sesquiterpene concentrations in the JPAC reaction chamber are 5 plotted in Fig. 5, together with isoprene and total monoterpene concentrations measured by the PTR-MS. The modelled isoprene concentrations were in a good agreement with the 6 measurements during the UV-off period on Day-1. During the rest of the simulated four days 7 8 period, the simulated isoprene concentration was generally lower than the observations. 9 However, considering the low isoprene concentrations and the uncertainties in the PTR-MS 10 measurements, we cannot draw any conclusions on weather the model actually underestimates 11 the isoprene concentration or not.

12 The summation of the modelled monoterpene isomer concentrations reached the same level as 13 the measurements. During the dark periods, the monoterpene concentrations in the reaction 14 chamber decreased to about one third of their concentrations measured in the plant chamber. 15 This was because of the in-flow dilution and chemical reactions with O₃. When the UV light 16 was switched on, the monoterpene concentration decreased sharply due to OH oxidation, as 17 seen both from the measurements and simulation results in Fig. 5. Before the UV-lights were turned on ~ 80 % of the reacting monoterpenes were oxidized by O₃. At the UV onset this 18 number dropped to ~ 10 % and in the end of the UV-on periods only ~ 2 % of the oxidized 19 20 monoterpenes were oxidized by O₃. The modelled monoterpene concentration shows a somewhat more pronounced decrease during the UV-on periods than measurement. However, 21 as with the isoprene concentration, the relative uncertainties in the PTR-MS measurements 22 23 increases at lower concentrations and part of the PTR-MS unity resolution mass peak that was 24 interpreted as monoterpenes may also have had small contributions from other VOCs.

The sesquiterpene concentrations were below the detection limit of the PTR-MS. Our modelled results show that the sesquiterpenes were strongly oxidized by O_3 once the tree emission from the plant chamber entered the reaction chamber (Fig. 3). Even when the UVlights were turned on, 40 to 60 % of the sesquiterpenes were oxidized by O_3 and the rest by OH. The sesquiterpene concentrations decreased rapidly to below 0.01 ppbv. In addition, the sesquiterpene concentrations showed a minor decrease during UV-on periods because of the oxidation by OH.

32 Although a large fraction of the monoterpenes and sesquiterpenes were oxidized in the JPAC

reaction chamber, nano-CN formation was only observed during the UV-on periods. This indicates that the terpene oxidation products formed during the dark periods (primarily by ozonolysis) either did not have the right properties or were not abundant enough to form, activate and/or grow nano-CN to sizes above the detection limit of the PSM-CPC setup.

5 **4.2** Simulations of the observed new particle formation events

6 The first simulations were designed in order to constrain the VOC wall losses and the mechanisms responsible for the observed particle growth (Sect. 4.2.1 and Sect. 4.2.2). For 7 these simulations, we used a fixed nano-CN formation ratio of 20 cm⁻³ s⁻¹ during the UV-light 8 on periods, except for the first 12 minutes with UV-light on for which we used $J = 80 \text{ cm}^{-3} \text{ s}^{-1}$, 9 for Day-1 and Day-2 and 60 cm⁻³ s⁻¹ for Day-3. During the UV-light off periods the nano-CN 10 formation rate was zero. The nano-CN had a dry diameter of 1.5 nm and was assumed to be 11 12 composed of equal number of sulphuric acid and ELVOC_{nucl} molecules. After, constraining the VOC wall losses and the potential particle growth mechanisms, we investigated several 13 14 nano-CN formation mechanisms (Eq. 1-6) and compared the results against the results from 15 the simulations with a fixed nano-CN formation rate (Sect. 4.2.3).

16 **4.2.1 Modelling the reversible VOC wall deposition**

17 Figure 6a shows the modelled total particle volume concentration (PV) starting 14 days prior to the start of the intensive experimental campaign. The measured PV during the campaign is 18 19 also displayed. The model results are from simulations with the 2D-VBS. As was described in Sect. 3.5, for the base case model simulation we used a constant species independent k_w equal 20 to 0.011 s⁻¹ and Eq. 9-13 to differentiate between the reversible VOC wall losses to the glass 21 22 and PTFE Teflon surfaces. The results in Fig. 6b and c which are from the base case 23 simulation illustrate how the 2D-VBS compounds with different C^* were distributed onto the glass walls and the PTFE Teflon floor, respectively. The LVOCs ($C_{298}^* < 10^{-1} \,\mu g \, m^{-3}$) deposited 24 both onto the glass and Teflon surfaces, while the SVOCs ($C_{298}^* > 10^{-1} \ \mu g \ m^{-3}$) were 25 26 exclusively found at the Teflon floor where they could absorb into C_w . Some of the compounds with $C_{298}^* = 10^{-1} \,\mu g \, m^{-3}$ deposited onto the glass walls shortly after the UV-light 27 28 was turned on, but when the condensation sink was increased and when the terpene inflow 29 concentration into the reaction chamber was lowered (Day-2 - Day-4), the gas-phase 30 concentrations of these compounds never reached above their pure liquid equilibrium 31 saturation concentrations, so they evaporated from the glass walls. Figure S8 in the 32 supplementary material shows similar results but for a simulation with the MCM compound 1 SOA formation mechanism and additional ELVOCs (R1) contribution to the particle growth.

Figure 6a also shows the modelled particle volume concentration when assuming that the 2 3 JPAC reaction chamber walls behave similarly to FEP Teflon walls (Eq. 8-9), in which case k_e was 4.2 s⁻¹ and the a_w parameterization was taken from Zhang et al. (2015). With this 4 parameterization, k_{wi} varies from 2×10^{-5} to 7×10^{-4} s⁻¹ for compounds with a molar mass of 5 300 g mol⁻¹ and vapour pressures in the range 10⁻² to 10⁻¹⁰ Pa. Thus, in these model 6 simulations the ELVOC wall losses were about 15 times lower than what was observed by 7 8 Ehn et al. (2014). Because of the lower VOC wall losses, the model overestimated the SOA 9 formation by a factor of 3-4 for Day-1, a factor of 1.5-2 for Day-2 and Day-3 and a factor of 10 2-3 for Day-4, there was no gradual increase in the SOA formation due to re-evaporation of 11 SVOCs from the walls and the correlation between the model and measured PV is 12 substantially worse. This illustrates that the wall losses in the JPAC chamber cannot be treated in the same way as in FEP smog chambers. 13

Figure 7 shows the 2D-VBS VOC composition for: (a) the SOA particles, (b) the gas-phase, 14 15 (c) the VOCs on the glass walls and (d) the VOCs on the PTEF Teflon floor, at Day-3 and after 5 hours with UV-lights on. The bar at O:C = 1, $\log_{10} C_{298}^*$ =-7 corresponds to the 16 17 ELVOCs which were assumed to be formed from ozonolysis and OH-oxidation of 18 monoterpenes and sesquiterpenes (molar yield of 7 % and 1 %, respectively). After 16.5 days of continuous experiments, 1.4 mg m⁻³ of SVOCs had deposited onto the PTEF Teflon floor 19 according to the model simulations. The largest fraction of the wall deposited VOCs were 20 21 first generation terpene oxidation products (O:C=0.4). A fraction of these SVOCs can re-22 evaporated and reacted with OH in the gas-phase. This explains why the modelled SOA formation was gradually increasing during the first 15 days (Fig. 6a), although all the other 23 24 model conditions were identical. A similar pattern in the modelled SOA particle volume was 25 also observed when the SOA formation was simulated with the MCM compound SOA 26 formation mechanism (Sect. 3.6) (Fig. S8a). The ELVOCs (that are formed as first generation oxidation products (R1)) and the LVOCs (formed from OH-oxidation of the first generation 27 28 oxidation products), primarily deposited onto the glass walls. However, in total the VOCs 29 deposited onto the glass walls only made up 5 % of the total amount of VOCs on the chamber 30 walls.

4.2.2 Evaluation of potential particle growth mechanisms

32 When using the 2D-VBS, the modelled SOA composition was dominated by LVOCs and

1 SVOCs formed from second- and multi-generation OH-oxidation products. This was the case 2 even if we considered that ELVOCs were formed as first generation products after the O₃-3 and OH-oxidation of all monoterpenes and sesquiterpenes (Fig. 7a). The reason for this is the 4 large ELVOC wall losses in the JPAC chamber and the small condensation sink during the 5 new-particle formation events. During the UV-light on periods, the gas-phase 2D-VBS VOC 6 composition was dominated by oxidized SVOCs formed by fragmentation of the first 7 generation oxidation products (Fig. 7b).

In Fig. 8 we compare the modelled (a) total particle number concentration and (b) total 8 9 particle volume concentration with the observations from the PSM-CPC and the SMPS. The 10 model results are from simulations with the 2D-VBS with or without ELVOC formation from 11 the ozonolysis and OH-oxidation of all monoterpenes and sesquiterpenes (R1-R2) (molar yield of 7 % and 1 %, respectively), as well as from a simulation with the MCM compound 12 13 SOA formation mechanism (Sect. 3.6) (including ELVOCs from the ozonolysis of α -pinene and Δ^3 -carene (R1)). For this simulation we also considered rapid peroxyhemiacetal dimer 14 formation in the particle phase using Eq. 7 and $B = 200 \text{ M}^{-1} \text{ s}^{-1}$. Without consideration of this 15 type of a rapid acid catalysed dimer formation process, the particle growth was substantially 16 17 underestimated and almost no SOA was formed when we simulated the SOA formation with the MCM compound SOA formation mechanism (Fig. S8a). Shown in Fig. 8 are also the 18 19 results from a simulation with a one-product model approach (Sect. 3.6). In the model, this non-volatile compound was formed as a first generation oxidation product from O₃ and OH 20 21 oxidation of all monoterpenes and sesquiterpenes with a molar yield of 25 % (mass yield of 22 almost 60 %).

23 In all the model simulations, the model overestimated the SOA particle volume formation during Day-1 but tended to underestimate, or to give values similar to the measurements, 24 during the period Day-2 - Day-4. The best agreement between the model and measured 25 particle volume concentration was found with the 2D-VBS method ($R^2 = 0.699$ with ELVOCs 26 formation and $R^2 = 0.697$ without ELVOC formation), even though the particle volume 27 28 concentration could nearly equally well be represented with the MCM compound SOA formation mechanism and acid catalysed PHA dimer formation ($R^2 = 0.672$). In the 29 30 simulation with the non-volatile one product model, the agreement between the modelled and measured particle volume concentration was worse ($R^2 = 0.634$), which indicates that the 31 particle growth cannot be purely explained by condensation of ELVOCs. This model 32

simulation also illustrates that even if the particle growth was only be kinetically limited (not 1 2 absorption partitioning limited) and if the yield of the formation of condensable organic compounds was be the same in all the days, the decrease in SOA volume concentration from 3 Day-1 to Day-2 should have been more pronounced because of the substantially smaller 4 5 condensation sink during Day-2 - Day-4 compared to Day-1. In the other model simulations, 6 this effect was partly but not fully compensated by the re-condensation of SVOCs from the walls, which contributed relatively more to the SOA formation during Day-2 - Day-4 7 8 compared to Day-1. The only remaining explanation we can find to why the model gave too 9 much SOA particle volume during Day-1 compared to Day-2 - Day-4 is that the BVOC 10 composition was substantially different during Day-1. It may be that ocimene which reacts 11 rapidly with O_3 may not form SOA to the same extent as e.g. α -pinene.

Figure S9 in the supplementary material compares the modelled total particle volume concentrations from simulations with B = 10 or 200, or using a constant PHA dimer formation rate of 12 M⁻¹ s⁻¹ based on the work by Shiraiwa et al. (2013). With B = 10 the modelled PV are in very good agreement with the measured PV for Day-1 but for the following days the model substantially underestimated the PV. With a constant PHA dimer formation rate of 12 M^{-1} s⁻¹ the model gives 3 times higher PV than the measurements for Day-1 but gives reasonable PV formation for Day-2 – Day-4.

Figure 9 shows the modelled SOA volatility distribution as a function of particle size. The results are from a simulation with the 2D-VBS, including ELVOC formation from all monoterpenes and sesquiterpens oxidized by O_3 and OH (R1-R2). As expected, the smallest particles contained the largest mole fraction of ELVOCs because of the Kelvin effect and because the uptake of the SVOCs is not kinetically limited but limited by absorption into the particle surface layer.

25 Figure 10 shows the modelled and measured particle number size distributions at $\frac{1}{2}$, 1, 2 and 5 hours of UV-lights on, for each day of the experimental campaign. The model results are 26 from a simulation with the 2D-VBS including ELVOC formation from all the monoterpenes 27 and sesquiterpenes oxidized by O₃ and OH (R1-R2). For Day-1, the model overestimated the 28 29 particle growth rate, which can also be seen from the overestimated SOA formation (Fig. 8b). 30 For the reminder of the experimental campaign, the modelled particle number size distributions were in good agreement with the SMPS measurements, except for the particles 31 having a diameter <30 nm, for which the model gave substantially higher concentrations. A 32

contributing explanation for this feature can be non-accounted diffusion losses of particles in
 the SMPS inlet.

In the supplementary we show how the volatility distribution of the MCM compounds and the SOA formation changed when the NO_x concentration in the inflow to the chamber was varied in the range 0.05 to 1 ppbv (Fig. S10). When the NO_x concentration was increased from 0.05 to 1 ppbv the particle SOA volume concentration was increased slightly (~10 %).

7 4.2.3 Evaluation of potential nano-CN formation mechanisms

8 In this section we evaluate the different nano-CN formation mechanisms described in Sect. 9 3.2. For these simulations we used the 2D-VBS to simulate the evolution of the condensable organic compounds. For all simulations in this section, ELVOCs were formed from all the 10 11 monoterpenes and sesquiterpenes oxidized by O_3 and OH (R1-R2). Because the exact vapour pressures, formation mechanisms and concentrations of ELVOCs still are very uncertain, we 12 cannot dismiss the possibility that the new particle formation (formation of particles with 13 $D_p > 1.6$ nm) was limited by the activation of nano-CN and not by the nano-CN formation rate 14 In this section we evaluate possible nano-CN formation mechanism with the 15 itself. 16 assumption that it was the nano-CN formation that primarily limited the observed new particle formation during the experiments. 17

Table 3 gives the coefficient of determination (R^2) between the modelled total particle 18 19 number concentration and measured total particle number concentration ($D_p > 1.6$ nm) from the PSM-CPC setup. The R^2 values are only given for those nano-CN mechanisms 20 (parameterizations) that gave a $R^2 > 0.75$. The R^2 values were calculated with the data from 21 22 hour 11 to hour 96, for which the measured total particle number concentration from the PSM-CPC setup was complete. The largest R^2 values (>0.95) between the modelled and 23 24 measured total particle number concentration were achieved with nano-CN formation mechanisms that involves both H₂SO₄ and ELVOCs. With Eq. 3, this were the case both if the 25 ELVOC_{nucl} molecules were formed from OH-oxidation or OH and O₃ oxidation of 26 27 monoterpenes and sesquiterpenes. With Eq. 4, this was only the case if the ELVOC_{nucl} molecules exclusively were formed from ozonolysis of sesquiterpenes. These are the only 28 mechanisms for which the values of R^2 were higher than in simulations with a fixed nano-CN 29 formation rate during the UV-light on periods. The relatively high R² value for the simulation 30 31 with a fixed nano-CN formation rate during the UV-light on period is an indication that the 32 variability in the dataset with respect to the compound(s) responsible for the nano-CN

formation was too small to fully constrain the dominating nano-CN formation mechanism during the experiments. However, the dataset is still useful to reject the less likely nano-CN mechanisms and to narrow down the list of possible mechanisms. For this, the R² value from the simulation with the fixed nano-CN formation rate was used as a benchmark.

5 In Fig. 11 we compare the modelled total particle number concentration when calculating Jusing Eq. 1, 2, 3 and 4. With Eq. 3, the ELVOC_{nucl} were either assumed to be formed from 6 ozonolysis and OH-oxidation of monoterpenes and sesquiterpens, or only from OH-oxidation. 7 8 With Eq. 4, the ELVOC_{nucl} were formed from sesquiterpenes oxidized by OH. Displayed is also the result from the simulation with constant $J = 20 \text{ cm}^{-3} \text{ s}^{-1}$ during the UV-light on 9 periods and J = 0 cm⁻³ s⁻¹ during the UV-light of periods. Based on Fig. 11 and the R² values, 10 11 it is evident that the new particle formation could not be captured very well with H₂SO₄ activation (R1) or H₂SO₄ kinetic (R2) type of nano-CN formation. The simulation with R1 12 had a relatively high R^2 value but substantially overestimated the total particle number 13 concentration during the UV-light off periods. The same is true for the kinetic type of nano-14 15 CN formation involving H₂SO₄ and ELVOC_{nucl} formed from oxidation of the BVOCs with O₃ 16 and OH. In this case, the model especially overestimated the total particle number 17 concentration during the first day of the UV-light off period when the BVOC emissions were the highest. The best agreement between the modelled and measured total particle number 18 concentration, both based on the R^2 value and Fig. 11, was achieved with the kinetic type of 19 nucleation involving H₂SO₄ and ELVOC_{nucl} (Eq. 3) formed from oxidation of monoterpenes 20 and sesquiterpenes with OH exclusively. 21

22

23 **5** Summary and conclusion

24

In this study, we used the ADCHAM model to simulate the nano-CN formation and growth during an experimental campaign in the Jülich Plant Atmosphere Chamber, which focused on new particle formation induced by photochemical reactions of VOCs emitted from real plants (Dal Maso et al., 2014). With the model we evaluated potential nano-CN formation and growth mechanisms and how the VOC chamber wall losses influenced the SOA formation and composition.

31 Our analysis on the JPAC experiments illustrates how complex it is to design smog chamber

1 experiments that mimic the conditions during new particle formation events in the 2 atmosphere, in addition to which it provides useful information about what need to be 3 considered when designing such experiments. In this work we showed that the relative 4 contribution of ELVOCs to the nano-CN formation and growth was effectively suppressed 5 due to their rapid and irreversible wall losses and the relatively high OH concentrations 6 during the UV-light on periods. With these respects, the conditions during the JPAC 7 experiments were not directly comparable with typical conditions during new particle 8 formation events in the atmosphere. In addition, the experimental conditions showed too 9 small variability to fully constrain the dominating nano-CN formation and growth 10 mechanisms. Despite these limitations, the JPAC experiments serve as a valuable dataset to 11 narrow down the list of potentially important nano-CN formation and growth mechanism over 12 the boreal forest region.

13 In the lack of ELVOCs in the gas-phase, one possible mechanism that can explain the initial particle growth is rapid heterogeneous dimer formation of SVOCs. In this work, we found out 14 15 that acid catalysed peroxyhemiacetal formation between aldehydes and hydroperoxides may explain the observed particle growth. However, the particle growth could be modelled equally 16 17 well if the SOA was formed by condensation of low-volatility second- and multi-generation OH gas-phase oxidation products simulated with a 2D-VBS approach. Thus, based on our 18 19 model simulations, we cannot conclude weather the observed particle growth primarily was 20 driven by low-volatility organic compounds formed in the gas-phase or by rapid dimer 21 formation in the particle phase. However, without the later mechanism the model fails to 22 reproduce the observed SOA formation when using the MCMv3.2 oxidation products as 23 condensable organic compounds.

Our results suggest that H_2SO_4 is one of the key compounds involved in the new particle formation, but cannot solely explain the new particle formation process. During the simulated experiments, the best agreement between the modelled and measured total particle number concentration was achieved when using a nano-CN formation rate of the form $J = K[H_2SO_4][ELVOC_{nucl}]$, where ELVOC_{nucl} was formed as an BVOC OH oxidation product.

29

30 Acknowledgements

P. Roldin would like to thank the Cryosphere-Atmosphere Interactions in a Changing Arctic
Climate (CRAICC) and the Swedish Research Council for Environment, Agricultural

1 Sciences and Spatial Planning FORMAS (Project No. 214-2014-1445) for financial support. 2 L. Liao wishes to thank the Maj and Tor Nessling foundation for financial support (grant No 3 2009362), as well as the Academy of Finland (project No. 128731). D. Mogensen would like 4 to thank the doctoral program in Atmospheric Sciences (ATM-DP) at the University of 5 Helsinki for financial support. We would like to acknowledge HENVI (Helsinki University 6 Centre for Environment), The FCoE (The Centre of Excellence in Atmospheric Science -7 From Molecular and Biological processes to the Global Climate (ATM)), the strategic 8 research area MERGE (Modelling the Regional and Global Earth system) and the PEGASOS 9 (Pan-European Gas-Aerosolsclimate interaction Study, project No FP7-ENV-2010-265148) 10 project. We would also like to thank Prof. Gordon McFiggans' research group at the 11 University of Manchester, and especially Dr. David Topping, for helpful discussions and for providing the Python script (now a publicly available function called Comp-SysProp: 12 http://ratty.cas.manchester.ac.uk/informatics/) to calculate Nannoolal-based sub-cooled liquid 13 14 equilibrium saturation vapour pressures for all organic compounds included in this paper.

15

16 **References**

17 Almeida, J., Schobesberger, S., Kürten, A., Ortega, I. K., Kupiainen-Määttä, O., Praplan, A.

P., Adamov, A., Amorim, A., Bianchi, F., and Breitenlechner, M.: Molecular understanding
of sulphuric acid-amine particle nucleation in the atmosphere, Nature, 502, 359–363, 2013.

Berndt, T., Stratmann, F., Sipilä, M., Vanhanen, J., Petäjä, T., Mikkilä, J., Grüner, A.,
Spindler, G., Lee Mauldin III, R., Curtius, J., Kulmala, M., and Heintzenberg, J.: Laboratory
study on new particle formation from the reaction OH+SO2: influence of experimental
conditions, H2O vapour, NH3 and the amine tert-butylamine on the overall process, Atmos.
Chem. Phys., 10, 7101–7116, doi:10.5194/acp-10-7101-2010, 2010.

Bonn, B. and Moortgat, G. K.: Sesquiterpene ozonolysis: origin of atmospheric new particle
formation from biogenic hydrocarbons, Geophys. Res. Lett., 30, 1585,
doi:10.1029/2003GL017000, 2003.

Boy, M., Rannik, Ü., Lehtinen, K. E. J., Tarvainen, V., Hakola, H., and Kulmala, M.:
Nucleation events in the continental boundary layer: long-term statistical analyses of aerosol
relevant characteristics, J. Geophys. Res., 108, 4667, doi:10.1029/2003JD003838, 2003.

31 Boy, M., Mogensen, D., Smolander, S., Zhou, L., Nieminen, T., Paasonen, P., Plass-Dülmer,

32 C., Sipilä, M., Petäjä, T., Mauldin, L., Berresheim, H., and Kulmala, M.: Oxidation of SO2

- 1 by stabilized Criegee intermediate (sCI) radicals as a crucial source for atmospheric sulfuric
- 2 acid concentrations, Atmos. Chem. Phys., 13, 3865-3879, doi:10.5194/acp-13-3865-2013,
- 3 2013.
- 4 Claeys, M., Graham, B., Vas, G., Wang, W., Vermeylen, R., Pashynska, V., Cafmeyer, J.,
- 5 Guyon, P., Andreae, M. O., Artaxo, P., and Maenhaut, W.: Formation of secondary organic
- aerosols through photooxidation of isoprene, Science, 303, 1173–1176, 2004, 27975 Crump, 6
- 7 J. G. and Seinfeld, J. H.: Turbulent deposition and gravitational sedimentation of an aerosol in
- 8 a vessel of arbitrary shape, J. Aerosol Sci., 12, 405–415, 1981.
- 9 Dal Maso, M., Liao, L., Wildt, J., Kiendler-Scharr, A., Kleist, E., Tillmann, R., Sipilä, M.,
- 10 Hakala, J., Lehtipalo, K., Ehn, M., Kerminen, V.-M., Kulmala, M., Worsnop, D., and 11 Mentel, T.: A chamber study of the influence of boreal BVOC emissions and sulphuric acid

12 on nanoparticle formation rates at ambient concentrations, Atmos. Chem. Phys. Discuss., 14,

- 13 31319-31360, doi:10.5194/acpd-14-31319-2014, 2014.
- 14 Damian, V., Sandu, A., Damian, M., Potra, F., and Carmichael, G. R.: The kinetic 15 preprocessor KPP-a software environment for solving chemical kinetics, Comput. Chem. 16 Eng., 26, 1567–1579, 2002.
- 17 Donahue, N. M., Epstein, S. A., Pandis, S. N., and Robinson, A. L.: A two-dimensional 18 volatility basis set: 1. organic-aerosol mixing thermodynamics, Atmos. Chem. Phys., 11, 19 3303-3318, doi:10.5194/acp-11-3303-2011, 2011.
- 20 Ehn, M., Thornton, J. A., Kleist, E., Sipilä, M., Junninen, H., Pullinen, I., Springer, M., Rubach, F., Tillmann, R., Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I.-H., Rissanen, M., 21 22 Jokinen, T., Schobesberger, S., Kangasluoma, J., Kontkanen, J., Nieminen, T., Kurtén, T., 23 Nielsen, L. B., Jorgensen, S., Kjaergaard, H. G., Canagaratna, M., Maso, M. D., Berndt, T., 24 Petäjä, T., Wahner, A., Kerminen, V.-M., Kulmala, M., Worsnop, D. R., Wildt, J., and 25 Mentel, T. F.: A large source of low volatility secondary organic aerosol, Nature, 506, 476-26 479, 2014.
- 27 Epstein, S., Riipinen, I., and Donahue, N. M.: A Semiempirical Correlation between Enthalpy
- 28 of Vaporization and Saturation Concentration for Organic Aerosol, Environ. Sci. Technol., 29 44, 743–748, 2010.
- 30
- Guenther, A., Hewitt, C. N., Erickson, D., Fall, R., Geron, C., Graedel, T., Harley, P., 31 Klinger, L., Lerdau, M., McKay, W. A., Pierce, T., Scholes, B., Steinbrecher, R., Tallamraju,

- R., Taylor, J., and Zimmerman, P.: A global-model of natural volatile organic-compound
 emissions, J. Geophys. Res., 100, 8873–8892, doi:10.1029/94JD02950, 1995.
- 3 Hallquist, M., Wenger, 5 J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M.,
- 4 Dommen, J., Donahue, N. M., George, C., Goldstein, A. H., Hamilton, J. F., Herrmann, H.,
- 5 Hoffmann, T., Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr, A.,
- 6 Maenhaut, W., McFiggans, G., Mentel, Th. F., Monod, A., Prévôt, A. S. H., Seinfeld, J. H.,
- 7 Surratt, J. D., Szmigielski, R., and Wildt, J.: The formation, properties and impact of
- 8 secondary organic aerosol: current and emerging issues, Atmos. Chem. Phys., 9, 5155–5236,
- 9 doi:10.5194/acp-9-5155-2009, 2009.
- 10 Hao, L. Q., Yli-Pirilä, P., Tiitta, P., Romakkaniemi, S., Vaattovaara, P., Kajos, M. K., Rinne,
- 11 J., Heijari, J., Kortelainen, A., Miettinen, P., Kroll, J. H., Holopainen, J. K., Smith, J. N.,
- 12 Joutsensaari, J., Kulmala, M., Worsnop, D. R., and Laaksonen, A.: New particle formation
- 13 from the oxidation of direct emissions of pine seedlings, Atmos. Chem. Phys., 9, 8121–8137,
- 14 doi:10.5194/acp-9-8121-2009, 2009.
- Heiden, A. C., Kobel, K., Langebartels, C., Schuh-Thomas, G., and Wildt, J.: Emissions of
 oxygenated volatile organic compounds from plants Part I: Emissions from lipoxygenase
- 17 activity, J. Atmos. Chem., 45, 143–172, 2003.
- 18 Hermansson, E., Roldin, P., Rusanen, A., Mogensen, D., Kivekäs, N., Väänänen, R., Boy, M.,
- 19 and Swietlicki, E.: Biogenic SOA formation through gas-phase oxidation and gas-to-particle

20 partitioning – a comparison between process models of varying complexity, Atmos. Chem.

- 21 Phys., 14, 11853-11869, doi:10.5194/acp-14-11853-2014, 2014.
- 22 Hoffmann, T., Odum, J. R., Bowman, F., Collins, D., Klockow, D., Flagan, R. C., and
- Seinfeld, J. H.: Formation of organic aerosols from the oxidation of biogenic hydrocarbons, J.
 Atmos. Chem., 26, 189–222, 1997.
- 25 Hoffmann, T., Bandur, R., Marggraf, U., and Linscheid, M.: Molecular composition of 26 organic aerosols formed in the α -pinene/O3 reaction: implications for new particle formation
- 27 processes, J. Geophys. Res., 103, 25569–25578, 1998.
- Jacobson, M. Z.: Numerical techniques to solve condensational and dissolutional growth
 equations when growth is coupled to reversible aqueous reactions, Aerosol Sci. Technol., 27,
 491–498, 1997.
- 31 Jacobson, M. Z.: A Solution to the Problem of Nonequilibrium Acid/Base Gas-Particle

- 1 Transfer at Long Time Step, Aerosol Sci. Technol., 39, 92–103, 2005.
- Jenkin, M. E., Saunders, S. M., and Pilling, M. J.: The tropospheric degradation of volatile
 organic compounds: a protocol for mechanism development, Atmos. Environ., 31, 81–104,
 1997.
- Jenkin, M. E., Wyche, K. P., Evans, C. J., Carr, T., Monks, P. S., Alfarra, M. R., Barley, M.
 H., McFiggans, G. B., Young, J. C., and Rickard, A. R.: Development and chamber
 evaluation of the MCM v3.2 degradation scheme for β-caryophyllene, Atmos. Chem. Phys.,
- 8 12, 5275–5308, doi:10.5194/acp-12-5275-2012, 2012.
- 9 Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H.,
- 10 DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. M.,
- 11 Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. A.,
- 12 Hueglin, C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara,
- 13 P., Ehn, M., Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J., Dunlea, E. J.,
- 14 Huffman, J. A., Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo, Y.,
- 15 Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D., Cottrell,
- 16 L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J. Y, Zhang, Y.
- 17 M., Dzepina, K., Kimmel, J. R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M.,
- 18 Williams, L. R., Wood, E. C., Middlebrook, A. M., Kolb, C. E., Baltensperger, U., and
- Worsnop, D. R.: Evolution of Organic Aerosols in the Atmosphere, Science, 326, 1525–1529,
 2009.
- 21 Kerminen, V.-M., Petäjä, T., Manninen, H. E., Paasonen, P., Nieminen, T., Sipilä, M., 22 Junninen, H., Ehn, M., Gagné, S., Laakso, L., Riipinen, I., Vehkamäki, H., Kurten, T., Ortega, I. K., Dal Maso, M., Brus, D., Hyvärinen, A., Lihavainen, H., Leppä, J., Lehtinen, K. E. J., 23 Mirme, A., Mirme, S., Hõrrak, U., Berndt, T., Stratmann, F., Birmili, W., Wiedensohler, A., 24 25 Metzger, A., Dommen, J., 5 Baltensperger, U., Kiendler-Scharr, A., Mentel, T. F., Wildt, J., Winkler, P. M., Wagner, P. E., Petzold, A., Minikin, A., Plass-Dülmer, C., Pöschl, U., 26 27 Laaksonen, A., and Kulmala, M.: Atmospheric nucleation: highlights of the EUCAARI 28 project and future directions, Atmos. Chem. Phys., 10, 10829-10848, doi:10.5194/acp-10-29 10829-2010, 2010.
- Kiendler-Scharr, A., Wildt, J., Dal Maso, M., Hohaus, T., Kleist, E., Mentel, T. F., Tillmann,
 R., Uerlings, R., Schurr, U., and Wahner, A.: New particle formation in forests inhibited by
 isoprene emissions, Nature, 461, 381–384, 2009.

- 1 Kokkola, H., Yli-Pirilä, P., Vesterinen, M., Korhonen, H., Keskinen, H., Romakkaniemi, S.,
- 2 Hao, L., Kortelainen, A., Joutsensaari, J., Worsnop, D. R., Virtanen, A., and Lehtinen, K. E. J.:
- 3 The role of low volatile organics on secondary organic aerosol formation, Atmos. Chem.
- 4 Phys., 14, 1689–1700, doi:10.5194/acp-14-1689-2014, 2014.
- 5 Korhonen, H., Lehtinen, K. E. J., and Kulmala, M.: Multicomponent aerosol dynamics model
- 6 UHMA: model development and validation, Atmos. Chem. Phys., 4, 757-771,
- 7 doi:10.5194/acp-4-757-2004, 2004.
- Kulmala, M., Pirjola, U., and Makela, J. M.: Stable sulphate clusters as a source of new
 atmospheric particles, Nature, 404, 66–69, 2000.
- 10 Kulmala, M., Vehkamäki, H., Petäjä, T., Dal Maso, M., Lauri, A., Kerminen, V. M., Birmili,
- 11 W., and McMurry, P. H.: Formation and growth rates of ultrafine atmospheric particles: a
- 12 review of observations, J. Aerosol Sci., 35, 143–176, 2004.
- 13 Kulmala, M., Lehtinen, K. E. J., and Laaksonen, A.: Cluster activation theory as an
- 14 explanation of the linear dependence between formation rate of 3nm particles and sulphuric
- 15 acid concentration, Atmos. Chem. Phys., 6, 787–793, doi:10.5194/acp-6-787-2006, 2006.
- 16 Kulmala, M., Riipinen, I., Sipilä, M., Manninen, H. E., Petäjä, T., Junninen, H., Dal Maso,
- 17 M., Mordas, G., Mirme, A., Vana, M., Hirsikko, A., Laakso, L., Harrison, R. M., Hanson, I.,
- 18 Leung, C., Lehtinen, K. E. J., and Kerminen, V. M.: Toward direct measurement of
- 19 atmospheric nucleation, Science, 318, 89–92, 2007.
- 20 Kulmala, M., Nieminen, T., Chellapermal, R., Makkonen, R., Bäck, J., and Kerminen, V.-M.:
- 21 Climate feedbacks linking the increasing atmospheric CO2 concentration, BVOC emissions,
- 22 aerosols and clouds in forest ecosystems, in: Biology, Controls and Models of Tree Volatile
- 23 Organic Compound Emissions, Springer, Springer Netherlands, 489–508, 2013.
- 24 Kulmala, M., Nieminen, T., Nikandrova, A., Lehtipalo, K., Manninen, H. E., Kajos, M. K.,
- 25 Kolari, P., Lauri, A., Petäjä, T., Krejci, R., Hansson, H.-C., Swietlicki, E., Lindroth, A.,
- 26 Christensen, T. R., Arneth, A., Hari, 5 P., Bäck, J., Vesala, T., and Kerminen, V.-M.: CO2 -
- 27 induced terrestrial climate feedback mechanism: from carbon sink to aerosol source and back,
- 28 Boreal Environ. Res., 19, 122–131, 2014.
- 29 Kurtén, T., Loukonen, V., Vehkamäki, H., and Kulmala, M.: Amines are likely to enhance
- 30 neutral and ion-induced sulfuric acid-water nucleation in the atmosphere more effectively
- 31 than ammonia, Atmos. Chem. Phys., 8, 4095–4103, doi:10.5194/acp-8-4095-2008, 2008.

- 1 Laaksonen, A., Kulmala, M., O'Dowd, C. D., Joutsensaari, J., Vaattovaara, P., Mikkonen, S.,
- 2 Lehtinen, K. E. J., Sogacheva, L., Dal Maso, M., Aalto, P., Petäjä, T., Sogachev, A., Yoon, Y.
- J., Lihavainen, H., Nilsson, D., Facchini, M. C., Cavalli, F., Fuzzi, S., Hoffmann, T., Arnold,
- 4 F., Hanke, M., Sellegri, K., Umann, B., Junkermann, W., Coe, H., Allan, J. D., Alfarra, M. R.,
- 5 Worsnop, D. R., Riekkola, M. -L., Hyötyläinen, T., and Viisanen, Y.: The role of VOC
- 6 oxidation products in continental new particle formation, Atmos. Chem. Phys., 8, 2657–2665,
- 7 doi:10.5194/acp-8-2657-2008, 2008.
- Lai, A. and Nazaroff, W. W.: Modelling indoor particle deposition from turbulent flow onto
 smooth surfaces, J. Aerosol Sci., 31, 463–476, 2000.
- 10 Matsunaga, A. and Ziemann, P. J.: Gas-wall partitioning of organic compounds in a Teflon
- 11 film chamber and potential effects on reaction product and aerosol yield measurements,
- 12 Aerosol Sci. Tech., 44, 881–892, 2010.
- 13 Mauldin, R., Frost, G., Chen, G., Tanner, D., Prevot, A., Davis, D., and Eisele, F.: OH
- 14 measurements during the First Aerosol Characterization Experiment (ACE 1): observations
- 15 and model comparisons, J. Geophys. Res., 103, 16713–16729, doi:10.1029/98JD00882, 1998.
- 16 McMurry, P. H., and Friedlander, S. K.: New particle formation in the presence of an aerosol,
- 17 Atmos. Environ., 13, 1635–1651, 1979.
- 18 McMurry P. H., and Grosjean, D.: Gas and aerosol wall losses in Teflon film smog chambers,
- 19 Environ. Sci. Technol. 19, 1176–1182, 1985.
- 20 McMurry, P. H. and Rader, D. J.: Aerosol wall losses in electrically charged chambers.
- 21 Aerosol Sci. Technol., 4, 249–268, 1985.
- McMurry P. H., Kulmala, M., Worsnop D. R.: Special Issue on Aerosol Measurements in the
 1 nm Range, Aerosol Sci. Technol. 45, i, 2011.
- 24 Mentel, T. F., Wildt, J., Kiendler-Scharr, A., Kleist, E., Tillmann, R., Dal Maso, M., Fisseha,
- 25 R., Hohaus, Th., Spahn, H., Uerlings, R., Wegener, R., Griffths, P. T., Dinar, E., Rudich, Y.,
- 26 and Wahner, A.: Photochemical production of aerosols from real plant emissions, Atmos.
- 27 Chem. Phys., 9, 4387–4406, doi:10.5194/acp-9-4387-2009, 2009.
- 28 Metzger, A., Verheggen, B., Dommen, J., Duplissy, J., Prevot, A. S., Weingartner, E.,
- 29 Riipinen, I., Kulmala, M., Spracklen, D. V., Carslaw, K. S., and Baltensperger U.: Evidence
- 30 for the role of organics in aerosol particle formation under atmospheric conditions, P. Natl.

- 1 Acad. Sci. USA, 107, 6646–6651, 2010.
- Mirme, S., Mirme, A., Minikin, A., Petzold, A., Hõrrak, U., Kerminen, V. -M., and Kulmala,
 M.: Atmospheric sub-3 nm particles at high altitudes, Atmos. Chem. Phys., 10, 437–451,
 doi:10.5194/acp-10-437-2010, 2010.
- Nannoolal, Y., Rarey, J., Ramjugernath, D., and Cordes, W.: Estimation of pure component
 properties Part 1, Estimation of the normal boiling point of non-electrolyte organic
 compounds via group contributions and group interactions, Fluid Phase Equilibr., 226, 45–63,
 2004.
- 9 Nannoolal, J., Rarey, J., and Ramjugernath, D.: Estimation of pure component properties Part
 10 3. Estimation of the vapour pressure of non-electrolyte organic compounds via group
 11 contributions and group interactions Fuild Phase Equilibria, 269, 117–133, 2008.
- Owen, S. M., Harley, P., Guenther, A., and Hewitt, C. N.: Light dependency of VOC
 emissions from selected Mediterranean plant species, Atmos. Environ., 36, 3147–3159, 2002.
- 14 Paasonen, P., Nieminen, T., Asmi, E., Manninen, H. E., Petäjä, T., Plass-Dülmer, C., Flentje,
- 15 H., Birmili, W., Wiedensohler, A., Hõrrak, U., Metzger, A., Hamed, A., Laaksonen, A.,
- 16 Facchini, M. C., Kerminen, V.-M., and Kulmala, M.: On the roles of sulphuric acid and low-
- 17 volatility organic vapours in the initial steps of atmospheric new particle formation, Atmos.
- 18 Chem. Phys., 10, 11223–11242, doi:10.5194/acp-10-11223-2010, 2010.
- Park, S. H., Kim, H. O., Han, Y. T., Kwon, S. B., and Lee, K. W.: Wall loss rate of
 polydispersed aerosols, Aerosol Sci. Tech., 35, 710–717, 2001.
- 21 Petäjä, T., Mauldin, III, R. L., Kosciuch, E., McGrath, J., Nieminen, T., Paasonen, P., Boy,
- 22 M., Adamov, A., Kotiaho, T., and Kulmala, M.: Sulfuric acid and OH concentrations in a
- 23 boreal forest site, Atmos. Chem. Phys., 9, 7435–7448, doi:10.5194/acp-9-7435-2009, 2009.
- Pierce, J. R., Engelhart, G. J., Hildebrandt, L., Weitkamp, E. A., Pathak, R. K., Donahue, N.
 M., Robinson, A. L., Adams, P. J., and Pandis, S. N.: Constraining Particle Evolution from
 Wall Losses, Coagulation, and Condensation-Evaporation in Smog- Chamber Experiments:
 Optimal Estimation Based on Size Distribution Measurements, Aerosol Sci. Technol., 42,
- 28 1001–1015, 2008.
- 29 Riccobono, F., Schobesberger, S., Scott, C. E., Dommen, J., Ortega, I. K., Rondo, L.,
- 30 Almeida, J. a., Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Downard, A., Dunne,

- 1 E. M., Duplissy, J., Ehrhart, S., Flagan, R. C., Franchin, A., Hansel, A., Junninen, H., Kajos,
- 2 M., Keskinen, H., Kupc, A., Kürten, A., Kvashin, A. N., Laaksonen, A., Lehtipalo, K.,
- 3 Makhmutov, V., Mathot, S., Nieminen, T., Onnela, A., Petäjä, T., Praplan, A. P., Santos, F.
- 4 D., Schallhart, S., Seinfeld, J. H., Sipilä, M., Spracklen, D. V., Stozhkov, Y., Stratmann, F.,
- 5 Tomé, A., Tsagkogeorgas, G., Vaattovaara, P., Viisanen, Y., Vrtala, A., Wagner, P. E.,
- 6 Weingartner, E., Wex, H., Wimmer, D., Carslaw, K. S., Curtius, J., Donahue, N. M., Kirkby,
- 7 J., Kulmala, M., Worsnop, D. R., and Baltensperger, U.: Oxidation products of biogenic
- 8 emissions contribute to nucleation of atmospheric particles, Science, 344, 717–721, 2014.
- 9 Riipinen, I., Sihto, S.-L., Kulmala, M., Arnold, F., Dal Maso, M., Birmili, W., Saarnio, K.,
- 10 Teinilä, K., Kerminen, V.-M., Laaksonen, A., and Lehtinen, K. E. J.: Connections between
- 11 atmospheric sulphuric acid and new particle formation during QUEST III-IV campaigns in
- 12 Heidelberg and Hyytiälä, Atmos. Chem. Phys., 7, 1899–1914, doi:10.5194/acp-7-1899-2007,
- 13 2007.
- 14 Riipinen, I., Pierce, J. R., Donahue, N. M., and Pandis, S. N.: Equilibration time scales of
- 15 organic aerosol inside thermodenuders: Evaporation kinetics versus thermodynamics, Atmos.
- 16 Environ., 44, 597–607, 2010.
- Riipinen, I., Yli-Juuti, T., Pierce, J. R., Petäjä, T., Worsnop, D. R., Kulmala, M., and
 Donahue, N. M.: The contribution of organics to atmospheric nanoparticle growth, Nat.
 Geosci., 5, 453–458, 2012.
- 20 Roldin, P., Swietlicki, E., Schurgers, G., Arneth, A., Lehtinen, K. E. J., Boy, M., and
- 21 Kulmala, M.: Development and evaluation of the aerosol dynamics and gas phase chemistry
- 22 model ADCHEM, Atmos. Chem. Phys., 11, 5867–5896, doi:10.5194/acp-11-5867- 2011,
 23 2011.
- Roldin, P., Eriksson, A. C., Nordin, E. Z., Hermansson, E., Mogensen, D., Rusanen, A., Boy,
 M., Swietlicki, E., Svenningsson, B., Zelenyuk, A., and Pagels, J.: Modelling non-equilibrium
 secondary organic aerosol formation and evaporation with the aerosol dynamics, gas and
 particle-phase chemistry kinetic multilayer model ADCHAM, Atmos. Chem. Phys., 14,
- 28 7953–7993, doi:10.5194/acp-14-7953-2014, 2014.
- 29 Ruuskanen, T. M., Taipale, R., Rinne, J., Kajos, M. K., Hakola, H., and Kulmala, M.:
- 30 Quantitative long-term measurements of VOC concentrations by PTR-MS: annual cycle at a
- 31 boreal forest site, Atmos. Chem. Phys. Discuss., 9, 81-134, doi:10.5194/acpd-9-81-2009,

- 1 2009.
- Saunders, S. M., Jenkin, M. E., Derwent, R. G., and Pilling, M. J.: Protocol for the
 development of the Master Chemical Mechanism, MCM v3 (Part A): tropospheric
 degradation of nonaromatic volatile organic compounds, Atmos. Chem. Phys., 3, 161–180,
 doi:10.5194/acp-3-161-2003, 2003.
- Schimang, R., Folkers, A., Kleffmann, J., Kleist, E., Miebach, M., and Wildt, J.: Uptake of
 gaseous nitrous acid (HONO) by several plant species, Atmos. Environ., 40, 1324–1335,
 2006.
- 9 Schobesberger, S., Junninen, H., Bianchi, F., Lönn, G., Ehn, M., Lehtipalo, K., Dommen, J.,
- 10 Ehrhart, S., Ortega, I. K., Franchin, A., Nieminen, T., Riccobono, F., Hutterli, M., Duplissy,
- 11 J., Almeida, J., Amorim, A., Breitenlechner, M., Downard, A. J., Dunne, E. M., Flagan, R. C.,
- 12 Kajos, M., Keskinen, H., Kirkby, J., Kupc, A., Kürten, A., Kurtén, T., Laaksonen, A., Mathot,
- 13 S., Onnela, A., Praplan, A. P., Rondo, L., Santos, F. D., Schallhart, S., Schnitzhofer, R.,
- 14 Sipilä, M., Tomé, A., Tsagkogeorgas, G., Vehkamäki, H., Wimmer, D., Baltensperger, U.,
- 15 Carslaw, K. S., Curtius, J., Hansel, A., Petäjä, T., Kulmala, M., Donahue, N. M., and
- 16 Worsnop, D. R.: Molecular understanding of atmospheric particle formation from sulfuric
- 17 acid and large oxidized organic molecules, P. Natl. Acad. Sci. USA, 110, 17223–17228, 2013.
- Sihto, S.-L., Kulmala, M., Kerminen, V.-M., Dal Maso, M., Petäjä, T., Riipinen, I., Korhonen,
 H., Arnold, F., Janson, R., Boy, M., Laaksonen, A., and Lehtinen, K. E. J.: Atmospheric
 sulphuric acid and aerosol formation: implications from atmospheric measurements for
 nucleation and early growth mechanisms, Atmos. Chem. Phys., 6, 4079–4091,
 doi:10.5194/acp- 6-4079-2006, 2006.
- 23 Sipilä, M., Berndt, T., Petäjä, T., Brus, D., Vanhanen, J., Stratmann, F., Patokoski, J.,
- 24 Mauldin, Roy, L., I., Hyvarinen, A.-P., Lihavainen, H., and Kulmala, M.: The role of sulfuric
- acid in atmospheric nucleation, Science, 327, 1243–1246, 2010.
- 26 Spracklen, D. V., Jimenez, J. L., Carslaw, K. S., Worsnop, D. R., Evans, M. J., Mann, G. W.,
- 27 Zhang, Q., Canagaratna, 5 M. R., Allan, J., Coe, H., McFiggans, G., Rap, A., and Forster, P.:
- 28 Aerosol mass spectrometer constraint on the global secondary organic aerosol budget, Atmos.
- 29 Chem. Phys., 11, 12109–12136, doi:10.5194/acp-11-12109-2011, 2011.
- 30 Surratt, J. D., Murphy, S. M., Kroll, J. H., Ng, N. L., Hildebrandt, L., Sorooshian, A.,
- 31 Szmigielski, R., Vermeylen, R., Maenhaut, W., Claeys, M., Flagan, R. C., and Seinfeld, J. H.:

- 1 Chemical composition of secondary organic aerosol formed from the photooxidation of 2 isoprene, J. Phys. Chem. A, 110, 9665-9690, 2006.
- Topping, D. O., McFiggans, G. B., Kiss, G., Varga, Z., Facchini, M. C., Decesari, S., and
 Mircea, M.: Surface tensions of multi-component mixed inorganic/organic aqueous systems
 of atmospheric significance: measurements, model predictions and importance for cloud
 activation predictions, Atmos. Chem. Phys., 7, 2371–2398, doi:10.5194/acp-7-2371-2007,
 2007.
- 8 Tsigaridis, K. and Kanakidou, M.: Global modelling of secondary organic aerosol in the 9 troposphere: a sensitivity analysis, Atmos. Chem. Phys., 3, 1849–1869, doi:10.5194/acp-3-10 1849- 2003, 2003.
- 11 Vanhanen, J., Mikkilä, J., Lehtipalo, K., Sipilä, M., Manninen, H. E., Siivola, E., Petäjä, T.,

12 and Kulmala, M.: Particle size magnifier for nano-CN detection, Aerosol Sci. Tech., 45, 533–

- 13 542, 2011.
- 14 VanReken, T. M., Greenberg, J. P., Harley, P. C., Guenther, A. B., and Smith, J. N.: Direct
- measurement of particle formation and growth from the oxidation of biogenic emissions,
 Atmos. Chem. Phys., 6, 4403–4413, doi:10.5194/acp-6-4403-2006, 2006.
- 17 Wang, Z. B., Hu, M., Mogensen, D., Yue, D. L., Zheng, J., Zhang, R. Y., Liu, Y., Yuan, B.,
- 18 Li, X., Shao, M., Zhou, L., Wu, Z. J., Wiedensohler, A., and Boy, M.: The simulations of 19 sulfuric acid concentration and new particle formation in an urban atmosphere in China,
- is surfaire delle concentration and new particle formation in an aroun autosphere in en
- 20 Atmos. Chem. Phys., 13, 11157–11167, doi:10.5194/acp-13-11157-2013, 2013.
- Zhang, R.: Getting to the critical nucleus of aerosol formation, Science, 328, 1366–1367,
 2010.
- Zhang, X., Cappa, C. D., Jathar, S. H., McVay, R. C., Ensberg, J. J., Kleeman, M. J., and
 Seinfeld, J. H.: Influence of vapor wall loss in laboratory chambers on yields of secondary
 organic aerosol, P. Natl. Acad. Sci. USA, 111, 5802–5807, 2014.
- 26 Zhang, X., Schwantes, R. H., McVay, R. C., Lignell, H., Coggon, M. M., Flagan, R. C., and
- 27 Seinfeld, J. H.: Vapor wall deposition in Teflon chambers, Atmos. Chem. Phys., 15, 4197-
- 28 4214, doi:10.5194/acp-15-4197-2015, 2015.
- 29 Zuend, A., Marcolli, C., Luo, B. P., and Peter, T.: A thermodynamic model of mixed organic-
- 30 inorganic aerosols to predict activity coefficients, Atmos. Chem. Phys., 8, 4559-4593,

1 doi:10.5194/acp-8-4559-2008, 2008.

Zuend, A., Marcolli C., Booth , A. M., Lienhard, D. M., Soonsin, V., Krieger, U. K., Topping,
D. O., McFiggans G., Peter, T., and Seinfeld, J. H.: New and extended parameterization of the
thermodynamic model AIOMFAC: calculation of activity coefficients for organic-inorganic
mixtures containing carboxyl, hydroxyl, carbonyl, ether, ester, alkenyl, alkyl, and aromatic
functional groups, Atmos. Chem. Phys., 11, 9155–9206, doi:10.5194/acp-11-9155-2011,
2011.

9 Table 1. The VOCs measured with GC-MS in the JPAC plant chamber. The measured 10 concentrations were used as input for the ADCHAM model. The "other MTs" and "other 11 SQTs" refer to other monoterpenes and sesquiterpenes than those specified in the table,

12 respectively.

Isoprene				
α-pinene	β-pinene	myrcene	sabinene	camphene
ocimene	Δ^3 -carene	α-terpinene	Δ -terpinene	α -phellandrene
β-phellandrene	terpinolene	tricyclene	other MTs	
farnesene	β-caryophyllene	α-longipinene	Δ -cardinene	other SQTs
2-butanol	hexanal	benzene	toluene	eucalyptol
nonanal	bornyl acetate	methyl salicylate		

1 Table 2. Summary of the model sensitivity tests that we performed in order to evaluate and

2 constrain possible mechanisms for the formation and growth of nano-CN during the JPAC

3 experiments.

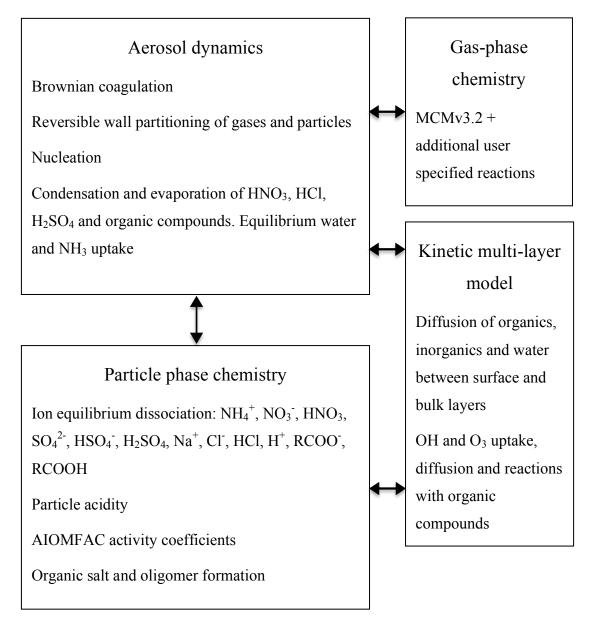
Sensitivity test category	Varied parameters (method) 1	Varied parameters (method) 2	
Condensable VOC properties method	MCMv3.2 + the Nannoolal vapour pressure method	2D-VBS	
2D-VBS assumptions	Functionalization, fragmentation and OH reaction rates	O:C of the first generation products	
Particle dry deposition loss rates	Varying u* in the range 0.1- 0.01 m s^{-1}		
Gas-phase chemistry	Uncertainties related to the assumed NO _x inflow concentration	Influence of the UV-light on the VOC composition	
Influence of ELVOCs on the particle growth	ELVOC formation from ozonolysis and OH-oxidation of terpenes		
Nano-CN formation and initial growth	Mechanisms (Eq. 1-6)	Surface tension 0.02-0.07 N m ⁻¹	
Oligomerization in the particle phase	Peroxyhemiacetal formation, possibly acid catalysed by co- condensing H ₂ SO ₄ (Eq. 7)		
Non-ideal mixing in the particle phase	Activity coefficients from AIOMFAC		
Reversible VOC wall losses	With or without absorptive uptake on the glass walls		
SOA phase-state	Liquid-like SOA	Solid-like SOA	



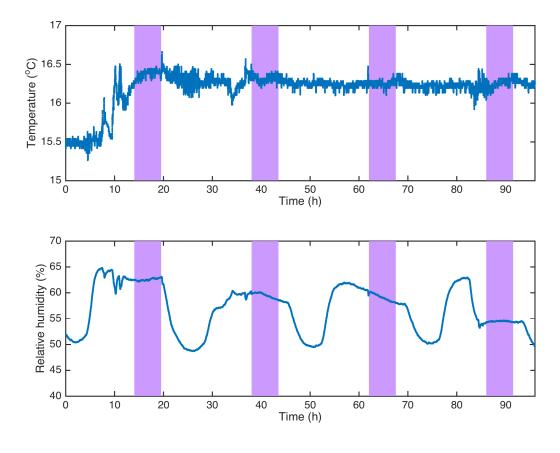
1 Table 3. Coefficient of determination (R^2) between the modelled and measured (PSM-CPC)

2	total particle number	concentration	during the	experiment	al campaign.	
---	-----------------------	---------------	------------	------------	--------------	--

Nano-CN formation mechanism	ELVOC _{nucl} source	\mathbf{R}^2
$J = 20 \text{ cm}^{-3} \text{ s}^{-1}$ when UV is on		0.937
$J = K_2 [H_2 S O_4]^2$		0.787
$J = A_1[\mathrm{H}_2\mathrm{SO}_4]$		0.887
$J = K_3[H_2SO_4][ELVOC_{nucl}]$	MT and SQT ox. by O_3 and OH (R3, R4)	0.951
$J = K_3[H_2SO_4][ELVOC_{nucl}]$	MT and SQT ox. OH (R4)	0.955
$J = K_3[H_2SO_4][ELVOC_{nucl}]$	SQT ox. by O_3 and OH (R3, R4)	0.891
$J = K_4 [H_2 SO_4]^2 [ELVOC_{nucl}]$	SQT ox. by O_3 and OH (R3, R4)	0.92
$J = K_4 [H_2 SO_4]^2 [ELVOC_{nucl}]$	SQT ox. by $O_3(R3)$	0.951
$J = A_5[\text{ELVOC}_{\text{nucl}}]$	MT and SQT ox. OH (R4)	0.815



3 Figure 1. Schematic picture of the ADCHAM model structure.



2 Figure 2. Measured temperature (a) and relative humidity (b) in the JPAC reaction chamber

- 3 during the first four days of the measurement campaign. The purple bars indicate UV-on
- 4 periods.

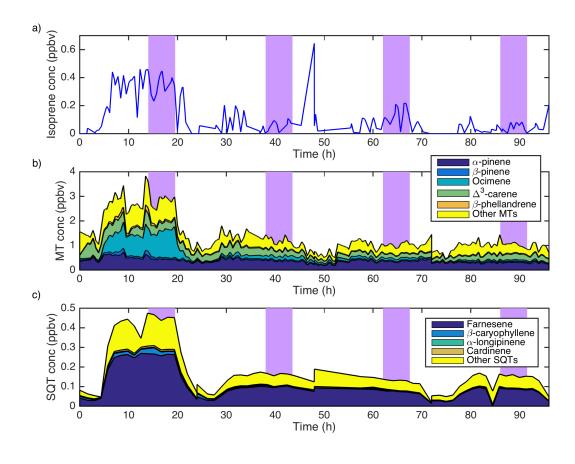
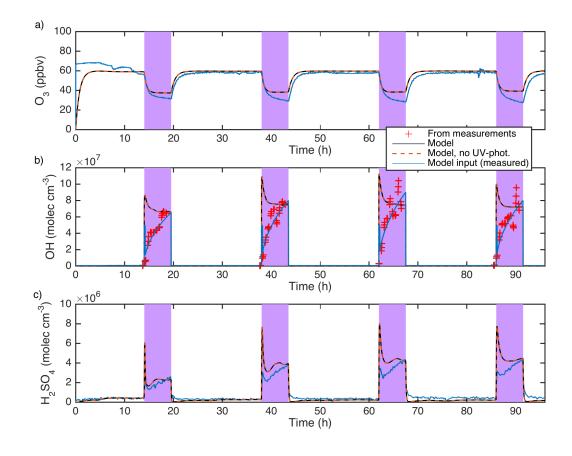


Figure 3. Isoprene (a), monoterpene (b), and sesquiterpene (c) concentrations measured from
the outlet air of the JPAC plant chamber. In Fig. b and c we left out those terpenes that had a
negligible contribution to the total terpene concentration. Purple bars indicate UV-on periods
during the measurements.



2 Figure 4. Measured and modelled (a) O₃ concentrations, (b) OH concentration, and (c) H₂SO₄ 3 concentration. The OH concentration (red +) was not measured directly but was derived from 4 the observed 2-butanol loss rate. The model results are from the simulations where the O₃, 5 OH and H₂SO₄ concentrations were simulated by the model instead of given as the input 6 concentration to the model. The results given by the orange dashed line corresponds to a 7 simulation were the 254 nm UV-light was only used to calculate the photolysis rates of O₃ but not influencing photolysis rates of the other compounds. In the model simulations, we used 8 the empirically determined JPAC reaction chamber-specific OH sink of 4 $\ensuremath{\text{s}}^{\text{-1}}$ and an O_3 9 concentration of 170 ppbv in the inflow to the chamber. The blue lines give the model input 10 11 concentrations that were used for the simulations presented in Sect. 4. For O₃ and H₂SO₄, the 12 model input concentration was taken directly from the measurements.

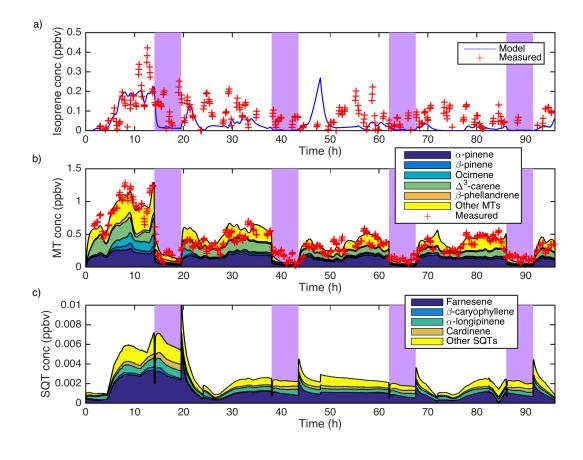
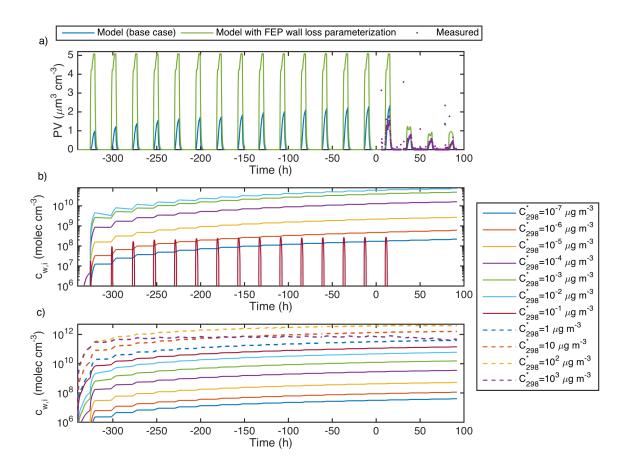
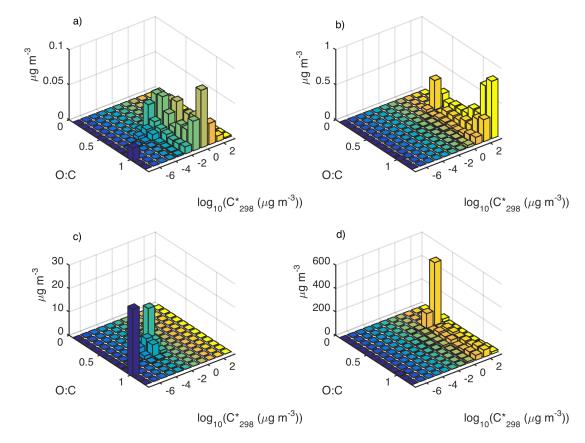


Figure 5. Modelled isoprene (a), monoterpene (b), and sesquiterpene (c) concentrations together with the measured isoprene and monoterpene concentrations in the JPAC reaction chamber. In Fig. b we left out those monoterpenes that had a negligible contribution to the total monoterpene concentration. The purple bars indicate UV-on periods during the measurements.



2 Figure 6. (a) Modelled and measured SOA volume concentration and 2D-VBS VOC wall 3 uptake onto (b) the glass walls (Eq. 10-11) and (c) the PTFE Teflon walls when considering 4 that the PTFE Teflon walls behave as FEB Teflon walls (Eq. 9, 12 and 13). The VOCs are summed over the all O:C but divided into the different C_{298}^* bins in the 2D-VBS. At time 0 h 5 6 the intensive measurement campaign started. Fig 6a also shows the modelled SOA volume 7 concentration when assuming that all chamber walls behave as FEP Teflon walls using Eq. 8-9, $k_e = 4.2 \text{ s}^{-1}$ and the VOC FEP Teflon wall mass accommodation coefficient 8 9 parameterization from Zhang et al. (2015). ELVOC formation from R1-R2 was not 10 considered for the simulations.

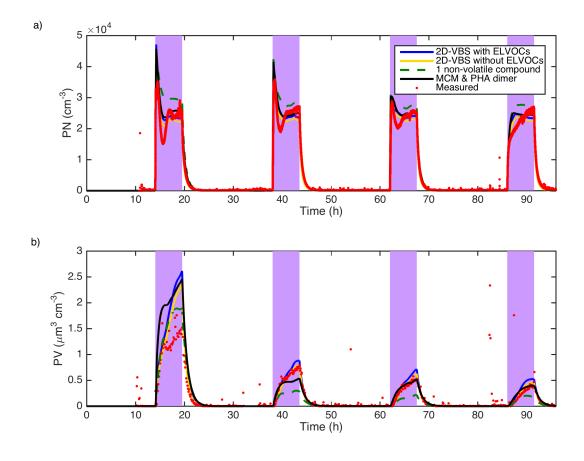
- 11
- 12



2 Figure 7. Modelled 2D-VBS distribution for (a) the SOA particles, (b) the gas-phase, (c) the

3 VOCs on the glass-walls and (d) the VOCs on the Teflon floor. The model results are from

- 4 Day-3 after 5 hours with UV-lights on. For this simulation R1 and R2 were used to simulate
- 5 the ELVOC formation.



2 Figure 8. Modelled and measured (a) total particle number concentration and (b) total particle 3 volume concentration. The model results are from simulations with the 2D-VBS and with or 4 without ELVOC formation via ozonolysis and OH-oxidation of monoterpenes and 5 sesquiterpenes (R1-R2), from a simulation when the MCM compounds were used as the 6 condensable organic compounds and considering PHA dimer formation with Eq. 7, and from 7 a simulation with only one condensable non-volatile compound. The mass yield of the non-8 volatile compound formed from O3 and OH oxidation of all monoterpenes and sesquiterpens 9 was 60 %.

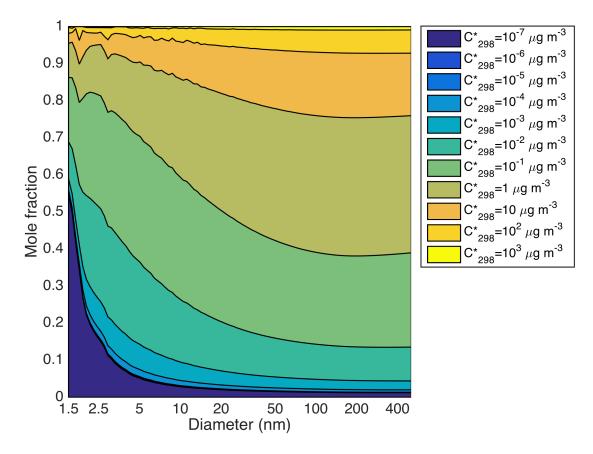




Figure 9. Modelled SOA volatility distribution as a function of the particle size from a
simulation with the 2D-VBS and ELVOC formation from ozonolysis and OH-oxidation of
monoterpenes and sesquiterpenes (R1-R2). The results are from Day-3 after 5 hours into the
UV-light period.

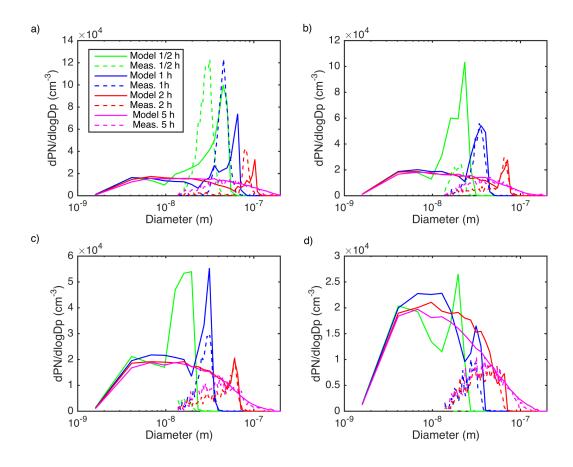




Figure 10. Modelled and measured particle number size distribution from (a) Day-1, (b) Day-2, (c) Day-3 and (d) Day-4 of the experimental campaign. The particle number size distributions are from ¹/₂, 1, 2 and 5 hours after the UV-light was turned on. The model results are from a simulation with the 2D-VBS (including ELVOC formation (R1-R2)) and the fullmoving size distribution method.

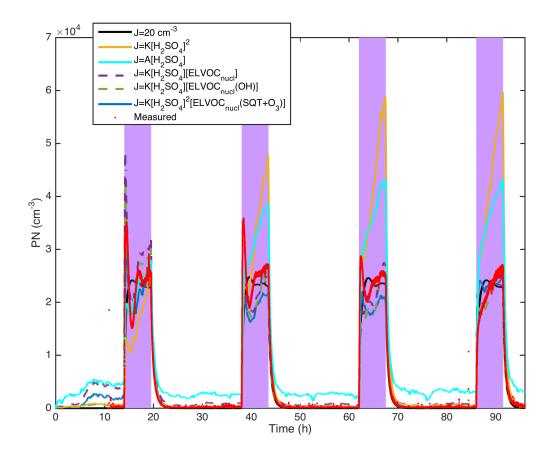


Figure 11. Measured (PSM-CPC) and modelled total particle number concentration with
different nano-CN formation mechanism. The SOA formation was modelled with the 2DVBS approach including ELVOC formation (R1-R4).