Modelling the contribution of biogenic volatile organic compounds to new particle formation in the Jülich plant atmosphere chamber

4

P. Roldin^{1,2}, L. Liao¹, D. Mogensen¹, M. Dal Maso³, A. Rusanen¹, V.-M.
Kerminen¹, T. F. Mentel⁴, J. Wildt⁵, E. Kleist⁵, A. Kiendler-Scharr⁴, R. Tillmann⁴,
M. Ehn¹, M. Kulmala¹, and M. Boy¹

- 8 [1]{Department of Physics, University of Helsinki, P.O. Box 64, 00014 Helsinki, Finland}
- 9 [2] {Division of Nuclear Physics, Lund University, P.O. Box 118, 221 00 Lund, Sweden}

10 [3]{Department of Physics, Tampere University of Technology, P.O. Box 692, 33101
11 Tampere, Finland}

- 12 [4] {Institute for Energy- and Climate Research (IEK-8), Forschungszentrum Jülich, 5242513 Jülich, Germany}
- 14 [5] {Institute of Biogeosciences (IBG-2), Forschungszentrum Jülich, 52425 Jülich, Germany}

15 Correspondence to: P. Roldin (pontus.roldin@nuclear.lu.se)

16

17 Abstract

We used the Aerosol Dynamics gas- and particle-phase chemistry model for laboratory 18 19 CHAMber studies (ADCHAM) to simulate the contribution of BVOC plant emissions to the observed new particle formation during photooxidation experiments performed in the Jülich 20 21 Plant-Atmosphere Chamber. ADCHAM couples the detailed gas-phase chemistry from Master Chemical Mechanism with a novel aerosol dynamics and particle phase chemistry 22 23 module. Our model simulations reveal that the observed particle growth either may have been 24 controlled by the formation rate of semi- and low-volatility organic compounds in the gas-25 phase or by acid catalyzed heterogeneous reactions between semi-volatility organic compounds in the particle surface layer (e.g. peroxyhemiacetal dimer formation). The 26 27 contribution of extremely low-volatility organic gas-phase compounds to the particle formation and growth was suppressed because of their rapid and irreversible wall losses. The 28

best correlation between the modelled and measured total particle number concentration (R^2 >0.97) was achieved if the nano-CN was formed by kinetic nucleation involving both sulphuric acid and BVOC oxidation products.

4

5 1 Introduction

6

7 New particle formation, including formation of nano condensation nucleii (nano-CN) 8 (McMurry et al., 2011) and their growth to larger sizes, has been observed world-widely in 9 continental boundary layers and free troposphere (Kulmala et al., 2004; Mirme et al., 2010). 10 Field observations, laboratory experiments and model simulations indicate that gaseous 11 sulphuric acid (H₂SO₄) plays an important role in atmospheric nano-CN formation, yet H₂SO₄ alone appears not to be able to explain all the steps of this process (Kulmala et al., 2000; Boy 12 et al., 2003; Sipilä et al., 2010; Riipinen et al., 2007; Sihto et al., 2006; Kerminen et al., 2010; 13 14 Kulmala et al., 2013, 2014). Basic compounds like ammonia and certain amines have been 15 proposed to act as stabilizing compounds in nano-CN clusters (Berndt et al., 2010; Almeida et 16 al., 2013; Kurtén et al., 2008), while subsequent steps of atmospheric new particle formation 17 seem to rely on the presence of low-volatility organic compounds (LVOCs) (e.g., Metzger et al., 2010; Paasonen et al., 2010; Riipinen et al., 2012; Ehn et al., 2014; Schobesberger et al., 18 19 2013).

20 Oxidation products of biogenic volatile organic compounds (BVOCs) constitute the largest source of secondary organic aerosol (SOA) in the global atmosphere (Tsigaridis and 21 22 Kanakidou, 2003; Hallquist et al., 2009; Spracklen et al., 2011), accounting for the main 23 composition of SOA condensational growth (VanReken et al., 2006; Hao et al., 2009; 24 Riipinen et al., 2012). BVOC oxidation also produces extremely low-volatility organic 25 compounds (ELVOCs) essential to the new particle formation process in the atmosphere (Ehn 26 et al., 2014). The most abundant group of BVOCs, accounting for more than half of their global emissions, are terpenoids (Guenther et al., 1995). Terpenoids include compounds 27 28 consisting of one to several isoprene units, e.g. isoprene (C_5H_8), monoterpenes ($C_{10}H_{16}$), and 29 sesquiterpenes ($C_{15}H_{24}$). Oxidation products of monoterpenes have substantial contribution to 30 SOA formation (Hoffmann et al., 1997, 1998; Laaksonen et al., 2008), and low-volatility 31 substances produced by sesquiterpene-ozone reactions may also initiate SOA formation 32 (Bonn and Moortgat, 2003). Oxidation of isoprene leads to the formation of SOA (Surratt et al., 2006; Claeys et al., 2004), yet isoprene may also suppress the new particle formation
process due to its high reactivity with OH (Kiendler-Scharr et al., 2009). Overall, the exact
contribution of BVOCs to SOA still remains uncertain, especially with respect to the initial
steps of atmospheric new particle formation.

5 In this study, we used the Aerosol Dynamics, gas- and particle-phase chemistry model for 6 laboratory CHAMber studies (ADCHAM) (Roldin et al., 2014), to investigate the new 7 particle formation during a measurement campaign conducted in the Jülich Plant Atmosphere 8 Chamber (JPAC). The in-depth analysis of the chamber measurements is discussed in another 9 paper (Dal Maso et al., 2014). Here, we use the full chamber dataset including gas and 10 particle phase measurements, either as model input or to evaluate the modelled performance 11 concerning the nano-CN formation and growth.

12

13 2 Measurement set up

14

15 The experiments were conducted in the JPAC located at Forschungszentrum Jülich, Germany. 16 Detailed description regarding the chamber facility was given in previous articles (e.g., Mentel et al., 2009; Schimang et al., 2006). In addition, more details about this measurement 17 18 campaign can be found from Dal Maso et al. (2014). In brief, the system consisted of two 19 borosilicate glass chambers with PTFE Teflon floors. The chambers were operated as 20 continuously stirred tank reactors (CSTR) with Teflon fans ensuring homogeneous air mixing. Each chamber housed adjustable temperature between 10 and 50 °C with a stability of ± 0.5 21 22 °C. The small chamber (1150 L) served as plant chamber and was connected to the larger chamber that worked as reaction chamber (1450 L, surface-area-to-volume ratio 4.87 m⁻¹). 23

24 Three small trees aging from three to four years were brought from Hyytiälä, Finland, and 25 included Scots pine (Pinus sylvestris), Norway spruce (Picea abies), and Silver birch (Betula 26 pendula). These trees, representing the main boreal forest species in Finland, were placed in 27 the JPAC plant chamber for almost two weeks before the intensive experiment campaign 28 started. This allowed them to adjust to the chamber environment. Starting from the day when 29 the trees were installed in the plant chamber, tree emissions were transferred into the reaction chamber, where O₃ was added together with water vapour directly and OH was generated 30 periodically by turning on and off the UV light, in the same fashion as applied in the intensive 31

phase. New particles were formed every day during the UV-light on periods and gases and particles deposited on the chamber walls. Discharge lamps (Osram HQI 400 W/D) were used for illumination to simulate the solar light spectrum in both chambers. Filters (OptoChem, type IR3) that reflect wavelengths between 750 and 1050 nm were used as heat shields to avoid infrared radiation inflicted plant overheating (Fig. S1 in the supplementary material shows the measured light spectrum from the discharge lamps in the reaction chamber).

The ambient air was purified by an adsorption dryer (Zander, KEA 70) and a palladium catalyst (450 °C). O_3 , NO, NO₂ and VOC levels decreased significantly after passing the purification system. The flow through the plant chamber was 115 L min⁻¹ from which a fraction of 20 L min⁻¹ was transferred to the reaction chamber. This flow was kept nearly constant by keeping the pressure drop constant. In addition, the 10 L min⁻¹ flow containing ozone was added, controlled by a second flow controller.

13 The concentrations of O₃, CO₂ and H₂O were measured by commercial analytical instruments. 14 Two Gas Chromatography Mass Spectrometer (GC-MS) systems were used, one to measure 15 the VOC concentrations from C_5 to C_{20} in the outflow air from the plant chamber (Heiden et al., 2003), and another to identify the OH concentration by determining the decrease in the 16 17 concentration of 2-butanol in the reaction chamber (Kiendler-Scharr et al., 2009). Meanwhile, the VOC concentration was continuously measured by an on-line Proton Transfer Reaction 18 19 Mass Spectrometer (PTR-MS) in the plant and reaction chamber. The gas phase H₂SO₄ 20 concentration in the reaction chamber was measured by a Chemical Ionization Mass 21 Spectrometer (CIMS) (Petäjä et al., 2009; Mauldin et al., 1998). A prototype Airmodus Particle Size Magnifier (PSM) coupled with a TSI condensation particle counter (CPC) was 22 used to count the total number concentration of particles larger than ca. 1.6 nm in diameter 23 (Vanhanen et al., 2011) and a TSI CPC (TSI3022A) was used to measure the total 24 concentration of particles larger than ca. 7 nm in diameter. A Scanning Mobility Particle Sizer 25 26 (SMPS TSI3071 + TSI3025A) was used to measure the particle size distribution in the size range of 14 – 600 nm. 27

The real plant emissions and the simulated day and night conditions make these experiments suitable for evaluation of methods used to describe the atmospheric transformation (ageing) of BVOCs and SOA beyond the first gas-phase VOC oxidation stage (e.g. the 2-Dimensional Volatility Basis Set (2D-VBS) Donahue et al., 2011).

32

1

3 Description of the ADCHAM model and its application

2

ADCHAM is a model primarily developed for simulations of laboratory chamber experiments on SOA formation and ageing. The model includes modules for reversible partitioning of organic compounds to and from the chamber walls, all fundamental aerosol dynamics processes, detailed gas- and particle-phase chemistry and a kinetic multilayer model which can be used to simulate mass transfer limited mixing of compounds in the particle phase (Fig 1). Below we describe how ADCHAM was set up in this work. For a more detailed description of ADCHAM we refer to Roldin et al. (2014).

10 **3.1 Gas-phase chemistry**

The gas-phase chemistry reactions were selected from the Master Chemical Mechanism (MCM) v3.2 (Jenkin et al., 1997, 2012; Saunders et al., 2003) via website: http://mcm.leeds.ac.uk/MCM/. The MCM is a near-explicit chemical degradation mechanism that simplifies the chemical path of compounds by lumping products beyond the second oxidation step. We used the Kinetic Pre-Processor (KPP) version 2.1 (Damian et al., 2002) to simulate the gas-phase chemistry.

Among all the compounds measured by GC-MS in the JPAC plant chamber, 28 organic 17 18 compounds were included in the gas-phase chemistry mechanism (see Table 1). In the table, 19 the "other MTs" equals to the concentration difference between the summation of the 20 concentrations of the 13 selected monoterpenes by GC-MS and the total monoterpene concentration measured by PTR-MS, while the "other SQTs" equals to the summation of 21 22 other sesquiterpene isomers besides the listed four sesquiterpenes measured by GC-MS. In 23 total, the gas-phase chemistry mechanism includes 2294 species and 6487 chemical and 24 photochemical reactions for the chemical calculations. This includes relevant inorganic 25 reactions and the full chemistry path for isoprene, α -pinene, β -pinene, limonene, β -26 caryophyllene, toluene, 2-butanol and hexanal. We also included the first-order reactions of OH, O₃, and NO₃ with the following organic compounds: myrcene, sabinene, camphene, 27 28 ocimene, Δ^3 -carene, "other MTs" (which assumed the same rate coefficients as α -pinene), 29 cineole (also known as eucalyptol), farnesene, "other SQTs" (which assumed the same rate 30 coefficients as β -caryophyllene), α -terpinene, Δ -terpinene, α -phellandrene, β -phellandrene, 31 and terpinolene. Furthermore, we included the first-order reactions between OH and the 32 following organic compounds: tricyclene, nonanal, and bornyl acetate. Finally, the first-order

reactions between O₃ and the following organic compounds were included: *α*-humulene, *α* longipinene, and Δ-cardinene were included.

3 Ocimene constituted a major fraction of MT on the first experiment day. The full chemistry 4 path for ocimene is not available in MCM. However, we included its chemistry by 5 approximating the chemistry beyond the first oxidation step with the chemistry path of 6 limonene, for which the full MCM chemistry path is available. We also approximated the 7 chemical path for Δ^3 -carene after the first oxidation step with that of α -pinene, since both are 8 bicyclic monoterpenes with an endocyclic double bond, as was done in Boy et al. (2013).

9 Based on the recent finding of rapid formation of extremely low-volatility organic compounds 10 (ELVOCs) from ozonolysis of monoterpenes containing endocyclic double bonds (Ehn et al., 2014) we also included a simplified ELVOC formation mechanism in the MCM gas-phase 11 chemistry code (R1), assuming that 7 mole % of the α -pinene + O₃ and Δ^3 -carene + O₃ 12 oxidation products were ELVOCs, with a molar mass of 325 g/mol and a vapour pressure of 13 10^{-10} Pa, which approximately corresponds to the VBS bin of $\log_{10}(C^*/\mu g m^{-3}) = -5$. We also 14 15 performed simulations where we considered that ELVOCs was formed from ozonolysis of any monoterpene or sesquiterpene, with the same ELVOC molar yield as for α -pinene. 16 MCM_{ox.prod} represents the MCMv3.2 oxidation products that was formed from the same 17 18 reactions as the ELVOCs but via a different reaction pathway.

19 VOC+ $O_3 \rightarrow 0.07$ ELVOC+0.93MCM_{ox,prod}

Ehn et al. (2014) also observed ELVOC formed from OH-oxidation of α -pinene with an estimated maximum molar yield of 1 %. In this work we evaluated the potential contribution of ELVOCs formed from OH-oxidation by using an ELVOC molar yield of 1 % for any monoterpene and sesquiterpene that reacts with OH (R2).

$$24 \quad \text{VOC+OH} \rightarrow 0.01 \text{ELVOC+} 0.99 \text{MCM}_{\text{ox,prod}} \tag{R2}$$

In order to evaluate the potential influence of specific ELVOCs (e.g. dimers) which may be involved in the nano-CN formation (here denoted ELVOC_{nucl}) we also included the possibility to add separate reactions where the monoterpene and sesquiterpene (VOC_{MT/SQT}) that react with O₃ or OH form trace amounts of ELVOC_{nucl} (R3-R4). In order to have little influence on the gas-phase chemistry or SOA formation, we used a very low ELVOC_{nucl} yield (γ_{ELVOC}) of 0.001 mole % and scaled the nucleation rate coefficient accordingly.

31
$$VOC_{MT/SQT} + O_3 \rightarrow \gamma_{ELVOC} ELVOC_{nucl} + (1 - \gamma_{ELVOC}) MCM_{ox.prod}$$
 (R3)

6

(R1)

1 $VOC_{MT/SQT} + OH \rightarrow \gamma_{ELVOC} ELVOC_{nucl} + (1 - \gamma_{ELVOC})MCM_{ox,prod}$ (R4)

2 3.2 Aerosol dynamics

3 The aerosol dynamics module in ADCHAM is based on the aerosol dynamics code from the 4 2-D Lagrangian model for Aerosol Dynamics, gas-phase CHEMistry and radiative transfer 5 (ADCHEM) (Roldin et al., 2011). It includes subroutines for nano-CN formation, 6 condensation/evaporation, Brownian coagulation and particle deposition onto the chamber 7 walls. ADCHAM simulates the condensation, dissolution and evaporation of sulphuric acid, 8 ammonia, nitric acid, hydrochloric acid and an unlimited number of organic compounds using 9 the analytic prediction of condensation scheme (Jacobson, 1997) and prediction of non-10 equilibrium growth scheme (Jacobson, 2005). The aerosol particle water content is calculated 11 with a thermodynamics model (Sect 3.4).

12 ADCHAM considers the deposition of particles onto the chamber walls and keep track of the amount of deposited material on the walls. In Roldin et al. (2014) we kept track of each 13 14 compound in each particle size bin that deposited on the chamber walls, and explicitly 15 simulated the mass-transfer-limited gas-particle partitioning between the gas-phase and the 16 wall deposited particles, assuming that the particles deposited on the walls remain as spherical particles on the walls. In this work, we instead assumed that the deposited SOA particles lose 17 18 their individual particle identity and merge into the VOC wall matrix together with the gas-19 phase VOCs that deposit directly to the chamber walls (Sect. 3.5).

20 During the JPAC experiments the particles were formed by nucleation and, as a result of 21 coagulation, wall losses and dilution, their average lifetime in the chamber was relatively 22 short (less than 45 minutes, wherein dilution generally is the dominant loss process). 23 Therefore, we expect that most of the formed particles were under charged with respect to the 24 Bolzmann charge equilibrium distribution so that we do not have to consider the enhanced 25 deposition rates induced by the particle charge (McMurry and Rader, 1985, Pierce et al., 2008 26 and Roldin et al., 2014). Thus, we calculated the first order wall deposition loss rate 27 assuming that all particles were non-changed using the indoor deposition loss rate model from 28 Lai and Nazaroff (2000), which accounts for different deposition loss rates on upward-, 29 downward- and vertically facing surfaces. In the model by Lai and Nazaroff (2000) the 30 fundamental (but unknown) parameter for the particle loss rate is the friction velocity, u*. In this work we found that with a relatively small value of u^* (0.02 m s⁻¹), the model best 31 32 captures the observed particle number and particle volume concentration loss rates after the

1 UV-light is turned off. With this value of u^* the deposition loss rate is equal to the observed 2 ELVOC molecule loss rate measured by Ehn et al. (2014) when the particle (molecule) 3 diameter is equal to 0.8 nm. We therefore used the value of $u^* = 0.02 \text{ m s}^{-1}$ for all the model 4 results presented in this work.

5 We evaluated six nano-CN formation parameterizations (Eq. 1-6) and compared them against 6 base-case simulations with a fixed new-particle formation rate (J). The first mechanism (Eq. 1) is sulphuric acid activation nucleation (Kulmala et al., 2006), Eq. 2 is used for kinetic 7 8 nucleation of two H₂SO₄ molecules (McMurry and Friedlander, 1979), Eq. 3 is used for kinetic nucleation of one H₂SO₄ molecule and one ELVOC_{nucl} Eq. 4 is the nano-CN 9 10 parameterization proposed by Riccobono et al., (2014) based on experiments in the CLOUD 11 chamber, Eq. 5 represents a mechanism were single ELVOCs serve as nano-CN (Ehn et al., 2014), and Eq. 6 represent a kinetic type of nucleation mechanism with ELVOC_{nucl}. 12

$$13 \quad J = A[\mathrm{H}_2\mathrm{SO}_4] \tag{1}$$

14
$$J = K[H_2SO_4]^2$$
 (2)

15
$$J = K[H_2SO_4][ELVOC_{nucl}]$$
(3)

16
$$J = K[H_2SO_4]^2[ELVOC_{nucl}]$$
(4)

$$17 \quad J = A[\text{ELVOC}_{\text{nucl}}] \tag{5}$$

$$18 J = K[ELVOC_{nucl}]^2 (6)$$

19 A (s⁻¹) and K (cm³ s⁻¹) in Eq 1-3 and Eq 5-6 are formation rate coefficients for activation type 20 and kinetic type of nucleation, respectively. For Eq 4 the formation rate coefficient has the 21 unit cm⁶ s⁻¹.

The composition of the nucleation clusters was chosen in order to match the respective new particle formation mechanism. The dry nano-CN volume was composed of equal mole fractions of H_2SO_4 and $ELVOC_{nucl}$ when we used Eq 1, 3 or 5, equal mole fractions of H_2SO_4 and NH_3 when we used Eq. 2, and pure $ELVOC_{nucl}$ when we used Eq. 6. The molar mass of $ELVOC_{nucl}$ was assumed to be 500 g/mol. The equilibrium water content of the nano-CN clusters was calculated with the thermodynamics model. The dry particle size of the nano-CN was assumed to be 1.5 nm.

29 Both $ELVOC_{nucl}$ and H_2SO_4 were assumed to be non-volatile. The $ELVOC_{nucl}$ first order wall

30 loss rate was 0.011 s⁻¹ according to Ehn et al. (2014). The ELVOC_{nucl} condensation sink was

modelled explicitly with the condensation algorithm in ADCHAM. Chemical degradation of
 ELVOC_{nucl} was not considered.

3 The ELVOC_{nucl} involved in the nano-CN formation were assumed to be formed 4 instantaneously after the first oxidation stage of monoterpenes and sesquiterpenes (R3-R4). 5 With Eq. 3 and 4 we investigated six different sources of $ELVOC_{nucl}$; (i) as a product formed exclusively from the ozonolysis of endocyclic monoterpenes (α -pinene and Δ^3 -carene), (ii) 6 7 from the ozonolysis of all monoterpene and sesquiterpene, (iii) from the ozonolysis of 8 sesquiterpenes, (iv) from the OH and O_3 oxidation of monoterpenes and sesquiterpenes, (v) 9 when monoterpenes and sesquiterpenes react with OH, or (vi) when sesquiterpenes react with 10 O_3 or OH. When using Eqs. 5 and 6 we only considered the ELVOC_{nucl} that was formed from 11 monoterpenes and sesquiterpenes reacting with OH.

12 **3.3 Size distribution structures**

ADCHAM can be operated with the full-moving, fixed-sections or moving-centre particle 13 14 size distribution approach (Roldin et al., 2011). In this work, we have tested both the fixed-15 sections approach and full-moving method using different numbers of size bins. The largest 16 advantage of the full-moving approach compared to the fixed-sections approach is that it does not introduce any numerical diffusion problems and allows the particles to growth into their 17 18 exact size during condensation. The main disadvantage is that new particle size bins need to 19 be introduced when new particles are formed. In this work, we handled this by only 20 introducing new particles formed by nucleation once every minute and at the same time as we 21 added a new size bin for the freshly nucleated particles, we also removed the size bin 22 containing the largest (oldest) particles. In order to not remove any particles within a time 23 frame substantially longer than their average residence time in the reaction chamber, which 24 was about 45 minutes, we used 400 size bins for the full-moving approach. Thus, with this 25 method we kept track of the nucleated particles for 400 minutes (almost 9 times the average residence time in the chamber), before they were removed from the modelled size 26 distribution. This method was compared with the fixed-section approach using a different 27 28 number of size bins in the diameter (D_n) range 1.5 to 800 nm. Based on these tests we could 29 conclude that the fixed-section approach was not suitable due to numerical diffusion if the 30 number of size bins were less than 400 (Fig. S2 in the supplementary material). With the fullmoving approach the particle volume (PV) and particle number concentrations (PN) were 31 well conserved although the last size bin was removed once every minute (Fig. S3). 32

Therefore, we decided to use the full-moving method to represent the particle number size
 distribution in this work.

One difficulty with the full-moving method is that the particle number size distribution needs to be mapped back onto a fixed diameter grid in order to illustrate it as a $dPN/dlogD_p$ distribution. This grid needs to be relatively coarse in order to not get a "jumpy" graphical representation of the particle number size distribution. For this we used 50 size bins between 1.5 and 360 nm in diameter.

8 3.4 Particle phase chemistry and phase-state

9 ADCHAM includes a detailed particle-phase chemistry module, which is used to calculate the particle equilibrium water content, the particle acidity, nitric acid and hydrochloric acid 10 11 saturation vapour pressures for each particle size bin, and the non-ideal interactions between 12 organic compounds, water and inorganic ions using the activity coefficient model AIOMFAC (Zuend et al., 2008, 2011). In this work, we did not model the specific interactions between 13 14 the organic and inorganic compounds but assumed a complete phase-separation of the 15 inorganic- and organic particle phase. We used AIOMFAC to calculate the equilibrium water 16 content in both the inorganic and organic particle phase and the individual compound activity coefficients. The organic compound activity coefficients were used when deriving the organic 17 18 compounds saturation vapour pressures above each particle size (Sect. 3.6).

19 The particle phase chemistry module also contains subroutines that can be used to calculate 20 organic salt formation, oligomerization and heterogeneous oxidation (Roldin et al., 2014). 21 Recently, Shiraiwa et al. (2013) illustrated what peroxyhemiacetal (PHA) formation between 22 organic compounds containing aldehydes and hydroperoxide functional groups may proceed 23 fast and contribute to a large and rapid increase of the formed SOA mass during photooxidation experiments. In this work we evaluate if this type of heterogeneous dimer 24 25 formation mechanism may explain the observed nano-particle growth during the JPAC experiment. For this we tested to use a constant value of the PHA formation rate (k_{PHA}) of 12 26 M^{-1} s⁻¹ adopted from Shiraiwa et al. (2013). We also tested to model k_{PHA} as a 27 parameterization of the sulphate particle mole concentration $(x_{S(VI)})$, assuming that the PHA 28 29 formation is acid catalysed by the co-condensing H_2SO_4 (Eq 7).

30
$$k_{PHA} = B \cdot x_{S(VI)} (M^{-1}s^{-1})$$
 (7)

Here *B* is a constant, the value of which we varied in the range 1-500 M^{-1} s⁻¹.

10

In this work we used the kinetic-multi-layer module in ADCHAM for investigating weather 1 2 the phase-state of the SOA particles might have influenced the evolution of the particle 3 number size distribution. In order to do this, we divided each particle into three layers (an approximately monolayer thick surface layer of 0.7 nm, and two bulk-layers). We considered 4 5 the two extreme conditions where the SOA particles either were considered to be completely 6 liquid-like (no concentration difference between the surface and bulk layers) or solid-like (no 7 molecule transport between the surface and the particle bulk layers). Still, this had only a 8 minor effect on the modelled particle growth (SOA formation), but by treating the SOA as 9 solid-like improved the agreement between the modelled and measured SOA particle volume 10 decay when the UV-light was turned off (Fig. S4). Thus, for the simulations used to produce 11 the results presented in Sect 4 we treated the SOA particles as solid-like with the assumption that the molecule transport between the particle bulk and the particle surface-layer is 12 13 relatively slow compared to the time scale it takes for the condensation to form a new 14 monolayer thick surface layer.

15 **3.5 Reversible VOC wall loss**

16 The JPAC reaction chamber was mixed with a Teflon fan with mixing times <2 min. The first 17 order VOC wall loss rate to the chamber walls (k_w) was therefore governed by the molecular 18 diffusion across the boundary layer near the chamber walls and by the uptake rate at the wall 19 surface. According to McMurry and Grosjean (1985), the first order VOC wall loss rate for 20 FEP Teflon films ($k_{w,FEP}$) can be modelled with Eq. 8, which has two key parameters; the 21 VOC wall mass accommodation coefficient (α_w) and the coefficient of eddy diffusion (k_e). Unfortunately neither α_w or k_e can be derived easily. Based on the observed wall losses of 22 particles, McMurry and Radar (1985) estimated k_e to be 0.12 s⁻¹ in a 60 m³ FEP Teflon film 23 chamber. Zhang et al. (2015) estimated k_e to be 0.075 s⁻¹ and 0.015 s⁻¹ in two not actively 24 25 mixed FEP Teflon film chambers with volumes of 24 and 28 m³.

26
$$k_{w,\text{FEP}} = \frac{A_w}{V_{chamber}} \left(\frac{\alpha_w \overline{v}/4}{1 + (\frac{\pi}{2})\alpha_w \overline{v}/(4\sqrt{(k_e D)})} \right)$$
(8)

Here A_w is the chamber wall surface area, $V_{chamber}$ is the chamber volume, \bar{v} is the mean thermal speed of the gas molecules and *D* is the molecular diffusion coefficient.

- In the JPAC reaction chamber Ehn et al. (2014) observed ELVOC first order wall loss rates in the range 0.013-0.011 s⁻¹. By inserting a value of 0.011 s⁻¹ for $k_{w,\text{FEP}}$ in Eq. 8 and assuming
- 31 that the surface wall uptake rate is not limiting the ELVOC wall loss rate ($\alpha_w > 10^{-3}$) we get a

1 k_e of 4.2 s⁻¹ for the JPAC reaction chamber. This value is substantially larger than what was 2 estimated by Zhang et al. (2015) and McMurry and Radar (1985) and is probably because the 3 JPAC reaction chamber was actively mixed and has a smaller volume.

From measurements in a 4 m³ FEP Teflon chamber Kokkola et al. (2014) observed that for 4 nopinone which has a pure-liquid saturation vapour pressure (p_0) of 53 Pa, the gas-wall 5 equilibrium was reached within a few minutes and $k_{w,\text{FEP}}$ was $\geq 0.03 \text{ s}^{-1}$. The observations by 6 7 Ehn et al. (2014) and Kokkola et al. (2014) indicate that gas-wall equilibration can be reached 8 rapidly for both volatile and low volatile VOCs and that their uptake on the chamber walls is 9 primarily limited by the transport to the chamber walls and not by surface uptake (α_w). In contrast, Zhang et al. (2014) observed that the gas-phase concentration of 25 different 10 oxidized VOCs ($p_0 = 6x10^{-6} - 20$ Pa) slowly decreased over more than 18 hours without 11 reaching gas-wall equilibrium. These experiments were performed in a 24 m³ FEP Teflon film 12 13 chamber that was not actively mixed. Based on Eq. 8, Zhang et al. (2014) concluded that the 14 VOC wall loss rate onto the FEB Teflon film walls was primarily limited by the surface uptake (α_w) and not by the molecule diffusion to the chamber walls. Zhang et al., (2014) also 15 16 derived a parameterization of α_w as a function of the compounds pure liquid saturation 17 vapour pressure.

Based on the rapid ELVOC wall losses observed by Ehn et al. (2014), we assumed that the VOC losses to the JPAC reaction chamber walls was primarily governed by the molecule diffusion to the chamber walls and used a constant k_w equal to 0.011 s⁻¹ for all condensable organic compounds. However, we also performed test simulations using Eq 8 and the α_w parameterization from Zhang et al. (2014).

According to Matsunaga and Ziemann (2010) the loss rate of VOCs from FEP Teflon chamber walls back to the gas-phase (k_a (s⁻¹)) can be represented by Eq 9,

25
$$k_{g,i} = \frac{k_w}{(RT/p_{0,i}c_w/\gamma_{w,i})}$$
 (9)

where C_w in Eq 9 is an effective mole concentration of organic compounds on the chamber walls (mol m⁻³) and $\gamma_{w,i}$ is the activity coefficient of compounds *i* in the organic film on the chamber walls, *T* is the temperature in Kelvin and *R* is the universal gas constant (J mol⁻¹ K⁻¹). According to Matsunaga and Ziemann (2010), the FEP Teflon film walls serve as a large organic pool where VOCs can absorb ($C_w/\gamma_{w,i}$ on the order of 100 µmol m⁻³). Thus, the actual VOCs deposited on the chamber walls is considered to have a negligible contribution to C_w .

We assumed that there was practically no C_w on the glass for the VOCs to dissolve into. We 1 2 also estimated that even if the deposited VOCs and SOA were distributed as a molecule 3 monolayer on the walls, the VOCs formed during the simulated experiment could only cover 4 maximum 2 % of the total glass wall surface area in the end of the measurement campaign 5 (see supplementary material). Thus, we assumed that the VOC losses onto the glass wall 6 surfaces could be treated as a condensation process but without dissolution (absorption) into 7 an organic matrix on the walls. With this approach the uptake of compound *i* is governed by 8 the difference between the concentration in the gas-phase $(c_{g,i})$ and the pure liquid saturation 9 concentration $(c_{0,i})$ (Eq 10-11). Thus, as long as $c_{g,i} < c_{0,i}$ the VOC will not condense on the glass walls. For many of the semi-volatility organic compounds (SVOCs), the limit at which 10 $c_{g,i}$ become larger than $c_{o,i}$ was never reached during the experiments. 11

12
$$\frac{dc_{g,i}}{dt} = -k_w (c_{g,i} - c_{0,i}) \qquad (if \ c_{g,i} > c_{0,i} \ or \ if \ c_{w,i} > 0)$$
(10)

13
$$\frac{dc_{w,i}}{dt} = k_w (c_{g,i} - c_{0,i})$$
(11)

For the PTFE Teflon floor, we used the theory developed for FEP Teflon films (Eq. 9, 12 and 13), assuming that the PTFE Teflon surface (in total 15 % of the total chamber surface area) has a $C_w/\gamma_{w,i} = 100 \ \mu\text{mol m}^{-3}$ and $k_{w,i} = 1.7 \times 10^{-3} \text{ s}^{-1} (0.15 \times 0.011 \text{ s}^{-1})$.

17
$$\frac{dc_{g,i}}{dt} = -k_{w,i}c_{g,i} + k_{g,i}c_{w,i}$$
(12)

18
$$\frac{dc_{w,i}}{dt} = k_{w,i}c_{g,i} - k_{g,i}c_{w,i}$$
(13)

In order to mimic the actual experiment procedure (Sect 2), we started the model simulations by running ADCHAM 14 days prior to the actual experiments started, with conditions similar to Day 1 of the experiment campaign (see Sect 3.7). This allowed the VOC concentration to build up on the chamber walls.

23 **3.6 Condensable organic compound properties**

In this work, we used two different approaches to simulate the SOA formation. In the first, the SOA formation was modelled by considering the gas-particle partitioning of all non-radical organic compounds from the MCMv3.2 gas-phase chemistry code with estimated pure-liquid saturation vapour pressure lower than 10^{-2} Pa at T=289 K (in total 488 compounds). The pureliquid saturation vapour pressures of these compounds were estimated using the boiling point and vapour pressure extrapolation method from Nannoolal et al. (2004; 2008), here after 1 referred to as the Nannooal method. MCMv3.2 only includes one sesquiterpene: β -2 caryophyllene. As an attempt to take into account the SOA formation from the other 3 sesquiterpens, for these simulations all sesquiterpenes were treated as β -caryophyllene.

The second method is the 2D-VBS approach (Jimenez et al., 2009; Donahue et al., 2011), which is a simplified approach to describe how the volatility distribution of the VOCs (as a function of their Oxygen to Carbon ratio (O:C)) evolves beyond the first oxidation step. A detailed description of the structure and general assumptions of the 2D-VBS method used in ADCHAM is given by Hermansson et al. (2014). Here we primarily describe how the 2D-VBS was modified in order to better capture the SOA formation observed during the JPAC experiments.

11 The 2D-VBS compounds were distributed across 11 volatility classes separated by powers of 10 in saturation concentration (at the standard temperature 298 K (C_{298}^*)), ranging from 10⁻⁷ to 12 10³ µg m⁻³, and 12 discrete O:C from 0.1 to 1.2, in steps of 0.1. VOCs were introduced into 13 14 the 2D-VBS by scaling the amount of reacted monoterpenes and sesquiterpenes with stoichiometric VBS yield parameterizations that represent the volatility distribution of the 15 16 first generation oxidation products. We used the parameterizations from Table 3 in Henry et al. (2012). These parameterizations represent the first generation product volatility 17 18 distributions formed during experiments where α -pinene was primarily oxidized by ozone (no 19 OH-scavenger) or where α -pinene was oxidized by OH. We used the former parameterization 20 to represent the first generation terpene (monoterpenes and sesquiterpenes) oxidation products during the UV-off periods and the later for the UV-light on periods. ELVOCs formed as first 21 generation oxidation products from R1 were introduced into the 2D-VBS assuming that they 22 had a O:C of 1 and $C_{298}^* = 10^{-7} \,\mu\text{g m}^{-3}$. 23

Hermansson et al. (2014) illustrated that the modelled SOA formation is sensitive to the volatility distribution and the assumed O:C of the first generation oxidation products. Moreover it depends on how the functionalization and fragmentation pattern is represented in the VBS and the reaction rate constant between the VOCs and OH. Equation 14 from Jimenez et al. (2009) was used to represent how the fraction of VOCs which fragmentize upon oxidation varies as a function of their O:C.

$$30 \qquad f_{\text{frag}} = (O:C)^{\text{y}} \tag{14}$$

31 Jimenez et al. (2009) assumed that y was equal to 1/6, that the 2D-VBS compounds were

oxidized by OH with a reaction rate coefficient (k_{OH}) of 3×10^{-11} cm³ s⁻¹, and that the 1 2 compounds that fragmentize had equal probabilities to split at any of the carbon bonds. 3 However, with this 2D-VBS setup, ADCHAM substantially overestimated the SOA formation (Fig. S5). This was mainly because a substantial fraction of the 2D-VBS fragmentation 4 5 products were still less volatile than the reacting VOCs. If we instead assumed that all 6 fragmenting compounds (on a mole fraction basis with equal proportions) end up into the VBS bins where C^*_{298} is at least 3 order of magnitude larger than the corresponding 7 8 functionalization products, and decreased the probability of fragmentation somewhat (y=1/3), 9 the modelled particle volume concentration agreed better with the measurements (Fig. S6). Thus, in the proceeding sections we used this fragmentation parameterization. We also 10 performed sensitivity tests where we varied the value of k_{OH} in the range of 3×10^{-11} to 5×10^{-11} 11 $cm^3 s^{-1}$ and the O:C of the first generation oxidation products in the range of 0.3 to 0.5 (Fig. 12 S6). Based on these tests, we decided to use the values of $k_{OH} = 5 \times 10^{-11} \text{ cm}^3 \text{ s}^{-1}$ and O:C = 0.4 13 for the first generation oxidation products in the simulations used to produce the results in 14 15 Sect. 4.

16 The temperature dependence of the 2D-VBS saturation concentrations were derived using the 17 Clausius Clapeyron equation and assuming the following form for the heat of vaporization 18 $(\Delta H;$ Epstein et al., 2010):

19
$$\Delta H = -11 \times \log_{10} C_{300}^* + 129 \text{ kJ mol}^{-1}$$
 (15)

Both with the 2D-VBS and when the MCM compounds were used to represent the condensable organic compounds, the Raoult's law and the Kelvin equation were used to calculate the saturation vapour pressure $(p_{s,i,j})$ for each compound (i) above each particle size (j) (Eq 16). With the 2D-VBS we assumed that the organic compounds were mixed ideally (unity activity coefficients (γ_i)). However, for the MCM compounds the molecule structure is known, so we calculated the organic compound activity coefficients in the SOA particle mixtures using AIOMFAC.

27
$$p_{s,i,j} = p_{0,i} x_{i,j} \gamma_{i,j} e^{\left(\frac{4M_i \sigma_i}{RT \rho_j D_{p,j}}\right)}$$
(16)

where $x_{i,j}$ is the mole fraction of compound *i* in the particle surface layer of particles in size bin *j*, M_i is the molar mass of compound *i* and ρ is the density of the absorbing phase. In this work the surface tension of the organic compounds (σ_i) was estimated to be equal to 0.05 N 1 m⁻¹ following Riipinen et al. (2010), even though we also tested the values of σ_i in the range 2 0.02 to 0.07 N m⁻¹. These sensitivity tests revealed that the modelled total particle number 3 concentration decreased by approximately 10 % and the total particle volume concentration 4 by approximately 30 % when the value of σ_i was increased from 0.02 to 0.07 N m⁻¹ (Fig. S7).

5 3.7 Optimizing model input parameters

6 We used the ADCHAM model to simulate the new particle formation and growth during one 7 measurement campaign conducted in the JPAC chamber. We chose to study the first four 8 continuous days from the measurement campaign because these days had continuous 9 measurements, coverage of the dataset was complete and UV light-on period was the same.

10 The measured temperature and relative humidity (RH) in the JPAC reaction chamber were 11 used directly in the model as input. The values from the first four days are illustrated in Fig. 2. 12 The purple bars in the figures illustrate the UV-on periods. The temperature stabilized around noon on Day-1, after that the temperature was ca. 16 °C for all days. The RH had minor 13 14 fluctuations during the experiment. RH was kept to ca. 60% during UV-on periods for the 15 first three days. On Day-4, the RH was around 55 %. In addition, 8 discharge lamps were 16 used to simulate solar illumination on Day-1 in the plant chamber, but only 4 lamps were used 17 on the remaining days to generate lower levels of VOC emissions from the trees.

18 Because the inflow of ambient air into the JPAC chamber was purified by an adsorption 19 dryer, by default, the concentrations of NO, NO₂, SO₂, and CO in the inflow to the reaction 20 chamber were assumed to be 0.03, 0.2, 0.015 and 15 ppbv, respectively. However, in order to 21 evaluate how sensitive the model results was, in particular the modelled VOC composition and SOA formation, we also performed sensitivity tests where we varied the inflow 22 23 concentrations of NO_x (NO+NO₂) within the range of 0.05-1 ppbv. Concentrations of VOCs 24 in the inflow to the reaction chamber were based on the GC-MS and PTR-MS data. The initial 25 concentrations of all VOCs were set to zero. The measured concentrations of isoprene, 26 monoterpenes and sesquiterpenes (in the plant chamber) are illustrated in Fig. 3. The 27 monoterpene and sesquiterpene concentrations displayed a certain diurnal distribution pattern, 28 because the discharge lamps mimicked sunlight in the plant chamber. The total monoterpene 29 concentration exceeded 4 ppbv on Day-1, and then decreased to a value below 2 ppbv on the 30 following three days. In particular, the ocimene concentration was the highest on Day-1, 31 which can be explained by the fact that ocimene emission is light dependent (Owen et al., 32 2002), and the simulated solar light intensity in the plant chamber from the discharge lamps 1 was two times higher on Day-1 than during Day 2-4.

The isoprene concentration and the total sesquiterpene concentration were about an order of magnitude lower than the monoterpene concentration. Overall, the measured concentrations of terpenes were at similar levels as ambient air concentrations in boreal forest environments (Ruuskanen et al., 2009).

At the first attempt, we tried to simulate the O₃ and OH concentrations with the gas-phase chemistry mechanism. An OH sink has been determined in the empty chamber by direct OH-LIF measurements (Diss. Sebastian Broch) and has been always found to be within 2 - 4 s⁻¹. However, even with this OH sink, the model could not capture the observed gradual decrease in the O₃ concentration and increase in OH concentration during the UV-light on periods without either varying the UV-light intensity or the OH sink.

12 Therefore, we decided to use the measured O₃ and H₂SO₄ concentrations, and estimated OH 13 concentration, as input to the model. Figure 4a shows the measured O₃ concentration during 14 the experiment, Fig. 4b shows the estimated OH concentration based on the observed loss rate 15 of 2-butanol and Fig. 4c shows the measured H₂SO₄ concentration. When the UV-light was 16 turned on the H₂SO₄ concentration increased very rapidly and had a distinct peak. This is most likely due to an initial peak in the OH concentration at the moment when the UV-lamps 17 was turned on. In the model we represented this by setting the OH concentration to 5×10^7 18 molecules cm⁻³ for the 12 first minutes after the UV-light were turned on. 19

20 The photolysis rates were simulated using the quantum yields and absorption cross sections 21 reported at the MCMv3.2 web site. The light spectrum in the reaction chamber was estimated using the measured spectrum for the discharge lamps in the wavelength range 280-650 nm 22 23 (Fig. S1) and with a single UV-light peak at a wavelength of 254 nm, which represents the UV-spectrum from the Philips, TUV 40W lamp. The UV-light source intensity corresponds to 24 a $O(^{1}D)$ photolysis rate of 2.9x10⁻³ s⁻¹ (Mentel et al., 2009). In order to evaluate the potential 25 influence of the 254 nm wavelength UV-light on the VOC composition and SOA formation 26 27 we performed a model simulation where the 254 nm UV-light was not considered when 28 calculating the photolysis rates. This test showed that the modelled condensable organic 29 composition and the SOA formation were not noticeable influenced by the 254 nm UV-light 30 source.

31

- 1
- 2

4 Results and discussion

Table 2 summarizes the different model tests that we performed in this work in order to constrain the VOC wall losses, the aerosol dynamics-, gas- and particle-phase chemistry mechanisms that can explain the nano-CN formation and growth observed during the JPAC experiments presented in Sect 2.

7 4.1 Time series of BVOC concentrations

8 To investigate the potential contribution of BVOCs to the nano-CN formation and growth, it 9 is essential to properly predict the time series of the BVOC concentrations. The modelled 10 isoprene, monoterpene, and sesquiterpene concentrations in the JPAC reaction chamber are plotted in Fig. 5, together with isoprene and total monoterpene concentrations measured by 11 the PTR-MS. The modelled isoprene concentrations were in a good agreement with the 12 13 measurements during the UV-off period on Day-1. During the rest of the simulated four days 14 period, the simulated isoprene concentration was generally lower than the observations. 15 However, considering the low isoprene concentrations and the uncertainties in the PTR-MS 16 measurements, we cannot draw any conclusions on weather the model actually underestimates 17 the isoprene concentration or not.

18 The summation of the modelled monoterpene isomer concentrations reached the same level as 19 the measurements. During the dark periods, the monoterpene concentrations in the reaction 20 chamber decreased to about one third of their concentrations measured in the plant chamber. 21 This was because of the in-flow dilution and chemical reactions with O₃. When the UV light 22 was switched on, the monoterpene concentration decreased sharply due to OH oxidation, as 23 seen both from the measurements and simulation results in Fig. 5. Before the UV-lights were 24 turned on ~ 80 % of the reacting monoterpenes were oxidized by O₃. At the UV onset this number dropped to ~ 10 % and in the end of the UV-on periods only ~ 2 % of the oxidized 25 26 monoterpenes were oxidized by O₃. The modelled monoterpene concentration shows a somewhat more pronounced decrease during the UV-on periods than measurement. However, 27 28 as with the isoprene concentration, the relative uncertainties in the PTR-MS measurements 29 increases at lower concentrations and part of the PTR-MS unity resolution mass peak that was 30 interpreted as monoterpenes may also have had small contributions from other VOCs.

31 The sesquiterpene concentrations were below the detection limit of the PTR-MS. Our

modelled results show that the sesquiterpenes were strongly oxidized by O_3 once the tree emission from the plant chamber entered the reaction chamber (Fig. 3). Even when the UVlights were turned on, 40 to 60 % of the sesquiterpenes were oxidized by O_3 and the rest by OH. The sesquiterpene concentrations decreased rapidly to below 0.01 ppbv. In addition, the sesquiterpene concentrations showed a minor decrease during UV-on periods because of the oxidation by OH.

Although a large fraction of the monoterpenes and sesquiterpenes were oxidized in the JPAC reaction chamber, nano-CN formation was only observed during the UV-on periods. This indicates that the terpene oxidation products formed during the dark periods (primarily by ozonolysis) either did not have the right properties or were not abundant enough to form, activate and/or grow nano-CN to sizes above the detection limit of the PSM-CPC setup.

12 **4.2** Simulations of the observed new particle formation events

13 The first simulations were designed in order to constrain the VOC wall losses and the mechanisms responsible for the observed particle growth (Sect 4.2.1 and Sect 4.2.2). For 14 these simulations, we used a fixed nano-CN formation ratio of 20 cm⁻³ s⁻¹ during the UV-light 15 on periods, except for the first 12 minutes with UV-light on for which we used $J = 80 \text{ cm}^{-3} \text{ s}^{-1}$, 16 for Day-1 and Day-2 and 60 cm⁻³ s⁻¹ for Day-3. During the UV-light off periods the nano-CN 17 18 formation rate was zero. The nano-CN had a dry diameter of 1.5 nm and was assumed to be 19 composed of equal number of sulphuric acid and ELVOC_{nucl} molecules. After, constraining 20 the VOC wall losses and the potential particle growth mechanisms, we investigated several 21 nano-CN formation mechanisms (Eq. 1-6) and compared the results against the results from 22 the simulations with a fixed nano-CN formation rate (Sect. 4.2.3).

4.2.1 Modelling the reversible VOC wall deposition

24 Figure 6a shows the modelled total particle volume concentration (PV) starting 14 days prior 25 to the start of the intensive experimental campaign. The measured PV during the campaign is 26 also displayed. The model results are from a simulation with the 2D-VBS. Figure 6b and c illustrate how the 2D-VBS compounds with different C^* were distributed onto the glass walls 27 and the PTFE Teflon floor, respectively. The LVOCs ($C_{298}^* < 10^{-1} \ \mu g \ m^{-3}$) deposited both onto 28 the glass and Teflon surfaces, while the SVOCs ($C_{298}^* > 10^{-1} \,\mu g \,m^{-3}$) were exclusively found at 29 the Teflon floor where they could absorb into C_{w} . Some of the compounds with $C_{298}^{*} = 10^{-1} \,\mu g$ 30 m⁻³ deposited onto the glass walls shortly after the UV-light was turned on, but when the 31

1 condensation sink was increased and when the terpene inflow concentration into the reaction 2 chamber was lowered (Day 2-4), the gas-phase concentrations of these compounds never 3 reached above their pure liquid saturation concentrations, so they evaporated from the glass 4 walls. Figure S9 in the supplementary material shows similar results but for a simulation 5 where the MCMv3.2 compounds and additional ELVOCs (R1) were used to represent the 6 condensable organic compounds.

7 Figure 7 shows the 2D-VBS VOC composition for: (a) the SOA particles, (b) the gas-phase, 8 (c) the VOCs on the glass walls and (d) the VOCs on the PTEF Teflon floor, at Day-3 and after 5 hours with UV-lights on. The bar at O:C = 1, $\log_{10} C_{298}^*$ =-7 corresponds to the 9 ELVOCs which were assumed to be formed from ozonolysis and OH-oxidation of 10 11 monoterpenes and sesquiterpenes (molar yield of 7 % and 1 %, respectively). After 16.5 days of continuous experiments, 1.4 mg m⁻³ of SVOCs had deposited onto the PTEF Teflon floor 12 13 according to the model simulations. The largest fraction of the wall deposited VOCs were first generation terpene oxidation products (O:C=0.4). A fraction of these SVOCs can re-14 15 evaporated and reacted with OH in the gas-phase. This explains why the modelled SOA formation was gradually increasing during the first 15 days (Fig. 6a), although all the other 16 17 model conditions were identical. A similar pattern in the modelled SOA particle volume was also observed when the SOA formation was simulated with the MCM compounds (Fig. S8a). 18 19 The ELVOCs (that are formed as first generation oxidation products (R1)) and the LVOCs 20 (formed from OH-oxidation of the first generation oxidation products), primarily deposited 21 onto the glass walls. However, in total the VOCs deposited onto the glass walls only made up 5 % of the total amount of VOCs on the chamber walls. 22

23 We also performed simulations assuming that the JPAC reaction chamber walls behave as FEP Teflon walls (Eq. 8), in which case $k_e = 4.2 \text{ s}^{-1}$ and the a_w parameterization was taken 24 from Zhang et al. (2014). For compounds with a molar mass of 300 g mol^{-1} and vapour 25 pressures in the range 10⁻² to 10⁻¹⁰ Pa the $k_{w,i}$ varies from 2×10⁻⁵ to 7×10⁻⁴ s⁻¹. Thus, for these 26 27 model simulations the ELVOC wall losses were about 15 times lower than what was observed by Ehn et al. (2014). Fig. S9 compares the measured total particle volume concentration with 28 29 the modelled total particle volume concentration from a simulation with this wall loss parameterization. With the lower VOC wall losses, the model overestimated the SOA 30 31 formation by a factor of 2-3 and there was no gradual increase in the SOA formation due to 32 re-evaporation of SVOCs from the walls during the days before the intensive measurement 1 campaign started.

2 **4.2.2** Evaluation of potential particle growth mechanisms

3 When using the 2D-VBS, the modelled SOA composition was dominated by LVOCs and 4 SVOCs formed from second- and multi-generation OH-oxidation products. This was the case 5 even if we considered that ELVOCs were formed as first generation products after the O₃-6 and OH-oxidation of monoterpenes and sesquiterpenes (Fig. 7a). The reason for this is the large ELVOC wall losses in the JPAC chamber and the small condensation sink during the 7 8 new-particle formation events. During the UV-light on periods, the gas-phase 2D-VBS VOC 9 composition was dominated by oxidized SVOCs formed by fragmentation of the first 10 generation oxidation products (Fig. 7b).

11 In Fig. 8 we compare the modelled (a) total particle number concentration and (b) total particle volume concentration with the observations from the PSM-CPC and the SMPS. The 12 model results are from simulations with the 2D-VBS with or without ELVOC formation from 13 ozonolysis and OH-oxidation of all monoterpenes and sesquiterpenes (molar yield of 7 % and 14 15 1 %, respectively), as well as from a simulation where the condensable organic compounds were represented by the MCMv3.2 compounds (in total 488 compounds with $p_0 < 10^{-2}$ Pa, 16 including ELVOCs from ozonolysis of α -pinene and Δ^3 -carene). For this simulation we also 17 18 considered rapid peroxyhemiacetal dimer formation in the particle phase using Eq. 7 and B =200 M⁻¹ s⁻¹. Without consideration of this type of a rapid acid catalysed dimer formation 19 20 process, the particle growth was substantially underestimated and almost no SOA was formed 21 when we simulated the SOA formation using the MCM compounds (Fig. S8a). Shown are also the results from a simulation with only one non-volatile condensable organic compound. 22 In the model, this compound was formed as a first generation oxidation product from O₃ and 23 OH oxidation of all monoterpenes and sesquiterpenes with a mass yield of 60 % (molar yield 24 25 of 25 %). Figure 9 shows results from the same simulations as in Fig 8, but here we instead compare the estimated SOA particle mass yields from the model with those derived from the 26 27 measurements. Both in the model and for the measurements the SOA particle density was assumed to be 1400 kg m³. The SOA mass yields were estimated by dividing the total SOA 28 29 particle mass with the amount of reacted terpenes (not including isoprene). The amount of 30 reacted terpenes was estimated as the difference between the in and outflow terpene concentrations. 31

32 For all model simulations, the model overestimated the SOA particle volume formation and

1 SOA mass vield during Dav-1 but tended to underestimate or give similar values as the 2 measurements for Day 2-4. The most likely explanation to this is that the BVOC composition 3 was substantially different during Day-1 (ocimene which reacts rapidly with O₃ may not form 4 SOA in the same extent as e.g. α -pinene). The best agreement between the model and measured particle volume concentration was found with the 2D-VBS method ($R^2 = 0.836$ 5 with ELVOCs formation and $R^2 = 0.835$ without ELVOC formation). For the simulation with 6 the non-volatile one product model, $R^2 = 0.796$, and for the simulation with MCMv3.2 and 7 acid catalysed PHA dimer formation, $R^2 = 0.820$. 8

9 Figure S10 in the supplementary material compares the modelled total particle volume 10 concentrations from simulations with B = 10 or 200, or using a constant PHA dimer formation 11 rate of 12 M⁻¹ s⁻¹ based on the work by Shiraiwa et al. (2013). With B = 10 the modelled PV 12 are in very good agreement with the measured PV for Day-1 but for the following days the 13 model substantially underestimated the PV. With a constant PHA dimer formation rate of 12 14 M⁻¹ s⁻¹ the model gives 3 times higher PV than the measurements for Day-1 but gives 15 reasonable PV formation for Day 2-4.

Figure 10 shows the modelled SOA volatility distribution as a function of particle size. The results are from a simulation with the 2D-VBS, including ELVOC formation from monoterpenes and sesquiterpens oxidized by O₃ and OH. As expected, the smallest particles contained the largest mole fraction of ELVOCs because of the Kelvin effect.

20 Figure 11 shows the modelled and measured particle number size distributions at $\frac{1}{2}$, 1, 2 and 21 5 hours of UV-lights on, for each day of the experimental campaign. The model results are 22 from a simulation with the 2D-VBS including ELVOC formation from monoterpenes and 23 sesquiterpenes oxidized by O₃ and OH. For Day-1, the model overestimated the particle 24 growth rate, which can also be seen from the overestimated SOA formation (Fig. 8b). For the 25 reminder of the experimental campaign, the modelled particle number size distributions were 26 in good agreement with the SMPS measurements, except for the particles having a diameter <30 nm. for which where the model gave substantially higher concentrations. A contributing 27 explanation for this feature can be non-accounted diffusion losses of particles in the SMPS 28 29 inlet.

30 In the supplementary we show how the volatility distribution of the MCM compounds and the 31 SOA formation changed when the NO_x concentration in the inflow to the chamber was 32 changed in the range 0.05 to 1 ppbv (Fig. S11). When the NO_x concentration was increased to 1 from 0.05 to 1 ppbv the particle SOA volume concentration was increased slightly (~10 %).

2 4.2.3 Evaluation of potential nano-CN formation mechanisms

3 In this section we evaluate the different nano-CN formation mechanisms described in Sect 4 3.2. For these simulations we used the 2D-VBS to simulate the evolution of the condensable 5 organic compounds. Because the exact vapour pressures, formation mechanisms and 6 concentrations of ELVOCs still are very uncertain, we cannot dismiss the possibility that the 7 new particle formation (formation of particles with $D_p > 1.6$ nm) was limited by the activation 8 of nano-CN and not by the nano-CN formation rate itself. In this section we evaluate possible 9 nano-CN formation mechanism with the assumption that it was the nano-CN formation that 10 primarily limited the observed new particle formation during the experiments.

Table 3 gives the coefficient of determination (R^2) between the modelled total particle 11 number concentration and the measured total particle number concentration ($D_p > 1.6$ nm) 12 from the PSM-CPC setup, and for the nano-CN mechanisms (parameterizations) that gave a 13 $R^2 > 0.85$. The best agreement ($R^2 > 0.97$) between the modelled and measured total particle 14 15 number concentration was achieved with a nano-CN formation mechanism that involves both 16 H₂SO₄ and ELVOCs formed from OH-oxidation or OH and O₃ oxidation of monoterpenes and sesquiterpenes (Eq. 3) or from Eq. 4 if the ELVOC_{nucl} molecules exclusively were formed 17 from ozonolysis of sesquiterpenes. These are the only mechanisms for which R^2 were higher 18 19 than with a fixed nano-CN formation rate during the UV-light on periods.

In Fig. 12 we compare the modelled total particle number concentration when calculating Jusing Eq. 1, 2 and 3. With Eq. 3, the ELVOC_{nucl} were either assumed to be formed from ozonolysis and OH-oxidation of monoterpenes and sesquiterpens or only from OH-oxidation. Displayed are also the results from simulations with constant J = 20 cm⁻³ s⁻¹ during the UVlight on periods and J = 0 cm⁻³ s⁻¹ during the UV-light of periods.

25

26 5 Summary and conclusion

27

In this study, we used the ADCHAM model to simulate the nano-CN formation and growth during an experimental campaign in the Jülich Plant Atmosphere Chamber, which focused on new particle formation induced by photochemical reactions of VOCs emitted from real plants (Dal Maso et al., 2014). With the model we evaluated potential nano-CN formation and growth mechanisms and how the VOC chamber wall losses influenced the SOA formation
 and composition.

3 Chamber wall losses can have a profound influence on the SOA formation. In this work we 4 showed that the contribution of ELVOCs to the nano-CN formation and growth was 5 effectively suppressed due to their rapid and irreversible wall losses. Thus, it is questionable 6 whether this type of smog chamber experiments is ideal for the evaluation of possible 7 mechanisms responsible for the observed new particle formation in the atmosphere.

8 In the lack of ELVOCs, one possible mechanism that can explain the initial particle growth is 9 rapid heterogeneous dimer formation of SVOCs. In this work, we found out that acid 10 catalysed peroxyhemiacetal formation between aldehydes and hydroperoxides may explain 11 the observed particle growth. However, the particle growth could be modelled equally well if 12 the SOA was formed by condensation of low-volatility second- and multi-generation OH gas-13 phase oxidation products simulated with a 2D-VBS approach. Thus, based on our model 14 simulations, we cannot conclude weather the observed particle growth primarily was driven 15 by low-volatility organic compounds formed in the gas-phase or by rapid dimer formation in the particle phase. However, without the later mechanism the model fails to reproduce the 16 17 observed SOA formation when using the MCMv3.2 oxidation products as condensable 18 organic compounds.

19 Our results suggest that H_2SO_4 is one of the key compounds involved in the new particle 20 formation, but cannot solely explain the new particle formation process. During the simulated 21 experiments, the best agreement between the modelled and measured total particle number 22 concentration was achieved when using a nano-CN formation rate of the form J =23 $K[H_2SO_4][ELVOC_{nucl}]$.

24

25 Acknowledgements

P. Roldin would like to thank the Cryosphere-Atmosphere Interactions in a Changing Arctic
Climate (CRAICC) and the Swedish Research Council for Environment, Agricultural
Sciences and Spatial Planning FORMAS (Project No. 214-2014-1445) for financial support.
L. Liao wishes to thank the Maj and Tor Nessling foundation for financial support (grant No
2009362), as well as the Academy of Finland (project No. 128731). D. Mogensen would like
to thank the doctoral program in Atmospheric Sciences (ATM-DP) at the University of
Helsinki for financial support. We would like to acknowledge HENVI (Helsinki University

1 Centre for Environment), The FCoE (The Centre of Excellence in Atmospheric Science -2 From Molecular and Biological processes to the Global Climate (ATM)), the strategic 3 research area MERGE (Modelling the Regional and Global Earth system) and the PEGASOS 4 (Pan-European Gas-Aerosolsclimate interaction Study, project No FP7-ENV-2010-265148) 5 project. We would also like to thank Prof. Gordon McFiggans' research group at the 6 University of Manchester, and especially Dr. David Topping, for helpful discussions and for providing the Python script (now a publicly available function called Comp-SysProp: 7 8 http://ratty.cas.manchester.ac.uk/informatics/) to calculate Nannoolal-based sub-cooled liquid 9 saturation vapour pressures for all organic compounds included in this paper.

10

11 References

- 12 Almeida, J., Schobesberger, S., Kürten, A., Ortega, I. K., Kupiainen-Määttä, O., Praplan, A.
- 13 P., Adamov, A., Amorim, A., Bianchi, F., and Breitenlechner, M.: Molecular understanding
- 14 of sulphuric acid-amine particle nucleation in the atmosphere, Nature, 502, 359–363, 2013.
- Berndt, T., Stratmann, F., Sipilä, M., Vanhanen, J., Petäjä, T., Mikkilä, J., Grüner, A.,
 Spindler, G., Lee Mauldin III, R., Curtius, J., Kulmala, M., and Heintzenberg, J.: Laboratory
 study on new particle formation from the reaction OH+SO2: influence of experimental
 conditions, H2O vapour, NH3 and the amine tert-butylamine on the overall process, Atmos.
- 19 Chem. Phys., 10, 7101–7116, doi:10.5194/acp-10-7101-2010, 2010.
- Bonn, B. and Moortgat, G. K.: Sesquiterpene ozonolysis: origin of atmospheric new particle
 formation from biogenic hydrocarbons, Geophys. Res. Lett., 30, 1585,
 doi:10.1029/2003GL017000, 2003.
- Boy, M., Rannik, Ü., Lehtinen, K. E. J., Tarvainen, V., Hakola, H., and Kulmala, M.:
 Nucleation events in the continental boundary layer: long-term statistical analyses of aerosol
 relevant characteristics, J. Geophys. Res., 108, 4667, doi:10.1029/2003JD003838, 2003.
- Boy, M., Mogensen, D., Smolander, S., Zhou, L., Nieminen, T., Paasonen, P., Plass-Dülmer,
 C., Sipilä, M., Petäjä, T., Mauldin, L., Berresheim, H., and Kulmala, M.: Oxidation of SO2
- 28 by stabilized Criegee intermediate (sCI) radicals as a crucial source for atmospheric sulfuric
- 29 acid concentrations, Atmos. Chem. Phys., 13, 3865–3879, doi:10.5194/acp-13-3865-2013,
- 30 2013.
- 31 Claeys, M., Graham, B., Vas, G., Wang, W., Vermeylen, R., Pashynska, V., Cafmeyer, J.,

- 1 Guyon, P., Andreae, M. O., Artaxo, P., and Maenhaut, W.: Formation of secondary organic
- 2 aerosols through photooxidation of isoprene, Science, 303, 1173–1176, 2004. 27975 Crump,
- 3 J. G. and Seinfeld, J. H.: Turbulent deposition and gravitational sedimentation of an aerosol in
- 4 a vessel of arbitrary shape, J. Aerosol Sci., 12, 405–415, 1981.
- 5 Dal Maso, M., Liao, L., Wildt, J., Kiendler-Scharr, A., Kleist, E., Tillmann, R., Sipilä, M.,
 6 Hakala, J., Lehtipalo, K., Ehn, M., Kerminen, V.-M., Kulmala, M., Worsnop, D., and
- 7 Mentel, T.: A chamber study of the influence of boreal BVOC emissions and sulphuric acid
- 8 on nanoparticle formation rates at ambient concentrations, Atmos. Chem. Phys. Discuss., 14,
- **9** 31319-31360, doi:10.5194/acpd-14-31319-2014, 2014.
- 10 Damian, V., Sandu, A., Damian, M., Potra, F., and Carmichael, G. R.: The kinetic
- 11 preprocessor KPP-a software environment for solving chemical kinetics, Comput. Chem.
- 12 Eng., 26, 1567–1579, 2002.
- Donahue, N. M., Epstein, S. A., Pandis, S. N., and Robinson, A. L.: A two-dimensional
 volatility basis set: 1. organic-aerosol mixing thermodynamics, Atmos. Chem. Phys., 11,
 3303–3318, doi:10.5194/acp-11-3303-2011, 2011.
- 16 Ehn, M., Thornton, J. A., Kleist, E., Sipilä, M., Junninen, H., Pullinen, I., Springer, M.,
- 17 Rubach, F., Tillmann, R., Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I.-H., Rissanen, M.,
- 18 Jokinen, T., Schobesberger, S., Kangasluoma, J., Kontkanen, J., Nieminen, T., Kurtén, T.,
- 19 Nielsen, L. B., Jorgensen, S., Kjaergaard, H. G., Canagaratna, M., Maso, M. D., Berndt, T.,
- 20 Petäjä, T., Wahner, A., Kerminen, V.-M., Kulmala, M., Worsnop, D. R., Wildt, J., and
- Mentel, T. F.: A large source of low volatility secondary organic aerosol, Nature, 506, 476–
 479, 2014.
- Epstein, S., Riipinen, I., and Donahue, N. M.: A Semiempirical Correlation between Enthalpy
 of Vaporization and Saturation Concentration for Organic Aerosol, Environ. Sci. Technol.,
- 25 44, 743–748, 2010.
- 26 Guenther, A., Hewitt, C. N., Erickson, D., Fall, R., Geron, C., Graedel, T., Harley, P.,
- 27 Klinger, L., Lerdau, M., McKay, W. A., Pierce, T., Scholes, B., Steinbrecher, R., Tallamraju,
- 28 R., Taylor, J., and Zimmerman, P.: A global-model of natural volatile organic-compound
- 29 emissions, J. Geophys. Res., 100, 8873–8892, doi:10.1029/94JD02950, 1995.
- 30 Hallquist, M., Wenger, 5 J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M.,
- 31 Dommen, J., Donahue, N. M., George, C., Goldstein, A. H., Hamilton, J. F., Herrmann, H.,

- Hoffmann, T., Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr, A.,
 Maenhaut, W., McFiggans, G., Mentel, Th. F., Monod, A., Prévôt, A. S. H., Seinfeld, J. H.,
 Surratt, J. D., Szmigielski, R., and Wildt, J.: The formation, properties and impact of
 secondary organic aerosol: current and emerging issues, Atmos. Chem. Phys., 9, 5155–5236,
 doi:10.5194/acp-9-5155-2009, 2009.
- Hao, L. Q., Yli-Pirilä, P., Tiitta, P., Romakkaniemi, S., Vaattovaara, P., Kajos, M. K., Rinne,
 J., Heijari, J., Kortelainen, A., Miettinen, P., Kroll, J. H., Holopainen, J. K., Smith, J. N.,
 Joutsensaari, J., Kulmala, M., Worsnop, D. R., and Laaksonen, A.: New particle formation
 from the oxidation of direct emissions of pine seedlings, Atmos. Chem. Phys., 9, 8121–8137,
 doi:10.5194/acp-9-8121-2009, 2009.
- Heiden, A. C., Kobel, K., Langebartels, C., Schuh-Thomas, G., and Wildt, J.: Emissions of
 oxygenated volatile organic compounds from plants Part I: Emissions from lipoxygenase
 activity, J. Atmos. Chem., 45, 143–172, 2003.
- 14 Hermansson, E., Roldin, P., Rusanen, A., Mogensen, D., Kivekäs, N., Väänänen, R., Boy, M.,
- 15 and Swietlicki, E.: Biogenic SOA formation through gas-phase oxidation and gas-to-particle
- 16 partitioning a comparison between process models of varying complexity, Atmos. Chem.
- 17 Phys., 14, 11853-11869, doi:10.5194/acp-14-11853-2014, 2014.
- Hoffmann, T., Odum, J. R., Bowman, F., Collins, D., Klockow, D., Flagan, R. C., and
 Seinfeld, J. H.: Formation of organic aerosols from the oxidation of biogenic hydrocarbons, J.
 Atmos. Chem., 26, 189–222, 1997.
- Hoffmann, T., Bandur, R., Marggraf, U., and Linscheid, M.: Molecular composition of
 organic aerosols formed in the α-pinene/O3 reaction: implications for new particle formation
 processes, J. Geophys. Res., 103, 25569–25578, 1998.
- Jacobson, M. Z.: Numerical techniques to solve condensational and dissolutional growth
 equations when growth is coupled to reversible aqueous reactions, Aerosol Sci. Technol., 27,
 491–498, 1997.
- Jacobson, M. Z.: A Solution to the Problem of Nonequilibrium Acid/Base Gas-Particle
 Transfer at Long Time Step, Aerosol Sci. Technol., 39, 92–103, 2005.
- 29 Jenkin, M. E., Saunders, S. M., and Pilling, M. J.: The tropospheric degradation of volatile
- 30 organic compounds: a protocol for mechanism development, Atmos. Environ., 31, 81-104,
- 31 1997.

- Jenkin, M. E., Wyche, K. P., Evans, C. J., Carr, T., Monks, P. S., Alfarra, M. R., Barley, M.
 H., McFiggans, G. B., Young, J. C., and Rickard, A. R.: Development and chamber
 evaluation of the MCM v3.2 degradation scheme for β-caryophyllene, Atmos. Chem. Phys.,
 12, 5275–5308, doi:10.5194/acp-12-5275-2012, 2012.
- 5 Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H.,
- 6 DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. M.,
- 7 Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. A.,
- 8 Hueglin, C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara,
- 9 P., Ehn, M., Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J., Dunlea, E. J.,
- 10 Huffman, J. A., Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo, Y.,
- 11 Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D., Cottrell,
- 12 L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J. Y., Zhang, Y.
- 13 M., Dzepina, K., Kimmel, J. R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M.,
- 14 Williams, L. R., Wood, E. C., Middlebrook, A. M., Kolb, C. E., Baltensperger, U., and
- 15 Worsnop, D. R.: Evolution of Organic Aerosols in the Atmosphere, Science, 326, 1525–1529,
- 16 2009.
- 17 Kerminen, V.-M., Petäjä, T., Manninen, H. E., Paasonen, P., Nieminen, T., Sipilä, M.,
- 18 Junninen, H., Ehn, M., Gagné, S., Laakso, L., Riipinen, I., Vehkamäki, H., Kurten, T., Ortega,
- 19 I. K., Dal Maso, M., Brus, D., Hyvärinen, A., Lihavainen, H., Leppä, J., Lehtinen, K. E. J.,
- 20 Mirme, A., Mirme, S., Hõrrak, U., Berndt, T., Stratmann, F., Birmili, W., Wiedensohler, A.,
- 21 Metzger, A., Dommen, J., 5 Baltensperger, U., Kiendler-Scharr, A., Mentel, T. F., Wildt, J.,
- 22 Winkler, P. M., Wagner, P. E., Petzold, A., Minikin, A., Plass-Dülmer, C., Pöschl, U.,
- 23 Laaksonen, A., and Kulmala, M.: Atmospheric nucleation: highlights of the EUCAARI
- 24 project and future directions, Atmos. Chem. Phys., 10, 10829-10848, doi:10.5194/acp-10-
- 25 10829-2010, 2010.
- 26 Kiendler-Scharr, A., Wildt, J., Dal Maso, M., Hohaus, T., Kleist, E., Mentel, T. F., Tillmann,
- 27 R., Uerlings, R., Schurr, U., and Wahner, A.: New particle formation in forests inhibited by
- 28 isoprene emissions, Nature, 461, 381–384, 2009.
- 29 Kokkola, H., Yli-Pirilä, P., Vesterinen, M., Korhonen, H., Keskinen, H., Romakkaniemi, S.,
- 30 Hao, L., Kortelainen, A., Joutsensaari, J., Worsnop, D. R., Virtanen, A., and Lehtinen, K. E. J.:
- 31 The role of low volatile organics on secondary organic aerosol formation, Atmos. Chem.
- 32 Phys., 14, 1689–1700, doi:10.5194/acp-14-1689-2014, 2014.

- 1 Korhonen, H., Lehtinen, K. E. J., and Kulmala, M.: Multicomponent aerosol dynamics model
- 2 UHMA: model development and validation, Atmos. Chem. Phys., 4, 757–771,
 3 doi:10.5194/acp-4-757-2004, 2004.
- Kulmala, M., Pirjola, U., and Makela, J. M.: Stable sulphate clusters as a source of new
 atmospheric particles, Nature, 404, 66–69, 2000.
- 6 Kulmala, M., Vehkamäki, H., Petäjä, T., Dal Maso, M., Lauri, A., Kerminen, V. M., Birmili,
- W., and McMurry, P. H.: Formation and growth rates of ultrafine atmospheric particles: a
 review of observations, J. Aerosol Sci., 35, 143–176, 2004.
- 9 Kulmala, M., Lehtinen, K. E. J., and Laaksonen, A.: Cluster activation theory as an
- 10 explanation of the linear dependence between formation rate of 3nm particles and sulphuric
- 11 acid concentration, Atmos. Chem. Phys., 6, 787–793, doi:10.5194/acp-6-787-2006, 2006.
- 12 Kulmala, M., Riipinen, I., Sipilä, M., Manninen, H. E., Petäjä, T., Junninen, H., Dal Maso,
- 13 M., Mordas, G., Mirme, A., Vana, M., Hirsikko, A., Laakso, L., Harrison, R. M., Hanson, I.,
- 14 Leung, C., Lehtinen, K. E. J., and Kerminen, V. M.: Toward direct measurement of
- 15 atmospheric nucleation, Science, 318, 89–92, 2007.
- 16 Kulmala, M., Nieminen, T., Chellapermal, R., Makkonen, R., Bäck, J., and Kerminen, V.-M.:
- 17 Climate feedbacks linking the increasing atmospheric CO2 concentration, BVOC emissions,
- 18 aerosols and clouds in forest ecosystems, in: Biology, Controls and Models of Tree Volatile
- 19 Organic Compound Emissions, Springer, Springer Netherlands, 489–508, 2013.
- 20 Kulmala, M., Nieminen, T., Nikandrova, A., Lehtipalo, K., Manninen, H. E., Kajos, M. K.,
- Kolari, P., Lauri, A., Petäjä, T., Krejci, R., Hansson, H.-C., Swietlicki, E., Lindroth, A.,
 Christensen, T. R., Arneth, A., Hari, 5 P., Bäck, J., Vesala, T., and Kerminen, V.-M.: CO2 –
- induced terrestrial climate feedback mechanism: from carbon sink to aerosol source and back,
 Boreal Environ. Res., 19, 122–131, 2014.
- 25 Kurtén, T., Loukonen, V., Vehkamäki, H., and Kulmala, M.: Amines are likely to enhance
- 26 neutral and ion-induced sulfuric acid-water nucleation in the atmosphere more effectively
- 27 than ammonia, Atmos. Chem. Phys., 8, 4095–4103, doi:10.5194/acp-8-4095-2008, 2008.
- 28 Laaksonen, A., Kulmala, M., O'Dowd, C. D., Joutsensaari, J., Vaattovaara, P., Mikkonen, S.,
- 29 Lehtinen, K. E. J., Sogacheva, L., Dal Maso, M., Aalto, P., Petäjä, T., Sogachev, A., Yoon, Y.
- 30 J., Lihavainen, H., Nilsson, D., Facchini, M. C., Cavalli, F., Fuzzi, S., Hoffmann, T., Arnold,
- 31 F., Hanke, M., Sellegri, K., Umann, B., Junkermann, W., Coe, H., Allan, J. D., Alfarra, M. R.,

- 1 Worsnop, D. R., Riekkola, M. -L., Hyötyläinen, T., and Viisanen, Y.: The role of VOC
- 2 oxidation products in continental new particle formation, Atmos. Chem. Phys., 8, 2657–2665,
- 3 doi:10.5194/acp-8-2657-2008, 2008.
- 4 Lai, A. and Nazaroff, W. W.: Modelling indoor particle deposition from turbulent flow onto
- 5 smooth surfaces, J. Aerosol Sci., 31, 463–476, 2000.
- 6 Matsunaga, A. and Ziemann, P. J.: Gas-wall partitioning of organic compounds in a Teflon
- 7 film chamber and potential effects on reaction product and aerosol yield measurements,
- 8 Aerosol Sci. Tech., 44, 881–892, 2010.
- 9 Mauldin, R., Frost, G., Chen, G., Tanner, D., Prevot, A., Davis, D., and Eisele, F.: OH
- 10 measurements during the First Aerosol Characterization Experiment (ACE 1): observations
- 11 and model comparisons, J. Geophys. Res., 103, 16713–16729, doi:10.1029/98JD00882, 1998.
- 12 McMurry, P. H., and Friedlander, S. K.: New particle formation in the presence of an aerosol,
- 13 Atmos. Environ., 13, 1635–1651, 1979.
- 14 McMurry P. H., and Grosjean, D.: Gas and aerosol wall losses in Teflon film smog chambers,
- 15 Environ. Sci. Technol. 19, 1176–1182, 1985.
- 16 McMurry, P. H. and Rader, D. J.: Aerosol wall losses in electrically charged chambers.
- 17 Aerosol Sci. Technol., 4, 249–268, 1985.
- 18 McMurry P. H., Kulmala, M., Worsnop D. R.: Special Issue on Aerosol Measurements in the
- 19 1 nm Range, Aerosol Sci. Technol. 45, i, 2011.
- 20 Mentel, T. F., Wildt, J., Kiendler-Scharr, A., Kleist, E., Tillmann, R., Dal Maso, M., Fisseha,
- 21 R., Hohaus, Th., Spahn, H., Uerlings, R., Wegener, R., Griffths, P. T., Dinar, E., Rudich, Y.,
- 22 and Wahner, A.: Photochemical production of aerosols from real plant emissions, Atmos.
- 23 Chem. Phys., 9, 4387–4406, doi:10.5194/acp-9-4387-2009, 2009.
- 24 Metzger, A., Verheggen, B., Dommen, J., Duplissy, J., Prevot, A. S., Weingartner, E.,
- 25 Riipinen, I., Kulmala, M., Spracklen, D. V., Carslaw, K. S., and Baltensperger U.: Evidence
- 26 for the role of organics in aerosol particle formation under atmospheric conditions, P. Natl.
- 27 Acad. Sci. USA, 107, 6646–6651, 2010.
- 28 Mirme, S., Mirme, A., Minikin, A., Petzold, A., Hõrrak, U., Kerminen, V. -M., and Kulmala,
- 29 M.: Atmospheric sub-3 nm particles at high altitudes, Atmos. Chem. Phys., 10, 437–451,
- 30 doi:10.5194/acp-10-437-2010, 2010.

- Nannoolal, Y., Rarey, J., Ramjugernath, D., and Cordes, W.: Estimation of pure component
 properties Part 1, Estimation of the normal boiling point of non-electrolyte organic
 compounds via group contributions and group interactions, Fluid Phase Equilibr., 226, 45–63,
 2004.
- Nannoolal, J., Rarey, J., and Ramjugernath, D.: Estimation of pure component properties Part
 3. Estimation of the vapour pressure of non-electrolyte organic compounds via group
 contributions and group interactions Fuild Phase Equilibria, 269, 117–133, 2008.
- 8 Owen, S. M., Harley, P., Guenther, A., and Hewitt, C. N.: Light dependency of VOC
 9 emissions from selected Mediterranean plant species, Atmos. Environ., 36, 3147–3159, 2002.
- 10 Paasonen, P., Nieminen, T., Asmi, E., Manninen, H. E., Petäjä, T., Plass-Dülmer, C., Flentje,
- 11 H., Birmili, W., Wiedensohler, A., Hõrrak, U., Metzger, A., Hamed, A., Laaksonen, A.,
- 12 Facchini, M. C., Kerminen, V.-M., and Kulmala, M.: On the roles of sulphuric acid and low-
- 13 volatility organic vapours in the initial steps of atmospheric new particle formation, Atmos.
- 14 Chem. Phys., 10, 11223–11242, doi:10.5194/acp-10-11223-2010, 2010.
- Park, S. H., Kim, H. O., Han, Y. T., Kwon, S. B., and Lee, K. W.: Wall loss rate of
 polydispersed aerosols, Aerosol Sci. Tech., 35, 710–717, 2001.
- 17 Petäjä, T., Mauldin, III, R. L., Kosciuch, E., McGrath, J., Nieminen, T., Paasonen, P., Boy,
- 18 M., Adamov, A., Kotiaho, T., and Kulmala, M.: Sulfuric acid and OH concentrations in a
- 19 boreal forest site, Atmos. Chem. Phys., 9, 7435–7448, doi:10.5194/acp-9-7435-2009, 2009.
- 20 Pierce, J. R., Engelhart, G. J., Hildebrandt, L., Weitkamp, E. A., Pathak, R. K., Donahue, N.
- 21 M., Robinson, A. L., Adams, P. J., and Pandis, S. N.: Constraining Particle Evolution from
- 22 Wall Losses, Coagulation, and Condensation-Evaporation in Smog- Chamber Experiments:
- Optimal Estimation Based on Size Distribution Measurements, Aerosol Sci. Technol., 42,
 1001–1015, 2008.
- 25 Riccobono, F., Schobesberger, S., Scott, C. E., Dommen, J., Ortega, I. K., Rondo, L.,
- 26 Almeida, J. a., Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Downard, A., Dunne,
- 27 E. M., Duplissy, J., Ehrhart, S., Flagan, R. C., Franchin, A., Hansel, A., Junninen, H., Kajos,
- 28 M., Keskinen, H., Kupc, A., Kürten, A., Kvashin, A. N., Laaksonen, A., Lehtipalo, K.,
- 29 Makhmutov, V., Mathot, S., Nieminen, T., Onnela, A., Petäjä, T., Praplan, A. P., Santos, F.
- 30 D., Schallhart, S., Seinfeld, J. H., Sipilä, M., Spracklen, D. V., Stozhkov, Y., Stratmann, F.,
- 31 Tomé, A., Tsagkogeorgas, G., Vaattovaara, P., Viisanen, Y., Vrtala, A., Wagner, P. E.,

- 1 Weingartner, E., Wex, H., Wimmer, D., Carslaw, K. S., Curtius, J., Donahue, N. M., Kirkby,
- 2 J., Kulmala, M., Worsnop, D. R., and Baltensperger, U.: Oxidation products of biogenic
- 3 emissions contribute to nucleation of atmospheric particles, Science, 344, 717–721, 2014.
- 4 Riipinen, I., Sihto, S.-L., Kulmala, M., Arnold, F., Dal Maso, M., Birmili, W., Saarnio, K.,
- 5 Teinilä, K., Kerminen, V.-M., Laaksonen, A., and Lehtinen, K. E. J.: Connections between
- 6 atmospheric sulphuric acid and new particle formation during QUEST III-IV campaigns in
- 7 Heidelberg and Hyytiälä, Atmos. Chem. Phys., 7, 1899–1914, doi:10.5194/acp-7-1899-2007,
- 8 2007.
- 9 Riipinen, I., Pierce, J. R., Donahue, N. M., and Pandis, S. N.: Equilibration time scales of
- 10 organic aerosol inside thermodenuders: Evaporation kinetics versus thermodynamics, Atmos.
- 11 Environ., 44, 597–607, 2010.
- 12 Riipinen, I., Yli-Juuti, T., Pierce, J. R., Petäjä, T., Worsnop, D. R., Kulmala, M., and
- 13 Donahue, N. M.: The contribution of organics to atmospheric nanoparticle growth, Nat.
- 14 Geosci., 5, 453–458, 2012.
- 15 Roldin, P., Swietlicki, E., Schurgers, G., Arneth, A., Lehtinen, K. E. J., Boy, M., and
- Kulmala, M.: Development and evaluation of the aerosol dynamics and gas phase chemistry
 model ADCHEM, Atmos. Chem. Phys., 11, 5867–5896, doi:10.5194/acp-11-5867- 2011,
 2011.
- Roldin, P., Eriksson, A. C., Nordin, E. Z., Hermansson, E., Mogensen, D., Rusanen, A., Boy,
 M., Swietlicki, E., Svenningsson, B., Zelenyuk, A., and Pagels, J.: Modelling non-equilibrium
 secondary organic aerosol formation and evaporation with the aerosol dynamics, gas and
 particle-phase chemistry kinetic multilayer model ADCHAM, Atmos. Chem. Phys., 14,
 7953–7993, doi:10.5194/acp-14-7953-2014, 2014.
- Ruuskanen, T. M., Taipale, R., Rinne, J., Kajos, M. K., Hakola, H., and Kulmala, M.:
 Quantitative long-term measurements of VOC concentrations by PTR-MS: annual cycle at a
 boreal forest site, Atmos. Chem. Phys. Discuss., 9, 81–134, doi:10.5194/acpd-9-81-2009,
 2009.
- Saunders, S. M., Jenkin, M. E., Derwent, R. G., and Pilling, M. J.: Protocol for the
 development of the Master Chemical Mechanism, MCM v3 (Part A): tropospheric
 degradation of nonaromatic volatile organic compounds, Atmos. Chem. Phys., 3, 161–180,
 doi:10.5194/acp-3-161-2003, 2003.

- 1 Schimang, R., Folkers, A., Kleffmann, J., Kleist, E., Miebach, M., and Wildt, J.: Uptake of
- 2 gaseous nitrous acid (HONO) by several plant species, Atmos. Environ., 40, 1324–1335,
- 3 2006.
- 4 Schobesberger, S., Junninen, H., Bianchi, F., Lönn, G., Ehn, M., Lehtipalo, K., Dommen, J.,
- 5 Ehrhart, S., Ortega, I. K., Franchin, A., Nieminen, T., Riccobono, F., Hutterli, M., Duplissy,
- 6 J., Almeida, J., Amorim, A., Breitenlechner, M., Downard, A. J., Dunne, E. M., Flagan, R. C.,
- 7 Kajos, M., Keskinen, H., Kirkby, J., Kupc, A., Kürten, A., Kurtén, T., Laaksonen, A., Mathot,
- 8 S., Onnela, A., Praplan, A. P., Rondo, L., Santos, F. D., Schallhart, S., Schnitzhofer, R.,
- 9 Sipilä, M., Tomé, A., Tsagkogeorgas, G., Vehkamäki, H., Wimmer, D., Baltensperger, U.,
- 10 Carslaw, K. S., Curtius, J., Hansel, A., Petäjä, T., Kulmala, M., Donahue, N. M., and
- 11 Worsnop, D. R.: Molecular understanding of atmospheric particle formation from sulfuric
- 12 acid and large oxidized organic molecules, P. Natl. Acad. Sci. USA, 110, 17223–17228, 2013.
- 13 Sihto, S.-L., Kulmala, M., Kerminen, V.-M., Dal Maso, M., Petäjä, T., Riipinen, I., Korhonen,
- 14 H., Arnold, F., Janson, R., Boy, M., Laaksonen, A., and Lehtinen, K. E. J.: Atmospheric
- sulphuric acid and aerosol formation: implications from atmospheric measurements for
 nucleation and early growth mechanisms, Atmos. Chem. Phys., 6, 4079–4091,
 doi:10.5194/acp- 6-4079-2006, 2006.
- Sipilä, M., Berndt, T., Petäjä, T., Brus, D., Vanhanen, J., Stratmann, F., Patokoski, J.,
 Mauldin, Roy, L., I., Hyvarinen, A.-P., Lihavainen, H., and Kulmala, M.: The role of sulfuric
- 20 acid in atmospheric nucleation, Science, 327, 1243–1246, 2010.
- 21 Spracklen, D. V., Jimenez, J. L., Carslaw, K. S., Worsnop, D. R., Evans, M. J., Mann, G. W.,
- 22 Zhang, Q., Canagaratna, 5 M. R., Allan, J., Coe, H., McFiggans, G., Rap, A., and Forster, P.:
- 23 Aerosol mass spectrometer constraint on the global secondary organic aerosol budget, Atmos.
- 24 Chem. Phys., 11, 12109–12136, doi:10.5194/acp-11-12109-2011, 2011.
- 25 Surratt, J. D., Murphy, S. M., Kroll, J. H., Ng, N. L., Hildebrandt, L., Sorooshian, A.,
- 26 Szmigielski, R., Vermeylen, R., Maenhaut, W., Claeys, M., Flagan, R. C., and Seinfeld, J. H.:
- 27 Chemical composition of secondary organic aerosol formed from the photooxidation of
- 28 isoprene, J. Phys. Chem. A, 110, 9665-9690, 2006.
- 29 Topping, D. O., McFiggans, G. B., Kiss, G., Varga, Z., Facchini, M. C., Decesari, S., and
- 30 Mircea, M.: Surface tensions of multi-component mixed inorganic/organic aqueous systems
- 31 of atmospheric significance: measurements, model predictions and importance for cloud

- activation predictions, Atmos. Chem. Phys., 7, 2371–2398, doi:10.5194/acp-7-2371-2007,
 2007.
- Tsigaridis, K. and Kanakidou, M.: Global modelling of secondary organic aerosol in the
 troposphere: a sensitivity analysis, Atmos. Chem. Phys., 3, 1849–1869, doi:10.5194/acp-31849-2003, 2003.
- Vanhanen, J., Mikkilä, J., Lehtipalo, K., Sipilä, M., Manninen, H. E., Siivola, E., Petäjä, T.,
 and Kulmala, M.: Particle size magnifier for nano-CN detection, Aerosol Sci. Tech., 45, 533–
 542, 2011.
- 9 VanReken, T. M., Greenberg, J. P., Harley, P. C., Guenther, A. B., and Smith, J. N.: Direct
 10 measurement of particle formation and growth from the oxidation of biogenic emissions,
 11 Atmos. Chem. Phys., 6, 4403–4413, doi:10.5194/acp-6-4403-2006, 2006.
- 12 Wang, Z. B., Hu, M., Mogensen, D., Yue, D. L., Zheng, J., Zhang, R. Y., Liu, Y., Yuan, B.,
- 13 Li, X., Shao, M., Zhou, L., Wu, Z. J., Wiedensohler, A., and Boy, M.: The simulations of
- 14 sulfuric acid concentration and new particle formation in an urban atmosphere in China,
- 15 Atmos. Chem. Phys., 13, 11157–11167, doi:10.5194/acp-13-11157-2013, 2013.
- 16 Zhang, R.: Getting to the critical nucleus of aerosol formation, Science, 328, 1366–1367,
 17 2010.
- 18 Zhang, X., Cappa, C. D., Jathar, S. H., McVay, R. C., Ensberg, J. J., Kleeman, M. J., and
- 19 Seinfeld, J. H.: Influence of vapor wall loss in laboratory chambers on yields of secondary
- 20 organic aerosol, P. Natl. Acad. Sci. USA, 111, 5802–5807, 2014.
- 21 Zhang, X., Schwantes, R. H., McVay, R. C., Lignell, H., Coggon, M. M., Flagan, R. C., and
- 22 Seinfeld, J. H.: Vapor wall deposition in Teflon chambers, Atmos. Chem. Phys., 15, 4197-
- 23 4214, doi:10.5194/acp-15-4197-2015, 2015.
- Zuend, A., Marcolli, C., Luo, B. P., and Peter, T.: A thermodynamic model of mixed organicinorganic aerosols to predict activity coefficients, Atmos. Chem. Phys., 8, 4559–4593,
 doi:10.5194/acp-8-4559-2008, 2008.
- Zuend, A., Marcolli C., Booth , A. M., Lienhard, D. M., Soonsin, V., Krieger, U. K., Topping,
 D. O., McFiggans G., Peter, T., and Seinfeld, J. H.: New and extended parameterization of the
 thermodynamic model AIOMFAC: calculation of activity coefficients for organic-inorganic
 mixtures containing carboxyl, hydroxyl, carbonyl, ether, ester, alkenyl, alkyl, and aromatic
 - 34

functional groups, Atmos. Chem. Phys., 11, 9155–9206, doi:10.5194/acp-11-9155-2011,
 2011.

Table 1. The VOCs measured with GC-MS in the JPAC plant chamber. The measured
concentrations were used as input for the ADCHAM model. The "other MTs" and "other
SQTs" refer to other monoterpenes and sesquiterpenes than those specified in the table,

7 respectively.

Isoprene				
α-pinene	β-pinene	myrcene	sabinene	camphene
ocimene	Δ^3 -carene	α-terpinene	Δ -terpinene	α-phellandrene
β-phellandrene	terpinolene	tricyclene	other MTs	-
farnesene	β-caryophyllene	α-longipinene	Δ -cardinene	other SQTs
2-butanol	hexanal	benzene	toluene	eucalyptol
nonanal	bornyl acetate	methyl salicylate		51

10			
11			
12			
13			
14			
15			
16			
17			
18			
19			
20			
21			

- 1 Table 2. Summary of the model sensitivity tests that we performed in order to evaluate and
- 2 constrain possible mechanisms for the formation and growth of nano-CN during the JPAC
- 3 experiments.

Condensable VOC properties method	MCMv3.2 + the Nannoolal vapour pressure method	2D-VBS		
2D-VBS assumptions	Functionalization, fragmentation and OH reaction rates	O:C of the first generation products		
Particle dry deposition loss rates	Varying u* in the range 0.1-0.01 m s ⁻¹			
Gas-phase chemistry	Uncertainties related to the assumed NO _x inflow concentration	Influence of the UV-light on the VOC composition		
Influence of ELVOCs on the particle growth	ELVOC formation from ozonolysis and OH-oxidation of terpenes			
Nano-CN formation and initial growth	MechanismsSurface tension(Eq. 1-6)0.02-0.07 N m ⁻¹			
Oligomerization in the particle phase	Peroxyhemiacetal formation, possibly acid catalysed by co- condensing H_2SO_4 (Eq. 7)			
Non-ideal mixing in the particle phase	Activity coefficients from AIOMFAC			
Reversible VOC wall losses	With or without absorptive uptake on the glass walls.			
SOA phase-state	Liquid-like SOA	Solid-like SOA		



- 1 Table 3. Coefficient of determination (R^2) between the modelled and measured (PSM-CPC)
- 2 total particle number concentration during the experimental campaign.

Nano-CN formation mechanism	ELVOC _{nucl} source	\mathbf{R}^2
Constant $J = 20 \text{ cm}^{-3} \text{ s}^{-1}$ when UV is on		0.968
H_2SO_4 activation (Eq. 1)		0.942
H ₂ SO ₄ kinetic (Eq. 2)		0.887
H ₂ SO ₄ – ELVOC kinetic (Eq. 3)	MT and SQT ox. by O_3 and OH	0.975
H ₂ SO ₄ – ELVOC kinetic (Eq. 3)	MT and SQT ox. OH	0.977
$H_2SO_4 - ELVOC$ kinetic (Eq. 3)	SQT ox. by O_3 and OH	0.944
$H_2SO_4 - ELVOC (Eq. 4)$	SQT ox. by O_3 and OH	0.959
$H_2SO_4 - ELVOC (Eq. 4)$	SQT ox. by O ₃	0.975
ELVOC (Eq. 5)	MT and SQT ox. OH	0.903

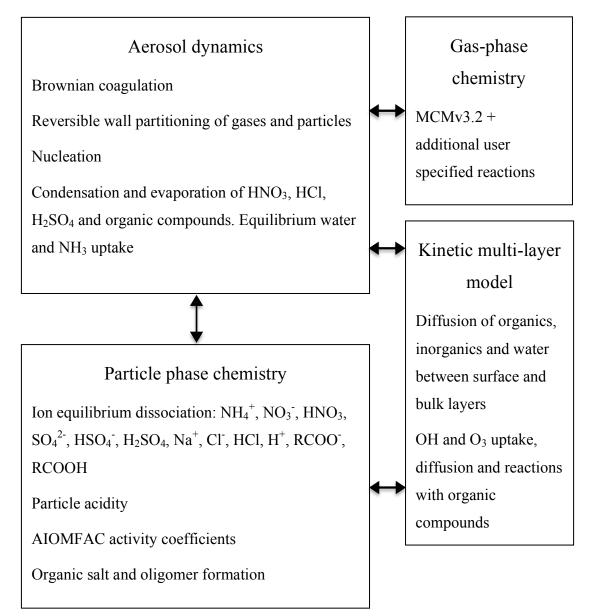
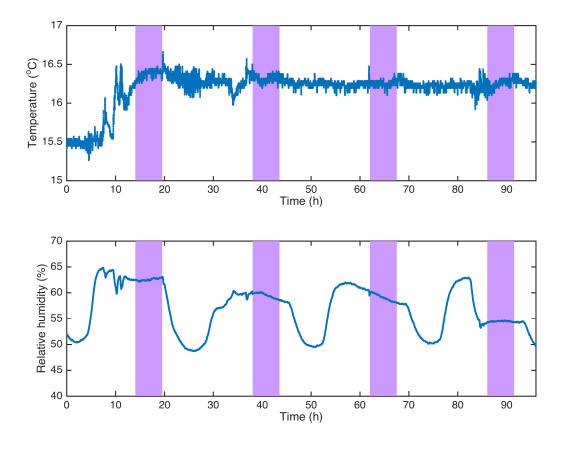


Figure 1. Schematic picture of the ADCHAM model structure.



2 Figure 2. Measured temperature (a) and relative humidity (b) in the JPAC reaction chamber

- 3 during the first four days of the measurement campaign. The purple bars indicate UV-on
- 4 periods.

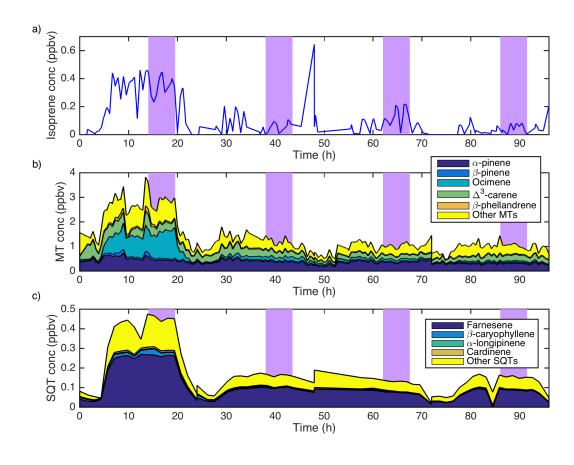


Figure 3. Isoprene (a), monoterpene (b), and sesquiterpene (c) concentrations measured from
the outlet air of the JPAC plant chamber. In Fig. b and c we left out those terpenes that had a
negligible contribution to the total terpene concentration. Purple bars indicate UV-on periods
during the measurements.

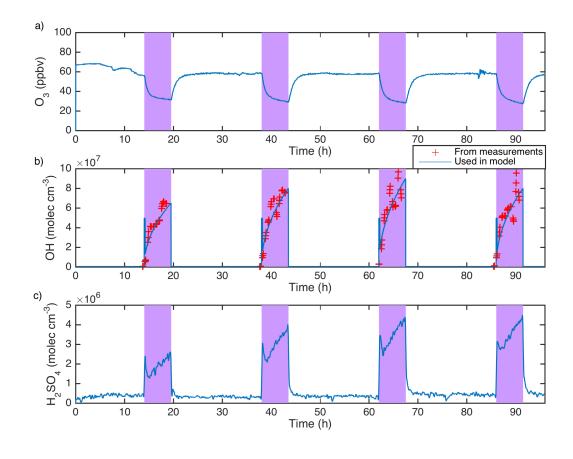


Figure 4. (a) Measured O_3 concentrations in the JPAC reaction chamber, (b) estimated OH concentration based on the observed 2-butanol loss rate and the OH concentration used as

 $4 \quad \ \ model \ input, \ and \ (c) \ measured \ H_2SO_4 \ concentration.$

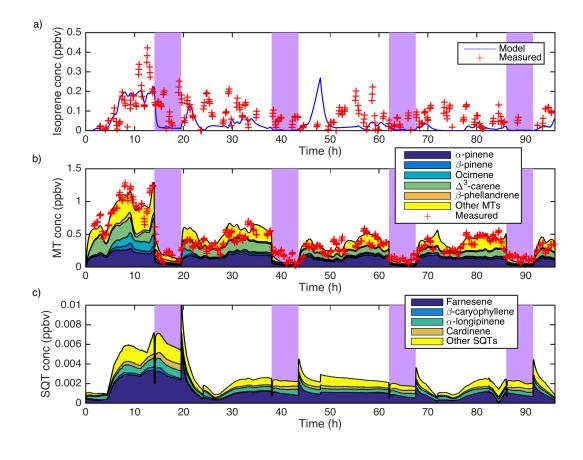


Figure 5. Modelled isoprene (a), monoterpene (b), and sesquiterpene (c) concentrations together with the measured isoprene and monoterpene concentrations in the JPAC reaction chamber. In Fig. b we left out those monoterpenes that had a negligible contribution to the total monoterpene concentration. The purple bars indicate UV-on periods during the measurements.

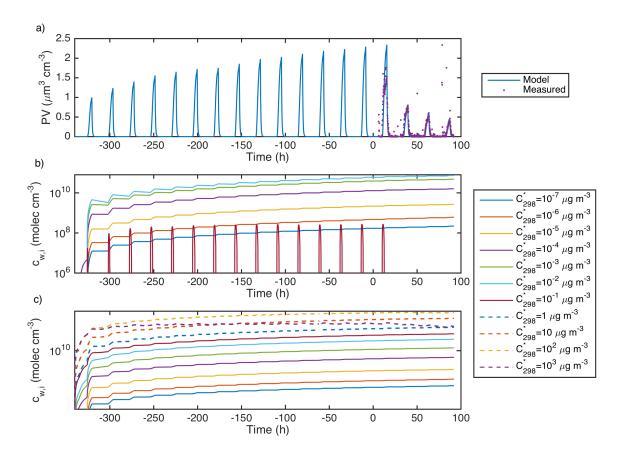
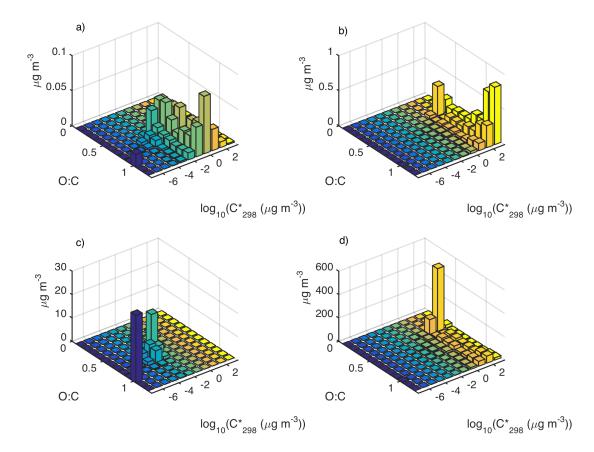


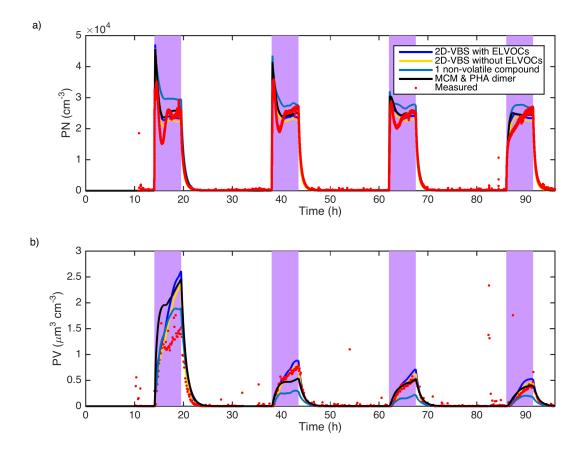
Figure 6. (a) Modelled and measured SOA volume formation and 2D-VBS VOC wall uptake onto (b) the glass walls and (c) the PTFE Teflon walls when considering that the PTFE Teflon walls behave as FEB Teflon walls (Eq. 8, 11 and 12). The VOCs are summed over the all O:C but divided into the different C_{298}^* bins in the 2D-VBS. At time 0 h the intensive measurement campaign started.



2 Figure 7. Modelled 2D-VBS distribution for (a) the SOA particles, (b) the gas-phase, (c) the

3 VOCs on the glass-walls and (d) the VOCs on the Teflon floor. The model results are from

⁴ Day-3 after 5 hours with UV-lights on.



2 Figure 8. Modelled and measured (a) total particle number concentration and (b) total particle 3 volume concentration. The model results are from simulations with the 2D-VBS and with or 4 without ELVOC formation via ozonolysis and OH-oxidation of monoterpenes and 5 sesquiterpenes, from a simulation when the MCM compounds were used as the condensable 6 organic compounds and considering PHA dimer formation with Eq. 7, and from a simulation 7 with only one condensable non-volatile compound. The mass yield of the non-volatile 8 compound formed from O₃ and OH oxidation of all monoterpenes and sesquiterpens was 60 9 %.

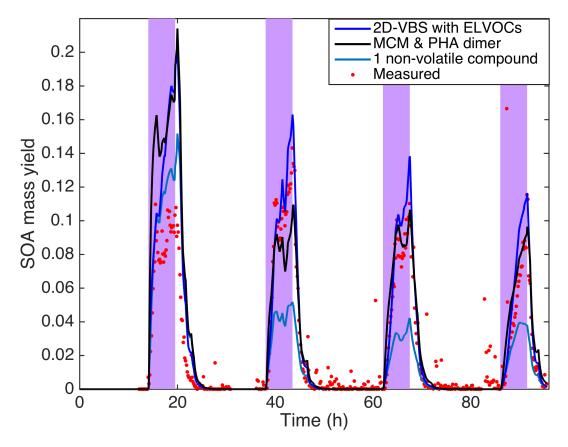


Figure 9. Modelled and measured SOA mass yields. The SOA density was assumed to be
1400 kg m³.

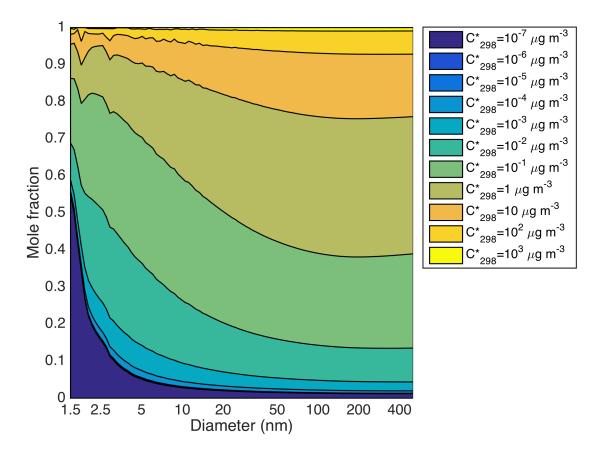




Figure 10. Modelled SOA volatility distribution as a function of the particle size from a
simulation with the 2D-VBS and ELVOC formation from ozonolysis and OH-oxidation of
monoterpenes and sesquiterpenes. The results are from Day-3 after 5 hours into the UV-light
period.

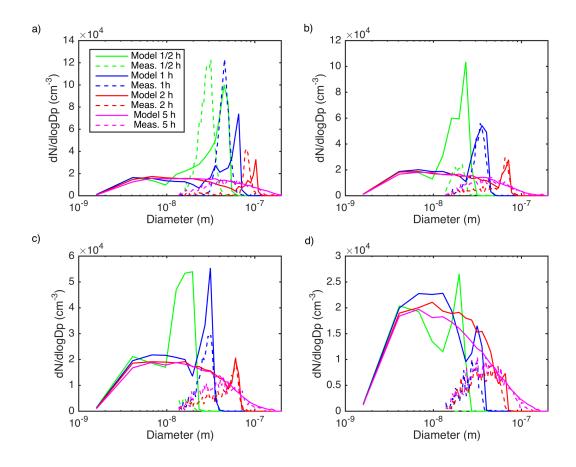


Figure 11. Modelled and measured particle number size distribution from (a) Day-1, (b) Day-2, (c) Day-3 and (d) Day-4 of the experimental campaign. The particle number size distributions are from ½, 1, 2 and 5 hours after the UV-light was turned on. The model results are from a simulation with the 2D-VBS (including ELVOC formation) and the full-moving size distribution method.

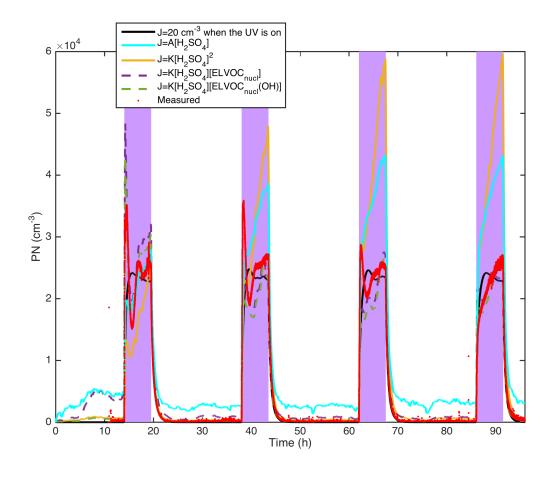


Figure 12. Measured (PSM-CPC) and modelled total particle number concentration with
different nano-CN formation mechanism.