

1 **SOURCES AND CONTRIBUTIONS OF WOOD SMOKE**
2 **DURING WINTER IN LONDON: ASSESSING LOCAL**
3 **AND REGIONAL INFLUENCES**
4

5 **Leigh R. Crilley¹, William J. Bloss¹, Jianxin Yin¹, David C. S. Beddows^{1†},**
6 **Roy M. Harrison^{1*†}, James D. Allan^{2‡}, Dominique E. Young², Mike Flynn²,**
7 **Paul Williams^{2‡}, Peter Zotter³, Andre S. H. Prevot³, Mathew R. Heal⁴,**
8 **Janet F. Barlow⁵, Christos H. Halios⁵, James D. Lee⁶, Sönke Szidat⁷ and**
9 **Claudia Mohr^{8,9}**

10
11 **¹School of Geography, Earth and Environmental Sciences**
12 **University of Birmingham, Edgbaston, Birmingham, B15 2TT, United Kingdom**

13
14 **² School of Earth, Atmospheric and Environmental Sciences**
15 **University of Manchester, Manchester, United Kingdom**

16
17 **³ Laboratory of Atmospheric Chemistry**
18 **Paul Scherrer Institute (PSI), 5232 Villigen PSI, Switzerland**

19
20 **⁴ School of Chemistry, University of Edinburgh, Edinburgh, United Kingdom**

21
22 **⁵ Department of Meteorology**
23 **University of Reading, Reading, United Kingdom**

24
25 **⁶ Department of Chemistry**
26 **University of York, York, United Kingdom**

27
28 **⁷ Department of Chemistry and Biochemistry &**
29 **Oeschger Centre of Climate Change Research**
30 **University of Bern, Bern, Switzerland**

31
32 **⁸ Department of Atmospheric Sciences,**
33 **University of Washington, Seattle, Washington 98195, United States**

34
35 **⁹now at: Institute for Meteorology and Climate Research, Atmospheric Aerosol**
36 **Research, Karlsruhe Institute of Technology, Karlsruhe, Germany**

37
38
39

† Also at: Department of Environmental Sciences / Center of Excellence in Environmental Studies, King Abdulaziz University, Jeddah, 21589, Saudi Arabia

* Corresponding author: r.m.harrison@bham.ac.uk

‡ Also at: National Centre for Atmospheric Science, United Kingdom

40 **ABSTRACT**

41 Determining the contribution of wood smoke to air pollution in large cities such as London is
42 becoming increasingly important due to the changing nature of domestic heating in urban areas.
43 During winter, biomass burning emissions have been identified as a major cause of exceedences of
44 European air quality limits. The aim of this work was to quantify the contribution of biomass
45 burning in London to concentrations of PM_{2.5} and determine whether local emissions or regional
46 contributions were the main source of biomass smoke. To achieve this, a number of biomass
47 burning chemical tracers were analysed at a site within central London and two sites in surrounding
48 rural areas. Concentrations of levoglucosan, elemental carbon (EC), organic carbon (OC) and K⁺
49 were generally well correlated across the three sites. At all the sites, biomass burning was found to
50 be a source of OC and EC, with the largest contribution of EC from traffic emissions, while for OC
51 the dominant fraction likely included contributions from secondary organic aerosols, primary
52 biogenic and cooking sources. Source apportionment of the EC and OC was found to give
53 reasonable estimation of the total carbon from non-fossil and fossil fuel sources based upon
54 comparison with estimates derived from ¹⁴C analysis. Aethalometer derived black carbon data were
55 also apportioned into the contributions from biomass burning and traffic, and showed similar trends
56 to that observed for EC. Mean wood smoke mass at the sites was estimated to range from 0.78-1.0
57 µg m⁻³ during the campaign in January-February 2012. Measurements on a 160 metre tower in
58 London suggested a similar ratio of brown to black carbon (reflecting wood burning and traffic
59 respectively) in regional and London air. Peaks in the levoglucosan and K⁺ concentrations were
60 observed to coincide with low ambient temperature, consistent with domestic heating as a major
61 contributing local source in London. Overall, the source of biomass smoke in London was
62 concluded to be a background regional source overlaid by contributions from local domestic
63 burning emissions. This could have implications when considering future emission control
64 strategies during winter and may be the focus of future work in order to better determine the
65 contributing local sources.

66 1. INTRODUCTION

67 A number of detrimental health effects have been associated with exposure to airborne particles
68 (Cohen et al., 2005; Kampa and Castanas, 2008). Therefore identifying the sources contributing to
69 air pollution is desirable in order to implement effective control strategies to reduce the health
70 burden associated with ambient particles. In an urban environment, the major primary sources of
71 ambient particles are typically road traffic, industrial emissions and biomass burning (Viana et al.,
72 2008). Of these primary sources, biomass burning is perhaps the more difficult to apportion or
73 quantify as its source profile is not well defined due to the large variation in actual sources that alter
74 the chemical nature of the emitted particles. These sources can include natural wild fires and
75 domestic heating, which will further vary depending on the heating device used (e.g. stove or
76 fireplace) and the type of wood burned (Schauer et al., 2001; Simoneit, 2002; Heringa et al., 2011).

77
78 A number of techniques have been developed to determine the contribution of biomass burning to
79 the total levels of ambient particles. These techniques include the chemical analysis of filter
80 samples for elemental carbon (EC) and organic carbon (OC) (Gelencsér et al., 2007), the
81 radioisotope of carbon (^{14}C , Heal, 2014; Heal et al., 2011; Szidat et al., 2006) and chemical tracers
82 such as levoglucosan and potassium (Harrison et al., 2012a; Viana et al., 2013; Wagener et al.,
83 2012). Instruments such as Aethalometers (Sandradewi et al., 2008a; Sandradewi et al., 2008b;
84 Herich et al., 2011) and aerosol mass spectrometers (Allan et al., 2010; Lanz et al., 2010; Crilley et
85 al., 2014) have also been useful for determining the contribution of biomass burning with the added
86 advantage of measurements at a high time resolution. Using these techniques, analytical methods
87 have been utilised to separate biomass burning from other urban sources such as vehicle emissions,
88 and so provide a quantitative estimate of the contribution of biomass burning to the overall particle
89 load. Recent work has shown that the contribution of biomass burning to the OC fraction in an
90 urban environment derived from radiocarbon, levoglucosan and aerosol mass spectrometry results

91 gave similar estimates, ranging from 0.3-0.7 $\mu\text{g m}^{-3}$ and indicates that these techniques can give
92 comparable results (Minguillón et al., 2011).

93
94 Alternatively, indirect measurements of wood smoke may be obtained using either a 2 or 7
95 wavelength Aethalometer. However, the reliability of wood smoke and traffic contribution
96 estimations obtained from Aethalometers has been called into question, particularly for locations
97 with low concentrations of wood smoke (Harrison et al., 2012a; Harrison et al., 2013). Estimating
98 the contribution of biomass burning relies on the enhanced UV absorption of wood smoke relative
99 to traffic emitted particles, as detected by an Aethalometer. Using an Aethalometer, the difference
100 between the UV wavelength channel (370 nm) and the 880 nm channel gives a measure of UV
101 absorbing particulate matter (UVPM) that has demonstrated strong correlation with levoglucosan
102 and hence been suggested as a qualitative tracer for biomass burning (Wang et al., 2011a; Wang et
103 al., 2011b). For a 7 wavelength Aethalometer, a two-component model has been proposed that can
104 separate out the traffic and wood burning contributions using two different wavelength channels
105 (Sandradewi et al., 2008a; Sandradewi et al., 2008b). However, the model relies on two
106 assumptions: 1) that only traffic and wood burning contribute to the absorption and 2) absorption
107 Ångstrom exponents (α) for both sources can be assigned a priori. Single α -values are assumed for
108 each source in this model, even though they, especially α for wood burning, can vary substantially
109 (Harrison et al., 2013). Thus the application of multiple techniques can enhance confidence in the
110 estimated contributions of biomass burning to the ambient particle mass concentrations.

111
112 Determining the contribution of wood smoke in large urban centres such as London is becoming
113 increasingly important due to the changing nature of domestic heating. In the UK for example, over
114 the last few decades the influence of wood smoke upon urban air quality has been thought to have
115 been minimal due to natural gas being the main combustion fuel used for domestic heating.
116 However, this is changing partly due to the installation of biomass burning boilers to meet

117 renewable energy targets and also a rise in so-called recreational burning for aesthetic reasons
118 (Fuller et al., 2013). Recent work in large urban centres (London, Paris and Berlin) has
119 demonstrated an increase in the contribution of wood smoke to ambient particles during winter, that
120 can at times exceed traffic emissions (Crippa et al., 2013; Fuller et al., 2013, Wagener et al., 2012).
121 With measures such as low emission zones (LEZ) reducing the impact of traffic emissions in urban
122 centres (See e.g. (Qadir et al., 2013)), wood smoke is becoming relatively more important to the
123 overall ambient particle concentration, particularly during winter (Fuller et al., 2014). In Europe,
124 biomass burning has been identified as a major cause of exceedences of European air quality limits
125 during winter (Fuller et al., 2013; Reche et al., 2011).

126
127 In light of the changing nature of emissions in urban areas there is a need for on-going
128 measurements to assess the impact of biomass burning in cities like London. Therefore we aimed to
129 determine quantitatively the contribution of biomass burning in London and surrounding rural
130 areas. As a consequence of the challenges in apportioning biomass burning, a number of different
131 methods and instruments were used and compared in the current work. This work also aimed to
132 compare the concentrations between London and rural sites and determine whether local emissions
133 or regional sources were the main source of biomass smoke in London.

134 **2. Method**

135 **2.1 Sampling Sites**

137 The measurements for this paper were conducted as a part of the NERC-funded ClearfLo project
138 (Clean Air for London, www.clearflo.ac.uk), to investigate boundary layer pollution in London. An
139 overview of the project can be found in Bohnenstengel et al. (2014). Four sites were selected for
140 chemical sampling in this particular study; an urban background and elevated site within London
141 and at two sites in surrounding rural areas, with details summarised in the following sections. Site
142 locations are shown in Figure 1. Furthermore, additional measurements with a scanning Doppler

143 lidar were also conducted for this study at sites within central London as indicated in Figure 1.

144 Details on the lidar measurements performed can be found in Section 2.2.3.

145

146 **2.1.1 North Kensington, London**

147 The North Kensington (NK) site is located within a school grounds in a residential area ($51^{\circ}31'15''$

148 N, $0^{\circ}12'49''$ W) and is classified as urban background. Central London lies 7 km to the east and a

149 major road is located approximately 500 m to the south of the site. A more complete description of

150 the site and air pollution climate is available in Bigi and Harrison (2010); the site is considered

151 representative of the background air quality within much of London. This site is part of the

152 Automatic Urban and Rural Network (AURN), with further details on the instrumentation and

153 sampling procedures available in Section 2.2.

154

155 **2.1.2 Detling, Kent**

156 This is a rural site located in the Detling show grounds, approximately 45 km southeast of central

157 London ($51^{\circ}18'7''$ N, $0^{\circ}35'22''$ E) and is bordered by fields. To the south of the site,

158 approximately 150 m lies a busy road, while the closest village (Detling) is to the southwest

159 (approximately 1.5km). The winter campaign site was not co-located with the long-term AURN

160 monitoring site, approximately 800 m to the south of the AURN site (Mohr et al., 2013).

161

162

163 **2.1.3 Harwell, Oxfordshire**

164 The second rural site was located approximately 80 km west of central London within the Harwell

165 Science Centre ($51^{\circ}34'16''$ N, $1^{\circ}19'31''$ W) and is part of the AURN. The closest village, Chilton

166 and the nearest main road (A34) are about 2 km to the east of the site. The site is surrounded by

167 fields and is considered as representative of a rural location (Charron et al., 2007).

168

169

170 **2.1.4 BT Tower, London**

171 The elevated site utilised in this study was the BT tower (51°31'17.31''N; 0°8'20.12''W) located at
172 60 Cleveland St in London. The main platform for the Aethalometer was a well-ventilated balcony
173 (level T35, 160 m above ground level), and the gas monitors sampled from a height of
174 approximately 180 m above ground level (CO and NO_x). Further description of the BT site can be
175 found in Harrison et al. (2012b).

176

177 **2.2 Sampling Methodology and Instrumentation**

178 In this study two datasets were collected, comprising a winter campaign and long-term
179 measurements. The winter campaign was part of the winter intensive observation period (IOP)
180 within the ClearfLo project, which ran from 06/01/2012 till 16/02/2012. The winter campaign
181 included both filter sampling and Aethalometer measurements at Detling, NK and Harwell. Long-
182 term measurements with an Aethalometer and gas monitors were conducted at the BT tower and
183 NK from 24 January 2012 until 24 June 2013. Details of the sampling methods for these two
184 datasets are summarised in the following sections.

185

186 **2.2.1 Winter campaign measurements**

187 During the winter campaign, filter samples were collected for the chemical analysis of selected
188 biomass burning tracers. Due to various logistical reasons, the sampling regimes for the different
189 filter and instrumental analysis utilised in this paper were not all run at the same time or over the
190 same sampling interval and the filter sampling and analysis methods are summarised in Table 1.
191 The levoglucosan sampling was similar at NK and Harwell but differed at Detling in terms of the
192 size fraction collected and sampling start time (Table 1). A Digitel DHA-80 sampler was used at the
193 NK and Harwell sites to collect PM_{2.5} particles on 150 mm diameter quartz fibre filter circles over a
194 24 hour sampling period. These filters were analysed for levoglucosan, total OC, EC, K⁺ and ¹⁴C
195 (the last only applies for NK). At Detling aerosols were collected on prebaked quartz fibre filters

196 (PALL 2500QAT-UP Tissuquartz) using a high-volume sampler (CAV-A/MSb) with a PM₁ inlet, a
197 flow rate of 500l/min and a sampling time of 24h. These filters were analysed for EC, OC, and ¹⁴C
198 in the total carbon fraction. Separate filter samples with a PM₁₀ inlet were collected for
199 levoglucosan analysis.

200
201 The different size fraction analysed for levoglucosan at Detling compared to the other sites is
202 unlikely to affect the results as most of the particle-associated levoglucosan is generally within the
203 PM_{2.5} (Saarikoski et al., 2008; Wagener et al., 2012; Yttri et al., 2005). Potassium samples were
204 collected at Detling as a part of this project but failed quality control and therefore were not
205 included in the current work.

206
207 In addition to the filter samples collected, measurements of BC were conducted using a Magee
208 Scientific 7 wavelength Aethalometer (Model AE-31) at NK and Detling during the winter
209 campaign. At NK, the 7 wavelength Aethalometer (7W AE) sampled continuously with a 5 min
210 sampling interval while at Detling the 7W AE sampled at a different sampling interval of 2 min.
211 Furthermore, at NK the 7W AE also sampled during the ClearfLo summer IOP (21/07/2012 –
212 23/08/2012) with a 5 min sampling interval.

213
214 As Harwell, NK and Detling are part of the AURN, BC and PM_{2.5} concentrations were obtained
215 from instruments in this network for the winter campaign. PM_{2.5} mass concentrations were
216 determined by the Filter Dynamic Measurement System (FDMS) method at an hourly interval and
217 were subjected to the AURN network quality control procedures (Ricardo-AEA,2013). In addition
218 to the 7W AE, data from a 2 wavelength Aethalometer (Magee Scientific, Model AE 22) was also
219 collected at 15 min intervals at each AURN site.

220
221
222
223

224 **2.2.2 Long-term measurements**

225 At the BT tower, measurements of BC were performed by a 7W AE (Magee Scientific, Model AE-
226 30), while concentrations of NO and NO₂ were analysed by chemiluminescence (Ecophysics, model
227 780TR fitted with photolytic blue light converter (BLC) for NO₂), all at a 5 min sampling interval.
228 Calibration was carried out using a 5ppm NO in nitrogen gas standard (BOC), diluted to 50ppb in
229 NO_x free air. Calibration of the BLC conversion efficiency was done using gas phase titration of the
230 NO calibration gas with ozone. An Aerolaser fast carbon monoxide analyser (Model AL5002) at a
231 10 Hz sampling rate was used at the BT tower to measure the concentration of CO. As 7W AE data
232 was only available for limited periods at NK compared to BT tower, 2 wavelength Aethalometer
233 (2W AE) data from the AURN network was also obtained at NK, along with concentrations of CO,
234 NO and NO₂ from 1 January 2012 until 30 June 2013.

235
236

237 **2.2.3 Meteorological measurements**

238 Meteorological data was obtained from the Met Office for three sites that were judged to be the
239 most suitable based upon relative location to the sampling sites. For NK and BT tower, wind data
240 from Heathrow airport station was used as it will not be affected by large buildings. The Met Office
241 stations Benson and Gravesend were used for Harwell and Detling, respectively, which were the
242 closest stations geographically with wind data. Benson is approximately 16 km to the east of
243 Harwell and Gravesend about 20 km northwest of Detling. The meteorological data obtained from
244 the Met Office had a 1 hour time resolution. The approximate times of sunrise and sunset
245 throughout the winter campaign were 07.40 and 16.45 UT (and local time).

246

247 During the ClearfLo summer and winter IOPs, a scanning Doppler lidar (Halo Photonics
248 Streamline) was used for monitoring the boundary layer depth and structure, with the locations
249 shown in Fig. 1. The lidar operates at 1.5 μm wavelength, integrated signals are outputted every 3.6
250 s, and the vertical resolution is 30 m. For the first part of the winter IOP period (6/1/2012 –

251 11/1/2012) the lidar was operating on the roof-top of the Westminster City Council's building
252 (WCC, 18m agl, 51^o 31' 17'' N, 0^o 09' 40'' W), and then it was moved to the Engineering
253 Building at Imperial College London (ICL, 33m agl, 51^o 29' 55'' N, 0^o 10' 29'' W) where it was
254 operating until the 8/2/2012. For the summer IOP period an identical lidar was operating at NK
255 (ground level, 51^o 31' 15'' N, 00^o 12' 49'' W) with a gate resolution of 18 m. The mixing height
256 was detected as the level where the vertical velocity variance measured with the lidar drops below a
257 threshold value, following the method described in Barlow et al. (2014) and originally proposed by
258 Barlow et al. (2011). The minimum observable height is different for each location, and depends on
259 the lidar gate and the height that the instrument is placed (above ground level) at each site. Thus it is
260 124m for WCC, 139m for ICL, and 63m for NK - all heights are measured above ground level.
261 Long term meteorological measurements were conducted at the top of BT tower (190.8m agl),
262 employing an instrumentation platform equipped with an eddy covariance system (Gill Instruments
263 R3-50 sonic anemometer, and Licor Li-7500 Hygrometer), a net radiometer (Kipp and Zonen
264 CNR4) and a weather station (Vaisala WXT520).

265

266

267 **2.3 Chemical Analysis**

268 **2.3.1 Levoglucosan**

269 The analytical method applied is similar to that described in Yin et al. (2010) and Wagener et al.
270 (2012) but modified for the detection of levoglucosan. In brief, one quarter of the quartz fibre filter
271 sample was spiked with an internal standard (IS), methyl-beta-D-xylopyranoside (Sigma-Aldrich
272 Ltd) and extracted with dichloromethane and methanol under mild sonication at room temperature.
273 The combined extracts were filtered and concentrated down to 50 μ l. One aliquot of the extract was
274 evaporated to nearly dryness and derivatised by addition of N,O-
275 bis(trimethylsilyl)trifluoroacetamide plus 1% trimethylchlorosilane (BSTFA + 1% TMCS) and
276 pyridine at 70^oC for 1h, and finally cooled in a desiccator. The quantification was based on the IS
277 and a six point authentic standard calibration curve, with selected ion monitoring mode on an

278 Agilent GC-MS instrument. The ions monitored were m/z 204 and 217 for the IS and m/z 204, 217
279 and 333 for levoglucosan.

280

281 **2.3.2 Elemental and organic carbon**

282 The quartz fiber filter samples from NK, Harwell and Detling were also analysed for OC and EC by
283 a Sunset Thermal-Optical Carbon analyser (Sunset Laboratory Inc.). The protocol used was the
284 EUSAAR 2 protocol (Cavalli et al., 2010), in which organic carbon is removed during an initial
285 non-oxidizing temperature ramp from 200°C to 650°C under a helium atmosphere, and then a
286 second temperature ramp from 500°C to 850°C is initialized with the carrier gas switched to a
287 helium/oxygen mixture for removing elemental carbon and organic carbon pyrolysis products.

288

289 **2.3.3 Radiocarbon (¹⁴C)**

290 **Detling:** Sample preparation for the ¹⁴C analysis was performed with two different techniques.
291 First, TC was isolated before the ¹⁴C analysis using the THEODORE system (see Szidat et al.
292 (2004) for more details). In brief, TC is oxidized to CO₂ from filter punches with at 650°C for 12
293 min in an oxygen stream. The evolving CO₂ is separated from interfering reaction gases, cryo-
294 trapped and sealed in glass ampoules for ¹⁴C measurements. Second, samples were also prepared
295 following the approach described in Fahrni et al. (2010). Several small filter pieces are put in quartz
296 tubes together with ~0.25 g of copper oxide (small rods for elemental analysis, Fluka, Switzerland),
297 evacuated, sealed and combusted for 4 hours at 800°C in a muffle furnace (model LE 14/11/B150,
298 Nabatherm, Germany).

299

300 The ¹⁴C measurements were carried out using the **MI**ni radio**CA**rbon **DA**ting System, MICADAS
301 (Synal et al., 2007) at the Swiss Federal Institute of Technology (ETH) Zürich using a gas ion
302 source (Ruff et al., 2007; Wacker et al., 2013), allowing the ¹⁴C analysis of the gaseous CO₂ without
303 prior graphitisation (Ruff et al., 2010). Results from the ¹⁴C measurement are expressed as fraction

304 of modern (f_M) denoting the ratio of the $^{14}\text{C}/^{12}\text{C}$ content of the samples related to the $^{14}\text{C}/^{12}\text{C}$ ratio in
305 1950, defined as 0.95 times the value of the standard SRM4990B (Stuiver and Polach, 1977). The
306 f_M values were corrected for $\delta^{13}\text{C}$ fractionation (Wacker et al., 2010) and for the ^{14}C decay between
307 1950 and the year of measurement. Furthermore, a mass depending blank correction was applied
308 following an isotopic mass balance as presented by Zapf et al. (2013). The obtained blank mass and
309 f_M of TC were $0.53 \pm 0.18 \mu\text{g cm}^{-2}$ and 0.50 ± 0.17 , respectively. To account for the excess ^{14}C
310 from the thermonuclear weapon tests in the late 1950s and early 1960s the blank corrected f_M values
311 were converted into non-fossil fractions using a reference value ($f_{NF,ref} = 1.082 \pm 0.04$) representing
312 the modern ^{14}C content of biogenic and biomass burning aerosols during the sampling period
313 compared to 1950 before the bomb testing (see e.g. Zotter et. al. 2014a).

314
315 **North Kensington:** Each filter of $\text{PM}_{2.5}$ was equally divided. The percent modern carbon (f_M) in
316 TC was derived from one half of the filter sample through the following two-step combustion,
317 similar to that described in Heal et al. (2011) and using the same quality assurance procedures.
318 Filters were first combusted in high purity O_2 at 375°C for 1 hour, then combusted in high purity O_2
319 at 800°C for 4 hours, with collection of the evolved CO_2 from each combustion. In both cases the
320 CO_2 was subject to off-line clean-up, graphitisation and AMS determination of f_M . The $f_M(\text{TC})$ was
321 calculated by mass balance of the two sub-samples. This was converted to fraction contemporary
322 carbon by dividing by a value that assumes an average age for the material comprising the
323 contemporary carbon. Here a divisor of 1.08 was used, as discussed in Heal et al. (2011).

324
325
326

2.3.4 Potassium

327 Small pieces of the quartz filters collected at Harwell and NK were extracted with distilled de-
328 ionised water (5 mL) by undergoing 40 minutes mechanical agitation. The resulting solutions were
329 filtered with 32 mm Acrodisc syringe filter (0.2 μm pore size) and analysed for K^+ with a Dionex
330 DX500 system which uses a GP-40 gradient pump but run with isocratic elution using 15.4 mM

331 MSA ($\text{CH}_3\text{SO}_3\text{H}$) as eluent. The sample concentrations were calibrated with a series of mixed
332 standards of known concentration (0.01-10 ppm).

333

334 **2.4 Data Analysis**

335 Rose plots, diurnal variations and conditional probability function (CPF) analyses were all
336 performed in R using the Openair package (Carslaw and Ropkins, 2012). CPF is a data analysis tool
337 to find the direction of source contributions and was applied to the Aethalometer and gas species
338 datasets as they had the necessary high time resolution. The difference in time intervals between
339 meteorological and the pollutant data meant that the hourly averaged Aethalometer and gas species
340 data were used. CPF analysis is presented here in the form of polar plots, which plots the number of
341 events where the concentration was greater than the 90th percentile as a function of both wind speed
342 and direction. As a result CPF polar plots present the probability that high concentrations of a
343 pollutant came from a particular wind direction and speed (Carslaw, 2014) and can give information
344 on the contributions from local and regional sources. One-way analysis of variance (ANOVA) was
345 used to compare mean values and was calculated in SPSS v19. Pearson's correlations between the
346 chemical species were calculated in SPSS using hourly or daily concentrations as appropriate to the
347 technique. Due to the different sampling start times at Detling (midnight) to the other sites (noon)
348 for levoglucosan, EC, OC and ^{14}C , the concentrations for consecutive days at Detling were
349 averaged to enable comparison between the sites.

350

351 **2.4.1 Source apportionment of the OC and EC**

352 A number of analytical methods were employed for apportioning the contribution of fossil fuel and
353 biomass burning emissions to the EC and OC fractions. All methods assume that these two sources,
354 biomass burning and fossil fuel emissions are the sole sources of either EC or BC. The first method,
355 referred to as the tracer method throughout this paper, is described in detailed in Gelencsér et al.
356 (2007). In this method, the primary OC concentration from biomass burning (OC_{bb}) is estimated

357 based upon the concentration of levoglucosan and the ratio of levoglucosan to OC measured for
358 biomass burning emissions from the literature. The EC concentration from biomass burning (EC_{bb})
359 is calculated from the OC_{bb} using previously observed OC/EC ratio for biomass burning emissions
360 and the remainder of EC is apportioned to EC from fossil fuel emissions (EC_{ff}). Primary OC from
361 fossil fuel emissions (OC_{ff}) is estimated using previously observed OC/EC ratios for vehicle
362 emissions. The remaining OC is from biogenic, cooking (both primary and secondary) or other
363 secondary sources that can include secondary organic aerosols formed from fossil fuel combustion
364 and this fraction is referred to as OC_{SOA} . One limitation of the tracer method is the assumption of
365 constant source ratios, which may not have been the case throughout the current work. Furthermore,
366 the value of these source ratios for vehicle emissions and biomass burning (both levoglucosan/OC
367 and OC/EC) are dependent on the combustion conditions and fuel type. As a consequence, a wide
368 range of ratios have been previously reported in the literature for vehicle emissions (e.g. El Haddad
369 et al., 2013; Gelencsér et al., 2007; Pio et al, 2011) and for biomass burning (e.g. Gelencsér et al.,
370 2007; Schmidl et al., 2008), and thus there is uncertainty regarding the correct source ratios for
371 emissions at the sites. In this study, we have initially applied the average source ratios for Europe as
372 reported in Gelencsér et al. (2007). As a result of the uncertainty regarding the correct source ratios
373 for biomass burning and vehicle emissions, a sensitivity study was carried out using a range of
374 source ratios from the literature to assess the variability in the source apportionment.

375

376 Further source apportionment of the total carbon (TC, sum of EC and OC concentration) was
377 carried out using the radioisotope of carbon (^{14}C) concentrations as described in Heal et al. (2011).
378 In this method, the TC is apportioned based on whether carbon had a fossil (TC_{ff}) or contemporary
379 (non-fossil, TC_{nf}) origin and will be referred to as the ^{14}C method throughout this paper. For
380 comparison to the tracer method, the sum of EC_{ff} and OC_{ff} is considered equal to the TC_{ff} while the
381 TC_{nf} is equivalent to the sum of OC_{SOA} , OC_{bb} and EC_{bb} . This approach ignores the fact that some
382 secondary OC is of fossil origin, although most published work indicates that the larger proportion

383 is from biogenic (non-fossil) precursors. Secondary OC of fossil origin has been estimated to be
384 $25\pm 13\%$ of secondary OC levels for several European cities in summer (Zotter et al., 2014a and
385 references therein). Recent results in the UK found that fossil EC and fossil OC comprised 27% and
386 20% of TC respectively (Heal et al., 2011). Using the average primary vehicle emissions OC/EC
387 ratio of 0.58 (Gelencsér et al., 2007) the fraction of fossil OC from primary vehicle emissions was
388 15.7% of TC. This implies that fossil secondary OC was only around 4% of TC and therefore any
389 differences in the calculated TC_{nf} assuming that the OC_{SOA} was totally non-fossil would be within
390 the uncertainties associated with the tracer method. As part of the sensitivity study reported in
391 Section 3.3.1, inclusion of a 20% fossil fraction of secondary OC generated results which were still
392 consistent with the ^{14}C data and within the variability of individual days. Furthermore, recent work
393 by Charron et al. (2013) suggests a large biogenic contribution to SOA at southern UK sites.

394

395 **2.4.2 BC corrections and source apportionment**

396 The 7 wavelength Aethalometer provides a real-time optical measurement of the light absorption
397 (b_{abs}) with seven different wavelengths (370 nm, 470 nm, 520 nm, 590 nm, 660 nm, 880 nm and
398 950nm) from which the equivalent black carbon concentration can be deduced (Petzold et al.,
399 2013). Aethalometers suffer from various systematic errors e.g. multiple scattering by the filter
400 fibres (scattering effect) and increased absorption by light absorbing particles accumulating in the
401 filter (shadowing or loading effect) (Bond et al., 1999, Lioussé et al., 1993; Petzold et al., 1997).
402 BC concentrations from the 2W AE and 7W AE were corrected for loading and scattering effects
403 according to the procedure described in Weingartner et al. (2003). In the Weingartner correction, a
404 value of 1.2 was used for the loading correction and the multiple scattering correction constant (C),
405 was determined based upon comparison to results from a Multi Angle Absorption Photometer
406 (MAAP) at 630 nm, as described in Collaud Coen et al. (2010). The results from this comparison
407 determined a C value of 3.336 and 3.095 for NK and Detling, respectively. As there was no MAAP
408 data available at Harwell or the BT tower, the average of the C values determined at NK and

409 Detling were applied to all 2W and 7W AE data to enable comparison between the sites. The mass
410 absorption cross-section, which is needed to convert the light absorption coefficient to BC, was
411 determined by plotting EC concentration (from thermal-optical analysis) against the corrected 880
412 nm b_{abs} at Detling, NK and Harwell. As there was no EC data at BT tower, the mass absorption
413 cross-section determined for NK was applied to the BT tower 7W AE data. The 2W AE mass
414 absorption cross-section values were calculated to be $7.3 \text{ m}^2\text{g}^{-1}$, $8.1 \text{ m}^2\text{g}^{-1}$ and $7.3 \text{ m}^2\text{g}^{-1}$ for Harwell,
415 NK and Detling, respectively while a mass absorption cross-section of $7.4 \text{ m}^2\text{g}^{-1}$ was found for the
416 7W AE at both NK and Detling.

417

418 Measurements from the 7W AE data were apportioned into the contribution from wood burning
419 (BC_{wb}) and traffic (BC_{tr}) according to Sandradewi et al. (2008a; 2008b), referred to as the
420 Aethalometer model. Aethalometer absorption values at 470 and 950 nm were used to determine
421 BC_{wb} and BC_{tr} , respectively and an absorption Ångström exponent (α) of 2 for wood smoke (α_{wb})
422 and 1 (α_{tr}) for traffic was used. As the UK AURN monitoring network employs 2W AE that
423 measure at 370 and 880 nm, UVPM (also referred to as delta-C) concentrations were also
424 determined as UVPM has been proposed as a marker for wood smoke. UVPM concentration was
425 calculated for both the 2W and 7W AE data by the difference of BC equivalent concentrations at
426 370 nm and 880 nm (Wang et al., 2011a; Wang et al., 2011b). Furthermore for 7W AE datasets at
427 the BT tower and NK, α for the total ambient absorbing particles was calculated using the 470 and
428 950 nm wavelength data according to the procedure described in Harrison et al. (2012a).

429

430 **3. RESULTS AND DISCUSSION**

431 **3.1 Meteorology**

432 Meteorological conditions were within expected seasonal variation for winter with low
433 temperatures during the sampling period. Similar temperatures were observed at the three Met
434 stations as demonstrated in Figure 2. Wind rose plots for the winter IOP are given in the Supporting

435 Information (Figures S1-3), with Gravesend and Heathrow stations having similar plots, with the
436 predominant wind direction being westerly. The wind rose plot for Benson differed slightly, with
437 southwest winds being slightly more frequent, as well as a stronger influence of winds from the
438 south.

439

440 **3.2 Overall results from the Filter Samples at Harwell, NK and Detling (Winter** 441 **Campaign)**

442 The summary statistics for the filter analysis results and the BC and PM_{2.5} concentrations are given
443 in Table 2. Similar average concentrations of levoglucosan were observed at the three sites but were
444 notably higher compared to previous measurements in London and the UK (Harrison et al., 2012a).
445 At NK, the mean levoglucosan concentration during the previous winter (2011) was 45 ng m⁻³ with
446 a 90th percentile value of 69 ng m⁻³ (Harrison et al., 2012a), notably lower than the average of 77.6
447 ng m⁻³ observed in this study. In the current work, the average temperature for the sampling period
448 was 4.4 ± 4.4°C while during the 2011 sampling period (Harrison et al., 2012a) it was 6.3 ± 3.1°C,
449 and this temperature difference was found to be statistically significant (p>0.05) using one-way
450 ANOVA. Thus the colder weather in 2012 may account for the observed higher concentrations of
451 levoglucosan due to an increase in the levels of domestic heating coupled with potentially increased
452 atmospheric stability. A similar relationship was observed in Switzerland (Zotter et al., 2014b).
453 Overall, levoglucosan concentrations observed in the current work are at the lower end of previous
454 winter observations across Europe for urban environments, which range from 40 to 570 ng m⁻³ in
455 the PM_{2.5} fraction (Reche et al., 2012 and references therein). Domestic heating by wood burning is
456 more common in mainland Europe than the UK and so would explain the observed lower
457 levoglucosan concentrations in the UK (Harrison et al., 2012a). Levoglucosan emissions from
458 biomass burning are also dependent on the burning conditions and type of fuel (Schauer et al., 2001;
459 Simoneit, 2002; Heringa et al., 2011) and thus may also explain the observed lower concentrations
460 in the UK.

461 The temporal variations in the concentration of levoglucosan, OC, EC and K^+ were similar at the
462 sites, as demonstrated in Figure 3, with peaks in concentration occurring on the same days. EC
463 concentrations were well correlated between all three sites ($r^2 > 0.73$) as were the concentration of
464 OC ($r^2 > 0.82$). Generally at the three sites the meteorology would be expected to be broadly similar
465 owing to their close proximity and the strong correlation of concentrations over time points
466 indicative of a possible regional source of carbonaceous aerosols at the sites, or similar local
467 sources. The intra-site concentrations of levoglucosan and K^+ were correlated at Harwell and North
468 Kensington (r^2 of 0.62 at both sites), suggesting a common source. Levoglucosan concentrations at
469 NK were well correlated with Harwell and Detling (r^2 of 0.70 and 0.71, respectively) but between
470 Harwell and Detling the concentrations were less correlated (r^2 of 0.50). That the concentrations of
471 levoglucosan were at least moderately correlated between all the sites points to a common source at
472 three sites. The weaker correlations between Harwell and Detling are probably a function of the
473 greater inter-site distance and hence greater variations in meteorological conditions, and may also
474 be related to local emissions.

475

476 It was also observed that periods of colder temperatures, such as 17-18th January and 2-4th February
477 (Figure 2) corresponded to peaks in the levoglucosan concentration (Figure 3), especially at
478 Harwell. This would implicate domestic heating as the main source, though it could also be due to
479 greater atmospheric stability during these colder periods resulting in reduced dispersion of the wood
480 smoke (Reche et al., 2012). However, these colder periods (e.g. 2-4th February) did not necessarily
481 equate to times of low wind speeds (Figure S4, Supporting Information), suggesting that the
482 levoglucosan concentrations were related more to changes in emissions. Overall, there appears to be
483 a notable contribution from wood smoke and the sources are further investigated in Section 3.4.

484

485

486 3.3 Contribution of Traffic and Biomass Burning to Carbonaceous Aerosols (Winter 487 Campaign)

488 3.3.1 Source apportionment of OC and EC

489 Using the tracer method, contributions of biomass burning (EC_{bb} and OC_{bb}) and fossil fuel (EC_{ff}
490 and OC_{ff}) to the EC and OC fractions were estimated first using the average of source ratios
491 reported for Europe in Gelencsér et al. (2007). These applied source ratios for biomass burning
492 were 0.136 for the ratio of levoglucosan to OC and an OC/EC ratio of 6, while for vehicle emissions
493 an OC/EC ratio of 0.58 was used. This average OC/EC ratio for vehicle emissions is notably similar
494 to that recently determined for vehicle emissions in London (0.63) (Pant et al., 2014). The average
495 estimated concentrations using these source ratios are shown in Figure 4 for the three sites, with the
496 results of the sensitivity study discussed later in this section. As expected the concentrations of EC_{ff}
497 and OC_{ff} were higher at NK ($p < 0.05$), as it is an urban site with higher traffic influence compared
498 to the rural sites. At all three sites, EC_{ff} accounted for the majority of the EC fraction, indicating
499 that the main source of EC was traffic emissions. Overall, the OC_{bb} and EC_{bb} concentrations were
500 similar across the sites, but OC_{bb} accounted for a higher average percentage of the total OC,
501 $30 \pm 16\%$ and $23 \pm 17\%$ at Harwell and Detling, respectively compared to NK ($16 \pm 7\%$). At all three
502 sites according to the tracer method, approximately 40% of the OC fraction was accounted for by
503 the sum of biomass burning and fossil fuel emissions (Figure 4). The remaining fraction, OC_{SOA}
504 which includes both primary biogenic, cooking and secondary organic aerosols therefore exceeds
505 the sum of primary emitted biomass burning and vehicle emissions.

506

507 A more direct analysis for apportioning the fossil and non-fossil fractions of TC is the ^{14}C method,
508 which was applied to samples from NK and Detling, with the results given in Table 3. At NK, the
509 percentage of TC_{ff} was higher compared to that observed at Detling (Table 3, with the daily TC_{nf}
510 fraction shown in Figure S5, Supporting Information) showing the greater influence of traffic
511 emissions at NK. As shown in Table 3, a very good agreement was observed between the

512 percentage of contemporary (TC_{nf}) and fossil (TC_{ff}) carbon as determined by both the tracer and ^{14}C
513 method, indicating that both methods were consistent in the TC apportionment. The high
514 proportions of TC_{nf} in $PM_{2.5}$ in the range ~50-70% derived in this work, even for urban background
515 (Table 3), are consistent with findings from elsewhere in Europe and worldwide (Heal, 2014;
516 Hodzic et al., 2010).

517

518 To examine whether the ratios used to apportion the TC_{nf} and TC_{ff} were reasonable, a sensitivity
519 study was carried out using a range of source ratios for biomass burning and vehicle emissions from
520 the literature. Gelencsér et al. (2007) proposed minimum and maximum source ratios based upon
521 the ranges in the published data and these have been employed to estimate the uncertainty for the
522 apportionment between contemporary and fossil carbon by the tracer method. It is worth noting that
523 Minguillón et al. (2011) used a similar range of levoglucosan/OC ratios to assess the uncertainty of
524 their source apportionment. Pio et al. (2011) examined $PM_{2.5}$ OC/EC ratios from vehicle emissions
525 and concluded that a 0.3-0.4 was typical of fresh vehicle emissions, and the lower value has also
526 been applied to assess the uncertainty in the apportionment of vehicle emissions. Harrison et al.
527 (2012a) recently proposed a mean levoglucosan/OC ratio of 0.09 for wood burning based on the
528 literature, which has also been applied in this sensitivity study. Ambient ratios for biomass burning
529 were recently determined using the non-fossil fraction of OC and EC during winter in Switzerland
530 (Zotter et al., 2014b), and the average ratios (0.098 and 8.2 for levoglucosan/OC and OC/EC,
531 respectively) have also been included in the sensitivity analysis.

532

533 The results of this sensitivity study are presented in Table 4, along with the values for the ratios
534 used to apportion the contributions of fossil and non-fossil sources. The calculated EC_{ff} , EC_{bb} ,
535 OC_{bb} , OC_{ff} and OC_{SOA} concentrations for different ratios can be found in the Supporting
536 Information, Table S1, while the results in Table 4 are shown in graphical form in Figure S6.
537 Overall, comparison between the calculated percentage TC_{nf} and TC_{ff} in Table 4 exhibit a wide

538 range of values and highlights the importance of selecting the correct source ratio. The closest
539 match to the ^{14}C results for TC_{nf} and TC_{ff} was found using the mean levoglucosan/OC ratio
540 proposed by Harrison et al. (2012a) (Table 4, E), the average Swiss ambient biomass burning ratios
541 (Table 4, G) and the average source ratios from Gelencsér et al. (2007) (Table 3). Therefore it
542 appears that with the application of average source ratios from the literature the split between
543 contemporary and fossil carbon was reasonably estimated, though these ratios may not always
544 accurately represent the sources (Pio et al., 2011). Inclusion of fossil secondary OC, not considered
545 in Table 4 will also give alternative possible fits to the data.

546

547 **3.3.2 Source apportionment of BC**

548 Diurnal cycles of the BC concentrations at Harwell, NK and Detling during the winter campaign
549 measured by 2W AE are shown in Figure 5, with the cycles for the 7W AE given in Figure S7,
550 Supporting Information. Similar cycles and concentrations were observed by the 2W AE and 7W
551 AE at NK while at Detling the diurnal cycles were slightly different, which may have been due the
552 instruments not being co-located. The classic bimodal cycle, coinciding with peak traffic times was
553 observed at both sites, indicating that traffic was the likely main source of BC at NK and Detling.
554 At Harwell the diurnal cycle of BC was also bimodal with a similar morning peak but the peak in
555 concentration in the evening was observed at a later time (20.00) more associated with domestic
556 heating. This difference in the diurnal cycles of BC suggests that local sources of biomass burning
557 may be more influential at Harwell than the other two sites.

558

559 UVPM, defined as the difference in Aethalometer determined concentrations for the 370 and 880
560 nm wavelength, has been proposed as a marker for biomass burning (Wang et al., 2011a). Hence the
561 diurnal cycles of UVPM for both the 2W AE and 7W AE were calculated and are shown in Figure
562 6. All of the five cycles presented in Figure 6 are broadly consistent with the diurnal cycles
563 expected for wood smoke from domestic heating, with a peak in concentration in the evening. At

564 Detling, the observed UVPM cycle is similar to the observed diurnal cycle for other biomass
565 burning markers at Detling, including a PMF-derived BBOA factor (Mohr et al., 2013). The timing
566 of the evening concentration peak varied between the sites, with the observed concentration peak
567 occurring earlier at the rural sites (18-19.00 and at 20.00 at Detling and Harwell, respectively) while
568 at NK the peak appeared later (22.00). This may point to differing source influences, with the
569 earlier peak point to more local emissions while the observed later peak at NK and to a lesser extent
570 at the rural sites suggestive of a regional influences. In addition to the evening concentration peak
571 observed as expected in the UVPM cycle at NK for both instruments, there was also a peak at
572 around 7.00, suggesting a traffic influence in the observed UVPM concentrations at NK.
573 Furthermore, the UVPM and levoglucosan concentrations were found to be correlated at all sites (r^2
574 between 0.8 and 0.9). Overall, based on the diurnal cycles for UVPM presented in Figure 6, it
575 suggests that UVPM may be a useful qualitative marker for wood smoke in areas where traffic
576 emissions are not dominant, in agreement with Harrison et al. (2013).

577

578 Another approach to apportioning the contributions of wood smoke and traffic emission to the
579 observed BC concentrations is the two component Aethalometer model (Sandradewi et al., 2008a;
580 Sandradewi et al., 2008b), which was applied to the 7W AE data. The concentrations of BC_{tr} at both
581 sites were well correlated with NO_x (r^2 of 0.92 and 0.65 for NK and Detling, respectively) and
582 demonstrated higher correlations than between BC_{wb} and NO_x (r^2 of 0.57 and 0.45 for NK and
583 Detling, respectively). At Detling, the concentrations of BC_{wb} recorded a higher correlation with
584 levoglucosan concentrations (r^2 of 0.69) compared to BC_{tr} (r^2 of 0.41), as expected. A similar trend
585 was also observed at NK, with BC_{wb} demonstrating a higher correlation with levoglucosan (r^2 of
586 0.95), than BC_{tr} (r^2 of 0.80). The high correlation observed between BC_{tr} and levoglucosan at NK is
587 surprising and may be due to dominance of traffic emissions in London.

588

589 Notable differences were also observed in the diurnal cycles of the BC_{wb} and BC_{tr} concentrations at
590 NK and Detling as shown in Figure 7. The diurnal cycles were as expected for BC_{tr} at both sites,
591 exhibiting a bimodal cycle as for traffic emissions. However, at Detling the BC_{tr} peak in the
592 morning is broader and the evening peak earlier than would normally be expected for a local traffic
593 source, however the NO_x diurnal cycle also demonstrated an evening peak at the same time.
594 Differences were observed in the BC_{wb} diurnal cycles at Detling and NK (Figure 7). At Detling, a
595 peak in concentration centred on 20.00 was observed, which is typical for emissions from domestic
596 heating. Meanwhile, the flat diurnal cycle in BC_{wb} concentrations observed at NK is suggestive of
597 regional rather than local sources. Overall, the estimated concentrations of wood smoke differed
598 between the methods and these differences are examined in the next section.

599

600 **3.3.3 Comparison of the estimated contributions from biomass burning by the different** 601 **methods**

602 Overall as shown in Figure 4 and Table 5, the main source of both BC and EC was traffic emissions
603 as opposed to biomass burning; however the relative contributions of each source varied depending
604 on the method employed and applied source ratio. EC_{bb} accounted for on average $24\pm 16\%$, $7\pm 2\%$
605 and $14\pm 12\%$ of the total EC at Harwell, NK and Detling, respectively. For the organic fraction, the
606 relative contribution of biomass burning also varied, and was generally higher than that from fossil
607 fuel emissions at the rural sites but not at NK (Figure 4). Using the average conversion factor for
608 levoglucosan to wood smoke particle mass of 11.2 adopted by Harrison et al. (2012a), the mean
609 concentrations of wood smoke mass were calculated to be 1.0, 0.87 and $0.78 \mu\text{g m}^{-3}$ for Harwell,
610 NK and Detling, respectively. From the tracer method, the total biomass burning contribution, taken
611 as the sum of the EC_{bb} and OM_{bb} (OC_{bb} converted to organic matter by conversion factor of 1.6;
612 Aiken et al., 2008) was 1.2, 1.0 and $0.88 \mu\text{g m}^{-3}$ for Harwell, NK and Detling, respectively. The
613 wood smoke mass concentrations as determined by both the levoglucosan conversion factor and the
614 tracer method were notably similar.

615 The two component Aethalometer model determined the contribution of wood burning to be an
616 average of $15\pm 12\%$ and $30\pm 13\%$ of the BC at NK and Detling, respectively. Fuller et al. (2014)
617 found that during December and January of 2009-11, wood burning accounted for an average of
618 23% of BC at NK, higher than the average found in the current work at NK though within the
619 associated variability. The difference is probably attributable to inter-annual variations in
620 concentration. In this study, the BC_{wb} followed the expected trend of higher contributions at the
621 rural compared to the urban site, as also observed for EC_{bb} . However, the percentage contributions
622 determined by the Aethalometer model were higher than that observed for the tracer model for EC_{bb}
623 (Table 5). One of the assumptions in the tracer method and Aethalometer model are constant source
624 composition and α , respectively. However, the α for wood burning is highly variable, ranging at
625 least from 1.4 to 2.2 and has been shown to be dependent on variables such as the fuel and amount
626 of atmospheric aging (Saleh et al., 2013; Kirchstetter et al., 2004). Therefore the applied α for wood
627 burning may not accurately represent the biomass burning aerosols across the sites in this study,
628 possibly resulting in higher calculated contributions. Similarly, the source composition for biomass
629 burning (OC/EC_{bb} and levoglucosan/ OC) are also variable and have been shown to have wide
630 range, which influences the calculated EC_{bb} percentage contributions, ranging from 3-32% and 6-
631 50% of the total EC at NK and Detling, respectively (Table 5 and S1). From Table 5, the closest
632 match to for EC_{bb} to the BC_{wb} percentage contributions was found using the mean levoglucosan/ OC
633 and OC/EC_{veh} and minimum OC/EC_{bb} (Table 4, F). However, this combination of source ratios did
634 not match ^{14}C source apportionment results as well as other combinations (See Table 4). There were
635 no combinations in the sensitivity analysis in Table 4 that were able to closely match both the ^{14}C
636 and Aethalometer model source apportionment and this may be due to the presence of a small
637 fraction of fossil secondary OC. Just as for α_{WB} , these source ratios may not have been selected
638 appropriately for the study and thus the differences between the Aethalometer model and tracer
639 method may be due to the uncertainties in the applied source ratios.

640

641 **3.4 Identification of the sources of Biomass Burning during the Winter Campaign**

642 **3.4.1 Diagnostic ratios**

643 There have been a number of diagnostic ratios utilised to differentiate between the sources of
644 biomass burning including the ratio of levoglucosan to both OC and K^+ (Harrison et al., 2012a;
645 Minguillón et al., 2011; Reche et al., 2012). While there are a number of sources of OC in an urban
646 environment in addition to biomass burning (e.g. vehicle emissions and cooking), observed
647 correlations between levoglucosan and OC in wood smoke for different fuels suggest that its ratio
648 can be used to determine the source (Sullivan et al., 2008). Average values for the diagnostic ratios,
649 levoglucosan/OC and levoglucosan/ K^+ at each site are given in Table 6. At the three sites the
650 average ratio of levoglucosan/OC were similar and are comparable to the characteristic ratio
651 determined for long-range transport of biomass burning in Spain (0.05 ± 0.01) (Reche et al., 2012).
652 This therefore suggests that the major source was regional biomass burning at all of the sites.
653 Furthermore, the observed high correlations between levoglucosan and OC at all the sites (r^2 of 0.71
654 – 0.88) along with the majority of the OC being found to be secondary in origin in Section 3.3.1,
655 suggests a common regional source, possibly including a contribution from mainland Europe
656 (Charron et al., 2013) for both levoglucosan and hence for OC_{bb} .

657

658 The levoglucosan/ K^+ ratios were determined by gradient to reduce the impact of other sources of
659 K^+ , such as airborne sea salt and crustal material particles. A higher gradient was observed at
660 Harwell (1.1 ± 0.14) than at NK (0.73 ± 0.09). Traditional fires (e.g. fireplaces, woodstoves and
661 garden bonfires) have been shown to have a higher levoglucosan/ K^+ ratio compared to more
662 modern heaters, with their higher burn efficiency (Harrison et al., 2012a; Heringa et al., 2011;
663 Khalil and Rasmussen, 2003). The results are suggestive of the use of more efficient devices in NK,
664 consistent with it being located in a Smoke Control Area where the use of fireplaces and
665 woodstoves is controlled.

666

667 At all of the sites, the concentrations of levoglucosan and K^+ showed some temperature
668 dependence, with the lowest concentrations observed during the warmer periods. However, above
669 about 5°C the concentrations did not appear to be as dependent on temperature and thus the
670 levoglucosan/ K^+ was largely unaffected by temperature. Higher concentrations during weekends
671 compared to weekdays are likely to be indicative of discretionary or recreational wood burning
672 (Fuller et al., 2014) and to determine if this influenced the observed ratios, weekend and weekday
673 levoglucosan/ K^+ gradients were compared. These plots are shown in the Supporting Information
674 Figure S8, with the weekday and weekend gradients similar at Harwell. In contrast, at NK, there
675 was a difference in the observed gradients for the weekend (1.35) compared to weekday (0.57). The
676 higher gradient on the weekend suggests more influence from recreational burning in fireplaces
677 during the weekend in London compared to the rural sites, where fireplaces may be utilised more
678 throughout the week as a means of domestic heating. However the small number of weekend
679 samples ($n = 8$) limits the conclusions that can be drawn in this regard.

680
681 Overall, the levoglucosan/ K^+ ratios are low compared to most published data directly measured at
682 the sources (e.g. fireplaces) (Caseiro et al., 2009; Puxbaum et al., 2007) but NK and Harwell were
683 within the range reported by Harrison et al. (2012a) for ambient measurements in the UK. NK had a
684 higher levoglucosan/ K^+ (0.73 ± 0.01) in the current study compared to the ratio of 0.15 observed by
685 Harrison et al. (2012a). The higher ratio observed in the present work may be due to more influence
686 from recreational burning at NK due to the colder winter. Previous work in London also found
687 higher concentrations of wood burning tracers on weekends, which was attributed to secondary or
688 recreational domestic heating (Fuller et al., 2014) supporting the trends observed in the current
689 work for the levoglucosan/ K^+ ratio. Thus based on the diagnostic ratios it appears that there was a
690 contribution from local sources dependent somewhat on ambient temperature, on top of a regional
691 background. Therefore in the next section, CPF analysis was applied to the BC datasets at Harwell,

692 NK and Detling during the winter campaign to analyse whether local or regional sources were the
693 largest contributors to biomass smoke.

694

695 **3.4.2 Direction of significant biomass burning contributions**

696 CPF analysis was applied to the Aethalometer datasets from the winter campaign in order to find
697 the direction of highest contributions at each site to aid the source identification. In Figure 8, CPF
698 analysis of the BC concentrations is presented as polar plots and indicates likely main sources at
699 each site. At NK the highest BC contributions were associated with low wind speeds (Figure 8),
700 which suggests that local emissions are the main source. This along with the diurnal cycles for BC
701 (Figure 5) further indicates that vehicle emissions were the main source of BC. Also from Figure 8,
702 high contributions were to the east of Harwell; in which direction lies the village of Chilton and the
703 A34 (both 2 km). Charron et al. (2013) observed a similar trend for EC concentrations and
704 attributed it to both local traffic emissions from the A34 and also regional sources from mainland
705 Europe. Studies have shown that the levels of BC and vehicle emissions diminish to background
706 levels within 150 m of a road (Massoli et al., 2012; Zhu et al., 2004), and thus the A34 was
707 probably too far away to have a strong influence. Therefore there was probably a regional
708 background of BC, with contributions from local domestic heating (Figure 8). A regional influence
709 was also evident at Detling, with the highest contributions to the BC concentrations associated with
710 high wind speeds from the east, in agreement with that observed at Harwell and probably influenced
711 by advection from mainland Europe (Charron et al., 2013). High contributions were also associated
712 with low wind speeds at Detling (Figure 8), indicative of an additional local source of BC along
713 with the regional source.

714

715 Polar plots of the mean concentrations of BC_{wb} and BC_{tr} concentrations were calculated in order to
716 further analyse whether the Aethalometer model could distinguish between the two sources. At
717 Detling different source directions were observed for the two components, however at NK the BC_{wb}

718 and BC_{tr} polar plots were similar (Figure 9). From Figure 9, the polar plot for BC_{tr} at NK showed a
719 slight trend in direction of the closest road (<100 m to the east), and further indicates local traffic
720 emissions as the source. In contrast, at Detling the BC_{tr} polar plot (Figure 9) indicated little
721 influence from the closest road, which was 150 m to the south and thus unlikely to have a large
722 influence (Massoli et al., 2012; Zhu et al., 2004). Rather, the highest concentrations were in the
723 direction of central London, suggesting that emissions from London may be a larger source of BC_{tr}
724 at Detling. Different source contributions between NK and Detling were also observed for BC_{wb}
725 (Figure 9). From Figure 9, the BC_{wb} polar plots indicated that local sources were the main source of
726 BC_{wb} at North Kensington whereas at Detling there were both local and regional influences. The
727 high concentrations of BC_{wb} at Detling (Figure 9) at low wind speeds to south/southwest may be
728 indicative of emissions from nearby villages. The highest concentrations were at high wind speeds
729 to the east at Detling (Figure 9), which is further indication of a strong regional biomass burning
730 source of BC at the sites.

731

732 **4. LONG-TERM MEASUREMENTS AT THE BT TOWER**

733 To examine the influence of regional sources on the levels of air pollution it is necessary to
734 distinguish between local and regional contributions. One method that has been used previously is
735 to sample from an elevated site such as the BT tower (160 m a.g.l.) (Harrison et al., 2012b). The
736 atmosphere at the tower height, under certain meteorological conditions can become ‘de-coupled’
737 from local sources at the surface when the top of the planetary boundary layer is below that of the
738 elevated site. During these times of de-coupling, local sources will not be contributing appreciably
739 to the observed particles at the elevated location and so regional sources will be the major
740 contributor, allowing for characteristics of these regional particles to be investigated. Previous
741 measurements at the BT tower in central London demonstrated that frequently at night the turbulent
742 mixing height can drop below the top of the tower (Barlow et al., 2011), indicating that the tower is
743 an ideal location for analysing regional sources of airborne particles during these de-coupling

744 periods. Because such periods mainly occur during nighttime hours, the derived results will reflect
745 the nighttime situation, although the regional air may have received its pollutant input largely
746 during daytime hours.

747

748 **4.1 Sources of BC at BT**

749 The two likely major sources of BC (traffic emissions and biomass burning) in urban areas have
750 distinct diurnal cycles, and hence the mean diurnal variation of BC for the London sites for the
751 long-term measurement period is shown in Figure 10. At each of the sites, the normalised BC
752 diurnal cycles (Figure 12) demonstrated similar cycles to that of NO_x and CO (Figure S9,
753 Supporting Information), thus implicating a similar source, likely traffic emissions based on the
754 high fraction of fossil EC found in Section 3.3.1. One feature of Figure 10 and Figure S9 is a one
755 hour lag in the morning concentration peak at the BT tower compared to the ground-level site,
756 possibly due to the time for the locally emitted pollutants to mix to the top of the tower (Despiou
757 and Croci, 2007). The main difference in the diurnal cycles shown in Figure 10 was in the absolute
758 concentrations of BC at each site, which varied depending on the distance from the main local
759 source, traffic emissions. BC concentrations at BT tower were, on average a factor of 2-3 times
760 lower than at NK (Figure 10). EC has been observed to follow similar trends in previous
761 measurements at BT tower and NK (Harrison et al., 2012b). Concentrations of BC at NK and BT
762 tower were correlated (r^2 of 0.6), indicating that the urban background as represented by the NK site
763 was the main source of pollutants at the elevated BT site.

764

765 CPF analysis was applied to the observed BC concentrations at NK and the BT tower for the long-
766 term measurements, to examine whether local or regional sources were the more significant
767 contributors, with the results shown in Figure 11, which are similar to the plots for NO_x and CO
768 (Figure S10, Supporting Information). In Figure 11, similar trends were observed at NK and the BT
769 tower, with the highest BC concentrations being associated with low wind speeds, which indicate

770 that the major contributions were from local sources, likely vehicle emissions based on the observed
771 diurnal cycles (Figure 12).

772 Overall, it appears that the main influences on BC concentrations at BT tower are local ground-level
773 sources (traffic emissions), with some influence from regional sources, investigated further in
774 subsequent sections.

775

776 **4.2 Absorption Ångstrom Coefficient at the Elevated (BT tower) and Ground Sites (NK)**

777 The frequency distributions of the absorption Ångstrom coefficient (α) calculated hourly for the
778 long-term measurements at the BT tower are given in Figure S11 (Supporting Information) and is
779 similar to the α distribution for NK calculated by Harrison et al. (2013). A mode of 1.1 – 1.2
780 indicates traffic emissions as the dominant source of BC at the BT tower (Harrison et al., 2013;
781 Sandradewi et al., 2008a). Little seasonal variation was observed in the distribution of α at the BT
782 tower, with a mean α value of 1.1 for all seasons (Figure S12, Supporting Information). The modest
783 seasonal variation in the frequency distributions of α indicates that traffic emissions were the main
784 source at the BT tower throughout the year. Furthermore, the α frequency distribution at the BT
785 tower was not found to be dependent on wind direction. Another observation from Figure S14 was
786 that a considerable number of measurements had an α value less than 1, as also observed by
787 Harrison et al. (2013) at NK and elsewhere (Gyawali et al., 2009; Herich et al., 2011). The observed
788 large number of α less than 1 at the BT tower may be a sampling artefact as result of aerosol
789 evaporation from the filters (Harrison et al., 2013). However, Gyawali et al. (2009) demonstrated
790 that particles consisting of collapsed BC core coated with secondary organic or inorganic aerosol
791 from vehicle emissions have an α less than 1, thus the presence of slightly aged vehicle emissions
792 may also explain the observed large number of α less than 1 at the BT tower.

793

794 As 7W AE data was only available at NK during the summer and winter IOPs, the distribution of
795 hourly α values was calculated for these two periods to allow direct comparison between the ground

796 and elevated sites, shown in Figure 12. The shape of the frequency distributions of α at both sites
797 was similar during each IOP, with a difference in the mode observed between the winter and
798 summer IOP. During the summer IOP, the mean α was 1.09 ± 0.12 at both sites, while during the
799 winter IOP the mean α was 1.19 ± 0.15 and 1.18 ± 0.17 at NK and BT tower, respectively and this
800 difference between seasons was found by one way ANOVA to be statistically significant ($p < 0.05$).
801 As wood smoke has a higher α than traffic emissions, it points to an increased influence of wood
802 smoke from domestic heating during the winter IOP. To examine this further, frequency
803 distributions of α were determined for the day and night times (21.00 until 05.00 inclusive, to avoid
804 peak traffic), as there would be expected to be higher wood smoke contribution during the evening
805 from domestic heating. As expected, there was no significant difference observed between day and
806 night α distributions during the summer IOP, due to the limited influence of domestic heating
807 (Figure S13, Supporting Information). During the winter IOP, different day/night trends were
808 observed at the sites, with a shift to larger α values at both sites in the night (Figure S14, Supporting
809 Information). At NK in Figure S14, while the α frequency distribution showed a notable shift for
810 night times to larger values, a mean α of 1.23 ± 0.19 suggests that traffic emissions were the
811 dominant factor even under these conditions (Harrison et al., 2013; Saleh et al., 2013).

812

813 **4.3 Identification of Periods when the BT Tower was De-coupled from Surface Emissions**

814 The turbulent mixing height (TMH) was calculated for the two IOPs and a number of occasions
815 were identified when the TMH was below the sampling height of the Aethalometer and gas
816 monitors at the BT tower, and thus the measurements were de-coupled from surface emissions, as
817 listed in Table S2, Supporting Information. As previously observed in London, extended periods of
818 de-coupling at the BT tower were observed at night (Barlow et al., 2011).

819

820 During only one of the de-coupling times listed in Table S2 both NO_x and BC data were available at
821 the BT tower, and the time series of the concentrations of these two pollutants are shown in Figure

822 13 along with the calculated TMH during this period. This event is used to determine whether the
823 times of de-coupling can be estimated based on the comparison between the NO_x and BC
824 concentration time series at the BT tower and NK. As can be seen in Figure 13a, when the sampling
825 platform at the BT tower is above the TMH (18.00 to 04.00), the trends in NO_x concentration at the
826 BT tower do not follow the trends in NO_x concentration at NK, unlike when the TMH is above the
827 tower height. Between 18.00 on 3/2/2012 and 04.00 on 4/2/2012 when de-coupling occurred, the
828 NO_x concentration was much less variable than that at NK, hinting at a more regional character.
829 Furthermore, in the morning when the TMH rises above the tower height (06.00), the BT tower
830 NO_x concentration is observed to peak along with that at NK as local ground-level emissions reach
831 the tower platform. Similar trends are also observed for the BC concentration (Figure 13b), and
832 would suggest that although there are no detectable TMH values between 22.00 to 03.00, the tower
833 was still de-coupled during this time.

834

835 To further confirm that it was likely that the BT tower measurements were decoupled, the sensible
836 heat flux (H) and turbulent kinetic energy normalized by the wind speed (tke/U^2) measured using
837 the sonic anemometer at 190 m on the BT tower are shown in Figure 13c. As the convective
838 boundary layer decays, and the BT Tower is thus in a residual layer, weakly stable (negative) heat
839 fluxes should be observed and low turbulence levels would be expected. Heat flux values change
840 sign early in the afternoon (~15:00) and remain negative until next morning (crossover from
841 positive to negative values occurs at 06:00). Correspondingly, the initially enhanced turbulence
842 (0.142 at 15:00) reduces to low values (around 0.05) later in the afternoon. Turbulence remains at
843 low levels throughout the night and increases the following morning (0.381 at 07:15). The two
844 relatively high values observed at 21:45 and 2:45 (0.123 and 0.105 respectively) might be attributed
845 to intermittent, locally driven turbulent events in the stable residual layer. The trends of both heat
846 flux and turbulence agree with the mixing height data and confirm the decoupling of the BT Tower
847 measurements during night-time when there is no mixing height data.

848 Overall, comparison of the observed trends in the NO_x and BC concentrations at BT and NK
849 allowed identification of times when the BT tower is de-coupled from ground-level emissions. Thus
850 visual inspection of the NO_x and BC concentrations was conducted to estimate times of de-coupling
851 by looking for the key trends including:

- 852 • Focus on night time only, outside of peak traffic times, as this is when it is more likely to occur
853 (Barlow et al., 2011).
- 854 • Divergence in trends between NK and the BT tower, with time series at NK typically more
855 variable than at the BT tower.
- 856 • Sharp peak in concentration in the morning at the BT tower as the TMH rises above the sampler
857 height.

858

859 By the application of these criteria, 16 additional de-coupling periods were identified and are
860 described in Table S3 (Supporting Information), along with the time series of the NO_x and BC
861 concentrations for selected events (Figures S15-S19, Supporting Information) demonstrating the
862 trends that are indicative of de-coupling. Using these times of de-coupling, the influences of local
863 and regional sources on the observed BC concentrations were analysed.

864

865 **4.4 Trends in the BC concentrations during De-coupling Events**

866 In addition to demonstrating the changes in the TMH, Figure 13b also revealed differences in the
867 BC concentration time series as measured by the two Aethalometer channels, 370 and 880 nm at
868 NK and BT during the period of de-coupling. While BC is the strongest light absorbing component
869 of aerosols (Moosmüller et al., 2009 and references therein) other aerosol components such as
870 certain organic compounds have been observed to absorb light, typically at the lower visible to UV
871 region of the spectrum (Zhang et al., 2013). This light absorbing organic aerosol components, also
872 referred to as brown carbon, are thought to arise from both biomass burning and secondary organic
873 aerosols (Andreae and Gelencsér, 2006; Lack et al., 2013). Using the event in Figure 13 when the

874 BT tower was de-coupled from ground level emissions, the BC concentrations at the BT tower were
875 analysed to characterise regional sources. Equally, the observed trends in the measured BC
876 concentrations at NK during de-coupling were likely due to local emissions with limited regional
877 influence, and these are investigated further in the following sections.

878

879 **4.4.1 Characterisation of Local BC Emissions at NK during De-coupling**

880 In Figure 13b while both Aethalometer channels follow the same temporal trends at NK, the
881 increased concentration at 370 nm relative to the 880 nm channel at around 20.00 suggests a change
882 in contributing sources. This timing is consistent with increased contributions from biomass burning
883 for domestic heating, while traffic emissions would be decreasing with reduced traffic flows. The
884 time series of the calculated α values further demonstrates the influence of wood smoke at NK
885 (Figure 14). Between 20.00 and 21.00, α increased from 1.08 to 1.41, which is indicative of a
886 change in the dominant emission source from traffic ($\alpha = 1$) to wood smoke, which has a highly
887 variable α ranging from 1.38 to 2 (Saleh et al., 2013; Sandradewi et al., 2008b). Furthermore the re-
888 emergence of traffic emission as the main source can be seen at 06.00 (when traffic counts starts to
889 rise) as α dropped to 1.16 (Figure 14).

890

891 The average α during the night was 1.32 ± 0.05 , consistent with a significant biomass burning
892 component of BC at NK. Similar temporal trends in the Aethalometer concentrations at 370 and 880
893 nm were also observed at NK during the other winter de-coupling events (e.g. Figure S15), while
894 during summer de-coupling events (e.g. Figures S17 and S18) there was little difference observed
895 between the concentrations at 370 and 880 nm.

896

897 **4.4.2 Characterisation of regional BC particles during de-coupled periods**

898 An increased concentration at 370 nm compared to 880 nm was also observed at the BT tower in
899 Figure 13, although the level of enhancement was not a large as at NK. The flat temporal trends

900 observed at the BT tower by the Aethalometer during this de-coupling period suggest
901 regional/background levels rather than the influence of local sources. A similar flat temporal trend
902 in BC concentrations was generally observed across the other de-coupling events (e.g. Figures S16,
903 S17 and S18) though without the higher concentrations at 370 nm. Elevated concentrations at 370
904 nm relative to the 880 nm channel were only observed during the winter de-coupling periods
905 (Figures 13 and S15), with an average increased concentration at 370 nm of 20-40% during these
906 de-coupled times. With little brown carbon observed during the summer/spring, it points to a
907 seasonal effect due to biomass burning on a regional scale.

908

909 In Figure 14, the constancy of α at the BT tower, with limited changes during this de-coupled period
910 suggests a well-mixed background of light absorbing aerosols with a relatively uniform
911 composition. To examine this further, the frequency distributions of α at the BT tower were
912 calculated for the identified de-coupling periods and compared to the remaining nights without
913 evidence of de-coupling (Figure 15). Similar distributions of α were found with and without de-
914 coupling from surface emissions, pointing to uniform background particles that were difficult to
915 distinguish from local sources of BC using UV wavelengths. Although most of the identified events
916 were during summer and spring, little seasonal variation in the α distributions at the BT tower was
917 observed which suggests that regional aerosols over the whole sampling period had a similar value
918 of α .

919

920 It should be noted that recent studies have demonstrated that Aethalometers may have difficulties
921 distinguishing brown carbon when its levels are low compared to BC, but when there are high
922 concentrations, for example from biomass burning, Aethalometers can detect these high levels (Liu
923 et al., 2013; Rizzo et al., 2011). Thus in the current work, apart from winter when there was a
924 substantial amount of biomass burning from domestic heating, the levels of regional brown carbon
925 may have been too low relative to BC for detection using an Aethalometer.

926 5. CONCLUSIONS

927 By estimation of spatial distributions, sources of wood smoke during winter in London were
928 investigated at an urban background site and two surrounding rural sites. An array of chemical
929 species that have previously been used as biomass burning tracers were analysed to determine the
930 wood smoke contribution. Concentrations of levoglucosan, EC, OC and K^+ were generally well
931 correlated, indicating similar sources of these species and similar meteorology at the three sites.
932 Based on the conversion factor for levoglucosan, mean wood smoke mass at Detling, NK and
933 Harwell was 0.78, 0.87 and $1.0 \mu\text{g m}^{-3}$, respectively. Source apportionment of the EC and OC by
934 the tracer method of Gelencsér et al. (2007) was found to give reasonable estimation of the TC_{nf} and
935 TC_{ff} based upon comparison to the ^{14}C method results. The choice of source ratios used in the tracer
936 method was found to be critical, with the averages of published data giving the most consistent split
937 between TC_{nf} and TC_{ff} . At all the sites, biomass burning was found to be the smallest of the major
938 sources of primary OC and EC, with the largest source of EC found to be traffic emissions, while
939 for OC the dominant fraction likely included contributions from secondary organic aerosols
940 (including SOA from biomass smoke), primary biogenic and cooking sources

941

942 Similar trends to EC were observed in the source apportionment of BC, though higher percentage
943 contributions of wood burning were determined using the Aethalometer model but were within the
944 uncertainties associated with the tracer method. The higher percentage contributions of wood
945 burning determined by the aethalometer model were possibly due to the chosen α which might not
946 perfectly represent the wood burning conditions, wood types used and atmospheric processing in
947 this study. The observed diurnal cycles of BC_{wb} differed between NK and Detling, with the former
948 demonstrating a flat diurnal cycle indicative of regional background while the cycle at Detling was
949 more characteristic of domestic wood burning. Generally, apart from at NK, the diurnal cycles for
950 UVPM were broadly consistent with expected trends for domestic heating emissions and suggests

951 that UVPM may be a useful qualitative marker for wood smoke in areas where traffic emissions are
952 not dominant, in agreement with a previous work (Harrison et al., 2013).

953

954 At all the sites, CPF analysis indicated high contributions from local sources, notably at NK. Peaks
955 in the levoglucosan and K^+ concentrations were observed to coincide with low ambient temperature,
956 suggesting domestic heating as a contributing source in London. Furthermore, levoglucosan
957 concentrations were higher in the present work compared to the previous winter, likely due to the
958 colder temperatures resulting in increased burning for domestic heating. Thus along with diurnal
959 cycles, these local sources were probably related to domestic heating.

960

961 Measurements aloft (160 metres) on the BT Tower showed BC to behave similar to CO (and NO_x),
962 with concentrations following those at NK, with a time lag of about one hour. Examination of
963 periods when the atmospheric boundary layer top was below the BT sampling platform allowed
964 comparison of air polluted by ground-level emissions (measured at NK) with regional air (sampled
965 at BT). This showed little difference in the absorption Ångström coefficient (α), even in winter,
966 between the two sites, although NK showed an increase in α of BC during periods with a low
967 boundary layer top in winter, consistent with the influence of wood burning emissions.

968 Overall, the inter-site correlations in concentration across southern England, the directionality
969 shown by the polar plots and the comparison of local ground-level and regional air composition are
970 all consistent with widespread emissions of biomass smoke across southern England and the near
971 continent, including emissions occurring within London itself. The contribution to PM mass is
972 significant, and any further changes in the pattern of domestic fuel usage towards greater wood
973 burning are a matter of concern.

974

975

976

977 **ACKNOWLEDGEMENTS**

978 The ClearfLo project was funded by the UK Natural Environment Research Council
979 (NE/H003142/1) and was coordinated by the UK National Centre for Atmospheric Science
980 (NCAS). We would like to thank the Sion Manning School for use of their grounds. We thank
981 Philip Naysmith, Gordon Cook and colleagues at the Scottish Universities Environment Research
982 Centre at East Kilbride, UK, for provision of ^{14}C values for North Kensington PM samples.
983 Furthermore, we are grateful to Lukas Wacker (ETH Zürich, Laboratory of Ion Beam Physics) for
984 making available ^{14}C analyses of the Detling PM samples. Raw data are available from the authors
985 on request.

986
987

988 **REFERENCES**

- 989
- 990 Aiken, A. C., DeCarlo, P. F., Kroll, J. H., Worsnop, D. R., Huffman, J. A., Docherty, K. S., Ulbrich,
991 I. M., Mohr, C., Kimmel, J. R., Sueper, D., Sun, Y., Zhang, Q., Trimborn, A., Northway, M.,
992 Ziemann, P. J., Canagaratna, M. R., Onasch, T. B., Alfarra, M. R., Prevot, A. S. H., Dommen, J.,
993 Duplissy, J., Metzger, A., Baltensperger, U. and Jimenez, J. L.: O/C and OM/OC ratios of primary,
994 secondary, and ambient organic aerosols with high-resolution time-of-flight aerosol mass
995 spectrometry, *Environ. Sci. Technol.*, 42, 4478-4485, 2008.
- 996 Allan, J. D., Williams, P. I., Morgan, W. T., Martin, C. L., Flynn, M. J., Lee, J., Nemitz, E.,
997 Phillips, G. J., Gallagher, M. W. and Coe, H.: Contributions from transport, solid fuel burning and
998 cooking to primary organic aerosols in two UK cities, *Atmos. Chem. Phys.*, 10, 647-668, 2010.
999
- 1000 Andreae, M. O. and Gelencsér, A.: Black carbon or brown carbon? The nature of light-absorbing
1001 carbonaceous aerosols, *Atmos. Chem. Phys.*, 6, 3131-3148, 2006.
1002
- 1003 Barlow, J. F., Dunbar, T. M., Nemitz, E. G., Wood, C. R., Gallagher, M. W., Davies, F., O'Connor,
1004 E. and Harrison, R. M.: Boundary layer dynamics over London, UK, as observed using Doppler
1005 lidar during REPARTEE-II, *Atmos. Chem. Phys.*, 11, 2111-2125, 2011.
1006
- 1007 Barlow, J. F., Halios, C. H., Lane, S. E. and Wood, C. R.: Observations of urban boundary layer
1008 structure during a strong urban heat island event, *Environ. Fluid Mech.*, 1-26, 2014.
1009
- 1010 Bigi, A. and Harrison R. M.: Analysis of the air pollution climate at a central urban background site,
1011 *Atmos. Environ.*, 44, 2004-2012, 2010.
1012
- 1013 Bohnenstengel, S. I., Belcher, S. E., Allan, J. D., Allen, G., Bacak, A., Bannan, T. J., Barlow, J. F.,
1014 Beddows, D. C. S., Bloss, W. J., Booth, A. M., Chemel, C., Coceal, O., Di Marco, C. F., Faloon, K.
1015 H., Fleming, Z., Furger, M., Geitl, J. K., Graves, R. R., Green, D. C., Grimmond, C. S. B., Halios,
1016 C., Hamilton, J. F., Harrison, R. M., Heal, M. R., Heard, D. E., Helfter, C., Herndon, S. C., Holmes,
1017 R. E., Hopkins, J. R., Jones, A. M., Kelly, F. J., Kotthaus, S., Langford, B., Lee, J. D., Leigh, R. J.,
1018 Lewis, A. C., Lidster, R. T., Lopez-Hilfiker, F. D., McQuaid, J. B., Mohr, C., Monks, P. S., Nemitz,
1019 E., Ng, N. L., Percival, C. J., Prévôt, A. S. H., Ricketts, H. M. A., Sokhi, R., Stone, D., Thornton, J.
1020 A., Tremper, A. H., Valach, A. C., Visser, S., Whalley, L. K., Williams, L. R., Xu, L., Young, D. E.
1021 and Zotter, P.: Meteorology, air quality, and health in London: The ClearLo project, *B. Am.*
1022 *Meteorol. Soc.*, doi:10.1175/bams-d-12-00245.1, 2014.
1023
- 1024 Bond, T. C., Anderson, T. L., and Campbell, D.: Calibration and intercomparison of filter-based
1025 measurements of visible light absorption by aerosols, *Aerosol Sci. Technol.*, 30, 582-600, 1999.
1026
- 1027 Carslaw, D. C.: The openair manual — open-source tools for analysing air pollution data, Manual
1028 for version 1.0, King's College London, 2014.
1029
- 1030 Carslaw, D. C. and Ropkins, K.: openair — An R package for air quality data analysis, *Environ.*
1031 *Modell. Softw.* 27–28, 52-61, 2012.
1032
- 1033 Caseiro, A., Bauer, H., Schmidl, C., Pio, C. A. and Puxbaum, H.: Wood burning impact on PM10
1034 in three Austrian regions, *Atmos. Environ.*, 43, 2186-2195, 2009.
1035
- 1036 Cavalli, F., Viana, M., Yttri, K.E., Genberg, J. and Putaud, J. P.: Toward a standardised thermal-
1037 optical protocol for measuring atmospheric organic and elemental carbon: the EUSAAR protocol,
1038 *Atmos. Meas. Tech.*, 3, 79-89, 2010.

1039 Charron, A., Birmili, W. and Harrison, R.M.: Factors influencing new particle formation at the
1040 rural site, Harwell, United Kingdom, *J. Geophys. Res.*, 112, D14210, 2007.
1041
1042 Charron, A., Degrendele, C., Laongsri, B. and Harrison, R.M.: Receptor modelling of secondary
1043 and carbonaceous particulate matter at a southern UK site, *Atmos. Chem. Phys.*, 13, 1879-1894,
1044 2013.
1045
1046 Cohen, A. J., Anderson, R. H., Ostro, B., Pandey, K. D., Krzyzanowski, M., Künzli, N.,
1047 Gutschmidt, K., Pope, A., Romieu, I., Samet, J. M. and Smith, K.: The Global Burden of Disease
1048 Due to Outdoor Air Pollution, *J. Toxicol. Environ. Health, Part A*, 68, 1301-1307, 2005.
1049
1050 Collaud Coen, M., Weingartner, E., Apituley, A., Ceburnis, D., Fierz-Schmidhauser, R., Flentje, H.,
1051 Henzing, J. S., Jennings, S. G., Moerman, M., Petzold, A., Schmid, O. and Baltensperger, U.:
1052 Minimizing light absorption measurement artifacts of the Aethalometer: evaluation of five
1053 correction algorithms, *Atmos. Meas. Tech.*, 3, 457-474, 2010.

1054 Crilley, L. R., Ayoko, G. A. and Morawska, L.: First measurements of source apportionment of
1055 organic aerosols in the Southern Hemisphere, *Environ. Pollut.*, 184, 81-88, 2014.
1056
1057 Crippa, M., DeCarlo, P. F., Slowik, J. G., Mohr, C., Heringa, M. F., Chirico, R., Poulain, L.,
1058 Freutel, F., Sciare, J., Cozic, J., Di Marco, C. F., Elsasser, M., Nicolas, J. B., Marchand, N., Abidi,
1059 E., Wiedensohler, A., Drewnick, F., Schneider, J., Borrmann, S., Nemitz, E., Zimmermann, R.,
1060 Jaffrezo, J. L., Prévôt, A. S. H. and Baltensperger, U.: Wintertime aerosol chemical composition
1061 and source apportionment of the organic fraction in the metropolitan area of Paris, *Atmos. Chem.*
1062 *Phys.*, 13, 961-981, 2013.
1063
1064 Despiiau, S. and Croci, D.: Concentrations and size distributions of fine aerosol particles measured
1065 at roof level in urban zone. *J. Geophys. Res.: Atmospheres*, 112, D09212, 2007, doi:
1066 10.1029/2006JD007228
1067
1068 El Haddad, I., D'Anna, B., Temime-Roussel, B., Nicolas, M., Boreave, A., Favez, O., Voisin, D.,
1069 Sciare, J., George, C., Jaffrezo, J. L., Wortham, H. and Marchand, N.: Towards a better
1070 understanding of the origins, chemical composition and aging of oxygenated organic aerosols: case
1071 study of a Mediterranean industrialized environment, Marseille, *Atmos. Chem. Phys.*, 13, 7875-
1072 7894, 2013.
1073
1074 Fahrni, S. M., Gäggeler, H. W., Hajdas, I., Ruff, M., Szidat, S. and Wacker, L.: Direct
1075 measurements of small ¹⁴C samples after oxidation in quartz tubes, *Nucl. Instr. Methods Phys.*
1076 *Res., Sec. B*, 268, 787-789, 2010.
1077
1078 Fuller, G. W., Sciare, J., Lutz, M., Moukhtar, S. and Wagener, S.: New Directions: Time to tackle
1079 urban wood burning?, *Atmos. Environ.*, 68, 295-296, 2013.
1080
1081 Fuller, G. W., Tremper, A. H., Baker, T. D., Yttri, K. E. and Butterfield, D.: Contribution of wood
1082 burning to PM₁₀ in London, *Atmos. Environ.*, 87, 87-94, 2014.
1083
1084 Gelencsér, A., May, B., Simpson, D., Sánchez-Ochoa, A., Kasper-Giebl, A., Puxbaum, H., Caseiro,
1085 A., Pio, C. and Legrand, M.: Source apportionment of PM_{2.5} organic aerosol over Europe:
1086 Primary/secondary, natural/anthropogenic, and fossil/biogenic origin, *J. Geophys. Res.*, 112,
1087 D23S04, 2007.
1088

1089 Gyawali, M., Arnott, W. P., Lewis, K. and Moosmüller, H.: In situ aerosol optics in Reno, NV,
1090 USA during and after the summer 2008 California wildfires and the influence of absorbing and non-
1091 absorbing organic coatings on spectral light absorption, *Atmos. Chem. Phys.*, 9, 8007-8015, 2009.

1092 Harrison, R. M., Beddows, D. C. S., Hu, L. and Yin, J.: Comparison of methods for evaluation of
1093 wood smoke and estimation of UK ambient concentrations, *Atmos. Chem. Phys.*, 12, 8271-8283,
1094 2012a.

1095
1096 Harrison, R. M., Dall'Osto, M., Beddows, D. C. S., Thorpe, A. J., Bloss, W. J., Allan, J. D., Coe, H.,
1097 Dorsey, J. R., Gallagher, M., Martin, C., Whitehead, J., Williams, P. I., Jones, R. L., Langridge, J.
1098 M., Benton, A. K., Ball, S. M., Langford, B., Hewitt, C. N., Davison, B., Martin, D., Petersson, K.
1099 F., Henshaw, S. J., White, I. R., Shallcross, D. E., Barlow, J. F., Dunbar, T., Davies, F., Nemitz, E.,
1100 Phillips, G. J., Helfter, C., Di Marco, C. F. and Smith, S.: Atmospheric chemistry and physics in
1101 the atmosphere of a developed megacity (London): an overview of the REPARTEE experiment and
1102 its conclusions, *Atmos. Chem. Phys.*, 12, 3065-3114, 2012b.

1103
1104 Harrison, R. M., Beddows, D. C. S., Jones, A. M., Calvo, A. and Alves, C., Pio, C.: An evaluation
1105 of some issues regarding the use of aethalometers to measure woodsmoke concentrations, *Atmos.*
1106 *Environ.*, 80, 540-548, 2013.

1107
1108 Heal, M. R., Naysmith, P., Cook, G. T., Xu, S., Duran, T. R. and Harrison, R. M.: Application of
1109 ¹⁴C analyses to source apportionment of carbonaceous PM_{2.5} in the UK, *Atmos. Environ.*, 45,
1110 2341-2348, 2011.

1111
1112 Heal, M.R.: The application of carbon-14 analyses to the source apportionment of atmospheric
1113 carbonaceous particulate matter: a review, *Anal. Bioanal. Chem.*, 406, 81-98, 2014.

1114
1115 Herich, H., Hueglin, C. and Buchmann, B.: A 2.5 year's source apportionment study of black
1116 carbon from wood burning and fossil fuel combustion at urban and rural sites in Switzerland,
1117 *Atmos. Meas. Tech.*, 4, 1409-1420, 2011.

1118
1119 Heringa, M. F., DeCarlo, P. F., Chirico, R., Tritscher, T., Dommen, J., Weingartner, E., Richter, R.,
1120 Wehrle, G., Prevot, A. S. H. and Baltensperger, U.: Investigations of primary and secondary
1121 particulate matter of different wood combustion appliances with a high-resolution time-of-flight
1122 aerosol mass spectrometer, *Atmos. Chem. Phys.*, 11, 5945-5957, 2011.

1123
1124 Hodzic, A., Jimenez, J. L., Prévôt, A. S. H., Szidat, S., Fast, J. D. and Madronich, S.: Can 3-D
1125 models explain the observed fractions of fossil and non-fossil carbon in and near Mexico City?,
1126 *Atmos. Chem. Phys.*, 10, 10997-11016, 2010.

1127
1128 Kampa, M. and Castanas, E.: Human health effects of air pollution, *Environm. Pollut.*, 151, 362-
1129 367, 2008.

1130
1131 Khalil, M. A. K. and Rasmussen, R. A.: Tracers of wood smoke, *Atmos. Environ.*, 37, 1211-1222,
1132 2003.

1133
1134 Kirchstetter, T. W., Novakov, T. and Hobbs, P. V.: Evidence that the spectral dependence of light
1135 absorption by aerosols is affected by organic carbon, *J. Geophys. Res.-Atmos.*, 109, D21208, doi:
1136 10.1029/2004JD004999, 2004.

1137

1138 Lack, D. A., Bahreini, R., Langridge, J. M., Gilman, J. B., Middlebrook, A. M.: Brown carbon
1139 absorption linked to organic mass tracers in biomass burning particles, *Atmos. Chem. Phys.*, 13,
1140 2415-2422, 2013.
1141
1142 Lanz, V. A., Prévôt, A. S. H., Alfarra, M. R., Weimer, S., Mohr, C., DeCarlo, P. F., Gianini, M. F.
1143 D., Hueglin, C., Schneider, J., Favez, O., D'Anna, B., George, C., and Baltensperger, U.:
1144 Characterization of aerosol chemical composition with aerosol mass spectrometry in Central
1145 Europe: an overview, *Atmos. Chem. Phys.*, 10, 10453-10471, 2010.
1146
1147 Lioussé, C., Cachier, H., and Jennings, S. G.: Optical and thermal measurements of black carbon
1148 aerosol content in different environments: Variation of the specific attenuation cross-section, σ
1149 (σ), *Atmos. Environ.*, 27, 1203-1211, 1993.
1150
1151 Liu, J., Bergin, M., Guo, H., King, L., Kotra, N., Edgerton, E. and Weber, R. J.: Size-resolved
1152 measurements of brown carbon in water and methanol extracts and estimates of their contribution to
1153 ambient fine-particle light absorption, *Atmos. Chem. Phys.*, 13, 12389-12404, 2013.
1154
1155 Massoli, P., Fortner, E. C., Canagaratna, M. R., Williams, L. R., Zhang, Q., Sun, Y., Schwab, J. J.,
1156 Trimborn, A., Onasch, T. B., Demerjian, K. L., Kolb, C. E., Worsnop, D. R., Jayne, J. T.: Pollution
1157 gradients and chemical characterization of particulate matter from vehicular traffic near major
1158 roadways: Results from the 2009 Queens College air quality study in NYC, *Aerosol Sci. Technol.*,
1159 46, 1201-1218, 2012.
1160
1161 Minguillón, M. C., Perron, N., Querol, X., Szidat, S., Fahrni, S. M., Alastuey, A., Jimenez, J. L.,
1162 Mohr, C., Ortega, A. M., Day, D. A., Lanz, V. A., Wacker, L., Reche, C., Cusack, M., Amato, F.,
1163 Kiss, G., Hoffer, A., Decesari, S., Moretti, F., Hillamo, R., Teinilä, K., Seco, R., Peñuelas, J.,
1164 Metzger, A., Schallhart, S., Müller, M., Hansel, A., Burkhardt, J. F., Baltensperger, U., Prévôt,
1165 A.S.H.: Fossil versus contemporary sources of fine elemental and organic carbonaceous particulate
1166 matter during the DAURE campaign in Northeast Spain, *Atmos. Chem. Phys.*, 11, 12067-12084,
1167 2011.
1168
1169 Mohr, C., Lopez-Hilfiker, F. D., Zotter, P., Prévôt, A. S. H., Xu, L., Ng, N. L., Herndon, S. C.,
1170 Williams, L. R., Franklin, J. P., Zahniser, M. S., Worsnop, D. R., Knighton, W. B., Aiken, A. C.,
1171 Gorkowski, K. J., Dubey, M. K., Allan, J. D. and Thornton, J. A.: Contribution of nitrated phenols
1172 to wood burning brown carbon light absorption in Detling, United Kingdom during winter time,
1173 *Environ. Sci. Technol.*, 47, 6316-6324, 2013.
1174
1175 Moosmüller, H., Chakrabarty, R. K. and Arnott, W. P.: Aerosol light absorption and its
1176 measurement: A review, *J. Quant. Spectrosc. Radiat. Transfer*, 110, 844-878, 2009.
1177
1178 Pant, P., Yin, J. and Harrison, R. M.: Sensitivity of a Chemical Mass Balance model to different
1179 molecular marker traffic source profiles, *Atmos. Environ.*, 82, 238-249, 2014.
1180
1181 Petzold, A., Kopp, C., and Niessner, R.: The dependence of the specific attenuation cross-section
1182 on black carbon mass fraction and particle size, *Atmos. Environ.*, 31, 661-672, 1997.
1183
1184 Petzold, A., Ogren, J. A., Fiebig, M., Laj, P., Li, S. M., Baltensperger, U., Holzer-Popp, T., Kinne,
1185 S., Pappalardo, G., Sugimoto, N., Wehrli, C., Wiedensohler, A., and Zhang, X. Y.:
1186 Recommendations for reporting "black carbon" measurements, *Atmos. Chem. Phys.*, 13, 8365-
1187 8379, 2013.
1188
1189

1190 Pio, C., Cerqueira, M., Harrison, R. M., Nunes, T., Mirante, F., Alves, C., Oliveira, C., Sanchez de
1191 la Campa, A., Artíñano, B., Matos, M.: OC/EC ratio observations in Europe: Re-thinking the
1192 approach for apportionment between primary and secondary organic carbon, *Atmos. Environ.*, 45,
1193 6121-6132, 2011.

1194
1195 Puxbaum, H., Caseiro, A., Sánchez-Ochoa, A., Kasper-Giebl, A., Claeys, M., Gelencsér, A.,
1196 Legrand, M., Preunkert, S. and Pio, C.: Levoglucosan levels at background sites in Europe for
1197 assessing the impact of biomass combustion on the European aerosol background, *J. Geophys. Res.*,
1198 112, D23S05, 2007.

1199
1200 Qadir, R. M., Abbaszade, G., Schnelle-Kreis, J., Chow, J. C. and Zimmermann, R.: Concentrations
1201 and source contributions of particulate organic matter before and after implementation of a low
1202 emission zone in Munich, Germany, *Environ. Pollut.*, 175, 158-167, 2013.

1203
1204 Reche, C., Querol, X., Alastuey, A., Viana, M., Pey, J., Moreno, T., Rodríguez, S., González, Y.,
1205 Fernández-Camacho, R., de la Rosa, J., Dall'Osto, M., Prévôt, A. S. H., Hueglin, C., Harrison, R.
1206 M. and Quincey, P.: New considerations for PM, Black Carbon and particle number concentration
1207 for air quality monitoring across different European cities, *Atmos. Chem. Phys.*, 11, 6207-6227,
1208 2011.

1209
1210 Reche, C., Viana, M., Amato, F., Alastuey, A., Moreno, T., Hillamo, R., Teinilä, K., Saarnio, K.,
1211 Seco, R., Peñuelas, J., Mohr, C., Prévôt, A.S.H., Querol, X.: Biomass burning contributions to
1212 urban aerosols in a coastal Mediterranean City, *Sci. Tot. Environ.*, 427-428, 175-190, 2012.

1213
1214 Ricardo-AEA: QA/QC Data Ratification Report for the Automatic Urban and Rural Network,
1215 January-March 2013, and Intercalibration Report, Winter 2013, Report for Department for
1216 Environment, Food and Rural Affairs, The Scottish Government, The Welsh Government, The
1217 Northern Ireland Department of Environment, Ricardo-AEA/R/3382, Issue 1, UK, 2013.

1218
1219 Rizzo, L.V., Correia, A.L., Artaxo, P., Procópio, A.S., Andreae, M.O.: Spectral dependence of
1220 aerosol light absorption over the Amazon Basin, *Atmos. Chem. Phys.*, 11, 8899-8912, 2011.

1221
1222 Ruff, M., Wacker, L., Gaggeler, H. W., Suter, M., Synal, H. A. and Szidat, S.: A gas ion source for
1223 radiocarbon measurements at 200 kV, *Radiocarbon*, 49, 307-314, 2007.

1224
1225 Ruff, M., Szidat, S., Gaggeler, H. W., Suter, M., Synal, H. A. and Wacker, L.: Gaseous
1226 radiocarbon measurements of small samples, *Nucl. Instr. Methods Phys. Res., Sec. B*, 268, 790-794,
1227 2010.

1228
1229 Saarikoski, S. K., Sillanpää, M. K., Saarnio, K. M., Hillamo, R. E., Pennanen, A. S. and Salonen, R.
1230 O.: Impact of biomass combustion on urban fine particulate matter in central and northern Europe,
1231 *Water Air Soil Poll.*, 191, 265-277, 2008.

1232
1233 Saleh, R., Hennigan, C. J., McMeeking, G. R., Chuang, W. K., Robinson, E. S., Coe, H., Donahue,
1234 N. M. and Robinson, A. L.: Absorptivity of brown carbon in fresh and photo-chemically aged
1235 biomass-burning emissions, *Atmos. Chem. Phys.*, 13, 7683-7693, 2013.

1236
1237 Sandradewi, J., Prévôt, A. S. H., Szidat, S., Perron, N., Alfarra, M. R., Lanz, V. A., Weingartner, E.
1238 and Baltensperger, U.: Using aerosol light absorption measurements for the quantitative
1239 determination of wood burning and traffic emission contributions to particulate matter, *Environ.*
1240 *Sci. Technol.*, 42, 3316-3323, 2008a.

1241

1242 Sandradewi, J., Prévôt, A. S. H., Weingartner, E., Schmidhauser, R., Gysel, M. and Baltensperger,
1243 U.: A study of wood burning and traffic aerosols in an Alpine valley using a multi-wavelength
1244 Aethalometer, *Atmos. Environ.*, 42, 101-112, 2008b.

1245
1246 Schauer, J. J., Kleeman, M. J., Cass, G. R., Simoneit, B. R. T.: Measurement of emissions from air
1247 pollution sources. 3. C1–C29 organic compounds from fireplace combustion of wood, *Environ. Sci.*
1248 *Technol.*, 35, 1716-1728, 2001.

1249
1250 Schmidl, C., Marr, I. L., Caseiro, A., Kotianová, P., Berner, A., Bauer, H., Kasper-Giebl, A. and
1251 Puxbaum, H.: Chemical characterisation of fine particle emissions from wood stove combustion of
1252 common woods growing in mid-European Alpine regions, *Atmos. Environ.*, 42, 126-141, 2008.

1253 Simoneit, B. R. T.: Biomass burning — a review of organic tracers for smoke from incomplete
1254 combustion, *Appl. Geochem.*, 17, 129-162, 2002.

1255
1256 Stuiver, M., and Polach, H. A.: Reporting of C-14 data – discussion, *Radiocarbon*, 19, 355-36,
1257 1977.

1258
1259 Sullivan, A. P., Holden, A. S., Patterson, L. A., McMeeking, G. R., Kreidenweis, S. M., Malm, W.
1260 C., Hao, W. M., Wold, C. E. and Collett, J. L.: A method for smoke marker measurements and its
1261 potential application for determining the contribution of biomass burning from wildfires and
1262 prescribed fires to ambient PM_{2.5} organic carbon, *J. Geophys. Res.*, 113, D22302,
1263 doi:10.1029/2008JD010216, 2008.

1264
1265 Synal, H. A., Stocker, M. and Suter, M.: MICADAS: A new compact radiocarbon AMS system,
1266 *Nucl. Instr. Methods Phys. Res., Sec. B*, 259, 7-13, 2007.

1267
1268 Szidat, S., Jenk, T. M., Gäggeler, H. W., Synal, H. A., Hajdas, I., Bonani, G. and Saurer, M.:
1269 THEODORE, a two-step heating system for the EC/OC determination of radiocarbon (¹⁴C) in the
1270 environment, *Nucl. Instr. Methods Phys. Res., Sec. B*, 223-224, 829-836, 2004.

1271
1272 Szidat, S., Jenk, T. M., Synal, H. A., Kalberer, M., Wacker, L., Hajdas, I., Kasper-Giebl, A. and
1273 Baltensperger, U.: Contributions of fossil fuel, biomass-burning, and biogenic emissions to
1274 carbonaceous aerosols in Zurich as traced by ¹⁴C, *J. Geophys. Res.*, 111, D07206, 2006.

1275
1276 Turpin, B. J. and Lim, H.-J.: Species contributions to PM_{2.5} mass concentrations: Revisiting
1277 common assumptions for estimating organic mass, *Aerosol Sci. Technol.*, 35, 602-610, 2001.

1278
1279 Viana, M., Kuhlbusch, T.A.J., Querol, X., Alastuey, A., Harrison, R. M., Hopke, P. K., Winiwarter,
1280 W., Vallius, M., Szidat, S., Prévôt, A. S. H., Hueglin, C., Bloemen, H., Wählén, P., Vecchi, R.,
1281 Miranda, A. I., Kasper-Giebl, A., Maenhaut, W. and Hitzenberger, R.: Source apportionment of
1282 particulate matter in Europe: A review of methods and results, *J. Aerosol Sci.*, 39, 827-849, 2008.

1283
1284 Viana, M., Reche, C., Amato, F., Alastuey, A., Querol, X., Moreno, T., Lucarelli, F., Nava, S.,
1285 Cazolai, G., Chiari, M. and Rico, M.: Evidence of biomass burning aerosols in the Barcelona urban
1286 environment during winter time, *Atmos. Environ.*, 72, 81-88, 2013.

1287
1288 Wacker, L., Christl, M. and Synal, H. A.: Bats: A new tool for AMS data reduction, *Nucl. Instr.*
1289 *Methods Phys. Res., Sec. B*, 268, 976-979, 2010.

1290 Wacker, L., Fahrni, S. M., Hajdas, I., Molnar, M., Synal, H. A., Szidat, S. and Zhang, Y. L.: A
1291 versatile gas interface for routine radiocarbon analysis with a gas ion source, *Nucl. Instr. Methods*
1292 *Phys. Res., Sec. B*, 294, 315-319, 2013.

1293
1294 Wagener, S., Langner, M., Hansen, U., Moriske, H.-J. and Endlicher, W. R.: Spatial and seasonal
1295 variations of biogenic tracer compounds in ambient PM₁₀ and PM₁ samples in Berlin, Germany,
1296 *Atmos. Environ.*, 47, 33-42, 2012.

1297 Wang, Y., Hopke, P. K., Rattigan, O. V., Xia, X., Chalupa, D. C.: Utell, M. J., Characterization of
1298 residential wood combustion particles using the two-wavelength aethalometer, *Environ. Sci.*
1299 *Technol.*, 45, 7387-7393, 2011a.

1300
1301 Wang, Y., Hopke, P. K., Rattigan, O. V., Zhu, Y.: Characterization of ambient black carbon and
1302 wood burning particles in two urban areas, *J. Environ. Monitor.*, 13, 1919-1926, 2011b.

1303
1304 Weingartner, E., Saathoff, H., Schnaiter, M., Streit, N., Bitnar, B. and Baltensperger, U.:
1305 Absorption of light by soot particles: determination of the absorption coefficient by means of
1306 aethalometers, *J. Aerosol Sci.*, 34, 1445-1463, 2003.

1307
1308 Yin, J., Harrison, R. M., Chen, Q., Rutter, A., Schauer, J. J.: Source apportionment of fine particles
1309 at urban background and rural sites in the UK atmosphere, *Atmos. Environ.*, 44, 841-851, 2010.

1310
1311 Yttri, K. E., Dye, C., Slørdal, L. H., Braathen, O.-A.: Quantification of monosaccharide anhydrides
1312 by liquid chromatography combined with mass spectrometry: Application to aerosol samples from
1313 an urban and a suburban site influenced by small-scale wood burning, *JAWMA*, 55, 1169-1177,
1314 2005.

1315
1316 Zapf, A., Nesje, A., Szidat, S., Wacker, L. and Schwikowski, M.: C-14 measurements of ice
1317 samples from the Juvfonne ice tunnel, Jotunheimen, southern Norway-validation of a C-14 dating
1318 technique for glacier ice, *Radiocarbon*, 55, 571-578, 2013.

1319
1320 Zhang, X., Lin, Y.-H., Surratt, J. D. and Weber, R. J.: Sources, composition and absorption
1321 ångström exponent of light-absorbing organic components in aerosol extracts from the Los Angeles
1322 Basin, *Environ. Sci. Technol.*, 47, 3685-3693, 2013.

1323
1324 Zhu, Y., Hinds, W. C., Shen, S. and Sioutas, C.: Seasonal trends of concentration and size
1325 distribution of ultrafine particles near major highways in Los Angeles Special Issue of *Aerosol*
1326 *Science and Technology on Findings from the Fine Particulate Matter Supersites Program*, *Aerosol*
1327 *Sci. Technol.*, 38, 5-13, 2004.

1328
1329 Zotter, P., El-Haddad, I., Zhang, Y., Hayes, P. L., Zhang, X., Lin, Y.-H., Wacker, L., Schnelle-
1330 Kreis, J., Abbaszade, G., Zimmermann, R., Surratt, J. D., Weber, R., Jimenez, J. L., Szidat, S.,
1331 Baltensperger, U. and Prévôt, A. S. H.: Diurnal cycle of fossil and non-fossil carbon using
1332 radiocarbon analyses during CalNex, *J. Geophys. Res.*, doi: 10.1002/2013JD021114, 2014a.

1333
1334 Zotter, P., Ciobanu, V. G., Zhang, Y. L., El-Haddad, I., Macchia, M., Daellenbach, K. R., Salazar,
1335 G. A., Huang, R. J., Wacker, L., Hueglin, C., Piazzalunga, A., Fermo, P., Schwikowski, M.,
1336 Baltensperger, U., Szidat, S. and Prévôt, A. S. H.: Radiocarbon analysis of elemental and organic
1337 carbon in Switzerland during winter-smog episodes from 2008 to 2012 – Part 1: Source
1338 apportionment and spatial variability, *Atmos. Chem. Phys. Discuss.*, 14, 15591-15643, 2014b.

1337 **TABLE LEGENDS**

- 1338 **Table 1:** Summary of sampling methodology for the filter analyses in the winter campaign.
1339
- 1340 **Table 2:** Summary statistics for selected species at Harwell, NK and Detling during the winter
1341 campaign. BC concentrations from a 2W AE. Variability given for the average is 1
1342 standard deviation
1343
- 1344 **Table 3:** Average of daily percentage fractions of non-fossil and fossil TC for the tracer and
1345 ^{14}C methods. Variability shown is 1 standard deviation. *Denotes the averages for
1346 the sampling days with both ^{14}C and EC OC data.
1347
- 1348 **Table 4:** Average of daily percentage fraction of non-fossil and fossil TC using different
1349 source ratios. $\text{OC}/\text{EC}_{\text{veh}}$ and $\text{OC}/\text{EC}_{\text{bb}}$ refers to the source ratio of primary vehicle
1350 and biomass burning emissions, respectively. Variability shown is 1 standard
1351 deviation.
1352
- 1353 **Table 5:** Percentage contributions of biomass burning and traffic emissions to total EC or BC
1354 concentration at 880 nm as determined by the tracer method and Aethalometer
1355 model, respectively. Variability shown is 1 standard deviation.
1356
- 1357 **Table 6:** Average of daily results for selected ratios. Variability shown is 1 standard deviation.
1358
- 1359 **Table 7:** Summary statistics of the long-term measurement at NK and the BT tower. Note the
1360 values given for 7W AE at NK is the combined data from summer and winter
1361 campaigns.
1362
1363

1364 **FIGURE LEGENDS**

- 1365 **Figure 1:** Map of the study area with the sampling sites (indicated with circles) and lidar
1366 locations (designated with triangles) indicated. Note for the start of the winter IOP
1367 the lidar operated from the Westminster City Council (WCC) building and then
1368 moved to Imperial College London (ICL). During the summer IOP the lidar operated
1369 from the North Kensington (NK) site.
1370
- 1371 **Figure 2:** Times series of the temperature at three Met stations used for the entire winter
1372 campaign.
1373
- 1374 **Figure 3:** Times series of the concentration of (A) levoglucosan, (B) OC, (C) EC and (D) K^+ at
1375 Harwell, NK and Detling. No potassium data is available for Detling.
1376
- 1377 **Figure 4:** Average of daily concentrations of EC_{bb} , EC_{ff} , OC_{bb} , OC_{ff} and OC_{SOA} apportioned
1378 using the tracer method (Gelencsér et al., 2007). Variability shown is 1 standard
1379 deviation.
1380
- 1381 **Figure 5:** Mean diurnal variations in the BC concentrations at 880 nm during the winter
1382 campaign as measured by a 2W AE with the shaded areas indicating the 95%
1383 confidence intervals. Note in the key HAR and DET represents Harwell and Detling,
1384 respectively.
1385

- 1386 **Figure 6:** Mean diurnal cycles of UVPM during the winter campaign as determined by 2W AE
 1387 (Column A) and 7W AE (Column B) for the sites where the data was available.
 1388 Shaded areas indicate the 95% confidence intervals. In the key HAR and DET
 1389 represents Harwell and Detling, respectively.
 1390
- 1391 **Figure 7:** Mean diurnal cycle of concentration of BC_{wb} and BC_{tr} at 880 nm during the winter
 1392 campaign at NK (A) and Detling (B) derived from the Aethalometer model. The
 1393 shaded areas represent the 95% confidence interval. Note the different y-axis scales
 1394 used between panels.
 1395
- 1396 **Figure 8:** Polar plots of CPF analysis (90th percentile) for the BC concentrations at 880 nm
 1397 measured with a 2W AE at the sites during the winter campaign. Wind speed (wind
 1398 spd) is in $m s^{-1}$.
 1399
- 1400 **Figure 9:** Polar plots of the mean concentration ($\mu g m^{-3}$) of BC from traffic (BC_{tr}) and wood
 1401 burning (BC_{wb}) calculated by the two component Aethalometer model at North
 1402 Kensington (top row) and Detling (bottom row) for the winter campaign. Wind speed
 1403 is in $m s^{-1}$.
 1404
- 1405 **Figure 10:** Mean diurnal cycles of the mean (A) and normalised (B) BC concentrations (880
 1406 nm) over the whole sampling period at the London sites. For plot B the values were
 1407 normalised to the mean. Note BC concentrations were measured by a 2W AE at NK
 1408 and by 7W AE at the BT tower (BT), with the same correction factor applied to both
 1409 instrument datasets.
 1410
- 1411 **Figure 11:** CPF analysis for BC concentrations at 880 nm over the whole sampling period at NK
 1412 and the BT tower (BT). Note the NK concentrations were determined with a 2W AE
 1413 and the BT tower concentrations with a 7W AE.
 1414
- 1415 **Figure 12:** Histogram of Ångstrom coefficients for winter and summer campaigns at NK and the
 1416 BT tower.
 1417
- 1418 **Figure 13:** Time series of NO_x concentration (A) and BC concentration measured at 370 and
 1419 880 nm (B) during a de-coupling event as marked by the turbulent mixing heights
 1420 (TMH). For times where there are no TMH data in the figure, this is due to the
 1421 turbulent layer lying below the lowest range gate of the Doppler lidar i.e. minimum
 1422 resolvable TMH is 139 m. Also shown for this de-coupling event is the sensible heat
 1423 flux (H) and turbulent kinetic energy normalized by the wind speed (tke/U^2) as
 1424 measured at the BT tower (C). Note the dashed line represents the height of the T35
 1425 sampling level of the aethalometers (160 m).
 1426
- 1427 **Figure 14:** Time series of the calculated absorption Ångstrom coefficient (α) during the de-
 1428 coupling event from Figure 13.
 1429
- 1430 **Figure 15:** Histogram of hourly of Ångstrom coefficient for BT tower for (A) all de-coupling
 1431 periods and (B) for the other nights.
 1432
 1433

1434 Table 1: Summary of sampling methodology for the filter analyses in the winter campaign.

		Levoglucosan	K ⁺	OC/EC	¹⁴ C
Harwell	Size Fraction	PM _{2.5}	PM _{2.5}	PM _{2.5}	N/A
	Sampling Interval	24 hr	24 hr	24 hr	
	Start time	Noon	Noon	Noon	
	Dates	12/01 – 9/02	12/01 – 9/02	12/01 – 9/02	
NK	Size Fraction	PM _{2.5}	PM _{2.5}	PM _{2.5}	PM _{2.5}
	Sampling Interval	24 hr	24 hr	24 hr	24 hr
	Start time	Noon	Noon	Noon	Noon
	Dates	12/01 – 8/02	12/01 – 8/02	12/01 – 8/02	31/01 – 07/02
Detling	Size Fraction	PM ₁₀	N/A	PM ₁	PM ₁
	Sampling Interval	24 hr	N/A	24 hr	24 hr
	Start time	Midnight	N/A	Midnight	Midnight
	Dates	12/01 – 9/02	N/A	12/01 – 9/02	12/01 – 9/02

1435

1436

1437 Table 2: Summary statistics for selected species at Harwell, NK and Detling during the winter
 1438 campaign. BC concentrations from a 2W AE. Variability given for the average is 1 standard
 1439 deviation

		Levoglucosan	K ⁺	OC	EC	BC	PM _{2.5}
		<i>ng m⁻³</i>		<i>μg m⁻³</i>			
Harwell	Average	92 ± 51.6	65 ± 47	2.3 ± 1.8	0.7 ± 0.5	0.7 ± 0.6	14 ± 11
	Range	18.3 - 219	7.9 - 189	0.4 - 6.1	0.03 - 2.2	0.06 – 3.8	0 - 63
	n	28	28	28	28	3131	831
NK	Average	77.6 ± 35.9	82.8 ± 48.6	3.5 ± 2.2	1.5 ± 0.9	1.4 ± 1.3	20 ± 12
	Range	27.8 - 155	25 - 210	1.0 - 8.6	0.4 - 4.1	0.07 – 15	2 - 65
	n	27	27	27	27	3345	798
Detling	Average	69.7 ± 62.7	N/A	2.2 ± 1.6	0.6 ± 0.4	0.9 ± 0.8	16 ± 11
	Range	5.5 - 264		0.3 - 5.9	0.2 - 1.3	0.07 – 5.0	0 - 58
	n	28		26	26	2286	441

1440

1441

1442 Table 3: Average of daily percentage fractions of non-fossil and fossil TC for the tracer and ^{14}C
 1443 methods. Variability shown is 1 standard deviation. *Denotes the averages for the sampling days
 1444 with both ^{14}C and EC OC data.

	Method	TC_{nf} (%)	TC_{ff} (%)	n
	Tracer	62 ± 18	38 ± 18	28
Detling	Tracer*	69 ± 13	31 ± 13	12
	^{14}C	64 ± 7	36 ± 7	12
NK	Tracer	54 ± 11	46 ± 11	27
	Tracer*	58 ± 9	42 ± 9	6
	^{14}C	53 ± 6	47 ± 6	6
Harwell	Tracer	72 ± 9	28 ± 9	26

1445

1446 Table 4: Average of daily percentage fraction of non-fossil and fossil TC using different source ratios. OC/EC_{veh} and OC/EC_{bb} refers to the source ratio
 1447 of primary vehicle and biomass burning emissions, respectively. Variability shown is 1 standard deviation.

<i>ID</i>	<i>Source ratio applied</i>			<i>Site</i>	<i>TC_{nf}(%)</i>	<i>TC_{ff}(%)</i>	
	Levoglucosan/ OC	OC/EC_{veh}	OC/EC_{bb}				
A	0.136	0.58	6	Detling	62 ± 18	38 ± 18	Average source ratio (Gelencsér et al., 2007)
				NK	54 ± 11	46 ± 11	
B	0.07	0.3	3	Detling	83 ± 13	17 ± 13	Min. source ratio
				NK	70 ± 9	30 ± 9	
C	0.17	1	10	Detling	48 ± 23	52 ± 23	Max. source ratio
				NK	39 ± 14	61 ± 14	
D	0.136	0.3	6	Detling	69 ± 14	31 ± 14	Pio et al. (2011) OC/EC _{veh}
				NK	62 ± 9	38 ± 9	
E	0.09	0.58	6	Detling	65 ± 17	35 ± 17	Harrison et al. (2012a) levoglucosan/OC
				NK	55 ± 11	45 ± 11	
F	0.09	0.3	6	Detling	71 ± 17	29 ± 14	
				NK	63 ± 9	37 ± 9	
G	0.098	0.58	8.2	Detling	62 ± 20	38 ± 18	Swiss biomass burning ratios (Zotter et al., 2014b)
				NK	54 ± 11	46 ± 11	

1448 Note: TC_{nf} = OC_{bb} + EC_{bb} + OC_{SOA} and TC_{ff} = EC_{ff} + OC_{ff}

1449

1450 Table 5: Percentage contributions of biomass burning and traffic emissions to total EC or BC
 1451 concentration at 880 nm as determined by the tracer method and Aethalometer model, respectively.
 1452 Variability shown is 1 standard deviation.

	EC_{bb}	BC_{wb}	EC_{ff}	BC_{tr}
Harwell	24% ± 16%	N/A	76% ± 16%	N/A
NK	7% ± 2%	15% ± 12%	93% ± 2%	85% ± 12%
Detling	14% ± 16%	30% ± 13%	82% ± 20%	70% ± 13%

1453

1454

1455 Table 6: Average of daily results for selected ratios. Variability shown is 1 standard deviation.

	Levoglucosan/K⁺	Levoglucosan/OC	OC/EC
Harwell	1.1 ± 0.14	0.05 ± 0.02	3.9 ± 1.8
NK	0.73 ± 0.09	0.02 ± 0.01	2.3 ± 0.8
Detling	N/A	0.03 ± 0.02	3.2 ± 1.5

1456

1457

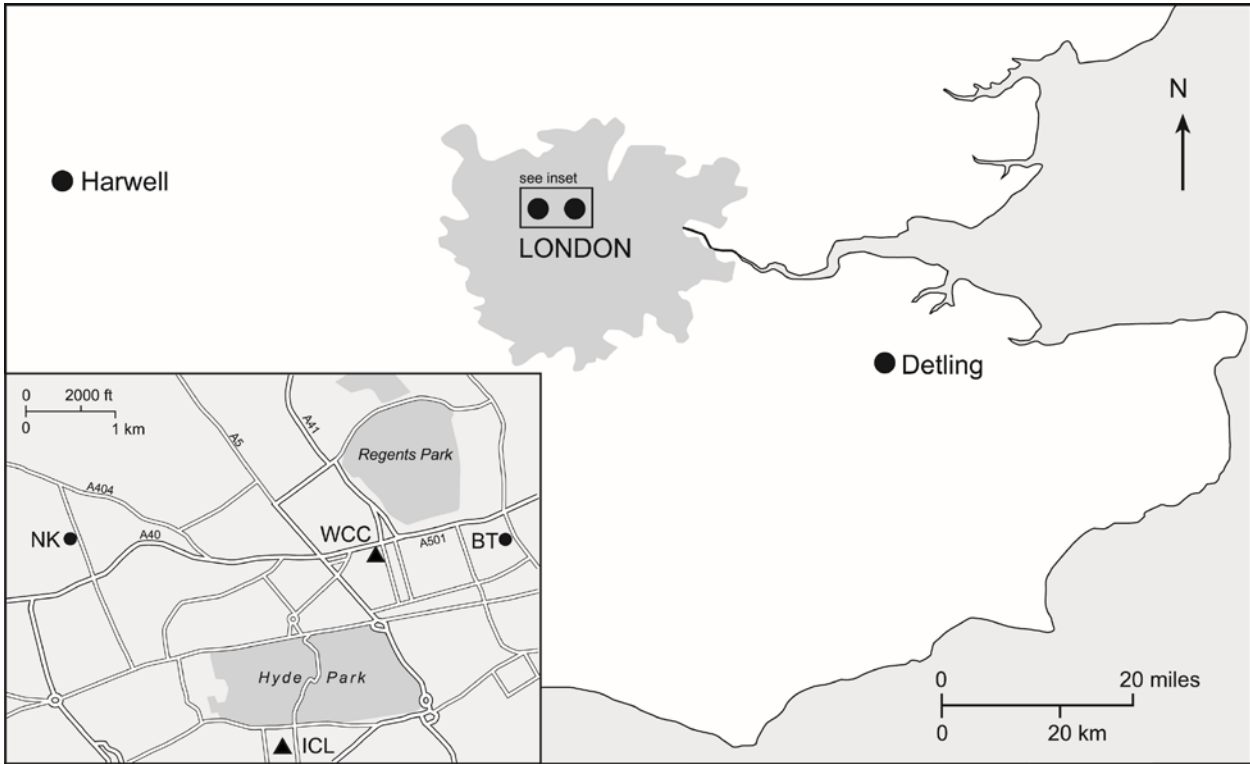
1458

1459 Table 7: Summary statistics of the long-term measurement at NK and the BT tower. Note the values
 1460 given for 7W AE at NK is the combined data from summer and winter campaigns.

		BC 7W AE	BC 2W AE	NO	NO₂	CO	O₃
		$\mu\text{g m}^{-3}$	$\mu\text{g m}^{-3}$	ppb	ppb	ppm	ppb
NK	Average	1.2	1.17	18.4	10.4	0.21	20.1
	Std. Dev.	1.1	1.11	11.6	24.5	0.13	13.2
	Min.	0	0	0	0	0	0
	Max.	24.7	15.6	121	471	1.85	96
	n	16386	50353	50784	50784	50791	51072
BT tower	Average	0.53		4.9	16.4	0.18	23.8
	Std. Dev.	0.42		12.8	17.1	0.075	13.8
	Min.	0	N/A	0	0	0	0
	Max.	6.23		355	249	0.72	99.6
	n	122394		90247	72806	7251	105643

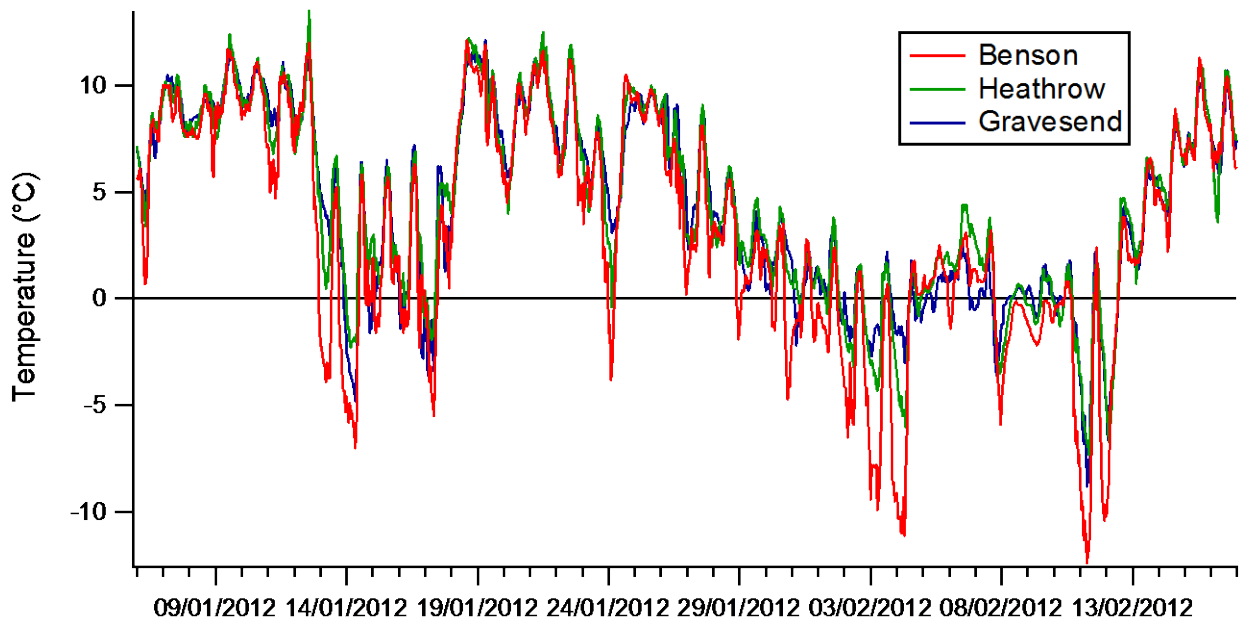
1461

1462



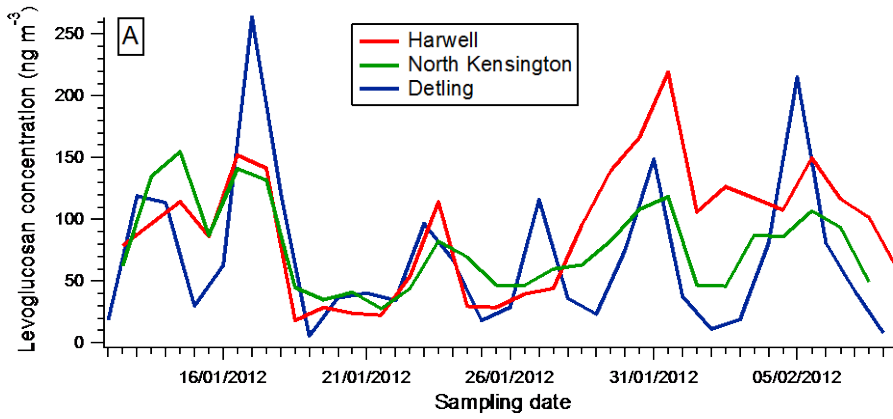
1464
 1465
 1466
 1467
 1468
 1469
 1470
 1471

Figure 1: Map of the study area with the sampling sites (indicated with circles) and lidar locations (designated with triangles) indicated. Note for the start of the winter IOP the lidar operated from the Westminster City Council (WCC) building and then moved to Imperial College London (ICL). During the summer IOP the lidar operated from the North Kensington (NK) site.

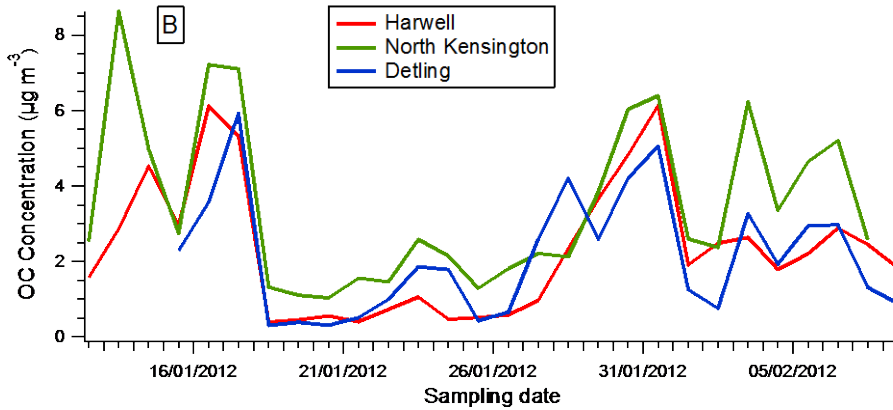


1472
1473

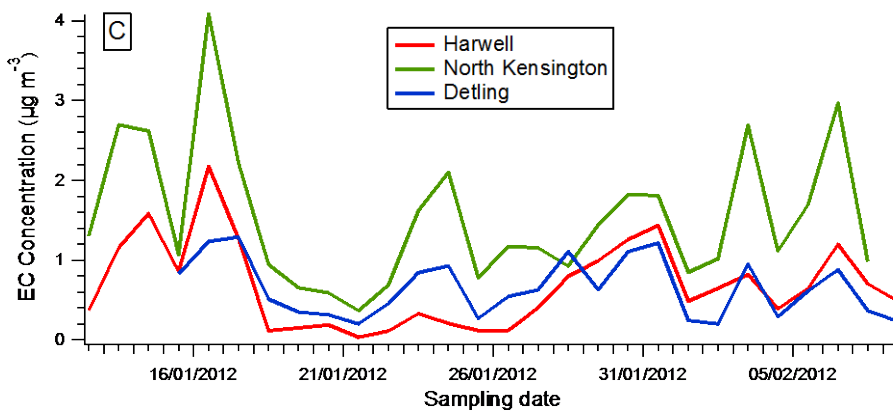
Figure 2: Times series of the temperature at three Met stations used for the entire winter campaign.



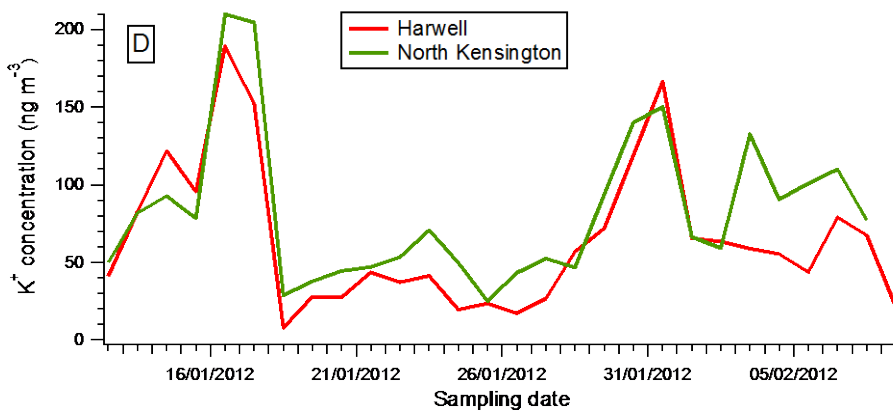
1474



1475



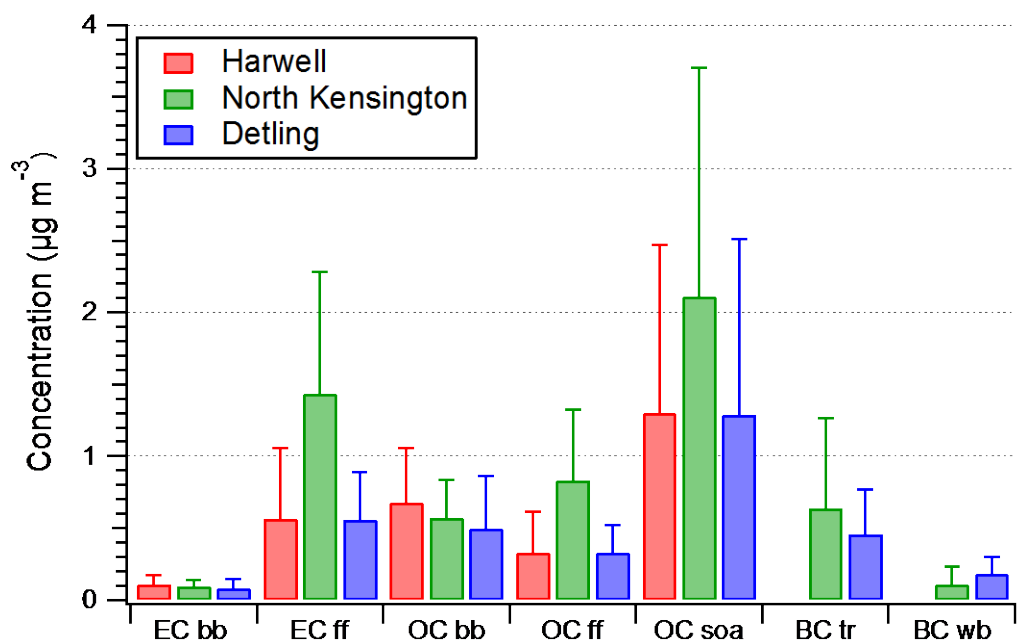
1476



1477

1478 Figure 3: Times series of the concentration of (A) levoglucosan, (B) OC, (C) EC and (D) K^+ at
 1479 Harwell, NK and Detling. No potassium data is available for Detling.

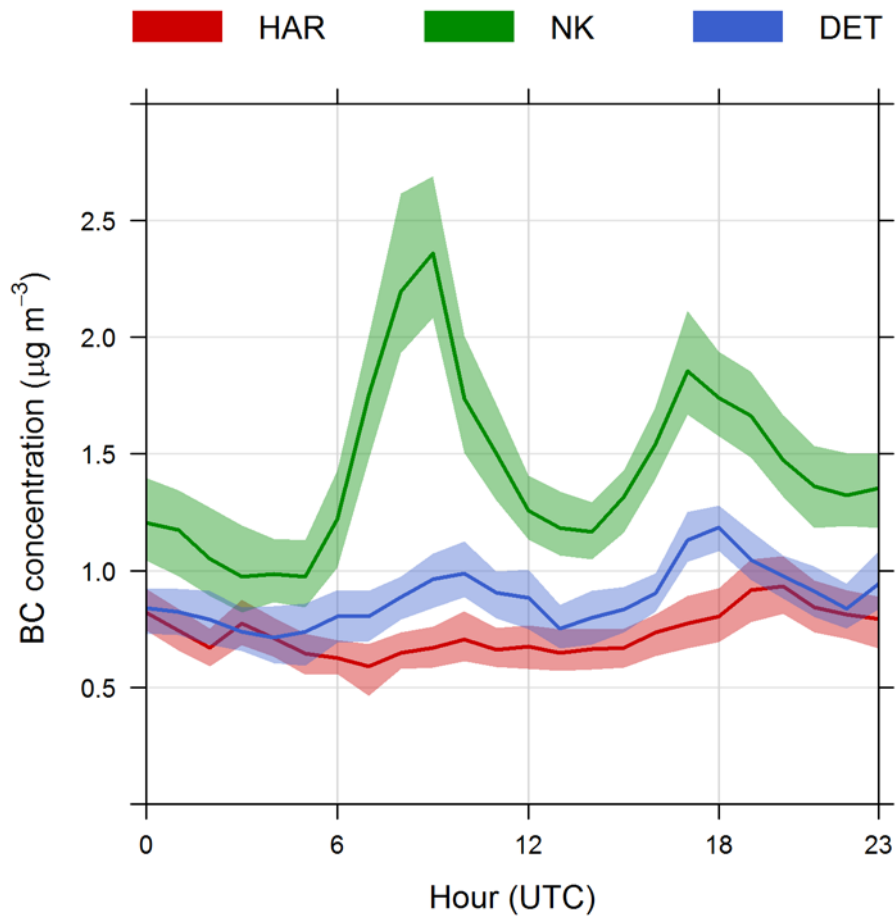
1480



1481

1482 Figure 4: Average of daily concentrations of EC_{bb}, EC_{ff}, OC_{bb}, OC_{ff} and OC_{SOA} apportioned using
 1483 the tracer method (Gelencsér et al., 2007). Variability shown is 1 standard deviation.

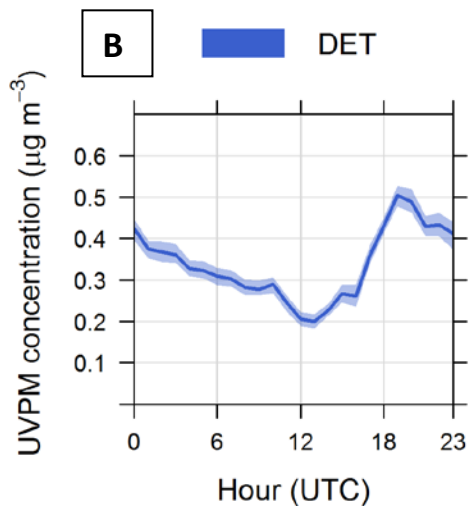
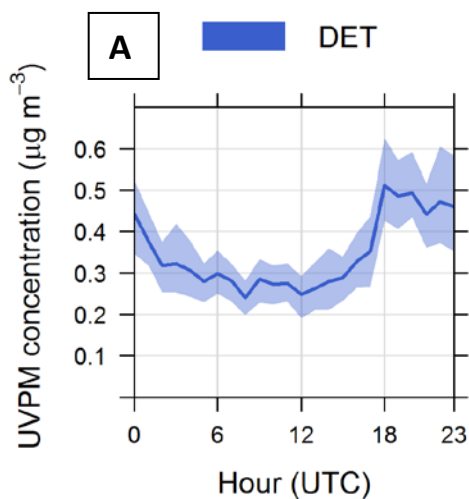
1484



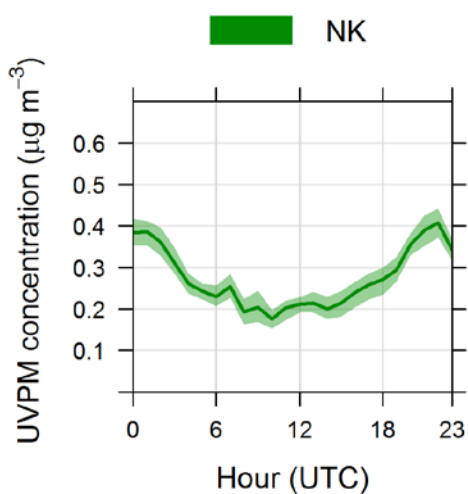
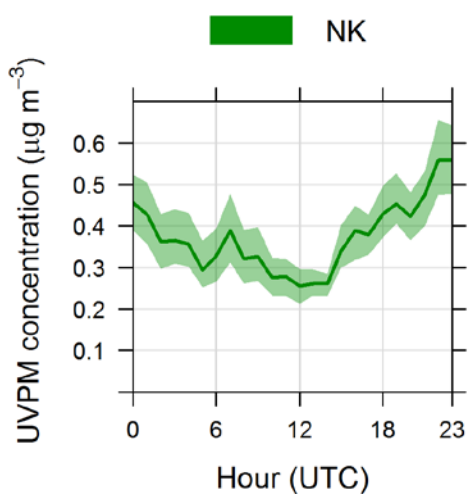
1485

1486 Figure 5: Mean diurnal variations in the BC concentrations at 880 nm during the winter campaign
 1487 as measured by a 2W AE with the shaded areas indicating the 95% confidence intervals. Note in the
 1488 key HAR and DET represents Harwell and Detling, respectively.

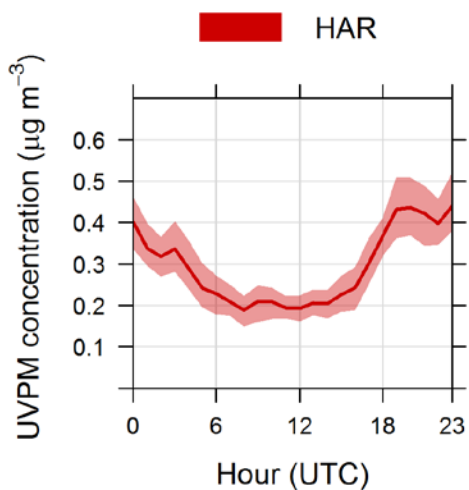
1489



1490

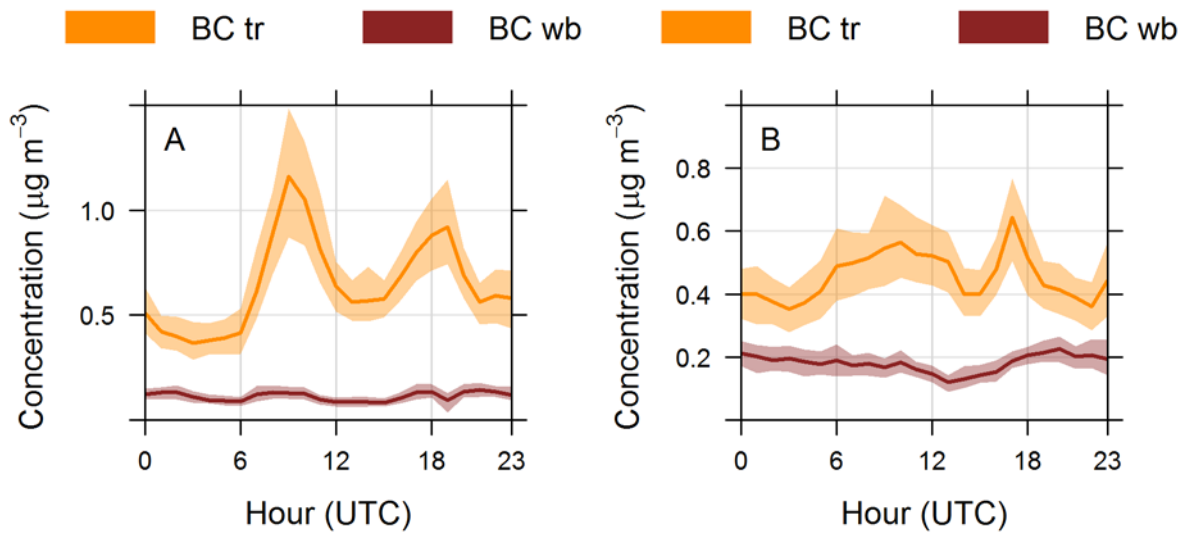


1491



1492

1493 Figure 6: Mean diurnal cycles of UVPM during the winter campaign as determined by 2W AE
 1494 (Column A) and 7W AE (Column B) for the sites where the data was available. Shaded areas
 1495 indicate the 95% confidence intervals. In the key HAR and DET represents Harwell and Detling,
 1496 respectively.

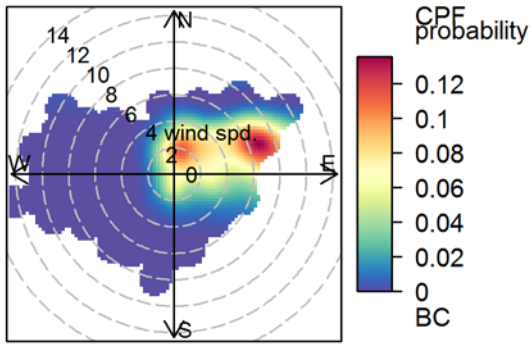


1497

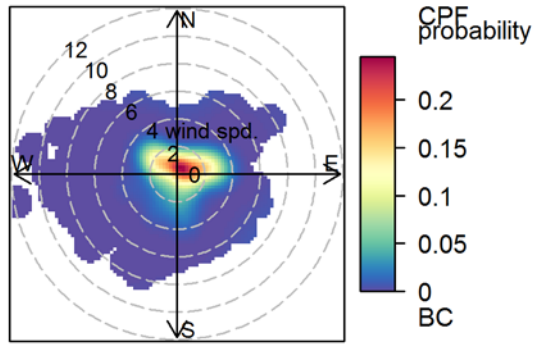
1498 Figure 7: Mean diurnal cycle of concentration of BC_{wb} and BC_{tr} at 880 nm during the winter
 1499 campaign at NK (A) and Detling (B) derived from the Aethalometer model. The shaded areas
 1500 represent the 95% confidence interval. Note the different y-axis scales used between panels.

1501

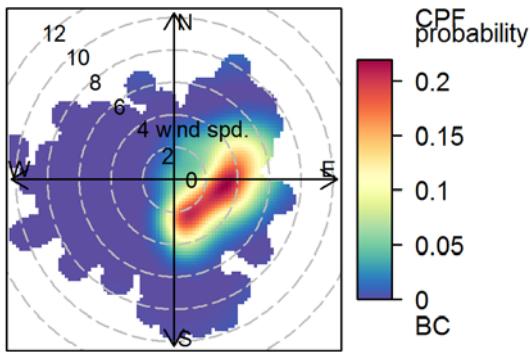
Detling



N. Kensington



Harwell

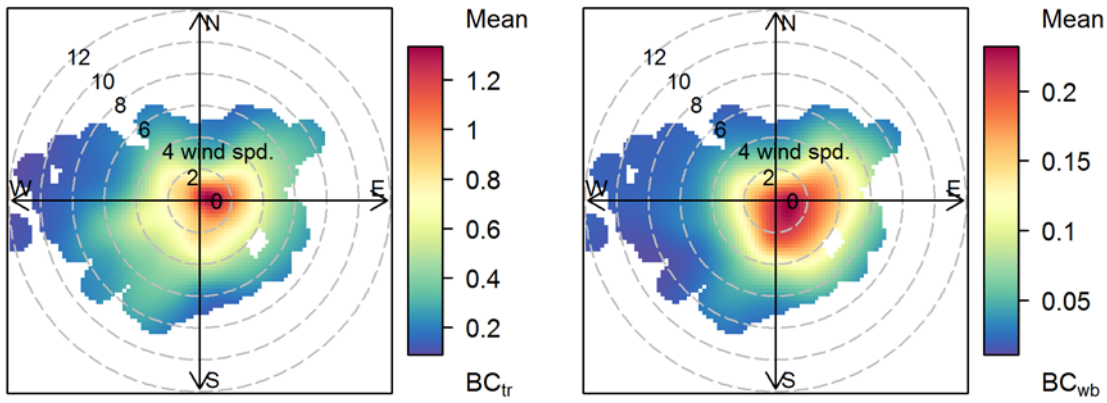


1502

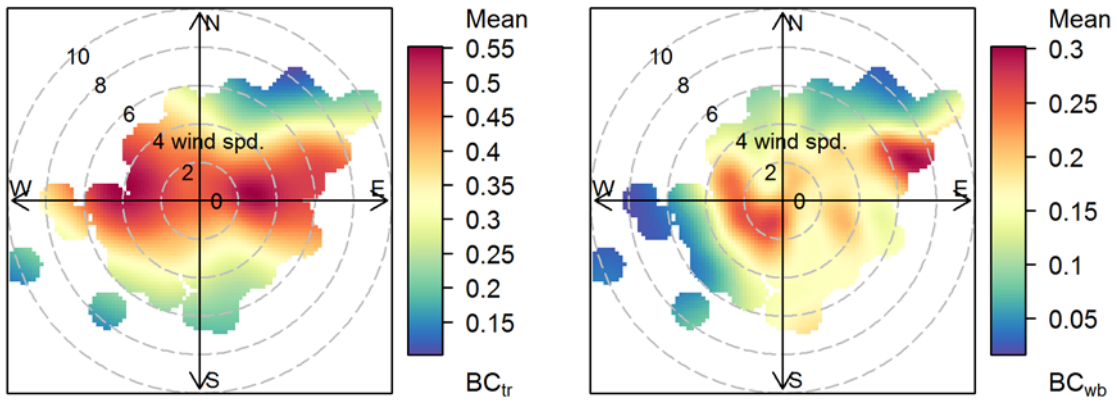
1503

1504 Figure 8: Polar plots of CPF analysis (90th percentile) for the BC concentrations at 880 nm
 1505 measured with a 2W AE at the sites during the winter campaign. Wind speed (wind spd) is in m s⁻¹.

1506



1507



1508

1509

1510

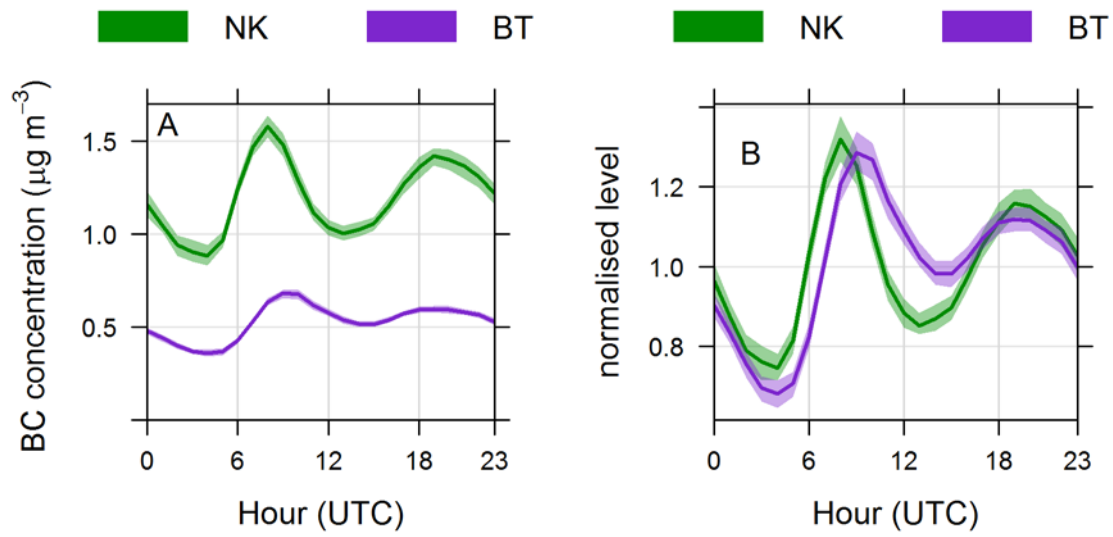
1511

Figure 9: Polar plots of the mean concentration ($\mu\text{g m}^{-3}$) of BC from traffic (BC_{tr}) and wood burning (BC_{wb}) calculated by the two component Aethalometer model at North Kensington (top row) and Detling (bottom row) for the winter campaign. Wind speed is in m s^{-1} .

1512

1513

1514

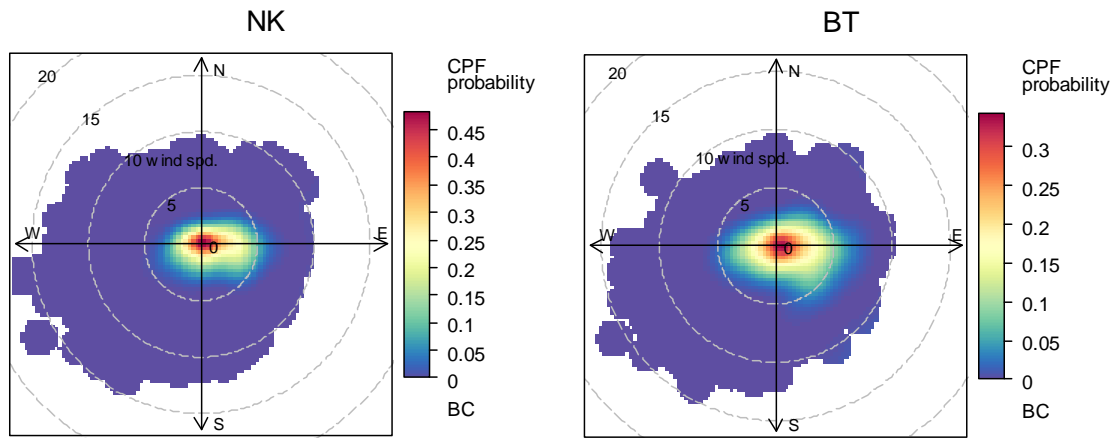


1515

1516 Figure 10: Mean diurnal cycles of the mean (A) and normalised (B) BC concentrations (880 nm)
1517 over the whole sampling period at the London sites. For plot B the values were normalised to the
1518 mean. Note BC concentrations were measured by a 2W AE at NK and by 7W AE at the BT tower
1519 (BT), with the same correction factor applied to both instrument datasets.

1520

1521

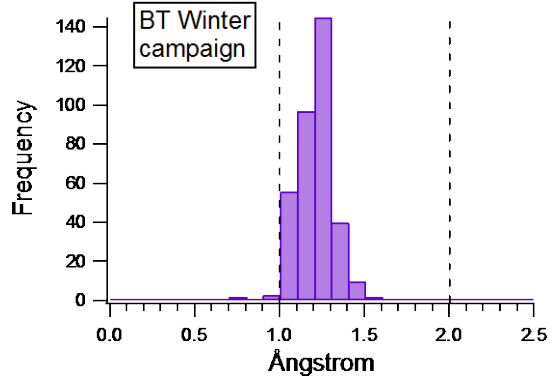
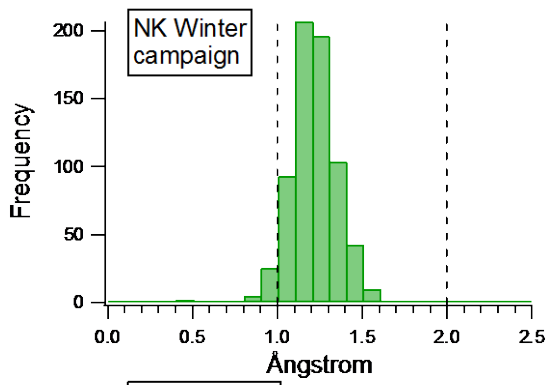


1522

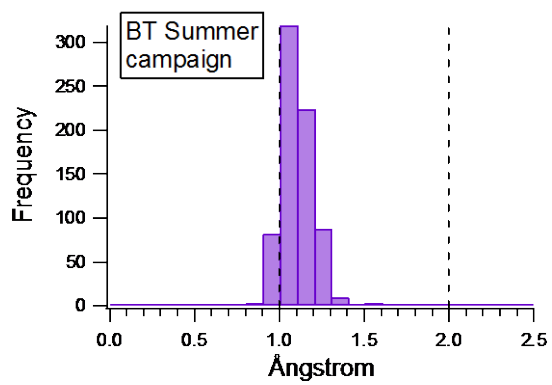
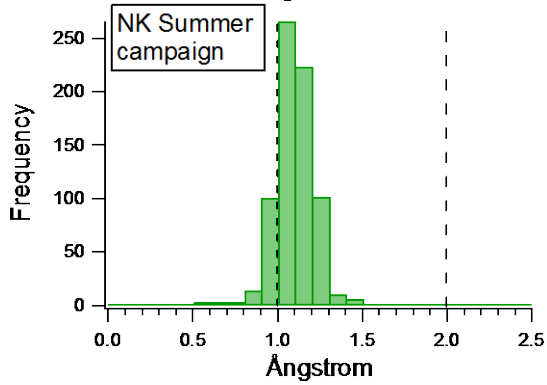
1523 Figure 11: CPF analysis for BC concentrations at 880 nm over the whole sampling period at NK
 1524 and the BT tower (BT). Note the NK concentrations were determined with a 2W AE and the BT
 1525 tower concentrations with a 7W AE.

1526

1527
1528

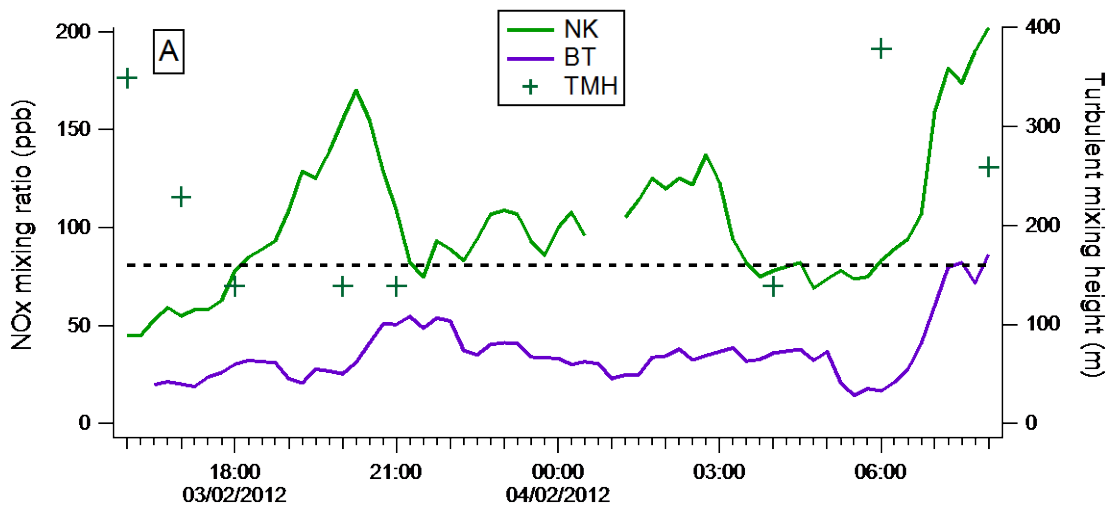


1529

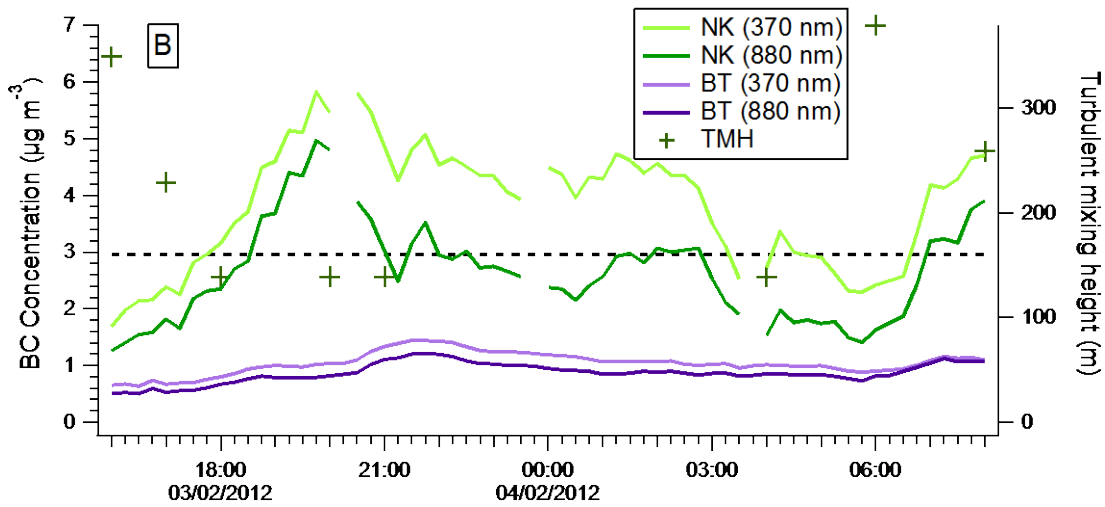


1530
1531
1532
1533
1534
1535

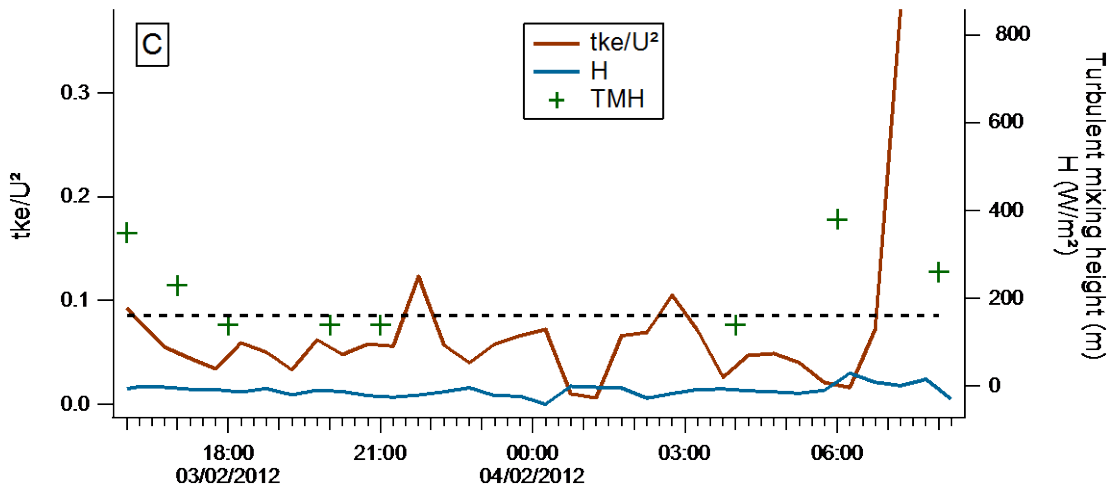
Figure 12: Histogram of Ångstrom coefficients for winter and summer campaigns at NK and the BT tower.



1536

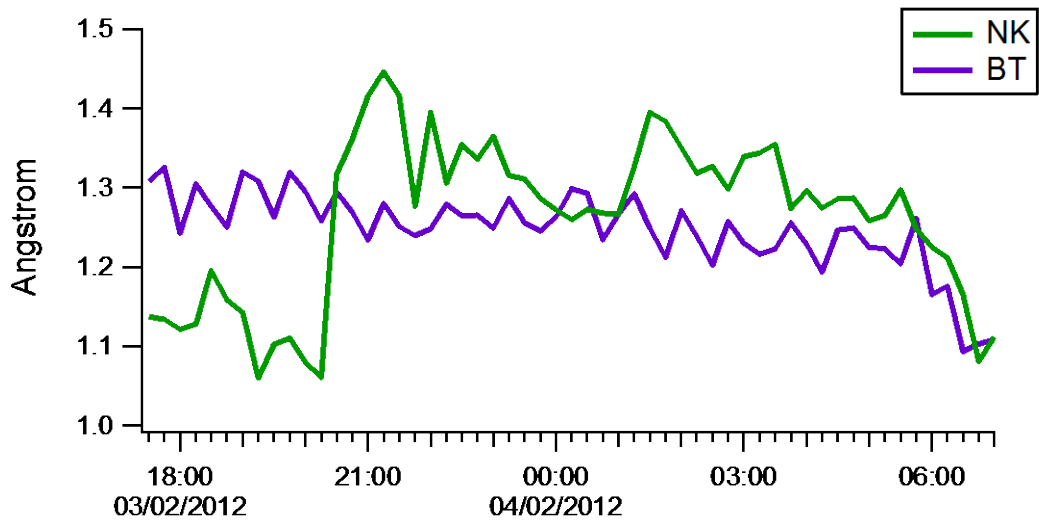


1537



1538

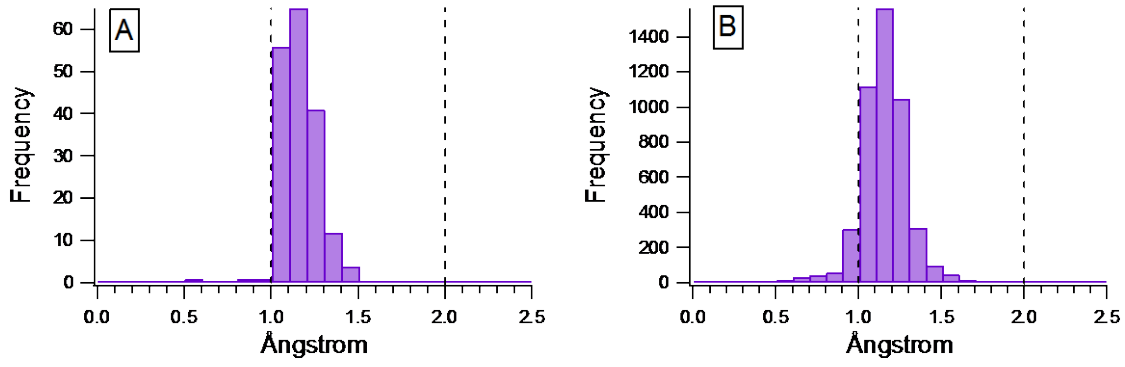
1539 Figure 13: Time series of NO_x concentration (A) and BC concentration measured at 370 and 880
 1540 nm (B) during a de-coupling event as marked by the turbulent mixing heights (TMH). For times
 1541 where there are no TMH data in the figure, this is due to the turbulent layer lying below the lowest
 1542 range gate of the Doppler lidar i.e. minimum resolvable TMH is 139 m. Also shown for this de-
 1543 coupling event is the sensible heat flux (H) and turbulent kinetic energy normalized by the wind
 1544 speed (tke/U^2) as measured at the BT tower (C). Note the dashed line represents the height of the
 1545 T35 sampling level of the aethalometers (160 m).



1546

1547 Figure 14: Time series of the calculated absorption Ångstrom coefficient (α) during the de-coupling
 1548 event from Figure 13.

1549



1550

1551

1552 Figure 15: Histogram of hourly of Ångstrom coefficient for BT tower for (A) all de-coupling
1553 periods and (B) for the other nights.

1554

1555

1556

1557