Strong Atmospheric New Particle Formation in Winter, Urban 1 Shanghai, China 2 Shan XIAO^{1,2}, Mingyi WANG^{1,2}, Lei YAO^{1,2}, Markku KULMALA³, Bin ZHOU^{1,2}, Xin YANG^{1,2}, 3 Jianmin CHEN^{1,2}, Dongfang WANG⁴, Qingyan FU⁴, Douglas R. WORSNOP⁵, Lin WANG^{1,2*} 4 5 ¹Shanghai Key Laboratory of Atmospheric Particle Pollution and Prevention (LAP³), Department 6 of Environmental Science & Engineering, Fudan University, Shanghai 200433, P. R. China ²Fudan Tyndall Centre, Fudan University, Shanghai 200433, P. R. China 7 ³Department of Physics, University of Helsinki, 00014 Helsinki, Finland 8 ⁴Shanghai Environmental Monitoring Centre, Shanghai 200030, P.R. China 9 ⁵Aerodyne Research, Billerica, MA 01821, USA 10 11 *Corresponding author: tel, +86-21-65643568; fax, +86-21-65642080, email, 12 13 lin wang@fudan.edu.cn 14 Abstract Particle size distributions in the range of 1.34-615 nm were recorded from 15 Nov. 25th, 2013 to Jan. 25th, 2014 in urban Shanghai, using a combination of one nano 16 Condensation Nucleus Counter system (nCNC), one nano-Scanning Mobility Particle 17 Sizer (SMPS), and one long-SMPS. Measurements of sulfur dioxide by an SO₂ 18 analyzer with pulsed UV fluorescence technique allowed calculation of sulfuric acid 19 20 proxy. In addition, concentrations of ammonia were recorded with a Differential Optical Absorption Spectroscopy (DOAS). During this 62-day campaign, 13 NPF 21 events were identified with strong burst of sub-3 nm particles and subsequent fast 22 growth of newly formed particles. The observed nucleation rate $(J_{1,34})$, formation rate 23 of 3 nm particles (J_3) , and condensation sink (CS) were 112.4-271.0 cm⁻³ s⁻¹, 2.3-24 19.2 cm⁻³ s⁻¹, and 0.030-0.10 s⁻¹, respectively. Subsequent cluster/nanoparticle growth 25 showed a clear size dependence, with average values of $GR_{1,35\sim1,39}$, $GR_{1,39\sim1,46}$, 26 $GR_{1.46\sim1.70}, GR_{1.70\sim2.39}, GR_{2.39\sim7}, \text{ and } GR_{7\sim20} \text{ being } 1.6\pm1.0, 1.4\pm2.2, 7.2\pm7.1, 9.0\pm1.0, 1.4\pm2.2, 7.2\pm7.1, 9.0\pm1.0, 1.4\pm2.2, 7.2\pm7.1, 9.0\pm1.0, 1.4\pm2.2, 7.2\pm7.1, 9.0\pm1.0, 1.4\pm1.0, 1.4\pm1.0$ 27 11.4, 10.9 \pm 9.8, and 11.4 \pm 9.7 nm h⁻¹, respectively. Correlation between nucleation 28 rate $(J_{1,34})$ and sulfuric acid proxy indicates that nucleation rate $J_{1,34}$ was proportional 29 to a 0.65 ± 0.28 power of sulfuric acid proxy, indicating that the nucleation of particles 30 can be explained by the activation theory. Correlation between nucleation rate $(J_{1,34})$ 31

and gas-phase ammonia suggests that ammonia was associated with NPF events. The
calculated sulfuric acid proxy was sufficient to explain the subsequent growth of 1.34nm particles, but its contribution became smaller when the particle size grew.
Qualitatively, NPF events in urban Shanghai likely occur on days with low levels of
aerosol surface area and sulfuric acid proxy is important only when aerosol surface
area is low.

38

39 **1 Introduction**

Aerosol particles can influence climate directly and indirectly (Andreae and Crutzen, 40 1997; Haywood and Boucher, 2000; IPCC, 2013), and have adverse impact on human 41 health (Dockery et al., 1993; Laden et al., 2006; Pope and Dockery, 2006). 42 Atmospheric nucleation of gas-phase precursors to clusters, and then further to 43 nanoparticles is the largest source of atmospheric aerosol particles (Kulmala et al., 44 2004; Zhang et al., 2012). This phenomenon has been observed in numerous locations 45 around the world, including areas with a pristine atmosphere, e.g., coastal areas 46 47 (O'Dowd et al., 2002), Antarctic/Arctic (Park et al., 2004), remote forest (Dal Maso et al., 2005), semi-rural locations with very low pollution levels such as Kent, OH 48 (Kanawade et al., 2012), and heavily polluted cities, such as Mexico City (Dunn et al., 49 2004). 50

51

The exact mechanism for atmospheric nucleation is still under active investigation. 52 Field measurements and laboratory studies have shown that sulfuric acid is a key 53 precursor species for atmospheric nucleation (Weber et al., 1996; Sipila et al., 2010), 54 55 and that atmospheric nucleation rate can be significantly promoted in presence of other precursors including ammonia (Ball et al., 1999; Benson et al., 2009), amines 56 (Berndt et al., 2010; Zhao et al., 2011), and organic acids (Zhang et al., 2004; Zhang 57 et al., 2009). At certain locations, ion-induced nucleation (Yu and Turco, 2001; Lee et 58 al., 2003) or iodine species (O'Dowd et al., 2002) may also help to explain the 59 60 observed new particle formation. Very recently progress has been made by the use of Particle Size Magnifier (PSM) and Chemical Ionization Atmospheric Pressure 61

interface Time-of-Flight (CI-APi-ToF) mass spectrometer, by combining the Cloud
(Cosmics Leaving OUtdoor Droplets) chamber experiments and ambient observations
including those at Hyytiälä, Finland, showing that oxidation products of biogenic
emissions, together with sulfuric acid, contribute to new particle formation in the
atmosphere (Schobesberger et al., 2013; Riccobono et al., 2014).

67

China suffers severe air pollution, especially high atmospheric particle loadings in 68 69 recent years (Chan and Yao, 2008). Among many potential sources of atmospheric particles, atmospheric nucleation has been suggested to be a significant source of 70 particles (Matsui et al., 2011; Yue et al., 2011). Correspondingly, a number of 71 extensive campaign or long-term observational studies have been carried out in the 72 Beijing area (e.g., Wu et al., 2007; Yue et al., 2009; Zhang et al., 2011; Gao et al., 73 2012) and Pearl River Delta, including Hong Kong (e.g., Guo et al., 2012; Yue et al., 74 2013). As one of the most industrialized area of China, one of the most populated 75 region in the world, and one of the hotspots for particle pollution, Yangtze River Delta 76 77 (YRD) has only seen a few research activities on atmospheric nucleation (Du et al., 2012; Herrmann et al., 2014). Among the few studies, measurements at the station for 78 Observing Regional Processes of the Earth System, Nanjing University (SORPES-79 NJU) offered a first insight for new particle formation in the western part of YRD 80 (Herrmann et al., 2014). On the other hand, atmospheric nucleation research in China 81 is still in its infant stage and the latest experimental techniques are yet to be applied in 82 China. For example, data on freshly nucleated particles are really sparse, except for 83 those from an air ion spectrometer employed at SORPES-NJU (Herrmann et al., 84 2014). To the best of our knowledge, the use of a Particle Size Magnifier (PSM), 85 which is able to study atmospheric nucleation at the size (mobility diameter) down to 86 1.5 ± 0.4 nm (Kulmala et al., 2012), has not been reported in a Chinese location in 87 literature. The lack of key information greatly hinders a better understanding of 88 nucleation mechanisms in China, where concentrations of sulfuric acid and basic 89 90 gases including ammonia and amines are high (Zheng et al., 2011; Zheng et al., 2015) but concentrations of extremely low volatility organic compounds formed from 91

92 biogenic emissions are yet to be determined.

93

Direct measurements of atmospheric nucleation rates down to 1.5 ± 0.4 nm provide a 94 better and more accurate characterization of atmospheric nucleation, since the indirect 95 calculation of atmospheric nucleation rates from the formation rates of 3 nm particles 96 97 leads to substantial uncertainty due to our incomplete understanding of condensational growth and coagulation scavenging of particles in the 1.5 to 3 nm range (Anttila et al., 98 99 2010; Korhonen et al., 2011). With the growing number of reports of real nucleation rates in clean atmosphere (e.g., Kulmala et al., 2012; Yu et al., 2014), it is ideal to 100 measure nucleation rates in a polluted urban atmosphere to find out how atmospheric 101 nucleation rates vary under different atmospheric conditions. In addition, the 102 nucleation mechanism in a polluted urban atmosphere that is vital to understand 103 atmospheric nucleation at a global scale and for atmospheric model development can 104 be preliminarily investigated by examining the relationship between the measured 105 atmospheric nucleation rates and the well-accepted precursor gases that exist in high 106 107 concentrations.

108

In this study, we measured atmospheric nucleation from Nov. 25th, 2013 to Jan. 25th, 2014 in urban Shanghai with nCNC and two sets of SMPS. Nucleation frequency, nucleation rate ($J_{1,34}$), condensation sink (*CS*), and growth rates (*GR*) are reported and compared with previous studies with similar or dissimilar atmospheric environments. In addition, the potential nucleation mechanism was explored by correlating sulfuric acid proxy calculated from sulfur dioxide precursor and gas-phase ammonia to nucleation rate ($J_{1,34}$).

116

117 **2 Experimental**

118 2.1 Nucleation Measurements

119 Nucleation measurements were carried out on the rooftop of a teaching building
120 (31°18'N, 121°30'E) that is about 20 m above ground on the campus of Fudan

University between Nov. 25th, 2013 and Jan. 25th, 2014. This monitoring site is mostly surrounded by commercial properties and residential dwellings. The Middle Ring Road, one of main overhead highways in Shanghai, lies about 100 m to the south of the site. Hence, the Fudan site can be treated as a representative urban site influenced by a wide mixture of emission sources (Wang et al., 2013; Ma et al., 2014).

126

Ambient particle size distributions in the range of 1.34-615 nm were measured using a combination of one nano Condensation Nucleus Counter system (model A11, Airmodus, Finland), one nano-SMPS (consisting of one DMA3085 and one CPC3776, TSI, USA), and one long-SMPS (consisting of one DMA3081 and one CPC3775, TSI, USA). The instruments were continuously running except for maintenance and minor instrument breakdown during the campaign.

133

Ambient air was drawn into a stainless steel manifold of 5.0 m length and 4 inch inner 134 diameter at a flow rate of 153 CFM using a blower (Model DJT10U- 25M, NUSSUN, 135 136 China). From this main manifold, 1.75 lpm ambient air was drawn through a 1/4 inch inner diameter stainless tube of 18 cm length, and diluted with a zero air flow 137 generated by a zero air generator (Model 111, Thermo, USA) at a ratio of 1:1 to 138 reduce the overall relative humidity (RH) and the number of particles entering PSM, 139 since high RH and particle loading had an impact on the saturation of diethylene 140 glycol in PSM and hence data quality. Subsequently 2.5 lpm diluted air was sampled 141 into nCNC. In addition, 30 lpm split flow was drawn from the main manifold through 142 a 1/4 inch inner diameter conductive silicon tubing of 50 cm length, and then 0.3 lpm 143 144 and 1.5 lpm ambient air from the split flow, respectively, were drawn into nano-SMPS and long-SMPS. The calculated diffusion loss is up to 29% for 1.35 nm particles with 145 the above setup, and has been taken into account in the entire size range during the 146 147 data processing.

148

The nCNC system consists of one PSM (model A10, Airmodus, Finland) and onebutanol Condensation Particle Counter (bCPC, model A20, Airmodus, Finland), and

5

was used to measure size distributions of 1.34-3 nm clusters/particles. Briefly, PSM 151 activates the smallest particles using diethylene glycol as a working fluid and 152 condensationally grows nanoparticles up to larger than 90 nm in mobility equivalent 153 diameter, after which an external CPC is used for further growing the particles to 154 optical sizes and counting the grown particles (Vanhanen et al., 2011). In this study, 155 PSM was used in the scanning mode in which the saturator flow rate is changed 156 continuously, giving an activation spectrum of the measured particles to obtain size 157 distribution of sub-3 nm clusters/particles. A scanning cycle of 100 steps between 158 saturator flow rates 0.1-1 lpm and a time resolution of 220 sec were chosen. Sub-3 nm 159 clusters/particles were classified into 5 bins, i.e., 1.34-1.37, 1.37-1.41, 1.41-1.52, 160 1.52-1.89, 1.89-3.0 nm, respectively. Geometric mean values of upper and lower 161 limits of the five bins, i.e., 1.35, 1.39, 1.46, 1.70, and 2.39 nm, respectively, were used 162 163 to refer to the five bins in the growth rate calculation.

164

The nano-SMPS measured particles in the size range from 3 to 64 nm and the long-165 SMPS recorded particles from 14 to 615 nm. For both SMPSs, 64 size bins and 5-min 166 time resolution were chosen. The sample flow to sheath flow ratios for both SMPSs 167 were set at 1:10. A comparison analysis on the total particle concentrations between 168 14 and 64 nm measured by both nano-SMPS and long-SMPS showed less than 10% 169 difference in the size range of 55-64 nm between two SMPSs. Hence, number 170 concentrations of particles in the size range of 3-615 nm, $N_{3\sim615}$, were calculated from 171 the sum of $N_{3\sim55}$ obtained from nano-SMPS, $N_{55\sim64}$ from the arithmetic average of 172 nano-SPMS and long-SMPS, and $N_{64\sim615}$ from long-SMPS. 173

174

At the same site, sulfur dioxide (SO₂) was measured by an SO₂ analyzer with pulsed UV fluorescence technique (Model 43i, Thermo, USA) with a time resolution of 5 min and calibration of this SO₂ analyzer was performed every month. A DOAS system was used to measure the integrated concentration of NH₃ along the optical path between a transmitter telescope using a 35W Deuterium lamp as the light source and a receiver telescope (53 m), and then to yield the average concentration of NH₃ through dividing the integrated concentration by the absorption length (Platt and Stutz,

182 2008). In this study, the concentration of NH₃ was determined by fitting the reference

spectra to the atmospheric spectra in a given window (205-220 nm) using a nonlinear

184 least-squares method, similarly to a previous measurement of HONO and NO₂ (Wang

et al., 2013). Detection limit of NH₃ was about 1 ppb with a 3 min integration time.

186

187 Solar radiation intensity measured by a Pyranometer (Kipp & Zonen CMP6,
188 Netherland) was obtained from Shanghai Pudong Environmental Monitoring Centre
189 (31°14'N, 121°32'E, about 8.78 km from the Fudan site).

190

191 **2.2 Data Processing**

192 2.2.1 Nucleation rate (J_{1,34}), formation rate of 3 nm particles (J₃), growth rate 193 (GR), and condensation sink (CS)

In this study, PSM allowed measurements of clusters/particles down to 1.34 nm. Hence, atmospheric nucleation rate, $J_{1.34}$, defined as the flux of particles growing over 1.34 nm, can be calculated by taking into account the coagulation losses and condensational growth out of the considered size range (Kulmala et al., 2012),

198
$$J_{1.34} = \frac{dN_{1.34-3}}{dt} + CoagS_{d_p=2nm} \cdot N_{1.34-3} + \frac{1}{1.66nm} GR_{1.34-3} \cdot N_{1.34-3}$$
(1)

where $CoagS_{d_p=2nm}$ represents coagulation sink of 2 nm particles, an approximation for that in the size interval of 1.34-3 nm; and $GR_{1.34-3}$ represents the apparent clusters/particle growth rate between 1.34 and 3 nm.

202

Formation rate of 3nm particles was calculated in a similar way (Sihto et al., 2006;
Kulmala et al., 2012), providing a comparison with previous studies,

205
$$J_3 = \frac{dN_{3\sim 6}}{dt} + CoagS_{d_p=4nm} \cdot N_{3\sim 6} + \frac{1}{3nm}GR_{3\sim 6} \cdot N_{3\sim 6}$$
(2)

where $CoagS_{d_p=4nm}$ represents coagulation sink of 4 nm particles, an approximation for that in the size interval of 3-6 nm. Growth rate (GR) is defined as the rate of change in the diameter of a growing particle population, using the maximum-concentration method (Kulmala et al., 2012),

211
$$GR = \frac{dd_p}{dt} = \frac{\Delta d_p}{\Delta t} = \frac{d_{p_2} - d_{p_1}}{t_2 - t_1}$$
 (3)

where d_{p_1} and d_{p_2} are the representative particle diameters at times t_1 and t_2 , respectively.

214

Condensation sink (*CS*) describes the condensing vapor sink caused by the particle
population (Kulmala et al., 2012),

217
$$CS = 4\pi D \int_{0}^{d_{p}} \beta_{m,d_{p}} d_{p} N_{d_{p}} dd_{p} = 4\pi D \sum_{d_{p}} \beta_{m,d_{p}} d_{p} N_{d_{p}}$$
(4)

where *D* is the diffusion coefficient of the condensing vapor, usually assumed to be sulfuric acid (0.104 cm² s⁻¹ used in this study); and β_{m,d_p} is the transitional regime correction factor.

221

222 **2.2.2 Sulfuric acid**

223 Sulfuric acid has been accepted as a key gas-phase precursor for atmospheric nucleation and contributes to the subsequent growth of newly-formed particles 224 (Weber et al., 1996; Sipila et al., 2010). The accurate measurement of gas-phase 225 sulfuric acid requires application of chemical ionization mass spectrometry using 226 nitrates as reagent ions (Eisele and Tanner, 1993), which is not possessed by this 227 research group during this study. Instead, the sulfuric acid proxy $[H_2SO_4]$ was 228 estimated based on local solar radiation level *Radiation*, SO_2 concentration $[SO_2]$, 229 condensation sink CS, and relatively humidity (Mikkonen et al., 2011), 230

231 $[H_2SO_4] = 8.21 \times 10^{-3} \cdot k \cdot Radiation \cdot [SO_2]^{0.62} \cdot (CS \cdot RH)^{-0.13}$ (5)

where k is the temperature dependent reaction rate constant. The relative error

between calculated sulfuric acid proxy and measured sulfuric acid concentration is

estimated to be 42% (Mikkonen et al., 2011). The time resolution of calculated

sulfuric acid proxy was 1 hour since those of temperature and relative humidity was 1hour.

237

238 Condensation of sulfuric acid contributes to the growth of newly-formed particles.

The growth of clusters/particles due to condensation of sulfuric acid, $GR_{H_2SO_4}$, can be estimated by the following equations (Nieminen et al., 2010),

241
$$GR_{H_2SO_4} = \frac{\gamma}{2\rho_v} \left(1 + \frac{d_v}{d_p}\right)^2 \left(\frac{8kT}{\pi}\right)^{1/2} \left(\frac{1}{m_p} + \frac{1}{m_v}\right)^{1/2} m_v [H_2SO_4]$$
(6)

242 and
$$\gamma = \frac{4}{3} \cdot Kn \cdot \beta_{m,d_p}$$
 (7)

where ρ_v and d_v are the condensed phase density and molecule diameter of H₂SO₄, respectively; m_p and m_v are particle and H₂SO₄ vapor molecule masses, respectively; γ is a correction factor; Kn is the Knudsen number (Lehtinen and Kulmala, 2003). For this calculation, particle density $\rho_p = 1.83$ g cm⁻³ was used.

247

The particle growth due to the hydration of H_2SO_4 is taken into account by assuming that sulfuric acid is instantaneously equilibrated with gas-phase water. During our campaign, daily average RH varied between 28.7%~60.0%. Hence, using the H₂SO₄hydrate distribution data given by Kurtén et al. (2007), the density and mass of the average hydrated H₂SO₄ molecule at 50% relative humidity is calculated and further used in equation (6).

254

255 **3 Results and Discussion**

256 **3.1 Classification of new particle formation (NPF) events**

Figure 1 presents a contour plot for particle size distributions of 3-615 nm and a number concentration plot of sub-3 nm clusters/particles $N_{1.34\sim3}$ during Nov. 25th, 2013 - Jan. 25th, 2014. Data were occasionally missing because of maintenance and minor breakdown of instruments. From the figure, frequent bursts of sub-3 nm

particles were evident, with concentrations up to 8.0×10^4 cm⁻³ around noon time. 261 However, similarly to previous studies (Kulmala et al., 2007; Kulmala et al., 2013; Yu 262 et al., 2014), not all sub-3 nm particles eventually underwent a continuous growth to 263 larger sizes. In this study, we define an observation day with appearances of sub-3 nm 264 clusters/particles over a time span of hours and subsequent growth to larger sizes for a 265 few hours as a NPF event day. In this case, a NPF day will present a banana-shaped 266 contour plot of particle size distributions obtained from SMPS (Dal Maso et al., 267 2005). We focus on characteristics and potential mechanisms of these events. 268

269

According to the classification, there were 13 event days during the 62-day campaign, as illustrated by the shadow in Figure 1. Although nCNC data were partially unavailable on Dec. 26th, 2013 and completely unavailable on Jan. 24th, 2014, these two days are still defined as NPF days since a distinctive banana-shaped contour plot for particle distributions between 3-615 nm existed. Dec. 18th, 2013 was not regarded as a NPF day since $N_{1.34~3}$ was not recorded and the growth of 3~20 nm particles was relatively short-lived.

277

Among these NPF events, 5 NPF events occurred in November, 3 in December, and 5 278 in January. The averaged frequency for NPF events was 21.0% during the 62-day 279 280 campaign. Our NPF frequency at Shanghai is larger than the average value in winter 1996-2003, SMEAR II station, Hyytiälä, Finland (Dal Maso et al., 2005), likely 281 because nucleation events at Hyytiälä have recently been related to oxidation products 282 of biogenic emissions (Kulmala et al., 1998; Schobesberger et al., 2013; Riccobono et 283 284 al., 2014) and photochemistry of volatile organic compounds is less intensive in winter months. This frequency is also higher than that in winter, semi-rural Kent, OH, 285 during which transport of sulfur dioxide from the east-southeast power plant to Kent 286 is not favored (Kanawade et al., 2012). NPF events occurred at a frequency of around 287 40% during Nov.-Dec. 2004 in Beijing (Wu et al., 2007), much more often than in 288 Shanghai. On the other hand, NPF frequency in Shanghai is remarkably close to that 289

290

measured in winter 2012, Nanjing, which is also located in Yangtze River delta (Herrmann et al., 2014).

292

291

Number concentrations of particles in different size ranges, i.e., $N_{1,34-3}$, N_{3-7} , and 293 $N_{7\sim30}$, respectively, on a NPF day (Dec. 11th, 2013) and an obvious non-NPF day 294 (Jan. 7th, 2014) are further examined to illustrate features of a NPF event, as shown 295 in Figure 2. On the NPF day, 1.34-3 nm particles appeared as early as 7 am in the 296 morning that was right after sunrise (6:42 am on Dec. 11th, 2013), reached its 297 maximum just before noontime, and spanned for almost the whole daytime (sunset at 298 4:52 pm, Dec. 11th, 2013), suggesting that photochemistry products likely contribute 299 to formation of smallest particles. This size distribution of atmospheric neutral and 300 charged clusters/particles by a scanning PSM is identical to that measured at 301 Hyytiälä, Finland (Kulmala et al., 2013). On the same NPF day, 3-7 nm and 7-30 nm 302 particles appeared much later, resembling previous findings only with SMPS 303 measurements (e.g., Yue et al., 2010). The lag in peaking times of $N_{1.34-3}$, N_{3-7} , and 304 $N_{7\sim30}$ on the NPF day clearly indicated the continuous growth of clusters/particles, 305 the reduction of particles due to coagulation during the growth, and the diverse 306 sources of particles in the size range of 7-30 nm. In contrast, $N_{\rm 1.34\sim3}$ and $N_{\rm 3\sim7}$ 307 showed a flat profile on the non-NPF day. The minor enhancement in $N_{7\sim30}$ between 308 10 am and 5 pm on the non-NPF day was not due to growth of newly formed 309 particles. Instead, direct emission of 7-30 nm particles from transportation activity 310 311 likely explained their appearance.

312

313 **3.2 General characteristics of NPF events**

Table 1 summaries characteristics of each NPF event observed in this campaign, including $J_{1.34}$, J_3 , $GR_{1.35\sim2.39}$ (from the bin of 1.34-1.37 nm to the bin of 1.89-3.0 nm), $GR_{2.39\sim7}$, $GR_{7\sim20}$, CS, $[H_2SO_4]$, $N_{1.34\sim3}$, and total number concentrations of

particles $N_{1,34-615}$, and compares the mean values to those in selected other studies. 317 Nucleation rate $J_{1.34}$ and formation rate of 3 nm particles J_3 were 112.4-271.0 and 318 2.3-19.2 cm⁻³ s⁻¹, respectively. Nucleation rate $J_{1,34}$ at Shanghai is obviously 319 significantly larger than 1.4 cm⁻³ s⁻¹ at Hyytiälä, Finland with a pristine atmosphere 320 (Kulmala et al., 2012) and 1.3 cm⁻³ s⁻¹ at Kent, OH with relatively lower levels of 321 pollutants (Yu et al., 2014). Direct comparison of our nucleation rate with that in a 322 Chinese location is not feasible because no previous reports are available. However, 323 Herrmann et al. (2014) reported a J_2 of 33.2 cm⁻³ s⁻¹ at the SORPES-NJU station, 324 Nanjing China. Together with their results, we conclude that strong nucleation events 325 326 occur geographically widely in the YRD region.

327

Formation rate of 3 nm particles J_3 has been more routinely reported. Similarly to $J_{1.34}$, J_3 at Shanghai is significantly larger than 0.61 cm⁻³ s⁻¹ at Hyytiälä, Finland (Kulmala et al., 2012), and generally comparable to 3.3-81.4 and 1.1-22.4 cm⁻³ s⁻¹ at Beijing (Wu et al., 2007; Yue et al., 2009), 3.6-6.9 cm⁻³ s⁻¹ at Hong Kong (Guo et al., 2012), and 2.4-4.0 cm⁻³ s⁻¹ in a back-garden rural site of Pearl River Delta (Yue et al., 2013). The fast reduction from $J_{1.34}$ to J_3 was likely due to the presence of a large background particle number as shown in Table 1.

335

The large background particle number concentrations corresponded to the high 336 condensation sink (CS of 0.030-0.10 s⁻¹) observed during the campaign. As shown in 337 Table 1, CS at Shanghai is much larger than $(0.05-0.35) \times 10^{-2}$ s⁻¹ at Hyytiälä, Finland 338 (Kulmala et al., 2012) and 0.8×10^{-2} s⁻¹ at Kent, OH (Yu et al., 2014), but comparable 339 to $(0.58-8.4) \times 10^{-2}$ s⁻¹ at Beijing (Wu et al., 2007; Yue et al., 2009; Zhang et al., 2011; 340 Gao et al., 2012), $(1.0-6.2) \times 10^{-2}$ s⁻¹ at Hong Kong (Guo et al., 2012), 2.4×10^{-2} s⁻¹ at 341 Nanjing (Herrmann et al., 2014), and $(3.5-4.6) \times 10^{-2}$ s⁻¹ in an urban site of Pearl 342 River Delta (Yue et al., 2013). High sulfuric acid proxy ($[H_2SO_4]$ of (2.3-6.4)×10⁷ 343

molecules cm⁻³) existed to promote the NPF events. Measurements of gas-phase 344 sulfuric acid by a chemical ionization mass spectrometer during the CAREBeijing 345 2008 Campaign, a time period with strict air quality control regulations, reported peak 346 concentrations of sulfuric acid up to the order of 10^7 molecules cm⁻³ (Zheng et al., 347 2011), smaller than our sulfuric acid proxy. Uncertainty may well exist for our 348 sulfuric acid proxy that was calculated from the concentrations of sulfur dioxide and 349 radiation intensity. However, Judging from CS and $[H_2SO_4]$ together, it is clear that 350 the condensable vapor at Shanghai is likely a main impetus for observed strong new 351 particle formation events. 352

353

 $GR_{1.35-2.39}$, $GR_{2.39-7}$, and GR_{7-20} were in the range of 0.49-8.1, 3.1-35.7, 4.5-38.3 nm h⁻ 354 ¹, respectively. The arithmetic average values of $GR_{1.35\sim2.39}$, $GR_{2.39\sim7}$, and $GR_{7\sim20}$ were 355 2.0 ± 2.7 (one standard deviation), 10.9 ± 9.8 and 11.4 ± 9.7 nm h⁻¹, respectively, 356 which are comparable to 3-20 nm h⁻¹ for nucleation mode particles in another sulfur-357 rich city, Atlanta, GA (Stolzenburg et al., 2005). In addition, $GR_{1.35-2.39}$ at Shanghai is 358 smaller than the growth rates (5.5-7.6 nm h⁻¹) for particles in 1~3 nm geometric 359 diameter range in Atlanta (Kuang et al., 2012). A closer examination of growth rates 360 was performed by dividing $GR_{1.35\sim2.39}$ into growth of clusters/particles from one bin to 361 another, i.e., $GR_{1.35\sim1.39}$ (1.6±1.0 nm h⁻¹ from the bin of 1.34-1.37 nm to the bin of 362 1.37-1.41 nm), $GR_{1.39\sim1.46}$ (1.4±2.2 nm h⁻¹ from 1.37-1.41 nm to 1.41-1.52 nm), 363 $GR_{1.46\sim1.70}$ (7.2±7.1 nm h⁻¹ from 1.41-1.52 nm to 1.52-1.89 nm), and $GR_{1.70\sim2.39}$ (9.0± 364 11.4 nm h⁻¹ from 1.52-1.89 nm to 1.89-3.0 nm). These growth rates show a clear size-365 dependent particle growth (Fig.3), owing to the nano-Köhler activation that suggests a 366 faster growth for activated nanoparticles due to a decreasing Kelvin effect and, thus, 367 an enhanced condensation flux (Kulmala et al., 2004), Kelvin effect, and surface or 368 volume-controlled reaction corrected for the Kelvin effect on surface or volume 369 concentrations (Kuang et al., 2012). Similar observations have been reported in 370

previous studies using nCNC (Kulmala et al., 2013) and DEG UCPC (Kuang et al.,

2012), respectively. Our $GR_{2.39-7}$ is larger than 6.3 nm h⁻¹ at Nanjing (Herrmann et al.,

2014), and our GR_{7-20} is close to the upper bound of those at urban Beijing (Wu et al.,

2007; Yue et al., 2009; Zhang et al., 2011; Gao et al., 2012), and generally larger than 1.5-8.4 nm h⁻¹ in Hong Kong (Guo et al., 2012), indicating that high concentrations of condensable vapors existed. In addition, our growth rates suggest that the smallest clusters (the bin of 1.34-1.37 nm), if not scavenged by larger particles, would grow to 3 nm within ~12 min, and to 20 nm within ~2 hr.

379

380 3.3 Potential mechanisms for NPF events

As shown in Table 1, nucleation rate $(J_{1.34})$ in this study is typically larger than 100 381 cm⁻³ s⁻¹, suggesting that the ion-induced nucleation was not a main mechanism for 382 observed fast nucleation (Hirsikko et al., 2011; Riccobono et al., 2014). The 2012 383 winter study at the SORPES-NJU station that is also located at YRD shows that the 384 ratio of J_2 between ions and total particles (ions plus neutral particles) was 0.002, 385 also indicating the minor role of ion-induced nucleation (Herrmann et al., 2014). 386 Hence, it is likely that nucleation of neutral precursor molecules actually largely 387 determined the observed NPF events. 388

389

Correlations between $\log J_{1,34}$ and $\log [H_2 SO_4]$ (Figure 4) and between $\log J_{1,34}$ and 390 $\log[NH_3]$ (Figure 5) were examined to elucidate potential mechanisms for our NPF 391 events. Since $J_{1,34}$ could not be accurately determined on some of the NPF days, the 392 number of data points in both figures was less than the actual number of NPF events 393 that have been observed. Daily peak concentration of sulfuric acid proxy and daytime 394 (6 am - 6 pm) averages of ammonia were used as approximations for their effective 395 concentrations on a NPF day since there was no peak concentration for ammonia. The 396 correlation (R²=0.62) between log $J_{1,34}$ and log [NH₃] is better than that (R²=0.38) 397

between log $J_{1.34}$ and log $[H_2SO_4]$, and slopes are 0.57 ± 0.17 and 0.65 ± 0.28 , respectively.

400

Most ambient studies showed that nucleation rate J is proportional to the first or 401 second power of the concentration of gas-phase sulfuric acid, i.e., $J = A \cdot [H_2 S O_4]^p$, 402 where P is equal to 1 or 2, conventionally interpreted as the number of sulfuric acid 403 404 molecules in the critical nucleus, and A is a pre-exponential factor (McMurry et al., 2005; Sihto et al., 2006; Erupe et al., 2010). Our P of 0.65 ± 0.28 is of a significant 405 uncertainty, which could come from the uncertainty during the calculation of sulfuric 406 acid proxy $[H_2SO_4]$ and the scarcity of our data points. The upper limit of our P 407 indicates that nucleation occurs after activation of clusters containing one molecule of 408 409 sulfuric acid, with subsequent growth involving other species (Kulmala et al., 2006). The lower limit, on the other hand, suggests that a less important role of sulfuric acid 410 in the critical nucleus during our campaign, which is unlikely to be true according to 411 412 numerous previous studies (Weber et al., 1996; Sipila et al., 2010; Yu and Hallar, Kupiainen-Määttä, et al. (2014) recently reported that the number of 413 2014). molecules in a critical cluster cannot be determined by a slope analysis in 414 atmospherically relevant applications, underscoring the need to further explore the 415 exact nucleation mechanism. Herrmann et al., (2014) also calculated the sulfuric acid 416 proxy, related it to observed nucleation rates, and speculated that the sulfuric acid 417 exponent might be well over 2, which underscores the reliability of calculation 418 methods in a Chinese location. Hence, our preliminary result should be further tested 419 420 with actual measurements of gas-phase sulfuric acid, but does indicate the key role of sulfuric acid in NPF events. On the other hand, linear correlation between $\log J$ and 421 $\log[NH_3]$ was observed in a previous nucleation study in Atlanta, GA, with a slope of 422 1.17 (McMurry et al., 2005), but a clear relationship was not perceived at Kent, OH 423 (Erupe et al., 2010). This discrepancy may come from the level of ammonia that has 424 been predicted to have a saturation threshold (Napari et al., 2002) and/or the co-425

existing sulfuric acid concentration (Benson et al., 2009). Nevertheless, our correlation between $\log J$ and $\log[NH_3]$ suggests that ammonia also participated in the nucleation. A recent CIMS study (Zheng et al., 2015) observed good correlations between NH₃ and amines at an urban site of Nanjing, China. Hence, it is plausible that amines may contribute to nucleation in our site at Shanghai, too.

431

The subsequent growth of newly formed particles can be partially attributed to 432 condensation of sulfuric acid. The theoretical maximum growth rate of 1.34-3 nm 433 clusters/particles due to condensation of hydrated sulfuric acid at 50% RH 434 $(GR_{H_2SO_4(1.34\times3)})$, calculated according to eq. 6 and 7, was 3.9 ± 1.3 nm h⁻¹ on average. 435 This rate is larger than the observed growth rates of clusters/particles from the bin of 436 1.34-1.37 nm to the bin of 1.89-3.0 nm ($GR_{1.35-2.39}$), being 2.0 ± 2.7 nm h⁻¹, indicating 437 that sulfuric acid proxy was enough to explain the observed growth for particles under 438 3 nm. Similar calculation of $GR_{H_2SO_4(3\times7)}$ and $GR_{H_2SO_4(7\times20)}$ yielded 2.8 ±0.94 and 2.2 439 440 \pm 0.74 nm h⁻¹, respectively. In figure 6, relative contributions of sulfuric acid to growth of particles in the range of 3-7 and 7-20 nm, respectively, on each NPF day is 441 presented. Since 7 nm particles reached their maximum earlier than 3 nm particles on 442 Jan. 9th and Jan. 15th, 2014, there was no calculated GR_{3-7} and hence no ratios 443 available on these two days. In addition, condensation of hydrated sulfuric acid was 444 more prominent for 3-7 nm particles on 6 NPF days (Nov. 25th, Nov. 26th, Nov. 30th, 445 Dec. 10th, Dec. 11th, and Dec. 12th 2013), whereas it was more significant for 7-20 nm 446 particles on the other 5 NPF days (Nov. 28th 2013, Nov. 29th 2013, Jan. 13th 2014, Jan. 447 21st 2014, and Jan. 24th 2014). On average, condensation of gas-phase hydrated 448 sulfuric acid explained 39.1% of $GR_{239\sim7}$, and 29.0% of $GR_{7\sim20}$, respectively. The rest 449 of growth might be largely attributed to condensation of extremely low volatility 450 organic compounds (Ehn et al., 2014), and potentially heterogeneous reactions of 451 organics on nanoparticle surface (Wang et al., 2010; Wang et al., 2011). 452

453

454 3.4 NPF and Aerosol Surface Area

NPF events in urban environment are of special interests since the pre-existing 455 particle surface may significantly scavenge the newly formed particles and change the 456 probability of NPF. We plot number concentrations of 1.34-10 nm particles ($N_{1.34-10}$), 457 sulfuric acid proxy ($[H_2SO_4]$), ammonia ($[NH_3]$), and aerosol surface area with 458 shadowed blocks representing NPF events in Figure 7. Note that $N_{1.34-10}$ was used as 459 an approximation for nucleation and subsequent growth while excluding particles 460 from direct emission. The average daytime (6 am-6 pm) $N_{1,34-10}$ on NPF days was 461 $(2.7\pm2.1)\times10^4$ cm⁻³, much larger than $(1.5\pm1.0)\times10^4$ cm⁻³ on the rest days of the 462 463 campaign, indicating that a stronger input of particles from nucleation processes on NPF days. However, daytime $[H_2SO_4]$ did not show an apparent difference between 464 on NPF days ((3.7±1.2) ×10⁷ molecules cm⁻³) and on the rest days ((3.9±2.5) ×10⁷ 465 molecules cm⁻³). For example, an episode with high daily sulfuric acid proxy during 466 Dec. 19th, 2013 and Jan. 16th, 2014 did not lead to any NPF events. Instead, observed 467 NPF events occurred on days with low aerosol surface area levels and moderate 468 $[H_2SO_4]$. During our campaign, NPF days were characterized with low aerosol 469 surface area ((7.8 \pm 3.8) ×10⁸ nm³ cm⁻³), whereas the average was (10.4 \pm 4.7) ×10⁸ nm³ 470 cm⁻³ on the rest days of the campaign. Ammonia varied dramatically, even within a 471 single day. NPF events occurred on days with around 10-fold difference in ammonia 472 concentrations. According to ammonia's profile and its positive correlation with $J_{1,34}$, 473 we speculate that ammonia was involved in nucleation but it is not the driving force. 474

475

Examination of these parameters from Jan. 12^{th} to Jan. 14^{th} , 2014 was performed since the three days were characterized with similar meteorological conditions. The average daytime concentration of sulfuric acid proxy was 2.8, 2.3, and 1.0×10^7 molecules cm⁻ ³ on Jan. 12^{th} , 13^{th} and 14^{th} , 2014, respectively. No NPF event was observed on Jan. 14^{th} at least partially because of the low sulfuric acid proxy. Appearance of a NPF

event on Jan. 13th and non-appearance on Jan, 12th could be explained by the high 481 aerosol surface area on Jan. 12th, with maximum aerosol surface area up to 1.8×10^9 482 nm³ cm⁻³, although similar sulfuric acid proxies existed between the two days. Hence, 483 we conclude that, qualitatively, NPF processes in urban Shanghai occurred with low 484 levels of aerosol surface and that high sulfuric acid favored NPF events when aerosol 485 surface area was low. This conclusion is identical to that drawn from a Mexico City 486 study where NPF events observed in the city correlated with elevated concentrations 487 of sulfur dioxide and low particulate matter mass concentrations in the afternoon 488 hours (Dunn et al., 2004). 489

490

491 **4 Summary and Conclusions**

Atmospheric new particle formation is a significant source of atmospheric aerosol 492 particles. Understanding NPF under the current levels of air pollution in China is of 493 special scientific interests because the exact nucleation mechanism under urban 494 environment remains elusive. From Nov. 25th, 2013 to Jan. 25th, 2014, a combination 495 496 of one nano-CNC, one nano-SPMS, and one long-SPMS has been utilized to investigate atmospheric nucleation by measuring particle size distributions in the 497 range of 1.34-615 nm at urban Shanghai, located in the east Yangtze River Delta. 498 During this 62-day campaign, 13 NPF events were identified with strong burst of sub-499 3 nm particles and subsequent fast growth of newly formed particles. Together with 500 nucleation frequency (21%), the obtained nucleation rate $J_{1.34}$ (112.4-271.0 cm⁻³ s⁻¹), 501 condensation sink CS (0.030-0.10 s⁻¹), and aerosol surface area ((7.8 \pm 3.8) \times 10⁸ nm³ 502 cm⁻³) on NPF event days clearly indicate that strong atmospheric new particle 503 formation occurred in winter, urban Shanghai with competition between promotion 504 from condensable vapors and scavenging by preexisting particles. The absolute values 505 of $J_{1,34}$ and CS are one to two orders of magnitude lager than those at locations with a 506 pristine atmosphere (e.g., Kulmala et al., 2012) and semi-rural locations with very low 507 pollution levels such as Kent, OH (Yu et al., 2014), as a reflection from the 508 significantly-altered atmospheric background. 509

Our preliminary exploration on nucleation mechanism indicates that nucleation rate $J_{1.34}$ was proportional to a 0.65 \pm 0.28 power of sulfuric acid proxy. It is hence likely that observed NPF events could be explained by the activation theory. As Herrmann et al., (2014) doubted reliability of sulfuric acid proxy, accurate measurements of gasphase sulfuric acid instead of calculation of a proxy is necessary to achieve an unambiguous conclusion. The positive correlation between $J_{1.34}$ and gas-phase ammonia hints the involvement of ammonia in new particle formation, but its exact

role cannot be determined without measurements of nucleating clusters, either.

519

520 A size-depend particle growth in the range of 1.34-20 nm has been observed in this study, consistent with predictions from nano-Köhler theory (Kulmala et al., 2004), and 521 Kelvin-limited diffusion, surface, and volume growth laws (Kuang et al., 2012). 522 Sulfuric acid proxy was enough to explain the observed growth for particles under 3 523 nm, and contributed to 39.1% of $GR_{2,39-7}$, and 29.0% of GR_{7-20} , respectively. The rest 524 of growth could be largely attributed to condensation of extremely low volatility 525 organic compounds (Ehn et al., 2014), and potentially heterogeneous reactions of 526 organics on nanoparticle surface (Wang et al., 2010; Wang et al., 2011). 527

528

529 Clearly, further long term measurements with integrated state-of-the-art measurement 530 techniques are necessary to draw a comprehensive picture of atmospheric nucleation 531 events in China. Currently, atmospheric sub-3 nm particle measurements are still 532 scarce in China, not mentioning measurements of gas-phase sulfuric acid, nucleating 533 clusters, and other potential precursors. Nevertheless, our study offers a very first 534 measurement of sub-3 nm particles in urban Shanghai, and provides some of the 535 preliminary mechanisms for NPF events in China.

536

537 Acknowledgments

538 This study was financially supported by National Natural Science Foundation of

19

539 China (No. 21107015, 21190053, 21277029 & 21222703), Ministry of Science &

- 540 Technology of China (2012YQ220113-4), and Science & Technology Commission of
- 541 Shanghai Municipality (12DJ1400100). LW thanks the Jiangsu Provincial 2011
- 542 Program (Collaborative Innovation Centre of Climate Change).
- 543

544 **References:**

- 545 Andreae, M. O., and Crutzen, P. J.: Atmospheric aerosols: Biogeochemical sources and role in 546 atmospheric chemistry, Science, 276, 1052-1058, 1997.
- Anttila, T., Kerminen, V.-M., and Lehtinen, K. E. J.: Parameterizing the formation rate of new particles:
 The effect of nuclei self-coagulation, Journal of Aerosol Science, 41, 621-636,
 10.1016/j.jaerosci.2010.04.008, 2010.
- Ball, S., Hanson, D., Eisele, F., and McMurry, P.: Laboratory studies of particle nucleation: Initial results
 for H₂SO₄, H₂O, and NH₃ vapors, Journal of Geophysical Research: Atmospheres (1984–2012), 104,
 23709-23718, 1999.
- Benson, D. R., Erupe, M. E., and Lee, S. H.: Laboratory-measured H₂SO₄-H₂O-NH₃ ternary
 homogeneous nucleation rates: Initial observations, Geophysical Research Letters, 36, L15818, doi:
 10.1029/2009GL038728, 2009.
- Berndt, T., Stratmann, F., Sipilä, M., Vanhanen, J., Petäjä, T., Mikkilä, J., Grüner, A., Spindler, G.,
 Mauldin III, L., and Curtius, J.: Laboratory study on new particle formation from the reaction OH+
 SO₂: influence of experimental conditions, H₂O vapour, NH₃ and the amine tert-butylamine on the
 overall process, Atmospheric Chemistry and Physics, 10, 7101-7116, 2010.
- 560 Chan, C. K., and Yao, X.: Air pollution in mega cities in China, Atmospheric Environment, 42, 1-42, 561 2008.
- 562 Dal Maso, M. D., Kulmala, M., Riipinen, I., Wagner, R., Hussein, T., Aalto, P. P., and Lehtinen, K. E.:
 563 Formation and growth of fresh atmospheric aerosols: eight years of aerosol size distribution data
 564 from SMEAR II, Hyytiälä, Finland, Boreal Environment Research, 10, 323-336, 2005.
- 565 Dockery, D. W., Pope, C. A., Xu, X., Spengler, J. D., Ware, J. H., Fay, M. E., Ferris Jr, B. G., and Speizer, F.
 566 E.: An association between air pollution and mortality in six US cities, New England journal of
 567 medicine, 329, 1753-1759, 1993.
- 568 Du, J., Cheng, T., Zhang, M., Chen, J., He, Q., Wang, X., Zhang, R., Tao, J., Huang, G., and Li, X.: Aerosol
 569 Size Spectra and Particle Formation Events at Urban Shanghai in Eastern China, Aerosol Air Qual.
 570 Res, 12, 1362-1372, 2012.
- 571 Dunn, M. J., Jiménez, J. L., Baumgardner, D., Castro, T., McMurry, P. H., and Smith, J. N.: Measurements
 572 of Mexico City nanoparticle size distributions: Observations of new particle formation and growth,
 573 Geophysical Research Letters, 31, L10102, doi:10.1029/2004GL019483, 2004.
- Ehn, M., Thornton, J. A., Kleist, E., Sipilä, M., Junninen, H., Pullinen, I., Springer, M., Rubach, F.,
 Tillmann, R., and Lee, B.: A large source of low-volatility secondary organic aerosol, Nature, 506,
 476-479, 2014.
- Eisele, F., and Tanner, D.: Measurement of the gas phase concentration of H₂SO₄ and methane sulfonic
 acid and estimates of H₂SO₄ production and loss in the atmosphere, Journal of Geophysical
 Research: Atmospheres (1984–2012), 98, 9001-9010, 1993.

- Erupe, M. E., Benson, D. R., Li, J., Young, L.-H., Verheggen, B., Al-Refai, M., Tahboub, O., Cunningham,
 V., Frimpong, F., Viggiano, A. A., and Lee, S.-H.: Correlation of aerosol nucleation rate with sulfuric
 acid and ammonia in Kent, Ohio: An atmospheric observation, Journal of Geophysical Research,
 115, D23216, doi:10.1029/2010jd013942, 2010.
- Gao, J., Chai, F., Wang, T., Wang, S., and Wang, W.: Particle number size distribution and new particle
 formation: New characteristics during the special pollution control period in Beijing, Journal of
 Environmental Sciences, 24, 14-21, 10.1016/s1001-0742(11)60725-0, 2012.
- Guo, H., Wang, D. W., Cheung, K., Ling, Z. H., Chan, C. K., and Yao, X. H.: Observation of aerosol size
 distribution and new particle formation at a mountain site in subtropical Hong Kong, Atmospheric
 Chemistry and Physics, 12, 9923-9939, 10.5194/acp-12-9923-2012, 2012.
- Haywood, J., and Boucher, O.: Estimates of the direct and indirect radiative forcing due to tropospheric
 aerosols: A review, Reviews of Geophysics, 38, 513, 10.1029/1999rg000078, 2000.
- Herrmann, E., Ding, A., Kerminen, V.-M., Petäjä, T., Yang, X., Sun, J., Qi, X., Manninen, H., Hakala, J.,
 and Nieminen, T.: Aerosols and nucleation in eastern China: first insights from the new SORPESNJU station, Atmospheric Chemistry and Physics, 14, 2169-2183, 2014.
- Hirsikko, A., Nieminen, T., Gagné, S., Lehtipalo, K., Manninen, H., Ehn, M., Horrak, U., Kerminen, V.-M.,
 Laakso, L., and McMurry, P.: Atmospheric ions and nucleation: a review of observations,
 Atmospheric Chemistry and Physics, 11, 767-798, 2011.
- IPCC: IPCC, 2013: Climate change 2013: The physical science basis, Contribution of working group I to
 the fourth assessment report of the intergovernmental panel on climate change Cambridge
 University Press, Cambridge, United Kingdom and New York, NY, USA, 2013.
- Kanawade, V. P., Benson, D. R., and Lee, S.-H.: Statistical analysis of 4-year observations of aerosol sizes
 in a semi-rural continental environment, Atmospheric Environment, 59, 30-38,
 10.1016/j.atmosenv.2012.05.047, 2012.
- Korhonen, H., Sihto, S. L., Kerminen, V. M., and Lehtinen, K. E. J.: Evaluation of the accuracy of analysis
 tools for atmospheric new particle formation, Atmospheric Chemistry and Physics, 11, 3051-3066,
 10.5194/acp-11-3051-2011, 2011.
- Kuang, C., Chen, M., Zhao, J., Smith, J., McMurry, P. H., and Wang, J.: Size and time-resolved growth
 rate measurements of 1 to 5 nm freshly formed atmospheric nuclei, Atmospheric Chemistry and
 Physics, 12, 3573-3589, 10.5194/acp-12-3573-2012, 2012.
- Kulmala, M., Toivonen, A., Mäkelä, J. M., and Laaksonen, A.: Analysis of the growth of nucleation
 mode particles observed in Boreal forest, Tellus B, 50, 449-462, 1998.
- Kulmala, M., Kerminen, V. M., Anttila, T., Laaksonen, A., and O'Dowd, C. D.: Organic aerosol formation
 via sulphate cluster activation, Journal of Geophysical Research: Atmospheres (1984–2012), 109,
 D04205, doi: 10.1029/2003JD003961, 2004.
- Kulmala, M., Vehkamäki, H., Petäjä, T., Dal Maso, M., Lauri, A., Kerminen, V. M., Birmili, W., and
 McMurry, P. H.: Formation and growth rates of ultrafine atmospheric particles: a review of
 observations, Journal of Aerosol Science, 35, 143-176, 10.1016/j.jaerosci.2003.10.003, 2004.
- Kulmala, M., Lehtinen, K., and Laaksonen, A.: Cluster activation theory as an explanation of the linear
 dependence between formation rate of 3nm particles and sulphuric acid concentration,
 Atmospheric Chemistry and Physics, 6, 787-793, 2006.
- Kulmala, M., Riipinen, I., Sipilä, M., Manninen, H. E., Petäjä, T., Junninen, H., Dal Maso, M., Mordas, G.,
 Mirme, A., and Vana, M.: Toward direct measurement of atmospheric nucleation, Science, 318, 8992, 2007.

- Kulmala, M., Petäjä, T., Nieminen, T., Sipilä, M., Manninen, H. E., Lehtipalo, K., Dal Maso, M., Aalto, P.
 P., Junninen, H., Paasonen, P., Riipinen, I., Lehtinen, K. E. J., Laaksonen, A., and Kerminen, V.-M.:
 Measurement of the nucleation of atmospheric aerosol particles, Nature Protocols, 7, 1651-1667,
 10.1038/nprot.2012.091, 2012.
- Kulmala, M., Kontkanen, J., Junninen, H., Lehtipalo, K., Manninen, H. E., Nieminen, T., Petaja, T., Sipila,
 M., Schobesberger, S., Rantala, P., Franchin, A., Jokinen, T., Jarvinen, E., Aijala, M., Kangasluoma, J.,
 Hakala, J., Aalto, P. P., Paasonen, P., Mikkila, J., Vanhanen, J., Aalto, J., Hakola, H., Makkonen, U.,
 Ruuskanen, T., Mauldin, R. L., Duplissy, J., Vehkamaki, H., Back, J., Kortelainen, A., Riipinen, I.,
 Kurten, T., Johnston, M. V., Smith, J. N., Ehn, M., Mentel, T. F., Lehtinen, K. E. J., Laaksonen, A.,
 Kerminen, V. M., and Worsnop, D. R.: Direct Observations of Atmospheric Aerosol Nucleation,
 Science, 339, 943-946, 10.1126/science.1227385, 2013.
- Kupiainen-Määttä, O., Olenius, T., Korhonen, H., Malila, J., Dal Maso, M., Lehtinen, K., and Vehkamäki,
 H.: Critical cluster size cannot in practice be determined by slope analysis in atmospherically
 relevant applications, Journal of Aerosol Science, 77, 127-144, doi:10.1016/j.jaerosci.2014.07.005,
 2014.
- Kurtén, T., Noppel, M., Vehkamaeki, H., Salonen, M., and Kulmala, M.: Quantum chemical studies of
 hydrate formation of H₂SO₄ and HSO₄⁻, Boreal Environment Research, 12, 431-453, 2007.
- Laden, F., Schwartz, J., Speizer, F. E., and Dockery, D. W.: Reduction in fine particulate air pollution and
 mortality: extended follow-up of the Harvard Six Cities study, American journal of respiratory and
 critical care medicine, 173, 667-672, 2006.
- Lee, S.-H., Reeves, J., Wilson, J., Hunton, D., Viggiano, A., Miller, T., Ballenthin, J., and Lait, L.: Particle
 formation by ion nucleation in the upper troposphere and lower stratosphere, Science, 301, 18861889, 2003.
- Lehtinen, K., and Kulmala, M.: A model for particle formation and growth in the atmosphere with
 molecular resolution in size, Atmospheric Chemistry and Physics, 3, 251-257, 2003.

Ma, Y., Xu, X., Song, W., Geng, F., and Wang, L.: Seasonal and diurnal variations of particulate
 organosulfates in urban Shanghai, China, Atmospheric Environment, 85, 152-160, 2014.

- Matsui, H., Koike, M., Kondo, Y., Takegawa, N., Wiedensohler, A., Fast, J. D., and Zaveri, R. A.: Impact of
 new particle formation on the concentrations of aerosols and cloud condensation nuclei around
 Beijing, Journal of Geophysical Research: Atmospheres, 116, D19208, doi: 10.1029/2011JD016025,
 2011.
- McMurry, P. H., Fink, M., Sakurai, H., Stolzenburg, M., Mauldin, R., Smith, J., Eisele, F., Moore, K.,
 Sjostedt, S., and Tanner, D.: A criterion for new particle formation in the sulfur-rich Atlanta
 atmosphere, Journal of Geophysical Research: Atmospheres, 110, D22S02, doi:
 10.1029/2005JD005901, 2005.
- Mikkonen, S., Romakkaniemi, S., Smith, J. N., Korhonen, H., Petäjä, T., Plass-Duelmer, C., Boy, M.,
 McMurry, P. H., Lehtinen, K. E. J., Joutsensaari, J., Hamed, A., Mauldin Iii, R. L., Birmili, W., Spindler,
 G., Arnold, F., Kulmala, M., and Laaksonen, A.: A statistical proxy for sulphuric acid concentration,
 Atmospheric Chemistry and Physics, 11, 11319-11334, 10.5194/acp-11-11319-2011, 2011.
- Napari, I., Noppel, M., Vehkamäki, H., and Kulmala, M.: Parametrization of ternary nucleation rates for
 H₂SO₄-NH₃-H₂O vapors, Journal of Geophysical Research: Atmospheres (1984–2012), 107, AAC 6-1 AAC 6-6, 2002.
- Nieminen, T., Lehtinen, K. E. J., and Kulmala, M.: Sub-10 nm particle growth by vapor condensation –
 effects of vapor molecule size and particle thermal speed, Atmospheric Chemistry and Physics, 10,

- 668 9773-9779, 10.5194/acp-10-9773-2010, 2010.
- O'Dowd, C. D., Jimenez, J. L., Bahreini, R., Flagan, R. C., Seinfeld, J. H., Hämeri, K., Pirjola, L., Kulmala,
 M., Jennings, S. G., and Hoffmann, T.: Marine aerosol formation from biogenic iodine emissions,
 Nature, 417, 632-636, 2002.
- Park, J., Sakurai, H., Vollmers, K., and McMurry, P. H.: Aerosol size distributions measured at the South
 Pole during ISCAT, Atmospheric Environment, 38, 5493-5500, 2004.
- Platt, U., and Stutz, J.: Differential Optical Absorption Spectroscopy-Principles and Applications,
 Springer Berlin Heidelberg, 2008.
- Pope, C. A., and Dockery, D. W.: Health effects of fine particulate air pollution: lines that connect,
 Journal of the Air & Waste Management Association, 56, 709-742, 2006.
- Riccobono, F., Schobesberger, S., Scott, C. E., Dommen, J., Ortega, I. K., Rondo, L., Almeida, J., Amorim,
 A., Bianchi, F., and Breitenlechner, M.: Oxidation Products of Biogenic Emissions Contribute to
 Nucleation of Atmospheric Particles, Science, 344, 717-721, 2014.
- Schobesberger, S., Junninen, H., Bianchi, F., Lönn, G., Ehn, M., Lehtipalo, K., Dommen, J., Ehrhart, S.,
 Ortega, I. K., and Franchin, A.: Molecular understanding of atmospheric particle formation from
 sulfuric acid and large oxidized organic molecules, Proceedings of the National Academy of
 Sciences of the United States of America, 110, 17223-17228, 2013.
- Sihto, S.-L., Kulmala, M., Kerminen, V.-M., Maso, M. D., Petäjä, T., Riipinen, I., Korhonen, H., Arnold, F.,
 Janson, R., and Boy, M.: Atmospheric sulphuric acid and aerosol formation: implications from
 atmospheric measurements for nucleation and early growth mechanisms, Atmospheric Chemistry
 and Physics, 6, 4079-4091, 2006.
- Sipila, M., Berndt, T., Petaja, T., Brus, D., Vanhanen, J., Stratmann, F., Patokoski, J., Mauldin, R. L.,
 Hyvarinen, A. P., Lihavainen, H., and Kulmala, M.: The Role of Sulfuric Acid in Atmospheric
 Nucleation, Science, 327, 1243-1246, 10.1126/science.1180315, 2010.
- Stolzenburg, M. R., McMurry, P. H., Sakurai, H., Smith, J. N., Mauldin, R. L., Eisele, F. L., and Clement, C.
 F.: Growth rates of freshly nucleated atmospheric particles in Atlanta, Journal of Geophysical
 Research, 110, 10.1029/2005jd005935, 2005.
- Vanhanen, J., Mikkilä, J., Lehtipalo, K., Sipilä, M., Manninen, H., Siivola, E., Petäjä, T., and Kulmala, M.:
 Particle size magnifier for nano-CN detection, Aerosol Science and Technology, 45, 533-542, 2011.
- Wang, L., Khalizov, A. F., Zheng, J., Xu, W., Ma, Y., Lal, V., and Zhang, R.: Atmospheric nanoparticles
 formed from heterogeneous reactions of organics, Nature Geoscience, 3, 238-242, 2010.
- Wang, L., Xu, W., Khalizov, A. F., Zheng, J., Qiu, C., and Zhang, R.: Laboratory investigation on the role
 of organics in atmospheric nanoparticle growth, The Journal of Physical Chemistry A, 115, 89408947, 2011.
- Wang, L., Du, H., Chen, J., Zhang, M., Huang, X., Tan, H., Kong, L., and Geng, F.: Consecutive transport
 of anthropogenic air masses and dust storm plume: Two case events at Shanghai, China,
 Atmospheric Research, 127, 22-33, 2013.
- Wang, S., Zhou, R., Zhao, H., Wang, Z., Chen, L., and Zhou, B.: Long-term observation of atmospheric
 nitrous acid (HONO) and its implication to local NO₂ levels in Shanghai, China, Atmospheric
 Environment, 77, 718-724, 10.1016/j.atmosenv.2013.05.071, 2013.
- Weber, R., Marti, J., McMurry, P., Eisele, F., Tanner, D., and Jefferson, A.: Measured atmospheric new
 particle formation rates: Implications for nucleation mechanisms, Chemical Engineering
 Communications, 151, 53-64, 1996.
- 711 Wu, Z., Hu, M., Liu, S., Wehner, B., Bauer, S., Ma ßling, A., Wiedensohler, A., Petäjä, T., Dal Maso, M.,

- and Kulmala, M.: New particle formation in Beijing, China: Statistical analysis of a 1-year data set,
 Journal of Geophysical Research-Atmospheres, 112, D09209, doi:10.1029/2006jd007406, 2007.
- Yu, F., and Turco, R. P.: From molecular clusters to nanoparticles: Role of ambient ionization in
 tropospheric aerosol formation, Journal of Geophysical Research: Atmospheres (1984–2012), 106,
 4797-4814, 2001.
- Yu, F. Q., and Hallar, A. G.: Difference in particle formation at a mountaintop location during spring and
 summer: Implications for the role of sulfuric acid and organics in nucleation, Journal of Geophysical
 Research-Atmospheres, 119, 12246-12255, 2014.
- Yu, H., Gannet Hallar, A., You, Y., Sedlacek, A., Springston, S., Kanawade, V. P., Lee, Y. N., Wang, J.,
 Kuang, C., and McGraw, R. L.: Sub-3 nm particles observed at the coastal and continental sites in
 the United States, Journal of Geophysical Research: Atmospheres, 119, 860-879, doi:
 10.1002/2013JD020841, 2014.
- Yue, D., Hu, M., Wu, Z., Wang, Z., Guo, S., Wehner, B., Nowak, A., Achtert, P., Wiedensohler, A., Jung, J.,
 Kim, Y. J., and Liu, S.: Characteristics of aerosol size distributions and new particle formation in the
 summer in Beijing, Journal of Geophysical Research, 114, D00G12, doi:10.1029/2008jd010894,
 2009.
- Yue, D. L., Hu, M., Zhang, R. Y., Wang, Z. B., Zheng, J., Wu, Z. J., Wiedensohler, A., He, L. Y., Huang, X. F.,
 and Zhu, T.: The roles of sulfuric acid in new particle formation and growth in the mega-city of
 Beijing, Atmospheric Chemistry and Physics, 10, 4953-4960, 10.5194/acp-10-4953-2010, 2010.
- Yue, D. L., Hu, M., Zhang, R. Y., Wu, Z. J., Su, H., Wang, Z. B., Peng, J. F., He, L. Y., Huang, X. F., Gong, Y.
 G., and Wiedensohler, A.: Potential contribution of new particle formation to cloud condensation
 nuclei in Beijing, Atmospheric Environment, 45, 6070-6077, 10.1016/j.atmosenv.2011.07.037,
 2011.
- Yue, D. L., Hu, M., Wang, Z. B., Wen, M. T., Guo, S., Zhong, L. J., Wiedensohler, A., and Zhang, Y. H.:
 Comparison of particle number size distributions and new particle formation between the urban
 and rural sites in the PRD region, China, Atmospheric Environment, 76, 181-188,
 10.1016/j.atmosenv.2012.11.018, 2013.
- Zhang, R., Suh, I., Zhao, J., Zhang, D., Fortner, E. C., Tie, X., Molina, L. T., and Molina, M. J.: Atmospheric
 new particle formation enhanced by organic acids, Science, 304, 1487-1490, 2004.
- Zhang, R., Wang, L., Khalizov, A. F., Zhao, J., Zheng, J., McGraw, R. L., and Molina, L. T.: Formation of
 nanoparticles of blue haze enhanced by anthropogenic pollution, Proceedings of the National
 Academy of Sciences of the United States of America, 106, 17650-17654, 2009.
- Zhang, R., Khalizov, A., Wang, L., Hu, M., and Xu, W.: Nucleation and Growth of Nanoparticles in the
 Atmosphere, Chemical Reviews, 112, 1957-2011, 10.1021/cr2001756, 2012.
- Zhang, Y., Zhang, X., Sun, J., Lin, W., Gong, S., Shen, X., and Yang, S.: Characterization of new particle
 and secondary aerosol formation during summertime in Beijing, China, Tellus B, 63, 382-394, 2011.
- Zhao, J., Smith, J., Eisele, F., Chen, M., Kuang, C., and McMurry, P.: Observation of neutral sulfuric acidamine containing clusters in laboratory and ambient measurements, Atmospheric Chemistry and
 Physics, 11, 10823-10836, 2011.
- Zheng, J., Hu, M., Zhang, R., Yue, D., Wang, Z., Guo, S., Li, X., Bohn, B., Shao, M., and He, L.:
 Measurements of gaseous H₂SO₄ by AP-ID-CIMS during CAREBeijing 2008 campaign, Atmospheric
 Chemistry and Physics, 11, 7755-7765, 2011.
- Zheng, J., Ma, Y., Chen, M., Zhang, Q., Wang, L., Khalizov, A. F., Yao, L., Wang, Z., Wang, X., and Chen,
 L.: Measurement of atmospheric amines and ammonia using the high resolution time-of-flight

- 756 chemical ionization mass spectrometry, Atmospheric Environment, 102, 249-259,
 757 10.1016/j.atmosenv.2014.12.002, 2015.

Date	$J_{1.34}$	J_3	<i>GR</i> _{1.35~2.39}	<i>GR</i> _{2.39~7}	$GR_{7\sim20}$	CS e	$[H_2SO_4]$ °	N _{1.34~3} ^f	N _{1.34~615} ^f	Ref.
	(cm ⁻³ s ⁻¹)	(cm ⁻³ s ⁻¹)	(nm h ⁻¹)	(nm h ⁻¹)	(nm h ⁻¹)	$(10^{-2} s^{-1})$	(10⁷ cm ⁻³)	$(10^4 \mathrm{cm}^{-3})$	(10^4 cm^{-3})	
Nov. 25 th , 2013	n.a. ^a	10.6	n.a.	12.4 ^b	38.3	4.7	2.7	3.3 ^g	6.3	this study
Nov. 26th, 2013	n.a.	2.3	n.a.	0.32 ^b	n.a.	5.9	2.6	n.a.	n.a.	this study
Nov. 28th, 2013	185.1	13.4	0.94	35.7	4.6	5.7	3.6	1.6	4.2	this study
Nov. 29th, 2013	271.0	3.9	1.7	10.6	4.5	6.3	4.3	2.1	4.5	this study
Nov. 30th, 2013	n.a	n.a.	0.82	3.4	10.2	n.a.	3.1	1.5	n.a.	this study
Dec. 10th, 2013	268.4	10.0	0.49	18.6	21.0	9.9	5.5	1.4	4.6	this study
Dec. 11th, 2013	219.0	19.2	n.a.	5.1	9.6	10.2	6.4	1.1	4.5	this study
Dec. 12 th , 2013	190.3	7.6	n.a.	3.1	12.3	8.8	4.5	1.1	4.1	this study
Jan. 9th, 2014	136.2	n.a.	8.1	n.a.	9.5	3.7	2.3	1.6	3.8	this study
Jan. 13 th , 2014	n.a.	2.7	n.a.	6.3	1.9	3.0	2.3	1.5	3.4	this study
Jan. 15 th , 2014	121.9	n.a.	0.56	n.a.	9.7	4.2	4.1	1.5	4.3	this study
Jan. 21st, 2014	112.4	9.2	1.5	11.9	7.5	4.9	3.7	1.1	3.9	this study
Jan. 24 th , 2014	n.a.	8.1	n.a.	12.2 ^b	7.8	4.7	3.4	1.7	4.2	this study
Mean	188.0	8.7	2.0	10.9	11.4	6.0	3.7	1.5	4.4	this study
Finland	1.4	0.61	1.4	3.9	4.9	0.05-0.35				Kulmala et al. (2012)
USA ^h	1.3					0.8		0.9		Yu et al. (2014)
Atlanta					3-20°					Stolzenburg et al. (2005)
Atlanta			5.5-7.6							Kuang et al. (2012)
Beijing					1.2-8.0 ^c	2.4-3.6				Gao et al. (2012)
Beijing		3.3-81.4			0.1-11.2 ^c	0.58-4.3				Wu et al. (2007)
Beijing					2.7-13.9°	0.6-8.4				Zhang et al. (2011)
Beijing		1.1-22.4			1.2-5.6 ^c	0.9-5.3				Yue et al. (2009)
Nanjing	33.2 ⁱ	1.1 ^j		6.3	8^d	2.4				Herrmann et al. (2014)
Hong Kong		3.6-6.9			1.5-8.4 ^c	1.0-6.2				Guo et al. (2012)
Pearl River Delta		2.4-4.0(rural)			4.0-22.7(rural) 10.1-18.9(urban)	2.3-3.3(rural) 3.5-4.6(urban)				Yue et al. (2013)

Table 1. Nucleation rate $(J_{1.34})$, formation rate of 3 nm particles (J_3) , particle growth rates $(GR_{1.35\sim2.39}, GR_{2.39\sim7})$, and $GR_{7\sim20})$, condensation sink (CS), sulfuric acid proxy $([H_2SO_4])$, number concentrations of 1.34-3 nm clusters/particles $(N_{1.34\sim3})$, and total number concentrations of particles $N_{1.34\sim615}$, of each NPF event during this campaign

^a Data were not available or could not be accurately determined; ^b Result were calculated from nano-SMPS data; ^c Shown here is $GR_{3\sim30}$; ^d Shown here is $GR_{7\sim30}$; ^e Daytime average (from 6:00 am to 6:00 pm); ^f 24hr-average; ^g Average values between 10 am and 4 pm; ^h Shown here are data in Kent, OH; ⁱ Shown here is J_2 ; ^j Shown here is J_6 .

Figure Captions

Figure 1. Contour plot for particle size distributions of 3-615 nm and plot of number concentrations of sub-3 nm clusters/particles ($N_{1.34-3}$) during Nov. 25th, 2013 - Jan. 25th, 2014. Data were occasionally missing because of the maintenance and minor breakdown of instruments. NPF events are illustrated with shadows.

Figure 2. Profiles of $N_{1,34-3}$, N_{3-7} , and N_{7-30} from 6 am to 6 pm on a NPF day (Dec. 11th, 2013) and a non-NPF day (Jan. 7th, 2014), respectively.

Figure 3. Averaged particle size evolution on NPF days. Arithmetic mean of particle growth rates are given with one standard deviation.

Figure 4. Correlation between $\log J_{1.34}$ and $\log [H_2SO_4]$. Daily peak concentration of sulfuric acid proxy was used as an approximation for its effective concentration on a NPF day. The error bar corresponds to a 42% uncertainty of sulfuric acid proxy according to Mikkonen et al., (2011).

Figure 5. Correlation between $\log J_{1.34}$ and $\log [NH_3]$. Daytime average of ammonia

was used as an approximation for its effective concentration on a NPF day. The error bar represents the standard deviation of the daytime average concentration of ammonia.

Figure 6. Relative contribution of sulfuric acid to growth of particles in the range of 3-7 and 7-20 nm, respectively, on each NPF day.

Figure 7. Number concentrations of 1.34-10 nm particles $(N_{1.34-10})$, sulfuric acid proxy $([H_2SO_4])$, concentrations of ammonia, and aerosol surface area during the campaign. NPF events are illustrated with shadows.



Figure 1.



Figure 2.

Figure 3.

Figure 4.

Figure 5.

Figure 6.

Figure 7