# Understanding the anthropogenic influence on formation of biogenic secondary organic aerosols via analysis of organosulfates and related oxidation products

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#### 16 Abstract

Anthropogenic emissions of sulfur dioxide  $(SO_2)$  and nitrogen oxides  $(NO_x)$  may affect 17 concentration levels and composition of biogenic secondary organic aerosols (BSOA) through 18 photochemical reactions with biogenic organic precursors to form organosulfates and nitrooxy 19 organosulfates. We investigated this influence in a field study from May 19 - June 22, 2011 at two 20 sampling sites in Denmark. Within the study, we identified a substantial number of organic acids, 21 22 organosulfates and nitrooxy organosulfates in the ambient urban curbside and semi-rural background air. A high degree of correlation in concentrations was found among a group of specific 23 organic acids, organosulfates and nitrooxy organosulfates, which may originate from various 24 precursors, suggesting a common mechanism or factor affecting their concentration levels at the 25 26 sites. It was proposed that the formation of those species most likely occurred on a larger spatial scale with the compounds being long-range transported to the sites on the days with highest 27

concentrations. The origin of the long-range transported aerosols was investigated using the Hybrid 1 2 Single Particle Lagrangian Integrated Trajectory (HYSPLIT) model in addition to modeled emissions of related precursors including isoprene and monoterpenes using the global Model of 3 Emissions of Gases and Aerosols from Nature (MEGAN) and SO<sub>2</sub> emissions using the European 4 Monitoring and Evaluation Program (EMEP) database. The local impacts were also studied by 5 examining the correlation between selected species which showed significantly enhanced 6 concentrations at the urban curbside site and the local concentrations of various gases including 7 SO<sub>2</sub>, ozone (O<sub>3</sub>), carbon monoxide (CO), NO<sub>x</sub>, aerosol acidity and other meteorological conditions. 8 This investigation showed that an inter-play of the local parameters such as the aerosol acidity, 9 NO<sub>x</sub>, SO<sub>2</sub>, relative humidity (RH), temperature and global radiation seemed to affect the 10 concentration level of those species, suggesting the influence of aqueous aerosol chemistry. The 11 local impacts however seemed minor compared to the regional impacts. The total concentrations of 12 organosulfates and nitrooxy organosulfates contributed to approximately 0.5 - 0.8 % of PM1 mass 13 on average at the two sampling sites. 14

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#### 16 **1 Introduction**

Volatile organic compounds (VOC) are emitted from both biogenic and anthropogenic sources and 17 are oxidized to form lower volatility products partitioning between the gas and particle phase, 18 leading to the formation of secondary organic aerosols (SOA) (Donahue et al., 2006; Kroll and 19 Seinfeld, 2008; Jimenez et al., 2009; Hallquist et al., 2009). In general, the production of biogenic 20 SOA (BSOA) is estimated to be significantly larger than anthropogenic SOA (ASOA) (Goldstein 21 and Galbally, 2007; Heald et al., 2010; Spracklen et al., 2011). For example, Hallquist et al. (2009) 22 estimated a many-fold higher annual global production of BSOA (88 TgC) compared to ASOA (10 23 TgC), which were however associated with high uncertainties. Other studies suggested that ASOA 24 might have been underestimated (De Gouw and Jimenez, 2009), while some local and regional 25 studies have indicated that ASOA could be more substantial than BSOA (Aiken et al., 2009; 26 27 Fushimi et al., 2011). It is also thought that a substantial proportion of aerosols is formed through condensation of low-volatility biogenic VOC onto existing particles of anthropogenic origin 28 29 (Carlton et al., 2010), thereby blurring the division between biogenic and anthropogenic SOA. 30 Several studies have indicated that anthropogenic activities could enhance the production of BSOA 31 via different mechanisms, for example by enhancing the incorporation of biogenic VOC products into the condensed phase due to pre-existing organic aerosol from anthropogenic activities; affecting the SOA yield by the complex effect of the anthropogenic species  $NO_x$  and nitrate ( $NO_3$ ) radical (a nighttime product of  $NO_2$  and ozone) or affecting new particle formation and growth where sulfuric acid has been identified as having an essential role in new particle formation (Aiken et al., 2009; Carlton et al., 2010; Szidat et al., 2006; Szidat et al., 2009; Hoyle et al., 2011; Kulmala et al., 2004a). Further research is thus required on the actual impact of anthropogenic emissions on the formation and growth of SOA formed from biogenic VOC.

Several classes of VOC precursors, which are either reactive or could form oxidation products for 8 9 SOA formation, have been identified. Cyclic compounds are particularly important, including compounds such as cycloalkanes, aromatic hydrocarbons, and terpenes. Terpenes are typically 10 oxidized by an addition mechanism, yielding oxidation products with more than two polar 11 functional groups, increasing the possibility of forming low-volatility products (Hallquist et al., 12 2009). A number of studies have focused on the oxidation pathways of monoterpenes by hydroxyl 13 radical (OH), ozone  $(O_3)$  or nitrate radical  $(NO_3)$  forming organic acids such as pinonic and pinic 14 acid as supposed first generation oxidation products (Glasius et al., 2000; Larsen et al., 2001; Yu et 15 16 al., 1999; Hoffmann et al., 1997; Librando and Tringali, 2005; Surratt et al., 2008b; Calogirou et al., 1999; Yasmeen et al., 2012). The first generation products are further oxidized to form more 17 oxidized products, for example pinonic acid is oxidized by the OH radical to form 3-methyl-1,2,3-18 butanetricarboxylic acid (MBTCA) (Szmigielski et al., 2007; Müller et al., 2012). Another example 19 20 is the photooxidation and ozonolysis of  $\alpha$ -pinene, one of the principal species of the monoterpene class, forming terpenylic acid and diaterpenylic acid acetate (DTAA) (Claevs et al., 2009). 21

Organosulfate formation has recently been demonstrated in both laboratory studies (Surratt et al., 22 2008b; Surratt et al., 2007a; Iinuma et al., 2007b) and ambient samples (Gómez-González et al., 23 2008; Kristensen and Glasius, 2011; Worton et al., 2011; Gómez-González et al., 2012) as SOA 24 produced from the oxidation of biogenic VOC, such as isoprene,  $\alpha$ -pinene,  $\beta$ -pinene and limonene 25 in the presence of acidic sulfate particles. In addition to representing an interesting link coupling the 26 formation of SOA from biogenic VOC precursors with the impact from anthropogenic pollutants, 27 organosulfates were also estimated to contribute up to 5-10% of the total organic mass (Tolocka and 28 Turpin, 2012) or even 30% of the total aerosol mass (Surratt et al., 2008b). Organosulfates and 29 30 nitrooxy organosulfates are also polar compounds, which can thus enhance the ability of aerosols to act as cloud condensation nuclei (CCN) with important climate implications (Hallquist et al., 2009; 31 IPCC, 2007). 32

Here we studied the occurrence, concentrations and trends of the acidic monoterpene oxidation 1 2 products (including for example *cis*-pinic acid, pinonic acid, terpenylic acid, MTBCA and DTAA), oxidation products from anthropogenic VOC precursors, benzoic acid and phthalic acid (typically 3 from aromatic hydrocarbons) or adipic acid and pimelic acid (from cyclic olefins), organosulfates, 4 and nitrooxy organosulfates in aerosols at two different locations in Denmark. These include an 5 urban curbside site and a semi-rural background site. The chemical analysis was performed 6 following an approach adapted from Kristensen and Glasius (2011), employing high-performance 7 liquid chromatography (HPLC) coupled with an electrospray ionization inlet (ESI) to a quadrupole 8 time-of-flight mass spectrometer (q-TOF-MS). The impacts of long-range transported SOA were 9 investigated using the Hybrid Single Particle Lagrangian Integrated Trajectory (HYSPLIT) model 10 assisted by the modeled emissions of the major biogenic precursors including isoprene and 11 monoterpene using the global Model of Emissions of Gases and Aerosols from Nature (MEGAN) 12 and the anthropogenic emissions of sulfur dioxide (SO<sub>2</sub>) using the long-range chemistry transport 13 model, the Danish Eulerian Hemispheric Model (DEHM), based on the European Monitoring and 14 Evaluation Program (www.EMEP.int) database. Additional support analyses were also performed, 15 including measurement of major inorganic ions to determine the aerosol acidity level, analysis of 16 local new particle formation events, local concentrations of atmospheric oxidants including ozone 17  $(O_3)$ , carbon monoxide (CO), sulfur dioxide  $(SO_2)$  and nitrogen oxides  $NO_x$   $(NO_x = NO + NO_2)$  and 18 other meteorological parameteres including the global radiation level, temperature, relative 19 20 humidity (RH), wind direction and wind speed.

#### 21 2 Experimental

#### 22 2.1 Field campaign sites

The campaign was conducted during the period May 19 - June 22, 2011 concurrently at two 23 sampling sites in Denmark (Figure 1). The sampling site HCAB (55°38' N, 12°34' E) is located in 24 central Copenhagen (~1.2 million inhabitants) on the curbside of H.C. Andersen's Boulevard with 25 considerable traffic influence (~50,000 vehicles per day) (Municipality of Copenhagen, 26 www.kk.dk). The second sampling site Risø (55°41' N, 12°05' E) is a semi-rural background site 27 located approximately 30 km west of Copenhagen and 7 km northeast of Roskilde, which houses 28 29 about 46,000 inhabitants. The semi-rural background area is characterized by agricultural land, 30 small villages, Roskilde Fjord located about 100 m to the west of the station, and a main road (A6) 31 located about 700 m east of the site. PM<sub>1</sub> samples were collected on quartz fiber filters (150 mm

diameter, Advantec QR-100) using Digitel DHA-80 High Volume Samplers (HVS) equipped with 1 2 PM<sub>1</sub> sampling heads. In order to assess the diurnal variation on PM<sub>1</sub> composition, sampling was performed in 12h daytime and nighttime intervals starting every day at 06:00 am and 06:00 pm 3 local time simultaneously at both sites. The sampling flow rate was set to 23.1 m<sup>3</sup> h<sup>-1</sup> providing a 4 nominal sampling volume of 277.2 m<sup>3</sup> per 12 h sample. While this 12 h sampling interval allowed 5 us to sample sufficient volume of the tracer compounds, the coarse time resolution complicates 6 investigation of any short-term chemistry or transport episodes. There were several gaps in the 7 8 dataset with one major gap at the Risø site from June 3 - 16 due to technical problems with the 9 HVS. In total, 64 and 37 filter samples (including both day and night samples) were collected at 10 HCAB and Risø, respectively. A field blank was collected in between every 12 samples or less, which showed concentrations below detection limit for the organic acids, organosulfates and 11 nitrooxy organosulfates. The field blanks were however used for correction of concentrations of 12 inorganic ions. 13

#### 14 2.2 Extraction

15 Extraction and analysis of organic acids, organosulfates and nitrooxy organosulfates followed the procedure of Kristensen and Glasius (2011). Each filter sample (excluding a circle punch of 28 mm 16 in diameter used for separate analysis of ions) was extracted in 75 mL of a solution consisting of 17 90% (v/v) acetonitrile and 10% (v/v) milli-Q water by ultra-sonication for 30 minutes in a water 18 bath, where ice was regularly added to the bath to avoid heating. The extract was filtered through a 19 20 nylon syringe filter (0.45 µm, VWR or Q-max), evaporated to dryness using a rotary evaporator and re-constituted twice in 500  $\mu$ L milli-Q water containing 3% (v/v) acetonitrile and 0.1% (v/v) acetic 21 acid. Prepared samples were kept at 3-5 °C prior to analysis. Camphoric acid was used as an 22 internal standard to calculate the recovery. Field blanks were extracted and analyzed using the same 23 procedure as for the samples. 24

#### 25 **2.3 HPLC-ESI-qTOF-MS analysis and quantification of compounds**

Samples were analyzed using a HPLC system (Dionex Ultimate 3000) coupled through an ESI inlet to a qTOF-MS (Bruker Daltonics GmbH, Bremen, Germany). A reversed-phase HPLC separation was achieved using a Waters Atlantis T3 C18 column (2.1 x 150 mm, 3  $\mu$ m). The mobile phase was comprised of eluent A (0.1% (v/v) acetic acid in milli-Q water) and eluent B (95% (v/v) acetonitrile in milli-Q water). The injection volume was 10  $\mu$ L and the flow rate of the mobile phase was 0.2 mL min<sup>-1</sup> with the following gradient: the mobile phase was kept at 3% of eluent B in 10 min prior to injection until 5 min after injection; eluent B was then increased to 20% (at 20 min), 60% (at 30 min) and 95% (at 35 min) where it was kept at a plateau (95%) until 45 min and increased to 100% at 46 min; the gradient was kept at 100% of eluent B until 51 min and decreased to 3% (at 54 min)
where it was further kept at 3% for 3 min. The MS scan range was set from 50 - 1000 m/z. The ESI-qTOF-MS was operated in negative ionization mode with nebulizer pressure at 3.0 bar, dry gas flow 8.0 L min<sup>-1</sup>, source voltage 4000 V and transfer time 50 µs. Acquired data was processed using the Bruker Compass software.

8 All organic acids were quantified using authentic standards with the exception of hydroxy-pinonic 9 acid, which was quantified using pinonic acid standard. The organosulfates and nitrooxy 10 organosulfates were quantified using three different surrogate standards depending on the retention 11 time of the compound. The surrogate standards included D-mannose sulfate (molecular weight = 12 260 g mol<sup>-1</sup> (MW 260)) for compounds eluting in the first 15 min, an organosulfate from  $\beta$ -pinene 13 (MW 250), which was synthesized in-house following Iinuma et al. (2009) for compounds eluting 14 during 15 - 40 min, and octyl sulfate (MW 210) for compounds eluting after 40 min.

#### 15 **2.4 Extraction and analysis of ions**

A circle punch (28 mm diameter) of the aerosol sample filter was sonicated in 4.00 ml milli-Q water 16 for 1/2 h. The extract was filtered through a polyethersulfone syringe filter (0.45 µm, PALL Life 17 Sciences IC Acrodisc) and analyzed by ion chromatography using a Metrohm 820 IC separation 18 center (with a Metrosep A Supp5-150 column for anion analysis and a Metrosep C4-150 column for 19 cation analysis) and an 819 IC conductivity detector. For anion analysis, an eluent of 3.2 mM 20 Na<sub>2</sub>CO<sub>3</sub>/1.0 mM NaHCO<sub>3</sub> was used. For cation analysis, an eluent of 0.7 mM dipicolinic acid/1.7 21 mM HNO<sub>3</sub> was used. An auto-sampler was used where 100 µL of each sample was introduced into 22 the eluent, which was kept at a flow rate of  $0.7 \text{ mL min}^{-1}$ . The inclusion of ion analysis in this work 23 is mainly to calculate the level of acidity ( $H^+$ ) from the charge balance, based on IC data of [SO<sub>4</sub><sup>2-</sup>], 24 25  $[NO_3]$  and  $[NH_4^+]$  following Zhang et al. (2012b).

#### 26 **2.5 Measurement of particle number size distributions**

Particle number size distributions in 15 channels from 6-700 nm were measured using a Differential
Mobility Particle Sizer (DMPS), which was composed of a Vienna-type Differential Mobility
Analyzer (DMA) (Winklmayer et al., 1991) employing re-circulating sheath air flow (Jokinen and
Makela, 1997) and a butanol Condensation Particle Counter (CPC) (TSI Inc. 3010, Shoreview,
MN). The data was corrected for reduced counting efficiency at the low size end (6-10 nm),

multiply charging, and losses during sampling following inversion algorithm by Wiedensohler
(1988). A ratio of sheath air to aerosol air of 6.8:1 was applied, with alternating up- and down-scans
of 150s for each size spectrum. Annual service of the CPC was performed at TSI to ensure correct
counting efficiency.

#### 5 2.6 Biogenic VOC and SO<sub>2</sub> emissions

6 Emissions of biogenic VOCs are estimated by using the global model MEGAN (Guenther et al., 2006; Zare et al., 2013; Zare et al., 2012). The empirical algorithm in MEGAN simulates the 7 emission rates based on relationships between ecosystem dependent emission factors and key 8 drivers that control emissions. The model takes into account the effects of temperature, radiation, 9 leaf area index, foliage age and soil moisture on the estimation of emissions. In MEGAN, plant 10 species are grouped into six plant functional types (PFTs): broadleaf tree, needle leaf evergreen tree, 11 needle leaf deciduous tree, shrub, crop and grass; and the model dataset provides the geographical 12 distribution of the standard emission factor of the six PFTs with high spatial resolution of 1 km<sup>2</sup>. 13 The required meteorological inputs are provided by the MM5v3.7 model (Grell et al., 1994). The 14 initial and boundary conditions for the MM5 model are derived from the National Center for 15 Environmental Prediction (NCEP) FNL (Final Analyses) data with 6 h temporal and  $1^{\circ} \times 1^{\circ}$  spatial 16 resolution (http://rda.ucar.edu/datasets/ds083.2/). 17

Sulfur dioxide emissions were modeled using the long-range chemistry transport model, the Danish
Eulerian Hemispheric Model (DEHM) (Brandt et al., 2012), based on the European Monitoring and
Evaluation Program (www.EMEP.int) database in Europe with a resolution of 50 km × 50 km.
These European anthropogenic sulfur dioxide emissions are mainly released by fuel combustion
(www.ceip.at), dominated by coal fired power plant and international ship traffic emissions.

#### 23 2.7 Gas and meteorological data

Meteorological data including temperature, RH and global radiation were obtained from a station 24 located on the roof-top of the H.C. Ørsted institute (HCOE, 55°42'N, 12°33'E), which represents 25 the urban background conditions of Copenhagen, Denmark. HCOE is located ~3.5 km north of 26 HCAB. O<sub>3</sub>, NO/NO<sub>x</sub> and CO were measured using photometry (Teledyne API 400, San Diego, CA, 27 USA), chemiluminescense (Teledyne API M200 A, San Diego, CA, USA) and infrared absorbance 28 (Teledyne API 300 E, San Diego, CA, USA), respectively at both field sites HCAB and Risø with 29 30 min resolution whereas measurement of SO<sub>2</sub> using UV fluorescence (Teledyne API M100E, San 30 Diego, CA, USA) with 30 min resolution was only obtained from HCAB. Gas and meteorological 31 32 data were collected within the Danish Air Quality Monitoring Program (Ellermann et al., 2012).

#### 1 2.8 Statistical analysis

Statistical analysis was performed on the data set to analyze the intraday variability between
daytime and nighttime samples as well as the possible inter-site variation between the two sites
HCAB and Risø. As the data was not normally distributed, the U Mann-Whitney test was used
instead of a two-way ANOVA to assess the intraday and inter-site variability (Wilks, 2011).

#### 6 3 Results and Discussion

# 3.1 Detection and characterization of organic acids, organosulfates and nitrooxy organosulfates

9 During the field campaign, a total of 39 species including 12 organic acids, 18 organosulfates (OS) and 9 nitrooxy organosulfates (NOS) were detected as summarized in Table 1-3. The organic acids 10 were identified using their authentic standards, with the exception of hydroxy-pinonic acid where 11 the identification was based on the characteristic mass-to-charge (m/z) fragments of the compound. 12 The organic acids included those of anthropogenic origin (i.e. benzoic acid, adipic acid, pimelic 13 acid and phthalic acid) and biogenic origin, specifically suberic acid and azelaic acid originating 14 from unsaturated fatty acids in addition to terpenylic acid, pinonic acid, pinic acid, hydroxy-pinonic 15 acid, DTAA and MTBCA originating from  $\alpha/\beta$ -pinene (Table 1). Among the anthropogenic organic 16 acids, benzoic acid is a photochemical degradation product of aromatic hydrocarbons, which 17 originate from anthropogenic automobile emissions (Suh et al., 2003), while adipic acid and pimelic 18 acid have been identified as photooxidation products from the ozonolysis of cyclic olefins 19 (Hatakeyama et al., 1985; Grosjean et al., 1978; Koch et al., 2000), and phthalic acid could be the 20 oxidation product of naphthalene and other polycyclic aromatic compounds (Fine et al., 2004). The 21 22 biogenic dicarboxylic acids with relatively long carbon chains including suberic acid  $(C_8)$  and azelaic acid  $(C_9)$  are believed to originate from the oxidation of unsaturated fatty acids (Mochida et 23 24 al., 2003), which are common in marine phytoplankton and leaves of terrestrial higher plants, though some contribution may originate from anthropogenic meat cooking (Rogge et al., 1991) or 25 wood burning (Rogge et al., 1998). The other biogenic organic acids (terpenylic acid, pinonic acid, 26 pinic acid, hydroxy-pinonic acid and DTAA) are oxidation products of  $\alpha/\beta$ -pinene (Claevs et al., 27 28 2009; Ma et al., 2007), with 3-methyl-1,2,3-butanetricarboxylic acid (MBTCA) being the second generation oxidation product of pinonic acid, which indicates aging by OH (Müller et al., 2012; 29 30 Szmigielski et al., 2007). The identified organic acids were all detected at both sites.

The organosulfates were identified using the characteristic fragment of m/z 97 (bisulfate anion 1 2  $HSO_4$ ) and m/z 80 (sulfur trioxide anion  $SO_3$ ). The nitrooxy organosulfates were identified based on a neutral loss of m/z 63 (HNO<sub>3</sub>) in addition to the characteristic fragments of HSO<sub>4</sub> and SO<sub>3</sub>. 3 Many organosulfates and nitrooxy organosulfates species detected in this study (Table 2-3) have 4 5 also been identified in previous smog chamber (Surratt et al., 2007a; Surratt et al., 2008b; Surratt et 6 al., 2010) and field studies (Gómez-González et al., 2008; Kristensen and Glasius, 2011; Gómez-González et al., 2012). Most organosulfates, which were identified based on the literature probably 7 8 originated from isoprene (OS molecular weight (MW) 154, 156, 170, 212, 214, 216 and 334) and monoterpenes including  $\alpha$ - and  $\beta$ -pinene precursors (OS 248, 250, 280 and 298). Two 9 10 organosulfates of MW 200 were identified, of which OS 200 1 could origin from isoprene via methacrylic acid epoxide (MAE) uptake onto acidified sulfate aerosol (Lin et al., 2013), while OS 11 200 2 could derive from 2-methyl-3-buten-2-ol (MBO) (Zhang et al., 2012b; Zhang et al., 2014). A 12 large share of the detected nitrooxy organosulfates could be the oxidation products from limonene 13 and/or monoterpenes (NOS 297, 311, 313, 327 and 329). The structures of several organosulfates 14 (OS 182, OS 208, and OS 210) and one nitrooxy organosulfate (NOS 331) were not identified. 15

#### 16 **3.2 Concentrations**

The ranges of concentration levels of the detected species are shown in Table 4. In general, a higher 17 number of valid measurements (N) were available from HCAB, mostly due to the gap in the set of 18 samples collected from the semi-rural background site Risø. The total concentrations of the detected 19 organic acids averaged per sampling interval (12h) during the sampling period was  $23 \pm 15$  ng m<sup>-3</sup> 20 and  $19 \pm 13$  ng m<sup>-3</sup> at HCAB and Risø, respectively, with the second generation product MBTCA 21 accounting for approximately 21% of total organic acids at HCAB and 20% at Risø. The median 22 value of MBTCA was 2.3 ng m<sup>-3</sup> at HCAB and 2.0 ng m<sup>-3</sup> at Risø, which are in range of the overall 23 median value of 2.7 ng m<sup>-3</sup> reported by Gómez-González et al. (2012) from a study at a Belgian 24 forest site with severe urban pollution impact. Among the anthropogenic organic acids, adipic acid 25 and phthalic acid seemed to dominate at both sites, whereas azelaic acid (4.0 ng m<sup>-3</sup> average) and 26 pinonic acid (2.2 ng m<sup>-3</sup> average) were the dominant first generation biogenic organic acid at HCAB 27 and Risø, respectively. 28

The organosulfates were detected in most samples from both sites (Table 4). The major organosulfates detected in high concentrations at both HCAB and Risø mostly originated from isoprene, such as OS 154, OS 156, OS 200 and OS 212 (Surratt et al., 2007a; Surratt et al., 2008b).

OS 216, which was suggested to derive from the reactive uptake of isoprene epoxydiols (IEPOX) in 1 the presence of acidic sulfate aerosol (Surratt et al., 2010; Lin et al., 2012), tends to be the most 2 abundant organosulfate observed in prior laboratory studies and field studies in the southeastern US 3 (Surratt et al., 2010; Lin et al., 2012; Surratt et al., 2007b; Lin et al., 2013; Surratt et al., 2008a). In 4 this study, a moderate mean concentration of OS 216 of 5 ng m<sup>-3</sup> at HCAB and 3.2 ng m<sup>-3</sup> at Risø 5 was observed. It should also be noted that we used acetonitrile as extraction solvent instead of 6 methanol as used by Surratt et al. (2007b, 2008, 2010) and Lin et al. (2012, 2013), which could 7 8 have affected the extraction yield of OS 216 in the present work. The total organosulfate concentration averaged per sampling interval (12h) during the sampling period was  $60 \pm 45$  ng m<sup>-3</sup> 9 at HCAB, whereas at Risø a slightly lower total concentration value of  $47 \pm 31$  ng m<sup>-3</sup> was obtained. 10 Nitrooxy organosulfates were detected in lower concentrations compared to organosulfates, with a 11 total concentration of  $4 \pm 1$  ng m<sup>-3</sup> at HCAB and  $3 \pm 1$  ng m<sup>-3</sup> at Risø. NOS 297, which was possibly 12 an oxidation product from limonene (Surratt et al., 2008b) was the major nitrooxy organosulfate 13 found at both sites. 14

The concentration levels of the detected compounds were compared with an earlier study performed 15 at a forest site in Denmark by Kristensen and Glasius (2011) during spring 2008, in which fewer 16 species (six organic acids, five organosulfate, five nitrooxy organosulfates) were included. The 17 18 comparison was thus only targeted at similar compounds. In general, the concentrations of most organic acids were higher in the present study, except *cis*-pinic acid and pinonic acid, which were 19 20 more abundant at the forest site. The concentration levels of the organosulfates and nitrooxy organosulfates in this study were many-fold higher compared to the previous study, especially with 21 22 regards to the organosulfate species. This, in addition to the differences among the sites and years 23 could be caused by differences in quantification standards, as Kristensen and Glasius (2011) used 24 camphor sulfonic acid as a surrogate standard, as opposed to the three different standards including 25 D-mannose sulfate, an in-house synthesized organosulfate from  $\beta$ -pinene and octyl sulfate used in this study (as described above). This highlights the analytical challenge of quantifying 26 organosulfates and comparing concentrations among different studies. 27

The total concentration of organic acids and organosulfates are significantly higher in our samples compared to the Arctic, where a total concentration of  $3 \pm 1$  ng m<sup>-3</sup> for organic acids and  $12 \pm 2$  ng m<sup>-3</sup> for organosulfates (quantified using the same organosulfate from  $\beta$ -pinene used in the present study) were measured at Zeppelin Mountain, Svalbard during summer of 2008 (Hansen et al., 2014). However, during late winter and early spring, total concentration levels up to 70 ng m<sup>-3</sup> and over 45 ng m<sup>-3</sup> were detected for all organic acid, organosulfate and nitrooxy organosulfate
compounds respectively at Station Nord, northeast Greenland and at Zeppelin (Hansen et al., 2014).
These values are in range with our concentration values. It must be noted that a total of 11 organic
acids, 12 organosulfates and one nitrooxy organosulfates were detected by Hansen et al. (2014).

#### 5 **3.3 Inter-site, day/night variability and inter-species correlation**

6 The non-parametric U Mann-Whitney test was performed on the data sets to analyze the inter-site variability between HCAB and Risø during the overlapping period (May 19 - June 1 and June 17 -7 22) where data from both sites were available. The 90% confidence interval was used for evaluation 8 of the statistical significance to account for the possible field sampling, filter extraction and 9 quantification uncertainties. The test results showed that 14 parameters out of the 41 tested 10 parameters were statistically different between the two sites at 90% confidence interval (p < 0.1) 11 (Supplementary Table 4). These 14 parameters included six organic acids (benzoic acid, pimelic 12 acid, phtalic acid, suberic acid, pinonic acid, and azelaic acid), the concentrations of total first 13 generation organic acids, five organosulfates (OS 182, OS 200, OS 210, OS 248 and OS 298), one 14 15 nitrooxy organosulfate (NOS 297) and the concentration of total nitrooxy organosulfates. The species with statistically higher concentrations were mostly from HCAB (12 out of 14) with the 16 exception of benzoic acid and pinonic acid, which showed significantly higher concentrations at 17 Risø. While a higher occurrence of benzoic acid at the semi-rural background site Risø was 18 unexpected, it seems that the urban curbside site HCAB was associated with enhanced 19 20 concentrations of both biogenic (suberic acid and azelaic acid) and anthropogenic (pimelic acid and phthalic acid) compounds in addition to the anthropogenic-biogenic coupling species (OS 182, OS 21 200, OS 210, OS 248, OS 298 and NOS 297), which probably indicated some local anthropogenic 22 enhancement impacts on the concentration levels of the compounds. In fact, the total concentrations 23 of organic acids and nitrooxy organosulfates were also found at significantly higher concentrations 24 at HCAB. In contrast, the second generation product of pinonic acid, MBTCA, which has been 25 regarded as an indicator of photochemical aging of pinonic acid by OH (Müller et al., 2012) was not 26 statistically different between the two sites. The fact of a statistically higher concentration level of 27 the pinonic acid precursor at Risø and the indifferent concentration of its product MBTCA at the 28 sites probably suggest that the degree of photochemical aging was not directly governed by the 29 30 concentrations of the precursor at the semi-rural background site Risø.

1 The variability between daytime and nighttime samples at each site during the whole field study 2 was also analyzed using a non-parametric U Mann-Whitney test at 90% confidence interval. The test results showed statistically indistinguishable daytime and nighttime concentrations for most 3 species (Supplementary Table 1-2), with the exception of higher daytime concentrations of phthalic 4 acid (p = 0.0385), OS 210 (p = 0.0622) and NOS 343 (p = 0.0825) at HCAB and higher nighttime 5 6 concentrations of NOS 297 at both sites (p = 0.0011). Phthalic acid, OS 210 and NOS 297 were associated with high numbers of valid measured daytime or nighttime concentrations (N  $\geq$  30) 7 whereas NOS 343 was associated with a lower number of valid N ( $N_{dav} = 10$ ;  $N_{night} = 13$ ). 8

Phthalic acid is widely used in plasticizer substrates and could be primarily emitted to the 9 atmosphere from off-gassing of plasticizers from plastics (Fraser et al., 2003), though previous 10 studies have also suggested that it correlated with SOA in ambient samples (Fraser et al., 2003; Fine 11 et al., 2004; Schauer et al., 2002). A recent study of the photooxidation of naphthalene has also 12 indicated the link between phthalic acid formation and photochemistry, while pointing towards 13 negligible primary emissions of the compound based on laboratory and field measurements 14 (Kleindienst et al., 2012). The findings support our observation of higher daytime concentration of 15 phthalic acid at HCAB, though phthalic acid was not necessarily formed directly at the site. 16 Meanwhile the indifferent daytime and nighttime concentrations of phthalic acid at Risø were 17 probably governed by a generally lower concentration level of the species at the site 18 (Supplementary Table 4). The other compound with significantly higher daytime concentrations at 19 20 HCAB was OS 210, which has recently been reported in aerosols in the Arctic (Hansen et al., 2014). 21

The compound NOS 297 (m/z fragments 296, 233, 177, 163 and 97) showed significantly higher 22 nighttime concentration at both sites, though its concentrations were statistically higher at HCAB 23 compared to Risø (Supplementary Table 4). This species was previously detected in ambient 24 aerosols (Gao et al., 2006; Reemtsma et al., 2006; Kristensen and Glasius, 2011) and has been 25 suggested as an oxidation product of limonene-like monoterpene precursors (Surratt et al., 2008b). 26 In this study, NOS 297 was identified as the major nitrooxy organosulfates (Table 4), accounting 27 for 36% and 34% of the total concentrations of nitrooxy organosulfates detected at HCAB and Risø, 28 respectively. The higher nighttime occurrence of NOS 297 could indicate the impacts of local or 29 30 regional nighttime chemistry on the concentration level of the compound. In fact, the formation of nitrooxy organosulfates has been suggested to involve nighttime nitrate (NO<sub>3</sub>) radical chemistry 31 32 (Iinuma et al., 2007a; Iinuma et al., 2007b; Surratt et al., 2008b). In contrast, NOS 343 showed

significantly higher daytime concentrations at HCAB, though the statistical calculation was 1 2 associated with a lower number of measurements, as mentioned above. Too few measurements of NOS 343 were available at Risø to compare its daytime and nighttime concentrations. Overall, 3 significantly higher nighttime concentrations were observed for total nitrooxy organosulfates at 4 both sites (Supplementary Table 1-2), which was attributed to the predominant contribution from 5 6 the nighttime-dominant species NOS 297. As a majority of the detected species did not show any immediate evidence of local site impacts on their concentration levels according to the U Mann-7 Whitney test results, the relations among the compounds were furthermore investigated using  $R^2$ 8 9 correlation coefficients to unveil any potentially driving mechanisms or factors influencing their 10 formation. The correlation coefficients were calculated for each pair of species at each site and subsequently averaged for both sites. The averaged  $R^2$  correlation coefficients are presented in 11 Figure 2, with each organic acid, organosulfate and nitrooxy organosulfate compound (lines) 12 correlating with other species (x-axis). Relative standard deviations (RSD, %) for the corresponding 13 averaged  $R^2$  correlation coefficient are shown in Supplementary Table 1. Only 36 out of the total 38 14 species were used to calculate the  $R^2$  correlation coefficient values as those with very few 15 measurement values were excluded. The results in Figure 2 were divided into panels of organic 16 acids (top), organosulfates (middle) and nitrooxy organosulfates (bottom). 17

Figure 2 shows a clear correlating group with high  $R^2$  correlation coefficient values (> 0.5) among 18 the species in the group, which is comprised of certain organic acids, organosulfates and nitrooxy 19 20 organosulfates. The correlating group included three organic acids, namely terpenylic acid, MBTCA and DTAA (dark blue), a vast number of 12 organosulfates (red) and four nitrooxy 21 22 organosulfates (green). These organosulfates and nitrooxy organosulfates among others included species closely related in molecular weight, such as the groups of OS (154, 156), OS (212, 214, 23 24 216), OS (252, 254) and NOS (327, 329, 331). The organic acids pinonic and pinic acid (light blue) and OS 248 (purple) seemed to only partly follow the correlation pattern as the degree of correlation 25 was somewhat lower, especially with regards to the correlation between these species and 26 organosulfates. 27

The degree of correlation (calculated as the  $R^2$  values) among the mentioned specific compounds was persistent at both study sites, as evident by the generally lower deviation with small RSD values (%) associated with averaging the R-squared values from both sites as opposed to the uncorrelated species (Supplementary Table 3). The organosulfates correlated consistently well with the organic acids and the other organosulfates within the correlating group, with RSD ranging from

typically a few to over 20% between sites. Correlation coefficients between the nitrooxy 1 2 organosulfates and the other species showed larger discrepancies between the two sites, with RSD ranging from a few percent (NOS 313) to considerably higher percentages (NOS 327, 329 and 331), 3 mostly due to a higher correlation level of these species at Risø compared to HCAB. Concentrations 4 5 of pinonic acid, pinic acid and phthalic acid, which only partly followed the correlation pattern 6 showed larger deviations between the two sites. In contrast, OS 248 which also partly followed the pattern (with slightly lower  $R^2$  coefficient ~0.5) showed a highly consistent correlating tendency at 7 both sites with low RSD values. The consistent correlation pattern among the group of organic 8 9 acids, organosulfates and nitrooxy organosulfates of different precursors detected at both sites 10 probably indicate a common source for the species at both sites.

#### 11 3.4 Regional impacts

Figure 3 shows the temporal variation of the total concentrations of organic acids, organosulfates, 12 nitrooxy organosulfates and PM<sub>1</sub> detected at HCAB and Risø during the field study. It is apparent 13 that all three classes of compounds showed similar temporal patterns, and the two sites do not differ 14 substantially, which were attributed to the large contribution of compounds belonging to the 15 correlating group affecting both sites. The observations strongly indicated that the major sources or 16 chemistry governing the total concentration levels of the compounds must occur at a spatial location 17 affecting both the urban curbside and semi-rural background sites, which represent quite different 18 environments (Figure 1). Elevated total concentrations of organic acids, organosulfates and nitrooxy 19 organosulfates seemed to occur at both sites during the same daytime or nighttime interval, 20 exemplified by the daytime samples on May 22, May 31 and June 18 or the nighttime sample on 21 May 23. Several other concentration maxima occurring on June 6 (daytime) and June 8 (nighttime) 22 were observed at HCAB (data from Risø was unavailable for these days). Organosulfate 200 2 23 24 (from MBO), which was mostly found at levels below the detection limit, was detected in considerably higher concentrations during these episodes of concentration maxima (for example in 25 26 May 31, June 18 (daytime) and May 23 (nighttime) samples).

The correlation between the total concentration level of organosulfates and the concentration of SO<sub>4</sub><sup>2-</sup> ion measured by IC was also examined. A relatively high correlation coefficient value ( $R^2 =$ 0.6) was obtained between the total organosulfates and sulfate at the HCAB site with the episodes of highest concentrations of organosulfates coinciding with peak concentrations of the SO<sub>4</sub><sup>2-</sup> ion (Figure 3), which is in line with previous observations of aged continental aerosols over the southeast Pacific Ocean (Hawkins et al., 2010). The correlation between organosulfates and 1 inorganic sulfate also agrees with previously proposed formation mechanisms for organosulfates 2 including reactive uptake of epoxides on the surface of acidic sulfate aerosols (Iinuma et al., 2009; Minerath and Elrod, 2009; Surratt et al., 2010; Lin et al., 2013; Worton et al., 2013) and sulfate 3 radical-initiated formation in the bulk phase (Nozière et al., 2010). However, while this observation 4 may support related formation processes of the organosulfates and the sulfate ion, the resemblance 5 6 in concentration trend among sulfate, organosulfates, organic acids, and nitrooxy organosulfates may also indicate that organic acids, organosulfates and nitrooxy organosulfates could also partition 7 8 into existing sulfate aerosol. It should be noted that the correlation was only visible at the urban curbside site HCAB, while such correlation was not found at Ris $(R^2 = 0.3)$ . 9

10 Although the formation of nitrooxy organosulfates could involve nighttime NO<sub>3</sub> radical chemistry as discussed above, no direct correlation was found between the total concentration level of 11 nitrooxy organosulfates and concentration of NO<sub>2</sub> at HCAB ( $R^2 = 0.1$ ). Meanwhile, a good 12 correlation ( $R^2 = 0.6$ ) of NO<sub>2</sub> and total concentration of nitrooxy organosulfates was found at Risø 13 (Figure 3). The difference could be attributed to the more complex in-situ contribution of NO<sub>2</sub> and 14 NO<sub>x</sub> at the urban curbside site, in addition to the complex mixture of locally-formed and long-range 15 transported nitrooxy organosulfates. The extent and impact of long-range transport and the impact 16 of NO<sub>2</sub> on the possible local species will be discussed in more detail below. 17

In order to investigate the regional impacts on the concentration levels of the detected compounds, 18 19 5-days air mass back trajectories were calculated using the Hybrid Single Particle Lagrangian 20 Integrated Trajectory (HYSPLIT) model for every 12h sample (Draxier and Hess, 1998). As HCAB and Risø are relatively close with a distance of ~30 km, HYSPLIT calculations were only 21 performed for the HCAB site. The results revealed a dominating westerly origin of air masses 22 during most of the campaign period. However a clear dominance of the south-easterly air masses 23 was found to be associated with the daytime or nighttime episodes of elevated concentrations, 24 including the daytime samples on May 22, 26 and 31 and June 6 and 18 and the nighttime sample 25 on June 8. The travelling path of the south-easterly air masses probably involved the region around 26 the borders between Germany, Poland and Czech Republic (Figure 1, 4). It must be noted that the 27 air masses following the south-easterly direction could have different origins (3-5 days backwards) 28 but all passed the south-easterly region 1-2 days immediately before arrival at the sampling site. An 29 example of the impacts of air masses arriving from the south-easterly direction is shown for three 30 consecutive days on May 30<sup>th</sup> (nighttime), May 31 (daytime with elevated concentrations) and June 31 1 (daytime) where data from both HCAB and Risø were available (Figure 4). To support the 32

interpretation of the HYSPLIT calculations, emissions of isoprene and monoterpenes (daily)
calculated using the MEGAN model and SO<sub>2</sub> (monthly) calculated using the DEHM model, are also
shown in Figure 4 to illustrate the magnitude of biogenic and anthropogenic emissions in the
region.

5 Figure 4 shows a sudden change of air mass origin from westerly (May 30) to south-easterly (May 6 31) back to westerly (June 1), which was coupled with a major episode on May 31 with elevated concentrations of the organic acids, organosulfates and nitrooxy organosulfates (Figure 3). The 7 modeled isoprene and monoterpene daily emissions did not vary noticeably over the corresponding 8 9 days (Figure 4), while emissions from the westerly direction affecting May 30 and June 1 seemed comparable to emissions from the southeasterly direction affecting May 31. This indicated that 10 VOC emissions were probably not the limiting factor in the formation of the investigated 11 compounds. 12

Figure 4 also shows a map of average SO<sub>2</sub> emission over the two months May - June 2011. Though 13 the SO<sub>2</sub> emission map could not be used to evaluate the daily change in anthropogenic emissions to 14 15 further investigate the observed elevated concentrations on May 31, it however indicated strong SO<sub>2</sub> emissions from hotspots in the region that were passed over by the south-easterly air mass shown in 16 the trajectories. This could probably influence the formation of the organosulfates and nitrooxy 17 organosulfates during long-range transport to the sites. The accompanying organic acids, which 18 were also found at high concentrations correlating to the organosulfates and nitrooxy 19 organosulfates, could also be long-range transported to the sampling sites. 20

It is difficult to determine the exact spatial location of the major sources or chemistry governing the 21 total concentrations of the compounds at the two sites, which is partly due to the coarse resolution 22 of the SO<sub>2</sub> emission map. However, it is unlikely that the major source location would be in the 23 local proximity of the two sites due to the highly different local conditions between the urban 24 curbside site HCAB and the semi-rural background site Risø (Figure 1). Such a local source, if 25 present, would be subject to immediate different local background conditions and also varying local 26 27 wind conditions, while in contrast, a highly similar temporal variation pattern of total observed concentrations was found at the sites during the whole campaign period (Figure 3). Furthermore, 28 29 any possible point source located southeasterly from HCAB and Risø, which could affect both 30 sampling sites would unlikely be in Denmark, as both sites are located only 10 - 20 km from Baltic

sea in the southeasterly direction (Figure 1). The source region therefore is possibly located at the
 broader regional scale across the Baltic sea extending to the southerly neighboring countries.

3 Particle mass concentrations were calculated based on integration of the calculated particle mass size distributions (from DMPS data), which were further based on measured particle number size 4 distributions assuming particle sphericity, using a SOA particle density of 1.7 g cm<sup>-3</sup> (Stock et al., 5 2011) (Figure 3). This density value is close to the bulk densities of ammonium sulfate and nitrate, 6 which are expected to be the major components of aged accumulation mode particles. The detected 7 and quantified organosulfates and nitrooxy organosulfates contributed to 0.1 - 1.0 % (approximately 8 0.5 % on average) of PM<sub>1</sub> mass at HCAB and 0.4 - 1.5 % (approximately 0.8 % on average) of PM<sub>1</sub> 9 mass at Risø. The values are considerably lower than previous estimates, including 30% (Surratt et 10 al., 2008b) or 5-10% contribution of organosulfates to total organic mass (Tolocka and Turpin, 11 2012). PM<sub>1</sub> concentrations at HCAB were in the lower range during the days under the influence of 12 southeasterly air mass (May 22, May 26, May 31, June 18 (daytime), with the exception of June 6 13 (daytime). The period of high PM<sub>1</sub> concentrations at HCAB (June 6 - 8) seemed to link to other 14 local conditions, including episodes of elevated concentrations of SO<sub>2</sub> and NO<sub>x</sub> (Figure 5 and 15 Supplementary Figure 1). The local impacts are further discussed in the section below. 16

#### 17 3.5 Local impacts

The possible local anthropogenic impacts were investigated by examining the species of 18 significantly higher concentrations at HCAB, which did not belong to the correlating group, 19 including OS 182 (m/z fragments 181, 97), OS 210 (m/z fragments 209, 153, 137, 97 and 79) and 20 21 NOS 297 (m/z fragments 233, 177, 163 and 97). As displayed in Figure 5, the patterns of changes in OS 182 and OS 210 concentrations in May were considerably similar, with a few episodes of 22 23 high daytime concentrations especially during the period May 20 - 25 with a lower concentration on May 22 (daytime). Such observation stands in contrast to the trend of the total organosulfate 24 25 concentrations during the same period, where a generally lower concentration level was observed except from elevated daytime concentrations on May 22 (Figure 3). Assuming that the trend of 26 27 variations the total detected species in Figure 3 was mostly governed by long-range transported compounds belonging to the highly correlated group, the trend of OS 182 and OS 210, which did 28 29 not belong to the correlated group, could probably reflect some local impact on the concentration levels of these species. However, while OS 182 and OS 210 seemed to correlate well in May, 30 concentrations of the species deviated towards the end of the field study (Figure 5), which could be 31 partly attributed to additional input from long-range transport to the site. For example, elevated 32

daytime concentration of OS 182 was observed on June 16, which is one of the characteristic days
of southeasterly arriving air masses associated with higher total concentrations of the detected
species (Figure 3).

Local impacts were studied to explain the significantly higher concentrations of OS 182, OS 210 4 5 and NOS 297 at the urban curbside site HCAB. Temporal variation of gases (NO, NO<sub>x</sub> and  $SO_2$ ), aerosol acidity level (calculated from charge balance using  $[SO_4^{2-}]$ ,  $[NO_3^{-}]$ , and  $[NH_4^{+}]$ 6 following Zhang et al. (2012b)) at HCAB, (Figure 5), and meteorological conditions including RH 7 8 (Figure 5), temperature (Temp) and global radiation (GR) (Supplementary Figure 1) at the urban 9 background site HCOE, which was located in close proximity to HCAB were analyzed. As can be seen from Figure 5, the period May 20 - 25 was accompanied by a generally higher level of aerosol 10 acidity which decreased on May 22 (daytime). Such variation of the aerosol acidity level perfectly 11 matched the trend of the varying concentrations of OS 210 and OS 182 during the same period 12 (Figure 5). In fact, the aerosol acidity level has been suggested as an essential parameter in the 13 formation of organosulfates (Zhang et al., 2012b). At the same time, the concentration level of the 14 organosulfates might rather be affected by an inter-play between the aerosol acidity level and other 15 factors, such as the global radiation and temperature. For example the May 20 and 21 (daytime) 16 samples showed high concentrations of OS 182 and OS 210, which was comparable to the 17 18 concentrations on May 24 or 25 (daytime) samples (Figure 5), whereas the aerosol acidity level on May 20 or 21 was considerably lower on May 20 and 21 (daytime) compared to the other days 19 (Figure 6). May 20 - 21 were however accompanied with a higher level of global radiation, 20 especially on May 21 (579 W m<sup>-2</sup>) compared to other latter days May 24-25 (daytime) 21 (Supplementary Figure 1). While a higher level of global radiation could probably enhance the 22 oxidation of SO<sub>2</sub> by the OH radical and thereby resulted in higher concentrations of  $SO_4^{2-}$  as 23 observed with the May 20 - 21 (daytime) samples,  $NH_4^+$  was also detected at higher concentrations 24 with these samples, resulting in an overall lower aerosol acidity calculated based on ion charge 25 balance on the days. In addition, the higher temperature observed on May 20 - 21 compared to May 26 24 - 25 may also have contributed to enhance the formation of supposed local organosulfates in the 27 former samples, by means of enhancing biogenic VOC emissions, or photochemistry in general. In 28 fact, a considerably higher level of O<sub>3</sub> oxidant was also found during the former days (May 20 - 21) 29 30 compared to the latter days (May 24 - 25), which would positively enhance the formation of SOA.

Higher daytime concentrations of OS 210 were also observed on June 7 and 8, whereas the aerosol acidity reached its highest level of 46 nmol m<sup>-3</sup> on June 7 (nighttime) with interestingly almost

1 neutral conditions on the adjacent June 7 (daytime) and June 8 (daytime) (Figure 5). At the same 2 time, during these two daytime intervals, higher average concentrations of SO<sub>2</sub> were also observed. At the same time, concentrations of NO and NO<sub>x</sub> in higher resolution (30 min) were also 3 considerably elevated (Supplementary Figure 1). Meanwhile, the O<sub>3</sub> level was in the lowest 4 5 concentration range presumably due to the lower level of sunlight in addition to titration by NO<sub>x</sub> 6 (figure not shown). A close examination revealed that the NO and NO<sub>x</sub> were highest during the early morning and late afternoon hours which are attributed to traffic emissions, though during the 7 8 day course both NO and NO<sub>x</sub> remained at a higher level compared to the other days during the campaign period. This observation could be partially attributed to lower wind speed during daytime 9 on June 7 (2.0 m s<sup>-1</sup>) and on June 8 (2.9 m s<sup>-1</sup>), which was considerably lower than the average wind 10 speed of 4.1 m s<sup>-1</sup> for the whole campaign period. It must also be noted that this period was coupled 11 with the highest RH range during the campaign period (Supplementary Figure 2), which resulted 12 from the corresponding very low global radiation during the days. The observations indicated an 13 inter-play between the anthropogenic gases, especially NO, NO<sub>x</sub> and probably aqueous aerosol 14 chemistry, leading to the occurrence of higher concentrations of OS 210 in these samples. A 15 previous study by Dommen et al. (2006) on the effect of RH under high NO<sub>x</sub> conditions indicated 16 that more volatile SOA were observed at higher RH rather than at lower RH. Additionally, SOA 17 formation including carboxylic acids and organosulfur compounds in cloud and fog droplets have 18 been hypothesized as plausible by Blando and Turpin (2000) while a recent smog chamber study by 19 Zhang et al. (2012a) suggested that elevated RH values could mediate the formation of 20 21 organosulfates, thereby emphasizing the role of wet sulfate aerosols in forming organosulfates in ambient aerosols. At the same time, though aerosol acidity has been found to enhance isoprene 22 23 SOA formation under low initial NO conditions (Surratt et al., 2010; Surratt et al., 2007c; Jaoui et al., 2010), seed aerosol acidity showed a negligible effect on SOA formation from isoprene under 24 high initial NO conditions (Surratt et al., 2006), which implied the less important role of seed 25 aerosol acidity on the formation of OS 210 on June 7 and 8 (daytime) where both NO and RH were 26 27 high.

Further investigation of the meteorological conditions during the period where data from both sites were available (May 19 - June 1 and June 17 - June 22, see Supplementary Figure 2) revealed a possible coupling between higher temperature and global radiation (508 Wm<sup>-2</sup>) with the occurrence of concentration maxima observed on days with the highest temperature May 31 (daytime) and June (daytime) (Figure 3). However, information was lacking on the spatial scale of high temperature and global radiation episodes to fully justify if concentration maxima were enhanced by high
 temperature and global radiation at source or at the sampling sites.

3 The third species used to study the local impacts was NOS 297, which showed significantly higher nighttime concentration as discussed above, though some possible long-range transport contribution 4 5 to its concentration at HCAB could not be completely eliminated, as shown by its elevated 6 concentration also on peak concentration days May 31 and June 16. To examine the local impacts on NOS 297, only nighttime concentrations of NO<sub>2</sub> and nighttime RH are shown (Figure 5), 7 whereas full trends of concentrations of NO<sub>x</sub> and RH are displayed in Supplementary Figure 1. 8 However with the exception of these days, elevated concentrations of NOS 297 were only observed 9 at nighttime on May 19, 20, 22, 26 and June 6, 8 and 10. In fact, nighttime concentration maxima of 10 NO<sub>2</sub> (34, 34 and 36 µg m<sup>-3</sup>) were also observed on May 19, 20 and June 10, respectively, compared 11 to an average nighttime NO<sub>2</sub> concentration of 25  $\mu$ g m<sup>-3</sup>, which would contribute to formation of the 12 nitrate (NO<sub>3</sub>) radical via the reaction between O<sub>3</sub> and NO<sub>2</sub>. It should also be noted that recent work 13 by Rollins et al. (2013) has also shown a nighttime correlation between nitrooxy organosulfates, 14 including NOS 297 and total particulate organic nitrate (RONO<sub>2</sub>), which results from oxidation of 15 organic compounds in high NO<sub>x</sub> environment. The corresponding nighttime RH was also in the 16 lower concentration range (~64% RH) compared to the average nighttime RH of 70% for the whole 17 18 campaign period, which probably reduced the available condensation sink for particle precursors and facilitated the formation of NOS 297 by nighttime NO3 radical chemistry. However, such 19 20 speculation should also be interpreted with caution as the RH difference was relatively small. In contrast, during the other nights (May 22, 26 and June 6, 8), where high concentrations of NOS 297 21 were also found, the RH was considerably higher ( $\sim$ 79%) whereas NO<sub>2</sub> concentrations were low. 22 Investigation of air mass back trajectories using HYSPLIT also revealed that the May 22, 26 and 23 24 June 6 nighttime were affected by air masses arriving from the south-south-easterly direction whereas on June 8 air masses arrived from easterly direction. Such observations suggest that the 25 detected concentrations of NOS 297 could be governed by both the local formation and the degree 26 of long-range transport of the compound to the site. As a result, no direct correlation between 27 nighttime concentrations of NOS 297 and NO<sub>2</sub> ( $R^2 = 0.3$  for both daytime and nighttime correlation 28 and  $R^2 = 0.2$  for only nighttime correlation) was found. On the other hand, it should be noted that 29 during this period of the year, the duration of the night is relatively short in Denmark, therefore 30 calculation of nighttime concentrations would be coupled with higher uncertainty Nevertheless, it 31 seems probable that occurrence of NOS 297 was favored by nighttime chemistry, regardless of 32

whether the species were formed in-situ or long-range transported to the site. It can be seen that the
local impacts could affect the concentration levels of the specific compounds as discussed above.
However it must be noted that the magnitude of the local impacts on the detected species seemed
rather modest in general.

5 Local new particle formation (NPF) events were also examined using the DMPS data following Dal 6 Maso et al. (2005) to investigate whether the local formation of organosulfates could be associated with particle nucleation at the sites. A limited number of NPF events were found occurring on the 7 same days at the semi-rural background site Risø, the urban background site HCOE and the urban 8 curbside site HCAB, though HCAB is heavily influenced by traffic emissions resulting in generally 9 short nucleation bursts rather than allowing a classic "banana" formation and growth. The NPFs 10 however seemed to occur on the days with low concentrations of most organic acid, organosulfate 11 and nitrooxy organosulfate species, including the possible local compounds OS 182 and OS 210, 12 which agree with previous indications that NPF may occur more frequently in cleaner environments 13 (Lyubovtseva et al., 2005; Kulmala et al., 2004b). As the occurrence of higher concentrations of the 14 detected organic acids, organosulfates and nitrooxy organosulfates did not correspond to NPF 15 16 events, which were coupled with elevated occurrence of smaller-sized particles; it is likely that the organic acids, organosulfates and nitrooxy organosulfates belong to the accumulation mode. At the 17 same time, any correlation to NPF should be investigated with size-segregated samples. 18

#### 19 4 Conclusion

A field study was conducted from May 19 to June 22, 2011 at two sites in Denmark, including the 20 urban curbside site HCAB and the semi-rural background site Risø to investigate the anthropogenic 21 impacts on the formation of biogenic SOA via analysis of organosulfates and related oxidation 22 products using HPLC-qTOF-MS. A substantial range of anthropogenic and biogenic organic acids, 23 24 organosulfates and nitrooxy organosulfates were detected, identified and quantified. Isoprene oxidation products comprised a large fraction of the detected organosulfates whereas a majority of 25 the detected nitrooxy organosulfates may originate from limonene and/or monoterpenes. The U 26 27 Mann-Whitney statistical test showed significantly higher concentrations of specific biogenic and anthropogenic species at HCAB, whereas an intraday statistical test only revealed significantly 28 higher daytime concentrations of phthalic acid and OS 210 and significantly higher nighttime 29 concentrations of NOS 297. Many detected compounds including a group of organic acids, 30 31 organosulfates and nitrooxy organosulfates of various precursors were found belonging to a

common highly correlated group, consistently affecting their concentration levels at both sites, suggesting a common source region or similarities in formation processes. The analysis of HYSPLIT back trajectories together with isoprene and monoterpene emissions calculated by the MEGAN model and SO<sub>2</sub> emission based on the EMEP monitoring database indicated that southeasterly air mass arriving from across the Baltic sea could relate to the corresponding episodes of the elevated total concentrations at the sites, implying the importance of the regional impacts on the occurrence of the compounds.

The local impacts were also investigated for several selected species which showed significantly 8 9 higher concentrations at HCAB compared to Risø and did not belong to the correlating group, including OS 182, OS 210 and NOS 297. The results showed that the local factors including aerosol 10 acidity, RH, NO and probably some wet aerosol chemistry could possibly lead to a higher 11 occurrence of the organosulfates, whereas the NO<sub>3</sub> radical chemistry could be important for the 12 formation of the major nitrooxy organosulfate NOS 297, in addition to temperature and global 13 radiation, which could also enhance the concentrations of the long-range transported species at the 14 sites. As only very few compounds were deemed as affected by the local conditions, the analysis of 15 16 the local impacts was restricted to a modest number of compounds. In general, the regional impacts seemed to considerably exceed the local impacts. It was estimated that organosulfates and nitrooxy 17 organosulfates contributed to approximately 0.5% and 0.8% of PM<sub>1</sub> mass, respectively at HCAB 18 and Risø. 19

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- 37 Tracers for Secondary Organic Aerosol (SOA) Formation from 2-Methyl-3-Buten-2-ol (MBO) in the Atmosphere Environ Sei Technol. 46, 0427-0446, Dei 10,1021/Eg2016487, 2012e
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- 4 centre of Copenhagen close nearby a busy street. Risø is located approximately 30 km west of
- 5 HCAB.

Figure 2. Correlations between organic acids, organosulfates and nitrooxy organosulfates with each other detected species in each sample throughout the sampling period (excluding those with too few measurements), expressed as correlation coefficient ( $R^2$ ). The results show  $R^2$  coefficient averaged from the two sites. High  $R^2$  values (> 0.5) were found among a specific group of species (yellow shade), suggesting common sources. A few other species including pinonic acid, pinic acid, phthalic acid, and OS 248 also partly followed the correlation pattern.

Figure 3. Temporal trends of total concentrations of organic acids, organosulfates, nitrooxy organosulfates and  $PM_1$  at the urban curbside site HCAB (solid line) and the semi-rural background site Risø (broken line). The figure shows a strong correlation between the two sites and among the three categories of compounds examined. Sulfate (HCAB) and NO<sub>2</sub> (Risø) are also shown in the organosulfate and nitrooxy organosulfate panels, respectively.

Figure 4. A sudden change in air mass back trajectories on May 31 to southeasterly direction was coupled with elevated concentrations of organosulfates, nitrooxy organosulfates, and organic acids at both HCAB and Risø. Modeled daily isoprene and monoterpene emissions did not vary significantly over the three days. The SO<sub>2</sub> emission map (May-June average) showed a higher abundance of emission hotspots in southeasterly direction (observed on May 31) compared to the westerly direction (observed on May 30 and June 1). HCAB is marked at the bottom tip of the black triangle.

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#### 8 Supplementary Information

- 9 Supplementary Table 1. Statistical results of U Mann-Whitney test performed on daytime and
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- 17 measurements (N < 5) were excluded from the test.
- Supplementary Table 3. Relative standard deviations (RSD, %), which were associated with 18 averaging R-square correlation coefficients (between pairs of compounds) at the two sites, were 19 20 calculated. The species belonging or partly belonging to a common correlating group are marked in blue shade with the corresponding RSD in red (those only partly belonging to the group 21 underlined). RSD values seemed low for the correlating organic acids and OS while larger for the 22 NOS, mostly due to higher correlation coefficients at Risø. The partly-correlating species pinic, 23 pinonic acid and phthalic acid showed relatively larger RSD compared to the other correlating 24 25 compounds, while another partly-correlating species (OS 248) showed a consistent partly correlating degree at both sites as evident by small RSD values. 26
- 27 Supplementary Table 4. Statistical results of U Mann-Whitney test performed on the organic acids,
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- 32 (30 min resolution) at HCAB during the campaign period. The x-axis shows a temporal scale from
- 33 00 24h for each day, with the white and grey shades indicating the intervals for day and night
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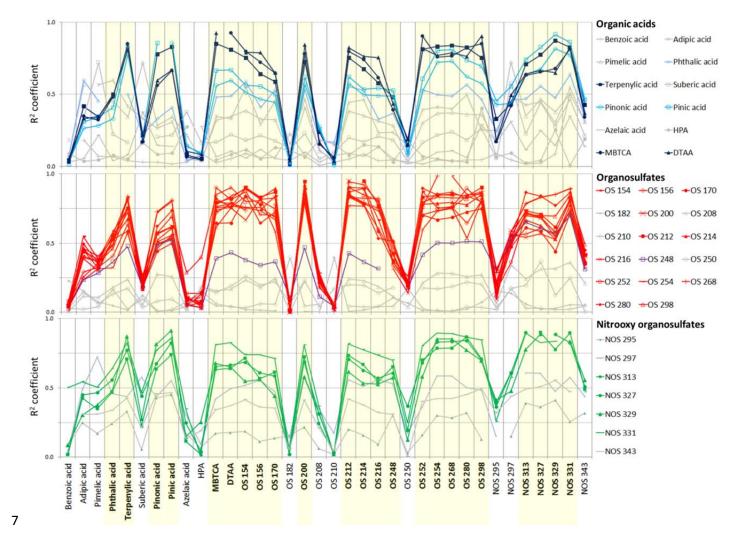
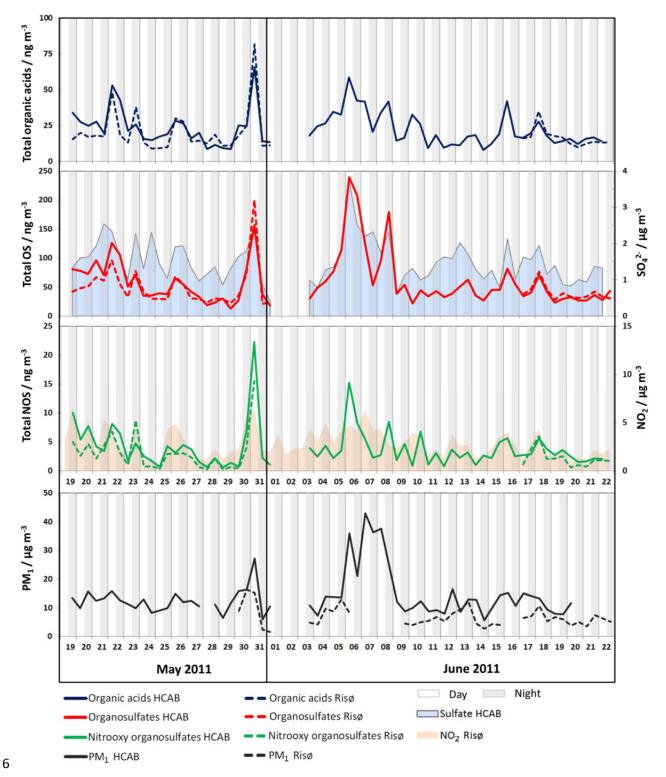
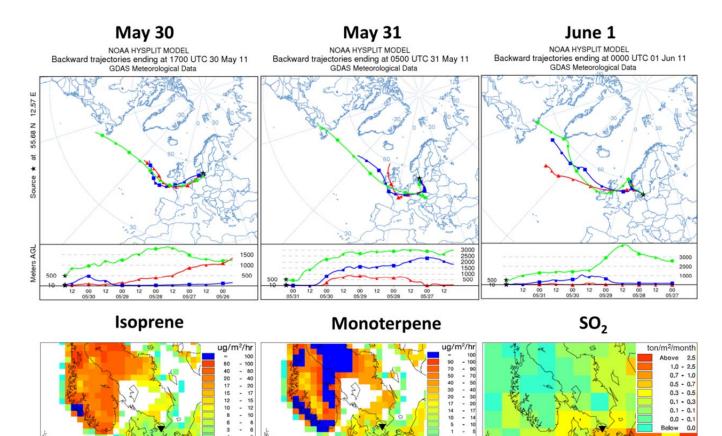


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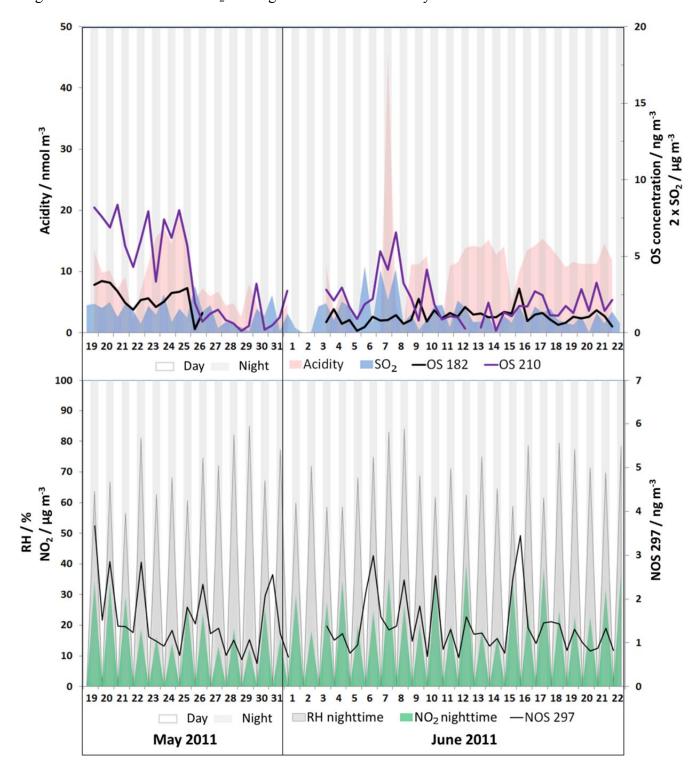
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## 1 Table

## Table 1. Detected organic acids during the campaign. Related references on suggested precursors are cited in the text.

		Retention time (min)	Proposed molecular formula	Proposed structure	Suggested precursors Aromatic hydrocarbons			
Benzoic acid	121.022	27.5	C <sub>7</sub> H <sub>6</sub> O <sub>2</sub>					
Adipic acid	145.058	14.9	$C_{6}H_{10}O_{4}$	HOH	Cyclic olefins			
Pimelic acid	159.061	20.3	$C_{7}H_{12}O_{4}$	НО ОН	Cyclic olefins			
Phthalic acid	165.123	19.0	$C_8H_6O_4$	ОНОНОН	Aromatic hydrocarbons			
Terpenylic acid	171.058	18.3	$C_8H_{12}O_4$	остон	α-pinene			
Suberic acid	173.074	25.4	$C_8H_{14}O_4$	но	Fatty acids			
Pinonic acid	183.100	27.5	$C_{10}H_{16}O_3$	но	α/β-pinene			
Pinic acid	HO       HO         thalic acid       165.123       19.0 $C_8H_6O_4$ rpenylic       171.058       18.3 $C_8H_{12}O_4$ d $\mathcal{I}$ $\mathcal{I}$ beric acid       173.074       25.4 $C_8H_{14}O_4$ ho $\mathcal{I}$ $\mathcal{I}$ boric acid       183.100       27.5 $C_{10}H_{16}O_3$ ho $\mathcal{I}$ $\mathcal{I}$ $\mathcal{I}$ ic acid       185.077       23.0; 24.0 $C_{9}H_{14}O_4$ elaic acid       187.091       28.1 $C_{9}H_{16}O_4$ $\mathcal{I}$		Х У Уон	α/β-pinene				
Azelaic acid	187.091	28.1	$C_9H_{16}O_4$	но он	Fatty acids			
Hydroxy- pinonic acid	v- 199.092 29.0 C <sub>10</sub> H <sub>16</sub> O <sub>4</sub> CH <sub>2</sub> C		CH <sub>2</sub> OH COOH	α/β-pinene				
3-Methyl- 1,2,3- butanetricarb oxylic acid MBTCA	203.050	16.2	C <sub>8</sub> H <sub>12</sub> O <sub>6</sub>	но от он	α/β-pinene			

(DTAA)	α-pinene	но он	$C_{10}H_{16}O_{6}$	24.3	231.220	Di-terpenylic acid acetate (DTAA)
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Organosulfate	ate[M-H]RetentionProposedProposed structureion (m/z)time (min)molecular formula				Suggested precursors	
OS 154 <sup>1,2</sup>	152.981	3.6; 6.7	$C_3H_6O_5S$	O O O O O H	hydroxyacetone (isoprene)	
OS 156 <sup>3</sup>	154.957	2.8	$C_2H_4O_6S$	но он	glyoxal	
OS 170 <sup>4</sup>	169.047	3.4	$C_3H_6O_6S$	HO SOH	methylglyoxal	
OS 182	181.063	34.8	$C_5H_{10}O_5S$	unknown	unknown	
OS 200_1 <sup>1</sup>	198.998	3.3	$C_4H_8O_7S$	но он	2-methylglyceric acid (isoprene)*	
OS 200_2 <sup>5,6</sup>	199.001	3.4	$C_5H_{12}O_6S$	он он он он	2-methyl-3-buten-2-ol (MBO)	
OS 208	207.101	28.4; 29.5	$C_6H_8O_6S$	unknown	unknown	
OS 210	209.086	28.0; 30.5	C <sub>7</sub> H <sub>14</sub> O <sub>5</sub> S	unknown	unknown	
OS $212^{2}$	210.983	4.8	$C_5H_8O_7S$	unknown	isoprene	
				unknown	isoprene <sup>2,5</sup>	
OS 214 <sup>2,7,8</sup>	213.001	3.5	C <sub>5</sub> H <sub>10</sub> O <sub>7</sub> S	озо <sub>3</sub> н ноон	fatty acid <sup>4</sup>	
OS 216 <sup>1,2</sup>	215.018	2.9	$C_5H_{12}O_7S$	HO O OH OH OH	2-methyltetrols (isoprene)**	
OS 248 <sup>2</sup>	247.061	32.2	C <sub>10</sub> H <sub>16</sub> O <sub>5</sub> S	unknown	α-pinene	
OS 250 <sup>2</sup>	249.074	31.6	$C_{10}H_{18}O_5S$	O-S-OH OHO	β-pinene***	
OS 252 <sup>2</sup>	251.017	15.3	$C_9H_{16}O_6S$	OH OSO3H	limonene	

1 Table 2. Detected organosulfates (OS) during the campaign.

OS 268 <sup>2</sup>	267.225	19.7	$C_9H_{16}O_7S$		limonene
OS 280 <sup>2</sup>	279.046	22.5; 25.2	$C_{10}H_{16}O_7S$	о он он он	α/β-pinene
OS 298 <sup>2</sup>	297.079	25.5	$C_{10}H_{18}O_8S$	OH O	α-pinene
OS 334 <sup>2</sup>	333.071	38.0	C <sub>10</sub> H <sub>22</sub> O <sub>10</sub> S	HO OH OH OH OH OH	Isoprene

Included references were <sup>1</sup> Surratt et al. (2007a); <sup>2</sup> Surratt et al. (2008b); <sup>3</sup> Galloway et al. (2009); <sup>4</sup> Olson et al. (2011); ;
<sup>5</sup> Zhang et al. (2012c) <sup>6</sup> Zhang et al (2014) <sup>7</sup> Gómez-González et al., (2008); <sup>8</sup> Kristensen and Glasius, (2011). \* OS 200
has been recently shown to specifically form from methacrylic acid epoxide (MAE) uptake onto acidified sulfate
aerosol (Lin et al., 2013). \*\* OS 216 has been shown to specifically derive from the reactive uptake of isoprene
epoxydiols (IEPOX) in the presence of acidic sulfate aerosol (Surratt et al., 2010; Lin et al., 2012). \*\*\* OS 250 can be
found in synthesis of both α and β-pinene derived organosulfates; however with different retention times. The OS 250

7 detected in this study was assigned to  $\beta$ -pinene precursor based on the corresponding retention time.

Nitrooxy[M-H]RetentionProposedorganosulfateion (m/z)time (min)molecular formula		Proposed structure	Suggested precursor			
NOS 295 <sup>2</sup>	294.063	42.1; 44.7	C <sub>10</sub> H <sub>17</sub> NO <sub>7</sub> S	HO-S-O O More isomers	α/β-pinene	
NOS 297 <sup>2</sup>	296.063	32.8	C <sub>9</sub> H <sub>15</sub> NO <sub>8</sub> S	OSO <sub>3</sub> H ONO <sub>2</sub>	limonene	
NOS 311 <sup>2</sup>	1 <sup>2</sup> 310.088 38.0 C <sub>10</sub> H		C <sub>10</sub> H <sub>17</sub> NO <sub>8</sub> S	ONO2 OSO3H	limonene, α-pinene	
NOS 313 <sup>2</sup>	312.209	28.0	C <sub>9</sub> H <sub>15</sub> NO <sub>9</sub> S	unknown	limonene	
NOS 327 <sup>2</sup> 326.279 26.8 C <sub>10</sub> H <sub>17</sub> NO <sub>9</sub> S		0 02N0 02N0	limonene, β-pinene, terpinolene			
NOS 329 <sup>2</sup>	OS $329^2$ 327.994 33.5 $C_{10}H_{19}NO_9S$		OH ONO2 OH OSO3H	limonene		
NOS 331 <sup>2</sup>	330.002	15.6	C <sub>9</sub> H <sub>17</sub> NO <sub>10</sub> S	unknown	unknown	
NOS 340 <sup>2</sup>	339.070	25.7	$C_{10}H_{16}N_2O_9S$	unknown	α-pinene	
NOS 343 <sup>2</sup>	342.043	38.8	C <sub>10</sub> H <sub>17</sub> NO <sub>10</sub> S	unknown	α-pinene; α-terpinen	

Table 3. Detected nitrooxy organosulfates (NOS) during the campaign. 

References include <sup>1</sup> Surratt et al., (2007a); <sup>2</sup> Surratt et al., (2008b); <sup>3</sup> Surratt et al., (2010); <sup>4</sup> Gómez-González et al.,<br/>(2008); <sup>5</sup> Kristensen and Glasius, (2011). 

Compound	HCAB Risø										
*	Mean	Stdev.	Median	Max	Ν	Mean	Stdev.	Median	Max	N	
Benzoic acid	0.4	0.6	0.1	2.0	14	1.4	1.9	0.4	7.2	26	
Adipic acid	3.3	1.4	2.8	8.4	64	2.9	1.1	2.8	5.9	38	
Pimelic acid	0.6	0.3	0.5	2.1	64	0.4	0.2	0.4	1.2	38	
Phthalic acid	5.6	2.3	5.1	13.0	64	4.3	1.8	4.4	10.4	38	
Terpenylic acid	1.6	1.6	1.1	8.9	64	1.4	1.7	0.9	8.5	38	
Suberic acid	1.1	0.6	1.0	3.7	64	0.9	0.5	0.8	2.9	38	
Pinonic acid	1.3	1.5	0.9	10.8	64	2.2	2.9	1.3	16.4	38	
cis-Pinic acid	0.5	0.5	0.4	3.9	64	0.6	0.9	0.4	5.6	37	
Azelaic acid	4.0	3.3	3.0	19.6	58	1.6	0.8	1.6	4.3	35	
Hydroxy-pinonic acid	0.1	0.1	0.1	0.2	12	0.1	0.0	0.1	0.1	4	
MBTCA	4.7	5.7	2.3	23.8	64	3.9	5.1	2.0	24.4	38	
DTAA	0.4	0.5	0.2	2.0	64	0.2	0.3	0.1	1.4	34	
Total organic acids	22.8	14.9	18.7	76.7	64	19.3	13.3	15.7	69.8	38	
OS 154	6.6	6.0	4.5	25.7	63	5.3	4.6	4.3	24.6	38	
OS 156	5.3	3.8	4.3	23.1	64	4.3	3.1	3.5	19.8	38	
OS 170	3.5	2.3	2.7	10.3	63	3.1	2.0	2.5	11.5	37	
OS 182	1.4	0.8	1.2	3.4	55	0.6	0.4	0.6	1.6	37	
OS 200_1	7.1	6.7	4.4	33.1	64	4.5	3.9	3.0	22.2	37	
OS 200_2	6.4	6.3	4.1	18.4	5	3.1	2.1	2.6	6.7	5	
OS 208	5.0	3.3	5.7	14.2	63	4.1	0.8	4.1	5.6	38	
OS 210	2.8	2.4	2.0	8.4	63	1.8	1.0	1.9	4.3	38	
OS 212	7.4	6.0	5.0	32.2	64	6.2	4.7	4.6	27.8	38	
OS 214	5.1	4.5	3.6	22.3	64	4.4	3.1	3.8	17.9	38	
OS 216	5.0	6.8	2.8	36.9	60	3.2	3.3	2.4	20.8	34	
OS 248	0.8	0.4	0.8	2.3	57	0.6	0.3	0.6	1.7	37	
OS 250	2.3	1.1	2.1	6.8	64	2.4	2.1	1.9	12.4	38	
OS 252	1.2	1.2	0.6	5.2	53	0.9	0.9	0.6	4.1	33	
OS 254	1.3	1.5	0.8	7.9	63	1.1	1.2	0.8	7.4	38	
OS 268	2.5	2.3	1.6	13.1	64	2.2	2.2	1.6	13.9	38	
OS 280	2.6	2.5	1.8	11.5	60	2.2	2.0	1.6	10.3	38	
OS 298	1.3	1.1	0.8	4.1	28	0.8	1.0	0.5	3.9	16	
OS 334	0.6	0.3	0.6	1.0	8	0.8	-	0.8	0.8	1	
Total organosulfates	59.8	44.5	43.8	130.6	64	46.9	31.1	36.3	144.0	38	
NOS 295	0.5	0.6	0.3	4.2	46	0.4	0.3	0.4	1.2	15	
NOS 297	1.4	0.7	1.3	3.7	64	0.9	0.6	0.7	3.4	38	
NOS 311	0.9	0.8	0.4	2.4	7	0.5	0.5	0.4	1.8	8	
NOS 313	0.7	0.7	0.5	3.4	56	0.6	0.6	0.5	3.0	30	
NOS 327	0.9	0.7	0.7	3.6	48	0.6	0.6	0.5	2.8	29	
NOS 329	0.8	0.9	0.4	3.6	19	0.5	0.6	0.3	2.6	17	
NOS 331	0.0	0.2	0.1	1.2	41	0.3	0.3	0.2	1.5	26	
NOS 340	0.2	0.1	0.2	0.2	3	-	-	-	-	-	
NOS 343	0.2	0.4	0.2	1.4	24	0.4	0.4	0.4	1.0	4	
Total nitrooxy organosulfates	3.9	1.3	2.8	3.7	64	2.8	1.2	2.1	1.8	36	

Table 4. Concentration range (ng m<sup>-3</sup>) of the detected species at HCAB and Risø, reported as mean, 1 standard deviation (stdev.), median, max and the number of detected samples (N) at each site.