



Deposition mode ice
nucleation
reexamined below
200 K

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Deposition mode ice nucleation reexamined at temperatures below 200 K

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Abstract

The environmental chamber of a molecular beam apparatus is used to study deposition nucleation of ice on graphite, alcohols and acetic and nitric acids at temperatures between 155 and 200 K. The critical supersaturations necessary to spontaneously nucleate water ice on six different substrate materials are observed to occur at higher supersaturations than are theoretically predicted. This contradictory result motivates more careful examination of the experimental conditions and the underlying basis of the current theories. An analysis based on classical nucleation theory supports the view that at these temperatures nucleation is primarily controlled by the rarification of the vapor and the strength of water's interaction with the substrate surface. The technique enables a careful probing of the underlying processes of ice nucleation and the substrate materials of study. The relevance of the findings to tropospheric temperatures is discussed.

1 Introduction

In the atmosphere water vapor is transformed into cloud particles and precipitation vis-à-vis processes of nucleation, whereby vapour becomes liquid droplets or solid ice particles. The nucleation and evolution of droplets and ice crystals is a keystone process for both the global radiative balance and the hydrologic cycle. In most of the atmospheric column the temperature and saturation conditions dictate that solid ice is the thermodynamically favorable phase. However, in the real world spontaneous ice formation is difficult to achieve and fundamentally quantifying atmospheric ice nucleation and growth remains an elusive scientific goal. The processes of ice nucleation are invariably influenced by the presence of other atmospheric constituents like particles and/or gases, meaning that the nucleation that is observed to be important in the atmosphere is largely heterogeneous. Heterogeneous *ice* nucleation may occur when (1) super-cooled liquid water contacts a foreign body and subsequently freezes, or (2) when H₂O

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vapour is directly deposited onto foreign material (Pruppacher and Klett, 1997). The later mechanism, also referred to as *deposition freezing*, is an important mechanism of ice formation at low atmospheric temperatures. Like other freezing mechanisms (e.g. immersion mode freezing, Marcolli et al., 2007; Möhler et al., 2005), measurements of deposition freezing show wide ranges of nucleation onset (Knopf and Koop, 2006; Möhler et al., 2006) and assessing the relative importance of different freezing modes remains an active area of inquiry (Hoose and Möhler, 2012), most critically at high temperatures and near water saturation where multiple freezing mechanisms may be active (Dymarska et al., 2006). That said, at low temperatures deposition freezing experiments are relatively straight forward and therefore provide an excellent template for detailed comparisons of observation and theory.

The fundamental character of nucleation as a statistical thermo-kinetic process makes it difficult to achieve an unambiguous mathematically predictive nucleation theory. Modelling nucleation is made difficult by the challenge of capturing the scope of nucleating systems; from molecular clustering to macroscopically observed solidification, and the observational barriers to detecting nucleation onset, efficiency, and particle growth rates (Fletcher, 1958; Hale and Plummer, 1974; Kashchiev, 2006). However, due to the fundamental importance of nucleation to phase behavior, much attention is paid to extracting and appraising the relevant kinetic and thermodynamic parameters (Chen et al., 2008; Liu, 1999; Niedermeier et al., 2011). For atmospheric ice nucleation the impediments to understanding are particularly formidable. Environmental temperature and saturation conditions, and the time and turbulent scales of the atmosphere are difficult to simulate in a laboratory. Field campaigns that yield spatially and temporally resolved free troposphere data are also difficult and resource intensive. Obstacles aside the current escalation of interest in atmospheric ice nucleation is and needs to be driven towards formulating true microphysical understanding that can be used in describing clouds and other atmospheric processes.

Low temperature deposition nucleation experiments are relatively straight forward to execute in a laboratory setting using idealized substrate surfaces and the results can

be applied to assessing the validity of ice nucleation parameterizations that seek to illuminate atmospheric processes. Here we describe a series of deposition ice nucleation experiments at temperatures ≤ 200 K, where initially surprising results led us to reexamine the classical formulation of nucleation theory. This examination has motivated us to probe the microphysics of nucleation at these temperatures and forces us to investigate what limits small ice embryo formation at low temperature. This makes plain the need for renewed investigation into the microphysics of nucleation and is important for understanding the formation of small surface water clusters and the behavior of a range of ice nucleating materials.

2 Experimental

2.1 Apparatus

These experiments utilize an Environmental Molecular Beam (EMB) chamber contained within a multi-chamber ultra-high vacuum (UHV) system that has been described previously (Kong et al., 2011). Within the environmental chamber highly oriented pyrolytic graphite (HOPG, produced by Advanced Ceramics Corp., grade ZYB) is utilized as a bare substrate and a surface upon which other molecular materials can be condensed. The 12 mm \times 12 mm surface is temperature controlled and gas inlets are positioned to facilitate the introduction of condensing species like water that form layers on the HOPG. Helium in the molecular beam provides a sensitive measure of nucleation on, and coverage of, the graphite surface (Kong et al., 2012). Macroscopic thickness, ice nucleation and growth, are also monitored from the reflected intensity of a 0.86 mW, 670 nm diode laser. A quadrupole mass spectrometer (QMS) is positioned within the primary UHV chamber in line with an opening in the environmental chamber. Thus the QMS provides a sensitive measure of the environmental chamber pressure and the molecular species leaving the cold HOPG surface.

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When investigating critical supersaturations for ice nucleation the experimental procedure begins by setting the surface temperature T_s (Kong et al., 2012). The graphite is then used as a bare surface or as a cold surface for condensing substrate materials of interest. Here we present results of ice nucleation on bare graphite and adlayers of methanol (MeOH), butanol (BuOH), hexanol (HxOH), acetic acid (AcOH) and nitric acid (HNO_3). For each condensed material, prior to inletting water vapor to induce nucleation, a stable substrate monolayer is adsorbed onto the graphite. On graphite monolayers of these materials are significantly more stable than equivalent macroscopic layers, and using the EMB can be characterized by adsorption isotherms determined from helium scattering (Kong et al., 2012; Papagiannakopoulos et al., 2014). The QMS is used to measure the water vapor intensity $\mathcal{J}_{\text{bk}}, \mathcal{J}_{\text{nuc}}, \mathcal{J}_{\text{eq}}$, where the subscripts refer to the states of the system (bk – background, nuc – nucleation, or eq – equilibrium). In this case the measured intensities directly represent the physical parameter of interest, pressure $P \propto \mathcal{J}_{\text{bk}}, \mathcal{J}_{\text{nuc}}, \mathcal{J}_{\text{eq}}$. After a background intensity \mathcal{J}_{bk} is identified, water vapor is systematically added to the system until ice nucleation is observed. Typically experimental time scales to reach the critical saturation for nucleation S_i^{nuc} are on the order of 10 to 15 min. Vapor is initially inlet to an estimated level of $0.9S_i^{\text{nuc}}$, and after the system has re-equilibrated at that level small incremental increases are made in H_2O vapor pressure at one minute intervals. The H_2O intensity required for ice nucleation \mathcal{J}_{nuc} is recorded and subsequently the ice layer is grown to a macroscopic thickness of $\approx 1 \mu\text{m}$. After achieving a macroscopic ice layer the vapor inlet is adjusted to maintain a constant, homogeneous ice layer in equilibrium with the vapor and the steady-state vapor pressure over the ice \mathcal{J}_{eq} is measured. Thus at each temperature all relevant vapor pressures are explicitly measured and the critical supersaturation required for ice nucleation can be calculated directly as described in Sect. 2.2.

2.2 Results

The direct linear relationship between the pressure within the environmental chamber and the intensity measured with the QMS, allows us to calculate the chamber super-

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5 saturations at the onset of nucleation. The range of temperature that we are able to explore using this method ($155 \text{ K} \leq T_s \leq 200 \text{ K}$) extends from the temperature at which the chamber background pressure exceeds S_i^{nuc} to the highest T_s at which an equilibrium ice surface is sensitive to the inlet vapor flux. At temperatures greater than 200 K the vapor inlet valve must be fully open to maintain equilibrium ice surfaces, which limits experimental control and precludes accurate pressure measurements.

The critical supersaturation for nucleation is defined as the ratio of the nucleation vapor pressure to the equilibrium vapor pressure above the condensed phase. Expressed in terms of the measured QMS intensities, after the chamber's background intensity is carefully removed yields,

$$S_i^{\text{nuc}} = \frac{J_{\text{nuc}}}{J_{\text{eq}}}. \quad (1)$$

15 This straightforward analysis was observed to be robust for all of the measured systems with the exception of the extreme temperatures measured for the pure ice on graphite nucleation. For that system at $T_s \geq 195 \text{ K}$, J_{eq} was observed to be affected by the high absolute vapor pressures. The high absolute pressures were observed to stress the UHV pumping capacity and therefore shift the background intensity. The shift was calculated by comparing the $T_s \geq 195 \text{ K}$ to expected results based on a Clausius-Clapyron fitting of the lower temperature data and used as a small correction to the data. The corrected results compare well with the independent results of Pratte et al. (2006). Likewise for graphite at $T_s \leq 160 \text{ K}$, the signal to noise ratio of the measured J_{eq} was nearly unity, forcing the results to be corrected based on higher temperature data. For all other substrate materials tuning the QMS control parameters increased the relative signal, thus allowing the experimentally measured data to be used directly over the entire range of temperature. The results of all measurements are presented in Fig. 2.

25 In Fig. 2 the critical supersaturations measured for vapor deposition ice nucleation on MeOH, BuOH, HxOH, AcOH, and HNO₃ adlayers, and bare graphite are

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plotted together with common theoretical curves. For all experimental measurements a strong increase in S_i^{nucl} is observed with decreasing temperature. The theoretical curves for homogeneous ice nucleation from solution (dashed, Koop et al., 2000, with $J = 5 \times 10^8 \text{ cm}^{-3} \text{ s}^{-1}$), water saturation (solid-black, Murphy and Koop, 2005) and gas phase homogeneous ice nucleation (upper-solid-green, Murray and Jensen, 2010) are also drawn. The two former curves, which are both based on low temperature extrapolations of liquid water properties have been widely used as references for discussing atmospheric ice nucleating processes (e.g., Hoose and Möhler, 2012), and generally it is assumed due to the energetic benefit of surfaces heterogeneous nucleation should occur at or below liquid saturation. However, comparing these curves with the measured values clearly has interesting ramifications. The supersaturations required for ice nucleation are high relative to the theoretical predictions based on liquid parameterizations, and the effect appears to increase exponentially with decreasing temperature – in a manner that appears more consistent with homogeneous nucleation from the gas phase. However, even when the effect is most pronounced, such as on graphite where S_i^{nucl} has increased by one order of magnitude at 155 K, the S_i^{nucl} are significantly enhanced relative to the prediction for homogeneous nucleation. Furthermore, the effect on nucleation is distinguishable between the individual surfaces, making it clear that the specific heterogeneity of each system does play a role.

Although, in absolute terms the onset of nucleation is observed to vary depending on substrate material, Fig. 3 where the data are replotted using an Arrhenius-type formulation illustrates that the trend with temperature is well represented by homogeneous nucleation theory. In Fig. 3 the data from Fig. 2 are re-rendered by taking the logarithm of S_i^{nucl} and plotting those values vs. inverse temperature. For comparison the data for each substrate are shifted by a single constant that is chosen to minimize the cumulative deviation from the theoretical curve for homogeneous nucleation (green line). The shift constants are listed in Table 1.

The observation of the homogeneous-like temperature trends in the presence of seemingly high supersaturation values leads us to revisit both classical nucleation theory and our basic understanding of how ice embryos are formed on surfaces.

3 Analysis

5 Unto themselves the measured S_i^{nucl} values are not completely anomalous. In the past such values have been measured in other systems (Trainer et al., 2009; Fortin et al., 2003; Shilling et al., 2006) and hypothesized as a potential explanation for the absence of ice particles in planetary atmospheres (Iraci et al., 2010). However, past attempts at explaining the few observations have had limited broad applicability, as they often
10 rely on physical parameterizations specific to the investigated systems. The observations we have described augment existing observations and demonstrate the need for a generalized description of deposition nucleation at low temperatures. It is rather unlikely that all observations can be explained by changes in contact angle, site-specific nucleation, or other empirical parameterizations that lack unambiguous chemical or
15 physical explanations at low temperature. For example, in the temperature range of this study graphite is well known to be hydrophobic (Andersson et al., 2007; Kong et al., 2012), while adsorption of methanol provides sites for efficient hydrogen-bonding, resulting in a highly hydrophilic surface (Thomson et al., 2011; Kong et al., 2012). Thus the question remains, how can surfaces with widely dissimilar hygroscopic behavior all
20 require high supersaturations for vapor deposition ice nucleation. To offer insight we first reexamine the basic edifice of nucleation theory.

3.1 Predicting freezing behavior at low temperature

The nucleation of ice has classically been treated using bulk thermodynamic theories of homogeneous and heterogeneous nucleation. Generally this “classical nucleation
25 theory” (CNT) is flexible enough that it can be used to robustly capture the freezing

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behavior of many materials including ice. From an analytical point of view the flexibility of CNT stems directly from its mathematical construct,

$$J = A \exp\left(\frac{-\Delta G^*}{k_b T}\right), \quad (2)$$

where the product of a pre-exponential term A and an exponential term that expresses the barrier to nucleation in terms of the free energy of formation of a critical nucleus ΔG^* , Boltzmann's constant k_b and the temperature T , is a nucleation rate J per second per unit volume or area depending upon whether the nucleation occurs in a volume or on a surface. Thus in its simplest form nucleation is described by a thermodynamic model with an activation barrier. However, as is typical the devil is in the details and much of CNT focuses on the details of the terms ΔG^* and A .

The standard derivation of ΔG^* begins with the surface/volume balance of the free energy of formation of an ice embryo.

$$\Delta G = -V_i \Delta g_v + A_i \gamma, \quad (3)$$

where the volume V_i and area A_i of a spherical ice embryo with radius r_i are $4/3\pi r_i^3$ and $4\pi r_i^2$ respectively, $\Delta g_v = k_b T \ln S/v$ is the bulk energy change per unit volume expressed in terms of H_2O supersaturation S , and γ is the surface energy of the embryo. Thus the radius of a critical ice embryo r_i^* is defined to be r_i where $\partial \Delta G / \partial r_i = 0$, or

$$r_i^* = \frac{2\gamma v}{k_b T \ln S}. \quad (4)$$

By substitution this yields a free energy of formation for a homogeneous spherical critical cluster,

$$\Delta G^* = \frac{16\pi\gamma^3 v^2}{3(k_b T \ln S)^2}. \quad (5)$$

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In the case of heterogeneous nucleation, when nucleation is facilitated by a foreign material surface, this term can be altered due to the shift in geometry and the change in surface energy caused by the presence of the foreign material. Elegant solutions that depend on the geometry and wettability of the surface–embryo contact have been derived for these considerations and can most simply be expressed as a single factor f that multiplies ΔG^* and can vary from $f = 0$ to $f = 1$ (Pruppacher and Klett, 1997). The result is that the free energy barrier to nucleation is easily scaled and a number of studies have used this to implement contact angle parameterizations that provide suitable statistical fits to experimental data (Chen et al., 2008; Niedermeier et al., 2011). Alternatively, if f is used as a fitting parameter and considered to represent a classical physical contact angle in the Young-Dupree sense the angle and therefore the “wettability” of the surface can be determined (Chen et al., 2008). Although surface energy and water contact angle do depend on material, temperature and pressure etc., this method of interpreting measurement data is problematic on several fronts. First, the variation in f that is required to fit experimental data is quite often significant even within small temperature ranges (Trainer et al., 2009). Thus to achieve a successful CNT fit, a functional dependence of f is often assumed (e.g., Chen et al., 2008; Trainer et al., 2009). These functional dependencies are difficult to verify because at low temperatures and for many materials, the temperature dependence of surface energy and/or contact angle is not independently constrained. Furthermore, the contact angle model is derived based on the assumption of an isotropic spherical ice/water cap. Ice germs are neither isotropic nor necessarily spherical and at low temperatures critical sizes on molecular scales may limit the applicability of bulk thermodynamic theories in general. Notwithstanding the fact that contact angle is a useful parameterization, particularly as temperatures approach bulk coexistence, it is plain that it is an empirical parameter and not strictly a physical quantity.

The precise constitution of the pre-exponential term A is also debated, but it is clear that this term represents the molecular fluxes to and from growing embryos. In early work this kinetic coefficient was considered to be an empirically derived constant (e.g.,

$10^{25} \text{ s}^{-1} \text{ cm}^{-2}$, Fletcher, 1958) to which many processes were shown to be rather insensitive over a few orders of magnitude. Again, at low temperatures such a treatment of A is problematic and an extra effort must be made to model the active physical processes. In the case of heterogeneous deposition freezing a more descriptive pre-factor $A = \alpha\beta Z N_{\text{ads}}$ is often used, where β represents the impingement rate of H_2O molecules onto the critical embryo, Z is the Zeldovich factor, N_{ads} is the number of adsorbate molecules on the surface, and α represents additional kinetic factors. As given by Winkler et al. (2008), Chen et al. (2008) and others,

$$\beta = \frac{Sp_s}{\sqrt{2\pi mk_b T}}, \quad Z = \frac{v}{2\pi r_i^2} \sqrt{\frac{\gamma}{k_b T}}, \quad N_{\text{ads}} = \frac{Sp_s}{v\sqrt{2\pi mk_b T}} \exp\left(\frac{E_{\text{ads}}}{k_b T}\right), \quad (6)$$

where p_s is the saturation vapor pressure, m and v are the molecular mass and volume, v is the vibrational frequency of an adsorbed molecule (10^{13} s^{-1} , Pruppacher and Klett, 1997), and E_{ads} is the adsorption energy of molecules on the surface. Thus the terms combine to describe the availability of seed molecules on the surface vis-à-vis the balance between monomer adsorption and desorption (N_{ads}), the deposition rate onto growing clusters (β), and a correction factor for the fluctuations that result in some critical clusters returning to sub-critical sizes (Z).

A further oversight at low temperature that comes about due to the classical assumption of bulk properties, is the existence of a finite difference in chemical potential between monomers in the condensed and gaseous phases. By inspection of Eqs. (2)–(5) one can see that the monomer formation energy will be greater than zero, which violates the basic edifice of phase equilibrium. Girshick and Chiu (1990) addressed this specific issue and to insure self-consistency calculated an additional pre-factor,

$$\frac{\exp(\Theta)}{S}, \quad \text{where } \Theta = \frac{\varphi_s \gamma}{k_b T}, \quad (7)$$

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with graphite are principally constrained by molecular simulations and thus a finite range of values exists that is based on the detail of simulated interaction potentials. However, in general there is close agreement between multiple studies (Marković et al., 1999; Rubeš et al., 2009; Lakhlifi and Killingbeck, 2010; Voloshina et al., 2011). In the case of MeOH, the data reported in Thomson et al. (2011) are affected by complex kinetics in the surface layer that likely involve at least two states with different binding energies. Recalculated binding energies based on a typical Arrhenius prefactor 10^{13} s^{-1} yield the range of MeOH binding energies presented in Table 2. In a recent study in which BuOH surfaces were investigated in detail, Papagiannakopoulos et al. (2013) used molecular dynamics (MD) simulations to show that water molecules that are able to locate and interact with hydroxyl groups have binding energies between 0.32 and 0.35 eV, but otherwise the interaction is somewhat weaker, 0.10 to 0.12 eV. Complementary EMB-BuOH experiments were consistent with BuOH having a close packed surface structure that precludes the chance for water binding (Papagiannakopoulos et al., 2013). Thus we suspect the lower values are the best approximation for the monolayers of butanol used in this study, which are expected to have a well-ordered close packed structure. Similar MD simulations do not exist for HxOH and with the present experimental system it is impossible to distinguish the expected H_2O binding energy with a HxOH monolayer from that of BuOH. However, as the alkyl tail elongates it may become more difficult for the layer to relax into its most stable structure and thus it is harder to assess the likelihood that the surface structures will be close packed. The result may be that more OH groups remain available for water binding. The EMB has also been used to study AcOH layers and the experimental results indicate that within a monolayer AcOH-AcOH bonds are strong and it is unlikely that adsorbed water molecules will induce bond breaking. On such surfaces water has a short residence time that effectively sets an upper bound for the H_2O binding energy ≤ 0.25 eV. This agrees with published simulations of water interactions with AcOH hydrophobic surfaces that yield a range of 0.13 to 0.20 eV (Allouche and Bahr, 2006).

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The expected water interaction with HNO_3 under our experimental conditions is perhaps the most difficult value to constrain and published binding energy values for water on HNO_3 surfaces are limited. Kofaski et al. (2011) found a value of ≈ 0.35 eV for a H_2O - HNO_3 dimer structure with efficient hydrogen bonding. That is close to the 0.31–0.32 eV determined from ab initio calculations that estimated the formation of stable cyclic monohydrate HNO_3 - H_2O complexes (Tao et al., 1996; Staikova and Donaldson, 2001). Our own EMB observations that were complementary to the supersaturation experiments suggest more strongly bound surface states with multiple hydrogen bonds. An Arrhenius plot of data from D_2O scattering from HNO_3 monolayers (Fig. 4) yields the binding energy 0.51 ± 0.06 eV listed in Table 2 and used in our calculations.

It is worthwhile noting that even with the limited number of substances studied here, we do observe a correlation between the binding energy (Table 2) and the absolute shifts (Table 1) required to collapse the data in Fig. 3. Of course this is also seen directly, as the binding energy increases S_i^{nuc} tends to decrease, which follows the physical intuition that those surfaces that more strongly bond with water, also more easily nucleate ice.

3.3 Contact parameter f

Unlike the binding energy, the geometric factor f cannot a priori be deduced and is therefore often treated as a fitting parameter. Following such a convention f can be considered to be single valued for individual substrates or some function of temperature $f \equiv f(T)$. In either case the observational data allow us to calculate individual values for each substrate at each temperature based on the measured S_i^{nuc} values.

The calculated f values serve as empirical fitting parameters for the observations, whose temperature dependence can be fitted using a least squares regression. Figures 5 and 6 show CNT solutions that demonstrate the temperature sensitivity of the f parameter, and the necessity of using a functional form at these temperatures. In Fig. 5 the experimental data is replotted with CNT solutions derived based on a single f values determined for bare graphite. The theoretical curves demonstrate that

with the correct choice of f individual data points can be modeled, but the trend with temperature deviates significantly from observations. At low temperatures the theory under-predicts S_i^{nuc} meaning nucleation is limited relative to the prediction, while at high temperatures nucleation is enhanced.

To capture the correct temperature trend it is necessary to assume a functional form $f(T)$. Examples of theoretical curves generated in this manner, using $f(T) = \exp(c_1 T) + c_2$ where c_1 and c_2 are constants produced by the least squares regressions done for each substrate material, are shown in Fig. 6. For each substrate the value of f decreases with increasing temperature, suggesting surfaces that are becoming more “wetter” in the traditional interpretation. The form of $f(T)$ is chosen to avoid the unphysical complication of a zero crossing that results from a simple linear fitting. However, within the experimental temperature range the details of the functional form have little effect on the global fitting. Thus, as expected Fig. 6 demonstrates that by combining the binding energy and a functional form of the geometric factor, CNT can be used to capture the observed behavior.

It is possible to proceed one step further and calculate a physical contact angle θ from $f(m) = (2 + m)(1 - m)^2/4$ where $m = \cos(\theta)$, which is an expression derived from purely geometrical considerations of a spherical nucleation cap (cf., Turnbull and Vonnegut, 1952; Pruppacher and Klett, 1997). However, at these temperatures such a wholly geometric picture of the freezing nucleus is, in fact, unphysical and thus the resulting values for contact angle θ are not necessarily intuitive. For example, based on fitting the S_i^{nuc} for graphite, where $E_{\text{ads}} = 0.13 \text{ eV}$, the f values range from 0.17 to 0.009 or $\theta \approx 61\text{--}27^\circ$. In contrast the contact angle of water on graphite is expected to be between 80° and 90° (Adamson and Gast, 1997). Although, the hydrophilicity of graphite may certainly change with temperature and concentration of water molecules, there exists no straightforward method for asymptotically matching contact angles derived from nucleation experiments with contact angles measured using bulk materials at temperatures near or above the liquidous. Furthermore, even if contact angle had a strong temperature dependence within a certain temperature range, there is no

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known physical reason why it would have a similar dependence irrespective of the underlying surface material. Hence the origin of the strong temperature dependence is unlikely to be a result of a changing contact parameter.

Ultimately we are left with a dilemma. Due to the nature of the system, the heterogeneous CNT treatment of the results presented here yields excellent theoretical fits to the data (Fig. 6) that remain to be explained physically (Sect. 4). Conversely, the results also appear to be well modeled empirically by a simple shift of the homogeneous nucleation curve (Fig. 3), which is difficult to explain for different reasons.

4 Discussion and atmospheric implications

The experimental results that we have presented demonstrate that at temperatures below 200 K high water vapor supersaturations are required to nucleate ice on substrate surfaces. Although the results can be well explained using the versatility of CNT, the theoretical results do not lend themselves to a ready physical explanation. From a theoretical standpoint, much of the unique behavior that seems to straddle the developed homo- and hetero-geneous CNT can be reasonably explained by close examination of the underpinnings of the heterogeneous theory presented in Sect. 3.1 and by recognizing that there is an underlying strong pressure dependence on temperature. The Ideal Gas law and the Clausius-Clapeyron relation guarantee that as H₂O vapor becomes rarified supersaturations must increase to compensate for the decreasing pressure. In the present heterogeneous CNT model the term for monomer adsorption N_{ads} contains a strong temperature dependence that acts to mitigate the effect of increasing S . However, it does so in a manner that is overemphasized, because in its present form N_{ads} accounts only for pure monomolecular adsorption and desorption. In fact it is more likely that monomers adsorb and diffuse on the surface and thus interact with other monomers and molecular clusters, thereby depleting the effective monomer concentration and reducing the nucleation rate (Kroon and Ford, 2011). In lieu of a strict bookkeeping of the complicated molecular kinetics the heterogeneous CNT introduces

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f , which compensates for the strong thermal dependence in the binding energy term but does not comply with physical intuition. Thus a more complete description of deposition nucleation from rarified gas is likely to be homologous with understanding thin-film nucleation, which explicitly reduces the dimensionality of the problem in the correct manner and eliminates geometric factors that are open to misinterpretation (Venables et al., 1984).

There are a number of other factors that also help to explain the difficulty of strict CNT treatments at these temperatures and encourage us to reexamine proposed alternatives for describing spontaneous nucleation. In a review of homogeneous nucleation Oxtoby (1992) gives a good summary of the shortcomings of, and myriad assumptions in CNT. In general the approximations of CNT, when considering nucleation as in this case forming solid from a dilute vapor where there exists a strong singularity in the density parameter at the phase boundary, are well controlled. However, at these experimental temperatures the critical cluster sizes that can be calculated vis-à-vis CNT are $\mathcal{O}(10)$ to $\mathcal{O}(100)$ molecules, and as a result the expected system behavior may not be consistent with what is predicted by bulk thermodynamics. The underlying question in nucleation, which remains open, is the applicability of macroscopic concepts to molecular-scale clusters (Oxtoby, 1992, 1999).

There does exist a small compendium of previous experimental deposition nucleation work which extends below 200 K. In the past this work has been motivated by studies of exoplanetary atmospheres like that of Mars, where temperatures are quite low but mineral dust and other aerosol particles may be abundant (e.g., Phebus et al., 2011; Cziczo et al., 2013). Trainer et al. (2009) and later Iraci et al. (2010) investigated deposition freezing on surrogate mineral dust materials. In both cases they discovered high requisite supersaturations, particularly for the limited measurements made below 170 K. Their analyses focused on using the empirically determined contact angle to parameterize the freezing and in the case of Trainer et al. (2009), they too observe a strong low-temperature dependence. However, once again the physical interpretation of the temperature dependence can only be described as the, “average surface

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ample, density functional theory provides a non-classical approach whereby molecular scale effects can be preserved while employing mean density fields. Although density functional methods are fundamentally exact, real calculations require approximations that are subject to choices like the free-energy functional. Even so such non-classical approaches can be used to calculate ΔG^* values that are augmented significantly from those of CNT. A higher nucleation barrier height ΔG^* results in significantly larger required undercooling or higher S_i^{nucl} (cf., Fig. 4 in Oxtoby, 1992). Functionals that can be applied to nucleation are continually improving and should be tested for the systems we have probed experimentally.

In the context of other phase change phenomena involving H₂O the strength and range of intermolecular interactions can strongly influence the character of the stable vs. metastable equilibrium. Premelting which refers to the metastable disordering of ice layers as the melting temperature is approached is an example of just such a phenomena (Dash et al., 2006). In premelting theory liquid melt evolves from crystalline surfaces as temperature and impurity change in ways that depend sensitively on the magnitude and fall-off of the relevant intermolecular interactions (Wettlaufer, 1999; Thomson et al., 2010; Bartels-Rausch et al., 2014). Predicted sensitivities have been experimentally probed and observations confirm the basic theoretical assertions (Elbaum et al., 1993; Thomson et al., 2013). Although, the processes of nucleation and melting are not reversible, in both cases capturing the details requires an intimate understanding of the governing interactions.

5 Conclusions

In this study we have presented results from deposition mode ice nucleation on substrate surfaces at temperatures below 200 K. Previously, there have been limited observations of ice nucleation at these temperatures, and thus these results expand the breadth of materials studied under such conditions. Our observations are consistent with previous studies and both demonstrate the necessity of very high supersaturations

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to initiate ice nucleation at ≤ 200 K. Observed supersaturations far exceed predicted critical saturations that are based on macroscopic theories and extrapolations from solid and liquid properties at higher temperatures. Thus the observations are indicative of an incomplete understanding of the key processes controlling systems under such conditions. Developing a better micro-physical understanding will have implications for how we understand nucleation in the atmosphere and more generally as a fundamental process important in many molecular systems.

An analysis of the experimental results based in CNT demonstrates that differences in the water adsorption energy on the different investigated substrate surfaces can partially but not completely explain the observed critical saturation ratios. Ice is observed to nucleate more easily on surfaces with larger H_2O affinities. To describe the sharp increase of S_i^{nuc} a temperature dependent contact parameter may be employed. For individual experiments the contact parameter can be determined explicitly, but a clear physical explanation for the parameter values is lacking. Due to its system specificity, this method of characterizing nucleation also suffers from limited predictive power.

Casting aside the issues of CNT, these new observations offer insight into the basic process of deposition nucleation. With the hope of robustly capturing such behavior, they also spur us to consider more detailed descriptions of the intermolecular interactions that lead from molecular clustering to nucleation. In addition to the development of more accurate theoretical models future work should focus on systematically determining the importance of material properties and detailed examinations of behavior at higher temperatures, where eventually macroscopic models will capture the system behavior.

Acknowledgements. This work is supported by the Swedish Research Council, and the Nordic Top-Level Research Initiative CRAICC. PP thanks the Wenner-Gren Foundation for providing funding for an extended stay at the University of Gothenburg. This work has benefited greatly from discussions with N. Marković.

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Table 1. Shift constants that minimize the cumulative deviation of substrate observations to the homogeneous reference state. Alternatively for each temperature and substrate, this can be expressed mathematically as, $\min[\sum(\ln S_i^{\text{nuc}} - \ln S_{i,\text{hom}}^{\text{nuc}})]$, where $\ln S_{i,\text{hom}}^{\text{nuc}}$ represents homogeneous nucleation from the gas phase (green line, Fig. 3).

Surface	Shift Constant
Graphite	2.58
Methanol	2.99
Butanol	3.01
Acetic Acid	2.71
Hexanol	2.78
Nitric Acid	3.19

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Table 2. Surface binding energies.

Surface	Binding Energy (eV)	Reference
Graphite	0.10–0.16	Marković et al. (1999); Rubeš et al. (2009); Lakhlifi and Killingbeck (2010); Voloshina et al. (2011)
Methanol	0.37–0.38	Thomson et al. (2011)
Butanol	0.10–0.12	Papagiannakopoulos et al. (2013)
Acetic Acid	0.13–0.20	Allouche and Bahr (2006); Papagiannakopoulos et al. (2014)
Hexanol	0.10–0.12 or 0.32–0.35	Papagiannakopoulos et al. (2013)
Nitric Acid	0.51 ± 0.06	Kong et al. (2014) and present study

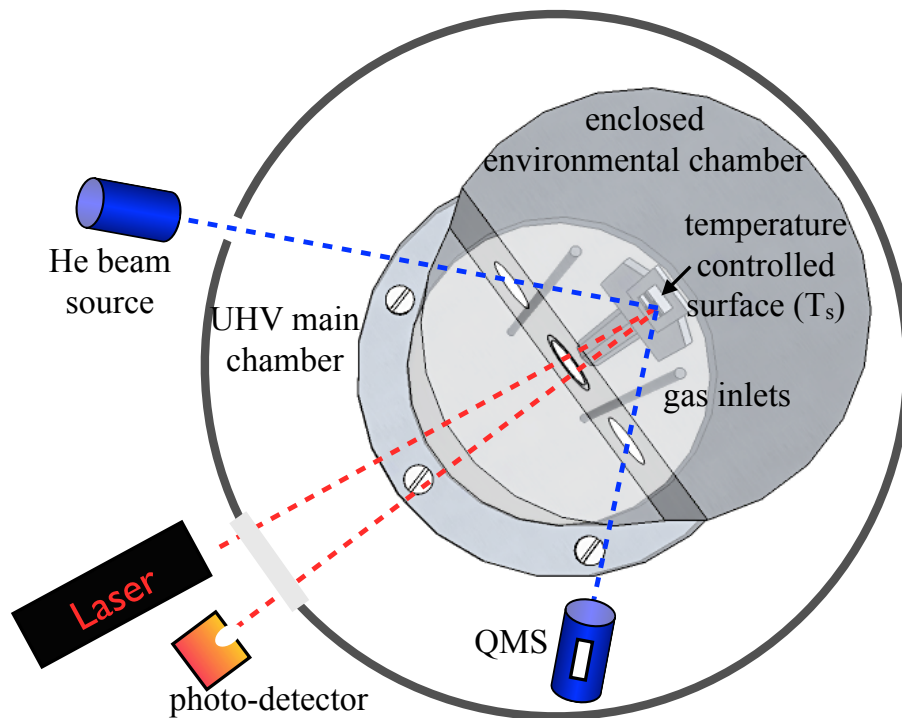


Figure 1. Schematic drawing of the principle components of the environmental chamber and nucleation and growth monitoring system.

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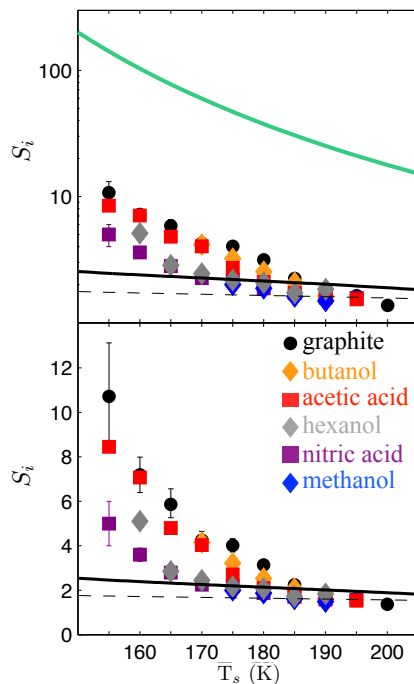


Figure 2. Ice supersaturations (S_i) plotted vs. surface temperature (T_s) with points indicating the critical supersaturations (S_i^{nucl}) required for vapor deposition ice nucleation on various materials. Error limits that are representative of all data are shown for the bare graphite and HNO_3 covered graphite. At $T_s > 165$ K the uncertainty is subsumed by the symbols. For comparison the theoretical curves for water saturation (solid-black, Murphy and Koop, 2005) and homogeneous ice nucleation (dashed, Koop et al., 2000) are also plotted. For the latter a nucleation rate $J = 5 \times 10^8 \text{ cm}^{-3} \text{ s}^{-1}$ that matches previous treatments (Hoose and Möhler, 2012) is modeled. The logarithmic vertical scale of the upper panel also allows the curve for homogeneous nucleation from the gas phase to be plotted (solid-green, Murray and Jensen, 2010).

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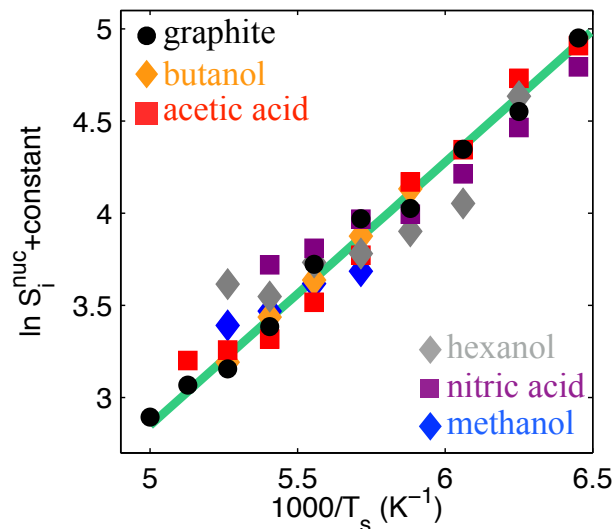


Figure 3. The logarithm of measured critical ice supersaturations S_i^{nuc} plotted vs. inverse temperature. For the purposes of comparison the nucleation data from the different materials are shifted and overlaid with the homogeneous reference case. Although in each case the magnitude of the shift is different, the shifted data all exhibit a temperature trend like that of homogeneous nucleation. The evidence of systematic changes in nucleation onset, suggests an underlying dependence on substrate material.

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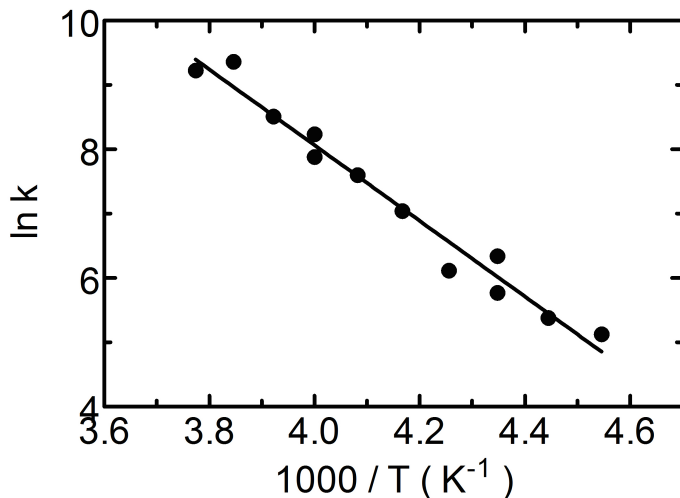



Figure 4. Plot of the logarithm of the desorption rate constant k vs. inverse temperature for D_2O desorbing from HNO_3 monolayers. A linear least squares Arrhenius fitting of the data yields a binding energy of 0.51 ± 0.06 eV with a pre-exponential factor $1 \times 10^{13.7 \pm 1.3}$ indicative of ordinary desorption. The temperature range of the desorption experiments 220 K to 265 K was chosen for experimental accessibility. At $T_s \leq 220$ K the desorption process was too slow to be observed within the experimental time window. However, HNO_3 monolayers on graphite are observed to be very robust with increasing temperature, and thus the water binding energetics are not expected to change. Kong et al. (2014) includes a detailed discussion of D_2O scattering experiments from HNO_3 surfaces.

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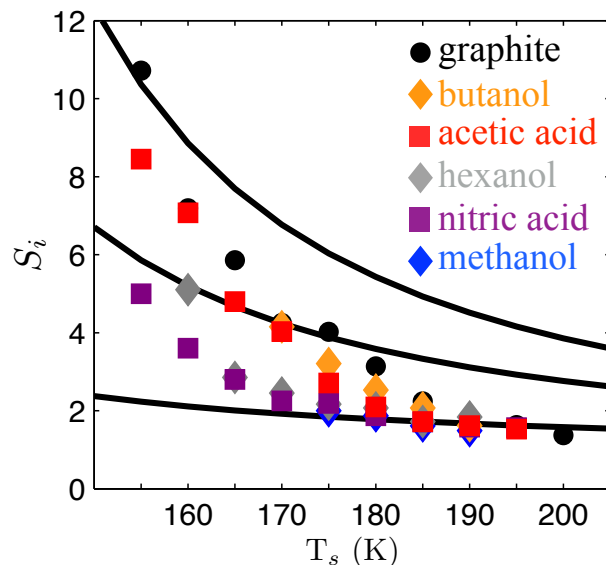


Figure 5. Measurement data plotted similarly to Fig. 2 overlaid with CNT solutions for graphite ($E_{\text{ads}} = 0.13$ eV) that assume constant f parameters for the full range of temperature. $f = 0.018, 0.09, 0.16$ (contact angle $\theta \approx 33^\circ, 51^\circ, 61^\circ$) for the lower, middle, and upper curves respectively.

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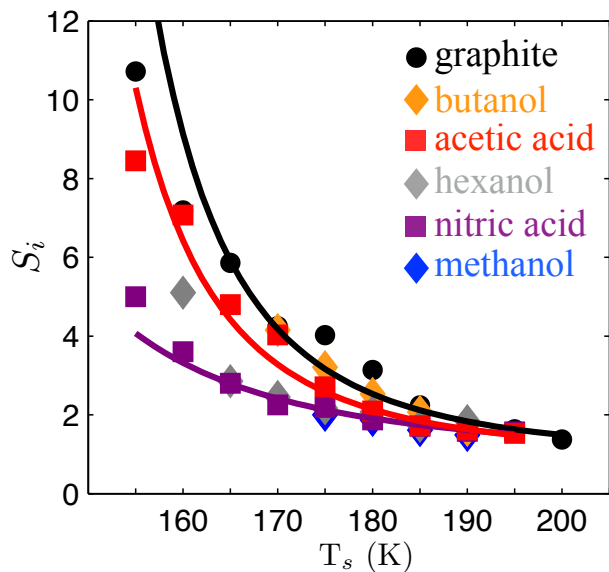


Figure 6. Measurement data plotted as in Fig. 2 with the CNT theoretical lines for the graphite (black), AcOH (red) and HNO₃ (violet) overlaid. The theoretical lines utilize least squares $f(T)$ fits to the calculated f parameters for each substrate, and illustrate the flexibility of this type of fitting – which performs equally well at the extrema of the measured data.

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