



Reactive oxygen species associated with water-soluble PM_{2.5}

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Reactive oxygen species associated with water-soluble PM_{2.5} in the southeastern United States: spatiotemporal trends and source apportionment

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Received: 10 July 2014 – Accepted: 17 July 2014 – Published: 30 July 2014

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Published by Copernicus Publications on behalf of the European Geosciences Union.

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property of ambient particulate matter suggests widespread population exposures to aerosol components that have the ability to catalyze the production of oxidants in vivo.

1 Introduction

Substantial research has been dedicated to understand the mechanisms by which ambient particulate matter (PM) causes adverse health effects in humans (Hoek et al., 2002; Samet et al., 2000; Gauderman et al., 2007; Stayner et al., 1998; Riediker et al., 2004; Sun et al., 1984; Sagai et al., 1993; Donaldson et al., 1996, 1997, 2003; Li et al., 2003, 2009a; Delfino et al., 2013). In many of these studies, a large number of PM health effects have been attributed to the oxidative or oxidant generating properties of ambient particles (Donaldson et al., 1996, 2003; Li et al., 2003, 2009a; Delfino et al., 2013). It has been hypothesized that several oxidative mechanisms can arise simultaneously, leading to a cascade of events that results in high concentrations of reactive oxygen species (ROS) in vivo (Tao et al., 2003; Castro and Freeman, 2001; Donaldson et al., 2003). ROS concentrations in excess of the anti-oxidant capacity to neutralize them leads to oxidizing other cellular components, which eventually translates into numerous health outcomes (Delfino et al., 2005; Li et al., 2009a; Peters et al., 2006).

Based on this possible mechanistic route linking aerosol chemistry to health, numerous studies have been initiated in the past several years focused on measuring the oxidative properties of ambient particles. A variety of probes have been developed to quantify different aspects of PM induced oxidative stress. These include chemical systems that mimic the loss of antioxidants, such as oxidation of dithiothreitol (DTT assay, Cho et al., 2005), glutathione (GSH, Godri et al., 2011), and ascorbic acid (AA, DiStefano et al., 2009; Mudway et al., 2004), covalent bonding with glyceraldehyde-3-phosphate dehydrogenase (GAPDH, Rodriguez et al., 2005), and hydroxyl radical generation in the presence of H₂O₂ (Shi et al., 2003). Other probes measure cellular responses when exposed to aerosols; such as macrophage ROS generation (Landreman et al., 2008), the induction of hemeoxygenase-1 (HO-1) and other stress protein

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expression (Li et al., 2003), and cytokine activation (Wilson et al., 2010). Each of these assays represents a plausible mode of PM toxicity and collectively they should be considered as the toolkit for deconstructing the complex mechanisms of PM ROS generation. One advantage of these assays, is that they are integrative of various aerosol chemical properties, and as pointed out by several researchers (Ayres et al., 2008; Venkatachari and Hopke, 2008; Kuenzli et al., 2004), many of them can serve as a screening step for assessing the PM samples/emission scenarios for more detailed chemical analysis and downstream health studies.

The DTT assay was devised to simulate the in-vivo generation of superoxide radicals, wherein DTT was used as a surrogate of the biological reducing agents (NADH and NADPH) (Kumagai et al., 1997, 2002). When incubated at 37 °C with a test PM sample, a decreasing DTT concentration over time is inferred as a measure of the ROS generating capability of the particles. A number of pure chemicals have been found to be associated with the response of this assay, e.g. quinones have been known to catalyze the transfer of electrons from DTT to oxygen (Kumagai et al., 2002). In a recent study, certain transition metals (e.g. Fe, Cu and Mn) have also been shown to be active in this assay (Charrier and Anastasio, 2012). However, evidences showing that any of these chemicals play a significant role in the DTT activity of ambient PM are sparse and limited in terms of their analytical efficacy. The major components of the ambient PM identified so far to be associated with the DTT activity are often bulk groups of species such as organic carbon (OC, both water-soluble (WSOC) and insoluble, WIOC) and water-soluble HULIS (Humic-like-substances) compounds (a class of WSOC characterized by strong hydrophobicity). However, most of these associations have been inferred based on statistical correlations (Ntziachristos et al., 2007; Verma et al., 2009a, b; Hu et al., 2008; Biswas et al., 2009b), which do not necessarily establish causation, while very few have used semi-mechanistic approaches such as physical separation of the organic compounds and metals (Verma et al., 2011; Lin and Yu, 2011; Charrier and Anastasio, 2012).

explain some of the differences reported previously on the association of DTT activity with chemical components in studies based on smaller sample sizes and provides a broader regional view of fine particle ROS characteristics.

2 Experimental methods

2.1 Sampling plan

The sampling plan for SCAPE involved paired simultaneous measurements using two instrumented sites, one of which was fixed at the Jefferson Street SEARCH (Southeastern Aerosol Research and Characterization) site (Edgerton et al., 2005, 2006; Hansen et al., 2003), the other was rotated among three different satellite sites, i.e. Yorkville, Roadside Atlanta, and Georgia Tech, all in Georgia. The sampling covered the period from June 2012 to September 2013; details of the sampling dates at each site are provided in Table 1. Each paired deployment lasted roughly one month, and was repeated in different seasons. Based on the temperature profile in Atlanta during the sampling period, shown in Supplement Fig. S1, data are classified into three seasons – summer (June–August 2012), fall (September 2012 and 2013), and winter (November 2012–March 2013). To provide a broader perspective, additional measurements (summer, June–July 2013) were made in Alabama, at the SEARCH rural-urban pair; Centreville (rural) and Birmingham (urban).

2.2 Sampling sites

A map showing the locations of all sites is shown in Fig. 1. Jefferson Street is located roughly 4 km northwest of downtown Atlanta and is representative of urban-Atlanta. The site is surrounded by commercial and residential buildings, with nearby street traffic patterns consistent with a mixed-use urban environment. The site has extensively been used in past studies characterizing urban Atlanta air quality (Hansen et al., 2006; Marmur et al., 2005) and the data used in epidemiological studies (Sarnat et al., 2008;

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summertime SOA than wintertime biomass burning on WSOC in the region is apparent from a gradually decreasing averaged concentration profile at JST from summer to winter ($3.8 \pm 2.2 \mu\text{g m}^{-3}$, $3.2 \pm 1.7 \mu\text{g m}^{-3}$ and $2.0 \pm 1.1 \mu\text{g m}^{-3}$ in summer, fall and winter). Similar decreases are also noted at other sites ($3.7 \pm 1.9 \mu\text{g m}^{-3}$ in summer to $1.8 \pm 0.8 \mu\text{g m}^{-3}$ in winter at YRK, and $3.3 \pm 1.1 \mu\text{g m}^{-3}$ and $3.9 \pm 2.3 \mu\text{g m}^{-3}$ in fall, 2012 and 2013, respectively to $1.6 \pm 1.3 \mu\text{g m}^{-3}$ in winter, 2013 at RS). These results are consistent with a more detailed analysis of the WSOC spatial and seasonal distribution conducted two years before in the southeast US (Zhang et al., 2010).

BrnC denotes the water-soluble organic compounds that absorb light in the UV-region of the UV-Vis spectrum (Hecobian et al., 2010). Biomass burning is by far the largest source of these compounds with additional contributions from vehicular sources and possibly aged SOA (Hecobian et al., 2010; Zhang et al., 2011). The spatial profile of BrnC (Fig. 2c) is intermediate between EC (heterogeneous) and WSOC (homogeneous), indicating the diversity of its sources. The moderate impact of vehicular emissions on BrnC is probably reflected in generally higher concentrations of BrnC at RS and other urban sites (JST, GT, BHM) than rural (CTR, YRK). However, the strong seasonality is evident from its extreme values in winter periods resulting from biomass burning (average levels at JST in winter = $0.97 \pm 0.78 \text{ M m}^{-1}$, compared to $0.42 \pm 0.25 \text{ M m}^{-1}$ in summer and $0.53 \pm 0.33 \text{ M m}^{-1}$ in fall).

3.1.2 Water-soluble metals

The selected elements (Fig. 3), have a wide range of concentrations at various sites and differing seasonal trends, pointing to heterogeneous sources. A more in-depth discussion on the distribution of the measured water-soluble elements will follow in a subsequent publication, here we only briefly discuss the major trends in selected metals. BHM, a site heavily impacted by industrial emissions, has very high concentrations of all metals, with Ca, Mn, and Zn the highest among all sites. After BHM, RS has the highest levels of most metals among all sites in the Atlanta region (i.e. JST, GT, YRK, RS), indicating either road-dust and/or direct vehicular emissions as their major source.

Rural sites, YRK and CTR, have generally lower metals concentrations, while GT and JST levels are moderate.

Higher rainfall in winter compared to summer months could also affect the metals concentrations since the re-entrainment of mineral dust, which is a major source of metals, is most effective in dry periods. The metals of predominantly crustal and vehicular origin (e.g. Ca, Fe, Mn and Cu) are generally higher in summer and fall compared to the winter. For example, the average Ca concentration at JST was 155 ± 116 , 160 ± 87 , and $67 \pm 37 \text{ ng m}^{-3}$ in summer, fall, and winter, respectively. Exceptions are K and Zn, which despite being associated with crustal dust (K; Coz et al., 2010) or vehicular sources (Zn; Liu et al., 2008) are also emitted in biomass burning (Echalar et al., 1995), which could account for a weaker seasonal variability due to multiple sources; average K concentration is 62 ± 25 , 59 ± 22 and $65 \pm 32 \text{ ng m}^{-3}$ in summer, fall and winter, while the respective Zn concentration is 9.6 ± 4.7 , 9.3 ± 4.8 and $9.9 \pm 5.8 \text{ ng m}^{-3}$ at JST.

3.2 PM ROS generation potential

Temporal and spatial differences in the water-soluble DTT activity are discussed to provide an overview of the PM ROS generation potential in Southeastern US.

Time series of daily average DTT activity from the sites in different months is shown in Fig. 4a and b. The rate of consumption of DTT is normalized by both the volume of sampled air (DTTv; expressed in units of $\text{nmol min}^{-1} \text{ m}^{-3}$, Fig. 4a) and the particulate mass (DTTm; expressed in units of $\text{nmol min}^{-1} \mu\text{g}^{-1}$, Fig. 4b). While volume normalized activity is a function of strength of the emission source, and thus an extrinsic property, mass normalized levels represent an intrinsic property of PM, independent of its mass concentration and a characteristic of the specific source.

Our measurements of both volume and mass normalized DTT activity are generally in the typical range ($0.1\text{--}1.5 \text{ nmol min}^{-1} \text{ m}^{-3}$ for DTTv and $0.005\text{--}0.1 \text{ nmol min}^{-1} \mu\text{g}^{-1}$ for DTTm, as summarized in Fang et al., 2014) observed for ambient particles at other locations. Figure 4 shows that in general, both the extrinsic (DTTv) and intrinsic (DTTm)

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activity of ambient PM is spatially uniform. There is no consistent and substantial difference in DTT activity of ambient PM at the two sites for any paired measurement, except the JST-YRK pair in winter [$DTTv(YRK) < DTTv(JST)$] and GT-RS pair in fall [$DTTm(RS) > DTTm(GT)$]. Intrinsic DTT activity is fairly uniform even on day-to-day scales for most sites, suggesting that the sources of ROS-associated aerosol components don't change rapidly. There is, however, a substantial daily variability in the extrinsic activity, which reflects variations in emissions, dispersion and sinks, factors that also control overall PM mass concentrations.

A summary of the spatial and seasonal trends in mean DTT activity (both $DTTv$ and $DTTm$) for the individual sampling periods and at various sites is shown in Fig. 5. A significant seasonal variability in both the volume and mass normalized DTT activity is evident; the levels are generally higher in the colder months than summer. A one-sample t -test showed that the $DTTv$ at JST in December was significantly ($p \leq 0.05$) higher from that in June–July ($51 \pm 34\%$), August ($48 \pm 22\%$), and September ($27 \pm 14\%$). These results are in agreement with a recent study conducted in the Los Angeles Basin, which showed a generally higher DTT activity (both mass and volume normalized) of quasi-ultrafine particles in cooler months compared to warmer periods (Saffari et al., 2014). The authors in that study attributed the higher DTT activity to an elevated concentration of redox-active semi-volatile organic compounds (SVOC) caused by their enhanced partitioning to the particulate phase and the lowered atmospheric mixing height in winter. Although, SVOC have been suggested to make a substantial contribution to the DTT activity of ultrafine particles (Verma et al., 2011), the $PM_{2.5}$ samples collected via HiVol filters in our study are not expected to contain a significant fraction of these species due to losses associated with large flow-rate and long sampling duration (Ashbaugh and Eldred, 2004).

The seasonal variability is even more pronounced in the intrinsic DTT activity. For example, differences in $DTTm$ levels in December vs. June–July ($119 \pm 48\%$), August ($86 \pm 31\%$) and September ($44 \pm 14\%$) are higher compared to the respective differences in $DTTv$ levels. Similar increases in the intrinsic DTT activity in

winter were observed at other sites too ($38 \pm 13\%$ at GT from average of August 2012 and September 2013, and $53 \pm 24\%$ at YRK), but not at RS, where average fall level ($0.032 \pm 0.009 \text{ nmol min}^{-1} \mu\text{g}^{-1}$) was nearly the same as in winter ($0.036 \pm 0.008 \text{ nmol min}^{-1} \mu\text{g}^{-1}$).

To further assess the spatiotemporal variability in DTTv activity in the region, correlation coefficients (R^2) were calculated for all of the site pairs and are shown in Fig. 6. To put the DTT activity comparison in a larger context, we have also included the corresponding coefficients for EC and WSOC. A high spatial correlation for a species indicates it has more spatially uniform emission sources in the region, on the scale of distances between the paired sites.

The site pair JST-GT shows high correlation coefficients ($R^2 > 0.5$) for many PM species (DTT, EC and WSOC), which was expected given their spatial proximity and absence of nearby strong sources, e.g., roadways. The other site pairs show disparate patterns for the different PM species depending upon their sources. For example, EC has relatively low correlation coefficients for the urban-rural site pairs, i.e. JST-YRK in both summer ($R^2 = 0.42$) and winter (0.45), and BHM-CTR ($R^2 = 0.39$), as expected. Despite relatively close proximity, the low correlation in fall for JST-RS pair ($R^2 = 0.10$) and GT-RS pair ($R^2 = 0.48$) for EC is due to the strong influence of freeway emissions (I-85) at the RS site. However, the JST-RS correlation is high in February ($R^2 = 0.74$), possibly suggesting some contributions from a common source of EC, such as biomass burning.

WSOC on the other hand exhibits high correlations for most site pairs (JST-YRK, JST-GT, CTR-BHM; $R^2 > 0.5$) in all seasons. The correlations are generally higher in summer ($R^2 > 0.70$), indicating a more regional influence of SOA formation as a source of WSOC in summertime, compared to biomass burning in winter ($R^2 < 0.65$). The correlations are generally weaker for the pairs involving RS ($R^2 = 0.08\text{--}0.46$), probably due to the added contribution of freeway emissions to WSOC.

Compared to WSOC and EC, the spatial correlations for DTT activity are moderate in all seasons. For example, R^2 for DTT are not as high as for EC for the sites in closer

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proximity ($R^2 = 0.68$ and 0.67 for JST-GT in fall and winter, respectively). Similarly, in summer when the correlations for WSOC are very high even for the far-apart sites (JST-YRK and CTR-BHM), correlations for DTT are only moderate ($R^2 = 0.47$ and 0.59). It appears that DTT activity is impacted by both regional sources and local emissions associated with individual site characteristics. The role of various emission sources in the DTT activity of ambient PM in different seasons and sites is the major point of discussion in the next section and also the subsequent papers from this study.

3.3 Sources of PM ROS generation potential

Sources of ROS generation potential were identified and apportioned using combination of linear regression, factor analysis and chemical mass balance techniques as described below.

3.3.1 Linear regression of DTT activity with PM chemical composition

Pearson's correlation coefficients for the linear regression between DTT activity and the measured chemical components were calculated for each site and season and are shown in Table 2. Considering the limited daily variability in mass normalized DTT activity at most sites in a given season, the regression was conducted on the volume normalized levels of DTT activity ($\text{nmol min}^{-1} \text{m}^{-3}$) and chemical components ($\mu\text{g m}^{-3}$). To consolidate the data for convenience, individual time-series at JST and GT sites from June-to-August were merged (named “JST-GT Summer”), given their similar concentration profiles. Similarly, JST and GT time series from December to March were merged as “JST-GT winter”. The complete regression matrix showing the correlation between all pairs of selected species, for each site and season, is provided in the Supplement (Table S1).

DTT activity is almost always correlated with OC in the present study. OC results from all major sources in the region (e.g. vehicular emissions, SOA, and biomass burning) and comprises the greatest fraction of PM mass ($\text{OM}/\text{PM}_{2.5} > 60\%$, not shown).

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Thus, the strong correlation of OC underlines the general contribution of PM organic compounds in the ROS generation potential, and also likely accounts for the DTT correlation with PM mass, as reported in Fang et al. (2014). However, it does not yield any novel information on the specific PM species or the emission sources associated with the DTT activity. Associations between water-soluble DTT activity and organic compounds are consistent with a number of previous studies (Biswas et al., 2009a; Verma et al., 2012; Cho et al., 2005; Yang et al., 2014), but in contrast to other reports showing metals as the major drivers in this reaction (Charrier and Anastasio, 2012).

In summer, DTT activity is well correlated with WSOC at all sites; JST-GT ($R = 0.81$), YRK ($R = 0.79$), CTR ($R = 0.78$) and BHM ($R = 0.67$). Inorganic ions such as SO_4^{2-} and NH_4^+ are also correlated with WSOC and DTT activity ($R > 0.60$ at most sites) indicating secondary photochemical formation as their common source in summer. This is consistent with the results of ambient (Verma et al., 2009a) and chamber studies (McWhinney et al., 2013; Li et al., 2009b) suggesting higher DTT activity of secondary organic compounds than their parent gases and primary particles.

In addition to WSOC, summertime DTT activity at urban sites, i.e. JST-GT and BHM, is also correlated with certain transition metals (Mn ($R = 0.72$), Fe (0.78), and Cu (0.71) at JST-GT, and Mn (0.75), Fe (0.66) and Zn (0.66) at BHM). At BHM, the iron foundry and coke oven plants are the likely sources of these DTT-correlated metals. None of these metals are strongly correlated with any specific emission markers at JST-GT. However, a moderate correlation of EC with DTT activity ($R = 0.69$) and also somewhat with metals ($R = 0.59, 0.75$ and 0.58 for Mn, Fe, and Cu; Supplement Table S1), suggests that vehicular emissions could be one of their common source.

BrnC is also correlated with DTT activity in summer at JST-GT, CTR, and BHM sites ($R > 0.65$). However, its correlation with both WSOC and EC ($R > 0.65$ at all sites; Table S1) does not allow isolating its predominant source. Both vehicular emissions and possibly aged SOA appear to contribute to these chromophores in summer.

In fall, DTT activity appears to be derived from a mixed contribution from primary vehicular emissions and secondary formation, as supported by its correlation with WSOC

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($R = 0.72$), and EC ($R = 0.90$) at JST. Metals such as Fe and Zn are also correlated with DTT activity ($R = 0.70$ and 0.82) and EC ($R = 0.90$ and 0.75) at JST. The stronger correlation of primary emissions (EC) than secondary species (WSOC, SO_4^{2-} , NH_4^+) indicates a diminishing effect of photochemistry on the DTT activity. This is even more apparent in fall 2013, when the highest correlation of DTT at GT is with EC ($R = 0.78$), while the correlation with WSOC is low ($R = 0.47$).

However, at the RS site, road-dust suspension also appears to be a significantly contributing factor to DTT activity in fall as indicated from the moderate to strong correlation ($R = 0.61$ – 0.84) of certain metals including those from crustal sources (Cu and Zn in fall 2012, and K, Ca, Mn, and Zn in fall 2013). The added contribution from dust probably dilutes the association of DTT activity with other emission sources. For example, DTT at RS is not as strongly correlated with EC ($R \leq 0.55$) or WSOC ($R = 0.71$ and 0.21 in Fall, 2012 and 2013, respectively) as at JST and GT.

In winter months, DTT activity is again correlated with WSOC at all sites; JST-GT ($R = 0.69$), YRK ($R = 0.84$), and RS ($R = 0.78$), however, this correlation is not accompanied by a similar correlation with inorganic ions ($R < 0.55$ for SO_4^{2-} and NH_4^+) as in summer months. It implies that another emission source of WSOC – biomass burning, starts contributing more to the DTT activity in winter. This is further reflected in the strong correlation of BrnC and K with DTT activity and WSOC ($R > 0.65$; Table 2 and Supplement Table S1) at all sites. In a study conducted in Los Angeles, the ambient concentrations of K and WSOC were elevated ($\sim 2 \times$) during the 2007 Southern California wildfires compared to the post-fire period (Verma et al., 2009b). More importantly, these wood-smoke particles were found to be at least two times more oxidative than the post-fire ambient PM_{2.5} typically dominated by vehicular emissions. The redox-active WSOC emitted in the biomass burning particles probably explain the higher intrinsic DTT activity of ambient PM in winter compared to other seasons in the present study (Figs. 4 and 5).

Interestingly, EC is also highly correlated with DTT activity ($R \geq 0.75$) in winter at all sites. However, the concurrent correlation of EC with BrnC and K ($R > 0.60$ at all sites;

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emissions. This factor is also most significant in summer and fall, which is consistent with Fig. 3 showing a generally lowered concentration of metals in winter. Factor 4 has a clear signature of biomass burning with very high concentrations of BrnC and K, and is prominent in winter. Lastly, the EC peak in factor 5 is a distinct feature suggesting it is linked to primary vehicular emissions. Also consistent with the vehicular sources, this factor does not have an obvious seasonal pattern.

The time-series of reconstituted DTT activity resolved by the five PMF factors, along with the residuals (not predicted), and a distribution of these factors in each season, is shown in Fig. 8. In summer, secondary oxidation processes (WSOC (29 %) + SO₄⁻² (17 %), factors 1 and 2) are the dominant contributors (46 %) to DTT activity, with additional contributions from biomass burning (24 %, factor 4) and vehicle emissions (19 %, factor 5). In fall, the contribution from secondary formation drops to 32 % (22 % associated with WSOC and only 10 % with SO₄⁻² factors) with an increase in the contribution from vehicle emissions (25 %). Biomass burning dominates the ROS generation potential in winter, contributing 47 % to the DTT activity, while relatively low contribution from secondary formation (20 %) and vehicle emissions (12 %) is observed. Mineral dust (factor 3), possibly including road-dust resuspension also makes a non-zero contribution to DTT activity, but is significant only in summer (13 %) and fall (11 %).

Relative loadings of individual species in various PMF factors, and the contributions of each of these factors to the DTT activity, could provide some insights on the specific chemical components in an emission source driving the aerosol oxidative potential. For example, WSOC has been associated with the DTT activity in previous studies (Verma et al., 2012; Biswas et al., 2009a) and our results showing the substantial DTT activity associated with factor 1 (Figs. 7 and 8) support these findings. A few laboratory studies have also noted the capability of transition metals, particularly Fe, Cu and Mn to transfer the electrons from DTT, (Charrier and Anastasio, 2012). In this study, metals were mostly concentrated in factor 3, i.e. mineral dust, which has the smallest contribution suggesting a lesser role for metals. However, the substantial contribution of biomass burning aerosols (factor 4) to DTT activity could be attributed in part to metals, along

The Supplement related to this article is available online at
doi:10.5194/acpd-14-19625-2014-supplement.

Acknowledgements. This publication was made possible by US EPA grant R834799. This publication's contents are solely the responsibility of the grantee and do not necessarily represent the official views of the US EPA. Further, US EPA does not endorse the purchase of any commercial products or services mentioned in the publication. The authors would like to thank R. Erik Weber who helped in sample collection and sample preparation for DTT, WSOC, BrnC and metals analysis, Janessa R. Rowland and Madhusudan Kamat for assistance with filter extractions and IC-analysis, and the SEARCH personnel for their many contributions supporting the field deployments.

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Table 1. Sampling plan and details of sampling dates at each site.

Season, Year	Sapling dates		Sampling Site			
	Start date	End date	Fixed site	Samples collected (N)	Mobile site	Samples collected (N)
Summer, 2012	8 Jun 2012	20 Jul 2012	JST	31	YRK	33
	24 Jul 2012	31 Aug 2012	JST	37	GT	38
Fall, 2012	6 Sep 2012	4 Oct 2012	JST	26	RS	29
Winter, 2012–2013	15 Nov 2012	30 Nov 2012	JST	13	JST	14
	6 Dec 2012	4 Jan 2013	JST	22	YRK	22
	27 Jan 2013	27 Feb 2013	JST	30	RS	31
	5 Mar 2013	27 Mar 2013	JST	23	GT	22
Summer, 2013	16 Jun 2013	16 Jul 2013	CTR	31	BHM	31
Fall, 2013	9 Sep 2013	3 Oct 2013	GT	25	RS	25
Total number of samples				238		245

Note: In the summer of 2013 (shown in bold) measurements were made at sites outside of the State of Georgia as part of other studies to provide a greater context and were not part of the paired sampling approach with JST as the central site.

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Table 2. Regression analysis (Pearson's R) between DTT activity and selected PM components.

Season	Site	WSOC	BrnC	Inorganic ions		OC	EC	Water-soluble metals					
				SO ₄ ²⁻	NH ₄ ⁺			K	Ca	Mn	Fe	Cu	Zn
Summer, 2012	<i>JST-GT</i>	0.81	0.67	0.68	0.73	0.81	0.69	0.49	0.20	0.72	0.78	0.71	0.62
	<i>YRK</i>	0.79	0.53	0.58	0.59	0.76	0.56	0.51	0.68	0.63	0.48	0.09	0.48
Fall, 2012	<i>JST</i>	0.72	0.81	0.49	0.52	0.83	0.90	0.61	-0.08	0.37	0.74	0.20	0.82
	<i>RS</i>	0.71	0.59	0.62	0.66	0.77	0.55	0.52	0.12	0.28	0.44	0.61	0.68
Winter, 2012–2013	<i>JST-GT</i>	0.69	0.78	0.15	0.13	0.85	0.82	0.78	-0.13	0.46	0.65	0.63	0.63
	<i>YRK</i>	0.84	0.88	0.34	0.45	0.82	0.80	0.75	0.31	0.68	0.04	0.43	0.72
	<i>RS</i>	0.78	0.86	0.47	0.52	0.86	0.75	0.69	0.31	0.49	0.57	0.35	0.60
Summer, 2013	<i>CTR</i>	0.78	0.88	0.71	0.78	0.77	0.72	0.66	0.23	0.36	0.41	-0.12	0.62
	<i>BHM</i>	0.67	0.74	0.66	0.66	0.85	0.68	0.41	0.55	0.75	0.66	0.23	0.66
Fall, 2013	<i>GT</i>	0.47	0.75	0.39	0.37	0.79	0.78	0.65	0.29	0.58	0.66	0.46	0.46
	<i>RS</i>	0.21	0.48	0.60	0.47	0.58	0.53	0.74	0.79	0.84	0.53	-0.09	0.74

Note: $R > 0.65$ are bolded.

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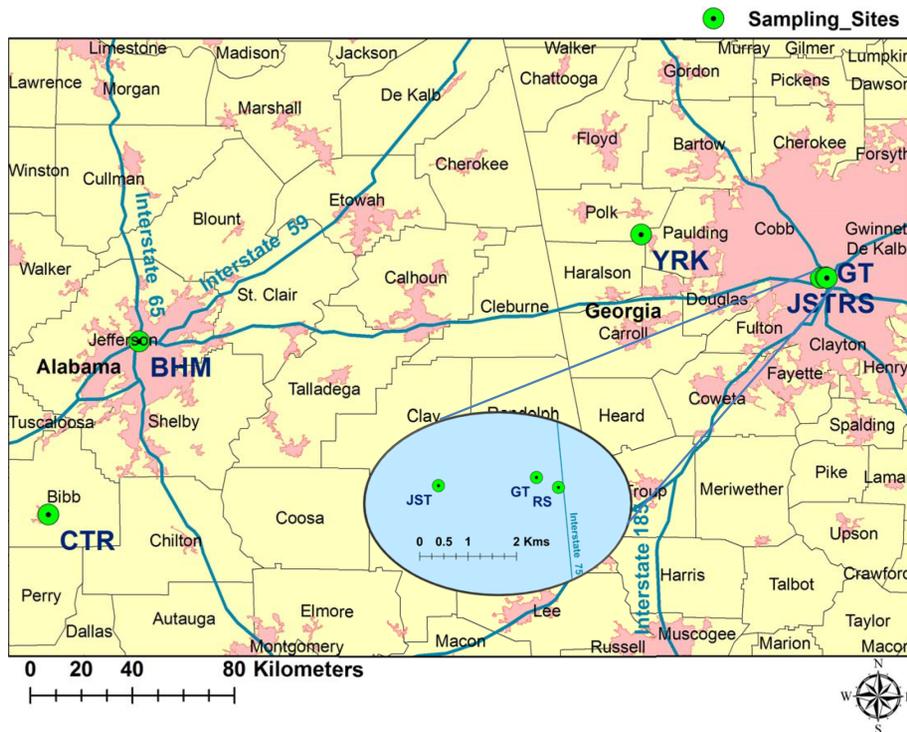


Figure 1. Map of sampling sites.

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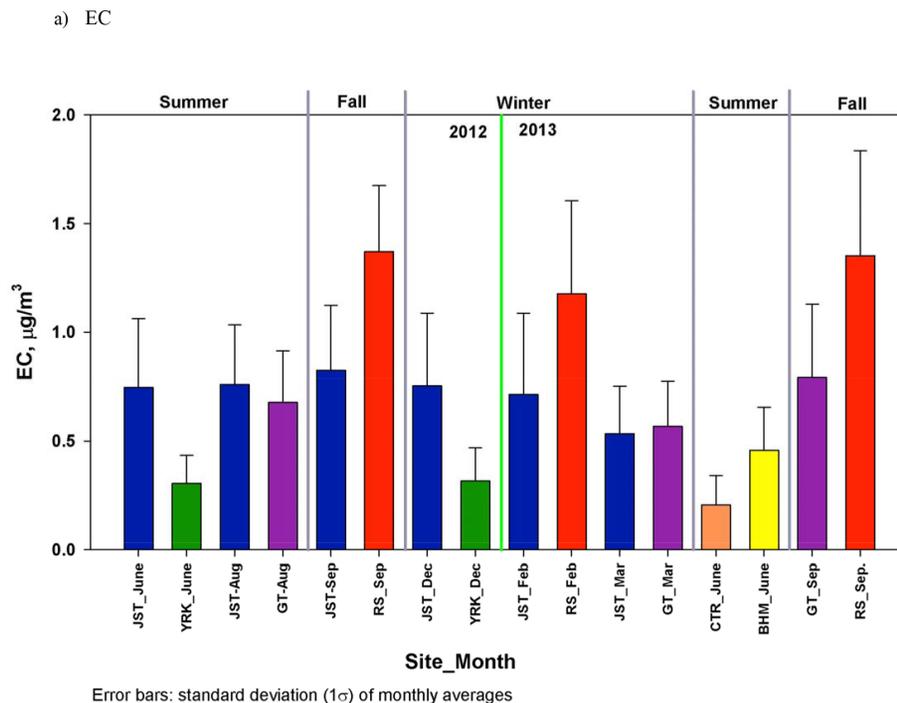


Figure 2. Monthly average ambient concentrations of EC (a) WSOC (b) and BrnC (c) at the sampling sites.

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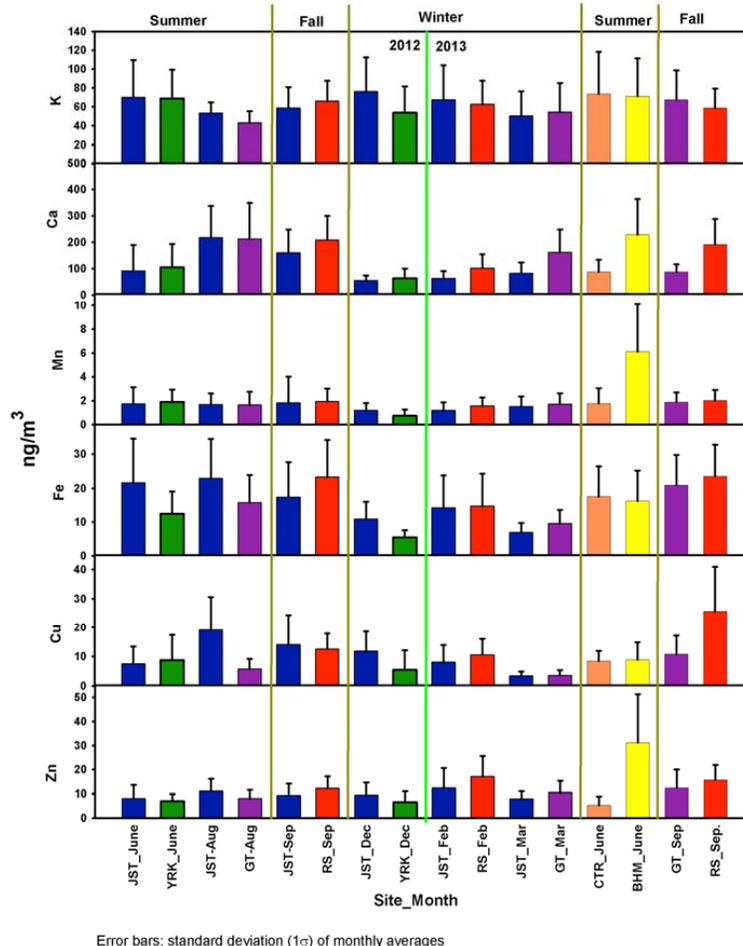


Figure 3. Monthly average ambient concentrations of water-soluble metals at the sampling sites.

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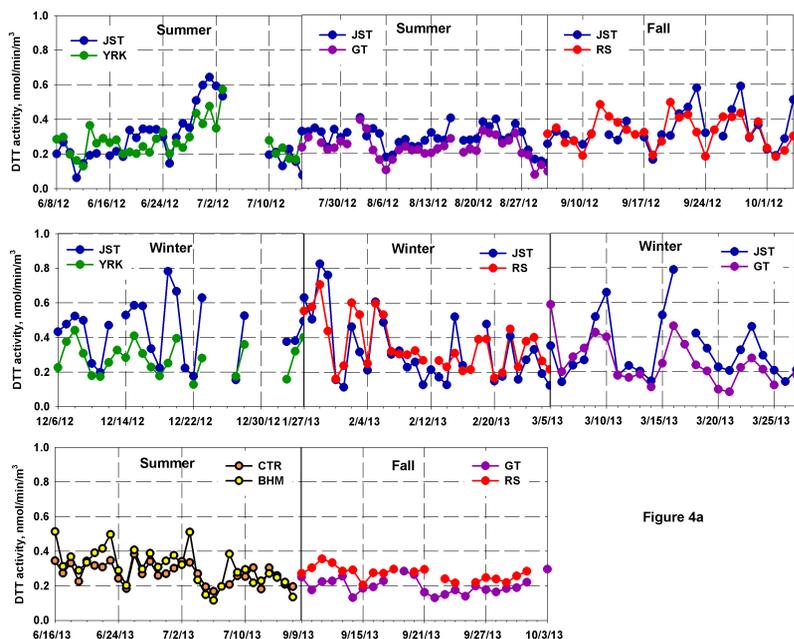


Figure 4a

Figure 4. Time-series of the daily average DTT activity (DTTv; **a** and DTTm; **b**) of ambient PM_{2.5} at the sampling site.

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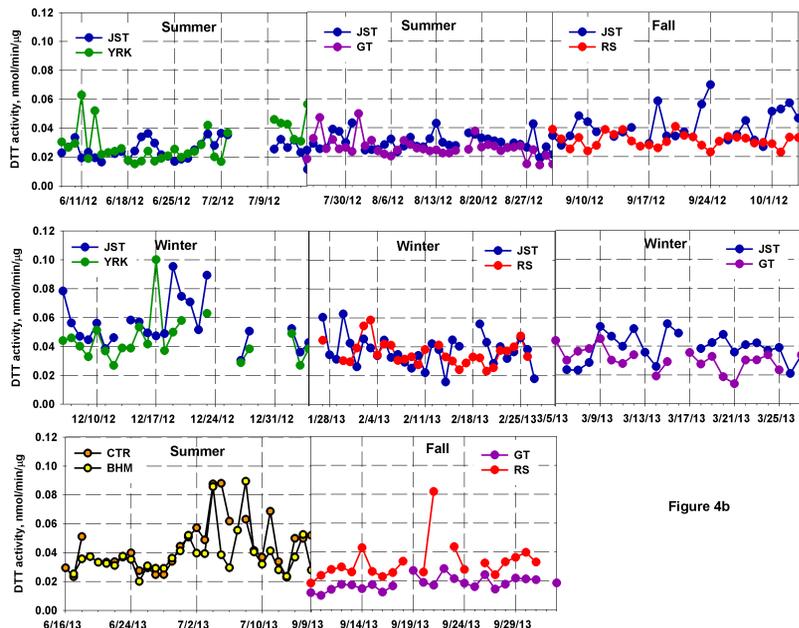


Figure 4b

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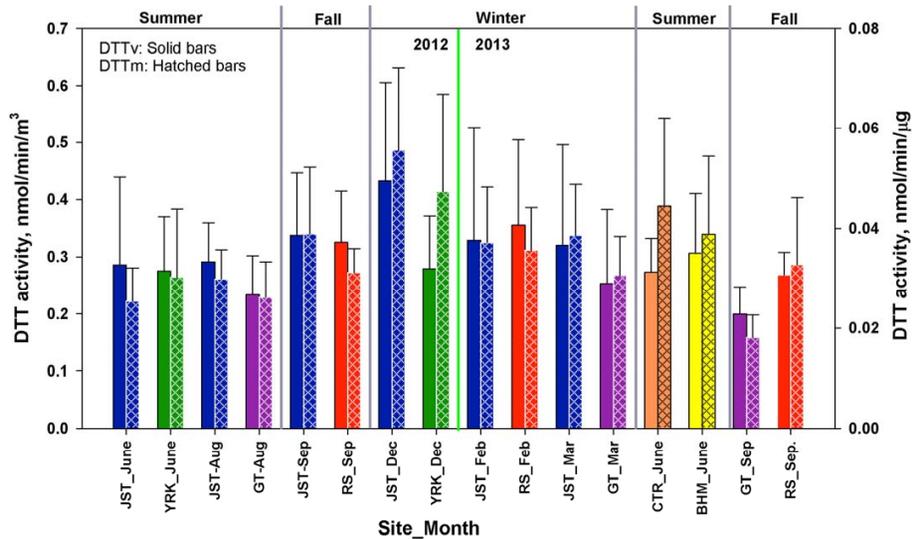


Figure 5. Monthly averages of volume- (DTTv) and mass-normalized (DTTm) DTT activity at the sampling sites.

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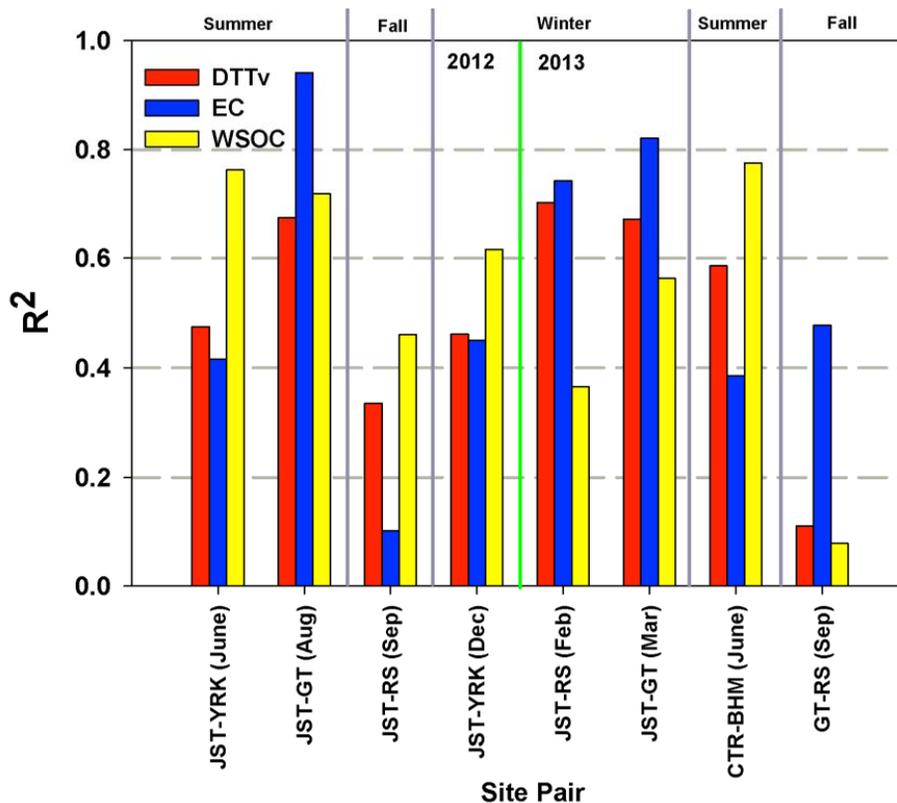


Figure 6. Site-to-site correlations (R^2) for volume normalized DTT activity (DTTv), elemental carbon (EC) and water-soluble carbon (WSOC) of PM_{2.5}.

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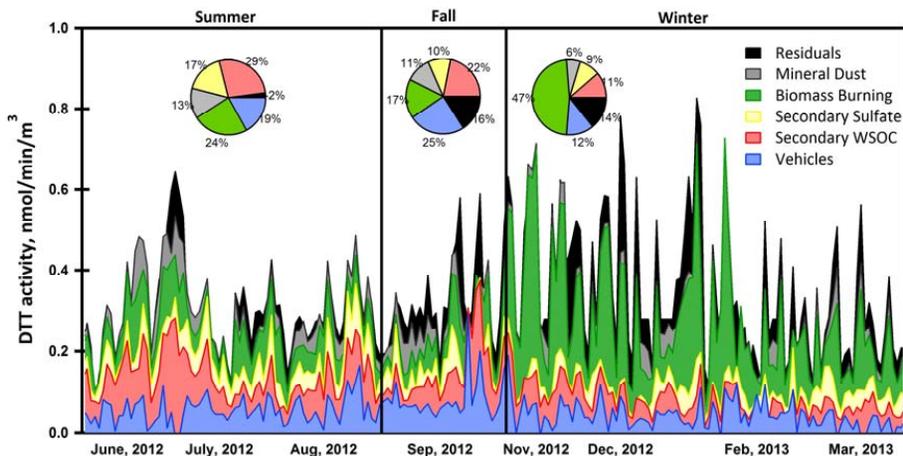


Figure 8. Time series of PM_{2.5} DTT activity segregated into different factors as predicted by PMF and their season-wise contributions.

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