

Explaining variance in black carbon’s aging timescale

Laura Fierce¹, Nicole Riemer², and Tami C Bond¹

¹Department of Civil and Environmental Engineering, University of Illinois at Urbana-Champaign, Urbana, Illinois, USA

²Department of Atmospheric Sciences, University of Illinois at Urbana-Champaign, Urbana, Illinois, USA

Correspondence to: N. Riemer
(nriemer@illinois.edu)

Abstract. The size and composition of particles contain-
ing black carbon (BC) are modified soon after emission
by condensation of semi-volatile substances and coagulation
with other particles, known collectively as “aging” processes.
Although this change in particle properties is widely recog-
nized, the timescale for transformation is not well con-
strained. In this work, we simulated aerosol aging with the
particle-resolved model PartMC-MOSAIC and extracted aging
timescales based on changes in particle cloud condensa-
tion nuclei (CCN). We simulated nearly 300 scenarios and,
through a regression analysis, identified the key parameters
driving the value of the aging timescale. We show that BC’s
aging timescale spans from hours to weeks, depending on
the local environmental conditions and the characteristics of
the fresh BC-containing particles. Although the simulations
presented in this study included many processes and particle
interactions, we show that 80% of the variance in the aging
timescale is explained by only a few key parameters. The
condensation aging timescale decreased with the flux of con-
densing aerosol and was shortest for the largest fresh par-
ticles, while the coagulation aging timescale decreased with
the total number concentration of large ($D > 100$ nm), CCN-
active particles and was shortest for the smallest fresh par-
ticles. Therefore, both condensation and coagulation play im-
portant roles in aging, and their relative impact depends on
the particle size range.

snow (Hansen and Nazarenko, 2004; Jacobson, 2004). Each
of these climate effects depends on the properties of individ-
ual BC-containing particles and their atmospheric residence
time. The dominant removal mechanism of BC mass from
the atmosphere is wet deposition (Cozic et al., 2007), with
one important pathway being the activation of BC-containing
particles into cloud condensation nuclei (CCN) and their sub-
sequent removal if the cloud precipitates. Although freshly
emitted BC-containing particles are too small and hydropho-
bic to activate (Maricq, 2007; Weingartner et al., 1997), their
morphology and chemical composition are altered soon af-
ter emission by condensation of semi-volatile gases and co-
agulation with pre-existing particles (Johnson et al., 2005;
Oshima et al., 2009; Zaveri et al., 2010).

These changes in particle characteristics, termed “aging”,
often increase the particles’ susceptibility to cloud droplet
nucleation and wet removal (Furutani et al., 2008; Cantrell
et al., 2001; Zuberi et al., 2005), so these processes must be
included in global models. However, a complex aerosol pop-
ulation that evolves with time is not easily simulated in cli-
mate models, so even sophisticated aerosol schemes do not
fully resolve aerosol properties on a per-particle level (Jacob-
son, 1997; Wexler et al., 1994; Bauer et al., 2008; Binkowski
and Roselle, 2003; McGraw, 1997; Jacobson, 2002; Aquila
et al., 2011; Matsui et al., 2013). The simplest representa-
tion of aging classifies BC mass as either hydrophobic or
hydrophilic, such that hydrophilic BC is susceptible to re-
moval by wet deposition and hydrophobic BC is not. In this
framework, BC is transferred from the hydrophobic (fresh)
category to the hydrophilic (aged) category according to a
first-order aging timescale (Cooke and Wilson, 1996; Croft
et al., 2005; Koch, 2001). Global models apply a fixed aging
timescale of 1–3 days. Global modeling studies have shown
that estimates of BC’s climate forcing are sensitive to the as-
sumed aging timescale (Koch et al., 2009), but its value is
not well constrained. *While some climate models have been*

1 Introduction

Particles containing black carbon (BC) alter the Earth’s en-
ergy balance by scattering and absorbing solar radiation (Mc-
Cormick and Ludwig, 1967; Rosen et al., 1978; Schulz et al.,
2006), by interacting with clouds (Twomey, 1977; Twomey
et al., 1984; Lohmann et al., 2005; Albrecht, 1989; Ackerman
et al., 2000), and by decreasing the albedo of ice and

moving toward aerosol modules that represent aerosol aging using several interacting modes (Aquila et al., 2011; Bauer et al., 2008; Wilson et al., 2001; Matsui et al., 2013), the practice of using a fixed aging timescale is still widespread (Jo et al., 2013; Chin et al., 2014; Schmidt et al., 2014).

To improve upon using one constant value for the aging timescale, several studies have developed parameterizations of BC's aging timescale that vary with environmental conditions. Liu et al. (2011) developed a parameterization of black carbon aging by condensation that depended on the condensation rate of sulfuric acid and overall BC surface area. They showed that, by allowing for slower aging in the winter, their parameterization was better able to represent seasonal variability in black carbon transport to the Arctic. Oshima and Koike (2013) extended this approach and developed a parameterization of aging timescales based on simulations with a box model. Their parameterization predicted the rate for BC to transition from a hydrophobic class to a hydrophilic class, expressed as a function of the mass-normalized coating rate and on parameters of the fresh BC size distribution. Riemer et al. (2004) showed that timescales for aging by coagulation decrease with the overall aerosol number concentration, which they parameterized using a simple power law, and this parameterization was applied by Croft et al. (2005). Pierce et al. (2009) parameterized size-resolved coagulation rates as a first-order loss process that depends on the overall size distribution. In an analysis of aging timescales in a specific urban environment using a particle-resolved model, Riemer et al. (2010) showed that timescales for particles to transition from CCN-inactive to CCN-active varied diurnally due to variations in condensation aging rates. Because the timescale from Riemer et al. (2010) is based on changes in particle CCN activity, it quantifies changes in particle characteristics that these first-order aging models are meant to represent.

This study builds on the work of Riemer et al. (2010) to generalize how the CCN-based aging timescale varies with *scenario-specific properties*. Unlike other aerosol schemes, which simplify the representation of particle composition, the particle-resolved model tracks the composition of each simulated particle and is, therefore, uniquely suited to study the impact of aging on per-particle CCN activity. *The focus of this paper is to identify the set of independent variables that best explain variance in BC's aging timescale for a large collection of simulations. Using the independent variables identified in this study, we will later introduce a simple aging parameterization for use in global models.*

2 Extracting aging timescales from particle-resolved model output

In a first-order model of aging, particles transition from fresh to aged according to an aging timescale, τ_{aging} . In this framework, a criterion must be applied to distinguish fresh and

aged particles. Particle CCN activity at a specified environmental supersaturation is the aging criterion applied in this work, so the aging timescale indicates changes in particles' susceptibility to removal by wet deposition. We define aged particles as those that are activated at a specified environmental supersaturation, and fresh particles are CCN-inactive at that supersaturation threshold. The first-order aging model is given by:

$$\left[\frac{dN_{\text{fresh}}}{dt} \right]_{\text{aging}} = -\frac{1}{\tau_{\text{aging}}} N_{\text{fresh}}, \quad (1)$$

where N_{fresh} is the number concentration of fresh particles.

Before discussing the full set of sensitivity simulations in Section 2.3, we describe the particle-resolved simulation of aerosol dynamics in a baseline scenario and show how the aging timescale is used to quantify changes in per-particle CCN activity. The particle-resolved model is described in Section 2.1, and the κ -Köhler model (Petters and Kreidenweis, 2007) for computing CCN activity from the PartMC-MOSAIC data is discussed in Section 2.2. We discuss methods for computing aging timescales from the particle-resolved model output in Section 2.3.

2.1 Particle-resolved simulation of aerosol aging

The Particle Monte Carlo model (Riemer et al., 2009) coupled to the Model for Simulating Aerosol Interactions and Chemistry (Zaveri et al., 2008), PartMC-MOSAIC, is a Lagrangian box model that simulates gas and aerosol chemistry, gas-aerosol mass transfer, aerosol coagulation, gas and aerosol emissions, and dilution with background air. *The boundary layer height varies temporally according to a prescribed profile. The treatment of dilution is the same as in (Riemer et al., 2009); a constant dilution rate of $1.5 \times 10^{-5} \text{ s}^{-1}$ is applied and additional dilution with background air occurs when the boundary layer height increases.* Coagulation events, particle emissions, and dilution with background air are simulated stochastically by PartMC. Gas- and aerosol-phase chemistry and gas-aerosol mass transfer are simulated deterministically by MOSAIC. MOSAIC includes modules for gas-phase photochemistry (Zaveri and Peters, 1999), particle-phase thermodynamics (Zaveri et al., 2005b,a), and gas-particle mass transfer (Zaveri et al., 2008). *MOSAIC treats secondary organic aerosol formation based on the SORGAM scheme (Schell et al., 2001).* The coupled model represents all atmospherically important aerosol species, including sulfate (SO_4), nitrate (NO_3), chloride (Cl), carbonate (CO_3), ammonium (NH_4), sodium (Na), calcium (Ca), methanesulfonic acid (MSA), black carbon (BC), primary organic aerosol (POA), and eight secondary organic aerosol (SOA) species. A full description of the coupled model can be found in Riemer et al. (2009). PartMC-MOSAIC represents changes in particle composition by condensation and coagulation; we do not consider changes in

particle shape or aging by photochemical oxidation. PartMC Version 2.1.4 was used to generate the results in this paper. Simulations were performed at a time step of 60 seconds, with approximately 10^5 computational particles.

We simulated 288 plume scenarios, varying meteorological conditions, emissions of gases and particles, and the background number concentration, with further description given in Section 2.3. The atmospheric composition and environmental conditions differed between the scenarios, but the general structure of all simulations was the same. In each case, we simulated a well-mixed air parcel that is advected over and away from large urban area. All scenarios started at 06:00 LST, at which time the parcel contained only background gas and aerosol without any freshly emitted particles. During transport over the urban area, the parcel received gas and aerosol emissions from 06:00 LST until 18:00 LST, after which all emissions ceased. In these scenarios, we simulate a well-mixed boundary layer during the day, and the parcel is assumed to be in the residual layer at night. The temperature, mixing height, and relative humidity were held constant. Before discussing the full set of scenarios in Section 2.3, we show changes in CCN activity and the diurnal evolution of aging timescales in a baseline scenario. *For this baseline scenario, Table 1 outlines background aerosol number concentration, aerosol emission intensity, and size distribution and composition information for both background and emitted aerosols. The background concentration and emission intensity of gas-phase species are provided in Table 2.*

2.2 κ -Köhler model for computing CCN activity

We determined aging timescales from the particle-resolved results by tracking changes in CCN activity over two consecutive time steps. *A particle's ability to activate cloud formation depends on its dry diameter $D_{\text{dry},i}$ and its hygroscopicity parameter κ_i .* The equilibrium saturation ratio (S_i) over an aqueous droplet is computed through the κ -Köhler model (Köhler, 1936; Petters and Kreidenweis, 2007) as:

$$S_i(D_i) = \frac{D_i^3 - D_{\text{dry},i}^3}{D_i^3 - D_{\text{dry},i}^3(1 - \kappa_i)} \exp\left(\frac{4\sigma_w M_w}{RT\rho_w D_i}\right), \quad (2)$$

where σ_w is the surface tension of water, M_w is the molecular weight of water, R is the universal gas constant, T is the ambient temperature, ρ_w is the density of water, D_i is the particle wet diameter, $D_{\text{dry},i}$ is the particle dry diameter, and κ_i is the hygroscopicity parameter introduced by Petters and Kreidenweis (2007). All other factors being equal, particles with a greater κ_i are more hygroscopic and more easily activated. The parameter κ has been determined empirically for a number of aerosol species (Table 3), and the effective hygroscopicity parameter κ_i for each particle is the volume-weighted average of κ for its constituent aerosol species. We denote the critical saturation ratio at which a particle activates and forms a cloud droplet with $S_{c,i}$ and the critical supersaturation as $s_{c,i} = (S(D_{c,i}) - 1) \times 100$.

Figure 1 shows the two-dimensional number density distribution as a function of the particle dry diameter ($D_{\text{dry},i}$) and the particle hygroscopicity parameter (κ_i). Only particles containing BC are shown in this figure. In the scenarios presented in this study, all BC originated from diesel or gasoline exhaust. Choosing a certain environmental supersaturation threshold allows us to classify the particles as fresh or aged. For example, all particles to the left of the line for $s_{c,i} = 1\%$ are considered “fresh” for environmental supersaturations of 1% or lower, and all particles to the right of the line for $s_{c,i} = 1\%$ are considered “aged” at supersaturations above 1%.

The number distributions corresponding to fresh emissions, prior to any aging, are shown in Figure 1.a, and changes in the distribution during two time periods are shown in Figures 1.b and 1.c. Freshly emitted combustion particles are small and hydrophobic, with geometric mean diameter $D_{\text{dry,gm}} = 0.5 \mu\text{m}$ and with a hygroscopicity parameter of $\kappa = 3 \times 10^{-4}$ or $\kappa = 8 \times 10^{-4}$ for particles from diesel or gasoline, respectively. Therefore, *most* BC-containing particles are initially unable to activate at any *environmental supersaturation* $s < 1\%$ (lines in Figure 1.a). As $D_{\text{dry},i}$ and κ_i for individual particles increase by condensation and coagulation, their critical supersaturation $s_{c,i}$ for CCN-active decreases, shown by particles crossing the lines of constant critical supersaturation in Figure 1. Secondary aerosol forms through photochemical reactions during the day, causing rapid changes in particles' size and hygroscopicity. At night aging by condensation rates are slow, so coagulation is the dominant aging mechanism. This diurnal variation in aging rates is consistent with observations (Rose et al., 2011; Cheng et al., 2012). We define particles that “age” over a specific time period as those that transition from CCN-inactive to CCN-active, that is the particles that move from below a supersaturation line (CCN-inactive) at time t to above supersaturation line (CCN-active) at $t + \Delta t$.

2.3 CCN-based aging timescale

For the entire particle population, this change in the particle properties is quantified using the first-order aging timescale defined in Equation 1. Because the time period Δt is short relative to the timescale τ_{aging} , Equation 1 can be approximated as:

$$\left[\frac{dN_{\text{fresh}}}{dt} \right]_{\text{aging}} \approx - \frac{\Delta N_{\text{f} \rightarrow \text{a}}}{\Delta t}, \quad (3)$$

where $\Delta N_{\text{f} \rightarrow \text{a}}$ is the number of discrete particles that transition from fresh at time t to aged at time $t + \Delta t$, calculated from changes in the number of fresh particles. In this study, aging timescales are computed over a time step $\Delta t = 10$ minutes. Combining Equations 1 and 3, the aging timescale is

computed as:

$$\tau_{\text{aging}}(t, s) \approx \Delta t \frac{N_{\text{fresh}}(t, s)}{\Delta N_{\text{f} \rightarrow \text{a}}(t, t + \Delta t, s)}. \quad (4)$$

We refer to this aging timescale as the “bulk aging timescale” because it corresponds to the entire fresh particle population, and the term $\Delta N_{\text{f} \rightarrow \text{a}}$ includes all particles that transition from fresh to aged, regardless of their size. Later, we will introduce an analogous “size-resolved aging timescale”. Further details on the derivation of the bulk aging timescale, including number balances for all processes affecting aging, are given in Riemer et al. (2010).

The temporal evolution of the timescale is shown for the baseline scenario in Figure 2.a at $s = 0.1\%$, $s = 0.3\%$, and $s = 1\%$. *The aging timescale is a simple metric for quantifying the effects of changes in per-particle size and hygroscopicity that are shown in Figure 1*, and the gray shading in Figure 2.a corresponds to the time periods shown in Figure 1. Particles must become highly hygroscopic to activate into cloud droplets at low s (e.g. $s = 0.1\%$) but require less processing to become CCN at higher values of s (e.g. $s = 1\%$), so the aging timescale tends to be shorter for higher values of s .

Any particle that transitions from fresh at t to aged at $t + \Delta t$ does so either by coagulation with a large, hygroscopic particle or by accumulating sufficient condensing material to become hygroscopic. The overall aging timescale τ_{aging} can be represented as the combination of separate timescales for aging by condensation τ_{cond} and by coagulation τ_{coag} :

$$\frac{1}{\tau_{\text{aging}}(t, s)} = \frac{1}{\tau_{\text{cond}}(t, s)} + \frac{1}{\tau_{\text{coag}}(t, s)}. \quad (5)$$

The contribution of condensation and coagulation to the overall aging timescale is shown by separate timescales for aging by condensation (τ_{cond}) and coagulation (τ_{coag}) at $s = 0.3\%$ in Figure 2.b. We computed the coagulation and condensation aging timescales by counting the number of particles that transition from fresh to aged after participating in a coagulation event, $\Delta N_{\text{f} \rightarrow \text{a}, \text{coag}}$, or that age only by condensation, $\Delta N_{\text{f} \rightarrow \text{a}, \text{cond}}$. Then, we applied Equation 4 to find the corresponding condensation and coagulation aging timescales. Figure 2.b shows that the overall aging timescale is shortest during the day (e.g. 1 h at $s = 0.3\%$) due to rapid condensation of semi-volatile substances, and it is considerably longer at night (e.g. 24 h at $s = 0.3\%$), when coagulation is the dominant aging mechanism. *The temporal evolution of τ_{aging} and τ_{cond} are shown for multiple supersaturation levels in Riemer et al. (2010).*

3 Ensemble of particle-resolved model scenarios

The aging timescales shown in Figure 2 are limited to only one scenario, and aging rates vary with local conditions. For

example, the number concentration and size distribution of background particles affect *coagulation rates and, thereby, the coagulation aging timescale*. In order to identify the set of independent variables that best explain variance in BC’s aging timescale under a range of atmospheric conditions, we simulated aerosol dynamics in a series of plume scenarios and extracted aging timescales for each scenario. *As we will show*, the environmental properties that affect aerosol dynamics varied diurnally and differed between scenarios, causing the aging timescale to range from less than an hour (a large portion of particles age per time interval) to longer than a week (few particles age per time interval).

The input parameters that were varied between the scenarios are shown in Table 4. These input parameters were selected to produce a range of environmental conditions, consistent with observations described by Jimenez et al. (2009) and references therein. Simulations were performed using every combination of input parameters given in Table 4, leading to a total of 288 scenarios. *In each scenario, aerosol concentrations and particle characteristics varied throughout the 24-hour simulation. The conclusions in this study are based on these simulations of urban air masses.*

Figure 3 shows the distribution of aerosol mass concentration for selected aerosol species for all scenarios simulated with PartMC-MOSAIC (black lines) and corresponding ambient observations compiled by Jimenez et al. (2009) (vertical colored lines). The range of conditions simulated in the ensemble of scenarios is representative of the distribution in concentrations *observed in these urban areas. The distribution in the number concentration of all particles and of BC-containing particles are shown in Figure 4.a and 4.b, respectively. The size and composition of BC-containing particles also varied over the course of an individual simulation, as we showed in Figure 1. Variation in the geometric mean diameter and in the geometric mean hygroscopicity parameter of BC-containing particles across all simulations are shown in Figures 4.c and 4.d, respectively. Figures 3 and 4 were constructed from data at 10-minute intervals in each of the 288 simulations, corresponding to 41,000 time steps.*

Variance in the aging timescale is shown by the probability density distribution in Figure 5, which includes each 10-minute time interval in each of the 288 simulations. Distributions are shown for timescales computed at $s = 0.1\%$, $s = 0.3\%$, and $s = 1\%$. The supersaturation threshold s specifies the degree of change in particle properties required to classify a particle as aged, and timescales tend to decrease as s increases. In the following sections, we show that most of the variance in black carbon’s aging timescale at a specific s is explained by only a few key variables.

4 Nonparametric regression analysis to quantify explained variance

Black carbon’s aging timescale ranges from minutes to weeks (Figure 5), depending on local conditions and characteristics of BC-containing particles. We evaluated how well different

375 combinations of independent variables explain variance in
black carbon's aging timescale by comparing predictions of
aging timescales from nonparametric regression with exact
aging timescales from PartMC-MOSAIC. A nonparametric
regression was chosen, rather than a parametric regression,
380 because we do not know a priori the shape of the predictor
surface.

The procedure in applying a nonparametric regression is as
follows: 1) select a set of candidate independent variables to
test; 2) use most (90% of simulations) of the data as the train-
385 ing set to find the expected value of the aging timescale as a
function of the independent variables, as will be explained
below; 3) evaluate this expected aging timescale using the
rest of the data (10% of simulations), called the testing set.
The timescale from the regression is assessed by how well it
390 predicts the values of the aging timescale in the testing set,
represented by R^2 . The purpose of this exploration is to find
the independent variables that explain most of the variance in
the aging timescale, indicated by the largest value of R^2 .
To ensure that our conclusions did not depend on the choice of
395 scenarios, we repeated the analysis several times with ran-
domly chosen testing and training sets and verified that R^2
was insensitive to the specific choice of testing and training
sets.

4.1 Kernel density regression applied to particle- 400 resolved model data

Figure 6 shows how the regression analysis is applied in this
study. For all times in all simulations in the testing set, a
405 particle that is fresh at time t may age between t and some
later time $t + \Delta t$ or it may remain fresh over that time pe-
riod. Because these two events are mutually exclusive, this
aging behavior in PartMC-MOSAIC may be represented by
a binary variable $Y_{\text{age},j}(t, t + \Delta t, s)$, where $Y_{\text{age},j} = 1$ if the
particle ages between t and $t + \Delta t$ and $Y_{\text{age},j} = 0$ if it remains
fresh. The aging timescale at each model time step can then
410 be computed as the average of $Y_{\text{age},j}$ across all fresh BC-
containing particles:

$$\tau_{\text{aging}}(t, s) = \Delta t \frac{N_{p,\text{fresh}}}{\sum_{i=1}^{N_{p,\text{fresh}}} Y_{\text{age},j}(t, t + \Delta t)}, \quad (6)$$

which is equivalent to Equation 4, computed from $N_{p,\text{fresh}}$ in-
dividual particles over a specific model time step.

415 Alternatively, the expected probability that a fresh parti-
cle will age, given its characteristics or the aging conditions
that it experiences, can be estimated from a nonparametric
regression. We applied the kernel density regression intro-
duced by Watson (1964) and Nadaraya (1964). The expected
420 value of $Y_{\text{age},j}$ for a specific particle in the testing set is pre-
dicted using the kernel density regression, using information
about the candidate variable x_j only. The candidate variable
 x_j may be a particle-level characteristic, which varies be-
tween particles and, for a specific particle, varies over time
425 (e.g. particle wet diameter). The candidate variable x_i may

also be a characteristic of the environment, which varies over
time but, at a specific time, is the same for all particles (e.g.
aerosol number concentration). All candidate variables ex-
plored in this study are outlined in Table 5. In this section,
we show how the nonparametric regression can be applied
to evaluate variance explained by a single candidate variable
at a time. Later, we show how this analysis can be extended
to evaluate combinations of independent variables.

At each time step in each simulation of the testing set,
the expected value of $Y_{\text{age}}(t, t + \Delta t, s)$ for each particle
was computed as a weighted average of $Y_{\text{age},i}(t, t + \Delta t, s)$
for millions of individual particles in the training set. Val-
ues for $Y_{\text{age},i}(t, t + \Delta t, s)$ in the training set are weighted
according to the kernel function $K_h(x - x_i)$, where x is
the independent variable of interest. The expected value of
 $Y_{\text{age}}(t, t + \Delta t, s)$ is given by:

$$E[Y_{\text{age}}|x_j(t), \Delta t, s] = \frac{\sum_{i=1}^{N_{p,\text{train}}} K_h(x_j(t) - x_i) Y_{\text{age},i}}{\sum_{i=1}^{N_{p,\text{train}}} K_h(x_j(t) - x_i)}, \quad (7)$$

where x_i is the value of the independent variable for each
particle in the training set, x_j is the value of the independent
variable for the target variable in the testing set, and $N_{p,\text{train}}$
is all particles in the training set, including all time steps in
all simulations.

Analogous to Equation 6, the expected value of the aging
timescale at a specific model time step, $\hat{\tau}_{\text{aging}}(t, s)$, is then
computed as the average of $E[Y_{\text{age}}|x_j(t), \Delta t, s]$ across all
BC-containing particles:

$$\hat{\tau}_{\text{aging}}(t, s) = \Delta t \frac{N_{p,\text{fresh}}}{\sum_{j=1}^{N_{p,\text{fresh}}} E[Y_{\text{age}}|x_j(t), \Delta t, s]}. \quad (8)$$

In Section 4.1, we provide further explanation on the in-
clusion of particle-level characteristics in the prediction of
 $\hat{\tau}_{\text{aging}}$.

In this study we used a Gaussian kernel function with stan-
dard deviation h :

$$K_h(x_j(t) - x_i) = \frac{1}{\sqrt{2\pi}h} \exp\left(-\frac{(x_j(t) - x_i)^2}{2h^2}\right). \quad (9)$$

The kernel function $K_h(x_j(t) - x_i)$ defines the weight ap-
plied to each model timescale $\tau_{\text{age},i}$ to compute the expected
timescale $\hat{\tau}_{\text{age}}$, such that timescales for conditions similar to
the conditions of the target point are weighted most heav-
ily in the regression. The regression function predicted by
the kernel regression depends on the prescribed value for h ,
where larger h results in smoother regression functions. We
applied Silverman's rule of thumb to select the value for h
(Silverman, 1986), such that h depends on the number of
independent variables, the standard deviation of each inde-
pendent variable, and the total number of data points in the
testing set.

If the candidate variable x strongly affects the value of the aging timescale, the expected aging timescale $\tau_{\text{aging},j}$ will accurately represent the actual aging timescale $\tau_{\text{aging},j}$, assuming a suitable kernel bandwidth h is applied in the kernel regression. Aging rates scale with the inverse of the aging timescale, so we quantified the variance explained by the regression function, R^2 , in terms of $1/\tau_{\text{aging}}$:

$$R^2(s) = 1 - \frac{\sum_{j=1}^{n_{\text{test}}} N_{\text{fresh},j}(s) \left(\frac{1}{\hat{\tau}_{\text{aging},j}(s)} - \frac{1}{\tau_{\text{aging},j}(s)} \right)^2}{\sum_{j=1}^{n_{\text{test}}} N_{\text{fresh},j}(s) \left(\frac{1}{\hat{\tau}_{\text{aging},j}(s)} - \frac{1}{\bar{\tau}_{\text{aging}}(s)} \right)^2}, \quad (10)$$

where $\tau_{\text{aging},j}$ is the timescale from PartMC-MOSAIC for each data point in the testing set, $\hat{\tau}_{\text{aging},j}$ is the expected timescale from the regression for each data point in the testing set, $\bar{\tau}_{\text{aging}}$ is the harmonic mean of the aging timescales across all data points in the testing set, and N_{test} is the number of data points in the testing set, where the data points include all time steps in all scenarios.

4.2 Inclusion of particle-level variables in the kernel regression

To illustrate our approach for including particle-level variables, we demonstrate the regression procedure using the wet diameter as the independent variable x . The resulting regression surface is a size-dependent timescale, which gives insight into the importance of aging processes as a function of particle wet diameter. For a given set of environmental conditions, some particles are more likely to age than others, and we find that a particles' tendency to age depends on their characteristics just prior to the aging period. We evaluated how aging rates vary with a number of per-particle characteristics, such as particles' diameter at emission, their dry diameter at the time when aging is evaluated, or their hygroscopicity parameter when aging is evaluated. We found that for given environmental conditions, per-particle aging rates were most correlated with the wet diameter of fresh (CCN-inactive) particles; that is, values of R^2 were greatest for regression functions that included the time-varying wet size distribution of fresh BC-containing particles.

It is therefore useful to introduce a size-resolved aging timescale that accounts for differences in aging rates between particles of different sizes. Size-resolved aging timescales were computed at each time t and supersaturation s using the kernel regression described in Section 4.1. The expected value of Y_{age} for a particle in the testing set with wet diameter $D_j(t)$ was computed as the weighted average of $Y_{\text{age},i}$ for rest particles the training set $i = 1, \dots, N_{\text{p,train}}$, computed at a specific t and s :

for a particle in the testing set of wet diameter D_j was

$$E[Y_{\text{age}}|D_j(t), \Delta t, s] = \frac{\sum_{i=1}^{N_{\text{p,train}}} K_{h_D}(D_j(t) - D_i) Y_{\text{age},i}}{\sum_{i=1}^{N_{\text{p,train}}} K_{h_D}(D_j(t) - D_i)}, \quad (11)$$

where the kernel weighting function $K_h(D_j(t) - D_i)$ is a Gaussian (Equation 9), such that fresh particles with D_i similar to the target diameter D_j are weighted most heavily in the regression. The size-resolved aging timescale can also be defined for a continuous size distribution of fresh particles $n_{\text{fresh}}(t, D, s)$. Similar to Equation 8, the size-resolved aging timescale, $\tau_{\text{aging}}(D)$, is computed as a function of $E[Y_{\text{age}}|D]$ and the time step Δt :

$$\tau_{\text{aging}}(t, D, s) = \frac{\Delta t}{E[Y_{\text{age}}|D(t), \Delta t, s]}. \quad (12)$$

For a particle-resolved population of fresh particles $j = 1, \dots, N_{\text{p,fresh}}$, where each particle has a unique wet diameter D_j , the bulk, population-level aging timescale can be estimated as the average of $E[Y_{\text{age}}|D_j(t), \Delta t, s]$ across all $N_{\text{p,fresh}}$ particles, as given in Equation 8 using $x_j = D_j$. Equivalently, the population-level aging timescale can be computed through the average of the continuous size-resolved aging timescale $\tau_{\text{aging}}(D, t, s)$, weighted by the size distribution of fresh particles $n_{\text{fresh}}(D, t, s)$: integration of the continuous size-resolved aging timescale $\tau_{\text{aging}}(D, t, s)$ over all D , continuous size distribution of fresh particle $n_{\text{fresh}}(D, t, s)$ as the average of $\tau_{\text{aging}}(D, t, s)$ across all D , weighted by the size distribution of fresh particles $n_{\text{fresh}}(D, t, s)$:

$$\tau_{\text{aging}}(t, s)^{-1} = \frac{\int_0^\infty \tau_{\text{aging}}(t, D, s)^{-1} n_{\text{fresh}}(t, D, s) dD}{\int_0^\infty n_{\text{fresh}}(t, D, s) dD}. \quad (13)$$

By this relationship, the bulk aging timescale under a specific set of environmental conditions also varies with the size distribution of CCN-inactive (fresh) BC.

The temporal evolution of the size-resolved aging timescale is shown for the baseline scenario in the middle column of Figure 7 for $s = 0.3\%$. The contributions of coagulation (Figure 7.b) and condensation (Figure 7.c) to the overall aging timescale (Figure 7.a) are shown by the separate size-resolved timescales for each process. The dominant mechanism driving the aging timescale depends on the time of day and the particle size.

A comparison between Figures 7.a and 7.c shows that condensation was the dominant process driving diurnal variation in the size-resolved aging timescale. This diurnal pattern in condensation aging conditions is reflected in the bulk aging timescale shown in Figure 2. The bulk condensation aging timescale was shorter than 4 h during the day for this scenario, and this was the dominant process affecting aging

560 rates at this time. However, Figure 7.c shows that these rapid transitions from CCN-inactive to CCN-active occurred only for the largest ($D > 50$ nm) fresh particles, although condensation also caused an increase in D for smaller fresh particles. The coagulation aging timescale, on the other hand, was short for the smallest fresh particles and varied only slightly over the course of the simulation.

4.3 Combining particle-level and population-level variables in kernel regression

570 *In this study, we performed a series of multivariate kernel regressions to identify the combination of independent variables that best explain variance in black carbon's aging timescale. In many cases, we extracted aging timescales that depend both on characteristics of individual particles, such as D , and on properties of entire particle populations or the environment, such as the overall aerosol number concentration N . One advantage of this approach is that both particle-level variables and population-level variables can be included in the prediction of $Y_{age,j}$. For example, the expected value of $Y_{age,j}$ for a particle with diameter D_j that is exposed to an aerosol number concentration N_j is computed with the bivariate kernel regression:*

$$E[Y_{age}|D_j, N_j] = \frac{\sum_{i=1}^{N_{p,train}} K_{h_D}(D_j(t) - D_i)K_{h_f}(N_j(t) - N_i)Y_{age,i}}{\sum_{i=1}^{N_{p,train}} K_{h_D}(D_j(t) - D_i)K_{h_f}(N_j(t) - N_i)}. \quad (14)$$

585 *The overall aging timescale for a particular size distribution exposed to a specific number concentration is the computed as the sum across individual particles (Equation 8) or, equivalently, by integrating over the size distribution (Equation 13). Equation 14 can easily be generalized to three or more independent variables.*

5 Independent variables that best explain variance in aging timescales

595 *We found that most variance in the aging timescale is explained by only a few independent variables. Explained variance R^2 is shown different combinations of independent variables as a function of the environmental supersaturation s at which CCN activity is evaluated. For all supersaturation levels, 90% of variance in the coagulation aging timescale (Figure 8.a) was explained by regression predictions that included the size distribution of fresh BC-containing particles and the number concentration of large, CCN-active particles ($N_{CCN,large}$). Three variables were needed to explain 85% of variance in the condensation aging timescale (Figure 8.b): the size distribution of fresh BC ($n_{fresh}(D)$); the flux of secondary aerosol (\dot{f}_{cond}), defined as the volume condensation rate of semi-volatile substances per particle surface area density; and the effective hygroscopicity parameter of secondary aerosol (κ_{cond}), where κ_{cond} is the*

volume-weighted average of κ for condensing semi-volatile species. The size distribution of fresh BC was included in each case by determining a regression for the size-resolved aging timescale before computing the bulk aging timescale according to Equation 13. Only 10-15% of variance remains unexplained, indicating that variables other than $n_{fresh}(D)$, \dot{f}_{cond} , κ_{cond} , and $N_{CCN,large}$ also weakly affect the value of the aging timescale.

Figure 8 shows the explained variance R^2 as a function of s for the independent variables that best explain variance in the coagulation and condensation aging timescales. Approximately 90% of variance in the coagulation aging timescale was explained by regressions in terms of $n_{fresh}(D)$ and $N_{CCN,large}$ (black line of Figure 8.a). Brownian coagulation events are most likely to occur between large and small particles, so the coagulation aging timescale decreases when there are more particles that are CCN-active and are also large enough to be good coagulation partners. The smallest fresh particles are likely to coagulate with large background particles, where we found the threshold for “large” to be $D > 100$ nm by identifying the threshold that resulted in the highest R^2 . A regression computed in terms of the number concentration of large particles (green line of Figure 8.a), rather than the number concentration of large and CCN-active particles, gave $R^2 \approx 85\%$ at high supersaturation thresholds ($s > 0.8\%$) but $R^2 < 10\%$ at low supersaturation thresholds ($s < 0.1\%$). This is because not all particles with $D > 100$ nm are CCN-active at this low s , but nearly all particles that are CCN-active at this low s have $D > 100$ nm. On the other hand, if the independent variable was the number concentration of CCN-active particles (blue line of Figure 8.a), rather than the number concentration of large and CCN-active particles, $R^2 \approx 90\%$ for timescales at low supersaturation thresholds ($s < 0.1\%$) and $R^2 \approx 70\%$ at high supersaturation thresholds ($s > 0.8\%$). *Only by considering the number concentration of particles that are both CCN-active and large, $N_{CCN,large}$ were we able to explain variance in the coagulation aging timescale at all supersaturation levels.* If the size distribution of fresh BC was neglected, R^2 ranged from 40% to 60%, depending on the supersaturation threshold (yellow line of Figure 8.a).

While the expected aging timescale computed in terms of $n_{fresh}(D)$, \dot{f}_{cond} , and κ_{cond} (black line of Figure 8.b) explained greater than 80% of variance in the condensation aging timescale, R^2 was less than 60% for regressions that did not include $n_{fresh}(D)$ (red line). Only 10-30% of variance was explained if κ_{cond} was not included in the regression (grey line). If \dot{f}_{cond} was not included, $R^2 \approx 0\%$ for all s , regardless of the other variables included in the regression (not shown). This suggests, not surprisingly, that the condensation rate is the key variable driving aging by condensation, but the condensation aging timescale also depends strongly on the hygroscopicity of condensing aerosol κ_{cond} and on the size distribution of fresh particles $n_{fresh}(D)$.

5.1 Aging timescales as a function of governing parameters

The regression surfaces $\hat{\tau}_{\text{coag}}(N_{\text{CCN,large}}, D)$ and $\hat{\tau}_{\text{cond}}(\dot{f}_{\text{cond}}, \kappa_{\text{cond}}, D)$ are shown in Figures 9.a and 9.b, respectively. Figure 9.a shows that timescales for aging by coagulation range from hours to weeks. The coagulation aging timescale decreases with the number concentration of “large”, CCN-active particles ($N_{\text{CCN,large}}$) and, for a given $N_{\text{CCN,large}}$, small BC-containing particles are more likely to age by coagulation than large BC-containing particles. On the other hand, condensation aging timescales are shortest for the largest fresh particles and, for these particles, the condensation aging timescale tends to decrease as \dot{f}_{cond} or κ_{cond} increase. The two panels of Figure 9.b show $\hat{\tau}_{\text{cond}}$ as a function of \dot{f}_{cond} and D for secondary aerosol with differing hygroscopicity, $\kappa_{\text{cond}} = 0.65$ on the left, representing secondary inorganic aerosol, and $\kappa_{\text{cond}} = 0.1$ on the right, representing secondary organic aerosol.

5.2 Sensitivity of aging timescale to aging conditions

In this section we apply the regression surfaces shown in Figure 9 to selected example cases to demonstrate how aging conditions and the fresh particle size distribution affect particle aging rates. Figure 10 shows how aging rates by condensation and coagulation can be reconstructed as a function of the size distribution of fresh particles (Figure 10.a) and the size-resolved aging timescale (Figure 10.b). We compare lognormal size distributions with geometric mean diameter (D_{gm}) of 30 nm (dashed line of Figure 10.a) and 60 nm (solid line of Figure 10.a). Timescales were computed for limiting environmental conditions, indicated by line colors in Figure 10.b): slow coagulation aging ($N_{\text{CCN,large}} = 500 \text{ cm}^{-3}$, red line) or fast coagulation aging ($N_{\text{CCN,large}} = 10,000 \text{ cm}^{-3}$, blue line) and slow condensation aging ($\dot{f}_{\text{cond}} = 0.01 \text{ nm h}^{-1}$, green line) or fast condensation aging ($\dot{f}_{\text{cond}} = 1 \text{ nm h}^{-1}$, yellow line). Size-resolved aging timescales are taken from the regression surfaces in Figure 9 for these values of \dot{f}_{cond} and $N_{\text{CCN,large}}$, assuming $\kappa_{\text{cond}} = 0.65$ in both cases. Then, the rate at which particles of a given size transition from fresh to aged (Figure 10.c) is computed as the product of $n_f(D)$ and $1/\tau_{\text{age}}(D)$. Figure 10.c shows aging rates for particle distributions with $D_{\text{gm}} = 30 \text{ nm}$ (dashed lines) and $D_{\text{gm}} = 60 \text{ nm}$ (solid lines) under these limiting environmental conditions that promote rapid (blue lines) or slow (red lines) aging by coagulation and rapid (yellow lines) or slow (green lines) aging by condensation.

For these two size distributions (Figure 10.a) and different combinations of aging conditions (Figure 10.b), bulk aging timescales at $s = 0.3\%$ were computed according to Equation 13, and the results are given in Table 6. The combinations of environmental conditions are as follows: 1) rapid condensation aging (yellow lines in Figure 10.b) and slow

coagulation aging (red lines), 2) slow condensation aging (green lines) and rapid coagulation aging (blue lines), 3) slow aging by both condensation and coagulation, and 4) rapid aging by both condensation and coagulation.

The sensitivity of the bulk aging timescales to \dot{f}_{cond} and $N_{\text{CCN,large}}$ depends strongly on the environmental supersaturation s , as shown in Figure 11. At each supersaturation, sensitivities are quantified as a logarithmic derivative, or relative change in τ_{age} to a relative change in \dot{f}_{cond} or $N_{\text{CCN,large}}$. Negative values of this metric indicate that increasing $N_{\text{CCN,large}}$ or \dot{f}_{cond} corresponds to a decrease in τ_{aging} .

While τ_{aging} is most sensitive to $N_{\text{CCN,large}}$ at low supersaturation levels, τ_{aging} shows the greatest sensitivity to \dot{f}_{cond} at high supersaturation levels. At low supersaturation levels, τ_{aging} is insensitive to \dot{f}_{cond} if the distribution contains a higher fraction of small particles ($D_{\text{gm}} = 30 \text{ nm}$), regardless of conditions for aging by coagulation. If particles are large ($D_{\text{gm}} = 60 \text{ nm}$), τ_{aging} at this low s is τ_{aging} sensitive to \dot{f}_{cond} only under conditions of slow aging by coagulation. At $s = 1\%$, τ_{aging} is sensitive to \dot{f}_{cond} in all cases, regardless of the fresh particle size distribution or conditions for aging by coagulation. Coagulation aging is relatively more important at low supersaturation compared to high supersaturation thresholds. Consistent with this fact, Figure 11a shows that the sensitivity of the aging time scale to $N_{\text{CCN,large}}$ generally decreases as s increases. The magnitude of the condensational flux \dot{f}_{cond} impacts the sensitivity towards $N_{\text{CCN,large}}$. Environments with lower \dot{f}_{cond} result in a larger sensitivity to $N_{\text{CCN,large}}$.

6 Discussion

Global models that employ first-order aging models assume a fixed timescale of 1–3 days, but observations show that aging timescales can be as short as a few hours in polluted areas (Zhang et al., 2008). Other modeling studies have suggested parameterizations that account for this variation in aging conditions. Riemer et al. (2004) evaluated aging timescales in a mesoscale model and parameterized timescales for aging by coagulation as a function of the overall number concentration. Pierce et al. (2009) developed an analytical expression that accounts for decreases in the number concentration of primary aerosol through coagulation events; for emitted particles of a specific size, the coagulation loss rate was computed by integrating the coagulation kernel over the entire background size distribution. However, the regression analysis applied in the current study reveals that 90% of the variance in the coagulation aging timescale can be explained using a relatively simple representation of the background size distribution. We showed that the variation in the size-resolved aging timescales can be attributed to the number concentration of particles that are both large ($D > 100 \text{ nm}$)

and CCN-active. Other characteristics of the background size distributions are not needed.

Oshima and Koike (2013) developed a parameterization of the condensation aging timescale based on results from a box model, and, similar to the present study, computed aging timescales based on changes in CCN activity. However, unlike the present study, Oshima and Koike (2013) did not consider differences in the hygroscopic properties of the condensing material, and their aging timescale varied with the mass condensation rate per total BC mass concentration. In contrast, the regression analysis in the present study reveals that the volume condensation rate per overall aerosol surface area is the variable that best explains variance in BC's condensation aging timescale, which is consistent with laboratory studies (Zhang et al., 2008; Khalizov et al., 2009). The present work also differs from Oshima and Koike (2013) in the representation of the aerosol size distribution. Whereas Oshima and Koike (2013) parameterized bulk aging timescales for lognormal size distributions, we presented a size-resolved aging timescale that can be applied to any arbitrary size distribution.

As in all relationships for BC's aging timescale, the value of the aging timescale depends strongly on the criterion used to distinguish fresh and aged particles. Particle activation at a specific environmental supersaturation is the aging criterion applied in this study, representing changes in particle characteristics that most affect their susceptibility to wet deposition. Table 6 shows that the value of the aging timescale depends strongly on the criterion supersaturation at which CCN activation was evaluated, consistent with Riemer et al. (2010) and Oshima and Koike (2013). Further, the relative importance of condensation versus coagulation as aging processes also depends on the supersaturation threshold.

7 Conclusions

This study identifies the minimal set of independent variables needed to explain variance in black carbon's aging timescale. We simulated the evolution of gases and aerosols in a series of urban scenarios with the particle-resolved model PartMC-MOSAIC and extracted time-dependent aging timescales based on the rate at which individual particles transition from CCN-inactive to CCN-active at a specified environmental supersaturation. The value of the aging timescale spanned orders of magnitude, depending on local environmental conditions and the supersaturation threshold at which CCN activity was evaluated. Aging timescales were shorter than an hour under conditions of rapid secondary aerosol formation, but on the order of days in the absence of secondary aerosol precursors. Condensation aging timescales exhibited more variation than coagulation aging timescales, and the relative importance of each aging mechanism depended on the size range of particles to be aged. We performed a non-parametric regression analysis on model data from 288 scenarios in or-

der to identify the independent variables with which aging timescales are best correlated and quantified the portion of variance explained by regressions in terms of these variables. *This paper is the groundwork for the development of aging parameterizations suitable for use in global models.*

To our knowledge, this is the first study to apply a regression analysis to identify the minimal set of parameters needed to explain variance in black carbon aging rates. After evaluating a number of independent variables, we found that the flux of secondary aerosol, the hygroscopicity of secondary aerosol, and the size distribution of CCN-inactive (fresh) BC-containing particles were the minimal set of parameters needed to explain 80% of variance in the condensation aging timescale. On the other hand, 90% of variance in the coagulation aging timescale was explained by only two variables: the size distribution of fresh BC-containing particles and the number concentration of particles that are both large ($D > 100$ nm) and CCN-active. This work distills the complex interactions captured by the particle-resolved model to a few input variables, all of which are tracked by existing global climate models, and is a first step toward developing physically-based parameterizations of aerosol aging.

References

- Ackerman, A., Toon, O., Stevens, D., Heymsfield, A., Ramanathan, V., and Welton, E. (2000). Reduction of tropical cloudiness by soot. *Science*, 288(5468):1042–1047.
- Albrecht, B. (1989). Aerosols, cloud microphysics, and fractional cloudiness. *Science*, 245(4923):1227–1230.
- Aquila, V., Hendricks, J., Lauer, A., Riemer, N., Vogel, H., Baumgardner, D., Minikin, A., Petzold, A., Schwarz, J., Spackman, J., et al. (2011). Made-in: a new aerosol microphysics submodel for global simulation of insoluble particles and their mixing state. *Geoscientific Model Development*, 4(2):325–355.
- Bauer, S., Wright, D., Koch, D., Lewis, E., McGraw, R., Chang, L., Schwartz, S., and Ruedy, R. (2008). MATRIX (multiconfiguration Aerosol TRacker of mIXing state): an aerosol microphysical module for global atmospheric models. *Atmospheric Chemistry and Physics*, 8(20):6003–6035.
- Binkowski, F. and Roselle, S. (2003). Models-3 community multiscale air quality (CMAQ) model aerosol component: 1. Model description. *Journal of Geophysical Research*, 108(4183):15.
- Cantrell, W., Shaw, G., Cass, G., Chowdhury, Z., Hughes, L., Prather, K., Guazzotti, S., and Coffee, K. (2001). Closure between aerosol particles and cloud condensation nuclei at Kaashidhoo Climate Observatory. *Journal of Geophysical Research*, 106:28.
- Cheng, Y., Su, H., Rose, D., Gunthe, S., Berghof, M., Wehner, B., Achtert, P., Nowak, A., Takegawa, N., Kondo, Y., et al. (2012). Size-resolved measurement of the mixing state of soot in the megacity Beijing, China: diurnal cycle, aging and parameterization. *Atmospheric Chemistry and Physics*, 12(10):4477–4491.
- Chin, M., Diehl, T., Tan, Q., Prospero, J., Kahn, R., Remer, L., Yu, H., Sayer, A., Bian, H., Geogdzhayev, I., et al. (2014). Multi-decadal aerosol variations from 1980 to 2009: a perspective from

- observations and a global model. *Atmospheric Chemistry and Physics*, 14(7):3657–3690.
- Clegg, S., Brimblecombe, P., and Wexler, A. (1998). Thermodynamic model of the system H^+ - NH_4^+ - SO_4^{2-} - NO_3^- - H_2O at tropospheric temperatures. *The Journal of Physical Chemistry A*, 102(12):2137–2154.
- Cooke, W. and Wilson, J. (1996). A global black carbon aerosol model. *Journal of Geophysical Research*, 101(D14):19395–19.
- Cozic, J., Verheggen, B., Mertes, S., Connolly, P., Bower, K., Petzold, A., Baltensperger, U., and Weingartner, E. (2007). Scavenging of black carbon in mixed phase clouds at the high alpine site jungfrauoch. *Atmospheric Chemistry and Physics*, 7(7):1797–1807.
- Croft, B., Lohmann, U., Von Salzen, K., et al. (2005). Black carbon ageing in the canadian centre for climate modelling and analysis atmospheric general circulation model. *Atmos. Chem. Phys.*, 5(7):1931–1949.
- Furutani, H., Dall’osto, M., Roberts, G., and Prather, K. (2008). Assessment of the relative importance of atmospheric aging on CCN activity derived from field observations. *Atmospheric Environment*, 42(13):3130–3142.
- Hansen, J. and Nazarenko, L. (2004). Soot climate forcing via snow and ice albedos. *Proceedings of the National Academy of Sciences of the United States of America*, 101(2):423.
- Jacobson, M. (1997). Development and application of a new air pollution modeling system—II. Aerosol module structure and design. *Atmospheric Environment*, 31(2):131–144.
- Jacobson, M. Z. (2002). Control of fossil-fuel particulate black carbon and organic matter, possibly the most effective method of slowing global warming. *Journal of Geophysical Research: Atmospheres (1984–2012)*, 107(D19):ACH–16.
- Jacobson, M. Z. (2004). Climate response of fossil fuel and biofuel soot, accounting for soot’s feedback to snow and sea ice albedo and emissivity. *Journal of Geophysical Research: Atmospheres (1984–2012)*, 109(D21).
- Jimenez, J., Canagaratna, M., Donahue, N., Prevot, A., Zhang, Q., Kroll, J. H., DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N., et al. (2009). Evolution of organic aerosols in the atmosphere. *Science*, 326(5959):1525–1529.
- Jo, D., Park, R., Kim, M., and Spracklen, D. (2013). Effects of chemical aging on global secondary organic aerosol using the volatility basis set approach. *Atmospheric Environment*, 81:230–244.
- Johnson, K., Zuberi, B., Molina, L., Molina, M., Iedema, M., Cowin, J., Gaspar, D., Wang, C., and Laskin, A. (2005). Processing of soot in an urban environment: case study from the Mexico City Metropolitan Area. *Atmospheric Chemistry and Physics*, 5:3033–3043.
- Khalizov, A. F., Zhang, R., Zhang, D., Xue, H., Pagels, J., and McMurry, P. H. (2009). Formation of highly hygroscopic soot aerosols upon internal mixing with sulfuric acid vapor. *Journal of Geophysical Research: Atmospheres (1984–2012)*, 114(D5).
- Koch, D. (2001). Transport and direct radiative forcing of carbonaceous and sulfate aerosols in the giss gcm. *Journal of Geophysical Research: Atmospheres (1984–2012)*, 106(D17):20311–20332.
- Koch, D., Schulz, M., Kinne, S., McNaughton, C., Spackman, J. R., Balkanski, Y., Bauer, S., Bernsten, T., Bond, T. C., Boucher, O., Chin, M., Clarke, A., De Luca, N., Dentener, F., Diehl, T., Dubovik, O., Easter, R., Fahey, D. W., Feichter, J., Fillmore, D., Freitag, S., Ghan, S., Ginoux, P., Gong, S., Horowitz, L., Iversen, T., Kirkevåg, A., Klimont, Z., Kondo, Y., Krol, M., Liu, X., Miller, R., Montanaro, V., Moteki, N., Myhre, G., Penner, J. E., Perlwitz, J., Pitari, G., Reddy, S., Sahu, L., Sakamoto, H., Schuster, G., Schwarz, J. P., Seland, Ø., Stier, P., Takegawa, N., Takemura, T., Textor, C., van Aardenne, J. A., and Zhao, Y. (2009). Evaluation of black carbon estimations in global aerosol models. *Atmospheric Chemistry and Physics*, 9(22):9001–9026.
- Köhler, H. (1936). The nucleus in and the growth of hygroscopic droplets. *Transactions of the Faraday Society*, 32:1152–1161.
- Liu, J., Fan, S., Horowitz, L. W., and Levy, H. (2011). Evaluation of factors controlling long-range transport of black carbon to the arctic. *Journal of Geophysical Research: Atmospheres (1984–2012)*, 116(D4).
- Lohmann, U., Feichter, J., et al. (2005). Global indirect aerosol effects: a review. *Atmospheric Chemistry and Physics*, 5(3):715–737.
- Maricq, M. M. (2007). Chemical characterization of particulate emissions from diesel engines: a review. *Journal of Aerosol Science*, 38(11):1079–1118.
- Matsui, H., Koike, M., Kondo, Y., Moteki, N., Fast, J. D., and Zaveri, R. A. (2013). Development and validation of a black carbon mixing state resolved three-dimensional model: Aging processes and radiative impact. *Journal of Geophysical Research: Atmospheres*, 118(5):2304–2326.
- McCormick, R. A. and Ludwig, J. H. (1967). Climate modification by atmospheric aerosols. *Science*, 156(3780):1358–1359.
- McGraw, R. (1997). Description of aerosol dynamics by the quadrature method of moments. *Aerosol Science and Technology*, 27(2):255–265.
- Nadaraya, E. A. (1964). On estimating regression. *Theory of Probability & Its Applications*, 9(1):141–142.
- Oshima, N. and Koike, M. (2013). Development of a parameterization of black carbon aging for use in general circulation models. *Geoscientific Model Development*, 6(2):263–282.
- Oshima, N., Koike, M., Zhang, Y., and Kondo, Y. (2009). Aging of black carbon in outflow from anthropogenic sources using a mixing state resolved model: 2. aerosol optical properties and cloud condensation nuclei activities. *Journal of Geophysical Research: Atmospheres (1984–2012)*, 114(D18).
- Petters, M., Prenni, A., Kreidenweis, S., DeMott, P., Matsunaga, A., Lim, Y., and Ziemann, P. (2006). Chemical aging and the hydrophobic-to-hydrophilic conversion of carbonaceous aerosol. *Geophysical Research Letters*, 33(24):L24806.
- Petters, M. D. and Kreidenweis, S. M. (2007). A single parameter representation of hygroscopic growth and cloud condensation nucleus activity. *Atmospheric Chemistry and Physics*, 7(8):1961–1971.
- Pierce, J., Theodoritsi, G., Adams, P., and Pandis, S. (2009). Parameterization of the effect of sub-grid scale aerosol dynamics on aerosol number emission rates. *Journal of Aerosol Science*, 40(5):385–393.
- Prenni, A., Petters, M., Kreidenweis, S., DeMott, P., and Ziemann, P. (2007). Cloud droplet activation of secondary organic aerosol. *Journal of Geophysical Research*, 112(D10):10223.
- Riemer, N., Vogel, H., and Vogel, B. (2004). Soot aging time scales in polluted regions during day and night. *Atmospheric Chemistry and Physics*, 4(7):1885–1893.

- 990 Riemer, N., West, M., Zaveri, R., and Easter, R. (2009). Simulating
the evolution of soot mixing state with a particle-resolved aerosol
model. *Journal of Geophysical Research*, 114(D9):D09202. 1050
- Riemer, N., West, M., Zaveri, R., and Easter, R. (2010). Estimating
black carbon aging time-scales with a particle-resolved aerosol
995 model. *Journal of Aerosol Science*, 41(1):143–158.
- Rose, D., Gunthe, S., Su, H., Garland, R., Yang, H., Berghof, M.,
Cheng, Y., Wehner, B., Achtert, P., Nowak, A., et al. (2011). 1055
Cloud condensation nuclei in polluted air and biomass burning
smoke near the mega-city guangzhou, china—part 2: Size-
resolved aerosol chemical composition, diurnal cycles, and ex-
ternally mixed weakly ccn-active soot particles. *Atmospheric
Chemistry and Physics*, 11(6):2817–2836. 1060
- Rosen, H., Hansen, A., Gundel, L., and Novakov, T. (1978). Identifi-
cation of the optically absorbing component in urban aerosols.
1000 *Applied Optics*, 17(24):3859–3861.
- Schell, B., Ackermann, I. J., Hass, H., Binkowski, F. S., and Ebel,
A. (2001). Modeling the formation of secondary organic aerosols 1065
within a comprehensive air quality model system. *Journal of
Geophysical Research. D. Atmospheres*, 106:28.
- 1010 Schmidt, G. A., Kelley, M., Nazarenko, L., Ruedy, R., Russell,
G. L., Aleinov, I., Bauer, M., Bauer, S. E., Bhat, M. K., Bleck,
R., et al. (2014). Configuration and assessment of the giss mod-1070
ele2 contributions to the cmip5 archive. *Journal of Advances in
Modeling Earth Systems*.
- 1015 Schulz, M., Textor, C., Kinne, S., Balkanski, Y., Bauer, S., Berntsen,
T., Berglen, T., Boucher, O., Dentener, F., Guibert, S., et al.
(2006). Radiative forcing by aerosols as derived from the ae-
rocom present-day and pre-industrial simulations. *Atmos. Chem.
Phys.*, 6(12):5225–5246. 1075
- 1020 Silverman, B. (1986). *Density estimation for statistics and data
analysis*, volume 26. Chapman & Hall/CRC.
- Svenningsson, B., Rissler, J., Swietlicki, E., Mircea, M., Bilde,
M., Facchini, M., Decesari, S., Fuzzi, S., Zhou, J., Mønster, J.,
et al. (2006). Hygroscopic growth and critical supersaturations
1025 for mixed aerosol particles of inorganic and organic compounds
of atmospheric relevance. *Atmospheric Chemistry and Physics*,
6(7):1937–1952.
- Twomey, S. (1977). The influence of pollution on the shortwave
albedo of clouds. *Journal of the Atmospheric Sciences*, 34:1149–
1030 1152.
- Twomey, S., Piepgrass, M., and Wolfe, T. (1984). An assessment
of the impact of pollution on global cloud albedo. *Tellus B*,
36(5):356–366.
- Watson, G. S. (1964). Smooth regression analysis. *Sankhyā: The
1035 Indian Journal of Statistics, Series A*, pages 359–372.
- Weingartner, E., Burtscher, H., and Baltensperger, U. (1997). Hy-
groscopic properties of carbon and diesel soot particles. *Atmo-
spheric Environment*, 31(15):2311–2327.
- 1040 Wexler, A., Lurmann, F., and Seinfeld, J. (1994). Modelling urban
and regional aerosols—i. model development. *Atmospheric Envi-
ronment*, 28(3):531–546.
- Wilson, J., Cuvelier, C., and Raes, F. (2001). A modeling study of
global mixed aerosol fields. *Journal of Geophysical Research:
Atmospheres (1984–2012)*, 106(D24):34081–34108.
- 1045 Zaveri, R., Easter, R., Fast, J., and Peters, L. (2008). Model for sim-
ulating aerosol interactions and chemistry (MOSAIC). *Journal
of Geophysical Research*, 113:D13204.
- Zaveri, R., Easter, R., and Peters, L. (2005a). A computationally ef-
ficient multicomponent equilibrium solver for aerosols (MESA).
Journal of Geophysical Research, 110(D24):D24203.
- Zaveri, R., Easter, R., and Wexler, A. (2005b). A new method
for multicomponent activity coefficients of electrolytes in aque-
ous atmospheric aerosols. *Journal of Geophysical Research*,
110(D2):D02201.
- Zaveri, R. and Peters, L. (1999). A new lumped structure pho-
tochemical mechanism for large-scale applications. *Journal of
Geophysical Research*, 104(D23):30387–30.
- Zaveri, R. A., Barnard, J. C., Easter, R. C., Riemer, N., and West,
M. (2010). Particle-resolved simulation of aerosol size, compo-
sition, mixing state, and the associated optical and cloud conden-
sation nuclei activation properties in an evolving urban plume.
Journal of Geophysical Research: Atmospheres (1984–2012),
115(D17).
- Zhang, R., Khalizov, A. F., Pagels, J., Zhang, D., Xue, H., and Mc-
Murry, P. H. (2008). Variability in morphology, hygroscopicity,
and optical properties of soot aerosols during atmospheric
processing. *Proceedings of the National Academy of Sciences*,
105(30):10291–10296.
- Zuberi, B., Johnson, K., Aleks, G., Molina, L., Molina, M., and
Laskin, A. (2005). Hydrophilic properties of aged soot. *Geo-
physical Research Letters*, 32(1).

Acknowledgements. This project is funded by NASA. N. Riemer also acknowledges US EPA grant 835042. Its contents are solely the responsibility of the grantee and do not necessarily represent the official views of the US EPA. Further, US EPA does not endorse the purchase of any commercial products or services mentioned in the publication.

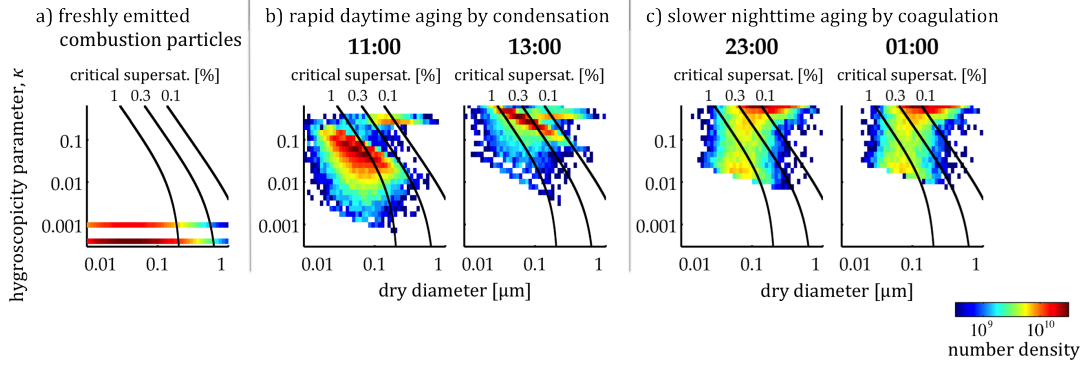


Fig. 1. Two-dimensional probability density distribution shows changes in particle properties. As particles increase in size (horizontal axis) and hygroscopicity (vertical axis), they are able to activate at lower critical supersaturation thresholds (superimposed lines). a) Freshly emitted particles are hydrophobic, with $\kappa = 3 \times 10^{-4}$ and $\kappa = 8 \times 10^{-4}$ for diesel and gasoline, respectively. b) During the daytime, particles age rapidly by condensation of semi-volatile substances that are produced through photochemical reactions. c) At night, condensation aging is slow, and particles age only by coagulation.

Table 1. *Aerosol Emissions and Initial Conditions for Baseline Simulation*

Initial/Background	N (m^{-3})	D_{gm} (μm)	σ_{g}	Composition by Mass
Aitken Mode	9×10^8	0.02	1.45	50% $(\text{NH}_4)_2\text{SO}_4$, 50% SOA
Accumulation Mode	7.5×10^8	0.116	1.65	50% $(\text{NH}_4)_2\text{SO}_4$, 50% SOA
Emissions	\dot{N}_{emit} ($\text{m}^{-2}\text{s}^{-1}$)	$D_{\text{gm,emit}}$ (μm)	σ_{g}	Composition by Mass
Meat cooking	9×10^6	0.0865	1.9	100% POA
Diesel vehicles	3.2×10^7	0.05	1.7	30% POA, 70% BC
Gasoline vehicles	1×10^7	0.05	1.7	80% POA, 20%BC

Table 2. *Gas-phase Initial Conditions and Emissions¹ for Baseline Simulation*

MOSAIC Species	Symbol	Initial Mole Fraction (ppb)	Emissions (nmol m ⁻² s ⁻¹)
Nitric oxide	NO	0.1	15.9
Nitrogen dioxide	NO ₂	1	0.84
Nitric acid	HNO ₃	1.0	
Ozone	O ₃	50	
Hydrogen peroxide	H ₂ O ₂	1.1	
Carbon monoxide	CO	21	291.3
Sulfur dioxide	SO ₂	0.8	7.53
Ammonia	NH ₃	0.5	6.11
Hydrogen chloride	HCl	0.7	
Methane	CH ₄	2200	
Ethane	C ₂ H ₆	1.0	
Formaldehyde	HCHO	1.2	1.68
Methanol	CH ₃ OH	0.12	0.28
Methyl hydrogen peroxide	CH ₃ OOH	0.5	
Acetaldehyde	ALD2	1.0	0.68
Paraffin carbon	PAR	2.0	96
Acetone	AONE	1.0	1.23
Ethene	ETH	0.2	7.3
Terminal olefin carbons	OLET	2.3×10^{-2}	2.42
Internal olefin carbons	OLEI	3.1×10^{-4}	2.42
Toluene	TOL		4.04
Xylene	XYL	0.1	2.41
Lumped organic nitrate	ONIT	0.1	
Peroxyacetyl nitrate	PAN	0.8	
Higher organic acid	RCOOH	0.2	
Higher organic peroxide	ROOH	2.5×10^{-2}	
Isoprene	ISOP	0.5	0.23
Alcohols	ANOL		3.45

Table 3. Hygroscopicity parameter assigned to aerosol species

aerosol species	K_i	citation
NO ₃	0.65	Clegg et al. (1998); Svenningsson et al. (2006); Petters and Kreidenweis (2007)
SO ₄	0.65	Clegg et al. (1998); Svenningsson et al. (2006); Petters and Kreidenweis (2007)
NH ₄	0.65	Clegg et al. (1998); Svenningsson et al. (2006); Petters and Kreidenweis (2007)
SOA	0.1	Prenni et al. (2007)
BC	0	Petters et al. (2006)
<i>POA</i>	0.001	Petters et al. (2006)

Table 4. *Input parameters varied in the ensemble of sensitivity simulations. Scenarios corresponding to the baseline conditions are indicated in bold. All combinations of scenarios were included in the full ensemble of 288 simulations.*

input parameter	scenario name	values assigned
gas-phase emissions	G₁	baseline gas emissions
	G ₂	increase SO ₂ emissions by 200% and turn of NH ₃ emissions
	G ₃	increase SO ₂ emissions by 300%
	G ₄	increase SOA precursor emissions by 200% and turn off NH ₃ emissions increase selected SOA-precursors by 100%
black carbon emissions	BC₁	baseline diesel and gasoline emissions
	BC ₂	increase diesel and gasoline emissions by a factor of five
	BC ₃	increase diesel and gasoline emissions by a factor of ten
background aerosol	BG₁	baseline background aerosol concentration
	BG ₂	decrease background aerosol concentration by 45%
relative humidity	RH₁	95%
	RH ₂	50%
meteorological conditions	M₁	5° N, July 19, 303 K
	M ₂	5° N, January 5, 303 K
	M ₃	40° N, July 19, 298 K
	M ₄	40° N, January 5, 275 K
	M ₅	60° N, July 19, 293 K
	M ₆	60° N, January 5, 263 K

Table 5. *Candidate variables included in the regression analysis.*

symbol	description	type
D	wet diameter	particle-level
D_{dry}	dry diameter	particle-level
κ	hygroscopicity parameter	particle-level
ε_{BC}	mass fraction BC	particle-level
ε_{SOA}	mass fraction SOA	particle-level
$\varepsilon_{\text{inorg}}$	mass fraction inorganic aerosol	particle-level
N	overall aerosol number concentration	population-level
A_{wet}	aerosol surface area concentration	population-level
\dot{V}_{cond}	secondary aerosol formation rate, volume	population-level
\dot{f}_{cond}	secondary aerosol flux, $\dot{V}_{\text{cond}}/A_{\text{wet}}$	population-level
\dot{m}_{cond}	secondary aerosol formation rate, mass	population-level
κ_{cond}	hygroscopicity parameter of secondary aerosol	population-level
N_{CCN}	number concentration of CCN-active particles	population-level
N_{large}	num. conc. of particles with $D > 100$ nm	population-level
$N_{\text{CCN,large}}$	num. conc. of CCN-active particles with $D > 100$ nm	population-level

Table 6. Bulk aging timescale for two fresh particle size distributions under different aging regimes. Condensation, coagulation, and overall aging timescales are given for $s = 0.3\%$. We assumed lognormal size distributions of fresh BC with a geometric standard deviation of 1.7.

input variables			expected aging timescale		
D_{gm} [nm]	\dot{f}_{cond} [nm h ⁻¹]	$N_{\text{CCN,large}}$ [cm ⁻³]	$s = 0.3\%$		
			$\hat{\tau}_{\text{cond}}$ [h]	$\hat{\tau}_{\text{coag}}$ [h]	$\hat{\tau}_{\text{age}}$ [h]
30	0.01	500	460	17	16
60	0.01	500	44	40	20
30	1	500	18	17	9
60	1	500	2.8	40	2.6
30	0.01	10,000	460	5	5
60	0.01	10,000	44	6	5.4
30	1	10,000	18	2.6	2.2
60	1	10,000	2.8	6	2

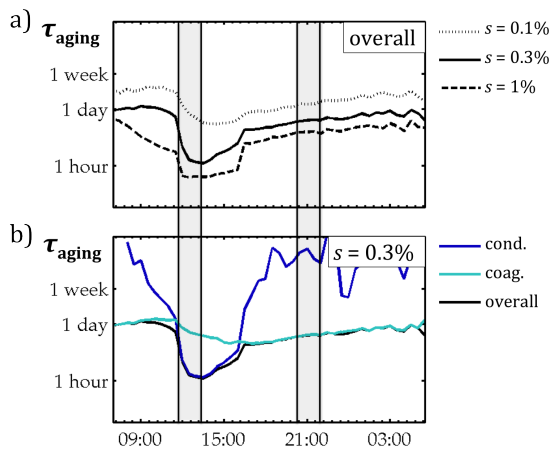


Fig. 2. For a single scenario, overall aging timescale for $s = 0.1\%$, $s = 0.3\%$, and $s = 1\%$ in Figure 2.a and the overall, condensation, and coagulation aging timescales for $s = 0.3\%$ in Figure 2.b. The shaded regions show how the value of the aging timescale reflects changes in per-particle characteristics, which correspond to Figure 1. Short aging timescales correspond to rapid increases in particle size and hygroscopicity (Figure 1.b), and long aging timescales correspond to slow changes in particle properties (Figure 1.c)

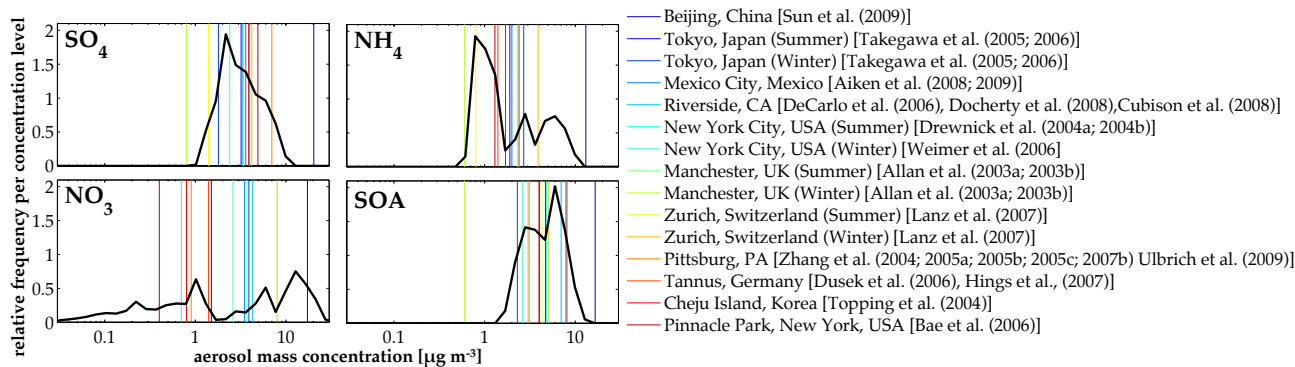


Fig. 3. Probability density function of aerosol mass species in simulations (black line in each graph) show that model cases represent variation in atmospheric conditions from ambient observations (vertical colored lines). Probability density functions include all output time steps in the full ensemble of sensitivity simulations.

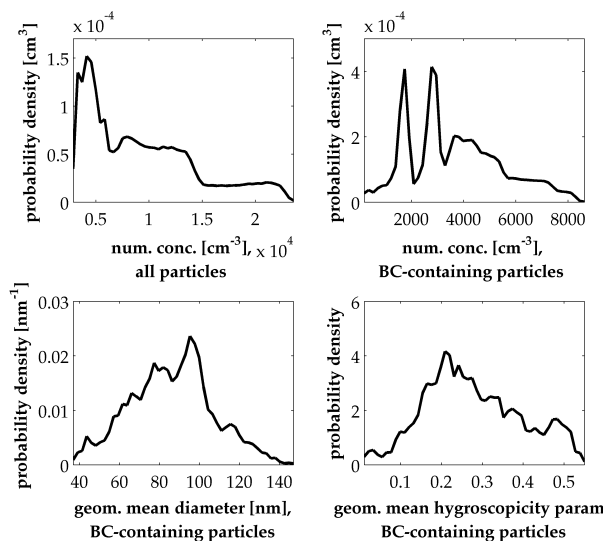


Fig. 4. Probability density function of a) total aerosol number concentration, b) total number concentration of BC-containing particles, c) geometric mean diameter of BC-containing particles, and d) geometric mean hygroscopicity parameter of BC-containing particles.

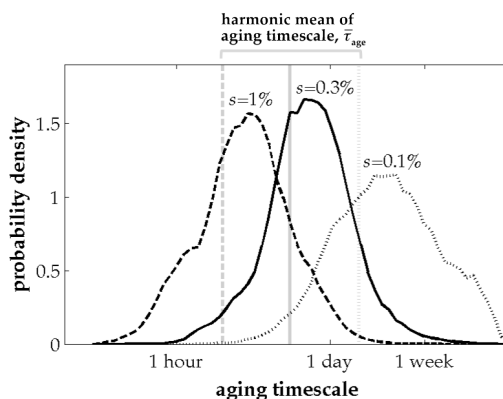


Fig. 5. Probability density function of aging timescales for the full ensemble of sensitivity simulations, computed at three environmental supersaturation levels: $s = 0.1\%$, $s = 0.3\%$, and $s = 1\%$.

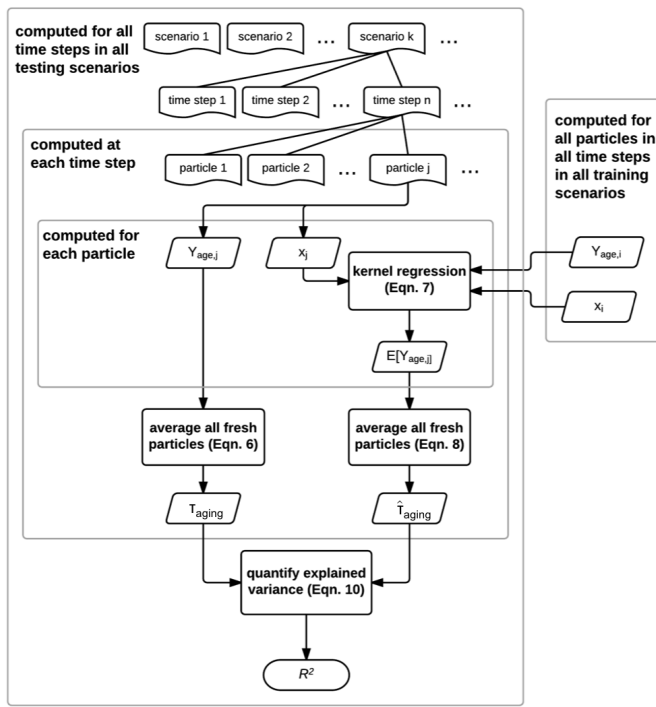


Fig. 6. Procedure for applying kernel regression to predict black carbon's aging timescale and quantifying the portion of variance explained by that prediction, shown for a hypothetical input variable x .

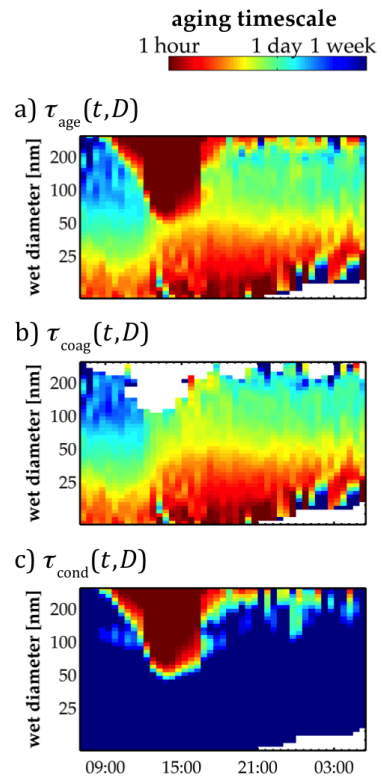


Fig. 7. For baseline scenario, a) overall size-dependent aging timescale, b) condensation aging timescale, and c) coagulation aging timescale. Values are shown for $s = 0.3\%$.

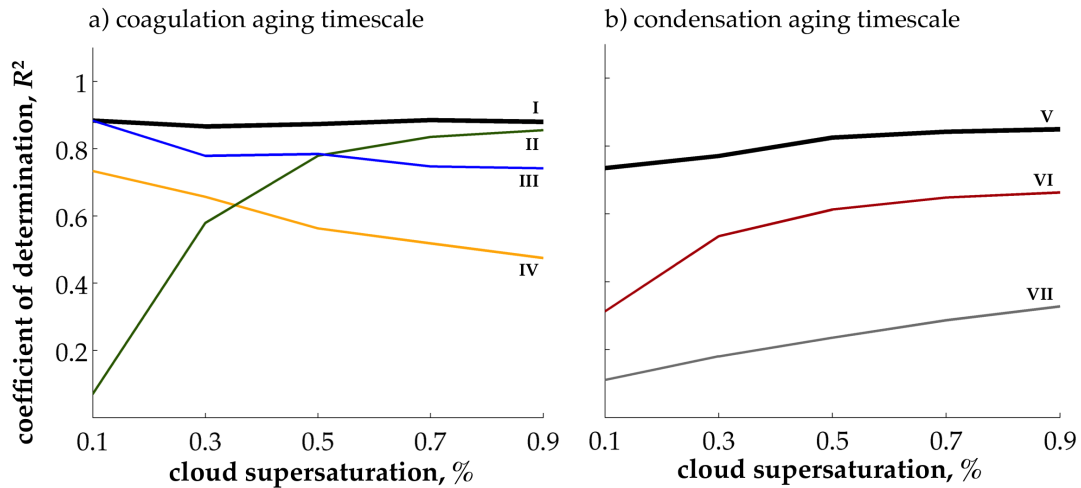


Fig. 8. Coefficient of determination R^2 for a) coagulation and b) condensation timescales as a function of supersaturation for selected combinations of independent variables, where the combination of variables that explain most of the variance are shown by the black lines in each graph. Regression analyses on the coagulation aging timescales are shown for four combinations of independent variables: (I) including wet diameter, D , of fresh BC-containing particles and the number concentration of large ($D > 100$ nm), CCN-active particles, $N_{CCN,large}$, (II) including D of fresh BC-containing particles and the number concentration of large particles, N_{large} , rather than $N_{CCN,large}$, (III) including D of fresh particles and the number concentration of CCN-active particles, N_{CCN} , rather than $N_{CCN,large}$, and (IV) including $N_{CCN,large}$ but without including D of fresh BC-containing particles. Regression analyses on the condensation aging timescale are shown for three combinations of independent variables: (V) including secondary aerosol flux, \dot{f}_{cond} , the hygroscopicity of secondary aerosol, κ_{cond} , and D of fresh BC-containing particles, (VI) including \dot{f}_{cond} and κ_{cond} but without including D of fresh BC-containing particles, and (VII) including \dot{f}_{cond} and D of fresh BC-containing particles but without including κ_{cond} . For all s , approximately 90% of variance in coagulation aging timescale is explained by two independent variables (black line in Figure 8.a), and 80% of variance in condensation aging timescale is explained by three independent (black line in Figure 8.b).

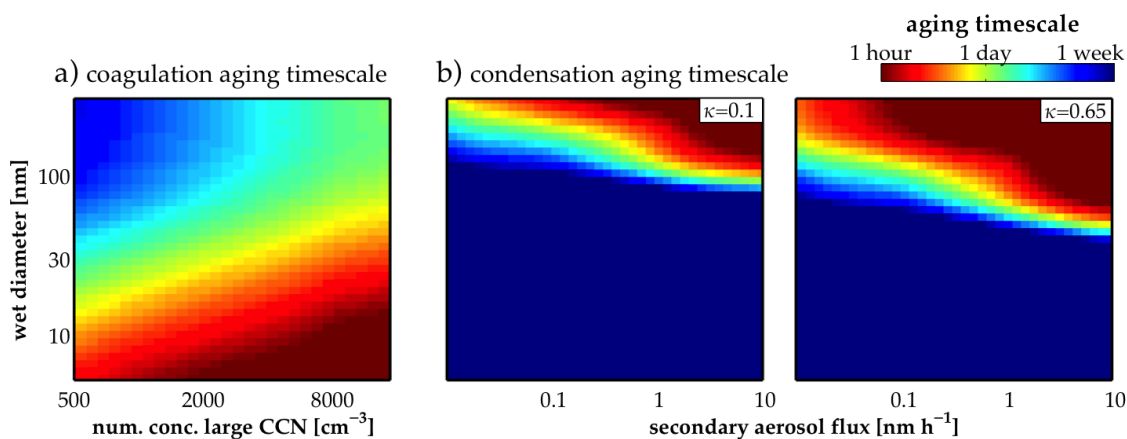


Fig. 9. Coagulation aging timescale as a function of wet diameter and number of large, CCN-active particles (Figure 9.a) and condensation aging timescale as a function of wet diameter, secondary aerosol flux, and hygroscopicity of secondary aerosol (Figure 9.b). Results are shown for a threshold supersaturation s of 0.3%.

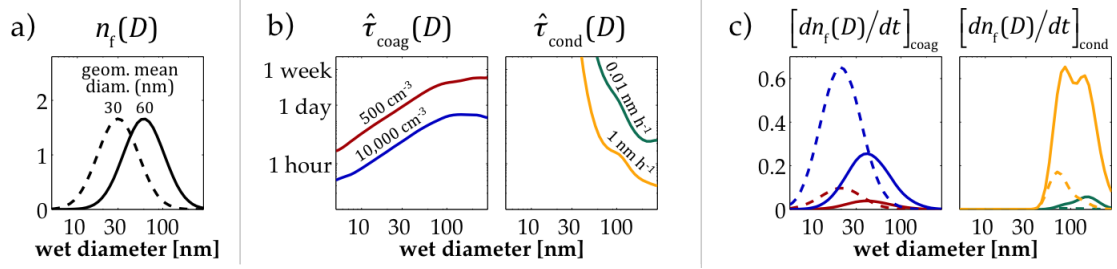


Fig. 10. Rate at which particles of specific size transition from fresh to aged (Figure 10.c) depends on size distribution of fresh BC (Figure 10.a) and size-resolved aging timescale (Figure 10.b). Results are shown at $s = 0.3\%$, where the size-resolved aging timescale under different conditions are determined from the regression function in Figure 9. The line colors in Figure 10.c correspond to the aging conditions shown in Figure 10.b, and the line style in Figure 10.c correspond to the size distributions shown in Figure 10.a.

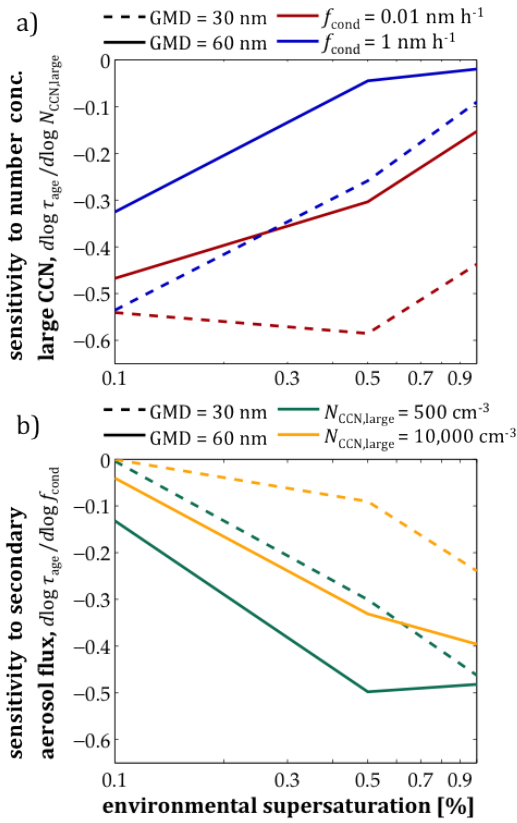


Fig. 11. Sensitivity of aging timescale to a) $N_{CCN,large}$ and b) f_{cond} as a function of supersaturation level, expressed as the logarithmic derivative of the timescale with respect to each variable. The value of $d\log \tau_{age}/d\log N_{CCN,large}$, for example, indicates the relative change in τ_{age} to a relative change in $N_{CCN,large}$. Shown for the size distributions and aging conditions in Figure 10.